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**Physiological responses of coastal and oceanic diatoms to diurnal fluctuations in seawater
carbonate chemistry under two CO₂ concentrations**

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Running head: ocean acidification influences diatoms under fluctuating pH

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19 **Abstract**

20 Diel or seasonal fluctuations in seawater carbonate chemistry are common in coastal waters, while in
21 the open ocean carbonate chemistry is much less variable. In both of these environments, ongoing ocean
22 acidification is being superimposed on the natural carbonate buffer system to influence the physiology
23 of phytoplankton. Here, we show that a coastal *Thalassiosira weissflogii* isolate and an oceanic diatom,
24 *Thalassiosira oceanica*, respond differentially to diurnal fluctuating carbonate chemistry in current and
25 ocean acidification (OA) scenarios. A fluctuating carbonate chemistry regime showed positive or
26 negligible effects on physiological performance of the coastal species. In contrast, the oceanic species
27 was significantly negatively affected, with higher respiration than cells grown under the corresponding
28 steady regime. The fluctuating regime reduced photosynthetic oxygen evolution rates of *T. oceanica*
29 under ambient CO₂ concentration, while in the OA scenario, the fluctuating regime depressed its growth
30 rate, chlorophyll *a* content, and elemental production rates. These contrasting physiological
31 performances of coastal and oceanic diatoms indicate that they differ in the ability to cope with dynamic
32 *p*CO₂. We propose that, in addition to the ability to cope with light, nutrient, and predation pressure, the
33 ability to acclimate to dynamic carbonate chemistry may act as one determinant of the spatial
34 distribution of diatom species. Habitat-relevant diurnal changes in seawater carbonate chemistry can
35 interact with OA to differentially affect diatoms in coastal and pelagic waters.

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41 **Key words:** diatom, growth, photosynthesis, elemental production rates, fluctuating carbonate
42 chemistry, CO₂

43

44 **1 Introduction**

45 Anthropogenic emissions of carbon dioxide (CO₂) since the industrial revolution have increased
46 atmospheric *p*CO₂ levels by 40% (Howes et al. 2015), mainly due to burning of fossil fuels and land use
47 changes (Ciais et al. 2014). The oceans absorb about 30% of the CO₂ emitted by human activities
48 (Sabine et al. 2004), leading to decreases in pH, concentration of carbonate ions, and saturation state of
49 calcium carbonate, along with increases of the concentrations of aqueous CO₂ and bicarbonate (i.e.,
50 ocean acidification). The global surface ocean mean pH has already decreased by about 0.1 units since
51 the industrial revolution (Orr et al. 2005; Doney 2010), and a further decrease of 0.3-0.4 units is
52 expected to happen by 2100 under the business as usual scenario (Orr et al. 2005; Gattuso et al. 2015).

53 For marine organisms, the reduced seawater mean pH caused by OA could be detectable on a
54 timescale of years to decades, while striking fluctuations in coastal seawater carbonate chemistry may
55 occur over much shorter timescales. The coastal zone plays a critical role in biogeochemical cycles, and
56 experiences great variability of physical and chemical factors (Drupp et al. 2011). In addition, it is the
57 area most impacted by anthropogenic pressures (Gattuso et al. 1998). Carbonate chemistry in coastal
58 seawater is affected by multiple drivers in addition to atmospheric CO₂ dissolution, such as tidal cycles
59 (Dai et al. 2009; Jiang et al. 2011; Wang et al. 2014), upwelling (Feely et al. 2008; Capone and Hutchins
60 2013), watershed processes, wind forcing (Drupp et al. 2011), anthropogenic nutrient inputs,



61 aquaculture activities, and changes in ecosystem structure and metabolism (Duarte et al. 2013;
62 Waldbusser and Salisbury 2014). Due to high biomass and sufficient or excess nutrients in coastal
63 waters, biological activities alter $p\text{CO}_2$, resulting in a diel cycle of pH. The diel range of pH variation in
64 some coastal ecosystems can be greater than 1 pH unit (Hinga 2002), which corresponds to a 900%
65 change in H^+ concentration.

66 During a diurnal cycle, organisms in coastal areas could experience pH values that may be lower than
67 the projected value for the surface ocean in the year 2100 (Hofmann et al. 2011; Hurd et al. 2011;
68 Waldbusser and Salisbury 2014). In contrast, pH in the open ocean is relatively stable, with a variation
69 range of only ~ 0.024 over a month (Hofmann et al. 2011). The buffering capacity will decrease as the
70 increase of dissolved inorganic carbon in both coastal and oceanic seawaters (Eggleston et al. 2010; Cai
71 et al. 2011; Denman et al. 2011; Wang et al. 2013), while the variation range of pH in coastal water may
72 be amplified, due to the multiple drivers mentioned above. Diurnal and seasonal variations in pH caused
73 by photosynthesis and respiration could be increased by more than 40% relative to the present extent of
74 variation (Eggleston et al. 2010).

75 Responses of fish (Dixon 2014), gastropods (Onitsuka et al. 2014), oysters (Keppel 2015), mussels
76 (Frieder et al. 2014), coral (Dufault et al. 2012; Comeau et al. 2014), canopy-forming kelp (Britton et al.
77 2016), and coralline algae (Gao et al. 1993; Cornwall et al. 2013; Noisette et al. 2013; Johnson et al.
78 2014) to diurnally fluctuating $p\text{CO}_2$ /pH have been studied recently. Dufault et al. (2012) hypothesized
79 that storage of dissolved inorganic carbon during the night-time high $p\text{CO}_2$ period fueled day-time
80 calcification (and perhaps photosynthesis), resulting in higher calcification and survival rate of coral
81 recruits. Thus, it appears that some marine organisms may benefit from $p\text{CO}_2$ fluctuations. In spite of



82 this body of literature, the responses of marine phytoplankton to fluctuating pH/ $p\text{CO}_2$ are still unclear.
83 To our knowledge, only one study has addressed the responses of the marine green alga *Ostreococcus* to
84 fluctuating $p\text{CO}_2$ (Schaum et al. 2016). However, how CO_2 variability affects other major marine
85 phytoplankton groups over either the short- or long-term remains unknown.

86 Coastal and open ocean species are distinguished by habitat-related differences in cell size, nutrient
87 utilization (Glibert and Ray 1990), photosynthetic architecture (Strzepek and Harrison 2004), and
88 photosynthetic performance (Lavaud et al. 2007; Li et al. 2011; Liu and Qiu 2012). Our study was
89 intended to understand whether coastal and oceanic species also differ in their capacity to respond to
90 fluctuating carbonate chemistry. A coastal *Thalassiosira weissflogii* isolate and an oceanic diatom,
91 *Thalassiosira oceanica*, were used in the present study. We manipulated $p\text{CO}_2$ to mimic diurnally
92 fluctuating carbonate chemistry and hypothesized that coastal diatoms would show better physiological
93 performance under fluctuating carbonate chemistry than oceanic ones, a difference that could
94 potentially be a key factor influencing the geographical distribution of diatoms.

95

96 **2 Materials and methods**

97 **2.1 Cultures and experimental setup**

98 *Thalassiosira weissflogii* (CCMP 1336, isolated from coastal Long Island, New York, USA in 1956)
99 and *Thalassiosira oceanica* (CCMP 1005, isolated from the Sargasso Sea in 1958) were incubated in
100 Aquil medium (Sunda et al. 2005), illuminated by cool white fluorescent light at an intensity of 115
101 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$. Cultures were maintained at 20 °C with a 12 h:12 h light and dark cycle. Cells
102 were maintained at exponential growth phase with maximal concentration $< 1.1 \times 10^4 \text{ mL}^{-1}$ (T .



103 *weissflogii*) or 3.5×10^4 mL⁻¹ (*T. oceanica*) in semi-continuous cultures.

104 *T. weissflogii* and *T. oceanica* were acclimated to four treatments: 1) steady carbonate chemistry at
105 ambient $p\text{CO}_2$ level (LCs); 2) diurnally carbonate chemistry fluctuated around ambient $p\text{CO}_2$ level
106 (LCf); 3) steady carbonate chemistry at elevated $p\text{CO}_2$ level (HCs); and 4) diurnally carbonate
107 chemistry fluctuated around elevated $p\text{CO}_2$ level (HCf) for 15 generations before sampling. Steady
108 regimes were bubbled with ambient air ($400 \pm 15 \mu\text{atm}$, LCs) or elevated ($1005 \pm 40 \mu\text{atm}$, HCs) $p\text{CO}_2$,
109 which was automatically achieved by mixing air/ CO_2 with a CO_2 Enricher (CE100B, RuiHua). The
110 fluctuating regimes were obtained by changing the CO_2 partial pressure every 12 h. Cells were aerated
111 with air of low $p\text{CO}_2$ (i.e., 0 or $557 \pm 15 \mu\text{atm}$ for LCF and HCF, respectively) during the photoperiod;
112 the aeration was changed to high $p\text{CO}_2$ (i.e., 870 ± 19 or $1949 \pm 35 \mu\text{atm}$ for LCF and HCF, respectively)
113 at the beginning of the dark period. Measurements showed that pH gradually increased and decreased,
114 similar to a natural diurnal cycle (see Results). Since pH increased quickly in the first few hours of the
115 photoperiod, the aeration rates were adjusted to make sure the fluctuating regimes reached similar pH
116 values with corresponding steady regimes in the middle of photoperiod and reached target values at the
117 end of photoperiod. The steady regimes were aerated with stable $p\text{CO}_2$ air at the same flow rate as the
118 fluctuating regimes. The pH was measured every 1.5 h by a pH meter (Orion 2 STAR, Thermo
119 Scientific) calibrated with standard National Bureau of Standards (NBS) buffers. Subsamples for
120 measurement of physiological parameters were always taken in the middle of the photoperiod, unless
121 otherwise noted.

122

123 **2.2 Growth rate and chlorophyll *a* content**



124 Cell concentration and mean cell size were measured by a Coulter Particle Count and Size Analyzer
125 (Z2, Beckman Coulter). Specific growth rate was calculated according the equation:

126 $\mu = (\ln N_1 - \ln N_0) / (t_1 - t_0)$, in which N_1 and N_0 represent cell concentrations at t_1 and t_0 . For the
127 chlorophyll *a* content determination, samples were filtered onto GF/F filters (25 mm, Whatman), and
128 extracted overnight at 4 °C in absolute methanol before centrifugation. The supernatants were analyzed
129 by a UV-VIS Spectrophotometer (DU800, Beckman Coulter) and the chlorophyll *a* content was
130 calculated according to the equation of Ritchie (2006).

131

132 **2.3 Elemental composition and production rate**

133 Samples for measuring particulate organic carbon (POC) and nitrogen (PON) were filtered onto pre-
134 combusted (450 °C for 6 h) GF/F filters (25 mm, Whatman). Filters were treated using HCl fumes to
135 remove any inorganic carbon and dried before analysis on a CHNS/O Analyzer (2400SeriesII,
136 PerkinElmer). 25 mm polycarbonate filters were used to determine biogenic silica (BSi) by the
137 spectrophotometric method of Brzezinski and Nelson (1995). Production rates of POC, PON, and BSi
138 were calculated by multiplying cellular content by specific growth rate.

139

140 **2.4 Chlorophyll *a* fluorescence**

141 The photochemical parameters were determined using a Xenon-Pulse Amplitude Modulated
142 fluorometer (Xe-PAM, Walz). Effective photochemical quantum yields were determined according to
143 the equation of Genty et al. (1989): $\Phi_{PSII} = (F'_m - F_t) / F'_m = \Delta F / F'_m$ for light-adapted samples, where
144 F'_m indicates maximum chlorophyll fluorescence of light-adapted samples, and F_t , steady chlorophyll



145 fluorescence of light-adapted samples. Non-photochemical quenching (NPQ) was calculated as:
146 $NPQ = (F_m - F'_m) / F'_m$, where F_m indicates maximum chlorophyll fluorescence of dark-adapted samples.
147 Φ_{PSII} and NPQ were measured under actinic light intensity similar to culture light level after 10 min
148 dark adaptation. Given the changing carbonate chemistry over a diurnal cycle, Φ_{PSII} and NPQ were
149 determined at three time points: 0.5, 6, and 11.5 h after illumination. NPQ versus irradiance curves were
150 determined by rapid light curves (RLCs) with 15 s duration for each light level. Although the values of
151 NPQ derived from RLCs were not as accurate as values from fluorescence induction curves, they
152 provide estimates of the kinetics of NPQ development with increasing light intensity.

153

154 **2.5 Photosynthetic oxygen evolution and dark respiration rates**

155 Net photosynthetic oxygen evolution and dark respiration rates were determined using a Clark-type
156 oxygen electrode (Oxygraph, Hansatech) at the experimental temperature. Oxygen evolution rates were
157 measured under $115 \mu\text{mol photons m}^{-2} \text{s}^{-1}$ at the same three time points as mentioned above. Oxygen
158 consumption rates were measured in the middle of photoperiod, when the steady and fluctuating
159 regimes reached similar pH values. Samples were gently filtered onto 47 mm cellulose acetate
160 membranes, and then re-suspended into 20 mmol L^{-1} Tris buffered medium. The re-suspended cells
161 were injected into an oxygen electrode chamber equipped with a magnetic stirrer. Rates of oxygen
162 evolution and consumption were derived from the linear portion of the slope of the oxygen record. The
163 pH values of Tris buffered medium were pre-adjusted to their corresponding culture medium values.
164 That is, pH values of the three time points in the LCf treatment were 7.84, 8.14, and 8.35, and those in
165 the HCF treatment were 7.54, 7.80, and 8.06. Values in the LCs and HCs treatments were set to 8.14 or



166 7.80 for all three time points, respectively.

167

168 **2.6 Statistical analyses**

169 Significant differences among treatments were tested using one-way analysis of variance (ANOVA)
170 with a significance level of $p < 0.05$. When necessary, the post hoc Duncan test was used to determine
171 the differences between treatments. All data are reported as mean value of triplicate samples \pm standard
172 deviation (SD).

173

174 **3 Results**

175 **3.1 Variation of pH in experimental regimes**

176 The variation ranges of pH in the LCf and HCf treatments were 0.52 ± 0.03 , and 0.53 ± 0.03 ,
177 respectively. For clarity, only mean pH values every 1.5 h are shown (Fig. 1). At the beginning of
178 photoperiod, pH of the LCf regime was 7.84 ± 0.02 , and then it increased to 8.15 ± 0.03 in the middle of
179 photoperiod, similar to the value of the LCs regime (8.13 ± 0.02). The pH value of the LCf regime
180 reached 8.35 ± 0.02 at the end of the photoperiod, and then decreased to 7.84 ± 0.02 . For the HCf
181 regime, pH ranged from 7.54 ± 0.01 to 8.06 ± 0.02 , and reached 7.82 ± 0.01 in the middle of
182 photoperiod, similar to the value of the HCs regime (7.79 ± 0.01).

183

184 **3.2 Specific growth rate and mean cell size**

185 Growth rates of *T. weissflogii* were not influenced by diurnally fluctuating carbonate chemistry in
186 either the current or the OA scenario (Fig. 2a). Likewise, there were no differences in growth rates



187 between steady and fluctuating regimes for *T. oceanica* under the ambient $p\text{CO}_2$ condition (Fig. 2b).
188 However, fluctuating regime reduced its growth rate by 9% under the elevated $p\text{CO}_2$ condition. OA
189 influenced the growth rate of *T. oceanica*, with rate of HCs cells being 16% lower than LCs cells. No
190 effects of OA on growth rate of *T. weissflogii* were detected. Additionally, growth rates of *T.*
191 *pseudonana* (CCMP 1335, isolated from Moriches Bay, New York, USA in 1958) were not influenced
192 by the fluctuating regime under both ambient and elevated $p\text{CO}_2$ conditions (data not shown).

193 Mean cell sizes were not affected by the fluctuating treatment under either ambient or elevated $p\text{CO}_2$
194 conditions in *T. weissflogii* (Table 1). *T. oceanica* cells showed minor but significant changes in cell size
195 in the fluctuating treatments. Cells in the LCf treatment cells were 1.2% larger than LCs cells, while
196 HCf cells were 1.4% smaller than cells in the corresponding steady treatments.

197

198 **3.3 Chlorophyll *a* content and elemental composition**

199 Chlorophyll *a* contents of *T. weissflogii* in the four treatments were not significantly different. For *T.*
200 *oceanica*, the fluctuating regime didn't influence chlorophyll *a* content under ambient CO_2 level.
201 However, in the HCf treatment chlorophyll *a* content decreased by 24% compared to the steady regime
202 (Table 1).

203 POC and PON quotas of both species were elevated in the OA scenario in both the steady and
204 fluctuating regimes, relative to present day $p\text{CO}_2$ levels (Table 1). However, no effects of fluctuating
205 regime on cellular POC and PON contents were detected in either species compared to the steady
206 treatments. The only exception was that POC increased by 9% in the LCf treatment relative to the LCs
207 treatment for *T. weissflogii*. Generally, elevated $p\text{CO}_2$ and fluctuating regime showed no effects on BSi



208 quota of either species, besides a slight decrease in the HCf treatment relative to that of the HCs
209 treatment for *T. weissflogii*.

210 The fluctuating regime increased the POC production rate of *T. weissflogii* at both ambient and
211 elevated $p\text{CO}_2$ levels, but had no effects on other elemental production rates of this species. By contrast,
212 the fluctuating regime decreased all of the elemental production rates in the OA scenario for *T. oceanica*
213 (Fig. 3). The C:N and Si:C ratios of *T. oceanica* and the Si:C ratio of *T. weissflogii* were lower in the
214 OA scenario, while C:N ratios of *T. weissflogii* were not significant different in the four treatments
215 (Table 1). Slight but significant decreases of the Si:C ratio in the fluctuating regime compared to the
216 steady regime were found at ambient $p\text{CO}_2$ for both species.

217

218 **3.4 Chlorophyll *a* fluorescence**

219 The effective photochemical quantum yields of both species varied little at different time points,
220 ranging from 0.54 ± 0.03 to 0.61 ± 0.03 among treatments (Table 2). Fluctuating regimes scarcely
221 influenced Φ_{PSII} of either species. The only exception was that Φ_{PSII} of HCf decreased by 8% relative to
222 that of the HCs for *T. weissflogii* at the beginning of the photoperiod. Elevated $p\text{CO}_2$ decreased Φ_{PSII} by
223 3% and 5% in the middle of the photoperiod for *T. weissflogii* and *T. oceanica*, respectively. Cells under
224 elevated $p\text{CO}_2$ showed 2% and 7% lower Φ_{PSII} compared to those under ambient $p\text{CO}_2$ 11.5 h after
225 illumination for *T. weissflogii* and *T. oceanica*, respectively. NPQ under culture light intensity ranged
226 from 0.06 ± 0.01 to 0.23 ± 0.05 at different time points. No detectable effects of fluctuating regime on
227 NPQ of either species were found, with exceptions of HCf cells of *T. weissflogii* at the beginning of the
228 photoperiod and LCf cells of *T. oceanica* 11.5 h after illumination. For steady regimes, elevated $p\text{CO}_2$



229 showed no detectable effect on NPQ of both species at the beginning of the photoperiod, while it
230 increased NPQ by 37.5% and 38.4% relative to values of LCs cells in the middle of the photoperiod for
231 *T. weissflogii* and *T. oceanica*, respectively. Values of NPQ of HCs cells were decreased by 25% and
232 33.3% relative to values of LCs cells at the end of the photoperiod for *T. weissflogii* and *T. oceanica*,
233 respectively.

234

235 **3.5 Photosynthetic oxygen evolution and dark respiration rates**

236 Chlorophyll normalized net oxygen evolution rates of these two species ranged from 0.39 ± 0.07 to
237 $0.55 \pm 0.07 \mu\text{mol O}_2 \mu\text{g chl } a^{-1} \text{ h}^{-1}$ in the middle of photoperiod. Neither elevated $p\text{CO}_2$ nor fluctuating
238 regime showed detectable effects on oxygen evolution rates per chlorophyll of *T. weissflogii* (Fig. 4a),
239 while *T. oceanica* cells under the LCf treatment had a 29% decrease of chlorophyll-normalized net
240 oxygen evolution rate relative to the LCs cells (Fig. 4b).

241 Both species, regardless of treatment, showed a similar diurnal rhythm of photosynthetic oxygen
242 evolution: oxygen evolution rates reached the highest values in the middle of the photoperiod (Fig. 4c,
243 d). For *T. weissflogii*, effects of fluctuating $p\text{CO}_2$ on net oxygen evolution per cell were only observed in
244 the middle of the photoperiod for cells at the ambient $p\text{CO}_2$ level, with 7% lower rates in the LCf
245 treatment than in the LCs (Fig. 4c). These effects were more obvious for *T. oceanica* cells at ambient
246 $p\text{CO}_2$ level. At 11.5 h after illumination, LCs cells of *T. oceanica* showed 41% higher net oxygen
247 evolution rates per cell than LCf cells (Fig. 4d). *T. oceanica* cells in the steady regime under elevated
248 $p\text{CO}_2$ evolved oxygen at 65% higher cell-specific rates than those in the fluctuating regime at the
249 beginning of the photoperiod.



250 Elevated $p\text{CO}_2$ increased dark respiration of *T. weissflogii* by 59% compared to that at ambient $p\text{CO}_2$
251 level, while fluctuating regime had no detectable effect (Fig. 5a). In contrast, dark respiration rates of *T.*
252 *oceanica* were stimulated by 44% and 55% for cells under the fluctuating regime compared to steady
253 one at ambient and elevated $p\text{CO}_2$ levels, respectively (Fig. 5b), while no effects of elevated $p\text{CO}_2$ were
254 observed. When dark respiration rates were normalized per cell, they generally showed the same
255 patterns as chlorophyll normalized rates, with different amplitudes of variation (Fig. 5c, d). The
256 exception was that no effects of fluctuating regime on dark respiration per cell of *T. oceanica* were
257 found in the OA scenario. The respiration to photosynthesis (R:P) ratios for *T. weissflogii* under elevated
258 $p\text{CO}_2$ was higher than at ambient $p\text{CO}_2$ by 73%, while no effects of fluctuating regime were detected.
259 R:P ratios for *T. oceanica* cells was higher by 104% in the fluctuating regime than for cells in the
260 corresponding steady regime at ambient $p\text{CO}_2$ level (Table 1).

261

262

263 3.6 NPQ versus irradiance curves

264 The development kinetics of NPQ with increasing light intensity differed in the two diatoms (Fig. 6).
265 No effects of fluctuating regimes on NPQ were found in either species, so for clarity, only steady
266 regimes are shown. *T. weissflogii* had higher NPQ values than *T. oceanica* above a light intensity of 330
267 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$, and its maximal extent of NPQ under the highest light of RLCs was 6-7 times
268 higher than that of *T. oceanica*.

269

270 4 Discussion



271 Both species were influenced by elevated $p\text{CO}_2$ in several ways, while they responded differentially
272 to fluctuating regime. In general, the fluctuating $p\text{CO}_2$ regime showed either positive (POC cellular
273 quota and production rate) or no obvious effects on physiological performance of the coastal diatom *T.*
274 *weissflogii*. In contrast, the oceanic diatom *T. oceanica* was significantly negatively affected by the
275 diurnal variation of carbonate chemistry, with higher dark respiration under the fluctuating regime than
276 under the steady regime. The fluctuating regime reduced photosynthetic oxygen evolution rates of *T.*
277 *oceanica* under ambient $p\text{CO}_2$ concentration, while in the OA scenario, the fluctuating regime depressed
278 its growth rate, chlorophyll *a* content, and elemental production rates.

279 OA depressed the growth of *T. oceanica*, consistent with results of a previous study (King et al.
280 2015), which showed a similar decrease (19%) to the present study (16%). No detectable effects of OA
281 on growth of *T. weissflogii* were found, as reported by previous studies (Burkhardt et al. 1999; Shi et al.
282 2009; Reinfelder 2012; King et al. 2015; Passow and Laws 2015; Taucher et al. 2015). However, the
283 growth responses of diatoms have also been shown to be affected by interactions between OA and other
284 abiotic factors. For instance, the energy saved from active inorganic carbon acquisition mechanisms due
285 to increased availability of CO_2 under OA conditions enhanced the growth of diatoms when daytime
286 mean light level was lower than 22-36% of sea surface solar light intensity. However, growth under OA
287 condition decreased when light exceeded 25-42% of incident irradiance (Gao et al. 2012). OA reduced
288 the growth rate of *T. weissflogii* under light and temperature stress, but no effects of OA were detected
289 in the absence of temperature stress (Passow and Laws 2015). Consequently, it appears that effects of
290 OA on phytoplankton species could be region-specific, depending on the local interactions with other
291 abiotic factors.



292 The silicified cell walls of diatoms act as mechanical protection to resist grazers (Hamm et al. 2003),
293 and also have potential roles in photoprotection (Raven and Waite 2004), as well as promotion of
294 catalysis by extracellular carbonic anhydrase (Milligan and Morel 2002). Si:C ratio of both species
295 decreased under the elevated $p\text{CO}_2$ condition, in accordance with results of Tatters et al. (2012) and
296 Mejia et al. (2013). This decreased ratio indicates that diatoms may have reduced silicon requirements
297 per carbon fixed under an OA scenario than under the ambient $p\text{CO}_2$ condition, and so has implications
298 for changes in local and global carbon and silicon budgets. In Si-replete regions, a lower ratio may
299 reduce the ballasting function of silica in carbon export by diatoms (Mejia et al. 2013). In Si-limited
300 waters, a consequence may be that the proportion of diatoms in phytoplankton communities may
301 increase due to reduced Si requirements (Mejia et al. 2013). However, diatom silicification is under a
302 complex set of controls. For instance, limitation by other nutrients such as, iron (Hutchins and Bruland
303 1998) and nitrogen (Flynn and Martin-Jézéquel 2000), may act to increase Si quotas and Si:C ratio.

304 Bicarbonate utilization has been suggested to be a general characteristic of marine diatoms, through
305 direct transport or conversion by extracellular carbonic anhydrase (eCA), while the fraction of direct
306 bicarbonate transport and eCA expression varies among species (Martin and Tortell 2008). Pathways
307 that can utilize HCO_3^- and provide CO_2 for Rubisco through C_4 (Reinfelder et al. 2000) or $\text{C}_3\text{-C}_4$
308 intermediate (Roberts et al. 2007) photosynthesis have been suggested for *T. weissflogii*. This species
309 takes up both CO_2 and HCO_3^- at a similar rate, and has the ability to adjust uptake rates to cope with a
310 wide range of inorganic carbon supplies (Burkhardt et al. 2001). Moreover, *T. weissflogii* has a
311 markedly higher fraction of direct bicarbonate transport and apparent eCA activity than *T. oceanica*
312 (Martin and Tortell 2008). In this study, *T. oceanica* showed significantly lower oxygen evolution rates



313 in the LCf treatment than in the LCs treatment before the end of the photoperiod, when the highest pH
314 and lowest CO₂ was reached. In contrast, no effects of fluctuating regime on oxygen evolution rates of
315 *T. weissflogii* were found at this time point. As shown here, cells of *T. weissflogii* benefited from their
316 inorganic carbon transport and uptake characteristics and were more tolerant of the high pH and low
317 CO₂ period under fluctuating carbonate chemistry than *T. oceanica*.

318 Under the fluctuating regime, *T. oceanica* showed higher respiration rates in both the current and OA
319 scenarios than under the corresponding steady regime. Just as *T. oceanica* makes a successful
320 compromise between iron requirements and capacity to acclimate to dynamic light regimes (Strzepek
321 and Harrison 2004), this oceanic diatom may also sacrifice its ability to acclimate to fluctuating
322 carbonate chemistry, since this is a characteristic of coastal rather than oceanic habitats. The higher
323 respiration rate may imply that this species needs more energy for maintaining its intracellular acid-base
324 balance under dynamic extracellular pH conditions, as dark respiration provides energy for growth and
325 metabolic processes (Beardall and Raven 2012). Moreover, the fluctuating regime reduced the
326 production rate of organic matter by *T. oceanica* at elevated *p*CO₂. Depressed biomass build-up has also
327 been found under dynamic light regimes (Wagner et al. 2006; Shatwell et al. 2012; Hoppe et al. 2015).
328 Together with our results, this may imply that organisms that are sensitive to fluctuating abiotic factors
329 maintain intracellular homeostasis under dynamic environments of light or *p*CO₂ at the expense of
330 reduced biomass production.

331 In contrast, either positive (POC production rate) or no obvious effects of fluctuating regime on
332 biomass production were found in the coastal species *T. weissflogii*. Coastal calcifying organisms have
333 shown the ability to achieve homeostasis within critical tissues to facilitate calcification under dynamic



334 pH/ $p\text{CO}_2$ condition, and this was suggested to be associated with diurnal and seasonal pH fluctuations
335 in coastal waters (Hendriks et al. 2015). Thus, some organisms could take advantage of the fluctuating
336 carbonate system regime to mitigate the negative effects of ocean acidification on physiological
337 performance. For instance, growth and calcification of corals benefit from oscillatory $p\text{CO}_2$ (Dufault et
338 al. 2012; Comeau et al. 2014). Organisms like *T. weissflogii* whose physiological performance were
339 enhanced or unaltered under dynamic carbonate chemistry conditions thus could be at a distinct
340 advantage in competing with species that showed negative responses to this condition (such as *T.*
341 *oceanica* in the present study).

342 Schaum et al. (2016) found that short-term plastic responses to high $p\text{CO}_2$ disappeared in a green
343 microalgae after extended experimental evolution at high $p\text{CO}_2$, particularly in fluctuating $p\text{CO}_2$
344 treatments. Whether a similar phenomenon may be operative in other algal groups such as diatoms
345 following exposures to high, fluctuating $p\text{CO}_2$ that are longer than those we employed, is currently
346 unknown. However, it is notable that growth rates and competitive abilities of all of the members of a
347 natural diatom community showed little change following one year of conditioning at two $p\text{CO}_2$ levels
348 and three temperatures, relative to the results of a short-term experiment conducted on the original
349 collected community (Tatters et al. 2013). Regardless of the responses of cell physiology to different
350 timescales of changes in $p\text{CO}_2$ concentrations, it is a significant observation that the fluctuating regime
351 reduced the production rate of organic matter in *T. oceanica* at elevated $p\text{CO}_2$.

352 Diatoms have an efficient dissipation of excess excitation energy through NPQ (Goss and Jakob
353 2010), which can be three to five times larger than that of higher plants (Ruban et al. 2004). NPQ
354 processes can be initiated in seconds to minutes (Müller et al. 2001; Eberhard et al. 2008), and so are



355 the first lines of defense for cells to respond to light stress (Lavaud et al. 2004; Lavaud et al. 2007).
356 Strzepak and Harrison (2004) found *T. weissflogii* had higher NPQ values than *T. oceanica* under both
357 low and high light conditions. Similarly, *T. weissflogii* showed 6-7 times higher NPQ than *T. oceanica* at
358 high light in this study. This difference may reflect their contrasting habitats, since species from
359 fluctuating light environments need a greater and more flexible capacity for photoprotection than those
360 from relatively stable light environments (Lavaud et al. 2007). In addition to greater and more flexible
361 capacity to dissipate excess excitation energy of coastal species, they were less sensitive to UV stress
362 than offshore ones. Inhibition of phytoplankton primary production induced by UV-A increases from
363 coastal to offshore waters (Li et al. 2011). Thus, NPQ capacity and UV sensitivity could be a major
364 factor influencing geographic distribution patterns of phytoplankton (Laviale et al. 2015).

365 The effect of fluctuating regime on *T. oceanica* was different under current and OA scenarios. Under
366 elevated rather than current $p\text{CO}_2$ condition, fluctuating carbonate chemistry decreased pigment content
367 and the production rate of organic matter. Although elevated CO_2 mitigated the limited availability of
368 $p\text{CO}_2$ that occurred at the end of photoperiod under the LCf condition, the effect of the fluctuating
369 regime under elevated $p\text{CO}_2$ tended to be negative, resulting in a decreased growth rate compared to the
370 steady regime. In our study, the same amplitude of pH variation (~ 0.5 units) was set under current and
371 elevated $p\text{CO}_2$ scenarios. Buffering capacity will decrease as the increase of dissolved inorganic carbon
372 in both coastal and oceanic seawater under projected elevated $p\text{CO}_2$ conditions (Eggleston et al. 2010;
373 Cai et al. 2011; Denman et al. 2011; Wang et al. 2013). With a larger diurnal pH variation range in the
374 future ocean, *T. oceanica* would be affected more than observed in the present study. Thus, based on our
375 results, the competitive disadvantage for organisms like *T. oceanica* would be amplified under elevated



376 $p\text{CO}_2$ condition.

377 Given the poor physiological performance of *T. oceanica* under fluctuating seawater carbonate
378 chemistry, and its limited ability to dissipate excess excitation energy through NPQ under high light,
379 this species is unlikely to be able to acclimate to coastal habitats. In contrast, *T. weissflogii* and *T.*
380 *pseudonana* appear to be insensitive to, even benefit from, fluctuating carbonate chemistry. This striking
381 contrast of physiological traits in coastal and oceanic diatoms suggests that the ability to cope with
382 fluctuating carbonate chemistry may play a role in influencing the geographic distributions of species. It
383 is possible that this ability, together with the abilities to cope with nutrient (Irwin et al. 2006), light
384 (Lavaud et al. 2007; Lavaud and Lepetit 2013; Laviale et al. 2015), and predation pressure (Irigoien et
385 al. 2005), are factors that will help to decide the spatial distribution patterns of species in both the
386 present and future oceans.

387

388 **Acknowledgements**

389 This study was supported by National Natural Science Foundation (41430967, 41120164007,
390 41206091), State Oceanic Administration (SOA, GASI-03-01-02-04), Joint project of NSFC and
391 Shandong province (Grant No. U1406403), Strategic Priority Research Program of Chinese Academy
392 of Sciences (Grant No. XDA11020302), the Fundamental Research Funds for the Central Universities
393 (20720150076), and U.S. National Science Foundation grant OCE 1538525 to F-X.F. and D.A.H. We
394 thank Prof. John Beardall for his suggestions to experimental design and Prof. Dalin Shi for providing
395 *Thalassiosira weissflogii*. We are grateful to Hangbin Miao and Dong Yan for their help in experiment.



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598 **Figure captions**

599 Figure 1. Measured pH variation over a diel cycle in the four experimental treatments (LCs, closed
600 squares; LCf, open squares; HCs, closed circles; HCf, open circles).

601
602 Figure 2. Specific growth rates of *T. weissflogii* (a) and *T. oceanica* (b) under steady (closed bars) and
603 fluctuating (open bars) regimes of ambient (LC) and elevated (HC) $p\text{CO}_2$ levels. Values are means \pm SD
604 of triplicate samples. The different letters indicate significant ($p < 0.05$) differences among treatments.

605
606 Figure 3. Production rates of chlorophyll *a* (a, b), POC (c, d), PON (e, f), and BSi (g, h) of *T. weissflogii*
607 (a, c, e, g) and *T. oceanica* (b, d, f, h) under steady (closed bars) and fluctuating (open bars) regimes of
608 ambient (LC) and elevated (HC) $p\text{CO}_2$ levels. Values are means \pm SD of triplicate samples. The
609 different letters indicate significant ($p < 0.05$) differences among treatments.

610
611 Figure 4. Chlorophyll-normalized net oxygen evolution rates in the middle of photoperiod of *T.*
612 *weissflogii* (a) and *T. oceanica* (b) under steady (closed bars) and fluctuating (open bars) regimes of
613 ambient (LC) and elevated (HC) $p\text{CO}_2$ levels. Oxygen evolution rates per cell of *T. weissflogii* (c) and *T.*
614 *oceanica* (d) of the four treatments determined 0.5, 6, and 11.5 h after illumination. Values are means \pm
615 SD of triplicate samples. The different letters indicate significant ($p < 0.05$) differences among
616 treatments.

617
618 Figure 5. Dark respiration rates per chlorophyll (a, b) or cell (c, d) in the middle of the photoperiod for



619 *T. weissflogii* (a, c) and *T. oceanica* (b, d) under steady (closed bars) and fluctuating (open bars) regimes
620 of ambient (LC) and elevated (HC) $p\text{CO}_2$ levels. Values are means \pm SD of triplicate samples. The
621 different letters indicate significant ($p < 0.05$) differences among treatments.

622

623 Figure 6. Non-photochemical quenching (NPQ) versus irradiance curves of *T. weissflogii* (solid lines)
624 and *T. oceanica* (dashed lines) measured at ambient (squares) and elevated (circles) $p\text{CO}_2$ levels. Values
625 are means \pm SD of triplicate samples. The maximum light intensities in RLCs were set as $1593 \mu\text{mol}$
626 photons $\text{m}^{-2} \text{s}^{-1}$ for *T. oceanica* of ambient $p\text{CO}_2$ level and $2130 \mu\text{mol photons m}^{-2} \text{s}^{-1}$ for the remaining
627 measurements.

628



629

630 Table 1. Cell size, respiration to photosynthesis ratio (R:P), cellular quotas of chlorophyll, particulate organic carbon (POC), particulate organic nitrogen
 631 (PON), and biogenic silica (BSi) and elemental ratios of *T. weissflogii* and *T. oceanica* under steady and fluctuating regimes of ambient (LC) and elevated
 632 (HC) $p\text{CO}_2$ levels. Values are means \pm SD of triplicate samples. The different letters indicate significant ($p < 0.05$) differences among treatments.

633

	<i>T. weissflogii</i>					<i>T. oceanica</i>						
	LCs	LCf	HCS	HCF	LCs	LCf	HCS	HCF	LCs	LCf	HCS	HCF
Cell size (μm)	12.17 \pm 0.05 ^a	12.17 \pm 0.05 ^a	12.20 \pm 0.04 ^a	12.18 \pm 0.04 ^a	5.58 \pm 0.01 ^A	5.65 \pm 0.03 ^B	5.71 \pm 0.02 ^C	5.63 \pm 0.02 ^B				
Cellular quotas												
Chl a (pg cell ⁻¹)	3.24 \pm 0.14 ^a	3.15 \pm 0.05 ^a	3.25 \pm 0.05 ^a	3.27 \pm 0.07 ^a	0.30 \pm 0.03 ^A	0.33 \pm 0.02 ^{AB}	0.38 \pm 0.06 ^B	0.29 \pm 0.02 ^A				
POC (pmol cell ⁻¹)	6.94 \pm 0.36 ^a	7.59 \pm 0.23 ^b	10.28 \pm 0.29 ^c	10.28 \pm 0.28 ^c	1.49 \pm 0.12 ^A	1.68 \pm 0.20 ^A	2.38 \pm 0.17 ^B	2.20 \pm 0.07 ^B				
PON (pmol cell ⁻¹)	1.21 \pm 0.14 ^a	1.34 \pm 0.12 ^a	1.94 \pm 0.11 ^b	1.80 \pm 0.06 ^b	0.25 \pm 0.03 ^A	0.29 \pm 0.01 ^A	0.49 \pm 0.03 ^B	0.44 \pm 0.03 ^B				
BSi (pmol cell ⁻¹)	1.11 \pm 0.01 ^{ab}	1.06 \pm 0.04 ^a	1.19 \pm 0.10 ^b	1.04 \pm 0.04 ^a	0.35 \pm 0.03 ^A	0.34 \pm 0.03 ^A	0.32 \pm 0.02 ^{AB}	0.29 \pm 0.01 ^B				
Ratios												
C:N	5.78 \pm 0.40 ^a	5.68 \pm 0.32 ^a	5.30 \pm 0.20 ^a	5.72 \pm 0.24 ^a	6.10 \pm 0.60 ^A	5.87 \pm 0.70 ^{AB}	4.90 \pm 0.16 ^B	5.05 \pm 0.36 ^B				
Si:C	0.16 \pm 0.01 ^a	0.14 \pm 0.01 ^b	0.12 \pm 0.01 ^c	0.10 \pm 0.01 ^c	0.24 \pm 0.02 ^A	0.20 \pm 0.02 ^B	0.14 \pm 0.01 ^C	0.13 \pm 0.01 ^C				
R:P	0.08 \pm 0.01 ^a	0.08 \pm 0.02 ^a	0.13 \pm 0.03 ^b	0.14 \pm 0.03 ^b	0.27 \pm 0.07 ^A	0.55 \pm 0.12 ^B	0.29 \pm 0.06 ^A	0.39 \pm 0.04 ^A				

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636 Table 2. Effective photochemical quantum yields (Φ_{PSII}) and non-photochemical quenching (NPQ) determined 0.5, 6, and 11.5 h after illumination of *T.*
 637 *weissflogii* and *T. oceanica* under steady and fluctuating regimes of ambient (LC) and elevated (HC) $p\text{CO}_2$ levels. Values are means \pm SD of triplicate
 638 samples. The different letters indicate significant ($p < 0.05$) differences among treatments.

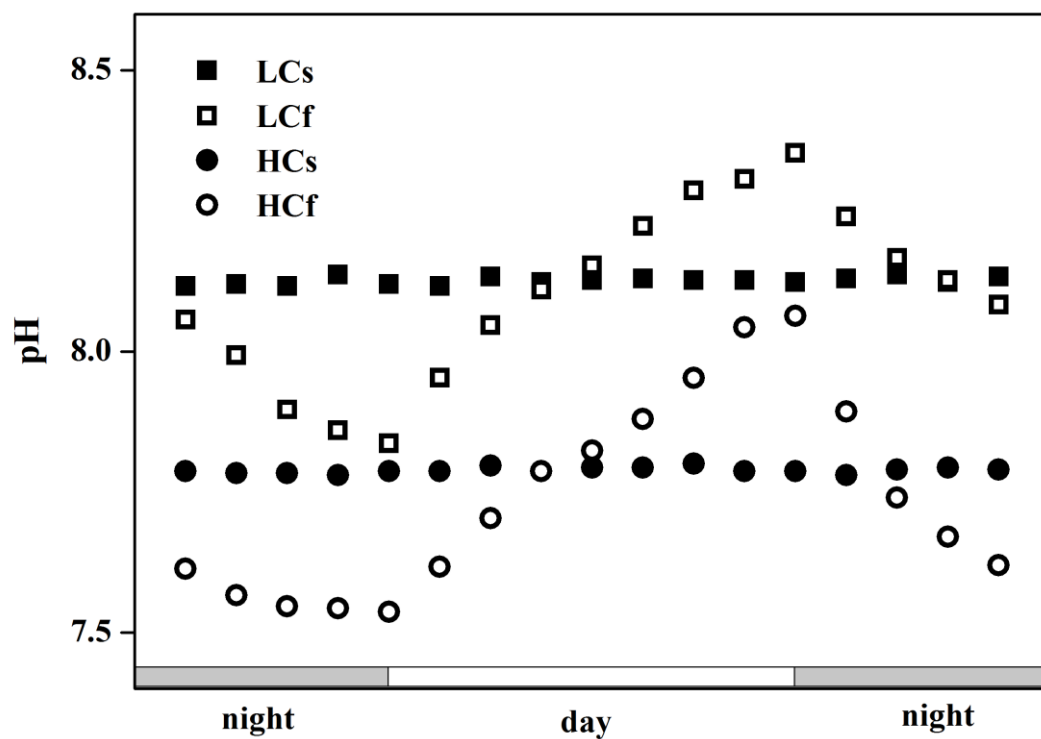
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		<i>T. weissflogii</i>								<i>T. oceanica</i>							
	Time	LCs	LCf	HCS	HCF	LCs	LCf	HCS	HCF	LCs	LCf	HCS	HCF	LCs	LCf	HCS	HCF
Φ_{PSII}	0.5	0.61 \pm 0.01 ^a	0.61 \pm 0.01 ^a	0.59 \pm 0.01 ^a	0.54 \pm 0.03 ^b	0.57 \pm 0.03 ^A	0.58 \pm 0.01 ^A	0.59 \pm 0.06 ^A	0.61 \pm 0.03 ^A	0.60 \pm 0.01 ^a	0.60 \pm 0.01 ^a	0.58 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}
	6	0.60 \pm 0.01 ^a	0.60 \pm 0.01 ^a	0.58 \pm 0.01 ^b	0.59 \pm 0.01 ^b	0.57 \pm 0.01 ^A	0.57 \pm 0.01 ^A	0.54 \pm 0.01 ^B	0.56 \pm 0.01 ^{AB}	0.61 \pm 0.01 ^a	0.60 \pm 0.01 ^a	0.58 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}
	11.5	0.58 \pm 0.01 ^a	0.58 \pm 0.01 ^a	0.57 \pm 0.01 ^b	0.57 \pm 0.01 ^{ab}	0.61 \pm 0.03 ^A	0.60 \pm 0.03 ^{AB}	0.57 \pm 0.01 ^B	0.57 \pm 0.01 ^{AB}	0.61 \pm 0.01 ^a	0.60 \pm 0.01 ^a	0.58 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^A	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}	0.57 \pm 0.01 ^{AB}
NPQ	0.5	0.13 \pm 0.02 ^a	0.13 \pm 0.01 ^a	0.13 \pm 0.01 ^a	0.23 \pm 0.05 ^b	0.10 \pm 0.02 ^A	0.06 \pm 0.02 ^A	0.08 \pm 0.05 ^A	0.12 \pm 0.09 ^A	0.13 \pm 0.01 ^a	0.13 \pm 0.01 ^a	0.11 \pm 0.01 ^b	0.09 \pm 0.01 ^{ab}	0.13 \pm 0.02 ^A	0.14 \pm 0.01 ^A	0.18 \pm 0.02 ^B	0.19 \pm 0.01 ^B
	6	0.08 \pm 0.03 ^a	0.08 \pm 0.02 ^{ab}	0.11 \pm 0.01 ^b	0.09 \pm 0.01 ^{ab}	0.13 \pm 0.02 ^A	0.14 \pm 0.01 ^A	0.18 \pm 0.02 ^B	0.19 \pm 0.01 ^B	0.08 \pm 0.01 ^a	0.07 \pm 0.01 ^a	0.06 \pm 0.01 ^b	0.07 \pm 0.01 ^{ab}	0.09 \pm 0.02 ^A	0.06 \pm 0.01 ^B	0.06 \pm 0.01 ^B	0.07 \pm 0.01 ^{AB}
	11.5	0.08 \pm 0.01 ^a	0.07 \pm 0.01 ^a	0.06 \pm 0.01 ^b	0.07 \pm 0.01 ^{ab}	0.09 \pm 0.02 ^A	0.06 \pm 0.01 ^B	0.06 \pm 0.01 ^B	0.07 \pm 0.01 ^{AB}	0.08 \pm 0.01 ^a	0.07 \pm 0.01 ^a	0.06 \pm 0.01 ^b	0.07 \pm 0.01 ^{ab}	0.09 \pm 0.02 ^A	0.06 \pm 0.01 ^B	0.06 \pm 0.01 ^B	0.07 \pm 0.01 ^{AB}

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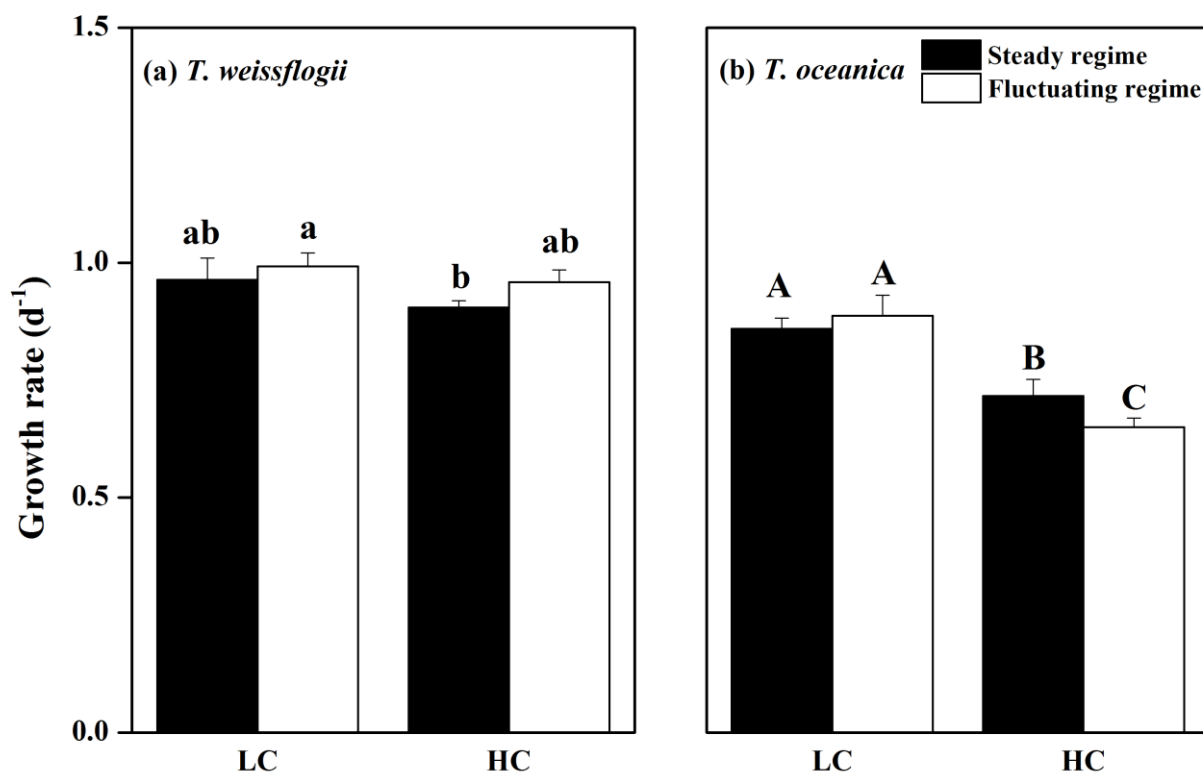


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Figure 1



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Figure 2

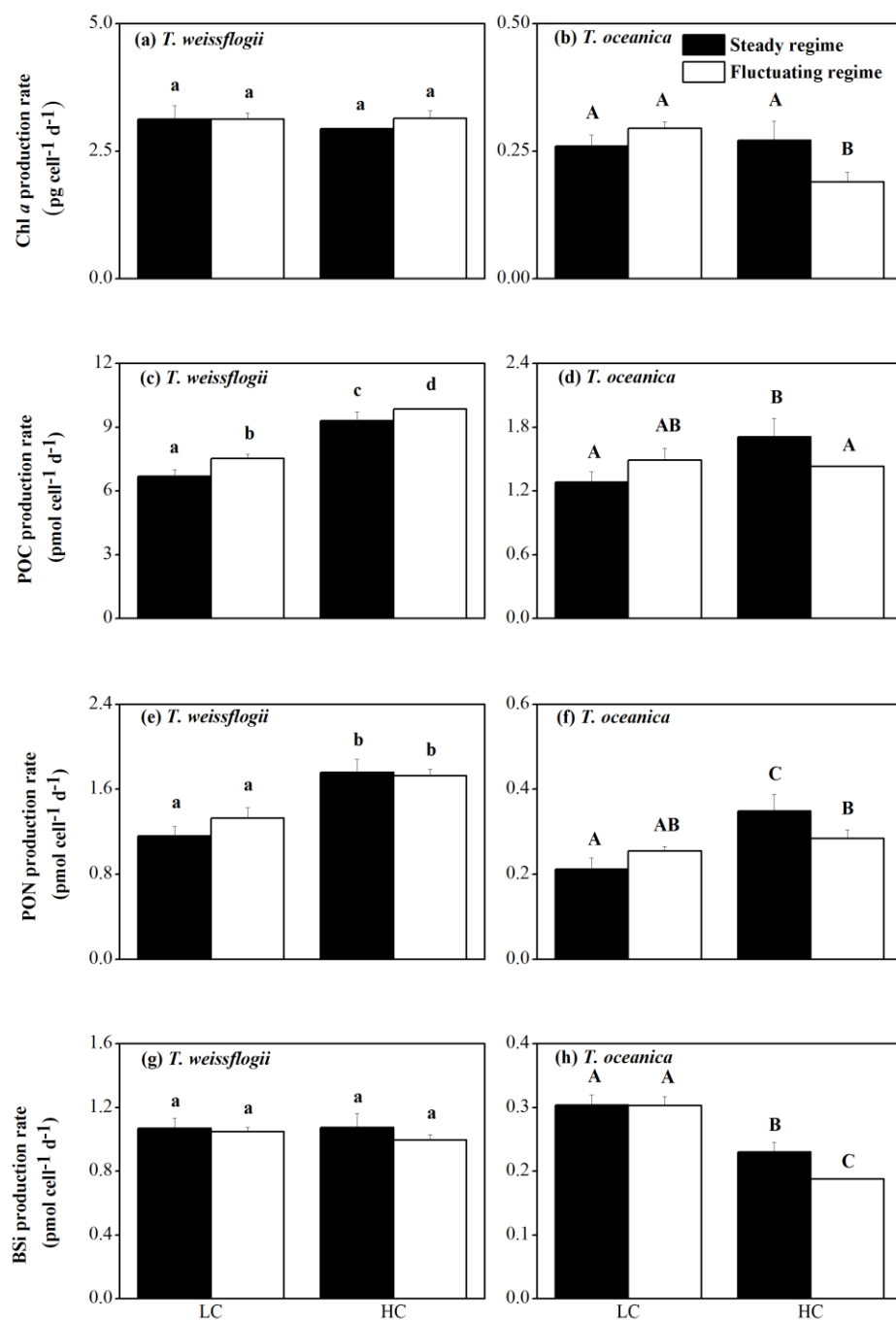
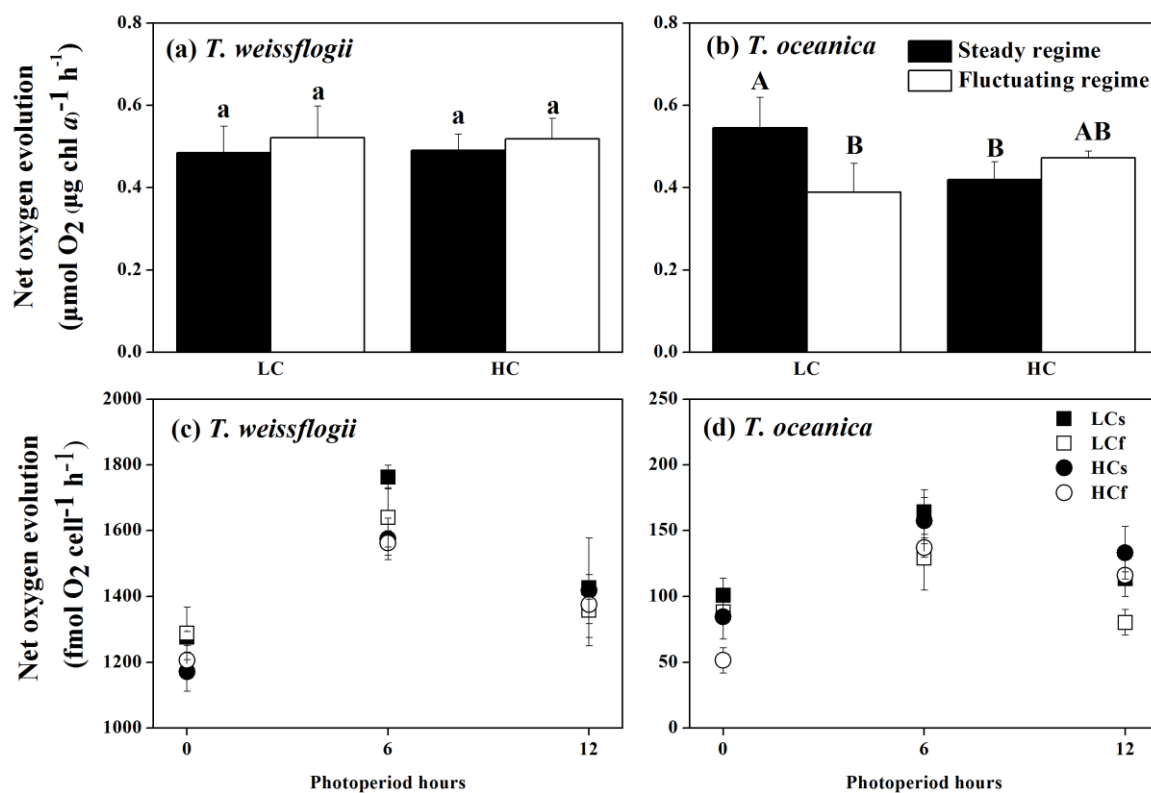


Figure 3

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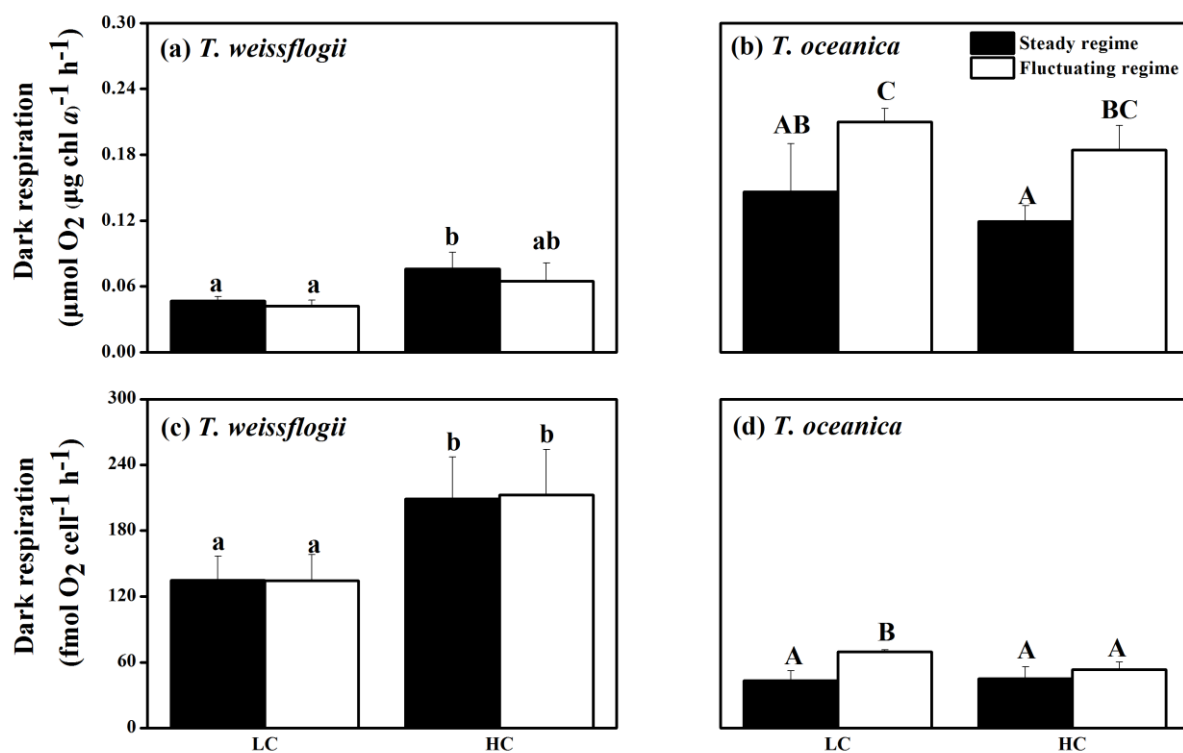
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Figure 4

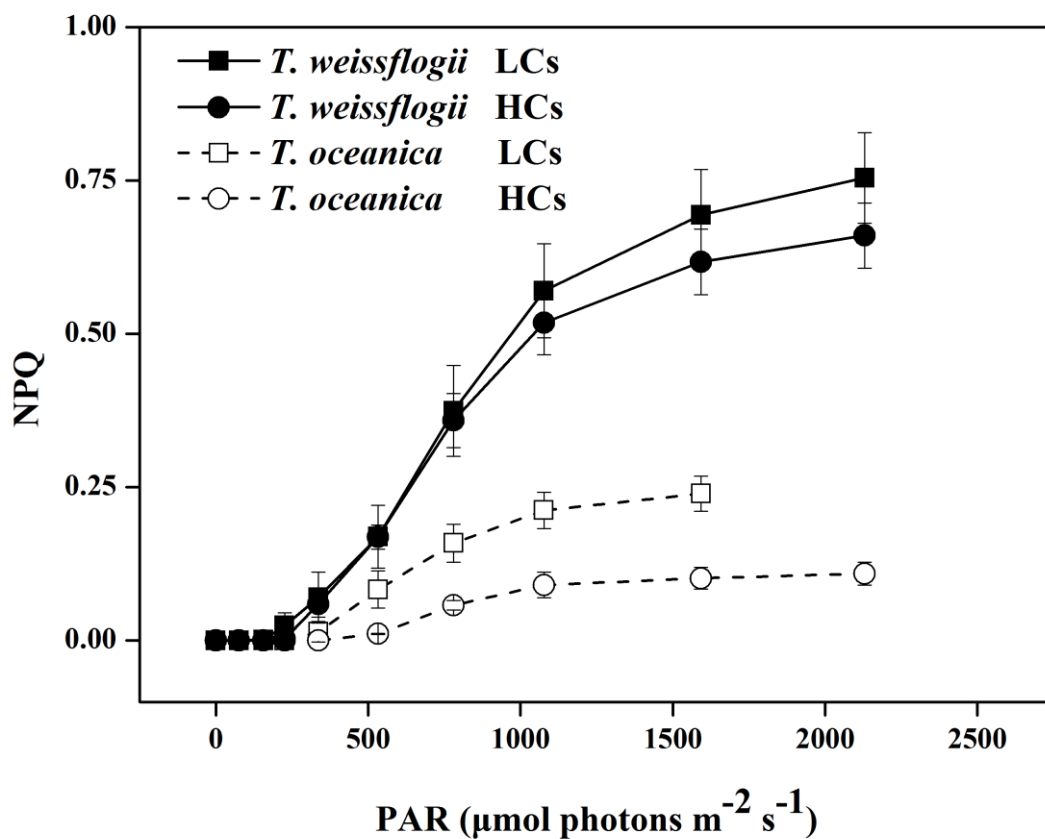


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Figure 5



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Figure 6