



1 **Modeling impacts of climate change and grazing effects on plant**
2 **biomass and soil organic carbon in the Qinghai–Tibetan**
3 **grasslands**

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13 **Abstract:**

14 The Qinghai Province supports over 40% of the human population but occupies about 29% of
15 the land area, and thus it plays an important role in the entire Qinghai–Tibetan Plateau (QTP).

16 The dominant land cover is grassland, which has been severely degraded over the last decade
17 due to a combination of increased human activities and climate change. Numerous studies

18 indicate that the plateau is sensitive to recent global climate change, but the drivers and
19 consequences of grassland ecosystem change are controversial, especially the effects of climate

20 change and grazing patterns on the grassland biomass and soil organic carbon (SOC) storage in
21 this region. In this study, we used the DeNitrification-DeComposition (DNDC) model and two

22 climate change scenarios (representative concentration pathways: RCP4.5 and RCP8.5) to
23 understand how the grassland biomass and SOC pools might respond to different grazing

24 intensities under future climate change scenarios. More than 1400 grassland biomass sampling
25 points and 46 SOC points were collected, which were then used to validate the simulated results.

26 The results showed that compared with the past 30 years, the biomass and SOC exhibited a
27 significant decreasing trend under all grazing intensities in the RCP4.5 and RCP8.5 scenarios,

28 and RCP8.5 had a more negative future effect on the biomass compared with RCP4.5. Thus,
29 future climate change could lead to greater temporal and spatial variations in the grassland



30 biomass and SOC. Overall, climate change may be the major factor that leads to fluctuations in
31 the inter-annual grassland biomass on the Qinghai Province, where the grazing intensity has
32 significantly affected the grassland vegetation dynamics. Therefore, urgent ecological
33 conservation of vulnerable grassland ecosystems is required to effectively regulate grazing
34 practices.

35 **Keywords:** Biogeochemical process; DNDC; Grazing intensity; Grassland management;
36 Degradation;

37 1 Introduction

38 Grassland is one of the most widespread terrestrial ecosystems and accounts for nearly 33% of
39 the land without ice cover (Ellis and Ramankutty, 2008), where it plays important roles in both
40 the global carbon cycle and terrestrial ecosystem processes (Li et al., 2013c). The Qinghai-
41 Tibetan Plateau (QTP) about 130 million hectares (ha), it counted 44% of China's grassland
42 area (Li et al., 2013a; Piao et al., 2012). This area plays a vital role for the ecological services
43 of China and Southeast Asian countries (Harris, 2010; Li et al., 2013b; Piao et al., 2012; Wang
44 et al., 2002; Zeng et al., 2015). Qinghai Province supports over 40% of the population but it has
45 about 29% of the total area, and thus it plays an important role in the whole QTP (Li et al.,
46 2013a; Piao et al., 2012). This area is recognized as one of the most ecologically fragile and
47 sensitive areas to global climate change and human disturbance (Harris, 2010; Li et al., 2013b;
48 Piao et al., 2012; Wang et al., 2002; Zeng et al., 2015). Moreover, this area is also the largest
49 animal husbandry production region in China, and it also contains the headwaters of the two
50 major rivers in China, i.e., the Yellow River and the Yangtze River, and thus it plays a vital role
51 in ecological conservation in China (Zeng et al., 2015).

52 In recent decades, due to climate change, increased human disturbances, the high altitude alpine
53 grassland ecosystems, which are the dominant grassland vegetation type, have been severely
54 degraded (Gao et al., 2010). The air temperature on the plateau has increased by 0.3°C per
55 decade, which is three times the global average (Li et al., 2008). Warming could significantly
56 increase the net primary productivity of alpine meadows (Chen et al., 2013; Du et al., 2004;
57 Fan et al., 2010). Other studies have found that warming also speeds up the decomposition rate



58 for litter and manure, and increases soil respiration (Luo et al., 2010; Xu et al., 2010), which
59 could cause significant losses of soil organic carbon (SOC) and affect the alpine grassland
60 ecosystem carbon pool balance (Pei et al., 2009; Tan et al., 2010). Although the ecological
61 impact of warming on the QTP alpine grassland ecosystem has not been fully elucidated in
62 previous studies, there is no doubt that warming will greatly accelerate the key processes in the
63 alpine grassland ecosystem carbon cycle (Luo et al., 2010). There are reported that both
64 precipitation amount and the number of precipitation days have increased significantly in QTP
65 (Li et al., 2010). As precipitation is another crucial climate factor in controlling the carbon cycle
66 of grassland ecosystems, how the higher variability precipitation impacts the SOC and biomass
67 in QTP need further investigation (Lehnert et al., 2016; Maussion et al., 2014).

68 Grazing is the most important biotic factor among the ecological processes that affect rapid
69 changes in the vegetation and soil, and it is the main method for deriving ecosystem services
70 from the QTP grassland (Tanentzap and Coomes, 2012). Moreover, grazing is one of the major
71 human disturbances to the grassland in this area. In general, overgrazing is considered to be one
72 of the main causes of carbon and nitrogen losses from the soil, thereby contributing to the
73 unsustainable use of grassland (McIntire and Hik, 2005). Therefore, sustaining a reasonable
74 grazing intensity has an indispensable role in maintaining the turnover of soil nutrients and
75 plant community stability (Klein et al., 2007).

76 Previous studies have shown that different types of vegetation and soil nutrient pools exhibit
77 significantly different responses to variations in the grazing intensity (Lavado et al., 1996).
78 However, there is still a lack of robust studies to evaluate the combined effect of grazing and
79 climate change, as well as their impact on the QTP grassland ecosystem at a large scale. Due to
80 the unique geographic characteristics and important ecological functions of the QTP grassland
81 ecosystem, it is necessary to evaluate the impacts of human management and climate change to
82 ensure that it continues to provide these ecosystem services.

83 In this study, using a well-calibrated DeNitrification-DeComposition (DNDC) model based on
84 long-term vegetation observations, we evaluated the response of the grassland ecosystem in
85 Qinghai Province in terms of both climate change and human management by analyzing the
86 grazing intensity. We also analyzed the interactions between grassland vegetation and soil
87 carbon storage with grazing intensity and climate change disturbances at a large scale in long-



88 term impact assessments.

89 **2 Materials and methods**

90 **2.1 Study area**

91 Qinghai Province (89°35′–103°04′ E, 31°39′–39°19′ N) is located in the northeast of QTP in
92 China (Fig. 1). This region has a typical plateau climate, with a mean annual temperature of
93 8.6°C (from –6°C to 9°C) and a mean annual precipitation of 424.7 mm (16.7–776.1 mm). In
94 general, the climate is cold and dry. The altitude of Qinghai province ranges between 1,650–
95 6,860 meters above sea level (m a.s.l.) and 67% of the land area is in the range of 3,000–5,000
96 masl. Grassland is the major land cover in the study area where alpine meadow and alpine
97 steppe are the dominant vegetation types, where they account for 60.5% of the total grassland
98 area.

99 Grazing is the primary human activity in the study area and livestock production is a key
100 industry in this region. Generally, natural grassland is the major food source for the livestock
101 in the QTP. Compared with 1949, the number of livestock has increased by almost three times
102 from 7.49×10^6 (Zhang, 2011) to the peak number 22.19×10^6 head in 2005 at the study area
103 (QPBS, 2015, 2005).

104 Since 2004, the Chinese government has implemented a series of ecological protection projects
105 and policies in Qinghai province, including reducing livestock and prohibiting grazing, building
106 fences to allow natural grassland recovery, as well as providing allowances and awards to local
107 herdsmen families to promote degraded pasture recovery and to balance the livestock rate
108 according to the forage productivity (Zeng et al., 2015). The core objective of these projects
109 and policies is changing the grazing intensity and achieving a balance between the livestock
110 intensity and grassland regenerability in order to construct a sustainable grassland ecosystem.
111 Due to new policies for ecological protection, the livestock numbers have declined in recent
112 years, but they have been maintained at the 2015 level of 19.42×10^6 head (supplementary
113 Table S1) (QPBS, 2015).



114 **2.2 DNDC model**

115 The DNDC 9.5 biogeochemical model, which was downloaded from the official web
116 (<http://www.dndc.sr.unh.edu/>), was employed in this study (Li et al., 2006; Li et al., 1992). The
117 model has been used widely in more than 20 countries to obtain accurate calibration and
118 verification results in various ecosystems (Abdalla et al., 2009; Chen et al., 2015;
119 Kariyapperuma et al., 2011; Li et al., 1996; Li et al., 2017; Li et al., 2014; Liu et al., 2006; Xu
120 et al., 2003; Zhang and Niu, 2016; Zhao et al., 2016).

121 The model has two major components. The first component can simulate the soil environmental
122 conditions, where it includes soil climate, vegetation growth, and decomposition submodels.
123 The second component includes three submodels for simulating nitrification, denitrification,
124 and fermentation processes, which are used to simulate biogeochemical production,
125 consumption, and emissions of CH₄, N₂O, NO, and NH₃, net ecosystem exchanges of CO₂, as
126 well as carbon and nitrogen losses due to leaching (Zhang et al., 2015).

127 The DNDC model simulates vegetation growth by tracking photosynthesis, respiration, water
128 demand, N demand, C allocation, crop yield, and litter production. The model predicts the SOC
129 dynamics mainly by quantifying the SOC input from crop litter incorporation and manure
130 amendment, as well as the SOC output through decomposition. More detailed information
131 about the model was given by Li (1996).

132 **2.3 Regional database**

133 In order to characterize the spatial heterogeneity of natural grasslands in the study area, we
134 collected the following geospatial data as inputs for the DNDC biogeochemical model:
135 grassland type and spatial distribution (Fig. 1), soil properties, and climate data.

136 **Grassland Database**

137 The vegetation parameters in the model were obtained from a grassland field monitoring project
138 implemented during 2005–2014 (ERSMC-a, 2016; ERSMC-b, 2016). This annual monitoring
139 project covered the major types of grassland within the project area. On average, 168
140 monitoring sites were sampled each year. For each monitoring site, the average value based on



141 six replicate sampling points was calculated to determine the aboveground biomass value for
142 the monitoring site. The aboveground biomass harvests used the quadrat method during the
143 plant growing season (July 10–August 20) in a 1 m × 1 m plot. A more detailed description of
144 the sampling method used to obtain the observation data can be found in reports by the
145 Ecological Environment Remote Sensing Monitoring Center of Qinghai Province (ERSMC-a,
146 2016; ERSMC-b, 2016). The grassland simulation based on the grassland functional group type
147 was categorized according to the grassland type map for the study area (Fig. 1). The detailed
148 grassland parameters used in the model were shown in Supplementary Table S2.

149 **Soil Database**

150 We used a 1:1,000,000 scale soil database developed by the Institute of Soil Science, Chinese
151 Academy of Sciences, which was compiled based on the second national soil survey conducted
152 in 1979–1994 for all the counties in China (Shi et al., 2004). The database had three attributes:
153 locations, soil attributes, and reference systems. It contained multi-layer soil properties (e.g.
154 organic matter, pH, and bulk density), soil texture (e.g. sand, silt and clay proportions), and
155 spatial information (Shi et al., 2004; Yu et al., 2007a; Yu et al., 2007b), which were used in the
156 model simulations.

157 **Climate Database**

158 Daily climate data were obtained from the China Meteorological Network for the study period,
159 and there were 39 stations inside the study areas (<http://data.cma.cn/>). The daily precipitation
160 and maximum/minimum temperatures between 1985–2014 were interpolated at 1-km
161 resolution grid for our model. Regression kriging and the inverse distance method were
162 employed for air temperature and precipitation interpolation, respectively (Fortin and Dale,
163 2005; Hengl et al., 2007).

164 **Model implementation**

165 All datasets were processed with ArcGIS version 10.2 (ESRI, Redlands, CA) to the formation
166 a georeferenced DNDC regional simulation database. The data processing flowchart could be
167 found in the supplementary Fig. S1. The county boundary data intersected with grassland type
168 map to formation the model simulation unit, meanwhile, the county based grazing intensity, soil



169 properties and climate information also assigned to the model simulation units. The DNDC was
170 running with regional simulation database based on individual model simulation units. The
171 detailed information of how to run the model could be found in Li (2012). The actual climate,
172 soil, grassland type and grazing intensity as the simulation baseline.

173 **2.4 Simulation scenarios**

174 **Grazing simulation scenarios**

175 The grazing period is all-year round and cattle (90% yaks), sheep, and goats are major livestock
176 types, while horses are a minor component in the study area. The grazing intensity data were
177 based on the annual national livestock statistical report provided by the National Bureau of
178 Statistics of China and the Bureau of Statistics for Qinghai Province. The detailed grazing data
179 are shown in Supplementary Table S3. In the DNDC model, grazing activity is defined by
180 specifying the grazing parameters, including the livestock type, grazing period, and grazing
181 intensity. The detailed parameters for the DNDC grazing model are shown in Supplementary
182 Table S4. The grazing intensity is defined according to Eq. 1 based on the grazing area in each
183 administrative region (Li et al., 2014):

$$GI = LP/GA, \quad (\text{Eq.1})$$

184 where GI is the grazing intensity (head ha^{-1}), LP is the livestock unit (head), and GA is the
185 grazing area (ha).

186 In order to test the responses of the grassland biomass and soil SOC to various grazing
187 intensities, we tested the following treatments: baseline, grazing intensity based on the actual
188 grazing intensity in 2005; G_0 , grazing intensity of zero; G_{-50} , 50% of the baseline intensity; and
189 G_{+50} , 50% higher than the baseline.

190 **Climate change scenarios**

191 The Intergovernmental Panel on Climate Change (IPCC) Fifth Report employed new stable
192 concentration-based scenarios in representative concentration pathways (RCPs) to project
193 future climate change (IPCC, 2013). The development of the RCP scenarios used a parallel
194 method, which combined climate, air, and the carbon cycle with emissions and the socio-
195 economic situation to assess the impact of climate change on a study area, as well as adaptation,



196 vulnerability, and mitigation analysis (Moss et al., 2010). The RCPs were named according to
197 their 2100 radiative forcing level and reported by individual modeling teams, i.e., 2.6–8.5 W/m².
198 The RCPs comprise four scenarios, i.e., RCP2.6, RCP4.5, RCP6.0, and RCP8.5 (Moss et al.,
199 2010). Each scenario provides a path affected by social and economic conditions and climate,
200 and each projection corresponds to the radiation force value predicted by 2100.

201 We considered RCP4.5 and RCP8.5 because these two scenarios have been used widely to
202 evaluate the potential impact of climate change on the environment (Di Vittorio et al., 2014; Li
203 et al., 2015; van Vuuren et al., 2011; Zhang et al., 2013). RCP4.5 represents a medium-low RCP
204 with stabilization of CO₂ emissions from 2150 onwards, and RCP8.5 represents a high RCP
205 with stabilizing CO₂ emissions post-2100 (Meinshausen et al., 2011). The projected climate
206 conditions in the present study under RCP4.5 and RCP8.5 were derived from the average values
207 of 25 Coupled Model Intercomparison Project Phase 5 (CMIP5) global climate models (Fu and
208 Feng, 2014).

209 Compared with 2014, the average temperature and precipitation increased by 0.72°C and
210 0.80°C, and by 11.81 mm and 12.50 mm under RCP4.5 and RCP8.5 in 2044, respectively, in
211 the study area (Table 1). The changes in the spatial distribution of precipitation are shown in
212 Supplementary Fig. S2. The pattern of increased precipitation was similar using RCP4.5 and
213 RCP8.5 for the period of 2014–2044, where it increased in the whole area and it increased
214 gradually from the north to the south of the study area. However, RCP8.5 obtained a higher
215 increase than RCP4.5 and the southwest part of the research area is projected to have a higher
216 temperature increase than the other regions. Moreover, the annual average temperature had a
217 similar distribution under the two climate change scenarios, where the temperature increase
218 using RCP4.5 (Supplementary Fig. S2c) was lower than that with RCP8.5 (Supplementary Fig.
219 S2d).

220 Three different periods were considered in the grassland simulations. First, a pretreatment
221 (1961–1984) period was used to initialize the soil climate conditions and SOC composition.
222 The pretreatment period represented the baseline climate with no increases in CO₂ or climate
223 change. The second period represented realistic climate scenarios (1985–2014) based on the
224 most recent climate. The third period comprised future climate scenarios (2015–2044), which
225 represented two future climates (RCP4.5, RCP8.5) scenarios with changes in temperature and



226 precipitation.

227 **2.5 Model validation**

228 The root mean squared error (RMSE) (Eq.2), coefficient of determination (R^2) (Eq.3) and model
229 efficiency (ME) (Eq.4) were employed for model validation. The RMSE estimates the scatter
230 between the simulated and measured data, where values close to zero indicate excellent
231 agreement and hence the good performance of the model (Araya et al., 2015). R^2 is used to test
232 the agreement between the modeled results and observations, where a value closer to 1 indicates
233 that the model provides a better explanation for the observed values (Willmott, 1982). The
234 positive ME value indicates that the model prediction is better than the mean of observations,
235 and the best model performance has ME value equal to 1 (Miehle, 2006). RMSE, R^2 and ME
236 were calculated as follows:

$$\text{RMSE} = \sqrt{\frac{\sum_{i=1}^n (P_i - O_i)^2}{n}} \quad (\text{Eq. 2})$$

$$R^2 = \left[\frac{\sum_{i=1}^n (O_i - \bar{O})(P_i - \bar{P})}{\sqrt{\sum_{i=1}^n (O_i - \bar{O})^2} \sqrt{\sum_{i=1}^n (P_i - \bar{P})^2}} \right]^2, \quad (\text{Eq. 3})$$

$$\text{ME} = 1 - \frac{\sum_{i=1}^n (P_i - O_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (\text{Eq. 4})$$

237 where P_i and O_i were modeled and observed values, and \bar{P} and \bar{O} are their averages. n is the
238 number of values.

239 The validation dataset included more than 1400 grassland biomass sampling points, which
240 covered the whole of the study area, and the field measurements were also fully representative
241 of the major grassland types in this area. In addition, 46 SOC observation points were sampled
242 between 2011–2012, which were randomly distributed among all of the simulation units
243 (county and grassland types). The grassland biomass was sampled in quadrat (1 m × 1m) with
244 3 replicates between mid-July and mid-August. Maximum biomass in each quadrat was
245 harvested and dried in an oven at 70 °C for 72 h, weighed and ground for analysis. The soil of
246 0–30 cm depth was sampled at 10-cm intervals with a soil drill (metal cylinder: diameter of 5
247 cm, length of 20 cm and the total length of the sampler 1.3 m). 3 samples were collected in each
248 replication plot. The ground soil samples passed a 0.15-mm sieve and wet oxidation method



249 was applied to determine SOC (Mebius, 1960). In general, every simulation unit had 1–2
250 validation points (ERSMC-a, 2016).

251 **2.6 Statistical analysis**

252 Two-way analysis of variance (ANOVA) was used to test the effects of climate and grazing
253 intensity on both the biomass and SOC. Mean values for the same treatments were compared
254 using Fisher's least significant difference (LSD) test with one-way ANOVA at $P = 0.05$. The
255 statistical analyses, including the test for normality (Shapiro-Wilk) and homogeneity of
256 variance (Levene), were performed using Origin 2016 version b9.3.1.273 (OriginLab
257 Corporation, MA, USA), and the multiple regression analysis was conducted with the Minitab
258 version 17 (Minitab Inc., State College, PA, USA).

259 **3 Results**

260 **3.1 Model validation**

261 The biomass simulation showed that the modeled total biomass was in good agreement with
262 the observations (Fig. 2). There was a significant linear relationship ($P < 0.001$) between the
263 measurements and the modeled total biomass ($R^2=0.71$, $ME=0.75$, $RMSE = 93.11 \text{ g C m}^{-2}$; P
264 < 0.001). The simulated SOC concentrations were in good agreement with the measured data
265 (Fig. 3). The calculated statistical indices indicated that the modeled SOC concentrations were
266 closely correlated with the measured data ($R^2 = 0.73$, $ME=0.69$, $RMSE = 21.51 \text{ g C kg}^{-1}$; $P <$
267 0.001).

268 **3.2 Sensitivity analysis**

269 A series of sensitivity tests were conducted to investigate the responses of the DNDC to
270 variation in climate factors (air temperature, precipitation) and grazing intensity. DNDC was
271 run with a 55-year baseline scenario that was based on the actual climate, soil and grazing
272 conditions of year 2005 in the study area. The ranges of values for alternative scenarios were



273 ± 10 , ± 20 and $\pm 30\%$ for precipitation, ± 1 , ± 2 and ± 3 °C for air temperature and ± 20 , ± 40 , ± 60 ,
274 ± 80 and $\pm 100\%$ for grazing intensity, respectively.

275 In the sensitivity analysis simulation, increases in precipitation resulted in elevated biomass and
276 SOC, however, the SOC was changed slightly compared to the biomass (Fig. 4A, B);
277 Temperature decrease induced the biomass decrease, and temperature increase could increase
278 the biomass. However, biomass change did not follow a simple linear relationship with change
279 in temperature. The 1 °C temperature increase could bring 24% of biomass increase, meanwhile,
280 1 °C temperature decrease could decrease 13% biomass (Fig. 4A) . Biomass was not susceptible
281 to the changes in precipitation. The biomass increased 7% and decreased 6% with precipitation
282 increased and decreased 30%, respectively. SOC had the reverse trend with increased or
283 decreased temperature, but there was a more complex relationship with temperature change.
284 The SOC had less sensitivity to temperature change compared to biomass. With a 1 °C
285 temperature increase, the SOC increased slightly with 0.26%, but when temperature increased
286 over 2 °C, the SOC decreased 0.26–0.83% (Fig. 4B). The modeled biomass was sensitive to
287 grazing intensity and biomass had a reverse trend with increased or decreased grazing intensity
288 (Fig. 4A). When grazing intensity changed from -100 to 100%, SOC increased rate from -0.22
289 to 0.40% (Fig. 4B). The sensitivity analysis demonstrated that the DNDC model was sensitive
290 to precipitation and temperature change and was useful for studying the biomass and SOC under
291 the grazing intensity change.

292 **3.3 Impact of grazing on biomass and SOC**

293 The biomass and SOC were significantly affected by climate change and the grazing intensity.
294 However, there were no significant interaction effects between climate and grazing intensity on
295 biomass and SOC during 1985–2044 throughout the study area (Table 2).

296 Under the same climate scenario, the grazing intensity change could significantly influence the
297 biomass, which had a negative relationship with the grazing intensity. The biomass differed
298 significantly under the four grazing intensities in the three climate scenarios. Among the
299 grazing intensity treatments, the biomass followed the order of: $G_0 > G_{-50} > \text{baseline} > G_{+50}$
300 (Table 3). Compared with the treatment without grazing, the grazing scenarios induced similar



301 changes in the biomass among the different grazing intensity treatments.
302 Grazing could increase the SOC storage. The SOC levels under various grazing intensities
303 followed the order of: $G_0 < G_{-50} < \text{baseline} < G_{+50}$ (Table 4). G_0 had the lowest SOC whereas
304 G_{+50} had the highest SOC under all the climate scenarios. Under the same climate scenario, a
305 reduction in the grazing intensity from the baseline could significantly decrease the SOC
306 concentration, but there was no significant change in the SOC when the grazing intensity
307 increased by 50% compared with the baseline.

308 **3.4 Impact of climate change on biomass and SOC**

309 The biomass exhibited a significant decreasing trend in the future climate scenarios compared
310 with the past 30 years under all the grazing intensities (Fig. 5), although precipitation increased
311 under both RCP4.5 and RCP8.5 (Table 1). Moreover, with the same grazing intensity, the
312 biomass was lower in RCP8.5 compared with RCP4.5. However, the biomass did not differ
313 significantly between RCP4.5 and RCP8.5 under the same grazing intensity (Table 3). This
314 suggests that RCP8.5 had a more negative effect on the biomass compared with RCP4.5 (Fig.
315 5).

316 The future climate could significantly decrease the SOC, and RCP8.5 had a more negative effect
317 than the RCP4.5 on the SOC. SOC exhibited a continuously decreasing trend according to the
318 RCP4.5 and RCP8.5 projections in the research area, where the changes in the SOC were similar
319 under the different grazing treatments (Fig. 6). A similar trend also occurred between 1985–
320 2014. The SOC was lower under RCP8.5 compared with that under RCP4.5 at all of the grazing
321 intensities. However, there were no significant differences between RCP4.5 and RCP8.5 (Table
322 4).

323 **3.5 The relationship between SOC and biomass change with** 324 **grazing and climate factors**

325 A multiple linear regression analysis indicated precipitation, air temperature and combined with
326 grazing intensity, can explain 33.2% of changes in biomass under the realistic climate scenarios



327 with a linear model. Meanwhile, precipitation, air temperature, and grazing intensity can
328 explain 52.3% of SOC variation (Table 5). Taking into account the prediction sum of squares
329 (PRESS) value, air temperature is the best predictor factor for biomass and SOC. It's suggested
330 that precipitation and grazing intensity with lower contributes to biomass and SOC change in
331 study region during past thirty years compared to temperature.

332 **4 Discussion**

333 **4.1 Effects of climate change on biomass and SOC**

334 Climate change is the main driver of the inter-annual fluctuations in the grassland biomass, as
335 observed in previous studies by Fan et al. (2010) and Gao et al. (2016). The unique climate
336 conditions such as precipitation and temperature on the QTP have a significant impact on the
337 grassland biomass (Fan et al., 2010; Yan et al., 2015). According to this study, the biomass of
338 alpine grassland could increase significantly in the short term as the temperature increases (Fig.
339 4), as also suggested by Chen et al. (2013) and Gao et al. (2016). However, under long-term
340 constant warming and without considering other meteorological factors, the alpine grassland
341 biomass will probably decrease (Zhu et al., 2016). This may be due to the higher temperature
342 increasing evaporation in the study area, thereby overcoming the benefits of increased
343 precipitation (Xu et al., 2009). The shortage of water will ultimately limit the increase in the
344 grassland biomass with significant warming and drying.

345 The decline of the SOC in our study indicates that climate warming will have more negative
346 effects and eliminated the positive effect of precipitation increasing in the study area. Riedo et
347 al. (2000) indicated that carbon storage may be lost from grazed grassland as the temperature
348 and precipitation increase. Tan et al. (2010) suggested that after a 2°C increase in temperature
349 in the QTP, the grassland ecosystem's net primary productivity will increase by 9%, but the
350 SOC will decrease by 10%. Temperature and precipitation are the main factors that affect the
351 SOC pools (Jobbagy and Jackson, 2000). Many studies have shown that sustained warming will
352 lead to increases in the SOC decomposition rate (Tan et al., 2010; Xu et al., 2012), especially
353 in the QTP region with high carbon storage at a low temperature in the high latitudes. Thus, the



354 SOC could be released by climate warming and become a more obvious carbon source
355 (Kirschbaum, 1995; Kvenvolden, 1993; Qin et al., 2014; Wang et al., 2008; Yang et al., 2008).
356 However, the effects of warming and precipitation on SOC storage remain a relatively complex
357 problem (Cao and Woodward, 1998; Schuur, 2003).

358 **4.2 Effects of grazing intensity on biomass and SOC**

359 The grazing intensity is most importance for the outcomes of grazing and it is the main external
360 factor that controls the influence of the grassland vegetation dynamics, as reported in the
361 previous studies (Guevara et al., 1996; McIntire and Hik, 2005; Pei et al., 2008; Veen et al.,
362 2012; Zeng et al., 2015). Indeed, an increase in the grazing intensity implies that more plants
363 would be removed by animals, which could eventually lead to a decline in the aboveground
364 biomass of the grassland (Yan et al., 2013).

365 Small differences in the SOC concentrations were observed after the grazing intensity increased.
366 However, there was a positive correlation between the grazing intensity with SOC. There is a
367 lack of consistent conclusions regarding the impact of grazing on the SOC concentration
368 according to previous studies. Thus, some studies showed that the grazing intensity and SOC
369 had a negative correlation (Bagchi and Ritchie, 2010; Derner et al., 1997; Wu et al., 2009) or
370 no relationship (Holt, 1997; Milchunas and Lauenroth, 1993). By contrast, many other studies
371 showed that grazing can increase the SOC (Li et al., 2011; Schuman et al., 1999; Wienhold et
372 al., 2001). This is partly because moderate grazing can increase the grassland below-ground
373 biomass, which is beneficial for the accumulation of SOC (López-Mársico et al., 2015). Some
374 studies have shown that increasing the plant root/shoot ratio and allocating more carbon to the
375 root system could induce different increases in the SOC (Derner et al., 1997). Nevertheless, the
376 main reason for the increase in the SOC in our study was the increasing number of grazing
377 animals, and thus the increased amount of manure returned after grazing on grassland (Hu et
378 al., 2015). Furthermore, the fertilizing effects of livestock excrement can increase the SOC
379 (Conant et al., 2001), especially in alpine grassland where the low temperature leads to the
380 relatively slow decomposition of litter (Davidson and Janssens, 2006). Moreover, increases in
381 the effects of hoof activity can accelerate the decomposition of litter and decaying roots, and



382 improve the contact with the soil, thereby accelerating the transfer of carbon to the soil to
383 improve the SOC concentration (Luo et al., 2010; Naeth et al., 1991).

384 **4.3 Patterns of regional change in the biomass and SOC**

385 From a spatiotemporal distribution perspective, the distribution of grassland biomass in
386 Qinghai Province is rather distinct due to the different constraints imposed by water and the
387 cumulative temperature. The biomass increased in the central and southwest of the research
388 region but decreased in the eastern and northern regions under RCP4.5 and RCP8.5,
389 respectively. However, the grassland biomass tended to decrease in more regions rather than
390 exhibiting an increasing trend (Fig. 7A). In particular, the vegetation activities are mainly
391 controlled by temperature in the eastern region, which may lead to greater negative effects than
392 the positive effects of increased precipitation (Zhou et al., 2007); therefore, the average regional
393 biomass may exhibit a significant decreasing trend.

394 In general, the SOC decreased from the low-temperature region to the high-temperature region,
395 where it followed the temperature distribution pattern in Qinghai Province and decreased from
396 the south to the north (Fig. 7B). The cold weather conditions would limit decomposition process
397 and there would be greater carbon storage over the years with accumulation in this area.
398 Furthermore, on the regional scale, although the SOC exhibited a decreasing trend in the whole
399 study area, the rate of change differed with a significant spatial distribution pattern.

400 **4.4 Uncertainty analysis**

401 Models are ideal tools for assessing the details of environment processes under various grazing
402 intensity. Furthermore, they can provide projections regarding the variations in grassland
403 biomass and SOC under alternative climate change scenarios. However, the uncertainty of the
404 data sources should be incorporated into the model outputs. The CMIP5 RCP scenarios were
405 used to provide the possible changes in climate in this study, but as a long-term climate
406 projection, the uncertainty of the projected climate will increase the time span (Moss et al.,
407 2010). The precipitation seasonal distribution pattern is critical to grassland growth (Shen et al.,
408 2011). In the present study, the precipitation distribution pattern of RCP scenarios was derived



409 from the year of 2014; this assumption may incorporate uncertainty for long term study.
410 In the present study, we assumed that the grassland type was the same in the scenarios, but the
411 grassland community structure could be modified by climate change. Therefore, the results of
412 this study only reflect the impacts of future climate change on the biomass and SOC storage.
413 The root/shoot ratio is one of the factors that are most sensitive to the simulated results and it
414 could be a potential source of uncertainty in biomass simulations. In this study, we used the
415 root/shoot ratio based on the grassland classes, which could generalize the spatial root/shoot
416 properties in the study area.
417 The grassland community structure could also be altered under both grazing and climate change
418 (Koerner and Collins, 2014), so the model assumed that the grassland community structure
419 remained the same throughout the whole simulation process, and thus it could incorporate
420 uncertainty in the simulation results. Due to a lack of data regarding the response of grassland
421 soil to animal trampling in the DNDC model, we ignored the trampling effect of the animals on
422 the soil structure, which may have led to some errors in the results.
423 The grazing rate can be another potential source of uncertainty. In most of the natural grassland
424 regions of the QTP, transhumance is usually practiced, which requires the transfer of livestock
425 from one pasture to another during different seasons, and staying in the same pasture for the
426 whole season. However, this grassland management practice was simplified in the present study
427 because we could not find specific statistical data to address this issue. Thus, we assumed that
428 livestock stayed in the same pasture for the whole year with 24 h d^{-1} of grazing and the stocking
429 rates were the same throughout the simulation unit and without yak dung remove (Zhang et al.,
430 2016). Furthermore, we assumed that all grasslands were useable. These assumptions could
431 have induced slight uncertainties in the simulation results.

432 **5 Conclusions**

433 In this study, we used the DNDC model to study the grassland biomass and SOC dynamics
434 under different climate change and grazing management scenarios. We found that the biomass
435 and SOC were significantly affected by climate change and grazing intensity. In the long term,
436 the total grassland biomass had a negative relationship and the SOC had a positive relationship



437 with the grazing intensity. The total biomass exhibited interannual fluctuations in the time series
438 and the SOC had a declining trend. All of the grazing scenarios obtained similar patterns of
439 change compared with the baseline scenario.

440 Future climate change could induce great uncertainty in the grassland dynamics. The total
441 grassland biomass and average SOC in the study area were reduced significantly under both the
442 RCP4.5 and RCP8.5 future climate change scenarios. However, there were significant
443 differences in the spatial distribution of the changing trends in the biomass and SOC. In the
444 eastern and northern regions of the study area, the biomass decreased, whereas it exhibited an
445 increasing trend in the southwest part of the research area. On a regional scale, the change in
446 the SOC had a significant spatial distribution pattern where it decreased from the south to the
447 north.

448 The grassland biomass and SOC will decline under sustained warming according to future
449 climate change projections. Therefore, grassland management should be adapted to potential
450 climate change to ensure sustainable grassland development in the study area. In the future,
451 suitable grazing intensity for the sustainable development of grasslands should be studied.
452 Moreover, greater human activity and management practices should be coupled according to
453 the model to develop more intelligent grassland management strategies.

454

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Table 1. Projected climatic changes (precipitation and maximum, minimum, and mean air temperature) under the RCP4.5 and RCP8.5 scenarios in 2044 compared with the corresponding values in the baseline data (2014).

Scenarios	Air temperature (°C)			Precipitation(mm)
	T _{max}	T _{min}	T _{mean}	
Baseline	3.63	-16.88	-3.56	279.24
RCP4.5	+0.99	+0.44	+0.72	+11.81
RCP8.5	+1.09	+0.51	+0.80	+12.50

Table 2. Summary of two-way analysis of variance for biomass and SOC relative to the climate, grazing intensity, and their interactions during 1985–2044. Degrees of freedom (d.f.), mean squares (M.S.), variance ratio (*F*-value), and level of significance (*P*-value) are shown.

Source of variation	d.f.	Biomass			SOC		
		M.S.	<i>F</i> -value	<i>P</i> -value	M.S.	<i>F</i> -value	<i>P</i> -value
Climate	2	16827.91	54.27	**	468.16	723.54	**
Grazing Intensity	3	22132.64	71.37	**	17.29	26.72	**
Climate*Grazing Intensity	6	2.63	0.01	n.s.	0.28	0.28	n.s.

**Significant effect; n.s., no significant effect.

Table 3. Simulated total biomass (mean ± SE) under the realistic, RCP4.5 and RCP8.5 scenarios for the G₀, G₋₅₀, baseline, and G₊₅₀ treatments.

Management practice	Total biomass (g C m ⁻²)		
	Realistic (1985–2014)	RCP4.5 (2015–2044)	RCP8.5 (2015–2044)
G ₀	223.6±10.8 ^{aA}	207.0±10.3 ^{aB}	199.5±10.2 ^{aB}
G ₋₅₀	213.3±10.5 ^{bA}	197.3±10.1 ^{bB}	189.7±10.1 ^{bB}
Baseline	202.1±10.4 ^{cA}	186.5±10.0 ^{cB}	178.9±10.0 ^{cB}
G ₊₅₀	190.2±10.4 ^{dA}	173.9±10.0 ^{dB}	166.3±10.0 ^{dB}

SE: Standard error. Significant differences among management practices are indicated by letters. Values within a column followed by the same lowercase letters or within a row followed by the same uppercase letter are not different at $P < 0.05$.

Table 4. Simulated SOC (mean ± SE) under the realistic, RCP4.5, and RCP8.5 scenarios for the G₀, G₋₅₀, baseline and G₊₅₀ treatments.

Management practice	Soil organic carbon (0–20 cm) concentrations (g C kg ⁻¹)		
	Realistic (1985–2014)	RCP4.5 (2015–2044)	RCP8.5 (2015–2044)
G ₀	65.84±2.30 ^{aA}	62.94±2.20 ^{aB}	62.87±2.20 ^{aB}
G ₋₅₀	66.04±2.30 ^{aA}	63.27±2.21 ^{aB}	63.20±2.20 ^{bB}
Baseline	66.32±2.31 ^{bA}	63.61±2.21 ^{bB}	63.53±2.21 ^{bB}
G ₊₅₀	66.59±2.31 ^{bA}	63.96±2.22 ^{bB}	63.88±2.21 ^{bB}

SE: Standard error. Significant differences among management practices are indicated by letters. Values within a column followed by the same lowercase letters or within a row followed by the same uppercase letter are not different at $P < 0.05$.



Table 5. Multiple linear regression analysis of grassland biomass and SOC change with relative factors.

	Variables numbers	R-square	PRESS	Temperature	Precipitation	Grazing Intensity
Biomass	1	26.4	273067.7	X		
	1	6.4	370402.4			X
	1	0.4	349337.6		X	
	2	26.4	287817.3	X	X	
	2	26.4	301908.4	X		X
	2	8.6	383224.5		X	X
	3	26.4	326183.5	X	X	X
SOC	1	47.6	179.2		X	
	1	2.3	310.9			X
	1	0.4	322.9	X		
	2	47.9	185.5		X	X
	2	47.7	189.5	X	X	
	2	4.7	328.8	X		X
	3	48.6	199.1	X	X	X

PRESS: The prediction sum of squares. The smaller the PRESS value, the better the model's predictive ability.

X: Indicates variable applied in the regression.

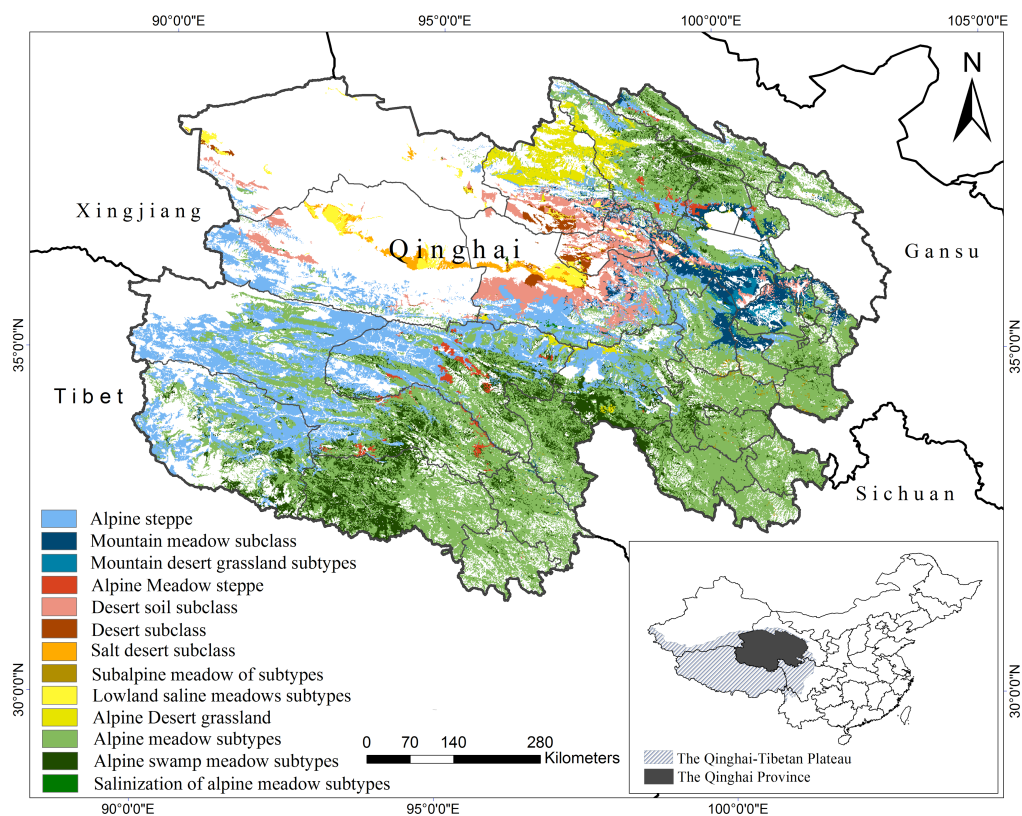


Fig. 1. Location of the study area and spatial distribution of the main grassland types. White areas are not covered by grassland.

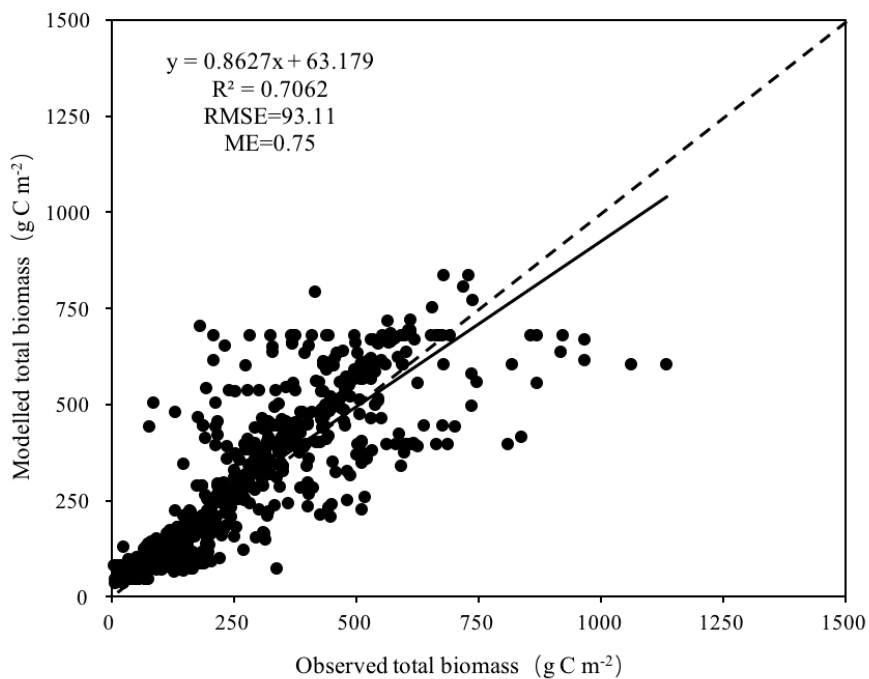


Fig. 2. Comparison of the modeled and observed total biomass values.

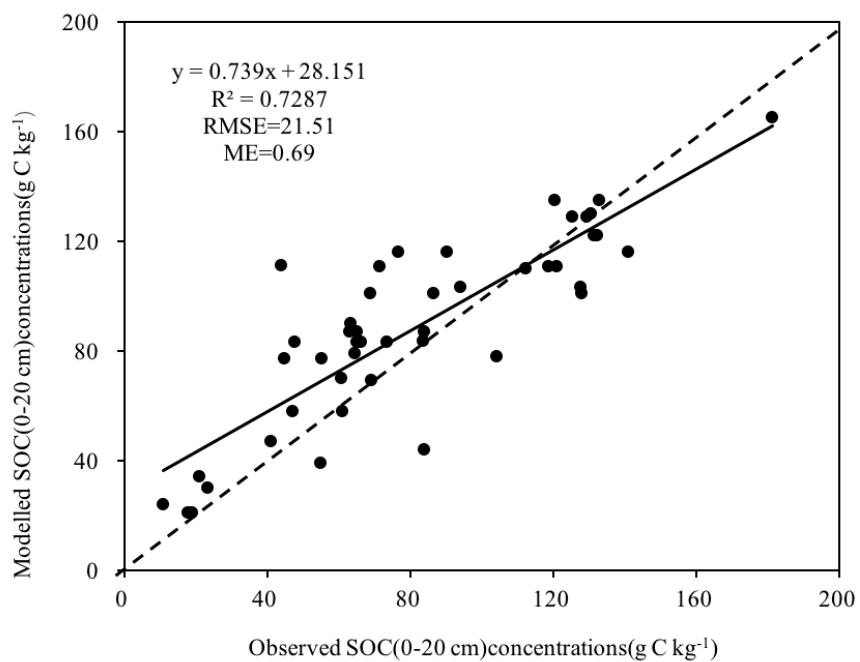


Fig. 3. Comparison of the modeled and observed SOC concentrations (0–20 cm).

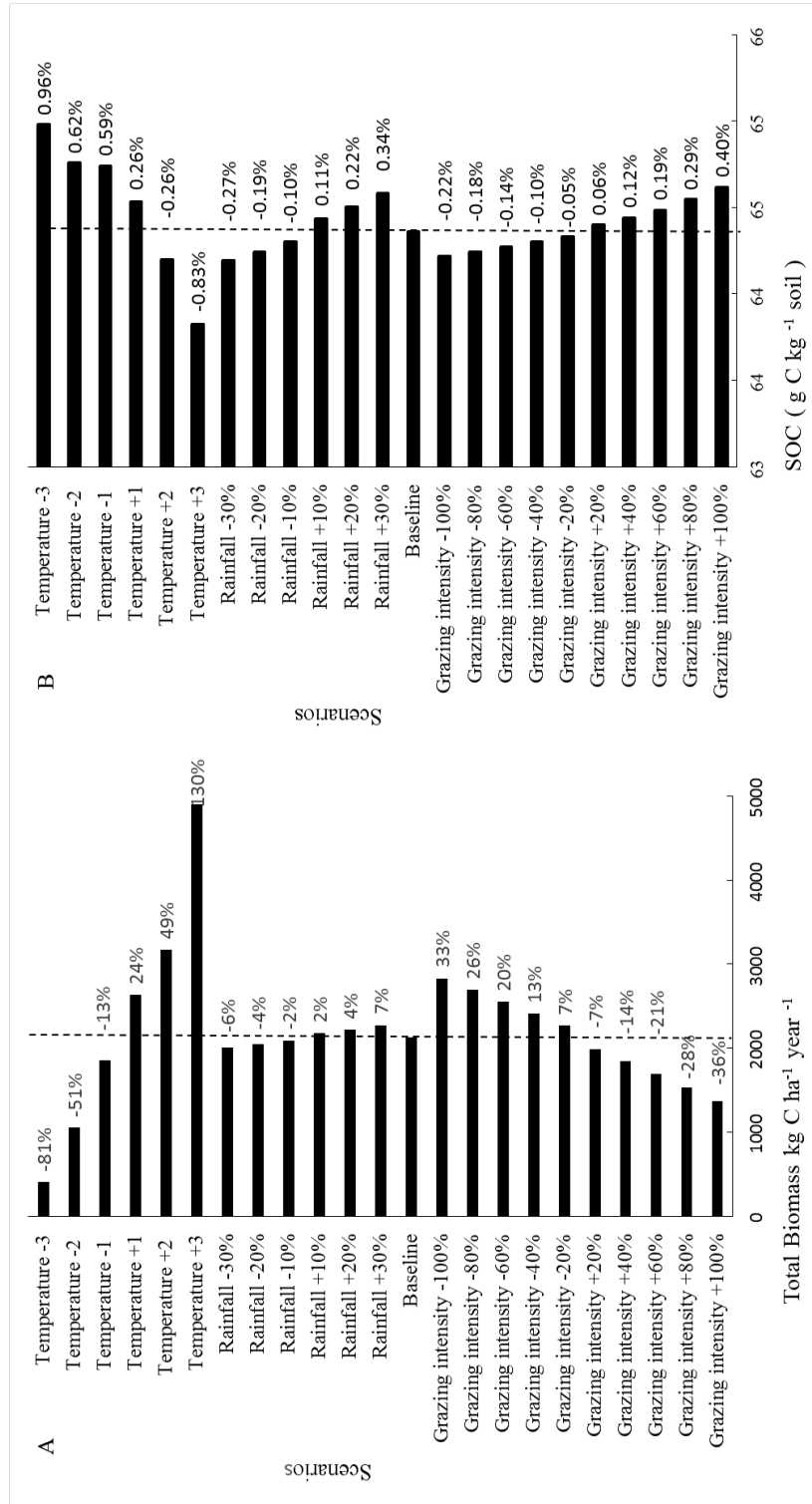


Fig. 4. Sensitivity analysis of model response to climate and grazing intensity change. The baseline biomass and SOC were the average value of a 55-year (1961-2014) simulation based on the actual climate and grazing conditions in the study area.

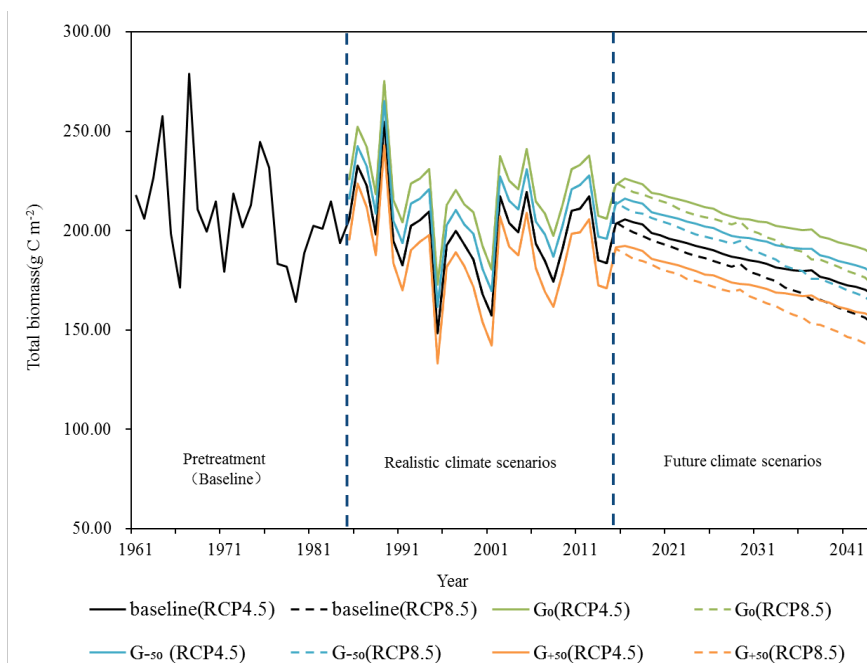


Fig. 5. Variations in the area-weighted mean biomass value under different scenarios. The stage on the left represents the preprocessing period from 1961 to 1984. The stage in the middle represents the realistic climate scenarios. The stage on the right represents future climate scenarios.

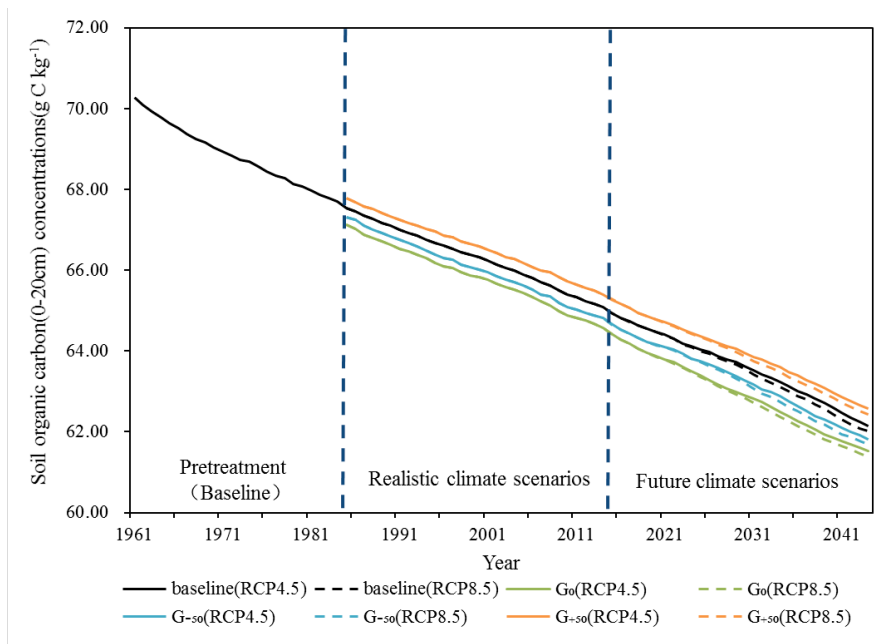


Fig. 6. Variations in the area-weighted mean SOC value under different scenarios. The stage on the left represents the preprocessing period from 1961 to 1984. The stage in the middle represents the realistic climate scenarios. The stage on the right represents future climate scenarios.

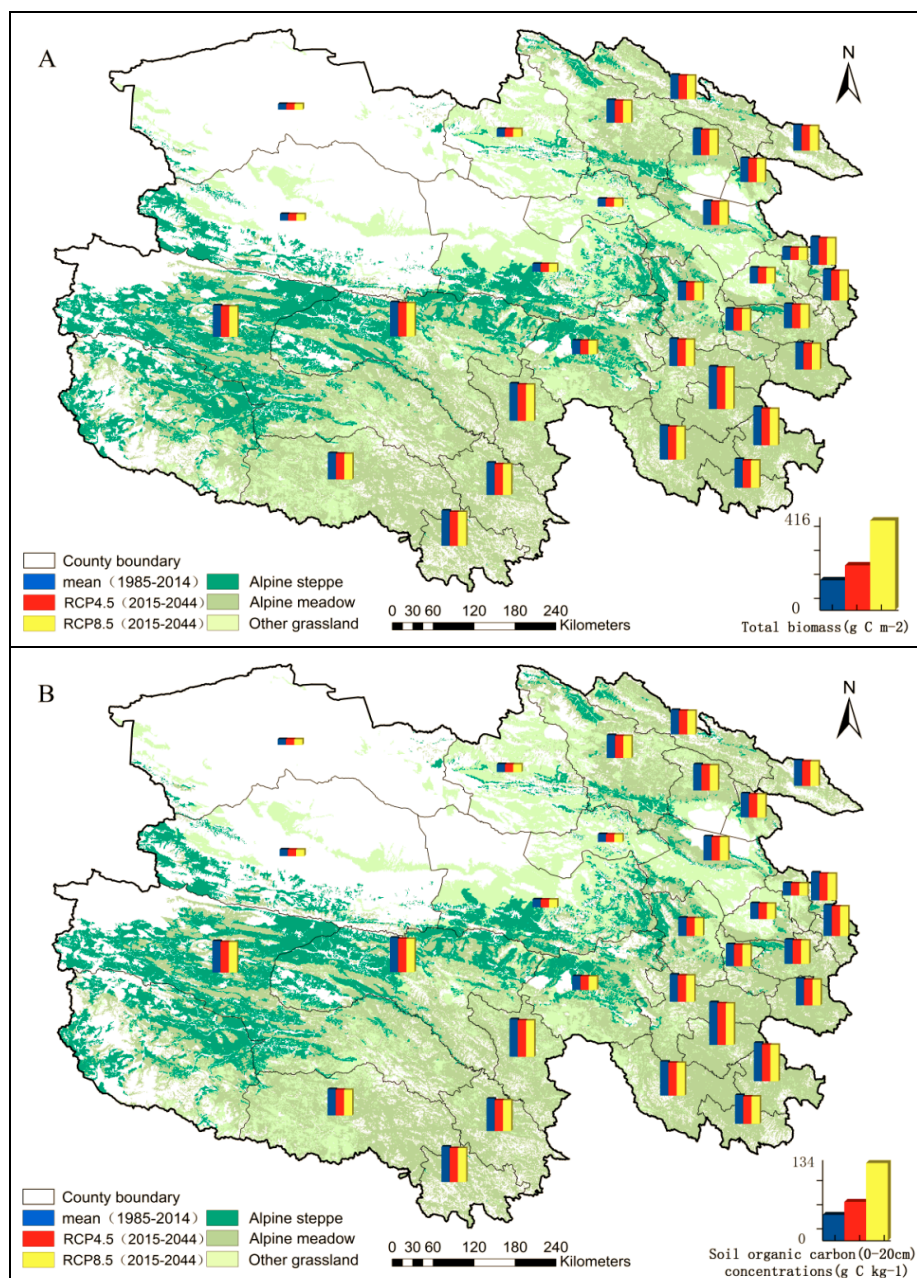


Fig. 7. Responses of the grassland biomass(A) and SOC(B) to climate change at a regional scale.