

# 1 **Measuring ecosystem nitrogen status: a comparison of proxies**

2 Maya Almaraz<sup>1</sup>, Stephen Porder<sup>1</sup>

3 <sup>1</sup>Department of Ecology and Evolutionary Biology, Brown University, Providence, 02912, USA

4 *Correspondence to:* Maya Almaraz (maya\_almaraz@brown.edu)

5 **Keywords:** nitrogen availability, nutrient limitation,  $\delta^{15}\text{N}$ , nitrogen mineralization, dissolved  
6 organic nitrogen

7  
8 **Abstract.** There are many proxies used to measure nitrogen (N) availability in watersheds, but  
9 the degree to which they do (or do not) correlate within a watershed has not been systematically  
10 addressed. We surveyed the literature for intact forest or grassland watersheds globally, in which  
11 several metrics of nitrogen availability have been measured. Our metrics included: foliar  $\delta^{15}\text{N}$ ,  
12 soil  $\delta^{15}\text{N}$ , net nitrification, net N mineralization, and the ratio of dissolved inorganic to organic  
13 nitrogen (DIN:DON) in soil solution and streams. Not surprisingly, the strongest correlation  
14 (Kendall's tau) was between net nitrification and N mineralization ( $\tau=0.71$ ,  $p<0.0001$ ). Net  
15 nitrification and N mineralization were each correlated with foliar and soil  $\delta^{15}\text{N}$  ( $p<0.05$ ). Foliar  
16 and soil  $\delta^{15}\text{N}$  were more tightly correlated across tropical sites ( $\tau=0.68$ ,  $p<0.0001$ ), than in  
17 temperate sites ( $\tau=0.23$ ,  $p=0.02$ ). To our surprise, the only significant correlations we found  
18 between terrestrial- and water-based metrics were that of net nitrification ( $\tau=0.48$ ,  $p=0.01$ ) and N  
19 mineralization ( $\tau=0.69$ ,  $p=0.0001$ ) with stream DIN:DON. The relationship between stream  
20 DIN:DON with both net nitrification and N mineralization was significant only in temperate, but  
21 not tropical regions. Given that both soil  $\delta^{15}\text{N}$  and stream DIN:DON are used to infer long-term  
22 N status, their lack of correlation in watersheds merits further investigation.

## 23 **1.0 Introduction**

24 Nitrogen (N) limitation to primary production is widespread in both terrestrial and  
25 aquatic ecosystems, and variation in N availability drives differences in ecosystem properties  
26 across space and time (Vitousek and Howarth, 1991; Elser et al., 2007; LeBauer and Treseder,  
27 2008). Yet quantifying N availability over timescales that are relevant in ecosystems is non-  
28 trivial. Short timescale measurements of N availability in soil are common (e.g. inorganic N  
29 pools, N mineralization and nitrification rates; Binkley and Hart, 1989; Sparks et al., 1996), but  
30 such short-term proxies are influenced by both short and long-term drivers, and thus it is difficult  
31 to know whether short-term proxies can be used to infer N status (i.e. the relative abundance of  
32 plant available N) over long timescales in relatively undisturbed ecosystems. For example,  
33 measured net mineralization and nitrification rates in arctic tundra are commonly less than  
34 annual plant uptake (Schimel et al., 1996; Schmidt et al., 1999), and annual N budgets based on  
35 short-term measurements are difficult to balance (e.g. Magill et al., 1997). While N status  
36 measured over longer temporal and larger spatial scales is relevant to many ecosystem properties  
37 and their response to global change, it is more difficult to measure.

38 Land-based investigations of N cycling commonly measure extractable N, N  
39 mineralization, and nitrification, which give a snapshot of N status over minutes to days (Binkley  
40 and Hart, 1989; Robertson et al., 1999). Some researchers also use lysimeters to quantify  
41 dissolved N losses from below the rooting zone (Hedin et al., 2003; McDowell et al., 2004;  
42 Lohse and Matson, 2005) on a similar timescale. Repeated measurements give longer timescale  
43 information, but even the longest studies are short relative to ecosystem development.

44 In addition to these short-term proxies, there are two relatively common measurements  
45 that are thought to average over space and/or time. The first is the ratio of dissolved inorganic

46 (DIN) to organic (DON) N concentration lost from ecosystems. Losses of DIN are considered  
47 controllable by biota, and thus should be low if soil N is in short supply. In contrast, most DON  
48 is not accessible to plants, and thus represents a loss beyond biotic control (Hedin et al., 1995;  
49 Figure 1). Thus low DIN:DON in streams has been used to infer relative N-poverty in  
50 watersheds (e.g. McDowell and Asbury, 1994; Perakis and Hedin, 2002; Brookshire et al., 2012).  
51 The few sites where such measurements have been made over decades (e.g. the Luquillo  
52 Mountains of Puerto Rico, Harvard Forest in Massachusetts, Hubbard Brook LTER in New  
53 Hampshire; McDowell et al., 1992, McDowell et al., 2004, Bormann and Likens 2012) suggest  
54 stream DIN:DON is not particularly variable over this timescale, and thus this metric may  
55 integrate over time as well as space (W.C. McDowell, pers. comm.). It is common that  
56 researchers using DIN:DON to infer ecosystem N status implicitly assume that a few  
57 measurements are indicative of longer-term patterns (e.g. Perakis and Hedin, 2002; Brookshire et  
58 al., 2012).

59 In contrast to stream DIN:DON, soil  $\delta^{15}\text{N}$  integrates solely over time, and at steady state  
60 reflects the isotopic signature associated inputs (N fixation and/or deposition) and fractionation  
61 associated with outputs (Handley and Raven, 1992). The major N loss pathways (primarily  
62 denitrification, and to a lesser extent nitrate leaching) discriminate against  $^{15}\text{N}$ , which thus  
63 remains in relative abundance in N-rich soils (Hogburg 1997; Martinelli et al., 1999; Craine et  
64 al., 2009; Houlton and Bai, 2009, Craine et al., 2015; Figure 1). To some degree foliar  $\delta^{15}\text{N}$   
65 reflects soil  $\delta^{15}\text{N}$  (Amundson et al., 2003), but there can be fractionation during nitrification,  
66 between bulk and soil solution N pools (Hogburg, 1997), as well as during N uptake by roots and  
67 mycorrhizae (Hobbie et al., 2009). For this reason, foliar  $\delta^{15}\text{N}$  may display greater variability

68 between species in a single site than the bulk soil  $\delta^{15}\text{N}$  (Vitousek et al., 1989; Nadlehoffer et al.,  
69 1996).

70         Given that these proxies for N availability function over different spatial and temporal  
71 scales, we asked which proxies correlated in watersheds where several measurements have been  
72 made in the same place and at roughly the same time. We were particularly interested in whether  
73 short-timescale measurements (nitrification, mineralization) correlated with the more temporally  
74 (foliar and soil  $\delta^{15}\text{N}$ ) and spatially (stream DIN:DON) integrated proxies. Unlike previous  
75 reviews (Sudduth et al., 2013) we focus solely on unmanaged systems where we were able to  
76 compare plant, soil, soil solution and stream proxies. This review assesses the correlation  
77 between common foliar, surface soil (i.e.  $\delta^{15}\text{N}$ , nitrification and mineralization), and nutrient loss  
78 (i.e. soil solution and stream N concentrations) metrics of N availability from unmanaged  
79 ecosystems globally (Figure 2).

80

## 81 **2.0 Methods**

82

### 83 **2.1 Literature Review**

84         We surveyed the literature (prior to 2013) and contacted individual investigators to gather  
85 data from forested and grassland watersheds where more than one proxy of long-term N  
86 availability had been measured. We focused on the most commonly-used proxies for N status:  
87 foliar (n=78) and surface soil  $\delta^{15}\text{N}$  (n=104; <20 cm depth), net nitrification (n=86; <20 cm  
88 depth), net N mineralization (n=88; <20 cm depth), the ratio of dissolved inorganic to organic N  
89 forms (DIN:DON) in soil solution below the rooting zone (n=43; >20 cm depth), and stream  
90 DIN:DON (n=32). We chose these metrics because 1) other authors have suggested that they are

91 indicative of soil nutrient status (Martinelli et al., 1999, Amundson et al., 2001, Brookshire et al.,  
92 2012; Figure 1), and 2) they are thought to integrate N fluxes on different timescales (e.g. soil  
93  $\delta^{15}\text{N}$  integrates N losses over decades while net N mineralization rates integrate inorganic N  
94 production over days; Binkley and Hart, 1989, Hogburg 1997).

95 We used the search engines Web of Science and Google Scholar and searched key words:  
96 “nitrogen”, “ $^{15}\text{N}$ ”, “natural abundance”, “mineralization”, and “dissolved organic nitrogen”,  
97 “*watershed name*”. References in papers that resulted from the keyword search were then used to  
98 gather additional data. We limited our search criteria to studies that took place in forest or  
99 grassland ecosystems that had not incurred any large disturbances that might impair their  
100 function.

101 We collected data from 154 watersheds across a broad climatic range (Figure 2), in which  
102 at least two of the six N proxies of interest had been measured (see *Supplemental Data*). We used  
103 DataThief II software (version 1.2.1) to extract data from figures when data were not available in  
104 text or tables. When necessary, data were converted to standardize units.

105 From each paper we collected the following site description data: country, site,  
106 watershed, biome, ecosystem type, latitude, longitude, elevation (m), mean annual temperature  
107 (MAT; °C), mean annual precipitation (MAP; mm yr<sup>-1</sup>), N deposition rate (kg N ha<sup>-1</sup> yr<sup>-1</sup>), soil  
108 depth (cm), soil solution (lysimeter) depth (cm), and N mineralization method. Site description  
109 data were gathered from other sources when they were not in the original publication.

110 In order to control for methodological differences, we limited our net nitrification and N  
111 mineralization methods to those which used intact soil core, buried bag, and laboratory  
112 incubations of unamended soils (Boone, 1992; Piccolo et al., 1994), and eliminated methods  
113 such as ion resin exchange beads or  $^{15}\text{N}$  tracer techniques (Binkley et al., 1986; Hart and

114 Firestone, 1989; Davidson et al., 1991; Templer et al., 2008). We did not limit net nitrification  
115 and N mineralization data based on the length of the incubation, as we see little change in rates  
116 between 1-7 days (Tietema et al., 1998), however we recognize that longer incubations may  
117 result in lower net rates. Soil values were from the mineral soil only, and were preferentially  
118 collected in the 0-10 cm range, however if soil samples were not in 10 cm increments, we  
119 selected the increment that was most similar (e.g. A horizon, 0-5 cm, 0-15 cm), and no deeper  
120 than 20 cm.

121         When data were missing, or we were uncertain about location or collection method, we  
122 contacted the authors to request unpublished data, elucidation of data collection, data reduction,  
123 or soil samples. Terrestrial metrics were typically gathered from different papers than that of  
124 water-based metrics, requiring validation of congruent watershed location. For five watersheds.  
125 (Puerto Rico's Pared, Sonadora, Bisley, Tronoja watersheds and Hubbard Brook's watershed 6)  
126 we collected soil that we analyzed for  $\delta^{15}\text{N}$ . In Puerto Rico, we collected mineral soil samples (0-  
127 10 cm) in replicates of five using an open side soil sampler from locations that were >3 m away  
128 from the stream. Replicate samples were combined in a Ziploc bag, air-dried and shipped to the  
129 Marine Biological Laboratory for analysis. Colleagues at Hubbard Brook collected three  
130 replicate horizon B samples for us from several soil pits dug across an elevation gradient in  
131 watershed 6 (Christopher Neill, pers. comm.), which were air-dried at the Marine Biological  
132 Laboratory prior to analysis.

133

## 134 **2.2 Soil Sample Analysis**

135         The few soils we analyzed in house for  $\delta^{15}\text{N}$  were homogenized, sieved (2 mm) and  
136 ground using a mortar and pestle. We analyzed samples at the Marine Biological Laboratory

137 Ecosystem Center Stable Isotope Laboratory for  $\delta^{15}\text{N}$  using a Europa 20-20 continuous-flow  
138 isotope ratio mass spectrometer interfaced with a Europa ANCA-SL elemental analyzer. The  
139 analytical precision based on replicate analyses of  $\delta^{15}\text{N}$  of isotopically homogeneous  
140 international standards was  $\pm 0.1 \text{ ‰}$ .

141

## 142 **2.3 Statistics**

143 Five of our six variables were not normally distributed, so we used a non-parametric  
144 Kendall tau rank test in R (version 2.11.1), to determine the significance of correlations.  
145 Kendall's tau evaluates the degree of similarity between two sets of ranked data and generates a  
146 smaller co-efficient as the number of discordant pairs between two ranking lists becomes greater  
147 (Abdi 2007). The Kendall tau rank test is well suited for these comparisons as it is not sensitive  
148 to missing data and outliers, it measures both linear and non-linear correlations, and generates a  
149 more accurate p-value with small sample sizes (Helsel and Hirsch, 1992; Raike et al., 2003). We  
150 corrected for multiple comparisons by reporting Bonferroni adjusted p-values for each of our 15  
151 comparisons (Bland and Altman, 1995). We removed a single stream DIN:DON value from  
152 Cascade Head, Oregon, as it was ~20 times higher than the mean of all other stream values  
153 (Compton et al., 2003); however removing this outlier had little effect on the correlations.

154

## 155 **3.0 Results**

156 All terrestrial-based proxies that integrate across long and short timescales were  
157 significantly correlated. Soil  $\delta^{15}\text{N}$  was positively correlated with both net nitrification ( $n=60$ ,  
158  $\tau=0.37$ ,  $p<0.0001$ ) and N mineralization ( $n=64$ ,  $\tau=0.41$ ,  $p<0.0001$ ). Foliar  $\delta^{15}\text{N}$  was also

159 positively correlated with net nitrification ( $n=43$ ,  $\tau=0.49$ ,  $p<0.0001$ ), and N mineralization  
160 ( $n=46$ ,  $\tau=0.34$ ,  $p=0.001$ ; Figure 2).

161 Not surprisingly, we found significant correlations between terrestrial-based proxies that  
162 measure nutrient availability on similar timescales. Foliar  $\delta^{15}\text{N}$  was positively correlated with  
163 soil  $\delta^{15}\text{N}$  ( $n=78$ ,  $\tau=0.40$ ,  $p<0.0001$ ). There was also a positive correlation between net  
164 nitrification and N mineralization ( $n=88$ ,  $\tau=0.71$ ,  $p<0.0001$ ; Figure 3).

165 Despite the correlation between all terrestrial-based measurements of N availability,  
166 terrestrial metrics did not exhibit similarly robust relationships with water-based proxies. No  
167 metric was significantly correlated with soil solution DIN:DON ( $n=53$ ,  $p>0.05$ ). Net nitrification  
168 ( $n=15$ ,  $\tau=0.48$ ,  $p=0.01$ ) and N mineralization ( $n=17$ ,  $\tau=0.69$ ,  $p=0.0001$ ) were the only metrics to  
169 correlate with stream DIN:DON. Soil solution and stream DIN:DON data were not correlated  
170 (Figure 3). All of the data in Figure 3, and their original sources, are available in Supplemental  
171 Table 1.

172 The lack of relationship between water-based and terrestrial-based metrics lead us to ask  
173 questions about variability of soil solution and stream DIN:DON across environmental gradients.  
174 We found that neither soil solution or stream DIN:DON were correlated with temperature,  
175 precipitation, elevation or N deposition ( $p>0.05$ ). To our surprise, solution DIN:DON was not  
176 correlated with lysimeter depth ( $p>0.05$ ).

177 Some relationships between proxies differed with latitude. Soil and foliar  $\delta^{15}\text{N}$  were more  
178 tightly correlated in the tropics ( $n=24$ ,  $\tau=0.68$ ,  $p<0.0001$ ) than in the temperate zone ( $n=49$ ,  
179  $\tau=0.23$ ,  $p=0.02$ ). Soil  $\delta^{15}\text{N}$  was correlated with net nitrification in tropical ( $n=17$ ,  $\tau=0.39$ ,  
180  $p=0.03$ ), but not temperate regions. Conversely, soil  $\delta^{15}\text{N}$  was correlated with net N  
181 mineralization ( $n=44$ ,  $\tau=0.34$ ,  $p=0.001$ ) in temperate but not tropical areas. Stream DIN:DON



182 was correlated with net nitrification ( $n=10$ ,  $\tau=0.63$ ,  $p=0.01$ ) and N mineralization ( $n=10$ ,  $\tau=0.78$ ,  
183  $p=0.002$ ) in the temperate zone, and not in the tropics ( $n=4$ ,  $p>0.05$ ).

184

#### 185 **4.0 Discussion**

186 The metrics presented here are typically interpreted to fall into one of three categories: 1)  
187 long-timescale (decades to centuries) integrators of soil N losses (foliar and soil  $\delta^{15}\text{N}$ ; Martinelli  
188 et al., 1999, Craine et al., 2015), 2) short-timescale direct measures of N transformations  
189 (mineralization, nitrification; Vitousek et al., 1982), and 3) short-medium timescale (weeks to  
190 years) measures of hydrologic N losses that are influenced by N availability in a catchment (soil  
191 solution and stream DIN:DON; Hedin et al., 1995; Perakis and Hedin, 2001). Our data suggest  
192 that category 1 and 2 metrics are correlated, and that short-term soil assays may capture similar  
193 patterns as inferred by long-term plant and soil-based proxies. However, the lack of correlation  
194 between long-term terrestrial proxies (plant and soil  $\delta^{15}\text{N}$ ) and both soil solution and stream  
195 DIN:DON is interesting, as several authors have suggested that both types of proxies give insight  
196 into ecosystem N status (Vitousek et al., 1982; Hedin et al., 1995; Martinelli et al., 1999; Perakis  
197 and Hedin, 2001; Amundson et al., 2003; Brookshire et al., 2012).

198 It is particularly interesting that stream DIN:DON was not correlated with soil  $\delta^{15}\text{N}$  as both  
199 are proxies used to infer long-term N status. There is a wealth of literature that uses stream  
200 DIN:DON to infer large spatial and temporal scale patterns in N availability (Hedin et al., 1995;  
201 Perakis and Hedin, 2002; McDowell et al., 2004; Fang et al., 2008). Similarly, many studies  
202 interpret soil  $\delta^{15}\text{N}$  as an integrator of N losses over time (Martinelli et al., 1999; Houlton et al.,  
203 2006; Houlton and Bai, 2009, Craine et al., 2015). These are the only two proxies for N status  
204 that integrate over relatively long timescales, and their lack of correlation warrants more careful

205 consideration. We note that stream DIN:DON is sensitive to N deposition, and that relatively  
206 pristine settings have a lower DIN:DON than polluted ones (Perakis and Hedin, 2001). In our  
207 dataset, N deposition was not correlated with stream DIN:DON ( $\tau=0.03$ ,  $p>0.05$ ), or any other  
208 metric. Thus our data do not support the idea that N deposition is responsible for the lack of  
209 correlation between these two long-term proxies.

210 Another surprise from our dataset is that soil solution DIN:DON was not significantly  
211 correlated with any other metric, not even with stream DIN:DON, despite ~40% of papers in our  
212 dataset reporting both soil solution and stream DIN:DON in the same watershed. While the  
213 correlation between soil solution DIN:DON below the rooting zone and N availability has been  
214 documented across gradients in soil age and fertility (Hedin et al., 1995), this correlation was not  
215 found across the range of sites examined here. We found no relationship between soil solution  
216 DIN:DON and lysimeter depth, suggesting that the majority of N transformations responsible for  
217 the discontinuity between soil solution DIN:DON and that of terrestrial metrics are likely  
218 occurring either within the rooting zone or in riparian zones. Neither soil solution or stream  
219 DIN:DON was sensitive to environmental variability (i.e. elevation, temperature, precipitation, N  
220 deposition), suggesting that processing along flow paths may be responsible for the disconnect  
221 between soil solution and stream N concentrations. From these data, at least, it does not seem  
222 that soil solution DIN:DON can be used to infer terrestrial N status across this suite of  
223 unmanaged sites. These data also do not support the idea that soil solution DIN:DON is  
224 representative of N forms that leach into streams (Binkley et al., 1992; Pregitzer et al., 2004;  
225 Fang et al., 2008).

226 While nitrate ( $\text{NO}_3^-$ ) removal along flow paths can reduce stream  $\text{NO}_3^-$  (Vidon et al., 2010),  
227 with higher percent removal in forested watersheds (Sudduth et al., 2013), DON has been shown

228 to be relatively resistant to removal by decomposition and biologic uptake along subsurface flow  
229 paths (Carreiro et al., 2000, Neff et al. 2003). We found no correlation between stream and soil  
230 solution DIN:DON, and suggest that variation in  $\text{NO}_3^-$  removal (relative to DON) along flow  
231 paths of undisturbed ecosystems may explain this lack of correlation. The extent to which  
232 riparian zones influence nutrients varies spatially with geomorphology, soil texture, vegetation,  
233 and riparian zone development (McDowell et al., 1992, Mayer et al., 2007); and soils with high  
234 rates of leaching to ground water may bypass riparian processing. As nutrients leach down the  
235 soil profile, denitrification, biologic uptake, and storage are all potential mechanisms that could  
236 alter soil solution and stream N species concentrations. Investigation of soil profile processes and  
237 riparian zone spatial variability may help determine where and when watershed-scale N status  
238 can be inferred from these proxies. Alternatively, varied land-use (e.g. pasture, N fixing plant  
239 species, etc.) upstream of undisturbed sites is typically not reported in the literature, but is  
240 another possible explanation for the break down between terrestrial and water-based proxies.

241 While most observed correlations were consistent across latitudes, a few differed between the  
242 tropics and the temperate zone. The correlations of soil  $\delta^{15}\text{N}$  with foliar  $\delta^{15}\text{N}$ , foliar  $\delta^{15}\text{N}$  with  
243 net nitrification, and net nitrification with N mineralization were consistent across both tropical  
244 and temperate regions. Net nitrification and N mineralization were correlated with stream  
245 DIN:DON only in temperate regions. These data suggest that while terrestrial proxies may be a  
246 useful across biomes, stream DIN:DON requires further research to understand the extent of its  
247 applicability across space. The correlation between foliar and soil  $\delta^{15}\text{N}$  also differs across  
248 latitudes, in that the correlation in the tropics was much tighter than in the temperate zone. Bias  
249 in the literature towards natural abundance isotopic data from the temperate zone may explain  
250 why previous research looking at this relationship has been noisy (Craine et al., 2009).

251 Although we found that temporal (soil  $\delta^{15}\text{N}$ ) and spatial (stream DIN:DON) integrators of  
252 watershed N were correlated with short-term proxies (net nitrification and net N mineralization),  
253 water-based metrics did not correlate very well with most of the soil-based metrics of N  
254 availability or each other. Explicit comparisons of these proxies to each other, with a focus on  
255 how they are influenced by hot-spots, hot-moments, biological diversity, and N transformation  
256 between the soil-stream interface, will enhance their utility for understanding N availability at  
257 the ecosystem scale.

258

## 259 **5.0 Conclusions**

260 The labor and expense associated with fertilization studies to assess nutrient limitation  
261 requires that we develop proxies to infer soil nutrient status. While nitrification and  
262 mineralization most frequently correlate with other metrics, they are short-term proxies that vary  
263 over short spatial and temporal scales. Soil  $\delta^{15}\text{N}$  and dissolved N losses from streams are long-  
264 term integrators of N loss that have been relied on to advance our understanding of N cycling at  
265 the global scale (Martinelli et al., 1999; Amundson et al., 2003; Hedin et al., 2003; Brookshire et  
266 al., 2012), however their lack of correlation brings to light a need to better understand how these  
267 terrestrial and stream-based metrics vary in relation to each other and with nutrient limitation.

268 Understanding ecosystem N status at the watershed and landscape scale is a first step towards  
269 projecting their response to climate change and environmental pollution (Aber et al., 1998; Oren  
270 et al., 2001; Reich et al., 2004). Soil N status can determine the rate at which detrimental N  
271 losses occur, such as  $\text{NO}_3^-$  (a drinking water contaminant) and nitrous oxide (a potent greenhouse  
272 gas). Furthermore, it is becoming more evident that projections regarding the potential for a  
273 terrestrial  $\text{CO}_2$  sink, and concomitant feedbacks to the trajectory of climate change, are

274 dependent on the nutrient status of soils (Thornton et al., 2007; Zaehle et al., 2010; Wieder et al.,  
275 2015). The health and environmental implications of soil N status heighten the need to develop  
276 methodology to adequately assess long-term soil N availability.

277

## 278 **6.0 Author contribution**

279 M. Almaraz and S. Porder conceived research and designed study. M. Almaraz collected data  
280 and performed statistical analyses. M. Almaraz and S. Porder wrote the manuscript.

281

## 282 **7.0 Acknowledgments**

283 We want to thank J. Campbell, C. Neill, and W. Wilcke for soil samples. M. Otter, C. Tamayo  
284 and C. Silva for help with analyses. MA and SP received funding from NIH IMSD

285 R25GM083270, NSF DDIG GR5260021 and NSF EAR 1331841.

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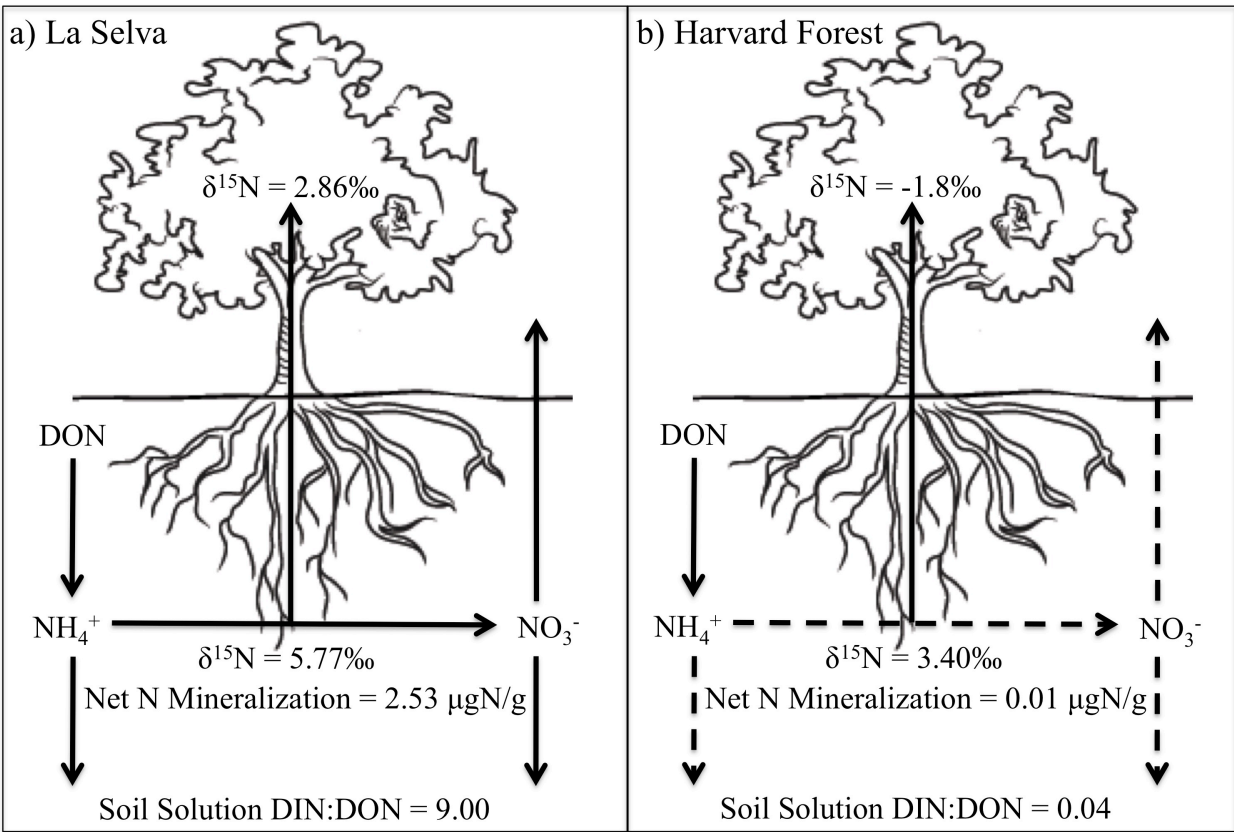
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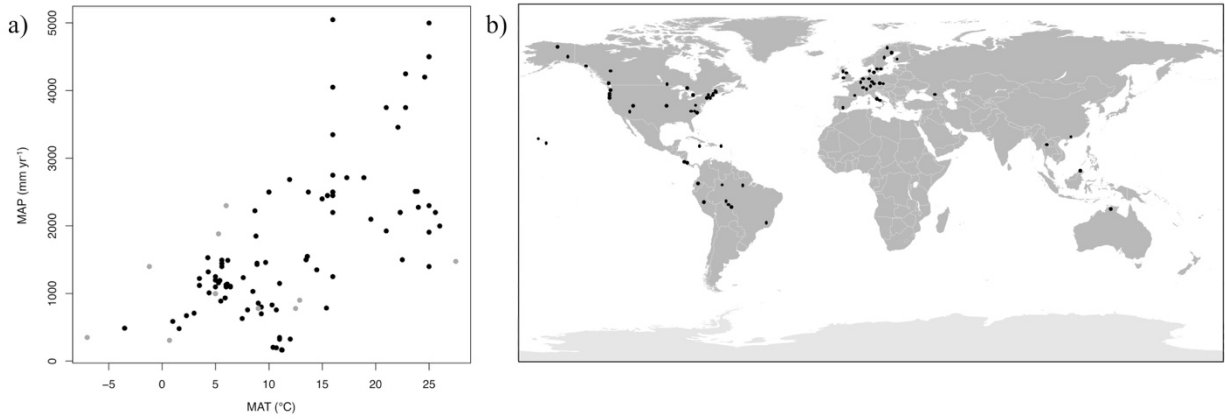
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**Figure 1.** Nitrogen availability values for a) a nitrogen rich tropical forest at the La Selva field station in Costa Rica, and for b) a nitrogen limited temperate pine forest at Harvard Forest, Massachusetts. Solid and dotted lines represent the relative magnitude of fluxes (i.e. net N mineralization, denitrification to the atmosphere, dissolved organic and inorganic nitrogen leaching), which are contingent on ecosystem nitrogen status.



515  
516 **Figure 2.** a) Distribution of grassland (grey) and forest (black) watershed mean annual  
517 temperature (MAT; °C) and mean annual precipitation (MAP; mm yr<sup>-1</sup>) included in meta-  
518 analysis (left), and b) location of 154 sites (some black dots represent multiple watersheds;  
519 right).  
520



529 **Supplemental Table 1.** Site description data, full citations, foliar and surface soil  $\delta^{15}\text{N}$  (<20 cm  
530 depth), net nitrification (<20 cm depth), net N mineralization (<20 cm depth), the ratio of  
531 dissolved inorganic to organic N forms (DIN:DON) in soil solution below the rooting zone (>20  
532 cm depth), and stream DIN:DON.  
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