

**bg-2016-404:** The riverine source of tropospheric CH<sub>4</sub> and N<sub>2</sub>O from the Republic of Congo, Western Congo Basin

Response to referee comments:

Three sets of reviewer comments were received for this manuscript. Below we reproduce these comments, italicised, and followed by our responses, which are un-italicised. Text that we have now deleted/modified is highlighted in grey, with text lines in the original manuscript identified. Replacement/modified text is highlighted in yellow, with new manuscript text lines identified. The replacement/modified text is also highlighted in yellow on the revised manuscript.

**NB. Please see additional Associate Editor Comments and the responses to these under the appropriate line numbers: line 357, Reviewer #1 and Line 70, Reviewer #2).**

### **Anonymous Reviewer #1**

*I suggest that the authors make their data-set public as a supplement of the paper. Considering the enormous range of (spatial and temporal) variability of CH<sub>4</sub> (and to a lesser extent of N<sub>2</sub>O) in freshwaters, there is a need to compile and aggregate available data-sets to revise and update CH<sub>4</sub> fluxes from inland waters. This is only possible if an open data access attitude is adopted by the community.*

We absolutely agree; however, we did not want to make our data public at the initial review stage, preferring to await full publication of the manuscript. We now include a supplementary data file that list all relevant data. We have included the source data as supplementary Table S1 and we now include the statement: “Source data for this paper are available as supplementary material (Table S1)”, in new text line 173.

*L47: Alternative explanations have been proposed, such as related to agriculture (Schaefer et al. 2016) or fossil fuel (Rice et al. 2016), that differ from the explanation from Bousquet et al.*

While this is true, very recent isotopic work (Nisbet et al. Rising atmospheric methane: 2007–2014 growth and isotopic shift. *Glob. Biogeochem. Cyc.* 30, 1356–1370, doi:10.1002/2016GB005406, 2016) now provides strong evidence for recently increased biogenic emissions, especially in the tropics, where these may reflect an expansion of tropical wetlands in response to positive rainfall anomalies. In view of this new evidence, we have removed the original text in lines 44-47:

“Following periods of declining and zero growth from the mid-1980s, tropospheric CH<sub>4</sub> progressively increased from the late 2000s (Rigby et al., 2008; Dlugokencky et al., 2009). Growth ~ 4-5 ppbv yr<sup>-1</sup> since 2009 (Sussmann et al., 2012) has been linked to increasing natural tropical emissions (Bousquet et al., 2011”.

This is now replaced with new text lines 46-52: “The early 1980s to the mid-2000s saw an overall decline in tropospheric CH<sub>4</sub> growth (Dlugokencky et al., 2009; 2011), which has been ascribed to declining fossil fuel emissions and/or variations in the OH radical sink (Rice et al., 2016; Schaefer et al., 2016), punctuated by episodic events such as the 1991-1992 Pinatubo eruption and an intense 1997–1998 El Niño (Nisbet et al., 2016). However, since 2007 increased growth has been sustained. Recent evidence from isotopic studies (Nisbet et al., 2016) and a box-model (Schaefer et al., 2016) implies this to be a consequence of increased biogenic emissions, particularly in the tropics, where these increased emissions have been linked to an expansion of tropical wetlands in response to positive rainfall anomalies (Nisbet et al., 2016), and/or growing emissions from agricultural sources (Schaefer et al., 2016).

References to Bousquet et al. (2011), Rigby et al. (2008) and Sussman et al. (2012) are now redundant and have been removed from the reference list. Nisbet et al. (2016), Dlugokencky et al. (2011), Rice et al. (2016) and Schaefer et al. (2016) have been added.

*L 59: This is not correct any more in the light of the paper by Stanley et al. 2016.*

*L 60 : It could be useful to mention that there's a discrepancy in bottom-up and top-down estimates of CH<sub>4</sub> fluxes (Saunois et al. 2016), hence the comparison of Bastviken et al. estimates with those of Kirschke et al. might be biased by the fact that they were derived by different and possibly incompatible methods.*

We considered these two comments together because they are linked in the text. In light of these we have revised the original text in 57-59: "The global freshwater CH<sub>4</sub> source could be  $\sim 10^{13}$ - $10^{14}$  g yr<sup>-1</sup> (Bastviken et al., 2011; Kirschke et al., 2013), the uncertainty reflecting data gaps, notably for major world river basins, and a sampling bias that has necessitated upscaling from exclusively temperate data (Bastviken et al., 2011)". Noting that the estimate of Stanley et al (2016) falls within the range set by Bastviken et al. (2011) and Kirschke et al. (2013), the new text reads (new text lines 61-66) now reads: "The global freshwater CH<sub>4</sub> source could be  $\sim 10^{13}$ - $10^{14}$  g yr<sup>-1</sup> (Bastviken et al., 2011; Kirschke et al., 2013; Stanley et al., 2016). This order of magnitude range largely reflects a discrepancy between "top-down" approaches based on atmospheric inversions (e.g. Kirschke et al., 2013) and "bottom-up" estimates that necessitate the upscaling of freshwater observations (e.g. Bastviken et al., 2011). For example, atmospheric constraints on top-down budgets imply that some component emissions of bottom-up approaches may be overestimates (Saunois et al., 2016)".

We have added the Saunois et al. (2016) to the reference list.

*L 171 : This explanation for the seasonal variations of DIN is surprising given these are near pristine watersheds due to the low population density in ROC and the absence of intensive agricultural practise (based on artificial fertilizer) and major industrial activities. Seasonal variations of DIN are likely due to surface run-off in the wet season and groundwater flow in deeper soil horizons in the dry season.*

We find this comment perplexing because we did not comment at all on seasonal variations in DIN. This is because we only have DIN data for the wet season, something that we made absolutely clear in the opening sentence of the results section (original text lines 162-163). We could only compare the three river "types" during the wet season and see no reason therefore, to speculate about wet vs dry season sources of DIN here, given that we have no data for the latter. Perhaps the reviewer does not mean to refer to seasonal variation at all but has simply become a little confused and has misinterpreted what it was we were intending to say in general about the importance of agricultural DIN sources. Our original text, lines 171-172: "Low wet season concentrations of dissolved inorganic nitrogen (DIN) components (Fig. 3) are consistent with low nitrogen input rates from agricultural, domestic and industrial sources (Clark and Decalo, 2012)" means to imply exactly what the reviewer argues, i.e. that agricultural impacts are negligible. To preclude any further misunderstanding however, we have modified this text so that it now reads (new text lines 181-182): "Low wet season concentrations of dissolved inorganic nitrogen (DIN) components (Fig. 3) are consistent with agricultural, domestic and industrial DIN sources all being negligible (Clark and Decalo, 2012)". We trust that this is even clearer than it was previously. In view of this we have also modified the corresponding text in the abstract: "total DIN concentrations ( $1.5$ - $45.3 \mu\text{mol L}^{-1}$ ) being consistent with small agricultural, domestic and industrial sources". "total DIN concentrations ( $1.5$ -

45.3  $\mu\text{mol L}^{-1}$ ) being consistent with negligible agricultural, domestic and industrial sources" (text lines 17-18).

*L 240-264: The existence, in rivers and wetlands, of high levels of CH<sub>4</sub> in oxygenated waters is not "enigmatic" nor "counterintuitive" as stated. This has been shown and explained for decades for instance in the Amazon (Richey et al. 1988), and is related to methane production in the anoxic sediments of river-beds and floodplains that diffuses into aerated river water. In shallow and low turbulent "swampy" waters such as those sampled, the diffusion of CH<sub>4</sub> from river sediments is stronger than loss terms in the water by oxidation or evasion to the atmosphere, leading to an accumulation of CH<sub>4</sub> in the water (even in the presence of more or less large quantities of O<sub>2</sub>). This is fairly straightforward and intuitive, in rivers with a probable depth between 1 and 5 m, hence, in close contact with organic rich sediments.*

We accept this point completely. We have therefore deleted the following original text (original text lines 240-249): "Notwithstanding this complexity, the coexistence of CH<sub>4</sub> with dissolved O<sub>2</sub> in rivers of the ROC (Fig. 2a) initially seems enigmatic. While dissolved O<sub>2</sub> was under-saturated in the majority of samples, being as low as 4% in one wet season swamp sample, it was always detectable and indeed was super-saturated in several savannah river samples in which CH<sub>4</sub> saturations ranged from ~4000-10000 % (Fig. 2a). These observations seem counterintuitive because the classical view of methanogenesis is that it is exclusively anoxic, carried out by severely O<sub>2</sub>-limited archaea (Bridgham et al., 2013). However, recent evidence is for a greater complexity of CH<sub>4</sub> production in river catchments. For example, methanogenesis in "anoxic microsites" within oxic soils is widely acknowledged (e.g. Teh et al., 2005; von Fisher & Hedin, 2007). Methanogens are now considered to be widespread in oxic soils and they are activated during flooding (Bridgham et al., 2013), their activity relating to soil carbon age and composition (Bridgham et al., 1998; Chanton et al., 2008) and likely involving substrate competition and other interactions".

This is now replaced by new text which is informed by the reviewer's comments (new text lines 250-263): "While dissolved O<sub>2</sub> was under-saturated in the majority of samples, being as low as 4% in one wet season swamp sample, it was always detectable and indeed was super-saturated in several savannah river samples in which CH<sub>4</sub> saturations ranged from ~4000-10000 % (Fig. 2a). Notwithstanding that methanogenesis is an exclusively anoxic process carried out by severely O<sub>2</sub>-limited archaea (Bridgham et al., 2013), the existence of high CH<sub>4</sub> concentrations in oxygenated rivers is well-known (e.g. Richey et al., 1988). It is a consequence of the diffusion of CH<sub>4</sub> produced in underlying river sediments, in adjacent floodplain soils and in adjacent wetlands, into aerated river water. Previous work showed that the CH<sub>4</sub> supply from groundwater to African rivers is generally comparatively low (Balagizi et al., 2015; Borges et al., 2015b) and that the spatial distribution of dissolved CH<sub>4</sub> is more closely related to the wetland distribution within the catchment, wetland water deriving principally from upland runoff (Borges et al., 2015a). In the Congo basin floating macrophytes, both in the centre of river channels and fringing their edges, are important additional sources of CH<sub>4</sub> (Borges et al., 2015a). These sources promote a unidirectional CH<sub>4</sub> flow, towards small and large river channels (Borges et al., 2015a). Where these river channels are shallow and have low levels of surface turbulence, as in the examples studied here, the CH<sub>4</sub> diffusion term evidently exceeds combined CH<sub>4</sub> losses via oxidation and water-to-air exchange, thereby further promoting the accumulation of high river CH<sub>4</sub> concentrations.

In addition to methanogenesis in fully anoxic sediment and soils however, CH<sub>4</sub> production can also occur in “anoxic microsites” within oxic soils (e.g. Teh et al., 2005; von Fisher & Hedin, 2007). Indeed, methanogens are now considered to be widespread in oxic soils and they are activated during flooding (Bridgham et al., 2013), their activity relating to soil carbon age and composition (Bridgham et al., 1998; Chanton et al., 2008) and likely involving substrate competition and other interactions”.

*There is no need to use exotic hypothesis related to DMS(P) cycling (Dam) or methylphosphanate (Karl) that were developed for the ocean, where the occurrence of CH<sub>4</sub> in oxygenated waters located hundredths to thousands of meters away from the seabed is indeed “enigmatic”, hence, the so called “oceanic CH<sub>4</sub> paradox”. Further, most of these hypothesis rely on a more or less direct production of CH<sub>4</sub> linked to phytoplankton (e.g. Grossart). However, phytoplankton is nearly absent in tributaries and wetlands of the Congo (Descy et al. 2016).*

The inclusion of this was a response to an initial request from an Associate Editor that we discuss ALL aquatic sources of CH<sub>4</sub>. We fully concur with the reviewer’s view and we are happy to have now deleted the offending text (original lines 255-258): “Additional production in oxic seawater may involve biological uptake of organic PO<sub>4</sub><sup>3-</sup> (Karl et al., 2008) and methylotrophic methanogenesis (Damm et al., 2010), both mechanisms being associated with nutrient stress, but neither has yet been identified in freshwaters”. Damm et al (2010) and Karl et al (2008) have been removed from the reference list.

We have also modified the preceding sentence (original lines 252-254): “Further, methanogenesis by photoautotroph-attached archaea has been detected in oxic lake water (Grossart et al., 2011), analogous to the anoxic micro-niches” associated with dead and living particles in oxic sea water (de Angelis and Lee, 1994; Oremland, 1979; Ditchfield et al., 2012)” replacing this by text lines 266-268: “Although methanogenesis by photoautotroph-attached archaea has been detected in oxic lake water (Grossart et al., 2011), this is unlikely in tributaries and wetlands of the Congo, where phytoplankton abundance is low (Descy et al. 2016)”. Consequently, de Angelis and Lee (1994), Oremland (1979) and Ditchfield et al. (2012) have been deleted from the reference list and Descy et al (2016) has been added.

*L 276: Please use the term “Pool Malebo”, the term “Stanley Pool” has been abandoned since colonial times.*

We note that “Stanley Pool” still appears on some maps and in some relatively recent documents. To provide continuity in this regard we have replaced “in the vicinity Stanley Pool (Fig. 1)” by (new text lines 287-288): “in the vicinity of Pool Malebo (Formerly known as Stanley Pool) (Fig. 1)”

*L 289-308 : Seasonal variations of CH<sub>4</sub> in floodplains has been relatively well described in the Amazon varzeas (Devol et al. 1990).*

We are unable to see how this comment refers to the text passage identified or where it could fit in, especially as no supporting argument is presented by the referee. Even so, we think this is a minor point that does not add to the thrust of our reasoning.

*L348-349: Rates of nitrification can exceed denitrification in NH<sub>4</sub> enriched temperate rivers such as the Mississippi studied by Richardson et al., however this does not necessarily apply in DIN poor tropical rivers where as stated NO<sub>3</sub> dominates the DIN pool.*

Although nitrate does indeed dominate the DIN pool it does not necessarily follow that nitrification would thus be unimportant. As we stated earlier, nitrate accounts for  $63 \pm 19\%$  of the sum of nitrate and ammonium so the latter is still important. We are therefore unsure what the referee is asking, especially as he/she uses the phrase “does not necessarily apply”. We do not believe this to be a strong case for dismissing the possibility of nitrification and as we argue, this explanation is compatible with our observations.

*L 357 : If the authors envisage all possible CH<sub>4</sub> sources in marine and freshwater environments then CH<sub>4</sub> has more diverse sources than N<sub>2</sub>O. However, the only documented CH<sub>4</sub> sources in tropical rivers are methanogenesis in riverbed and floodplain sediments.*

[Additional Associate Editor’s comment} Authors should provide evidence/statement for most possible source of CH<sub>4</sub> emission in Western Congo Basin].

In view of the additional Associate Editors comment we have further revised new text lines 252-260, inserting: Previous work showed that the CH<sub>4</sub> supply from groundwater to African rivers is generally comparatively low (Balagizi et al., 2015; Borges et al., 2015b) and that the spatial distribution of dissolved CH<sub>4</sub> is more closely related to the wetland distribution within the catchment, wetland water deriving principally from upland runoff (Borges et al., 2015a). In the Congo basin floating macrophytes, both in the centre of river channels and fringing their edges, are important additional sources of CH<sub>4</sub> (Borges et al., 2015a). These sources promote a unidirectional CH<sub>4</sub> flow, towards small and large river channels (Borges et al., 2015a)”. Therefore this complete section (new text lines 250-264) now makes clear both the principal processes and the major sites of CH<sub>4</sub> generation relating to ROC rivers. In view of this we have revised the concluding paragraph: To conclude, while our data have allowed us to draw some conclusions regarding the production and cycling of CH<sub>4</sub> and N<sub>2</sub>O in contrasting rivers of the ROC, we consider these to be more robust for N<sub>2</sub>O given that its aquatic sources are the least diverse. However, for both gases an unequivocal identification of the primary controls of their riverine distributions would require additional detailed measurements. So that it now reads: (new text lines To conclude, while our data have allowed us to draw some conclusions regarding the production and cycling of CH<sub>4</sub> and N<sub>2</sub>O in contrasting rivers of the ROC, for both gases an unequivocal identification of the primary controls of their riverine distributions would require additional detailed measurements. We have thus deleted the contentious section contending that N<sub>2</sub>O sources are the least diverse, because we agree that this cannot really be justified.

*L 363-365 : The cited range of CH<sub>4</sub> and N<sub>2</sub>O fluxes correspond to the basin average values and not the full range of individual CH<sub>4</sub> and N<sub>2</sub>O flux estimates for each of the 12 river basins.*

*The range of CH<sub>4</sub> fluxes for all individual estimates across the 12 rivers studied by Borges et al. (2015) is 0 to 274,600  $\mu\text{mol}/\text{m}^2/\text{d}$  for Aufdenkampe K estimate and 0 to 461,967  $\mu\text{mol}/\text{m}^2/\text{d}$  for Raymond K estimate. This range in fact corresponds to the one of the Congo that encompasses the data from all other African rivers. The range of N<sub>2</sub>O fluxes for all individual estimates across the 12 rivers studied by Borges et al. (2015) is -30 to 299  $\mu\text{mol}/\text{m}^2/\text{d}$  for Aufdenkampe K estimate and -37 to 377  $\mu\text{mol}/\text{m}^2/\text{d}$  for Raymond K estimate. This range in fact corresponds to the one of the Congo that encompasses the data from all other African rivers.*

We apologise for the error and we have substituted the correct ranges. The original text in lines 361- 365: “For 12 sub-Saharan African rivers Borges et al. (2015b) give ranges of 502 - 18019  $\mu\text{mol CH}_4 \text{m}^{-2} \text{d}^{-1}$  and 2 - 16  $\mu\text{mol N}_2\text{O m}^{-2} \text{d}^{-1}$  using  $k_w$  from Aufdenkampe et al. (2011), and 583 - 28579  $\mu\text{mol CH}_4 \text{m}^{-2} \text{d}^{-1}$  and 2 - 28  $\mu\text{mol N}_2\text{O m}^{-2} \text{d}^{-1}$  using  $k_w$  from Raymond et al. (2013)” now reads: “For 12 sub-Saharan African rivers Borges et al. (2015b) give ranges of 0 to 274,600  $\mu\text{mol CH}_4 \text{m}^{-2} \text{d}^{-1}$  and -30 to 299  $\mu\text{mol N}_2\text{O m}^{-2} \text{d}^{-1}$  using  $k_w$  from Aufdenkampe et al. (2011), and 0 to 461,967  $\mu\text{mol CH}_4 \text{m}^{-2} \text{d}^{-1}$  and -37 to 377  $\mu\text{mol N}_2\text{O m}^{-2} \text{d}^{-1}$  using  $k_w$  from Raymond et al. (2013).” (new text lines 374-377).

In view of this it has been necessary to modify the text in original lines 385-386: “The overall ranges of CH<sub>4</sub> and N<sub>2</sub>O emissions from rivers of the ROC (33-48705 μmol CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>; 1-67 μmol N<sub>2</sub>O m<sup>-2</sup> d<sup>-1</sup>) are somewhat wider than these earlier estimates for African and temperate rivers, the maximum values (Table 3) being around twice as high as previously reported”.

This is now changed to: “The overall ranges of CH<sub>4</sub> and N<sub>2</sub>O emissions from rivers of the ROC (33 to 48705 μmol CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>; 1 to 67 μmol N<sub>2</sub>O m<sup>-2</sup> d<sup>-1</sup>, Table 3) fall within the ranges encompassed by these earlier estimates for African and temperate rivers” (new text lines 385-386). We also deleted “Nevertheless” from the start of the following sentence (new text line 387). We have also modified the corresponding line in the abstract, changing wider than previously estimated to “within the range previously estimated” (line 36).

## Anonymous Reviewer #2

*..the dataset is valuable and merits publication – but I strongly agree with Reviewer #1 that it is crucial that the full data are made available in digital form, there is little point in collecting these data and leaving a legacy of just average or mean values in a summary Table.*

We agree entirely; please see our response to Reviewer #1.

*-Title: I find the title a little awkward, the ‘riverine source of tropospheric CH<sub>4</sub> and N<sub>2</sub>O’ is somewhat misleading as the paper does not focus on budgeting sources of tropospheric CH<sub>4</sub> and N<sub>2</sub>O, it looks at CH<sub>4</sub> and N<sub>2</sub>O exchange (source, or sink insome cases for N<sub>2</sub>O) between surface waters and atmosphere. I suggest to rephrasethe title.*

While we do not agree at all that the title as it is implies some aspect of budgeting sources of tropospheric CH<sub>4</sub> and N<sub>2</sub>O, we agree to remove the contentious word “tropospheric”, so that the revised title now reads “The riverine source of CH<sub>4</sub> and N<sub>2</sub>O from the Republic of Congo, Western Congo Basin”. We contend that this is now clear and unambiguous.

*-L19-20: ‘predominantly supersaturated’: with an average of 100% saturation, why call this ‘predominantly oversaturated’ ? -L22: The abbreviation ROC is used here for the first time – spell it out here.*

This is because they were predominantly supersaturated in savannah rivers: most of the samples were >100% saturation but that is not to say that the average saturation cannot be 100%. It is simple statistics. Consider a simple example of three river samples, of 70%, 116% and 115% saturation respectively. The average ± standard deviation of these is 100 ± 26% but based on the samples collected, these river waters as a group are predominantly supersaturated because two thirds of them have > 100% saturation.

We have added “(ROC)” immediately following “Republic of Congo” in new text line 15.

*-L32-33: awkward sentence, not clear what the authors are trying to say here (‘was coincident with’)*

We see absolutely nothing awkward with the sentence. It is grammatically sound and we see no reason to change it.

*-L70: what do you mean with ‘seasonal emissions’?*

[Associate editor comment] vL-70: Are Authors referring to seasonal variability in the emission fluxes of CH<sub>4</sub> and N<sub>2</sub>O? Or are these emissions highly seasonal (occurring only during a specific

season)?

On reflection and taking account of the Associate Editor's additional comment we believe that the term "seasonal emissions" does not add anything to the sentence: "Borges et al (2015b) recently reported seasonal emissions  $\sim 3\text{-}4 \times 10^{12}$  g CH<sub>4</sub> yr<sup>-1</sup> and  $\sim 10^{10}$  g N<sub>2</sub>O yr<sup>-1</sup> for twelve large river systems in sub-Saharan Africa, including the three largest by catchment area (Congo, Niger, Zambezi)". Because Borges et al. present annual emissions data we have changed this to "annual emissions" and removed "yr<sup>-1</sup>" from each of the estimates so that the revised sentence now reads (new lines 78-80) "Borges et al (2015b) recently reported annual emissions  $\sim 3\text{-}4 \times 10^{12}$  g CH<sub>4</sub> and  $\sim 10^{10}$  g N<sub>2</sub>O for twelve large river systems in sub-Saharan Africa, including the three largest by catchment area (Congo, Niger, Zambezi)".

-L124: *inoculated: poisoned*

We believe this to be a rather trivial criticism. We have used both terms in previous papers in this context and both are acceptable. However, we have replaced *inoculated* by *poisoned* (new text line 134) to satisfy the reviewer's preference.

-L240-264: *I suggest deleting this entire section, not relevant for freshwater systems as pointed out by Reviewer #1.*

This has indeed been modified to remove material that is not relevant to freshwaters; please see response to Reviewer #1. However, as our modified text shows, we do not agree that the entire deletion of this section is at all appropriate because some aspects of it are indeed relevant to freshwaters and it is therefore important to retain the modified text that we discuss above, in our response to reviewer #1, i.e. new text lines 259-264.

## L. Golston

L41-42: *Hartmann et al. 2013 does not describe the impacts of CH<sub>4</sub> on oxidizing capacity and other atmospheric species. Myhre et al. 2013 is a suitable reference.*

We have now substituted (*Myhre et al., 2013*) for (*Hartmann et al., 2013*) in new text line 42

L47-48, also L53-54: *This is the global surface mean, not the tropospheric mean, which would be slightly lower. Chapter 2 of AR5 (Hartmann et al. 2013) is a better citation than Chapter 6 (Ciais et al. 2013) since it details the data sources where these numbers are derived from. Also, simply say "150%" rather than "more than 150%".*

We have substituted (*Hartmann et al., 2013*) for (*Ciais et al., 2013*) as suggested (new text lines 53 and 58). We have also deleted "more than" in original text line 48 (new text line 52).

L48-49: *It is not clear why "soils, freshwaters and coastal waters" is given. Freshwater and coastal waters were not explicitly discussed in the reference, and this excludes important sources they do give including animal waste management, biomass burning, oceans, and energy systems.*

On reflection we have decided that in order to avoid any unnecessary confusion it is best to remove the sentence in question, in lines 48-50: "Increasing tropospheric N<sub>2</sub>O largely reflects its enhanced emission from soils, freshwaters and coastal waters via the accelerated mobilisation of reactive nitrogen (Syakila and Kroeze, 2011)". Doing so does not materially affect our overall argument. We have also deleted the citation to (*Syakila and Kroeze, 2011*) from the reference list as it is not cited elsewhere in the manuscript.

L50-51: This has several issues and does not follow from the previous sentence. 1) rivers, estuaries, and coastal zones was already included as a source category for N<sub>2</sub>O in IPCC AR4 (see Denman et al. 2007), though the sentence seems to imply it was added in AR5. 2) it is not that they switched the classification from natural to anthropogenic, but that sufficiently reliable data became available on the anthropogenic component.

We have removed the previous sentence so the first comment no longer applies. We would like to retain this information about the switch in classification from natural to anthropogenic but we accept this point and so we have modified the sentence: “Consequently the IPCC now classifies river, estuary and coastal zone N<sub>2</sub>O sources as anthropogenic (Ciais et al., 2013)” (original text lines 50-51) and we have replaced it with: “Of particular note in regard to N<sub>2</sub>O, the availability of sufficiently reliable data on the anthropogenic components of river, estuary and coastal zone sources resulted in a change in their classification in the IPCC AR4 synthesis, from “natural” to “anthropogenic” (Denman et al., 2007) (new text lines 53-55). We have also added (Denman et al., 2007) to the reference list.

L60-61: “based on warming potentials and atmospheric lifetimes” -> “based on a 100 year Global Warming Potential” would be more descriptive

We have amended the text as suggested. “Converting it to CO<sub>2</sub> equivalents based on warming potentials and atmospheric lifetimes” (original text lines 61-62) now becomes “Converting these to CO<sub>2</sub> equivalents based on a 100 year Global Warming Potential” (new text line 70).

L60-62: Kirschke et al. 2013 gives a bottom-up value of 40 [8-73] Tg/yr for freshwater sources, and a total source of 678 [542-852] Tg/yr, giving a mean freshwater contribution of about ~6% along with a wide uncertainty range. The estimates given of ~30-47% for natural and ~12-20% total sources seem to be miscalculated. In addition, it is unclear why we are asked to ignore the range of 10 – 100 Tg/yr given on L57 for freshwater, which would yield an even larger range of uncertainty.

Considering the first statement that Kirschke et al. (2013) give a bottom-up value of 40 [8-73] Tg/yr for freshwater sources, this is  $4 \times 10^{13}$  g yr<sup>-1</sup> and the range (8-73 Tg/yr) is  $0.8-7.3 \times 10^{13}$  g yr<sup>-1</sup>. Therefore this is already encapsulated in our statement that “The global freshwater CH<sub>4</sub> source could be ~10<sup>13</sup>-10<sup>14</sup> g yr<sup>-1</sup> (Bastviken et al., 2011; Kirschke et al., 2013; Stanley et al., 2016)” (new text lines 61-62) so we see no issue here, although we have now modified this statement in light of information discussed below.

The assertion that: a total source of 678 [542-852] Tg/yr, giving a mean freshwater contribution of about ~6% is based entirely on the estimates presented in Kirschke et al., (2013), whereas we have additionally accounted for the estimates presented in Bastviken et al. (2011) and Stanley et al. (2016). However, we do accept that our calculated ranges based on these are a bit out; the lower ends are an order of magnitude too high, which must have been a transcription error, because the upper ends are about right given the errors. To do this calculation we took the maximum and minimum estimates for total and natural emissions presented in Kirschke et al. (2013) (Natural emissions  $179-484 \times 10^{12}$  g yr<sup>-1</sup>; Total emissions  $526-852 \times 10^{12}$  g yr<sup>-1</sup>). We then divided the minimum of our freshwater range ( $10^{13}$  g yr<sup>-1</sup>) by the maximum estimate in each case to give a lower boundary to the percentage contribution and we divided the maximum of our freshwater range ( $10^{14}$  g yr<sup>-1</sup>) by the minimum estimate in each case to give an upper boundary to the percentage contribution. This is quite usual practice that does not need spelling out in an Introduction. Our revised contributions are thus 2-56% for natural sources and 1-19% for total sources, i.e. the lower ends are an order of magnitude reduced but the upper ends only change marginally from the original estimates.



The comment: *In addition, it is unclear why we are asked to ignore the range of 10 – 100 Tg/yr given on L57 for freshwater, which would yield an even larger range of uncertainty* is misplaced because as we state above, these were taken account of in the calculation so are not ignored This comment seems to stem from a misreading of what we actually said.

Based on the above issues we have replaced the original text, incorporating the estimates of Kirschke et al (2103) for clarity. We have thus deleted” *Notwithstanding the uncertainty, this freshwater source estimate is ~30-47 % of natural CH<sub>4</sub> emissions and ~12-20% of total CH<sub>4</sub> emissions (Kirschke et al., 2013)*” (original text lines 60-61) and replaced it with” *Notwithstanding this uncertainty, the global freshwater CH<sub>4</sub> source can be evaluated in the light of natural and total global CH<sub>4</sub> source estimates of 179-484 x 10<sup>12</sup> g yr<sup>-1</sup> and 526-852 x 10<sup>12</sup> g yr<sup>-1</sup> respectively (Kirschke et al., 2013). Quantifying the global freshwater contribution thus has high inherent uncertainty but based on these estimates it could be ~ 2-56 % of natural CH<sub>4</sub> emissions and ~1-19% of total CH<sub>4</sub> emissions*” (new text lines 66-70).

*L63: Please double check this. It does not seem to be consistent with Figure 2 in Le Quéré et al. 2015 which suggests a combined sink of ~5.5 x 10<sup>15</sup> g C yr<sup>-1</sup>*

This seems to be a typographic error that we have now corrected (new line 72) but the conclusion of a significant offset is unchanged.

*L159-160: Same comment about being surface rather than tropospheric mean CH<sub>4</sub> and N<sub>2</sub>O. Also unclear how to access the referenced data using the link given at <http://www.eea.europa.eu/data-and-maps/>*

We have again deleted “tropospheric” (new text line 169), although we feel this a rather minor point. We have also replaced the original web-link with a more accessible data gateway: <http://cdiac.esd.ornl.gov/tracegases.html> (new text line 170).

*L269: supersaturations -> super-saturations (for consistency with usage in the rest of the text)*

This has been changed as requested (new text line 279).

*L383: Melack et al. 2004 and Bastkviken et al. 2010 both missing from References list*

“Bastkviken et al. 2010” (original text line 384) is a typographic error. It should be *Bastviken et al., 2011*. This is now corrected (new text line 394). *Melack et al., 2004* is now added to the reference list.

*L396-398: It is clear there are issues with emissions estimation that need to be addressed in the future, but it is not clear with the issue is with CH<sub>4</sub> and N<sub>2</sub>O measurement and data calibration. It seems that standard, accurate, techniques are available for collection of samples and subsequently measuring CH<sub>4</sub> and N<sub>2</sub>O concentrations please remove or clarify this point.*

We strongly disagree! To state that *standard, accurate, techniques are available for collection of samples and subsequently measuring CH<sub>4</sub> and N<sub>2</sub>O concentrations* is both naïve and too simplistic. The issue of N<sub>2</sub>O and CH<sub>4</sub> calibration is extremely important and it is far from resolved. There are no currently agreed international calibration standards for these gases, individuals currently making or obtaining their own by various means. To rectify this two of us (RCUG and JB) are members of an international SCOR working group on N<sub>2</sub>O and CH<sub>4</sub>, one of the aims of which is the full standardization of calibration gases and associated protocols. A recent inter-calibration exercise

(results in prep.) demonstrates clearly the discrepancies that can in large part be attributed to errors in calibration. Having said that we have modified the text in question in order to emphasize our point and to publicize the SCOR effort, which encourages dialogue with interested parties. We have therefore deleted: "This is an issue that is yet to be adequately addressed, not only for freshwaters but for aquatic systems more generally". (original text lines 398-399) and replaced it with "There are currently no internationally agreed calibration standards for CH<sub>4</sub> or N<sub>2</sub>O but this is now being addressed via an international SCOR (Scientific Committee on Oceanic Research) Working Group (WG-143: <https://portal.geomar.de/web/scor-wg-143/home>), which is engaged in inter-laboratory calibration and the dissemination of high quality calibration gases. WG-143 welcomes additional interest from the wider aquatic CH<sub>4</sub> and N<sub>2</sub>O research community" (new text lines 408-413).