

## Abstract (with insertions/deletions):

We report on concentrations of dissolved CH<sub>4</sub>, N<sub>2</sub>O, O<sub>2</sub>, NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup>, and corresponding CH<sub>4</sub> and N<sub>2</sub>O emissions for river sites in savannah, swamp forest and tropical forest, along the Congo main stem and in several of its tributary systems of the Western Congo Basin, Republic of Congo (ROC), during November 2010 (41 samples; “wet season”) and August 2011 (25 samples; “dry season”; CH<sub>4</sub> and N<sub>2</sub>O only). Dissolved inorganic nitrogen (DIN: NH<sub>4</sub><sup>+</sup> + NO<sub>3</sub><sup>-</sup>; wet season; NH<sub>4</sub><sup>+</sup> + NO<sub>3</sub><sup>-</sup>) was dominated by NO<sub>3</sub><sup>-</sup> (63 ± 19% of DIN), total DIN concentrations (1.5-45.3 μmol L<sup>-1</sup>) being are consistent with negligible near absence of agricultural, domestic and industrial sources. *Question: Is this true for all the three land types?* Dissolved O<sub>2</sub> (wet season) was mostly under-saturated in swamp forest (36 ± 29%) and tropical forest (77 ± 36%) rivers but predominantly super-saturated in savannah rivers (100 ± 17%). The dissolved concentrations of CH<sub>4</sub> and N<sub>2</sub>O were are within previously reported ranges the range of values reported earlier for sub-Saharan African rivers. While Dissolved CH<sub>4</sub> was always found to be super-saturated (11.2 - 9553 nmol L<sup>-1</sup>; 440-354400% *Comment: Check this number*); whereas N<sub>2</sub>O ranged from strong under-saturation to strong super-saturation (3.2-20.6 nmol L<sup>-1</sup>; 47-205%). Evidently, rivers of the ROC are persistent local sources of tropospheric CH<sub>4</sub> but and can be small a minor sources or sinks for N<sub>2</sub>O. During the dry season, concentration means and ranges of CH<sub>4</sub> and N<sub>2</sub>O concentrations were indistinguishable quite similar for all the three land types; and whereas seasonal differences in the means and ranges were not significant for N<sub>2</sub>O concentration for in any land type or for CH<sub>4</sub> in savannah rivers. The latter observation is consistent with seasonal buffering of river discharge by an underlying sandy-sandstone aquifer. By In contrast, CH<sub>4</sub> concentration in swamp and forest rivers CH<sub>4</sub> was significantly higher in the wet season, possibly reflecting suggesting that CH<sub>4</sub> can be derived from floating macrophytes during flooding and/or enhanced methanogenesis in adjacent flooded soils in flood bank. Swamp rivers also exhibited both low (47%) and high (205%) N<sub>2</sub>O saturations but wet season values were overall significantly lower than in either tropical forest or savannah rivers. These which rivers were always super-saturated (103-266%) and for which the overall means and ranges of N<sub>2</sub>O were not significantly different. In swamp and forest rivers % O<sub>2</sub>-saturation (%) co-varied negatively inversely with log % CH<sub>4</sub> saturation (log %) and positively linearly with % N<sub>2</sub>O saturation (%). The strong A significant positive correlation for N<sub>2</sub>O - O<sub>2</sub> correlation saturation in swamp rivers was coincident is consistent with strong N<sub>2</sub>O and O<sub>2</sub> under-saturation, indicating N<sub>2</sub>O consumption by during denitrification in the sediments. denitrification. In savannah rivers persistent N<sub>2</sub>O super-saturation and a negative N<sub>2</sub>O - O<sub>2</sub> correlation may indicate suggest N<sub>2</sub>O production mainly by nitrification, consistent with a stronger significant correlation between N<sub>2</sub>O and NH<sub>4</sub><sup>+</sup> than between N<sub>2</sub>O and NO<sub>3</sub><sup>-</sup>. Our range in of values for CH<sub>4</sub> and N<sub>2</sub>O emissions fluxes (33-48705 μmol CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>; 1-67 μmol N<sub>2</sub>O m<sup>-2</sup> d<sup>-1</sup>), is are within the range previously estimated for sub-Saharan African rivers but it includes and associated with uncertainties deriving arising from our use of “basin-wide” values for CH<sub>4</sub> and N<sub>2</sub>O gas transfer velocities. Even so, because Furthermore, as we did not account for any contribution from ebullition, which is quite likely for CH<sub>4</sub> is likely to be (at least 20%), our emissions fluxes estimates for CH<sub>4</sub> are probably rather conservative estimates.

**Abstract** (with all make up):

We report on concentrations of dissolved CH<sub>4</sub>, N<sub>2</sub>O, O<sub>2</sub>, NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup>, and corresponding CH<sub>4</sub> and N<sub>2</sub>O emissions for river sites in savannah, swamp forest and tropical forest, along the Congo main stem and in several of its tributary systems of the Western Congo Basin, Republic of Congo (ROC), during November 2010 (41 samples; “wet season”) and August 2011 (25 samples; “dry season”; CH<sub>4</sub> and N<sub>2</sub>O only). Dissolved inorganic nitrogen (DIN: NH<sub>4</sub><sup>+</sup> + NO<sub>3</sub><sup>-</sup>; wet season) was dominated by NO<sub>3</sub><sup>-</sup> (63 ± 19% of DIN), total DIN concentrations (1.5-45.3 μmol L<sup>-1</sup>) are consistent with near absence of agricultural, domestic and industrial sources. *Question: Is this true for all the three land types?* Dissolved O<sub>2</sub> (wet season) was mostly under-saturated in swamp forest (36 ± 29%) and tropical forest (77 ± 36%) rivers but predominantly super-saturated in savannah rivers (100 ± 17%). The dissolved concentrations of CH<sub>4</sub> and N<sub>2</sub>O are within the range of values reported earlier for sub-Saharan African rivers. Dissolved CH<sub>4</sub> was found to be super-saturated (11.2 - 9553 nmol L<sup>-1</sup>; 440-354400% *Comment: Check this number*); whereas N<sub>2</sub>O ranged from strong under-saturation to super-saturation (3.2-20.6 nmol L<sup>-1</sup>; 47-205%). Evidently, rivers of the ROC are persistent local sources of CH<sub>4</sub> and can be a minor source or sink for N<sub>2</sub>O. During the dry season, mean and range of CH<sub>4</sub> and N<sub>2</sub>O concentrations were quite similar for all the three land types; whereas seasonal differences in the mean and range were not significant for N<sub>2</sub>O concentration in any land type or for CH<sub>4</sub> in savannah rivers. The latter observation is consistent with seasonal buffering of river discharge by an underlying sand-sandstone aquifer. In contrast, CH<sub>4</sub> concentration in swamp and forest rivers was significantly higher in the wet season, suggesting that CH<sub>4</sub> can be derived from floating macrophytes during flooding and/or enhanced methanogenesis in adjacent flooded soils. Swamp rivers also exhibit both low (47%) and high (205%) N<sub>2</sub>O saturation but wet season values were overall significantly lower than in either tropical forest or savannah rivers. These rivers were always super-saturated (103-266%) and for which the overall mean and range of N<sub>2</sub>O were not significantly different. In swamp and forest rivers O<sub>2</sub>-saturation (%) varied inversely with CH<sub>4</sub> saturation (log %) and linearly with N<sub>2</sub>O saturation (%). A significant positive correlation for N<sub>2</sub>O - O<sub>2</sub> saturation in swamp rivers is consistent with N<sub>2</sub>O and O<sub>2</sub> under-saturation, indicating N<sub>2</sub>O consumption during denitrification in the sediments. In savannah rivers persistent N<sub>2</sub>O super-saturation and a negative N<sub>2</sub>O - O<sub>2</sub> correlation suggest N<sub>2</sub>O production mainly by nitrification, consistent with significant correlation between N<sub>2</sub>O and NH<sub>4</sub><sup>+</sup> than between N<sub>2</sub>O and NO<sub>3</sub><sup>-</sup>. Our range of values for CH<sub>4</sub> and N<sub>2</sub>O emission fluxes (33-48705 μmol CH<sub>4</sub> m<sup>-2</sup> d<sup>-1</sup>; 1-67 μmol N<sub>2</sub>O m<sup>-2</sup> d<sup>-1</sup>) are within the range previously estimated for sub-Saharan African rivers and associated with uncertainties arising from our use of “basin-wide” values for CH<sub>4</sub> and N<sub>2</sub>O gas transfer velocities. Furthermore, as we did not account for any contribution from ebullition, which is quite likely for CH<sub>4</sub> (at least 20%), our emission fluxes for CH<sub>4</sub> are rather conservative estimates.