

Response to Editors comments:

Dear Editor

Thank you very much for your careful reading of our manuscript. We have changed the language edits you suggested and the manuscript has been proofread by a professional language editor. We hope that you are pleased with our revised manuscript.

Regarding the redundancy you mentioned in the method section. As we used bootstrap simulations both in our analysis of coupling between temporal and spatial dynamics in F_{opt} and a , and in the parameterisation and evaluation of the GPP model, we have to mention this in both instances. We have, however, removed as much redundant text as we feel is possible.

Yours sincerely,

Torbern Tagesson and co-authors

1 **Modelling spatial and temporal dynamics of GPP in the Sahel from
2 earth observation based photosynthetic capacity and quantum
3 efficiency**

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20 **Abstract.** It has been shown that vegetation growth in semi-arid regions is important to the global terrestrial CO₂ sink,
21 which indicates the strong need for improved understanding and spatially explicit estimates of CO₂ uptake (gross
22 primary production (GPP)) in semi-arid ecosystems. This study has three aims: 1) to evaluate the MOD17A2H GPP
23 (collection 6) product against eddy covariance (EC) based GPP for six sites across the Sahel; 2) to characterize
24 relationships between spatial and temporal variability in EC based photosynthetic capacity (F_{opt}) and quantum
25 efficiency (α) and earth observation (EO) based vegetation indices (normalized difference vegetation index (NDVI);
26 renormalized difference vegetation index (RDVI); enhanced vegetation index (EVI); and shortwave infrared water
27 stress index (SIWSI)); and 3) to study the applicability of EO up-scaled F_{opt} and α for GPP modelling purposes.
28 MOD17A2H GPP (collection 6) drastically underestimated GPP strongly, most likely because maximum light use
29 efficiency is set too low for semi-arid ecosystems in the MODIS algorithm. Intra-annual dynamics in F_{opt} was were
30 closely related to SIWSI being sensitive to equivalent water thickness, whereas α was closely related to RDVI being
31 affected by chlorophyll abundance. Spatial and inter-annual dynamics in F_{opt} and α were closely coupled to NDVI and
32 RDVI, respectively. Modelled GPP based on F_{opt} and α up-scaled using EO based indices reproduced in situ GPP well
33 for all except a cropped site that was strongly impacted by anthropogenic land use. Up-scaled GPP for the Sahel 2001-
34 2014 was 736±39 g C m⁻² y⁻¹. This study indicates the strong applicability of EO as a tool for spatially explicit estimates
35 of GPP, F_{opt} and α ; incorporating EO-based F_{opt} and α in to-dynamic global vegetation models could improve global
36 estimates of vegetation production, ecosystem processes and biogeochemical and hydrological cycles.

37

38 **Keywords:** Remote sensing, Gross Primary Productivity, MOD17A2H, light use efficiency, photosynthetic capacity,
39 quantum efficiency

40 **1 Introduction**

41 Vegetation growth in semi-arid regions is an important sink for fossil fuel emissions. Mean carbon dioxide (CO₂)
42 uptake by terrestrial ecosystems is dominated by highly productive lands, mainly tropical forests, whereas semi-arid
43 regions are the main biome driving its inter-annual variability (Ahlström et al., 2015; Poulter et al., 2014). Semi-arid
44 regions even contribute to 60% of the long-term trend in the global terrestrial C sink (Ahlström et al., 2015). It is thus
45 important to understand long-term variability of vegetation growth in semi-arid areas and ~~their~~ the response of
46 vegetation to environmental conditions to better quantify and forecast effects of climate change.

47 The Sahel is a semi-arid transition zone between the dry Sahara desert in the North and the humid Sudanian savanna
48 in the South. The region has experienced numerous severe droughts ~~during~~over the last decades, ~~which~~ that resulted in
49 region-wide famines in 1972–1973 and 1984–1985 and localized food shortages across the region in 1990, 2002, 2004,
50 2011 and 2012 (Abdi et al., 2014; United Nations, 2013). Vegetation production is thereby an important ecosystem
51 service for livelihoods in the Sahel, but it is under threat. The region ~~is~~ experiencing ~~a~~ strong population growth,
52 increasing the demand on ecosystem services due to cropland expansion, increased pasture stocking rates and fuelwood
53 extraction (Abdi et al., 2014).

54 At the same time as we have reports of declining vegetation production, we have contradicting reports of the greening
55 of the Sahel based on earth observation (EO) data (Dardel et al., 2014; Fensholt et al., 2013). The greening of the Sahel
56 has mainly been attributed to alleviated drought stress conditions due to increased precipitation since the mid-1990s
57 (Hickler et al., 2005). Climate is thus another important factor regulating vegetation production. Semi-arid regions, such
58 as the Sahel, are particularly vulnerable to climate fluctuations due to their dependency on moisture.

59 Estimation of gross primary production (GPP), i.e. uptake of atmospheric CO₂ by vegetation, is still a major challenge
60 ~~within~~for the remote sensing of ecosystem services. Gross primary production is a main driver of ecosystem services
61 such as climate regulation, carbon (C) sequestration, C storage, food production, ~~or~~and livestock grassland production.
62 Within EO, spatial quantification of GPP generally involves light use efficiency (LUE), defined as the conversion
63 efficiency of absorbed solar light into CO₂ uptake (Monteith, 1972, 1977). It has been shown that LUE varies in space
64 and time due to factors such as plant functional type, drought and temperature, nutrient levels and physiological
65 limitations of photosynthesis (Garbulsky et al., 2010; Paruelo et al., 2004; Kergoat et al., 2008). The LUE concept has
66 been applied ~~using~~through various methods, either by using a biome-specific LUE constant (Ruimy et al., 1994); or by
67 modifying a maximum LUE using meteorological variables (Running et al., 2004).

68 An example of ~~a~~ an LUE based model is the standard GPP product from the Moderate Resolution Imaging
69 Spectroradiometer (MODIS) sensor (MOD17A2). Within the model, absorbed photosynthetically active radiation
70 (PAR) is estimated as a product of the fraction of PAR absorbed by green vegetation (FPAR from MOD15A2)
71 multiplied with daily PAR from the meteorological data of the Global Modeling and Assimilation Office (GMAO). A
72 set of maximum LUE parameters specified for each biome are extracted from a Biome Properties Look-Up Table
73 (BPLUT). Then maximum LUE is modified depending on air temperature (T_{air}) and vapour pressure deficit (VPD)
74 (Running et al., 2004). Sjöström et al. (2013) evaluated the MOD17A2 product (collection 5.1) for Africa~~s~~ and showed

75 that it underestimated GPP for semi-arid savannas in the Sahel. Explanations for this underestimation were that the
76 assigned maximum LUE from BPLUT was set too low and that there were uncertainties in the FPAR product
77 (MOD15A2). Recently, a new collection of MOD17A2 at a 500 m spatial resolution was released (MOD17A2H;
78 collection 6) with an updated BPLUT, updated GMAO meteorological data, improved quality control and gap-filling
79 of the FPAR data from MOD15A2 (Running and Zhao, 2015).

80 It has been shown that the LUE method does not perform well in arid conditions and at agricultural sites (Turner et
81 al., 2005). Additionally, the linearity assumed by the LUE model is not usually not found as the response of GPP to
82 incoming light follows more of an asymptotic curve (Cannell and Thornley, 1998). Investigating other methods for
83 remotely determining GPP is thus of great importance, especially for semi-arid environments. Therefore, instead of
84 LUE, we focus on the light response function of GPP at the canopy scale, and spatial and temporal variation of its two
85 main parameters: maximum GPP under light saturation (canopy-scale photosynthetic capacity; F_{opt})⁵ and the initial
86 slope of the light response function (canopy-scale quantum efficiency; α) (Falge et al., 2001; Tagesson et al., 2015a).
87 Photosynthetic capacity is a measure of the maximum rate at which the canopy can fix CO₂ during photosynthesis
88 ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$), whereas α is the amount of CO₂ fixed per incoming PAR ($\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$). ~~Just~~⁴ To clarify
89 the difference in LUE and α in this study⁶, LUE ($\mu\text{mol CO}_2 \mu\text{mol APAR}^{-1}$) is the slope of a linear fit between CO₂
90 uptake and absorbed PAR, whereas α ($\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$) is the initial slope of an asymptotic curve against
91 incoming PAR.

92 It has been proven that F_{opt} and α are closely related to chlorophyll abundance due to their coupling with the electron
93 transport rate (Ide et al., 2010). Additionally, in semi-arid ecosystems⁷, water availability is generally considered to be
94 the main limiting factor affecting intra-annual dynamics of vegetation growth (Fensholt et al., 2013; Hickler et al.,
95 2005; Tagesson et al., 2015b). Several remote sensing studies have established relationships between remotely sensed
96 vegetation indices and ecosystem properties such as chlorophyll abundance and equivalent water thickness (Yoder and
97 Pettigrew-Crosby, 1995; Fensholt and Sandholt, 2003). In this study⁸, we will analyse ~~if~~whether EO vegetation indices
98 can be used ~~for~~to up-scaling F_{opt} and α and investigate ~~if~~whether this could offer a promising way to map GPP in semi-
99 arid areas. This potential will be analysed by the use of detailed ground observations from six eddy covariance (EC)
100 flux tower sites across the Sahel.

101 The three aims of this study are:

- 102 1) To investigate ~~if~~whether the recently released MOD17A2H GPP (collection 6) product is better at capturing
103 GPP for the Sahel than collection 5.1. We hypothesise that the MOD17A2H GPP (collection 6) product will
104 estimate GPP well for the six Sahelian EC sites⁹, because of major changes ~~done~~made in comparison to
105 collection 5.1 (Running and Zhao, 2015).
- 106 2) To characterize the relationships between spatial and temporal variability in F_{opt} and α and remotely sensed
107 vegetation indices. We hypothesise that EO vegetation indices that are closely related to chlorophyll
108 abundance will be most strongly coupled with spatial and inter-annual dynamics in F_{opt} and α , whereas
109 vegetation indices closely related to equivalent water thickness will be most strongly coupled with intra-annual
110 dynamics in F_{opt} and α across the Sahel.
- 111 3) To evaluate the applicability of a GPP model based on the light response function using EO vegetation indices
112 and incoming PAR as input data.

114 **2 Materials and Methods**

115 **2.1 Site description**

116 The Sahel stretches from the Atlantic Ocean in the west to the Red Sea in the east. The northern border towards the
117 Sahara and the southern border towards the humid Sudanian Savanna are defined by the 150 and 700 mm isohyets,
118 respectively (Fig. 1) (Prince et al., 1995). Tree and shrub canopy cover is now generally low (< 5%) and dominated by
119 species of *Balanites*, *Acacia*, *Boscia* and *Combretaceae* (Rietkerk et al., 1996). Annual grasses such as *Schoenfeldia*
120 *gracilis*, *Dactyloctenium aegypticum*, *Aristida mutabilis*, and *Cenchrus biflorus* dominate the herbaceous layer, but
121 perennial grasses such as *Andropogon gayanus*, *Cymbopogon schoenanthus* can also be found (Rietkerk et al., 1996; de
122 Ridder et al., 1982). From the FLUXNET database (Balocchi et al., 2001) we selected the six available measurement
123 sites with EC based CO₂ flux data from the Sahel (Table 1; Fig. 1). The sites represent a variety of ecosystems present
124 in the region, from dry fallow bush savanna to seasonally inundated acacia forest. For a full description of the
125 measurement sites, we refer to Tagesson et al. (2016a) and references in Table 1.

126 <Table 1>

127 <Figure 1>

128

129 **2.2 Data collection**

130 **2.2.1 Eddy covariance and hydrometeorological in situ data**

131 Eddy covariance and hydrometeorological data originating from the years between 2005 and 2013 were collected from
132 the principal investigators of the measurement sites (Tagesson et al., 2016a). The EC sensor set-up consisted of open-
133 path CO₂/H₂O infrared gas analysers and 3-axis sonic anemometers. Data were collected at 20 Hz rate and statistics
134 were calculated for 30-minute periods. For a full description of the sensor set-up and post processing of EC data, see the
135 references in Table 1. Final fluxes were filtered according to quality flags provided by FLUXNET and outliers were
136 filtered according to Papale et al. (2006). We extracted the original net ecosystem exchange (NEE) data without any
137 gap-filling or partitioning of NEE to GPP and ecosystem respiration. The collected hydrometeorological data were: air
138 temperature (T_{air}; °C), rainfall (P; mm), relative air humidity (Rh; %), soil moisture at 0.1 m depth (SWC; % volumetric
139 water content), incoming global radiation (R_g; W m⁻²), incoming photosynthetically active radiation (PAR; μmol m⁻² s⁻¹),
140 VPD (hPa), peak dry weight biomass (g dry weight m⁻²), C3/C4 species ratio, and soil conditions (nitrogen and C
141 concentration; %). For a full description of the collected data and sensor set-up, see Tagesson et al. (2016a).

142

143 **2.2.2 Earth Observation data and gridded ancillary data**

144 Composite products from MODIS/Terra covering the Sahel were acquired at Reverb ECHO (NASA, 2016). Collected
145 products were GPP (MOD17A2H; collection 6), nadir bidirectional reflectance distribution function adjusted
146 reflectance (NBAR) (8-day composites; MCD43A4; collection 5.1) at 500×500 m² spatial resolution, the normalized
147 difference vegetation index (NDVI) and the enhanced vegetation index (EVI) (16-day composites; MOD13Q1;
148 collection 6) at 250×250 m² spatial resolution. The NBAR product was preferred over the reflectance product
149 (MOD09A1), in order to avoid variability caused by varying sun and sensor viewing geometry (Huber et al., 2014;
150 Tagesson et al., 2015c). We extracted the median of 3x3 pixels centred at the location of each EC tower. Time series of
151 EO products were filtered according to MODIS quality control data; MOD17A2H is a gap-filled and filtered product,

152 QC data from MCD43A2 were used for filtering of MCD43A4; and bit 2-5 (highest –decreasing quality) was used for
153 MOD13Q1. Finally, data were gap-filled to daily values using linear interpolation.

154 We downloaded ERA Interim reanalysis PAR at the ground surface (W m^{-2}) with a spatial resolution of $0.25^\circ \times 0.25^\circ$
155 accumulated for each 3-hour period from 2000-2015 from the European Centre for Medium-Range Weather Forecasts
156 (ECMWF) (Dee et al., 2011; ECMWF, 2016a).

157

158 **2.3 Data handling**

159 **2.3.1 Intra-annual dynamics in photosynthetic capacity and quantum efficiency**

160 To estimate daily values of EC based F_{opt} and α , the asymptotic Mitscherlich light-response function was fitted
161 between daytime NEE and incoming PAR using a 7-day moving window with a 1-day time step:

$$162 \text{NEE} = -(F_{\text{opt}}) \times (1 - e^{\left(\frac{-\alpha \times \text{PAR}}{F_{\text{opt}}} \right)}) + R_d \quad (1)$$

163 where F_{opt} is CO_2 uptake at light saturation (photosynthetic capacity; $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$), R_d is dark respiration
164 ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$), and α is the initial slope of the light response curve (quantum efficiency; $\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$)
165 (Falge et al., 2001). By subtracting R_d from Eq. 1, the function was forced through zero and GPP was thereby
166 estimated. To assureensure a high quality of fitted parameters, parameters were excluded from the analysis when fitting
167 was insignificant (p -value > 0.05), and when they were out of range (F_{opt} and $\alpha >$ peak value of the rainy season times
168 1.2). Additionally, outliers were filtered following the method by Papale et al. (2006) using a 30-day moving window
169 with a 1-day time step.

170

171 **2.3.2 Vegetation indices**

172 The maximum absorption in red wavelengths generally occurs at 682 nm as this is the peak absorption for chlorophyll a
173 and b (Thenkabail et al., 2000), which makes vegetation indices that include the red band sensitive to chlorophyll
174 abundance. By far the most common vegetation index is NDVI (Rouse et al., 1974):

$$175 \text{NDVI} = \frac{(\rho_{\text{NIR}} - \rho_{\text{red}})}{(\rho_{\text{NIR}} + \rho_{\text{red}})} \quad (2)$$

176 where ρ_{NIR} is the reflectance factor in the near infrared (NIR) band (band 2) and ρ_{red} is the reflectance factor in the red
177 band (band 1). Near infrared radiance is reflected by leaf cells since absorption of these wavelengths would result in
178 overheating of the plant, whereas red radiance is absorbed by chlorophyll and its accessory pigments (Gates et al.,
179 1965). Normalization is done to reduce effects of atmospheric errors, solar zenith angles, and sensor viewing geometry,
180 as well as to increasing the vegetation signal (Qi et al., 1994; Inoue et al., 2008).

181 A well-known deficiency of NDVI is problems of index saturation at high biomass because absorption of red light at
182 ~ 670 nm peaks at higher biomass loads, whereas NIR reflectance continues to increase due to multiple scattering effects
183 (Mutanga and Skidmore, 2004; Jin and Eklundh, 2014). By reducing atmospheric and soil background influences, EVI
184 is designed to increase the signal from the vegetation and maintain sensitivity in high biomass regions (Huete et al.,
185 2002).

$$186 \text{EVI} = G \frac{(\rho_{\text{NIR}} - \rho_{\text{red}})}{(\rho_{\text{NIR}} + C_1 \rho_{\text{red}} - C_2 \rho_{\text{blue}} + L)} \quad (3)$$

187 where ρ_{blue} is the reflectance factor in the blue band (band 3). The coefficients $C_1=6$ and $C_2=7.5$ correct for atmospheric
188 influences, while $L=1$ adjusts for the canopy background. The factor $G=2.5$ is a gain factor.

189 Another attempt to overcome problems of NDVI saturation was proposed by Roujean and Breon (1995), who
190 suggested the renormalized difference vegetation index (RDVI), ~~that which~~ combines advantages of DVI (NIR-red) and
191 NDVI for low and high vegetation cover, respectively:

$$192 \quad RDVI = \frac{(\rho_{NIR} - \rho_{red})}{\sqrt{(\rho_{NIR} + \rho_{red})}} \quad (4)$$

193 As a non-linear index, RDVI is not only less sensitive to variations in geometrical and optical properties of unknown
194 foliage but also less affected by solar and viewing geometry (Broge and Leblanc, 2001). The vegetation index RDVI
195 was calculated based on NBAR bands 1 and 2.

196 The NIR and SWIR bands are affected by the same ground properties, except that SWIR bands are also strongly
197 sensitive to equivalent water thickness. Fensholt and Sandholt (2003) proposed a vegetation index, the shortwave
198 infrared water stress index (SIWSI), using NIR and SWIR bands to estimate drought stress for vegetation in semi-arid
199 environments:

$$200 \quad SIWSI_{12} = \frac{(\rho_{NIR} - \rho_{SWIR12})}{(\rho_{NIR} + \rho_{SWIR12})} \quad (5)$$

$$201 \quad SIWSI_{16} = \frac{(\rho_{NIR} - \rho_{SWIR16})}{(\rho_{NIR} + \rho_{SWIR16})} \quad (6)$$

202 where ρ_{swir12} is NBAR band 5 (1230-1250 nm) and ρ_{swir16} is NBAR band 6 (1628-1652 nm). As the vegetation water
203 content increases, reflectance in SWIR decreases, indicating that low and high SIWSI values point to sufficient water
204 conditions and drought stress, respectively.

205 | 2.3.3 Incoming PAR across ~~the~~ Sahel

206 A modified version of the ERA Interim reanalysis PAR was used in the current study as there was an error in the code
207 producing these PAR estimates; the estimates were generally too low (ECMWF, 2016b). Accordingly, incoming PAR
208 at the ground surface from ERA Interim was systematically underestimated even though it followed the pattern of PAR
209 measured at the six Sahelian EC sites (Fig. S1 in supplementary material). In order to correct for this error, we fitted
210 and applied an ordinary least squares linear regression between in situ PAR and ERA Interim PAR (Fig. S1). The PAR
211 produced from this relationship is at the same level as in situ PAR and should be at a correct level even though the
212 original ERA Interim PAR is actually produced from the red and near infrared part of the spectrum.

213 | 2.4 Data analysis

214 | 2.4.1 Coupling temporal and spatial dynamics in photosynthetic capacity and quantum efficiency with 215 | explanatory variables

216 The coupling between intra-annual dynamics in F_{opt} and α and the vegetation indices for the different measurement sites
217 were studied using Pearson correlation analysis. As part of the correlation analysis, we used a bootstrap simulation
218 methodology with 200 iterations from which ~~the~~ mean and ~~the~~ standard deviation of the correlation coefficients were
219 calculated (Richter et al., 2012). Relationships between intra-annual dynamics in F_{opt} and α and the vegetation indices

222 for all sites combined were also analysed. In the analysis for all sites, data were normalized ~~in order~~ to avoid influence
 223 of spatial and inter-annual variability. Time series of ratios of F_{opt} and α (F_{opt_frac} and α_{frac}) against the annual peak
 224 values (F_{opt_peak} and α_{peak} ; see below for calculation of annual peak values) were estimated for all sites:

$$225 \quad F_{opt_frac} = \frac{F_{opt}}{F_{opt_peak}} \quad (7)$$

$$226 \quad \alpha_{frac} = \frac{\alpha}{\alpha_{peak}} \quad (8)$$

227 The same standardization procedure was used for all vegetation indices (VI_{frac}):

$$228 \quad VI_{frac} = \frac{VI}{VI_{peak}} \quad (9)$$

229 where VI_{peak} is the annual peak values of the vegetation indices (14-day running mean with highest annual value). The
 230 α_{frac} and F_{opt_frac} were correlated with the different VI_{frac} to investigate the coupling between intra-annual dynamics in
 231 F_{opt} and α and the vegetation indices for all sites.

232 Regression trees were used to fill gaps in the daily estimates of F_{opt} and α . One hundred tree sizes were chosen based
 233 on 100 cross-validation runs, and these trees were then used ~~for~~ to estimate F_{opt} and α following the method in
 234 De'ath and Fabricius (2000). We used SWC, VPD, T_{air} , PAR, and the vegetation index with the strongest correlation
 235 with intra-annual dynamics as explanatory variables in the analysis. In the analysis for all sites, the same
 236 standardization procedure as done for F_{opt} , α , and the vegetation indices was done for the hydrometeorological
 237 variables. The 100 F_{opt} and α output subsets from the regression trees were averaged and used for filling gaps in the
 238 times series of F_{opt} and α . From these time-series, we estimated annual peak values of F_{opt} and α (F_{opt_peak} and α_{peak}) as
 239 the 14-day running mean with the highest annual value. To investigate spatial and inter-annual variability in F_{opt} and α
 240 across the measurement sites of the Sahel, F_{opt_peak} and α_{peak} were correlated with the annual sum of P_{eq} , yearly means of
 241 T_{air} , SWC, RH, VPD, and R_g ; annual peak values of biomass, soil nitrogen and C concentrations; the C3/C4 ratio;
 242 and VI_{peak} .

243

244 **2.4.2 Parameterization and evaluation of the GPP model and evaluation of the MODIS GPP**

245 ~~Based on~~ On the basis of Eq. 1 and the outcome of the statistical analysis previously described under subsection 2.4.1 (for
 246 results, see subsect. 3.2), a model for estimating GPP across the Sahel was created:

$$247 \quad GPP = -F_{opt} \times \left(1 - e^{\left(\frac{-\alpha \times PAR}{F_{opt}}\right)}\right) \quad (10)$$

248 Firstly, F_{opt_peak} and α_{peak} were estimated spatially and inter-annually using linear regression functions fitted against the
 249 vegetation indices with strongest relationships to spatial and inter-annual variability in F_{opt_peak} and α_{peak} for all sites.
 250 Secondly, exponential regression functions were established for F_{opt_frac} and α_{frac} with the vegetation index with the
 251 strongest relationships to intra-annual variability of F_{opt_frac} and α_{frac} for all sites. By combining these relationships, F_{opt}
 252 and α can be calculated for any day of year and for any point in space across the Sahel:

$$253 \quad F_{opt} = F_{opt_peak} \times F_{opt_frac} = \left(k_{F_{opt}} \times NDVI_{peak} + m_{F_{opt}} \right) \left(n_{F_{opt}} \times e^{(l_{F_{opt}} \times RDVI_{frac})} \right) \quad (11)$$

254
$$\alpha = \alpha_{\text{peak}} \times \alpha_{\text{frac}} = (k_a \times \text{RDVI}_{\text{peak}} + m_a) (n_a \times e^{(l_a \times \text{RDVI}_{\text{frac}})}) \quad (12)$$

255 where k_{Fopt} and k_a are slopes and m_{Fopt} and m_a are intercepts of the linear regressions giving $F_{\text{opt_peak}}$ and α_{peak} ,
 256 respectively; l_{Fopt} and l_a are coefficients and n_{Fopt} and n_a are intercepts of the exponential regressions giving $F_{\text{opt_frac}}$ and
 257 α_{frac} , respectively. Equations 11 and 12 were inserted into Eq. 10, and GPP ~~were~~ was thereby estimated as:

258
$$\text{GPP} = - (F_{\text{opt_peak}} \times F_{\text{opt_frac}}) \times (1 - e^{\left(\frac{-(\alpha_{\text{peak}} \times \alpha_{\text{frac}}) \times \text{PAR}}{F_{\text{opt_peak}} \times F_{\text{opt_frac}}} \right)}) = - \left((k_{\text{Fopt}} \times \text{NDVI}_{\text{peak}} + m_{\text{Fopt}}) (n_{\text{Fopt}} \times e^{(l_{\text{Fopt}} \times \text{RDVI}_{\text{frac}})}) \right) \times 1 - e^{\left(\frac{-(k_a \times \text{RDVI}_{\text{peak}} + m_a) (n_a \times e^{(l_a \times \text{RDVI}_{\text{frac}})}) \times \text{PAR}}{(k_{\text{Fopt}} \times \text{NDVI}_{\text{peak}} + m_{\text{Fopt}}) (l_{\text{Fopt}} \times \text{RDVI}_{\text{frac}} + n_{\text{Fopt}})} \right)} \quad (13)$$

259 The bootstrap simulation methodology was used when fitting the least-squares regression functions for
 260 parameterization of the GPP model (Richter et al., 2012). For each of the iterations, some of the EC sites were included
 261 and some were ~~left out~~ omitted. The bootstrap simulations generated 200 sets of k_{Fopt} , k_a , m_{Fopt} , m_a , l_{Fopt} , l_a , n_{Fopt} , n_a ,
 262 and coefficient of determination (R^2). Possible errors (e.g. random sampling errors, aerosols, electrical sensor noise,
 263 filtering and gap-filling errors, clouds, and satellite sensor degradation) can be present in both the predictor and the
 264 response variables. Hence, we selected reduced major axis regressions to account for errors in both predictor and
 265 response variables when fitting the regression functions. The regression models were validated against the ~~left-~~
 266 ~~out~~ omitted sites within the bootstrap simulation methodology by calculating the root-mean-square error (RMSE), and
 267 by fitting an ordinary least squares linear regression between modelled and independent variables.

268 Similarly, the MODIS GPP product (MOD17A2H~~1~~ collection 6) was evaluated against independent GPP from the EC
 269 sites by calculating ~~the~~ RMSE, and by fitting an ordinary least squares linear regression.

270 **3 Results**

271 **3.1 Evaluation of the MODIS GPP product**

272 There was a strong linear relationship between the MODIS GPP product (MOD17A2H; collection 6) and independent
 273 GPP (slope=0.17; intercept=0.11 g C m⁻² d⁻¹; R²=0.69; n=598). However, MOD17A2H strongly underestimated
 274 independent GPP (Fig. 2), resulting in a high RMSE (2.69 g C m⁻² d⁻¹). It can be seen that some points for the Kelma
 275 site were quite low for MOD17A2H, whereas they were relatively high for the independent GPP (Fig. 2). Kelma is an
 276 inundated Acacia forest located in a clay soil depression. These differentiated values were found in the beginning of the
 277 dry season, when the depression was still inundated, whereas the larger area was turning dry.

278 <Figure 2>

279 **3.2 Intra-annual dynamics in photosynthetic capacity and quantum efficiency**

280 Intra-annual dynamics in F_{opt} and α differed in amplitude, but were otherwise similar across the measurement sites in
 281 ~~the~~ Sahel (Fig. 3). There was no green ground vegetation during the dry season, and the low photosynthetic activity was
 282 due to few evergreen trees. This resulted in low values for both F_{opt} and α during the dry season. The vegetation
 283 responded strongly to rainfall, and both F_{opt} and α increased during the early phase of the rainy season. Generally, F_{opt}
 284 peaked slightly earlier than α (average \pm 1 standard deviation: 7 \pm 10 days) (Fig. 3).

287 <Figure 3>
288 All vegetation indices described intra-annual dynamics in F_{opt} reasonably well at all sites (Table 2). The vegetation
289 index SIWSI₁₂ had the highest correlation for all sites except Wankama Millet, where it was RDVI. When all sites were
290 combined, all indices described well seasonality in F_{opt} , but RDVI had the strongest correlation (Table 2).

291 Intra-annual dynamics in α were also closely coupled to intra-annual dynamics in the vegetation indices for all sites
292 (Table 2). For α , RDVI was the strongest index describing intra-annual dynamics, except for Wankama Fallow, where it
293 was EVI. When all sites were combined, all indices described well intra-annual dynamics in α , but RDVI was still the
294 index with strongest relationship (Table 2).

295 <Table 2>
296 The regression trees used for gap-filling explained the intra-annual dynamics in F_{opt} and α well for all sites (Table 3;
297 Fig. S2 in Supplementary material). The regression trees explained intra-annual dynamics in F_{opt} better than in α , and
298 multi-year sites were better predicted than single year sites (Fig. S2). The main explanatory variables coupled to intra-
299 annual dynamics in F_{opt} for all sites across the Sahel were in the order of RDVI, SWC, VPD, T_{air} , and PAR; and for α
300 they were RDVI, SWC, VPD and T_{air} (Table 3). The strong relationship to SWC and VPD indicates drought stress
301 during periods of low rainfall. For all sites across the Sahel, incorporating hydrometeorological variables increased the
302 ability to determine intra-annual dynamics in F_{opt} and α compared to the ordinary least squares linear regressions against
303 vegetation indices (Table 2, data given as r ; Table 3; Fig. 3 and Fig. S2). For all sites, incorporation of these variables
304 increased R^2 from 0.81 to 0.87 and from 0.74 to 0.84 for F_{opt} and α , respectively.

305 <Table 3>

306
307 **3.3 Spatial and inter-annual dynamics in photosynthetic capacity and quantum efficiency**
308 Large spatial and inter-annual variability in F_{opt_peak} and α_{peak} were found across the six measurement sites in the Sahel;
309 F_{opt_peak} ranged between 10.1 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (Wankama Millet 2005) and 50.0 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ (Dahra 2010), and
310 α_{peak} ranged between 0.020 $\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$ (Demekeya 2007) and 0.064 $\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$ (Dahra 2010)
311 (Table 4). The average two-week running mean peak values of F_{opt} and α for all sites were 26.4 $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ and
312 0.040 $\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$, respectively. All vegetation indices determined spatial and inter-annual dynamics well in
313 both F_{opt_peak} and α_{peak} (Table 5); F_{opt_peak} was most closely coupled with $NDVI_{peak}$, whereas α_{peak} was coupled more
314 closely with $RDVI_{peak}$ (Fig. 4). F_{opt_peak} also correlated well with peak dry weight biomass, C content in the soil, and
315 RH, whereas α_{peak} also correlated well with peak dry weight biomass, and C content in the soil (Table 5).

316 <Table 4>

317 <Table 5>

318 <Figure 4>

319
320 **3.4 Spatially extrapolated photosynthetic capacity, quantum efficiency, and gross primary production across the**
321 **Sahel and evaluation of the GPP model**

322 The spatially extrapolated F_{opt} , α and GPP averaged over the Sahel for 2001-2014 were $22.5 \pm 1.7 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$,
323 $0.030 \pm 0.002 \mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$, and $736 \pm 39 \text{ g C m}^{-2} \text{ y}^{-1}$, respectively. At a regional scale, it can be seen that F_{opt} ,
324 α , and GPP decreased substantially with latitude (Fig. 5). The highest values were found in south-eastern Senegal,
325 western Mali, in parts of southern Sudan and on the border between Sudan and South Sudan. Lowest values were found

326 | along the northernmost parts of the Sahel on the border to the Sahara in Mauritania, in northern Mali, and in northern
327 | Niger.

328 | Modelled GPP was similar to independent GPP on average, and there was a strong linear relationship between
329 | modelled GPP and independent GPP for all sites (Fig. 6; Table 6). However, when separating the evaluation between
330 | measurement sites, it can be seen that the model reproduced some sites better than others (Fig. 7; Table 6). Wankama
331 | Millet was generally overestimated, whereas the model worked well on average for Demokeya but underestimated high
332 | values (Fig. 7; Table 6). Variability of independent GPP at the other sites was well reproduced by the model (Fig. 7;
333 | Table 6). The final parameters of the GPP model (Eq. 13) are givenshown in Table 7.

334 | <Figure 5>

335 | <Figure 6>

336 | <Figure 7>

337 | < Table 6>

338 | < Table 7>

339 |

340 | 4 Discussions

341 | Our hypothesis that vegetation indices closely related to equivalent water thickness (SIWSI) would be most strongly
342 | coupled with intra-annual dynamics in F_{opt} and α was not rejected for F_{opt} , since this was the case for all sites except for
343 | Wankama Millet (Table 2). However, our hypothesis was rejected for α , since it was more closely related to vegetation
344 | indices related to chlorophyll abundance (NDVI and EVI). In the Sahel, soil moisture conditions in the early rainy
345 | season are important for vegetation growth and during this phase vegetation is especially vulnerable to drought
346 | conditions (Rockström and de Rouw, 1997; Tagesson et al., 2016a; Mbow et al., 2013). Photosynthetic capacity (F_{opt})
347 | peaked earlier in the rainy season than α did (Fig. 3), thereby explaining the close relationship of F_{opt} to SIWSI. Leaf
348 | area index increased over the growing season and leaf area index is closely coupled with vegetation indices related to
349 | chlorophyll abundance (Tagesson et al., 2009). The increase in leaf area index increased canopy level quantum
350 | efficiency (α), which thereby explainsing the closer relationship of α to RDVI.

351 | Our hypothesis that vegetation indices closely related to chlorophyll abundance would be most strongly coupled with
352 | spatial and inter-annual dynamics in F_{opt} and α was not rejected for either F_{opt} or α ; NDVI, EVI, and RDVI all
353 | correlated all-with spatial and inter-annual dynamics in F_{opt} and α (Table 5). However, it was surprising that NDVI_{peak}
354 | had the strongest correlation with spatial and inter-annual variability for F_{opt} (Table 5). Both EVI and RDVI should be
355 | less sensitive to saturation effects than NDVI (Huete et al., 2002; Roujeau and Breon, 1995), and based on this it can be
356 | assumed that peak values of these indices should have stronger relationships to peak values of F_{opt} and α . However,
357 | vegetation indices with a high sensitivity to changes in green biomass at high biomass loads, getbecome less sensitive
358 | to green biomass changes at low biomass loads (Huete et al., 2002). The Ppeak leaf area index for ecosystems across the
359 | Sahel is generally $\sim 2 \text{ m}^2 \text{ m}^{-2}$ or less, whereas the saturation issue of NDVI generally starts at an leaf area index of about
360 | $2\text{--}5 \text{ m}^2 \text{ m}^{-2}$ (Haboudane et al., 2004).

361 | The F_{opt_peak} estimates from Agoufou, Demokeya, and the Wankama sites were similar, whereas Dahra and Kelma
362 | values were high in relation to previously reported canopy-scale F_{opt_peak} from the Sahel (~ 8 to $23 \mu\text{mol m}^{-2} \text{ sec}^{-1}$)
363 | (Hanafi et al., 1998; Merbold et al., 2009; Moncrieff et al., 1997; Boulain et al., 2009; Levy et al., 1997; Monteny et
364 | al., 1997). These previous studies reported much lower F_{opt} at canopy scale than at leaf scale (e.g. Levy et al. (1997): 10 vs.

365 | 44 $\mu\text{mol m}^{-2} \text{ sec}^{-1}$; Boulain et al. (2009): 8 vs. 50 $\mu\text{mol m}^{-2} \text{ sec}^{-1}$). The Leaf area index at Dahra and Kelma peaked at
366 | 2.1 and 2.7, respectively (Timouk et al., 2009; Tagesson et al., 2015a), and it was substantially higher than at the above-
367 | mentioned sites. A possible explanation ~~to~~for high F_{opt} estimates at Dahra and Kelma could ~~there~~foreby be the higher
368 | leaf area index. Tagesson et al. (2016b) performed a quality check of the EC data due to the high net CO_2 exchange
369 | measured at the Dahra field site and explained the high values by a combination of moderately dense herbaceous C4
370 | ground vegetation, high soil nutrient availability, and a grazing pressure resulting in compensatory growth and
371 | fertilization effects. Another possible explanation could be that the West African Monsoon brings a humid layer of
372 | surface air from the Atlantic, possibly increasing vegetation production for the most western part of ~~the~~ Sahel (Tagesson
373 | et al., 2016a).

374 | Our model substantially overestimated GPP for Wankama Millet (Fig. 7f). Being a crop field, this site differed from
375 | the other sites ~~by~~in its species composition, and ecosystem structure, as well as land and vegetation management. Crop
376 | fields in southwestern Niger are generally characterized by ~~a~~rather low production, resulting from decreased fertility
377 | and soil loss caused by intensive land use (Cappelaere et al., 2009). These specifics of the Wankama Millet site may
378 | cause the model, parameterized with observations from the other study sites without this strong anthropogenic
379 | influence, to overestimate GPP at this site. Similar results were found by Boulain et al. (2009) when applying an up-
380 | scaling model using leaf area index for Wankama Millet and Wankama Fallow. It worked well for Wankama fallow,
381 | whereas it was less conclusive for Wankama Millet. The main explanation for this difference was low leaf area index in
382 | millet fields because of a low density of millet stands due to agricultural practice. There is extensive savanna clearing
383 | for food production in ~~the~~ Sahel (Leblanc et al., 2008; Boulain et al., 2009; Cappelaere et al., 2009). To further
384 | understand impacts of this land cover change on vegetation production and land-atmosphere exchange processes, ~~it is~~
385 | ~~of~~there is an urgent need for more study sites covering cropped areas in this region.

386 | In Demokeya, GPP was slightly underestimated for the year 2008 (Fig. 7c) because modelled F_{opt} was much lower
387 | than the actual measured value in 2008 (the thick black line in Fig. 4). An improvement of the model could be to
388 | incorporate some parameters that constrain or enhance F_{opt} depending on environmental stress. Indeed, the regression
389 | tree analysis indicated that incorporating hydrometeorological variables increased the ability to predict both F_{opt} and α .
390 | On the other hand, for spatial upscaling purposes, it has been shown that including modelled hydrometeorological
391 | constraints on LUE decreases the ability to predict vegetation production due to the incorporated uncertainty in these
392 | modelled variables (Fensholt et al., 2006; Ma et al., 2014). For spatial upscaling to regional scales, it is therefore better
393 | to simply use relationships to EO data. This is particularly the case for ~~the~~ Sahel, one of the largest dryland areas in the
394 | world, ~~that~~which includes only a few sites of hydrometeorological observations.

395 | The pattern seen in the spatially explicit GPP budgets (Fig. 5c) may be influenced by a range of biophysical and
396 | anthropogenic factors. The clear North-South gradient is expected given the strong North-South rainfall gradient in ~~the~~
397 | Sahel. The West African Monsoon mentioned above could also be an explanation of high GPP values in the western
398 | part of ~~the~~ Sahel, where values were relatively high in relation to GPP at similar latitudes in the central and eastern
399 | Sahel (Fig. 5c). The areas with highest GPP are sparsely populated woodlands or shrubby savanna with a relatively
400 | dense tree cover (Brandt et al., 2016). However, the maps produced should be used with caution as they are based on
401 | up-scaling of the only six available EC sites that exist in the region; especially given the issues related to the cropped
402 | fields discussed earlier. Still, the average GPP budget for the entire Sahel 2001-2014 was close to an average annual
403 | GPP budget as estimated for these six sites ($692 \pm 89 \text{ g C m}^{-2} \text{ y}^{-1}$) (Tagesson et al., 2016a). The range of GPP budgets in

404 Fig. 5c is also similar to previous annual GPP budgets reported from other savanna areas across the world (Veenendaal
405 et al., 2004; Chen et al., 2003; Kanniah et al., 2010; Chen et al., 2016).

406 Although MOD17A2 GPP has previously been shown to capture GPP in several ecosystems types well (Turner et al.,
407 2006; Turner et al., 2005; Heinsch et al., 2006; Sims et al., 2006; Kanniah et al., 2009), it has been shown to
408 underestimate it for others (Coops et al., 2007; Gebremichael and Barros, 2006; Sjöström et al., 2013). GPP of Sahelian
409 drylands have not been captured ~~well~~ by MOD17A2-~~well~~ (Sjöström et al., 2013; Fensholt et al., 2006), and as we have
410 shown, this underestimation persists in the latest MOD17A2H GPP (collection 6) product (Fig. 2). The main reason for
411 this pronounced underestimation is that maximum LUE is set to 0.84 g C MJ⁻¹ (open shrubland; Demokeya) and 0.86 g
412 C MJ⁻¹ (grassland; Agoufou, Dahra, Kelma; Wankama Millet and Wankama Fallow) in the BPLUT, i.e. much lower
413 than maximum LUE measured at the Sahelian measurement sites of this study (average: 2.47 g C MJ⁻¹; range: 1.58-3.50
414 g C MJ⁻¹) (Sjöström et al., 2013; Tagesson et al., 2015a), a global estimate of ~1.5 g C MJ⁻¹ (Garbulsky et al., 2010)~~+~~
415 and a savanna site in Australia (1.26 g C MJ⁻¹) (Kanniah et al., 2009).

416 Several dynamic global vegetation models have been used for decades to quantify GPP at different spatial and
417 temporal scales (Dickinson, 1983; Sellers et al., 1997). These models are generally based on the photosynthesis model
418 of Farquhar et al. (1980), a model particularly sensitive to uncertainty in photosynthetic capacity (Zhang et al., 2014).
419 This and several previous studies have shown that both photosynthetic capacity and efficiency (both α and LUE) can
420 vary considerably between seasons as well as spatially, and both within and between vegetation types (Eamus et al.,
421 2013; Garbulsky et al., 2010; Ma et al., 2014; Tagesson et al., 2015a). This variability is difficult to estimate using
422 broad values based on land cover classes, yet most models apply a constant value₂ which can cause substantial
423 inaccuracies in the estimates of seasonal and spatial variability in GPP. This is particularly a problem in savannas that
424 consists of several plant functional types (C3 and C4 species, and a large variability in tree/herbaceous vegetation
425 fractions) (Scholes and Archer, 1997). This study indicates the strong applicability of EO as a tool for parameterizing
426 spatially explicit estimates of plant physiological variables, which could improve our ability to simulate GPP. Spatially
427 explicit estimates of GPP at a high temporal and spatial resolution are essential for environmental change studies in ~~the~~
428 Sahel and can contribute to increased knowledge regarding changes in GPP, its relationship to climatic change and
429 anthropogenic forcing, and estimations of ecosystem processes and biochemical and hydrological cycles.

430
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438
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721 **Tables**722 **Table 1.** Description of the six measurement sites including location, soil type, ecosystem type and dominant species.

Measurement site	Coordinates	Soil type	Ecosystem	Dominant species
Agoufou ^a (ML-AgG, Mali)	15.34°N, 1.48°W	Sandy ferruginous Arenosol	Open woody savannah (4% tree cover)	Trees: <i>Acacia spp.</i> , <i>Balanites aegyptiaca</i> , <i>Combretum glutinosum</i> Herbs: <i>Zornia glochidiata</i> , <i>Cenchrus biflorus</i> , <i>Aristida mutabilis</i> , <i>Tragus berteronianus</i>
Dahra ^b (SN-Dah, Senegal)	15.40°N, 15.43°W	Sandy luvic arenosol	Grassland/shrubland Savanna (3% tree cover)	Trees: <i>Acacia spp.</i> , <i>Balanites aegyptiaca</i> Herbs: <i>Zornia latifolia</i> , <i>Aristida adscensionis</i> , <i>Cenchrus biflorus</i>
Demokeya ^c (SD-Dem, Sudan)	13.28°N, 30.48°E	Cambic Arenosol	Sparse acacia savannah (7% tree cover)	Trees: <i>Acacia spp.</i> , Herbs: <i>Aristida pallida</i> , <i>Eragrostis tremula</i> , <i>Cenchrus biflorus</i>
Kelma ^a (ML-Kem, Mali)	15.22°N, 1.57°W	Clay soil depression	Open acacia forest (90% tree cover)	Trees: <i>Acacia seyal</i> , <i>Acacia nilotica</i> , <i>Balanites aegyptiaca</i> Herbs: <i>Sporobolus hevolvus</i> , <i>Echinochloa colona</i> , <i>Aeschynomene sensitiva</i> , <i>Guiera senegalensis</i>
Wankama Fallow ^d (NE-WaF, Niger)	13.65°N, 2.63°E	Sandy ferruginous Arenosol	Fallow bush	
Wankama Millet ^e (NE-WaM, Niger)	13.64°N, 2.63°E	Sandy ferruginous Arenosol	Millet crop	<i>Pennisetum glaucum</i>

723 ^a(Timouk et al., 2009)724 ^b(Tagesson et al., 2015b)725 ^c(Sjöström et al., 2009)726 ^d(Velluet et al., 2014)727 ^e(Boulain et al., 2009)

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Table 2. Correlation between intra-annual dynamics in photosynthetic capacity (F_{opt} ; $F_{\text{opt_frac}}$ for all sites), quantum efficiency (α ; α_{frac} for all sites), and the different vegetation indices for the six measurement sites (Fig. 1). Values are averages ± 1 standard deviation from 200 bootstrapping runs. The bold values are the indices with the strongest correlation. EVI is the enhanced vegetation index, NDVI is the normalized difference vegetation index, RDVI is the renormalized difference vegetation index, SIWSI is the shortwave infrared water stress index. SIWSI₁₂ is based on the MODIS Bidirectional Reflectance Distribution Functions (NBAR) band 2 and band 5, whereas SIWSI₁₆ is based on MODIS NBAR band 2 and band 6.

Measurement site	F_{opt}					α				
	EVI	NDVI	RDVI	SIWSI ₁₂	SIWSI ₁₆	EVI	NDVI	RDVI	SIWSI ₁₂	SIWSI ₁₆
ML-AgG	0.89 \pm 0.02	0.87 \pm 0.02	0.95 \pm 0.01	-0.95\pm0.01	-0.93 \pm 0.02	0.92 \pm 0.02	0.91 \pm 0.01	0.96\pm0.01	-0.94 \pm 0.01	-0.88 \pm 0.02
SN-Dah	0.92 \pm 0.005	0.91 \pm 0.01	0.96 \pm 0.003	-0.96\pm0.004	-0.93 \pm 0.01	0.89 \pm 0.01	0.90 \pm 0.01	0.93\pm0.01	-0.92 \pm 0.01	-0.87 \pm 0.01
SD-Dem	0.81 \pm 0.01	0.78 \pm 0.01	0.91 \pm 0.01	-0.93\pm0.01	-0.90 \pm 0.01	0.76 \pm 0.02	0.73 \pm 0.02	0.86\pm0.01	-0.82 \pm 0.02	-0.79 \pm 0.02
MA-Kem	0.77 \pm 0.02	0.83 \pm 0.02	0.95 \pm 0.01	-0.95\pm0.01	-0.90 \pm 0.02	0.69 \pm 0.05	0.73 \pm 0.04	0.80\pm0.03	-0.77 \pm 0.03	-0.76 \pm 0.03
NE-WaF	0.87 \pm 0.02	0.81 \pm 0.02	0.78 \pm 0.02	-0.90\pm0.01	-0.80 \pm 0.02	0.89\pm0.01	0.84 \pm 0.01	0.85 \pm 0.01	-0.88 \pm 0.01	-0.79 \pm 0.01
NE-WaM	0.41 \pm 0.05	0.50 \pm 0.04	0.72\pm0.03	-0.55 \pm 0.04	-0.43 \pm 0.05	0.72 \pm 0.02	0.76 \pm 0.02	0.81\pm0.01	-0.75 \pm 0.01	-0.72 \pm 0.01
All sites	0.86 \pm 0.0	0.79 \pm 0.0	0.90\pm0.0	0.75 \pm 0.0	0.70 \pm 0.0	0.83 \pm 0.01	0.80 \pm 0.01	0.86\pm0.01	0.62 \pm 0.01	0.54 \pm 0.01

Table 3. Statistics for the regression tree analysis. The regression tree analysis was used ~~for~~ to study ~~ing~~ relationships between intra-annual dynamics in the ~~the~~ photosynthetic capacity (F_{opt} ; F_{opt_frac} for all sites) and quantum efficiency (α ; α_{frac} for all sites) and the explanatory variables for the six measurement sites (Fig. 1). The pruning level is the number of splits of the regression tree and an indication of complexity of the system.

Measurement site	Explanatory variables:					Pruning level	R^2
	1	2	3	4	5		
F_{opt}							
ML-AgG	SIWSI ₁₂	Tair	PAR	SWC		16	0.98
SN-Dah	SIWSI ₁₂	SWC	VPD	Tair	PAR	84	0.98
SD-Dem	SIWSI ₁₂	VPD	SWC	Tair	PAR	33	0.97
ML-Kem	SIWSI ₁₂	PAR	Tair	VPD		22	0.98
NE-WaF	SIWSI ₁₂	SWC	VPD	Tair		14	0.92
NE-WaM	RDVI	SWC	VPD	Tair		18	0.75
All sites	RDVI	SWC	Tair	VPD		16	0.87
α							
ML-AgG	RDVI					3	0.95
SN-Dah	RDVI	VPD	SWC	Tair	PAR	21	0.93
SD-Dem	RDVI	SWC	PAR	Tair		16	0.93
ML-Kem	RDVI	Tair				4	0.75
NE-WaF	EVI	SWC	VPD			10	0.90
NE-WaM	RDVI	SWC	VPD	Tair		15	0.86
All sites	RDVI	SWC	VPD	Tair		16	0.84

Table 4. Annual peak values of quantum efficiency (α_{peak} ; $\mu\text{mol CO}_2 \mu\text{mol PAR}^{-1}$) and photosynthetic capacity ($F_{\text{opt,peak}}$; $\mu\text{mol CO}_2 \text{m}^{-2} \text{s}^{-1}$) for the six measurement sites (Fig. 1). The peak values are the 2-week running mean with highest annual value.

Measurement site	Year	α_{peak}	$F_{\text{opt,peak}}$
ML-AgG	2007	0.0396	24.5
SN-Dah	2010	0.0638	50.0
	2011	0.0507	42.3
	2012	0.0480	39.2
	2013	0.0549	40.0
SD-Dem	2007	0.0257	16.5
	2008	0.0327	21.0
	2009	0.0368	16.5
ML-Kem	2007	0.0526	33.5
NE-WaF	2005	0.0273	18.2
	2006	0.0413	21.0
NE-WaM	2005	0.0252	10.6
	2006	0.0200	10.1
Average		0.0399	26.4

Table 5. Correlation matrix between annual peak values of photosynthetic capacity (F_{opt_peak}) and quantum efficiency (α_{peak}) and measured environmental variables. P is annual rainfall; T_{air} is yearly averaged air temperature at 2 m height; SWC is yearly averaged soil water content (% volumetric water content) measured at 0.1 m depth; Rh is yearly averaged relative humidity; VPD is yearly averaged vapour pressure deficit; R_g is yearly averaged incoming global radiation; N and C cont. are soil nitrogen and carbon contents; NDVI_{peak} is annual peak normalized difference vegetation index (NDVI); EVI_{peak} is annual peak enhanced vegetation index (EVI); RDVI_{peak} is annual peak renormalized difference vegetation index (RDVI); SIWSI_{12peak} is annual peak short-wave infrared water stress index based on MODIS NBAR band 2 and band 5; and SIWSI_{16peak} is annual peak short-wave infrared water stress index based on MODIS NBAR band 2 and band 6. Sample size was 13 for all except the marked explanatory variables.

Explanatory variable	F_{opt_peak}	α_{peak}
Meteorological data		
P (mm)	0.24±0.26	0.13±0.27
T_{air} (°C)	-0.07±0.25	-0.01±0.25
SWC (%) ^a	0.33±0.25	0.16±0.27
Rh (%)	0.73±0.16*	0.60±0.19
VPD (hPa)	0.20±0.26	0.15±0.30
R_g (W m ⁻²)	-0.48±0.21	-0.41±0.24
Biomass and edaphic data		
Biomass (g DW m ⁻²) ^a	0.77±0.15*	0.74±0.14*
C3/C4 ratio	-0.05±0.26	0.06±0.30
N cont. (%) ^b	0.22±0.11	0.35±0.14
C cont. (%) ^b	0.89±0.06**	0.87±0.07**
Earth observation data		
NDVI _{peak}	0.94±0.05**	0.87±0.07*
EVI _{peak}	0.93±0.04**	0.87±0.07**
RDVI _{peak}	0.93±0.04**	0.89±0.07**
SIWSI _{12peak}	0.85±0.08**	0.84±0.08**
SIWSI _{16peak}	0.67±0.12*	0.65±0.15*
Photosynthetic variables		
F_{opt}	-	0.94±0.03**

^asample size equals 11.

^bsample size equals 9.

* significant at 0.05 level.

** significant at 0.01 level

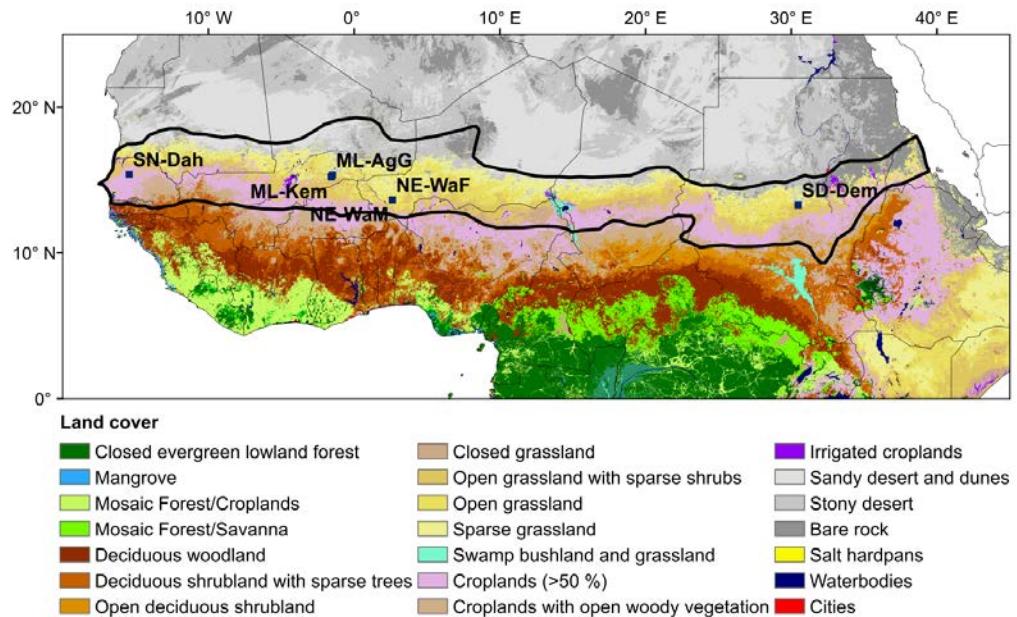
Table 6. Statistics regarding the evaluation of the gross primary production (GPP) model for the six measurement sites (Fig. 1). In situ and modelled GPP are averages ± 1 standard deviation. RMSE is the root-mean-square error, and slope, intercept and R^2 are from the fitted ordinary least squares linear regression.

Measurement site	In situ GPP ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$)	Modelled GPP ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$)	RMSE ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$)	slope	Intercept ($\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$)	R^2
ML-AgG	5.35 \pm 6.38	5.97 \pm 5.80	2.48 \pm 0.10	0.84 \pm 0.003	1.46 \pm 0.01	0.86 \pm 0.002
SN-Dah	9.39 \pm 10.17	8.87 \pm 9.67	3.99 \pm 1.34	0.88 \pm 0.002	0.62 \pm 0.01	0.85 \pm 0.001
SD-Dem	4.26 \pm 4.55	3.98 \pm 3.90	3.15 \pm 1.06	0.63 \pm 0.003	1.31 \pm 0.007	0.54 \pm 0.02
ML-Kem	11.16 \pm 8.02	10.52 \pm 9.22	4.35 \pm 1.23	1.02 \pm 0.003	-0.82 \pm 0.03	0.78 \pm 0.002
NE-WaF	5.77 \pm 4.17	6.63 \pm 3.53	2.47 \pm 1.05	0.70 \pm 0.005	2.58 \pm 0.02	0.69 \pm 0.003
NE-WaM	3.04 \pm 1.93	6.35 \pm 3.47	4.12 \pm 0.99	1.31 \pm 0.004	2.37 \pm 0.02	0.53 \pm 0.003
Average	6.73 \pm 7.72	7.02 \pm 7.39	3.68 \pm 0.55	0.83 \pm 0.07	1.34 \pm 0.82	0.84 \pm 0.07

Table 7. The parameters for Eq. 13 that ~~was~~^{were} used in the final gross primary production (GPP) model. RMSE is the root mean square error, and R^2 is the coefficient of determination for the regression models predicting the different variables.

Parameter	Value	RMSE	R^2
k_{Fopt}	79.6 ± 6.3		
m_{Fopt}	-7.3 ± 3.2	5.1 ± 1.3	0.89 ± 0.05
l_{Fopt}	3.51 ± 0.19		
n_{Fopt}	0.03 ± 0.006	0.15 ± 0.02	0.88 ± 0.06
a	0.16 ± 0.02		
m_a	-0.014 ± 0.007	0.0069 ± 0.0021	0.81 ± 0.10
l_a	3.75 ± 0.27		
n_a	0.02 ± 0.007	0.20 ± 0.02	0.80 ± 0.10

Figures



5 **Figure 1.** Land cover classes for the Sahel and the location of the six measurement sites included in the study. The land cover classes are based on multi-sensor satellite observations (Mayaux et al., 2003). The sites are Agoufou (ML-AgG), Dahra (SN-Dah), Demokeya (SD-Dem), Kelma (ML-Kem), Wankama Fallow (NE-WaF), and Wankama Millet (NE-WaM). The thick black line ~~is~~are the borders of the Sahel based on the isohytes 150 and 700 mm of annual precipitation (Prince et al., 1995).

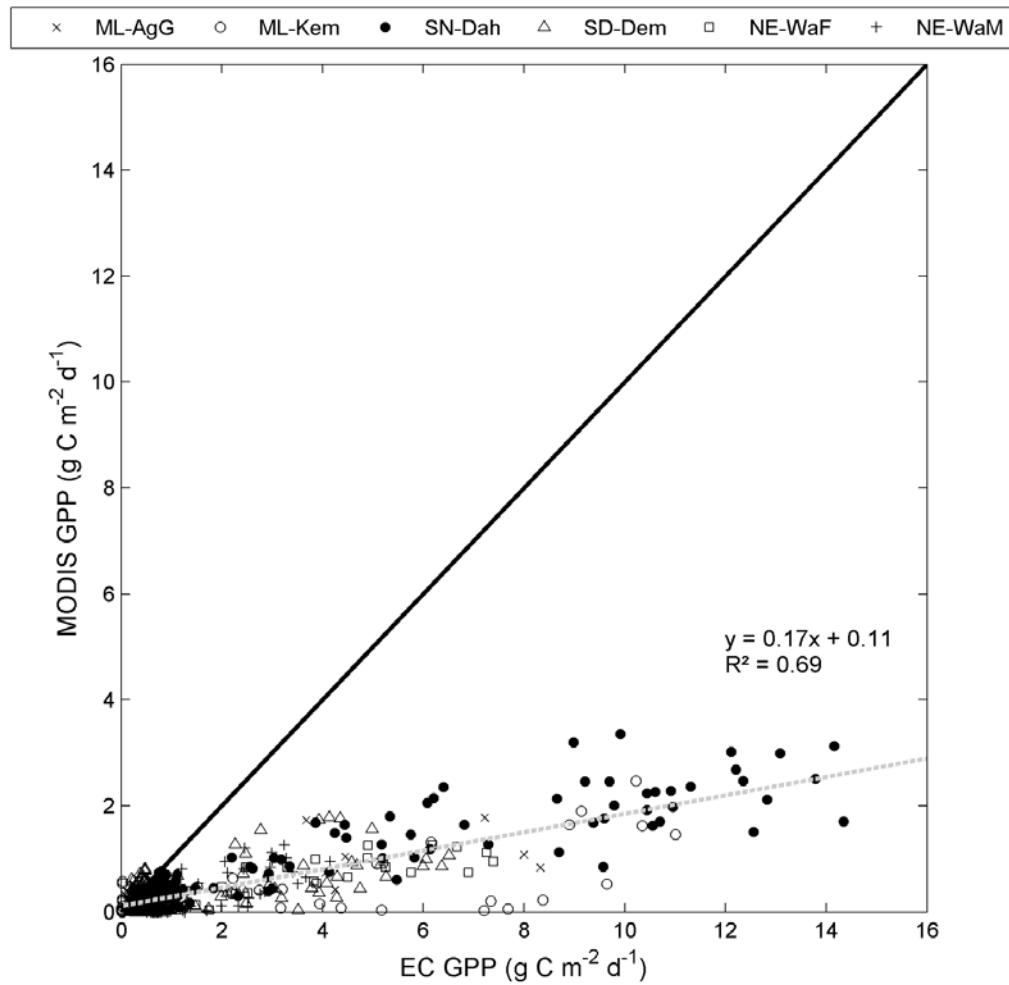


Figure 2. Evaluation of the MODIS based GPP product MOD17A2H (collection 6) against eddy covariance based GPP from the six measurement sites (Fig. 1) across the Sahel. The thick black line shows the one-to-one ratio, and the grey dotted line is the fitted ordinary least squares linear regression.

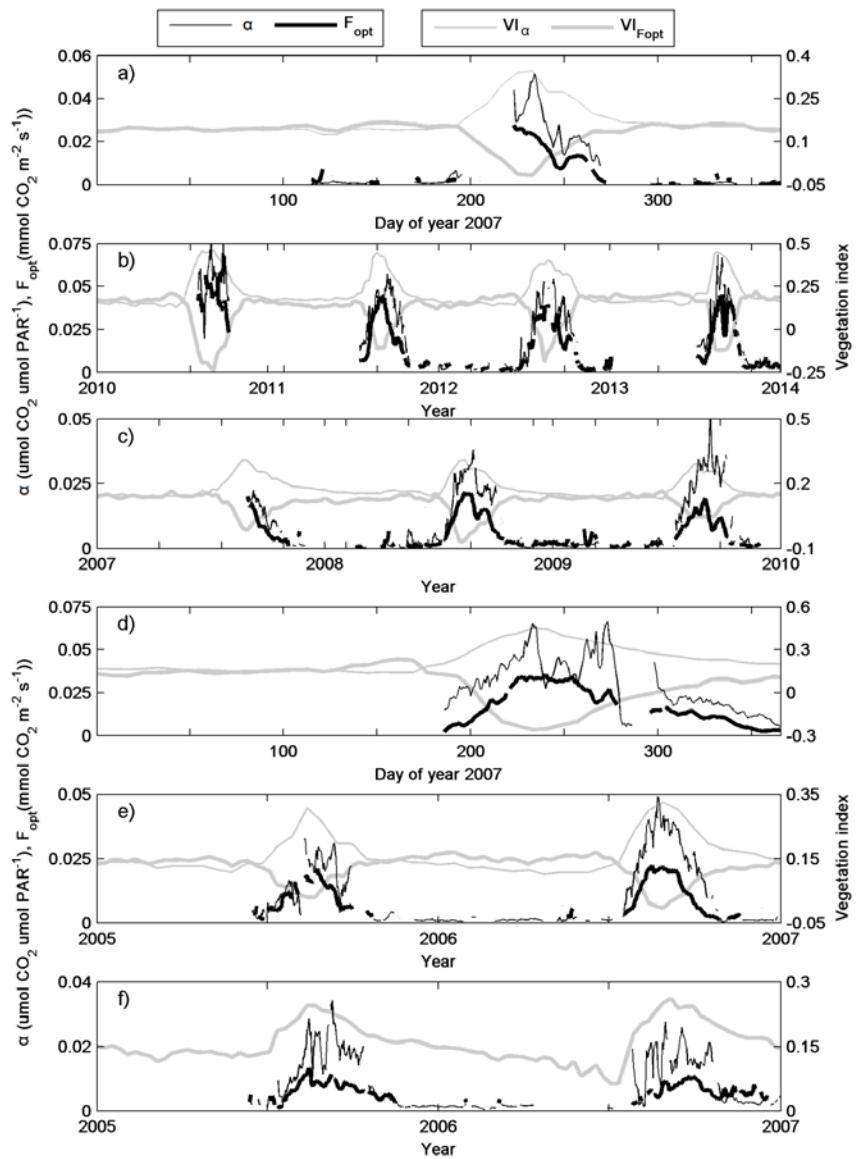


Figure 3. Dynamics in photosynthetic capacity (F_{opt}) and quantum efficiency (α) for the six measurement sites. Also included ~~is also are~~ are dynamics in the vegetation indices with highest correlation to the intra-annual dynamics in F_{opt} ($VI_{F_{\text{opt}}}$) and to quantum efficiency (VI_{α}) (Table 2). The sites are a) Agoufou (ML-AgG), b) Daha (SN-Dah), c) Demokeya (SD-Dem), d) Kelma (ML-Kem), e) Wankama Fallow (NE-WaF), and f) Wankama Millet (NE-WaM).

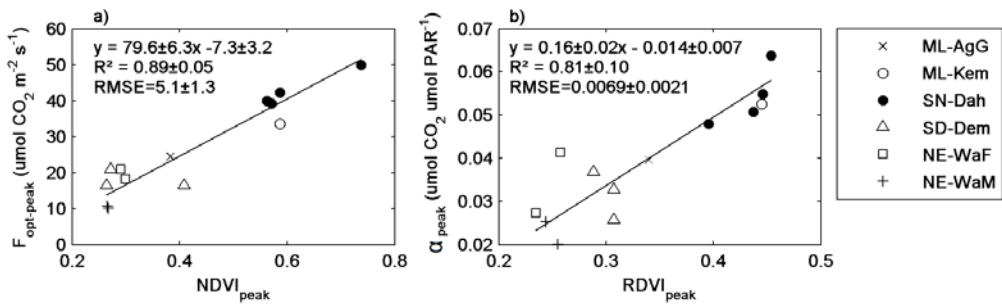


Figure 4. Scatter plots of annual peak values for the six measurement sites (Fig. 1) of a) photosynthetic capacity ($F_{\text{opt_peak}}$) and b) quantum efficiency (α_{peak}) against peak values of normalized difference vegetation index ($\text{NDVI}_{\text{peak}}$) and renormalized difference vegetation index ($\text{RDVI}_{\text{peak}}$), respectively. The annual peak values were estimated by taking the annual maximum of a ~~two~~²-week running mean.

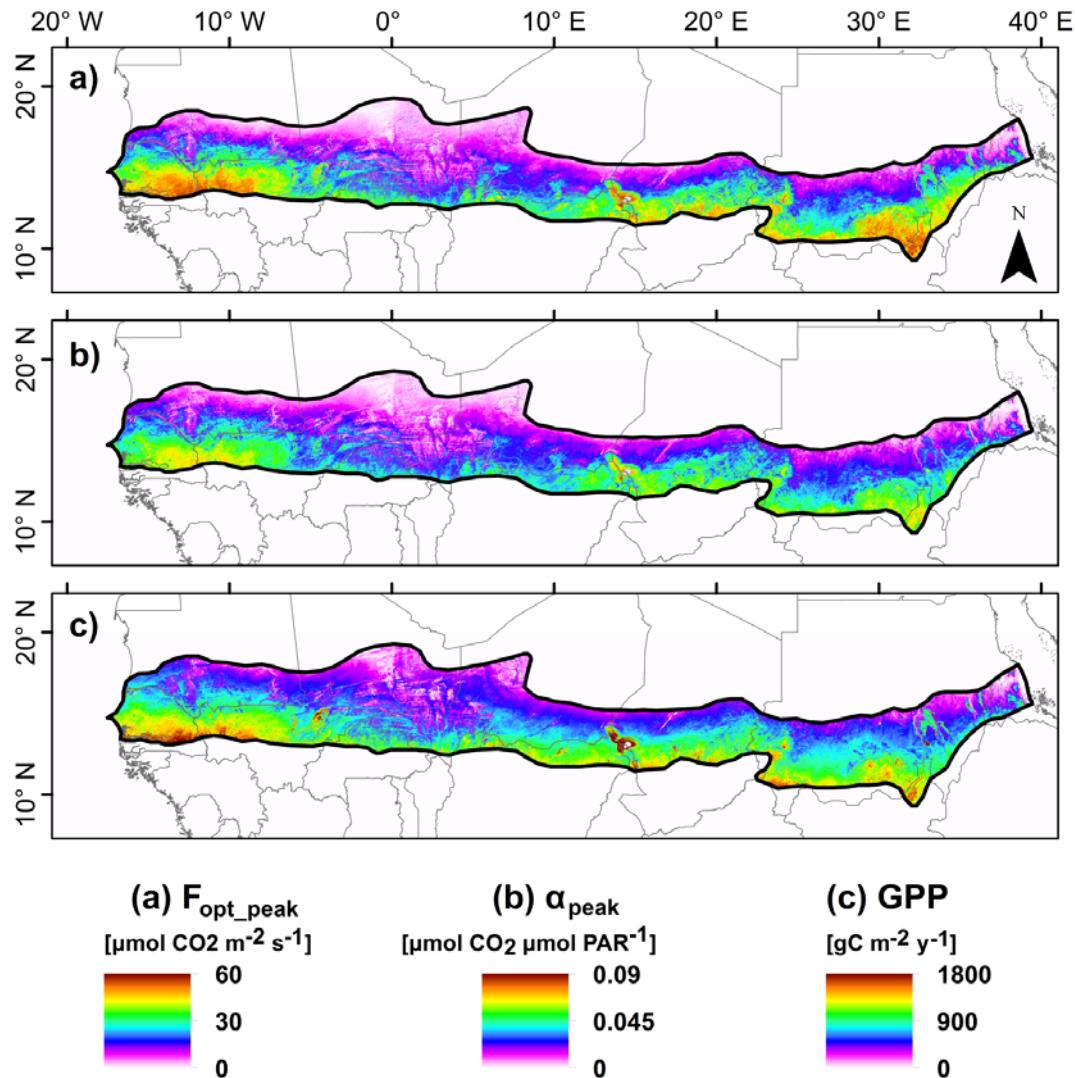


Figure 5. Maps of a) peak values of photosynthetic capacity ($F_{\text{opt_peak}}$) averaged for 2001-2014, b) peak values of quantum efficiency (α_{peak}) averaged for 2001-2014, and c) annual budgets of GPP averaged for 2001-2014.

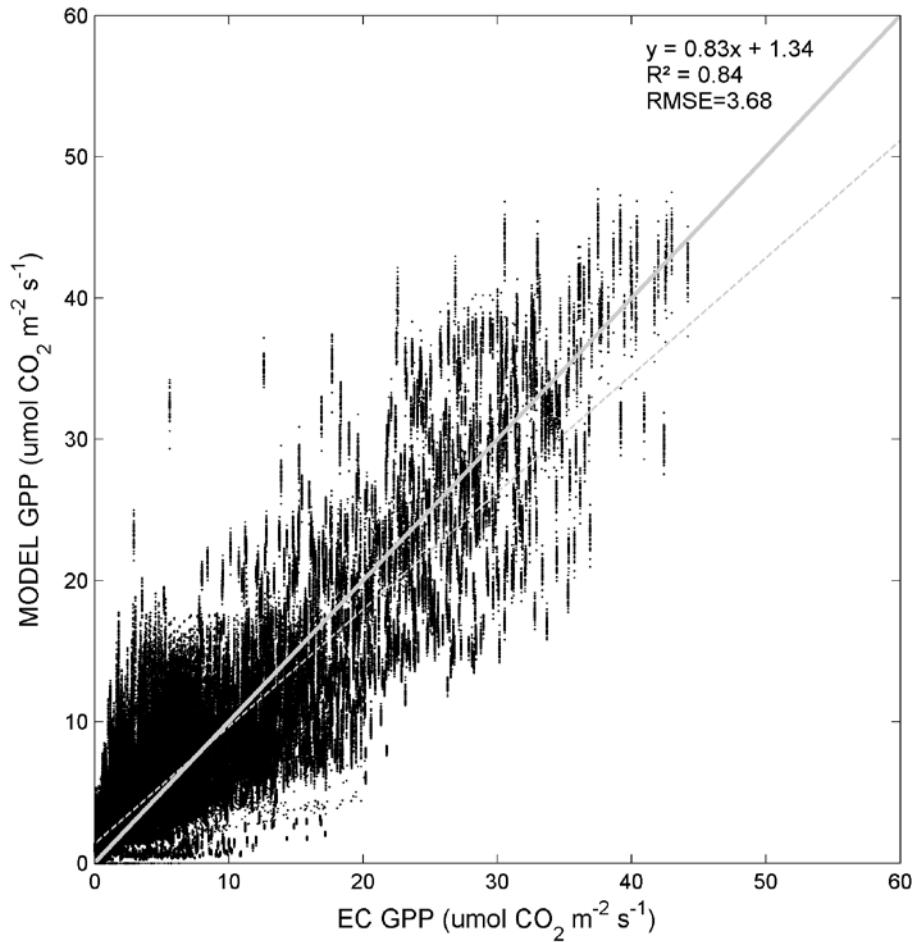


Figure 6. Evaluation of the modelled gross primary production (GPP) (Eq. 13) against in situ GPP from all six measurement sites across the Sahel. The thick grey line shows the one-to-one ratio, whereas the dotted-thin dotted grey line is the fitted ordinary least squares_s linear regression.

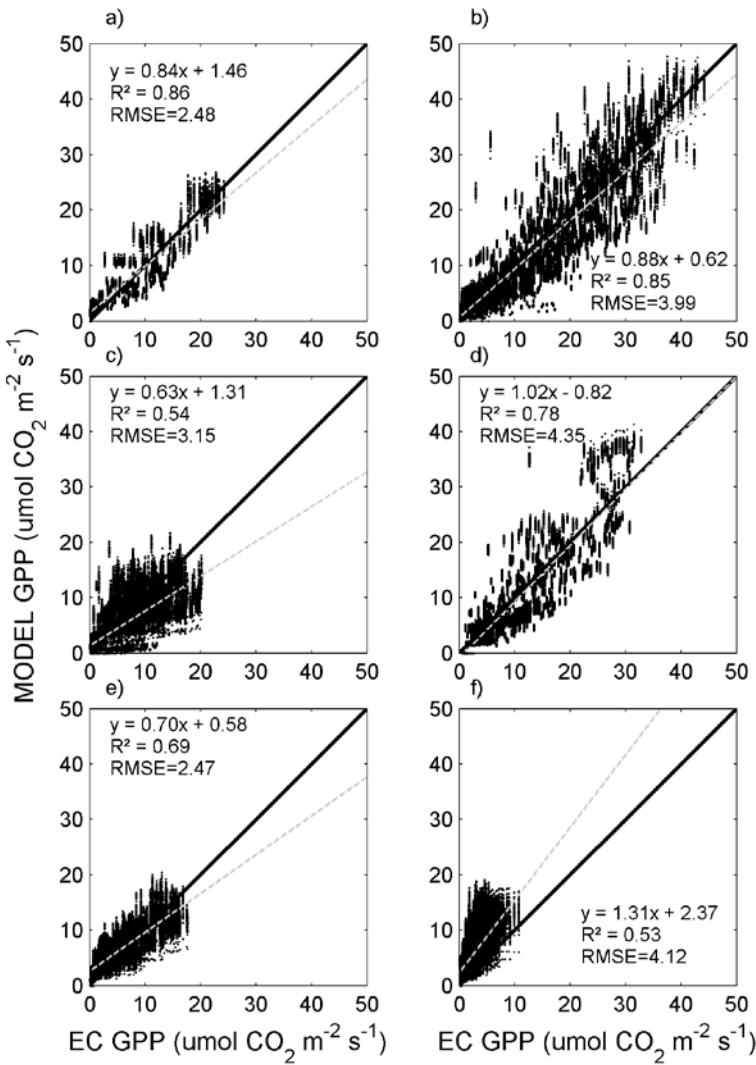


Figure 7. Evaluation of the modelled gross primary production (GPP) (Eq. 13) against in situ GPP for the six sites across Sahel (Fig. 1). The thick black line shows the one-to-one ratio, whereas the dotted thin grey line is the fitted ordinary least

squares linear regression. The sites are a) Agoufou (ML-AgG), b) Dahra (SN-Dah), c) Demokeya (SD-Dem), d) Kelma (ML-Kem), e) Wankama Fallow (NE-WaF), and f) Wankama Millet (NE-WaM).