

AUTHOR RESPONSES TO REVIEWER COMMENTS

Title: Technical Note: Comparison of methane ebullition modelling approaches used in terrestrial wetland models

Author(s): Olli Peltola et al.

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We thank the reviewers for their constructive criticism and thorough comments which helped improve the manuscript. The main criticism from both reviewers was directed towards the comparison of modelling results against eddy covariance (EC) data. They argued that EC fluxes are not a good measure of the bubble flux and hence are not ideal for this kind of study. We partly agree with this notion and will modify the text accordingly. Please see more details below.

The reviewer comments are addressed individually below. The comments are shown in bold and responses with regular font.

On behalf of the authors,
Olli Peltola

REFEREE #1: Prof. Nigel Roulet

Overall impression

This is an interesting paper. The problem of methane transport has vexed empiricists and modellers alike since the exchange of methane from wetlands and the atmosphere was identified as a serious problem. This manuscript presents three models of methane ebullition from peatlands and then assesses how they perform against annual totals and a time series of half-hourly methane fluxes. They show that all the models give approximately the same annual flux but the proportion for the emissions transported via bubbles varies among the models. They conclude, based on comparison with half-hour eddy covariance data that the model they developed, the free-phase gas volume model (EBG), gives more realistic results than the other two models. I am not convinced by the evidence presented that the EBG more is better because I do not see how half-hour average EC fluxes are a good measure of the bubble flux. My concern is the boundary layer mixing blurs the bubble signal and that the post-processing removes concentrations from the high frequency data (e.g. 10 Hz) that indicate bubbles. I have wondered whether a comparison of the co-spectra of the momentum and concentrations could be used to identify the frequency of bubbles? Bubbles are very hard to measure well. In flooded wetlands funnel traps give a time integrated measure to the magnitude of the bubble flux but they do not work in peatlands. When auto-chambers are used one can see bubbles but examining the trace of concentration over time. At our site, Mer Bleue, we saw evidence of bubbles in less than 1% of our fluxes but in more fen systems such as Sallies fen (Goodrich in the references) bubbles appear in most chamber closers. Ideally the authors would have higher frequency records or used chambers where they can actually see the bubbles but it has to be an automated systems. Probably as good a test of the models is measures of the changes in methane storage. The authors point out this is not easy to do with interfering with the concentrations but diffusion samplers such as peppers or continuous flow samples could

give evidence for the changes in storage – this is the state variable in all the models. I think there is value in this manuscript (see below) but I think the authors should outline the ways they think the models could be tested. This would help others recognize the data they are sitting on could be used as evidence to attempt to refute the models. It is only through this testing we will gain confidence in the models and differentiate which model is more appropriate. I do think the manuscripts serves a useful purpose in setting out the three approaches to modelling ebullition. It is the clearest explanation of these kinds of models I have seen and on that basis alone I think it serves a very useful addition to the literature. On theoretical arguments I do think their EBG model stands up better than the other two models but the authors should be inviting or stimulating the community to test the models. However we are still faced a significant problem with estimating methane production: R_{prod} in equation (1). It does not really matter how elegant our models of the transport mechanisms are if production is poorly estimated.

RESPONSE: We want to thank Prof. Roulet for his comments. We do agree with him that the boundary layer mixing does blur the ebullition signal in the fluxes derived using the eddy covariance (EC) technique, yet it still does capture the methane (CH_4) flux caused by the released bubbles accurately. We argue that the ebullition signal can be seen in the 10-Hz CH_4 concentration time series as events with elevated CH_4 concentrations. These events are related to air parcels carrying the CH_4 released by bubbles and transported from the surface upwards and eventually past the EC measurement devices which then record these CH_4 emission bursts. We carefully post-processed the data in order to retain these events throughout the EC data processing chain. In fact, we are currently investigating whether these events could be separated from the turbulent time series using e.g. wavelet transforms to derive ecosystem scale bubble flux (topic of another study).

Another way to frame the problem raised by the reviewer is that the EC system estimates the fluxes at ecosystem scale whereas the model estimates the CH_4 fluxes from a single peat column. Hence there is an obvious mismatch in horizontal scale. EC source area contains locations where ebullition is taking place and locations where ebullition is at the same time absent. The conventionally estimated 30-min averaged EC fluxes combine all the different emission pathways (diffusion, plant transport and ebullition) at ecosystem scale and hence the modelled and measured fluxes are not fully comparable. Therefore, we will modify the manuscript so that we will mention this scale mismatch in all relevant locations (abstract, Sect. 3.3 & 4). We will emphasize more the comparison between the ebullition modelling methods instead of the validation against EC measurements. As suggested by the reviewer, we will also add a small section to Discussion where we will mention the more appropriate datasets against which this kind of models should be tested (auto-chambers, changes in CH_4 stored in the peat).

Minor points

Pg 1 Ln 25 Not sure of this number. Ebullition is important where it is important but when plants are present that root below the water table plant mediated is important. Diffusion can't be that large. Also think there is steady bubble flux. Regardless it does not matter because it is important.

RESPONSE: Agreed. We will remove the percentages and simply just mention that in some studies ebullition has been observed to contribute a significant fraction to the overall CH_4 flux to the atmosphere.

Pg 1 Ln 30 This is an understatement which is why I question the 0 to 70% claim above.

RESPONSE: We will emphasize the lack of knowledge in this matter more.

Pg 5 Ln 21 Subscript 4 on CH_4

RESPONSE: Thanks, will be corrected.

Pg 10 Ln 11-13 This is fine but it my experience that most peatlands do not have significant confining layers. Beaver ponds are a good example - they a constantly bubbling and the bubbles are shallow.

RESPONSE: This is a good point and many recent studies have shown that the bubbles are mostly formed close to the surface, as already mentioned later on in the Discussion.

Pg 10 Ln 35 Patrick Crill is measuring concentrations in peat with circulating diffusion samplers that equilibrate with soil concentrations. They would not sample bubbles but they would tell you the depth and duration that the concentrations approach and exceed saturation. Also peepers would tell you the same thing.

RESPONSE: We are using similar diffusion samplers at our lake site to estimate CO₂ concentration profiles in the lake. The plan is to set up similar system at our peatland site to measure CH₄ concentrations in the peat. However, such measurements average the soil pore water concentrations over a relatively large area, since long sampling tubes are needed before the air in the tubes is in equilibrium with the surrounding pore water concentrations (diffusion through the tube walls is slow). Hence, most likely the small scale variability in the pore water concentration cannot be captured with such systems. This variability might be crucial for creating favorable conditions for bubble formation at depths where on average low pore water concentrations inhibit the bubble formation. These issues are already mentioned in the Discussion section of the manuscript.

Pg 11 Ln 4 Yes but the easiest thing to measure is the concentration profiles in some non-destructive manner that does not require putting a negative pressure to extract the samples. My guess is diffusion samples circulating through a CH₄ analyzer would be the best route to get at this problem. At least then you are measuring the temporal variability of the state variable. If the diffuser were of a sufficient length, say several meters, they could obtain spatial averages. What do you recommend as a test?

RESPONSE: We are not sure what the referee is meaning with this comment, since the text on page 11, line 4 does not discuss the concentration profiles. Maybe due to an oversight this comment is refers to wrong location in the manuscript. In any case, we agree with the referee that the diffusion samplers are a convenient way to measure concentrations in the water, however they cannot capture the small scale variability in the pore water concentrations which could be vital for bubble formation. Moreover, we agree that the pore water concentration is the state variable against which in general this kind of models should be ideally tested.

Pg 11 Ln 14-26 I am not sure how one can use EC to determine a bubble flux. What EBG shows is that it matches the temporal pattern of the EC fluxes but it can't see bubbles. I have always wondered in the high frequency data the concentration and momentum spectra should see departures that would indicate bubbles being mixed in the boundary layer. Automated chambers see bubble events - the time trace of concentrations show step changes. You refer to Goodrich et al. And they saw thus at Sallies fen in NH.

RESPONSE: As mentioned above, we argue that the bubble flux can be seen in the 10-Hz CH₄ concentration time series as events with significantly elevated CH₄ concentrations. These events contribute to the covariance between the vertical wind speed and the gas concentrations as any other fluctuation in the time series and hence EC does capture the ebullition signal. However, it is just mixed with all the other emission pathways taking place within the EC ecosystem-scale source area whereas the model works on a single peat column scale. As mentioned before, we will emphasize this point more in the revised version of the manuscript.

Spectral analysis is not necessarily the best tool for detecting these events as the spectra are calculated from the complete time series which combines all the emission signals. Instead, wavelets and wavelet transforms allow separating individual events from the high frequency time series and hence possibly enable the estimation of ecosystem scale bubble flux. We are currently investigating whether the ebullition signal could be disentangled from the 10-Hz time series using wavelets, but that is a topic for another study.

REFeree # 2: Dr. Jorge Ramirez

Comments

This paper does address a subject that is relevant for BG and provides a good overview of existing methane ebullition models. The explanation of the models is clear and background information from the literature is provided. The aim of the paper is straightforward and the study intends to test ebullition models using observed data. Furthermore, the models are used to produce outputs that cannot be directly verified using observations. This is fine, and is one of the major points of modelling systems. What is unclear, is how the performance of the models are gauged against the observed data.

The main topic of this study is methane emissions via ebullition and the authors do not clearly describe how ebullition events are measured using Eddy Covariance (EC) data. In the end, the authors compare model output from all methane transport mechanisms combined (ebullition, plant mediated, and diffusion) against the EC data. Unfortunately, this comparison does not allow the authors to derive any definitive conclusions about modelling methane ebullition. The paper can also benefit from more explanation on the collection and processing of the field data. On the modelling side, the authors should provide information on calibration of the models and provide reasons if calibration was not performed and how this could affect model results. A brief model sensitivity analysis is presented, and it would be interesting to expand this section with a more in depth, systematic sensitivity analysis that includes a figure. The majority of the paper is well written, but some sentences need restructuring (see comments below). To improve clarity the manuscript should be read by a native English speaker.

RESPONSE: We thank Dr. Ramirez for his comments. As already discussed above, we acknowledge the shortcomings of EC data and will revise the manuscript accordingly by mentioning the spatial scale mismatch in all relevant locations and shifting the focus of the manuscript more into the modelling approach comparison. We will also add more details about the measurement setup and data processing to the Sect. 2.2, however this is a modelling study and hence we would like to keep this section as short as possible.

The HIMMELI model was calibrated in Susiluoto et al. (2017) and this will be mentioned in the Sect. 2.1. However, we would like to emphasize that the main aim of this study is to compare different ebullition modelling approaches, not to model the CH₄ emissions from the peatland with high accuracy. We argue that for this kind of comparison model calibration is not crucial, albeit in general for CH₄ cycling modelling it should be done with care.

The parameters used in the three ebullition modelling approaches were not calibrated, instead typical values found in the literature were used. We opted for this approach since the model parameters are often correlated with each other (e.g. Susiluoto et al., 2017), making careful calibration time-consuming and sensitive process. Furthermore, full model calibration with different ebullition approaches would make it difficult to evaluate whether the differences in modelling results stem from ebullition modelling approach or from the model calibration. Hence, we believe that using parameter values found from the literature should provide robust results for the readers to gauge the differences between the ebullition modelling approaches. We will modify

the manuscript and try to make it clearer that parameter values from the literature are used, not calibrated ones.

As requested by the reviewer, we will expand the model sensitivity section of the study and include an appropriate figure detailing the most relevant results from this analysis.

Technical corrections

Pg 1 Line 25: Ebullition is not important in all wetlands. perhaps change sentence by adding, "in some wetlands"

RESPONSE: Thanks, we will modify this part.

Pg 1 Line 28: Ebullition is not only sporadic in space, but also in time and you should provide background evidence for this (see introduction of Ramirez et al. [2017]).

RESPONSE: In this sentence the word "events" is supposed to emphasize that ebullition is also sporadic in time. We will modify this sentence so that it is clearer that the temporal variability is also significant and add an appropriate reference.

Pg 1 Line 36: Total volume of what?

RESPONSE: Of the peat volume (see Rosenberry et al., 2006). We will modify this sentence so that this will be clear.

Pg 2 Line 8: Mention that both increasing and decreasing atmospheric pressure trigger ebullition and provide references.

RESPONSE: Increasing pressure is already mentioned on page 2 line 11 and a reference is also provided.

Pg2 Line 19: Avoid starting sentences with a number, instead spell out the number. This sentence also needs to be restructured.

RESPONSE: Yes, we agree that sentences should not begin with a number. We will fix this part.

Pg2 Line 20: Replace questioned with questionable?

RESPONSE: The sentence here means that this modelling approach is controversial given the current knowledge of the ebullition process. Hence we argue that "questioned" is correct.

Pg 2 Line 28: Replace inflict with produce.

RESPONSE: Yes, produce sounds better.

Pg3 Line 1: Is this upward transport, and/or also lateral transport? Be specific.

RESPONSE: Thanks, we will specify that it is upward transport.

Pg3: Does the model have a spatial resolution? How many layers exist in the peat column and how thick are the layers? Introducing this information early on helps the reader visualise how the model operates.

RESPONSE: The model is a 1D-column model and hence does not as such have a spatial resolution. We run the model with 2 m of peat with ten layers, therefore each layer was 0.2 m thick. We will add these pieces of information to the Sect. 2.1.

Pg4 Line 8: Can you explain better what the is meant by the lowest air layer. Here you mention the model time step, can you provide the actual time step early in the model introduction (e.g. hrs or days).

RESPONSE: The lowest air layer is the model layer which is the lowest model layer which is above water level and below peat level. The model output time step is one day, but the model uses a shorter internal time step in order to ensure numerical stability. We will mention these details in the model description section (Sect. 2.1.).

Pg 4 Line 20: Again, the lowest air layer is vague, perhaps define it earlier. Replace ascend with ascent.

RESPONSE: We will define the lowest air layer in Sect. 2.1. Thanks, we will replace ascend with ascent.

Pg 6 Line 9: You introduce a layer thickness 0.2 m, is this the case for all three models (see comments above regarding spatial resolution).

RESPONSE: Yes, the model settings were kept constant and only the ebullition modelling approach was varied in this study. We will mention the layer thickness in Sect. 2.1.

Pg7 Line 1: replace gapfilled with gap-filled

RESPONSE: Thanks, will be replaced.

Pg 7 Line 4: Can you provide evidence in the literature that ebullition events from peat can be measured using EC.

RESPONSE: At least Hargreaves et al. (2001) and Sachs et al. (2008) argue that their EC CH₄ flux measurements were significantly affected by ebullition. However, a study where ebullition is recorded with multiple independent measurement systems (e.g. chambers and EC) at the same is yet to be published.

Pg 7 Line 10: Can you provide further explanation as to how ebullition was detected using the EC data. Include details about the post-processing of the data.

RESPONSE: We argue that the ebullition events can be seen in the 10 Hz EC data as events with significantly elevated CH₄ concentrations. These fluctuations in the concentration time series contribute to the flux the same way as any other fluctuation. However, a problem arises from the scale mismatch between EC and the model results. See the discussion above. During data post-processing we made sure that these events were not flagged as erroneous measurements, e.g. by de-spiking. We will add details about data post-processing and mention the scale mismatch in all relevant locations.

Pg 7 Line 33: Avoid starting sentences with a number, instead spell out the number. Apply to other instances within the manuscript.

RESPONSE: Thanks, we will remove numbers from the beginning of sentences.

Figure 4: Dashed line is not clearly visible.

RESPONSE: We will plot the dashed lines with darker colors.

Pg 8 Line 36: Change "got stuck" and "alive" with formal words.

RESPONSE: We will reformulate the sentence as: "Bubble volumes released from deeper layers and which were re-attached during their ascent maintained the gas volume at -1.1 m depth even though the c-term on average decreased the volume (cf. Fig 4a)."

Pg 8 Line 38 and 39: "Inflicted" is not the correct word choice.

RESPONSE: We will replace "inflicted" with "generated".

Figure 5a: It is difficult to distinguish between the four datasets plotted. Consider restructuring this figure with a separate panel for each model output, with the observed data superimposed. Additional figures you may consider are histograms and scatterplots between observed and modelled emissions (with the later the differences are clearly noticeable).

RESPONSE: We decided to plot all the lines to the same panel, since we wanted to have also the different emissions pathways in the same figure (subplots b-d). However, based on this comment we will modify the figure so that it will have six panels: one for comparing against measurements and one for showing the different emissions pathways for each ebullition modelling approach.

We will not add the additional figures suggested by the reviewer (histograms or scatter plots), due to the scale mismatch between model results and EC derived CH₄ fluxes (see the discussion above). Moreover, in the revised manuscript we will emphasize more the comparison between modelling approaches and less the validation against EC measurements.

Pg 9 Line 4: Does figure 5a compare modelled and observed CH₄ emissions considering all three transport mechanisms (diffusion, plant mediated, and ebullition)? If this is the case, it is quite confusing because the manuscript up to now was focusing on ebullition, and I was expecting a comparison between observed and modelled ebullition emissions. Please provide further explanation.

RESPONSE: Yes, it includes emissions via all three emission pathways. There are at least three reasons for this:

- 1) In the HIMMELI model bubbles released to lowest air layer are not recorded as ebullition flux since diffusion is still needed to transport the released CH₄ to the atmosphere. Hence the diffusion flux contains the ebullition flux released to the lowest air layer which is usually below the peat surface. Therefore ebullition flux cannot be completely separated from the diffusion flux. This is also mentioned in the Fig. 5 caption.
- 2) EC measurements contain all the three emission pathways and there is no robust method to separate the emissions pathways from each other.
- 3) The ebullition modelling approach alters the model CH₄ cycling in a fundamental way (see e.g. the CH₄ concentration profiles in Fig. 2 in the manuscript) and hence also the other emission pathways are affected. For instance the magnitude of plant transport is quite different between the EPT and the EBG

approach (see Table 2 in the manuscript). Hence we argue that it would not be sufficient to compare only the ebullition flux against measurements.

Pg 9 Line 10 and 11: This variability is difficult to see in Figure 5a, see comment above.

RESPONSE: We will modify the figure, see above.

Pg 9 Line 14: Can you provide more explanation how the R² were derived for each model.

RESPONSE: Thanks for this comment, since it allowed us to find an oversight in our analysis. Previously the R² value was calculated as the squared Pearson correlation coefficient, however a more appropriate way to estimate the coefficient of determination (R²) in our case would be:

$$R^2 = 1 - \frac{\sum_i (y_i - f_i)^2}{\sum_i (y_i - \bar{y})^2}, \quad (1)$$

where y are the measurements and f the modelled CH₄ fluxes. The overbar denotes averaging and subscript i a running index over the whole time series. The equation above produces the following R² values for each model run: 0.63 for EBG, 0.56 for ECT and 0.35 for EPT. We will update these numbers to the manuscript.

Pg 9 Line 22: Replace the phrase “complicates the picture” with formal wording and explanation.

RESPONSE: We will reformulate this sentence as “although CH₄ oxidation in the peat column might be different between the modelling approaches”.

Pg 9 Section 3.4: This a model sensitivity section and should be presented before the results section.

RESPONSE: We would like to keep the sensitivity section here, since at this point the reader is already familiar with the different concepts used in this study (for instance ebullition event and its profile). Hence it is easier for the reader to understand the differences between the sensitivity runs (see Table 1) without an elaborate explanation of the different concepts.

We want to emphasize that the default parameter values for the EBG approach were derived from literature and not from this sensitivity analysis. This section is meant merely for future studies using the EBG approach to understand which parameters are the most crucial for this modelling approach.

Pg 10 Line 11-13: How do the models support the rupturing of confining layers if peat micro structure (e.g. woody layers) is not entirely represented within the models? Can you provide explanation for this conclusion.

RESPONSE: This conclusion is based on the depth of the ebullition events (mostly below 1 m, see Fig. 1) and on the temporal variability of the modelled ebullition flux, which follows from the threshold logic (concentration, pressure or volume) used in all approaches (ECT, EPT and EBG, respectively). The reviewer is correct that the model does not have as such a detailed description of peat microstructure and we will mention this at the end of the sentence which ends on line 13:

"Hence the results resemble the "deep peat" hypothesis put forward by Glaser et al. (2004) in which the free-phase gas is produced in deep (> 1 m) peat and trapped under semiconfined layers, even though peat micro structure (e.g. woody layers) was not described in the model."

Pg 11 Line 5: Sentence needs restructuring.

RESPONSE: The sentence will be replaced with sentence: "Clearly the effect sedges on ebullition should be studied more prior to these effects can be implemented in terrestrial CH₄ models".

Pg 11 Line 23: You mention calibration of methane models performed in other studies, can you explain why you chose not to perform a calibration prior to model testing and how this affects your results?

RESPONSE: As mentioned above, the HIMMELI model was already calibrated by Susiluoto et al. (2017) and this will be mentioned in Sect 2.1. In general, model calibration is a large task worth its own publication and was not included here. The ebullition modelling approaches were run using parameter values found from the literature, which has at least two advantages:

- 1) The differences between modelling results do not follow from differences in the model calibration. Model parameters are often correlated with each other (see e.g. Susiluoto et al., 2017) and hence if the model would have been calibrated separately for each ebullition modelling approach, then also other parameters not directly related to ebullition (e.g. plant transport related parameters) might change. Then it would be difficult to evaluate whether the differences between model runs would be related to differences between ebullition modelling approaches or to model calibration.
- 2) When using the literature values, the modelling results can be directly compared against other studies using the same parameter values. For instance ECT results are comparable with Wania et al. (2010) ebullition modelling, since similar approach with the same parameter values were used.

Pg 11 Discussion section: Assuming that you cannot definitively identify ebullition events in the EC data, it is possible that the EC data contains few ebullition events. Could this further explain the mismatch between observed and modelled CH₄ emissions? Please address this possibility in the discussion section.

RESPONSE: Yes, we agree that this is a possibility. We argue that this is related to the scale mismatch: the model estimates fluxes from a single peat column, whereas EC measures fluxes at ecosystem scale. The EC source area may contain locations where ebullition is taking place and locations where at the same it is absent. Therefore, the ebullition events are diluted in the conventionally calculated EC fluxes, yet their impact on ecosystem scale fluxes is captured accurately. As mentioned above, we will mention this in all relevant locations of the manuscript.

References

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LIST OF RELEVANT CHANGES IN THE MANUSCRIPT

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Author(s): Olli Peltola et al.

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We modified the manuscript based on the referee comments. A list of the most relevant changes made to the manuscript is given below. Small changes are not listed, but they can be seen in the modified manuscript below.

Additionally, the estimates of annual CH₄ emissions based on measurements in the Table 2 were updated.

On behalf of the authors,
Olli Peltola

1. The mismatch in horizontal scale between the measurements and the model results is now mentioned in all relevant locations: (abstract, Sect 2.2. (page 7 lines 16-18), Sect. 3.3 (p. 9 l. 36-37), Sect. 4 (p. 12 l. 18-24), Sect. 5 (p. 13 l. 14-15))
2. A more detailed model sensitivity analysis was done and included in the manuscript. Hence the text in Sect. 3.4. is now longer than before, a new figure was added (Fig. 6) and the new sensitivity runs were added to the Table 1.
3. Information about model calibration was added to Sect. 2.1 (p. 3 l. 24-31), in addition to details about model spatial and temporal resolution (p. 3 l. 31-33).
4. More details about the measurement setup and data processing and a reference were added to the Sect 2.2 (p. 3. l. 22-24, l. 28-30).
5. Based on the reviewer comment the more appropriate datasets against which this kind of models should be tested are now discussed in the Sect. 4 (p. 12 l. 24-28).
6. Figures 4 and 5 were modified based on referee comments

Technical Note: Comparison of methane ebullition modelling approaches used in terrestrial wetland models

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Abstract. Emission via bubbling, i.e. ebullition, is one of the main methane (CH₄) emission pathways from wetlands to the atmosphere. Direct measurement of gas bubble formation, growth and release in the peat-water matrix is challenging and in consequence these processes are relatively unknown and are coarsely represented in current wetland CH₄ emission models. In
10 this study we aimed to evaluate three ebullition modelling approaches and their effect on model performance. This was achieved by implementing the three approaches in one process based CH₄ emission model. All the approaches were based on some kind of threshold: either on CH₄ pore water concentration (ECT), pressure (EPT) or free-phase gas volume (EBG) threshold. The model was run using four years of data from a boreal sedge fen and the results were compared against eddy covariance measurements of CH₄ fluxes.

15 Modelled annual CH₄ emissions were largely unaffected by the different ebullition modelling approaches, however temporal variability of CH₄ emissions varied an order of magnitude between the approaches. Hence the ebullition modelling approach drives the temporal variability of modelled CH₄ emissions and therefore significantly impacts for instance high-frequency (daily scale) model comparison and calibration against measurements. The modelling approach based on the most recent knowledge of the ebullition process (volume threshold, EBG) agreed the best with the measured fluxes ($R^2=0.7463$) and hence
20 produced the most reasonable results-, although there was a scale mismatch between the measurements (ecosystem scale with heterogeneous ebullition locations) and model results (single horizontally homogeneous peat column). The approach should be favoured over the two other more widely used ebullition modelling approaches and researchers are encouraged to implement it into their CH₄ emission models.

1 Introduction

25 A large fraction of methane (CH₄) emitted from wetlands to the atmosphere is released in rapid bubbling events, during which part of the biogenic gas bubbles trapped below the surface are released and transported quickly to the atmosphere. This emission route, called ebullition, accounts for 0...70 % of has been observed in some wetlands to make a significant contribution to the total CH₄ emissions (e.g. Tokida et al., 2007; Yu et al., 2014; Christensen et al., 2003), whereas the rest are emitted more steadily via diffusion or plant-mediated transport (e.g. Le Mer and Roger, 2001). Rapid emission of CH₄ in
30 ebullition events allows the emitted CH₄ to bypass an oxic zone where the transported CH₄ might have been otherwise oxidized prior to reaching the atmosphere (Rosenberry et al., 2006). A characteristic feature of ebullition is that it takes place in sporadic events which are irregularly distributed in space- and time (e.g. Klapstein et al., 2014; Tokida et al., 2007). Thus their long-term measurement and quantification in the field has proven to be very challenging and hence a complete understanding of the mechanisms controlling ebullition is still lacking.

35 Based on the current knowledge an episodic ebullition event takes place when a gaseous-phase gas volume at a certain level below the surface reaches a critical threshold, after which the excess gas is released in an ebullition event. This follows from the fact that for big enough gas bubbles the buoyancy forces exceed the retention forces that have been keeping the entrapped bubbles in place and hence they start to ascend towards the surface (Fechner-Levy and Hemond, 1996). Generally, the critical volumetric threshold for triggering ebullition is considered to be around 10 % (at maximum 20 %) of the total peat volume

(Rosenberry et al., 2006). Based on the existing empirical evidence the released gas bubbles consist largely of CH₄ (20...80 % of the bubble gas) (Waddington et al., 2009;Kellner et al., 2006;Tokida et al., 2005;Glaser et al., 2004;Walter et al., 2008) and the rest is nitrogen (N₂), oxygen (O₂) and carbon dioxide (CO₂) (Tokida et al., 2005;Kellner et al., 2006;Walter et al., 2008) which suggest that the bubbles originate from conditions with relatively high pore water CH₄ concentrations. The gas bubbles may form under anaerobic conditions where high pore water concentrations facilitate the formation and build-up of gas-phase bubbles. Existence of bubble formation nuclei is not considered to limit bubble formation.

Considering the ideal gas law and Henry's law, the gas-phase bubble volume can be modified by 1) pressure, 2) temperature or 3) pore water concentration changes (Fechner-Levy and Hemond, 1996). The effect of pressure can be further divided into atmospheric and hydrostatic (i.e. the weight of the water column above) pressure. Hence, if the bubble volume threshold is also considered, an ebullition event may be triggered by decreasing atmospheric pressure, water table depth (WTD), increasing peat temperature or pore water CH₄ concentration. Out of these, decreasing atmospheric pressure has been most often reported to trigger ebullition (Tokida et al., 2007a;Tokida et al., 2005;Waddington et al., 2009;Yu et al., 2014;Strack et al., 2005;Kellner et al., 2006), whereas some studies have also reported the effect temperature (Waddington et al., 2009;Kellner et al., 2006;Yu et al., 2014). Also other forcings (e.g. increasing pressure, wind speed) have been linked with ebullition (Goodrich et al., 2011;Comas et al., 2011).

According to a recent review by Xu et al. (2016) most (24 out of 40) of the process-based models focusing on CH₄ cycling incorporate some kind of an approach to model ebullition. In a seminal modelling paper by Walter and Heimann (2000) a simple approach was adopted: If pore water CH₄ concentration exceeded a certain threshold value, then the excess CH₄ was directly released to the atmosphere. Thereafter, the approach has been adopted with slight modifications to several CH₄ models and can be regarded as the most widespread method to model ebullition. The most common alteration of the approach is to estimate the concentration threshold based on CH₄ solubility (e.g. Wania et al., 2010;Riley et al., 2011).

~~67 % (16 out of 24) of Pore water concentration threshold was the models that incorporated in the review by Xu et al. (2016) some kind of most widely used ebullition modelling approach used a pore water concentration threshold in the models reviewed by Xu et al. (2016). It was incorporated in 16 out 24 models (67 %) that included this transport route.~~ However, based on current knowledge this approach can be questioned, since it lacks almost all of the details outlined above about the ebullition process and hence could result in unrealistically modelled process. Other ebullition modelling approaches have been implemented as well. For instance, Grant (1998), Tang et al. (2010) and Raivonen et al. (2017) triggered ebullition if the total pressure of water-dissolved gases exceeded the ambient pressure at a given depth, whereas in some studies (Segers et al., 2001; Granberg et al., 2001; Zhang et al., 2012) ebullition was modelled using a threshold for gaseous volume at a certain depth below surface.

This study was motivated by the fact that in many process-based models ebullition is modelled in a manner which does not agree with the current knowledge of the process. Hence the models possibly produce erroneous ebullition fluxes and thus may for instance bias the modelled annual CH₄ emissions and ~~inhibit~~ produce a mismatch between modelled and measured CH₄ fluxes. The aim of this study is to compare three ebullition modelling approaches which are based on pore water concentration (Sect. 2.1), pressure (Sect. 2.2) and gaseous volume (Sect. 2.3) thresholds. This is achieved by implementing the three approaches in one process-based model called HIMMELI (Raivonen et al., 2017) and running the model with the same input data and only altering the ebullition modelling approach. We aim to characterise the differences in 1) amount, timing and depth below peat surface of the modelled ebullition events, 2) variables causing the events and 3) modelled CH₄ flux to the atmosphere. In addition, the performance of the volume threshold approach using different model parameters is evaluated. Hypothetically, the approach based on gas volume threshold should produce the most reasonable results since it resembles the current knowledge of ebullition the most. The main aim of this technical note is to report the differences between ebullition modelling approaches and to promote the usage of physically sound methods in the coming CH₄ modelling studies.

2 Materials and methods

In this study a model concentrated on CH₄ cycling is run using data from a boreal fen to study the differences between three ebullition modelling approaches. The model and measurement data are described below. The following nomenclature is used throughout this study: “Ebullition event” is used to define an episode during which concentration (ECT), pressure (EPT) or bubble volume (EBG) threshold was exceeded at a certain depth, “ebullition” is defined as the upward transport of CH₄ via the ebullition process to the lowest air filled pore space and “direct ebullition to the surface” is the ebullition flux directly to the atmosphere. Direct ebullition to the surface can take place only when the water table level is at the surface or above it.

2.1 Process-based model HIMMELI

In this study a process-based model HIMMELI provided a framework used to compare the three ebullition modelling approaches. The model is described elsewhere (Raivonen et al., 2017) and thus the description is not repeated in detail here. In short, the 1-D model estimates sources, sinks and interactions between three substances, namely CO₂, O₂ and CH₄ in a vertically layered peat-water-air column. HIMMELI incorporates the following reaction-diffusion equation to model the temporal evolution of CH₄ concentration (c_w , unit mol m⁻³) in the peat pore water:

$$\frac{\partial c_w(t, z)}{\partial t} = -Q_{diff}(t, z) - Q_{plant}(t, z) - Q_{ebu}(t, z) + R_{prod}(t, z) - R_{oxi}(t, z), \quad (1)$$

where Q_{diff} is the diffusive flux of CH₄ in the peat, Q_{plant} is the transport rate of CH₄ via plants roots, Q_{ebu} is the transport rate of CH₄ via ebullition, R_{prod} is the production rate of CH₄ and R_{oxi} is the rate at which oxidation removes CH₄ from the pore water. The units for the terms on the right hand side of Eq. (1) are mol m⁻³ s⁻¹. In this study we altered Q_{ebu} between the runs, the other terms were not modified.

The model is driven with peat temperature (T_p), atmospheric pressure, water table depth (WTD), leaf area index (LAI) of gas-transporting vegetation, and rate of anoxic soil respiration. We ran the model for the Siikaneva peatland site (Rinne et al. 2007) using measured T, atmospheric pressure and WTD and simulated LAI and anoxic respiration as input. The latter was simulated similarly to the study by Susiluoto et al. (2017) in which HIMMELI was combined with a model of anoxic respiration. The model simulated respiration as a fraction of NPP plus temperature-dependent peat decay. The former was distributed vertically according to the root distribution while the latter was distributed evenly into the peat layers below water level, therefore, the main factor driving the anoxic respiration was net primary productivity (NPP) modelled for Siikaneva. ~~We used the same parameter values as in Raivonen et al. (2017).~~ We used the same calibrated parameter values as in the discussion paper of Raivonen et al. (2017). The model has been calibrated in Susiluoto et al. (2017), however parameters used in the different ebullition modelling approaches (see below) were not calibrated. Instead values found from the literature were used for the ebullition modelling approaches, since this way the results are more comparable across studies (e.g. with Wania et al., 2010). Furthermore, when using literature values the results are truly related to differences between the approaches, and not related to differences between calibrations. In this study HIMMELI was run using a 2 m thick peat column with 10 layers, therefore each layer was 0.2 m thick. The model output time step is one day, but the model uses a shorter internal time step in order to ensure numerical stability.

In HIMMELI, WTD divides the peat column into oxic (air-filled) and anoxic (water-saturated) parts. Anoxic respiration is a source of CH₄ and the main part of the anoxic respiration was distributed vertically along an exponential function that describes the vertical distribution of root mass, as root exudates are known to be an important substrate for methanogens (Ström et al., 2003). In contrast to Raivonen et al. (2017) in this study anaerobic decomposition of litter and old peat below water level was also included as a source of CH₄. It was modelled with a simple Q₁₀-model according to Schuldt et al. (2013).

The model simulates transport of CH₄, O₂ and CO₂ between peat and the atmosphere by diffusion in air-filled and water-saturated peat and through aerenchymatous wetland plant roots, as well as by ebullition. In case ebullition occurs when WTD is below the peat surface, the ebullited gases are not released directly into the atmosphere but they are transported into the bottom air-filled peat layer. This is called as the lowest air layer throughout the study. The plant transport capacity depends on LAI that determines the root mass available for gas transport. Methanotrophy oxidizing CH₄ to CO₂ is modelled as a dual-substrate Michaelis-Menten process in which both CH₄ and O₂ concentrations control the oxidation rate. Consequently, the simulated O₂ concentrations affect the CH₄ loss rate in the peat.

2.1.1 Ebullition based on concentration threshold (ECT)

Walter and Heimann (2000) adopted a simple approach to model ebullition: in their approach if the CH₄ concentration at a certain depth exceeded a certain threshold concentration, the excess CH₄ was directly transported to the air layer above the water table depth. In this study we follow Wania et al. (2010) and approximate the threshold pore water CH₄ concentration to equal the equilibrium concentration calculated based on Henry's law. Thus the rate of dissolved CH₄ concentration (C_w) change due to ebullition at a certain depth can be calculated as:

$$Q_{ebu} = k \left(\frac{C_w}{C_w^e} - H^{cc} \frac{P_{tot}}{RT} \right) \quad (2)$$

where k is ebullition half-life ($\frac{\ln(2)}{1800 \text{ s}}$), H^{cc} is the dimensionless Henry solubility of CH₄ calculated based on Sander (2015), P_{tot} is the sum of atmospheric and hydrostatic pressure, i.e. total pressure (Pa), R is the universal gas constant (8.3145 J mol⁻¹ K⁻¹) and T is temperature (K). An equal amount of CH₄ that was removed from pore water based on Eq. (2) is immediately released to the lowest air layer which implies that the bubbles ascend fast enough in order to reach the lowest air layer within the model time step.

2.1.2 Ebullition based on pressure threshold (EPT)

Tang et al. (2010) criticised the ECT approach since it uses a CH₄ concentration threshold to trigger ebullition, whereas in numerous studies decreasing atmospheric pressure has been shown to lead to bubble release event (Green and Baird, 2012; Kellner et al., 2006; Strack et al., 2005; Tokida et al., 2007; Waddington et al., 2009; Yu et al., 2014). Thus they devised a modelling approach which triggers ebullition at a certain depth if the partial pressures of CH₄, CO₂, O₂ and N₂ combined exceed the sum of hydrostatic and atmospheric pressure (P_{tot}). Tang et al. (2010) took into account also the capillary forces when estimating the pressure threshold, however in this study these forces were neglected. In HIMMELI partial pressures of CH₄, CO₂ and O₂ are explicitly calculated using Henry's law and pore water gas concentrations, whereas partial pressure of N₂ is assumed constant (40 % of atmospheric pressure) throughout the peat column. This value for N₂ agrees with empirical evidence for instance by Tokida et al. (2005) and Walter et al. (2008). The excess moles are then released directly to the lowest air layer, unlike in Tang et al. (2010) where a re-dissolution of bubbles back to water is allowed during their ascend_{ascent}. The rate of dissolved CH₄ concentration change due to ebullition can then be calculated for each layer as:

$$Q_{ebu} = k \frac{f_{ss} C_w}{H^{cc}}, \quad (3)$$

where f_{ss} is the relative supersaturation (dimensionless) calculated as

$$f_{ss} = \begin{cases} \frac{p_{tot} - \dot{a}_i p_i}{\dot{a}_i p_i}, & \dot{a}_i p_i > p_{tot} \\ 0, & \dot{a}_i p_i < p_{tot} \end{cases} \quad (4)$$

where p_i denotes the partial pressure of i th gas calculated based on Henry's law. The same amount of CH_4 that was removed from pore water is then released to the lowest air layer. With this approach ebullition of CH_4 may take place even though based on Henry's law water is not supersaturated with CH_4 . Thus ebullition may originate from depths with relatively low pore water CH_4 concentrations when compared to the ECT approach.

2.1.3 Ebullition based on bubble growth (EBG)

Fechner-Levy and Hemond (1996) devised a mathematical framework that describes how temperature, pressure and mass transfer to/from a bubble suspended in water alter the bubble volume. They applied this framework to analyse their peatland data. Later Kellner et al. (2006) applied it to model CH_4 ebullition and Zhang et al (2012) slightly modified the approach and implemented it to a larger process-based model (NEST-DNDC). In this approach bubble volumes are calculated and updated constantly based on ideal gas law and Henry's law and if bubble volume at a certain depth exceeds a predefined threshold, then the excess volume is released to the atmosphere. This approach is supported by laboratory and field experiments which have shown that ebullition occurs only if large enough fraction of the overall pore space volume consists of free-phase gas. This fraction, denoted as $f_{V_{\max}}$ in this study, is commonly reported to be approximately 0.1.

Initially a bubble volume is formed at a certain depth if CH_4 concentration exceeds the concentration that the water can withhold based on Henry's law and assuming that the bubble CH_4 mixing ratio (C) is fixed at 50 % (see Table 1). The excess CH_4 is transferred into a gaseous volume calculated based on the amount of excess moles and the predefined bubble CH_4 mixing ratio. At each model layer this volume is divided evenly between N_{bubble} spherical bubbles. The bubbles do not have any interaction between each other (no coalescence etc) and they remain stationary in the peat-water matrix. In principle N_{bubble} is merely a tuning parameter which controls the rate of mass flux between the gas volume and the pore water. Once the bubbles have been formed the CH_4 exchange between the stationary bubbles and the pore water is calculated based on Epstein and Plesset (1950):

$$Q_{\text{ebu}} = \frac{4prDN_{\text{bubble}}}{V_w} \frac{C - C_w}{C} - \frac{H^{cc} cp_{\text{tot}}}{RT} \quad (5)$$

where r is the radius of one bubble (m), D is temperature dependent effective diffusion coefficient of CH_4 in water ($\text{m}^2 \text{s}^{-1}$) and V_w is the volume of pore water in this model layer (m^3). D was calculated based on Arah and Stephen (1998) and in order to take into account the fact that the media did not consist solely of water, but it was a peat-water mixture, the calculated value for D was multiplied with 0.9 prior usage (Raivonen et al., 2017). For simplicity, temperature and pressure inside the bubble volume were assumed to be equal to their pore water counterparts and the mass transfer was assumed to be stationary within the model time step. In order to keep the model in balance the rate of bubble CH_4 (n_b , unit mol) change at a specific model layer is

$$\frac{dn_b}{dt} = - Q_{\text{ebu}} V_w \quad (6)$$

Thus in this modelling approach CH₄ can be also transferred from the bubbles back to the pore water surrounding the bubbles. This kind of feedback is missing from the other ebullition modelling approaches used in this study.

The bubble volume is updated after every model time step based on Eq. (7):

$$DV = \frac{Dn_b}{c_b} + \frac{DT}{T}V - \frac{Dp_{tot}}{p_{tot}}V, \quad (7)$$

5 where V is the combined volume (m³) of all bubbles at a specific model layer and c_b is the bubble CH₄ concentration (mol m⁻³). The terms on the right hand side of Eq. (7) represent the change in the volume due to diffusion of CH₄ to/from the bubbles (cf. Eq. (6) above), change in volume due to temperature or total pressure change. From now on they are called as the c-, T- and p-term, respectively. The T- and p-terms can be readily calculated from the model input data, whereas the c-term can be determined based on Eq. (6).

10 As stated before, an ebullition event occurs only if bubble volume at a certain depth exceeds a predefined threshold (V_{max}):

$$F_{ebu} = \begin{cases} 0, & V + DV \leq V_{max} \\ c_b \frac{V + DV - V_{max}}{Dt}, & V + DV > V_{max} \end{cases}, \quad (8)$$

where F_{ebu} is the ebullition flux of CH₄ to the lowest air layer (mol s⁻¹), V_{max} can be calculated as a product of V_w and $f_{V_{max}}$ and Dt is the model time step. Finally the updated bubble volume at a certain depth is

$$V = \begin{cases} V + DV, & V + DV \leq V_{max} \\ V_{max}, & V + DV > V_{max} \end{cases}, \quad (9)$$

15 and the amount of CH₄ moles in the bubbles at a certain model layer after each model time step is

$$n_b = \frac{cp_{tot}V}{RT} \quad (10)$$

In order to take into account bubble movement in the peat column after it has been released (i.e. $V + DV > V_{max}$) a simple approach was adopted: excess bubble volumes are released starting from the bottom of the peat column and while the gaseous volume is ascending it will get stuck at certain 0.2 m thick model layer with a probability of P , which was set to 0.3. Thus
 20 for instance bubble released from 1 m below the lowest air layer will reach the air layer with a probability of 0.17 ($(1 - 0.3)^5 \gg 0.17$). Otherwise it will stay at the depth where it got stuck and its volume will be updated during the next time step with the procedure described above. This process is repeated for each layer where $V + DV > V_{max}$ and thus at the end of the time step bubble volumes are always smaller or equal to V_{max} . This kind of approach produces somewhat similar bubble movement as the “inverted bubble avalanches” modelled with the approach suggested by Coulthard et al. (2009) and used by
 25 Ramirez et al. (2015). By setting the probability P to 0 the released excess bubble volume will always reach the lowest air layer, similarly as with the other ebullition modelling approaches included in this study. The performance of the EBG approach using different values for P and other parameters are evaluated in Sect. 3.4.

The effect of changing water table level was taken into account in the following way: if the water table dropped below a layer which contained bubbles, the CH₄ in the bubbles was immediately released to the newly formed air filled pore space. If the
 30 water table rose, the new water clogged pore space did not initially contain any bubbles.

2.2 Site and measurements

The measurements were carried out at an oligotrophic open fen part of the Siikaneva wetland complex which is situated in the southern part of Finland ($61^{\circ}50'N, 24^{\circ}12'E$, 162 m a.s.l.). The site is in the boreal region with an annual average temperature of $3.3^{\circ}C$ and rainfall of 710 mm (Drebs et al., 2002). The vegetation composition is dominated by sedges (*C. rostrata*, *C. limosa*, *E. vaginatum*), Rannoch-rush (*Scheuchzeria palustris*) and peat mosses (*Sphagnum balticum*, *S. majus*, *S. papillosum*). Peat depth at the site varies between 2 and 4 meters. See more details in Riutta et al. (2007) and Rinne et al. (2007).

The HIMMELI model was driven using measured peat temperature, atmospheric pressure, WTD and simulated LAI and anoxic respiration. See the details about LAI and anoxic respiration simulation in Raivonen et al. (2017) and Susiluoto et al. (2017).

Peat temperature was measured at five depths below the surface (-5 cm, -10 cm, -20 cm, -35 cm and -50 cm) and the measurements were interpolated linearly in time and space in order to match every model time step and depth. Temperatures below -50 cm were obtained by assuming that at 3 m the peat temperature is constant ($7^{\circ}C$, average temperature at -50 cm). Also WTD and atmospheric pressure time series were ~~gapfilled~~gap-filled using linear interpolation. Air pressure was not measured at the site and hence measurements made 5 km away at SMEAR II (Hari and Kulmala, 2005) site were used.

In this study the modelled daily CH_4 fluxes are compared against fluxes obtained with the eddy covariance (EC) method at the Siikaneva site. Data from years 2008...2011 are used. EC measurements estimate the ecosystem scale emissions of measured compounds (CH_4 in this study) and hence the measured fluxes integrate all the three CH_4 emission pathways (plant transport, diffusion and ebullition) at ecosystem scale. The EC setup consisted of a sonic anemometer (USA-1, METEK GmbH) which measured the three wind components and air temperature and a fast response gas analyser used to measure CH_4 . The instruments were measuring the turbulent fluctuations 2.75 m above the peat surface. There were some changes in the CH_4 instrumentation during the years. The CH_4 analysers used were RMT-200 by Los Gatos Research (2008...2011), TGA-100 by Campbell Scientific (04/2010...08/2010) and G1301-f by Picarro (04/2010...10/2011). The gas analysers were located in a separate housing which protected the instruments from the outdoor conditions. A sampling line consisting of filters and heated Teflon tubing was used to sample air to the closed-path gas analysers. See more details in Peltola et al. (2013).

The EC measurements were made at 10 Hz from which the fluxes were calculated as a covariance between vertical wind and gas concentration using 30-min averaging time. Coordinates were rotated with sector-wise planar fitting and high frequency losses were corrected using empirical procedures- (cf. Peltola et al., 2013). All these EC data post-processing steps were done using EddyUH (Mammarella et al., 2016). The EC data were carefully processed keeping in mind the nature of the ebullition process (rapid release events with significantly elevated concentrations) and processing steps such as de-spiking were tuned so that these events were not flagged as erroneous measurements.

3 Results

3.1 Timing and depth of ebullition events

The EPT approach resulted in the highest total amount of ebullition events (866 events), followed by EBG (797) and ECT (389) (Fig. 1). An ebullition event is defined as a time period and depth where concentration (ECT approach), pressure (EPT) or volume (EBG) threshold was exceeded (cf. Sect. 2). As stated in Sect. 2.2 EPT approach does not require that the pore water CH_4 concentration reaches supersaturation, it uses the total pressure of water-dissolved gases to trigger an ebullition event, which is a less strict requirement for ebullition. Hence higher amount of events were observed with EPT approach. ECT approach requires high pore water CH_4 concentrations in order to trigger an ebullition event which limits the amount of ebullition events, whereas EBG may trigger an ebullition event due to four reasons: increased CH_4 concentration in pore water (c-term of Eq. (7)), increased temperature (T-term of Eq. (7)), decreased WTD or atmospheric pressure (p-term of Eq. (7)).

Most of the events took place in July, for ECT, EPT and EBG approaches 43 %, 35 % and 36 % of all events happened during that particular month, respectively (Fig. 1). The vertical distributions of the ebullition events in Fig. 1 show that for EBG and EPT approaches the events usually originated from below 1 m depth and for ECT below 1.4 m depth. For EBG this means that below 1 m depth the conditions were favourable for bubble volume increase which lead the volume to exceed the maximum volume allowed (V_{max}) and hence to trigger an ebullition event. In order to sustain and grow bubbles with the EBG approach relatively high pore water CH_4 concentrations are needed (cf. Eq. (5)) and thus the ebullition events originated from depths below rooting depths and layers with oxic pore water (Fig. 1d). For the same reason ECT approach resulted in ebullition events only from the deep pore water. CH_4 in the pore water below rooting depth resulted mostly from decay of old peat, since root exudates were not present and if this source of CH_4 was omitted in the model, all the approaches stopped producing ebullition events altogether.

Less than half (41 %%) of the ebullition events triggered by the EBG approach coincided with a co-located event triggered by the ECT approach, whereas 78 % of the events took place at the same depth and time as with the EPT approach. Most (70 % and 78 %%, respectively) of the events happened at the same time as the events modelled with the ECT and EPT approaches, respectively, regardless of depth. On the other hand, 72 % of the events triggered with the EPT approach matched EBG events, while they matched only 43 % of the ECT events. Hence EPT and EBG triggered ebullition events mostly at the same time and location, however ECT differed more from the other two approaches.

Decreasing WTD (i.e. decreasing hydrostatic pressure) and air pressure triggered approximately 66 % and 67 % of ebullition events when using the EPT approach, respectively. Almost all (94 %%) of the ebullition events observed with EPT approach coincided with either decreasing WTD or decreasing atmospheric pressure, the rest 6 % were triggered solely by increasing partial pressures, i.e. increase in CH_4 , CO_2 or O_2 pore water concentrations (cf. Sect. 2.2). For EBG approach large proportion of the events were triggered simultaneously due to all the three terms, namely the c-, p- and the T-term (47 %), only 1 % of the events were caused solely by decreasing total pressure, i.e. the p-term. Individually examined, almost all of the events took place due to the c-term (99 %) or the p-term of Eq. (7) (86 %), whereas the T-term had slightly smaller impact (57 % of all events). Most (72 %%) of the ebullition events triggered by the p-term coincided with decreasing atmospheric pressure and 59 % with decreasing WTD.

3.2 CH_4 and bubble volume profiles and dynamics

EPT approach showed the lowest pore water CH_4 concentrations (0.39 mol m^{-3}), EBG and ECT calculated on average 1.7 and 2.5 times higher values (0.66 and 0.98 mol m^{-3} , respectively) and the difference was emphasized at the deepest layers (cf. Fig. 2). If the concentrations were converted to gaseous phase partial pressures using Henry's law and scaled with the total pressure (p_{CH_4}/p_{tot}), the reason for the differences could be found. With EBG approach the diffusive mass transfer between the bubbles and the pore water was directed from the water to the bubbles if p_{CH_4}/p_{tot} exceeded the predefined bubble CH_4 mixing ratio \mathcal{C} which was set to 50 % in this study. Conversely, the CH_4 flux was directed from the bubbles to the pore water if the ratio was smaller than \mathcal{C} . This follows directly from the Eq. (5), more specifically from the difference between C_w and the

equilibrium concentration ($\frac{H^{cc} \varphi_{tot}}{RT}$). Hence, the CH_4 concentrations rose until p_{CH_4}/p_{tot} exceeded \mathcal{C} after which part of the produced CH_4 went into the bubbles instead of staying in the pore water, which limited the increase of the CH_4 pore water concentrations. If the effect of surface tension would have been included in the pressure inside the bubbles, higher p_{CH_4}/p_{tot} values would have been needed (i.e. higher pore water CH_4 concentrations) to transfer CH_4 from the pore water to the bubbles. The build-up of pore water CH_4 concentrations took place only below 1 m depth since closer to the surface the plant roots effectively removed CH_4 from the pore water via aerenchyma (compare Figs. 1d and 3). For ECT approach the pore water CH_4 concentrations increased until p_{CH_4}/p_{tot} equalled unity (i.e. supersaturation of CH_4 in pore water) after which the excess

CH₄ was removed in ebullition events. This implies that with ECT approach the bubbles consist 100 % of CH₄. On the other hand, for EPT approach the increase in pore water CH₄ concentrations were limited by the fact that the sum of CH₄, CO₂, N₂ and O₂ partial pressures was allowed to be at maximum p_{tot} and since p_{O_2} values were low and p_{N_2} was 40 % of the atmospheric pressure throughout the peat column, then the sum of p_{CO_2} and p_{CH_4} below the rooting depth could be around 50...65 % of p_{tot} , depending on depth. Hence the low pore water CH₄ concentrations ~~and $p_{\text{CH}_4}/p_{\text{tot}}$ values~~ in Fig. 2c. $p_{\text{CH}_4}/p_{\text{tot}}$ (i.e. the bubble CH₄ mixing ratio) during ebullition events with the EPT approach was generally between 25 % and 38 %, whereas $p_{\text{CO}_2}/p_{\text{tot}}$ (i.e. the bubble CO₂ mixing ratio) was between 28 % and 37 %.

Bubble volumes modelled with the EBG approach mostly resided below 1 m depth (Fig. 3) where the conditions were favourable for bubble growth. Bubble volumes were affected by three terms: CH₄ transfer between the bubbles and pore water (c-term) and expansion/contraction due to temperature (T-term) or pressure changes (p-term). Also input of gaseous volume released from deeper layers increased the bubble volumes, whereas bubble release limited the volume to be at maximum V_{max} (Fig.4).

The c-term showed clear seasonality (decreasing bubble volumes during late summer and autumn and increasing volumes in spring and early summer), since it was closely related to temperature (cf. Eq. (5)): CH₄ solubility decreases with temperature (i.e. H^{cc} decreases) and due to enhanced CH₄ production, pore water CH₄ concentrations (c_w) increase with temperature. Bubble volumes released from deeper layers and which ~~got stuck~~ were re-attached during their ascent ~~kept maintained~~ the ~~bubbles gas volume~~ at -1.1 m depth ~~alive~~ even though the c-term on average decreased ~~their~~ the volume (cf. Fig 4a).

The T- and p-terms ~~inflited generated~~ only temporal variation in the volumes without any permanent increase/decrease. The T-term ~~inflited generated~~ seasonality to the bubble volumes, although mostly the seasonality was controlled by the c-term. The p-term caused strong short-term variation and a small seasonal cycle due to the annual cycle of WTD. Altogether the combination of all these terms, in addition to the input and release of gaseous volume led the bubble volumes to reach their maxima in July and minima in April and generated the ebullition event profile and seasonality shown in Fig. 1a.

3.3 CH₄ emissions to the atmosphere

The EPT approach resulted in significantly higher temporal variability in the CH₄ emission to the atmosphere than the other ebullition modelling approaches (Fig. 5) due to the fact that it produced more ebullition events which caused the short term variability in the CH₄ emissions. In order to quantify the short term variability, the amount of variance in the time series at high frequencies were estimated by calculating the time series' power spectral densities and integrating them over the frequency range of interest. Variability at shorter than one week time scale contributed approximately 24, 12 and 7 % to the total time series variance obtained with the EPT, ECT and EBG approaches, respectively, whereas for the measured CH₄ flux time series the contribution was 3 % to the total variance. There were also differences in the total time series variance: EPT and ECT gave 92 and 17 % higher variances for the CH₄ flux time series than the measurements, whereas EBG estimated 13 % lower variance for the CH₄ emission time series. These results can be qualitatively observed in the Fig. 5: the EPT approach showed clearly highest variability, especially at short time scales, followed by ECT and finally EBG and the measured CH₄ flux time series. EBG approach explained the variability in the measured CH₄ emissions the best (R^2 was 0.7463), followed by ECT (0.5856) and EPT (0.35). These results are consistent with the differences in time series variance outlined above. However, the model results represent CH₄ exchange from a single horizontally homogeneous peat column, whereas the measured fluxes correspond to ecosystem scale CH₄ exchange. Hence they are not fully comparable with each other.

Even though the CH₄ flux time series variances were affected by the ebullition modelling approach, the annual net CH₄ emissions were not largely different between the three modelling approaches (Table 2). At maximum the modelling approaches diverged in year 2010 by 1.0 g(CH₄) m⁻² yr⁻¹ (8 % of the annual emission). In general all the approaches showed similar decreasing trend from year 2008 to 2010 in the annual CH₄ emissions and none of them showed always the lowest or highest emissions. Similar annual CH₄ emissions could have been expected given the fact that ebullition does not directly alter the

CH₄ production and hence it only provides another transport pathway for emissions which would happen at annual time scale regardless of the transport route, although CH₄ oxidation in the peat column ~~complicates this picture~~ might be different between the modelling approaches.

The relative magnitude of different emission pathways varied between the modelling approaches (Fig. 5b, ~~5e5d~~ and ~~5d5f~~ and Table 2). On annual scale for EBG approach the CH₄ emissions via plant aerenchyma contributed approximately 90 and diffusion 10 % of the total CH₄ emissions, for ECT the percentages were 80 and 20 % and for EPT 60 and 40 %, respectively. In HIMMELI the ebullition flux is released to the lowest air layer which is often below peat surface and hence the diffusion flux to the atmosphere contains also the ebullition flux signal. Due to the same reason the direct ebullition to the surface is rare (Fig. 5), since ebullition events usually take place when WTD is below surface and thus ebullition flux is not released directly to the atmosphere. The ebullition flux to the lowest air layer was 4 to 9 times higher in EPT approach than in EBG approach and also approximately double that of calculated with the ECT approach (Table 2) which was related to the overall higher amount of ebullition events modelled by the EPT approach (see Sect. 3.1). In addition with EBG approach many of the ebullition events were stuck during their ascent and then dissolved back to the pore water and thus they did not reach the lowest air layer and contribute to the ebullition flux.

15 3.4 Testing the EBG approach with different model parameter values

~~Testing~~ The EBG approach was tested with different parameter values (cf. Table 1) in order to evaluate the sensitivity of the results on the used parameters. This analysis revealed that the EBG approach was the most sensitive to the bubble CH₄ mixing ratio. If the mixing ratio was set to 20 % ~~(parameter set 2.1)~~, instead of the 50 % used in the default run, 2164 events were triggered with the EBG approach which is approximately three times the amount observed with the default run ~~(cf. Fig. 6b)~~. Consequently, more CH₄ was transported to the lowest air layer via ebullition (annual mean: 1.2 g(CH₄) m⁻² yr⁻¹, with default run: 0.7 g(CH₄) m⁻² yr⁻¹, Fig. 6d) and hence slightly larger fraction of CH₄ was emitted to the atmosphere via diffusion (14 % ~~(parameter set 2.1)~~, Fig. 6e). However, the total amount of emitted CH₄ to the atmosphere was approximately 5 % lower (Fig. 6c), possibly due to the fact that the large amount of ebullition events transported CH₄ to upper oxic peat layers where the CH₄ was oxidized rather than released to the atmosphere. Also, bubbles were formed closer to the surface than in the EBG default run (cf. Fig. 3 for the default run). These results were reasonable given the fact that the lower mixing ratio facilitated the CH₄ transfer from the pore water to the bubbles and hence the bubble growth ~~(parameter set 2.1)~~. On the other hand, higher bubble CH₄ mixing ratio (parameter set 2.2) decreased the amount of ebullition events to 394 (50 % decrease) and limited the bubble formation to depths below 1.2 m, which was deeper than in the default run (Fig. 3).

~~If the threshold fraction for gaseous volume ($f_{v_{max}}$) was increased~~ The variance of the modelled CH₄ flux to 15 % (parameter set 1, cf. Table 1) the amount of ebullition events was reduced by 26 % (586 events observed), since larger bubble volumes were needed before an ebullition event was triggered. ~~Increasing atmosphere was~~ the amount of bubbles in a model layer (N_{bub}) ~~accelerated the CH₄ transfer between the bubbles and pore water (cf. c term in Eq. 4), but on the whole it had a minimal effect on the modelled ebullition.~~ most sensitive to the probability that the released bubble volume will get stuck while it is ascending (Fig. 6a). If the probability was increased from the default value (parameter set 4.2), the flux time series variance was slightly decreased. If the probability was set to 0 (parameter set 4.1, cf. Table 1), the EBG results resembled results obtained with the ECT approach since the ebullition flux was more directly linked with bubble production which was in turn driven by changes in CH₄ pore water concentration (cf. Fig. 4). Ebullition flux to the lowest air layer was approximately three times higher (Fig. 6d), since all the ebullition events reached the layer and were not stuck during their ascent. This also lead the diffusion pathway to have a higher fraction of the overall emission of CH₄ to the atmosphere (28 % vs. 8 % with the default run, Fig. 6e).

If the threshold fraction for gaseous volume ($f_{v_{max}}$) was decreased to 5 % (parameter set 1.1, cf. Table 1) the amount of ebullition events was increased by 59 % (1268 events observed), since smaller bubble volumes were needed before an ebullition event was triggered. Larger value for $f_{v_{max}}$ (parameter set 1.2) decreased the amount of ebullition events by 26 % (Fig. 6b). Otherwise changing the value for $f_{v_{max}}$ had a negligible impact on the model results. Increasing/decreasing the amount of bubbles in a model layer (N_{bub}) accelerated/decelerated the CH₄ transfer between the bubbles and pore water (cf. c-term in Eq. 4), but on the whole it had a minimal effect on the modelled ebullition (see the results for parameter sets 3.1 and 3.2 in Fig. 6).

4 Discussion

All of the modelling approaches produced ebullition events only deep below the surface. Hence the results resemble the “deep peat” hypothesis put forward by Glaser et al. (2004) in which the free-phase gas is produced in deep (> 1 m) peat and trapped under semiconfined layers-, even though peat micro structure (e.g. woody layers) was not described in the model.. According to the hypothesis, these layers episodically rupture due to changes in gas volume buoyancy inflicted by for instance pressure changes. Hence this hypothesis links ebullition mostly to the processes that take place in the deep peat and it has been supported by some field studies. For instance in a relatively recent study Bon et al. (2014) observed high CH₄ pore water concentrations below 2 m depth which they claim to be an indication of free-phase gas and hence ebullition from the deep peat.

There is, however, mounting evidence that bubble formation and release is also taking place close to the surface, which is in contrast to the “deep peat” hypothesis. Hence, Coulthard et al. (2009) argued that the bubble formation and ebullition is more directly linked with processes that take place close to the surface, primarily because most of the labile fresh carbon is located in the rooting zone and hence CH₄ production is higher in the shallow peat than deeper below the surface. The “shallow peat” hypothesis was supported for instance by Klapstein et al. (2014) who showed that in their field study over 90 % of ebullition occurred in the surface peat layer and the carbon in the bubble CH₄ was recently fixed from the atmosphere. As mentioned

before for the EBG modelling approach pore water CH₄ concentrations need to be high enough (above $\frac{H^{cc} \varphi_{tot}}{RT}$) to form and grow a bubble at a certain model layer. This sets a strict limit where bubbles may exist and the presence of vascular plants’ roots effectively prevented bubble formation close to the surface by limiting the pore water CH₄ concentrations. Thus there is clearly a conflict between the “shallow peat” hypothesis and the EBG modelling approach. It could be slightly alleviated by lowering the bubble CH₄ mixing ratio (C), which would allow bubbles to exist at lower pore water CH₄ concentrations, or by using a predefined profile for C (increase with depth), instead of one constant fixed value. In general the reported bubble CH₄ mixing ratios are at minimum 10 %, which equals at $T = 15^{\circ}\text{C}$ and $p = 1013\text{hPa}$ based on Henry’s law equilibrium CH₄ pore water concentration of 0.15 mol m⁻³. However, for instance Baird et al. (2004) have shown bubble build-up initiation at approximately 5-times lower average CH₄ pore water concentrations. This could be explained by strong small scale variability in pore water CH₄ concentration, which is undetectable with the current measurement methods and hence by the methods used by Baird et al. (2004). Such variability could create small “pockets” of high CH₄ concentrations where bubble formation and growth could take place. That kind of variability cannot be however readily implemented in 1-D column models and alternate ways of dealing with this issue should be developed.

It is also possible that the vascular plants in reality did hinder bubble formation and growth at the study site, deeming the EBG results plausible. However, this cannot be confirmed since free-phase gas content was not measured at the site. Coulthard et al. (2009) claim that vascular plants do not necessarily inhibit ebullition from shallow peat due to their strong influence on

CH₄ production, whereas Chanton et al. (2005) suggest the opposite. Also the empirical studies are not conclusive in this matter. For instance Klapstein et al. (2014) did not observe any negative effect of sedge cover on ebullition, whereas Green and Baird (2011) showed that the percentage of CH₄ emitted in episodic ebullition events decreased with the presence of sedges. Green and Baird (2011) conclude that the sedge effects on CH₄ emissions via ebullition may be species dependent.

5 Clearly the effect of sedges on ebullition should be studied more ~~studies are needed~~ prior to these ~~processes effects~~ can be implemented in terrestrial CH₄ models.

While the bubble movement was implemented in the EBG approach in a relatively simple way note that in most of the modelling studies it has been ignored completely. Ramirez et al. (2016) showed that the peat pore structure had a significant influence on the bubble size distribution and release and their results suggest that peat structure might be more important than
10 production rate in controlling ebullition. Hence emphasis on future ebullition modelling should be on describing the bubble movement in a simplistic, yet accurate way. The reduced complexity model MEGA developed by Ramirez et al. (2015) is a step in the right direction in this respect. In the EBG approach used in this study the effect of peat pore structure on bubble movement can be controlled by modifying the probability at which bubbles adhere at certain levels while they ascent (P) and by changing the volume threshold (V_{\max}) after which the bubbles are released. A profile for both of these parameters
15 would allow to take into account the vertical variation in peat pore structure and its effect on bubble mobility and accumulation (Chen and Slater, 2015).

Despite the obvious shortcomings of the EBG approach discussed above it still produced the best match against measured CH₄ fluxes with a relatively high coefficient of determination ($R^2=0.71$)–63). However, the EC system estimated fluxes at ecosystem scale whereas the model results represent fluxes from a single horizontally homogeneous peat column and hence they are not fully comparable. The EC source area (i.e. footprint) may contain locations where ebullition is taking place and locations where at the same it is absent. Therefore, the ebullition events are presumably averaged with other sources (e.g. plant transport) in the conventionally processed EC fluxes, yet their impact on ecosystem scale fluxes is captured accurately. Hence, while EC does estimate fluxes on a continuous basis, EC derived fluxes are not a perfect measure of bubble flux from a single peat column. Ideally these kind of models should be tested against autochambers that are continuously estimating the CH₄ flux at a smaller horizontal scale than EC. Moreover, ebullition events can be detected and quantified from such measurements (e.g. Goodrich et al., 2011) and hence could be directly compared against the modelled ebullition. Besides autochambers, the models could also be validated against changes in CH₄ storage (liquid and gas phase) in the peat, since ebullition modelling approach largely alters the amount of CH₄ stored below the surface during the model run (cf. Fig. 2).
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The process-based models are often evaluated by comparing with measurements and possibly optimised to match observed
30 CH₄ emissions by minimising some statistic or objective function, for instance root mean square error (RMSE) (Wania et al., 2010) or Nash-Sutcliffe efficiency (NS) (van Huissteden et al., 2009), or by using for instance Markov chain Monte Carlo methods (Susiluoto et al. 2017). Poorly represented processes, for instance ebullition, hinder this comparison and may yield erroneous values for many of the model parameters when the models are calibrated, even though other processes would be described in a realistic manner. This is simply because the poorly described process causes apparent mismatch between models
35 and measurements. This might be especially true for ebullition, since it strongly influences CH₄ flux time series variance which in turn has a direct impact on many metrics, such as RMSE, NS and R^2 . Hence, slightly more realistic ebullition modelling, such as the EBG approach, would allow better tuning of the CH₄ models and ultimately more accurate CH₄ emission estimates. Finally it should be mentioned that there is also merit in simplicity. While the ECT approach lacks feedback to many ebullition drivers (e.g. pressure changes), the approach is simple and hence adds only a minimal amount of degrees of freedom to the
40 model and therefore possibly provides more robust modelling results. The approach could be further modified to take into account the fact that bubbles do not consist 100 % of CH₄ (e.g. Riley et al., 2011) in order to make it more realistic. On the other hand the merit of the EPT approach, besides being a simple approach, is that it does not use a predefined CH₄ mixing

ratio in the bubbles and hence it is a viable modelling approach for models which explicitly calculate also CO₂ and O₂ concentrations in the peat column.

5 Conclusions

In this study three approaches to model CH₄ transport via bubbling, i.e. ebullition, were compared by implementing them in one peatland CH₄ cycling model called HIMMELI. The model was run using forcing data from a boreal sedge fen. The study was motivated by the fact that ebullition is modelled rather crudely in many models and hence there is clearly a need for improvement and comparison of methods. All the three approaches were based on thresholding on some variable, either pore water CH₄ concentration (ECT approach), pressure (EPT) or free-phase gas volume (EBG). The ECT approach is commonly used in process-based CH₄ models even though it describes the physical processes behind ebullition in a simplistic manner, whereas EBG approach resembles most closely the current knowledge on the process.

EPT simulated the highest amount of ebullition events and hence produced the highest ebullitive CH₄ fluxes to the surface. All the modelling approaches triggered ebullition events only well below the surface, which was caused by the effect of vascular plant root distribution on pore water CH₄ concentrations. The modelled CH₄ fluxes were also compared against eddy covariance (EC) measurements of CH₄ fluxes and EBG produced the best match against measurements ($R^2=0.7463$), although a horizontal scale mismatch is evident between the model results (single peat column) and EC measurements (ecosystem scale). EBG incorporates most of the ebullition drivers observed in different studies (temperature, pressure, CH₄ production, water level changes), whereas the other modelling approaches, especially ECT, are missing a link to many of the drivers listed. While simple modelling approach, such as ECT, may yield robust results without many tunable parameters, overly simplified processes in models may hinder model comparison against measurements. Hence modellers and researchers are encouraged to incorporate a realistic description of the ebullition pathway to their models.

Code availability

The Fortran codes describing the EBG modelling approach are included as supplementary material to this article.

Competing interests

The authors declare that they have no conflict of interest.

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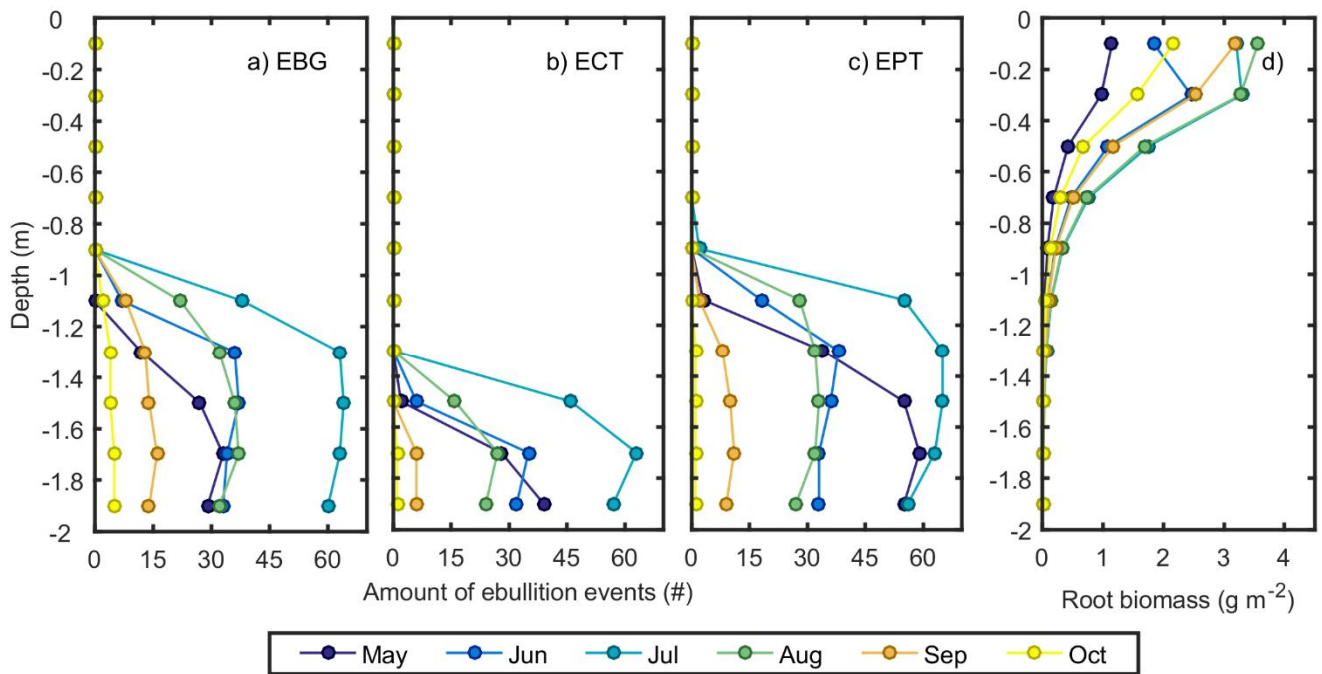


Figure 1: Amount of ebullition events at different months and depths (subplots a...c) and distribution of plant roots under water table (subplot d). During other months (November...April) there were only 1 % of all ebullition events and thus data from those months are not shown.

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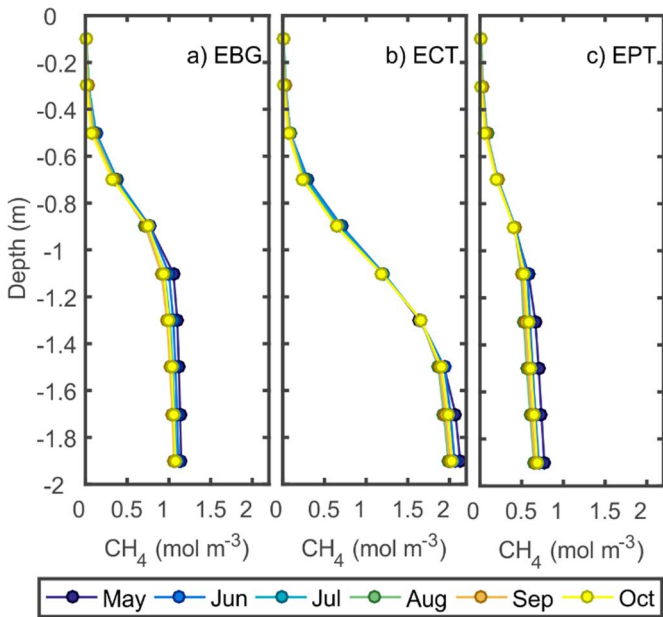


Figure 2: Average CH₄ pore water concentrations at different depths and months.

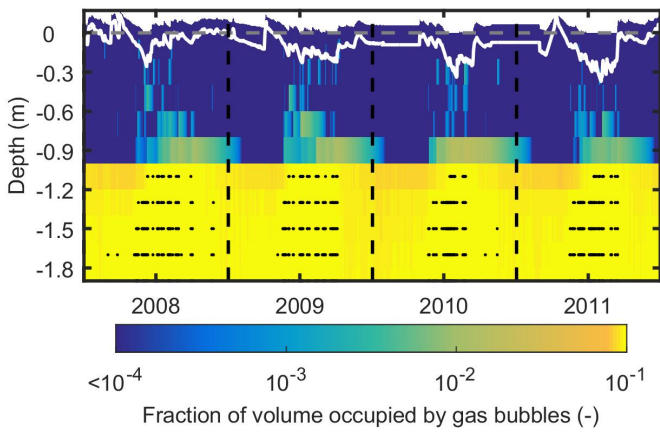


Figure 3: Bubble volume profiles calculated with the EBG approach. Black dots highlight periods and depths with ebullition events. White line shows the water table depth.

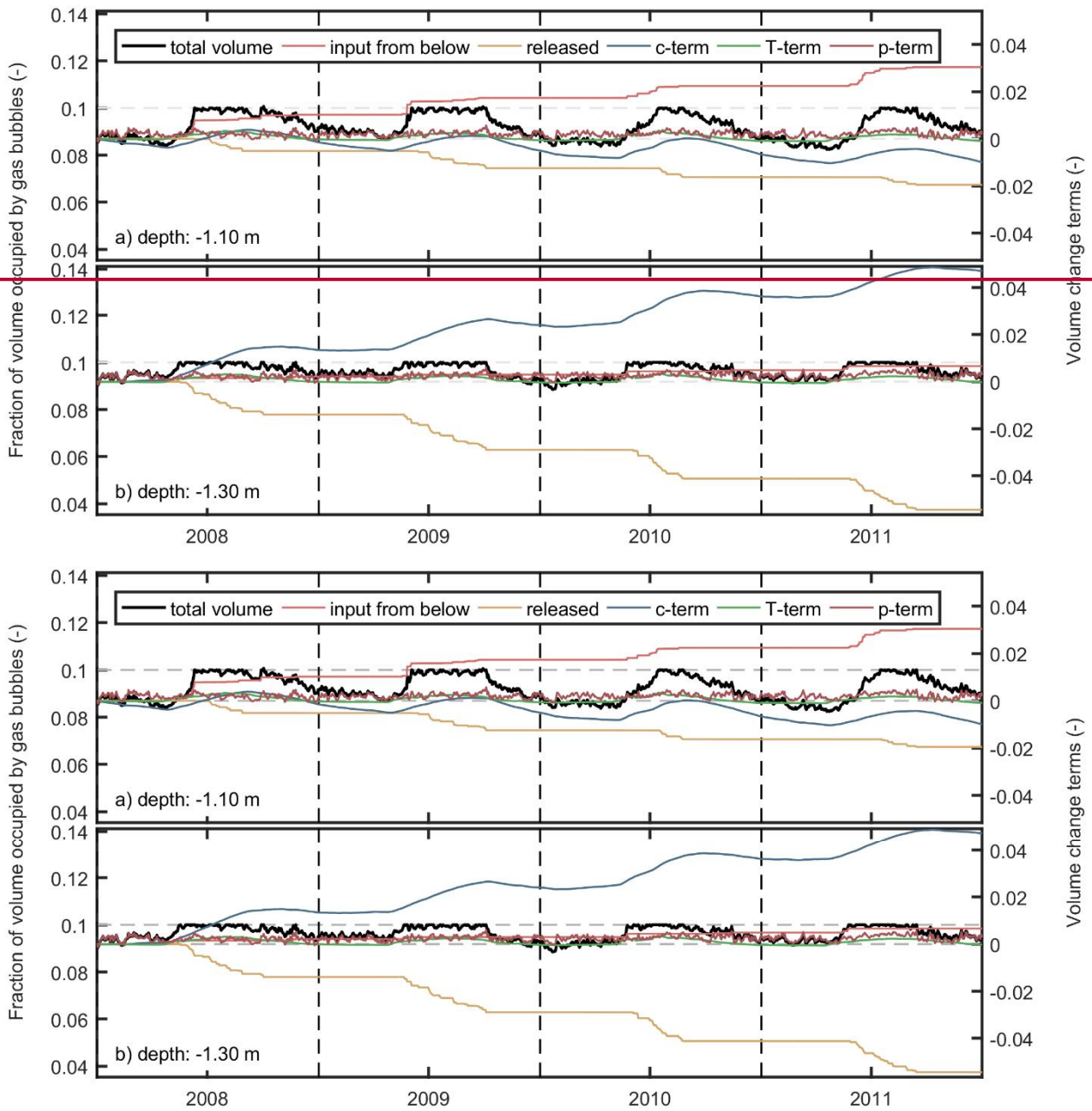
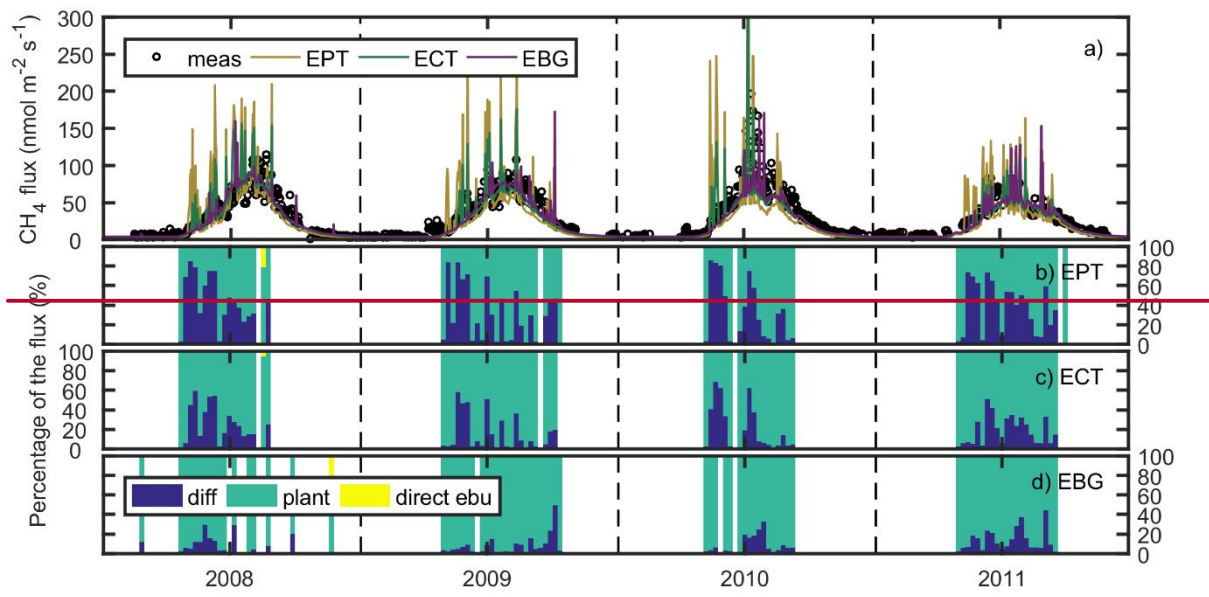


Figure 4: Fraction of volume occupied by gas bubbles at two example depths calculated with the EBG approach. Black lines show the modelled volume (left y-axis) and the other continuous lines (right y-axis) show the terms causing the volume changes. Therefore

sum of the lines in colour yield the changes in the black line at every time step. Light grey dashed line highlights the volume threshold ($f_{V_{max}} = 0.1$), after which additional volume increase was released in an ebullition event.



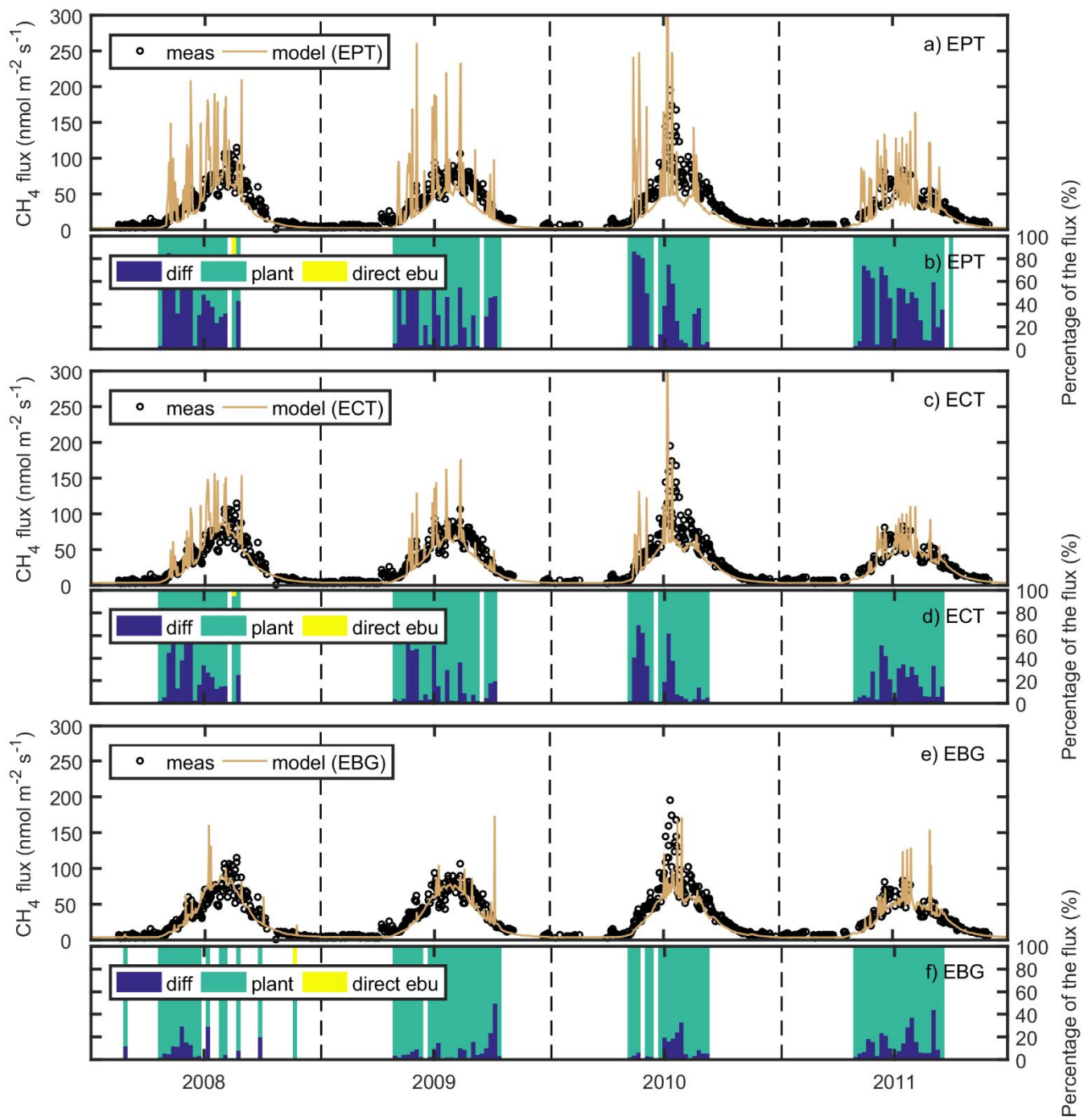


Figure 5: Daily CH₄ fluxes from the fen to the atmosphere obtained using different approaches for modelling ebullition (subplots a, c and e) and weekly contribution of three emission pathways to flux (subplots b, d and f). Measured fluxes are shown with circles. diff = diffusion, plant = plant transport, direct ebu = ebullition directly to the surface. The diffusion flux contains the ebullition flux released to the lowest air layer which is usually below the peat surface. White areas in the subplots b, d and f correspond to periods when over 99 % of the flux was related to the plant-mediated transport.

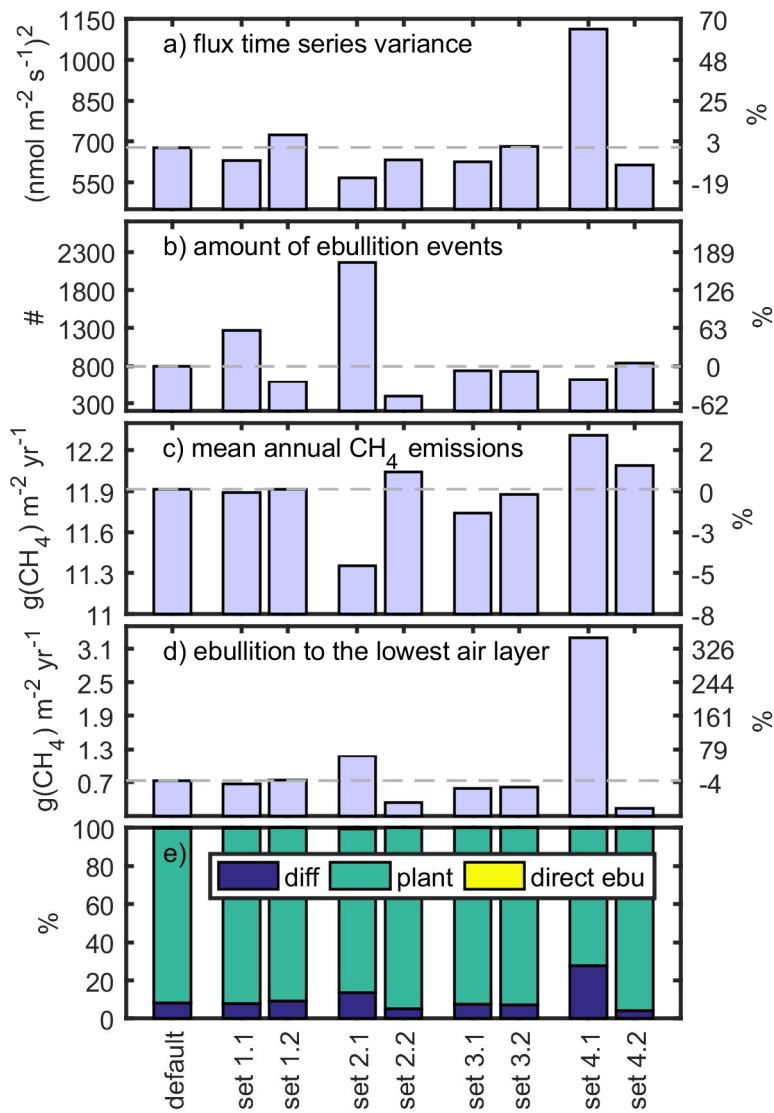


Figure 6: Results from the EBG sensitivity analysis (see parameter values in Table 1): variance of the CH₄ flux time series (subplot a), total amount of ebullition events (b), mean annual CH₄ emissions (c), mean annual ebullition flux to the lowest air layer (d) and the average contribution of different emission pathways to the overall CH₄ flux to the atmosphere (e). Direct ebullition to the surface is negligible on annual scale and hence not visible in the subplot e. Absolute values are on the y-axis on the left and change relative to the default run are on the right side. Grey dashed lines highlight the values obtained with the default run.

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Table 1: Parameters needed in the EBG ebullition module. Default values are given, along with different values used in the sensitivity analysis.

<u>Coefficient</u> <u>#Parameter</u> <u>†</u>	Description	Default values	Parameter set 1.1 (<u>$f_{V_{max}}$</u> <u>low</u>)	Parameter set 1.2 ($f_{V_{max}}$ high)	Parameter set 2.1 (C low)	Parameter set 2.2 (<u>C</u> <u>high</u>)	Parameter set 3.1 (<u>N_{bub}</u> <u>low</u>)	Parameter set 3.2 (N_{bub} high)	Parameter set 4.1 (<u>P</u> <u>low</u>)	Parameter set 4.2 (P high)
<u>$f_{V_{max}}$</u> <u>(dimensionless)</u>	Threshold fraction of pore space filled with gas bubbles needed for ebullition	0.1	0.05	0.15	0.1	0.1	0.1	0.1	0.1	0.1
<u>C</u> (mol mol ⁻¹)	CH ₄ mixing ratio in the bubbles	0.5	0.5 mol mol⁻¹	0.5	0.2 mol mol⁻¹	0.8	0.5 mol mol⁻¹	0.5 mol mol⁻¹	0.5	0.5
<u>N_{bub}</u> (#)	Amount of bubbles in one 0.2 m thick model layer	100	100	100	100	100	10	1000	100	100
<u>P</u> <u>(dimensionless)</u>	Probability that the released bubble volume will get stuck at one 0.2 m thick model layer while it is ascending	0.3	0.3	0.3	0.3	<u>0.3</u>	0.3	0.3	0	0.5

Table 2: Annual CH₄ emissions (g(CH₄) m⁻² yr⁻¹) estimated with HIMMELI model using different ebullition modules and measured with eddy covariance. Values in parentheses show the relative contribution (%) of each CH₄ emission pathway to the total annual emission estimate.

		EBG	ECT	EPT	Measured
2008	Total	12.9	13.7	13.9	13.0 <u>11.9</u>
	Plant-transport	12.2 (94)	11.3 (82)	9.0 (65)	
	Diffusion	0.7 (6)	2.4 (18)	4.7 (34)	
	Ebullition	0 (0)	0 (0)	0.1 (1)	
	Ebullition to the lowest air filled pore space ^a	0.5	2.4	4.8	
2009	Total	12.4	12.3	12.6	13.4 <u>12.8</u>
	Plant-transport	11.5 (93)	10.2 (83)	8.0 (63)	
	Diffusion	0.8 (7)	2.1 (17)	4.6 (37)	
	Ebullition	0 (0)	0 (0)	0 (0)	
	Ebullition to the lowest air filled pore space ^a	0.6	1.9	4.5	
2010	Total	11.6	12.4	12.4	14.7
	Plant-transport	10.6 (91)	9.9 (80)	7.6 (61)	
	Diffusion	1.0 (9)	2.5 (20)	4.8 (39)	
	Ebullition	0 (0)	0 (0)	0 (0)	
	Ebullition to the lowest air filled pore space ^a	0.7	2.3	4.7	
2011	Total	10.8	10.7	10.7	12 <u>11.6</u>
	Plant-transport	9.5 (87)	8.8 (82)	6.6 (61)	
	Diffusion	1.4 (13)	2.0 (18)	4.2 (39)	
	Ebullition	0 (0)	0 (0)	0 (0)	
	Ebullition to the lowest air filled pore space ^a	1.0	1.7	4.0	

^a Not included in the total annual CH₄ emission.