Revised version of Cooper et al (2017), 'Ferrihydrite associated organic matter (OM) stimulates reduction by *Shewanella oneidensis* MR-1 and a complex microbial consortia,' *Biogeosciences Discuss.*, doi.org/10.5194/bg-2017-270

This document contains 1) point-by-point responses to the reviews by referee #1, 2) point-by-point responses to the reviews by referee #2, 3) point-by-point responses to the reviews by the editor, and 4) the marked-up manuscript version including track changes.

Note 1: the line numbers in the responses to the referee and editor comments refer to line numbers in the marked-up manuscript.

Point-by-point responses to the reviews by referee #1 (Sarah Hayes):

-It seems like the bulk of the section 4.0 could be redistributed into the other sections of the paper. This section was reorganized/rewritten in a way that the information provided can now stand alone as an independent section instead of needing to be re-distributed to other sections of the manuscript. Please note, this is also addressed in our response to your comments regarding Line 329-345, 358-365 on the next page.

-Section 4.3 could probably be absorbed by the results section. And any mention in the discussion might be best integrated into another section.

Section 4.3 was removed from the discussion section. This information was integrated into the results section 3.4 as suggested.

-The conclusions section could also be streamlined to quickly revisit the key findings of the study.

We did some changes to the Summary and Conclusions (now section 4.3) to better streamline, however, we feel the way we presented our summary and overall conclusions in this section is effective in getting our message across and would prefer not to make significant changes to the text here.

Specific comments:

L11: should be "this study investigated to what extent"

Corrected (Line 11).

L14: should be "OM content" not "contents"

Corrected (line 14).

L48: should be "coprecipitation is the more common process"

Corrected (line 48).

L49-51: This sentence isn't super clear. Consider breaking it into two sentences or maybe just clarifying the end, something like "reducing agents potentially differ between ferrihydrite precipitated in the presence of OM and purely adsorbed OM.

This sentence was modified for clarity. The following sentence was added: 'The different properties of pure ferrihydrite and OM-ferrihydrite coprecipitates may lead to different behaviors during microbial reduction. Due to their smaller crystal size and more defective crystal structure, coprecipitates might faster dissolve. The associated organic material will change the mineral's surface properties, e.g., the surface charge, with consequences for the accessibility of Fe(III) to microbes, redox-active shuttling compounds, or extracellular enzymes.' (Line 49-52)

L61: Again, this sentence isn't quite right. Consider something like: "dissolved humic acid and the mineral/humic acid ratios have been shown by some studies to increase Fe (III) reduction rates (ref) while other studies have not reproduced this result (ref)." But I'm not sure it keeps the meaning of what you are trying to communicate.

This line was modified as suggested (Line 65-66).

L85: Consider revising to: "Further, this study found that reduction rates by Shewanella were correlated..."

This sentence was revised as suggested (Line 90).

L86-97: This section could flow a bit better. All the pieces are there, but it feels disjointed.

Modifications to this section were included to improve the flow (Line 94-101)

L94-97: This is an excellent way to close this paragraph!

Thank you!

L105-7: I think this surprising result (which is way cool!) should be moved to later in the paper, maybe the results section. It doesn't really fit here in the introduction.

We considered your suggestion to remove this sentence, but we prefer to leave this as the conclusion to the introduction.

L136: Add the information about the organic matter here (from lines L146-147 and from the discussion L379-82). Then maybe make the rest of the paragraph here with the information about the soil organic matter extraction first and then

Information about organic matter moved to Line 144.

L152: Should be "These cultures were..."

Corrected (Line 167).

L326: Add mineral formulas to the list of minerals observed.

Mineral formulas added to the list of minerals observed in XRD spectra (Line 360-361).

L331-4: This sentence is a little difficult to understand, please reword to improve clarity.

This section was modified to improve clarity. Please see section 3.4.

L352: I'm not sure of your exact meaning, please consider revising to something similar to these suggestions, depending on your meaning: "amendment with each of these substrates stimulates Fe(III) reduction..." or "amendment with all of these substrates stimulates Fe(III) reduction..."

Sentence modified as suggested ('amendment with each of these substrates stimulates Fe(III) reduction...') (Line 401-402).

L356: Consider "the substrate mixture used in these ferrihydrite..."

Sentence modified to reflect your suggestions (line 405-406).

L329-345: Would some or all of this material be better suited for introductory material? If not, try to make a stronger discussion point. L358-365: It seems like there is a topic change here that feels disjointed. The summary comments beginning on L 360 seem like they should precede the content earlier in the paragraph and are important to discuss in their own right.

Thanks for pointing out that this first section of the Discussion was a bit disjointed. We took some time to read through and decided the best way to address this was to rewrite/reorganize the information presented in an effort to better link the first paragraph with the overall outcome

of our study. The 'second paragraph' is now divided into two separate paragraphs which, in our opinion, are easier and more clear to read and provide a better link to the first paragraph as well as the subsequent sections in the discussion (Line 397-420).

L367: Consider: "OM stimulates growth and Fe(III) reduction by S. oneidensis MR-1 and the microbial consortia."

Sentence modified as suggested (line 442-443).

L373-4: Consider revising: "reducing conditions, which showed that S. putrefaciens was capable of using humic acids as an electron shuttle to enhance Fe(III) reduction. However, enhanced Fe reduction was only observed when the concentration of organic..."

Sentence modified as suggested. See lines 451-453.

L377: Consider revising: "which has lower aromatic moieties, and...

Sentence modified as suggested (Line 456).

L378: Consider revising: "electron shuttling capacity to..."

Sentence modified as suggested (Line 456-457).

L379-82: I think the info about the organic matter seems like it should be in the methods section.

Information removed from Section 4.1.

L383-386: seems like these details should also be in the methods, although they are also important here.

Sentence modified so that sufficient information remains here (Line 463), and more detailed information can be found in the methods (Section 2.2).

L394: The colon should probably be a period.

Corrected (Line 473).

L397: The comma between "directly" and "because" does not seem necessary.

Corrected (Line 476).

L405: Consider revising: "oxide surfaces: via direct contact..."

Sentence modified as suggested (line 484).

L406: The "or" should be an "and"

Corrected (line 485).

L428: Consider revising "Fe(III) oxides. However, ..."

Sentence modified as suggested (Line 508).

L429: "abundant" should be "abundance"

Corrected (Line 509).

L433: "inked" should be "linked"

Corrected (line 513).

L441: Should be "the question of whether"

Corrected (Line 521).

L441: consider revising: "bacteria remains unanswered."

Sentence modified as suggested (Line 4521-522).

L442: the citation at the end of the sentence is redundant of in text citation.

Citation at the end of the sentence removed (Line 522).

L474: Should be "formation from pure..."

Corrected (Line 372). Note, information in section 4.3 incorporated into results section 3.4.

L480-482: This is kinda new information, which should be mentioned earlier (maybe with the other content in section 4.3). Further, you need a citation because this study does not provide any direct evidence to support this claim (which is likely the case).

This information is now introduced in Section 4.1 (Line 447-449). References were added to support this claim.

Point-by-point responses to the reviews by anonymous referee #2:

Specific Comments:

Introduction

Line 33: should be "The majority of which is dispersed.."

Corrected (Line 33).

Line 44: should be "The coprecipitation of OM with Fe results in..."

Corrected (Line 44).

Line 49-51: This sentence is confusing. It seems like you are trying to make two separate arguments (1) coprecipitated ferrihydrite has different properties than pure ferrihydrite, and (2) coprecipitated ferrihydrite has different reactivity compared to ferrihydrite with sorbed OC. Maybe split into two sentences?

Sentence revised for clarification as follows: 'The different properties of pure ferrihydrite and OM-ferrihydrite coprecipitates may lead to different behaviors during microbial reduction. Due to their smaller crystal size and more defective crystal structure, coprecipitates might faster dissolve. The associated organic material will change the mineral's surface properties, e.g., the surface charge, with consequences for the accessibility of Fe(III) to microbes, redox-active shuttling compounds, or extracellular enzymes.' (Line 49-52).

Line 74-93: This paragraph has important information that you reply on in your discussion. Lines 94-97 clearly state the overall findings, but I found it a bit difficult to follow lines 74-93.

Modifications to this section were included to improve the flow (Line 94-101)

Line 101: should be "is the dominant electron accepting process"

Corrected (Line 104).

Methods

Line 110-112: This first sentence is confusing.

Sentence modified for clarification as follows: 'Peat cores were obtained from the Schlöppnerbrunnen fen (Northern Bavaria, Germany; 50°07'55''N, 11°52'52''E) using a Pürkheimer soil corer in May 2016. This minerotrophic, slightly acidic (pH ~5) fen has been previously described in detail (Blodau et al., 2004; Eusterhues et al., 2014; Hausmann et al., 2016; Küsel et al., 2008; Loy et al., 2004; Pester et al., 2012).' (Line 119-121).

Line 117: what is Corg?

'Solid organic carbon' added for clarification (Line 127).

Line 129: This first sentence seems like it should be last sentence of the previous paragraph.

The sentence is now the concluding sentence of the previous paragraph as suggested (Line 139-140).

Line 130-133: These sentences are out of place. I don't think you need to talk about DNA extraction or PCR here. Maybe move them to the "DNA extractions" and "Quantitative PCR" sections.

The descriptions regarding DNA extraction and subsequent qPCR analyses were moved to section 2.5 (Line 196-202).

Line 146: Move this up so that the reader knows that you performed the forest floor extract solutions.

Information about forest floor extract solution moved to Line 144 (detailed descriptions can be found in Esterhues, et al. 2014).

Line 149-156: You may want to add this paragraph to a section called "Preparation of pure cultures".

We appreciate the suggestion, however, we have decided to keep the information regarding pure culture preparation and the subsequent incubation set-up with both *S. oneidensis* and the microbial consortia extracted from peat in one section instead of separating the information into separate sections (2.4 Microbial OM ferrihydrite reduction experiments).

Line 156: I'm not sure why you write about "Defined freshwater medium" here if you also have a much more detailed description in lines 162-167.

Sentence modified to prevent repetition of detailed description as follows: 'Cells were washed twice with a defined medium, resuspended in 2 ml of medium and diluted to a final concentration of 2×10^5 cells ml⁻¹ in culture tubes, containing a defined medium.' (Line 168-170).

Line 176: If you are not showing any data (not even in the supplemental information), then why mention that you did this?

Secondary mineral formation during reduction is of interest in soil and sediment research. Many previous studies address this question. We therefore believe it is important to mention that we also paid attention to this aspect, although we did not find any new minerals. However, a table or a figure with 21 very similar diffractogramms does not seem necessary to us to document our findings.

179: Did you also collect XRD spectra of the starting materials?

Yes, XRD spectra was collected from the starting material. Section 2.10 (Line 264-265) and 3.4 (Line 360-361) were both modified to reflect this.

Line 199: move this sentence to the end so that it follows the order in which things were done. It will also make it easier for the reader to understand that you did the PCR and prepared the libraries yourself and just did the sequencing at the LGC Genomic GmbH.

PCR, library preparation, and sequencing was all done by LGC Genomic GmbH. Including this type of detailed information (i.e. primers, etc.) in the materials and methods is commonly done. Sentence added to this section to make it clear that LCG did the PCR, library prep, and subsequent sequencing.

Line 224: Did you normalize the data using qiime or any other method? If you did, you should mention this because it is important when comparing the abundance of different taxa across various samples.

Yes, the data was normalized using qiime. An additional sentence was added to explain how differentially abundant taxonomic groups were identified after the metadata was assembled for downstream analysis in qiime (Line 252-256).

Results The order of these sections has a great flow. However, I would consider merging the Results and Discussion sections into one section. This will avoid many repetitive statements.

We opted to keep the results and discussion sections as independent section and based on comments you provided as well as the other reviewer we hope the repetitiveness is minimized while maintaining the flow of the results and subsequent discussion section.

Section 3.1: The reduction rates stated in this paragraph are not the same as the ones in Table 1.

Reduction rates stated in paragraph corrected to reflect those found in Table 1.

Section 3.3: When reading this section and flipping to the figures, I found it difficult to determine if I was looking at the correct figure.

Please double check the Figure numbers that you refer to in the text. For example, in line 285 you mention beta diversity and Figure 5b and I think you meant to reference Figure 6b.

Figure numbers double-checked. Thanks for pointing this out. Additionally, it's important to note that all figures were re-numbered due to merging of Figures 1 and 2, 3 and 4.

Discussion

Line 373: Add reference.

References added (Line 452).

Line 394: ":" should be "."

Corrected (Line 473).

Line 404: Add reference. Line 405: Add reference.

References added (Line 486).

Line 435: Have you considered removing singletons and see what the data looks like?

Yes, we did consider this and added now the following sentences to section 2.8 (Microbiome profiling): 'After alpha diversity analyses, the dataset was stripped by removing OTUs that were represented by less than three sequences. Relative abundances were calculated and standardized based on median sequencing depth across all samples. Datasets were not rarefied to prevent a loss of sensitivity in downstream sample-wise comparisons (McMurdie and Holmes, 2014). The dataset was stripped by removing OTUs that were only identified in single samples (i.e. singletons).' (Line 243-247). Additionally, the impact on the inclusion of singletons is addressed in Lines 513-515.

Section 4.3: Include mineral formulas and consider citing the following paper: Influence of Coprecipitated Organic Matter on Fe2+(aq)Catalyzed Transformation of Ferrihydrite: Implications for Carbon Dynamics Chunmei Chen, Ravi Kukkadapu, and Donald L. Sparks Environmental Science & Technology 2015 49 (18), 10927-10936 DOI: 10.1021/acs.est.5b02448

Mineral formulas added to the list of minerals observed in XRD spectra (Line 360-363). Citation also added as suggested.

Line 480: Where is the evidence of this finding? Maybe consider placing this information in the section about mineral composition.

We introduced this information Section 4.1 (Line 447-453) and references were added to support this claim.

Line 500-503: This last sentence is not clear.

The last sentence was rewritten for clarity as follows: 'This study also revealed that natural soil OM on the surface of ferrihydrites provides enough redox active groups to maintain microbial Fe(III) reduction processes, instead of passivating the surface. While at the same time the OM enhanced overall microbial growth and shaped the microbial community structure.' (Line 580-585).

References

Consider citing more recent studies such as: Influence of Coprecipitated Organic Matter on Fe2+(aq)-Catalyzed Transformation of Ferrihydrite: Implications for Carbon Dynamics Chunmei Chen, Ravi Kukkadapu, and Donald L. Sparks Environmental Science & Technology 2015 49 (18), 10927-10936 DOI: 10.1021/acs.est.5b02448

Dinesh Adhikari, Qian Zhao, Kamol Das, Jacqueline Mejia, Rixiang Huang, Xilong Wang, Simon R. Poulson, Yuanzhi Tang, Eric E. Roden, Yu Yang, Dynamics of ferrihydrite-bound organic carbon during microbial Fe reduction, Geochimica et Cosmochimica Acta, Volume 212, 2017, Pages 221-233, ISSN 0016-7037,

http://dx.doi.org/10.1016/j.gca.2017.06.017.(http://www.sciencedirect.com/science/article/pii/S001670 3717303666)

Weinan Pan, Jinjun Kan, Shreeram Inamdar, Chunmei Chen, Donald Sparks, Dissimilatory microbial iron reduction release DOC (dissolved organic carbon) from carbon-ferrihydrite association, Soil Biology and Biochemistry, Volume 103, 2016, Pages 232-240, ISSN 0038-0717, http://dx.doi.org/10.1016/j.soilbio.2016.08.026.

(http://www.sciencedirect.com/science/article/pii/S0038071716302115) Keywords: Carbon sequestration; DOC-Mineral stability; Shewanella oneidensis MR-1; Dissimilatory iron reduction; Anaerobic conditions

The first two suggested references (Chen, et al. and Adhikari, et al.) were added to the manuscript. Thank you for pointing out that they were missing.

Figures Consider merging Figures 1 and 2 into a 4-panel Figure. This will allow the reader to compare the results obtained from pure cultures and natural inoculum.

Figures 1 and 2 combined into a single 4-panel figure (see Figure 1). Additionally, we modified the figure caption to reflect the information presented in the merged figure.

Consider merging Figures 3 and 4 into a 4-panel Figure (make sure the y-axis is the same for all). This will allow the reader to compare the results obtained from pure cultures and natural inoculum. You may also want to consider changing T0, T middle and Tend to day 0, day X and day 288. This will keep the axis consistent with the other figures.

Figures 3 and 4 combined into a single 4-panel figure (see Figure 2). Figure caption modified to reflect information presented in the merged figure.

Figure 5C. The results illustrated in this Figure were not clear to me. Perhaps state the taxonomic level of the x-axis? Or add different shapes to differentiate between CFh vs Fh and AFh vs Fh?

The taxonomic level is for each portion of Figure 3 (previously Figure 5) is described in the figure caption (i.e. 3A – prokaryotic phyla level; 3B – family level; 3C – genus level).

Figure 6. I understand the importance of using various alpha diversity parameters. However, I am not sure illustrating them all in a main figure is necessary. Perhaps choose one parameters and talk about that one in the paper. The rest could go in the supplemental materials.

This figure was reconfigured to only include Observed OTUs, Chao1 and Shannon indices for simplicity (Fig. 4A). The main text (Section 3.3) and figure caption was modified to reflect these changes.

Point-by-point responses to the reviews by the editor (Denise Akob):

L. 41, 56 and elsewhere: correct "n.d." for the in text citation and update the reference in the bibliography

Reference removed from manuscript.

L. 120: specific here that you are extracting live organisms for cultivation

Corrected (L 137-138).

L. 129: what do you mean by "duplicated triplicates"?

Text changed for clarification. Sentence now states we used six replicates, instead of duplicated triplicates. (L 140).

L. 138-139: I'm not understanding what "AFhA 0.4, AFhB 1.3, and AFhD 4.2" means—are the letters extraction names? If so, define this (applies to L. 141 too). –

Abbreviations defined as requested (L 148, 151).

L. 143: were these actually hungate tube or Balch tubes? The ones with the big blue rubber stoppers are actually Balch tubes.

Thanks for pointing out this mistake. These are Balch tubes. The text has been changed to reflect this (L 155).

L. 144-145: define all abbreviations at first use

Abbreviations defined as requested (L 155-159, 182).

L. 154 and elsewhere: please give all centrifugation speeds as "x g"

Corrected throughout text where necessary.

L. 185: add a space between Bacterial and 16S

Space added between Bacterial and 16S (L 208).

L. 201: µl

Corrected (L 225).

L. 325: can you provide any more information about the database?

A more detailed description of the databased name and source is now provided in the text, along with a reference (L 265-266).

L. 883: remove the duplicate citation

Duplicate citation removed.

L. 265: add a space between Bacterial and 16S

Space added between Bacterial and 16S (L 300).

L. 298: it should be Gammaproteobacteria

Text changed to Gammaproteobacteria (L 330).

L. 392 and 413: remove the comma after the author name

Commas removed after author names (L 471, 492).

L. 402-403: provide a reference for Geobacter and nanowires.

References added (L 482-484).

Ferrihydrite associated organic matter (OM) stimulates reduction by Shewanella oneidensis MR-1 and a complex microbial consortia

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Abstract. The formation of Fe(III) oxides in natural environments occurs in the presence of natural organic matter (OM), resulting in the formation of OM-mineral complexes that form through adsorption or coprecipitation processes. Thus, microbial Fe(III) reduction in natural environments most often occurs in the presence of OM-mineral complexes rather than pure Fe(III) minerals. In this study we investigated to whatish extent the content of adsorbed or coprecipitated OM on ferrihydrite influences the rate of Fe(III) reduction by Shewanella oneidensis MR-1, a model Fe(III)-reducing microorganism, in comparison to a microbial consortium extracted from the acidic, Fe-rich Schlöppnerbrunnen fen. We found that increased OM contents led to increased rates of microbial Fe(III) reduction by S. oneidensis MR-1 in contrast to earlier findings with the model organism Geobacter bremensis. Ferrihydrite-OM coprecipitates were reduced slightly faster than ferrihydrites with adsorbed OM. Surprisingly, the complex microbial consortia stimulated by a mixture of electrons donors (lactate, acetate, and glucose) mimics S. oneidensis under the same experimental Fe(III)-reducing conditions suggesting similar mechanisms of electron transfer whether or not the OM is adsorbed or coprecipitated to the mineral surfaces. We also followed potential shifts of the microbial community during the incubation via 16S rRNA gene sequence analyses to determine variations due to the presence of adsorbed or coprecipitated OM-ferrihydrite complexes in contrast to pure ferrihydrite. Community profile analyses showed no enrichment of typical model Fe(III)-reducing bacteria, such as Shewanella sp. or Geobacter sp., but an enrichment of fermenters (i.e. Enterobacteria) during pure ferrihydrite incubations which are known to use Fe(III) as an electron sink. Instead, OM-mineral complexes favoured the enrichment of microbes including *Desulfobacteria* and *Pelosinus* sp., both of which can utilize lactate and acetate as an electron donor under Fe(III) reducing conditions. In summary, this study shows that increasing concentrations of OM in OM-mineral complexes determines microbial Fe(III) reduction rates and shapes the microbial community structure involved in the reductive dissolution of ferrihydrite. Similarities observed between the complex Fe(III)-reducing microbial consortia and the model Fe(III)-reducer S. oneidensis MR-1 suggest electron shuttling mechanisms dominate in OM-rich environments, including soils, sediments, and fens, where natural OM interacts with Fe(III) oxides during mineral formation.

1 Introduction

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Iron is the fourth most abundant element in the Earth's crust, indicating the role of Fe-cycling in nature is both ubiquitous and widespread. The majority of which is are dispersed as iron minerals in natural environments, including soils and sediments with ferrous iron [Fe(II)] and ferric iron [Fe(III)] as the primary redox states (Davison, 1993; Stumm and Sulzberger, 1992; Weber et al., 2006a). Microbially mediated and abiotic redox cycling of Fe play crucial roles in global biogeochemistry, for example, preservation and degradation of organic matter, and the fate of nutrients as well as contaminants for living organisms (Borch et al., 2010; Melton et al., 2014). Iron minerals such as Fe(III) oxides exhibit a high reactivity towards dissolved

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organic matter (OM), thus Fe(III) oxides are either partially or completely covered by OM in natural environments (Eusterhues et al., 2005; Kaiser and Zech, 2000; Lalonde et al., 2012; Torn et al., 1997). Fe(III) oxides enveloped by OM potentially lead to changes in the surface properties in comparison to non-modified Fe(III) oxides, which ultimately may influence mobility, solubility, and aggregation (Eusterhues et al., 2014; Narvekar et al., 2017; Pédrot et al., 2011; Vindedahl et al., 2016). The poorly crystalline ferrihydrite is one of the most common Fe(III) oxides and typically forms aggregates of nanometer-sized individual crystals (Bigham et al., 2002; Cornell and Schwertmann, 2003; Jambor and Dutrizac, 1998). The coprecipitation of OM with Fe results in the adsorption and occlusion of organic molecules within the interstices between individual ferrihydrite crystals, while adsorption of OM occurs on pre-existing ferrihydrite surfaces (Eusterhues et al., 2014). Moreover, the presence of dissolved OM impedes ferrihydrite growth (Cismasu et al., 2011; Eusterhues et al., 2008; Mikutta et al., 2008; Schwertmann et al., 2005), therefore, ferrihydrite with coprecipitated OM have smaller crystal sizes and more crystallographic defects. Since ferrihydrite is formed usually in OM-rich environments coprecipitation is the more common process in nature (Eusterhues et al., 2008). The different properties of pure ferrihydrite and OM-ferrihydrite coprecipitates may lead to-a different behaviours during microbial reduction. Because of tDue to their smaller crystal size and more defective crystal structure, coprecipitates might faster dissolve. The associated organic material will change the mineral's surface properties, e.g., the surface charge, with consequences for the accessibility of Fe(III) to Many properties differ between coprecipitated ferrihydrite and compared to pure ferrihydrite, thus theineluding the accessibility and solubility of ferrihydrite surfaces, along with the In addition the accessibility of the adsorbed/occluded OM to microbes, redox-active shuttling compounds, or extracellular enzymes, or reducing agents potentially differ is similarly affected from ferrihydrites with purely adsorbed OM.

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Not only do the mechanisms through which OM interacts with Fe(III) oxides, e.g. adsorbed or coprecipitated, influence the reactivity of the mineral-OM complex, the composition of OM may also effect its reactivity. The redox reactivity of OM is attributed mostly to quinones. Past experiments have often been performed using humic acids, i.e., alkaline extracts of natural organic material, which usually contain a higher concentration of the aforementioned redox active functional groups, than natural OM from soil (Klüpfel et al., 2014; Lehmann and Kleber, 2015; Piepenbrock et al., 2014). Humic acids should therefore be more effective in electron shuttling than natural soil organic matter. Previous studies have shown that the presence of dissolved humic acids leads to complexation of Fe(II) (Royer et al., 2002), and complexation and dissolution of Fe(III) (Jones et al., 2009), but also potentially enhance Fe(III) reduction via electron shuttling (Hansel et al., 2004; Jiang and Kappler, 2008; Lovley et al., 1996; Roden et al., 2010). Furthermore, the concentrations of dissolved humic acid or the mineral/humic acid ratios have been shown by some studies to increase Fe(III) reduction rates while other studies have not reproduced this resultwhere shown to influence the possibility of humic acids increasing Fe(III) reduction rates or not (Amstaetter et al., 2012; Jiang and Kappler, 2008). Despite the aforementioned studies, the influence of mineral bound OM on Fe(III) reduction rates and mineral transformation is not fully understood in regards to the broad diversity of Fe(III) reducing microorganisms present in nature (Weber et al., 2006a, 2006b).

Fe(III) oxides as well as OM-mineral complexes are not able to pass through the outer membrane of Fe(III)-reducing microorganisms, therefore different electron transfer pathways have evolved to alleviate this physiological problem

(Poggenburg et al., 2016; Weber et al., 2006b). Several mechanisms used as electron transfer pathways have been characterized including (i) direct transfer of electrons via enzymes anchored to the outer membrane of the bacterial cell (Myers and Nealson, 1988; Nevin and Lovley, 2000) or conductive nanowires (El-Naggar et al., 2010; Gorby et al., 2006; Reguera et al., 2005) and (ii) mediated electron transfer via exogenous electron shuttling compounds found in natural OM (Lovley et al., 1996), endogenous electron shuttling compounds, for example flavins (von Canstein et al., 2008; Kotloski and Gralnick, 2013; Marsili et al., 2008; Nevin and Lovley, 2002; Newman and Kolter, 2000), or the use of Fe(III)-chelating compounds, such as siderophores (Kraemer, 2004; Lovley et al., 1994; Nevin and Lovley, 2002).

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Shewanella and Geobacter species represent two of the best studied model Fe(III) reducing microorganisms. Shewanella species have been shown to utilize endogenous electrons shuttles, such as flavins, rather than transferring electron via direct contact between the mineral and the microbe (von Canstein et al., 2008; Kotloski and Gralnick, 2013). Not surprisingly, microbial reduction of Fe(III) oxides like nanometer-sized lepidocrocite and ferrihydrite by Shewanella species can be boosted by the presence of humic acids (Adhikari et al., 2017; Amstaetter et al., 2012; Pédrot et al., 2011; Shimizu et al., 2013) and quinone moieties (Newman and Kolter, 2000) when the amount of dissolved organic carbon is greater than a threshold concentration of 5-10 mg C L⁻¹ (Jiang and Kappler, 2008). Conversely, a study by Eusterhues, et al (2014) reported high C/Fe ratios in coprecipitated and adsorbed ferrihydrite complexes inhibited the reduction of ferrihydrite by G. bremensis. Poggenburg et al. (2016) produced ferrihydrite OM complexes with different organic materials (extracted extracellular polymeric substances, water extracts of soil litter) and used S. putrefaciens and G. metallireducens to reduce them. Their results confirmed that ferrihydrite-bound OM increases reduction rates of Shewanella, whereas it protects the mineral surface from reduction by Geobacter. Further, this study found that reduction rates For reduction by Shewanella this study found that reduction rates were correlated to the aromaticity of the mineral-bound OM. Geobacter species have been shown to use direct contact and conductive nanowire structures for electron transfer (Boesen and Nielsen, 2013; Malvankar et al., 2011). Unlike Shewanella, Geobacter do not produce endogenous electron shuttling or chelating compounds (Nevin and Lovley, 2000). however, in the absence of Fe(III) oxides —Geobacter are capable of reducing both humic acid (Jiang and Kappler, 2008; Loyley et al., 1996) and extracellular quinone moieties (Scott et al., 1998) in the absence of Fe(III) oxides. Furthermore, tThe decreased rates of Fe(III) reduction of OM ferrihydrite complexes by G. bremensis in incubations containing OM-ferrihydrite complexes compared to pure ferrihydrite minerals suggest that Geobacter do not use external electron shuttles in natural environments rich in OM and Fe(III) oxides. The aforementioned results implicate the Since the effects of OM on Fe(III) reduction rates vary between the two model organisms, *Shewanella* and *Geobacter*, presumably due to the various mechanisms of electron transfer used. 3 we should move forward by studying the response of a diverse community of Fe(III) reducers living in a habitat rich in Fe(III) oxides and OM. Taken together, these studies indicate (i) that it depends on the microorganisms present and their Fe respiratory pathways, whether OM on Fe(III) oxide mineral surfaces hinders or promotes Fe(III) oxide reduction. (ii) When electron shuttling is the dominant electron accepting processtransfer mechanism, a high content of redox active groups in the mineral-associated OM speeds up reduction.

In this study, we used the same experimental design described in the earlier study of Eusterhues et al., 2014, but replaced *G. bremensis* with *S. oneidensis* MR-1 to test the effect of increasing amounts of adsorbed and co-precipitated OM on microbial Fe(III) reduction. In a parallel set of incubation studies we used a microbial consortia as inoculum derived from an iron-rich peatland where the microbial reduction of Fe(III) is the dominant electron accepting processes for the degradation of OM under anoxic conditions. Thus, examining the respiratory behaviour of *S. oneidensis* MR-1 compared to a microbial consortia allows us to study the response of a diverse community of Fe(III) reducers living in a habitat rich in Fe(III) oxides and OM. This Schlöppnerbrunnen fen has been studied in great detail during the last 10 years, however, the impact of OM on the microbially mediated Fe(III) reduction processes in this fen are not fully understood. We compared rates of Fe(III) reduction of ferrihydrite adsorbed and co-precipitated with OM with pure ferrihydrite and followed qualitative and quantitative changes in the microbial community during incubation. To our great surprise, rates observed with *Shewanella* matched nearly exactly those observed with the complex consortia derived from Schlöppnerbrunnen fen, although *Shewanella* was not the dominant Fe(III)-reducing microorganism detected at the beginning and end of incubation.

2 Methods

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2.1 Study site and sampling procedure

Peat cores were obtained from the Schlöppnerbrunnen fen (Northern Bavaria, Germany; 50°07'55''N, 11°52'52''E); using a Pürkheimer soil corer in May 2016. This a-minerotrophic, slightly acidic (pH ~5) fen has been previously described in detail (Blodau et al., 2004; Eusterhues et al., 2014; Hausmann et al., 2016; Küsel et al., 2008; Loy et al., 2004; Pester et al., 2012), using a Pürkheimer soil corer in May 2016. Briefly, granite bedrock is covered by a fibric histosol, often referred to as peat, with a thickness between 40 cm and 70 cm. This fen is fed by anoxic, Fe(II)-rich groundwater (Küsel et al., 2008). The sampling site within the fen used in the current study (site "M") is positioned southwest of previously described drying-rewetting experimental plots (Knorr and Blodau, 2009; Reiche et al., 2008). Previous studies have reported that high concentrations of Fe(II) (up to 7.4 mM, mean 2001-2004 0.3 mM), dissolved organic carbon (DOC; -(up to 48 mM, mean 2001-2004 5.1 mM) and-solid organic carbon (Corg (; mean 37%) are routinely found at site "M" (Küsel et al., 2008; Reiche et al., 2008). Peat cores were transported on ice under anoxic conditions (N₂) and peat samples were either processed or stored at -20°C until use.

2.2 Microbial extraction from peat core samples

The peat microbial community was extracted using a standard protocol adapted to soil-microbial extractions. Briefly, a 20 mL syringe, with tip removed, was used as a manual short core sampler to remove 20 mL fractions (~25 g) from bulk peat cores and added to sterile, anoxic 125 mL serum bottles containing 100 mL 0.85% NaCl solution and 8 g sterile glass beads. The slurry was flushed with Ar and shaken at 225 rpm at 4 °C overnight. Samples were transferred to 50 mL Falcon tubes and centrifuged at increasing speeds to observe sedimentation rates. The samples were initially centrifuged at 900 x g for 5 minutes.

The sample was decanted into a fresh Falcon tube and the supernatant was centrifuged again at 1200 x g for 10 min. Next, the solid phase was resuspended in 10 mL anoxic 0.85% NaCl solution. The extract-, containing live organisms subsequently used for cultivation experiments, was transferred to 30 mL serum bottles and flushed with Ar to ensure anoxic conditions (Supplementary Fig. 1). The presence of bacteria in the peat microbial extract was confirmed via SYTO (50 µM) staining and microscopy. The peat microbial extractions were performed in duplicated trisix replicates to ensure sufficient inoculum was available for subsequent incubations. Genomic DNA was obtained from the peat microbial extract using the PowerSoil DNA isolation kit (MO BIO Laboratories, Carlsbad, CA, USA) according to manufacturer's instructions and stored at -20°C. DNA extracts were subsequently used as a template for quantitative PCR (qPCR) reactions targeting bacterial 16S rRNA genes to quantify abundance of bacteria over time or microbial community analysis via Illumina sequencing.

2.3 Synthesis of ferrihydrite and organic matter (OM)-mineral complexes

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A water extract of the forest floor (Oa and Oe horizons) of a Podzol under spruce and 2-line ferrihydrite was produced as previously described by Eusterhues, et al. (2014). Briefly, the 2-line ferrihydrite was produced via titration of 0.01 M Fe(NO₃)₃ solution with 0.1 M NaOH to pH 5 with constant stirring. Ferrihydrites with varying concentrations of adsorbed OM (AFh) were produced by mixing forest floor extract solutions containing increasing C concentrations with suspensions of freshly precipitated 2-line ferrihydrite at pH 5. Molar C/Fe ratios of these initial solutions were AFhA 0.4, AFhB 1.3, and AFhD 4.2. Coprecipitated ferrihydrites (CFh) were produced by dissolving Fe(NO₃)₃ in forest floor extract solutions of increasing concentrations and adding 0.1 M NaOH under vigorous stirring until the solution reached pH 5. Molar C/Fe ratios of these initial solutions were CFhA 0.4, CFhB 1.3, and CFhD 4.2. The solid/solution (g/L) ratio for all synthesized minerals ranged between 0.3 and 1.6. The solid products were separated using centrifugation, washed two times with deionized H₂O, freeze dried, and stored in hungate-Balch-type tubes under anoxic conditions (N₂). The C content was measured using a Carbon-Nitrogen (CN)-Analyzer (Vario EL, Elementar-Analysensysteme, Hanau, Germany) and the Fe content by Inductively coupled plasma optical emission spectrometry (ICP-OES) after total digestion (Spectroflame, Spectro, Kleve, Germany). A characterization of the forest floor extract and ferrihydrite-OM complexes by Fourier-transform infrared spectroscopy (FTIR) and X-ray photoelectron spectroscopy (XPS) can be found in Eusterhues et al. (2014). Note, the extraction of soil organic matter was also performed as described by Eusterhues et al. (2014) using the Oa and Oe layers of a Podzol under spruce located near Freising, Germany. The organic matter was shown to have an electron accepting capacity of EAC = 32 ±1 mmol e⁻ (mol C) and an electron donating capacity of EDC = 24.7±1.2 mmol e⁻ (mol C) (see WEOMwinter) (Fritzsche et al., n.d.).

2.4 Microbial OM ferrihydrite reduction experiments

We used *S. oneidensis* MR-1 as model Fe(III)-reducer since it is well-characterized and able to utilize a variety of carbon substrates as electron donors (Myers and Nealson, 1988, 1990; Venkateswaran et al., 1999). Liquid aerobic pre-cultures were grown in 50 mL of lysogeny broth (Luria Bertani medium; LB medium) containing 10 g tryptone, 10 g NaCl, 5 g yeast extract. These cultures were incubated on a rotary shaker at 150 rpm and the cell concentration was determined by optical density

(OD) measurements at 600 nm. For the Fe(III) reduction experiments 2 ml of LB-grown pre-culture was harvested after 14 h at late exponential growth phase and centrifuged at 10,000 x g for 5 min. Cells were washed twice with a defined freshwater medium, resuspended in 2 ml of medium and diluted to a final concentration of 2 x 10⁵ cells ml⁻¹ in culture tubes_z containing a defined freshwater medium. which was amended with Defined freshwater medium (pH 6.8; Geobacter medium ATCC 1957) was prepared with a combination of electron donors (2 mM lactate, 2mM acetate, 2 mM glucose) and NaHCO₃ to adjust the pH. For all incubations a defined freshwater medium (pH 6.8; Geobacter medium ATCC 1957) containing 1.5 g L⁻¹ NH₄Cl and 0.1 g L⁻¹ KCl, was used. After autoclaving and cooling under anoxic conditions (N₂/CO₂ (80 / 20 v/v) atmosphere), 30 mL L⁻¹ of 1 M NaHCO₃ (autoclaved, CO₂), 10 mL L⁻¹ Wolfe's vitamin solution (ATCC 1957), and 10 mL L⁻¹ modified Wolfe's minerals (ATCC 1957). A combination of Ffilter-sterilized (0.2 μm) electron donors (sodium-acetate (2 mM), sodium lactate (2mM), and glucose (2mM)) were added as carbon sources to all incubations to ensure standardization. NaH₂PO₄ was not added to circumvent interaction of PO₄³⁻ with ferrihydrite.

In the current study the microbial community extracted from peat cores collected in the Schlöppnerbrunnen fen as well as *S. oneidensis* MR-1 were used as the inoculum in incubations containing either ferrihydrite, OM-adsorbed ferrihydrite, or OM-co-precipitated ferrihydrite. For OM-adsorbed ferrihydrite incubations, three different concentrations of organic matter were used: 44 mg/g, 105 mg/g. and 181 mg/g total organic carbon (TOC). Similarly, three different concentrations of organic matter were used in OM-co-precipitated ferrihydrite incubations: 44 mg/g, 98 mg/g, and 182 mg/g TOC.

10 mL medium was added to pre-sterilized (6 h, 180 °C) 21 mL hungate tubes containing pre-weighed ferrihydrite, OM-adsorbed ferrihydrite (TOC: 44 mg/g C, 105 mg/g C, 181 mg/g C), or OM-co-precipitated ferrihydrite (TOC: 44 mg/g C, 98 mg/g C, 182 mg/g C) with constant flushing (N₂). The concentration of Fe per tube ranged from 0.63-0.85 mmol. The hungate tubes were sealed with butyl rubber stoppers and flushed with sterile N₂/CO₂ (80%/20% v/v) for ~2 h to ensure anoxic conditions. Either peat microbial extract (5% v/v) or *S. oneidensis* MR-1 were used as inoculum in separate incubations containing ferrihydrite, OM-adsorbed ferrihydrite (TOC: 44 mg/g, 98 mg/g, 182 mg/g), or OM-co-precipitated ferrihydrite (TOC: 44 mg/g, 105 mg/g, 181 mg/g). All treatments were performed in triplicate and incubated horizontally at room temperature in the dark and shaken periodically. Fe(III) reduction was monitored approximately every 24 hours using the phenanthroline assay (Tamura et al., 1974) (data not shown). Fe(II) formation kinetics were used as analogues for Fe(III) reduction. Apparent initial reaction rates were estimated by fitting linear regression lines to Fe(II) [mM]) versus time [hrs] for all data points. The slope of the line represents the initial reaction rate. At the end of incubation (~12 days), solid remnants were freeze dried and stored under N₂ until X-ray diffraction (XRD) measurements were performed (see description below).

2.5 DNA extraction

Genomic DNA was obtained from the peat microbial extract and from the incubation studies using the PowerSoil DNA isolation kit (MoBio Laboratories, Catalog # 12888-50, Carlsbad, CA, USA) according to manufacturer's instructions and stored at -20°C. The DNA was extracted from three biological replicates at 3 different time points (T₀, T_{middle} – 136 h, T_{end} – 296 h) from all incubations inoculated with either peat microbial extract or *S. oneidensis* MR-1. DNA extracts were

subsequently used as a template for quantitative PCR (qPCR) reactions targeting bacterial 16S rRNA genes to quantify abundance of bacteria over time or microbial community analysis via Illumina sequencing. Genomic DNA from samples taken at T₀ and T_{end} (296 h) were used for Illumina sequencing, while DNA from all three time points were used as the template for qPCR. Genomic DNA was extracted from three biological replicates at 3 different time points (T₀, T_{middle}—136 h, T_{end}—296 h) from all incubations inoculated with either peat microbial extract or *S. oncidensis* MR 1 using the PowerSoil DNA isolation kit (MoBio: Catalog # 12888-50). The genomic DNA was stored at 20°C.

2.6 Quantitative PCR

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Bacterial_16S rRNA gene copy and transcript numbers of the total peat microbial community were determined by quantitative PCR using a Mx3000P instrument (Agilent) with Maxima SYBR Green Master Mix (Fermentas). Bacterial 16S rRNA genes were amplified using primer pairs Bac8Fmod/Bac338Rabc (Daims et al., 1999; Loy et al., 2002) according to cycling conditions and standards previously described (Herrmann et al., 2012). *S. oneidensis* MR-1 16S rRNA gene copy numbers were determined with primer pairs She12F/She220R (Himmelheber et al., 2009). 2 to 20 ng of genomic DNA was used as template. Standard curves were produced using serial dilutions of representative plasmid mixtures and were linear for both primer sets from 5x10⁸ to 5x10² copies with R² values of 0.999-1.000 and the qPCR performed with efficiencies between 80-90%. To monitor temporal changes in transcriptional activity on gPCR was performed on-subsamples from T₀, T_{middle} (136 h), and T_{end} (296 hr)_T, transcript numbers were normalized to copy numbers of the corresponding gene copy number per genome: 4.02 for bacterial 16S rRNA and 9 for *S. oneidensis* 16S rRNA gene (https://rrndb.umms.med.umich.edu, as retrieved on June 13th 2017). qPCR was performed on subsamples from T₀, T_{middle} (136 h), and T_{end} (296 hr). One-way analysis of variance (ANOVA) was used to compare treatments using the calculated quantification results and determine statistical significance (p<0.05).

2.7 Illumina MiSeq Amplicon sequencing

16S rRNA gene-targeting amplicon sequencing was done using a bacteria-specific primer pair (341f [S-D-Bact-0341-b-S]/805r [S-D-Bact-0785-a-A]) (Klindworth et al., 2013). Amplicon sequencing was performed by LGC Genomics GmbH (Berlin, Germany). PCR products were generated starting from about 5 ng of extracted genomic DNA extract, 15 pmol of forward and reverse primer in 20 μeL of 1 × MyTaq buffer containing 1.5 units MyTaq DNA polymerase (Bioline, City, Country) and 2 μl of BioStabII PCR Enhancer (Sigma-Aldrich, Munich, Germany). For each sample, the forward and reverse primers had the same 10-nt barcode sequence. PCRs were carried out for 30 cycles using the following parameters: 2 min 96°C predenaturation; 96°C for 15 s, 50°C for 30 s, 70°C for 90 s. DNA concentrations of amplicons of interest were quantified semi-quantitatively via gel electrophoresis. ~20 ng amplicon DNA of each sample were pooled for up to 48 samples carrying different barcodes. If needed PCRs showing low yields were further amplified for an additional 5 cycles. The amplicon pools were purified with one volume AMPure XP beads (Beckmann-Coulter, Krefeld, Germany) to remove primer dimer and other small mispriming products, followed by an additional purification on MinElute columns (Qiagen, Dusseldorf, Germany).

About 100 ng of each purified amplicon pool DNA was used to construct sequencing libraries using the Ovation Rapid DR Multiplex System 1–96 (NuGEN, Leek, Netherlands). Illumina libraries were pooled and size selected by preparative gel electrophoresis. Sequencing was done on an Illumina MiSeq using V3 chemistry (Illumina) in paired-end mode (2 x 300 bp). Amplicon sequencing was performed genomic DNA samples taken at T₀ and T_{end} (296 hr).

2.8 Microbiome profiling

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Sequence characteristics were checked using prinseq (v. 0.20.4) (Schmieder and Edwards, 2011). Adaptor sequences still present were subsequently removed with *cutadapt* (v. 1.7) (Martin, 2011). Paired-end assembly and subsequent quality control were done using vsearch (v. 1.10.2) (Rognes et al., 2016) using the implemented -fastq mergepairs (-fastq minmergelen = expected amplicon size) and -fastq filter (-fastq maxee = 0.5) functions. Chimera removal was done with uchime (v. 4.2.40) (Edgar et al., 2011). OTU clustering with a defined threshold was omitted, instead sequences were dereplicated to streamline the dataset. Deduced dereplicated sequences were treated as being representative for sequence populations featuring an identity of 100% (in the following referred to as OTU). After alpha diversity analyses, the dataset was stripped by removing OTUs that were represented by less than three sequences. Relative abundances were calculated and standardized based on median sequencing depth across all samples. Datasets were not rarefied to prevent a loss of sensitivity in downstream sample-wise comparisons (McMurdie and Holmes, 2014). The dataset was stripped by removing OTUs that were only identified in single samples. Sequences from this simplified dataset were taxonomically assigned against SILVA (release 123, as of September 2016) using the uclust algorithm (Edgar, 2010) implemented in vsearch applying a minimum similarity of 90%. An OTU table including determined taxonomic affiliations and available metadata was assembled for downstream analysis in qiime (v. 1.9.2) (Caporaso et al., 2010) and phyloseg (v. 1.10) (McMurdie et al., 2013). Statistical analyses were carried out using the R statistical software framework (R core team, 208) and the following packages, including respective dependencies, phyloseq, vegan (v. 2.4-3) (Oksanen et al., 2017), and DESeq2 (v. 1.10.1) (Love et al., 2014). Differentially abundant taxonomic groups were identified in sample-wise comparisons assuming negative binomial distributions for taxa read count data and using Wald significance tests, and parametric fitting for determined dispersions. These reflect biological variation better as for instance Poisson distributions (Love et al., 2014) and thus prevent overestimation due to false-positives. Differential abundances of taxonomic groups against defined parameters (ferrihydrite complexation, C/Fe ratio) were done using negative binomial distributions combined with Wald significance tests, and parametric fitting for determined dispersions.

2.9 Sequence data deposition

Generated amplicon sequencing data was deposited as raw data at NCBI SRA and are available under the following BioProject accession number: PRJNA391563

2.10 X-ray diffraction (XRD) measurements

X-ray diffraction measurements were performed with a D8 Advance DaVinci diffractometer (Bruker AXS, Karlsruhe,

Germany) using Cu-Kα radiation at 40 kV and 40 mA, scanning from 5 to 80 °2θ with increments of 0.02 °2θ. Mineral identification was based on the <u>International Centre for Diffraction Data (ICDD) powder diffraction file (PDF)</u> database of 2011 (ICDD, 2011). —XRD measurements were performed on both the starting material and the end products from all incubation experiments.

3 Results

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3.1 Microbial Fe(III) reduction by Shewanella oneidensis MR-1 and by the microbial consortia extracted from peat

270 We compared the Fe(III) reduction capacity of the model organism S. oneidensis MR-1 with a peat microbial consortia extracted from an anoxic iron-rich peatland for incubations containing pure ferrihydrite (Fh control), coprecipitated OMferrihydrite complexes (CFhA (44 mg/g C), CFhB (98 mg/g C), CFhD (182 mg/g C)), or adsorbed OM-ferrihydrite complexes (AFhA (44 mg/g C), AFhB (105 mg/g C), AFhD (181 mg/g C)). The Fe(III) reduction patterns of S. oneidensis MR-1 and the peat microbial consortia were nearly identical. FRates of Fe(III) reduction in incubations with the peat microbial consortia and 275 S. oneidensis MR-1 increased with increasing TOC concentrations (and C/Fe ratios) regardless of whether adsorbed OMferrihydrite complexes or coprecipitated OM-ferrihydrite complexes were used as the terminal electron acceptor (Fig. 1, Fig. 2. Table 1). -The lowest Fe(III) reduction rate was observed in S. oneidensis MR-1 incubations with pure ferrihydrite with a reduction rate of 0.0224 mmol hr⁻¹ compared to 0.031 mmol hr⁻¹ for incubations with our peat microbial consortia (Fig. 1, 5 Fig. 2. Table 1). The highest Fe(III) reduction rates were observed in incubations containing coprecipitated OM-ferrihydrite 280 complexes, with increasing rates of reduction corresponding to increased C/Fe ratios (from 44 to 182 mg g C). Moreover, the highest rates of Fe(III) reduction were observed in incubations containing coprecipitated OM ferrihydrite (182 mg/g C) and either the peat microbial consortia or S. oneidensis MR 1 (0.055 mmol hr⁻¹ and 0.050 mmol hr⁻¹, respectively) (Fig. 1a, Fig. 2a, Table 1). Similar patterns occurred in incubations with adsorbed OM-ferrihydrite complexes, such that the higher the C content Fe ratios the higher the rate of Fe(III) reduction (from 44 to 181 mg g C). For samples with the same C content The 285 rates of Fe(III) reduction was were slightly higher in incubations with peat microbial extract compared to S. oneidensis MR-1 (0.03753-0.058 mmol hr⁻¹ and 0.0430.032-0.051 mmol hr⁻¹, respectively) (Fig. 1, Table 1). At the same time, rates of Fe(III) reduction -were slightly higher for OM-ferrihydrite coprecipitates compared to OM-ferrihydrite adsorption complexes (0.038-0.058 mmol hr⁻¹ compared to 0.032-0.050 mmol hr⁻¹ Fig. 1b, Fig. 2b, Table 1) when all incubations are considered (Fig. 1, Table 1). Surprisingly, the rates of Fe(III) reduction in incubations with the peat microbial consortia and S. oneidensis MR-1 increased with increasing TOC concentrations (and C/Fe ratios) regardless of whether adsorbed OM-ferrihydrite complexes 290 or coprecipitated OM ferrihydrite complexes were used as the terminal electron acceptor(Fig. 1, Fig. 2, Table 1).

3.2 Quantitative changes of 16S rRNA genes in incubation experiments

We followed the abundance of bacterial 16S rRNA gene copies over time in incubations inoculated with either *S. oneidensis* MR-1 or the microbial consortia extracted from peat to determine potential effects of coprecipitated or adsorbed OM-

295 ferrihydrite complexes on microbial abundances. In general, bacterial 16S rRNA copy numbers increased over time in all incubations with noticeable increases in gene abundances at the end of incubation and slight variations dependent on the TOC concentration of the OM coprecipitated or adsorbed to the surface of ferrihydrite (Fig. 3, Fig. 42). In incubations with pure ferrihydrite, S. oneidensis MR-1 16S rRNA gene copy numbers increased four orders of magnitude from 7.90 x 10⁷ cells mL ¹ at T₀ to 1.99 x 10¹¹ cells mL⁻¹ at T_{end} (296 h). S. oneidensis MR-1 16S rRNA gene copy numbers increased approximately five orders of magnitude in incubations with coprecipitated or adsorbed OM-ferrihydrite complexes with the highest C/Fe 300 ratios (8.10 x 10^7 cells mL⁻¹ at T₀ to 8.47 x 10^{12} cells mL⁻¹ at T_{end} and 8.10 x 10^7 cells mL⁻¹ at T₀ to 1.24 x 10^{12} cells mL⁻¹ at T_{end}, respectively). In incubations with pure ferrihydrite and the peat microbial consortia, bacterial 16S rRNA gene copy numbers only increased two orders of magnitude from 6.23 x 10¹⁰ cells mL⁻¹ at T₀ to 2.02 x 10¹² cells mL⁻¹ at T_{end}. Bacterial 16S rRNA gene copy numbers increased roughly three orders of magnitude in incubations using the peat microbial consortia as inoculum and adsorbed OM-ferrihydrite complexes with the highest C/Fe ratios (6.23 x 10¹⁰ cells mL⁻¹ at T₀ to 2.58 x 10¹³ 305 cells mL⁻¹ at T_{end}), however, the gene copy numbers increased approximately four orders of magnitude in incubations with coprecipitated OM-ferrihydrite with the highest C/Fe ratios (6.23 x 10¹⁰ cells mL⁻¹ at T₀ to 1.36 x 10¹⁴ cells mL⁻¹ at T_{end}, respectively). Thus, increases in gene copy numbers correlate to the increased Fe(III) reduction rates observed across all incubations with OM-ferrihydrite complexes and pure ferrihydrite.

3.3 Microbial community profiling of the microbial consortia in incubation experiments

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Incubations with absorbed and coprecipitated OM-ferrihydrite and varying C/Fe ratios led to distinct shifts in microbial community structure when compared to pure ferrihydrite incubations and the initial microbial consortia extracted from peat (Fig. 3, Fig. 45). Alpha diversity indices addressed intra-sample diversity either based on the observed number of taxa according to the applied OTU definition (observed OTUs) or based on extrapolations of observed taxa numbers putting special emphasis on rare taxa (Chao1, ACE). Comparing the different incubations with the initial peat microbial extract (Observed OTUs: 52064, Chao1: 801599, ACE: 819056, except AFhB) revealed decreased diversities in all cases (Observed OTUs: 23120-46770, Chao1: 334438-545083, ACE: 266190-455099) except AFhB (Fig. 45a). Intra-sample diversity measures were complemented by assessing the evenness of datasets by determining the entropy of respective datasets (Shannon)—and determining the probability of not randomly subsampling sequences of same taxonomic origin when two sequences are sampled (Simpson). Respective values matched observed intra-sample diversities. Calculated Shannon and Simpson-indices showed strongly increased evenness for the initial peat microbial extract (Shannon: 9.47, Simpson: 0.991) against all incubations (Shannon: 6.77-7.48, Simpson: 0.88 0.966)f, excluding AFhB. Beta diversity analysis is a common approach to gain insights about inter-sample diversity when dealing with complex datasets such as microbial community profiles (Fig. 45b). Corresponding analyses based on Jensen-Shannon divergence highlighted a minor clustering of microbial communities linked to coprecipitated OM-ferrihydrite incubations in case of low and medium C/Fe ratios (CFhA, CFhB). Community profiles originating from incubations supplemented with OM-ferrihydrite having a high C/Fe ratio revealed a clustering independent of ferrihydrite complexation (AFhD, CFhD). The observation of community profiles from incubations clustered distinctly from communities present in the initial peat microbial extract matched the abovementioned differences with respect to intra-sample diversity and evenness (Fig. 45a).

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Looking at underlying community profiles from a taxonomic perspective revealed a dominance of few phyla when considering determined community profiles as a whole (Fig. 3a6A). Dominant phyla included especially PGammaproteobacteria (57.9-95.4% relative abundance), Firmicutes (3.1-41.6%), and to a lesser extent Bacteroidetes (up to 0.6% [AFhD, CFhD]) and Acidobacteria (4.5% [peat microbial extract]). Looking at the taxonomic composition at a more resolved level (family-level) showed that carried out incubations led in all cases to a strong increase of Pseudomonadaceae in relative abundance within the Gamma-Proteobacteria (e.g. AFhA 58.1%, in comparison to peat microbial extract 5.3%) (Fig. 36b). Pronounced increases were also seen for Veillonellaceae (AFhA, CFhA, CFhB [24.6-35.5%], in comparison to peat microbial extract [< 0.00001%], and Fh [< 0.00001%]), Bacillaceae (AFhA, AFhB, CFhA [5.4-12.0%], in comparison to peat microbial extract [< 0.00001%], and Fh [3.6%]), and Desulfovibrionaceae (CFhA, CfhB [1.8-2.5%], in comparison to peat microbial extract [< 0.001%], and Fh [< 0.00001%]). Increased abundances of Veillonellaceae were correlating with low and medium C/Fe ratios, while incubations with high C/Fe ratios favoured elevated relative abundances of Enterobacteriaceae, when compared with low and medium C/Fe ratios. Groups that appeared to be enriched in the incubation with pure ferrihydrite, for example Burkholderiaceae, were found to show only minor abundances in any of the other setups.

Considering the results from beta diversity analysis and community profiling, we were interested in assessing the differential abundance of genus-level taxonomic groups to check for the potential enrichment of taxonomic groups that are linked to iron cycling when comparing the pure ferrihydrite incubation and OM-ferrihydrite incubations. Differential abundances were examined by assuming negative binomial distributions for taxa read counts (Fig 36c). It was apparent that numerous genus-level taxonomic groups were underrepresented when CFh and AFh are contrasted against Fh. This was consistent with reduced intra-sample diversities for adsorbed and co-precipitated incubations in comparison to the pure ferrihydrite incubation (Fig. 45a). Most underrepresented genus-level taxonomic groups were linked to Proteobacteria and especially Enterobacteriales/Enterobacteriaceae including: Ewingella, Rahnella, Serratia, and Yersinia. This pattern was similar for CFh against Fh, as well as AFh against Fh. In most cases log2 fold changes ranged between -5 and -10 for underrepresented groups. The number of overrepresented taxonomic groups was in general lower than the number of underrepresented groups. This finding was much more pronounced for AFh vs Fh than CFh vs Fh. Only three overrepresented genus-level taxonomic groups were identified for AFh vs Fh, namely *Pelosinus*, *Pseudomonas*, and *Telmatospirillum* with log2 fold changes above 5. The same groups were also identified to be enriched in case of CFh vs Fh. However, overall the number of enriched taxonomic groups was higher including various Firmicutes such as Bacillus, Clostridium, Lachnoclostridium, and Paenibacillus. Few genus-level groups comprised over- and underrepresented OTUs. This was the case for Bacillus, Paenibacillus, and Pseudomonas. In the case of Pseudomonas, one particular OTU was found to be enriched. Consequently, this indicates that pronounced increases in the abundance of *Pseudomonaceae* in adsorbed and co-precipitated OM-ferrihydrite are apparently mediated by a well-defined sub population of *Pseudomonas*.

3.4 XRD measurements

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X-ray diffraction patterns of the starting material of the solid remnants after reduction showed the two very broad peaks of ferrihydrite (Fe₁₀O₁₄(OH)₂). After reduction, diffractograms of the solid remnants -still revealed the presence of ferrihydrite asbut additionally showed well as overlying sharp peaks of different salts such as (halite (NaCl), nahcolite (NaHCO₃), and salammoniac (NH₄Cl), which derived from the medium. Goethite (α-FeOOH), other Fe(III) oxides or Fe(II) minerals were not found. The newly formed Fe(II) was either effectively chelated by the present organic material or dissolved Fe concentrations were too low to exceed the solubility product of Fe oxides or other common secondary Fe(II) minerals. The latter seems likely as the reduction rates of the present experiment were in fact low $(0.022 - 0.058 \text{ mM h}^{-1}; \text{ Table 1})$. In previous experiments, using the same OM-ferrihydrites, the same medium, but Geobacter bremensis as the model Fe(III) reducer, we detected goethite and siderite for samples without OM or only low OM contents (Fh, AFhA, CFhA, CFhB), where reduction rates where high (0.119-0.148 mM h⁻¹). Other sSimilar reduction experiments reported no mineral transformation for OM ferrihydrites (Henneberry et al., 2012) or observed-concluded that secondary mineral formation is depended on the OM content of the coprecipitates: Shimizu et al. (2013) observed goethite formation fromer pure ferrihydrite, while OMHA-ferrihydrites transformed into green rust and magnetite (Fe₃O₄) after reduction by Shewanella (Shimizu et al., 2013). Abiotically reducing Abiotic reduction of -OM- ferrihydrite coprecipitates by aqueous Fe(II), Chen et al. (2015) found that OM hindered transformation into goethite and magnetite but favoured transformation into lepidocrocite (γ-FeOOH). Likewise, Adhikari et al. (2017) reduced HA-ferrihydrite coprecipitates by Shewanella and found complete transformation into goethite and magnetite for pure ferrihydrite, while the amount of secondary minerals (goethite, magnetite, lepidocrocite, siderite (FeCO₃), akaganeite (β-FeO(OH, Cl))) decreased with the content of humic acids in the coprecipitates.

4 Discussion

Dissolved, particulate and mineral associated OM plays a crucial role in many biogeochemical processes in soils and sediments. The ability of OM to serve as a terminal electron acceptor for microbial respiration in anoxic environments is of special importance and has been well characterized (Lovley et al., 1996). However, most experiments utilizing OM as a terminal electron acceptor have been performed with alkaline extracts: Reduced humic acids have been shown to transfer electrons abiotically to terminal electron acceptors, such as metals, thereby functioning as electron shuttles between microbes and electron acceptors (Kappler et al., 2004). Humic acids, therefore, possess a redox-mediating ability which positively affects the cycling of redox-active minerals, including Fe(III) and Mn(IV)/Mn(III), in conjunction with transformation of organic and inorganic compounds. Subsequent studies have shown that reduced forms of humic substances can serve as electron donors for anaerobic organisms growing on a variety of terminal electron acceptors, including nitrate and fumarate (Coates et al., 2002). Under these conditions, microorganisms utilize the reduced humic substances as an energy source, while using a readily degradable carbon source, such as acetate, for growth. This unique ability to use humic substances as an electron donor provides a possible competitive advantage over other heterotrophs in the environment which are only capable of utilizing

a single organic compound as both the carbon and energy source. Previous studies have revealed that some Fe(III)-reducing microorganisms, such as *S. alga, G. metallireducens*, and *Geothrix fermentans*, have the ability to undergo dissimilatory humic reduction as well as the ability to couple reduced humic oxidation with nitrate reduction. Moreover, *Paracoccus denitrificans*, a known denitrifier, is unable to utilize humics or Fe(III) as a terminal electrons acceptors, however, these microorganisms are capable of using reduced humics as an electron donor for denitrification (Coates et al., 2002; Lovley et al., 1999).

This Schlöppnerbrunnen fen has been studied in great detail during the last 10 years, however, the impact of OM on the microbially-mediated Fe(III) reduction processes in this site are not fully understood. Earlier incubation studies with peat soil from the Schlöppnerbrunnen fen demonstrated that amendment with each of these individual substrates (lactate, acetate, glucose) stimulates Fe(III) reduction, with glucose showing the most rapid increase of Fe(II) production (Küsel et al., 2008). Glucose was completely consumed within 3 days of incubation, yielding acetate and CO₂ as end products, and H₂ and ethanol as transient products. Supplemental lactate yielded acetate as a main product (Küsel et al., 2008). Both acetate and H₂ are preferred substrates for many Fe(III)-reducing microorganisms, including *Shewanella* and *Geobacter* species. Thus, amendment of incubations containing a microbial consortia extracted from Schlöppnerbrunnen fen with a substrate mixture comprised of the carbon sources used in the aforementioned study (i.e. lactate, acetate, glucose) and ferrihydrite as the terminal electron acceptor should augment growth of many peat microorganisms involved in different steps of anaerobic organic matter degradation using Fe(III) as terminal electron acceptor.

The current study provides clear evidence that the microbial community found in the fen not only has the potential to reduce Fe(III)-oxides complexed with organic matter to the same extent as the model Fe(III)-reducer, *S. oneidensis* MR-1, but the microbes also increase in abundance under the experimental growth conditions used. We utilized a mixture of electron donors (glucose, lactate, and acetate) to stimulate a great variety of peat microorganisms capable of dissimilatory Fe(III) reduction including those that couple Fe(III) reduction with growth, but also fermenting microorganisms and others that just use Fe(III) oxides as electron sink (Hori et al., 2015; Lentini et al., 2012). Differences in the gene copy numbers at the end of the incubation suggested that not only did the supplemented substrates support growth but also OM adsorbed or coprecipitated to the surface of the ferrihydrite minerals. Based on these results, we can infer that the mineral associated OM was at least partially degradable and not totally protected by its association to minerals or by its molecular structure and size. Furthermore, it is likely that a portion of the ferrihydrite associated OM became re-dissolved over the course of the incubation period, which may have facilitated its utilization by the microbes. Thus, the enhanced growth observed in both the *S. oneidensis* MR-1 and the microbial consortia can be attributed to the presence of additional complex carbon sources in the form of OM-ferrihydrite complexes compared to pure ferrihydrite.

The current study provides clear evidence that the microbial community found in the Fe rich Schlöppnerbrunnen fen not only has the potential to reduce Fe(III) oxides complexed with organic matter to the same extent as the model Fe(III) reducer, S. oncidensis, but the microbes also increase in abundance under the experimental growth conditions used. We used a mixture of electron donors (glucose, lactate, and acetate) to stimulate a great variety of peat microorganisms capable of dissimilatory Fe(III) reduction including those that couple Fe(III) reduction with growth, but also fermenting microorganisms

and others that just use Fe(III) oxides as electron sink (Hori et al., 2015; Lentini et al., 2012). Earlier incubation studies with peat soil from the Schlöppnerbrunnen fen demonstrated that amendment with all these individual substrates stimulates Fe(III) reduction with glucose showing the most rapid increase of Fe(II) production (Küsel et al., 2008). Glucose was completely consumed within 3 days of incubation, yielding acetate and CO₂ as end products, and H₂ and ethanol as transient products. Supplemental lactate yielded acetate as a main product (Küsel et al., 2008). Both acetate and H₂ are preferred substrates for many Fe(III) reducing microorganisms including *Shewanella* and *Geobacter* species. Thus, the substrate mixture used in these ferrihydrite incubation experiments should have stimulated growth of many peat microorganisms involved in different steps of anaerobic organic matter degradation using Fe(III) as terminal electron acceptor. Differences in the gene copy numbers at the end of the incubation suggested that not only did the supplemented substrates support growth but also OM adsorbed or coprecipitated to the surface of the ferrihydrite minerals. Based on these results, we can infer that the mineral associated OM was at least partially degradable and not totally protected by its association to minerals or by its molecular structure and size. Furthermore, it is likely that a portion of the ferrihydrite associated OM became re dissolved over the course of the incubation period, which may have facilitated its utilization by the microbes. Thus, the enhanced growth observed in both the *S. oneidensis* and the microbial consortia can be attributed to the presence of additional complex carbon sources in the form of OM ferrihydrite complexes compared to pure ferrihydrite.

4.1 Mechanisms of OM-ferrihydrite reduction by S. oneidensis MR-1 and the microbial consortia

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The results of this study demonstrate that OM ferrihydrite associated OM stimulates growth and Fe(III) reduction by S. oneidensis MR-1 and the microbial consortia, as well as growth. Fe(II) production increased as the C/Fe ratios increased in all incubations. Both the reduction rates and the total amount of Fe(II) produced at the end of incubation (~12 days) increased in the order Fh < adsorbed OM-ferrihydrite < coprecipitated ferrihydrite. (Fig. 1, Fig. 2, Fig. 57, and Table 1). Samples containing coprecipitated OM-ferrihydrite complexes were more reactive than samples containing adsorbed OM-ferrihydrite complexes with similar OM loadings. The latter can be explained by a generally smaller crystal size of OM-ferrihydrite coprecipitates compared to pure ferrihydrite (Mikutta et al., 2008; Eusterhues et al., 2008) and was also observed during reduction by Geobacter bremensis (Eusterhues et al., 2014). These findings that mineral associated OM may enhance ferrihydrite reduction supports the the results presented in previous studies examining the role of humic acids as potential electron shuttles by S. putrefaciens under Fe(III) reducing conditions, which showed — These studies showed that S. putrefaciens was capable of using humic acids as an electron shuttle to enhance Fe(III) reduction (Amstaetter et al., 2012; Shimizu et al., 2013). However, enhanced Fe(III) reduction was only observed when as long as the concentration of organic material was high enough to overcome any slowing in the reduction processes caused by OM sorption to the surface reduction sites of the mineral (Amstaetter et al., 2012; Shimizu et al., 2013). Our study revealed that a water extract of natural soil organic matter, which has lower contents of aromatic moieties groups, and usually lower electron shuttling accepting and electron donating capacityies to than humic acids (Fritzsche et al., n.d.; Piepenbrock et al., 2014) can also provide enough electron shuttling compounds to stimulate ferrihydrite reduction by S. oneidensis as well as by the peat microbial consortium. The organic matter used to produce the OM_ferrihydrites for this study was extracted by H_2O from a Podzol litter layer and showed an electron accepting capacity of EAC = 32 ± 1 mmol e^- (mol C)⁻¹ and an electron donating capacity of EDC = 24.7 ± 1.2 mmol e^- (mol C)⁻¹ (see WEOM_{WINTER}) (Fritzsche et al., n.d.).

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Furthermore, since the S. oneidensis MR-1 pre-culture used in the present study was grown aerobically and the peat microbial consortia was resuspended in 10 mL anoxic 0.85% NaCl salt solution following extraction from peat cores, we can assume no electron shuttling molecules were secreted by the microorganisms themselves prior to inoculating the incubation experiments. Thus, the absence of any electron shuttles in our pure ferrihydrite incubations might explain the lowest Fe(III) reduction rates out of all the experiments conducted. On the contrary, we suggest the presence of electron shuttling compounds, such as quinones, in the incubations with coprecipitated or adsorbed OM. Quinone moieties have been described as electron shuttling compounds in previous studies (Scott et al., 1998). In our experiments even small contents of OM on ferrihydrite (44 mg g⁻¹, i.e., small molar C/Fe ratios of 0.4) lead to increased reduction rates. This is in contrast to Shimizu et al. (2013) and Adhikari et al. (2017), who reported surface passivation at low C/Fe ratios, but increased reduction rates at molar ratios C/Fe > 1.8. A subsequent study by Poggenburg et al. (2016) reported enhanced Fe(III) reduction rates in incubations containing OM-ferrihydrite complexes with molar C/Fe ratios of 0.9 - 2.1 and using S. putrefaciens as model reducer. Here reduction rates differed mostly due to OM-composition, i.e. aromaticity. -Lowest reduction rates were found for ferrihydrite coprecipitates produced with extracellular polymeric substances, while highest reduction rates were found for ferrihydrite coprecipitates produced with a water extract from Podzol litter. The latter is likely comparable in composition to the OM of our experiment. However, absolute reduction rates cannot be compared directly, because Poggenburg et al. (2016) attributed their high rates of Fe(III) reduction partly to the desorption of NOM from coprecipitates due to large concentration of phosphate in their nutrient solutions. In the present study, our C/Fe molar ratios ranged from 0.4 to 4.2, yet we observed an increase in Fe(III) reduction rates for all incubations containing either coprecipitated or adsorbed OM-ferrihydrite complexes compared to incubations with pure ferrihydrite.

In *Geobacter* sp., direct contact with oxide surfaces is required for Fe(III) reduction_(Nevin and Lovley, 2000). *Geobacter* produces nanowires, typically described as conductive extracellular appendages, which facilitate electron transfer by serving as an electrical conduit to the Fe(III) oxide surface_(Alves et al., 2016; Reguera et al., 2005). Conversely, *Shewanella* species are capable of using three primary strategies to allow for electron transfer between itself and solid Fe(III) oxide surfaces; via direct contact. *Shewanella* has been shown to transfer electrons via direct contact with the oxide surfaces, production of a complexing ligand (i.e. siderophore), ander the utilization of endogenously or exogenously produced electron shuttles_(Cooper et al., 2016; Melton et al., 2014; Shi et al., 2012). A previous study by Eusterhues, et al. (2014) showed the rates of Fe(III) reduction by *G. bremensis* decreased as the concentrations of OM adsorbed or co-precipitated to the surface of ferrihydrite increased (Eusterhues, *et al.* 2014). However, under identical reduction experiments and the same ferrihydrite material, Fe(III) reduction by *G. bremensis* resulted in considerably higher reduction rates of (0.088-0.148 mM h⁻¹) compared to *S. oneidensis* (0.022-0.051 mM h⁻¹) and the microbial consortia (0.031-0.058 mM h⁻¹) (Table 1, Eusterhues et al., 2014). Although favoured at high concentrations of ferrihydrite-associated OM, electron shuttling is slower than direct electron

transfer via nanowires. Similar results were reported in a study by Poggenburg; et al. (2016), in which the rates of Fe(III) reduction by *G. metallireducens* were greatest with pure ferrihydrite compared to ferrihydrite coprecipitates produced with different natural organic materials. These findings suggest that the reduction rates are not only influenced by organic matter composition and the C/Fe molar ratios, instead they also reflect species dependent capabilities. It was hypothesized that OM may passivate Fe(III) oxide surfaces for direct contact requiring reducers such as *Geobacter*, whereas it may promote reduction by providing exogenous electron shuttles for non-direct contact requiring reducers such as *Shewanella* sp. (Eusterhues et al., 2014). In the current study, we observed a similar behaviour for *S. oneidensis* and the complex microbial consortia extracted from peatland suggesting that non direct-contact requiring microorganisms dominate reduction in the anoxic, Fe-rich fen, making use of the exogenous electron shuttling compounds of the natural OM (Fig. 57). The similarity of the patterns observed point to the suitability of *S. oneidensis* as model organism to study the effect of OM on ferrihydrite reduction in natural environments.

4.2 Potential key players of OM-ferrihydrite reduction

Though the microbial consortia behaves similarly to *S. oneidensis* MR-1 with regards to Fe(III) reduction capacities and growth over time, the community profiles at the end of incubation did not reveal an enrichment of *Shewanella* species. Differential abundance analyses between adsorbed OM-ferrihydrite incubations and the pure ferrihydrite incubation even revealed reduced abundances for *Shewanella* in case of adsorbed OM-ferrihydrite incubations. The enriched community is indeed comprised of microorganisms capable of reducing Fe(III) oxides₂, Hhowever, *Shewanella* sp. but also and *Geobacter* sp. did not respond to these growth conditions and are noticeably low abundancet. But the lack of enrichment of typical dissimilatory Fe(III)-reducer cannot be explained by the lack of suitable electron donors. Previous work using ¹³C-labelled acetate in paddy soils revealed that especially *Geobacter*-related organisms respond to acetate amendment (Hori et al., 2010).

We saw distinct changes in microbial community structure over time as reflected by alpha- and beta diversity analyses which appeared to be linked to several main drivers shaping the microbial communities in the incubations. The unexpectedly high alpha-diversity observed for the adsorbed OM-ferrihydrite incubation featuring a medium C/Fe ratio (AFhB) is a result of over-estimating diversity, due to a high frequency of singletons (OTUs that are only represented by single sequences). The overall anaerobic incubation conditions with ferrihydrite offered as terminal electron acceptor and a mixture of the carbon substrates glucose, lactate, and acetate activated only a subset of the original community extracted from peat. Fermentative Enterbacteriaceae (*Ewingella, Rahnella, Serratia*) were overrepresented in the pure ferrihydrite incubation which is reflective of their ability to use glucose as carbon source and available Fe(III) as electron sink. Enterbacteriaceae are known for a mixed acid fermentation profile similar to the products observed in earlier peat soil incubation studies amended with glucose_(Reiche et al., 2008). The question of whether Fe(III) can be also used for energy conservation by these bacteria remainsed unanswered thus far. Early work by Ottow (1970) suggested *Serratia* to be involved in hematite reduction___(Ottow, 1970). The strong reduction in their relative abundance in OM-ferrihydrite incubations is presumably a result of Fe(III) being less accessible.

Community profiles of OM-ferrihydrite incubations differed tremendously from the one of the pure ferrihydrite incubation. Pseudomonadaceae were highly enriched in all OM-ferrihydrite incubations. The enrichment of Veillonellaceae in adsorbed and coprecipitated OM-ferrihydrite incubations with low C concentrations could be mostly attributed to *Pelosinus*. *Pelosinus* sp. have been characterized as fermentative bacteria which utilize Fe(III) as an electron sink (Shelobolina et al., 2007). These bacteria can use H₂ as an electron donor and acetate as a carbon source, and are also capable of fermenting citrate, lactate, and organic impurities. *Pelosinus* has also been shown to reduce small concentrations of anthraquinone-2.6-disulfonate (AQDS), a synthetic quinone, in the presence of citrate or lactate (Shelobolina et al., 2007). The ability of these microorganisms to use AQDS as an electron sink under reducing conditions might explain their enrichment during our incubations. Quinone moieties are ubiquitous in soils and peat, thus providing the perfect breeding ground for these microorganisms.

Interestingly, Fe(III)-reduction was previously found to be correlated with the presence of Desulfovibrionaceae, when lactate is available as carbon source (Lentini et al., 2012). This partially matches our observations for incubations with coprecipitated ferrihydrite where *Desulfovibrio* was enriched at low and medium C/Fe ratios. However, *Desulfovibrio* were nearly absent in adsorbed OM-ferrihydrite incubations. Differences in Fe(III) accessibility in adsorbed and coprecipitated OM-ferrihydrite or the higher solubility of coprecipitates might explain this difference in community structure. The presence and ecological role of sulfate-reducing bacteria in the Schlöppnerbrunnen fen has been documented several times (Loy et al., 2004; Pester et al., 2010). This acidic (pH 4), low-sulfate (20-200 µM) peatland shows significant sulfate-reducing capacities (Loy et al., 2004) which is explained by a rapid recycling mechanism (Küsel et al., 2008). Recycling of sulfate can proceed by the aerobic oxidation of sulfide, e.g., in regions where oxygen penetration and anoxic micro-niches overlap (Knorr and Blodau, 2009). The sulfate-reducing activity is mainly attributed to a member of the 'rare biosphere' in the peat, a *Desulfosporosinus* species (Pester et al., 2010). Our study suggests that other members of the sulfate reducers are involved in Fe(III) reduction in this fen to overcome the temporal and/or spatial limitation of the terminal electron acceptor sulfate.

4.3 Mineral transformation during reduction (XRD)

X ray diffraction after reduction showed the presence of remaining ferrihydrite as well as salts which precipitated from the medium. Goethite, other Fe(III) oxides or Fe(II) minerals were not found. The newly formed Fe(II) was either effectively chelated by the present organic material or dissolved Fe concentrations were too low to exceed the solubility product of Fe oxides or other common secondary Fe(II) minerals. The latter seems likely as the reduction rates of the present experiment were in fact low (0.022 – 0.058 mM h⁻¹; Table 1). In previous experiments, using the same OM ferrihydrites, the same medium, but *Geobacter bremensis* as the model Fe(III) reducer, we detected goethite and siderite for samples without OM or only low OM contents (Fh, AFhA, CFhA, CFhB), where reduction rates where high (0.119 0.148 mM h⁻¹). Other similar experiments reported no mineral transformation (Henneberry et al., 2012) or observed goethite formation from or pure ferrihydrite, while OM ferrihydrites transformed into green rust and magnetite (Shimizu et al., 2013).

4.4-3 Summary and Conclusions

In the present study mineral-bound soil OM is shown to increase microbial Fe(III) reduction by *S. oneidensis* MR-1 and by a microbial consortia extracted from Schlöppnerbrunnen fen peat cores. The reactivity of ferrihydrites with adsorbed OM differed from ferrihydrites with coprecipitated OM, such that higher extents of Fe(III) reduction were observed for coprecipitated ferrihydrites in comparison to adsorbed ferrihydrites. This can be explained by the smaller crystal size and more defective crystal structure of coprecipitated ferrihydrites compared to normal ferrihydrites (with or without adsorbed organic matter). Additionally, reduction rates were slightly higher for the microbial consortia than for *S. oneidensis* MR-1. The fact that reduction rates for both *S. oneidensis* MR-1 and the microbial consortia increased systematically with increasing contents of OM in the OM-ferrihydrite complexes tempts us to speculate about how reduction rates are related to electron transfer mechanisms. *Shewanella* sp. are capable of reducing Fe(III) oxides using one of three mechanisms known for reductions, including direct contact, electron shuttles, or production of complexing ligands. Thus, the enhanced Fe(III) reduction rates in our incubations led us to the conclusion that under these conditions electron shuttling using exogenous moieties of the mineral-bound OM is likely the preferred mechanism used for reduction (Fig. 57). Since In-natural environments, coprecipitation of OM to mineral surfaces is a common process in nature, thus, we can hypothesize that electron shuttling mechanisms are the dominant mechanism for microbial Fe(III) reduction in environments with high concentration of OM, such as soils, sediments, and peat.

Experiments with the microbial consortia extracted from peat showed that the presence of mineral associated OM influenced the microbial community composition: Whereas *Enterobacteriaceae* and *Burkholderiaceae* dominated in incubations with pure ferrihydrite, *Pseudomonadaceae* were the dominant microbes present when ferrihydrites were (partly) covered by OM. When ferrihydrites had low organic matter loadings *Veillonellaceae* occurred in relatively high concentrations, conversely, *Rhodospirillaceae* are found only in samples with high OM contents. In incubations with OM-mineral complexes, dominant microbes enriched included *Desulfobacteria* and *Pelosinus* sp., both of which can utilize lactate and acetate as an electron donor under Fe(III) reducing conditions.

Thus, our observations imply OM coverage does not protect the underlying Fe mineral from microbial reduction in soils dominated by microorganisms that predominantly use electron shuttles or ligands for Fe(III) reduction. Our Thise study also revealed that natural soil organic matter OM on the surface of ferrihydrites provides enough redox active groups to maintain microbial Fe(III) reduction processes, (instead of passivating the surface, ?). At the same time t. While at the same time the OM while the additional carbon in the form of Om enhances overall microbial growth and shapes shaped the microbial community structure.

Data availability. Generated amplicon sequencing data can be found as raw data at NCBI SRA and are available under the following BioProject accession number: PRJNA391563.

Author contributions. K. Eusterhues, K. U. Totsche, and K. Küsel designed the experiments. R.E. Cooper performed the experiments. K. Eusterhues synthesized the pure ferrihydrite and OM-ferrihydrite complexes. C-E Wegner carried out sequence analyses. R.E. Cooper prepared the manuscript with contributions from all co-authors.

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References

- Adhikari, D., Zhao, Q., Das, K., Mejia, J., Huang, R., Wang, X., Poulson, S. R., Tang, Y., Roden, E. E. and Yang, Y.: Dynamics of ferrihydrite-bound organic carbon during microbial Fe reduction, Geochim. Cosmochim. Acta, 212, 221–233, doi:10.1016/j.gca.2017.06.017, 2017.
- Alves, M. N., Fernandes, A. P., Salgueiro, C. A. and Paquete, C. M.: Unraveling the electron transfer processes of a nanowire protein from Geobacter sulfurreducens, BBA Bioenerg., 1857, 7–13, doi:10.1016/j.bbabio.2015.09.010, 2016.
 - Amstaetter, K., Borch, T. and Kappler, A.: Influence of humic acid imposed changes of ferrihydrite aggregation on microbial Fe(III) reduction, Geochim. Cosmochim. Acta, 85, 326–341, doi:10.1016/j.gca.2012.02.003, 2012.
- Benjamini, Y. and Hochberg, Y.: Controlling the false discovery rate: A Practical and powerful approach to multiple testing, J. R. Stat. Soc. Ser. B, 57(1), 289–300 [online] Available from: http://www.jstor.org/stable/2346101 (Accessed 27 June 2017), 1995.
 - Bigham, J. M., Fitzpatrick, R. W. and Schulze, D. G.: Iron oxides in: Soil Mineralology with Environmental Applications, in SSSA Book Set. No. 7, edited by D. G. Dixon, J. B. and Schulze, Soil Science Society of America, Madison, WI., 2002.
 - Blodau, C., Basiliko, N. and Moore, T. R.: Carbon turnover in peatland mesocosms exposed to different water table levels,
- 615 Biogeochemistry, 67(3), 331–351, doi:10.1023/B:BIOG.0000015788.30164.e2, 2004.
 - Boesen, T. and Nielsen, L. P.: Molecular dissection of bacterial nanowires., MBio, 4(3), e00270-13, doi:10.1128/mBio.00270-13, 2013.
- Borch, T., Kretzschmar, R., Kappler, A., Cappellen, P. Van, Ginder-Vogel, M., Voegelin, A. and Campbell, K.: Biogeochemical Redox Processes and their Impact on Contaminant Dynamics, Environ. Sci. Technol., 44(1), 15–23, doi:10.1021/es9026248, 2010.
 - von Canstein, H., Ogawa, J., Shimizu, S. and Lloyd, J. R.: Secretion of flavins by Shewanella species and their role in extracellular electron transfer., Appl. Environ. Microbiol., 74(3), 615–23, doi:10.1128/AEM.01387-07, 2008.

- Caporaso, J. G., Kuczynski, J., Stombaugh, J., Bittinger, K., Bushman, F. D., Costello, E. K., Fierer, N., Peña, A. G., Goodrich, J. K., Gordon, J. I., Huttley, G. A., Kelley, S. T., Knights, D., Koenig, J. E., Ley, R. E., Lozupone, C. A., McDonald, D.,
- Muegge, B. D., Pirrung, M., Reeder, J., Sevinsky, J. R., Turnbaugh, P. J., Walters, W. A., Widmann, J., Yatsunenko, T., Zaneveld, J. and Knight, R.: QIIME allows analysis of high-throughput community sequencing data, Nat. Methods, 7(5), 335–336, doi:10.1038/nmeth.f.303, 2010.
 - Chao, A.: Nonparametric estimation of the number of classes in a population, Source Scand. J. Stat., 11(4), 265–270 [online] Available from: http://www.jstor.org/stable/4615964 (Accessed 27 June 2017), 1984.
- Cismasu, A. C., Michel, F. M., Tcaciuc, A. P., Tyliszczak, T. and Brown, Jr, G. E.: Composition and structural aspects of naturally occurring ferrihydrite, Comptes Rendus Geosci., 343(2–3), 210–218, doi:10.1016/j.crte.2010.11.001, 2011.
 Coates, J. D., Cole, K. A., Chakraborty, R., O 'connor, S. M. and Achenbach, L. A.: Diversity and Ubiquity of Bacteria Capable of Utilizing Humic Substances as Electron Donors for Anaerobic Respiration, Appl. Environ. Microbiol., 68(5), 2445–2452, doi:10.1128/AEM.68.5.2445–2452.2002, 2002.
- Cooper, R. E., Goff, J. L., Reed, B. C., Sekar, R. and DiChristina, T. J.: Breathing Iron: Molecular Mechanism of Microbial Iron Reduction by Shewanella oneidensis, in Manual of Environmental Microbiology, 4th Edition, p. 5.2.1-1-5.2.1-13, American Society of Microbiology., 2016.
 - Cornell, R. M. and Schwertmann, U.: The Iron Oxides, Wiley-VCH Verlag GmbH & Co. KGaA, Weinheim, FRG., 2003.
 - Daims, H., Brühl, A., Amann, R., Schleifer, K.-H. and Wagner, M.: The Domain-specific Probe EUB338 is Insufficient for
- the Detection of all Bacteria: Development and Evaluation of a more Comprehensive Probe Set, Syst. Appl. Microbiol., 22(3), 434–444, doi:10.1016/S0723-2020(99)80053-8, 1999.
 - Davison, W.: Iron and manganese in lakes, Earth-Science Rev., 34(2), 119–163, doi:10.1016/0012-8252(93)90029-7, 1993. Edgar, R. C., Haas, B. J., Clemente, J. C., Quince, C. and Knight, R.: UCHIME improves sensitivity and speed of chimera detection, Bioinformatics, 27(16), 2194, doi:10.1093/bioinformatics/btr381, 2011.
- El-Naggar, M. Y., Wanger, G., Leung, K. M., Yuzvinsky, T. D., Southam, G., Yang, J., Lau, W. M., Nealson, K. H. and Gorby, Y. A.: Electrical transport along bacterial nanowires from Shewanella oneidensis MR-1, doi:10.1073/pnas.1004880107, 2010.
 Eusterhues, K., Rumpel, C. and Kogel-Knabner, I.: Organo-mineral associations in sandy acid forest soils: importance of specific surface area, iron oxides and micropores, Eur. J. Soil Sci., 0(0), 50912034650049, doi:10.1111/j.1365-2389.2005.00710.x, 2005.
- Eusterhues, K., Wagner, F. E., Häusler, W., Hanzlik, M., Knicker, H., Totsche, K. U., Kögel-Knabner, I. and Schwertmann, U.: Characterization of ferrihydrite-soil organic matter coprecipitates by X-ray diffraction and Mössbauer spectroscopy., Environ. Sci. Technol., 42(21), 7891–7 [online] Available from: http://www.ncbi.nlm.nih.gov/pubmed/19031877 (Accessed 1 June 2017), 2008.
- Eusterhues, K., Hädrich, A., Neidhardt, J., Küsel, K., Keller, T. F., Jandt, K. D. and Totsche, K. U.: Reduction of ferrihydrite with adsorbed and coprecipitated organic matter: microbial reduction by Geobacter bremensis vs. abiotic reduction by Nadithionite, Biogeosciences, 11(18), 4953–4966, doi:10.5194/bg-11-4953-2014, 2014.

- Fritzsche, A., Sander, M., Ritschel, T., Steffens, M., Kleiner, R., Frosch, T., Popp, J., Eusterhues, K. and Totsche, K. .: Composition and redox properties of mobile organic matter from anoxic topsoils in comparison to common organic references, Geochim. Cosmochim. Acta. n.d.
- Gorby, Y. A., Yanina, S., McLean, J. S., Rosso, K. M., Moyles, D., Dohnalkova, A., Beveridge, T. J., Chang, I. S., Kim, B. H., Kim, K. S., Culley, D. E., Reed, S. B., Romine, M. F., Saffarini, D. A., Hill, E. A., Shi, L., Elias, D. A., Kennedy, D. W., Pinchuk, G., Watanabe, K., Ishii, S., Logan, B., Nealson, K. H. and Fredrickson, J. K.: Electrically conductive bacterial nanowires produced by Shewanella oneidensis strain MR-1 and other microorganisms., Proc. Natl. Acad. Sci. U. S. A., 103(30), 11358–63, doi:10.1073/pnas.0604517103, 2006.
- Hansel, C. M., Benner, S. G., Nico, P. and Fendorf, S.: Structural constraints of ferric (hydr)oxides on dissimilatory iron reduction and the fate of Fe(II), Geochim. Cosmochim. Acta, 68(15), 3217–3229, doi:10.1016/j.gca.2003.10.041, 2004.
 Hausmann, B., Knorr, K.-H., Schreck, K., Tringe, S. G., Glavina del Rio, T., Loy, A. and Pester, M.: Consortia of low-abundance bacteria drive sulfate reduction-dependent degradation of fermentation products in peat soil microcosms, ISME J., 10(10), 2365–2375, doi:10.1038/ismej.2016.42, 2016.
- Henneberry, Y. K., Kraus, T. E. C., Nico, P. S. and Horwath, W. R.: Structural stability of coprecipitated natural organic matter and ferric iron under reducing conditions, Org. Geochem., 48, 81–89, doi:10.1016/j.orggeochem.2012.04.005, 2012.

 Herrmann, M., Hädrich, A. and Küsel, K.: Predominance of thaumarchaeal ammonia oxidizer abundance and transcriptional activity in an acidic fen, Environ. Microbiol., 14(11), 3013–3025, doi:10.1111/j.1462-2920.2012.02882.x, 2012.
- 675 Sediment Cap and Correlation to Stratified Redox Zones, Environ. Sci. Technol., 43(1), 66–74, doi:10.1021/es801834e, 2009. Hori, T., Müller, A., Igarashi, Y., Conrad, R. and Friedrich, M. W.: Identification of iron-reducing microorganisms in anoxic rice paddy soil by 13C-acetate probing, ISME J., 4(2), 267–278, doi:10.1038/ismej.2009.100, 2010.
 - Hori, T., Aoyagi, T., Itoh, H., Narihiro, T., Oikawa, A., Suzuki, K., Ogata, A., Friedrich, M. W., Conrad, R. and Kamagata, Y.: Isolation of microorganisms involved in reduction of crystalline iron(III) oxides in natural environments., Front.

Himmelheber, D. W., Thomas, S. H., Löffler, F. E., Taillefert, M. and Hughes, J. B.: Microbial Colonization of an In Situ

- 680 Microbiol., 6, 386, doi:10.3389/fmicb.2015.00386, 2015.
 - International Centre for Diffraction Data (ICDD): PDF-4/Organics 2011 (Database), edited by Dr. Soorya Kabekkodu, Newtown Square, PA, USA, 2010.
 - Jambor, J. L. and Dutrizac, J. E.: Occurrence and Constitution of Natural and Synthetic Ferrihydrite, a Widespread Iron Oxyhydroxide, , doi:10.1021/CR970105T, 1998.
- Jiang, J. and Kappler, A.: Kinetics of Microbial and Chemical Reduction of Humic Substances: Implications for Electron Shuttling, Environ. Sci. Technol., 42(10), 3563–3569, doi:10.1021/es7023803, 2008.
 Jones, A. M., Collins, R. N., Rose, J. and Waite, T. D.: The effect of silica and natural organic matter on the Fe(II)-catalysed
 - transformation and reactivity of Fe(III) minerals, Geochim. Cosmochim. Acta, 73(15), 4409–4422, doi:10.1016/j.gca.2009.04.025, 2009.
- 690 Kaiser, K. and Zech, W.: Dissolved organic matter sorption by mineral constituents of subsoil clay fractions, J. Plant Nutr.

- Soil Sci., 163(5), 531–535, doi:10.1002/1522-2624(200010)163:5<531::AID-JPLN531>3.0.CO;2-N, 2000.
- Kappler, A., Benz, M., Schink, B. and Brune, A.: Electron shuttling via humic acids in microbial iron(III) reduction in a freshwater sediment, FEMS Microbiol. Ecol., 47(1), 85–92, doi:10.1016/S0168-6496(03)00245-9, 2004.
- Klindworth, A., Pruesse, E., Schweer, T., Peplies, J., Quast, C., Horn, M. and Glöckner, F. O.: Evaluation of general 16S ribosomal RNA gene PCR primers for classical and next-generation sequencing-based diversity studies., Nucleic Acids Res., 41(1), e1, doi:10.1093/nar/gks808, 2013.
 - Klüpfel, L., Piepenbrock, A., Kappler, A. and Sander, M.: Humic substances as fully regenerable electron acceptors in recurrently anoxic environments, Nat. Geosci., 7(3), doi:10.1038/ngeo2084, 2014.
- Knorr, K.-H. and Blodau, C.: Impact of experimental drought and rewetting on redox transformations and methanogenesis in mesocosms of a northern fen soil, Soil Biol. Biochem., doi:10.1016/j.soilbio.2009.02.030, 2009.
 - Koren, O., Knights, D., Gonzalez, A., Waldron, L., Segata, N., Knight, R., Huttenhower, C. and Ley, R. E.: A Guide to enterotypes across the human body: meta-analysis of microbial community structures in human microbiome datasets, edited by J. A. Eisen, PLoS Comput. Biol., 9(1), 1–16, doi:10.1371/journal.pcbi.1002863, 2013.
- Kotloski, N. J. and Gralnick, J. A.: Flavin Electron Shuttles Dominate Extracellular Electron Transfer by Shewanella oneidensis, doi:10.1128/mBio.00553, 2013.
 - Kraemer, S. M.: Iron oxide dissolution and solubility in the presence of siderophores, Aquat. Sci, 6604, 3–18, doi:10.1007/s00027-003-0690-5, 2004.
 - Küsel, K., Blöthe, M., Schulz, D., Reiche, M. and Drake, H. L.: Microbial reduction of iron and porewater biogeochemistry in acidic peatlands, Biogeosciences, 5(6), 1537–1549, doi:10.5194/bg-5-1537-2008, 2008.
- Lalonde, K., Mucci, A., Ouellet, A. and Gélinas, Y.: Preservation of organic matter in sediments promoted by iron, Nature, 483, doi:10.1038/nature10855, 2012.
 - Lehmann, J. and Kleber, M.: The contentious nature of soil organic matter, Nature, 528, 60–68, doi:10.1038/nature16069, 2015.
- Lentini, C. J., Wankel, S. D. and Hansel, C. M.: Enriched iron(III)-reducing bacterial communities are shaped by carbon substrate and iron oxide mineralogy, Front. Microbiol., 3, 1–19, doi:10.3389/fmicb.2012.00404, 2012.
 - Love, M. I., Huber, W. and Anders, S.: Moderated estimation of fold change and dispersion for RNA-seq data with DESeq2, Genome Biol., 15, doi:10.1186/s13059-014-0550-8, 2014.
 - Lovley, D. R., Woodward, J. C. and Chapelle, F. H.: Stimulated anoxic biodegradation of aromatic hydrocarbons using Fe(III) ligands, Nature, 370(6485), 128–131, doi:10.1038/370128a0, 1994.
- Lovley, D. R., Coates, J. D., Blunt-Harris, E. L., Phillips, E. J. P. and Woodward, J. C.: Humic substances as electron acceptors for microbial respiration, Nature, 382(6590), 445–448, doi:10.1038/382445a0, 1996.
 - Lovley, D. R., Fraga, J. L., Coates, J. D. and Blunt-Harris, E. L.: Humics as an electron donor for anaerobic respiration, Environ. Microbiol., 1(1), 89–98, doi:10.1046/j.1462-2920.1999.00009.x, 1999.
 - Loy, A., Lehner, A., Lee, N., Adamczyk, J., Meier, H., Ernst, J., Schleifer, K.-H. and Wagner, M.: Oligonucleotide microarray

- for 16S rRNA gene-based detection of all recognized lineages of sulfate-reducing prokaryotes in the environment., Appl. Environ. Microbiol., 68(10), 5064–81 [online] Available from: http://www.ncbi.nlm.nih.gov/pubmed/12324358 (Accessed 2 June 2017), 2002.
 - Loy, A., Küsel, K., Lehner, A., Drake, H. L. and Wagner, M.: Microarray and functional gene analyses of sulfate-reducing prokaryotes in low-sulfate, acidic fens reveal cooccurrence of recognized genera and novel lineages., Appl. Environ.
- 730 Microbiol., 70(12), 6998–7009, doi:10.1128/AEM.70.12.6998-7009.2004, 2004.
 - Malvankar, N. S., Vargas, M., Nevin, K. P., Franks, A. E., Leang, C., Kim, B.-C., Inoue, K., Mester, T., Covalla, S. F., Johnson, J. P., Rotello, V. M., Tuominen, M. T. and Lovley, D. R.: Tunable metallic-like conductivity in microbial nanowire networks, Nat. Nanotechnol., 6, 573–579, doi:10.1038/NNANO.2011.119, 2011.
- Marsili, E., Baron, D. B., Shikhare, I. D., Coursolle, D., Gralnick, J. A. and Bond, D. R.: Shewanella secretes flavins that mediate extracellular electron transfer., Proc. Natl. Acad. Sci. U. S. A., 105(10), 3968–73, doi:10.1073/pnas.0710525105, 2008.
 - Martin, M.: Cutadapt removes adapter sequences from high-throughput sequencing reads, EMBnet.journal, 17(1), 10–12, doi:http://dx.doi.org/10.14806/ej.17.1.200, 2011.
 - McMurdie, P. J. and Holmes, S.: Waste Not, Want Not: Why Rarefying Microbiome Data Is Inadmissible, PLoS Comput Biol,
- 740 10(4) [online] Available from: http://journals.plos.org/ploscompbiol/article/file?id=10.1371/journal.pcbi.1003531&type=printable (Accessed 1 September 2017), 2014.
 - McMurdie, P. J., Holmes, S., Kindt, R., Legendre, P. and O'Hara, R.: phyloseq: An R Package for Reproducible Interactive Analysis and Graphics of Microbiome Census Data, edited by M. Watson, PLoS One, 8(4), e61217,
- 745 doi:10.1371/journal.pone.0061217, 2013.
 - Melton, E. D., Swanner, E. D., Behrens, S., Schmidt, C. and Kappler, A.: The interplay of microbially mediated and abiotic reactions in the biogeochemical Fe cycle, Nat. Rev. Microbiol., 12(12), 797–808, doi:10.1038/nrmicro3347, 2014.
 - Mikutta, C., Mikutta, R., Bonneville, S., Wagner, F., Voegelin, A., Christl, I. and Kretzschmar, R.: Synthetic coprecipitates of exopolysaccharides and ferrihydrite. Part I: Characterization, Geochim. Cosmochim. Acta, 72(4), 1111–1127,
- 750 doi:10.1016/j.gca.2007.11.035, 2008.
 - Myers, C. R. and Nealson, K. H.: Bacterial Manganese Reduction and Growth with Manganese Oxide as the Sole Electron Acceptor, Science (80-.)., 240(4857), 1319–1321, doi:10.1126/science.240.4857.1319, 1988.
 - Myers, C. R. and Nealson, K. H.: Respiration-linked proton translocation coupled to anaerobic reduction of manganese(IV) and iron(III) in Shewanella putrefaciens MR-1., J. Bacteriol., 172(11), 6232–8 [online] Available from:
- 755 http://www.ncbi.nlm.nih.gov/pubmed/2172208 (Accessed 27 June 2017), 1990.
 - Narvekar, S. P., Ritschel, T. and Totsche, K. U.: Colloidal stability and mobility of extracellular polymeric substance amended hematite nanoparticles, Vandose Zo. J., 16(8), doi:10.2136/vzj2017.03.0063, 2017.
 - Nevin, K. P. and Lovley, D. R.: Lack of production of electron-shuttling compounds or solubilization of Fe(III) during

- reduction of insoluble Fe(III) oxide by Geobacter metallireducens., Appl. Environ. Microbiol., 66(5), 2248–51, doi:10.1128/AEM.66.5.2248-2251.2000, 2000.
 - Nevin, K. P. and Lovley, D. R.: Mechanisms for Fe(III) Oxide Reduction in Sedimentary Environments, Geomicrobiol. J., 19(2), 141–159, doi:10.1080/01490450252864253, 2002.
 - Newman, D. K. and Kolter, R.: A role for excreted quinones in extracellular electron transfer, Nature, 405(6782), 94–97, doi:10.1038/35011098, 2000.
- Oksanen, J., Blanchet, F. G., Friendly, M., Kindt, R., Legendre, P., McGlinn, D., Minchin, P. R., O'Hara, R. B., Simpson, G. L., Solymos, P., Stevens, H., Szöcs, E. and Wagner, H. H.: vegan: Community Ecology Package. Ordination methods, diversity analysis and other functions for community and vegetation ecologists. Version 2.4-3, , 1–293 [online] Available from: https://cran.r-project.org/package=vegan (Accessed 15 June 2017), 2017.
- Pédrot, M., Boudec, A. Le, Davranche, M., Dia, A. and Henin, O.: How does organic matter constrain the nature, size and availability of Fe nanoparticles for biological reduction?, J. Colloid Interface Sci., 359(1), 75–85, doi:10.1016/j.jcis.2011.03.067, 2011.
 - Pester, M., Bittner, N., Deevong, P., Wagner, M. and Loy, A.: A "rare biosphere" microorganism contributes to sulfate reduction in a peatland, ISME J., 4(12), 1591–1602, doi:10.1038/ismej.2010.75, 2010.
 - Pester, M., Knorr, K.-H., Friedrich, M. W., Wagner, M. and Loy, A.: Sulfate-reducing microorganisms in wetlands fameless actors in carbon cycling and climate change, Front. Microbiol., 3, 72, doi:10.3389/fmicb.2012.00072, 2012.
 - Piepenbrock, A., Schröder, C. and Kappler, A.: Electron Transfer from Humic Substances to Biogenic and Abiogenic Fe(III) Oxyhydroxide Minerals, Environ. Sci. Technol., 48(3), 1656–1664, doi:10.1021/es404497h, 2014.
 - Poggenburg, C., Mikutta, R., Sander, M., Schippers, A., Marchanka, A., Dohrmann, R. and Guggenberger, G.: Microbial reduction of ferrihydrite-organic matter coprecipitates by Shewanella putrefaciens and Geobacter metallireducens in comparison to mediated electrochemical reduction, Chem. Geol., 447, 133–147, doi:10.1016/j.chemgeo.2016.09.031, 2016.
- Reguera, G., McCarthy, K. D., Mehta, T., Nicoll, J. S., Tuominen, M. T. and Lovley, D. R.: Extracellular electron transfer via microbial nanowires, Nature, 435(7045), 1098–1101, doi:10.1038/nature03661, 2005.

780

4, doi:10.7717/peerj.2584, 2016.

- Reiche, M., Torburg, G. and Küsel, K.: Competition of Fe(III) reduction and methanogenesis in an acidic fen, FEMS Microbiol. Ecol., 65(1), 88–101, doi:10.1111/j.1574-6941.2008.00523.x, 2008.
- Roden, E. E., Kappler, A., Bauer, I., Jiang, J., Paul, A., Stoesser, R., Konishi, H. and Xu, H.: Extracellular electron transfer through microbial reduction of solid-phase humic substances, Nat. Geosci., 3(6), 417–421, doi:10.1038/ngeo870, 2010.
 Rognes, T., Flouri, T., Nichols, B., Quince, C. and Mahé, F.: VSEARCH: a versatile open source tool for metagenomics, PeerJ,
- Royer, R. A., Burgos, W. D., Fisher, A. S., Jeon, B.-H., Unz, R. F. and Dempsey, B. A.: Enhancement of Hematite Bioreduction by Natural Organic Matter, doi:10.1021/ES015735Y, 2002.
 - Schmieder, R. and Edwards, R.: Quality control and preprocessing of metagenomic datasets, Bioinformatics, 27(6), 863–864, doi:10.1093/bioinformatics/btr026, 2011.

- Schwertmann, U., Wagner, F. and Knicker, H.: Ferrihydrite-Humic Associations, Soil Sci. Soc. Am. J., 69(4), 1009, doi:10.2136/sssaj2004.0274, 2005.
- Scott, D. T., McKnight, D. M., Blunt-Harris, E. L., Sarah E. Kolesar and Lovley, D. R.: Quinone Moieties Act as Electron Acceptors in the Reduction of Humic Substances by Humics-Reducing Microorganisms, doi:10.1021/ES980272Q, 1998. Shannon, C. E.: A Mathematical theory of communication, Bell Syst. Tech. J., 27(3), 379–423, doi:10.1002/j.1538-7305.1948.tb01338.x, 1948.
- Shelobolina, E. S., Nevin, K. P., Blakeney-Hayward, J. D., Johnsen, C. V., Plaia, T. W., Krader, P., Woodard, T., Holmes, D. E., VanPraagh, C. G. and Lovley, D. R.: Geobacter pickeringii sp. nov., Geobacter argillaceus sp. nov. and Pelosinus fermentans gen. nov., sp. nov., isolated from subsurface kaolin lenses, Int. J. Syst. Evol. Microbiol., 57(1), 126–135, doi:10.1099/ijs.0.64221-0, 2007.
 - Shi, L., Rosso, K. M., Clarke, T. A., Richardson, D. J., Zachara, J. M. and Fredrickson, J. K.: Molecular Underpinnings of Fe(III) Oxide Reduction by Shewanella Oneidensis MR-1., Front. Microbiol., 3, 50, doi:10.3389/fmicb.2012.00050, 2012.
- Shimizu, M., Zhou, J., Schröder, C., Obst, M., Kappler, A. and Borch, T.: Dissimilatory Reduction and Transformation of Ferrihydrite-Humic Acid Coprecipitates, Environ. Sci. Technol., 47(23), 13375–13384, doi:10.1021/es402812j, 2013.
 Stumm, W. and Sulzberger, B.: The cycling of iron in natural environments: Considerations based on laboratory studies of heterogeneous redox processes, Geochim. Cosmochim. Acta, 56(8), 3233–3257, doi:10.1016/0016-7037(92)90301-X, 1992.
 Tamura, H., Goto, K., Yotsuyanagi, T. and Nagayama, M.: Spectrophotometric determination of iron(II) with 1,10-phenanthroline in the presence of large amounts of iron(III), Talanta, 21(4), 314–318, doi:10.1016/0039-9140(74)80012-3,
 - Torn, M. S., Trumbore, S. E., Chadwick, O. A., Vitousek, P. M. and Hendricks, D. M.: Mineral control of soil organic carbon storage and turnover, Nature, 389(6647), 170–173, doi:10.1038/38260, 1997.
- Venkateswaran, K., Moser, D. P., Dollhopf, M. E., Lies, D. P., Saffarini, D. A., MacGregor, B. J., Ringelberg, D. B., White,
 D. C., Nishijima, M., Sano, H., Burghardt, J., Stackebrandt, E. and Nealson, K. H.: Polyphasic taxonomy of the genus Shewanella and description of Shewanella oneidensis sp. nov., Int. J. Syst. Bacteriol., 49(2), 705–724, doi:10.1099/00207713-49-2-705, 1999.
 - Vindedahl, A. M., Strehlau, J. H., Arnold, W. A. and Penn, R. L.: Organic matter and iron oxide nanoparticles: aggregation, interactions, and reactivity, Environ. Sci. Nano, 3(3), 494–505, doi:10.1039/C5EN00215J, 2016.
- Weber, K. A., Urrutia, M. M., Churchill, P. F., Kukkadapu, R. K. and Roden, E. E.: Anaerobic redox cycling of iron by freshwater sediment microorganisms., Environ. Microbiol., 8(1), 100–13, doi:10.1111/j.1462-2920.2005.00873.x, 2006a. Weber, K. A., Achenbach, L. A. and Coates, J. D.: Microorganisms pumping iron: anaerobic microbial iron oxidation and reduction, Nat. Rev. Microbiol., 4(10), 752–764, doi:10.1038/nrmicro1490, 2006b.

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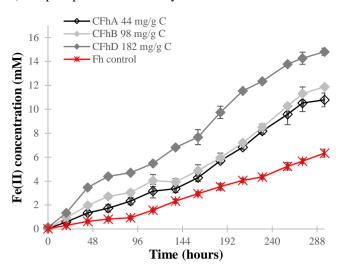
| | | | | | S. on | eidensis | MR-1 | microbial co | | sortia | G. bremensis* | | |
|----------------|------|--------------------|--------------------|-----------------------|--------------------|----------|-------------------------------|------------------|----------------|-------------------------------|--------------------|----------------|-------------------------------|
| | | | | | linear fit | | | linear fit | | | linea | ır fit | |
| | | С | Fe | C/Fe | k | r² | Fe(II) at T _{end} | k | r ² | Fe(II) at T _{end} | k | r ² | Fe(II) at T _{end} |
| | | mg g ⁻¹ | mg g ⁻¹ | mol mol ⁻¹ | mM h ⁻¹ | | % of Fe | ${\rm mMh^{-1}}$ | | % of Fe | mM h ⁻¹ | | % of Fe |
| control | Fh | - | 560 | - | 0.022 | 0.98 | 10 | 0.031 | 0.95 | 15 | 0.119 | 0.96 | 59 |
| adsorbed OM | AFhA | 44 | 527 | 0.39 | 0.032 | 0.96 | 15 | 0.037 | 0.97 | 17 | 0.148 | 0.96 | 55 |
| | AFhB | 105 | 468 | 1.04 | 0.038 | 0.98 | 14 | 0.044 | 0.97 | 16 | 0.096 | 0.96 | 35 |
| | AFhD | 181 | 342 | 2.46 | 0.041 | 0.97 | 16 | 0.050 | 0.96 | 19 | 0.088 | 0.94 | 32 |
| coprecip. | CFhA | 44 | 498 | 0.41 | 0.038 | 0.97 | 17 | 0.046 | 0.98 | 19 | 0.142 | 0.93 | 66 |
| | CFhB | 98 | 428 | 1.06 | 0.039 | 0.97 | 17 | 0.048 | 0.95 | 20 | 0.142 | 0.97 | 49 |
| | CFhD | 182 | 299 | 2.83 | 0.051 | 0.99 | 20 | 0.058 | 0.97 | 22 | 0.091 | 0.95 | 39 |

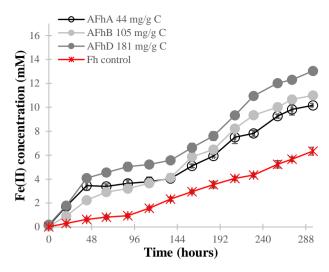
^{*} data from Eusterhues et al., (2014)

Table 1. C and Fe concentration and molar C/Fe ratio of ferrihydrite-organic matter associations and results of microbial reduction experiments (reduction rate k and Fe(II) concentration after 296.5 hours of incubation in % of total Fe).

A) Coprecipitated OM-Ferrihydrite

B) Adsorbed OM-Ferrihydrite



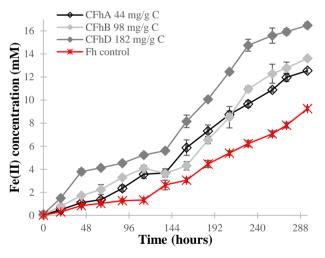


C) Coprecipitated OM-Ferrihydrite

855

860

D) Adsorbed OM-Ferrihydrite



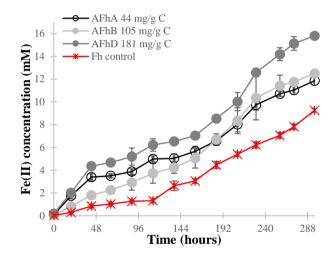
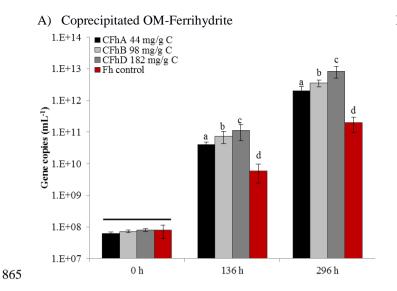
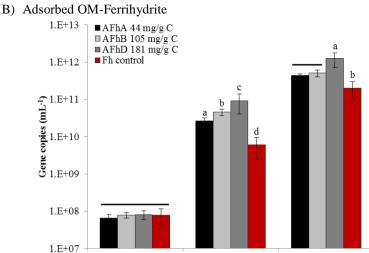


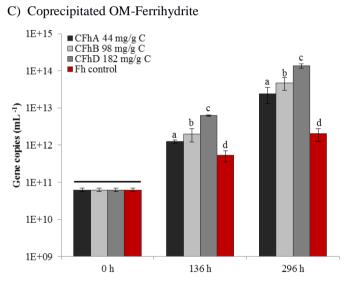
Figure 1. Microbial reduction <u>ferrihydrite</u> and <u>organic matter</u> (OM) <u>ferrihydrite complexes-of by Shewanella oneidensis MR-1 and a microbial consortia extracted from peatland. Panels A and B depict S. oneidensis MR-1-mediated reduction of ferrihydrite (Fh control), and coprecipitated OM-ferrihydrite (CFh) (A), and adsorbed OM-ferrihydrite (AFh) (B). Panels C and D depict reduction of ferrihydrite (Fh control), coprecipitated OM-ferrihydrite (CFh) (C), and adsorbed OM-ferrihydrite (AFh) (D) in incubations with a microbial consortia extracted from peatland organic matter (OM) ferrihydrite complexes (coprecipitated OM ferrihydrite: CFh; adsorbed OM ferrihydrite: AFh) in incubations with Shewanella oneidensis MR 1. Error bars represent standard deviation in triplicate incubations.</u>

27



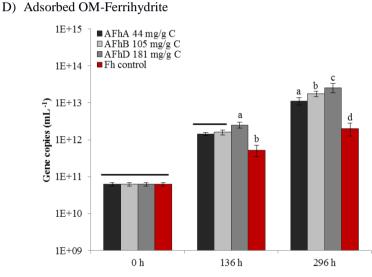


0 h



870

875



136 h

296 h

Figure 23. Abundances of <u>bacterial</u> *S. oneidensis* MR-1-16S rRNA gene in subsamples taken from ferrihydrite incubations containing <u>either</u> *S. oneidensis* MR-1 <u>or peat microbial consortia extracted from peat</u> at T₀ (0 h), T_{middle} (136 h), T_{end} (296 h). Bars represent mean abundances of triplicate samples from <u>incubations</u> with <u>S. oneidensis</u> MR-1 incubations with coprecipitated OM-ferrihydrite (CFh) (A) or adsorbed OM-ferrihydrite (AFh) (B) and peat microbial consortia incubations with coprecipitated OM-ferrihydrite (CFh) (C) or adsorbed OM-ferrihydrite (AFh) (D). Each sample was measured in triplicate qPCR reactions with applicable standard deviations. One-way ANOVA was applied to compare individual datasets for statistical significant differences between ferrihydrite incubations. Different letters (a-d) indicate statistically significant differences, while horizontal black lines indicate non-significant differences.

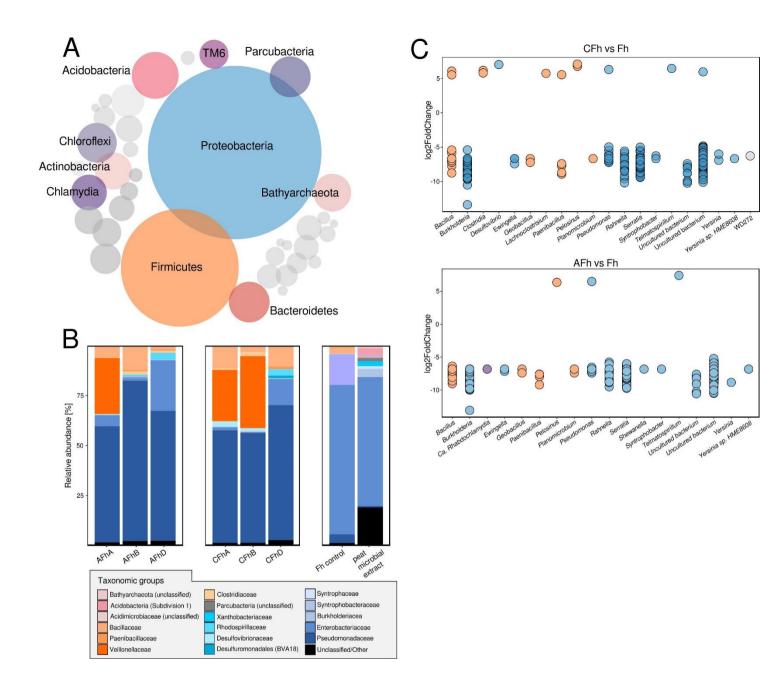


Figure 53. Microbial community profiling from incubations with coprecipitated OM-ferrihydrite (CFh) or adsorbed OM-ferrihydrite (AFh) compared to microbial communities with pure ferrihydrite (Fh control) at the end of incubation and microbial consortia extracted from peatland at beginning of incubation (peatland microbial extract). Bubble plot indicating the relative abundance of prokaryotic phyla when individual datasets are pooled (A). Bubbles are sized according to respective relative abundances. The ten most abundant phyla are highlighted by color. Grey bubbles refer to less abundant phyla. Family-

level taxonomic profiles are given for individual datasets. Family-level groups showing abundances less than 1%, as well as unclassified OTUs are summarized as "Unclassified/Other"(B). Differentially abundant genus-level groups in sample-wise comparisons were identified assuming negative binomial distributions for read count data and applying Wald significance tests, and parametric fitting for determined dispersions. Overrepresented genus level groups were identified by negative binomial distributions. Tests were carried using ferrihydrite complexation as decisive parameter. Multiple inference correction was done according to Benjamini-Hochberg (1995) (Benjamini and Hochberg, 1995). *p*-values less than 0.05 were considered significant (C). The color scheme in panels B and C is according to panel A. Family-level groups belonging to the Firmicutes are for instance shown in shades of orange in case of panel B. In panel B, AFhA-C refers to increasing TOC concentrations (44 mg/g C, 105 mg/g C, 181 mg/g C) in adsorbed OM-ferrihydrite incubations and CFhA-C refers to increasing TOC concentrations (TOC: 44 mg/g C, 98 mg/g C, 182 mg/g C) in incubations with coprecipitated ferrihydrite.

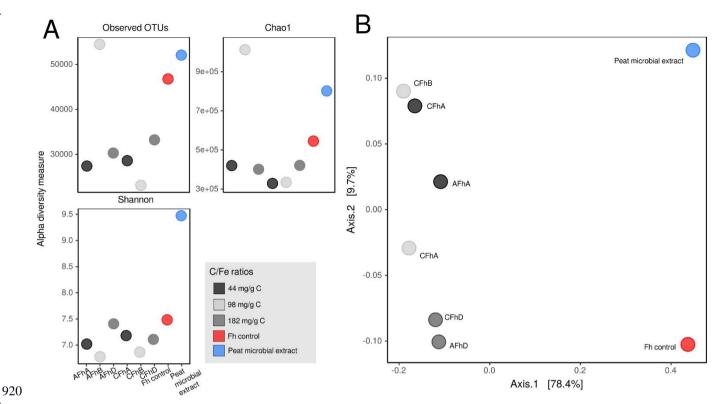


Figure 46. Alpha- and beta-diversity analyses. Alpha diversity analysis was done using various indices indicative for diversity (Observed OTUs, Chao1, ACE) and evenness (Shannon, Simpson) (A). Alpha diversity in general describes within dataset diversity. Observed OTUs refers to OTU counts based on the OTU definition given in the materials and methods._-Chao1-and ACE (Chao, 1984) is an index , are two indices that puts special emphasis on rare taxa to extrapolate the real number of OTUs based on observed OTUs. The Shannon index (Shannon, 1948) describes the entropy in a community profile, and reaches its maximum when taxonomic groups are evenly distributed in a dataset. The Simpson index (Simpson, 1949) refers to the probability that two randomly sampled sequences from a dataset belong to the same OTU. Beta diversity (= differences in diversity between samples) analysis was based on Jensen-Shannon divergences (Koren et al., 2013) (B).

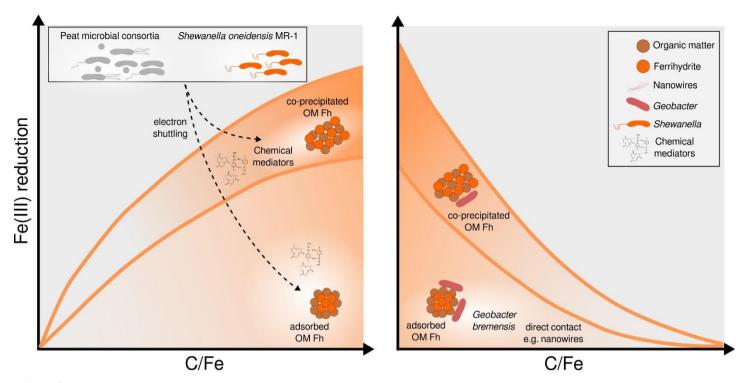


Figure 57. Microbial reduction of ferrihydrite with both adsorbed and coprecipitated OM by Shewanella oneidensis MR-1 and the microbial community consortia extracted from Schlöppnerbrunnen peat samples. For comparison, microbial reduction of ferrihydrite with both adsorbed and coprecipitated OM by *Geobacter bremensis* as described in Eusterhues, et al. 2014 is depicted in the right panel.