



1 **Soil moisture control on sap-flow response to biophysical factors in a desert-shrub**
2 **species, *Artemisia ordosica***

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16 **Short title: Sap flow in *Artemisia ordosica***

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21 **Author Contribution Statement:**

22 Dr. Tianshan Zha and Dr. Duo Qian contributed equally to the design and implementation of
23 the field experiment, data collection and analysis, and writing the first draft of the manuscript.

24 Dr. Xin Jia gave helpful suggestions concerning the analysis of the field data and contributed
25 to the scientific revision and editing of the manuscript.

26 Prof. Bin Wu contributed to the design of the experiment.

27 Dr.'s Charles P.-A. Bourque and Heli Peltola contributed to the scientific revision and editing
28 of the manuscript.

29 Yujie Bai, Jingyong Ma, Wei Feng, and Yun Tian were involved in the implementation of
30 the experiment and in the revision of the manuscript.

31

32 **Key Message:** This study provides a significant contribution to the understanding of
33 acclimation processes in desert-shrub species to drought-associated stress in dryland
34 ecosystems

35

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44



45 **Abstract:** Current understanding of acclimation processes in desert-shrub species to drought
46 stress in dryland ecosystems is still incomplete. In this study, we measured sap flow in
47 *Artemisia ordosica* and associated environmental variables throughout the growing seasons
48 of 2013-2014 (May-September period of each year) to better understand the environmental
49 controls on the temporal dynamics of sap flow. We found that the occurrence of drought in
50 the dry year of 2013 during the leaf-expansion and leaf-expanded periods caused sap flow
51 per leaf area (J_s) to decline significantly, resulting in a sizable drop in transpiration. Sap flow
52 per leaf area correlated positively with radiation (R_s), air temperature (T), and vapor pressure
53 deficit (VPD), when volumetric soil water content (VWC) was $> 0.11 \text{ m}^3 \text{ m}^{-3}$. Diurnal J_s was
54 generally ahead of R_s by as much as 6 hours. This lag time, however, decreased with
55 increasing VWC. Relative response of J_s to the environmental variables (i.e., R_s , T, and VPD)
56 varied with VWC, J_s being more biologically controlled with low decoupling coefficient and
57 thus being less sensitive to the environmental variables during dry periods. According to this
58 study, soil moisture is shown to control sap-flow (and, therefore, plant-transpiration)
59 response in *Artemisia ordosica* to diurnal variations in biophysical factors. The findings of
60 this study add to the knowledge of acclimation processes in desert-shrub species under
61 drought-associated stress. This knowledge is essential to model desert-shrub-ecosystem
62 functioning under changing climatic conditions.

63 **Keywords:** sap flow; transpiration; cold-desert shrubs; environmental stress; volumetric soil
64 water content

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66



67 1. Introduction

68 Due to the low amount of precipitation and high potential evapotranspiration in desert
69 ecosystems, low soil water availability limit both plant water- and gas-exchange and, as a
70 consequence, limit vegetation productivity (Razzaghi et al., 2011). Therefore, it is important
71 to understand the mechanisms controlling the vegetation-water dynamics under rapidly
72 changing environments (Jacobsen et al., 2007). Grass species have already been replaced by
73 shrub and semi-shrub species in arid and semi-arid areas of northwestern China. This
74 progression is predicted to continue under a changing climate (Asner et al., 2003; Houghton
75 et al., 1999; Pacala et al., 2001). This is mostly because desert shrubs are able to adapt to hot-
76 dry environments by modifying their morphological characteristics, e.g., by (1) minimizing
77 plant-surface area directly exposed to sun and hot air, (2) producing thick epidermal hairs,
78 (3) thickening cuticle, (4) recessing stomata into leaves (Yang and Zhu, 2011), and (5)
79 increasing root-to-shoot ratios (Eberbach and Burrows, 2006; Forner et al., 2014). Also,
80 acclimation of physiological characteristics of plants under water stress, e.g., water potential,
81 osmotic regulation, anti-oxidation, and photosynthetic characteristics, assist the plants to
82 maintain a hydrological balance (Huang et al., 2011a). Changes in stomatal conductance and,
83 thus, transpiration may likewise affect plant water use efficiency (Pacala et al., 2001;
84 Vilagrosa et al., 2003).

85 In recent studies, sap flow in *Tamarix elongate* has been observed to be controlled by
86 solar radiation and air temperature, whereas in *Caragana korshinskii* vapor pressure deficit
87 and solar radiation appear to be more important (Jacobsen et al., 2007; Xia et al., 2008). In
88 *Elaeagnus angustifolia*, transpiration is observed to peak at noon, i.e., just before stomatal
89 closure at mid-day under water-deficit conditions (Liu et al., 2011). In contrast, transpiration
90 in *Hedysarum scoparium* peaks multiple times during the day (Xia et al., 2007). Sap flow has
91 been observed to decrease rapidly when the volumetric soil water content (VWC) is lower
92 than the water loss through evapotranspiration (Buzkova et al., 2015). In general, desert
93 shrubs can close their stomata to reduce transpiration when exposed to dehydration stress
94 around mid-day. However, differences exist among shrub species with respect to their
95 stomatal response to changes in soil and air moisture deficits (Pacala et al., 2001). For some



96 shrubs, sap-flow response to precipitation varies from an immediate decline after a heavy
97 rainfall to no observable change after a small rainfall event (Asner et al., 2003; Zheng and
98 Wang, 2014). Sap flow has been found to increase with increasing rainfall intensity (Jian et
99 al., 2016). Drought-insensitive shrubs have relatively strong stomatal regulation and,
100 therefore, tend to be insensitive to soil water deficits and rainfall unlike their drought-
101 sensitive counterparts (Du et al., 2011).

102 *Artemisia ordosica*, a shallow-rooted desert shrub, is the dominant plant species in the
103 Mu Us desert of northwestern China. The shrubs have an important role in combating
104 desertification and in stabilizing sand dunes (Li et al., 2010). Increases in air temperature and
105 precipitation variability and associated shorter wet periods and longer intervals of periodic
106 drought are expected to ensue with projected climate change (Lioubimtseva and Henebry,
107 2009). During dry periods of the year, sap flow in *Artemisia ordosica* has been observed to
108 be controlled by VWC at about 30-cm depth in the soil (Li et al., 2014). Sap-flow rate is
109 known to be affected by variation in precipitation patterns. Soil water content, in combination
110 with other environmental factors, may have a significant influence on sap-flow rate (Li et al.,
111 2014; Zheng and Wang, 2014). Thus, understanding the controlling mechanisms of sap flow
112 in desert shrubs as a function of variations in biotic and abiotic factors is greatly needed (Gao
113 et al., 2013; Xu et al., 2007).

114 In this study, we measured stem sap flow in *Artemisia ordosica* and associated
115 environmental variables throughout the growing seasons of 2013-2014 (May-September
116 period of each year) to better understand the environmental controls on the temporal
117 dynamics of sap flow. We believe that our findings will provide further understanding needed
118 on acclimation processes in desert-shrub species under dehydration stress.

119

120 **2. Materials and Methods**

121 **2.1 Experimental site**

122 Continuous sap-flow measurements were made at the Yanchi Research Station (37°42' 31"
123 N, 107°13' 47" E, 1530 m above mean sea level), Ningxia, northwestern China. The
124 research station is located between the arid and semi-arid climatic zones along the southern
125 edge of the Mu Us desert. The sandy soil in the upper 10 cm of the soil profile has a bulk



126 density of $1.54 \pm 0.08 \text{ g cm}^{-3}$ (mean \pm standard deviation, $n=16$). Mean annual precipitation
127 in the region is about 287 mm, of which 62% falls between July and September. Mean annual
128 potential evapotranspiration and air temperature are about 2,024 mm and 8.1°C based on
129 meteorological data (1954-2004) from the Yanchi County weather station. Normally, shrub
130 leaf-expansion, leaf-expanded, and leaf-coloration stages begin in April, June, and
131 September, respectively.

132

133 2.2 Measurements of sap flow, leaf area and stomatal conductance

134 The experimental plot ($10 \text{ m} \times 10 \text{ m}$) was located on the western side of Yanchi Research
135 Station in an *Artemisia ordosica*-dominated area. Mean age of the *Artemisia ordosica* was
136 10-years old. Five stems of *Artemisia ordosica* were randomly selected within the plot as
137 replicates for sap-flow measurement. Mean height and sapwood area of sampled shrubs was
138 84 cm and 0.17 cm^2 , respectively. Sampled stems represented the average size of stems in
139 the plot. A heat balance sensor (Flow32-1K, Dynamax Inc., Houston, USA) was installed at
140 about 15 cm above the ground surface on each of the five stems (after Dynamax, 2005). Half-
141 hourly data were recorded by a Campbell CR1000 data logger from May 1 to September 30,
142 2013-2014 (Campbell Scientific, Logan, UT, USA).

143 Leaf area was estimated for each stem every 7-10 days by sampling about 50-70 leaves
144 from five neighbouring shrubs with similar characteristics as the shrubs used for sap flow
145 measurement to avoid damaging them. Leaf area was measured immediately at the station
146 laboratory with a portable leaf-area meter (LI-3000, Li-Cor, Lincoln, NE, USA). Also
147 stomatal conductance (g_s) was measured *in situ* for three to four leaves on each of the sampled
148 shrubs with a LI-6400 portable photosynthesis analyzer (Li-Cor Inc., Lincoln, USA).
149 Stomatal conductance measurements were made every two hours from 7:00 to 19:00 h every
150 ten days from May to September, 2013-2014.

151 The degree of coupling between the ecosystem surface and the atmospheric boundary
152 layer was estimated with the decoupling coefficient (Ω). The decoupling coefficient varies
153 from 0 (i.e., leaf transpiration is mostly controlled by g_s) to 1 (i.e., leaf transpiration is mostly
154 controlled by radiation). The decoupling coefficient was calculated as described by Jarvis
155 and McNaughton (1986):



$$156 \quad \Omega = \frac{\Delta + \Upsilon}{\Delta + \Upsilon \left(1 + \frac{g_a}{g_s} \right)}, \quad (1)$$

157 where Δ is the rate of change of saturation vapor pressure vs. temperature (kPa K^{-1}), γ is the
 158 psychrometric constant (kPa K^{-1}), and g_a is the aerodynamic conductance (m s^{-1} ; Monteith
 159 and Unsworth, 1990):

$$160 \quad g_a = \left(\frac{u}{u_*^2} + 6.2u_*^{-0.67} \right)^{-1}, \quad (2)$$

161 where u is wind speed (m s^{-1}) at 6 m above the ground and u_* is the friction velocity (m s^{-1}).

162

163 **2.3 Environmental measurements**

164 Shortwave radiation (R_s in W m^{-2} ; CMP3, Kipp & Zonen, Netherland), air temperature (T in
 165 $^{\circ}\text{C}$), wind speed (u in m s^{-1} , 034B, Met One Instruments Inc., USA), and relative humidity
 166 (RH in %; HMP155A, Vaisälä, Finland) were measured simultaneously near the sap-flow
 167 measurement plot. Half-hourly data were recorded by a data logger (CR3000 data logger,
 168 Campbell Scientific Inc., USA). VWC at 10- and 30-cm depths were monitored with three
 169 ECH₂O-5TE soil moisture probes (Decagon Devices, USA). In the analysis, we used half-
 170 hourly averages of VWC from the three soil moisture probes. Vapor pressure deficit (VPD
 171 in kPa) was calculated from recorded RH and T .

172

173 **2.4 Data analysis**

174 In the analysis, March-May represented spring, June-August summer, and September-
 175 November autumn. Duration and severity of ‘drought’ were defined based on mean VWC <
 176 $0.08 \text{ m}^3 \text{ m}^{-3}$ measured at 10- and 30-cm depths over the spring-autumn period, from March
 177 through November (Fig. 1a). An extreme dry-soil period was defined as the time period when
 178 VWC < $0.08 \text{ m}^3 \text{ m}^{-3}$ for both 10- and 30-cm depths for at least 10 consecutive days.

179 Data analysis was conducted using hourly mean data from five sensors. Linear and non-
 180 linear regression was used to analyze abiotic control on sap-flow rate. In order to minimize
 181 the effects of different phenophases and rainfall, we used data only from mid-growing season,
 182 non-rainy day, daytime measurements (8:00-20:00), i.e., from June 1 to August 31, with



183 hourly shortwave radiation $> 10 \text{ W m}^{-2}$. Regression slopes were used to identify the
184 sensitivity of sap flow (degree of response) to the environmental variables (see e.g., Zha et
185 al., 2013). All statistical analyses were performed with SPSS v. 17.0 for Windows software
186 (SPSS Inc., USA). Significance level was set at 0.05.

187

188 **3. Results**

189 **3.1 Seasonal variations in environmental factors and sap flow**

190 Range of daily means (24-hour mean) for T, R_s, VPD, and VWC during the 2013 growing
191 season (May-September) were 8.8-24.4°C, 31.1-364.9 W m^{-2} , 0.05-2.3 kPa, and 0.06-0.17
192 $\text{m}^3 \text{m}^{-3}$ (Fig. 1a, b, c); annual means for the same year were 17.7°C, 224.8 W m^{-2} , 1.03 kPa,
193 and 0.08 $\text{m}^3 \text{m}^{-3}$. Corresponding range of daily means for 2014 were 7.1-25.8°C, 31.0-369.9
194 W m^{-2} , 0.08-2.5 kPa, and 0.06-0.16 $\text{m}^3 \text{m}^{-3}$ (Fig. 1a, b, c); annual means were 17.2°C, 234.9
195 W m^{-2} , 1.05 kPa, and 0.09 $\text{m}^3 \text{m}^{-3}$.

196 Total precipitation and number of rainfall events during the 2013 measurement period
197 (257.2 mm and 46 days) was lower than that during 2014 (272.4 mm and 51 days; Fig. 1d).
198 In 2013, more irregular rainfall events occurred than in 2014. The measurement period in
199 2013 had many irregular rainfall events compared to 2014, with 45.2% of rainfall falling in
200 July and 8.8% in August, 2013.

201 Extreme dry-soil periods with $\text{VWC} < 0.08 \text{ m}^3 \text{m}^{-3}$ (at 10- and 30-cm depths) for at least
202 10 consecutive days occurred in May, June, and August of 2013 and in May and June of 2014
203 (Fig. 1d). Both years had dry springs. A nearly one-month long summer drought occurred in
204 August of 2013.

205 Range of daily means of J_s during the growing season was 0.01-4.36 $\text{kg m}^{-2} \text{d}^{-1}$ in 2013
206 and 0.01-2.91 $\text{kg m}^{-2} \text{d}^{-1}$ in 2014 (Fig. 1e); annual means were 0.89 $\text{kg m}^{-2} \text{d}^{-1}$ in 2013 and 1.31
207 $\text{kg m}^{-2} \text{d}^{-1}$ in 2014. Mean J_s over the growing season of 2013 was 32% lower than that of
208 2014. Seasonal fluctuations in J_s corresponded with the seasonal pattern in VWC (Fig. 1d, e).
209 Daily mean J_s decreased or remained nearly constant during dry-soil periods (Fig. 1d, e), with
210 the lowest J_s observed in spring and mid-summer (August) of 2013.

211

212 **3.2 Sap flow response to environmental factors**



213 In summer, J_s increased with increasing VWC (Fig. 2), saturating at VWC $\sim 0.11 \text{ m}^3 \text{ m}^{-3}$ in
214 2013 and $\sim 0.09 \text{ m}^3 \text{ m}^{-3}$ in 2014, then decreasing with VWC when VWC is greater than the
215 thresholds of $0.11 \text{ m}^3 \text{ m}^{-3}$ in 2013 and $0.09 \text{ m}^3 \text{ m}^{-3}$ in 2014. Soil water is shown to modify the
216 response of J_s to environmental factors (Fig. 2). Sap flow J_s increased more rapidly with
217 increases in R_s , T, and VPD under high VWC (i.e., VWC $> 0.11 \text{ m}^3 \text{ m}^{-3}$ in 2013, VWC $>$
218 $0.09 \text{ m}^3 \text{ m}^{-3}$ in 2014) compared with periods with lower VWC (i.e., VWC $< 0.11 \text{ m}^3 \text{ m}^{-3}$ in
219 2013, VWC $< 0.09 \text{ m}^3 \text{ m}^{-3}$ in 2014). Sap flow J_s was more sensitive to R_s , T, and VPD under
220 high VWC (Fig.2), which coincided with a larger regression slope under high VWC
221 conditions.

222 Sensitivity of J_s to environmental variables (in particular, R_s , VPD, T, and VWC) varied
223 depending on the time of a day (Fig. 3). Regression slopes for the relations of J_s - R_s , J_s -T, and
224 J_s -VPD were greater in the morning before 11:00 h, and lower during mid-day and early
225 afternoon (12:00-16:00 h). In contrast, regression slopes of the relation of J_s -VWC were
226 lower in the morning (Fig. 3), increasing thereafter, peaking at $\sim 13:00$ h, and subsequently
227 decreasing in late afternoon. Regression slopes of the response of J_s to R_s , T, and VPD in
228 2014 were greater than those in 2013.

229

230 3.3 Diurnal changes and hysteresis between sap flow and environmental factors

231 Diurnal patterns of J_s were similar in both years (Fig.4), initiating at 7:00 h and increasing
232 thereafter, peaking before noon (12:00 h), and subsequently decreasing thereafter and
233 remaining near zero from 20:00 to 6:00 h. Diurnal changes in stomatal conductance (g_s) were
234 similar to J_s , but peaking about 2 and 1 h earlier than J_s in July and August, respectively (Fig.
235 4).

236 There were pronounced time lags between J_s and R_s over the two years (Fig. 5), J_s
237 peaking earlier than R_s and, thus, earlier than either VPD or T. These time lags differed
238 seasonally. For example, mean time lag between J_s and R_s was 2 h during July, 5 h during
239 May, and 3 h during June, August, and September of 2013. However, the time lags in 2014
240 were generally shorter than those observed in 2013 (Table 1).

241 Use of normalized variables may remove the influence of J_s and R_s from the data. As a
242 result, clockwise hysteresis loops between J_s and R_s during the growing period were observed



243 (Fig.5). As R_s increased in the morning, J_s increased until it peaked at ~10:00 h. Sap-flow
244 rate declined with decreasing R_s during the afternoon. Sap flow J_s was higher in the morning
245 than in the afternoon, forming a clockwise hysteresis loop.

246 Diurnal time lag in the relation of J_s - R_s were influenced by VWC (Fig. 6, 7). For
247 example, J_s peaked about 2 h earlier than R_s on days with low VWC (Fig. 6a), 1 h earlier than
248 R_s on days with moderate VWC (Fig. 6b), and at the same time as R_s on days with high VWC
249 (Fig. 6c). Lag hours between J_s and R_s over the growing season were negatively and linearly
250 related to VWC [Fig.7: Lag (h) = -133.5×VWC+12.24, $R^2=0.41$]. Effect of VWC on time lags
251 between J_s and R_s was smaller in 2014, with evenly distributed rainfall during the growing
252 season, than in 2013, with a pronounced summer drought (Fig. 7).

253

254 4. Discussion and conclusions

255 4.1 Sap flow response to environmental factors

256 Drought tolerance of some plants may be related to lower overall sensitivity of the plants'
257 physiological attributes to environmental stress and/or stomatal regulation (Huang et al.,
258 2011b; Naithani et al., 2012). In this study, large regression slopes of linear relationships
259 between J_s and the environmental variables (R_s , VPD, and T) in the morning indicated that
260 sap flow was more sensitive to variations in R_s , VPD, and T during the less dry and hot period
261 of the day (Fig.3). Stomatal conductances were the largest in the morning (Fig.4), which led
262 to increases in water fluxes to the atmosphere as a result of increased R_s , T, and VPD. When
263 R_s peaked during mid-day (13:00-14:00 h), there was often insufficient soil water to meet the
264 atmospheric demand of water, causing stomatal conductance to be limited by available soil
265 moisture and making J_s more responsive to VWC at noon, but less responsive to R_s and T.

266 Synergistic interactions among environmental factors influencing sap flow is complex.
267 In general, VWC has an influence on physiological processes of plants in water-limited
268 ecosystems (Lei et al., 2010; She et al., 2013). Our findings regarding lower sensitivity in J_s
269 to environmental factors (R_s , T and VPD) during dry periods was in line with a previous
270 study of boreal grasslands (Zha et al., 2010). Our finding that soil water regulates the response
271 of other environmental factors, suggests that VWC is the most important factor controlling
272 sap flow in *Artemisia ordosica*. This is in contrast to other shrub species, where for example



273 it has been found that sap flow in *Haloxylon ammodendron* was mainly controlled by T
274 (Zhang et al., 2003), while sap flow in *Cyclobalanopsis glauca* was controlled by R_s and T,
275 when VWC was not limiting (Huang et al. 2009).

276 Precipitation, being the main source of VWC at our site, affected transpiration directly.
277 In this sense, frequent small rainfall events (< 5 mm) were important to the survival and
278 growth of the desert plants (Zhao and Liu, 2010). Variations in J_s were clearly associated
279 with the intermittent supply of water to the soil during rainfall events, as we found at our site
280 (Fig.1d, e). Reduced J_s during rainy days can be explained by a reduction in incident R_s and
281 water-induced saturation on the leaf surface, which led to a decrease in leaf turgor and
282 stomatal closure. After each rainfall event, J_s increased quickly when soil water was
283 replenished. This finding is related to a positive response in J_s to R_s , T, and VPD under high
284 VWC (Fig. 2). Schwinning and Sala (2004) showed previously for similar research sites that
285 VWC contributed the most to the response in plant transpiration to post-rainfall events. We
286 showed in this study that *Artemisia ordosica* responded in a different way to wet and dry
287 conditions. In the mid-growing season, high J_s in July were related to rainfall-fed VWC,
288 which increased the rate of transpiration. However, dry soil conditions combined with high
289 T and R_s , led to a reduction in J_s in August of 2013 (Fig. 1). In some desert shrubs,
290 groundwater may replenish water lost by transpiration by having deep roots (Yin et al., 2014).
291 *Artemisia ordosica* roots are generally distributed in the upper 60 cm of the soil (Zhao et al.,
292 2010), and as a result the plant usually depends on water directly supplied by precipitation
293 because groundwater levels in drylands can be well below the rooting zone, typically, at
294 depths ≥ 10 m.

295

296 **4.2 Hysteresis between sap flow and environmental factors**

297 Diurnal patterns in J_s corresponded with those of R_s from sunrise until diverging later in the
298 day (Fig.5), suggesting that R_s was a primary controlling factor of diurnal variation in J_s .
299 According to O'Brien et al. (2004), diurnal variation in R_s could cause change in the diurnal
300 variation in the consumption of water. As an initial energy source, R_s can force T and VPD
301 to increase, causing a phase difference in time lags among the relations J_s - R_s , J_s -T, and J_s -
302 VPD.



303 We found a consistent clockwise hysteresis loop between J_s and R_s over a diurnal cycle
304 (Fig. 5), indicating that R_s lagged J_s , and the response of J_s to R_s varied both diurnally and
305 seasonally. A large g_s in the morning promoted higher rates of transpiration (Fig. 4). In dry
306 and hot conditions, stomatal conductance decreased, causing the control of the stomata on J_s
307 to increase relative to changes in environmental factors. Diurnal trends in J_s and g_s occurred
308 together, both peaking earlier than R_s . Stomatal conductance peaked 3-4 h earlier than R_s ,
309 leading to reduction in J_s and an increase in R_s and a clockwise hysteresis loop. Contrary to
310 our findings, counterclockwise hysteresis has been observed to occur between transpiration
311 (J_s) and R_s in tropical and temperate forests (Meinzer et al., 1997; O'Brien et al., 2004; Zeppel
312 et al., 2004). A possible reason for this difference may be due to differences in VWC
313 associated with the different regions. According to Zheng and Wang (2014) favorable water
314 conditions after rainfall could render clockwise hysteresis loops between J_s and R_s under dry
315 conditions to counterclockwise loops. In this study, due to a large incidence of small rainfall
316 events, soil water supply by rainfall pulses could not meet the transpiration demand under
317 high mid-day R_s , resulting in clockwise loops even though rainfall had occurred.

318 In semi-arid regions, low VWC restricts plant transpiration more than VPD. Water
319 vapor deficits tend to restrict transpiration in forest species in wet regions to a greater extent.
320 According to Zheng et al. (2014), high water availability in alpine shrubland meadows may
321 contribute to weakened hysteresis between evapotranspiration and the environmental
322 variables. Our results showed that hysteresis between J_s and R_s decreased as VWC increased
323 (Fig. 6, 7). The result that stomatal conductance increased with VWC (Fig. 8a), along with
324 the harmonization of J_s and g_s , suggests that J_s is less sensitive to stomatal conductance (g_s)
325 in high VWC and more so to R_s . Temporal patterns in J_s became more coherent with those
326 in R_s as VWC increased, leading to a weakened hysteresis between the two variables. This is
327 further supported by a large decoupling coefficient, when VWC is high (Fig. 8b). The larger
328 the decoupling coefficient is, the greater is the influence of R_s on J_s . The effect of VWC on
329 time lag varied between 2013 and 2014.

330

331 4.3. Conclusions



332 Drought during the leaf-expansion and leaf-expanded periods led to a greater decline in J_s ,
333 causing J_s to be lower in 2013 than in 2014. The relative influence of R_s , T, and VPD on J_s
334 in *Artemisia ordosica* was modified by soil water content, indicating J_s 's lower sensitivity to
335 environmental variables (R_s , T and VPD) during dry periods. Sap flow J_s was constrained by
336 soil water deficiency, causing J_s to peak several hours prior to R_s . Diurnal hysteresis between
337 J_s and R_s varied seasonally, because of the control by stomatal conductance under low VWC
338 and R_s under high VWC. According to this study, soil moisture controlled sap-flow response
339 in *Artemisia ordosica*. This species is capable to tolerate and adapt to soil water deficiencies
340 and drought conditions during the growing season. Altogether, our findings add to our
341 understanding of acclimation in desert-shrub species under stress of dehydration. The
342 knowledge gain can assist in modeling desert-shrub-ecosystem functioning under changing
343 climatic conditions.

344

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353

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464 **Table 1** Mean monthly diurnal cycles of sap-flow rate (J_s) response to shortwave radiation (R_s), air
 465 temperature (T), and vapor pressure deficit (VPD), including time lags (h) in J_s as a function of R_s , T, and
 466 VPD.

Pattern	May		June		July		August		September	
	2013	2014	2013	2014	2013	2014	2013	2014	2013	2014
J_s - R_s	5	2	3	0	2	1	3	1	3	2
J_s -T	8	6	7	4	4	4	6	5	6	6
J_s - VPD	8	5	7	4	6	4	6	5	6	5

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470 **Figure captions:**

471

472 **Fig. 1** Seasonal changes in daily (24-hour) mean air temperature (T; a), shortwave radiation (R_s; b), vapor
473 pressure deficit (VPD; c), volumetric soil water content (VWC; d) at 10- and 30-cm depths, daily total
474 precipitation (PPT; d), and sap-flow rate per leaf area (J_s; e) from May to September for both 2013 and
475 2014. Shaded bands in plot (d, e) indicate periods of drought.

476

477 **Fig. 2** Sap-flow rate per leaf area (J_s) in non-rainy, daytime hours during the mid-growing period of June
478 1-August 31 for both 2013 and 2014 as a function of volumetric soil water content (VWC) at a 30-cm
479 depth, shortwave radiation (R_s), air temperature (T), vapor pressure deficit (VPD) under high VWC (>
480 0.11 m³ m⁻³ in 2013 and > 0.09 m³ m⁻³ in 2014) and low VWC (< 0.11 m³ m⁻³ in 2013 and < 0.09 m³ m⁻³
481 in 2014); J_s is given as binned averages according to VWC, R_s, T, and VPD, based on increments of 0.005
482 m³ m⁻³, 100 W m⁻², 1°C, and 0.2 kPa, respectively. Bars indicate standard error.

483

484 **Fig. 3** Regression slopes of linear fits between sap-flow rate per leaf area (J_s) in non-rainy days and
485 shortwave radiation (R_s), vapor pressure deficit (VPD), air temperature (T), and volumetric soil water
486 content (VWC) between 8:00-20:00 h during the mid-growing period of June 1-August 31 for both 2013
487 and 2014.

488

489 **Fig. 4** Mean monthly diurnal changes in sap-flow rate per leaf area (J_s) and stomatal conductance (g_s) in
490 *Artemisia ordosica* during the growing season (May-September) for both 2013 and 2014. Each point is
491 given as the mean at specific times during each month.

492

493 **Fig. 5** Seasonal variation in hysteresis loops between sap-flow rate per leaf area (J_s) and shortwave
494 radiation (R_s) using normalized plots for both 2013 and 2014. The y-axis represents the proportion of
495 maximum J_s (dimensionless) and the x-axis represents the proportion of maximum R_s (dimensionless).
496 The arrow indicates the clockwise direction of response during the day.

497

498 **Fig. 6** Sap-flow rate per leaf area (J_s) and shortwave radiation (R_s) over consecutive 3 days in 2013, i.e.,
499 (a) under low volumetric soil water content (VWC) and high vapor pressure deficit (VPD, 0.063 m³ m⁻³



500 $^3 < \text{VWC} < 0.064 \text{ m}^3 \text{ m}^{-3}$, $\text{VPD}_{\text{mean}} = 2.07 \text{ kPa}$), (b) moderate VWC and VPD ($0.067 \text{ m}^3 \text{ m}^{-3} < \text{VWC} < 0.072 \text{ m}^3$
501 m^{-3} , $\text{VPD}_{\text{mean}} = 1.95 \text{ kPa}$) before rainfall, and (c) high VWC and low VPD ($0.094 < \text{VWC} < 0.098 \text{ m}^3 \text{ m}^{-3}$,
502 $\text{VPD}_{\text{mean}} = 1.20 \text{ kPa}$) after rainfall; VPD_{mean} is the mean value of the three days.

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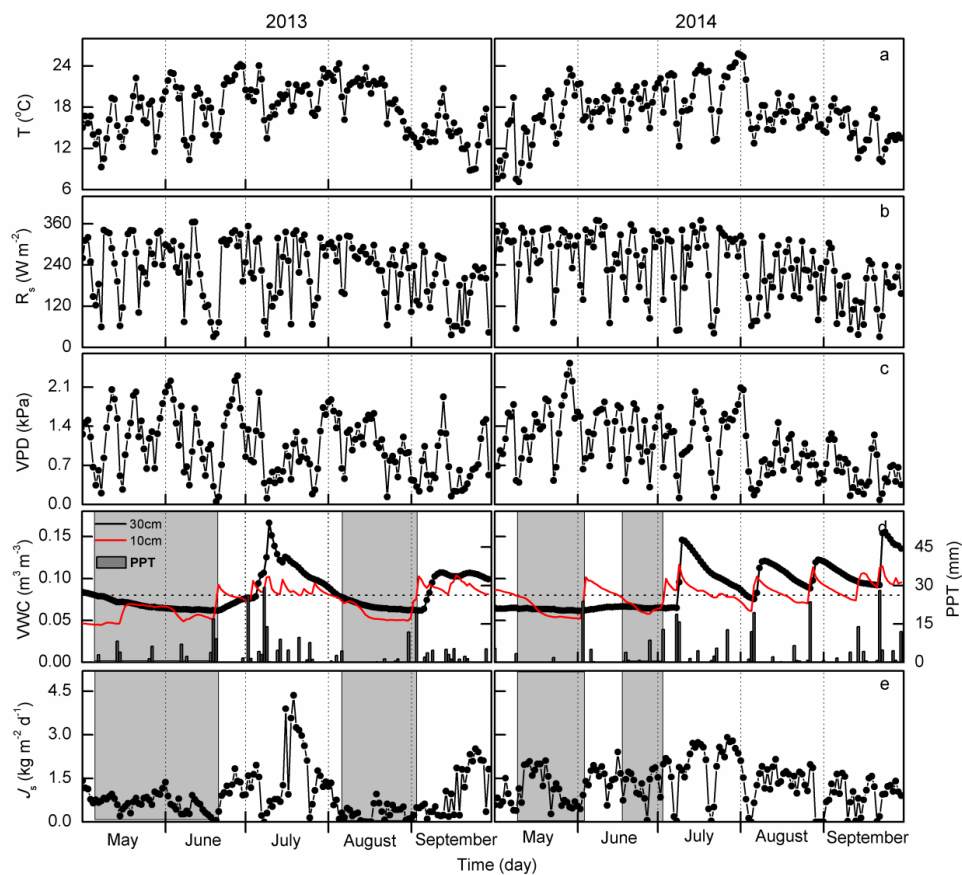
504 **Fig. 7** Time lag between sap-flow rate per leaf area (J_s) and short wave radiation (R_s) in relation to
505 volumetric soil water content (VWC) at a 30-cm depth. Hourly data in non-rainy days during the mid-
506 growing period of June 1-August 31 for both 2013 and 2014. The lag hours were calculated by a cross-
507 correlation analysis using a three-day moving window with a one-day timestep. Rainy days were excluded.
508 The solid line is based on exponential regression.

509

510 **Fig. 8** Relationship between volumetric soil water content (VWC) and (a) stomatal conductance (g_s) in
511 *Artemisia ordosica*, and (b) decoupling coefficient (Ω) for both 2013 and 2014. Hourly values are given
512 as binned averages based on a VWC-increment of $0.005 \text{ m}^3 \text{ m}^{-3}$. Bars indicate standard error.

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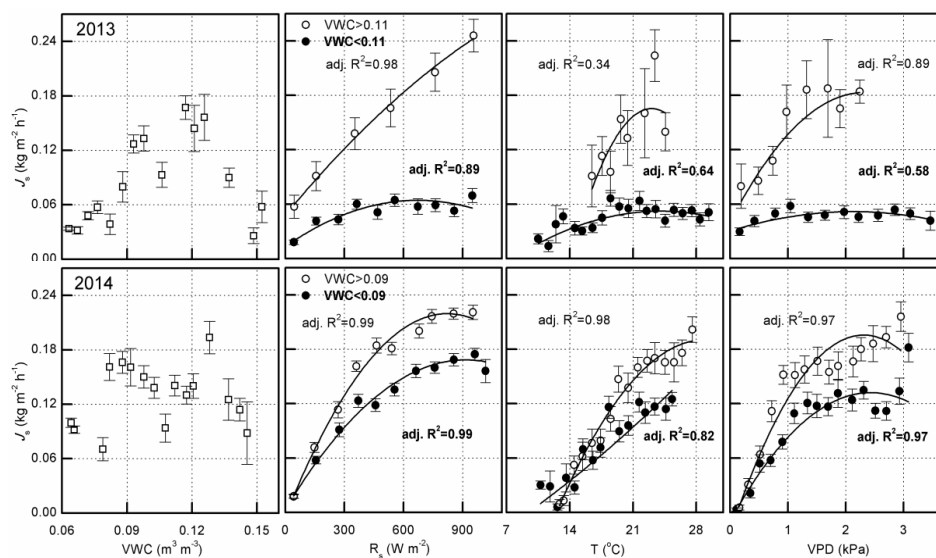
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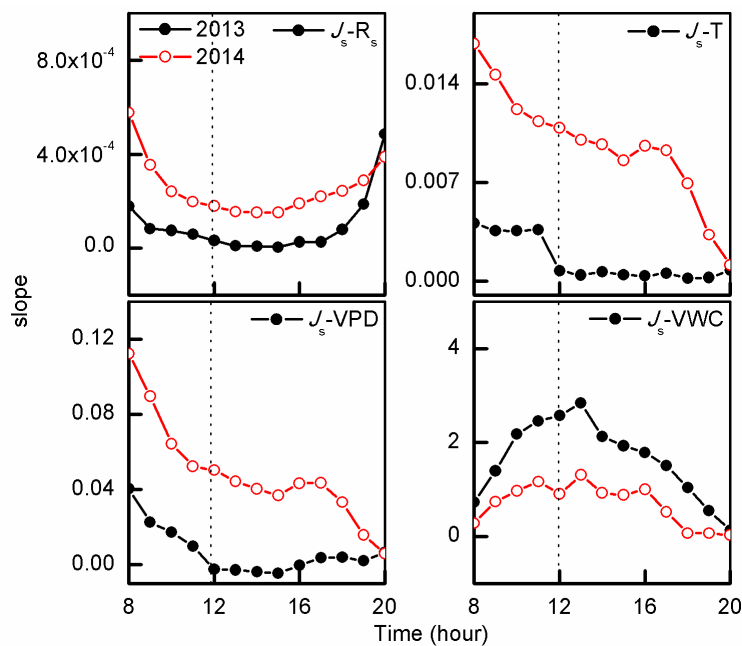
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 526 2013 and > 0.09 m³ m⁻³ in 2014) and low VWC (< 0.11 m³ m⁻³ in 2013 and < 0.09 m³ m⁻³ in 2014); J_s is given
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 528 and 0.2 kPa, respectively. Bars indicate standard error.

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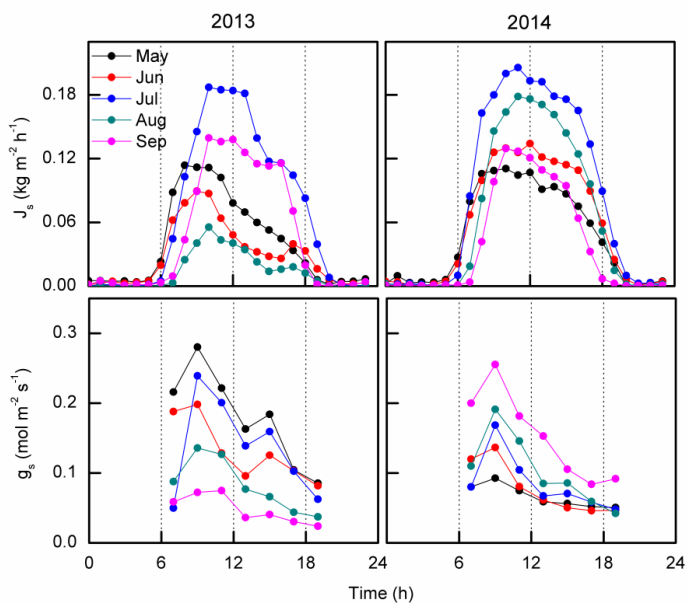
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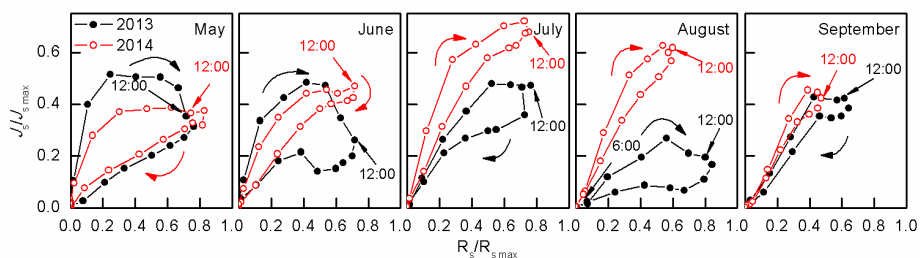
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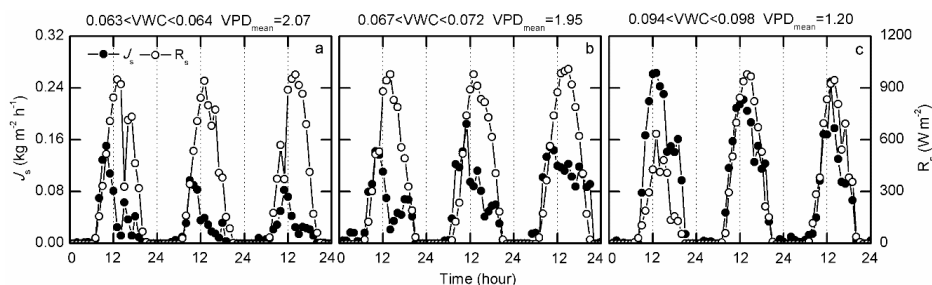
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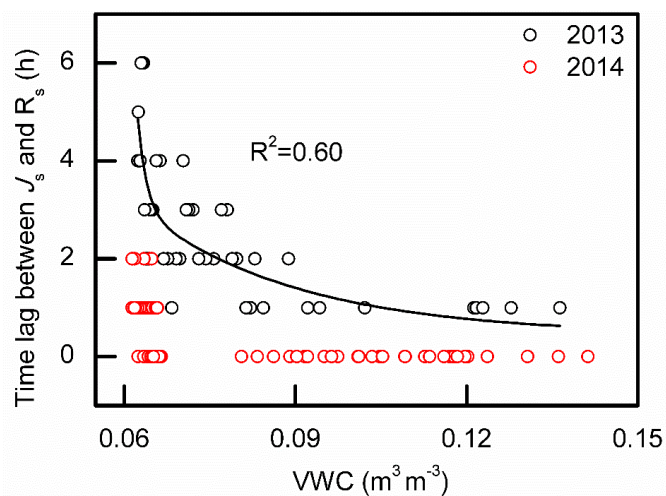
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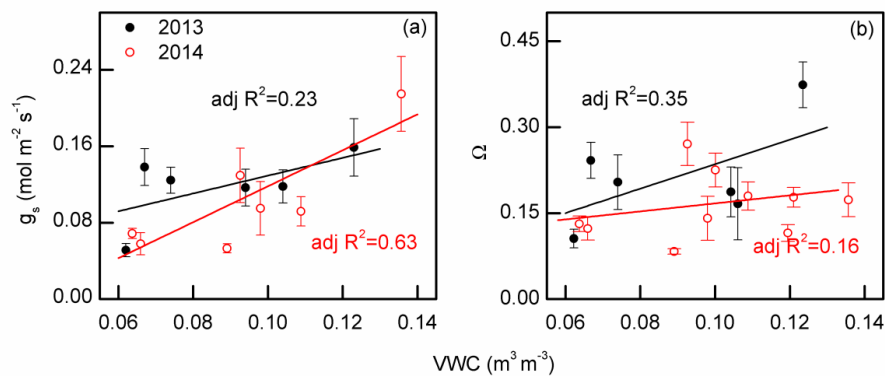
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