Tidal variability of nutrients in a coastal coral reef system influenced by groundwater

Guizhi Wang[1,2], Shuling Wang[1], Zhangyong Wang[1], Wenping Jing[1], Yi Xu[1], Zhouling Zhang[1], Ehui Tan[1], Minhan Dai[1,2]

[1]State Key Laboratory of Marine Environmental Science, Xiamen University, Xiamen, 361102, China
[2]College of Ocean and Earth Sciences, Xiamen University, Xiamen, 361102, China

Correspondence to: Guizhi Wang (gzhwang@xmu.edu.cn)

Abstract. To investigate variations in nitrite, nitrate, phosphate and silicate in a spring-neap tide in a coral reef system influenced by groundwater discharge, we carried out a time-series observation of these nutrients and ^{228}Ra, a tracer of groundwater discharge, in the Luhuitou fringing reef at Sanya Bay in the South China Sea. The maximum ^{228}Ra, 45.3 dpm 100 L^{-1}, appeared at a low tide and the minimum, 14.0 dpm 100 L^{-1}, showed up during a flood tide in the spring tide. The activity of ^{228}Ra was significantly correlated with water depth and salinity in the spring-neap tide, reflecting the tidal-pumping feature of groundwater discharge. Concentrations of all nutrients exhibited strong diurnal variations with a maximum in the amplitude of the diel change for nitrite, nitrate, phosphate and silicate in the spring tide of 0.46 μM, 1.54 μM, 0.12 μM, and 2.68 μM, respectively. Nitrate and phosphate were negatively correlated with water depth during the spring tide, but showed no correlation during the neap tide. Nitrite was positively correlated with water depth in the spring and neap tide due to mixing of nitrite-deplete groundwater and nitrite-rich offshore seawater. They were also significantly correlated with salinity ($R^2>=0.9$ and $P<0.05$) at the ebb flow of the spring tide, negative for nitrate and phosphate and positive for nitrite, indicating the mixing of nitrite-deplete, nitrate and phosphate-rich less saline groundwater and nitrite-rich, nitrate and phosphate-deplete saline offshore seawater. We quantified variations in oxidized nitrogen (NO_x) and phosphate contributed by biological processes based on deviations from mixing lines of these nutrients. During both the spring and neap tide biologically contributed NO_x and phosphate were significantly correlated with regression slopes of 4.60 ($R^2=0.16$) in the spring tide and 13.4 ($R^2=0.75$) in the neap tide, similar to the composition of these nutrients in the water column, 5.43 ($R^2=0.27$) and 14.2 ($R^2=0.76$), respectively. This similarity indicates that the composition of nutrients in the water column of the reef system was closely related with biological processes during both tidal periods, but the biological influence appeared to be less as inferred from the less significant correlations ($R^2=0.16$) during the spring tide when groundwater discharge was more prominent. Thus, the variability of nutrients in the coral reef system was regulated mainly by biological uptake and release in a spring-neap tide and impacted by mixing of tidally-driven groundwater and offshore seawater during spring tide.

1 Introduction

Coral reefs are considered to be one of the most sensitive and stressed ecosystems occupying the coastal zone (Ban et al., 2014). Groundwater input to coral reefs was shown to be globally important and carry a significant amount of terrestrially derived nutrients to the reef systems (D'Elia et al., 1981; Paytan et al., 2006; Houk et al., 2013). Groundwater discharge is usually enriched in N relative to P with an N:P ratio higher than the Redfield ratio, 16:1 (Redfield, 1960), because of more efficient immobilization of P than N in coastal aquifers (Slomp and Van Cappellen, 2004). Such groundwater characterized by a high N:P ratio thus could have significant impacts on coastal reef ecosystems considering that benthic marine plants are much more depleted in P with an N:P ratio of about 30:1 (Atkinson and Smith, 1983). Cuet et al. (2011) have found that the net community production in a coral-dominated fringing reef at La Réunion, France is sustained by net uptake of new nitrogen from groundwater and net uptake of phosphate from the ocean.

Groundwater flux onto coral reefs was found to fluctuate with the tidal cycle (Lewis, 1987; Santos et al., 2010). The contribution of groundwater discharge to the nutrient budget of adjacent marine waters of coral reefs varies greatly from one site to another around the globe and at each site varies from one tidal state to another (Paytan et al., 2006). However, there is no study to reveal variations in the composition of nutrients from spring to neap tide in reef systems influenced by groundwater. Then, questions are posed: a) in coral reef systems influenced by groundwater how do the abundance and composition of nutrients vary from spring to neap tide? b) what contributes to the tidal variation of nutrients in such a system?

To address these questions, this study examined the nutrient variability in a spring-neap tidal cycle in the Luhuitou fringing reef in Sanya Bay, China during a dry season. Our previous study showed that tidally-driven groundwater discharge affected the carbonate system in the Luhuitou fringing reef (Wang et al., 2014). In this reef system, groundwater discharge played a predominant role during the spring tide and biological activities (including photosynthesis/respiration and calcification/dissolution) dominated during the neap tide in regulating diurnal variations of the carbonate parameters. Time-series observations of nutrients carried out at the same time as for the carbonate parameters in this reef system made this study possible. The naturally occurring radioactive radium isotope, ^{228}Ra, was utilized as a tracer of groundwater discharge in this study.

2 Materials and Methods

2.1 Site description

Sanya Bay is a tropical bay situated at the southern tip of Hainan Island, China in the northern South China Sea under the influence of the Southeast Asian monsoon (Fig. 1). Seasonal monsoons dominate Hainan Island with northeast winds in November to March and southwest winds in May to September. Rainfall ranges from 961 to 2439 mm yr^{-1} in 1994-2011 with about 80 % precipitation occurring during May to October (Zhang et al., 2013). The coastal reef time-series station CT is located at the Luhuitou fringing reef in the southeast of Sanya Bay. There was no rain in the two weeks before our

sampling starting on Feb. 2, 2012 and during our 11-day long sampling period based on data from the nearby meteological station in the Hainan Tropical Marine Biology Research Station, Chinese Academy of Science. No surface runoff was present during these periods in this area. Surface salinity in Sanya Bay in our sampling period ranged 33.60-33.89 (Wang et al., 2014). Irregular diurnal tides prevail in Sanya Bay, with a mean tidal range of 0.90 m and the largest of 2.14 m (Zhang, 2001). The Luhuitou fringing reef is a leeward coast with low wave energy in winter (Zhang, 2001). In summer coastal upwelling off eastern Hainan Island mainly induced by the southeast monsoon may extend to this area (Wang et al., 2016). The Holocene deposits of coral debris and biogenic carbonate sands (secondary reef) form the surfacial unconfined aquifer around the fringing reef (Zhao et al., 1983), making groundwater a diffuse source of nutrients for the reef system. Macroalgae cover about 60 %, on average, of the bottom hard substrates in the Luhuitou fringing reef (Titlyanov and Titlyanova, 2013). Living scleractinian corals were observed in the lower intertidal zone and subtidal zone with coverage of 5-40 % (Titlyanov and Titlyanova, 2013; Titlyanov et al., 2014; 2015). Cyanobacteria and Rhodophyta prevailed in the upper intertidal zone, while Rhodophyta and Chlorophyta were the most abundant in the mid and lower intertidal zones (Titlyanov et al., 2014). Rhodophyta dominated the benthic macroalgal community, 54 % in the upper subtidal zone (Titlyanov and Titlyanova, 2013). The number of species in the marine flora has increased by 28 % from 1990 to 2010 with a displacement of slow-growing species likely due to anthropogenic influences and coral bleaching (Titlyanov et al., 2015). The mean coral cover has decreased in the Luhuitou fringing reef from 90 % in the 1960s to 12 % in 2009 (Zhao et al., 2012), likely owing to a combination of regional anthropogenic impacts and climate change (Li et al., 2012).

To the north of the Luhuitou fringing reef, the Sanya River flows into Sanya Bay with an annual average discharge of 5.86 m^3 s^{-1} (Wang et al., 2005). The river is fed mainly by southwest monsoons from May to October. There is no dam in the upstream to regulate the river. During our sampling period the Sanya River plume was confined in the northeast of the bay and the coastal reef station CT was outside the influence of the Sanya River plume (Fig. 1) (Wang et al., 2014). Investigations of nutrients, Chl *a* and phytoplankton in the bay have been conducted seasonally for several years (Dong et al., 2010; Wu et al., 2011; Wu et al., 2012a; 2012b) and demonstrate that the inner bay is influenced by the discharge of the Sanya River with its relatively high nutrient levels, and the central and outer bay are dominated by oceanic exchange with the South China Sea (Wu et al., 2012c). Nutrients carried by submarine groundwater discharge into Sanya Bay account for at least 38 % phosphate, 90 % inorganic nitrogen, and 83 % silicate of the nutrients into the bay in our sampling period (Wang et al., submitted).

2.2 Sampling and measurements

The setup of the sampling platform at the time-series station CT is provided in detail in Wang et al. (2014). Briefly, water was collected using a submersible pump and depth and salinity were measured with a conductivity-temperature-depth system (Citadel, RDI Co., USA) attached on a buoy. Discrete nutrient and radium samples were taken every 3 hours during February 6-13, 2012, except on February 7-8 when the maximum tidal range of 1.4 m occurred (Wang et al., 2014) and the

samples were collected every 2 hours. A mapping cruise was conducted in Sanya Bay during February 2-3, 2012 (Fig. 1) to evaluate the influence of the Sanya River and to constrain the end-member of the offshore water. Nutrient samples for nitrate, nitrite, phosphate and silicate were collected in Sanya Bay at surface and bottom depths using 5 L Niskin bottles. Temperature and salinity were measured using a multi-parameter sonde YSI 6600. The salinity was reported using the Practical Salinity Scale.

Nutrient samples were filtered with 0.45 μm cellulose acetate membranes and poisoned with 1-2 ‰ chloroform. One filtrate was preserved at 4 °C for dissolved silicate determination, and one was frozen and kept at -20 °C for nitrate, nitrite, and phosphate measurements. In the laboratory, nutrients were measured with an AA3 Auto-Analyzer (Bran-Luebbe, GmbH) following the same methods in Han et al. (2012). The analytical precision was better than 1 % for nitrate and nitrite, 2 % for phosphate, and 2.8 % for silicate. The detection limit was 0.04 μM for nitrate and nitrite, 0.08 μM for phosphate, and 0.16 μM for silicate. Blanks were directly set up as baselines during the measurements and subtracted. Radium samples were passed through a 1 μm cartridge filter before through a MnO_2-impregnated acrylic fiber (Mn-fiber) column to extract dissolved radium (Rama and Moore, 1996). The Mn-fibers were leached with 1 M solutions of hydroxylamine hydrochloride and HCl to release ^{226}Ra and ^{228}Ra, which were then co-precipitated with $BaSO_4$ and measured in a germanium gamma detector (GCW4022, Canberra) (Moore, 1984) with an error less than 7 %.

2.3 Linear regression and contour plotting

To gain insight into factors affecting nutrients from spring to neap tide, linear regressions were conducted between water depth, salinity, and ^{228}Ra activity, between water depth, salinity, and nutrients concentration, and between biologically contributed nutrients during the spring and neap tide. A linear curve-fitting, $y=ax+b$, was applied using least-square minimization algorithm to find the coefficients (a, b) of the independent variable that gave the best fit between the linear equation and the data (e.g., Press et al., 1986). A significance level of 0.05 was taken. The data was fit using SigmaPlot (Systat Software, San Jose California USA, www.systatsoftware.com). In plotting contours in Sanya Bay, Surfer 11 was utilized with kriging interpolation due to its good linear unbiased prediction of the intermediate values in spatial analysis (Papritz and Stein, 2002).

3 Results

3.1 Time-series observations of nutrients and radium at the coastal coral reef station

Time-series observations of salinity, ^{226}Ra, and water depth at Station CT were reported in Wang et al. (2014), which demonstrated that the water depth at Station CT varied from 0.7 to 2.1 m and the salinity ranged from 33.43 to 33.67 during February 6-13, 2012 (Table S1). The greatest tidal range occurred on February 7, 2012 (Wang et al., 2014), the 16th of the

lunar month. To separate neap tide from spring tide days, the daily variance of water depth and salinity were plotted (Fig. 2). The daily variance of a variable was calculated using Microsoft Excel (2007) as

$$\sigma^2 = \frac{n\sum x^2 - \left(\sum x\right)^2}{n(n-1)} \qquad (1)$$

where x is the average of the variable in a day and n is the number of samples of the variable in that day. A sharp decrease in the variance of salinity occurred on February 10, 2012 and the variance remained low (<0.001) afterwards. Thus, two distinctive groups stood out, with one group in the period of February 6-9, 2012 having greater variance of water depth and salinity and the other in the period of February 10-13, 2012 having less variance. Therefore, we took February 6-9, 2012 as the spring tide period and February 10-13, 2012 as the neap tide period in this work.

The concentration of nutrients varied with different patterns from spring to neap tide (Fig. 3). Nitrite varied from 0.11 to 0.71 μM during the spring tide and from 0.12 to 0.74 μM in the neap tide with the maximum diel variation of 0.46 μM present during the spring tide (Fig. 3a). The diurnal variation was 0.24-0.46 μM during the spring tide and 0.34-0.45 μM in the neap tide. Daily peaks of nitrite usually appeared at high tides from the spring to neap tide. The concentration was positively correlated with water depth (P<0.05) during both the spring and neap tide, but the correlation was less significant during the neap tide (Fig. 4a). Nitrate and phosphate, however, showed an opposite pattern. During the spring tide, nitrate and phosphate were negatively correlated with water depth (P<0.05)(Fig. 4b,c). They reached their peak concentrations of 1.91 μM and 0.22 μM, respectively in the late afternoon and their minima of 0.37 μM and 0.10 μM, respectively at night on February 7, 2012 (Fig. 3b,c). The diurnal variation fell in the range of 0.44-1.54 μM for nitrate and 0.04-0.12 μM for phosphate. During the neap tide, the concentrations of nitrate and phosphate varied from 0.27 to 1.32 μM for nitrate and 0.084 to 0.18 μM for phosphate with less diurnal variation in the range of 0.35-0.52 μM for nitrate and 0.04-0.05 μM for phosphate. The correlation with water depth was not significant for both nutrients (P>0.15). Nitrate is the dominant species (>50 %) of oxidized nitrogen (NO_x) during the spring-neap tidal period except at 2 O'clock on February 12, 2012 when the concentrations of nitrite and nitrate were almost equal. The NO_x:P ratio varied from 4.78 to 12.9 in the spring-neap tide (Fig. 3c). Silicate showed a trend different from either nitrite or nitrate and phosphate (Fig. 3d). It was not significantly correlated with water depth during either spring or neap tide (P>0.2). The concentration of silicate, in general, decreased from spring to neap tide. During the spring tide, the concentration of silicate fell in the range of 4.57-7.25 μM. The daily peak concentration of silicate appeared almost at the daily lowest salinity. The diurnal variation in silicate was 1.91-2.68 μM. During the neap tide, however, silicate ranged from 2.89 to 5.59 μM and showed less diurnal variability, 1.44-2.09 μM.

The diurnal variation in the activity of [228]Ra at Station CT was 16.5-27.4 dpm 100 L^{-1} (i.e., 2.75-4.56 Bq m^{-3}) during the spring tide, the maximum of which appeared on February 7, and 5.31-10.6 dpm 100 L^{-1} around the neap tide (Fig. 3e). The maximum [228]Ra, 45.3 dpm 100 L^{-1}, appeared at the lowest tide on February 8 during the spring tide and the minimum, 14.0 dpm 100 L^{-1}, showed up during the flood tide of the spring tide on February 7. The activity of [228]Ra was significantly correlated with water depth in the spring-neap tidal period (P=0.002)(Fig. 5a). This pattern reflected the variation in the

groundwater discharge induced by tidal pumping in this coral reef system (Wang et al., 2014), which is also observed in other coastal regions (Burnett and Dulaiova, 2003; Santos et al., 2010).

3.2 Distributions of nutrients in Sanya Bay

In Sanya Bay the highest concentration of nutrients appeared near the Sanya River estuary and the concentration, in general,
5   decreased from the northeast coast, where the influence of the Sanya River plume is apparent in winter (Wang et al., 2014), to the south and west, where the South China Sea water intrudes (Fig. 6). At stations far offshore (Stations J4-5 and W3-4), the concentrations of nitrite, nitrate and phosphate were all below the detection limit and the concentration of silicate was about 4.00 μM. At other stations, the concentration of all the nutrients remained low, but was nonetheless detectable. For example, the maximum concentration of only 0.43 μM for nitrite, 0.70 μM for nitrate, 0.18 μM for phosphate and 7.92 μM
10   for silicate were recorded at Station P1, the station closest to the Sanya River estuary. The small islands in Sanya Bay did not show apparent influence on the nutrients in the bay since nutrients were below their detection limits or remained low around these islands (Fig. 6). The water depth at these mapping stations was no less than 5 m and the concentration of nutrients at the bottom depth differed little from that at the surface at most of these offshore stations (Table 1). This vertical distribution confirms that the water in Sanya Bay is relatively homogenous in February (Wang et al., 2014). The NO_x:P ratio was less
15   than 7 in Sanya Bay, except at Stations P2 and L6 where the NO_x:P ratio was around 9.

4 Discussion

4.1 What affects tidal variations in nutrients at the reef station CT?

The time-series observation of salinity at Station CT suggests that more freshwater input into the reef system occurred during the ebb flow of the spring tide as inferred from lower salinity than during that of the neap tide (Wang et al., 2014). The
20   distribution of salinity in Sanya Bay demonstrated that the surface salinity was slightly lower in the northeast off the Sanya River mouth than in southern Sanya Bay (Fig. 1a). At Stations P1 and P2 the surface salinity was less than 33.70, while at stations around Station CT, i.e., Stations L1, L2, L3 and P3, the surface salinity was greater than 33.80 (Table 1). The verticle profiles of salinity at these stations also showed similar difference with salinity below 33.80 throughout the water column at river-influenced stations (Fig. 1b). This indicates that the Sanya River plume affected the northeast of the bay with
25   little impact on Station CT and the only source of freshwater at Station CT in February would be groundwater discharge (Wang et al., 2014) since in the two weeks before our sampling and during our sampling period there were no rainfall and consequent surface runoff in this area. The coincidence of the daily minimum salinity with the highest activity of ^{228}Ra during the ebb flow of the spring tide (Fig. 3e) and the significant correlation between the activity of ^{228}Ra and salinity during the spring-neap tidal period (P<0.0001)(Fig. 5b) confirms that the tidally-driven groundwater discharge occurred at

the coral reef station CT. Greater groundwater discharge appeared during the ebb flow in the spring tide than in the neap tide as indicated by the higher activity of ^{228}Ra, bringing more groundwater into the reef system.

Under the influence of tidally-driven groundwater discharge, variations in nitrite, nitrate, phosphate and silicate during the spring tide followed a tidal pattern. Inferred from the significant correlation between nutrients and water depth during the spring tide (Fig. 4), the groundwater discharge was characterized by higher nitrate and phosphate and lower nitrite than the offshore seawater. Since nitrate dominated NO_x during the time-series observation (Fig. 3), groundwater discharge was characterized by higher NO_x and phosphate than the offshore seawater. Because groundwater discharge was greater at low tides than at high tides due to its tidal pumping feature, higher NO_x and phosphate would appear at low tides. From Fig. 4, the daily maximum concentration of NO_x, phosphate, and silicate appeared in the day time at relatively low tides, while the minimum showed up mostly at night at high tides, indicating the mixing of tidally-driven groundwater and offshore seawater. During the neap tide, however, NO_x and phosphate showed less diurnal variations. The daily maximum concentration of NO_x and phosphate appeared around the mid-night, when a flood tide appeared. This pattern of daily maximum at a flood tide at night reflected dominance of biological processes, because at a flood tide there were less groundwater-associated nutrients and nutrients were released the most at night by biological processes. This pattern is consistent with the time-series observation of dissolved oxygen at this site (Wang et al., 2014). The daily minimum showed up for NO_x and phosphate in the afternoon or between mid-night and dawn at high tides, reflecting the dominance of nutrient-deplete offshore seawater. Adsorption/desorption from particles might be a factor influencing the phosphate concentration, as proposed for estuaries (e.g., Froelich et al., 1982; van der Zee et al., 2007). At the reef station the salinity was close to the seawater (>33) and the water was clear (the total suspended matter was low, about 15 mg L^{-1}), which makes adsorption/desorption negligible. The clear water, as well as low wave energy in the reef in winter (Zhang, 2001), also limits the possibility of sediment re-suspension as a source of radium and nutrients.

Under the controls of tidally-driven groundwater discharge and biological processes, the composition of nutrients in the reef system also differed from the spring tide to the neap tide. During the spring tide when groundwater discharge played a predominant role on regulating the concentration of nutrients in the reef system, the concentration of NO_x was positively correlated with the concentration of phosphate, with a regression slope of 5.43 and R^2 of 0.27 (Fig. 7a). The concentration of silicate was not significantly correlated with the concentration of NO_x (Fig. 7b). During the neap tide when groundwater discharge was less prominent, the correlation between the concentrations of NO_x and phosphate was more significant, with a regression slope of 14.2 and R^2 of 0.76. The NO_x:P ratio was closer to the Redfield ratio than during the spring tide. The concentration of silicate showed significant correlation with the concentration of NO_x in the water column, with a regression slope of 1.24 and R^2 of 0.58. Diatoms dominate the phytoplankton community in Sanya Bay (Zhou et al., 2009). The elemental ratio of Si:N is 0.80±0.35 for nanoplankton and 1.20±0.37 for netplankton (Brzezinski, 1985). The similarity of the composition of silicate and NO_x in the water column to the elemental ratio of diatoms implies a biological control. Unfortunately, no information is available on particular reef primary producers and sponges that may take up/release silicate

7

in this reef system to further the discussion. The activity of ^{228}Ra, however, was not significantly correlated with the NO$_x$:P ratio in the water column from spring to neap tide ($P>0.05$)(Fig. 5c), indicating that the composition of nutrients in the water column was not predominantly controlled by groundwater discharge. Therefore, we propose that biological processes predominantly controlled the composition of nutrients in the reef system, but the impact was less due to groundwater discharge.

4.2 The generation and consumption of NO$_x$ and phosphate at the reef station CT

N and P are the general limiting nutrients for the abundance of phytoplankton in coastal ecosystems (Jickells et al., 1998). To quantify the contribution of biological processes to the variations in the NO$_x$ and phosphate at Station CT, a closer look was taken at the behaviors of nitrite, nitrate and phosphate with salinity during the falling and rising phases on Feb. 7, the day with the greatest tidal range in the spring tide period. Fig. 8 shows that these nutrients behaved differently during the two phases. During the ebb flow, with a fast falling speed as indicated by the sharp slope of water depth (Fig. 3), nitrite, nitrate and phosphate behaved conservatively, i.e., their concentrations were significantly correlated with salinity ($P<0.05$). Nitrite was positively correlated with salinity ($R^2=0.94$), while nitrate and phosphate were negatively correlated with salinity ($R^2=0.91$ and 0.90, respectively) (Fig. 8). These conservative behaviors indicated mixing between the groundwater discharge and the offshore seawater. During the flood tide, with a relatively slow speed as indicated by a less sharp slope of water depth (Fig. 3), however, nitrite showed an apparent removal signal relative to the conservative mixing line while additions of nitrate and phosphate showed up. This consumption of nitrite and generation of nitrate and phosphate were due to biological processes in this period. Based on the conservative mixing lines shown in Fig. 8, we could estimate nitrite, nitrate and phosphate owing to mixing of the offshore seawater and groundwater discharge using the salinity measured at Station CT (S_{CT}), designated as NO$_{2mix}$, NO$_{3mix}$ and P$_{mix}$.

$$NO_{2mix} = 1.3696 \times S_{CT} - 45.7520 \tag{2},$$

$$NO_{3mix} = -1.7797 \times S_{CT} + 60.5024 \tag{3},$$

$$P_{mix} = -0.3565 \times S_{CT} + 12.1176 \tag{4}.$$

Two assumptions were made before setting up these equations: (a) there was no other water mass into the reef system besides offshore seawater and groundwater, and (b) mixing of offshore seawater and groundwater from spring to neap tide followed the relation derived from data on the day with the greatest tidal range. The differences between the measured concentrations of nutrients and the nutrient concentrations resulting from mixing represented nutrients contributed by biological processes, designated as ΔNO_{2bio}, ΔNO_{3bio} and ΔP_{bio},

$$\Delta NO_{2bio} = NO_{2CT} - NO_{2mix} \tag{5},$$

$$\Delta NO_{3bio} = NO_{3CT} - NO_{3mix} \qquad\qquad (6),$$

$$\Delta P_{bio} = P_{CT} - P_{mix} \qquad\qquad (7).$$

where the subscripts 'CT' represents the measured value at Station CT. The oxidized nitrogen contributed by biological processes, ΔNO_{xbio}, is the sum of ΔNO_{2bio} and ΔNO_{3bio}. Positive values represent regeneration and release of nutrients in the water column and negative values reflect uptake of nutrients by marine flora (including phytoplankton and benthic flora in this system). Benthic release due to remineralization of organic matter contributes to the positive values.

The nutrients contributed by biological processes showed the greatest diurnal variation in nitrate and phosphate on February 7, 2012, which is in the spring tide, while the maximum of biologically contributed nitrite appeared on February 12, 2012, which is in the neap tide (Fig. 9). Nitrite contributed by biological processes ranged from -0.15 to 0.39 μM during the spring tide and from -0.20 to 0.40 μM during the neap tide (Fig. 9a). From 6 pm on February 8 to 6 pm on February 11, 2012, biologically contributed nitrite was positive throughout the period, indicating production of nitrite. For nitrate it was produced throughout the period from 4 am on February 8 to the midnight on February 11, 2012. During the spring tide biologically contributed nitrate varied from -0.24 to 1.25 μM and during the neap tide it fell in the range of -0.38 to 0.70 μM. Net NO_x production occurred from 6 pm on February 8 to 8 am on February 12, 2012 and ΔNO_{xbio} was negative afterwards on February 12-13, 2012, indicating net consumption (Fig. 9b). The biological contribution of phosphate had greater diurnal variations during the spring tide than during the neap tide (Fig. 9c). The greatest diel variation during the spring tide in ΔP_{bio} appeared on February 7, 2012 when ΔP_{bio} varied from -0.027 to 0.088 μM, while during the neap tide the greatest variation occurred on February 10, 2012 when ΔP_{bio} ranged from 0.009 to 0.056 μM. Net phosphate consumption occurred throughout the period of February 12-13, 2012.

The relationship between ΔNO_{xbio} and ΔP_{bio} during the spring tide differed from that during the neap tide. Note that NO_x is not equivalent to dissolved inorganic nitrogen (DIN). Ammonium data is not available in this study for us to discuss the relationship between DIN and phosphate. However, tidal variations in ammonium are expected at this site and may have been related to the change in the relationship between ΔNO_{xbio} and ΔP_{bio}. During the spring tide there was significant correlation between ΔNO_{xbio} and ΔP_{bio}, with a regression slope of 4.60 and R^2 of 0.16 (Fig. 10). During the neap tide, however, the correlation was much more significant with a regression slope of 13.4 and R^2 of 0.75. The regression slope of the regression between biologically contributed NO_{xbio} and phosphate was similar to that of the significant regression between NO_{xbio} and phosphate in the water column, which was 5.43 during the spring tide and 14.18 during the neap tide. This similarity indicates that the composition of nutrients in the water column was closely related with biological processes during both tidal periods, but the biological effect appeared to be less during the spring tide as inferred from the less significant correlations. The net release of nutrients during the neap tide with a very Redfield-like ratio suggests that the net nutrient fluxes in this system were likely to be dominated by the uptake and remineralization of plankton/oceanic organic particles by benthic filter feeders as observed in other reefs (e.g., Ayukai, 1995; Ribes et al., 2005; Southwell et al., 2008; Genin et al., 2009; Monismith et al., 2010). The net uptake of nitrate and phosphate was mainly made by reef primary

producers. Thus, the composition of nutrients in the water column seemed to be directly related with biological contributions from the spring to neap tide. The biological influence was less during the spring tide most likely due to groundwater discharge. This confirms our proposal that biological processes predominantly controlled the composition of nutrients in the reef system, but the impact was less due to groundwater discharge.

5   Successive uptake rates of NO_x were approximated by the depth-integration of the biologically contributed NO_x divided by the sampling time interval from the spring to neap tide. The uptake rate ranged from -9.04 to 19.1 mmol m^{-2} d^{-1}, which compares well with the sum of nitrate and nitrite fluxes over Ningaloo Reef, a fringing reef in Australia, -24 to 15 mmol m^{-2} d^{-1} (Wyatt et al., 2012). It is significantly correlated with the concentration of NO_x in the water column (Fig. 11), with a slope of 14.5 and R^2 of 0.94 (P<0.0001), indicating the mass-transfer limitation of NO_x uptake. The slope (in m d^{-1}) falls in

10   the range of the typical uptake rate coefficient for dissolved inorganic nitrogen reported in Falter et al. (2004).

4.3 Seasonal and regional extrapolations

This study was carried out in winter. Seasonal variations are present in the river discharge as inferred from precipitation (Wang et al., 2005) and there might be increase in the groundwater discharge and associated nutrient fluxes in summer as in other coastal systems (e.g., Lewis, 1987; Costa et al., 2006; Kelly and Moran, 2002; Wang et al., 2015). However, the

15   relative changes in the groundwater discharge and associated nutrient fluxes would be much smaller than those of the river. The tidally-driven feature of the groundwater discharge in this reef system might make our conclusions applicable to other seasons. But it is likely that what we observed in a dry season might be different from what would happen in a wet season due to the involvement of other forces, e.g., upwelling in summer (Wu et al., 2012a; Wang et al., 2016), which merits further studies.

20   In relatively oligotrophic coastal systems with coral reefs, such groundwater-associated nutrient fluxes may sustain the reef community production (Cuet et al., 2011), result in increases in diversity and occurrence of algae and sponge where relatively low salinity is present (Houk and Starmer, 2010), or induce the proliferation of diatom and cyanobacteria (Blanco et al., 2011). In addition, tidally-driven groundwater into nearshore ecosystems was found to be negatively correlated with seagrass habitat condition (Houk et al., 2013). Nutrients loads via groundwater discharge may affect the community structure

25   to move towards macroalgal blooms via bottom-up control (Lapointe, 1997) and likely play a role in the displacement of slow-growing benthic flora with fast-growing species observed in Sanya Bay in the last two decades (Titlyanov et al., 2015). Future changes in these fluxes, likely caused by climate change and human activities, might make the situation worse and need to be monitored in reef protection programs and be considered in assessing the environmental health of coral reef systems, especially in regions with expected higher inputs of anthropogenic nutrients into the groundwater.

4 Conclusions

The variability of nutrients in a spring-neap tidal cycle in a coral reef system in winter was revealed for the first time under the synergistic control of tidally-driven groundwater discharge and biological processes. The activity of ^{228}Ra was significantly correlated with water depth and salinity, indicating tidally-driven groundwater discharge at this site. Nitrate and phosphate were negatively correlated with salinity at the ebb flow of the spring tide, indicating that groundwater discharge was enriched in nitrate and phosphate. Nitrate, phosphate and silicate in the water column showed greater diurnal variations during the spring tide than during the neap tide, while the diel change in the concentration of nitrite demonstrated no consistent pattern. The nutrient composition in the water column seemed to differ between the spring tide and neap tide, but was similar to their biological uptake/release in either tidal period for oxidized nitrogen (NO_x) and phosphate. This similarity indicates that variations in nutrients in the water column in the reef system were mainly regulated by biological processes. However, correlations between NO_x and phosphate in the water column and between biologically contributed NO_x and phosphate were less significant during the spring tide when groundwater discharge was more prominent. The concentration of silicate in the water column was significantly correlated with that of NO_x during the neap tide, but they were not significantly correlated during the spring tide. This indicates that the composition of nutrients in the water column was also affected by tidally-driven groundwater discharge, especially during the spring tide. Therefore, biological processes predominantly controlled the composition of nutrients in the reef system, but the impact was less due to groundwater discharge.

The stoichiometric relationship of NO_x and phosphate from the spring to neap tide in this reef system is important in understanding how biologically processes predominantly affected these nutrients variations under the influence of tidally-driven groundwater discharge. The composition of silicate and NO_x during the neap tide when groundwater discharge was less was comparable to the elemental ratio of diatoms. The release/consumption ratio of NO_x:P by biological processes followed a Redfield-like ratio during the neap tide, but about one third as much during the spring tide. Whether this change in the biological release/uptake ratio of NO_x:P is associated with a change in the community structure needs further study.

Supplement Time-series data are provided in Table S1.

Author contribution Guizhi Wang and Minhan Dai wrote the main text of the manuscript. Guizhi Wang, Shuling Wang, Zhangyong Wang, Wenping Jing, Yi Xu, and Zhouling Wang collected samples in the field and measured the parameters. Guizhi Wang analyzed the data and did the calculations. Ehui Tan drew some of the figures.

*Competing interests* The authors declare that they have no conflict of interest.

References

Atkinson, M. J. and Smith, S. V.: C-N-P ratios of benthic marine plants, Limnol. Oceanogr., 28(3), 568-574, 1983.

Ayukai, T.: Retention of phytoplankton and planktonic microbes on coral-reefs within the great-barrier-reef, australia, Coral Reefs, 14(3), 141-147, 1995.

Ban, S. S., Graham, N. A., and Connolly, S. R.: Evidence for multiple stressor interactions and effects on coral reefs, Global Change Biology, 20(3), 681-697, 2014.

Blanco, A. C., Watanabe, A., Nadaoka, K., Motooka, S., Herrera, E. C., and Yamamoto, T.: Estimation of nearshore groundwater discharge and its potential effects on a fringing coral reef, Mar. Pollution Bull., 62(4), 770-785, 2011.

Brzezinski, M. A.: The Si-C-N ratio of marine diatoms - interspecific variability and the effect of some environmental variables, J. Phycology, 21(3), 347-357, 1985.

Burnett, W. C. and Dulaiova, H.: Estimating the dynamics of groundwater input into the coastal zone via continuous radon-222 measurements, J. Environ. Radioactivity, 69(1), 21-35, 2003.

Costa Jr., O. S., Attrill, M. J., and Nimmo, M.: Seasonal and spatial controls on the delivery of excess nutrients to nearshore and offshore coral reefs of Brazil, J. Mar. Sys., 60, 63-74, 2006.

Cuet, P., Atkinson, M. J., Blanchot, J., Casareto, B. E., Cordier, E., Falter, J., Frouin, P., Fujimura, H., and others: CNP budgets of a coral-dominated fringing reef at La Réunion, France: coupling of oceanic phosphate and groundwater nitrate, Coral Reefs, 30(S1), 45-55, 2011.

D'Elia, C. F., Webb, K. L., and Porter, J. W.: Nitrate-rich groundwater inputs to Discovery Bay, Jamaica: a significant source of N to local coral reefs, Bull. Mar. Sci., 31(4), 903-910, 1981.

Dong, J.-D., Zhang, Y.-Y., Wang, Y. S., Wu, M.-L., Zhang, S., and Cai, C.-H.: Chemometry use in the evaluation of the sanya bay water quality, Brazilian J. Oceanogr., 58(4), 339-352, 2010.

Falter, J. L., Atkinson, M. J., and Merrifield, M. A.: Mass-transfer limitation of nutrient uptake by a wave-dominated reef flat community, Limnol. Oceanogr., 49(5), 1820-1831, 2004.

Froelich, P. N, Bender, M. L, and Luedtke, N A.: The marine phosphorus cycle, Am. J. Sci., 282, 474-511, 1982.

Genin, A., Monismith, S. G., Reidenbach, M. A., Yahel, G., and Koseff, J. R.: Intense benthic grazing of phytoplankton in a coral reef, Limnol. Oceanogr., 54(3), 938-951, 2009.

Han, A. Q., Dai, M. H., Kao, S. J., Gan, J. P., Li, Q., Wang, L. F., Zhai, W. D., and Wang L.: Nutrient dynamics and biological consumption in a large continental shelf system under the influence of both a river plume and coastal upwelling, Limnol. Oceanogr., 57(2), 486-502, 2012.

Houk, P., and Starmer, J.: Constraints on the diversity and distribution of coral-reef assemblages in the volcanic Northern Mariana Islands, Coral Reefs, 29, 59-70, 2010.

Houk, P., Golbuu, Y., Gorong, B., Gorong, T., and Fillmed, C.: Watershed discharge patterns, secondary consumer abundances, and seagrass habitat condition in Yap, Micronesia, Mar. Pollution Bull., 71, 209-215, 2013.

Jickells, T. D.: Nutrient biogeochemistry of the coastal zone, Science, 281(5374), 217-222, 1998.

Kelly, R. P., and Moran, S. B.: Seasonal changes in groundwater input to a well-mixed estuary estimated using radium isotopes and implications for coastal nutrient budgets, Limnol. Oceanogr., 47, 1796–1807, 2002.

Lapointe, B. E.: Nutrient thresholds for bottom-up control of macroalgal blooms on coral reefs in Jamaica and southeast Florida, Limnol. Oceanogr., 42(5), 1119-1131, 1997.

Lewis, J. B.: Measurements of groundwater seepage flux onto a coral reef: spatial and temporal variations, Limnol. Oceanogr., 32(5), 1165-1169, 1987.

Li, X., Liu, S., Huang, H., Huang, L., Jing, Z., and Zhang C.: Coral bleaching caused by an abnormal water temperature rise at Luhuitou fringing reef, Sanya Bay, China, Aquatic Ecosystem Health & Management, 15(2), 227-233, 2012.

Monismith, S. G., Davis, K. A., Shellenbarger, G. G., Hench, J. L., Nidzieko, N. J., Santoro, A. E., Reidenbach, M. A., Rosman, J. H., and others: Flow effects on benthic grazing on phytoplankton by a Caribbean reef, Limnol. Oceanogr., 55(5), 1881-1892, 2010.

Moore, W. S.: Radium isotope measurements using germanium detectors, Nucl. Instru. Methods in Physics Res., 223(2), 407-411, 1984.

Papritz, A. and Stein, A.: Spatial prediction by linear kriging, in: Stein, A., Van der Meer, F., and Gorte, B. (eds) Spatial Statistics for Remote Sensing, Remote Sensing and Digital Image Processing, Vol. 1, Springer, Dordrecht, 1999. DOI: 10.1007/0-306-47647-9_6.

Paytan, A., Shellenbarger, G. G., Street, J. H., Gonneea, M. E., Davis, K., Young, M. B., and Moore, W. S.: Submarine groundwater discharge: an important source of new inorganic nitrogen to coral reef ecosystems, Limnol. Oceanogr., 51(1), 343-348, 2006.

Press, W. H., Flannery, B. P., Teukolsky, S. A., and Vetterling, W. T.: Numerical Recipes, Cambridge, Cambridge University Press, 1986.

Rama, and Moore, W. S.: Using the radium quartet for evaluating groundwater input and water exchange in salt marshes, Geochim. Cosmochim. Acta., 60(23), 4645-4652, 1996.

Redfield, A. C.: The biological control of chemical factors in the environment, Science progress, 11, 150-170, 1960.

Ribes, M., Coma, R., Atkinson, M. J., and Kinzie, R. A.: Sponges and ascidians control removal of particulate organic nitrogen from coral reef water, Limnol. Oceanogr., 50(5), 1480-1489, 2005.

Santos, I. R., Erler, D., Tait, D., and Eyre, B. D.: Breathing of a coral cay: Tracing tidally driven seawater recirculation in permeable coral reef sediments, J. Geophys. Res., 115, C12010, 2010.

Slomp, C. P., and Van Cappellen, P.: Nutrient inputs to the coastal ocean through submarine groundwater discharge: controls and potential impact, J. Hydrology, 295, 64-86, 2004.

Southwell, M. W., Weisz, J. B., Martens, C. S., and Lindquist, N.: In situ fluxes of dissolved inorganic nitrogen from the sponge community on Conch Reef, Key Largo, Florida, Limnol. Oceanogr., 53(3), 986-996, 2008.

Titlyanov, E. A., and Titlyanova, T. V.: Changes in the species composition of benthic macroalgal communities of the upper subtidal zone on a coral reef in Sanya Bay (Hainan Island, China) during 2009–2012, Russian J. Mar. Biology, 39(6), 413-419, 2013.

Titlyanov, E. A., Titlyanova, T. V., Li, X., Hansen, G. I., and Huang, H.: Seasonal changes in the intertidal algal communities of Sanya Bay (Hainan Island, China), J. Mar. Biological Association UK, 94(05), 879-893, 2014.

Titlyanov, E. A., Titlyanova, T. V., Belous, O. S., and Kalita, T. L.: Inventory change (1990s-2010s) in the marine flora of Sanya Bay (Hainan Island, China), J. Mar. Biological Association UK, 95(3), 461-470, 2015.

van der Zee, C., Roevros, N., Chou, L.: Phosphorus speciation, transformation and retention in the Scheldt estuary (Belgium/The Netherlands) from the freshwater tidal limits to the North Sea, Mar. Chem., 106, 76-91, 2007.

Wang, Y., Jing, Z., and Qi, Y.: Coastal upwelling off eastern Hainan Island observed in the summer of 2013. Chinese J. Tropical Oceanography, 35, 40-49, 2016.

Wang, G., Jing, W., Wang, S., Xu, Y., Wang, Z., Zhang, Z., Li, Q., and Dai, M.: Coastal acidification induced by tidal-driven submarine groundwater discharge in a coastal coral reef system, Environ. Sci. Technol., 48, 13069-13075, 2014.

Wang, G., Wang, S., Wang, Z.: Significance of submarine groundwater discharge in nutrients budget in tropical Sanya Bay, China, submited to J. Mar. Systems.

Wang, G., Wang, Z., Zhai, W., Moore, W. S., Li, Q., Yan, X., Qi, D., and Jiang, Y.: Net subterranean estuarine export fluxes of dissolved inorganic C, N, P, Si, and total alkalinity into the Jiulong River estuary, China, Geochim. Cosmochim. Acta, 149, 103-114, 2015.

Wang, H., Dong, J., Wang, Y., Chen, G., and Zhang, Y.: Variations of nutrient contents and their transportation estimate at Sanya Bay, J. Tropical Oceanogr., 25, 90-95, 2005.

Wu, M.-L., Zhang, Y.-Y., Dong, J.-D., Wang, Y.-S., and Cai, C.-H.: Identification of coastal water quality by self-organizing map in Sanya Bay, South China Sea, Aquatic Ecosystem Health & Management, 14(3), 291-297, 2011.

Wu, M.-L., Zhang, Y.-Y., Dong, J.-D., Cai, C.-H., Wang, Y.-S., Long, L.-J., and Zhang, S.: Monsoon-driven dynamics of environmental factors and phytoplankton in tropical Sanya Bay, South China sea, Oceanological and Hydrobiological Studies, 41(1), 57-66, 2012a.

Wu, M.-L., Zhang, Y.-Y., Long, L.-J., Zhang, S., Wang, Y.-S., Ling, J., and Dong, J.-D.: Identification of coastal water quality, including heavy metals, in the South China Sea, Polish J. Environ. Studies, 21(5), 1445-1552, 2012b.

Wu, M.-L., Ling, J., Long, L.-J., Zhang, S., Zhang, Y.-Y., Wang, Y.-S., and Dong, J.-D.: Influence of human activity and monsoon dynamics on spatial and temporal hydrochemistry in tropical coastal waters (Sanya Bay, South China Sea), Chem. Ecol., 28(4), 375-390, 2012c.

Wyatt, A. S. J., Falter, J. L., Lowe, R. J., Humphries, S., and Waite, A. M.: Oceanographic forcing of nutrient uptake and release over a fringing coral reef, Limnol. Oceanogr., 57(2), 401-419, 2012.

Zhang, J., Wang, D. R., Jennerjahn, T., and Dsikowitzky, L.: Land-sea interactions at the east coast of Hainan Island, South China Sea: a synthesis, Cont. Shelf Res., 57, 132-142, 2013.

Zhang, Q.: On biogeomorphology of Luhuitou fringing reef of Sanya city, Hainan Island, China, Chin. Sci Bull., 46, 97-101, 2001.

Zhao, M., Yu, K., Zhang, Q., Shi, Q., and Price, G. J.: Long-term decline of a fringing coral reef in the northern South China Sea, J. Coastal Res., 28(5), 1088-1099, 2012.

Zhao, X., Zhang, J., and Li, G.: Development of the Holocene coral reefs along the southern coast of Hainan Island, Scientia Geologica Sinica, 2, 150-160, 1983.

Zhou, W., Li, T., Cai, C., Huang, L., Wang, H., Xu, J., Dong, J., and Zhang, S.: Spatial and temporal dynamics of phytoplankton and bacterioplankton biomass in Sanya Bay, northern South China Sea, J. Environ. Sci., 21(5), 595-603, 2009.
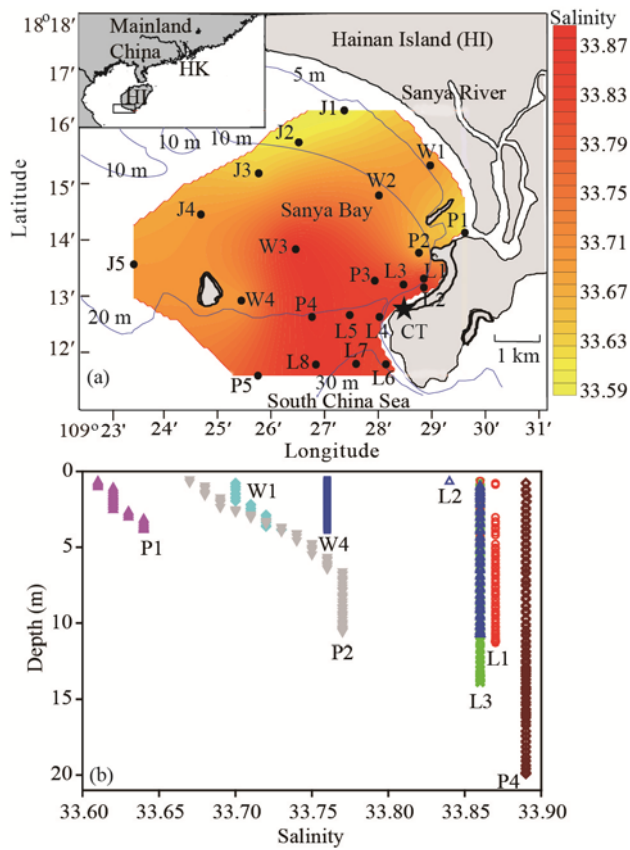
Figure 1: Study area, sampling stations and salinity distribution (a) and vertical profiles of salinity (b) in February 2012 in Sanya Bay, Hainan Island (HI) in the South China Sea. HK represents Hong Kong. CT is the coastal reef time-series station.
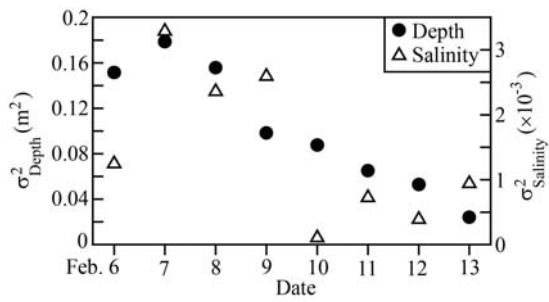
5 **Figure 2: Daily variance of water depth (σ^2_{Depth}) and salinity ($\sigma^2_{Salinity}$) at the coastal reef station CT during February 6-13, 2012.**
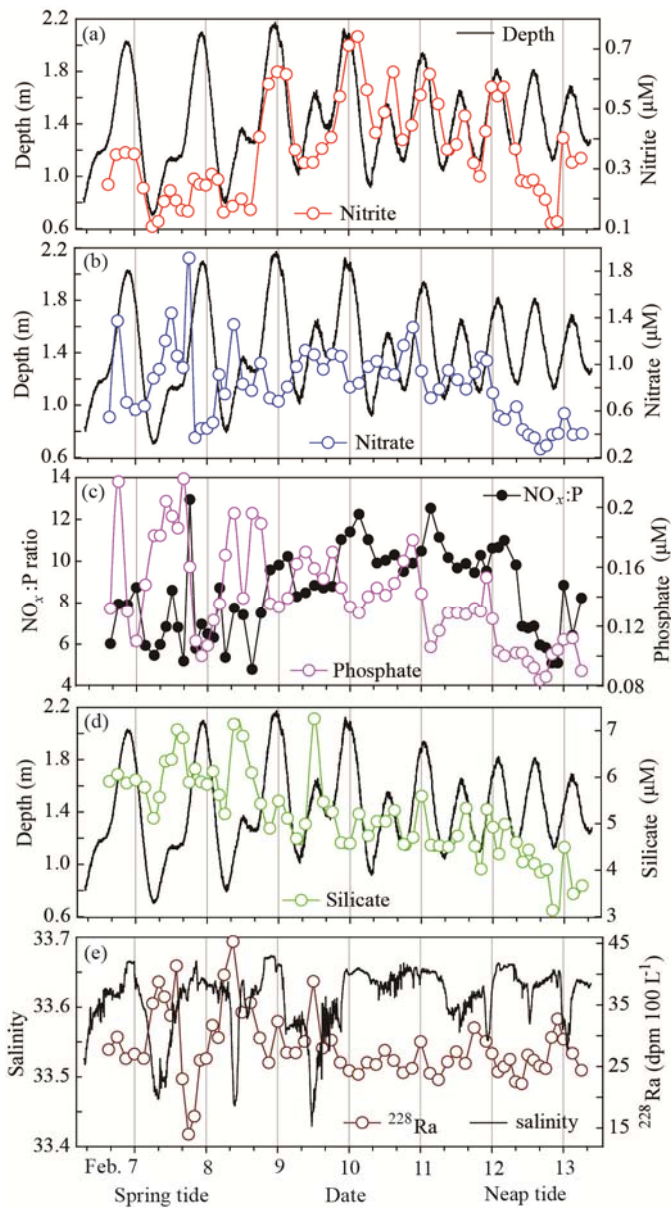
Figure 3: Time-series observations of nutrients and ^{228}Ra at Station CT in the Luhuitou reef of Sanya Bay, China during February 6-13, 2012. (a) Nitrite; (b) nitrate; (c) phosphate and NO$_x$:P ratio; (d) silicate; and (e) ^{228}Ra. Lines connecting the symbols are to show trends. Water depth and salinity were reported in Wang et al. (2014).
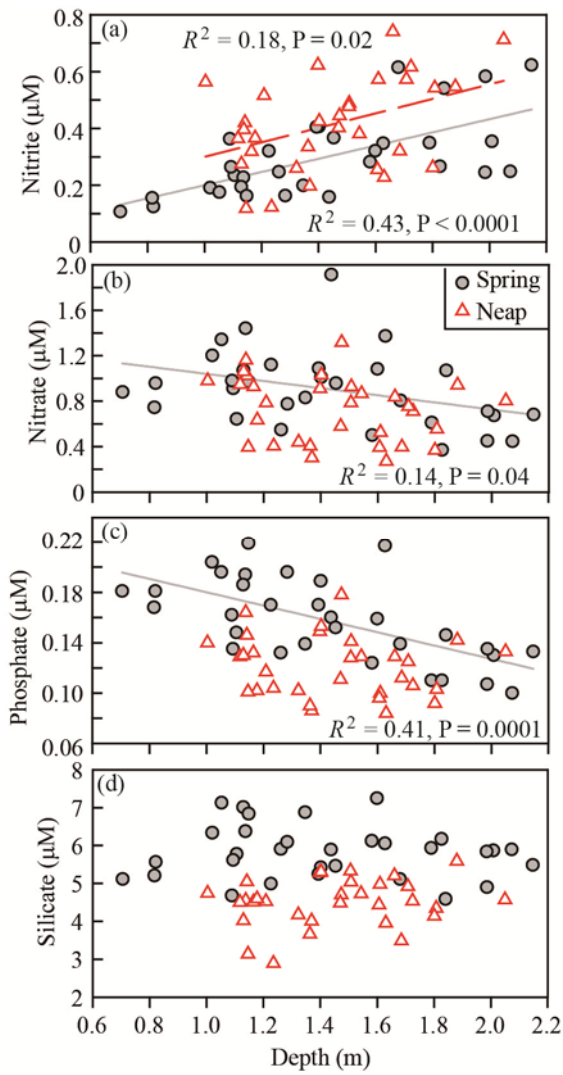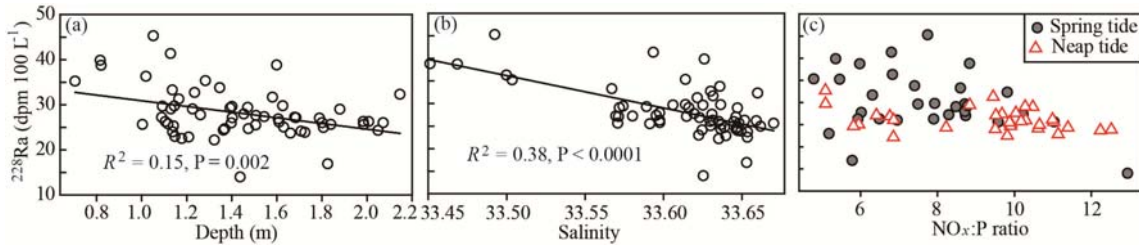
Figure 4: Concentrations of nutrients in the water column against water depth during the spring tide and neap tide at Station CT in the Luhuitou reef during February 6-13, 2012. (a) nitrite; (b) nitrate; (c) phosphate; and (d) silicate.

5   **Figure 5: The activity of [228]Ra against (a) water depth, (b) salinity, and (c) the NO$_x$:P ratio in the water column at Station CT during February 6-13, 2012.**
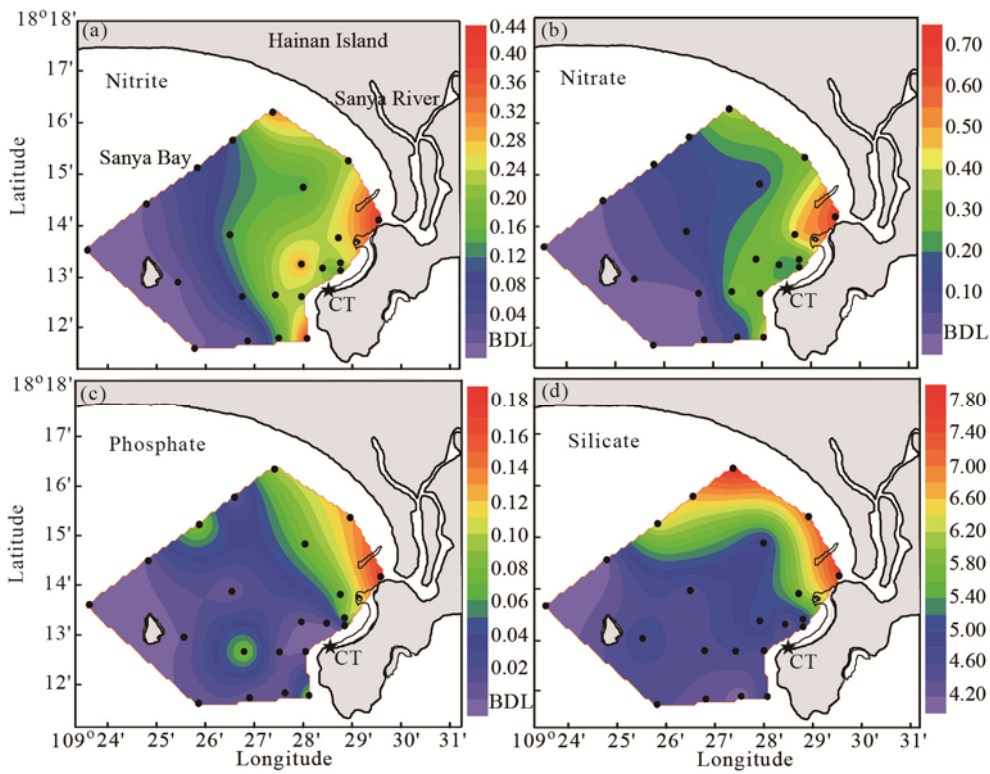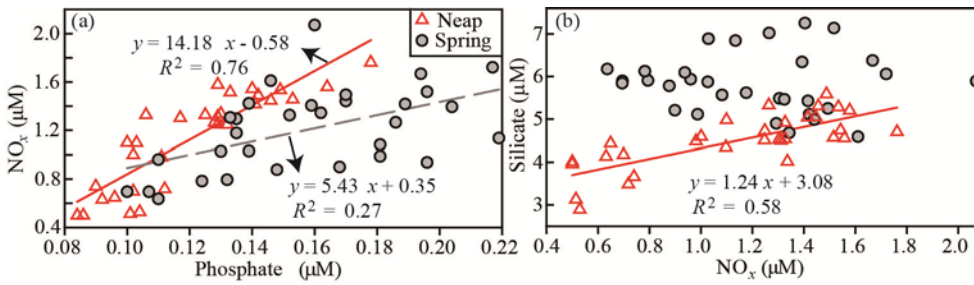
Figure 6: Surface distributions of nutrients in Sanya Bay in February 2012. (a) Nitrite; (b) nitrate; (c) phosphate; and (d) silicate. The units are in μM. BDL is below the detection limit, which is 0.04 μM for nitrate and nitrite and 0.08 μM for phosphate.

Figure 7: Concentrations of (a) NO$_x$ against phosphate and (b) silicate against NO$_x$ in the water column during the spring tide and neap tide at Station CT during February 6-13, 2012.

19
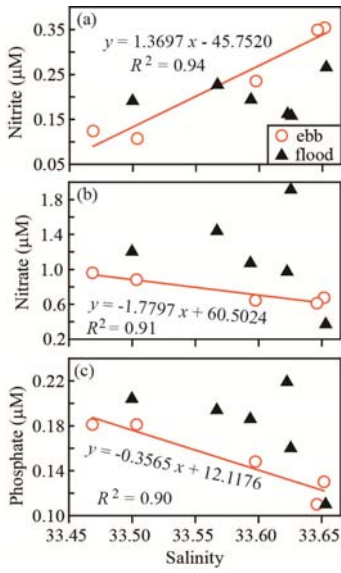
Figure 8: Behaviours of nutrients with salinity during the ebb flow and flood tide of the spring tide at Station CT. (a) nitrite; (b) nitrate; and (c) phosphate.
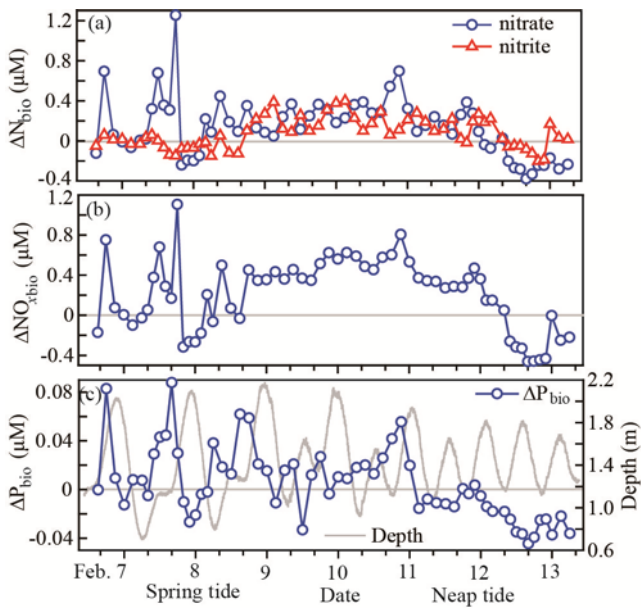
5 **Figure 9: Variations of nutrients contributed by biological processes in a spring-neap tide during February 6-13, 2012 at the coastal reef station CT. (a) nitrite and nitrate; (b) NO_x; and (c) phosphate (P). Water depth was reported in Wang et al. (2014).**
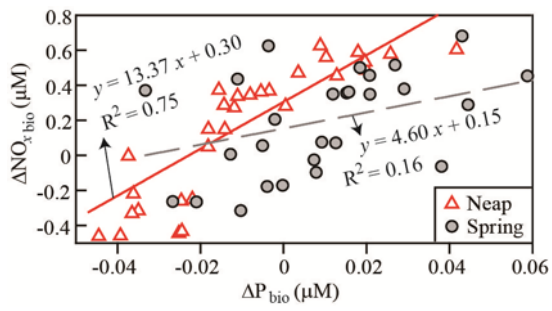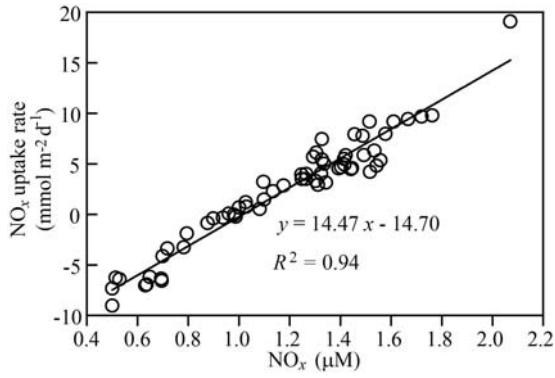
Figure 10: Relationship between biologically contributed NO$_x$ and phosphate during the spring tide and neap tide at Station CT in the Luhuitou fringing reef in February 6-13, 2012.

5   **Figure 11: Uptake rate of NO$_x$ against the concentration of NO$_x$ in the water column at reef Station CT in a spring-neap tide during February 6-13, 2012.**

Table 1. Sampling stations and data collected in Sanya Bay in February 2012.

Station	Latitude (°N)	Longitude (°E)	Bottom Depth (m)	Sample Depth (m)	Temperature (°C)	Salinity	NO_2^- (µM)	NO_3^- (µM)	PO_4^{3-} (µM)	SiO_3^{2-} (µM)
J1	18.2718	109.4565	8	0.5	22.80	33.60	0.328	0.410	0.104	7.916
				6.5	22.74	33.60	0.298	0.343	0.098	7.485
J2	18.2623	109.4423	9	0.5	22.66	33.62	0.103	0.149	BDL	6.708
				8.0	22.64	33.63	0.124	0.162	BDL	6.531
J3	18.2531	109.4298	12	0.5	22.70	33.64	0.073	0.104	0.090	6.472
				11.0	22.69	33.65	0.104	0.067	0.108	6.318
J4	18.2409	109.4118	11	0.5	22.81	33.70	BDL	BDL	BDL	4.069
				12.4	22.81	33.70	BDL	BDL	BDL	4.095
J5	18.2261	109.3909	15	0.5	22.90	33.70	BDL	BDL	BDL	4.058
				14.0	22.88	33.74	BDL	BDL	BDL	4.126
W4	18.2154	109.4244	17	0.5	22.90	33.70	BDL	BDL	BDL	4.768
				17.5	22.73	33.75	BDL	BDL	BDL	4.760
W3	18.2306	109.4413	16	0.5	22.97	33.89	0.136	0.112	BDL	4.476
				16.0	23.40	33.60	0.063	0.098	BDL	5.188
W2	18.2466	109.4672	12	0.5	22.93	33.72	0.147	0.158	0.081	4.724
				9.5	22.76	33.73	0.075	0.127	BDL	5.179
W1	18.2555	109.4832	5	0.5	23.12	33.70	0.228	0.299	0.131	7.136
				3.0	22.92	33.73	0.228	0.234	0.102	6.317
P3	18.2213	109.4660	16	0.5	22.75	33.84	0.300	0.309	BDL	5.172
				16.0	22.87	33.76	0.132	0.144	BDL	4.655
P2	18.2296	109.4797	11	0.5	23.01	33.67	0.262	0.496	0.082	6.035
				11.0	22.90	33.77	0.206	0.204	BDL	4.569
P1	18.2355	109.4940	5	0.5	22.98	33.62	0.426	0.699	0.178	7.726
				2.8	22.97	33.64	0.350	0.525	0.157	7.671
P4	18.2105	109.4464	12	0.5	22.71	33.89	0.108	0.002	0.081	4.519
				19.0	22.67	33.89	0.200	0.013	0.130	4.935
P5	18.1931	109.4296	26	0.5	22.69	33.81	BDL	BDL	BDL	4.428
				26.0	22.74	33.87	0.054	0.076	BDL	4.522
L8	18.1964	109.4476	25	0.5	22.78	33.88	BDL	BDL	BDL	4.282
				25.5	22.75	33.88	0.191	0.005	0.082	4.528
L7	18.1966	109.4601	32	0.5	22.83	33.86	0.171	0.092	BDL	4.093
				30.7	22.78	33.87	0.077	0.081	BDL	4.400
L6	18.1965	109.4694	23	0.5	22.79	33.82	0.420	0.516	0.097	4.859
				27.0	22.77	33.87	0.405	0.431	0.112	4.839
L5	18.2111	109.4582	21	0.5	22.74	33.85	0.231	0.326	BDL	4.643
				18.0	22.79	33.86	0.355	0.392	0.097	4.480
L4	18.2105	109.4674	20	0.5	22.76	33.85	0.219	0.248	BDL	4.484
				21.0	22.77	33.87	0.285	0.309	BDL	4.645
L3	18.2201	109.4749	12	0.5	22.79	33.84	0.194	0.193	BDL	4.315
				12.8	22.79	33.86	0.202	0.183	BDL	4.444
L2	18.2193	109.4812	11	0.5	22.81	33.85	0.192	0.309	BDL	5.006
				11.0	22.80	33.86	0.195	0.218	BDL	4.639
L1	18.2219	109.4812	11	0.5	22.76	33.84	0.244	0.253	0.101	4.887
				11.0	22.81	33.87	0.353	0.235	0.107	5.252

Note: BDL is below the detection limit.