



14 **Abstract**

15 Biochemical composition of particulate organic matter (POM) through phytoplankton
16 photosynthesis is important to determine food quality for planktonic consumers as well as
17 physiological conditions of phytoplankton. Major environmental factors controlling for the
18 biochemical composition were seasonally investigated in Gwangyang Bay which has only natural
19 conditions (e.g., no artificial dams) in South Korea. Water samples for the biochemical compositions
20 were obtained from three different light depths (100%, 30%, and 1%) mainly at 3 sites in Gwangyang
21 bay from April 2012 to April 2013. Different biochemical classes (carbohydrates [CHO], proteins
22 [PRT], and lipids [LIP]) were extracted and then the concentrations were determined by the optical
23 density measured with a spectrophotometer. The highest and lowest of PRT compositions among the
24 three biochemical classes were in April 2012 (58.0%) and August 2012 (21.2%), whereas the highest
25 and lowest LIP compositions were in August 2012 (49.0%) and April 2012 (24.8%), respectively.
26 CHO composition was recorded high in January 2013 and maintained above 25% during the study
27 period. The calorific contents of food material (FM) ranged from 1.0 Kcal m⁻³ to 6.1 Kcal m⁻³ (annual
28 mean ± S.D. = 2.8 Kcal m⁻³ ± 1.1 Kcal m⁻³). Based on Pearson's correlation coefficient analysis, a
29 major governing factor for biochemical composition of POM was dissolved inorganic nitrogen
30 loading from river-input in Gwangyang bay. In conclusion, relatively larger amount of FM and higher
31 calorific contents of POM found in this study compared to other regions reflected good nutritive
32 conditions for sustaining productive shellfish and fish populations in Gwangyang bay. Continuous
33 observations are needed for monitoring marine ecosystem response to potential environmental
34 perturbations in Gwangyang bay.

35

36 **Key words:**

37 Particulate organic matter, biochemical composition, phytoplankton, nitrogen source



38 1. Introduction

39 Particulate organic matter (POM) mostly from phytoplankton photosynthesis in the euphotic
40 layer is an important food source for planktonic consumers in water columns (Cauwet, 1978) and their
41 biochemical contents reaching the benthic environments are largely utilized by benthic organisms
42 (Nelson and Smith, 1986; Rice et al., 1994). Therefore, POM is an essential link between surface and
43 benthic ecosystems (Graf, 1992). Previous studies showed that the biochemical composition of the
44 POM such as protein (PRT), lipid (LIP) and carbohydrate (CHO) levels could provide useful
45 information on the nutritional value which is potentially available to consumers (Mayzaud et al, 1989;
46 Navarro et al., 1993; Navarro and Thompson, 1995). However, previous studies mainly focused on
47 the occurrence in the different patterns of biochemical composition of POM. It is noteworthy to
48 investigate how biochemical composition of POM responds to changes in various environmental
49 factors, such as nutrients, light, temperature, and salinity and to assess food quantity for higher trophic
50 levels.

51 The coastal areas represent one of the world's most vital aquatic resources, supporting and
52 providing food resources and habitats for large numbers of fish and shellfish species (Kwak et al.,
53 2012; Wetz and Yoskowitz, 2013; references therein). In Gwangyang bay, the southern coast of Korea
54 (Fig. 1), coastal fisheries and shellfish farming have been prevalence. Over the past decades, the bay
55 have become industrialized such as the construction of steel mill company, power plant and industrial
56 complex and environmental disturbances have been predicted. Also, estuaries have a high short-term
57 variability depending on many episodic events, such as freshwater inputs, tidal cycles (neap-spring),
58 and wind (storms) (Cloern and Nichols, 1985). These anthropogenic forces and environmental
59 changes drastically affect the estuarine habitat properties which can cause different biochemical
60 compositions of POM. Unfortunately, little information is yet available on the biochemical
61 composition of POM in the bay, South Korea. Hence, this study tested the question of the main
62 environmental factors determining seasonal variation and of biochemical composition POM and



63 assessed quantity of food material (FM) in the bay. Physical (temperature, salinity, irradiance, river-
64 input and rainfall data), chemical (nutrients), and biological (chlorophyll-*a* [chl-*a*], particulate organic
65 carbon [POC] and nitrogen [PON]) parameters were measured in order to both characterize the origin
66 of POM and understand their effects on the biochemical composition of POM. The aims of this study
67 were to: (1) investigate seasonal variation of biochemical composition of POM, (2) identify the origin
68 of POM, and (3) determine a major governing environmental factor for biochemical composition of
69 POM.

70 **2. Materials and methods**

71 **2.1. Study site and sampling procedure**

72 The study site was located in Gwangyang Bay (34.9 ° N, 127.8 ° E), the southern coast of
73 Korea (Fig. 1). The total area of the bay is 230 km² at mean sea level (Kang et al., 2003). The bay is
74 characterized by semidiurnal tides with a maximal range of about 4.8 m at spring tides (Korea
75 Hydrographic and Oceanographic Administration). Freshwater flows into the bay from the Seomjin
76 River at the northern part of the bay (mean flow 27 m³ s⁻¹ and annually 1.9 × 10⁹ t during the study
77 period; the National Institute of Environmental Research) and seawater enters through the narrow
78 southern channel (Yeosu Channel).

79 To obtain data for seasonal variation of POM in the euphotic depth, the field samplings were
80 undertaken at 3 stations of the bay (St.1 or St. 2A, St. 4, and St. 5; see Fig. 1) on a seasonal basis April,
81 June, August, and October in 2012 and January and April in 2013. St. 1 was changed to St. 2A after
82 April 2012 because of logistic problems. Both stations have similar environmental conditions at a
83 relatively close distance. Using a 5 L Niskin water sampler, water samples were collected at different
84 depths of 3 light intensities (100%, 30%, and 1% of surface irradiances; hereafter 3 light depths)
85 which were determined by a secchi disk and transferred to brown sample bottles which were
86 previously washed with a solution of 0.1 N HCl. To obtain *in situ* physical parameters, water



87 temperature and salinity were measured with YSI-30 (YSI incorporated) and photosynthetically active
88 radiation (PAR) was measured by a quantum sensor (LI-190SA, LI-COR) with a data logger (LI-1400,
89 LI-COR). Rainfall and river input data during the study period were obtained from the Korea
90 Meteorological Administration (<http://www.kma.go.kr/index.jsp>) and the National Institute of
91 Environmental Research (<http://water.nier.go.kr/main/mainContent.do>).

92 **2.2. Chl-*a* and major inorganic nutrient analysis**

93 In order to determine chl-*a* concentration, water samples from 3 light depths were filtered
94 through 25 mm GF/F (Whatman, 0.7 μm) which were kept frozen immediately and returned to the
95 laboratory at Pusan National University, Korea for a further analysis. The filters for chl-*a*
96 concentration were extracted in 90% acetone in a fridge (4 °C) for 24 h and centrifuged for 20
97 minutes at 4000 rpm. Using a fluorometer (Turner Designs, 10-AU) which had been calibrated with
98 commercially purified chl-*a* preparations, chl-*a* concentrations were measured and calculated (Parsons
99 et al., 1984). Water samples for inorganic nutrient concentrations from surface and bottom waters
100 were obtained from Niskin bottles. The samples were kept frozen (-70 °C) and sent for analysis to the
101 laboratory in the East Sea Fisheries Research Institute (QUAATRO, Seal Analytical).

102 **2.3. Particulate organic carbon and nitrogen analysis**

103 The water samples were filtered through pre-combusted (450 °C) 25 mm GF/F (Whatman,
104 0.7 μm). The filters for POC, PON, and $\delta^{13}\text{C}$ values were preserved frozen (-20 °C) and determined
105 using a Finnigan Delta + XP mass spectrometer at the stable isotope laboratory of the University of
106 Alaska Fairbanks, USA.

107 **2.4. Biochemical composition analysis**

108 The water samples for the biochemical composition (carbohydrates, proteins, and lipids) of
109 POM were collected from 3 light depths. The water samples were filtered through 47 mm GF/F



110 (Whatman, 0.7 μm pore), which were immediately frozen at $-70\text{ }^{\circ}\text{C}$ and preserved for biochemical
111 composition analysis at the home laboratory.

112 *Protein analysis*

113 Protein (PRT) concentrations were assessed according to a modified method of Lowry et al.
114 (1951). The filters for PRT analysis were transferred into 12 mL centrifuge tubes with 1 mL DH_2O ,
115 respectively. The filters were grounded (using a glass rod) in the tubes with a 5 ml alkaline copper
116 solution (a mixture of 2% Na_2CO_3 in 0.1 N NaOH with 0.5% $\text{CuSO}_4 \cdot 5\text{H}_2\text{O}$ in 1 % sodium or
117 potassium tartrate; 50:1, v/v). The solutions for PRT concentrations were mixed well (using a vortex)
118 and allowed to stand for 10 min at room temperature in the hood. After 10 min, 0.5 mL of diluted
119 Folin-Ciocalteu phenol reagent (1:1, v/v) was added into the solution, mixed occasionally with a
120 vortex mixer, and allowed to sit for 1 h 30 min. The solutions with a blue color were centrifuged at
121 3,000 rpm for 10 min. Absorbance of the supernatant was measured at 750 nm. Bovine Serum
122 Albumin (2 mg mL^{-1} , SIGMA) was used as a standard for the PRT concentration.

123 *Lipid analysis*

124 Lipid (LIP) concentrations were extracted according to a column method modified from
125 Bligh and Dyer (1959), and Marsh and Weinstein (1966). The filters for LIP analysis were transferred
126 into 16 mL glass tubes with 3 mL of chloroform-methanol (1:2, v/v). The filters in the tubes were
127 grounded, and then the mixtures were mixed using a vortex mixer. For LIP extraction, glass tubes
128 with samples were stored in the fridge ($4\text{ }^{\circ}\text{C}$) to prevent the solvents from evaporating. After 1 h, the
129 solvents were centrifuged at 2,000 rpm for 10 min and the supernatants were collected and stored in
130 new tubes. This extraction procedure was performed once again immediately. When the extractions
131 were completed, 4 mL of DH_2O was added to the solution in the new tubes, and the solution was
132 homogenized using a vortex mixer. After mixing, the tubes were centrifuged at 2,000 rpm for 10 min,
133 and the solvents were separated into two phases (the chloroform phase for lipids and methanol +



134 DH₂O phase). The methanol + DH₂O phase was removed from the solvent using a Pasteur pipette.
135 The chloroform phase was placed in a dry oven at 40 °C for 48 h. After it totally dried for
136 carbonization analysis (Marsh and Weinstein 1966), 2 mL of H₂SO₄ was added to the tubes and they
137 were placed in a heating block at 200 °C for 15 min. After this heating procedure, the tubes were
138 quickly placed in a water bath at room temperature; 3 ml of DH₂O was added to the tubes and the
139 solvents were homogenized (with a vortex mixer) and stood for 10 min or until all bubbles had
140 disappeared. Absorbance of the supernatant was measured at 375 nm. Tripalmitin solutions were used
141 as a standard for the LIP concentration.

142 *Carbohydrate analysis*

143 Carbohydrate (CHO) concentrations were measured according to Dubois et al. (1956). The
144 POM samples for carbohydrate analysis were transferred individually into 15 mL polypropylene (PP)
145 tubes. After 1 mL of DH₂O was added to the PP tubes, the samples were grounded using a glass rod.
146 One ml of 5 % phenol for CHO extraction was added additionally, and the solutions were allowed to
147 stand for 40 min at room temperature in the hood. After the extraction, 5 mL of sulfuric acid (H₂SO₄)
148 was added to the solutions, mixed using a vortex mixer, and allowed to stand for 10 min. The
149 solutions with an orange-yellow color were centrifuged at 3,500 rpm for 10 min. Absorbance of the
150 supernatant was measured at 490 nm using UV spectrophotometer (Labomed, Germany). D (+) -
151 glucose solutions (1 mg mL⁻¹, SIGMA) were used as a standard for the CHO concentration.

152 **2.5. Statistical analyses and calorific value calculation**

153 Statistical tests were carried out using the statistic software “SPSS” (*t*-test, ANOVA and
154 Pearson’s Correlation Coefficient). The level of significance was set at $p < 0.05$. The calorific value
155 (Kcal g⁻¹) of the food material (FM) (FM was defined by Danovaro et al. (2000); PRT + LIP + CHO
156 concentrations; hereafter FM) and the calorific content of FM (Kcal m⁻³ = Kcal g⁻¹ × g FM m⁻³) were
157 calculated using the Winberg (1971) equation.

158 **3. Results**159 **3.1. Seasonal distribution and variation of environmental factors and chl-*a* concentrations**

160 The values of environmental factors were summarized in Table 1. The temperature ranged
161 from 5.5 to 26.1 °C. The temperature increased from April to August (the highest temperature in
162 August 2012 at St. 4: 26.1 °C) and decreased from August to January (the lowest temperature in
163 January 2013 at St. 2A: 5.5 °C). The salinity ranged from 14.5 to 32.9 ‰. Generally, the salinity
164 increased from St. 1 or St. 2A to St. 5. Relatively lower salinity, which is mainly affected by fresh
165 water input from the Seomjin River, was observed at St. 2A. The annual average euphotic depth was
166 6.5 ± 3.4 m, ranging from 2 to 12 m.

167 The highest nutrient concentrations were measured in April 2012, when the concentrations of
168 $\text{NO}_2 + \text{NO}_3$, SiO_2 , NH_4 , and PO_4 were above 5.0 μM , 2.0 μM , and 0.2 μM , respectively, except at 1%
169 light depth at St. 4. All inorganic nutrients except SiO_2 were nearly depleted in August 2012 (Table 1).
170 During the rest of our study period, $\text{NO}_2 + \text{NO}_3$ and SiO_2 concentrations were observed with similar
171 decreasing patterns from St.1 or St. 2A to St. 5. NH_4 concentrations averaged from October 2012 to
172 April 2013 were $1.1 \mu\text{M} \pm 0.4 \mu\text{M}$, ranging from 0.5 μM to 1.9 μM . PO_4 concentrations (average \pm
173 S.D. = $0.1 \pm 0.1 \mu\text{M}$) ranged from 0 to 0.4 μM throughout the water columns at all stations except at
174 St. 2A in April 2012 during the study period.

175 Monthly rainfall and river-input in the study location ranged from 15.6 mm to 559.0 mm
176 (annual mean \pm S.D. = $151.0 \text{ mm} \pm 155.5 \text{ mm}$) and 42.3 to $447.2 \times 10^6 \text{ t}$ (annual mean = $144.4 \times 10^6 \text{ t}$),
177 respectively. Rainfall and river-input were recorded as high during summer and low during winter
178 (Table 2). Average irradiance during our incubation hour ranged from 167.9 ± 133.5 to 1593.3 ± 414.5
179 $\mu\text{mol m}^{-2} \text{ s}^{-1}$ (average \pm S.D.) from April 2012 to April 2013. The highest and lowest irradiance were
180 recorded in April 2013 and April 2012, respectively.

181 Chl-*a* concentrations in the euphotic depth ranged from 0.8 $\mu\text{g L}^{-1}$ to 14.2 $\mu\text{g L}^{-1}$ during the



182 study period (annual mean \pm S.D. = $3.4 \mu\text{g L}^{-1} \pm 2.8 \mu\text{g L}^{-1}$; Table 1). There were no significant
183 differences of chl-*a* concentrations among 3 light depths and spatial distribution. However, there was
184 seasonal variation of chl-*a* concentrations during study period. Chl-*a* concentrations were increased
185 from April to August and decreased from August to October in 2012 and increased slightly again in
186 January and April 2013.

187 3.2. $\delta^{13}\text{C}$ values and carbon to nitrogen ratios of POM

188 $\delta^{13}\text{C}$ values of sea surface POM ranged from - 23.1 ‰ to - 16.5 ‰ and the annual average
189 $\delta^{13}\text{C}$ value was -20.9 ‰ (S.D. = ± 3.2 ‰). The annual average carbon to nitrogen (C:N) ratio of POM
190 was 7.0 ± 0.4 (average \pm S.D.), ranging from 6.8 to 7.7 (Table 3).

191 3.3. Seasonal variation of biochemical composition

192 The contents of CHO, PRT, and LIP of POM in the water column ranged from $14.2 \mu\text{g L}^{-1}$ to
193 $412.3 \mu\text{g L}^{-1}$ ($129.5 \pm 87.2 \mu\text{g L}^{-1}$), from $22.8 \mu\text{g L}^{-1}$ to $382.4 \mu\text{g L}^{-1}$ ($155.0 \pm 73.3 \mu\text{g L}^{-1}$), and from
194 $21.4 \mu\text{g L}^{-1}$ to $401.4 \mu\text{g L}^{-1}$ ($154.9 \pm 78.9 \mu\text{g L}^{-1}$), respectively (Table 4). The FM contents of POM
195 ranged from $170.9 \mu\text{g L}^{-1}$ to $915.7 \mu\text{g L}^{-1}$ ($435.5 \pm 175.5 \mu\text{g L}^{-1}$). Since there were no significant
196 differences in biochemical concentrations of POM and FM among 3 light depths and spatial
197 distributions, we averaged each biochemical compound and FM on monthly basis. The CHO and LIP
198 concentrations increased from April to August and decreased from August to October in 2012. In
199 contrast, the PRT concentrations decreased from April to October in 2012 and increased from October
200 in 2012 to April in 2013. The seasonal pattern of FM concentrations was similar to the pattern of chl-*a*
201 concentrations.

202 In order to estimate the biochemical composition as food quality, we obtained relative
203 contributions of each biochemical concentration of POM to FM, based on percentage basis. The
204 biochemical composition of each class (CHO, PRT and LIP) ranged from 8.3% to 59.1%, from 6.8%
205 to 74.9% and from 9.4% to 68.3%, respectively (annual mean \pm S.D. of CHO, PRT, and LIP



206 composition = $26.4 \pm 9.4\%$, $37.8 \pm 16.1\%$, and $35.7 \pm 13.9\%$, respectively; Table 4). We found the
207 seasonal variation of biochemical composition based on monthly basis of biochemical composition
208 (Fig. 2). To illustrate these variations of biochemical composition of POM, the highest and lowest
209 PRT compositions were in April 2012 and August 2012. In contrast to PRT compositions, the highest
210 and lowest LIP compositions were in August 2012 and April 2012. CHO composition was recorded
211 high in January 2013, but to compare CHO composition to PRT and LIP composition, CHO
212 composition was not strong varied during the study period.

213 3.4. Seasonal variations of the calorific values and contents of FM

214 The calorific values and contents of FM ranged from 5.4 Kcal g^{-1} to 7.9 Kcal g^{-1} (annual
215 mean \pm S.D. = $6.6 \text{ Kcal g}^{-1} \pm 0.6 \text{ Kcal g}^{-1}$) and 1.0 Kcal m^{-3} to 6.1 Kcal m^{-3} (annual mean \pm S.D. = 2.8
216 $\text{Kcal m}^{-3} \pm 1.1 \text{ Kcal m}^{-3}$), respectively (Table 4). The calorific values of FM had no apparent seasonal
217 pattern, whereas the calorific contents of FM had a seasonal pattern similar to the seasonal variation
218 of FM concentrations.

219 3.5. Relationship between biochemical pools and environmental conditions

220 Relationships between biochemical pools and environmental conditions were performed
221 using Pearson's correlation matrix. Based on the results, we found a significant, positive relationships
222 between PRT composition and river-input ($r = 0.84$, $p < 0.01$, Table 5, Fig. 3) and PRT composition
223 and dissolved nitrogen concentrations (NH_4 : $r = 0.69$, $p < 0.01$; $\text{NO}_2 + \text{NO}_3$: $r = 0.54$, $p < 0.01$, Table
224 5). Lipid composition had an inverse relationships with river-input ($r = -0.63$, $p < 0.01$) and dissolved
225 nitrogen concentrations (NH_4 : $r = -0.59$, $p < 0.01$; $\text{NO}_2 + \text{NO}_3$: $r = -0.53$, $p < 0.01$). These
226 relationships led to a significant reverse relationship between PRT composition and LIP composition
227 ($r = -0.81$, $p < 0.01$, Fig. 4). PRT composition was negatively correlated with temperature ($r = -0.52$, p
228 < 0.01), whereas LIP composition was positively correlated with temperature ($r = 0.72$, $p < 0.01$).
229 There were no significant relationships between PRT composition and irradiance and LIP composition



230 and irradiance.

231 4. Discussion

232 4.1. Environmental conditions and chl-*a* concentration

233 The annual average chl-*a* concentration during the research period was $3.4 \mu\text{g L}^{-1}$ (S.D.=
234 $\pm 2.8 \mu\text{g L}^{-1}$) with a range from 0.8 to $14.2 \mu\text{g L}^{-1}$ which is in a similar range of chl-*a* concentrations
235 reported previously in Gwangyang bay, although it varied across different seasons and sampling
236 depths (Cho et al., 1994; Choi et al., 1998; Lee et al., 2001; Kwon et al., 2001; Jang et al., 2005; Yang
237 et al., 2005; Beak et al., 2011; Min et al., 2011; Beak et al., 2015). Previous studies reported that chl-*a*
238 concentration was influenced mainly by salinity, temperature, and nutrients (nitrate and phosphate)
239 depending on freshwater input from the Seomjin River. Our results in this study were similar to
240 former studies ($r = 0.34$ and -0.41 , $p < 0.05$, $n = 48$ and 28 for salinity and NH_4 , respectively).
241 However, high chl-*a* concentrations were previously recorded in spring and fall, whereas the highest
242 concentrations were observed in summer (August 2012) from this study. In fact, Baek et al. (2015)
243 reported that high chl-*a* concentrations were found in summer similarly, although there was difference
244 between environmental factors and chl-*a* concentrations as compared with our results. The high levels
245 of chl-*a* were observed with high nutrient concentrations and low salinity levels in the surface water
246 by Baek et al. (2015), whereas the high values existed with low nutrient concentrations and high
247 salinity levels in our results.

248 Despite this dissimilarity of environmental factors with high chl-*a* concentrations, we also
249 found the highest chl-*a* concentrations observed in summer. According to Shaha and Cho (2009),
250 there is a tendency with increasing precipitation and river-input in Gwangyang Bay during summer.
251 This trend could increase loading nutrients from freshwater for maintaining phytoplankton growth in
252 summer. In addition, a strong light intensity during summer could be favorable for phytoplankton
253 growth since our study area was extremely turbid conditions during almost all seasons due to



254 freshwater discharge and a strong spring-neap tidal oscillation. As a result, the combination of these
255 factors is believed to enhance chl-*a* concentration and primary production of phytoplankton during
256 summer in Gwangyang Bay.

257 **4. 2. POM characterization**

258 In general, POM consists of a mixture of living as well as detritus materials (phytoplankton,
259 bacteria, zooplankton, fecal pellets, terrestrial matters, etc.) originating from freshwater and estuarine
260 and marine environments. POM samples can be characterized or determined for source of the major
261 contributor(s). The C:N ratio generally ranges between 6 and 10 for phytoplankton, whereas the ratios
262 are between 3 and 6 for zooplankton and bacteria (Savoye et al, 2003; references therein). For
263 terrestrial organic matters, the C:N ratios are normally over 12 (Savoye et al, 2003; references therein).
264 Therefore, it is useful to classify phytoplankton from heterotrophs and terrestrial materials (Lobbes et
265 al., 2000; Savoye et al., 2003; Lee and Whitledge, 2005). In this study, the mean C:N ratios of POM
266 was 7.0 (S.D. = ± 0.4), which indicates that this POM is mainly phytoplankton (Table 3). However,
267 the C:N ratio must be used with caution because of its variation in the process of POM degradation
268 (Savoye et al, 2003). For example, PON is preferentially degraded compared to POC of
269 phytoplankton, which causes an increase of the C:N ratio. Terrestrial organic matters (high C:N ratio)
270 colonized by bacteria (low C:N ratio) lowers their initial high C:N ratio (Savoye et al, 2003;
271 references therein). Therefore, similar C:N ratios of POM could be produced by degraded
272 phytoplankton and bacteria-colonized terrestrial organic matters (Lancelot and Billen 1985; Savoye et
273 al, 2003). In addition to C:N ratios, $\delta^{13}\text{C}$ of POM can be used for determining their origin. Kang et al.
274 (2003) reported that the mean $\delta^{13}\text{C}$ signature of phytoplankton in Gwangyang Bay was -20.8‰ (S.D.
275 = $\pm 1.1\text{‰}$). In this study, our mean $\delta^{13}\text{C}$ signature of POM was -20.9‰ (S.D. = $\pm 3.2\text{‰}$), which also
276 indicates that POM was mostly phytoplankton during the study periods (Table 3). Based on our C:N
277 ratio and $\delta^{13}\text{C}$ value in this study, we confirmed that our POM samples were primarily comprised of
278 phytoplankton in Gwangyang Bay.



279 4. 3. Environmental conditions and biochemical pools

280 Biochemical pools of POM originating from phytoplankton are influenced by various
281 environmental factors, such as temperature, salinity, nutrients, and light conditions (Morris et al., 1974;
282 Smith and Morris, 1980; Rivkin and Voytek, 1987; Boëchat and Giani, 2008; Cuhel and Lean, 1987;
283 Mock and Kroon, 2002; Khotimchenko and Yakoleva, 2005; Ventura et al., 2008; Sterner et al. 1997).
284 In this study, significant relationships were found between environmental conditions and biochemical
285 pools, especially PRT and LIP (Table 5). Temperature was positively and negatively correlated with
286 LIP and PRT. Previous studies reported that higher temperature stress mainly affects nitrogen
287 metabolism (Kakinuma et al., 2006) which is related to significant decrease of PRT with increases of
288 LIP and CHO content (Tomaselli et al., 1988; Oliveira et al., 1999). In a high temperature-stressed
289 condition of phytoplankton, the decrease in PRT content is related to breakdown of protein structure
290 and interference with enzyme regulators (Pirt, 1975), whereas LIP is predominant because LIP is more
291 closely associated with cell structure such as thickened cell walls (Smith et al., 1989; Kakinuma et al.,
292 2001, 2006). Our results are in agreement with other works, as described above.

293 The relationships between nutrients and biochemical pools could be explained by nutrient
294 limitation and the characteristics of each biochemical compound. A combination of nutrient
295 concentrations and ratios can be used to assess nutrient limitation (Dortch and Whitedge, 1992; Justić
296 et al., 1995). Dortch and Whitedge (1992) suggested that nutrient limitations are existed in the
297 Mississippi river plume and Gulf of Mexico, if the dissolved inorganic phosphorus (DIP), dissolved
298 inorganic nitrogen (DIN), and dissolved silicon (DSi) concentrations in water column are less than 0.2,
299 1.0 and 2.0 μM , respectively. In addition, molar ratios of the DIN:DIP and DSi:DIN can be indicators
300 of nutritional status and the physiological behavior of phytoplankton (Redfield et al., 1963; Goldman
301 et al., 1979; Elrifi and Turpin, 1985; Roelke et al. 1999). The following criteria of their molar ratios
302 were (a) DSi:P ratio >16 , and DSi:N ratio >16 for phosphorus (P) limitation; (b) DSi:DIP ratio >16
303 and DSi:N ratio >1 for nitrogen (N) limitation; (c) DSi:DIP ratio <16 and DSi:DIN ratio <1 for



304 silicate (Si) limitation. In this study, nutrient limitation conditions were observed by absolute nutrient
305 concentrations or/and their molar ratios depending on seasons (Table 6). Previous studies of
306 biochemical composition in relation to nutrient limitation reported that PRT production of
307 phytoplankton was enhanced under abundant N conditions (Fabiano et al., 1993; Lee et al., 2009). In
308 contrast, LIP production and storage were dominant (Shifrin and Chisholm, 1981; Harrison et al.,
309 1990) and PRT contents decreased (Kilham et al., 1997; Lynn et al., 2000; Heraud et al., 2005) under
310 N-depleted conditions. High LIP contents have also been detected in phytoplankton under P or/and Si
311 limitation (Lombardi and Wangersky, 1991; Lynn et al. 2000; Heraud et al., 2005; Sigeo et al., 2007).
312 Under N or P-limited conditions, triglyceride content (energy storage) increases and shifts from PRT
313 to LIP metabolism since proteins are nitrogenous compounds whereas LIP and CHO are non-
314 nitrogenous substrates (Lombardi and Wangersky, 1991; Smith et al., 1997; Takagi et al., 2000). In
315 our study, Si and P concentrations may not significantly impact on biochemical composition of
316 phytoplankton. Si concentrations were almost above 2.0 μM except in April 2013 during the study
317 period. P limitation was observed based on the absolute concentration and molar ratios during study
318 period. However, under P limitation, phytoplankton can relocate the cellular P pool to maintain their P
319 requirements for the maximum growth rate (Cembella et al., 1984; Ji and Sherrell, 2008). In this
320 respect, we suggest that DIN could be significantly impact on biochemical composition of
321 phytoplankton in our study area. DIN was initially believed to be the most important limiting factor
322 for phytoplankton growth in marine ecosystems (Ryther and Dunstan, 1971; Howarth, 1988). In fact,
323 DIN was strongly positively correlated with PRT composition, whereas it was negatively correlated
324 with LIP composition. The most of DIN loading came from freshwater input of the Seomjin River
325 (Table 5, river-input vs NH_4 and NO_2+NO_3 ; $r = 0.91$ and 0.55 , $p < 0.01$, respectively) influences on
326 PRT and LIP synthesis and subsequently macromolecular composition of phytoplankton. As a result,
327 the amount of river-input was also strongly correlated with PRT composition (Table 5 and Fig. 3).
328 Therefore, DIN is an important controlling factor for biochemical composition, especially PRT and
329 LIP composition of phytoplankton in Gwangyang bay.



330 Although irradiance is also known for an important governing factor for biochemical
331 composition, irradiance was not significantly correlated with biochemical pools in this study (Table 5).
332 We measured irradiance during our incubation time (4~5h) for phytoplankton productivity. This short
333 time of measured irradiance might be not enough to reflect and detect the change of biochemical
334 composition in phytoplankton with irradiance

335 The structure and composition of phytoplankton assemblages and species could have a
336 significant influence on the seasonal variation of biochemical composition. Although we did not
337 conduct a study of phytoplankton community structure, there is seasonal succession of phytoplankton
338 community structure in the bay. Previous studies showed that the dominant phytoplankton community
339 was diatoms and dominant diatom species were *Skeletonema spp.* during summer and winter in
340 Gwangyang bay (Choi et al., 1998; Baek et al., 2015). Kim et al. (2009) also reported that diatom and
341 dinoflagellate communities have experienced a considerable change because of increased nutrient
342 loadings from both domestic sewage and industrial pollution during summer. Therefore, the seasonal
343 change of phytoplankton species composition and community structure could lead to determining
344 different biochemical pools on seasonal basis.

345 However, other studies in different regions reported that environmental conditions, such as
346 temperature, nutrients and irradiance are more important controlling factors in biochemical
347 composition than variation of phytoplankton community and species composition (Lindqvist and
348 Lingnell, 1997; Suárez and Marañón, 2003). In this study, we also concluded that DIN from river-
349 input was a primary governing factor for the seasonal variation of biochemical composition of
350 phytoplankton in Gwangyang Bay as discussed above.

351 **4.4. Total FM and energy content of POM in a global context**

352 The annual average of FM was $434.5 \mu\text{g L}^{-1}$ (S.D. = $\pm 175.5 \mu\text{g L}^{-1}$) in this study. Since there
353 were no comparable data available in South Korea, we compared our results with other regions (Table



354 7), although they were conducted in different seasons and sampling depths. PRT contents in this study
355 were as high as in the Ross Sea (Fabiano and Pusceddu, 1998; Fabiano et al., 1999a), the Amundsen
356 Sea (Kim et al., 2015) and the Humboldt Current System (Isla et al., 2010). A similar range of LIP
357 contents was observed in Bedford Basin (Mayzaud et al., 1989), Yaldad Bay (Navarro et al., 1993)
358 and the Humboldt Current System (Isla et al., 2010). CHO contents were comparatively higher in this
359 study than other studies except Bedford Basin (Mayzaud et al., 1989) and Yaldad Bay (Navarro et al.,
360 1993). One of the highlights is that the calorific contents of FM were generally higher than those of
361 other areas except several regions. The FM values were comparatively higher than other regions such
362 as the northern Chuckchi Sea (Kim et al., 2014; Yun et al., 2014), Ross Sea (Fabiano et al., 1996;
363 Fabiano and Pusceddu, 1998; Fabiano et al., 1999a; Pusceddu et al., 1999), Amundsen Sea (Kim et al.,
364 2015) and the northern part of the East/Japan Sea (Kang et al., unpublished) or similar to the
365 Humboldt Current System which is known as an important spawning sites for pelagic fishes and the
366 highest abundance of anchovy eggs (Isla et al., 2010). Actually, the southern coastal sea (including our
367 study area) in Korea represents calm seas, an indented coastline, and numerous bays, which have
368 high diversities of habitat for fishes and shellfishes (Kwak et al., 2012) and give a favorable condition
369 for mariculture (Kwon et al., 2004). The high quantity of FM and the calorific contents of POM found
370 in this study reflected good nutritive conditions of primary food materials mainly provided by
371 phytoplankton for the maintenance of productive shellfish and fish populations in Gwangyang bay.

372 **5. Summary and Conclusion**

373 This study is the first report that was investigated the biochemical composition of POM on
374 seasonal basis in Gwangyang Bay, South Korea and we determined major controlling factors for
375 biochemical composition which is influenced by various environmental factors (Morris et al., 1974;
376 Smith and Morris, 1980; Rivkin and Voytek, 1987; Boëchat and Giani, 2008; Cuhel and Lean, 1987;
377 Mock and Kroon, 2002; Khotimchenko and Yakoleva, 2005; Ventura et al., 2008; Sterner et al. 1997).
378 Among different factors, temperature was positively correlated with LIP whereas negatively



379 correlated with PRT in this study (Table 5), which is consistent with previous works. In addition, we
380 found that PRT and LIP compositions were strongly correlated with DIN loading largely depending on
381 the amount of river-input from the Seomjin river which influences on PRT and LIP synthesis and
382 subsequently macromolecular composition of phytoplankton in Gwangyang bay. The concentrations
383 and the calorific contents of FM found in this study were relatively higher in comparison to previous
384 studies in various regions, which reflecting that POM (mainly from phytoplankton) provides a good
385 nutritive condition to maintain this highly productive estuarine ecosystem in Gwangyang bay.

386 Recently, significant environmental perturbations in their watersheds and externally from
387 climatic forcings have been reported in various estuaries (Wetz and Yoskowitz, 2013). More intense
388 but less frequent tropical cyclones are expected over the coming century (e.g., Elsner et al., 2008;
389 Knutson et al., 2010) and many changes in drought and flood cycles have been proceeding globally
390 (e.g., Min et al., 2011; Pall et al., 2011; Trenberth and Fasullo, 2012; Trenberth, 2012). The
391 cumulative effects of these perturbations could alter the quantity and quality of biochemical
392 composition of POM and cause subsequent changes in ecosystem structure and trophic dynamics in
393 estuaries (Cloern, 2001; Paerl et al., 2006; Rabalais et al., 2009; Wetz and Yoskowitz, 2013).
394 Therefore, continuous field measurements and observations on biochemical composition of POM as
395 food quality are needed to monitor for better understanding future response of marine ecosystem on
396 potential environmental perturbations in Gwangyang Bay.

397

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656 **Table captions**

657 Table 1. Environmental factors and chl-*a* concentrations in Gwangyang bay during the research period

658 (- : no data).

659 Table 2. Rainfall and river input.

660 Table 3. $\delta^{13}\text{C}$ values and C:N ratios of POM in Gwangyang bay (surface).

661 Table 4. Biochemical concentrations and composition, calorific values and contents in Gwangyang

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663 Table 5. Significant correlation coefficient (*r*) among proteins (PRT), lipids (LIP) and environmental

664 factors (ns ; no significance, **: $p < 0.01$).

665 Table 6. Observed nutrient limitations during the study period.

666 Table 7. Comparison of biochemical quantity of POM, FM and the calorific contents.



667 **Figure captions**

668 Fig. 1. Sampling location in Gwangyang bay, Korea ; Maps of Korea (a), Southern Coastal Sea (b)

669 and main sampling stations (c).

670 Fig. 2. Seasonal variation of biochemical composition in Gwangyang bay.

671 Fig. 3. The positive relationship between river-input and protein composition.

672 Fig. 4. The inverse relationship between lipid compositions and protein compositions.

673



674

Table 1. Environmental factors and chl-*a* concentrations in Gwangyang Bay during the research period (- : no data).

675

Year	Date	Irradiance ($\mu\text{mol m}^{-2} \text{s}^{-1}$)	Station	Light depth (%)	Temperature (°C)	Salinity (‰)	Depth (m)	NH ₄ (μM)	NO ₂ +NO ₃ (μM)	SiO ₂ (μM)	PO ₄ (μM)	Chl- <i>a</i> ($\mu\text{g L}^{-1}$)	
2012	April	167.9 ± 133.5 (average ± S.D.)	St.1	100	13.9	14.5	0	3.6	56.4	26.0	80.9	1.89	
				30	13.3	25.6	1	-	-	-	-	1.95	
				1	13.5	28.0	3	2.4	16.0	9.8	0.2	2.08	
			St.4	100	15.0	24.4	0	2.6	15.1	16.3	0.2	1.81	
				30	13.6	31.4	1	-	-	-	-	-	
				1	12.3	32.9	5	1.9	2.1	2.1	0.1	2.03	
	St.5	100	12.6	31.7	0	3.1	9.5	7.1	0.3	2.07			
		30	12.3	31.6	1	-	-	-	-	-			
		1	12.2	32.4	5	3.0	6.4	5.1	0.3	2.04			
	June	1158.1 ± 627.6	St.2A	100	22.9	27.6	0	-	-	-	-	-	1.77
				30	22.8	27.6	1	-	-	-	-	-	0.76
				1	22.9	28.7	3	-	-	-	-	-	0.76
			St.4	100	23.6	31.5	0	-	-	-	-	-	1.00
				30	22.6	31.9	3	-	-	-	-	-	1.67
				1	22.1	32.3	11	-	-	-	-	-	1.02
	August	1320.0 ± 316.9	St.4	100	25.8	30.6	0	0.1	0.1	10.6	0.1	0.1	8.11
				30	25.7	31.6	2	-	-	-	-	-	8.49
				1	25.7	31.7	8	0.1	0.1	11.9	0.1	0.1	5.99
St.5			100	25.6	31.6	0	0.7	0.3	8.2	0.0	0.0	14.20	
			30	26.1	31.5	2	-	-	-	-	-	9.85	
			1	25.7	31.7	8	0.1	0.1	10.1	0.1	0.1	3.19	
October	-	St.2A	100	20.6	29.8	0	1.4	3.0	11.3	0.1	0.1	1.07	
			30	20.5	29.8	1	-	-	-	-	-	1.30	
			1	21.9	30.2	3	1.3	1.3	8.1	0.1	0.1	1.24	
		St.4	100	20.9	30.3	0	1.6	3.1	14.0	0.1	0.1	2.69	
			30	20.7	30.3	1	-	-	-	-	-	2.93	
			1	20.6	30.6	5	1.1	0.6	7.4	0.1	0.1	1.74	
St.5	100	19.1	30.4	0	1.0	0.4	6.5	0.1	0.1	2.47			
	30	18.5	30.5	2	-	-	-	-	-	1.98			
	1	18.1	30.4	8	1.2	0.2	5.3	0.0	0.0	2.20			



676

Table 1. (continued)

Year	Date	Irradiance ($\mu\text{mols m}^{-2} \text{s}^{-1}$)	Station	Light depth (%)	Temperature ($^{\circ}\text{C}$)	Salinity (‰)	Depth (m)	NH_3 (μM)	NO_2+NO_3 (μM)	SiO_2 (μM)	PO_4 (μM)	Chl- <i>a</i> ($\mu\text{g L}^{-1}$)
2013	January	297.4 ± 310.5	St.2A	100	5.5	20.5	0	0.5	4.2	4.0	0.1	1.39
				30	7.0	28.0	1	-	-	-	-	1.52
				1	7.3	29.4	4	0.5	3.7	3.6	0.1	1.48
			St.4	100	7.7	31.1	0	1.0	3.8	3.4	0.1	2.79
				30	7.4	31.3	4	-	-	-	-	3.41
				1	7.3	32.8	12	0.6	3.1	2.5	0.0	5.37
			St.5	100	6.3	31.8	0	0.8	3.3	2.6	0.1	5.79
				30	6.6	31.9	3	-	-	-	-	5.25
				1	6.4	32.5	11	1.0	3.0	3.6	0.2	5.33
	April	1593.3 ± 414.5	St.2A	100	14.3	26.2	0	1.9	3.7	3.1	0.1	1.81
				30	14.4	27.5	1	-	-	-	-	1.72
				1	14.3	29.1	3	1.5	2.5	2.3	0.1	2.06
			St.4	100	14.7	32.0	0	1.6	2.0	2.5	0.1	2.24
				30	15.3	32.0	1	-	-	-	-	4.41
				1	15.2	32.6	5	1.5	1.7	1.6	0.1	7.39
St.5	100	16.1	31.9	0	1.1	1.3	1.3	0.1	4.39			
	30	16.1	32.0	3	-	-	-	-	5.22			
	1	16.6	32.3	11	1.1	0.7	1.0	0.1	5.90			



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Table 2. Rainfall and river input

Year	Date	Rainfall (mm)	River input (10⁶ t)
2012	April	195.5	149.4
	May	44.4	148.9
	June	69.6	42.3
	July	235.8	223.3
	August	559.0	228.9
	September	360.1	447.2
	October	38.0	98.5
	November	52.5	83.4
	December	96.7	89.4
2013	January	15.6	79.3
	February	116.4	94.6
	March	79.9	91.5
	April	99.1	100.3



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Table 3. $\delta^{13}\text{C}$ values and C:N ratios of POM in Gwangyang Bay (surface)

Year	Date	$\delta^{13}\text{C}$ (‰)	C:N ($\mu\text{g } \mu\text{g}^{-1}$)
2012	April	-22.8	7.0
	June	-23.1	6.8
	August	-16.5	6.7
	October	-17.1	6.9
2013	January	-22.5	7.7
	April	-23.1	6.8
(average \pm S.D.)		-20.9 \pm 3.2	7.0 \pm 0.4



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Table 4. Biochemical concentrations and composition, calorific values and contents in Gwangyang Bay (- : no data).

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Year	Date	Station	Light depth (%)	CHO ($\mu\text{g L}^{-1}$)	PRT ($\mu\text{g L}^{-1}$)	LIP ($\mu\text{g L}^{-1}$)	FM ($\mu\text{g L}^{-1}$)	CHO/FM (%)	PRT/FM (%)	LIP/FM (%)	Kcal g^{-1}	Kcal m^{-3}	
2012	April	St.1	100	45.0	144.2	22.9	212.1	21.2	68.0	10.8	5.6	1.2	
			30	53.1	218.6	51.9	323.6	16.4	67.6	16.0	5.9	1.9	
			1	53.1	220.4	84.2	357.6	14.8	61.6	23.5	6.2	2.2	
		St.4	100	14.2	128.1	28.6	170.9	8.3	74.9	16.7	6.1	1.0	
			30	50.0	155.1	21.4	226.5	22.1	68.5	9.4	5.6	1.3	
			1	20.2	146.0	37.3	203.5	9.9	71.8	18.3	6.1	1.2	
		St.5	100	60.2	198.0	143.0	401.2	15.0	49.3	35.7	6.7	2.7	
			30	132.4	198.0	42.8	373.2	35.5	53.1	11.5	5.5	2.0	
			1	146.7	265.3	210.0	622.1	23.6	42.7	33.8	6.5	4.1	
	June	St.2A	100	170.7	99.7	233.5	503.8	33.9	19.8	46.3	6.9	3.5	
			30	135.5	108.0	251.9	495.4	27.3	21.8	50.9	7.2	3.5	
			1	163.5	85.0	225.1	473.7	34.5	17.9	47.5	6.9	3.3	
		St.4	100	99.1	44.6	199.5	343.2	28.9	13.0	58.1	7.4	2.5	
			30	133.4	142.4	203.5	479.3	27.8	29.7	42.4	6.8	3.3	
			1	91.6	110.8	232.3	434.6	21.1	25.5	53.5	7.3	3.2	
		August	St.4	100	69.3	73.9	213.5	356.7	19.4	20.7	59.9	7.6	2.7
				30	61.2	56.5	173.8	291.5	21.0	19.4	59.6	7.6	2.2
				1	127.2	77.9	162.2	367.3	34.6	21.2	44.2	6.8	2.5
	St.5		100	155.5	289.4	204.7	649.6	23.9	44.6	31.5	6.4	4.2	
			30	412.3	102.0	401.4	915.7	45.0	11.1	43.8	6.6	6.1	
			1	83.3	22.8	228.3	334.4	24.9	6.8	68.3	7.9	2.6	
October	St.2A	100	71.0	82.2	104.1	257.3	27.6	32.0	40.5	6.7	1.7		
		30	42.7	62.4	100.3	205.4	20.8	30.4	48.8	7.2	1.5		
		1	74.3	111.6	98.5	284.4	26.1	39.2	34.6	6.5	1.9		
	St.4	100	51.6	105.2	105.3	262.2	19.7	40.1	40.2	6.8	1.8		
		30	119.4	121.9	144.4	385.6	31.0	31.6	37.4	6.6	2.5		
		1	78.5	169.0	134.4	381.9	20.6	44.2	35.2	6.6	2.5		
	St.5	100	37.2	70.0	86.5	193.6	19.2	36.1	44.7	7.0	1.4		
		30	42.3	92.5	112.0	246.7	17.2	37.5	45.4	7.1	1.7		
		1	33.9	108.4	97.3	239.7	14.2	45.2	40.6	6.9	1.7		



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Table 4. (continued)

Year	Date	Station	Light depth (%)	CHO ($\mu\text{g L}^{-1}$)	PRT ($\mu\text{g L}^{-1}$)	LIP ($\mu\text{g L}^{-1}$)	FM ($\mu\text{g L}^{-1}$)	CHO/FM (%)	PRT/FM (%)	LIP/FM (%)	Kcal g^{-1}	Kcal m^{-3}	
2013	January	St.2A	100	150.3	139.3	115.5	405.2	37.1	34.4	28.5	6.1	2.5	
			30	347.0	131.1	109.2	587.3	59.1	22.3	18.6	5.4	3.2	
			1	331.3	127.1	-	-	-	-	-	-	-	-
		St.4	100	171.6	164.0	-	-	-	-	-	-	-	-
			30	183.5	168.7	139.7	491.9	37.3	34.3	28.4	6.1	3.0	
			1	115.9	182.3	107.1	405.2	28.6	45.0	26.4	6.2	2.5	
		St.5	100	113.6	212.0	133.4	459.0	24.7	46.2	29.1	6.3	2.9	
			30	264.1	204.8	120.5	589.4	44.8	34.8	20.4	5.7	3.4	
			1	99.3	195.5	104.2	399.0	24.9	49.0	26.1	6.2	2.5	
	April	St.2A	100	237.7	262.9	189.9	690.5	34.4	38.1	27.5	6.1	4.2	
			30	185.5	308.0	198.7	692.3	26.8	44.5	28.7	6.3	4.3	
			1	274.8	382.4	180.3	837.5	32.8	45.7	21.5	5.9	4.9	
		St.4	100	115.0	141.9	181.4	438.4	26.2	32.4	41.4	6.8	3.0	
			30	116.4	187.0	191.0	494.5	23.5	37.8	38.6	6.7	3.3	
			1	205.2	222.1	185.7	612.9	33.5	36.2	30.3	6.2	3.8	
St.5	100	160.4	176.3	289.1	625.7	25.6	28.2	46.2	7.0	4.4			
	30	146.9	217.8	253.3	618.0	23.8	35.2	41.0	6.8	4.2			
	1	171.3	204.9	272.6	648.8	26.4	31.6	42.0	6.8	4.4			



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Table 5. Significant correlation coefficient (r) among proteins (PRT), lipids (LIP) and environmental factors (ns ; no significance, **; $p < 0.01$).

Variables	r	p	n
%PRT \times Temp.	- 0.52	**	46
%LIP \times Temp.	0.72	**	46
%PRT \times NH ₄	0.69	**	28
%LIP \times NH ₄	-0.59	**	28
%PRT \times NO ₂ +NO ₃	0.54	**	28
%LIP \times NO ₂ +NO ₃	-0.53	**	28
%PRT \times River-input	0.84	**	46
%LIP \times River-input	-0.63	**	46
NH ₄ \times River-input	0.91	**	28
NO ₂ +NO ₃ \times River-input	0.55	**	28
%PRT \times %LIP	-0.81	**	46
%PRT \times Irradiance	-0.22	ns	39
%LIP \times Irradiance	0.24	ns	39



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Table 6. Observed nutrient limitations during the study period.

Year	Date	Based on absolute concentrations (μM)				Based on molar ratios		
		DIN	SiO_2	PO_4	Limitation	Si:N	Si:P	Limitation
2012	April	20.3 ± 20.2	11.1 ± 8.8	13.6 ± 32.9	nd	0.6 ± 0.2	37.5 ± 36.9	nd
	June	-	-	-	-	-	-	-
	August	0.4 ± 0.4	10.2 ± 1.5	0.1 ± 0.0	N, P	42.7 ± 23.7	173.4 ± 56.5	N, P
	October	2.7 ± 1.5	8.8 ± 3.3	0.1 ± 0.0	P	3.6 ± 0.8	142.2 ± 74.0	N
2013	January	4.2 ± 0.4	3.3 ± 0.6	0.1 ± 0.1	P	0.8 ± 0.1	50.6 ± 41.4	nd
	April	3.4 ± 1.3	2.0 ± 0.8	0.1 ± 0.0	P	0.6 ± 0.1	15.5 ± 5.5	Si



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Table 6. Comparison of biochemical quantity of POM, FM and the calorific contents.

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Regions (depth)	PRT ($\mu\text{g l}^{-1}$)	LIP ($\mu\text{g l}^{-1}$)	CHO ($\mu\text{g l}^{-1}$)	FM ($\mu\text{g l}^{-1}$)	Kcal m^{-3} (average \pm S.D.)	Authors
Gwangyang Bay, South Korea (Euphotic depth)	23-382	21-401	14-412	171-916	2.8 ± 1.1	This study
Arctic regions						
Bedford Basin, Canada(2.5 m)	200-650	130-440	160-630	660-1570		Mayzaud et al. (1989)
Logy Bay, Newfoundland (6 m)	80-740	20-75	8-120	130-1030	2.7 ± 2.8	Navarro & Thompson (1995)
The Northern Chukchi Sea, 2011 (Euphotic depth)	1-86	50-105	22-147	94-246	1.0 ± 0.2	Kim et al. (2015)
The Northern Chukchi Sea, 2012 (Euphotic depth)	9-183	37-147	16-253	90-373	1.2 ± 0.2	Yun et al. (2015)
Antarctic regions						
Pacific Sector Antarctic Ocean (0-1500 m)	14-100	3-60	3-66	25-220		Tanoue (1985)
Off Princess Astrid Coast, Antarctica (0-100m)	24-200	15-174	22-147	148-393		Dhargalkar et al. (1996)
Ross Sea, Antarctica (10m)	11-402	91	91-187	193-680	2.6 ± 1.8	Fabiano and Pusceddu (1998)
Ross Sea, Antarctica (0-200 m)	40-406	18-115	22-251	110-660		Fabiano et al. (1999)
Terra Nova Bay, Antarctica (0-750 m)	10-620	2-77	8-144	19-885	1.3 ± 1.0	Fabiano et al. (1996)
Terra Nova Bay, Antarctica (under pack ice)	96-201	38-112	10-68	145-382	1.7 ± 1.1	Pusceddu et al. (1999)
Amundsen Sea (Euphotic depth)	6-396	13-37	3-216	43-639	1.2 ± 0.8	Kim et al. (2015)
Other regions						
W-Mediterranean (0-200 m)	72-105	37-51	33-88	143-246		Fabiano et al. (1984)
W-Mediterranean submarine cave (10m)	4-77	4-104	1-75	15-220	0.4 ± 0.2	Fichez (1991b)
Mediterranean seagrass (4 m)	25-135	50-180	40-110	125-395		Danovaro et al. (1998)
Ligurian Sea (10 m) NW-Mediterranean	32-107	21-140	21-131	74-378	1.5 ± 1.4	Danovaro & Fabiano (1997)
Mediterranean (30m)	70-90	90-110	10-20	177-213	1.4 ± 0.2	Modica et al. (2006)
Cretan Sea (0-1500 m)	7-92	4-63	13-149	54-200	0.6 ± 0.2	Danovaro et al. (2000)
Bay of Biscay, 2000 (0-30m)	109-2426	26-2037	2-345	961 (a.v.)	6.7 ± 5.0	Díaz et al. (2007)
Yaldad Bay, Chile (10 cm a.b.)	300-2250	30-560	50-1050	3310-2960	10.0 ± 10.9	Navarro et al. (1993)
The Humboldt current system, Northern Chile (5-89m)	40-470	60-390	70-510	24-1282	3.5 ± 3.3	Isla et al. (2010)
Magellan Strait (0-50m)	60-150	30-70	20-40	110-256	1.0 ± 0.5	Fabiano et al. (1999)
The northern part of the East Sea (Euphotic depth)	28-425	12-180	19-206	109-810	1.5 ± 0.6	Kang et al. (unpublished)



689

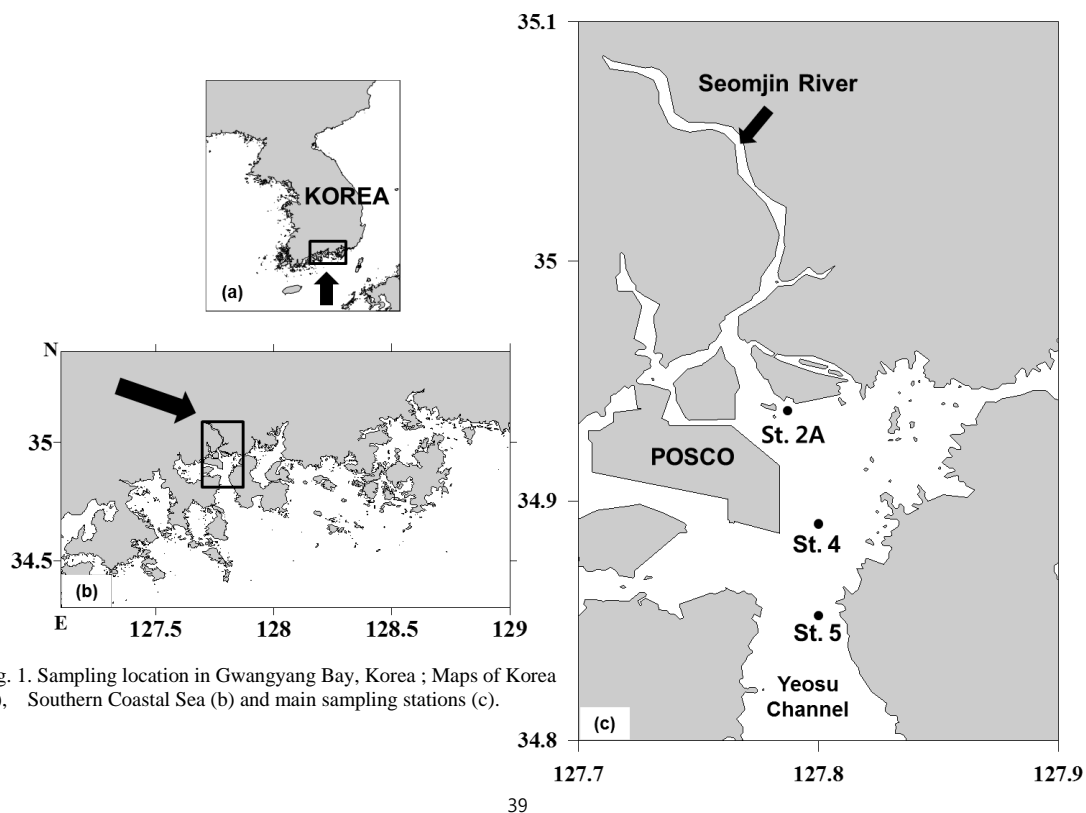


Fig. 1. Sampling location in Gwangyang Bay, Korea ; Maps of Korea (a), Southern Coastal Sea (b) and main sampling stations (c).



690

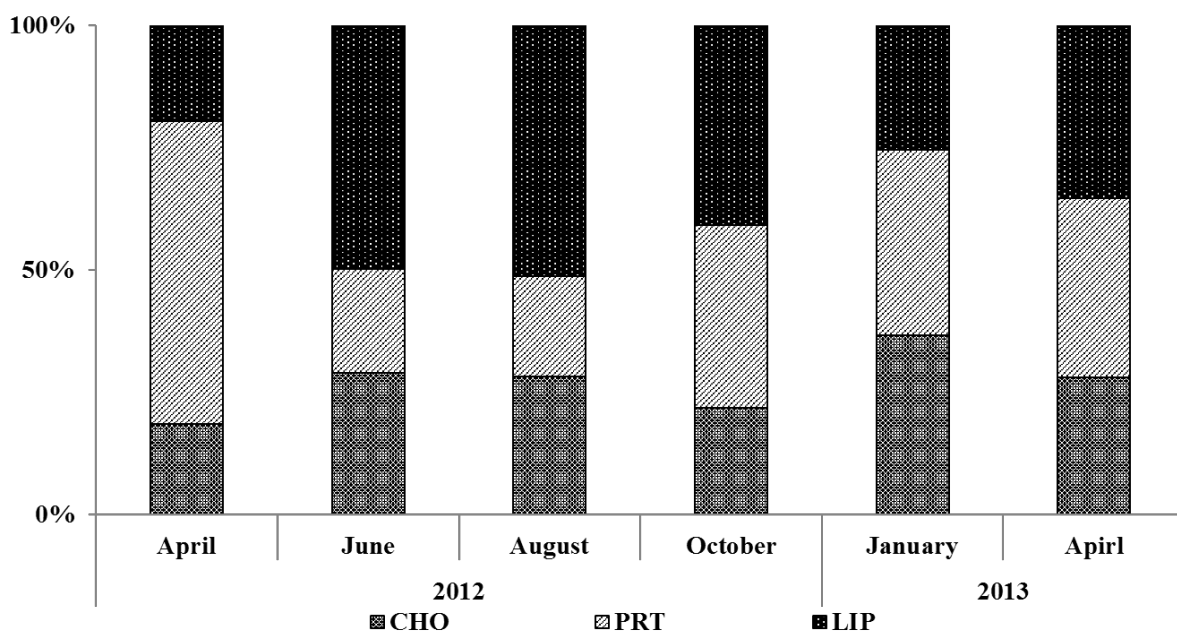


Fig. 2. Seasonal variation of biochemical composition in Gwangyang Bay.



691

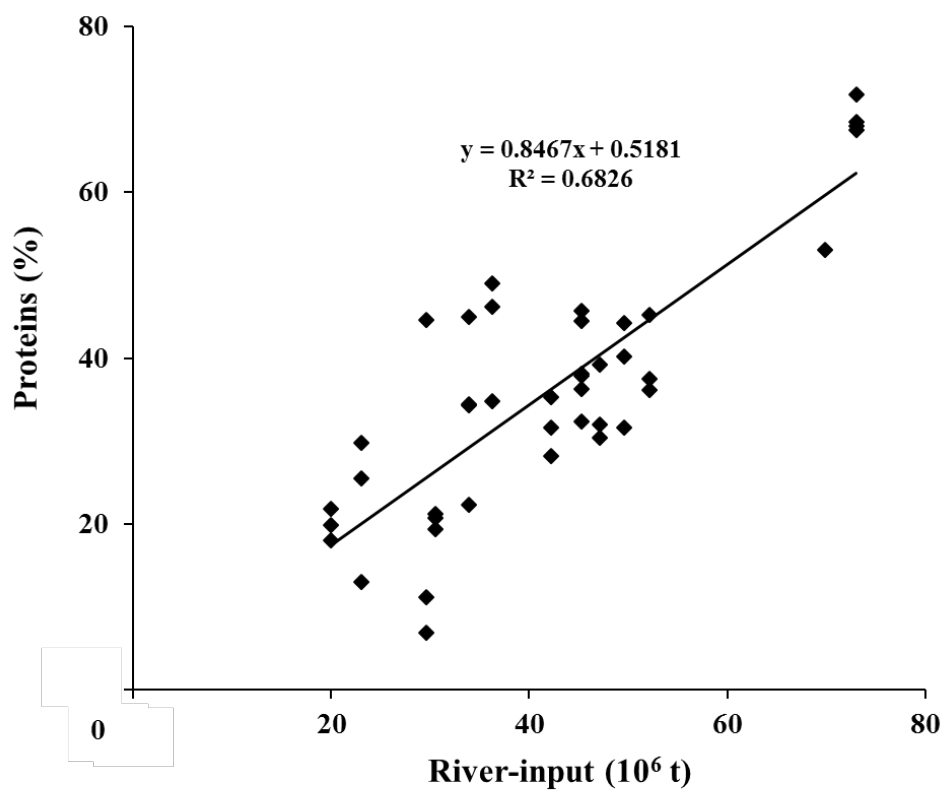


Fig. 3. The positive relationship between river-input and protein composition.



692

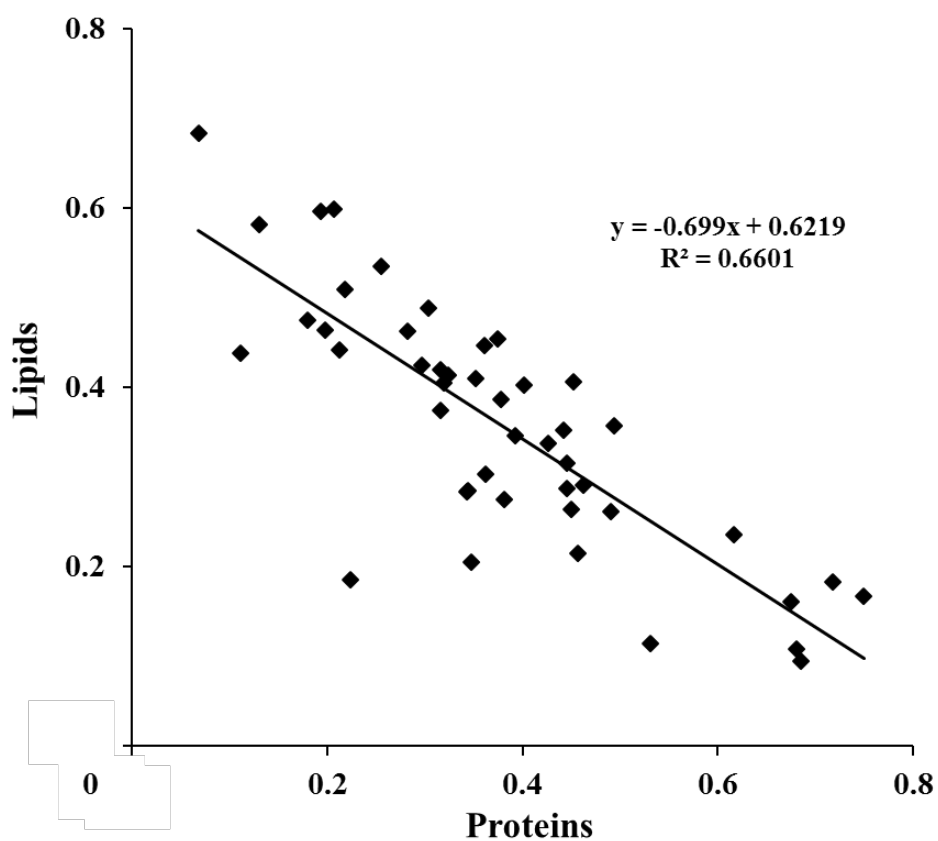


Fig. 4. The inverse relationship between lipid compositions and protein compositions