

Supporting Information for

Quantifying Soil Carbon Accumulation in Alaskan Terrestrial Ecosystems during the Last 15,000 Years

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Introduction

This supporting information provides the description of model parameters of soil carbon accumulation.

Model Description

The core carbon and nitrogen dynamics module of TEM was developed for upland ecosystems (Zhuang et al., 2003), where NEP is calculated at a monthly time step:

$$NEP = NPP - R_H \quad (1)$$

Soil organic carbon (SOC) heterotrophic respiration (R_H) is calculated as (Table S1):

$$R_H = K_d C_S f(M_V) e^{0.069 H_T} \quad (2)$$

where $f(M_V)$ is a nonlinear relationship that describes the effect of soil moisture in the unsaturated zone on microbial activity for decomposition. Soil moisture affects oxygen level in soils. K_d is the logarithm of heterotrophic respiration rate at 0°C. C_S is the total amount of upland mineral SOC above the plant rooting depth. H_T is mean monthly temperature of the organic layer.

We revised the decomposition to include both aerobic heterotrophic respiration above the water table which produces CO₂ and anaerobic respiration below water table, which produces both CO₂ and methane (CH₄). The soil organic carbon accumulation rate (ΔSOC) is equal to NEP, where NEP is calculated:

$$NEP = NPP - R'_H - R_{CH_4} - R_{CWM} - R_{CM} - R_{COM} \quad (3)$$

R_{CH_4} represents the monthly methane emission after methane oxidation and R_{CWM} represents the CO₂ emission due to methane oxidation (Zhuang et al., 2015). A ratio of 1:1 is assumed to calculate the CO₂ release (R_{CM}) accompanied with the methanogenesis (Tang et al., 2010; Conrad, 1999). R_{COM} represents the CO₂ release from other anaerobic processes (e.g.,

fermentation, terminal electron acceptor (TEA) reduction) (Keller and Bridgham, 2007; Keller and Takagi, 2013). The ratio of $R_{COM}:R_{CH_4}$ varies largely according to previous studies. The molar ratios ($CO_2:CH_4$) of the emission rates under inundated conditions were 4-173 for the fen and bog, respectively (Moore and Knowles, 1989), while Freeman et al. (1992) and Yavitt et al. (1987) estimated this ratio as 1. Here we assumed $R_{COM}:R_{CH_4}$ to be 5 so that the simulated $CO_2:CH_4$ of the emission rates from the anaerobic processes is ~ 10 for a fen. R'_H now represents the monthly aerobic respiration related to the variability of water-table depth (WTD) (Table S1):

$$R'_H = K_d C_{S1} f(M'_V) e^{0.069H_T} \times \frac{WTD}{LWB} \quad (4)$$

where M'_V represents the soil water content in the unsaturated zone above the WTD . The peat SOC throughout the rooting zone (fixed as 1 m) is subject to both aerobic and anaerobic decomposition and the factor (F_{sub}) determines the substrate availability was set to 1.0 within the rooting zone. The effect of F_{sub} was assumed to decrease exponentially with depth below the rooting zone (Walter and Heimann, 2000) as follows:

$$F_{sub} = e^{\frac{-(z-\frac{1}{100})}{10.0}} \quad (5)$$

where z is the peat depth (centimeters).

The initial waterlogging event was assumed to occur 2000 years before peat started to form, providing the necessary hydrological conditions for peatland initiation. The SOC between the lowest water-table boundary (LWB , a fixed model parameter, the soil below which is set saturated) and soil surface in the transient condition (C_{S1}) was obtained after a 2000-year equilibrium run, providing a stable SOC at 15 ka.

Table S1. Variables and model parameters used for calculating heterotrophic respiration in this study.

Variables	Description	Unit
R_H	Monthly heterotrophic respiration of soil organic carbon (upland soils)	$\text{g C m}^{-2} \text{ mon}^{-1}$
R'_H	Monthly aerobic heterotrophic respiration of soil organic carbon (peatland soils)	$\text{g C m}^{-2} \text{ mon}^{-1}$
R_{CH_4}	Monthly methane emission	$\text{g C m}^{-2} \text{ mon}^{-1}$
R_{CWM}	Monthly CO_2 emission due to methane oxidation	$\text{g C m}^{-2} \text{ mon}^{-1}$
R_{CM}	Monthly CO_2 emission due to methane production	$\text{g C m}^{-2} \text{ mon}^{-1}$
R_{COM}	Monthly CO_2 emission due to other anaerobic processes	$\text{g C m}^{-2} \text{ mon}^{-1}$
K_d	Logarithm of heterotrophic respiration rate at 0°C	$\text{g C g}^{-1} \text{ mol}^{-1}$
C_S	Quantity of the state variable describing total amount of soil organic carbon (SOC)	g C m^{-2}
C_{S0}	SOC between the lowest water-table boundary and soil surface (Equilibrium)	g C m^{-2}
C_{S1}	SOC between the lowest water-table boundary and soil surface (Transient)	g C m^{-2}
M_V	Soil water content (upland soils)	%
M'_V	Soil water content in the unsaturated zone (peatland soils)	%
H_T	Mean monthly temperature of the organic soil layer	$^\circ\text{C}$
LWB	Lowest water-table boundary (fixed model parameter)	mm
WTD	Water-table depth	mm
F_{Sub}	Factor (from 0 to 1) determining the substrate availability	unitless

References

- Conrad, R. (1999). Contribution of hydrogen to methane production and control of hydrogen concentrations in methanogenic soils and sediments. *FEMS Microbiology Ecology*, 28(3), 193-202.
- Freeman, C., Lock, M. A., & Reynolds, B. (1992). Fluxes of CO₂, CH₄ and N₂O from a Welsh peatland following simulation of water table draw-down: potential feedback to climatic change. *Biogeochemistry*, 19(1), 51-60.
- Keller, J. K., & Bridgham, S. D. (2007). Pathways of anaerobic carbon cycling across an ombrotrophic–minerotrophic peatland gradient.
- Keller, J. K., & Takagi, K. K. (2013). Solid-phase organic matter reduction regulates anaerobic decomposition in bog soil. *Ecosphere*, 4(5), 1-12.
- Moore, T. R., & Knowles, R. (1989). The influence of water table levels on methane and carbon dioxide emissions from peatland soils. *Canadian Journal of Soil Science*, 69(1), 33-38.
- Tang, J., Zhuang, Q., Shannon, R. D., & White, J. R. (2010). Quantifying wetland methane emissions with process-based models of different complexities. *Biogeosciences*, 7(11), 3817-3837.
- Walter, B. P., & Heimann, M. (2000). A process-based, climate-sensitive model to derive methane emissions from natural wetlands: Application to five wetland sites, sensitivity to model parameters, and climate. *Global Biogeochemical Cycles*, 14(3), 745-765.
- Yavitt, J. B., Lang, G. E., & Wieder, R. K. (1987). Control of carbon mineralization to CH₄ and CO₂ in anaerobic, Sphagnum-derived peat from Big Run Bog, West Virginia. *Biogeochemistry*, 4(2), 141-157.
- Zhuang, Q., McGuire, A. D., Melillo, J. M., Clein, J. S., Dargaville, R. J., Kicklighter, D. W., ... & Hobbie, J. E. (2003). Carbon cycling in extratropical terrestrial ecosystems of the Northern Hemisphere during the 20th century: a modeling analysis of the influences of soil thermal dynamics. *Tellus B*, 55(3), 751-776.
- Zhuang, Q., Zhu, X., He, Y., Prigent, C., Melillo, J. M., McGuire, A. D., ... & Kicklighter, D. W. (2015). Influence of changes in wetland inundation extent on net fluxes of carbon dioxide and methane in northern high latitudes from 1993 to 2004. *Environmental Research Letters*, 10(9), 095009.