Interactive comment on "Small phytoplankton contribution to the total
primary production in the Amundsen Sea" by Sang H. Lee et al.
Anonymous Referee #1
Received and published: 17 November 2016
The manuscripts "Small phytoplankton contribution to the 1 total primary production in the Amundsen
Sea" by Lee et al. presents observational data in Amundson during 1-14 January 2014 cruise and
discussed an important issue on the small phytoplankton contribution to the total primary production. I
found the data and discussion deserved for publication in BG with minor revision. I suggest the authors to
improve description of the differences between non-polynya and polynya regions, maybe a regroup those
stations in order to make the conclusions stronger.
\rightarrow Since our study region was separated into polynya and non-polynya areas based on sea ice
concentration data from National Snow & Ice Data Center during the cruise period (Fig. 1) as we
mentioned that in line 88-92, page 4, regrouping those stations based on the result outcome is rather
arbitrary. Therefore, we would like to stick with the previous two groups based on sea ice concentration.
It is also important to include time period of measurements when discuss comparison with other studies in
many places in the manuscript. Here are some details: L223-225: "our total 223 daily carbon uptake rate
in 224 polynya region (0.84 g C m-2 d-1) is within the range between Lee et al. (2012; 2.2 g C m-2 d-1)
and Kim et 225 al. (2015; 0.2 g C m-2 d-1)." The wide range of carbon uptake rates are mainly due to the
different measurement timing (or location). This is an example where it is necessary to add which month
(not just year) the data were measured when comparing those numbers.
\rightarrow Yes, the different carbon uptake rates among different studies are mainly due to the different
measurement timing. We indicated the time period of the measurement for each study for the comparison
of the rates in line 225-227 and line 239-240, page 10.
L274 states "small phytoplankton were higher in non-polynya region (Table 1)". L281 states 'diatoms are
relatively dominant in the non-polynya regions (Lee et al., 2012)'. Please explain why they are different
as we normally think diatom is large phytoplankton.
\rightarrow We are not saying diatom is small phytoplankton in this paragraph. As we mentioned in line 143-145,
the average contributions of small phytoplankton to the total chlorophyll-a concentration were 42.4 $\%$

33 (S.D. = \pm 37.2 %) for non-polynya based on different sizes of chlorophyll-a concentration which

34	indicating there were still some large amount of small phytoplankton (< 5μ m) although they were not
35	dominant group. Since it is rather confused, we rephrased it in line 284, page 12.
36	
37	In Fig 2-4, small phytoplankton were lower in non-polynya stations 3 and 3-1, higher in 1 and 2. Stations
38	1 and 2 had very low production and its ratio may not represent the ratio when bloom occurs in those
39	locations. It is necessary to note whether the ratios in Table 1 is the average of ratios in each station or
40	calculated from the average of chl-a, PP.
41	→ The ratios in Table 1 are the euphotic water column values averaged from all stations, non-polynya
42	station, and polynya stations. We clarified that in the caption of Table 1.
43	
44	L315: 'anticipating small-dominant ecosystem under warming oceans'. We have found increasing small
45	phytoplankton due to warming Arctic, but in Amundson, small phytoplankton contribution was found to
46	be higher under ice (cold) rather than in polynya (warm) in this study. It looks like we are heading to
47	large-dominant ecosystem under warming ocean in Amundson.
48	→ Polynya and non-polynya regions are different systems with different environmental conditions so that
49	we can not simply say that. That is a main reason for why we separated them in this study. Actually, the
50	data in Figure 7 included all stations from polynya and non-polynya regions.
51	

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4:11

1.4

- 141-

52 Interactive comment on "Small phytoplankton contribution to the total

53 primary production in the Amundsen Sea" by Sang H. Lee et al.

54 Anonymous Referee #2

- 55 Received and published: 12 December 2016
- 56

57 General Comments:

- 58 The manuscript "Small phytoplankton contribution to the total primary production in the Amundsen Sea"
- 59 by Lee et al. presents size-fractionated chlorophyll, particulate organic carbon/nitrogen, and
- 60 carbon/nitrogen uptake rates in the Amundsen Sea to characterize the contribution of small phytoplankton.
- As the authors state, this type of data is lacking in the Amundsen Sea, yet is invaluable for understanding
- 62 how the region might be altered by climate change. I commend the authors on the collection of a unique
- dataset, and given the importance of the data, would be excited to see this manuscript published in
- 64 Biogeosciences. However, it is my opinion that it should be reconsidered after major revisions for the
- 65 following reasons:
- There are strong statements re. the future importance and driving mechanism of small phytoplankton in
- 67 the Amundsen Sea based on limited evidence from that region, and rather extrapolated from other regions
- 68 (more northernWestern Antarctic Peninsula and Arctic Ocean). Ultimately I feel that the focus should
- 69 primarily be on establishing a baseline dataset for the region on small phytoplankton, rather than
- predictions that cannot be supported by the data presented (i.e. data from one year) and instead are based on data from other regions.
- We agree with the reviewer's opinion. So, we modified our manuscript to delete the prediction parts in
 line 36-38, page 2 and line 309-314, page 13.
- There is seemingly an inconsistency (or at best, a lack of explanation) between the demonstrated
- 75 importance of small phytoplankton outside the polynya region and the claim that small phytoplankton
- will grow in importance with climate change (won't the non-polynya region decrease in size with
- 77 increased warming?).
- 78 → Polynya and non-polynya regions are different systems with different environmental conditions so that
- 79 we can not simply say that. That is a main reason for why we separated them in this study. Actually, the
- 80 data in Figure 7 included all stations from polynya and non-polynya regions. Anyway, we modified our
- 81 manuscript to delete the prediction parts.
- There is a large focus on the comparison of data inside and outside of the polynya region, but with
- 83 limited justification for this comparison, or discussion of how the polynya may be altered by climate

84	change. Ultimately I agree that this comparison is valuable, but primarily in the context of establishing a
85	baseline dataset for the region.
86	\rightarrow We agree with the reviewer's opinion. So, we modified our manuscript to delete the prediction parts in
87	line 36-38, page 2 and line 309-314, page 13.
88	
89	- The Results section needs to be reorganized (see suggestions below).
90	\rightarrow We reorganized as reviewer suggested throughout the result section.
91	- There are numerous grammatical errors, some of which I have identified in the "Technical
92	Corrections" section.
93	\rightarrow We checked and revised the grammatical errors throughout the text.
94	
95	Specific Comments:
96	- Lines 54-62: I think it is important to indicate that Ducklow et al. (2007) and Montes-Hugo et al. (2009)
97	detail the western Antarctic Peninsula (WAP) that is a focus of the LTER (north of _68S), and do not
98	include the Amundsen Sea region.
99	→ We indicated that in line 56-57, page 3.
100	- Line 71: "in response to a regional warming trend" - I think this wording is too strong. Moline et al.
101	(2004) note the association between cryptophytes and low salinity water (likely glacial meltwater), and
102	hypothesize that cryptophytes will increase in importance given the predicted regional warming trend.
103	Regarding the association between cryptophytes and glacial meltwater, Moline et al. (2004) suggest that
104	this is salinity driven (they cite studies demonstrating cryptophytes tolerate/prefer lower salinity water), a
105	point that Moline made nearly a decade earlier (Moline and Prezelin 1996, MEPS).
106	→ We deleted the sentence.
107	- Lines 71-73: re. an example of food web alteration due to a shift in phytoplankton community
108	composition to smaller cells at least provide the example that krill do not feed efficiently on cryptophytes
109	(see Moline et al. 2004 for references).
110	\rightarrow We further discussed on that in line 321-325, page 13.
111	- Line 79: "environmental conditions" - could this not simply be referred to as climate change?
112	→Yes, it could. We revised it in line 81-82, page 4.
113	- Line 82: Consider renaming, e.g. "Water samples".
114	→We renamed it in line 85, page 4.
115	- Results section: this section is very tedious to read. Perhaps that is unavoidable given the results
116	presented (essentially a long list of averages and standard deviations). However, I think it would benefit

117 tremendously from some reorganization. All statistics should be reported in a consistent manner, e.g.

range followed by mean +/- SD in parentheses. Additionally, each topic has the same info presented, e.g. total/small cells, % contribution, inside/outside polynya. I think it would help guide the reader if this info

120 was presented in a consistent order for each topic.

121 \rightarrow We revised the result section as suggested.

- 122 Lines 273-275: The authors present strong evidence that small phytoplankton contribute more in the
- 123 non-polynya region than the polynya region. How might we expect the polynya to be altered with climate
- 124 change? It seems reasonable to expect that the non-polynya region will decrease in size, and thus reduce
- 125 the contribution of small phytoplankton. This is inconsistent with the stated motivation and implications
- of the paper (i.e. an increase in the contribution of small phytoplankton, and resulting decrease in primary
- 127 production), and needs to be addressed.
- 128 → Actually, polynya and non-polynya regions are different systems with different environmental
- 129 conditions so that we can not simply expect that. Actually, our non-polynya stations were not an ice free
- 130 open ocean in this study (see Figure 1). Increasing polynya region altered with climate change could
- 131 cause different conditions from previous original conditions. That is a main reason for why we separated
- them in this study. The data in Figure 7 included all stations from polynya and non-polynya regions.
- 133 Lines 299-304: the prediction of Moline et al. (2004) for an increase in the contribution of smaller
- 134 phytoplankton with expanding meltwater is for the portion of the WAP that is a focus of the LTER (north
- 135 of _68S), and did not explicitly include the Amundsen Sea region. Do the authors have any evidence
- 136 specific to their region of interest for a potential shift to smaller phytoplankton, as well as a driving
- 137 mechanism? If not, I do not think they can make strong statements re. the future of Amundsen Sea
- 138 phytoplankton community composition, as well as its impact on primary production (using the
- relationship in Fig. 7).
- 140 \rightarrow We deleted the sentence.
- Lines 305-315: this discussion should include the fact that krill do not efficiently feed on small
- 142 phytoplankton (see Moline et al. 2004 for references).
- 143 \rightarrow We further discussed on that in line 323-325, page 13.
- 144
- 145 Technical Corrections:
- 146 Lines 18-19: "Small-sized phytoplankton : : : ocean condition." rephrase.
- 147 \rightarrow We rephrased that in line 18-19, page 2.
- 148 Line 45: refer to Fig. 1.
- 149 \rightarrow We referred to Fig. 1 in line 45, page 3.
- 150 Line 65: "In an expecting : : :" rephrase.
- 151 \rightarrow We rephrased that in line 67, page 3.

- 152 Line 67: "In consistent : : :" rephrase.
- 153 \rightarrow We rephrased that in line 69-71, page 3.
- 154 Line 73: "higher trophic levels" and "subsequent food web" are redundant.
- 155 \rightarrow We deleted subsequent food web in line 322, page 13.
- 156 Line 76: "what extend" rephrase.
- 157 \rightarrow We rephrased that in line 79, page 4.
- 158 Line 78: "marine ecosystem : : : ongoing changes" rephrase.
- 159 \rightarrow We rephrased that in line 81-82, page 4.
- 160 Lines 83-88: refer to Fig. 1 in here somewhere.
- 161 \rightarrow We referred to Fig. 1 in line 87, page 4.
- 162 Lines 91-22: "were belong" rephrase.
- 163 \rightarrow We rephrased that in line 95, page 5.
- 164 Line 95: "biological and chemical property" please be specific.
- 165 Actually I tried to mention that other researchers collected water samples for their own biological and
- 166 chemical research. We deleted that since it might be confused.
- 167 Lines 109-113: the information re. the isotope tracer technique, light depths, and light
- 168 sensor was already provided.
- 169 \rightarrow We deleted the same information in line 113-117, page 5.
- 170 Lines 137-138: "integrated from six different light depths" change to "depth integrated"?
- →We changed it to depth integrated in line 141, page 6, line 163, page 7, and line 183, page 8.
- 172 Lines 141-143: "In the Amundsen Sea : : : 2014 : : :" unnecessary info (the cruise location and date
- 173 has already been specified).
- 174 \rightarrow We deleted unnecessary info in line 146-147, page 7.
- 175

177	Small phytoplankton contribution to the total primary production in
178	the Amundsen Sea
179	
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192	

193 Abstract

Small-sized phytoplankton isare anticipated to be more important for phytoplankton community 194 195 in a recently changing ocean condition. However, little information on the contribution of small-sized 196 phytoplankton to overall phytoplankton production is currently available in the Amundsen Sea. To 197 determine the contributions of small-sized phytoplankton to total biomass and primary production, carbon and nitrogen uptake rates of total and small-sized phytoplankton were obtained from 12 productivity 198 199 stations in the Amundsen Sea. The daily carbon uptake rates of total phytoplankton averaged in this study were 0.42 g C m⁻² d⁻¹ (S.D. = \pm 0.30 g C m⁻² d⁻¹) and 0.84 g C m⁻² d⁻¹ (S.D. = \pm 0.18 g C m⁻² d⁻¹) whereas 200 the daily total nitrogen (nitrate and ammonium) uptake rates were 0.12 g N m⁻² d⁻¹ (S.D. = \pm 0.09 g N m⁻² 201 d^{-1}) and 0.21 g N m⁻² d^{-1} (S.D. = \pm 0.11 g N m⁻² d^{-1}), respectively for non-polynya and polynya regions, 202 203 which were within the ranges reported previously. Small phytoplankton contributed 26.9 % and 27.7 % to 204 the total carbon and nitrogen uptake rates of phytoplankton in this study, respectively, which were relatively higher than the chlorophyll-a contribution (19.4 %) of small phytoplankton. For a comparison 205 206 of different regions, the contributions for chlorophyll-a concentration and primary production of small phytoplankton averaged from all the non-polynya stations were 42.4 % and 50.8 %, which were 207 208 significantly higher than those (7.9 % and 14.9 %, respectively) in polynya region. A strong negative correlation ($r^2 = 0.790$, p < 0.05) was found between the contributions of small phytoplankton and the 209 total daily primary production of phytoplankton in this study. This finding implies that daily primary 210 211 production decreases as small phytoplankton contribution increases, which is mainly due to the lower 212 carbon uptake rate of small phytoplankton than large phytoplankton. Under ongoing environmental 213 changes caused by global warming, a potential decrease of total primary production would be led by increasing contribution of small phytoplankton in the Amundsen Sea. 214

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217 Keywords: Phytoplankton, Primary production, Polynya, Amundsen Sea

219 1. Introduction

The Amundsen Sea is located in the West Antarctica between the Ross Sea and Bellingshausen 220 221 Sea (Fig. 1), which is one of the least-biologically studied regions in the Southern Ocean. Recently several international research programs (KOPRI Amundsen project, iSTAR, ASPIRE, and DynaLiFe) 222 223 were launched to understand this remote area. Field-measurement data revealed that annual primary production of phytoplankton reaching to 220 g C m⁻² y⁻¹ in the Amundsen Sea polynya is as high as that 224 of Ross Sea polynya (200 g C m⁻² y⁻¹) which was previously known for the highest productivity region in 225 the Southern Ocean (Lee et al., 2012). Given the fact that the ehl-achlorophyll-a concentration averaged 226 from all the chlorophyll-achl a-measured stations was twice higher than that of the only productivity-227 228 measured stations, Lee et al., (2012) argued that the annual production in the Amundsen Sea polynya could be even two-fold higher than that of Ross Sea polynya. 229

Over the past several decades a rapid climate change has been detected and subsequently 230 physical changes have occurred in the marine ecosystem in the western Antarctic Peninsula (WAP) 231 232 mainly based on the results from Palmer Antarctic Long-Term Ecological Research project which focused on the north of ~69 °S (Rückamp et al. 2011; Ducklow et al. 2007; Montes-Hugo et al. 2009). Recent 233 234 studies revealed that the Thwaites Glacier in Pine Island Bay is retreating fast and the ice volume loss in the nearby Getz Ice shelf is accelerating (Joughin et al., 2014; Paolo et al., 2015). Shoaling warm 235 Circumpolar Deep Water is believed to be a main reason for the ice sheet mass loss largely caused by ice 236 237 shelf basal melt underside of the ice shelves (Yager et al. 2012; Schmidtko et al. 2014). The climate 238 change from a cold-dry polar-type to a warm-humid sub-Antarctic-type drives subsequent changes in 239 ocean biological productivity along the WAP shelf over the recent three decades (Montes-Hugo et al. 240 2009).

241 Phytoplankton as the base of oceanic food webs can be an indicator for changes in marine
242 ecosystems responding to <u>current climategnvironmental</u> changes (Moline et al., 2004; Wassman et al.,
243 2011; Arrigo et al., 2015). For example, a shift in phytoplankton community structure from large diatoms
244 to relatively small cryptophytes could be tightly associated with changes in glacial melt-water runoff and

245	reduced surface water salinity (Moline et al., 2004). In an expecting warmer ocean condition, small sized
246	phytoplankton is anticipated to contribute more to total phytoplankton community and thus marine
247	ecosystems (Morán et al., 2010; Li et al., 2009; Lee et al., 2013). In consistent, Li et al. (2009) found
248	increasing small sized phytoplankton in the Canada Basin in the Arctic Ocean under freshening surface
249	waters which results in a stronger stratification and lower nutrient supply in the upper water column.
250	Moreover, in the Antarctic Ocean, Moline et al. (2004) found a consistent transition from large diatoms to
251	small cryptophytes associated with glacial melt water in the coastal waters along the Antarctic Peninsula
252	in response to a regional warming trend. This change in dominant phytoplankton community from large
253	to small cells will likely cause further alteration of higher trophic levels and subsequent food web (Moline
254	et al., 2004)HoweverTo date, little information on the contribution of small-sized phytoplankton to
255	primary production is available in the Antarctic Ocean (Saggiomo et al. 1998), especially in the
256	Amundsen Sea with a rapid melting of ice shelf (Yager et al. 2012; Schmidtko et al. 2014). Thus, the
257	main objective in this study is to determine that to what extends small-sized phytoplankton contributes to
258	overall total biomass and primary production in the Amundsen Sea to lay the groundwork for the future
259	monitoring of marine ecosystem change responding to ongoing changes in environmental conditions.

261 2. Materials and methods

262 2.1. <u>Water Ss</u>ampl<u>eing</u>s

Water samples were collected for carbon and nitrogen uptake measurements of phytoplankton in 263 264 the Amundsen Sea (Fig. 1) during the KOPRI Amundsen cruise from 1 to15 January, 2014 onboard the Korean Research Icebreaker ship Araon. Using a dual stable isotope technique (Lee et al., 2012; Kim et 265 al., 2015), the experiments of carbon and nitrogen uptake rates of phytoplankton were conducted at 12 266 selected productivity stations including 2 revisited-stations (St. 3-1 and St. 19-1) when on-deck 267 incubations were available during daytime at oceanographic survey stations. Based on sea ice 268concentration data from National Snow & Ice Data Center during the cruise period (Fig. 1), our study 269 region was further separated into polynya and non-polynya areas for comparison based on sea ice 270

distribution and concentration during the cruise period. Four stations (St. 1, St. 2, St. 3, and St. 3-1) among the 12 stations were-belong to non-polynya region and the rest of the stations were-belong to polynya region.

After 6 light depths (100, 50, 30, 12, 5, and 1% penetration of the surface irradiance, PAR) were determined with an LI-COR underwater 4π light sensor, water samples for the uptake experiments as well as biological and chemical property analysis were obtained from a CTD-rosette sampler system equipped with 24 10-L Niskin bottles.

278 2.2. Total and size-fractionated chlorophyll-a concentration

279 Water samples for total and size-fractionated chlorophyll-a concentrations of phytoplankton were 280 obtained at the 12 productivity stations. Total chlorophyll-a concentrations were measured at six different light depths (100, 50, 30, 12, 5 and 1% of PAR). For size-fractionated chlorophyll-a concentrations, water 281 282 samples were collected at three light depths (100, 30, and 1 %). Water samples (0.3-0.5 L) for total 283 chlorophyll-a concentrations were filtered using Whatman glass fiber filters (GF/F; 25 mm). For different size-fractionated chlorophyll-a concentrations water samples (0.7-1 L) were passed sequentially through 284 285 20 and 5 μ m Nucleopore filters (47 mm) and 0.7 μ m GF/F filters (47 mm). After the filters were extracted 286 using the method described by Kim et al. (2015), all chlorophyll-a concentrations were subsequently 287 determined onboard using a Trilogy fluorometer (Turner Designs, USA). The methods and calculations 288 for chlorophyll-a were based on Parsons et al. (1984).

289 2.3. Carbon and nitrogen uptake experiments

290 Carbon and nitrogen uptake experiments of phytoplankton were executed by a 13 C 15 N dual 291 isotope tracer technique previously applied for the Amundsen Sea (Lee et al. 2012; Kim et al. 2015). In 292 this study, basically we followed same procedure of Lee et al. (2012). In brief, six light depths (100, 50, 293 30, 12, 5, and 1%) were determined with an LI-COR underwater 4π light sensor (LI-COR Inc., Lincoln, 294 Nebraska, USA) lowered with CTD/rosette sampler. Water sample from each light depth was transferred 295 into different screened polycarbonate incubation bottle (1 L) which matches with each light depth. The

productivity bottles were incubated in large polycarbonate material incubators cooled with running 296 297 surface seawater on deck under natural light conditions, after the water samples were inoculated with labeled carbon (NaH13CO3) and nitrate (K15NO3) or ammonium (15NH4Cl) substrates. After 4-5 h 298 299 incubations, the incubated waters were well mixed and distributed into two filtration sets for the carbon and nitrogen uptake rates of total (> 0.7 μ m) and small-sized cells (< 5 μ m). The incubated waters (0.3 L) 300 for total uptake rates were filtered through pre-combusted GF/F filters (24 mm diameter), whereas waters 301 samples (0.5 L) for the uptake rates of small-sized cells were passed through 5 µm Nuclepore filters (47 302 mm) to remove large-sized cells (> 5 µm) and then the filtrate was passed through GF/F (24 mm) for the 303 small-sized cells (Lee et al., 2013). The values for large phytoplankton in this study were obtained from 304 305 the difference between small and total fractions (Lee et al., 2013). The filters were immediately preserved 306 at -80°C until mass spectrometric analysis. After acid fuming overnight to remove carbonate, the concentrations of particulate organic carbon (POC) and nitrogen (PON) and the abundance of ¹³C and 307 308 ¹⁵N were determined by a Finnigan Delta+XL mass spectrometer at the Alaska Stable Isotope Facility, 309 USA.

All contribution results of small phytoplankton in this study were estimated from comparison of small phytoplankton to total phytoplankton integral values from 100 to 1 % light depth at each station based on the trapezoidal rule. Daily carbon and nitrogen uptake rates of phytoplankton were based on our hourly uptake rates measured in this study and a 24-h photoperiod per day during the summer period in the Amundsen Sea (Lee et al., 2012).

315

316 3. Results

317 3.1. Chlorophyll-a, POC, and PON contributions of small phytoplankton

The <u>depth-integrated</u> total (large + small phytoplankton) chlorophyll-a concentration integrated from six different light depths ranged from 11.1 to 80.3 mg chl-a m⁻² (mean \pm S.D. = 57.4 \pm 25.2 mg chl-a m⁻²), whereas small (< 5 µm) chlorophyll-a concentration ranged from 3.9 to 9.4 mg chl-a m⁻² (mean \pm

321	S.D. = 5.7 ± 1.7 mg chl-a m ⁻²) in this study (Fig. 2). The average-contribution of small phytoplankton to
322	the total chlorophyll-a concentration was $\frac{4.9-76.5 \% 19.4 \%}{19.4 \%} (S.D. = 19.4 \pm 26.0 \%)$ -ranging from 4.9 to
323	$\frac{76.5 \text{ \%}}{1000}$. In the Amundsen Sea, <u>lL</u> arge phytoplankton (> 5 µm) were generally predominant
324	(approximately 80 %) during the cruise period in 2014 based on different-sized chlorophyll-a
325	concentrations. For a regional comparison, the average contributions of small phytoplankton to the total
326	chlorophyll-a concentration were 42.4 % (S.D. = \pm 37.2 %) and 7.9 % (S.D. = \pm 3.5 %) for non-polynya
327	and polynya regions, respectively (Table 1). The chlorophyll-a contribution of small phytoplankton was
328	larger in the non-polynya regionstations than the polynya regionstations although they were no
329	significantly different (t-test, $p = 0.16$).

330 The total integral POC concentration of phytoplankton displayed no large spatial variation ranging from 4.72 to 9.22 mg C m⁻² (Fig. 3). In comparison, the total integral PON concentration of 331 332 phytoplankton ranged from was 0.76 to _1.74 mg C m⁻². The POC contribution of small phytoplankton 333 similarly from <u>was</u> 30.8 to 67.2 % (mean ± S.D. = 41.3 ± 11.5 %) in the Amundsen Sea (Fig. 3). 334 335 Specifically, the POC and PON contributions of small phytoplankton averaged from all the productivity stations in the polynya regionstations were 36.9 % (S.D. = ± 4.6 %) and 37.0 % (S.D. = ± 6.9 %), 336 respectively. In comparison, the POC and PON contributions of small phytoplankton averaged from the 337 non polynya stations, whereas they were 49.5 % (S.D. = \pm 14.4 %) and 50.0 % (S.D. = \pm 15.1 %), 338 339 respectively in the non-polynya region (Table 1).

340 3.2. Carbon uptake rate contributions of small phytoplankton

The <u>depth-integrated</u> total daily carbon uptake rates of phytoplankton (large + small phytoplankton) integrated from six different light depths ranged from was 150.4 to _1213.4 mg C m⁻² d⁻¹ with an average of (696.5 mg C m⁻² d⁻¹ (S.D. = _1± 298.4 mg C m⁻² d⁻¹) in this study (Fig. 4). In contrast, the rates of small phytoplankton ranged between 58.6 and 266.4 mg C m⁻² d⁻¹ with an average of (124.9 mg C m⁻² d⁻¹ (S.D. = _1± 62.4 mg C m⁻² d⁻¹). Small phytoplankton contributed 26.9 % (S.D. = ± 29.3%) to total daily carbon uptake rate of total phytoplankton.

347	Specifically, the total daily carbon uptake rates of phytoplankton ranged from was 150.4 to _796.4
348	mg C m ⁻² d ⁻¹ with an average of (415.0 mg C m ⁻² d ⁻¹ (S.D. =± 298.2 mg C m ⁻² d ⁻¹) in the non-polynya
349	region, whereas they ranged from <u>it was</u> 654.8-to-1213.4 mg C m ⁻² d ⁻¹ with an average of (837.3-mg C m ⁻²
350	d^+ (S.D. = ± 184.1 mg C m ⁻² d ⁻¹) in <u>the</u> polynya region. The total daily carbon uptake rates of
351	phytoplankton were significantly higher (t-test, $p < 0.05$) in <u>the polynya regionstations</u> than <u>the non-</u>
352	polynya regionstations. The rates of small phytoplankton-ranged between was_58.6-and_193.6 mg C m ⁻²
353	d^{-1} with an average of (126.5 mg C m ² - d^{-1} (S.D. = _± 55.2 mg C m ⁻² d^{-1}) in the non-polynya region,
354	whereas they ranged from it was 62.2-to $_{2}$ 266.4 mg C m ⁻² d ⁻¹ with an average of (124.1-mg C m ⁻² -d ⁻¹ (S.D.
355	$=$ _± 69.3 mg C m ⁻² d ⁻¹) in <u>the</u> polynya region. The daily carbon uptake rates of small phytoplankton were
356	not significantly different (t-test, $p > 0.05$) between <u>the polynya</u> and non-polynya stations. The average
357	contributions of small phytoplankton to total daily carbon uptake rates were 50.8 % (S.D. = \pm 42.8 %) and
358	14.9 % (S.D. = \pm 8.4 %), respectively for <u>the</u> non-polynya and polynya regions (Table 1). The average
359	contributions were largely different between the polynya and non-polynya regions but they were not
360	statistically significant (t-test, $p > 0.05$).

361 3.3. Nitrogen uptake rate contributions of small phytoplankton

362 The <u>depth-integrated</u> total daily nitrate uptake rates of phytoplankton (large + small phytoplankton) ranged from was 34.0-to-174.2 mg N m⁻² d⁻¹ with an average of (93.7-mg N m⁻² d⁺ (S.D. = 363 \pm 43.2 mg N m⁻² d⁻¹), whereas the rates of small phytoplankton ranged from was 6.1-to _40.9 mg N m⁻² d⁻¹ 364 with an average of (19.0 mg N m² d⁻¹ (S.D. = \pm 11.3 mg N m² d⁻¹) in this study (Fig. 5). Small 365 phytoplankton contributed 21.5 % (S.D. = ± 11.1 %) to total daily nitrate uptake rates. In comparison, the 366 total daily ammonium uptake rates of phytoplankton ranged from 12.4 to 173.8 mg N m⁻² d⁻¹ with an 367 average of (86.7-mg N m⁻²-d⁻¹-(S.D. = \pm 75.9 mg N m⁻² d⁻¹), whereas the rates of small phytoplankton 368 ranged from 9.1 to 81.1 mg N m⁻² d⁻¹ with an average of (25.7 mg N m⁻² d⁻¹ (S.D. = \pm 21.1 mg N m⁻² d⁻¹) 369 in this study (Fig. 6). Small phytoplankton contributed 38.7 % (S.D. = ± 24.9 %) to total daily ammonium 370 uptake rates. The contributions of small phytoplankton were significantly higher in ammonium uptake 371 rate than nitrate uptake rate (t-test, p < 0.05). 372

373	Specifically fFor different regions, the total daily nitrate uptake rates of phytoplankton ranged
374	$\frac{\text{from were}}{\text{from were}} 34.0 \text{-to-}_{1}142.1 \text{ mg N m}^{-2} \text{ d}^{-1} \text{ with an average of } (71.9 \text{-mg N m}^{-2} \text{-d}^{-1} \text{(S.D.}_{-}\text{-}_{-} \pm 48.4 \text{ mg N m}^{-2} \text{ d}^{-1})$
375	in <u>the</u> non-polynya region whereas they ranged from<u>and</u> 44.2-to-174.2 mg N m⁻² d⁻¹ with an average of
376	(104.6 $\frac{\text{mg N m}^2 \text{d}^4}{(\text{S.D.} - \pm 39.0 \text{ mg N m}^2 \text{d}^{-1})}$ in <u>the</u> polynya region, respectively. In comparison, the
377	daily nitrate uptake rates of small phytoplankton ranged from were 7.5-to-26.6 mg N m ⁻² d ⁻¹ with an
378	average of (16.7 mg N m ⁻² d ⁻⁴ (S.D. = \pm 7.8 mg N m ⁻² d ⁻¹) and from 6.1 to \pm 40.9 mg N m ⁻² d ⁻¹ with an
379	average of (20.1 mg N m ⁻² d ⁺ (S.D. = \pm 13.1 mg N m ⁻² d ⁻¹), respectively for the non-polynya and polynya
380	regions. The contributions of small phytoplankton to the total daily nitrate uptake rates were 28.2 % (S.D.
381	= \pm 15.9 %) in <u>the</u> non-polynya region and 18.1 % (S.D. = \pm 6.8 %) in <u>the</u> polynya region, respectively
382	(Table 1). The total daily ammonium uptake rates of total phytoplankton ranged betweenwere 12.3 and
383	106.1 mg N m ⁻² d ⁻¹ (mean \pm S.D. = 49.7 \pm 41.2 mg N m ⁻² d ⁻¹) in the non-polynya region and between 18.1
384	and _269.3 mg N m ⁻² d ⁻¹ (mean \pm S.D. = 105.2 \pm 84.6 mg N m ⁻² d ⁻¹) in the polynya region. In comparison,
385	the rates of small phytoplankton ranged between 9.1 and 22.4 mg N m ⁻² d ⁻¹ (mean \pm S.D. = 15.8 \pm 6.4 mg
386	N m ⁻² d ⁻¹) in <u>the</u> non-polynya region and between 9.9 and 81.1 mg N m ⁻² d ⁻¹ (mean \pm S.D. = 30.7 \pm 24.5
387	mg N m ⁻² d ⁻¹) in <u>the</u> polynya region. Small phytoplankton contributed 52.8 % (S.D. = \pm 40.5 %) and 31.6 %
388	$(\frac{\text{S.D.}=\pm 10.1 \text{ \%})$ to the total daily ammonium uptake rates in <u>the</u> non-polynya and polynya regions,
389	respectively which were not significantly different (t-test, $p = 0.37$).

390 The total integral daily nitrogen uptake rate (nitrate + ammonium uptake rates) of phytoplankton ranged from was 46.4 to 443.5 mg N m⁻² d⁻¹ (mean \pm S.D. = 180.4 \pm 106.7 mg N m⁻² d⁻¹) in this study. 391 For <u>the</u> non-polynya and polynya regions, they ranged from were 46.4 to 248.1 mg N m⁻² d⁻¹ (mean \pm 392 **S.D.** = $121.6 \pm 89.3 \text{ mg N m}^2 \text{ d}^{-1}$) and from $91.7 \text{ to }_{-}443.5 \text{ mg N m}^2 \text{ d}^{-1}$ (mean $\pm \text{ S.D.} = 209.8 \pm 107.3 \text{ mg}$) 393 N m⁻² d⁻¹), respectively. In comparison, the total integral daily nitrogen uptake rates of small 394 phytoplankton ranged from were 16.6 to 46.6 mg N m⁻² d⁻¹ (mean \pm S.D. = 32.5 \pm 13.2 mg N m⁻² d⁻¹) and 395 from 17.6 to _122.0 mg N m⁻² d⁻¹ (mean \pm S.D. = 50.8 \pm 32.4 mg N m⁻² d⁻¹) for the non-polynya and 396 polynya regions, respectively. Small phytoplankton contributed 36.2 % (S.D. = ± 23.0 %) to the total 397 integral daily nitrogen uptake rates in the non-polynya region, whereas they contributed 23.5 % (S.D. = ± 398

6.0 %) for <u>the polynya region (Table 1)</u>. The integral daily nitrogen uptake rates and contributions of small phytoplankton were not statistically different between <u>the non-polynya and polynya regions</u>.

401

402 4. Discussion and conclusion

403 The total daily carbon uptake rates of phytoplankton averaged for the non-polynya and polynya regions were 0.42 g C m⁻² d⁻¹ (S.D. = \pm 0.30 g C m⁻² d⁻¹) and 0.84 g C m⁻² d⁻¹ (S.D. = \pm 0.18 g C m⁻² d⁻¹), 404 405 respectively in this study. According to the previous reports in the Amundsen Sea (Lee et al., 2012; Kim et al., 2015), the total daily carbon uptake rates ranged from 0.2 to 0.12 g C m⁻² d⁻¹ in the non-polynya 406 region. Our rate (0.42 g C m⁻² d⁻¹) in the non-polynya region is somewhat higher than those reported 407 408 previously but they are not significantly different (t-test, p = 0.77). In comparison, our total daily carbon uptake rate in the polynya region (0.84 g C m⁻² d⁻¹) is within the range between Lee et al. (2012; 2.2 g C 409 m⁻² d⁻¹) and Kim et al. (2015; 0.2 g C m⁻² d⁻¹). The carbon uptake rates of phytoplankton in Lee et al. 410 (2012) and Kim et al. (2015) were measured during December 21, 2010-January 23, 2011 and February 411 412 11 to March 14, 2012, respectively. Our measurements in this study were executed mainly during January 1-15, 2014. For the Amundsen polynya region, a large seasonal variation in the total daily carbon uptake 413 414 rate of phytoplankton was already reported by Kim et al. (2015) and Arrigo et al. (2012) based on filedmeasured data and satellite-derived approach, respectively. It is appeared that this seasonal variation 415 largely depends on the bloom stage of phytoplankton which peaks during the late December-January and 416 417 terminates at late February (Arrigo and van Dijken 2003; Arrigo et al., 2012; Kim et al., 2015). The 418 carbon uptake rates of phytoplankton in Lee et al. (2012) and Kim et al. (2015) were measured during 419 December 21 January 23, 2010 and February 11 to March 14, 2012, respectively. Our measurements in this study were executed mainly during January 1-15, 2014. 420

421 The total daily nitrogen uptake rates of phytoplankton were 0.12 g N m⁻² d⁻¹ (S.D. = \pm 0.09 g N m⁻ 422 2 d⁻¹) and 0.21 g N m⁻² d⁻¹ (S.D. = \pm 0.11 g N m⁻² d⁻¹) for non-polynya and polynya regions, respectively 423 in this study. Previous studies reported that the total daily nitrogen uptake rates in non-polynya region 424 were 0.24 g N m⁻² d⁻¹ during Dec. 21, 2010-Jan. 23, 2011 in 2010/2011 and 0.04 g N m⁻² d⁻¹ during Feb. 11

to Mar. 14, 2012 in 2012 whereas the uptake rates in polynya region were 0.93 g N m⁻² d⁻¹ in 2010/2011 425 and 0.06 g N m⁻² d⁻¹ in 2012 in the Amundsen Sea (Lee et al., 2012; Kim et al., 2015). Our total daily 426 427 nitrogen uptake rates of phytoplankton in non-polynya and polynya regions were between the two 428 previous studies (Lee et al., 2012; Kim et al., 2015). Based on the nitrate and ammonium uptake rates in 429 this study, f-ratios (nitrate uptake rate/nitrate+ammonium uptake rates) averaged for non-polynya and 430 polynya regions were 0.62 (S.D. = \pm 0.08) and 0.54 (S.D. = \pm 0.20), respectively. These ratios also were 431 between the two previous studies. Although they were not significant different because of a large spatial variation, larger f-ratios in non-polynya than in polynya region are consistent with the results of the 432 previous studies (Lee et al., 2012; Kim et al., 2015). At this point, we do not have a solid explanation for 433 434 that but a further future study is needed for the higher *f*-ratio mechanism in non-polynya region.

435 The percent contributions of small phytoplankton in terms of chlorophyll-a, POC/PON, daily 436 carbon and nitrogen uptake rates are shown in Table 1. The overall contribution of small phytoplankton to the total chlorophyll-a concentration for all the productivity stations was 19.4 % ($\frac{S.D.}{2} \pm 26.0$ %) which 437 438 is significantly (t-test, p < 0.05) lower than the POC contribution (41.1 ± 10.6 %). This is consistent with 439 the result in the Chukchi Sea, Arctic Ocean reported by Lee et al. (2013). They explained that higher POC 440 content per chlorophyll-a unit of small phytoplankton could cause their higher POC contribution (Lee et al., 2013). Given C/N ratio (mean \pm S.D. = 6.6 \pm 0.6) and δ^{13} C (mean \pm S.D. = -25.9 \pm 1.0 ‰) of sample 441 442 filters attained for POC and PON in this study, our filtered samples are believed to be mainly phytoplankton-originated POC and PON (Kim et al., 2016). Thus, a significant potential overestimated 443 POC contribution of non-phytoplankton materials could be excluded for the higher POC contribution than 444 445 chlorophyll-a contribution of small phytoplankton. Therefore, small phytoplankton contributions based on conventional assessments of chlorophyll-a concentration might lead an underestimated contribution of 446 447 small phytoplankton (Lee et al., 2013). In fact, several authors argued that chlorophyll-a concentration might be not a good index for phytoplankton biomass since it largely depends on environmental factors 448 449 such as nutrient and light conditions as well as dominant groups and physiological status of 450 phytoplankton (Desortová 1981; Behrenfeld et al., 2005; Kruskopf and Flynn, 2006; Behrenfeld and Boss

2006). However, the effects of non-phytoplankton carbon materials such as extracellular carbon mucilagecan not be completely excluded for the POC contribution as discussed below.

453 The overall contributions of carbon and nitrogen (nitrate and ammonium) uptake rates of small phytoplankton at all the productivity stations in this study are similar with 26.9 % and 27.7 %, 454 455 respectively. These contributions are relatively higher than the chlorophyll-a contribution of small phytoplankton but they are not statistically different (t-test, p > 0.05). In general, the contribution of daily 456 457 ammonium uptake rate of small phytoplankton are significantly (t-test, p < 0.05) higher than the contribution of daily nitrate uptake rate of small phytoplankton at all the stations in this study. It is well-458 known for the ammonium preference of small phytoplankton in various regions (Koike et al., 1986; 459 460 Tremblay et al., 2000, Lee et al., 2008; Lee et al., 2013).

In terms of the contributions in different regions, all the contributions (Echlorophyll-a, POC/PON, 461 462 carbon and nitrogen uptake rates) of small phytoplankton were higher in the non-polynya region than in the polynya region (Table 1). In addition, the chlorophyll-a contribution of small phytoplankton (mean ± 463 S.D. = 7.9 \pm 3.5 %) was significantly (t-test, p < 0.01) lower than the POC contribution (mean \pm S.D. = 464 36.9 ± 4.6 %) in the polynya region, whereas they were not statistically different in the non-polynya 465 466 region (Table 1). This indicates that small phytoplankton contributed more to the total POC than the 467 chlorophyll-a concentration in the polynya region. We do not have species compositions of phytoplankton in this study, but previous results reported that Phaeocystis sp. are dominant in the Amundsen Sea 468 polynya region whereas diatoms are relatively dominant in the non polynya regions (Lee et al., 2012). 469 470 Generally, Phaeocystis spp. release a large portion (up to 46 %) of extracellular carbon mucilage which 471 makes their colonial form (Matrai et al., 1995). This non-phytoplankton carbon material without 472 chlorophyll-a might cause a higher POC contribution of small phytoplankton in the polynya region during this study. In fact, the contribution of the daily carbon uptake rates of small phytoplankton $(14.9 \pm 8.4 \%)$ 473 was not as high as the POC contribution (36.9 \pm 4.6 %) in the polynya region. The chlorophyll-a 474 475 contributions of small phytoplankton were lower than that those of the daily carbon uptake rate in this 476 study, which is consistent with the results from polynya and marginal ice zone stations in the Ross Sea,

Antarctica during austral spring and summer (Saggiomo et al., 1998). They reported that the chlorophyll-a and primary production contributions of pico-phytoplankton (< 2 μ m) were 29 % and 40 % at polynya stations whereas the contributions were 17 % and 32 % at marginal ice zone stations, respectively. In <u>the</u> polynya region, they found much higher contributions in chlorophyll-a and primary production of small phytoplankton than those in this study although their size of small phytoplankton is somewhat smaller than our size (< 5 μ m).

483 We found a strong negative correlation ($r^2 = 0.790$, p < 0.05) between the productivity contributions of small phytoplankton and total daily carbon uptake rates of phytoplankton in the 484 Amundsen Sea (Fig. 7), which implies that daily primary production decreases as small phytoplankton 485 contribution increases. This is mainly because of the relatively lower carbon uptake rate of small 486 487 phytoplankton than large phytoplankton in the Chukchi Sea, Arctic Ocean reported by Lee et al. (2013). Moline et al. (2004) suggested that further warming air temperatures will increase inputs of glacial 488 water and subsequently increase the contributions of small phytoplankton over large melting 489 phytoplankton community (Moline et al. 2004). If these small phytoplankton were dominant under 490 ongoing more melting conditions of glaciers, a potential increasing contribution of small phytoplankton 491 492 might cause a subsequent decrease in the total primary production of phytoplankton in the Amundsen Sea based on this study in Figure 7. 493

494 In respect to food quality of small phytoplankton as a basic food source to herbivores, 495 macromolecular compositions such as proteins, lipids, and carbohydrates as photosynthetic-end products 496 will be needed for better understanding alterations of small cells-dominant marine ecosystem in response 497 to ongoing environmental changes (Lee et al., 2013). According to Kang et al. (accepted), small phytoplankton- assimilate more food materials and calorific contents per unit of chlorophyll-a 498 499 concentration and thus provide more contributions in respect to energy aspect than other phytoplankton community in the East/Japan Sea. However, Fthis change in dominant phytoplankton community from 500 501 large to small cells will likely cause further alteration of higher trophic levels because of prey size itself available to higher trophic grazers and subsequent food web (Moline et al., 2004).- In conclusion, 502

503	monitoring the contributions of small-sized phytoplankton to total biomass and primary production of
504	total phytoplankton community could be important as a valuable indicator to sense futureenvironmental
505	changes in marine ecosystem-under ongoing various climate-associated environmental changes. Moreover,
506	further detailed studies for macromolecular compositions of small phytoplankton will be necessary for the
507	anticipating small-dominant ecosystem under warming oceans.
508	

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513

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590 Table caption

501	Table 1 Contributions (%)) of chloroph	villa DOC	DOM and carbon an	d nitrogan u	ntaka rates) of small
391	Table 1. Contributions (70	or emotoph	yn a, 100	, 1 Orv, and Carbon an	a maogen a	plake fates) of shall

- 592 phytoplankton in the Amundsen Sea. <u>Contributions of chlorophyll-a, POC, PON, and carbon and nitrogen</u>
- 593 uptake rates were derived from water euphotic column-integrated values averaged from stations.

595 Figure captions

596	Fig. 1. Sampling locations in the Amundsen Sea. Red closed circles represent productivity stations. Sea
597	ice concentration data during the cruise period from Nimbus-7 SMMR and DMSP SSM/I-
598	SSMIS Passive Microwave data provided by National Snow & Ice Data Center.

- Fig. 2. Water column-integrated chlorophyll-a concentration at the productivity stations in the AmundsenSea.
- 601 Fig. 3. Water column-integrated concentrations of POC and PON of small and large phytoplankton.
- 602 Fig. 4. Water column-integrated daily carbon uptake rates of small and large phytoplankton.
- 603 Fig. 5. Water column-integrated daily nitrate uptake rates of small and large phytoplankton.
- Fig. 6. Water column-integrated daily ammonium uptake rates of small and large phytoplankton.
- Fig. 7. Relationship between productivity contributions of small phytoplankton and the total daily carbon
 uptake rates of phytoplankton (large + small). The total daily carbon uptake rates were
 transformed into natural logs for a linear regression.
- 608
- 609

	Chlorophyll-a	РОС	PON	Daily carbon uptake rate	Daily nitrate uptake rate	Daily ammonium uptake rate	Total nitrogen uptake rate
All stations	19.4 ± 26.0	41.1 ± 10.6	41.3 ± 11.5	26.9 ± 29.3	21.5 ± 11.1	38.7 ± 24.9	27.7 ± 14.4
Non-polynya	42.4 ± 37.2	49.5 ± 14.4	50.0 ± 15.1	50.8 ± 42.8	28.2 ± 15.9	52.8 ± 40.5	36.2 ± 23.0
Polynya	7.9 ± 3.5	36.9 ± 4.6	37.0 ± 6.9	14.9 ± 8.4	18.1 ± 6.8	31.6 ± 10.1	23.5 ± 6.0

Table 1. Contributions (%) of small phytoplankton in the Amundsen Sea	<u>Contributions</u>	of chlorophyll-a	<u>, POC, PON,</u>	and carbon

and nitrogen uptake rates were derived from water euphotic column-integrated values averaged from stations.



Fig. 1. Sampling locations in the Amundsen Sea. Red closed circles represent productivity stations. Sea ice concentration data during the cruise period from Nimbus-7 SMMR and DMSP SSM/I-SSMIS Passive Microwave data provided by National Snow & Ice Data Center.



Fig. 2. Water column-integrated chlorophyll-a concentration at the productivity stations in the Amundsen Sea.



Fig. 3. Water column-integrated concentrations of POC and PON of small and large phytoplankton.



Fig. 4. Water column-integrated daily carbon uptake rates of small and large phytoplankton.



Fig. 5. Water column-integrated daily nitrate uptake rates of small and large phytoplankton.



Fig. 6. Water column-integrated daily ammonium uptake rates of small and large phytoplankton.



Fig. 7. Relationship between productivity contributions of small phytoplankton and the total daily carbon uptake rates of phytoplankton (large + small). The total daily carbon uptake rates were transformed into natural logs for a linear regression.