

Dear Associate Editor and Reviewers,

please find below a detailed response to all Reviewers comments and questions regarding our manuscript. Since the Reviewers gave us a lot of useful hints and comments, we decided to completely rewrite the Discussion and Conclusions. We also significantly revised method, results (including figures) and supplementary materials to improve our manuscript. We hope that the changes we made will increase the quality of our manuscript in order to fulfill the requirements for publication in Biogeosciences.

Sincerely,

Petr Kotas and co-authors

Reviewer 1

General comments

The authors investigate the effect of horizontal (across a valley) and vertical (altitude) gradients on microbial community structure (PLFA), biomass and activity in High Arctic. They found that both gradient affect microbial parameters, with shift in the dominance of bacteria and fungi related to the chemistry of the bedrock. The study target interesting question, is relevant for publication in Biogeosciences and is overall well done. My main criticisms are the method used to measure microbial activity (main issue), and too many assumptions made outside the variables measured, going beyond what the results can show.

The measure of microbial activity seems unrealistic. First, 2 mm soil was used which was frozen and thaw prior to incubation in the lab. So, the microbial community and soil endure 1 freeze-thaw cycle + sieving that will affect OM availability and the microbial community. Then you left the samples for 14 days at 6 degrees before measuring the CO₂ for 24h, which you define as “basal respiration”. Fourteen days represent 1/4 of the summer (> 5 degrees) in the Arctic (low altitude) or even your entire summer for the high-altitude site (Table 1), this is a significant amount of time in the Arctic. There is no justification and references used to explain why you made these choices. Overall, we can doubt that the high altitude produce high CO₂ emissions in in-situ conditions and we can ask the values of your results regarding microbial activity. You did not discuss at any time the limitations of such measurement. We can imagine that the microbial community adapted better or took longer to adapt to incubation condition in high altitude soil explaining the higher CO₂ emissions at 14 days. We can also imagine that at low altitude, because of the higher TOC, the CO₂ emissions are high rapidly after thawing and after 14 there is not much activity, while for high altitude it took longer to mineralize more complex OM. In other words, your results of microbial activity could be just the results of your incubation/sample preparation. You need to fully acknowledge this in the article, and avoid any conclusion stating that high altitude is a hot spot of microbial activity because your data can't fully support this. You need to be much more conscious about the microbial activity result. Have you measure CO₂ emissions over time?

Author response: We admit that it was inappropriate to call the measured respiration “basal”. The characteristic we measured is rather the “potential respiratory activity”. We agree with the reviewer that our methodological approach should be explained better. We also did measure the CO₂ emissions during the incubation. We included the explanation in the Methods section (L130-140) and commented on this issue in the results (L238-244, Fig. 2) and discussed the results on L317-325.

The discussion and conclusions are too long and go far beyond what you can say based on your results. There are many sections you discuss about the dynamic of microbial community but you only did one sampling time. You can't make big conclusions about dynamic of the system, such as L362-376. You need to just briefly mention potential dynamic but don't go much further. Similarly, you

speak a lot about the effect of plant cover even you did not measure any parameters to characterize the plant cover (you also forgot to mention anything about mosses and lichens despite their importance in the Arctic) such as above ground biomass, root biomass, percentage cover (did you properly assessed it?), diversity. You just described the main vascular plants. So your section 4.4, is simply too long and not fully supported with your data. This entire section could be reduced in few sentences and focus on presence/absence of plants and not linked to microbial dynamic.

Author response: We agree with the reviewer's concern that discussion was too long and sometimes not organized and confusing. We completely rewrote the whole discussion. We also added information about soil crust, mosses and plant cover in the supplements and comment o it more specifically throughout the manuscript.

On the other hand, your discussion lack of putting your results in perspective with the literature, for example other studies investigating microbial community in bare/unvegetated soil (there is several article on this in the Arctic). It would be interesting to see if F/B ratio is similar in unvegetated a low altitude compared to other study.

Author response: We extensively searched the WOS again and included new sources in the manuscript. The discussion was completely rewritten.

Your main conclusion should be the absence of plant rather than altitude effect, especially if we consider that high altitude soil in the Arctic (especially High Arctic) is likely to be rare, and also extremely shallow (only few cm) and not a massive stock of C. So we can wonder of their importance?

Author response: We don't fully understand these remarks. Isn't the decreasing plant abundance due to the altitude effect? Please consider our sampling strategy (L 88-92). What connection is there with rareness of high altitude soils? We don't think that the high altitude soils doesn't deserve scientific interest (despite of their low C stocks and shallow soil profile), especially in context of proceeding global warming and future development of ecosystems in the Arctic. We would like to point out that the high elevation habitats form significant part of the non-glaciated landscape not only in Svalbard, but also in other parts of the northern circumpolar region. However, we completely revised our conclusions and discussion to make it more straightforward.

Also you should more conclude on the site effect you have (Gr1) as site is clearly an effect on microbial biomass (consistently lower) and structure for the high altitude site.

Author response: We agree with reviewer opinion. This issue was revised throughout the manuscript including result (section 3.3) and discussion.

Another point you miss in your discussion is the fact that the soil you study is always alkaline (are alkaline soils prevalent on Svalbard or acidic?). You need to discuss or conclude if your results would be similar on acidic soil in the Arctic and compare with the relevant literature as pH is a big driver of microbial community including in the Arctic. This is an important point to make and be more critical about your results.

Author response: We thank the reviewer for pointing out this problem. Soils on Svalbard are neutral or alkaline. The soil pH in other parts of the northern circumpolar region is more variable, ranging from acidic pH (mainly on granite and gneiss bedrock) to highly alkaline pH. We are aware of pH effect on microorganisms. This issue is discussed more thoroughly in the current manuscript version.

Finally, you find an effect of altitude on microbial biomass mainly when you divide the data by TOC. This bias the data and don't reveal hidden effect of altitude. Yes, you have less C at high altitude but you don't have more microbial biomass. The fact there is more biomass per unit of C is not of a major interest. Dividing your results by TOC is not important and bias your results. Focus on the altitude

effect on microbial community structure and ratio, and acknowledge that there is not a major effect of altitude on biomass but rather a site effect.

Author response: We agree with the reviewer that we should emphasize the spatial variability in the microbial biomass and activity data. The normalized data were removed from the manuscript.

Specific comments

Altitude and transect are both gradient and not only transect. This is really confusing in the text when you speak about gradient, as it is unclear if you speak about vertical or horizontal. You can't refer to the horizontal gradient as "gradient" and the vertical one as altitude. Decide if you speak about vertical or horizontal gradient, or altitude and transect. Change in the entire text, but don't go from "gradient" for horizontal and then use gradient also for "vertical". Be consistent.

Author response: We agree with the reviewer comment. We clarified this throughout the whole manuscript. We clearly distinguished between the effects of altitude (vertical aspect) and transect (horizontal aspect) in the revised manuscript.

Introduction

L20: true but it is simply related to less C from a plant origin, it does not mean that there is more biomass. This is not true as well for GR1. You are telling only one part of the story here.

Author response: Normalized microbial characteristics were removed from the manuscript.

L21: "the 2 dominant microbial groups" it sounds like it is unusual or a result on its own but in the same time with PLFA you have access to fungi and bacteria only. Just speak about fungi and bacteria.

Author response: We agree with this comment. Abstract was rewritten.

L23: you didn't measure microbial dynamic over time, only the change in soil temperature. So, I would focus on what you measured and not make assumption, especially in the abstract. Keep this for the discussion

Author response: Assumptions were removed from the abstract.

L25-26: the conclusion is an overstatement. In general, unvegetated area should be considered as previous studies showed. Speaking about high elevation as hotspots of microbial activity based on 1 measurement is an overstatement (see main comment).

Author response: We removed such conclusions from abstract.

L36-41: this is normal as high altitude usually have no soil present or are extremely shallow (few cm) and may not be as important in their distribution and volume than low altitude (see main comment).

Author response: We completely agree with the reviewer that the high altitude soils are more important in distribution and volume compared to high elevation soils. However, we wanted to point out in these lines that the soil microbial properties were not thoroughly studied yet along the altitudinal gradients in the Arctic. We consider the information given in these lines relevant.

L42: not true, you can assess the effect of changing microclimate not using altitudinal climate as you can be using different sampling time (which you did not do), different exposition, open top chambers etc. So, just focus on altitudinal but you can't say that other studies can't assess change of microclimate.

Author response: The sentence was removed.

L47: are the ranges of altitude comparable and are the ecosystems comparable?

Author response: No, they are not. We are writing here about general altitudinal trends. The referred studies investigated microbial communities in different latitudes and across different altitudinal ranges (please see our next response).

L51: you cite articles on complete different ecosystems, such as Fierer et al 2011 (tropical), Meng et al 2013 (forest)... Focus on the Arctic and no other biomes, and same ecosystems (i.e. tundra and not forest) as you will not expect to have the same trends. Clearly state the location and ecosystems the studies you cited are based on.

Author response: We are aware of that. General latitudinal, but also altitudinal trends in diversity of animals and plants are one of the most widely recognized patterns in ecology. However, they are not valid for the microbial diversity as we wanted to show here. The number of references about microbial diversity or community structure along altitudinal gradients from the Arctic is strongly limited.

L59: this is not true. Your study is also true at your sites and at other sites the effects will differ. The number of studies help us to determine the common drivers across different sites. Your study is not better than others at that level. For generalisation, you could have cited Chu et al 2010 as global study of microbial diversity across the Arctic (Soil bacterial diversity in the Arctic is not fundamentally different from that found in other biomes)

Author response: We completely agree with the reviewer opinion – our study is not better than others and the effect is site specific, which does not allow generalization – as we have written here. What is not true then? We didn't want to highlight our study here. However, we rewrote the sentence for clarity (L62-64).

L60: “Fundamental” this is a strong word, please rephrase.

Author response: Sentence was removed..

Materials and Methods

L80-84: any idea of the percentage plant cover? Did you measure it? You don't mention mosses and lichens, but they represent a large part of the plant cover in Arctic tundra and are completely missing from your description.

Author response: Regarding the primary producers, we have data about plant and lichenized soil crust percentage cover at the sampling sites. Mosses were relatively scarce at these locations. These data were added to the result section (L228-231, Fig. S5, Table S1) and discussed. However, we didn't assessed the plant diversity and biomass at the sampling sites.

L85: you speak about the bedrock in the entire article but you never define/describe it. Could you give some information on it.

Author response: We would like to thank the reviewer for pointing out this deficit. We have detailed information about geology of the Petunia Bay. The information was added to the site description (L83-86, see also L379-380).

L86: was an organic horizon present in the low altitude soils?

Author response: No, the soil profile is poorly developed. Based on our experience there is rather litter layer on the soil surface and then relatively homogeneous mineral soil layer overlaying coarse gravel.

L93: “kept frozen” at which temperature?

Author response: At -20 °C (L95)

L122: section 2.4 there are no references and no justification of your measurements choices: temperature, duration of incubation...?

Author response: The information was added (section 2.4) and the methodological approach was commented in results (L238-244, Fig. 2) and discussion (L317-325).

L131: why did you adjust the amount of soil based on TOC and in which way? Could this bring a bias if you have more soil for example in high altitude to compare you results?

Author response: We optimized the PLFA extraction protocol in our laboratory to suit wide range of different soil types (with respect to the amount of lipids and size of the SPE cartridge for lipid fractionation, analytical procedure and stock sample aliquots for eventual reanalysis). As we use 0.7-1g of soil with TOC content around 5% and the microbial biomass is usually proportional to TOC content, we adjust the sample size for C “poor” soils. This modification could not bias our results since the extraction efficiency is not affected and the PLFA yield is thus comparable for all samples. Rather the opposite is true – not accounting for low TOC content could lead to concentrations of particular PLFAs below the detection limit.

L154, 156: change “mL” to “ml”

Author response: The abbreviation mL was used in all recent articles published in BGS.

L161: I guess you checked also for homoscedasticity?

Author response: Yes, we checked the data also for homoscedasticity (L174).

L161: state clearly if you transformed or not the PLFA data.

Author response: The relative data were long-transformed (L174-175).

L163-164: the horizontal and vertical transect are both gradient (one vertical one horizontal). See comment at the beginning.

Author response: We agree with the reviewer. We clearly distinguished between the effects of altitude and transect in the revised manuscript.

L166: how was the forward selection done?

Author response: It was performed using CANOCO 5.0 software. The soil geochemical parameters were used as explanatory variables while the relative abundances of microbial groups (MCS) were used as the dependent variables (RDA). The test offer list of candidate variables sorted according to their contribution to total explained variation in the dependent variables, together with their significance. After selection of the best candidate, the contribution and significance of remaining candidates is recalculated to explain the remaining variability in the dependent data. Then the next candidate can be selected (of course only if significant).

L167: why did you use only P values adjusted by Holms corrections. Any reference for that?

Author response: We used the significant values adjustment to reflect the multiple tests performed on the same dataset. The Holm’s correction follows the approach described *in* Holm (1979): A simple sequentially rejective multiple test procedure. *Scand. J. Stat.* **6**: 65-70. This procedure is slightly less conservative compared to the often recommended Bonferroni correction. On the other hand, it is a sequential procedure and takes into account that the candidate predictors with stronger effect were selected first. Thus it suits better for the

forward selection procedure (please see above our comment on the forward selection). Reference given in L183.

L172-173: it is really confusing when you speak about whole-plots vs splits-plots when you don't have a plot experiment. I am not sure what you refer to here.

Author response: As we mentioned in lines 185-187, we could assume that the characteristics of each sample will be auto-correlated with characteristics from other two samples taken from the same site (otherwise we had 9 independent transects and not three, which is not the case). In other words, the samples from the triplicate cannot be considered as independent samples due to relatively low inter-sample distance. The sampling design was in this context hierarchical with repeated measurements for each sampling site. We clarified this on L 185-187.

L176: what type of correlation did you use? Why there is no direct reference to correlation in the previous sentence?

Author response: We used the Pearson correlations to find out how tightly were two variables related to each other (L190-191).

Results

L190-192: this is a repetition of L 204. L192, it is also wrong what you say as the low altitude site show higher soil moisture than high altitude. Delete the sentence.

Author response: The results were largely rewritten and the duplications were removed (see sections 3.1 and 3.2).

L187, 203: which gradient are you talking about, be clear. L214: cite the Table you refer to L217-218: say if the correlation is positive or negative when you mention correlation even if it is given in brackets. L220: finish the sentence by "while increased in Gr2 and Gr3".

Author response: The changes were done according to reviewers comments throughout the manuscript.

L223-225: this should be given in the materials and methods and justify why you should use it. This problematic for me and can bias your results as mentioned in the main comment.

Author response: The normalization per TOC content was removed from the results.

L233: what is the "whole PLFA profile"? You did not use all the biomarkers in previous tests, it is not the same than MCS?

Author response: Comments on PLFA profile were removed.

L229-237: should cite figure 5? For example, L230 which figure you refer to?

Author response: There is no figure showing these results (L 249-251 in the current manuscript version). We think that figure showing the relation between selected environmental variables and MCS (Fig. 3 in the current manuscript version) are more important.

L250: change "typical" by "characterized"

Author response: Sentence was revised.

Discussion

L256: do you mean "did not" or "did" correspond. Looking at your plot, you have the same trend between soil and atmospheric, just few degrees' differences. Nothing surprising here. Your

explanation is not logical, snow will insulate the soil from air temperature, so having less snow should make the soil temperature more similar to air temperature but at the beginning you say they don't correspond. So, what are you trying to say?

L255-264: this whole section I am not sure what message you are trying to deliver. Nothing is really new here, and could be condensed in a shorter section.

L281: OM does not grow but increase. Change "growing" by "increasing" L285: "documented growing contribution of", rephrase this is difficult English to understand

L288- 290: ok there is no vascular plant, but what about mosses and lichens? Could they be partly responsible for presence of sitosterol? You have to discuss about mosses and lichens in the article, you completely omitted to mentioned them, and I don't think they are not present on the soil. Lichens have a distribution up to high-altitude and you often find them on Svalbard on top of mountain even without any soil.

L306: this reference is a bit old, is there any more recent references done on a larger number of bacteria?

L307: what is the parental material?

L309-314: of course if you divide microbial biomass and activity by TOC you will find an altitude effect, because there is less plant input at higher altitude so lower TOC. It does not mean your microbial biomass is higher at high altitude or there is higher microbial activity. Is it really important or interesting to know that there is higher proportion of living microbial biomass per soil TOC content? Your site effect is stronger on microbial biomass (PLFA /g soil) than altitude (only present for GR1) which is an interesting result on its own and show the variability of the supposed vertical gradient.

L316-324: This is repetition of L309-324. You need to merge both section and make it shorter, and again that you have higher microbial biomass per TOC is not of a major interest.

L330-331: keep in mind that you work on alkaline soil and it might be difficult to compare to other soil which are acidic...L348: you are not working on "dynamics" because you only have one sampling time. Remove dynamic from the title as you can only make some assumption

L349: Why do you think it is low quality litter? Compare to what? The fact you have a high C/N does not mean low quality but just more OM and less degraded. Low OM quality usually have low C/N. This also could contradict what you say L354 "which released easily assimilable". Is low quality litter easily assimilable?

L352: So what? Ok it reduces the microbial biomass per unit of C, it just means that there is more C in the soil. You describe this as an issue but I don't see why it should be an issue? The microbial biomass does not decrease with increasing altitude! Sorry to repeat myself but you use the results you prefer to support your theory without considering your entire results.

L356-357: this is a strange wording which make the sentence difficult to understand. What do you mean by "inverse consequences from soil MCS compared to development of microbial communities"? In which way this is "inverse" and how do you have an inverse microbial community structure (or just different) and why you speak about development or young soil when you don't measure microbial growth or dynamic and the age of the soil?

L358: Are you talking of your study or not? You are not working on a "succession". You are again making too many assumption and extrapolation based on your results. This is also contradictory to what you say in the previous sentence. Here you say at the "maximal plant biomass" the fungi dominate, while you say bacteria dominate in the previous section and in general in the article.

L360: so, any conclusion? Can you really make this comparison based on 1 study? L362-376: your results do not support what you say. You have one sampling time point, don't make assumption on what you don't measure: dynamic. Focus on your results. You can say that the presence of ergosterol coincide with continuous dominance of fungi at high altitude sites, but you don't need an entire section about it. Delete most of this section into one or few sentences.

[Author response: We agreed with most of the above-mentioned remarks. Based on these numerous comments, we decided to completely rewrite the whole discussion.](#)

Conclusion:

L379: move "were" just before "characterized" L380: this is not true. Unless you divide by TOC, there is no consistent effect of altitude on biomass and activity. L381: can you really say that there is negligible effect of microclimatic conditions over the summer with only one date of sampling? Do you think you have enough resolution with your sampling strategy to assess the effect of summer

microclimate? L382: again, you use gradient without saying which gradient you are talking about and when you define in the material and methods “gradient” to refer to horizontal not vertical. L383: you need to clearly state the decrease in pH. The decrease is less than a pH unit and the soil remains slightly alkaline. This is important because your results are likely to be completely different on acidic soil. L384-385: again, what is the bedrock at your sampling site? L386-388: well, there is plenty of unvegetated area at low altitude and even when there are plants. Your thinking must be developed to unvegetated area not only at high altitude. Do plants will colonise high altitude soil which are only few cm thick with global warming? Also give a reference for the potential increase in plant cover in the Arctic as several articles were recently published. L389: you can't really say that it diminishes the variability because it depends on plant species colonizing new area, the bedrock as you say. You don't measure variability with PLFA, the resolution in the method you use is not high enough. L390: you don't measure microbial diversity, how do you know this could have a negative effect? L393: again not true, you don't have a considerable microbial biomass and your measure of microbial activity is questionable. You just can't make this conclusion L379-394: there is no mention of the site effect even if you clearly have a site effect on microbial biomass and activity. This should be clearly stated as the vertical gradient is directly affect by the horizontal one in relation (in your study) to bedrock.

[Author response: The conclusions were completely revised.](#)

Reviewer 2

General comments:

The study by Kotas et al. was focused on changes in microbial biomass, activity, and broad community structure (based on PFLA) along altitudinal gradients in the Arctic. This question has great significance concerning the implications of global warming on these ecosystems. The study consists of 3 different transects represented by 4 different elevations, and for each sample the authors collected substantial amounts of data representing soil type, soil chemistry (pH, ion content and concentrations, TOC, TN, moisture content, and temperature ranges), and very briefly mention vegetation coverage. The authors try to disentangle the impacts of all these along with elevation on microbes using partial redundancy analysis as well as several other statistical approaches. They have a robust sample design with good replication to try and address this question.

I did have several issues with the manuscript. First, I found it very confusing that the authors kept referring to two different gradients, altitudinal (the main gradient of interest), and horizontal. However, this horizontal aspect is never discussed in the methods section and I assume it is referring to the south to north orientation of the 3 transects along the Petunia Bay. This needs to be clarified explicitly and its significance needs to be discussed. Is it expected there is a strong S-N effect? I assumed these 3 gradients were expected to be replicates of each other, but they have strong differences in soil characteristics and microbial community (particularly Gr1). This becomes more apparent in the Discussion, but the author's need to make this clear early on.

Author response: We agree with the reviewer opinion. We clarified this throughout the whole manuscript and clearly distinguished between the effects of altitude (vertical aspect) and transect (horizontal aspect). The 3 transects were expected to be replicates of each other. We didn't expect any variability in soil geochemical or microbial characteristics which could be ascribed to the differences in orientation of the selected transects. Opposite was true - we did our best to select similarly oriented transects (slopes on the western coast of Petunia Bay) in order to minimize the effect of distinct slope orientation.

I also had concerns with their microbial respiration data and the authors need to justify their choice of a 2 week pre-incubation at 6 C. The pre-incubation will burn off all the labile carbon and drastically alters this respiration rate. This needs discussed as it can substantially alter the conclusions of a large portion of the paper.

Author response: We agree with the reviewer that we have to justify and discuss our methodological approach. The methodology was chosen according to available knowledge and our experiences with similar experiments. We insist that the presented respiration data corresponds to in situ microbial activity. We measured the CO₂ emissions during the incubation and included these data in the manuscript. We also included justification of our methodological choices in the Methods section (L130-140) and commented on this in result section (L238-244, Fig. 2) and discussed the results on L317-325.

The discussion is too long and wordy. I found it difficult to understand the main points the authors were trying to convey. It seemed to be rushed relative to the excellent writing of the rest of the manuscript and has multiple grammar issues. I also think that there was too much superfluous material that distracts from the main message. The authors spend a great deal of time discussing impacts due to plant biomass, but have no data presented quantitatively examining plant communities, biomass, root biomass, etc. A lot of this can be safely removed, especially in sections 4.1 and 4.4, as the degree of detail discussed doesn't add too much to the broader implications of the study.

Author response: We agree with these comments. Discussion was revised, shortened, and previous sections were merged into two main parts. We also provided the information about plant and lichenized soil crust percentage cover as these data are recently available for the sampling sites (please see L228-231, Fig. S5, Table S1).

With some mostly editorial changes focusing on clarifying the findings I think this paper represents a significant contribution towards Arctic research and understanding the environmental parameters shaping microbial communities in this sensitive area.

Specific Comments

L124: I was interested in why the authors decided to pre-incubate the soils at 6 °C (far above the mean of -3.8°C, and below the max of 16.2, as well as different from the 5 °C cut-off used in L186)?

Author response: The incubation temperature of 6 °C was chosen as it represents mean summer soil temperature along the whole elevational gradient (mean summer temperatures for particular elevational levels ranged from 5.3 to 7.1, see Table 1; the mean summer temperature across the whole gradient is 6 °C). Justification is given in section 2.4.

Also, why did the authors choose to pre-incubate for 2 weeks at this temperature? Is this typical for these kinds of measurements? I would think you want to minimize the pre-incubation time to prevent a strong bottle effect, as well as removing all your labile carbon.

Author response: Our methodological choices and more detailed description of our incubation experiment are given in section 2.4. We further commented implications of our measurement in L238-245 and L317-326.

L126: Is the specific respiration ratio typical to compare with the field? Is it possible to convert PLFA to a more generalizable unit (such as per cell, per g biomass etc.) using conversion factors?

Author response: We excluded the specific respiration rate from results. However, we don't think that conversion of soil PLFA content to microbial biomass carbon (or per cell) could add any value. The conversion factors vary in the literature sources and are inevitably affected by cell morphology (membrane area versus cell biovolume). There is different PLFA to microbial biomass ratio not only for fungi and bacteria, but also for bacterial cells differing in size and shape. As the fungi to bacteria ratios varied significantly between sites, we consider any recalculation using a single conversion factor as speculative and hardly employable for comparison with other studies based on measurements of soil microbial carbon content (e.g. by chloroform fumigation method).

L144: Is there a reference to support this sum? Are you not overcounting the bacterial contribution by summing general bacterial biomarkers with specific bacterial group biomarkers (Actinos, G-, G+)? Would it not be preferable to us general fungal : general bacterial only?

Author response: The bacterial abundance is in majority (if not all) of papers using PLFA as quantitative measure of microbial biomass calculated as a sum of all markers specific to bacteria. The specific bacterial groups (Actinobacteria, G-, G+) belongs to bacteria and they need to be considered when calculating the F/B ratio. Considering only general bacterial markers, which are specific to bacteria but cannot be ascribed to one of the above mentioned bacterial groups, would lead to significant overestimation of fungal presence in the soil (references e.g. Frostegård and Bååth 1996, Biol. Fert. Soils 22, 59-65; Bååth and Anderson 2003 SBB 35, 955-965; Kaiser et al. 2010, New Phytologist 187, 843-858).

L189: Maybe change "In contrary" to "In contrast".

Author response: Sentence was rewritten.

L214: Maybe add at the end "and was instead transect specific". I realize this is implied, but I feel it makes it clearer.

Author response: Sentence was rewritten.

L213 – L227: This section is confusing to me. It is very surprising that microbial activity (as you assayed it) is not related to carbon or nitrogen content and is instead related to positively with Ca and negatively with Mg. I worry the trend in increasing respiration with altitude is due to the pre-incubation.

Author response: This relationship between respiration and base cation availabilities was surprising also for us. However, the microbial activity (respiration in this case) doesn't have to correspond with biomass as was shown previously (Šantrůčková and Straškraba, 1991, SBB 23, 525-532). Based on the background data from our respiration measurements (please see above our response to general comments), we insist that the presented respiration data are not a result of our pre-incubation step and can be used as potential respiratory activity of soil microbes. We thus believe that soil geochemical properties such as high magnesium availability can be very important drivers of microbial activity and abundance in these arctic soils. Moreover, the studies of Webb (Webb, 1949; reference in the manuscript) support the assumption, that parent material with very high Mg²⁺ content could have such negative effect on microbes.

L228: Write out "Microbial Community Structure" in the header of this section.

Author response: Done

L229: Gradient here is the transect? Does this mean there is a continuous change along the S-N transects or that each is different?

Author response: Yes, gradient is transect here. The results mean that there is a significant shift in the MCS not only between elevations, but also significant differences between transects in horizontal direction. The use of horizontal (transect) and vertical (altitude) aspects was emphasized throughout the manuscript.

L230: Nice to see so much explained due to altitude!

L231: Which gradients? Elevation or between the transects? Please fix or clarify this terminology!

Author response: Terminology was clarified throughout the manuscript.

L229 – L233: These few sentences are quite confusing and I think readers would be helped if you clarify. If I understand, the microbial community structure is impacted by elevation, but even more so by how the soils change with elevation? You ran multiple different tests to parse out these effects at different levels? Also, is microbial community structure here a relative score or absolute values?

Author response: We will clarify these statements. Let us to offer brief explanation: the microbial community structure significantly changed along the elevational gradients and between transects (ie. both factors, transect and elevation, were significant). The significant effects of transect and elevation can be well explained by spatial variability in the soil geochemical properties which were determined (ie. horizontal and altitudinal variability in the soil properties). The MCS used here and in general throughout the manuscript are relative abundances of microbial groups (not scores, see L175-176 in Method section).

L237: Re-running the analysis with the selected variables was non-significant? Can you clarify this statement? Why do you want to run the forward selection if the variables selected do not significantly explain the microbial community composition? Is the main message of this part, that these variables are not significant while altitude is?

Author response: These results were removed from the manuscript.

L240 – L251: Nice results! I think this is more interesting than the previous paragraph. However, there are a lot of grammar mistakes here, some listed below. Maybe re-write this section for clarity.

Author response: Section was rewritten and clarified.

L243: missing a space

L247: “A similarly significant trend”

L248: Change to PFLAs.

L249: change discrepant to disparate

Author response: Was corrected.

L249: Consider re-writing, this is a very long sentence that can be shortened, maybe “The most disparate site in terms of MCS was the highest elevation sampled along Gr1. It was typified by a high abundance of PLFAs specific to Actinobacteria and a lower abundance of fungal PFLAs compared to analogous sites along Gr2 and Gr3.”

Author response: Sentence was rewritten.

L255: What does this sentence mean?

Author response: The whole section was shortened and clarified.

L265: “positive surface energy balance had a strong..”

Author response: Corrected

L273: This is an incredibly important but difficult to decipher sentence. I think a lot of the sentences above it can be shortened or removed, but this should be clarified. Do you mean that “Mean temperatures and temperature stability did not change with altitude in this study”? [Therefore, variations in your parameters due to altitude are not simply due to temperature differences?] Here I would start off with a stronger statement of what you mean, and then offer your support.

Author response: We mean that mean temperatures and temperature stability (diurnal temperature fluctuation) does not change with elevation as we expected – ie. temperature will decrease with increasing elevation and the microclimate will be less stable in higher altitudes. We also expected generally higher fluctuation of soil moisture. However, we found very similar temperature conditions in the lowest and highest elevations, while the mid-elevated sites experienced warmer but less stable summer soil microclimate. The most important microclimatic parameter thus seemed to be the length of vegetation season and its effect on vegetation. The whole section was rewritten and clarified.

L277: Extremely important to clarify what gradient you are talking about here.

Author response: Clarified.

L277: Are you missing a “not”. This is a confusing sentence.

Author response: Discussion was completely rewritten.

L281 – L296: Simplify this! It is too wordy and difficult to follow. E.G. “We explain this discrepancy by the proximity of glacier stream, which could wash away the upper soil organic layer during abnormal spring-melt events in the past”, can be changed to “The only exception was the lowest site of Gr2 which had similar OM content to higher elevation sites along the other transects. This is likely due to the proximity of a glacier stream, which would wash away the topsoil during a flood.”

Author response: We agree that the paragraph is too wordy. Paragraph was completely revised.

L284: “vascular plants also influenced”

L286: Please provide a citation for this.

Author response: citation provided (L337).

L288 – L290: Is this important for your findings?

Author response: Rewritten

L290: Lots of grammar issues.

Author response: Rewritten

L292: Or high lichen components at high elevation?

Author response: We agree that the importance of lichens must be thoroughly discussed. However, lichens contain algal and cyanobacterial photobionts so there is not a conflict with our statement.

L298 – L314: You need to discuss the implications of your pre-incubation step in this section. It can also be clarified or simplified for the readers.

Author response: We discussed the implications of our pre-incubation step. The whole paragraph was revised (see our comments to incubation experiment above).

L304-L308: Please include relevant concentrations of the Mg inhibitory effect here.

Author response: We would like to thank the reviewer for this comment. The inhibitory concentrations of Mg^{2+} in solution were above 5 p.p.m and 50 p.p.m. for G- and G+ bacterial species, respectively (Webb 1949, Microbiology 3, 410–424). The limiting concentrations will be mentioned in the discussion.

L309 – L314: This is a nice summary. However, the normalized characteristics are inherently dependent on the soil OM, so isn't their increase directly due to the OM decrease?

Author response: The normalized microbial characteristics were removed from the manuscript.

L323 – L324: Please clarify this statement. What shift in resources lead to the slow accumulation of low quality OM? What are the ramifications of your pre-incubation when you are suggesting some samples are enriched in more recalcitrant OM?

Author response: Rewritten

L327 – L336: A lot of speculation. Is all this necessary

L337 – L347: Very speculative.

Author response: We believe that Mg^{2+} availability is very important factor shaping MCS along the transects. It largely explained the trends in G-/G+ bacteria ratios (compare Table 3 and Fig. 6c, d in the manuscript). It was shown that growth of G- and G+ bacteria is limited at very different Mg^{2+} concentration levels (difference of one order of magnitude, see our response to comments on L304-308). The Mg^{2+} availability in the investigated soils exceeded these limiting concentrations, especially for G- bacteria (considering all available Mg^{2+} in soil solution and average soil moisture content 30%, the Mg^{2+} concentrations ranged approximately from 50-420 p.p.m.). We thus consider the given interpretation of observed shifts in MCS due to Mg^{2+} availability (Mg^{2+} availability was retained by RDA with forward selection of explanatory variables) as critical evaluation of relevant literature. However, we

admit that statements about substitution of fungi by Actinobacteria are speculative and will be removed. The section was completely revised.

L384: “bedrock chemistry were recognized as the main factors”

Author response: Rewritten

L387 – L388: A confusing sentence, consider revising.

Author response: Rewritten

Figure2: Consider moving either this figure, or Table1 to the supplemental information to shorten the main paper.

Author response: We would like to keep Table 1 in the main text. Figure 2 was moved to supplements.

Figure 4: How much variation is there between altitude replicates? Maybe add a supplementary figure showing ellipsoids or individual sample points.

Author response: We agree with reviewer comment on Fig. 4. New version of the figure showing the variability between altitude replicates and transects (Fig. 3 in the current version of our manuscript).

1 Soil microbial biomass, activity and community composition along 2 altitudinal gradients in the High Arctic (Billefjorden, Svalbard)

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12 **Abstract** The unique and fragile High Arctic ecosystems are vulnerable to proceeding global climate warming. Elucidation
13 of factors driving microbial distribution and activity in Arctic soils is essential for comprehensive understanding of the
14 ecosystem functioning and its response to environmental change. The goals of this study were to investigate the microbial
15 biomass, activity, microbial community structure (MCS) and their environmental controls in soils along three elevational
16 transects in coastal mountains of Billefjorden, Central Svalbard. Soils from four different altitudes (25, 275, 525, and 765 m
17 above sea level) were analysed for a suite of characteristics including temperature regimes, organic matter content, base
18 cation availability, moisture, pH, basal respiration, and microbial biomass and community structure using phospholipid fatty
19 acids (PLFA). We observed significant spatial heterogeneity of edaphic properties among transects, resulting in transect-
20 specific effect of altitude on most soil parameters. We did not observed any clear elevation pattern in the microbial biomass
21 and the microbial activity revealed contrasting elevational patterns between transects. We found relatively large horizontal
22 variability in MCS, mainly due to different composition of bacterial PLFAs, but also systematic altitudinal shift in MCS
23 related with different habitat preferences of fungi and bacteria, resulting in high fungi to bacteria ratios at the most elevated
24 sites. Our data further showed that the biological soil crusts on these most elevated, unvegetated sites can host microbial
25 assemblages of the size and activity comparable with the arctic tundra ecosystem. The key environmental factors
26 determining horizontal and vertical changes in soil microbial properties were soil pH, organic carbon content, soil moisture
27 and Mg²⁺ availability.

28

29

30 1 Introduction

31 Knowledge about the spatial distribution and activity patterns of soil microbial communities is essential to understand
32 ecosystem functioning as the soil microbes play fundamental role in biogeochemical cycling and drive productivity in
33 terrestrial ecosystems (van de Heijden et al., 2008). The soil microbial diversity in the Arctic is comparable to that in other
34 biomes (Chu et al., 2010) and the spatiotemporal variability in microbial community composition is large (Lipson, 2007;
35 Blaud et al., 2015; Ferrari et al., 2016). However, it is still uncertain which environmental factors drive the heterogeneity of
36 soil microbial properties in the Arctic.

37 Altitudinal transects offer great opportunity to study a distribution of microbial communities adapted to local
38 habitats and explain the patterns by natural gradients of soil conditions, [vegetation occurrence](#) and climate regimes over short
39 spatial distances (Ma et al., 2004; Körner et al., 2007). The proceeding climate change will further affect environmental
40 conditions in the Arctic (Collins et al., 2013) including expected upward migration of the vegetation and increasing plant
41 cover (Vuorinen et al., 2017; Yu et al. 2017). Therefore, the knowledge of current microbial distribution and activity patterns
42 along the altitudinal gradients together with identifying their controlling factors can help to predict future development of
43 ecosystems in this region. However, such studies are scarce despite the fact that Arctic tundra comprises 5% of the land on
44 Earth (Nemergut et al., 2005) and most coastal areas in the northern circumpolar region have mountainous character. So far,
45 only few studies assessing altitudinal trends in soil microbial properties were conducted in the Scandinavian Arctic (Löffler
46 et al., 2008; Männistö et al., 2007). The research on spatial variation in microbial community composition and activity in
47 polar regions was conducted mainly at narrow elevation range (Oberbauer et al., 2007; Trevors et al., 2007; Björk et al.,
48 2008; Chu et al., 2010; Van Horn et al., 2013; Blaud et al., 2015; Tytgat et al., 2016) or was focused on initial soil
49 development following glacier retreat (Bekku et al., 2004; Yoshitake et al., 2007; Schütte et al., 2010). Majority of studies on
50 the elevational patterns in microbial community structure (MCS) and activity has been done in mountain regions of lower
51 latitudes from tropics to temperate zone. The studies commonly show that the microbial activity decreases with increasing
52 elevation (Schinner, 1982; Niklińska and Klimek, 2007; Margesin et al., 2009), while there are no general altitudinal patterns
53 in soil microbial diversity and community structure. For example, the microbial community composition did not change
54 along elevational gradients in Swiss Alps (Lazzaro et al., 2015), while other studies have documented decreasing bacterial
55 (Ma et al., 2004; Lipson, 2007; Shen et al., 2013) and fungal (Schinner and Gstraunthaler, 1981) diversity with an increasing
56 altitude, and several studies reported the mid-altitudinal peak in microbial diversity (Fierer et al., 2011; Singh et al., 2012;
57 Meng et al., 2013). Beside the fungal and bacterial diversity, the relative abundance of these main microbial functional
58 groups is also variable. For example, Djukic et al. (2010), Xu et al. (2014) and Hu et al. (2016) found decreasing fungi to
59 bacteria (F/B) ratio with an increasing elevation, while Margesin et al. (2009) reported opposite trend in Central Alps.

60 The research focusing on environmental controls over microbial communities in polar and alpine regions
61 recognized many significant factors, including vegetation, litter C : N stoichiometry, organic carbon content, soil pH,
62 nutrient availability, microclimatic conditions, and bedrock chemistry. However, the effect of these variables was site- and
63 scale-specific (Van Horn et al., 2013; Blaud et al., 2015; Ferrari et al., 2016), which highlights the need for further research
64 on environmental controls of microbial community size, activity and structure at local and regional scales. To extend our
65 knowledge about microbial ecology and soil functioning in the arctic alpine ecosystems, we conducted study aiming to
66 assess the activity, biomass and structure of soil microbial communities and to determine their controlling environmental
67 factors along three altitudinal transects located in Central Svalbard. [These transects spanned from the vegetated tundra](#)
68 [habitats at the narrow areas at the sea level to unvegetated soils at the top of the coastal mountains.](#) The specific objectives of
69 our study were (i) to describe gradients of microclimatic and geochemical soil properties; (ii) to assess microbial activity
70 (soil respiration) and abundance of main microbial groups (fungi, Gram-negative and Gram-positive bacteria,
71 Actinobacteria, phototrophic microorganisms) using phospholipid fatty acid (PLFA) analysis; and (iii) to identify
72 environmental factors explaining the trends in soil microbial parameters along these altitudinal gradients.

73

74 2 Materials and methods

75 2.1 Study area and soil sampling

76 The Petunia bay (Billefjorden; 78° 40' N, 16° 35' E) is located in the center of Svalbard archipelago and represents typical
77 High Arctic ecosystem in the northern circumpolar region. The mean, minimum and maximum air temperatures recorded in
78 the area at 25 m above the sea level (a.s.l.) were -3.7, -28.3 and 17 °C in the period of 2013–2015, respectively, and stayed
79 permanently below 0 °C for eight months a year (Ambrožová and Láška, 2017). The mean annual precipitation in the Central
80 Svalbard area is only 191 mm (Svalbard Airport, Longyearbyen, 1981–2010) and is equally distributed throughout the year
81 (Førland et al., 2010).

82 In August 2012, we collected soils from three altitudinal transects (Tr1–3) on the east coast of Petunia bay. Each
83 transect was characterized by four sampling sites at altitudes 25, 275, 525 and 765 m a.s.l. (± 5 m). Transects were located
84 on slopes with similar exposition (Tr1 W–E, Tr2 WNW–ESE, Tr3 WSW–ENE; Fig. 1) and lithostratigraphy. Soils at the
85 lowest elevations developed from Holocene slope (Tr1 and Tr3) and marine shore deposits (Tr2), while the bedrock at more
86 elevated sites is formed by dolomite and limestone with units of basal calcareous sandstone (Dallmann et al., 2004). The
87 soils were classified as Leptic Cryosols (Jones et al., 2010) with loamy texture and clay content increasing with altitude
88 (Table 2), and were from 0.15–0.2 m to only few cm deep at 25 and 765 m a.s.l., respectively. The poorly developed organic
89 horizon was present only at the lowest elevation. The sampling locations were selected in geomorphologically stable areas
90 with a similar slope ($20\pm 5^\circ$). On each sampling site, nine soil cores (4 cm deep, 5.6 cm diameter) were collected and mixed
91 into three representative samples. Each representative sample was mixed from one soil core taken from the edge of the
92 vegetation tussocks (if vegetation was present) and two other cores taken in increasing distance from the vegetation to
93 maintain the consistency with respect to heterogeneity of vegetation cover and soil surface. The triplicates were collected
94 approximately 5 m apart from each other. Immediately after sampling, the soil was sieved (2 mm) to remove larger rocks and
95 roots, sealed in plastic bags and kept frozen at -20 °C till further processing. Soil subsamples for biomarker analysis were as
96 soon as possible freeze-dried and stored at -80 °C until extraction.

97 Transects represented climosequences from high Arctic tundra to unvegetated bare soil. Vegetation of two lowest
98 sites was dominated by *Dryas octopetala*, with significant contribution of *Saxifraga oppositifolia*, and variable contribution
99 of *Cassiope tetragona*, *Salix polaris* and grasses (*Carex nardina*, *C. rupestris*, *C. misandra*; Prach et al., 2012; personal
100 observations). The vascular plants species formed scattered vegetation patches at the altitude of 525 m a.s.l. with *Salix*
101 *polaris* and *Saxifraga oppositifolia* being the most abundant species. The soils at the most elevated sites were covered
102 mainly by soil crusts with scarcely occurring *Saxifraga oppositifolia* and *Papaver dahlianum* (personal observations). The
103 percentage cover of main surface types (i.e. stones, bare soil, vegetation, crusts and mosses) was estimated on each sampling
104 site from approximately 1m² area in a close vicinity of coring sites (Table S1, Fig. S6).

105 2.2 Monitoring of microclimatic characteristics

106 To describe the soil microclimatic conditions along the altitudinal transects, we continuously measured soil temperature at -
107 5 cm from 2012–2013 directly at the sampling sites of Tr1 using dataloggers (Minikin Ti Slim, EMS Brno, CZ). The soil
108 water content at the time of sampling was determined in soil subsamples by drying to constant weight at 105 °C. The
109 temperature regimes at particular altitudinal levels were characterized by 10 climatic variables (Table 1). The period of
110 above-zero daily mean ground temperatures is referred to as summer season throughout the text. We also considered number
111 of days with daily mean ground temperatures above 5 °C, characterizing a period with conditions suitable for vascular plant
112 growth (Kleidon and Mooney, 2000). The positive soil surface energy balance was calculated as a sum of daily mean
113 summer temperatures. The records from three years (2011–2013) continuous measurements at two automated weather
114 stations located at 25 and 455 m a.s.l. approximately 3 km apart from the observed transects (hereafter referred as AWS₂₅

115 and AWS₄₅₅, respectively; Fig. 1; see Ambrožová and Láška, 2017 for detailed description) were used to evaluate seasonal
116 variation of soil temperature and moisture regimes (Figs. S2, S3, respectively), and coupling of soil and atmospheric
117 temperatures (measured at -5 cm and 2 m above terrain, respectively; Fig. S2). Even though we were not able to
118 continuously measure soil moisture directly at the sampling sites, we regarded data from both AWS locations as
119 representative for the evaluation of seasonal moisture regimes.

120 **2.3 Soil characteristics**

121 The particle size distribution was assessed using aerometric method (Lovelland and Whalley, 2001), the soil type was
122 classified according to U.S. Department of Agriculture. The soil pH was determined in soil–water mixture (1:5, w/v) using
123 glass electrode. The cation exchange capacity (CEC) was considered to be equal to the sum of soil exchangeable base cations
124 Mg²⁺, Ca²⁺, Na⁺, K⁺ extracted with 1M NH₄Cl (Richter et al., 1992). The amount of H⁺ and Al³⁺ ions was neglected due to
125 the high soil pH. Base cations accessible for plant and microbial uptake (Mg²⁺, Ca²⁺, Na⁺, K⁺) were extracted by the Mehlich
126 3 reagent (Zbíral and Němec, 2000). Cations were measured by atomic absorption spectroscopy (AA240FS instrument,
127 Agilent Technologies, USA). Total soil organic carbon (TOC) and nitrogen (TN) contents were measured in HCl fumigated
128 samples (Harris et al., 2001) using elemental analyser (vario MICRO cube, Elementar, Germany).

129 **2.4 Microbial respiration**

130 Since we were not able to measure soil respiration on site or immediately after soil collection, we measured the potential
131 respiratory activity (soil CO₂ production) in the laboratory incubation experiment. We stored and transported the soils frozen
132 because it was previously demonstrated that freezing-thawing has a weaker effect on microbial activity than long-term
133 refrigeration (Stenberg et al., 1998) and comparable effect as drying-rewetting (Clein and Schimel, 1994). We then measured
134 microbial respiration in slowly melted field-moist soils twice during the adaptation period (day 4 and 12), which allowed
135 stabilization of the microbial activity after a respiratory flush following freeze-thaw events (Schimel and Clein, 1996), and at
136 day 13, when we expected a stabilized microbial activity. Briefly, soil subsamples (10 g) were incubated in 100 mL flasks at
137 6 °C, which corresponds to the mean summer soil temperature of all sites along Tr1. At days 4, 12 and 13, a cumulative CO₂
138 production from the soils was measured using Agilent 6850 GC system (Agilent technologies, CA, USA). The flasks were
139 then thoroughly ventilated and sealed again. Due to high soil pH, the total amount of produced CO₂ was corrected for its
140 dissolution and dissociation in soil solution according to Henderson-Hasselbach equation (Sparling and West, 1990) and
141 expressed as the microbial respiration rate per day. The daily microbial respiration rates measured between days 4-12 and
142 after stabilization (day 13) were not significantly different in any soil samples, therefore, we present only the later one.

143 **2.5 Microbial biomass and community structure**

144 The soil microbial community structure was defined using PLFA analysis according to modified protocol of Frostegård et al.
145 (1993). Briefly, 1-3 g (according to TOC content) of freeze-dried soil samples was extracted twice with a single-phase
146 extraction mixture consisting of chloroform, methanol and citrate buffer. After overnight phase separation achieved by
147 adding more chloroform and buffer, the organic phase was purified on silica columns (SPE-SI Supelclean 250mg/3 mL;
148 Supelco®, PA, USA) using chloroform, acetone and methanol. The polar fraction was trans-esterified to the fatty acid
149 methyl esters (FAME) (Bossio and Scow, 1998). All FAMES were quantified by an internal standard calibration procedure
150 using methyl-nonadecanoate (19:0) as an internal standard. To identify the FAMES, retention times and mass spectra were
151 compared with those obtained from standards (Bacterial Acid Methyl Esters standard, the 37-component FAME Mix,

Komentář [P.K.1]: NO₃- and reactive phosphorus (SRP) were excluded from the results

152 PUFA-2, and PUFA-3; Supelco, USA). The ISQ mass spectrometer (MS) equipped with Focus gas chromatograph (GC)
153 (Thermo Fisher Scientific, USA) was used for chromatographic separation and detection.

154 Only specific PLFAs were used to assess the microbial community structure: a14:0, i15:0, a15:0, i16:0, i17:0, a17:0
155 were used as markers of Gram-positive bacteria (G+); 16:1 ω 9, 16:1 ω 5, cy17:0, 18:1 ω 11, 18:1 ω 7, cy19:0 as markers of
156 Gram-negative bacteria (G-); 10Me16:0 and 10Me18:0 as markers of Actinobacteria (Kroppenstedt, 1985), 18:1 ω 9,
157 18:2 ω 6,9 as fungal markers (Frostegård and Bååth, 1996) and polyunsaturated fatty acids 18:4 ω 3, 20:5 ω 3 were used as
158 markers of phototrophic microorganisms (Hardison et al., 2013; Khotimchenko et al., 2002). A sum of Actinobacterial
159 markers, PLFAs specific to G+ and G- bacteria and general bacterial markers 15:0, 17:0 and 18:1 ω 5 was used to calculate
160 bacterial biomass and fungi to bacteria (F/B) ratio. The sum of all lipid markers mentioned above and nonspecific PLFAs
161 14:0, 16:0, 18:0 and 16:1 ω 7 was used as proxy for microbial biomass (PLFA_{tot}).

162

163 2.6 Sterol analyses

164 The β -sitosterol and brassicasterol were used as biomarkers of plant (Sinsabaugh et al., 1997) and microalgal (Volkman,
165 1986; 2003) residues in organic matter (OM), respectively. Sterols were simultaneously determined using microwave
166 assisted extraction adapted from Montgomery et al. (2000) and GC/MS (ISQ MS equipped with Focus GC, Thermo Fisher
167 Scientific, USA) analysis. Briefly, 0.5 g of freeze-dried soil was treated with 6 mL of methanol and 2 mL of 2 M NaOH.
168 Vials were heated twice at the centre of a microwave oven (2450 MHz and 540 W output) for 25 s. After cooling, the
169 contents were neutralized with 1 M HCl, treated with 3 mL of methanol and extracted with hexane (3 \times 4 mL). Extracts were
170 spiked by an internal standard (cholesterol), evaporated and derivatized by adding of pyridine and 1 % BSTFA at 60 °C for
171 30 min prior analysis. Sterols were quantified by an internal standard calibration procedure.

172

173 2.7 Statistical analyses

174 All data were checked for normality and homoscedasticity, and log-transformed if necessary. The relative PLFA data
175 (mol%) were log-transformed in all statistical tests. The significance of environmental gradients and corresponding shifts in
176 MCS (mol% of summed PLFA specific for fungi, G- and G+ bacteria, Actinobacteria and soil phototrophic microorganisms)
177 in horizontal direction (ie. effect of transect) and vertical direction (ie. effect of altitude) were tested using the partial
178 redundancy analyses (RDA) with covariates. Variation partitioning was subsequently performed to quantify the unique and
179 shared effects of transect and altitude on variability of MCS. Forward selection procedure was used to identify the soil
180 geochemical parameters best explaining the shifts in MCS. During the forward selection procedure, only *P* values adjusted
181 by Holms correction were considered. This procedure is slightly less conservative compared to the often recommended
182 Bonferroni correction, but it is a sequential procedure and takes into account that the candidate predictors with stronger
183 effect were selected first (Holm, 1979). The multivariate tests were performed without standardization by samples, but with
184 centering and standardization by variables (because the variables were not always measured at the same scale, see Šmilauer
185 and Lepš 2014) and Monte Carlo test with 1999 permutations. Only adjusted explained variation is referred throughout the
186 text. Since the samples from each triplicate cannot be considered as independent observations due to relatively low inter-
187 sample distance (otherwise we had 9 independent transects), only the sampling sites were freely permuted while the
188 individual samples were exchangeable only within the sampling sites. The differences in particular soil and microbial
189 parameters between respective transects and altitudes were addressed by ANOVA complemented with Tukey-HSD post hoc
190 test. To find out how tightly were variables related to each other, Pearson correlation coefficient was used. All statistical tests
191 were considered significant at $P < 0.05$. Multivariate statistical analyses were performed with CANOCO for Windows

Komentář [P.K.2]: Ergosterol was excluded from the results

192 version 5.0 (Ter Braak and Šmilauer 2012), for ANOVA, Tukey-HSD test and correlations between soil and/or microbial
193 parameters, Statistica 13 was used (StatSoft, USA).

194

195 3 Results

196 3.1 Altitudinal changes in soil microclimate

197 The soil microclimate at the studied sites was characterized by two distinct periods respecting the air temperature dynamics
198 (compare Fig. S2a with S2b). The winter period lasted typically from the middle of September to early June. The winter soil
199 temperatures were stratified according to the elevation and the temperature means decreased from $-4\text{ }^{\circ}\text{C}$ at 25 m a.s.l. to -10
200 $^{\circ}\text{C}$ at 765 m a.s.l. (Table 1, Fig. S2). In contrast, a short summer period was characterized by a significant diurnal fluctuation
201 of soil temperatures and weak altitudinal temperature stratification (Fig. S2). The length of the summer season more than
202 doubled at the lowest elevations compared to the most elevated study sites, while the period with daily mean soil
203 temperatures above $5\text{ }^{\circ}\text{C}$ shortened almost four times. Correspondingly, the positive surface energy balance gradually
204 decreased with an increasing altitude (Table 1). The maximum daily mean temperatures and diurnal temperature fluctuation
205 were highest at the mid-elevated sites, with the highest mean summer soil temperature reached at 275 m a.s.l. In contrast, the
206 least and most elevated sites experienced lower summer maximum daily means and soil temperature amplitudes (Table 1).
207 The effect of altitude on soil moisture was significant along Tr1 and Tr3 ($P < 0.001$ and 0.01 , $F = 22.76$ and 7.39 ,
208 respectively) with soil moisture content decreasing along with increasing elevation, but nonsignificant along Tr2. *Continual*
209 *volumetric measurements of soil water content at AWS₂₅ and AWS₄₅₅ showed that the soil moisture was relatively stable*
210 *during the summer season and desiccation events did not occur during the summer periods 2011–2013 (for more*
211 *information, see Fig. S3).*

212 3.2 Gradients of soil geochemical properties and surface vegetation cover

213 Both factors, *transect and altitude*, significantly affected soil geochemical properties (partial RDA, pseudo- $F = 8.3$, $P <$
214 0.001) and explained 61% of the total variation in soil characteristics. The RDA ascribed most of the explained variability
215 (73%) to vertical zonation. Accordingly, the effect of altitude was significantly reflected in all soil parameters (Table 2, 3,
216 Fig. S4), *but the significant interactive effect between transect and altitude indicated that the elevational trends were in most*
217 *cases specific for particular transects (Tables 2, 3). Especially the CEC and availabilities of Ca^{2+} , Mg^{2+} , K^{+} and Na^{+} were*
218 *spatially variable, reflecting complicated geology of the Petunia bay area. The soils along Tr1 were significantly richer in*
219 *available Mg^{2+} and K^{+} than soils from other two transects (Table 2). The Mg^{2+} availability also significantly increased with*
220 *increasing elevation along the Tr1 (Table 2). Other soil properties showed more systematic altitudinal patterns. The mean*
221 *soil pH ranged from 7.8 to 9.0 and increased with altitude along all transects (Table 2, Fig. S4). Oppositely, the soil TOC and*
222 *TN contents declined towards higher elevations along all transects; the exception was the lowest site along Tr2 with lower*
223 *soil OM content compared to the respective sites from Tr1 and Tr3. The OM poorest soil occurred at the highest site of Tr1*
224 *(Table 3). The soil C/N ratio, sitosterol content in TOC and the ratio between plant-derived sitosterol and brassicasterol of*
225 *algal origin were solely affected by the altitude. Their values systematically decreased with an increasing elevation*
226 *irrespective of the soil OM content (Table 3), indicating an altitudinal shift in the OM quality and origin. The percentage of*
227 *plant cover also continuously decreased with an increasing elevation along Tr1 and Tr3 but was comparable on the three*
228 *lower sites along Tr2 (Fig. S5), which significantly resembled the trends in soil OM content ($r = 0.53$; $P = 0.001$). The*
229 *lichenized soil crusts were predominant type of soil surface cover at all sites, while mosses covered very small proportion of*

230 surface area. The bare surface without any vegetation (bare soil) occurred only at the two most elevated sites (Fig. S5, Table
231 S1).

232 3.3 Soil microbial biomass and activity

233 The soil PLFA content, used here as a measure of soil microbial biomass, was significantly correlated with soil TOC and TN
234 contents ($r = 0.773$ and 0.719 , respectively; both $P < 0.0001$) and soil moisture ($r = 0.772$; $P < 0.0001$), and negatively
235 affected by Mg^{2+} availability ($r = -0.775$; $P < 0.0001$). Despite these relations, the soil PLFA content did not show any
236 altitudinal pattern. The soil PLFA amounts were comparable among differently elevated sites along particular transect (Fig.
237 2a). Only the most elevated site of Tr1 had significantly lower soil PLFA content than other sites, which corresponded with
238 its very low stock of OM (Table 3). Similarly, neither the flush of microbial respiration measured after soil thawing (day 4 of
239 incubation) nor the respiration measured after stabilization (day 12, not shown, and day13) showed any systematic altitudinal
240 pattern (Fig. 3b, c). Generally, the flush respiration rate was closely related ($r = 0.74$, $P < 0.0001$, $n = 36$) to microbial
241 respiration after stabilization and ca 2.3 ± 0.3 times faster, showing similar freezing-thawing effect on the whole set of
242 samples independently of altitude and transect. Along each transect, the three lower sites (from 25 to 525 m a.s.l.) had after
243 stabilization comparable microbial respiration rates, but the most elevated sites always differed - along Tr1 had the most
244 elevated site significantly lower microbial respiration rate, whilst the most elevated sites along Tr2 and Tr3 produced
245 markedly more CO_2 compared to remaining sites along these transects (Fig. 2b). The respiration rate was related neither to
246 PLFA nor to TOC contents, but significant positive correlation with soil Ca^{2+} availability and F/B ratio, and negative
247 correlation with Mg^{2+} availability ($r = 0.489$, 0.661 and -0.545 ; $P = 0.003$, < 0.001 and 0.001 , respectively) was observed.

248 3.4 Microbial community structure

249 The partial RDA revealed significant interactive effect of altitude and transect on MCS (pseudo- $F = 4.8$, $P < 0.001$). Both
250 factors explained 51% of the total variation in the MCS, with 66 % of explained variability ascribed to altitude, 26% to
251 transect, and 8% of explained variability shared by both factors. The soil geochemical variables explained 72% of the
252 variation in the MCS (pseudo- $F = 7.1$; $P < 0.001$) indicating that the interactive effect of altitude and transect on MCS was
253 largely driven by vertical and horizontal variability in soil properties. The forward selection of explanatory variables retained
254 four geochemical parameters: Mg^{2+} availability, pH, moisture and TOC content, all together accounting for 55% of variation
255 in the data (pseudo- $F = 11.6$, $P < 0.001$). The most pronounced shift in the MCS was given by different altitudinal
256 preferences of bacteria and fungi. The bacteria were consistently more abundant in the soils from lower elevations, having
257 lower pH and higher TOC and moisture contents (Fig. 3). In general, PLFAs specific to G- bacteria were more abundant
258 than PLFAs of G+ bacteria (Fig. 4a; mean G-/G+ ratio \pm SD = 1.76 ± 0.17 ; $n = 36$). Oppositely, the fungal contribution to
259 microbial community increased with an increasing altitude, at the sites having TOC poorer soils and higher pH (Fig. 3).
260 Therefore, the F/B ratio gradually increased with an increasing altitude along all three transects (Fig. 4b). The significant
261 interactive effect of altitude and transect on MCS was mainly connected with a strong effect of soil Mg^{2+} availability, which
262 was higher along the whole Tr1 and differentiated its microbial communities from sites located along Tr2 and Tr3, where
263 microbial communities of respective sites were more similar. The differences in MCS among the respective sites along Tr1
264 and other two transects further increased towards higher elevations in coincidence with an increasing soil Mg^{2+} availability
265 along Tr1 (Fig. 3). In result, the TOC poorest and Mg^{2+} richest soil at the highest site on Tr1 had the most distinct MCS from
266 all the sites. Its microbial community was characterized by higher abundance of Actinobacteria and PLFAs of phototrophic
267 microorganisms and much lower contribution of G- bacteria compared to communities of all other sampling sites (Fig. 3,
268 4a).

269

270 4 Discussion

271 4.1 Climatic and soil edaphic conditions along altitudinal transects

272 The coastal area of the Petunia Bay in Svalbard is characteristic by ca four months lasting summer, long winter period
273 (Ambrožová and Láška, 2017) and very low precipitations (Førland et al., 2010). Our measurements in this area further
274 showed that soils along an elevation gradient from 25 to 765 m a.s.l. face significantly different microclimatic regimes.
275 During winter, when the air temperatures varied a lot in time but less with elevation (Fig. S2b, data from AWS₂₅ and
276 AWS₄₅₅), the soil temperatures were relatively stable but significantly stratified with altitude (Fig. S1, S2a). The mean winter
277 soil temperatures decreased from -4 to -10 °C along the elevation gradient from 25 to 765 m a.s.l. (Table 1), which can
278 strongly reduce winter soil microbial activity at high altitudes (Drotz et al., 2010; Nikrad et al., 2016). In contrast, the mean
279 summer soil temperatures did not reflect the site elevation (Table 1) and the comparison of temperature fluctuations, mean
280 and maximum daily mean temperatures showed that the lowest and highest sites experienced during summer on average
281 colder, but more stable soil microclimate compared to the mid-elevated sites (Table 1). However, the summer season
282 prolonged with decreasing elevation and the increasing number of days with mean temperature above 5 °C and a rising
283 positive surface energy balance (Table 1) positively affected the occurrence and spreading of vascular plants (Kleidon and
284 Mooney, 2000; Klimeš and Doležal, 2010), which had strong implications for a transition of edaphic conditions along
285 studied elevation transects. Together with increased litter inputs and stocks of soil OM with lower C/N ratio (Table 3) was
286 the plant growth associated with root respiration, cation uptake, and release of H⁺ and organic acids from roots, all together
287 accounting for decreased soil pH (van Breemen et al., 1984). The increasing soil OM content was further positively related
288 to soil moisture (Fig. 3). Interestingly, the soils in general did not suffer from desiccation (Fig. S3), commonly identified
289 among the most stressing factors in polar and alpine ecosystems (Ley et al., 2004; Van Horn et al., 2013; Tytgat et al., 2016),
290 probably due to high cloudiness and fog occurrence (Sawaske and Freyberg, 2015) in the maritime climate.

291 The alkaline bedrock material resulted in high soil pH (7.8–9) and high availabilities of basic cations, which were,
292 however, spatially variable due to diverse geology of the studied area (Dallmann et al., 2004; Table 2). Beside clear
293 altitudinal trends in soil edaphic conditions connected mostly with the soil OM content, the Mg²⁺ availability was recognized
294 as main factor driving differences in soil microbial properties between transects (Fig. 3). In result, the character of the parent
295 substrate mostly controlled soil microbial properties at the most elevated sites, which had generally low OM content and the
296 most divergent MCS compared to lower located sites (Fig. 3). The highest site along Tr1 was the most extreme habitat
297 among all the chosen sites, with the highest proportion of bare unvegetated soil surface (Fig. S5), the lowest OM and
298 moisture contents, highest Mg²⁺ availability and soil pH, and consequently also the most distinct microbial characteristics
299 (Fig. 2, 3). Towards lower elevations, the soil OM content became increasingly important and the microbial characteristics
300 of the sites on different transects were more similar.

301 4.2 Soil microbial properties along altitudinal transects

302 The altitudinal shifts in soil edaphic properties were not significantly reflected in the soil microbial biomass and potential
303 microbial respiration. Generally, the soil PLFA contents were comparable between all the sites along particular elevation
304 transects, with the exception of very low soil PLFA concentration on the highest site of the Tr1 (Fig. 2a). There are no other
305 studies from the High Arctic ecosystems reporting about altitude effect on soil microbial biomass. However, other studies
306 conducted on alpine gradients in the temperate and boreal zones documented weak or absent altitudinal trends in the

307 microbial biomass (Djukic et al. 2010, and Xu et al., 2014 using PLFA; Löffler et al., 2008 using cell counts) but also a
308 negative effect of elevation in the Alps (Margesin et al., 2009) and northwestern Finland (Väre et al., 1997). Importantly,
309 none of the studies considered unvegetated habitats and all of them were conducted in soils with acidic or neutral soil pH.

310 Microbial respiration also did not change systematically with increasing elevation. The three lowest sites along each
311 transect always had comparable soil microbial respiration rates (Fig. 2b), while soil microbial activities of the highest sites
312 differed. The most elevated site on the Tr1 showed significantly lower respiration rates than the lower sites on this transect,
313 which was in line with the lowest OM content as well as soil PLFA content. However, the soils from the highest sites on
314 both Tr2 and Tr3 respired significantly more than the soils from lower sites on these transects, irrespective of relatively
315 stable microbial biomass. This is in contrast to other studies, which reported decreasing microbial activity with increasing
316 elevation (Schinner, 1982; Väre et al., 1997; Niklińska and Klimek, 2007). However, these studies were conducted in lower
317 latitudes and the studied altitudinal gradients did not include unvegetated habitats. To comment on and justify our results, we
318 are aware that microbial activities were measured in freeze-stored and not fresh samples (see section 2.3 for details) and,
319 therefore, the respiration rates measured after thawing show the potential activity of soil microbial communities in the soils.
320 However, the respiration rates in three subsequent measurements (after flush, during adaptation and after stabilization) were
321 positively correlated ($r = 0.93$ and 0.74 , both $P < 0.0001$, $n = 36$), the ratios between the flush and stabilized respiration rates
322 were comparable across all the soils (compare Fig. 2b and c) and the above-described differences in microbial activities
323 among the sites were consistent. Our data are in accord with the study of Larsen et al., (2002), who found comparable
324 response to freeze-thaw events between two different arctic ecosystem types. We thus suggest that the soils responded
325 similarly to the storage treatment independently of site location and that observed differences in soil microbial activities are
326 representative for the studied transects. Therefore, the higher soil microbial respiration at the most elevated sites point to a
327 higher lability of the present OM (Lipson et al., 2000; Uhlřřová et al., 2007) and/or to a shift in microbial communities
328 towards groups with higher potential to mineralize the OM (Gavazov, 2010; Djukic et al., 2013). Previous studies,
329 considering either bare soil or vegetated habitats, reported rather increasing complexity of soil OM with elevation (Ley et al.,
330 2004; Xu et al., 2014). However, in this study was majority of OM and microbial biomass at the most elevated sites
331 associated with biological soil crusts with high algal and cyanobacterial abundance (Table S1, Fig. S5), known for their high
332 microbial activity (Pushkareva et al., 2017; Bastida et al., 2014). The high microbial activity in the most elevated sites could
333 be ascribed to prevalence of compounds of algal/cyanobacterial origin with very low portion of complex and slowly
334 decomposable lignin and lignified compounds and protective waxes (like cutin and suberin) mainly derived from vascular
335 plants. In accord, the sitosterol to brassicasterol ratio gradually decreasing with increasing elevation (Table 3) and increasing
336 sitosterol content in the TOC pool at lower elevations pointed to growing importance of microalgal sources of OM in high
337 elevation habitats (Sinsabaugh et al., 1997; Rontani et al., 2012). Even though both sterols can be found in higher plants and
338 microalgae, the changing ratio indicates shift in the origin of OM (reviewed by Volkman, 1986, see also Volkman, 2003).
339 Changes within microbial communities, which can also help to explain higher soil microbial respiration at the most elevated
340 sites are discussed below.

341 Although the soil PLFA content did not change along the studied elevation transects, we have found a systematic
342 altitudinal shift in the PLFA composition, resulting in significantly increasing F/B ratio towards higher elevations. This shift
343 was best explained by a decreasing soil OM content and soil moisture and increasing pH (Fig. 3). Reports about soil F/B
344 ratios and their altitudinal changes from the High Arctic are missing, but studies from lower latitudes showed either a similar
345 trend of increasing F/B ratio with an altitude in the Alps (Margesin et al., 2009) or the opposite altitudinal effect in the Alps
346 (Djukic et al., 2010) and Himalayas (Xu et al., 2014; Hu et al., 2016). Such divergent results indicate that altitude alone is
347 not the key driving factor of the soil F/B ratio. In contrast to our observation, these studies reported very low soil F/B ratios
348 of 0.05-0.2, which may indicate important role of fungi in functioning of the Arctic habitats. Soil pH was previously

349 identified as the main driver of fungal-bacterial dominance in the soil (Baath and Anderson, 2003; Högberg et al., 2007;
350 Rousk et al., 2009; Siles and Margesin, 2016). Fungi have been found more acid tolerant than bacteria, leading to higher F/B
351 ratio in acidic soils (Högberg et al., 2007; Rousk et al., 2009; reviewed by Strickland and Rousk, 2010). However, here we
352 report high F/B ratios in the alkaline soils (pH 7.8-9.0) and increasing F/B ratios with an increasing soil pH. Similar trend
353 was reported also by Hu et al., (2016), but the authors found F/B ratios one order of magnitude lower compared to our study.
354 The possible explanation of generally high fungal abundance and increasing F/B ratio at more elevated sites, which are
355 typical by unfavourable edaphic conditions and severe winter microclimate, could be higher competitiveness of fungi
356 compared to bacteria in suboptimal conditions due to their wider pH (Wheeler et al., 1991) and lower temperature (Margesin
357 et al., 2003) growth optima. We further found that the increasing F/B ratio was significantly coupled with an increasing soil
358 respiration ($r = 0.649$; $P < 0.001$). Indeed, such relationship can be related to higher fungal ability either to prosper in the soil
359 conditions at the most elevated sites, or to utilize more efficiently available C sources (Ley et al., 2004; Bardgett et al., 2005;
360 Nemergut et al., 2005; van der Heijden et al., 2008). In turn, the higher bacterial contribution at lower elevations may be
361 associated with more benign soil conditions and bacterial preference for utilization of labile root exudates released by
362 vascular plants (Lipson et al., 1999; Lipson et al., 2002). Since the projected warming in the Arctic (Collins et al., 2013) will
363 likely cause an upward migration of the vegetation and increasing plant cover in detriment of lichens and biological soil
364 crusts (Vuorinen et al., 2017; Yu et al. 2017; de Mesquita et al., 2017), the soil microbial communities will likely respond by
365 decreasing F/B ratios at higher elevations.

366 Apart from the systematic altitudinal shift in the F/B ratio, we observed a strong shift in the bacterial composition,
367 which differentiated the altitudinal trends in the soil MCS along Tr1 from trends along Tr2 and Tr3. This difference between
368 transects increased towards higher elevations and was best explained by Mg^{2+} availability (Fig. 3). The soils from Tr1, except
369 the lowest site, had a lower G- to G+ bacterial ratios within microbial communities than soils from other two transects.
370 Further, the microbial community of the most elevated site along Tr1 was significantly more contributed by actinobacteria
371 and phototrophic microorganisms compared to all other sites (Fig. 3, 4a). It is known that the high Mg^{2+} availability inhibits
372 growth of many soil bacterial species. The observed inhibitive Mg^{2+} levels were 5 and 50 p.p.m for G- and G+ bacteria,
373 respectively (Webb 1949), indicating that these bacterial groups significantly differ in their tolerance for enhanced Mg^{2+}
374 levels. Considering half of available Mg^{2+} in soil solution and average soil moisture content 20%, the Mg^{2+} concentrations
375 ranged approximately from 16-140 p.p.m., which could explain decreased abundance of G- bacteria in sites with high Mg^{2+}
376 availability. This inhibitive Mg^{2+} effect further corresponds with the negative correlations between Mg^{2+} availability and soil
377 microbial biomass and respiration found in our study, and could explain the lower microbial biomass and respiration in the
378 soils from Tr1. Our data thus indicate that beside the traditionally identified drivers of microbial activity and MCS such as
379 soil OM content, moisture and pH, Mg^{2+} availability is an important factor shaping the microbial environment along the arctic
380 altitudinal transects on dolomitic parent materials.

381 **5 Conclusions**

382 The results obtained in this study have shown significant altitudinal zonation of most edaphic properties, but also significant
383 spatial heterogeneity in horizontal direction, resulting in transect-specific effect of altitude on abiotic soil properties. Our
384 data demonstrated that soils on the most elevated, unvegetated sites around the Petunia Bay can host microbial assemblages
385 comparable in size and activity with the tundra ecosystem. The high microbial biomass and activity at the most elevated sites
386 were almost exclusively associated with biological soil crusts, largely contributed by fungi. However, their development was
387 retarded on some sites by high pH, low moisture and high Mg availability, resulting in pronouncedly low OM content,
388 microbial biomass and distinct MCS. Despite the ubiquitous occurrence of soil crusts, the gradually increasing plant
389 productivity and litter inputs down along transects were associated with decreasing soil pH, increasing OM content and soil

390 moisture. Concurrently, the soil edaphic and microbial properties become more uniform. As the rise in temperatures and
391 humidity predicted by climatic models will likely cause an upward migration of the vegetation and increasing plant cover,
392 the higher plant litter inputs will overreach the influence of parent material and entail an increasing abundance of bacteria
393 and decreasing F/B ratio in the summer microbial assemblages.

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396 **Author contribution**

397 P. Kotas and E. Kaštovská analysed the data and wrote the manuscript with assistance of all coauthors. P. Kotas and J. Elster
398 designed the study and performed sampling. The microbial community structure and environmental parameters were
399 assessed by P. Kotas, E. Kaštovská and H. Šantrůčková.

400

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407

408 **Competing interests**

409 The authors declare that they have no conflict of interest.

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421 **References**

- 422 Ambožová, K. and Láska, K.: Air temperature variability in the vertical profile over the coastal area of Petuniabukta, central
423 Spitsbergen, *Pol. Polar Res.*, 38, 41–60, 2017.
- 424
- 425 Bååth, E. and Anderson, T.–H.: Comparison of soil fungal/bacterial ratios in a pH gradient using physiological and PLFA
426 based techniques. *Soil Biol. Biochem.*, 35, 955–63, 2003.
- 427
- 428 Bardgett, R. D., Bowman, W. D., Kaufmann, R., and Schmidt, S.K.: A temporal approach to linking aboveground and
429 belowground ecology, *Trends Ecol. Evol.*, 20, 634–641, 2005.
- 430
- 431 Bastida, F., Jehmlich, N., Ondoño, S., von Bergen, M., García, C., and Moreno, J. L.: Characterization of the microbial
432 community in biological soil crusts dominated by *Fulgensia desertorum* (Tomin) Poelt and *Squamarina cartilaginea* (With.)
433 P. James and in the underlying soil, *Soil Biol. Biochem.*, 76, 70–79, 2014.
- 434
- 435 Bekku, Y. S., Kume, A., Masuzawa, T., Kanda, H., Nakatsubo, T., and Koizumi, H.: Soil respiration in a high arctic glacier
436 foreland in Ny-Ålesund, Svalbard, *Polar Bioscience*, 17, 36–46, 2004.
- 437
- 438 Björk, R. G., Björkman, M. P., Andersson M. X., and Klemetsson L.: Temporal variation in soil microbial communities in
439 Alpine tundra, *Soil Biol. Biochem.*, 40, 266–268, 2008.
- 440
- 441 Blaud, A., Lerch, T. Z., Phoenix, G. K., and Osborn M. A.: Arctic soil microbial diversity in a changing world, *Res.*
442 *Microbiol.*, 166, 796–813, 2015.
- 443
- 444 Bossio, D.A. and Scow, K.M.: Impacts of carbon and flooding on soil microbial communities: phospholipid fatty acid
445 profiles and substrate utilization patterns, *Microb. Ecol.*, 35, 265–278, 1998.
- 446
- 447 Clein, J. S. and Schimel, J. P.: Reduction in microbial activity of litter due to drying and rewetting events, *Soil Biol.*
448 *Biochem.*, 26, 403–406, 1994.
- 449
- 450 Collins, M., Knutti, R., Arblaster, J., Dufresne, J. L., Fichet, T., Friedlingstein, P., Gao, X., Gutowski, W. J., Johns, T.,
451 Krinner, G., Shongwe, M., Tebaldi, C., Weaver, A. J., and Wehner, M.: Long-term climate change. Projections,
452 commitments and irreversibility. In: Stocker, T. F., Qin, D., Plattner, G. K., Tignor, M. M. B., Allen, S. K., Boschung, J.,
453 Nauels, A., Xia, Y., Bex, V., and Midgley, P. M. (eds.): *Climate Change 2013. The physical science basis. Contribution of*
454 *working group I to the fifth assessment report of the intergovernmental panel on climate change. Cambridge University*
455 *Press. Cambridge. p. 1029-1136, ISBN: 9781107057991, 2013.*
- 456
- 457 Chu, H., Fierer, N., Lauber, Ch. L., Caporaso, J. G., Knight, R., and Grogan, P.: Soil bacterial diversity in the Arctic is not
458 fundamentally different from that found in other biomes, *Environ. Microb.*, 12, 2998–3006, 2010.
- 459
- 460 Dallmann, W. K., Piepjohn, K., and Blomeier, D.: Geological map of Billefjorden, Central Spitsbergen, Svalbard, Temakart
461 Nr. 36, Norsk Polarinstitut, 2004.
- 462

463 de Mesquita, C. P. B., Knelman, J. E., King, A. J., Farrer, E. C., Porazinska, D. L., Schmidt, S. K., and Suding, K. N.: Plant
464 colonization of moss-dominated soils in the alpine: Microbial and biogeochemical implications, *Soil Biol. Biochem.*, 111,
465 135–142, 2017.

466

467 Djukic, I., Zehetner, F., Mentler, A., and Gerzabek, M. H.: Microbial community composition and activity in different
468 Alpine vegetation zones, *Soil Biol. Biochem.*, 42, 155–161, 2010.

469

470 Djukic, I., Zehetner, F., Watzinger, A., Horacek, M., and Gerzabek, M. H.: In situ carbon turnover dynamics and the role of
471 soil microorganisms therein: a climate warming study in an Alpine ecosystem, *FEMS Microbiol. Ecol.*, 83, 112–124, 2013.

472

473 Drotz, S. H., Sparrman, T., Nilsson, M. B., Schleucher, J., and Öquist, M. G.: Both catabolic and anabolic heterotrophic
474 microbial activity proceed in frozen soils, *PNAS*, 107, 21046–21051, 2010.

475

476 Ferrari, B. C., Bissett, A., Snape, I., van Dorst, J., Palmer, A. S., Ji, M., Siciliano, S. D., Stark, J. S., Winsley, T., and Brown,
477 M. V.: Geological connectivity drives microbial community structure and connectivity in polar, terrestrial ecosystems,
478 *Environ. Microb.*, 18, 1834–1849, 2016.

479

480 Fierer, N., McCain C. M., Meir, P., Zimmermann, M., Rapp, J. M., Silman, M. R., and Knight, R.: Microbes do not follow
481 the elevational diversity patterns of plants and animals, *Ecology*, 92, 797–804, 2011.

482

483 Førland, E.J., Benestad, R., Hanssen-Bauer, I., Haugen, J.E., and Skaugen, T.: Temperature and Precipitation Development
484 at Svalbard 1900–2100, *Advances in Meteorology*, doi:10.1155/2011/893790, 2010.

485

486 Frostegård, A. and Bååth, E.: The use of phospholipid fatty acid analysis to estimate bacterial and fungal biomass in soil,
487 *Biol. Fert. Soils*, 22, 59–65, 1996.

488

489 Frostegård, Å., Bååth, E., and Tunlid, A.: Shifts in the structure of soil microbial communities in limed forests as revealed by
490 phospholipid fatty acid analysis, *Soil Biol. Biochem.*, 25, 723–730, 1993.

491

492 Gavazov, K. S.: Dynamics of alpine plant litter decomposition in a changing climate, *Plant Soil*, 337, 19–32, 2010.

493

494 Hardison, A. K., Canuel, E. A., Anderson, I. C., Tobias, C. R., Veuger, B., and Waters, M. N.: Microphytobenthos and
495 benthic macroalgae determine sediment organic matter composition in shallow photic sediments, *Biogeosciences*, 10, 5571–
496 5588, 2013.

497

498 Harris, D., Horwath, W. R., and van Kessel, Ch.: Acid fumigation of soils to remove carbonates prior to total organic carbon
499 or carbon-13 isotopic analysis, *Soil Sci. Soc. Am. J.*, 65, 1853–1856, 2001.

500

501 Holm, S.: A simple sequentially rejective multiple test procedure. *Scand. J. Stat.*, 6, 65–70, 1979.

502

503 Högberg, M. N., Högberg, P., and Myrold, D. D.: Is microbial community composition in boreal forest soils determined by
504 pH, C-to-N ratio, the trees, or all three? *Oecologia*, 150, 590–601, 2007.

505

506 Hu, L., Xiang, Z., Wang, G., Rafique, R., Liu W., and Wang, C.: Changes in soil physicochemical and microbial properties
507 along elevation gradients in two forest soils, *Scandinavian Journal of Forest Research*, 31, 242–253, 2016.

504
505 Jones, A, Stolbovoy, V., Tarnocai, C., Broll, G., Spaargaren, O., and Montanarella, L. (eds.): Soil Atlas of the Northern
506 Circumpolar Region. European Commission, Publications Office of the European Union, Luxembourg, 144pp., 2010.
507
508 Khotimchenko, S. V., Vaskovsky, V. E., and Titlyanova, T. V.: Fatty acids of marine algae from the pacific coast of North
509 California. *Bot. Mar.*, 45, 17–22, 2002.

510 Kleidon, A. and Mooney, H. A.: A global distribution of biodiversity inferred from climatic constraints: results from a
511 process-based modelling study, *Global Change Biol.*, 6, 507–523, 2000.

512 Klimeš, L. and Doležal J.: An experimental assessment of the upper elevational limit of flowering plants in the western
513 Himalayas, *Ecography*, 33, 590–596, 2010.

514
515 Körner, Ch.: The use of ‘altitude’ in ecological research, *Trends in Ecol. And Evol.*, 22, 569–574, 2007.
516
517
518 Kroppenstedt, R. M.: Fatty acid and menaquinone analysis of actinomycetes and related organisms. In: Goodfellow, M.,
519 Minikin, D. E., Eds. *Chemical methods in bacterial systematics*. Academic Press, London, pp. 173–194, 1985.

520 Larsen, K. S., Jonasson, S., and Michelsešn, A.: Repeated freeze–thaw cycles and their effects on biological processes in two
521 arctic ecosystem types, *Appl. Soil Ecol.*, 21, 187–195, 2002.
522

523 Lazzaro, A., Hilfiker, D., and Zeyer, J.: Structure of microbial communities in alpine soils: seasonal and elevational effects,
524 *Frontiers in Microbiology*, doi: 10.3389/fmicb.2015.01330, 2015.

525
526 Ley, R. E., Williams, M. W., and Schmidt, S. K.: Microbial population dynamics in an extreme environment: controlling
527 factors in talus soils at 3750m in the Colorado Rocky Mountains, *Biogeochemistry*, 68, 313–335, 2004.
528

529 Lipson, D. A.: Relationships between temperature responses and bacterial community structure along seasonal and
530 altitudinal gradients, *FEMS Microbiol. Ecol.*, 59, 418–427, 2007.
531

532 Lipson, D. A., Schmidt, S. K., and Monson, R. K.: Links between microbial population dynamics and nitrogen availability in
533 an alpine ecosystem, *Ecology*, 80, 1623–1631, 1999.
534

535 Lipson, D. A., Schmidt, S. K., and Monson, R. K.: Carbon availability and temperature control the post-snowmelt decline in
536 alpine soil microbial biomass, *Soil Biol. Biochem.*, 32, 441–448, 2000.

537 Lipson, D. A., Schadt, C. W., and Schmidt, S. K.: Changes in soil microbial community structure and function in an alpine
538 dry meadow following spring snow melt, *Microb. Ecol.*, 43, 307–314, 2002.
539

540 Lovelland P. J. and Whalley W. R.: Particle size analysis. In: *Soil and Environmental Analysis Physical Method* (Eds Smith,
541 K.A. and Mullins, Ch.E.), Dekker Press, New York, USA, 2001.
542

543 Löffler, U. C. M., Cypionka, H., and Löffler, J.: Soil microbial activity along an arctic-alpine altitudinal gradient from a
544 seasonal perspective, *Eur. J. Soil Sci.*, 59, 842–854, 2008.

545 Ma, X., Chen, T., Zhang, G., and Wang R.: Microbial community structure along an altitude gradient in three different
546 localities, *Folia Microbiol.*, 49, 105–111, 2004.

547

548 Männistö, M., Tirola, M., and Häggblom, M. M.: Bacterial communities in Arctic fields of Finnish Lapland are stable but
549 highly pH-dependent, *FEMS Microbiol. Ecol.*, 59, 452–465, 2007.

550

551 Margesin, R., Gander, S., Zacke, G., Gounot, A. M., and Schinner, F.: Hydrocarbon degradation and enzyme activities of
552 cold-adapted bacteria and yeasts, *Extremophiles*, 7, 451–458, 2003.

553

554 Margesin, R., Jud, M., Tschirko, D., and Schinner, F.: Microbial communities and activities in alpine and subalpine soils,
555 *FEMS Microbiol. Ecol.*, 67, 208–218, 2009.

556

557 Meng, H., Li, K., Nie, M., Wan, J.-R., Quan, Z.-X., Fang, C.-M., Chen, J.-K., Gu, J.-D., and Li, B.: Responses of bacterial
558 and fungal communities to an elevation gradient in a subtropical montane forest of China, *Appl. Microbiol. Biot.*, 97, 2219–
559 2230, 2013.

560

561 Montgomery, H. J., Monreal, C. M., Young, J. C., and Seifert, K. A.: 2000. Determination of soil fungal biomass from soil
562 ergosterol analyses, *Soil Biol. Biochem.*, 32, 1207–1217, 2000.

563

564 Nemergut, D. R., Costello, E. K., Meyer, A. F., Pescador, M. Y., Weintraub, M. N., and Schmidt, S. K.: Structure and
565 function of alpine and arctic soil microbial communities, *Res. Microbiol.*, 156, 775–784, 2005.

566 Niklińska, M. and Klimek, B.: Effect of temperature on the respiration rate of forest soil organic layer along an elevation
567 gradient in the Polish Carpathians, *Biol. Fert. Soils*, 43, 511–518, 2007.

568

569 Nikrad, M. P., Kerkhof, L. J., and Häggblom, M. M.: The subzero microbiome: microbial activity in frozen and thawing
570 soils, *FEMS Microbiol. Ecol.*, 92, doi:10.1093/femsec/fiw08, 2016.

571

572 Oberbauer, S. F., Tweedie, C. E., Welker, J. M., Fahnestock, J. T., Henry, G. H. R., Webber, P. J., Hollister, R. D., Walker,
573 M. D., Kuchy, A., Elmore, E., and Starr, G.: Tundra CO₂ fluxes in response to experimental warming across latitudinal and
574 moisture gradients. *Ecol. Monogr.*, 77, 221–238, 2007.

575

576 Prach, K., Klimešová, J., Košnar, J., Redčenko O., and Hais M.: Variability of contemporary vegetation around
577 Petuniabukta, central Spitsbergen, *Pol. Polar Res.*, 33, 383–394, 2012.

578

579 Pushkareva, E., Kviderová, J., Šimek, M., and Elster, J.: Nitrogen fixation and diurnal changes of photosynthetic activity in
580 Arctic soil crusts at different development stage, *Europ Journal of Soil Biology*, 79, 21–30, 2017.

581

582 Richter, D. D., Johnson, D. W., and Dai, K. H.: Cation exchange reactions in acid forested soils: effects of atmospheric
583 pollutant deposition. In: Johnson, D.W., Lindberg, S.E. (eds.), *Atmospheric Deposition and Nutrient Cycling in Forest*
584 *Ecosystems*, Springer-Verlag, New York, pp 339–358, 1992.

585

586 Rontani, J.-F., Charriere, B., Petit, M., Vaultier, F., Heipieper, H. J., Link, H., Chaillou, G., and Sempéré, R.: Degradation
587 state of organic matter in surface sediments from the Southern Beaufort Sea: a lipid approach, *Biogeosciences*, 9, 3513–
588 3530, 2012.

589 Rousk, J., Brookes, P. C., and Bååth, E.: Contrasting soil pH effects on fungal and bacterial growth suggest functional
590 redundancy in carbon mineralization, *Appl. Environ. Microb.*, 75, 1589–1596, 2009.

592

593 Sawaske, S. R. and Freyberg, D. L.: Fog, fog drip, and streamflow in the Santa Cruz Mountains of the California Coast
594 Range, *Ecohydrology*, 8, 695–713, 2015.

595

596 Schimel, J. P. and Clein, J. S.: Microbial response to freeze-thaw cycles in tundra and taiga soils, *Soil Biol. Biochem.*, 28,
597 1061–1066, 1996.

598

599 Schinner, F.: Soil microbial activities and litter decomposition related to altitude, *Plant Soil*, 65, 87–94, 1982.

600

601 Schinner, F. and Gstraunthaler, G.: Adaptation of microbial activities to the environmental conditions in alpine soils,
602 *Oecologia*, 50, 113–116, 1981.

603

604 Schütte, U. M. E., Abdo, Z., Foster, J., Ravel, J., Bunge, J., Solheim, B., and Forney, L. J.: Bacterial diversity in a glacier
605 foreland of the high Arctic, *Mol. Ecol.*, 19, 54–56, 2010.

606

607 Shen, C., Xiong, J., Zhang, H., Feng, Y., Lin, X., Li, X., Liang, W., and Chu, H.: Soil pH drives the spatial distribution of
608 bacterial communities along elevation on Changbai Mountain, *Soil Biol. Biochem.*, 57, 204–211, 2013.

609

610 Siles, J. S. and Margesin, R.: Abundance and diversity of bacterial, archaeal, and fungal communities along an altitudinal
611 gradient in alpine forest soils: What are the driving factors? *Soil Microbiology*, 72, 207–220, 2016.

612

613 Singh, D., Takahashi, K., Kim, M., Chun, J., and Adams, J. M.: A hump-backed trend in bacterial diversity with elevation on
614 Mount Fuji, Japan, *Microb. Ecol.* 63, 429–437, 2012.

615

616 Sinsabaugh, R. L., Antibus, R. K., Jackson, C. R., Karpanty, S., Robinson, M., Liptak, M., and Franchini, P.: A β -sitosterol
617 assay for fine-root mass in soil, *Soil Biol. Biochem.*, 29, 39–44, 1997.

618

619 Sparling, G. P. and West, A. W.: A comparison of gas chromatography and differential respirometer methods to measure soil
620 respiration and to estimate the soil microbial biomass, *Pedobiologia*, 34, 103–112, 1990.

621

622 Stenberg, B., Johansson, M., Pell, M., Sjö Dahl-Svensson, K., Stenström, J., and Torstensson, L.: Microbial biomass and
623 activities in soil as affected by frozen and cold storage, *Soil Biol. Biochem.*, 30, 393–402, 1998.

624

625 Strickland, M. S. and Rousk, J.: Considering fungal:bacterial dominance in soils – Methods, controls, and
626 ecosystem implications, *Soil Biol. Biochem.*, 42, 1385–1395, 2010.

627 Šmilauer, P. and Lepš, J.: *Multivariate analysis of ecological data using CANOCO5*. Cambridge University Press,
628 Cambridge, 2014.

629 Ter Braak, C. J. F. and Šmilauer, P.: Canoco reference manual and user's guide: software for ordination, version 5.0.
630 Microcomputer Power, Ithaca, USA, 496 pp, 2012.

631

632 Trevors, J. T., Kevan, P. G., and Tam, L.: Microbial diversity across a Canadian sub-Arctic, isostatically rebounding, soil
633 transect. *Polar Science*, 4, 81–91, 2010.

634

635 Tytgat, B., Verleyen, E., Sweetlove, M., D'hondt, S., Clercx, P., Van Ranst, E., Peeters, K., Roberts, S., Namsaraev, Z.,
636 Wilmotte, A., Vyverman, W., and Willems, A.: Bacterial community composition in relation to bedrock type and macrobiota
637 in soils from the Sør Rondane Mountains, East Antarctica. *FEMS Microbiol. Ecol.*, 92, doi: 10.1093/femsec/fiw126, 2016.

638

639 Uhlířová, E., Šantrůčková, H., and Davidov, S. P.: Quality and potential biodegradability of soil organic matter
640 preserved in permafrost of Siberian tussock tundra. *Soil Biol. Biochem.*, 39, 1978–1989, 2007.

641

642 van Breemen, N., Driscoll, C. T., and Mulder, J.: Acidic deposition and internal proton sources in acidification of soils and
643 waters. *Nature*, 307, 599–604, 1984.

644

645 van de Heijden, M. G. A., Bardgett, R. D., and van Straalen, N. M.: The unseen majority: soil microbes as drivers of plant
646 diversity and productivity in terrestrial ecosystems. *Ecol. Lett.*, 11, 296–310, 2008.

647

648 Van Horn, D. J., Van Horn, M. L., Barrett, J. E., Gooseff, M. N., Altrichter, A. E., Geyer, K. M., Zeglin, L. H., and Takacs-
649 Vesbach, K.: Factors controlling soil microbial biomass and bacterial diversity and community composition in a cold desert
650 ecosystem: Role of geographic scale. *PLoS One*, 8, doi: 10.1371/journal.pone.0066103, 2013.

651

652 Väre, H., Vestberg, M., and Ohtonen, R.: Shifts in mycorrhiza and microbial activity along an oroarctic altitudinal gradient
653 in northern Fennoscandia. *Arctic Alpine Res.*, 29, 93–104, 1997.

654

655 Volkman, J. K.: Sterols in microorganisms. *Appl. Microbiol. Biot.*, 60, 495–506, 2003.

656

657 Volkman, J. K.: A review of sterol markers for marine and terrigenous organic matter. *Org. Geochem.*, 9, 83–99, 1986.

658

659 Vuorinen, K. E. M., Oksanen, L., Oksanen, T., Pyykönen, A., Olofsson, J., and Virtanen, R.: Open tundra persist, but arctic
660 features decline—Vegetation changes in the warming Fennoscandian tundra. *Global Change Biol.*, 23, 3794–3807, 2017.

661

662 Webb, M.: The influence of magnesium on cell division. 2. The effect of magnesium on the growth and cell division of
663 various bacterial species in complex media. *Microbiology*, 3, 410–417, 1949a.

664

665 Wheeler, K. A., Hurdman, B. F., and Pitt, J. I.: Influence of pH on the growth of some toxigenic species
666 of *Aspergillus*, *Penicillium* and *Fusarium*. *Int. J. Food Microbiol.*, 12, 141–150, 1991.

667

668 Xu, M., Li, X., Cai, X., Gai, J., Li, X., Christie, P., and Zhang, J.: Soil microbial community structure and activity along a
669 montane elevational gradient on the Tibetan Plateau. *Eur. J Soil Biol.*, 64, 6–14, 2014.

670

671 Yoshitake, S., Uchida, M., Koizumi, H., and Nakatsubo, T.: Carbon and nitrogen limitation of soil microbial respiration in a
672 High Arctic successional glacier foreland near Ny-Ålesund, Svalbard. *Pol. Res.*, 26, 22–30, 2007.

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Yu, Q., Epstein, H., Engstrom, R., and Walker, D.: Circumpolar arctic tundra biomass and productivity dynamics in response to projected climate change and herbivory, *Global Change Biol.*, 23, 3895–3907, 2017.

Zbiral, J. and Němec, P.: Data presentation, interpretation, and communication: Integrating of Mehlich 3 extractant into the Czech soil testing scheme, *Communications in Soil Science and Plant Analysis*, 31, 2171–2182, 2000.

707 **Tables**

708

709 **Table 1. Climatic variables; temperatures given in °C**

Sites [m a.s.l.]	Means Summer	Means Winter	Means Year	Min daily means Winter	Max daily means Summer	Mean daily amplitude Summer	Max daily amplitude Summer	Number of days with daily mean > 0 °C	Number of days with daily mean > 5 °C	Positive soil surface energy balance
25	5.8	-3.6	-0.8	-7.0	11.2	5.2	10.9	110	62	615
280	7.1	-5.7	-2.7	-10.3	14.5	8.5	18.2	96	54	571
520	5.8	-8.9	-4.9	-15.8	14.7	8.1	17.7	91	40	480
765	5.3	-9.5	-6.6	-17.1	11.6	5.5	14.0	51	11	290

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734 **Table 2. Geochemical characteristics of soils along the studied altitudinal transects (Tr1-Tr3). Means \pm SD (n = 3) are given in the**
 735 **upper part of the table. Results of two-way ANOVAs (F-values) of the effects of transect (Tr), altitude (Alt) and their interaction**
 736 **(Tr x Alt) are presented in the lower part of the table.**

transect	altitude [m.a.s.l.]	soil type	soil moisture [%]	pH	CEC [meq/100g ⁻¹]	Ca ²⁺ [mg g ⁻¹]	Mg ²⁺ [mg g ⁻¹]	K ⁺ [μg g ⁻¹]	Na ⁺ [μg g ⁻¹]
Tr1	25	sandy loam	a 28.4 ± 2.5	b 7.8 ± 0.1	a 35.8 ± 0.4	b 4.9 ± 0.2	c 0.50 ± 0.03	b 104 ± 2.3	a 16.0 ± 1.4
	275	sandy loam-loam	b 18.0 ± 0.5	b 7.9 ± 0.2	b 27.4 ± 2.3	b 5.2 ± 0.6	c 0.55 ± 0.08	b 81 ± 8.8	bc 8.4 ± 1.3
	525	loam	b 18.6 ± 2.5	b 8.1 ± 0.1	b 30.3 ± 0.7	b 4.3 ± 0.4	b 0.85 ± 0.04	a 160 ± 18.1	b 11.3 ± 1.1
	765	clay-loam	c 12.1 ± 1.8	a 9 ± 0.0	b 26.8 ± 2.3	a 19.8 ± 1.0	a 1.25 ± 0.06	c 11 ± 2.7	c 7.3 ± 0.0
Tr2	25	sandy loam	a 21.1 ± 2.4	c 7.8 ± 0.1	b 25.6 ± 2.7	b 14.7 ± 2.6	c 0.19 ± 0.01	ab 52 ± 4.0	a 13.2 ± 1.7
	275	sandy loam-loam	a 21.1 ± 2.4	c 7.9 ± 0.1	b 30.3 ± 1.7	ab 16.5 ± 1.1	b 0.26 ± 0.01	a 59 ± 4.3	ab 10.1 ± 1.7
	525	sandy loam-loam	a 21.7 ± 5.3	b 8.4 ± 0.1	b 30.8 ± 1.1	c 7.8 ± 1.6	a 0.34 ± 0.01	a 69 ± 3.3	ab 9.6 ± 1.8
	765	loam	a 22.5 ± 1.7	a 8.8 ± 0.1	a 45.1 ± 0.5	a 27.9 ± 9.3	b 0.25 ± 0.01	b 41 ± 8.8	b 8.1 ± 1.4
Tr3	25	sandy loam	a 39.5 ± 1.4	b 8.1 ± 0.1	a 49.4 ± 2.1	c 7.7 ± 0.3	a 0.20 ± 0.03	b 52 ± 5.3	a 17.1 ± 1.1
	275	sandy loam-loam	ab 31.9 ± 2.9	b 8.1 ± 0.1	b 39.2 ± 5.4	b 10.8 ± 0.6	a 0.21 ± 0.01	ab 59 ± 1.9	a 18.5 ± 0.5
	525	loam	ab 28.2 ± 6.5	b 8 ± 0.1	b 34.9 ± 3.0	ab 13.0 ± 4.6	a 0.22 ± 0.00	a 66 ± 6.6	a 18.4 ± 3.1
	765	loam	b 22.5 ± 1.7	a 8.8 ± 0.1	b 30.6 ± 3.9	a 14.2 ± 0.1	b 0.16 ± 0.00	b 52 ± 1.6	b 9.9 ± 0.2
d.f.									
Tr	2		31.4 ***	0.10	22.1 ***	6.43 **	63.4 ***	51.7 ***	36.2 ***
Alt	3		11.1 ***	98 ***	4.61 *	14.1 ***	66.9 ***	74.9 ***	18.7 ***
Tr x Alt	6		5.07 **	5.6 ***	20.5 ***	0.83	60.6 ***	31.6 ***	3.94 **

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 738 Different letters indicate significant differences between sampling sites along particular transects ($P < 0.05$; upper part of the table). Statistically significant
 739 differences are indicated by: * $P < 0.05$, ** $P < 0.01$, *** $P < 0.001$ (lower part of the table).
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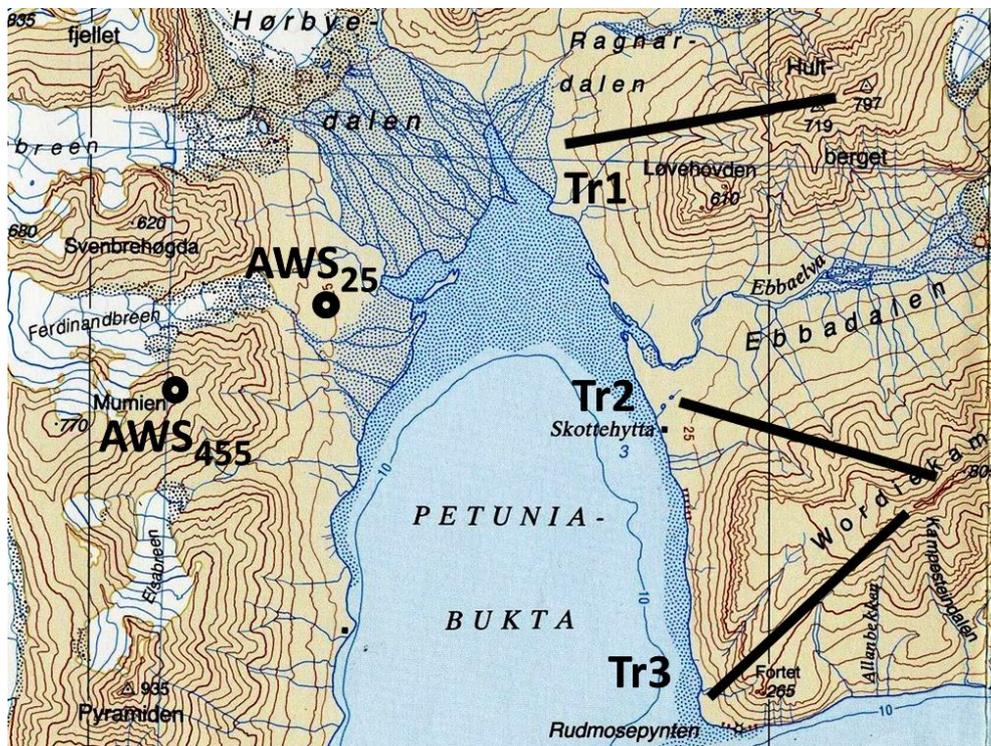
764 **Table 3. Total soil carbon (TOC) and nitrogen (TN) contents, their molar ratios, contents of sitosterol in TOC and sitosterol /**
 765 **brassicasterol ratios and soil PLFA contents in soils along the altitudinal transects (Tr1-Tr3). Means \pm SD (n = 3) are given in the**
 766 **upper part of the table. Results of two-way ANOVAs (F-values) of the effects of transect (Tr), altitude (Alt) and their interaction**
 767 **(Tr x Alt) are presented in the lower part of the table.**

transect	altitude [m.a.s.l.]	TOC [mg g ⁻¹]	TN [mg g ⁻¹]	TOC/TN	Sitosterol [μ g g ⁻¹ TOC]	Sitosterol / Brassicasterol
Tr1	25	c 70.6 \pm 13.4	b 5.0 \pm 1.01	b 12.1 \pm 0.2	c 534 \pm 62.8	b 5.5 \pm 0.4
	275	b 21.1 \pm 1.9	a 2.0 \pm 0.29	ab 9.0 \pm 0.7	bc 521 \pm 140	b 5.3 \pm 0.8
	525	b 18.5 \pm 4.2	a 1.8 \pm 0.31	ab 8.8 \pm 0.7	ab 293 \pm 66.5	b 4.7 \pm 1.0
	765	a 4.4 \pm 1.5	a 0.5 \pm 0.07	a 7.9 \pm 2.6	a 81.1 \pm 2.7	a 2.3 \pm 0.4
Tr2	25	ab 30.6 \pm 4.8	a 1.9 \pm 0.40	c 13.7 \pm 0.9	bc 515 \pm 44.9	b 6.7 \pm 0.7
	275	b 37.2 \pm 5.0	a 3.0 \pm 0.26	b 10.7 \pm 0.7	c 616 \pm 143	b 5.6 \pm 1.2
	525	a 24.4 \pm 7.8	a 1.9 \pm 0.64	b 9.8 \pm 1.2	ab 299 \pm 73.3	a 2.9 \pm 0.4
	765	a 21.6 \pm 3.6	a 2.8 \pm 0.20	a 6.7 \pm 0.6	a 161 \pm 36.9	a 2.7 \pm 0.7
Tr3	25	c 81.1 \pm 8.7	b 6.1 \pm 0.38	b 11.5 \pm 0.7	b 587 \pm 144	b 6.4 \pm 2.1
	275	b 62.2 \pm 9.1	ab 4.8 \pm 0.32	b 11 \pm 0.7	ab 370 \pm 42.9	a 4.2 \pm 0.7
	525	ab 39.6 \pm 11.4	a 4.8 \pm 0.32	b 10.6 \pm 0.6	a 270 \pm 112	a 3.3 \pm 1.0
	765	a 23.1 \pm 3.9	a 2.5 \pm 0.37	a 7.9 \pm 0.2	a 151 \pm 37.8	a 3.1 \pm 0.9
d.f.						
Tr	2	27.8 ***	31.5 ***	1.57	0.79	1.04
Alt	3	42.4 ***	26.4 ***	23.6 ***	28.4 ***	14.4 ***
Tr x Alt	6	8.33 ***	11.3 ***	1.96	1.34	2.17

768 Different letters indicate significant differences between sampling sites along particular transects ($P < 0.05$; upper part of the table). Statistically significant
 769 differences are indicated by: * $P < 0.05$, ** $P < 0.01$, *** $P < 0.001$ (lower part of the table).
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785 **Figures**

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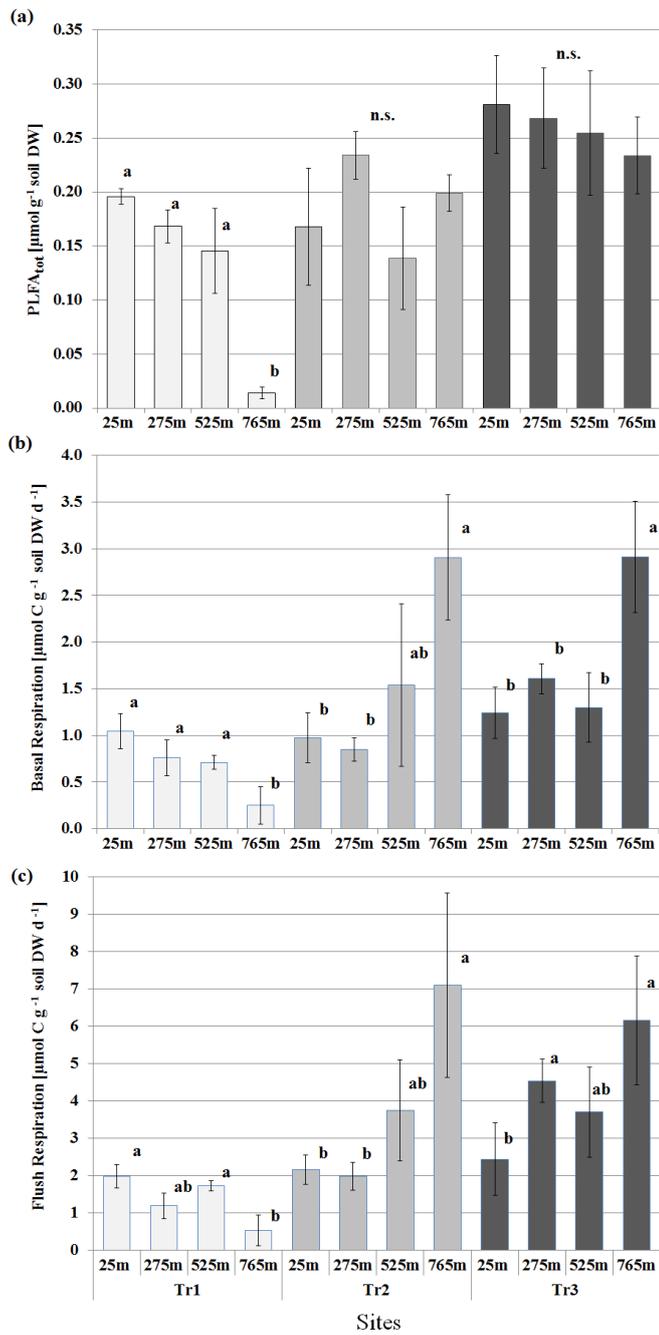


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788 **Figure 1.** Location of the three investigated transects Tr1–Tr3 and automated weather stations (AWS) in Petunia bay,
789 Billefjorden, Central Spitsbergen. Map source: map sheet C7, Svalbard 1:100 000, Norwegian Polar Institute 2008.

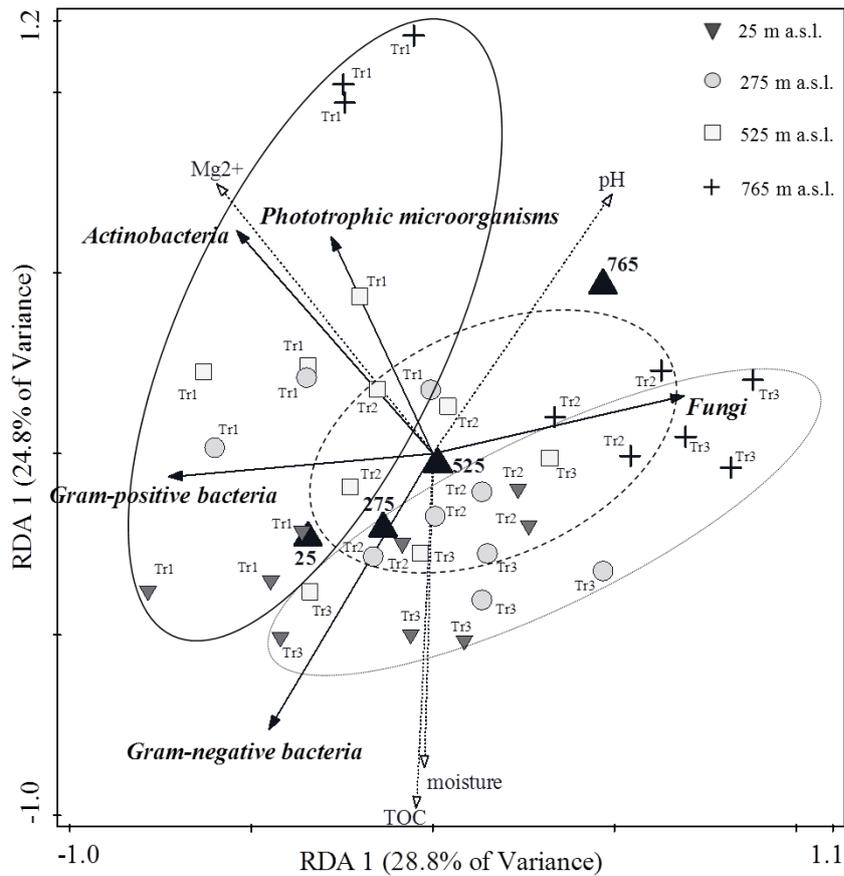
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793 **Figure 2.** The soil PLFA contents (a), the potential respiration rates (b) and the flush respiration rates (c) in the soils along
 794 altitudinal transects (Tr1-Tr3). Error bars indicate mean \pm SD (n = 3). Small case letters denote significant differences among
 795 altitudes within particular transects ($P < 0.05$; One-way ANOVA combined with Tukey post hoc test).



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797 **Figure 3. The correlation between abundance of main microbial groups (bold italic) and soil geochemical parameters retained by**
 798 **forward selection of explanatory variables. Results of RDA. Altitude of sampling sites was used as supplementary variable. Arrows**
 799 **indicate the direction in which the respective parameter value increases, solid lines indicate microbial groups, dotted lines indicate**
 800 **selected environmental variables. Up triangles are centroids of sites with corresponding elevation (n = 9), numbers indicate**
 801 **elevation (m a.s.l.). The thin solid line encases sites along the Transect 1 (Tr1), the dashed line encases sites along the T transect 2**
 802 **(Tr2), and the dotted line encases sites along the Transect 3 (Tr3). The numbers in parentheses are the portions of the variation**
 803 **explained by each axis.**

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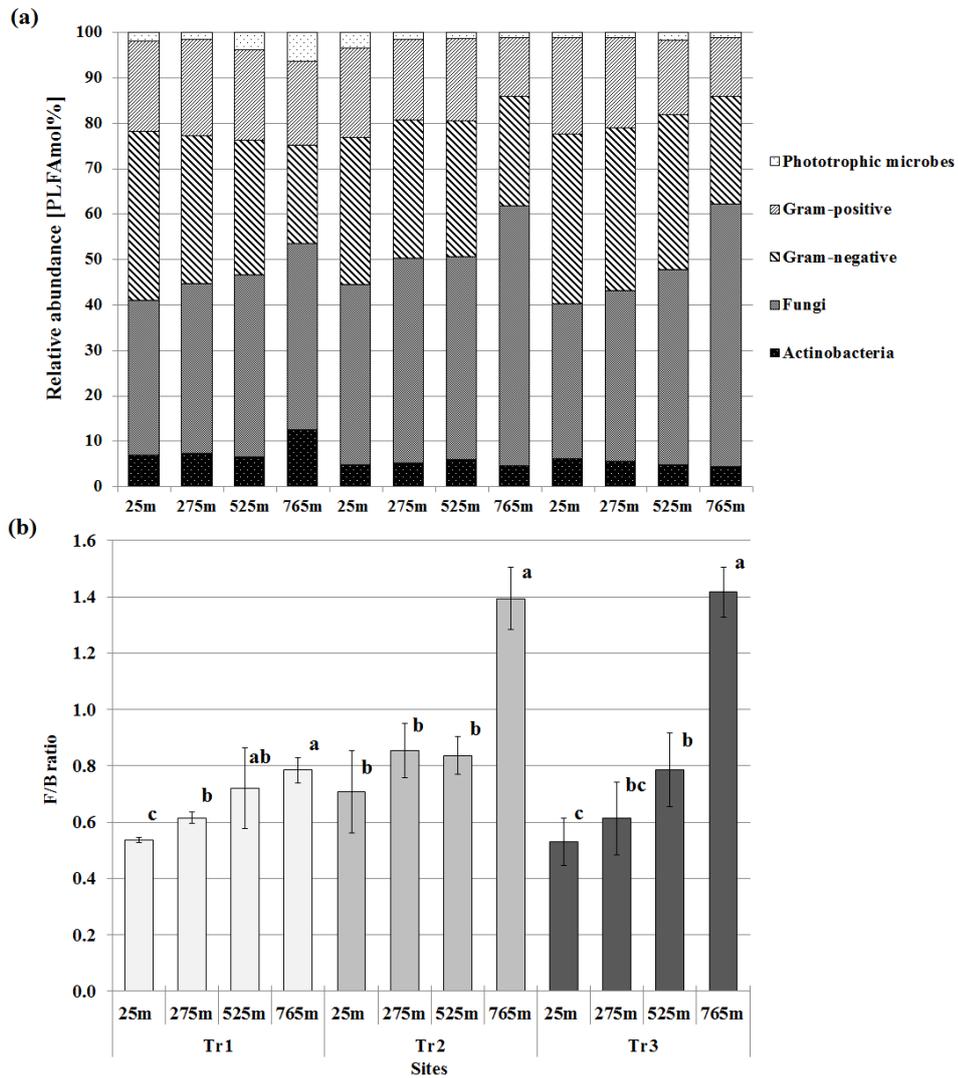
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Figure 4. Relative abundance of specific PLFAs within the microbial community (a), and fungi to bacteria (F/B) ratios (b) along altitudinal transects (Tr1-Tr3). Error bars indicate mean \pm SD ($n = 3$). Small case letters denote significant differences between altitudes within particular transects ($P < 0.05$; One-way ANOVA combined with Tukey post hoc test).