Detecting small-scale spatial heterogeneity and temporal dynamics of soil organic carbon (SOC) stocks: a comparison between automatic chamber-derived C budgets and repeated soil inventories

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Abstract

Carbon (C) sequestration in soils plays a key role in the global C cycle. It is therefore crucial to adequately monitor dynamics in soil organic carbon (ΔSOC) stocks when aiming to reveal underlying processes and potential drivers. However, small-scale spatial (10-30 m) and temporal changes in SOC stocks, particularly pronounced on arable lands, are hard to assess. The main reasons for this are limitations of the well-established methods. On the one hand, repeated soil inventories, often used in long-term field trials, reveal spatial patterns and trends in ΔSOC but require a longer observation period and a sufficient number of repetitions. On the other hand, eddy covariance measurements of C fluxes towards a complete C budget of the soil-plant-atmosphere system may help to obtain temporal ΔSOC patterns but lack small-scale spatial resolution.

To overcome these limitations, this study presents a reliable method to detect both short-term temporal dynamics as well as small-scale spatial differences of ΔSOC. Therefore, a combination of automatic chamber (AC) measurements of CO_2 exchange and empirically modeled aboveground biomass development (NPP_{shoot}) was used. To verify our method, results were compared with ΔSOC observed by soil resampling.

Soil resampling and AC measurements were performed from 2010 to 2014 at a colluvial depression located in the hummocky ground moraine landscape of NE Germany. The measurement site is characterized by a variable groundwater level (GWL) and pronounced small-scale spatial heterogeneity regarding SOC and nitrogen (Nt) stocks. . Tendencies and magnitude of ΔSOC values derived by AC-measurements and repeated soil inventories corresponded well. The period of maximum plant growth was identified as being most important for the development of spatial differences in annual ΔSOC. Hence, we were able to confirm that AC-based C budgets are able to reveal small-scale spatial differences and short-term temporal dynamics of ΔSOC.

Keywords

Net ecosystem exchange (NEE), net primary productivity (NPP), biomass modeling, soil resampling

1. Introduction

Soils are the largest terrestrial reservoirs of organic carbon (SOC), storing two to three times as much C as the atmosphere and biosphere (Chen et al., 2015; Lal et al., 2004). In the context of climate change mitigation as well as soil fertility and food security, there has been considerable interest in the development of SOC, especially in erosion-affected agricultural landscapes (Berhe and Kleber, 2013; Conant et al., 2011; Doetterl et al., 2016; Stockmann et al., 2015; Van Oost et al., 2007; Xiong et al., 2016). Detecting the development of soil organic carbon stocks (ΔSOC) in agricultural landscapes needs to consider three major challenges: First, the high small-scale spatial heterogeneity of SOC (e.g., Conant et al., 2011; Xiong et al., 2016). Erosion and land use change reinforce natural spatial and temporal variability, especially in hilly landscapes such as hummocky ground moraines where correlation lengths in soil parameters of 10-30 m are very common. Second, pronounced short-term temporal dynamics, caused by, e.g., type of cover crop, frequent crop rotation and soil cultivation practices. Third, the rather small magnitude of ΔSOC compared to total SOC stocks (e.g., Conant et al., 2011; Poeplau et al., 2016).

However, information on the development of SOC is an essential precondition to improve the predictive ability of terrestrial C models (Luo et al., 2014). As a result, sensitive measurement techniques are required to precisely assess short-term temporal and small-scale (10-30 m) spatial dynamics in ΔSOC (Batjes and van Wesemael, 2015). To date, the assessment of ΔSOC is typically based on two methods, namely (i) destructive, repeated soil inventories through soil resampling and (ii) non-destructive determination of ecosystem C budgets by measurements of gaseous C exchange, C import and C export (Leifeld et al., 2011).

The first method is usually used during long-term field trials (Batjes and van Wesemael, 2015; Chen et al., 2015; Schrumpf et al., 2011). Given a sufficient time horizon of 5 to 10 years, the soil resampling method is generally able to reveal spatial patterns and trends within ΔSOC

(Batjes and van Wesemael, 2015; Schrumpf et al., 2011). Most repeated soil inventories are designed to study treatment differences in the long-term. As a result, short-term temporal dynamics in C exchange remain concealed (Poeplau et al., 2016; Schrumpf et al., 2011). A number of studies tried to overcome this methodical limitation by increasing (e.g., monthly) the soil sampling frequency (Culman et al., 2013; Wuest, 2014). This allows for the detection of seasonal patterns of ΔSOC but still mixes temporal and spatial variability of SOC because every new soil sample represents not only a repetition in time but also in space. Temporal differences observed through repeated soil sampling are therefore always spatially biased.

By contrast, the net ecosystem carbon budget (NECB; Smith et al. 2010) and thereon based temporal dynamics of ΔSOC can be easily derived through the eddy covariance (EC) technique as a common approach to obtain gaseous C exchange (Alberti et al., 2010; Leifeld et al., 2011; Skinner and Dell, 2015). However, C fluxes based on EC measurements are integrated over a larger, altering footprint area (several hectares). As a result, small-scale (< 20 m) spatial differences in ΔSOC are not detected.

Accounting for the above-mentioned methodical limitations, a number of studies investigated spatial patterns in gaseous C exchange by using manual chamber measurement systems (Eickenscheidt et al., 2014; Pohl et al., 2015). Compared to EC measurements, these systems are characterized by a low temporal resolution, where the calculated net ecosystem CO_2 exchange (NEE) is commonly based on extensive gap filling (Gomez-Casanovas et al., 2013; Savage and Davidson, 2003) conducted, e.g., using empirical modeling (Hoffmann et al., 2015). Therefore, management practices and different stages in plant development that are needed to precisely detect NEE often remain unconsidered (Hoffmann et al., 2015).

Compared to mentioned approaches for detecting ΔSOC by either repeated soil sampling or observations of the gaseous C exchange, automatic chamber (AC) systems combine several advantages. On the one hand flux measurements of the same spatial entity avoid the mixing of spatial and temporal variability, as done in case of point measurements by repeated soil inventories. On the other hand, AC measurements combine advantages of EC and manual chamber systems because they not only increase the temporal resolution compared to manual chambers but also allow for the detection of small-scale spatial differences and treatment comparisons regarding the gaseous C exchange (Koskinen et al., 2014).

To date hardly any direct comparisons between AC-derived C budgets and soil resampling-based ΔSOC values have been reported in the literature. Leifeld et al. (2011) and Verma et al. (2005) compared the results of repeated soil inventories with EC-based C budgets over 5- and 3-year study periods, respectively. Even though temporal dynamics in ΔSOC were shown e.g. for grazed pastures and intensively used grasslands (Skinner and Dell 2015; Leifeld et al., 2011), no attempt was made to additionally detect small-scale differences in ΔSOC. In our study, we introduce the combination of AC measurements and empirically modeled aboveground biomass production (NPP_{shoot}) as a precise method to detect small-scale spatial differences and short-term temporal dynamics of ΔSOC. Measurements were performed from 2010 to 2014 under a *silage maize/winter fodder rye/sorghum-Sudan grass hybrid/alfalfa* crop rotation at an experimental plot located in the hummocky ground moraine landscape of NE Germany.

We hypothesize that the AC-based C budget method is able to detect small-scale spatial and short-term temporal dynamics of ΔSOC in an accurate and precise manner. Therefore, we compare ΔSOC values measured by soil resampling with ΔSOC values derived through AC-based C budgets (Fig. 1).

2. Materials and methods

2.1 Study site and experimental setup

Measurements were performed at the 6-ha experimental field “CarboZALF-D”. The site is located in a hummocky arable soil landscape within the Uckermark region (NE-Germany; 53°23`N, 13°47`E, ~50-60 m a.s.l.). The temperate climate is characterized by a mean annual air temperature of 8.6°C and annual precipitation of 485 mm (1992–2012, ZALF research station, Dedelow). Typical landscape elements vary from flat summit and depression locations with a gradient of approximately 2 %, across longer slopes with a medium gradient of approx. 6 %, to short and rather steep slopes with a gradient of up to 13 %. The study site shows complex soil patterns mainly influenced by erosion, relief and parent material, e.g., sandy to marly glacial and glaciofluvial deposits. The soil type inventory of the experimental site consists of non-eroded Albic Luvisols (Cutanic) at the flat summits, strongly eroded Calcic Luvisols (Cutanic) on the moderate slopes, extremely eroded Calcaric Regosols on the steep slopes, and a colluvial soil, i.e., Endogleyic Colluvic Regosols (Eutric), over peat in the depression (IUSS Working Group WRB, 2015).

During June 2010, four automatic chambers and a WXT520 climate station (Vaisala, Vantaa, Finland) were set up at the depression (Sommer et al., 2016) (see 2.2.1). The chambers were arranged along a topographic gradient (upper (A), upper middle (B), lower middle (C), and lower (D) chamber position; length ~30 m; difference in altitude ~1 m) within in a distance of approx. 5 m of each other (Fig. 2). As part of the CarboZALF project, a manipulation experiment was carried out at the end of October 2010, i.e., after the vegetation period. Topsoil material from a neighboring hillslope was incorporated into the upper soil layer of the depression (Ap horizon). The amount of translocated soil was equivalent to tillage erosion of a decennial time horizon (Sommer et al., 2016). The change in SOC for each chamber was monitored by three topsoil inventories, carried out (I) prior to soil manipulation during April 2009, (II) after soil manipulation during April 2011, and (III) during December 2014. ΔSOC derived through soil resampling and AC-based C budgets, was compared for the period between April 2011 and December 2014 (Fig. 1).

Records of meteorological conditions (1 min frequency) include measurements of air temperature at 20 cm and 200 cm height, PAR (inside and outside the chamber), air humidity, precipitation, air pressure, wind speed and direction. Soil temperatures at depths of 2 cm, 5 cm, 10 cm and 50 cm were recorded using thermocouples, installed next to the climate station (107, Campbell Scientific, UT, USA).

The groundwater level (GWL) was measured using tensiometers assuming hydrostatic equilibrium. The tensiometers were installed at a soil depth of 160 cm, at soil profile locations in the upper and lower end of the transect. The average GWL of both profiles was used for further data analysis. Data gaps < 2 days were filled using simple linear interpolation. Larger gaps in GWL did not occur. The measurement site was cultivated with five different crops during the study period, following a practice-orientated and erosion-expedited farming procedure. The crop rotation was silage maize (*Zea mays*) - winter fodder rye (*Secale cereale*) - sorghum-Sudan grass hybrid (*Sorghum bicolor* x *sudanese*) - winter triticale (*Triticosecale*) - alfalfa (*Medicago sativa*). Cultivation and fertilization details are presented in Tab. A.1. Aboveground biomass (NPP_{shoot}) development was monitored using up to four biomass sampling campaigns during the growing season, covering the main growth stages. Additional measurements of leaf area index (LAI) started in 2013. Collected biomass samples were chopped and dried to a constant weight (48 h at 105°C). The C, N, K and P contents were determined using elementary analysis (C, N: TruSpec CNS analyzer, LECO Ltd., Mönchengladbach, Germany) and Kjehldahl digestion (P, K; AT200, BeckmanCoulter (Olympus), Krefeld, Germany and AAS-iCE3300, ThermoFisher-SCIENTIFIC GmbH, Darmstadt, Germany). To assess the potential impact of chamber placement on plant growth, chemical analyses were carried out for the final harvests of each chamber and compared to biomass samples collected next to each chamber.

2.2 C budget method

2.2.1 Automatic chamber system

Automatic flow-through non-steady-state (FT-NSS) chamber measurements (Livingston and Hutchinson, 1995) of CO_2 exchange were conducted from January 2010 until December 2014. The AC system consists of 4 identical, rectangular, transparent polycarbonate chambers (thickness of 2 mm; light transmission ~70 %). Each chamber has a height of 2.5 m and covers a surface area of 2.25 m^2 (volume: 5.625 m^3). To adapt for plant height (alfalfa), the chamber volume was reduced to 3.375 m^3 in autumn 2013. Airtight closure during measurements was ensured by a rubber belt that sealed at the bottom of each chamber. A 30-cm open-ended tube on the slightly concave top of the chambers guided rain water into the chamber and additionally assured pressure equalization. Two small axial fans (5.61 m^3 min^{-1}) were used for mixing the chamber headspace. The chambers were mounted onto steel frames with a height of 6 m and lifted between measurements using electrical winches at the top. For controlling the AC system and data collection, a CR1000 data logger was used (Campbell Scientific, UT, USA). The CO_2 concentration changes over time were measured within each chamber using a carbon dioxide probe (GMP343, Vaisala, Vantaa, Finland) connected to a vacuum pump (0.001 m^3 min^{-1}; DC12/16FK, Fürgut, Tannheim, Germany). All CO_2 probes were calibrated prior to installation using ± 0.5 % accurate gases containing 0 ppm, 200 ppm 370 ppm, 600 ppm, 1000, ppm and 4000 ppm CO_2. The operation schedule of the AC system, decisively influenced by agricultural treatments, is presented in A.2. The chambers closed in parallel at an hourly frequency, providing one flux measurement per chamber and hour. The measurement duration was 5-20 minutes, depending on season and time of day. Nighttime measurements usually lasted 10 min during the growing season and 20 min during the non-growing season (due to lower concentration increments). The length of the daytime measurements was up to 10 min, depending on low PAR fluctuations (< 20 %). CO_2 concentrations (inside the chamber) and general environmental conditions, such as PAR (SKP215, Skye, Llandridad Wells, UK) and air temperatures (107, Campbell Scientific, UT, USA), were recorded inside and outside the chambers at a 1 min frequency from 2010 to 2012 and a 15 sec frequency from October 2012.

2.2.2 CO_2 flux calculation and gap filling

An adaptation of the modular R program script, described in detail by Hoffmann et al. (2015), was used for stepwise data processing. The atmospheric sign convention was used for the components of gaseous C exchange (ecosystem respiration (R_{eco}; sum of autotrophic and heterotrophic respiration), gross primary production (GPP) and NEE), whereas positive values for ΔSOC indicate a gain and negative values a loss in SOC. Based on records of environmental variables and CO_2 concentration change within the chamber headspace, CO_2 fluxes were calculated and parameterized for R_{eco} and GPP within an integrative step. Subsequently, R_{eco}, GPP, and NEE were modeled for the entire measurement period using climate station data. Statistical analyses, model calibration and comprehensive error prediction were provided for all steps of the modeling process.

CO_2 fluxes (F, µmol C m^{-2} s^{-1}) were calculated according to the ideal gas law (Eq. 1).

$$F = \frac{pV}{RTA} * \frac{\Delta c}{\Delta t} \quad \text{[Eq. 1]}$$

where $\Delta c/\Delta t$ is the concentration change over measurement time, A and V denote the basal area and chamber volume, respectively, and T and p represent the air temperature inside the chamber (K) and air pressure. Because plants below the chambers accounted for < 0.2 % of the total chamber volume, a static chamber volume was assumed. R is a constant (8.3143 m^3 Pa K^{-1} mol^{-1}). To calculate $\Delta c/\Delta t$, data subsets based on a variable moving window with a minimum length of 4 minutes were used (Hoffmann et al., 2015). $\Delta c/\Delta t$ was computed by applying a linear regression to each data subset, relating changes in chamber headspace CO_2 concentration to measurement time (Leiber-Sauheitl et al., 2013; Leifeld et al., 2014; Pohl et al., 2015). In the case of the 15-sec measurement frequency, a death-band of 5 % was applied prior to the moving window algorithm. Thus, data noise that originated from either turbulence or pressure fluctuation caused by chamber deployment or from increasing saturation and canopy microclimate effects was excluded (Davidson et al., 2002; Kutzbach et al., 2007; Langensiepen et al., 2012). Due to the low measurement frequency, no data points were discarded for records with 1-min measurement frequency (2010-2012). The resulting CO_2 fluxes per measurement (based on the moving window data subsets) were further evaluated according to the following exclusion criteria: (i) range of within-chamber air temperature not larger than ± 1.5 K (R_{eco} and NEE fluxes) and a PAR deviation (NEE fluxes only) not larger than ± 20 % of the average to ensure stable environmental conditions within the chamber throughout the measurement; (ii) significant regression slope ($p \leq 0.1$, *t*-test); and (iii) non-significant tests ($p > 0.1$) for normality (Lillifor´s adaption of the Kolmogorov-Smirnov test), homoscedasticity (Breusch-Pagan test) and linearity of CO_2 concentration data. Calculated CO_2 fluxes that did not meet all exclusion criteria were discarded. In cases where more than one flux per measurement met all exclusion criteria, the CO_2 flux with the steepest slope was chosen.

To account for measurement gaps and to obtain cumulative NEE values, empirical models were derived based on nighttime R_{eco} and daytime NEE measurements following Hoffmann et al. (2015). For R_{eco}, temperature-dependent Arrhenius-type models were used and fitted for recorded air as well as soil temperatures in different depths (Lloyd and Taylor 1994; Eq. 2).

$$R_{eco} = R_{ref} * e^{E_0\left(\frac{1}{T_{ref}-T_0}-\frac{1}{T-T_0}\right)} \qquad \text{[Eq. 2]}$$

where R_{eco} is the measured ecosystem respiration rate [µmol^{-1} C m^{-2} s^{-1}], R_{ref} is the respiration rate at the reference temperature (283.15 K; T_{ref}); E_0 is an activation energy like parameter; T_0 is the starting temperature constant (227.13 K) and T is the mean air or soil temperature during the flux measurement. Out of the four R_{eco} models (one model for air temperature, soil temperature in 2 cm, 5 cm and 10 cm depth) obtained for nighttime R_{eco} measurements of a certain period, the model with the lowest Akaike Information Criterion (AIC) was used.

GPP fluxes were derived using a PAR-dependent, rectangular hyperbolic light response function based on the Michaelis-Menten kinetic (Elsgaard et al., 2012; Hoffmann et al., 2015; Wang et al., 2013; Eq. 3). Because GPP was not measured directly, GPP fluxes were calculated as the difference between measured NEE and modeled R_{eco} fluxes.

$$GPP = \frac{GP_{max} * \alpha * PAR}{\alpha * PAR + GP_{max}} \qquad \text{[Eq. 3]}$$

where GPP is the calculated gross primary productivity [µmol^{-1} CO_2 m^{-2} s^{-1}]; GP_{max} is the maximum rate of C fixation at infinite PAR [µmol CO_2 m^{-2} s^{-1}]; α is the light use efficiency [mol CO_2 mol^{-1} photons] and PAR is the photon flux density (inside the chamber) of the photosynthetically active radiation [μmol^{-1} photons m^{-2} s^{-1}]. In cases where the rectangular hyperbolic light response function did not result in significant parameter estimates, a non-rectangular hyperbolic light-response function was used (Gilmanov et al. 2007, 2013; Eq. 4).

$$GPP = \alpha * \mathrm{PAR} + \mathrm{GP_{max}} - \sqrt{(\alpha * \mathrm{PAR} + \mathrm{GP_{max}})^2 - 4 * \alpha * \mathrm{PAR} * \mathrm{GP_{max}} * \theta} \quad \text{[Eq. 4]}$$

where θ is the convexity coefficient of the light-response equation (dimensionless).

Due to plant growth and season, parameters of derived R_{eco} and GPP models may vary with time. To account for this, a moving window parameterization was performed, by applying fluxes of a variable time window (2-21 consecutive measurement days) to Eq.2-4. Temporally overlapping R_{eco} and GPP model sets were evaluated and discarded in case of positive (GPP), negative (R_{eco}) or insignificant parameter estimates. Finally, the model set with the lowest AIC (R_{eco}) was used. If no fit or a non-significant fit was achieved, averaged flux rates were applied for R_{eco} and GPP. The length of the averaging period was thereby selected by choosing the variable moving window with the lowest standard deviation (SD) of measured fluxes. This procedure was repeated until the whole study period was parameterized.

Based on continuously monitored temperature and PAR (outside the chamber), R_{eco}, GPP and NEE were modeled in half-hour steps for the entire study period. Because GPP was parameterized based on PAR records inside but modeled with PAR records outside the chamber, no PAR correction in terms of reduced light transmission was needed. Uncertainty of annual CO_2 exchange was quantified using a comprehensive error prediction algorithm described in detail by Hoffmann et al. (2015).

2.2.3 Modeling aboveground biomass dynamics

Aboveground biomass development (NPP_{shoot}) was predicted using a logistic empirical model (Yin et al., 2003; Zeide, 1993). From 2010 to 2012, modeled NPP_{shoot} was based on the relationship between sampling date and the C content of harvested dry biomass measured during sampling campaigns (three to four times per year following plant development). For alfalfa in 2013 and 2014, NPP_{shoot} was modeled based on biweekly measurements of LAI because no additional biomass sampling was performed between the multiple cuts per year. To calculate the C content corresponding to the measured LAI, the relationship between LAI prior to the chamber harvest and the C content measured in the chamber harvest of all six alfalfa cuts was used. Daily values of C stored within NPP_{shoot} were calculated using derived logistic functions.

2.2.4 Calculation of ΔSOC

Annual ΔSOC for each chamber was determined as the sum of annual NEE and NPP_{shoot}, representing C removal due to the chamber harvest (Eq. 4; Leifeld et al., 2014). Temporal dynamics in ΔSOC were calculated as the sum of daily NEE and NPP_{shoot}.

$$\Delta SOC_n = \sum_{i=1}^{n}[NEE_i + CH_4 + (NPP_{shoot_i} - C_{import}) + \Delta DOC_i + \Delta DIC_i] \qquad \text{[Eq. 5]}$$

Several minor components of Eq. 5 were not considered (see also Hernandez-Ramirez et al., 2011). First, C import (C_{import}) due to seeding and fertilization, which was close to zero because the measurement site was fertilized by a surface application of mineral fertilizer throughout the entire study period. Second, methane (CH_4-C) emissions, which were measured manually at the same experimental field but did not exceed a relevant order of magnitude (-0.01 g C m^{-2} y^{-1}) and were therefore not included in the ΔSOC calculation. Third, lateral C fluxes, originating from dissolved organic (DOC) and inorganic carbon (DIC) as well as particulate soil organic carbon

(SOC_p). In addition to the rather small magnitude of the subsurface lateral C fluxes in soil solution (Rieckh et al., 2012), it was assumed that their C input equaled C output at the plot scale. Lateral SOC_p transport along the hillslope was excluded by grassland stripes established between experimental plots in 2010 (Fig. 1 in Sommer et al., 2016).

2.3 Soil resampling method

To obtain ΔSOC using the soil resampling method, soil samples were collected three times during the study period. Initial SOC along the topographic gradient was monitored prior to soil manipulation during April 2009 at two soil pits, which were sampled by pedogenetic horizons. After soil manipulation, a 5-m raster sampling of topsoils (Ap horizons) was performed during April 2011. Each Ap horizon was separated into an upper (0-15 cm) and lower segment (15-25 cm), which were analyzed separately for bulk density, SOC, Nt and coarse fraction (< 2 mm) (data not shown). From these data, SOC and Nt mass densities were calculated separately for each segment and finally summed up for the entire Ap-horizon (0-25 cm). The mean SOC and Nt content for the Ap horizon of each raster point was calculated by dividing SOC or Nt mass densities (0-25 cm) through the fine-earth mass (0-25 cm). In December 2014, composite soil samples of the Ap horizon were collected. The composite samples consist of samples from four sampling points in a close proximity around each chamber. Prior to laboratory analysis coarse organic material was discarded from collected soil samples (Schlichting et al. 1995). Thermogravimetric desiccation at 105°C was performed in the laboratory for all samples to determine bulk densities ($Mg\ m^{-3}$). Bulk soil samples were air dried, gently crushed and sieved (2 mm) to obtain the fine fraction (particle size < 2 mm). The total carbon and total nitrogen contents were determined by elementary analysis (TruSpec CNS analyzer, LECO Ltd., Mönchengladbach, Germany) as carbon dioxide via infrared detection after dry combustion at

1250°C (DIN ISO10694, 1996), in duplicate. As the soil horizons did not contain carbonates, total carbon was equal to SOC.

2.4 Uncertainty prediction and statistical analysis

Uncertainty prediction for ΔSOC derived by the C budget method was performed according to Hoffmann et al. (2015), following the law of error propagation. To test for differences in topsoil SOC (SOC_{Ap}) and total nitrogen (Nt) stocks between soil resampling performed after soil manipulation in 2010 and 2014, a paired *t*-test was applied. Computation of uncertainty prediction and calculation of statistical analyses were performed using R 3.2.2.

3. Results

3.1 C budget method

3.1.1 NEE and NPP_{shoot} dynamics

NEE and its components R_{eco} and GPP were characterized by a clear seasonality and diurnal patterns. Seasonality followed plant growth and management events (e.g., harvest; Fig. 3), Highest CO_2 uptake was thus observed during the growing season, whereas NEE fluxes during the non-growing season were significantly lower. Diurnal patterns were more pronounced during the growing season and less obvious during the non-growing season. In general R_{eco} fluxes were higher during daytime, whereas GPP and NEE, in case of present cover crops, were lower or even negative, representing a C uptake during daytime by the plant-soil system. Annual NEE was crop dependent, ranging from -1600 g C m^{-2} y^{-1} to -288 g C m^{-2} y^{-1}. Highest annual uptakes were observed for maize and sorghum during 2011 and 2012, whereas alfalfa cultivation showed lower annual NEE (Tab. 1). From 2010 to 2012, annual NEE followed the topographic gradient, with higher NEE in the direction of the depression and lower NEE away from the depression. These small-scale spatial differences in gaseous C exchange changed with alfalfa cultivation. As a result, only minor differences between the chamber positions were observed, showing no clear trend or tendency (Tab. 1).

C in living biomass (due to biomass sampling campaigns and LAI measurements) and C removals due to harvest were in general well reflected by modeled NPP_{shoot} (Fig. 4). Annual C removal due to harvest was clearly crop dependent, with highest NPP_{shoot} for maize and sorghum ranging from 420 g C m^{-2} to 1238 g C m^{-2}, and lower values in the case of winter fodder rye and alfalfa. Similar to NEE from 2010 to 2012, annual sums of NPP_{shoot} followed the topographic gradient, with lower values close to the depression (Tab. 1). Again, lower differences in annual NPP_{shoot} between the chambers and no spatial trends were found for alfalfa in 2013 and 2014.

3.1.2 ΔSOC dynamics

Temporal and spatial dynamics of continously cumulated daily ΔSOC values during the four years after soil manipulation are shown in Fig. 5. Differences in ΔSOC were in general less pronounced during the non-growing season compared to the growing season. During the non-growing season, differences were mainly driven by differences in R_{eco} rather than GPP or NPP_{shoot}. This changed at the beginning of the growing season, when ΔSOC responded to changes in cumulative NEE and NPP_{shoot}. Hence, up to 79 % of the standard deviation of estimated annual ΔSOC developed during the period of maximum plant growth. Except for the lower middle chamber position, alfalfa seemed to counterbalance spatial differences in ΔSOC that developed during previous years (Fig. 5).

Annual ΔSOC values derived by the C budget method are presented in Tab. 1. Highest annual SOC gains were obtained in 2012 for winter fodder rye and sorghum-Sudan grass, reaching an average of 474 g C m^{-2} y^{-1}. In contrast, maize cultivation during 2011 was characterized by C losses between 59 g C m^{-2} y^{-1} and 169 g C m^{-2} y^{-1}. However, prior to soil manipulation, maize showed an average SOC gain of 102 g C m^{-2} y^{-1}.

3.2 Soil resampling method

As a result of soil translocation in 2010, initially measured SOC_{Ap} stocks increased by an average of 780 g C m^{-2}. However, due to the lower C content of the translocated topsoil material (0.76 %), the SOC_{Ap} content of the measurement site dropped by 10 - 14 % after soil manipulation (Tab. 1). Significant differences (paired *t*-test; t = -2.48, $p < 0.09$), which showed an increase in SOC_{Ap} of up to 11 %, were found between SOC_{Ap} stocks measured in 2010 and 2014. Three out of the four chamber positions showed a C gain during the 4 measurement years following soil manipulation. C gains were similar for the upper and lower chamber positions, but lower for the upper middle position. No change in SOC was obtained in the case of the lower middle (Fig. 5; Fig. 6) chamber position.

3.3 Method comparison

Average annual ΔSOC values for the soil resampling and C budget method are shown in Fig. 6. ΔSOC based on these methods showed a good overall agreement, with similar tendencies and magnitudes (Fig. 6). Irrespective of the applied method, significant differences were found between SOC stocks measured directly after soil manipulation in 2010 and SOC stocks measured in 2014. Following soil manipulation, both methods revealed similar tendencies in site and chamber-specific ΔSOC (Fig. 6). Both methods indicated a clear C gain for three out of the four chamber positions. C gains derived by the C budget method were similar for the upper, upper middle and lower chamber positions. By contrast, C gains derived by the soil resampling method were slightly but not significantly lower (paired *t*-test; t = -1.23, $p > 0.30$). This was most pronounced for the upper middle chamber position. No change in ΔSOC and only a minor gain in C was observed for the lower middle chamber position according to both methods. Differences between chamber positions indicate the presence of small-scale spatial ΔSOC dynamics typical of soils.

4. Discussion

4.1 Accuracy and precision of applied methods

Despite the similar magnitude and tendencies of the observed ΔSOC values, both methods were subject to numerous sources of uncertainty. These errors affect the accuracy and precision of observed ΔSOC values differently, which might help to explain differences between the soil resampling and the C budget method.

The soil resampling method is characterized by high measurement precision, which allows for the detection of relatively small changes in SOC. Related uncertainty in derived spatial and temporal ΔSOC dynamics is therefore mainly attributed to the measurement accuracy, affected by sampling strategy and design (Batjes and van Wesemael, 2015; De Gruijter et al., 2006). This includes (i) the spatial distribution of collected samples, (ii) the sampling frequency, (iii) the sampling depth and (iv) whether different components of soil organic matter (SOM) are excluded prior to analyses. The first aspect determines the capability to detect the inherent spatial differences in SOC stocks. This allows the conclusion that point measurements do not necessarily represent AC measurements, which integrate over the spatial variability within their basal area. The second aspect defines the temporal resolution, even though the soil resampling method is not able to perfectly separate spatial from temporal variability because repeated soil samples are biased by inherent spatial variability of the measurement site. The third aspect sets the vertical system boundary, which is often limited because only topsoil horizons are sampled within a number of soil monitoring networks (Van Wesemael et al., 2011) and repeated soil inventories (Leifeld et al., 2011). Similarly, the fourth aspect defines which components of SOM are specifically analyzed. Usually, coarse organic material is discarded prior to analysis (Schlichting et al., 1995) and therefore, total SOC is not assessed (e.g., roots, harvest residues, etc.).

In comparison, the C budget method considers any type of organic material present in soil by integrating over the total soil depth. As a result, both methods have a different validity range and area, which makes direct quantitative comparison more difficult. This may explain the higher uptake reported for three out of four chamber positions in the case of the C budget method.

In contrast to the soil resampling method, we postulate a higher accuracy and a lower precision in the case of the AC-based C budget method. The reasons for this include a number of potential errors affecting especially the measurement precision of the AC system, whereas over a constant area and maximum soil depth, integrated AC measurements increase measurement accuracy. First, it is currently not clear whether microclimatological and ecophysiological disturbances due to chamber deployment, such as the alteration of temperature, humidity, pressure, radiation, and gas concentration, may result in biased C flux rate estimates (Juszczak et al., 2013; Kutzbach et al., 2007; Lai et al., 2012; Langensiepen et al., 2012). Second, uncertainties related to performed flux separation and gap-filling procedures may influence the obtained annual gaseous C exchange (Gomez-Casanovas et al., 2013; Görres et al., 2014; Moffat et al., 2007; Reichstein et al., 2005). Although continuous operation of the AC system should allow for direct derivation of C budgets from measured CO_2 exchange and annual yields, in practice, data gaps always occur. To fill the measurement gaps, temperature- and PAR-dependent models are derived and used to calculate R_{eco} and GPP, respectively (Hoffmann et al. 2015). Due to the transparent chambers used, modeled R_{eco} is solely based on nighttime measurements. Hence, systematic differences between nighttime and daytime R_{eco} will yield an over- or underestimation of modeled R_{eco}. Because modeled R_{eco} is used to calculate GPP fluxes, GPP will be affected in a similar manner. However, the systematic over- or underestimation of fluxes in both directions may counterbalance the computed NEE, and estimated C budgets may be unaffected. Third, the development of NPP_{shoot} underneath the chamber might be influenced by the permanently installed AC system. Fourth, several minor components such as leaching losses of dissolved inorganic and organic carbon (DIC and DOC), C transport via runoff and atmospheric C deposition were not considered within the applied budgeting approach (see also 2.7).

Despite the uncertainties mentioned above, error estimates for annual NEE in this study are within the range of errors presented for annual NEE estimates derived from EC measurements (30 to 50 g C m^{-2} y^{-1}) (e.g., Baldocchi, 2003; Dobermann et al., 2006; Hollinger et al., 2005) and below the minimum detectable difference (MDD) reported for most repeated soil inventories (e.g., Batjes and Van Wesemael, 2015; Knebl et al., 2015; Necpálová et al., 2014; Saby et al., 2008; Schrumpf et al., 2011; VandenBygaart, 2006).

4.2 Plausibility of observed ∆SOC

Both the soil resampling and the C budget method showed C gains during the four years following soil manipulation. A number of authors calculated additional C sequestration due to soil erosion (Berhe et al., 2007; Dymond, 2010; VandenBygaart et al., 2015; Yoo et al., 2005), which was explained by the burial of replaced C at depositional sites and dynamic replacement at eroded sites (e.g., Doetterl et al., 2016). This is in accordance with erosion-induced C sequestration postulated by, e.g., Berhe and Kleber (2013) and Van Oost et al. (2007). In addition, observed C sequestration could also be a result of the manipulation-induced saturation deficit in SOC. By adding topsoil material from an eroded unsaturated hillslope soil, the capacity and efficiency to sequester C was theoretically increased (Stewart et al., 2007). Hence, additional

C was stored at the measurement site. This might be due to physicochemical processes, such as physical protection in macro- and micro aggregates (Six et al., 2002) or chemical stabilization by clay and iron minerals (Kleber et al., 2015).

Irrespective of the similar C gain observed by both methods, crop-dependent differences in ΔSOC were only revealed by the C budget method. The reason is the higher temporal resolution of AC-derived C budgets, displaying daily C losses and gains. Observed crop-dependent differences in ΔSOC are in accordance with, e.g., Kutsch et al. (2010), Jans et al. (2010), Hollinger et al. (2005) and Verma et al. (2005), who reported comparable EC-derived C balances for inter alia, maize, sorghum and alfalfa.

In 2012, substantial positive annual ΔSOC values were observed. Due to low precipitation during May and June, germination and plant growth of sorghum-Sudan grass was delayed (Fig. 4). As a result, the reproductive phenological stage was drastically shortened. This reduced C losses prior to harvest due to higher R_{eco}:GPP ratios (Wagle et al., 2015). In addition, the presence of cover crops during spring and autumn could have increased SOC, as reported by Lal et al. (2004), Ghimire et al. (2014) and Sainju et al. (2002). No additional C sequestration was observed for alfalfa in 2013 and 2014 or for the lower middle chamber position, which acted neither as a net C source nor sink (Tab. 1; Fig. 5). This opposes the assumption of increased C sequestration by perennial grasses (Paustian et al., 1997) or perennial crops (Zan et al., 2001). However, NEE estimates of alfalfa were within the range of -100 to -400 g C m^{-2}, which is typical for forage crops (*Lolium*, alfalfa, etc.) in different agro-ecosystems (Bolinder et al., 2012; Byrne et al., 2005; Gilmanov et al., 2013; Zan et al., 2001). In addition, Alberti et al. (2010) reported a soil C loss of > 170 g C m^{-2} after crop conversion from continuous maize to alfalfa, concluding that no effective C sequestration occurs in the short-term.

Regardless of the crop type, the AC-derived dynamic ΔSOC values showed that up to 79 % of the standard deviation of estimated annual ΔSOC occurred during the growing season and the main plant growth period from the beginning of July to the end of September.

5. Conclusions

We confirmed that AC-based C budgets are in principle able to detect small-scale spatial differences and might be thus used to detect spatial heterogeneity of ΔSOC similar to the soil resampling method. However, compared to soil resampling AC-based C budgets also reveal short-term temporal dynamics. AC-derived C budgets showed not only pedon-scale differences but also pronounced temporal dynamics in ΔSOC (Fig. 5). In addition, AC-based ΔSOC values corresponded well with the tendencies and magnitude of the results observed in the repeated soil inventory. The period of maximum plant growth was identified as being most important for the development of spatial differences in annual ΔSOC. For upscaling purposes of the presented results, further environmental drivers, processes and mechanisms determining C allocation in space and time within the plant-soil system need to be identified. This type of an approach will be pursued in future within the CarboZALF experimental setup (Sommer et al., 2016; Wehrhan et al., 2016). Moreover, the AC-based C budget method opens new prospects for clarifying unanswered questions, such as the influence of plant development or erosion on ΔSOC.

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The soil profile is shown on the right. Soil horizon-specific SOC (%) and Nt (%) contents are indicated by solid and dashed vertical white lines, respectively. Spatial differences in ΔSOC and the basic principle of the C budget method are shown as the scheme within the picture.

Fig. 3.: Time series of CO_2 exchange (A-D) for the four chambers of the AC system during the study period from 2010 to 2014. R_{eco} (black), GPP (light gray) and NEE (dark gray) are shown as daily sums (y-axis). NEE_{cum} is presented as a solid line, representing the sum of continuously accumulated daily NEE values (secondary y-axis). The presented values display cumulative NEE following soil manipulation to the end of 2014. Note the different scales of the y-axes. The grey shaded area represents the period prior to soil manipulation. The dashed vertical line indicates the soil manipulation. Dotted lines represent harvest events.

Fig. 4.: Time series of modeled aboveground biomass development (NPP_{shoot}) (A-D) for the four chambers of the AC system during the study period from 2010 to 2014. NPP_{shoot} is shown as cumulative values. The presented values display cumulative NPP_{shoot} following soil manipulation to the end of 2014. The biomass model is based on biomass sampling (2010-2012) and biweekly LAI measurements (2013-2014) during crop growth (grey dots). C removal due to chamber harvests is shown by black dots. The grey shaded area represents the period prior to soil manipulation. The dashed vertical line indicates the soil manipulation. Dotted lines represent harvest events.

Fig. 5.: Temporal and spatial dynamics in cumulative ΔSOC throughout the study period based on (A) the C budget method (measured/modeled; black lines) and (B) the soil resampling method (linear interpolation; gray lines). The grey shaded area represents the period prior to soil manipulation. The dashed vertical line indicates the soil manipulation. Dotted lines represent harvest events. Temporal dynamics revealed by the C budget method allow for the identification

of periods being most important for the development of ΔSOC. Major spatial deviation occurred during the maximum plant growth period (May to September). The proportion (%) of these periods with respect to the standard deviation of estimated annual ΔSOC accounted for up to 79 %.

Fig. 6.: Average annual ΔSOC observed after soil manipulation (April 2011 to December 2014) by soil resampling and the C budget method for (A) the entire measurement site and (B) single chamber positions within the measured transect. ΔSOC represents the change in carbon storage, with positive values indicating C sequestration and negative values indicating C losses. Error bars display estimated uncertainty for the C budget method and the analytical error of ± 5 % for the soil resampling method. A perfomed Wilcoxon rank-sum test showed no significant difference between ΔSOC values obtained by both methodological approaches for all four chambers (p-value=0.25).

A.3.: Time series of recorded environmental conditions throughout the study period from 2010 to 2014. Daily Precipitation and GWL are shown for the upper (solid line) and lower (dashed line) chamber position in the upper panel (A). The lower panel (B) shows the mean daily air temperature. The grey shaded area represents the period prior to soil manipulation. The dashed vertical line indicates the soil manipulation.

Tab.1

Year	Crop rotation	Position	R_{eco}	GPP	NEE	ΔSOC (C budget)	NPP_{shoot}		NPP_{shoot}			SOC to 1 m depth	SOC in Ap horizon	ΔSOC (soil inventory)	Nt to 1 m depth	Nt in Ap horizon	Precip.	GWL
							harvested	modeled	N	P	K							
			(g C m^{-2})				(g C m^{-2})		(g m^{-2})			(Kg m^{-2} 1 m^{-1})	(Kg m^{-2} 0.3 m^{-1})	(g C m^{-2})	(Kg m^{-2} 1 m^{-1})	(Kg m^{-2} 0.3 m^{-1})	(mm)	(cm)
		A (upper)	1014 ±9	-1845 ±8	-831^{a} ±12	86 ±66	744	745^{a}±65	28.1	5.0	25.6	11.6	5.1		1.3	0.6	516	135
		B (upper middle)	987 ±11	-1970^{a} ±8	-983 ±13	251 ±66	727	732^{a} ±64	24.7	4.1	18.0	9.1	4.2		0.9	0.4		103
2010	**maize**	**C (lower middle)**	1064 ±38	-2000^{a} ±11	-935^{a} ±40	190 ±77	744	745^{a} ±65	25.5	4.2	16.9	9.1	4.2	-	0.9	0.4		95
		D (lower)	1110 ±21	-1737 ±10	-627^{a} ±23	-118 ±69	744	745^{a} ±65	25.0	4.2	18.2	12.8	5.0		1.3	0.5		69
		A (upper)	891 ±13	-2022 ±18	-1131^{a} ±22	-149 ±103	1238	1280^{a} ±101	29.5	5.4	30.2	10.5	3.5		1.1	0.4	618	129
		B (upper middle)	855^{a} ±10	-1894 ±13	-1039^{a} ±16	-169 ±96	1167	1208^{a} ±95	36.4	5.9	32.7	8.7	3.4		0.9	0.4		97
2011	**maize**	**C (lower middle)**	980 ±14	-2062 ±25	-1082 ±28	-79 ±95	1115	1161^{a} ±91	33.7	5.6	32.9	9.0	3.7		0.9	0.4		87
		D (lower)	843^{a} ±31	-1730 ±8	-888 ±32	-59 ±80	900	947^{a} ±73	35.0	5.7	31.8	12.2	4.0		1.3	0.4		61
		A (upper)	1058 ±86	-2659 ±12	-1600 ±87	648 ±104	297*/634	952^{a} ±56	36.3	6.3	42.6						585	139
	winter wheat	**B (upper middle)**	1075 ±8	-2591 ±11	-1516 ±13	472 ±65	310*/727	1044^{a} ±64	33.3	5.8	37.5							107
2012		**C (lower middle)**	1286 ±8	-2617 ±9	-1331 ±12	346 ±60	310*/665	985^{a} ±59	32.7	5.4	35.5		-	-				87
	sorghum	**D (lower)**	1044 ±10	-2194 ±9	-1150 ±13	430 ±39	299*/420	720^{a} ±37	33.9	5.8	40.4							61
		A (upper)	1140 ±83	-1583 ±9	-443 ±83	43 ±91	290	400^{a,b} ±37	14.0	1.7	11.6				-		499	154
		B (upper middle)	1283 ±80	-1819 ±8	-536 ±80	93 ±86	304	443^{b} ±32	14.7	1.8	12.1							122
2013		**C (lower middle)**	1438 ±20	-1726 ±7	-288 ±22	-107 ±36	324	395^{a} ±29	15.6	1.9	12.9							94
		D (lower)	1587 ±80	-2036 ±8	-448 ±80	6 ±87	329	442^{b} ±34	15.9	2.0	13.2							68
	alfalfa	**A (upper)**	1161 ±15	-1615 ±7	-455^{a} ±16	-126 ±26	605	581^{a} ±20	29.2	3.6	24.2	10.9	3.9	376	1.2	0.5	591	181
		B (upper middle)	1443 ±18	-2063 ±7	-619^{a} ±19	52 ±28	635	567^{a} ±20	30.7	3.8	25.4	8.9	3.5	156	0.9	0.4		149
2014		**C (lower middle)**	1683 ±18	-2111 ±6	-428 ±19	-36 ±26	632	535^{a} ±18	30.5	3.8	25.3	9.0	3.7	0	0.9	0.5		121
		D (lower)	1584 ±12	-2113 ±14	-528 ±19	-52 ±28	587	580^{a} ±21	28.3	3.5	23.5	12.5	4.2	276	1.3	0.4		95
		A (upper)	1063 ±49	-1970 ±12	-901 ±52	98 ±43	766	803 ±54	27.3	4.3	27.2			94 ±43				151
		B (upper middle)	1164 ±29	-2092 ±10	-919 ±32	104 ±37	786	815 ±53	28.8	4.3	26.9			39 ±43				119
annual average (2011-2014)		**C (lower middle)**	1347 ±15	-2129 ±12	-779 ±20	10 ±30	762	769 ±49	28.1	4.2	26.7		-	0 ±46	-		573	97
		D (lower)	1265 ±33	-2018 ±10	-739 ±38	67 ±32	634	672 ±41	28.3	4.3	27.2			69 ±47				71
		site	1209 ±32	-2052 ±11	-843 ±36	78 ±18	737	765 ±49	28.1	4.3	27.0			51 ±18				156

*NPP_{shoot} is based on biomass samples collected next to each chamber because no chamber harvest was performed for *winter fodder rye* in 2012; superscript letter indicate non-significant differences (Wilcoxon rank sum test; p-value > 0.05) between measured CO_2 fluxes and NPP_{shoot}.

Fig. 1

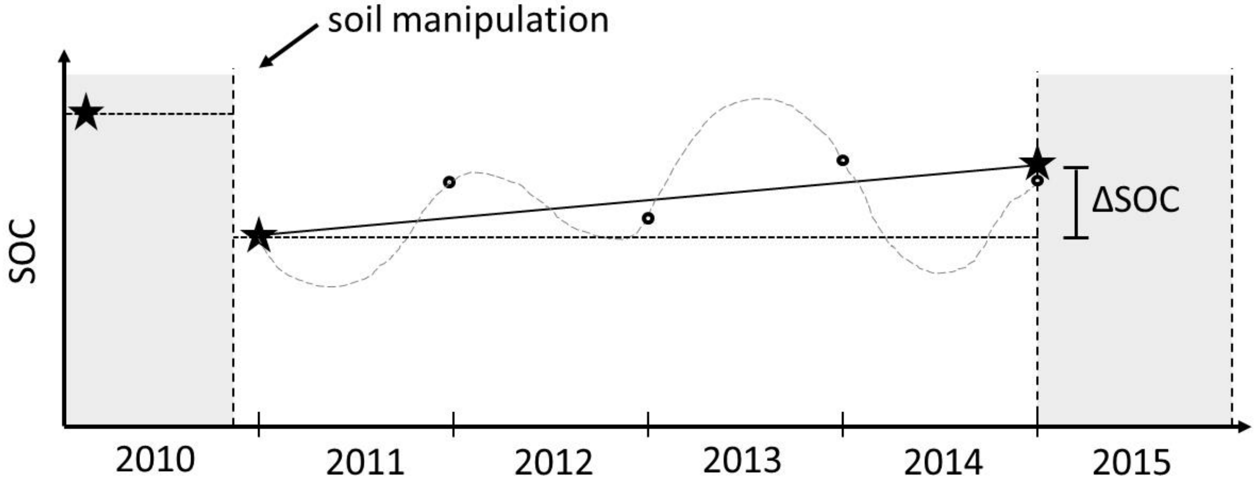
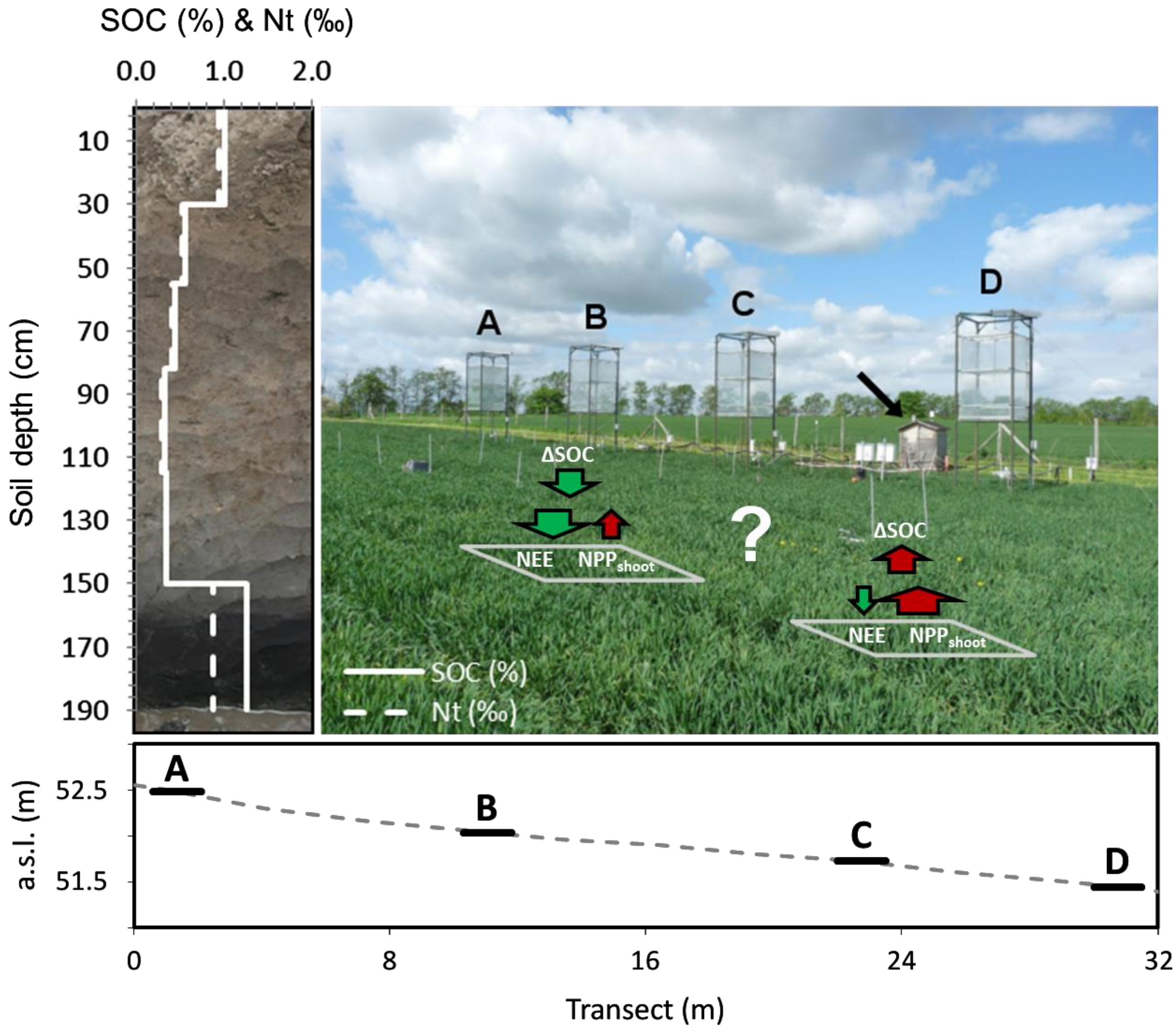


Fig. 2



SOC (%) & Nt (‰)
0.0
1.0
2.0
Soil depth (cm)
10
30
50
70
90
110
130
150
170
190
A
B
C
D
ΔSOC
NEE
NPPshoot
?
ΔSOC
NEE
NPPshoot
SOC (%)
Nt (‰)
a.s.l. (m)
52.5
51.5
A
B
C
D
0
8
16
24
32
Transect (m)

Fig. 3

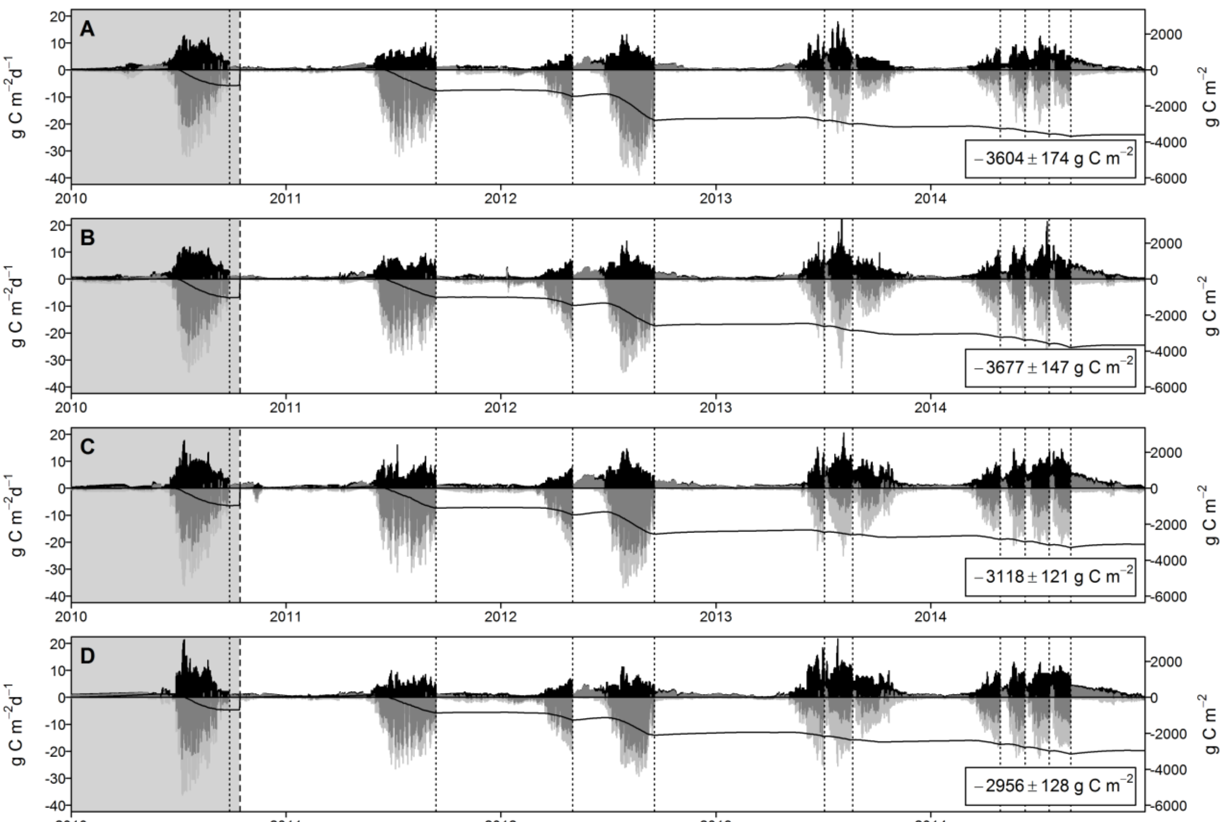
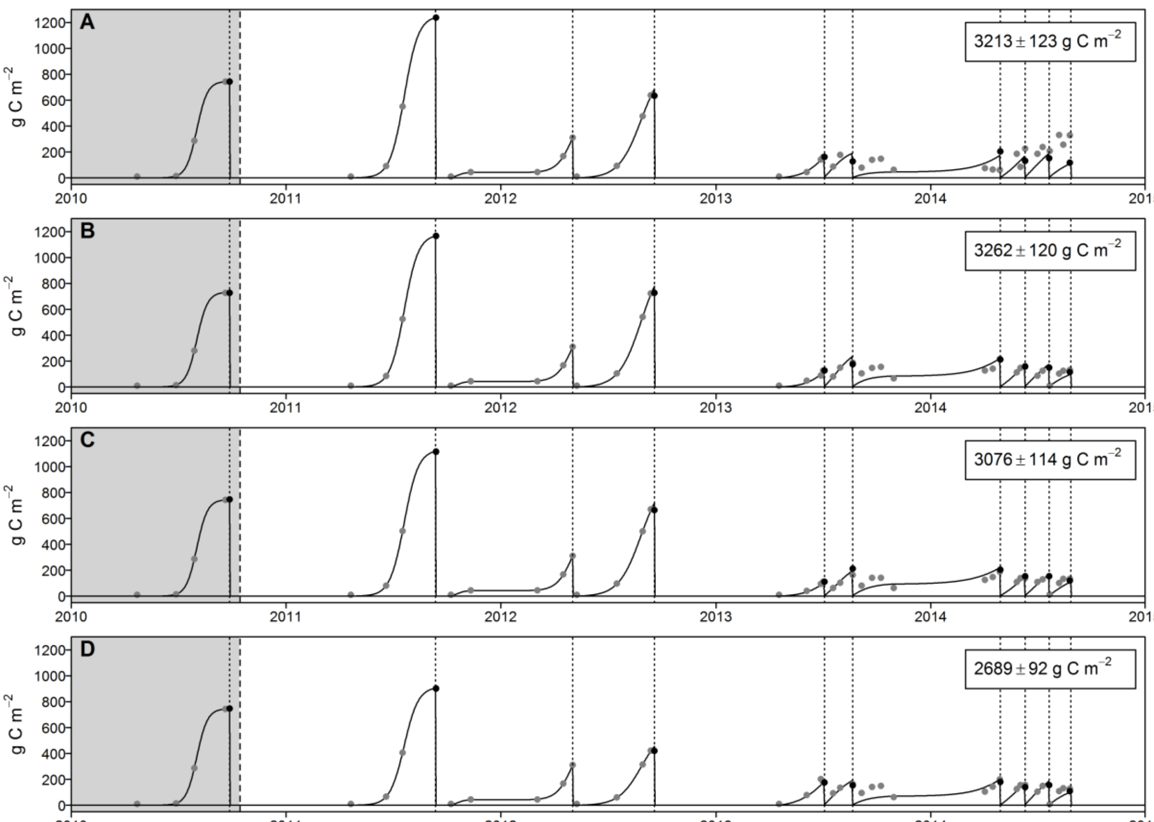

A
B
C
D
g C m−2 d−1
g C m−2
−3604 ± 174 g C m−2
−3677 ± 147 g C m−2
−3118 ± 121 g C m−2
−2956 ± 128 g C m−2
2010
2011
2012
2013
2014
20
10
0
-10
-20
-30
-40
2000
0
-2000
-4000
-6000

Fig. 4



A
3213 ± 123 g C m⁻²
B
3262 ± 120 g C m⁻²
C
3076 ± 114 g C m⁻²
D
2689 ± 92 g C m⁻²
g C m⁻²
0
200
400
600
800
1000
1200
2010
2011
2012
2013
2014
2015

Fig. 5

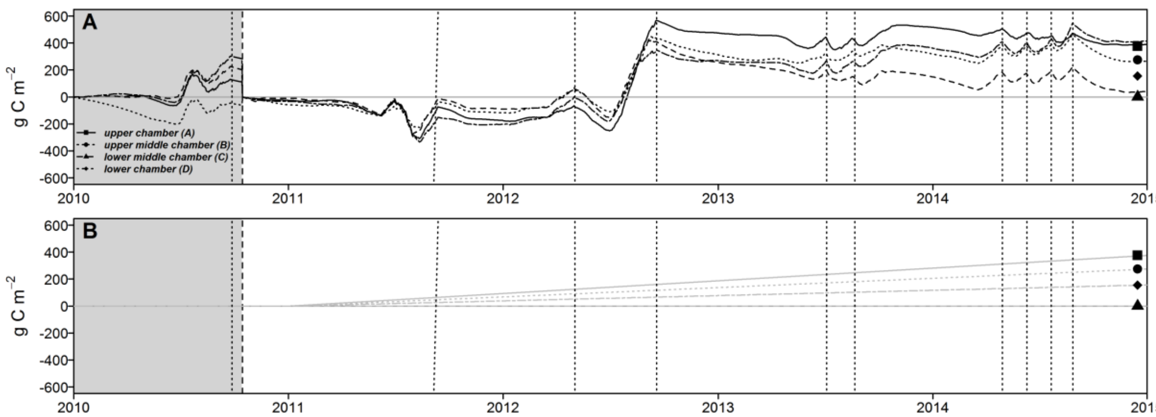
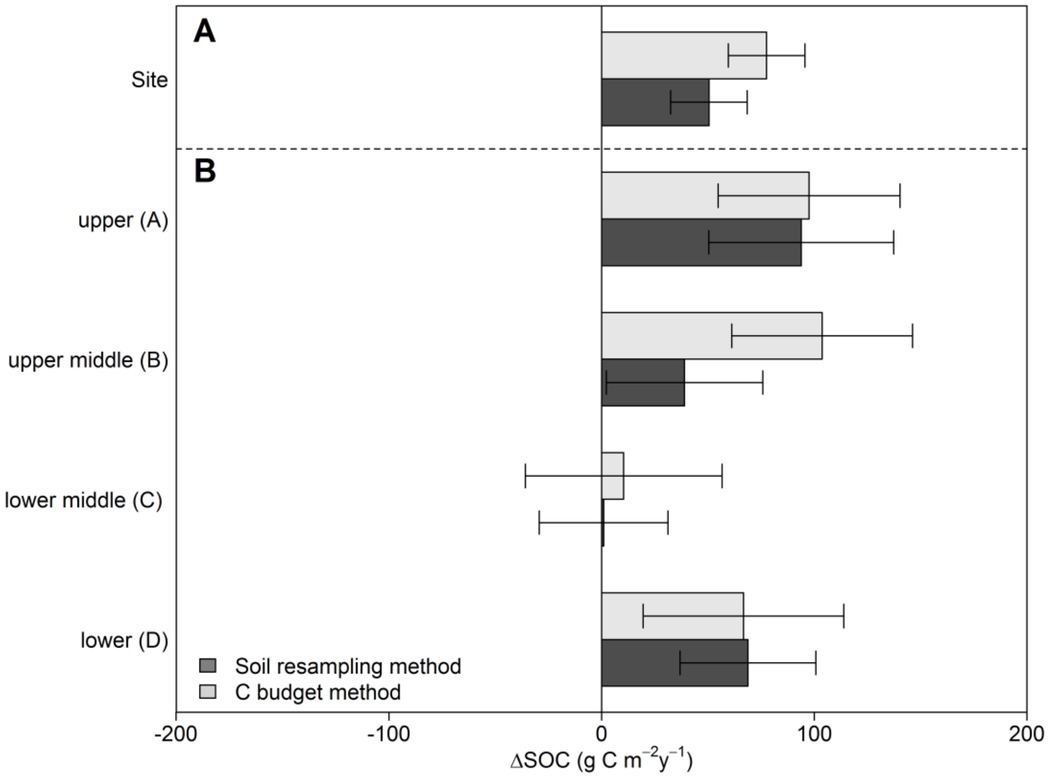

A
B
g C m^{-2}
600
400
200
0
-200
-400
-600
2010
2011
2012
2013
2014
2015
upper chamber (A)
upper middle chamber (B)
lower middle chamber (C)
lower chamber (D)

Fig. 6


A
B
Site
upper (A)
upper middle (B)
lower middle (C)
lower (D)
Soil resampling method
C budget method
-200
-100
0
100
200
ΔSOC (g C m−2y−1)

Appendices

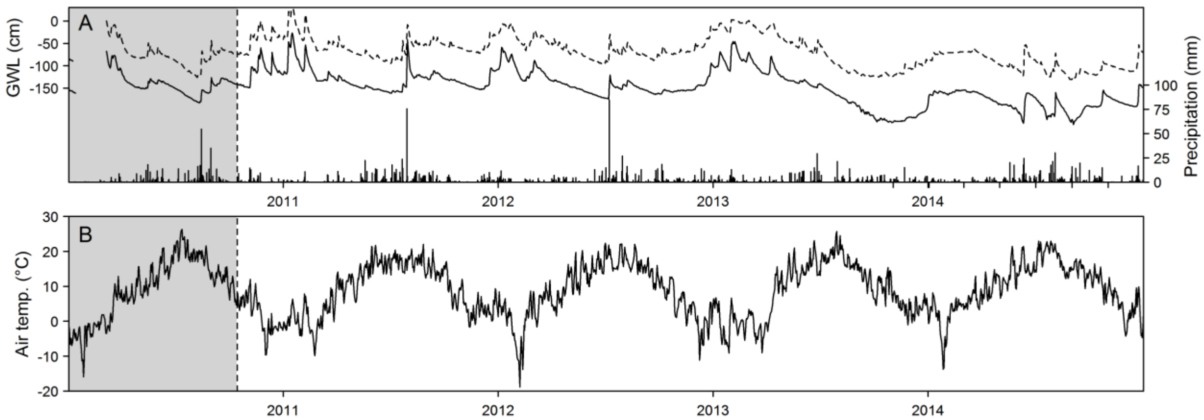
A.1

Crop	Treatment	Details	Date
	Chamber dismounting		10/04/2010
Winter fodder rye (*Secale cereale*)	Herbicide application	Roundup (2 l/ha)	19/04/2010
	Fertilization	KAS (160 kg/ha N), 110 kg/ha P2O5, 190 kg/ha K2O, 22 kg/ha S and 27 kg/ha MgO	23/04/2010
	Ploughing	Chisel Plough	23/04/2010
	Sowing	10 seeds/m^2	23/04/2010
Silage maize (*Zea mays*)	Chamber installation		04/05/2010
	Herbicide application	Zintan Platin Pack	26/05/2010
	Harvest		19/09/2010
	Chamber dismounting		20/09/2010
Bare soil	Chamber installation		27/10/2010
	Chamber dismounting		05/04/2011
	Fertilization	110 kg/ha P2O5, 190 kg/ha K2O, 22 kg/ha S and 27 kg/ha MgO	06/04/2011
	Ploughing	Chisel Plough	21/04/2011
	Sowing	10 seeds/m^2	21/04/2011
Silage maize (*Zea mays*)	Herbicide application	Gardo Gold Pack, 3.5 l/ha	27/04/2011
	Fertilization	KAS (160 kg/ha N)	03/05/2011
	Chamber installation		04/05/2011
	Harvest		13/09/2011
Bare soil	Chamber dismounting		13/09/2011
	Ploughing	Chisel Plough	30/09/2011
	Sowing	270 seeds/m^2	30/09/2011
Winter fodder rye (*Secale cereale*)	Chamber installation		05/10/2011
	Fertilization	KAS (80 kg/ha N)	06/03/2012
	Harvest		02/05/2012
Bare soil	Chamber dismounting		02/05/2012
	Ploughing		08/05/2012
	Sowing	30 seeds/m^2	09/05/2012
	Fertilization	KAS (100 kg/ha N), Kieserite (100 kg/ha), 220 kg/ha P2O5, 190 kg/ha K2O	14/05/2012
Sorghum-Sudan grass (*Sorghum bicolor* x *sudanese*)	Chamber installation		22/05/2012
	Replanting		29/05/2012
	Herbicide application	Gardo Gold Pack (3 l/ha), Buctril (1.5 l/ha)	12/07/2012
	Harvest		18/09/2012
Bare soil	Chamber dismounting		19/09/2012
	Ploughing	Chisel Plough	09/10/2012
	Sowing	400 seeds/m^2	09/10/2012
Winter triticale (*Triticosecale*)	Chamber installation		19/10/2012
	Chamber dismounting		20/09/2012
	Chamber installation		17/10/2012
	Ploughing; fertilization	Chisel Plough; 44 kg/ha K2O, 48.4 kg/ha P40	15/04/2013
	Sowing	22 kg/ha	18/04/2013
	Harvest (first cut)		04/07/2013
	Fertilization	88 kg/ha K2O	10/07/2013
	Harvest (second cut)		21/08/2013
Luzerne (*Medicago sativa*)	Fertilization	200 kg/ha K2O, 110 kg/ha P2O5	27/02/2014
	Harvest (first cut)		29/04/2014
	Harvest (second cut)		10/06/2014
	Harvest (third cut)		21/07/2014
	Harvest (fourth cut)		27/08/2014
	Chamber dismounting		28/08/2014

A.2 Weather and soil conditions

A.3 shows the development of important environmental variables throughout the study period (January 2010 – December 2014). In general, weather condition were similarly warm (8.7°C) but also wetter (562 mm) compared to the long-term average (8.6°C; 485 mm). Temperature and precipitation were characterized by distinct inter- and intra-annual variability. The highest annual air temperature was measured in 2014 (9°C). The highest annual precipitation was recorded during 2011 (616 mm). Lower annual mean air temperature and comparatively drier weather conditions were recorded in 2010 (7.7°C; 515 mm) and 2013 (8.5°C; 499 mm). Clear seasonal patterns were observed for air temperature. The daily mean air temperature at a height of 200 cm varied between -18.8°C in February 2012 and 26.3°C in July 2010. Rainfall was highly variable and mainly occurred during the growing season (55 % to 93 %), with pronounced heavy rain events during summer periods, exceeding 50 mm d^{-1}. Despite a rather wet summer, only 67 mm was measured in March and April 2012, the driest spring period within the study, resulting in late germination and reduced plant growth. Annual GWL differed by up to 77 cm along the chamber transect and followed precipitation patterns. Seasonal dynamics were characterized by a lower GWL within the growing season (1.10 m) and enhanced GWL during the non-growing season (0.85 m). From a short-term perspective, GWL was closely related to single rainfall events. Hence, a GWL of 0.10 m was measured immediately after a heavy rainfall event in July 2011, whereas the lowest GWL occurred during the dry spring in 2010. From August 2013 to December 2014, the GWL was too low to apply the principal of hydrostatic equilibrium; therefore, the groundwater table depth (> 235 cm) had to be used as a proxy.

925 A.


A
GWL (cm)
0
-50
-100
-150
Precipitation (mm)
100
75
50
25
0
2011
2012
2013
2014
B
Air temp. (°C)
30
20
10
0
-10
-20
2011
2012
2013
2014

926