



1 **Simultaneous shifts in stoichiometric and fatty acid composition**
2 **of *Emiliania huxleyi* in response to environmental changes**

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23 **Abstract**

24 Climate-driven changes in environmental conditions have significant and complex
25 effects on marine ecosystems. Variability in phytoplankton elements and biochemicals
26 can be important for global ocean biogeochemistry and ecological functions, while
27 there is currently limited understanding on how elemental stoichiometry and
28 biochemicals respond to the changing environments in key coccolithophore species
29 such as *Emiliania huxleyi*. We investigated responses of stoichiometric C:N:P ratios,
30 PIC:POC contents and ratios, and fatty acid (FA) composition in a strain of *E. huxleyi*
31 under three temperatures (12, 18 and 24 °C), three N:P supply ratios (10:1, 24:1 and
32 63:1 mol mol⁻¹) and two *p*CO₂ levels (560 and 2400 µatm). Overall, C:N:P biomass
33 ratios showed the most pronounced response to N:P supply ratios, with low N:C and
34 N:P biomass ratios in low N-media, and low P:C and high N:P biomass ratios in low
35 P-media. PIC:POC ratios and polyunsaturated FA proportions strongly responded to
36 temperature and *p*CO₂, both being lower under high *p*CO₂ and higher with warming.
37 We observed synergistic interactions between warming and nutrient deficiency (and
38 high *p*CO₂) on PIC and POC cellular contents in most cases, indicating the enhanced
39 effect of warming on *E. huxleyi* calcification and photosynthesis under nutrient
40 deficiency (and high *p*CO₂). Our results suggest differential sensitivity of elements
41 and FAs to the changes in temperature, nutrient availability and *p*CO₂ in *E. huxleyi*,
42 which is to some extent unique compared with non-calcifying algal classes. Thus,
43 simultaneous changes of elements and FAs should be considered when predicting
44 future roles of *E. huxleyi* in the biotic-mediated connection between biogeochemical



45 cycles, ecological functions and climate change.

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47 **Key words:** Coccolithophores; elements; biochemicals; warming; nutrients; CO₂

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67 **1 Introduction**

68 Climate change and intensive anthropogenic pressures have pronounced and
69 diverse effects on marine ecosystems. Physical and chemical properties in marine
70 ecosystems are changing simultaneously such as the concurrent shifts in temperature,
71 CO₂ and oxygen concentrations, and nutrient availability (Boyd et al., 2015). These
72 changes have altered trophic interactions in both bottom-up and top-down directions
73 and thus result in changes in community structure and ecosystem functions (Doney et
74 al., 2012). Phytoplankton are the base of marine food webs and major drivers of ocean
75 biogeochemical cycling, and thus quantifying their responses to changing oceanic
76 conditions is a major challenge in studies of food web structure and ocean
77 biogeochemistry.

78 Coccolithophores are a key phytoplankton group in the ocean because of their
79 production of calcified scales called coccoliths. They are not only important primary
80 producers, but also play predominant roles in the oceanic calcification via releasing
81 CO₂ to the atmosphere (Rost and Riebesell, 2004). Thus, coccolithophores have a
82 complex and significant influence on global carbon cycle, by playing an important
83 role in ocean-atmosphere exchange of CO₂ (Rost and Riebesell, 2004). Of all
84 coccolithophores, *Emiliania huxleyi* is the most widely distributed and the most
85 abundant species (Winter et al., 2014), with the capacity to form spatially extensive
86 blooms in mid- to high-latitudes (Raittos et al., 2006; Tyrrell and Mérion, 2004).
87 Evidence from *in situ* and satellite observations indicates that *E. huxleyi* is
88 increasingly expanding its range poleward in both hemispheres over the last two



89 decades, and contributing factors to this poleward expansion may differ between
90 regions and hemispheres (Winter et al., 2014). For example, nutrients and dissolved
91 inorganic carbon (DIC) were positively correlated with the increase in
92 coccolithophore abundance in the subtropical North Atlantic (Krumhardt et al., 2016),
93 while temperature and irradiance were best able to explain variability in *E.*
94 *huxleyi*-dominated coccolithophore community composition and abundance across the
95 Drake Passage (Southern Ocean) (Charalampopoulou et al., 2016). Hence, empirical
96 data on the responses of *E. huxleyi* to different environmental drivers would be critical
97 for fully understanding the roles of this prominent coccolithophore species in marine
98 ecosystems.

99 Extensive experimental studies have shown highly variable responses of *E. huxleyi*
100 to rising atmospheric CO₂ (reviewed by Feng et al., 2017; Meyer and Riebesell, 2015),
101 while other studies focused on the influence of other environmental factors such as
102 temperature (Rosas-Navarro et al., 2016; Sett et al., 2014; Sorrosa et al., 2005), light
103 intensity (Nanninga and Tyrrell, 1996; Xing et al., 2015) and nutrient availability
104 (Oviedo et al., 2014; Paasche, 1998). Responses of *E. huxleyi* to the interactions
105 between these different factors have recently received more attention (De Bodt et al.,
106 2010; Feng et al., 2008; Milner et al., 2016; Perrin et al., 2016; Rokitta and Rost,
107 2012). Many of these studies above focused on the physiological, calcification and
108 photosynthetic responses of *E. huxleyi* due to its considerable role in global carbon
109 cycle. However, biogeochemical cycles of the major nutrient elements (nitrogen and
110 phosphorus) and carbon are tightly linked (Hutchins et al., 2009), and thus variability



111 in stoichiometric C:N:P ratios (cellular quotas and ratios of C, N and P) in *E. huxleyi*
112 can also be a core feature of ocean biogeochemistry. Moreover, element budgets in
113 organisms are primarily determined by the physiology and biochemistry of
114 biochemicals such as proteins and fatty acids (FAs) (Anderson et al., 2004; Sterner
115 and Elser, 2002). Thus, studying simultaneous changes of elements and biochemicals
116 enables the connection between climate change and ecosystem functions such as
117 elemental cycles; however, the role of nutrient content of food is often overlooked in
118 climate change ecology (Rosenblatt and Schmitz, 2016). Recently, Bi et al. (2017)
119 investigated responses of stoichiometric C:N:P and FAs to the interactions of three
120 environmental factors in the diatom *Phaeodactylum tricornutum* and the cryptophyte
121 *Rhodomonas* sp., showing dramatic effects of warming and nutrient deficiency, and
122 modest effects of increased $p\text{CO}_2$. However, for the key coccolithophore species *E.*
123 *huxleyi* much less is known about the simultaneous changes in stoichiometric C:N:P
124 ratios, calcification, photosynthesis and FAs in response to multiple environmental
125 factor changes.

126 In the present study, we conducted semi-continuous cultures of *E. huxleyi* to
127 disentangle potential effects of temperature, N:P supply ratios and $p\text{CO}_2$ on *E. huxleyi*
128 stoichiometric C:N:P ratios, particulate inorganic carbon (PIC) and particulate organic
129 carbon (POC) contents and their ratios, and FA composition. As the physiological (i.e.,
130 cellular) PIC and POC variations cannot directly be up scaled to total population
131 response (Matthiessen et al., 2012), responses of PIC and POC contents in our study
132 were shown both on the cellular (as pg cell^{-1}) and the population (as $\mu\text{g ml}^{-1}$) levels.



133 FA data were expressed as a percentage of total fatty acids (TFAs) (FA proportion, %
134 of TFAs) to better compare our results with those in previous studies. FAs were also
135 quantified on a per unit biomass ($\mu\text{g mg C}^{-1}$), which is an ideal approach when
136 considering nutritional quality of phytoplankton for herbivores (Piepho et al., 2012).
137 Specifically, we addressed the following two questions in the present study, (i) what is
138 the sensitivity of stoichiometric C:N:P ratios, PIC:POC contents and their ratios, and
139 FAs of *E. huxleyi* to changes in temperature, N:P supply ratios and $p\text{CO}_2$? (ii) how do
140 stoichiometric C:N:P ratios, PIC:POC and FAs of *E. huxleyi* respond to the
141 interactions between the three environmental factors?

142 **2 Material and methods**

143 **2.1 Experimental setup**

144 To address our questions on how multiple environmental drivers influence
145 elemental and FA composition in *E. huxleyi*, we performed a semi-continuous culture
146 experiment crossing three temperatures (12, 18 and 24 °C), three N:P supply ratios
147 (10:1, 24:1 and 63:1 mol mol⁻¹) and two $p\text{CO}_2$ levels (560 and 2400 μatm) with a
148 strain of *E. huxleyi* originating from waters off Terceira Island (Azores, North
149 Atlantic). The target values of temperature, N:P supply ratios and $p\text{CO}_2$ were chosen
150 to reflect a present natural regime and future ocean projections of each factor
151 (Boersma et al., 2016; Lewandowska et al., 2014; Moore et al., 2013; IPCC 2014;
152 Peñuelas et al., 2012).

153 All cultures were exposed to a light intensity of 100 $\mu\text{mol photons} \cdot \text{m}^{-2} \cdot \text{s}^{-1}$ at a
154 16:8 h light:dark cycle in temperature-controlled rooms. The culture medium was



155 prepared with sterile filtered (0.2 μm pore size, Sartobran[®] P 300; Sartorius,
156 Goettingen, Germany) North Sea water with a salinity of 37 psu. Macronutrients were
157 added as sodium nitrate (NaNO_3) and potassium dihydrogen phosphate (KH_2PO_4) to
158 achieve three N:P supply ratios, i.e., $35.2 \mu\text{mol} \cdot \text{L}^{-1}$ N and $3.6 \mu\text{mol} \cdot \text{L}^{-1}$ P (10:1 mol
159 mol^{-1}), $88 \mu\text{mol} \cdot \text{L}^{-1}$ N and $3.6 \mu\text{mol} \cdot \text{L}^{-1}$ P (24:1 mol mol^{-1}) and $88 \mu\text{mol} \cdot \text{L}^{-1}$ N and
160 $1.4 \mu\text{mol} \cdot \text{L}^{-1}$ P (63:1 mol mol^{-1}). Vitamins and trace metals were added based on the
161 modified Provasoli's culture medium (Ismar et al., 2008; Provasoli, 1963). Initial
162 $p\text{CO}_2$ of the culture medium was manipulated by bubbling with air containing the
163 target $p\text{CO}_2$. Three replicates were set up for each treatment, resulting in 54
164 experimental units. Each culture was kept in a sealed cell culture flask with 920 mL
165 culture volume. Culture flasks were carefully rotated twice per day at a set time to
166 minimize sedimentation.

167 First, batch culture experiments were performed to obtain an estimate of the
168 observed maximal growth rate (μ_{max} , day $^{-1}$) under three temperatures, three N:P
169 supply ratios and two $p\text{CO}_2$ levels. μ_{max} was calculated based on the changes of
170 population cell density within exponential phase (Bi et al., 2012). Once batch cultures
171 reached the early stationary phase, semi-continuous cultures were started with the
172 algae from batch cultures. The specific growth rate of 20% of μ_{max} (μ , day $^{-1}$) was
173 applied. The equivalent daily renewal rate (D , day $^{-1}$) can be calculated according to
174 the equation $D = 1 - e^{-\mu t}$, where t is renewal interval (day) (here $t = 1$ day). The
175 incubation water was exchanged with fresh filtered seawater enriched by
176 macronutrients and micronutrients according to the target N:P supply ratios, as well as



177 pre-acclimated to the desired $p\text{CO}_2$ level. To counterbalance the biological
178 CO_2 -drawdown, the required amount of CO_2 -saturated seawater was also added.
179 Renewal of the cultures was carried out at the same hour every day. The steady state
180 in semi-continuous cultures was assessed based on the net growth rate (r). When r
181 was zero (at steady state), μ was equivalent to D .

182 **2.2 Sample analysis**

183 Sampling took place at steady state for the following parameters: cell density, DIC,
184 total alkalinity (TA), pH, total particulate carbon (TPC), POC, particulate organic
185 nitrogen and phosphorus (PON and POP), and FAs. Cell density was counted daily in
186 batch and semi-continuous cultures. pH measurements were conducted daily in
187 semi-continuous cultures (Fig. S1), and the electrode was calibrated using standard
188 pH buffers (pH 4 and pH 7; WTW, Weilheim, Germany).

189 DIC water samples were gently filtered using a single-use syringe filter (0.2 μm ,
190 Minisart RC25; Sartorius, Goettingen, Germany) which was connected to the intake
191 tube of a peristaltic pump. Samples were collected into 10 ml glass vials, and all vials
192 were immediately sealed after filling. DIC was analyzed following Hansen et al.
193 (2013) using a gas chromatographic system (8610C; SRI-Instruments, California,
194 USA). Samples for TA analysis were filtered through GF/F filters (Whatman GmbH,
195 Dassel, Germany) and analyzed with the Tirino plus 848 (Metrohm, Filderstadt,
196 Germany). The remaining carbonate parameter $p\text{CO}_2$ was calculated using CO2SYS
197 (Pierrot et al., 2006) and the constants supplied by Hansson (1973) and Mehrbach et
198 al. (1973) that were refitted by Dickson and Millero (1987) (Table S1).



199 TPC, POC, PON and POP samples were filtered onto pre-combusted and
200 pre-washed (5%~10% HCl) GF/F filters (Whatman GmbH, Dassel, Germany). For
201 POC samples, PIC was removed by exposing filters containing TPC to fuming
202 hydrochloric acid for 12h. Before analysis, filters were dried at 60 °C and stored in a
203 desiccator. POC and PON was simultaneously determined by gas chromatography in
204 an organic elemental analyzer (Thermo Flash 2000; Thermo Fisher Scientific Inc.,
205 Schwerte, Germany) after Sharp (1974). POP was analyzed colorimetrically by
206 converting organic phosphorus compounds to orthophosphate (Hansen and Koroleff,
207 1999). PIC was determined by subtracting POC from TPC. PIC and POC production
208 were estimated by multiplying μ with cellular PIC or POC content, respectively.

209 FA samples were taken on pre-combusted and hydrochloric acid-treated GF/F
210 filters (Whatman GmbH, Dassel, Germany), and stored at -80 °C before measurement.
211 FAs were measured as fatty acid methyl esters (FAMEs) using a gas chromatograph
212 (Trace GC-Ultra; Thermo Fisher Scientific Inc., Schwerte, Germany) according to the
213 procedure described in detail in Arndt and Sommer (2014). The FAME 19:0 was
214 added as internal standard and 21:0 as esterification control. The extracted FAs were
215 dissolved with n-hexane to a final volume of 100 μ L. Sample aliquots (1 μ L) were
216 given into the GC by splitless injection with hydrogen as the carrier gas. Individual
217 FAs were integrated using Chromcard software (Thermo Fisher Scientific Inc.,
218 Schwerte, Germany) and identified with reference to the standards Supelco 37
219 component FAME mixture and Supelco Menhaden fish oil.

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221 **2.3 Statistical analysis**

222 Generalized linear mixed models (GLMMs) were applied to test the best model
223 explaining the variations in elemental and FA composition. In our study, response
224 variables included stoichiometric C:N:P ratios (as mol mol⁻¹), PIC and POC contents
225 per cell (as pg cell⁻¹) and per ml (as µg ml⁻¹), PIC and POC production (as pg cell⁻¹
226 d⁻¹), PIC:POC ratio, FA proportion (as % of TFAs) and contents (as µg mg C⁻¹), with
227 temperature, N:P supply ratios and *p*CO₂ as fixed effects. Target distributions were
228 tested and link functions were consequently chosen. For all response variables, we
229 tested models containing first order effects, and second and third order interactions of
230 the three factors. The model that best predicted targets was selected based on the
231 Akaike Information Criterion corrected (AICc), i.e., a lower AICc value representing
232 a better fit of the model. Changes of 10 units or more in AICc values were considered
233 as a reasonable improvement in the fitting of GLMMs (Bolker et al., 2009). In case
234 AICc values were comparable (<10 units difference), the simpler model was thus
235 chosen, unless there were significant second or third order interactions detected.
236 Differences in AICc values between different models were more than 10 for most
237 variables, with the exception for N:P biomass ratio, cellular PIC and POC contents,
238 and saturated fatty acid (SFA) and DHA proportions, which were less than 10 between
239 different models (Table S2).

240 Two factorial ANOVA was used to test the effects of temperature and *p*CO₂ on μ_{\max} .
241 As μ_{\max} should be equal across different nutrients (Cherif and Loreau, 2010), the
242 effect of N:P supply ratios on μ_{\max} was not tested. The normality of dependent



243 variables was checked using the Shapiro-Wilk's *W*-test. For the significant factors, the
244 magnitude of effect ($\omega^2 = (\text{effect sum of squares} - \text{effect degree of freedom} \times \text{error}$
245 $\text{mean square}) / (\text{total sum of squares} + \text{error mean square})$) was calculated to
246 determine the variance in a response variable and to relate this to the total variance in
247 the response variable (Graham and Edwards, 2001; Hughes and Stachowicz, 2009).

248 Nested models were applied to test whether the response pattern to one factor (a
249 nested factor) was significant within another factor, in case significant second order
250 interactions were detected in GLMM. Also, the nature (antagonistic, additive, or
251 synergistic) of significant second order interactions was analysed according to
252 Christensen et al. (2006). The observed combined effect of two factors was compared
253 with their expected net additive effect [e.g., $(\text{factor}_1 - \text{control}) + (\text{factor}_2 - \text{control})$],
254 which was based on the sum of their individual effects. If the observed combined
255 effect exceeded their expected additive effect, the interaction was defined as
256 synergism. In contrast, if the observed combined effect was less than the additive
257 effect, the interaction was defined as antagonism.

258 All statistical analyses were conducted using SPSS 19.0 (IBM Corporation, New
259 York, USA). Significance level was set to $p < 0.05$ in all statistical tests.

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265 **3 Results**

266 **3.1 Maximal growth rate (μ_{\max})**

267 The observed maximal growth rate of *E. huxleyi* responded significantly to
268 temperature and the interactions between temperature and $p\text{CO}_2$ (Bold letters in Table
269 1), with temperature causing 82% of the effect magnitude. Increasing temperature
270 stimulated μ_{\max} in each $p\text{CO}_2$ treatment, causing μ_{\max} to be two to three times higher
271 at the highest temperature than those at the lowest temperature (Fig. 1). In contrast,
272 the effect of elevated $p\text{CO}_2$ on μ_{\max} showed no consistent pattern, indicating a slight
273 positive effect at 12°C, a strong negative effect at 18°C and a weak negative effect at
274 24°C (Fig. 1). Moreover, the trend of μ_{\max} to increase with increasing temperature
275 differed between the two $p\text{CO}_2$ treatments, showing that at the low $p\text{CO}_2$, the slope of
276 μ_{\max} response to increasing temperature was higher from 12°C to 18°C and it became
277 lower from 18°C to 24°C, while the slope of μ_{\max} response showed no clear difference
278 between three temperatures at the high $p\text{CO}_2$ treatment.

279 **3.2 C:N:P stoichiometry**

280 GLMMs results showed that N:C, P:C and N:P biomass ratios responded
281 significantly to N:P supply ratios (Bold letters in Table 2), while only N:C biomass
282 ratios showed significant responses to temperature, with non-significant effect of
283 $p\text{CO}_2$ detected. Increasing N:P supply ratios caused an increased trend in N:C biomass
284 ratios (Fig. 2a) and a decrease in P:C biomass ratios (Fig. 2b), resulting in a positive
285 relationship between N:P biomass ratios and N:P supply ratios (Fig. 2c). The response
286 of N:C biomass ratios to increasing temperature was complex, showing a U-shaped



287 response under N deficiency (N:P supply ratio = 10:1 mol mol⁻¹), and positive
288 responses under higher N:P supply ratios (Fig. 2a).

289 All three PIC-related variables (cellular PIC contents, population yields of PIC and
290 PIC production) responded significantly to temperature and $p\text{CO}_2$; and there were also
291 significant interactions between temperature and $p\text{CO}_2$ (and N:P supply ratios) on
292 cellular PIC contents, and significant effects of N:P supply ratios on population yields
293 of PIC and PIC production (Table 2). Increasing temperature and N deficiency
294 affected cellular PIC contents antagonistically, while increasing temperature and P
295 deficiency (N:P supply ratio = 63:1 mol mol⁻¹) had synergistic interactions on cellular
296 PIC contents (Table S3). As a result, cellular PIC contents showed a trend to decrease
297 slightly with increasing temperature under N deficiency and a trend to increase under
298 higher N:P supply ratios (Fig. 3a). Increasing temperature and enhanced $p\text{CO}_2$
299 affected cellular PIC contents synergistically (Table S3), with the negative response of
300 cellular PIC contents to enhanced $p\text{CO}_2$ being significantly weaker as temperature
301 increased (Fig. 3d; Nested model, $p < 0.001$). Population yields of PIC decreased
302 (under higher temperatures), but PIC production increased with increasing N:P supply
303 ratios (Fig. 3b, c), while both PIC variables increased with increasing temperature and
304 decreased with enhanced $p\text{CO}_2$ (Fig. 3e, f).

305 The three POC-related variables (cellular POC contents, population yields of POC
306 and POC production) responded significantly to temperature; and there were also
307 significant interactions between temperature and N:P supply ratios on cellular POC
308 contents, and significant effects of N:P supply ratios on population yields of POC and



309 POC production (Table 2). For cellular POC contents, increasing temperature and N
310 (and P) deficiency showed synergistic interactions, resulting in the highest values at
311 the lowest temperature under lower N:P supply ratios and an increasing trend with
312 increasing temperature under P deficiency (Fig. 4a; Table S3). Population yields of
313 POC showed an unimodal response and POC production showed a trend to increase in
314 response to increasing temperature and N:P supply ratios (Fig. 4b, c).

315 PIC:POC responded significantly to temperature and $p\text{CO}_2$ (Table 2), showing a
316 clear increase with increasing temperature and a decrease with enhanced $p\text{CO}_2$ (Fig.
317 5).

318 **3.3 Fatty acids**

319 The most abundant FA group were polyunsaturated fatty acids (PUFAs) (33%-54%
320 of TFAs), followed by SFAs (22%-46%) and monounsaturated fatty acids (MUFAs)
321 (13%-27%), across the entire tested gradients of temperature, N:P supply ratios and
322 $p\text{CO}_2$ (Table S4). The high proportion of PUFAs was predominantly caused by high
323 amounts of DHA (12%-31%) and 18:4n-3 (3%-13%), and SFAs was mainly
324 represented by 14:0 (13%-23%) and 16:0 (5%-11%). The major individual MUFA
325 was 18:1n-9 (8%-21%).

326 GLMMs results showed significant effects of temperature and $p\text{CO}_2$ on the
327 proportions of both MUFAs and PUFAs, and significant interactions between N:P
328 supply ratios and $p\text{CO}_2$ on SFAs (Table 2). Increasing temperature caused a decrease
329 in the proportion of MUFAs and an increase in PUFAs (Fig. 6a, b). In contrast,
330 enhanced $p\text{CO}_2$ resulted in an increase in MUFAs and a decrease in PUFAs at higher



331 temperatures (Fig. 6d, e). Moreover, enhanced $p\text{CO}_2$ and N (and P) deficiency
332 affected SFA proportion synergistically (Table S3), with the unimodal response of
333 SFA proportion to increasing N:P supply ratios being more pronounced at the high
334 $p\text{CO}_2$ (Fig. S2; Nested model, $p < 0.001$).

335 The proportion of major individual PUFAs (DHA) showed significant responses to
336 temperature and N:P supply ratios, and the interactions between temperature and N:P
337 supply ratios (and $p\text{CO}_2$) (Table 2). Increasing temperature caused an overall increase
338 in DHA, and DHA also had higher values under N and P deficiency (Fig. 6c). The
339 interactions between increasing temperature and N deficiency (and P deficiency and
340 enhanced $p\text{CO}_2$) affected DHA synergistically (Table S3), and the positive effect of
341 temperature became more pronounced at lower N:P supply ratios (Nested model, $p <$
342 0.001) and at the low $p\text{CO}_2$ (Nested model, $p < 0.001$) (Fig. 6c, f).

343

344 **4 Discussion**

345 Our study scales the impacts of temperature, N:P supply ratios and $p\text{CO}_2$ on
346 elemental and FA composition of the ubiquitously important calcifier *E. huxleyi*, while
347 accounting for their interactive effects. Overall, C:N:P biomass ratios changed
348 markedly in response to N:P supply ratios, resulting in up to a 62% increase in N:P
349 biomass ratios (Fig. 7). PIC:POC showed an increase of 41% with warming and a
350 decrease of 35% under high $p\text{CO}_2$. PUFA proportions showed an increase of 13%
351 with warming and a decrease of 7% with enhanced $p\text{CO}_2$, indicating a partial
352 compensation by $p\text{CO}_2$ of a predominantly temperature-driven response. The overall



353 response patterns of C:N:P stoichiometry and PUFAs in our study are consistent with
354 those on the global scale (Martiny et al., 2013), and conform with the meta-analysis
355 results on haptophytes (Hixson and Arts, 2016). In line with these studies, we also
356 detected significant interactions between temperature, N:P supply ratios and $p\text{CO}_2$,
357 indicating variable response patterns of elemental and FA composition in *E. huxleyi*
358 under any given constellation of environmental factors. Our results thus underscore
359 the importance of simultaneous consideration of multiple environmental drivers,
360 demonstrating differential effects of the three environmental factors on elemental and
361 FA composition of *E. huxleyi*.

362 **4.1 Responses of maximal growth rate**

363 Increasing temperature (12–24°C) significantly accelerated μ_{max} of *E. huxleyi* in our
364 study (Fig. 1). This positive correlation between increasing temperature and growth
365 rate is typical for many *E. huxleyi* strains within the range of temperature 12 to 24°C
366 (Feng et al., 2008; Rosas-Navarro et al., 2016; Sett et al., 2014; van Bleijswijk et al.,
367 1994). However, the extent to which growth rate of *E. huxleyi* increases with
368 increasing temperature varies between *E. huxleyi* strains, which may contribute to
369 specific biogeographic distribution of different strain (Paasche, 2002). For example,
370 growth rate of *E. huxleyi* from the Gulf of Maine (~42 °N) was 1.2 times higher at 26°C
371 than at 16°C, while growth rate of *E. huxleyi* from the Sargasso Sea (~20–35 °N) was
372 1.6 times higher at the higher temperature (Paasche, 2002). Also, our results showed
373 that μ_{max} of *E. huxleyi* was two to three times higher at the highest temperature than at
374 the lowest temperature, suggesting that temperature is an important environmental



375 factor in controlling the distribution of this strain of *E. huxleyi*.

376 Also, temperature showed significant interactions with $p\text{CO}_2$ on μ_{max} of *E. huxleyi*
377 in this study. On the one hand, elevated $p\text{CO}_2$ led to an increase in μ_{max} at the lowest
378 temperature (12°C) but a decrease at higher temperatures (18 and 24°C) (Fig. 1).

379 Similarly, in response to increasing $p\text{CO}_2$, either an increase (Feng et al., 2008) or a
380 decrease (Milner et al., 2016) in the growth of *E. huxleyi* was observed in previous
381 studies. Such a diverse response of phytoplankton growth rate to elevated $p\text{CO}_2$ has
382 been widely observed within functional groups such as coccolithophores and diatoms,
383 between taxa and even between strains of the same species (Dutkiewicz et al., 2015).
384 Significant interactions between $p\text{CO}_2$ and temperature on *E. huxleyi* growth in our
385 study suggest that experimental temperature can be an important factor resulting in
386 diverse responses of phytoplankton growth to rising $p\text{CO}_2$ in previous studies.

387 On the other hand, we observed different slopes of μ_{max} in response to increasing
388 temperature at two $p\text{CO}_2$ levels, showing a decrease in the slope with increasing
389 temperature at the low $p\text{CO}_2$ and a relatively constant slope at the high $p\text{CO}_2$ (Fig. 1).

390 Our results are consistent with a conceptual graph proposed by Sett et al. (2014). The
391 graph showed a clear increase in metabolic rates from low to intermediate temperature
392 and a slight increase from intermediate to high temperature at the low $p\text{CO}_2$ (~560
393 μatm), while the changes of metabolic rates are similar from low to intermediate
394 temperature and from intermediate to high temperature at the high $p\text{CO}_2$ (~2400 μatm)
395 (Sett et al., 2014). While the physiological mechanisms governing adaptation of
396 phytoplankton to rising $p\text{CO}_2$ are still unclear, one possible explanation is that



397 increasing temperature may modulate the balance between a fertilizing effect of ocean
398 carbonation and a metabolic repression by ocean acidification (Bach et al., 2011; Sett
399 et al., 2014).

400 **4.2 Responses of C:N:P biomass ratios**

401 **4.2.1 Effects of N:P supply ratios**

402 N:P supply ratios showed highly significant effects on N:C, P:C and N:P biomass
403 ratios in *E. huxleyi* in this study, contributing up to 62% changes in C:N:P biomass
404 ratios (Table 2; Fig. 7). The strong contribution of nutrient availability to the
405 elemental ratios of marine phytoplankton community biomass was also found on the
406 global scale (Daines et al., 2014; Martiny et al., 2013), with nitrate concentration as a
407 proxy of nutrient availability explaining 36% of variation in N:P ratio and 42% of
408 variation in C:P ratio (Martiny et al., 2013). Similarly, previous lab experiments
409 reported that nutrient availability prevailed the governing effect on stoichiometric
410 ratios in *E. huxleyi* (Skau, 2015).

411 Overall, N deficiency caused low N:C and N:P biomass ratios, while P deficiency
412 resulted in low P:C biomass ratios and high N:P biomass ratios in *E. huxleyi* in this
413 and most previous studies (Langer et al., 2013; Leonardos and Geider, 2005b; Perrin
414 et al., 2016). An important biogeochemical question is the extent to which N:P
415 biomass ratios change under N and P deficiency, respectively. Our results showed that
416 changes in N:P biomass ratios of *E. huxleyi* were mostly due to P deficiency (a 62%
417 increase), with smaller changes (a 36% decrease) induced by N deficiency (Fig. 7).
418 This observation is consistent with the high variability of P:C in response to changes



419 in phosphate and the less variable N:C to changes in nitrate based on global
420 suspended particle measurements (Galbraith and Martiny, 2015). Indeed, conflicting
421 response patterns of C:N:P biomass ratios of *E. huxleyi* to nutrient deficiency were
422 observed between different studies (Borchard and Engel, 2012; Matthiessen et al.,
423 2012; Perrin et al., 2016), which could be due to strain-specific responses and
424 interactions between nutrients and other environmental drivers. Nevertheless, the
425 nutrient-dependence of C:N:P biomass ratios of *E. huxleyi* in our study is consistent
426 with responses of marine plankton on the global scale, which may reflect the capacity
427 of this species to thrive under a wide range of environmental conditions. This capacity
428 was largely revealed by a pan-genome assessment, which distributed genetic traits
429 variably between strains and showed a suit of core genes for the uptake of inorganic
430 nitrogen and N-rich compounds such as urea (Read et al., 2013). In spite of strain
431 diversity within *E. huxleyi*, a recent study suggested that the global physiological
432 response of this species to nutrient environments is highly conserved across strains
433 and may underpin its success under a variety of marine environments (Alexander,
434 2016).

435 **4.2.2 Effects of temperature**

436 Temperature had a weaker effect (5-8% changes) than N:P supply ratios on
437 variation of N:P and P:C ratios (Table 2; Fig. 7), consistent with results in marine
438 plankton communities on the global scale (Martiny et al., 2013). While both N:C and
439 P:C biomass ratios increased with increasing temperature in our study, the changes in
440 N:C ratios (8%) were larger than those in P:C ratios (5%). As a result, the observed



441 N:P ratios also increased in response to warming, in accordance with the positive
442 relationship between temperature and N:P ratios of marine phytoplankton sampled
443 from different biogeochemical provinces (Martiny et al., 2016; Toseland et al., 2013;
444 Yvon-Durocher et al., 2015). These responses are consistent with proposed
445 physiological mechanisms (Toseland et al., 2013), which showed that eukaryotic
446 phytoplankton at higher temperatures required less P-rich ribosomes and thus
447 produced higher N:P ratios.

448 **4.2.3 Effects of $p\text{CO}_2$**

449 Partial CO_2 pressure showed a non-significant effect on *E. huxleyi* C:N:P biomass
450 ratios in our study (Table 2), being consistent with the previous findings of the less
451 important effect of $p\text{CO}_2$ than nutrients and temperature on *E. huxleyi* (Boyd et al.,
452 2010; Feng, 2015). Several experimental studies also showed non-significant changes
453 in *E. huxleyi* C:N:P stoichiometry in response to rising $p\text{CO}_2$ (Engel et al., 2014;
454 Lefebvre et al., 2012; Olson et al., 2017); however, some studies reported clear
455 changes in C:N or C:P ratio (Engel et al., 2005; Leonardos and Geider, 2005a;
456 Matthiessen et al., 2012). Both experimental and model studies have suggested that
457 rising $p\text{CO}_2$ seems to change phytoplankton stoichiometry under specific conditions,
458 e.g., at high light condition (Feng et al., 2008) and at low nutrient loads (Leonardos
459 and Geider, 2005a; Verspagen et al., 2014). Moreover, the sensitivity of phytoplankton
460 to CO_2 may also depend on culture systems, either transient or sufficiently stable,
461 influencing optimal allocation of energy and resources (Engel et al., 2014).

462 Taken together, our results indicate that N:P supply ratios are reflected in elemental



463 make-up of cell biomass in *E. huxleyi*, across different temperatures and $p\text{CO}_2$ levels,
464 showing the absence of significant interactions between the three environmental
465 factors. However, for two algal species from non-calcifying classes (the diatom *P.*
466 *tricornutum* and the cryptophyte *Rhodomonas* sp.) temperature had the most
467 consistent significant effect on N:C and P:C biomass ratios and showed significant
468 interactions with N:P supply ratios and $p\text{CO}_2$ in our previous work (Bi et al., 2017).
469 Both temperature and N:P supply ratios were also ranked as important factors in
470 regulating phytoplankton stoichiometry in previous studies (Boyd et al., 2010; Feng,
471 2015). Differential C:N:P responses to environmental drivers between phytoplankton
472 groups suggest that taxonomic composition may explain the variable C:N:P ratios in
473 surface phytoplankton community (Martiny et al., 2013).

474 **4.3 Responses of PIC:POC contents and ratios**

475 Both partial CO_2 pressure and temperature had highly significant effects on
476 PIC:POC in our study, with enhanced $p\text{CO}_2$ resulting in an overall 49% decrease in
477 PIC:POC and warming resulting in a 41% increase in PIC:POC, while N:P supply
478 ratios showed no significant effect (Table 2; Fig. 7). This result is in agreement with
479 rankings of the importance of environmental drivers on PIC:POC in a Southern
480 Hemisphere strain of *E. huxleyi*, showing the order of $p\text{CO}_2$ (negative effect) >
481 temperature (positive effect) and non-significant effect of nitrate or phosphate (Feng,
482 2015).

483 The negative effect of enhanced $p\text{CO}_2$ on PIC:POC was also observed for different
484 strains of *E. huxleyi* in most previous studies (Meyer and Riebesell, 2015 and



485 references therein). Negative responses of PIC:POC to increasing $p\text{CO}_2$ in our study
486 were driven by the significant decrease in the three PIC-related response variables
487 (calcification) and no significant change in all three variables of POC (photosynthesis)
488 (Table 2; Fig. 7). To date, studies and reviews also showed a greater impact of ocean
489 acidification on calcification than on photosynthesis in coccolithophores (De Bodt et
490 al., 2010; Feng et al., 2017; Meyer and Riebesell, 2015). Feng et al. (2017) suggested
491 that the decreased calcification in *E. huxleyi* may be caused by the increased
492 requirement of energy to counteract intracellular acidification. And the increased
493 activity of carbonic anhydrase (CA) at low $p\text{CO}_2$ may explain the lack of a significant
494 effect of $p\text{CO}_2$ on the photosynthetic or growth rate (Feng et al., 2017), as
495 up-regulation of CA at low DIC was previously observed (Bach et al., 2013).

496 A positive response or no clear change in PIC:POC was also observed for *E.*
497 *huxleyi* in response to high $p\text{CO}_2$ in previous studies (Feng et al., 2017), with the
498 influence of other environmental drivers proposed as a potential driver (De Bodt et al.,
499 2010; Feng et al., 2017; Feng et al., 2008). Indeed, we observed that the negative
500 relationship between cellular PIC contents and enhanced $p\text{CO}_2$ became weaker at the
501 highest temperature (Fig. 3d). This result is in agreement with the modulating effect
502 of temperature on the CO_2 sensitivity of key metabolic rates in coccolithophores, due
503 to the shift of the optimum CO_2 concentration for key metabolic processes towards
504 higher CO_2 concentrations from intermediate to high temperatures (Sett et al., 2014).

505 Temperature causes diverse responses of calcification and photosynthesis within *E.*
506 *huxleyi* species in the literature (Rosas-Navarro et al., 2016 and references therein)



507 and the present study. Overall, our study showed that the positive response of
508 PIC:POC to increasing temperature was driven by a marked increased PIC
509 (28%~161%), and by a less pronounced change in POC (-8%~68%) (Table 2; Fig. 7).
510 The overall responses of PIC:POC contents and ratios to increasing temperature was
511 consistent with those in other *E. huxleyi* strains (Matson et al., 2016; Sett et al., 2014),
512 indicating carbon allocation to calcification rather than photosynthesis at high
513 temperatures (Sett et al., 2014). Specifically, we observed that the interactions of
514 warming and nutrient deficiency (and high $p\text{CO}_2$) synergistically affected both PIC
515 and POC cellular contents in most cases, suggesting that nutrient deficiency and high
516 $p\text{CO}_2$ are likely to enhance the effect of warming on *E. huxleyi* calcification and
517 photosynthesis efficiency.

518 In summary, our results showed an overall reduced PIC:POC in *E. huxleyi* under
519 future ocean scenarios of warming and higher CO_2 (Fig. 5b), consistent with the
520 reduced ratio of calcium carbon production to organic carbon during the *E. huxleyi*
521 bloom in previous mesocosm experiments (Delille et al., 2005; Engel et al., 2005). It
522 is worth noting that cellular PIC and POC contents are a measure for physiological
523 response and cannot be used to infer population response and thus potential carbon
524 export (Matthiessen et al., 2012), as different responses between cellular and
525 population yields of PIC (and POC) to environmental changes were evident in
526 previous work (Matthiessen et al., 2012) and the present study (Fig. 7). Moreover,
527 significant interactions between temperature and $p\text{CO}_2$ (and N:P supply ratios) on
528 cellular particulate carbon contents suggest that the effects of multiple environmental



529 drivers should be tested to better generalize the dynamic of particulate carbon from
530 laboratory observations to that in natural conditions.

531 **4.4 Responses of fatty acids**

532 **4.4.1 Effects of temperature**

533 Temperature had highly significant effects on the proportions of MUFA, PUFA
534 and DHA, with increasing temperature causing a 20% decline in MUFA and a 13%
535 increase in PUFA in our study (Table 2; Fig. 7). This result is consistent with the
536 negative response of MUFA proportions and positive response of PUFA to warming
537 in other haptophytes based on a meta-analysis on 137 FA profiles (Hixson and Arts,
538 2016), showing an opposite response to general patterns of phytoplankton FAs to
539 warming. Although warming is expected to have a negative effect on the degree of
540 fatty acid unsaturation to maintain cell membrane structural functions (Fuschino et al.,
541 2011; Guschina and Harwood, 2006; Sinensky, 1974), variable FA responses to
542 warming were widely observed in different phytoplankton groups (Bi et al., 2017;
543 Renaud et al., 2002; Thompson et al., 1992). Contradictory findings were even
544 reported in meta-analyses on large FA profiles such as the absence (Galloway and
545 Winder, 2015) or presence (Hixson and Arts, 2016) of the negative correlation
546 between temperature and the proportion of long-chain EFAs in freshwater and marine
547 phytoplankton. While the underlying mechanisms of variable FA responses are still
548 unclear, it is known that both phylogeny and environmental conditions determine
549 phytoplankton FA composition (Bi et al., 2014; Dalsgaard et al., 2003; Galloway and
550 Winder, 2015). In our study, we found significant interactions between temperature



551 and $p\text{CO}_2$ (and N:P supply ratios) on the individual FA component DHA, showing that
552 $p\text{CO}_2$ and nutrient availability may alter the effect of warming on *E. huxleyi* FA
553 composition.

554 **4.4.2 Effects of $p\text{CO}_2$**

555 Partial CO_2 pressure had significant effects on the proportion of MUFA and
556 PUFA in *E. huxleyi* in our study, with enhanced $p\text{CO}_2$ causing an overall 7% increase
557 in MUFA and a 7% decrease in PUFA (Table 2; Fig. 7), consistent with FA response
558 patterns in another strain of *E. huxleyi* (Riebesell et al., 2000) and other
559 coccolithophores (Fiorini et al., 2010). Also in a natural plankton community
560 (Raunefjord, southern Norway), PUFA proportion was reduced at high $p\text{CO}_2$ level in
561 the nano-size fraction, suggesting a reduced Haptophyta (dominated by *E. huxleyi*)
562 biomass and a negative effect of high $p\text{CO}_2$ on PUFA proportion (Bermúdez et al.,
563 2016). To date, several mechanisms have been suggested to explain the reduced
564 PUFA at high $p\text{CO}_2$ in green algae (Pronina et al., 1998; Sato et al., 2003; Thompson,
565 1996), with much less work conducted in other phytoplankton groups. One possible
566 mechanism was demonstrated in the study on *Chlamydomonas reinhardtii*, showing
567 that the repression of the CO_2 -concentrating mechanisms (CCMs) was associated with
568 reduced FA desaturation at high CO_2 concentration (Pronina et al., 1998). Similarly in
569 our study, both the proportion and contents of PUFA decreased at high $p\text{CO}_2$ (Fig. 7),
570 which may be attributed to the repression of CCMs at high $p\text{CO}_2$ in *E. huxleyi*.

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573 **4.4.3 Effects of N:P supply ratios**

574 Significant effects of N:P supply ratios was only observed for the proportion of
575 DHA in this study (Table 2), with N and P deficiency causing a 14% and 22%
576 increase in DHA proportion, respectively (Fig. 7). While nutrients often play a major
577 role on phytoplankton lipid composition (Fields et al., 2014; Hu et al., 2008), the less
578 pronounced effects of nutrient deficiency in our study indicate a unique lipid
579 biosynthesis in *E. huxleyi*. Indeed, Van Mooy et al. (2009) suggested that *E. huxleyi*
580 used non-phosphorus betaine lipids as substitutes for phospholipids in response to P
581 scarcity. Genes are also present in the core genome of *E. huxleyi* for the synthesis of
582 betaine lipids and unusual lipids used as nutritional/feedstock supplements (Read et
583 al., 2013). Therefore, the lack of significant nutrient effects on most FA groups in *E.*
584 *huxleyi* in our study may be caused by the functioning of certain lipid substitutions
585 under nutrient deficiency.

586 In summary, this study showed the strongest effects of temperature and $p\text{CO}_2$, a
587 weaker effect of N:P supply ratios, and non-significant interactions between the three
588 environmental factors, on the proportions of unsaturated fatty acids in *E. huxleyi*.
589 Specifically, we observed that PUFA proportions in *E. huxleyi* responded positively to
590 warming and negatively to enhanced $p\text{CO}_2$. It should be noted that using different
591 units to quantify FA composition may cause contradictory results, e.g., an increase in
592 PUFA proportion (% of TFAs) but an overall decline in PUFA contents per biomass
593 ($\mu\text{g mg C}^{-1}$) with increasing temperature in our study (Fig. 7; Table S4, S5). Such
594 biochemical quality differences can translate to higher trophic levels (Rossoll et al.,



595 2012) and refer to direct effects of environmental changes on low trophic level
596 consumers, which can be modified by indirect bottom-up driven impacts through the
597 primary producers (Garzke et al., 2016; Garzke et al., 2017). Similar to *E. huxleyi*,
598 PUFA contents in two species of non-calcifying classes (*P. tricornutum* and
599 *Rhodomonas* sp.) responded negatively to warming and positively to N (and P)
600 deficiency (Bi et al., 2017). However, differential responses were also observed, e.g.,
601 a significant negative effect of enhanced $p\text{CO}_2$ on PUFA contents in *E. huxleyi*, but a
602 non-significant effect of $p\text{CO}_2$ on PUFA contents in *P. tricornutum* and *Rhodomonas*
603 sp. (Bi et al., 2017). This different response between phytoplankton groups is in
604 agreement with findings in mesocosm studies (Bermúdez et al., 2016; Leu et al.,
605 2013), suggesting that changes in taxonomic composition can cause different
606 relationship between PUFAs and $p\text{CO}_2$ in natural phytoplankton community.

607 5 Conclusions

608 Our study showed that N:P supply ratios had the strongest effect on C:N:P biomass
609 ratios, while $p\text{CO}_2$ and temperature played more influential roles on PIC:POC and
610 PUFA proportions in *E. huxleyi*. The specific response patterns of elemental ratios and
611 PUFAs have important implications for understanding biogeochemical and ecological
612 functioning of *E. huxleyi*. The overall low PIC:POC under future ocean scenarios
613 (warming and enhanced $p\text{CO}_2$) indicates that carbon production by the strain *E.*
614 *huxleyi* in our study acts as a carbon sink. This argument is consistent with the finding
615 that the calcification of coccolithophores decreases with increasing $p\text{CO}_2$ in the
616 present ocean and over the past forty thousand years (Beaufort et al., 2011). It should



617 be noted that carbon production discussed so far are based on responses of cellular or
618 population PIC:POC in a single strain of *E. huxleyi*, and thus it cannot be used to infer
619 potential carbon flux between surface ocean and the atmosphere. From the ecological
620 point of view, the contradictory response patterns of N:C and P:C biomass ratios (an
621 overall increase) and PUFA contents (an overall decrease) under future ocean
622 scenarios in our study make the extraction of a generic response signal of nutritional
623 food quality difficult. Nonetheless, the observations presented here suggest
624 differential responses of elements and FAs to rising temperature, enhanced $p\text{CO}_2$ and
625 nutrient deficiency in *E. huxleyi*, being to some extent unique compared with algal
626 species from non-calcifying classes. Thus, the role of multiple environmental drivers
627 under the biodiversity context should be considered to truly estimate the future
628 functioning of phytoplankton in the changing marine environments.

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639 **Data availability:** data sets are available upon request by contacting Meixun Zhao

640 (maxzhao@ouc.edu.cn and maxzhao04@yahoo.com).

641 **Author contribution:** R. Bi, S. Ismar, U. Sommer and M. Zhao designed the
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645

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661 **References**

662 Alexander, H.: Defining the ecological and physiological traits of phytoplankton across
663 marine ecosystems, Ph.D. thesis, Woods Hole Oceanographic Institution, Woods Hole,
664 USA, 179 pp., 2016.

665 Anderson, T. R., Boersma, M., and Raubenheimer, D.: Stoichiometry: linking
666 elements to biochemicals, *Ecology*, 85, 1193-1202, doi: 10.1890/02-0252, 2004.

667 Arndt, C. and Sommer, U.: Effect of algal species and concentration on development
668 and fatty acid composition of two harpacticoid copepods, *Tisbe* sp. and *Tachidius*
669 *discipes*, and a discussion about their suitability for marine fish larvae, *Aquac. Nutr.*,
670 20, 44-59, doi: 10.1111/anu.12051, 2014.

671 Bach, L. T., Mackinder, L. C. M., Schulz, K. G., Wheeler, G., Schroeder, D. C.,
672 Brownlee, C., and Riebesell, U.: Dissecting the impact of CO₂ and pH on the
673 mechanisms of photosynthesis and calcification in the coccolithophore *Emiliania*
674 *huxleyi*, *New Phytol.*, 199, 121-134, doi: 10.1111/nph.12225, 2013.

675 Bach, L. T., Riebesell, U., and Schulz, K. G.: Distinguishing between the effects of
676 ocean acidification and ocean carbonation in the coccolithophore *Emiliania huxleyi*,
677 *Limnol. Oceanogr.*, 56, 2040-2050, doi: 10.4319/lo.2011.56.6.2040, 2011.

678 Beaufort, L., Probert, I., de Garidel-Thoron, T., Bendif, E. M., Ruiz-Pino, D., Metzl,
679 N., Goyet, C., Buchet, N., Coupel, P., Grelaud, M., Rost, B., Rickaby, R. E. M., and
680 de Vargas, C.: Sensitivity of coccolithophores to carbonate chemistry and ocean
681 acidification, *Nature*, 476, 80-83, doi: 10.1038/nature10295, 2011.

682 Bermúdez, J. R., Riebesell, U., Larsen, A., and Winder, M.: Ocean acidification
683 reduces transfer of essential biomolecules in a natural plankton community, *Sci. Rep.-*
684 *UK*, 6, 27749, doi: 10.1038/srep27749, 2016.

685 Bi, R., Arndt, C., and Sommer, U.: Stoichiometric responses of phytoplankton species
686 to the interactive effect of nutrient supply ratios and growth rates, *J. Phycol.*, 48,
687 539-549, doi: 10.1111/j.1529-8817.2012.01163.x, 2012.

688 Bi, R., Arndt, C., and Sommer, U.: Linking elements to biochemicals: effects of
689 nutrient supply ratios and growth rates on fatty acid composition of phytoplankton
690 species, *J. Phycol.*, 50, 117-130, doi: 10.1111/jpy.12140, 2014.

691 Bi, R., Ismar, S. M. H., Sommer, U., and Zhao, M.: Environmental dependence of the
692 correlations between stoichiometric and fatty acid-based indicators of phytoplankton
693 food quality, *Limnol. Oceanogr.*, 62, 334-347, doi: 10.1002/lno.10429, 2017.

694 Boersma, M., Gruener, N., Signorelli, N. T., Gonzalez, P. E. M., Peck, M. A., and
695 Wiltshire, K. H.: Projecting effects of climate change on marine systems: is the mean



696 all that matters?, P. Roy. Soc. B-Biol. Sci., 283, 20152274, doi:
697 10.1098/rspb.2015.2274, 2016.

698 Bolker, B. M., Brooks, M. E., Clark, C. J., Geange, S. W., Poulsen, J. R., Stevens, M.
699 H. H., and White, J.-S. S.: Generalized linear mixed models: a practical guide for
700 ecology and evolution, Trends Ecol. Evol., 24, 127-135, doi:
701 10.1016/j.tree.2008.10.008, 2009.

702 Borchard, C. and Engel, A.: Organic matter exudation by *Emiliania huxleyi* under
703 simulated future ocean conditions, Biogeosciences, 9, 3405-3423, doi:
704 10.5194/bg-9-3405-2012, 2012.

705 Boyd, P. W., Lennartz, S. T., Glover, D. M., and Doney, S. C.: Biological
706 ramifications of climate-change-mediated oceanic multi-stressors, Nat. Clim. Change,
707 5, 71-79, doi: 10.1038/nclimate2441, 2015.

708 Boyd, P. W., Strzepek, R., Fu, F., and Hutchins, D. A.: Environmental control of
709 open-ocean phytoplankton groups: Now and in the future, Limnol. Oceanogr., 55,
710 1353-1376, doi: 10.4319/lo.2010.55.3.1353, 2010.

711 Charalampopoulou, A., Poulton, A. J., Bakker, D. C. E., Lucas, M. I., Stinchcombe, M.
712 C., and Tyrrell, T.: Environmental drivers of coccolithophore abundance and
713 calcification across Drake Passage (Southern Ocean), Biogeosciences, 13, 5717-5735,
714 doi: 10.5194/bg-13-5917-2016, 2016.

715 Cherif, M. and Loreau, M.: Towards a more biologically realistic use of Droop's
716 equations to model growth under multiple nutrient limitation, Oikos, 119, 897-907,
717 doi: 10.1111/j.1600-0706.2010.18397.x, 2010.

718 Christensen, M. R., Graham, M. D., Vinebrooke, R. D., Findlay, D. L., Paterson, M. J.,
719 and Turner, M. A.: Multiple anthropogenic stressors cause ecological surprises in
720 boreal lakes, Glob. Change Biol., 12, 2316-2322, doi:
721 10.1111/j.1365-2486.2006.01257.x, 2006.

722 Daines, S. J., Clark, J. R., and Lenton, T. M.: Multiple environmental controls on
723 phytoplankton growth strategies determine adaptive responses of the N : P ratio, Ecol.
724 Lett., 17, 414-425, doi: 10.1111/ele.12239, 2014.

725 Dalsgaard, J., St. John, M., Kattner, G., Müller-Navarra, D., and Hagen, W.: Fatty acid
726 trophic markers in the pelagic marine environment, Adv. Mar. Biol., 46, 225-340, doi:
727 10.1016/S0065-2881(03)46005-7, 2003.

728 De Bodt, C., Van Oostende, N., Harlay, J., Sabbe, K., and Chou, L.: Individual and
729 interacting effects of $p\text{CO}_2$ and temperature on *Emiliania huxleyi* calcification: study
730 of the calcite production, the coccolith morphology and the coccospHERE size,
731 Biogeosciences, 7, 1401-1412, doi: 10.5194/bg-7-1401-2010, 2010.



732 Delille, B., Harlay, J., Zondervan, I., Jacquet, S., Chou, L., Wollast, R., Bellerby, R. G.
733 J., Frankignoulle, M., Borges, A. V., Riebesell, U., and Gattuso, J. P.: Response of
734 primary production and calcification to changes of $p\text{CO}_2$ during experimental blooms
735 of the coccolithophorid *Emiliania huxleyi*, Global Biogeochem. Cy., 19, GB2023, doi:
736 10.1029/2004gb002318, 2005.

737 Dickson, A. and Millero, F.: A comparison of the equilibrium constants for the
738 dissociations of carbonic acid in seawater media, Deep-Sea Res., 34, 1733-1741, doi:
739 10.1016/0198-0149(87)90021-5, 1987.

740 Doney, S. C., Ruckelshaus, M., Duffy, J. E., Barry, J. P., Chan, F., English, C. A.,
741 Galindo, H. M., Grebmeier, J. M., Hollowed, A. B., Knowlton, N., Polovina, J.,
742 Rabalais, N. N., Sydeman, W. J., and Talley, L. D.: Climate change impacts on marine
743 ecosystems, Annu. Rev. Mar. Sci., 4, 11-37, doi:
744 10.1146/annurev-marine-041911-111611, 2012.

745 Dutkiewicz, S., Morris, J. J., Follows, M. J., Scott, J., Levitan, O., Dyhrman, S. T.,
746 and Berman-Frank, I.: Impact of ocean acidification on the structure of future
747 phytoplankton communities, Nature Clim. Change, 5, 1002-1006, doi:
748 10.1038/nclimate2722, 2015.

749 Engel, A., Novoa, C. C., Wurst, M., Endres, S., Tang, T. T., Schartau, M., and Lee, C.:
750 No detectable effect of CO_2 on elemental stoichiometry of *Emiliania huxleyi* in
751 nutrient-limited, acclimated continuous cultures, Mar. Ecol. Prog. Ser., 507, 15-30, doi:
752 10.3354/meps10824, 2014.

753 Engel, A., Zondervan, I., Aerts, K., Beaufort, L., Benthien, A., Chou, L., Delille, B.,
754 Gattuso, J. P., Harlay, J., Heemann, C., Hoffmann, L., Jacquet, S., Nejstgaard, J., Pizay,
755 M. D., Rochelle-Newall, E., Schneider, U., Terbrueggen, A., and Riebesell, U.:
756 Testing the direct effect of CO_2 concentration on a bloom of the coccolithophorid
757 *Emiliania huxleyi* in mesocosm experiments, Limnol. Oceanogr., 50, 493-507, doi:
758 10.4319/lo.2005.50.2.0493, 2005.

759 Feng, Y.: Environmental controls on the physiology of the marine coccolithophore
760 *Emiliania huxleyi* strain NIWA 1108, Ph.D. thesis, University of Otago, Dunedin,
761 New Zealand, 194 pp., 2015.

762 Feng, Y., Roleda, M. Y., Armstrong, E., Boyd, P. W., and Hurd, C. L.: Environmental
763 controls on the growth, photosynthetic and calcification rates of a Southern
764 Hemisphere strain of the coccolithophore *Emiliania huxleyi*, Limnol. Oceanogr., 62,
765 519-540, doi: 10.1002/lno.10442, 2017.

766 Feng, Y., Warner, M. E., Zhang, Y., Sun, J., Fu, F.-X., Rose, J. M., and Hutchins, D. A.:
767 Interactive effects of increased $p\text{CO}_2$, temperature and irradiance on the marine
768 coccolithophore *Emiliania huxleyi* (Prymnesiophyceae), Eur. J. Phycol., 43, 87-98,
769 doi: 10.1080/09670260701664674, 2008.



770 Fields, M. W., Hise, A., Lohman, E. J., Bell, T., Gardner, R. D., Corredor, L., Moll, K.,
771 Peyton, B. M., Characklis, G. W., and Gerlach, R.: Sources and resources: importance
772 of nutrients, resource allocation, and ecology in microalgal cultivation for lipid
773 accumulation, *Appl. Microbiol. Biot.*, 98, 4805-4816, doi:
774 10.1007/s00253-014-5694-7, 2014.

775 Fiorini, S., Gattuso, J.-P., van Rijswijk, P., and Middelburg, J.: Coccolithophores lipid
776 and carbon isotope composition and their variability related to changes in seawater
777 carbonate chemistry, *J. Exp. Mar. Biol. Ecol.*, 394, 74-85, doi:
778 10.1016/j.jembe.2010.07.020, 2010.

779 Fuschino, J. R., Guschina, I. A., Dobson, G., Yan, N. D., Harwood, J. L., and Arts, M.
780 T.: Rising water temperatures alter lipid dynamics and reduce N-3 essential fatty acid
781 concentrations in *Scenedesmus obliquus* (Chlorophyta), *J. Phycol.*, 47, 763-774, doi:
782 10.1111/j.1529-8817.2011.01024.x, 2011.

783 Galbraith, E. D. and Martiny, A. C.: A simple nutrient-dependence mechanism for
784 predicting the stoichiometry of marine ecosystems, *Proc. Natl. Acad. Sci. U. S. A.*,
785 112, 8199-8204, doi: 10.1073/pnas.1423917112, 2015.

786 Galloway, A. W. E. and Winder, M.: Partitioning the relative importance of phylogeny
787 and environmental conditions on phytoplankton fatty acids, *Plos One*, 10, e0130053,
788 doi: 10.1371/journal.pone.0130053, 2015.

789 Garzke, J., Hansen, T., Ismar, S. M. H., and Sommer, U.: Combined effects of ocean
790 warming and acidification on copepod abundance, body size and fatty acid content,
791 *Plos One*, 11, e0155952, doi: 10.1371/journal.pone.0155952, 2016.

792 Garzke, J., Sommer, U., and Ismar, S. M. H.: Is the chemical composition of biomass
793 the agent by which ocean acidification influences on zooplankton ecology?, *Aquat.
794 Sci.*, doi: 10.1007/s00027-017-0532-5, 2017.

795 Graham, M. H. and Edwards, M. S.: Statistical significance versus fit: estimating the
796 importance of individual factors in ecological analysis of variance, *Oikos*, 93,
797 505-513, doi: 10.1034/j.1600-0706.2001.930317.x, 2001.

798 Guschina, I. A. and Harwood, J. L.: Mechanisms of temperature adaptation in
799 poikilotherms, *Febs Lett.*, 580, 5477-5483, doi: 10.1016/j.febslet.2006.06.066, 2006.

800 Hansen, H. P. and Koroleff, F.: Determination of nutrients, in: *Methods of Seawater
801 Analysis*, Grasshoff, K., Kremling, K., and Ehrhardt, M. (Eds.), WILEY-VCH,
802 Weinheim, Germany, 159–228, 1999.

803 Hansen, T., Gardeler, B., and Matthiessen, B.: Technical Note: Precise quantitative
804 measurements of total dissolved inorganic carbon from small amounts of seawater
805 using a gas chromatographic system, *Biogeosciences*, 10, 6601-6608, doi:



806 10.5194/bg-10-6601-2013, 2013.

807 Hansson, I.: A new set of acidity constants for carbonic acid and boric acid in
808 seawater, Deep-Sea Res., 20, 661-678, doi: 10.1016/0011-7471(73)90100-9, 1973.

809 Hixson, S. M. and Arts, M. T.: Climate warming is predicted to reduce omega-3,
810 long-chain, polyunsaturated fatty acid production in phytoplankton, Glob. Change
811 Biol., 22, 2744-2755, doi: 10.1111/gcb.13295, 2016.

812 Hu, Q., Sommerfeld, M., Jarvis, E., Ghirardi, M., Posewitz, M., Seibert, M., and
813 Darzins, A.: Microalgal triacylglycerols as feedstocks for biofuel production:
814 perspectives and advances, Plant J., 54, 621-639, doi, 2008.

815 Hughes, A. R. and Stachowicz, J. J.: Ecological impacts of genotypic diversity in the
816 clonal seagrass *Zostera marina*, Ecology, 90, 1412-1419, doi: 10.1890/07-2030.1,
817 2009.

818 Hutchins, D. A., Mulholland, M. R., and Fu, F.: Nutrient cycles and marine microbes
819 in a CO₂-enriched ocean, Oceanography, 22, 128-145, doi: 10.5670/oceanog.2009.103,
820 2009.

821 IPCC: Climate change 2014: Synthesis report. Contribution of working groups I, II
822 and III to the fifth assessment report of the intergovernmental panel on climate change,
823 IPCC, Geneva, Switzerland, 2014.

824

825 Ismar, S. M. H., Hansen, T., and Sommer, U.: Effect of food concentration and type of
826 diet on *Acartia* survival and naupliar development, Mar. Biol., 154, 335-343, doi:
827 10.1007/s00227-008-0928-9, 2008.

828 Krumhardt, K. M., Lovenduski, N. S., Freeman, N. M., and Bates, N. R.: Apparent
829 increase in coccolithophore abundance in the subtropical North Atlantic from 1990 to
830 2014, Biogeosciences, 13, 1163-1177, doi: 10.5194/bg-13-1163-2016, 2016.

831 Langer, G., Oetjen, K., and Brenneis, T.: Coccolithophores do not increase particulate
832 carbon production under nutrient limitation: A case study using *Emiliania huxleyi*
833 (PML B92/11), J. Exp. Mar. Biol. Ecol., 443, 155-161, doi:
834 10.1016/j.jembe.2013.02.040, 2013.

835 Lefebvre, S. C., Benner, I., Stillman, J. H., Parker, A. E., Drake, M. K., Rossignol, P.
836 E., Okimura, K. M., Komada, T., and Carpenter, E. J.: Nitrogen source and *pCO₂*
837 synergistically affect carbon allocation, growth and morphology of the
838 coccolithophore *Emiliania huxleyi*: potential implications of ocean acidification for
839 the carbon cycle, Glob. Change Biol., 18, 493-503, doi:
840 10.1111/j.1365-2486.2011.02575.x, 2012.

841 Leonardos, N. and Geider, R. J.: Elemental and biochemical composition of



842 *Rhinomonas reticulata* (Cryptophyta) in relation to light and nitrate-to-phosphate
843 supply ratios, *J. Phycol.*, 41, 567-576, doi: 10.1111/j.1529-8817.2005.00082.x, 2005a.

844 Leonardos, N. and Geider, R. J.: Elevated atmospheric carbon dioxide increases
845 organic carbon fixation by *Emiliana huxleyi* (Haptophyta), under nutrient-limited
846 high-light conditions, *J. Phycol.*, 41, 1196-1203, doi:
847 10.1111/j.1529-8817.2005.00152.x, 2005b.

848 Leu, E., Daase, M., Schulz, K. G., Stuhr, A., and Riebesell, U.: Effect of ocean
849 acidification on the fatty acid composition of a natural plankton community,
850 *Biogeosciences*, 10, 1143-1153, doi: 10.5194/bg-10-1143-2013, 2013.

851 Lewandowska, A. M., Boyce, D. G., Hofmann, M., Matthiessen, B., Sommer, U., and
852 Worm, B.: Effects of sea surface warming on marine plankton, *Ecol. Lett.*, 17,
853 614-623, doi: 10.1111/ele.12265, 2014.

854 Martiny, A. C., Pham, C. T. A., Primeau, F. W., Vrugt, J. A., Moore, J. K., Levin, S. A.,
855 and Lomas, M. W.: Strong latitudinal patterns in the elemental ratios of marine
856 plankton and organic matter, *Nat. Geosci.*, 6, 279-283, doi: 10.1038/ngeo1757, 2013.

857 Martiny, A. C., Talarmin, A., Mouginot, C., Lee, J. A., Huang, J. S., Gellene, A. G.,
858 and Caron, D. A.: Biogeochemical interactions control a temporal succession in the
859 elemental composition of marine communities, *Limnol. Oceanogr.*, 61, 531-542, doi:
860 10.1002/lno.10233, 2016.

861 Matson, P. G., Ladd, T. M., Halewood, E. R., Sangodkar, R. P., Chmelka, B. F., and
862 Iglesias-Rodriguez, D.: Intraspecific differences in biogeochemical responses to
863 thermal change in the coccolithophore *Emiliana huxleyi*, *Plos One*, 11, e0162313, doi:
864 10.1371/journal.pone.0162313, 2016.

865 Matthiessen, B., Eggers, S. L., and Krug, S. A.: High nitrate to phosphorus regime
866 attenuates negative effects of rising $p\text{CO}_2$ on total population carbon accumulation,
867 *Biogeosciences*, 9, 1195-1203, doi: 10.5194/bg-9-1195-2012, 2012.

868 Mehrbach, C., Culberson, C., Hawley, J., and Pytkowicz, R.: Measurement of the
869 apparent dissociation constants of carbonic acid in seawater at atmospheric pressure,
870 *Limnol. Oceanogr.*, 18, 897-907, doi: 10.4319/lo.1973.18.6.0897, 1973.

871 Meyer, J. and Riebesell, U.: Reviews and Syntheses: Responses of coccolithophores
872 to ocean acidification: a meta-analysis, *Biogeosciences*, 12, 1671-1682, doi:
873 10.5194/bg-12-1671-2015, 2015.

874 Milner, S., Langer, G., Grelaud, M., and Ziveri, P.: Ocean warming modulates the
875 effects of acidification on *Emiliana huxleyi* calcification and sinking, *Limnol.*
876 *Oceanogr.*, 61, 1322-1336, doi: 10.1002/lno.10292, 2016.

877 Moore, C. M., Mills, M. M., Arrigo, K. R., Berman-Frank, I., Bopp, L., Boyd, P. W.,



878 Galbraith, E. D., Geider, R. J., Guieu, C., Jaccard, S. L., Jickells, T. D., La Roche, J.,
879 Lenton, T. M., Mahowald, N. M., Maranon, E., Marinov, I., Moore, J. K., Nakatsuka,
880 T., Oschlies, A., Saito, M. A., Thingstad, T. F., Tsuda, A., and Ulloa, O.: Processes and
881 patterns of oceanic nutrient limitation, *Nat. Geosci.*, 6, 701-710, doi:
882 10.1038/ngeo1765, 2013.

883 Nanninga, H. J. and Tyrrell, T.: Importance of light for the formation of algal blooms
884 by *Emiliania huxleyi*, *Mar. Ecol. Prog. Ser.*, 136, 195-203, doi: 10.3354/meps136195,
885 1996.

886 Olson, M. B., Wuori, T. A., Love, B. A., and Strom, S. L.: Ocean acidification effects
887 on haploid and diploid *Emiliania huxleyi* strains: Why changes in cell size matter, *J. Exp. Mar. Biol. Ecol.*, 488, 72-82, doi: 10.1016/j.jembe.2016.12.008, 2017.

889 Oviedo, A. M., Langer, G., and Ziveri, P.: Effect of phosphorus limitation on coccolith
890 morphology and element ratios in Mediterranean strains of the coccolithophore
891 *Emiliania huxleyi*, *J. Exp. Mar. Biol. Ecol.*, 459, 105-113, doi:
892 10.1016/j.jembe.2014.04.021, 2014.

893 Paasche, E.: Roles of nitrogen and phosphorus in coccolith formation in *Emiliania*
894 *huxleyi* (Prymnesiophyceae), *Eur. J. Phycol.*, 33, 33-42, doi:
895 10.1017/s0967026297001480, 1998.

896 Paasche, E.: A review of the coccolithophorid *Emiliania huxleyi* (Prymnesiophyceae),
897 with particular reference to growth, coccolith formation, and
898 calcification-photosynthesis interactions, *Phycologia*, 40, 503-529, doi:
899 10.2216/i0031-8884-40-6-503.1, 2002.

900 Peñuelas, J., Sardans, J., Rivas-ubach, A., and Janssens, I. A.: The human-induced
901 imbalance between C, N and P in Earth's life system, *Glob. Change Biol.*, 18, 3-6, doi:
902 10.1111/j.1365-2486.2011.02568.x, 2012.

903 Perrin, L., Probert, I., Langer, G., and Aloisi, G.: Growth of the coccolithophore
904 *Emiliania huxleyi* in light- and nutrient-limited batch reactors: relevance for the
905 BIOSOPE deep ecological niche of coccolithophores, *Biogeosciences*, 13, 5983-6001,
906 doi: 10.5194/bg-13-5983-2016, 2016.

907 Piepho, M., Arts, M. T., and Wacker, A.: Species-specific variation in fatty acid
908 concentrations of four phytoplankton species: does phosphorus supply influence the
909 effect of light intensity or temperature?, *J. Phycol.*, 48, 64-73, doi:
910 10.1111/j.1529-8817.2011.01103.x, 2012.

911 Pierrot, D., Lewis, E., and Wallace, D.: MS Excel program developed for CO₂ system
912 calculations: ORNL/CDIAC-105a. Carbon Dioxide Information Analysis Centre, Oak
913 Ridge National Laboratory, US Department of Energy, Oak Ridge, TN, 2006.



914 Pronina, N. A., Rogova, N. B., Furnadzhieva, S., and Klyachko-Gurvich, G. L.: Effect
915 of CO₂ concentration on the fatty acid composition of lipids in *Chlamydomonas*
916 *reinhardtii* cia-3, a mutant deficient in CO₂-concentrating mechanism, Russ. J. Plant
917 Physiol., 45, 447-455, 1998.

918 Provasoli, L.: Growing marine seaweeds., in: Proc. 4th Internatl. Seaweed Symp., De
919 Virville, A. D. and Feldmann, J. (Eds.), Pergamon Press, Oxford, UK, 9-17, 1963.

920 Raitsos, D. E., Lavender, S. J., Pradhan, Y., Tyrrell, T., Reid, P. C., and Edwards, M.:
921 Coccolithophore bloom size variation in response to the regional environment of the
922 subarctic North Atlantic, Limnol. Oceanogr., 51, 2122-2130, doi:
923 10.4319/lo.2006.51.5.2122, 2006.

924 Read, B. A., Kegel, J., Klute, M. J., Kuo, A., Lefebvre, S. C., Maumus, F., Mayer, C.,
925 Miller, J., Monier, A., Salamov, A., Young, J., Aguilar, M., Claverie, J. M.,
926 Frickenhaus, S., Gonzalez, K., Herman, E. K., Lin, Y. C., Napier, J., Ogata, H., Sarno,
927 A. F., Shmutz, J., Schroeder, D., de Vargas, C., Verret, F., von Dassow, P., Valentin, K.,
928 Van de Peer, Y., Wheeler, G., Allen, A. E., Bidle, K., Borodovsky, M., Bowler, C.,
929 Brownlee, C., Cock, J. M., Elias, M., Gladyshev, V. N., Groth, M., Guda, C., Hadaegh,
930 A., Iglesias-Rodriguez, M. D., Jenkins, J., Jones, B. M., Lawson, T., Leese, F.,
931 Lindquist, E., Lobanov, A., Lomsadze, A., Malik, S. B., Marsh, M. E., Mackinder, L.,
932 Mock, T., Mueller-Roeber, B., Pagarete, A., Parker, M., Probert, I., Quesneville, H.,
933 Raines, C., Rensing, S. A., Riano-Pachon, D. M., Richier, S., Rokitta, S., Shiraiwa, Y.,
934 Soanes, D. M., van der Giezen, M., Wahlund, T. M., Williams, B., Wilson, W., Wolfe,
935 G., Wurch, L. L., Dacks, J. B., Delwiche, C. F., Dyhrman, S. T., Gloeckner, G., John,
936 U., Richards, T., Worden, A. Z., Zhang, X. Y., and Grigoriev, I. V.: Pan genome of the
937 phytoplankton *Emiliania* underpins its global distribution, Nature, 499, 209-213, doi:
938 10.1038/nature12221, 2013.

939 Renaud, S. M., Thinh, L.-V., Lambrinidis, G., and Parry, D. L.: Effect of temperature
940 on growth, chemical composition and fatty acid composition of tropical Australian
941 microalgae grown in batch cultures, Aquaculture, 211, 195-214, doi:
942 10.1016/S0044-8486(01)00875-4, 2002.

943 Riebesell, U., Revill, A. T., Holdsworth, D. G., and Volkman, J. K.: The effects of
944 varying CO₂ concentration on lipid composition and carbon isotope fractionation in
945 *Emiliania huxleyi*, Geochim. Cosmochim. Ac., 64, 4179-4192, doi:
946 10.1016/s0016-7037(00)00474-9, 2000.

947 Rokitta, S. D. and Rost, B.: Effects of CO₂ and their modulation by light in the
948 life-cycle stages of the coccolithophore *Emiliania huxleyi*, Limnol. Oceanogr., 57,
949 607-618, doi: 10.4319/lo.2012.57.2.0607, 2012.

950 Rosas-Navarro, A., Langer, G., and Ziveri, P.: Temperature affects the morphology
951 and calcification of *Emiliania huxleyi* strains, Biogeosciences, 13, 2913-2926, doi:
952 10.5194/bg-13-2913-2016, 2016.



953 Rosenblatt, A. E. and Schmitz, O. J.: Climate change, nutrition, and bottom-up and
954 top-down food web processes, *Trends Ecol. Evol.*, 31, 965-975, doi:
955 10.1016/j.tree.2016.09.009, 2016.

956 Rossoll, D., Bermúdez, R., Hauss, H., Schulz, K. G., Riebesell, U., Sommer, U., and
957 Winder, M.: Ocean acidification-induced food quality deterioration constrains trophic
958 transfer, *Plos One*, 7, e34737, doi: 10.1371/journal.pone.0034737, 2012.

959 Rost, B. and Riebesell, U.: Coccolithophores and the biological pump: responses to
960 environmental changes, in: *Coccolithophores: From molecular processes to global*
961 *impact*, Thierstein, H. R. and Young, J. R. (Eds.), Springer, Heidelberg, Germany,
962 99-125, 2004.

963 Sato, N., Tsuzuki, M., and Kawaguchi, A.: Glycerolipid synthesis in *Chlorella kessleri*
964 11h - II. Effect of the CO₂ concentration during growth, *BBA-Mol. Cell Biol. L.*, 1633,
965 35-42, doi: 10.1016/s1388-1981(03)00070-2, 2003.

966 Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.:
967 Temperature modulates coccolithophorid sensitivity of growth, photosynthesis and
968 calcification to increasing seawater pCO₂, *PLoS ONE*, 9, e88308, doi:
969 10.1371/journal.pone.0088308, 2014.

970 Sharp, J.: Improved analysis for particulate organic carbon and nitrogen from
971 seawater., *Limnol. Oceanogr.*, 19, 984-989, doi: 10.4319/lo.1974.19.6.0984, 1974.

972 Sinensky, M.: Homeoviscous adaptation - a homeostatic process that regulates the
973 viscosity of membrane lipids in *Escherichia coli*, *Proc. Natl. Acad. Sci. U. S. A.*, 71,
974 522-525, doi: 10.1073/pnas.71.2.522, 1974.

975 Skau, L. F.: Effects of temperature and phosphorus on growth, stoichiometry and size
976 in three haptophytes, M.S. thesis, Centre for Ecological and Evolutionary Synthesis
977 (CEES), Section for Aquatic Biology and Toxicology (AQUA), University of Oslo,
978 Oslo, Norway, 64 pp., 2015.

979 Sorrosa, J. M., Satoh, M., and Shiraiwa, Y.: Low temperature stimulates cell
980 enlargement and intracellular calcification of Coccolithophorids, *Mar. Biotechnol.*, 7,
981 128-133, doi: 10.1007/s10126-004-0478-1, 2005.

982 Sterner, R. W. and Elser, J. J.: *Ecological stoichiometry: The biology of elements from*
983 *molecules to the biosphere*, Princeton University Press, Princeton, U.S.A., 2002.

984 Thompson, G. A.: Lipids and membrane function in green algae, *BBA-Lipid Lipid*
985 *Met.*, 1302, 17-45, doi: 10.1016/0005-2760(96)00045-8, 1996.

986 Thompson, P. A., Guo, M.-x., Harrison, P. J., and Whyte, J. N. C.: Effects of variation
987 in temperature. II. On the fatty acid composition of eight species of marine
988 phytoplankton, *J. Phycol.*, 28, 488-497, doi: 10.1111/j.0022-3646.1992.00488.x, 1992.



989 Toseland, A., Daines, S. J., Clark, J. R., Kirkham, A., Strauss, J., Uhlig, C., Lenton, T.
990 M., Valentin, K., Pearson, G. A., Moulton, V., and Mock, T.: The impact of
991 temperature on marine phytoplankton resource allocation and metabolism, *Nat. Clim. Change*, 3, 979-984, doi: 10.1038/nclimate1989, 2013.

993 Tyrrell, T. and Merico, A.: *Emiliania huxleyi*: bloom observations and the conditions
994 that induce them, in: *Coccolithophores: From molecular processes to global impact*,
995 Thierstein, H. R. and Young, J. R. (Eds.), Springer, Heidelberg, Germany, 75-97,
996 2004.

997 van Bleijswijk, J. D. L., Kempers, R. S., Veldhuis, M. J., and Westbroek, P.: Cell and
998 growth characteristics of types A and B of *Emiliania huxleyi* (Prymnesiophyceae) as
999 determined by flow cytometry and chemical analyses, *J. Phycol.*, 30, 230-241, doi:
1000 10.1111/j.0022-3646.1994.00230.x, 1994.

1001 Van Mooy, B. A. S., Fredricks, H. F., Pedler, B. E., Dyhrman, S. T., Karl, D. M.,
1002 Koblizek, M., Lomas, M. W., Mincer, T. J., Moore, L. R., Moutin, T., Rappe, M. S.,
1003 and Webb, E. A.: Phytoplankton in the ocean use non-phosphorus lipids in response to
1004 phosphorus scarcity, *Nature*, 458, 69-72, doi: 10.1038/nature07659, 2009.

1005 Verspagen, J. M. H., Van de Waal, D. B., Finke, J. F., Visser, P. M., and Huisman, J.:
1006 Contrasting effects of rising CO₂ on primary production and ecological stoichiometry
1007 at different nutrient levels, *Ecol. Lett.*, 17, 951-960, doi: 10.1111/ele.12298, 2014.

1008 Winter, A., Henderiks, J., Beaufort, L., Rickaby, R. E. M., and Brown, C. W.:
1009 Poleward expansion of the coccolithophore *Emiliania huxleyi*, *J. Plankton Res.*, 36,
1010 316-325, doi: 10.1093/plankt/fbt110, 2014.

1011 Xing, T., Gao, K., and Beardall, J.: Response of growth and photosynthesis of
1012 *Emiliania huxleyi* to visible and UV irradiances under different light regimes,
1013 *Photochem. Photobiol.*, 91, 343-349, doi: 10.1111/php.12403, 2015.

1014 Yvon-Durocher, G., Dossena, M., Trimmer, M., Woodward, G., and Allen, A. P.:
1015 Temperature and the biogeography of algal stoichiometry, *Global Ecol. Biogeogr.*, 24,
1016 562-570, doi: 10.1111/geb.12280, 2015.

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1023 Table 1. Results of ANOVA testing for the responses of the observed maximum
1024 growth rate to temperature and $p\text{CO}_2$ in *Emiliania huxleyi*. Significant p values are
1025 shown in bold.

Factor	df	F	p	ω^2
Temperature	2	89.842	<0.001	0.82
$p\text{CO}_2$	1	3.808	0.075	
Temperature \times $p\text{CO}_2$	2	10.638	0.002	0.09

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1044 Table 2. Results of the selected GLMMs testing for the effects of temperature, N:P
 1045 supply ratios and $p\text{CO}_2$ on stoichiometric C:N:P biomass ratios, PIC and POC
 1046 contents and their ratios, PIC and POC production, and fatty acid proportions in
 1047 *Emiliania huxleyi*. Significant p values are shown in bold; T: temperature; N:P: N:P
 1048 supply ratio; TFA: total fatty acid; SFA: saturated fatty acid; MUFA: monounsaturated
 1049 fatty acid; PUFA: polyunsaturated fatty acid; DHA: docosahexaenoic acid (22:6n-3).

Variable	Factor	Coefficient \pm SE	t	p
N:C biomass ratio (mol mol ⁻¹)	Intercept	0.062 \pm 0.007	8.659	<0.001
	T	0.001 \pm <0.001	2.641	0.011
	$p\text{CO}_2$	<0.001 \pm <0.001	-0.240	0.811
	N:P	<0.001 \pm <0.001	4.717	<0.001
P:C biomass ratio (mmol mol ⁻¹)	Intercept	4.481 \pm 0.451	9.946	<0.001
	T	0.016 \pm 0.021	0.773	0.443
	$p\text{CO}_2$	<0.001 \pm <0.001	0.628	0.553
	N:P	-0.037 \pm 0.005	-8.154	<0.001
N:P biomass ratio (mol mol ⁻¹)	Intercept	2.702 \pm 0.145	18.590	<0.001
	T	0.001 \pm 0.007	0.157	0.876
	$p\text{CO}_2$	<0.001 \pm <0.001	-0.169	0.866
	N:P	0.016 \pm 0.001	11.200	<0.001
PIC (pg cell ⁻¹)	Intercept	3.293 \pm 0.406	8.122	<0.001
	T	-0.067 \pm 0.021	-3.193	0.003
	$p\text{CO}_2$	-0.001 \pm <0.001	-5.519	<0.001
	N:P	-0.003 \pm 0.009	-0.292	0.772
	T \times $p\text{CO}_2$	<0.001 \pm <0.001	4.584	<0.001
	T \times N:P	0.001 \pm <0.001	2.340	0.024
	$p\text{CO}_2$ \times N:P	<0.001 \pm <0.001	0.111	0.912
PIC (μg ml ⁻¹)	Intercept	6.922 \pm 0.968	7.149	<0.001
	T	0.201 \pm 0.045	4.442	<0.001
	$p\text{CO}_2$	-0.002 \pm <0.001	-8.955	<0.001
	N:P	-0.034 \pm 0.010	-3.404	0.001
PIC production (pg cell ⁻¹ d ⁻¹)	Intercept	-0.689 \pm 0.105	-6.581	<0.001
	T	0.047 \pm 0.005	9.589	<0.001
	$p\text{CO}_2$	<0.001 \pm <0.001	-5.294	<0.001
	N:P	0.007 \pm 0.001	6.339	<0.001
POC (pg cell ⁻¹)	Intercept	3.683 \pm 0.377	9.779	<0.001
	T	-0.089 \pm 0.020	-4.577	<0.001
	$p\text{CO}_2$	<0.001 \pm <0.001	-0.929	0.358
	N:P	-0.008 \pm 0.008	-0.996	0.324



	T \times $p\text{CO}_2$	<0.001 \pm <0.001	1.886	0.066
	T \times N:P	0.001 \pm <0.001	3.477	0.001
	$p\text{CO}_2$ \times N:P	<0.001 \pm <0.001	-0.359	0.721
POC ($\mu\text{g ml}^{-1}$)	Intercept	13.456 \pm 1.007	13.360	<0.001
	T	-0.096 \pm 0.047	-2.045	0.046
	$p\text{CO}_2$	<0.001 \pm <0.001	-0.361	0.719
	N:P	-0.035 \pm 0.010	-3.436	0.001
POC production ($\text{pg cell}^{-1} \text{ d}^{-1}$)	Intercept	-0.261 \pm 0.101	-2.587	0.013
	T	0.023 \pm 0.005	4.895	<0.001
	$p\text{CO}_2$	<0.001 \pm <0.001	1.631	0.109
	N:P	0.007 \pm 0.001	6.899	<0.001
PIC:POC	Intercept	0.460 \pm 0.066	7.010	<0.001
	T	0.025 \pm 0.003	8.184	<0.001
	$p\text{CO}_2$	<0.001 \pm <0.001	-12.837	<0.001
	N:P	<0.001 \pm 0.001	-0.166	0.869
SFA proportion (% of TFAs)	Intercept	3.506 \pm 0.145	24.178	<0.001
	T	-0.012 \pm 0.008	-1.538	0.131
	$p\text{CO}_2$	<0.001 \pm <0.001	-0.238	0.813
	N:P	-0.004 \pm 0.003	-1.248	0.218
	T \times $p\text{CO}_2$	<0.001 \pm <0.001	1.816	0.076
	T \times N:P	<0.001 \pm <0.001	1.657	0.104
	$p\text{CO}_2$ \times N:P	<0.001 \pm <0.001	-2.487	0.016
MUFA proportion (% of TFAs)	Intercept	30.259 \pm 1.344	22.518	<0.001
	T	-0.579 \pm 0.063	-9.240	<0.001
	$p\text{CO}_2$	0.001 \pm <0.001	2.269	0.028
	N:P	-0.014 \pm 0.014	-1.050	0.299
PUFA proportion (% of TFAs)	Intercept	32.264 \pm 2.300	14.028	<0.001
	T	0.638 \pm 0.107	5.949	<0.001
	$p\text{CO}_2$	-0.002 \pm 0.001	-2.769	0.008
	N:P	0.034 \pm 0.023	1.453	0.152
DHA proportion (% of TFAs)	Intercept	2.204 \pm 0.185	11.887	<0.001
	T	0.054 \pm 0.010	5.611	<0.001
	$p\text{CO}_2$	<0.001 \pm <0.001	1.874	0.067
	N:P	0.010 \pm 0.004	2.735	0.009
	T \times $p\text{CO}_2$	<0.001 \pm <0.001	-2.946	0.005
	T \times N:P	-0.001 \pm <0.001	-2.898	0.006
	$p\text{CO}_2$ \times N:P	<0.001 \pm <0.001	1.249	0.218

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1055 **Fig. 1** Responses of the observed maximal growth rate (μ_{\max} ; mean \pm SE) to
1056 temperature and $p\text{CO}_2$ in *Emiliania huxleyi*.

1057 **Fig. 2** Responses of (a) N:C, (b) P:C and (c) N:P biomass ratios (mean \pm SE) to
1058 temperature and N:P supply ratios in *Emiliania huxleyi*.

1059 **Fig. 3** Responses of PIC content (a, d) per cell and (b, e) per ml and (c, f) PIC
1060 production (mean \pm SE) to temperature, N:P supply ratios and $p\text{CO}_2$ in *Emiliania*
1061 *huxleyi*.

1062 **Fig. 4** Responses of POC content (a, d) per cell and (b, e) per ml and (c, f) POC
1063 production (mean \pm SE) to temperature, N:P supply ratios and $p\text{CO}_2$ in *Emiliania*
1064 *huxleyi*.

1065 **Fig. 5** Responses of PIC:POC ratio (mean \pm SE) to temperature, N:P supply ratios and
1066 $p\text{CO}_2$ in *Emiliania huxleyi*.

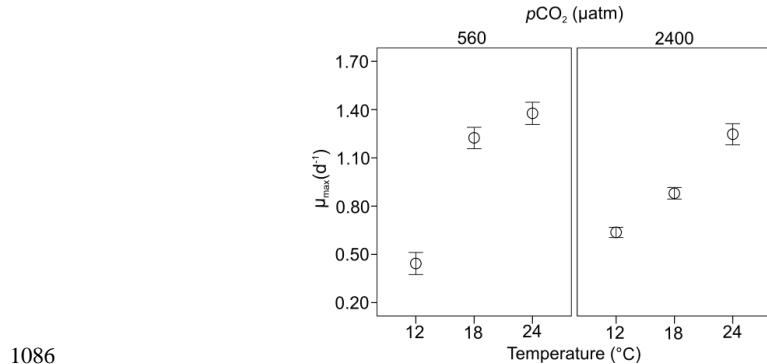
1067 **Fig. 6** Responses of the proportions of (a, d) monounsaturated fatty acids (MUFA),
1068 (b, e) polyunsaturated fatty acids (PUFA) and (c, f) docosahexaenoic acid (DHA)
1069 (mean \pm SE) to temperature, N:P supply ratios and $p\text{CO}_2$ in *Emiliania huxleyi*.

1070 **Fig. 7** The changes in C:N:P stoichiometry, PIC and POC contents and their ratios,
1071 PIC and POC production, and the proportions and contents of major fatty acid groups
1072 and DHA in response to warming, N and P deficiency and enhanced $p\text{CO}_2$ in
1073 *Emiliania huxleyi*. Here, not only significant effects are depicted, but also
1074 non-significant and substantial effects on response variables. Significant interactions
1075 are presented based on GLMM results in Table 2. Red and blue arrows indicate a
1076 mean percent increase and decrease in a given response, respectively.

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1085 **Fig. 1**



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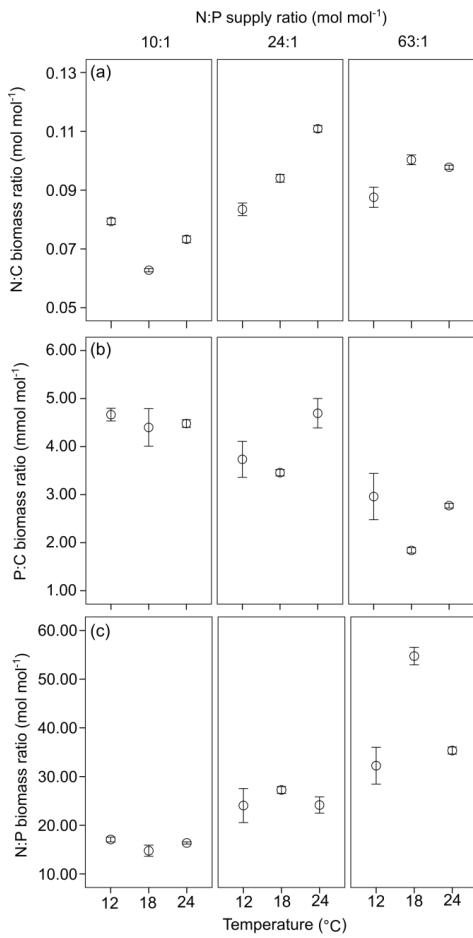
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1103 **Fig. 2**



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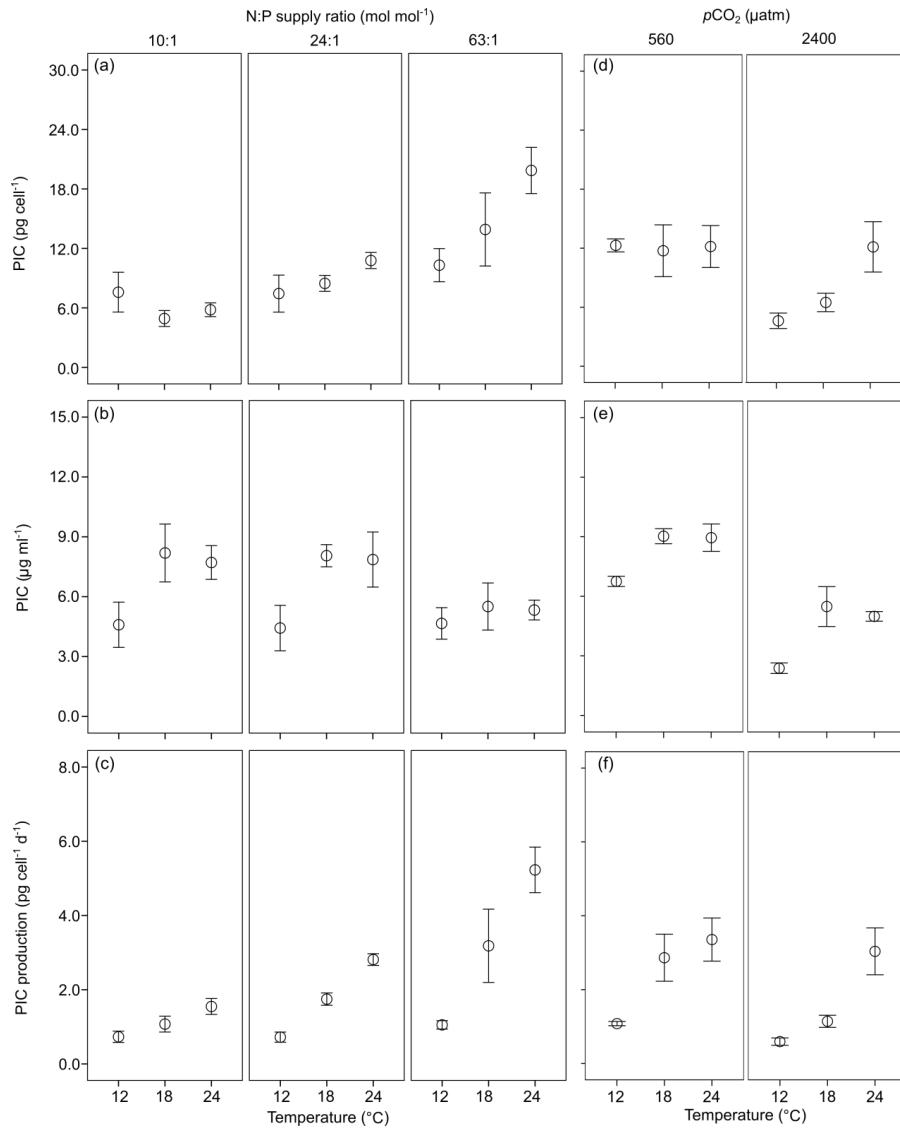
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1112 **Fig. 3**



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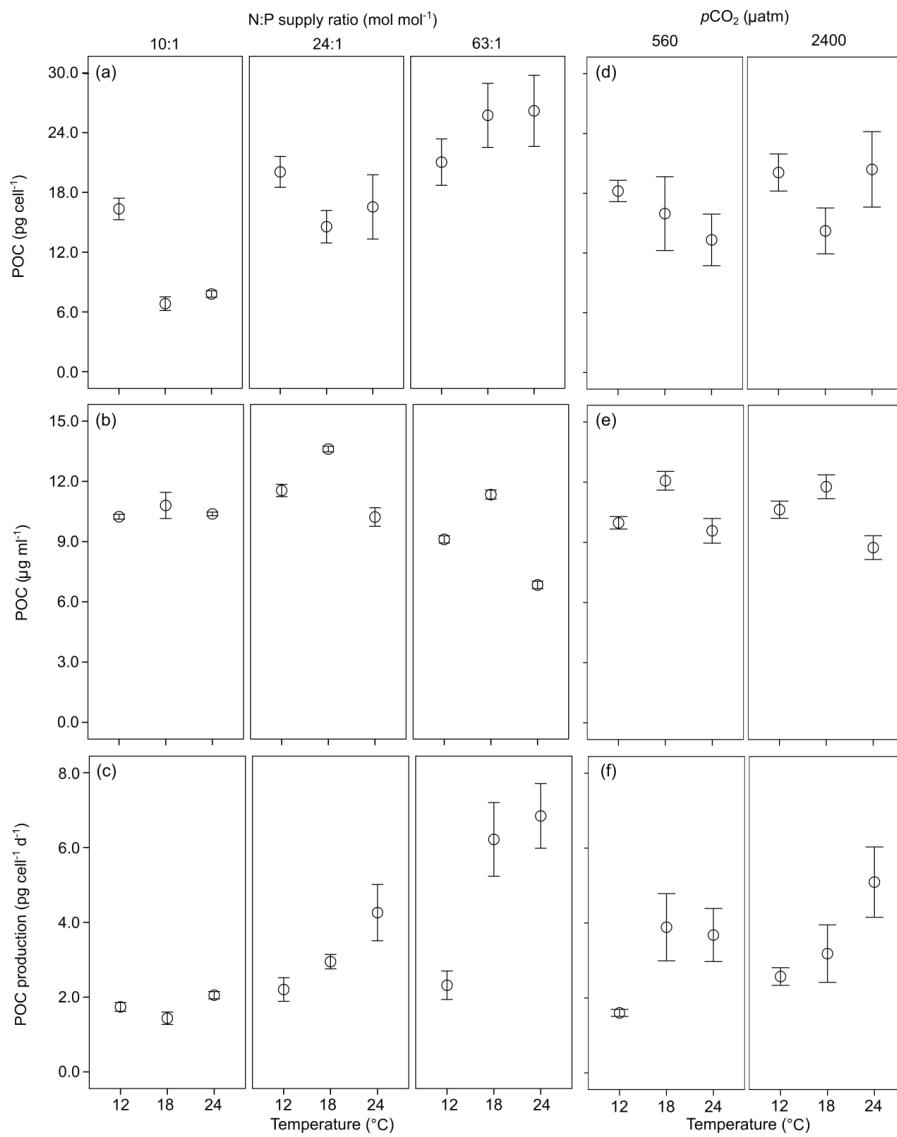
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1118 **Fig. 4**



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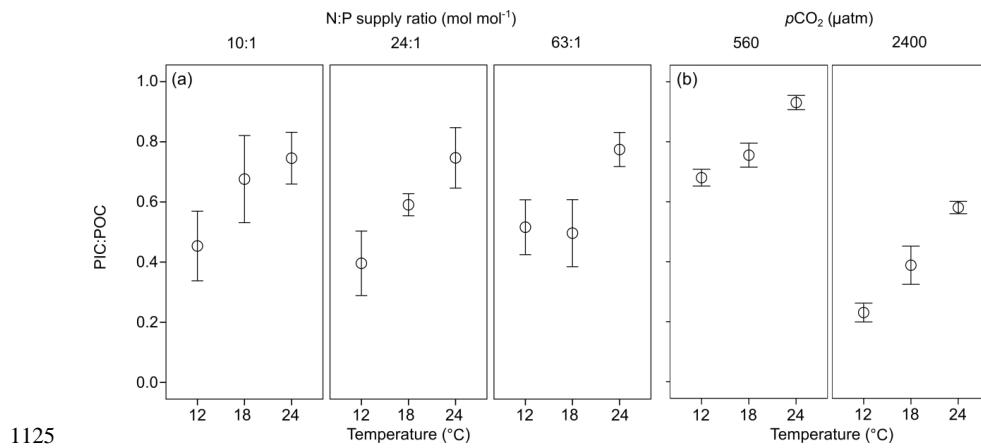
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1124 **Fig. 5**



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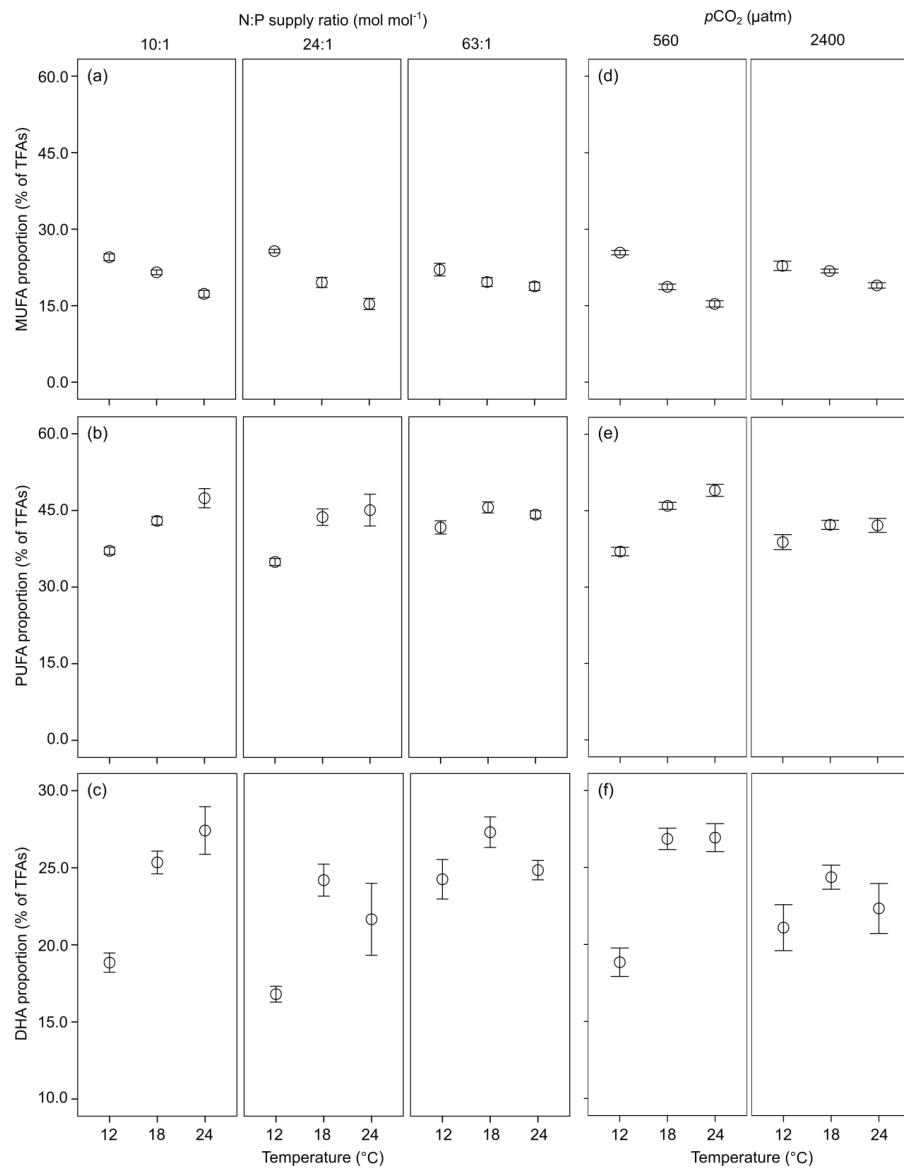
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1140 **Fig. 6**



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1145 **Fig. 7**

Response	Effect				
	Warming	-N	-P	Enhanced $p\text{CO}_2$	Interactions
N:C biomass ratio	↑ 8%	↓ -25%	-	-	
P:C biomass ratio	↑ 5%	↑ 14%	↓ -36%	-	
N:P biomass ratio	↑ 5%	↓ -36%	↑ 62%	-	
PIC content (per cell)	↑ 28%	↓ -31%	↑ 65%	↓ -36%	T×N:P supply T× CO_2
PIC content (per ml)	↑ 36%	↓ -48%	↑ 1%	↓ -24%	
PIC production	↑ 161%	↓ -37%	↑ 79%	↓ -35%	
POC content (per cell)	↓ -8%	↓ -39%	↑ 50%	-	T×N:P supply
POC content (per ml)	↓ -6%	↓ -11%	↓ -23%	-	
POC production	↑ 68%	↓ -44%	↑ 63%	-	
PIC:POC ratio	↑ 41%	-	-	↓ -49%	
SFAs (% of TFAs)	↑ 5%	↓ -7%	↓ -15%	↑ 7%	N:P supply× CO_2
MUFAs (% of TFAs)	↓ -20%	-	-	↑ 7%	
PUFAs (% of TFAs)	↑ 13%	-	-	↓ -7%	
DHA (% of TFAs)	↑ 16%	↑ 14%	↑ 22%	↓ -7%	T×N:P supply T× CO_2
SFAs (per biomass)	↓ -21%	↑ 36%	↑ 11%	↓ -14%	
MUFAs (per biomass)	↓ -35%	↑ 56%	↑ 35%	↓ -11%	
PUFAs (per biomass)	↓ -9%	↑ 44%	↑ 29%	↓ -24%	
DHA (per biomass)	↓ -5%	↑ 53%	↑ 42%	↓ -24%	

1146 Changes $\geq 25\%$ Changes $< 25\%$ - No clear changes

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