

**Soil properties determine the elevational patterns of base cations and
micronutrients in plant-soil system up to the upper limits of trees and shrubs**

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Abstract

20 To understand whether base cations and micronutrients in the plant-soil system
change with elevation, we investigated the patterns of base cations and micronutrients
in both soils and plant tissues along three elevational gradients in three climate zones
in China. Base cations (Ca, Mg and K) and micronutrients (Fe, Mn and Zn) were
determined in soils, trees and shrubs growing at lower and middle elevations as well
25 as at their upper limits on Balang (subtropical, SW China), Qilian (dry-temperate, NW
China) and Changbai (wet-temperate, NE China) mountains. No consistent elevational
patterns were found for base cation and micronutrient concentrations in both soils and
plant tissues (leaves, roots, shoots and stem sapwood). Soil pH, soil organic carbon
(SOC), total soil nitrogen (TN), the SOC to TN ratio (C:N), and soil extractable
30 nitrogen (NO_3^- and NH_4^+) determined the elevational patterns of soil exchangeable Ca
and Mg and available Fe, Mn and Zn. However, the controlling role of soil pH and
SOC was not universal as revealed by their weak correlations with soil base cations
under tree canopies at the wet-temperate mountain and with micronutrients under both
tree and shrub canopies at the dry-temperate mountain. In most cases, soil base cation
35 and micronutrient availabilities played fundamental roles in determining the base
cation and micronutrient concentrations in plant tissues. An exception existed for the
decoupling of leaf K and Fe with their availabilities in the soil. Our results highlight
the importance of soil physicochemical properties (mainly SOC, C:N and pH) rather
than elevation (i.e., canopy cover and environmental factors, especially temperature),
40 in determining base cation and micronutrient availabilities in soils and subsequently

their concentrations in plant tissues.

Key words base cation, micronutrient, plant tissue, soil physicochemical property, shrubline, treeline

45 **1 Introduction**

Base cations and micronutrients are essential for soil fertility and plant
physioecological processes of photosynthesis, metabolism, growth and productivity
(Salisbury and Ross, 1992). For instance, exchangeable calcium (Ca) and magnesium
(Mg) are predominant base cations responsible in buffering soil acidity, and
50 deficiency of these nutrients can occur in terrestrial ecosystems (Naples and Fisk,
2010; Baribault et al., 2012; Sardans and Peñuelas, 2015), particularly when they are
exposed to acid rain. Micronutrient deficiency, on the other hand, occurs more
frequently, for instance, when replenishment of micronutrients via litter
decomposition does not keep pace with output processes of plant uptake and leaching
55 (White and Zasoski, 1999; Hernandez-Apaolaza, 2014). High soil pH can limit the
availability of micronutrients of iron (Fe), manganese (Mn) and zinc (Zn) (Reisenauer,
1988; Lucena, 2000; Rengel, 2007), while low soil pH can induce toxicities of trace
metals constraining terrestrial net primary productivity (He et al., 2005; Reisenauer,
1988; Tian et al., 2016). The concentrations of soil base cations and available
60 micronutrients were suggested to be positively and negatively correlated with soil pH,
respectively, but both positively correlated with soil organic matter (SOM)
concentration (Reisenauer, 1988; Wang et al., 2017). Concentrations of base cation
and micronutrient can differ among plant tissues as a result of their slow translocation
from the roots to the canopy (van der Heijden et al., 2015), distinct mobility among
65 plant tissues (Warnock, 1970), and occurrence of remobilization during different
physiological stages (Maillard et al., 2015). Quantifying base cation and micronutrient

concentrations in soils and plant tissues (leaves, roots, shoots and stems) can help understand the nutritional status and potential deficiencies of micronutrients during plant growth (Richardson, 2004). However, little attention has been paid to base
70 cation and micronutrient availabilities in soils as well as their variations among plant roots, leaves, shoots and stem sapwood under changing environmental conditions (Rengel, 2007).

The plant distribution and growth along elevational gradients reflect changes in environmental conditions (Li et al., 2003, 2004, 2006, 2008a,b; Zhu et al., 2012a,b).
75 Plants growing at high elevation, especially close to their upper limits, are expected to be highly sensitive to climate change, in particular to global warming (Noble, 1993). Physiological studies of treeline trees have mainly focused on macronutrients such as nitrogen (N), phosphorus (P), and Ca (Richardson, 2004; Liptzin et al., 2013; Mayor et al., 2017), while there are few data available for micronutrients in plant-soil
80 systems along elevational gradients (Wang et al., 2009). Two hypotheses have been proposed to explain nutrient accumulation and/or nutrient deficiency in plant tissues at high elevations (Oleksyn et al., 2002; Richardson, 2004). First, the decrease in temperature with increasing elevation declines soil microbial activity and plant metabolism, and thus constrains soil nutrient cycling and plant uptake processes
85 (Körner & Paulsen, 2004; Thibault et al., 2014). At the alpine treeline, low temperature slows down microbial-mediated litter decomposition and thus reduces nutrient supply to plants (van den Driessche, 1974; Richardson, 2004). Second, another paradigm exists that plants retain higher nutrient concentrations in their

tissues to maintain metabolic capacity and to avoid cold injury at higher elevations
90 with cold growth conditions (Oleksyn et al., 2002). These two hypotheses are mainly
tested on treeline trees, and little attention has been put on other plant life forms, such
as shrubline shrubs. Whether trees and shrubs growing at high elevations have higher
or lower base cation and micronutrient concentrations is still unclear.

The change in environmental conditions along elevational gradients, including
95 temperature and growing season length (Barry, 1981), provides a unique opportunity
to examine the spatial distribution of base cations and micronutrients in plant-soil
systems. Uncertainties still exist whether soil properties or environmental factors
determine the base cations and micronutrients in plant-soil systems. We, therefore,
studied the base cation (Ca, Mg, and K) and micronutrient (Fe, Mn, and Zn)
100 concentrations in plant-soil systems along elevational gradients up to the alpine
treeline and shrubline in sub-tropic, dry temperate and wet temperate climate zones in
China. Soil base cation and micronutrient concentrations can increase through soil
weathering and decomposition of organic matter, but can decrease with plant uptake
and loss through leaching. We therefore hypothesized that soil base cation and
105 micronutrient concentrations increase with increasing elevation because plant uptake
decreases more than the supply through weathering and decomposition with elevation.
We also expected that plants of both trees and shrubs at higher elevations would have
greater base cation and micronutrient concentrations in their tissues (leaves, roots,
shoots, and stem sapwood) to maintain physio-ecological processes in a colder
110 environment. To test these hypotheses, we collected soil and plant samples along three

elevational gradients from lower elevations up to the alpine treeline or shrubline in three climate zones in China, and studied the Ca, Mg, K, Fe, Mn, and Zn concentrations in plant-soil systems.

115 **2 Materials and methods**

2.1 Site description and sample collection

Study sites were located in three climate zones (summarized in Table S1): Balang mountain with a subtropical climate located in Wolong Nature Reserve (“subtropical Mt. ”, 102°52'-103°24'E, 30°45'-31°25'N) in southwestern China, Qilian mountain
120 located in the dry-temperate climate zone (“dry-temperate Mt. ”, 102°58'-103°01'E, 37°14'-37°20'N) in northwestern China, and Changbai mountain with a wet-temperate climate (“wet-temperate Mt.”, 126°55'-129°00'E, 41°23'-42°36'N) located in northeastern China (see Fig. S1). Three distinct sites were chosen to find the general patterns for base cations and micronutrients along elevational gradients across climate
125 scales rather than to investigate the comparability among study sites. The subtropical Mt. is influenced by warm-wet monsoon masses in summer and continental air masses in winter (Li et al., 2012). The mean annual precipitation (MAP) of the subtropical Mt. is about 846 mm monitored by Dengsheng Meteorological Station at 2730 m (Li et al., 2012). For the dry-temperate Mt., the MAP is 435 mm, which is monitored by the
130 Qilian weather station at 2787 m altitude (Qiang et al., 2003). The wet-temperate Mt. is located in a typical continental temperate monsoon climate zone with MAP increasing from 632 to 1154 mm along the elevational gradient from 530 to 2200 m

(Shen et al., 2013).

In this study, the alpine treeline and shrubline are defined as the upper limit of
135 obvious trees and shrubs, respectively. The trees that were investigated are *Abies*
faxoniana (elevation range 2860-3670 m) for the subtropical Mt., *Picea crassifolia*
(elevation range 2540-3250 m) for the dry-temperate Mt., and *Betula ermanii*
(elevation range 1700-2030 m) for the wet-temperate Mt.. The shrubs are *Quercus*
aquifolioides (elevation range 2840-3590 m) for the subtropical, *Salix gilashanica*
140 (elevation range 3020-3540 m) for the dry-temperate, and *Vaccinium uliginosum*
(elevation range 1430-2380 m) for the wet-temperate Mts. The targeted treeline trees
and shrubline shrubs are dominant and common species for each study site. The soils
from the three sampling sites of subtropical, dry-temperate, and wet-temperate Mts.
were classified as Umbric Cryic Cambisols, Calcaric Ustic Cambisols and Andic
145 Gelic Cambisols, respectively (IUSS Working Group WRB, 2014).

Plant tissue samples of current-year mature leaves, roots (< 2 mm), stem sapwood,
and shoots (twigs) from trees and shrubs were collected at lower and middle
elevations, as well as at the upper limits. At each elevation, 6 independent plots (10 m
× 10 m) were selected to serve as 6 replicates on southern slopes with 25 m distance
150 between adjacent plots. Within each plot, 6-10 trees or shrubs of similar height were
randomly selected for tissue sampling. Sampling elevations for trees and shrubs were
different for each site (Table S1). Soils (0-10 cm) were directly collected under the
canopy of trees or shrubs sampled for each plot using a 3-cm diameter corer. Both
plant and soil samples were homogenized and composited within each plot. Samples

155 were collected at the middle of July for subtropical Mt., at the beginning of August for
dry-temperate Mt., and at end of August for wet-temperate Mt. in 2014. The main
characteristics of the three study sites are summarized in Table S1.

2.2 Chemical analysis

160 The soil samples were separated into two subsamples with one subsample being
air-dried to constant weight and the other one stored at 4 °C for further analyses. For
subtropical and wet-temperate Mts., soil organic carbon (SOC) and total nitrogen (TN)
was determined on ground soils using an elemental analyzer (Vario MACRO Cube,
Elementar, Germany). For dry-temperate Mt., the ground soil samples were treated
165 with 12 M HCl according to Wang et al. (2015) to remove inorganic C before organic
C determination on the elemental analyzer. Soil NO_3^- -N and NH_4^+ -N were extracted
from fresh soils with a 2 M KCl solution and measured using an AutoAnalyser III
continuous Flow Analyzer (Bran & Luebbe, Norderstedt, Germany). Soil total
inorganic nitrogen (TIN) was the sum of extractable NO_3^- -N and NH_4^+ -N. Soil Olsen
170 phosphorus (P) was quantified by colorimetric analysis after extraction with a 0.5 M
 NaHCO_3 solution (Olsen et al., 1954).

A subsample of 5 g soil was used to determine soil pH in a 1:5 (w/v) soil-to-water
suspension. Soil exchangeable base cations were extracted with a 1 M ammonium
acetate solution (Wang et al., 2017). Soil available micronutrients were extracted by
175 diethylenetriamine pentaacetic acid (DTPA) according to Lüt et al. (2016). Briefly, 10
g of soil was extracted by 20 ml 0.005 M DTPA + 0.01 M CaCl_2 + 0.1 M TEA

(triethanolamine) (pH 7.0). The soil-solution suspension was shaken for 2 h at 180 rpm and then filtered through ash-free filter paper. The concentrations of base cations and micronutrients were determined using an atomic absorption spectrometer (AAS, Shimadzu, Japan).

Plant samples of leaves, roots, shoots and stem sapwood were oven-dried at 60 °C for 48 h and ground for base cation and micronutrient analyses. Root samples were washed prior to being oven-dried. To determine total base cation and trace element concentrations, 0.2 g plant samples were digested with a mixture of acids of HNO₃ and HClO₄ (5:1, v/v) on a hot plate. After the mixture turned into clear solution, the digests were decanted into 50 ml volumetric flasks and the volume was adjusted to 50 ml. The concentrations of Ca, Mg, K, Fe, Mn, and Zn were determined by the AAS (Shimadzu, Japan).

190 **2.3 Statistical analyses**

Normality of data was determined using the Kolmogorov-Smirnov test, and homogeneity of variances using Levene's test. Generalized linear mixed models (GLMM) were executed to determine the effects of plant life form (tree or shrub), elevation position and their interactions on soil pH, SOC, soil exchangeable base cations and available micronutrients, and total base cations and micronutrients in plant tissues. We assigned sampling site as a random factor in the statistics, as this study aimed to test the general elevational patterns instead of site-specific heterogeneity of base cations and micronutrients in plant-soil system across three sites. The GLMM

was performed using R version 3.2.3 (<http://www.r-project.org>).

200 Within each site, the effect of elevation on measured parameters was determined by multiple comparisons with Duncan's multiple range test for soils and each life form. Pearson correlation analysis was performed to determine the relationships between measured parameters using SPSS 16.0 (SPSS, Inc., Chicago, IL, USA). The soil pH, SOC, TN, ratio of SOC to TN (C:N), NH_4^+ and Olsen P were analyzed as
205 factors explaining the variability of soil base cation and micronutrient concentrations by structural equation modeling (SEM). Before constructing models, simple linear regressions between all parameters were examined in order to select factors explaining the maximum variability. The SEM analyzed the direct, indirect and total effect of soil parameters on targeted variables. We fitted the models using R statistics
210 and determined the best-fit models using Akaike information criteria. Statistical significance was accepted at $P < 0.05$ for all the analyses.

3 Results

3.1 Soil pH and SOC

215 Soil pH was significantly different among elevational positions (Table 1). For both subtropical and dry-temperate Mts., soil pH decreased with increasing elevation under tree canopy, while it was the opposite trend under shrub canopy (Fig. 1a). For wet-temperate Mt., the upper limit of shrubs had significantly higher soil pH (Fig. 1a).

For all three sites, SOC concentration showed a hump-shaped trend with the
220 highest value at the middle elevation under tree canopy (Fig. 1b). Under shrubs, SOC

concentration significantly increased with increasing elevation for subtropical and wet-temperate Mts., while it was the lowest at the upper limit of shrubs for dry-temperate Mt. (Fig. 1b).

225 **3.2 Changes in soil base cations and available micronutrients**

Soil exchangeable Ca and Mg decreased with increasing elevation under tree canopy of subtropical and wet-temperate Mts. and under shrubs of dry-temperate Mt. (not for Mg) (Fig. 2a,b). However, they showed the opposite trend under shrubs of subtropical (not for Mg) and wet-temperate Mts. and under trees of dry-temperate Mt. (Fig. 2a,b).

230 Soil exchangeable K decreased with increasing elevation under tree and shrub canopies at subtropical Mt. and under trees at dry-temperate Mt. (Fig. 2c).

Soil available Fe was significantly affected by elevation position (Table 1 & Fig. 2d). The upper limit had the lowest concentration under both tree and shrub canopies for subtropical Mt. and under shrub canopy for wet-temperate Mt. (Fig. 2d). For 235 dry-temperate Mt., soil available Fe significantly increased with increasing elevation under both tree and shrub canopies (Fig. 2a). For subtropical Mt., soil available Mn was significantly higher at the middle elevation under tree canopies. Soil available Mn decreased with increasing elevation under shrub and tree canopies for subtropical and wet-temperate Mts., respectively, while it showed the opposite trend under both tree 240 and shrub canopies of dry-temperate Mt. and under shrubs of wet-temperate Mt. (Fig. 2e). Soil available Zn was significantly affected by plant life form and the interactive effect between life form and elevation position (Table 1). Specifically, soils at the

upper limit had the highest available Zn under shrubs at wet-temperate Mt..

245 **3.3 Base cations in plants**

For subtropical Mt., a significant decrease of Ca concentration was detected in the leaves of trees and shrubs (Fig. 3a), roots and shoots of trees (Fig. 3b,c), and stem sapwood of shrubs with increasing elevation (Fig. 3d). For dry-temperate Mt., Ca concentration decreased with increasing elevation in roots, shoots and stem sapwood of trees (Fig. 3b,c,d), and in shoots and stem sapwood of shrubs (Fig. 3c,d). For wet-temperate Mt., shoot Ca concentration decreased with increasing elevation for trees (Fig. 3c). Along with increasing elevation, a significant decrease of Mg was found in shrub leaves, tree roots, shrub shoots and stem sapwood at sub-tropical Mt. (Fig. 3e,f,g,h), and in roots, shoots and stem sapwood of trees at dry-temperate Mt. (Fig. 3f,g,h), and in leaves and shoots of both trees and shrubs at wet-temperate Mt. (Fig. 3e,g). With the increase in elevation, K concentration significantly decreased in leaves of trees, roots and stem sapwood of both trees and shrubs at subtropical Mt. (Fig. 3i,j,l), in tree shoots of dry-temperate Mt. (Fig. 3k), and in leaves of both trees and shrubs at wet-temperate Mt. (Fig. 3i).

260

3.4 Micronutrients in plants

For subtropical Mt., Fe concentrations in leaves (Fig. 4a) and roots (Fig. 4b) showed a similar trend with soil available Fe, with the highest values at the middle elevation for both trees and shrubs. For dry-temperate Mt., the highest Fe concentrations were

265 found at the lower elevation in leaves (Fig. 4a), roots (Fig. 4b), shoots (Fig. 4c) and stem sapwood (Fig. 4d) for trees and in shoots (Fig. 4c) and stem sapwood (Fig. 4d) for shrubs. For wet-temperate Mt., Fe concentration was the highest in tree shoots at the middle elevation (Fig. 4c), in shrub leaves at lower elevation (Fig. 4a), and in roots, shoots and stem sapwood of shrubs at the upper limit of trees (Fig. 4b,c,d).

270 The Mn concentration decreased with increasing elevation in leaves and shoots of both trees and shrubs at subtropical Mt. (Fig. 4e,g), in stem sapwood of shrubs at both subtropical and dry-temperate Mts. (Fig. 4h). The Mn concentration increased with increasing elevation in leaves of trees at dry-temperate Mt. (Fig. 4e), in roots of both trees and shrubs at wet-temperate Mt. (Fig. 4f), in shoots of shrubs and stem sapwood
275 of trees and shrubs at wet-temperate Mt. (Fig. 4g,h).

The Zn concentration was the highest at middle elevation for trees in leaves at wet-temperate Mt. (Fig. 4i), in roots at dry-temperate Mt. (Fig. 4j), in shoots at wet-temperate Mt. (Fig. 4k) and in stem sapwood at subtropical Mt. (Fig. 4l). With the increase in elevation, a decrease of Zn concentration was found in roots of trees at
280 subtropical Mt. (Fig. 4j) and in stem sapwood of shrubs at dry-temperate Mt. (Fig. 4l); however, an increase of Zn was found in shrub roots at wet-temperate Mt., in shoots of trees at dry-temperate Mt. and shrubs at wet-temperate Mt. (Fig. 4k), and in stem sapwood of trees at dry-temperate Mt. and shrubs at wet-temperate Mt. (Fig. 4l).

285 **3.5 Correlations between plant and soil parameters**

Across all sampling sites and plant life forms, both soil exchangeable Ca and Mg were

positively correlated with soil pH (Fig. S2a,b) and TN (Table 2), while they were negatively correlated with soil C:N, NO_3^- , and NH_4^+ (Table 2). For wet-temperate Mt., both soil pH and SOC showed no relationship with soil exchangeable Ca and Mg
290 under tree canopies, although SOC was positively related to exchangeable K (Table S2). Negative correlations were found for both Mg and K concentrations between stems and leaves (both $p < 0.01$; Table S3). However, Mg and K concentrations in roots showed no correlation with that in leaves (Table S3).

When analyzing data across sampling sites and plant life forms, soil available Fe,
295 Mn and Zn were negatively correlated with soil pH ($p < 0.01$; Table 2, Fig. S2c,d,e), and soil available Fe and Zn were positively correlated with SOC ($p < 0.01$; Table 2). However, available micronutrients had no relationships with both soil pH and SOC at dry-temperate Mt., except for a positive correlation between soil pH and available Mn under shrub canopies (Table S2). For both Mn and Zn concentrations, significant and
300 positive correlations were found between soil and plant tissues (Table S3). Soil available Fe was negatively correlated with Fe concentrations in shoots and stems (both $p < 0.01$; Table S3).

According to the SEM analyses, soil pH and TN positively, while C:N negatively affected exchangeable Ca (Fig. 5a). Exchangeable Mg was positively affected by soil
305 pH and TN (Fig. 5b). Soil pH, NH_4^+ , and Olsen P explained 83% of the variance in exchangeable K (Fig. 5c). Soil available Fe was negatively affected by soil pH and positively affected by SOC and C:N (Fig. 5d). Soil pH had negative effects on soil available Fe (Fig. 5d), Mn (Fig. 5e) and Zn (Fig. 5f). A positive effect was detected

for C:N on soil available Fe (Fig. 5d) and for NH_4^+ on both available Mn (Fig. 5e) and
310 Zn (Fig. 5f). Soil pH indirectly affected exchangeable Ca and Mg as well as available
Fe and Mn through changing soil C:N (Fig. 5).

4 Discussion

4.1 Elevational patterns of base cations and available micronutrients in soils and 315 relationships with pH and SOC

Contrary to our first hypothesis, no consistent elevational patterns were detected for
soil exchangeable base cations and available micronutrients under either trees or
shrubs. Inconsistent elevational patterns of soil base cations and available
micronutrients indicated that plant uptake of these nutrients did not necessarily
320 decrease more than nutrient supply at higher elevation due to more open canopies.
Our results suggest that soil physiochemical parameters were the dominant
contributors and more important than environmental gradients affecting elevational
patterns of soil exchangeable base cations and available micronutrients (Fig. 5 & Fig.
S2). For instance, soil available Fe, Mn and Zn followed patterns of SOC under trees
325 along the elevational gradient at subtropical Mt. (Table S2), while for shrubs at
subtropical Mt., soil pH, instead of SOC, regulated elevational patterns of soil
available Fe, Mn and Zn (Table S2). Our findings are consistent with a vast amount of
previous studies confirming the pivotal role of soil pH and SOC concentration in
determining soil base cation and micronutrient availabilities (Sharma et al., 2004; Lü
330 et al., 2016; Wang et al., 2017). However, the fundamental roles of SOC and soil pH

in controlling soil base cation and micronutrient availabilities, was not universal as suggested by the relatively weak relationships of soil pH and SOC with soil base cations under tree canopies at wet-temperate Mt. and with micronutrients under both tree and shrub canopies at dry-temperate Mt. (Table S2). This could indicate species- and life form-specific effects on soil base cation and micronutrient availabilities.

Other soil parameters, such as C:N and extractable NO_3^- and NH_4^+ also influenced the availability of base cations (Table 2 & Fig. 5). The soil C:N ratio serves as an indicator of SOM decomposition status where more decomposed SOM possesses a lower C:N ratio (Sollins et al., 2009) and a higher content of negatively charged functional groups (i.e. phenolic, carboxyl, and hydroxyl groups) (Haberhauer et al., 1998). In this study, negative correlations between soil C:N and base cations (Table 2) suggest that more decomposed SOM is beneficial for the retention of soil base cations. Furthermore, soil with a higher level of extractable NO_3^- predisposes cations to leach accompanied by loss of NO_3^- (Cremer and Prietzel, 2017). Therefore, significant negative correlations were detected between soil NO_3^- and base cations in this study (Table 2). Soil extractable NH_4^+ was also negatively correlated with exchangeable Ca and Mg, possibly because NH_4^+ can exchange with base cations on surface soil colloids into soil solution thereby enhancing their loss (Wang et al., 2015; Cusack et al., 2016).

A negative correlation between soil pH and soil available micronutrients (Table 2 & Fig. 5) might be due to precipitation of micronutrient cations at higher soil pH (Rengel, 2007). Indeed, solubility of micronutrients was suggested to decrease from

100-fold (for Mn and Zn) to 1000-fold (for Fe) with one-unit increase of soil pH (Rengel, 2001). Soil organic matter plays an important role in micronutrient retention due to its negative charge (He et al., 2005; Wang et al., 2015, 2017). This may be a reason for the positive relationships between SOC and micronutrients (although not always significant, Table 2). While no general patterns were found for distribution of micronutrients under both tree and shrub canopies with elevation, our results suggest that the determinants of soil micronutrient availabilities were predominantly soil pH and SOC concentration, which are reflections of long-term climatic conditions, plant-soil interactions, and biogeochemical processes (Sinsabaugh et al., 2008).

4.2 Elevational patterns of base cations and micronutrients in plants and plant-soil system

In contrast to our second hypothesis, both trees and shrubs at higher elevation did not necessarily contain higher base cation and micronutrient concentrations in their tissues. No general patterns were found for base cations and micronutrients in both trees and shrubs along elevational gradients across the three sites (Fig. 3 & Fig. 4). Even normalizing the data to per unit concentration of soil available nutrients, there were still no consistent elevational patterns for both base cations and micronutrients in plant tissues (Fig. S3). This suggests that base cation and micronutrient concentrations in plants are influenced by other factors besides elevation-induced changes in temperature, precipitation, specific nutrient absorption characteristics of different tissues, soil base cation and micronutrient availabilities and other edaphic properties

375 (Campo-Alves, 2003; Richardson, 2004). Another explanation could be that initial differences in soil properties (*e.g.*, parent material) among climate zones were larger than the effects of elevation. Soil base cation and micronutrient availabilities were an important factor influencing their concentrations in plant tissues across all plant species and sampling sites (Table S3). Similar results were found for macronutrients 380 (i.e. nitrogen and phosphorus) suggesting that “plants are what they root in” (Elser et al., 2010; Han et al., 2014). However, plant nutrients did not covary with soil nutrients along a 2200 km-long climatic gradient in grasslands of northern China (Luo et al., 2015, 2016). The discrepancy of our study with Luo et al. (2015, 2016) might be driven by different ecosystem types (forest vs. grassland), dominant climatic factor 385 gradients (temperature vs. precipitation), and different soil properties. The studies of Luo et al. (2015, 2016) were conducted in grassland ecosystems where precipitation played an essential role in nutrient concentrations in plant-soil systems. Moreover, base cation and micronutrient cycling processes are likely to be different between high-organic and fine-grained forest soils in our study versus low-organic and sandy 390 grassland soils in Luo et al. (2016). Inconsistently elevational patterns in plant nutrient concentrations could also be derived from the fact that individual plant species reinforced patterns of soil nutrient availabilities in their vicinity causing a positive feedback between plant and soil (Hobbie, 1992).

The topic of base cation and micronutrient translocation in intact plant is 395 important as it deals with the movement of micronutrients from root to the leaves for physiological activities, such as photosynthesis (Welch and Shuman, 1995). Also, it is

an important process in determining plant chemical composition and subsequently litter quality, litter decomposition and nutrient release (Sun et al., 2016). Given earlier findings that transport of base cations from roots to the leaves in woody plants is slow
400 (van der Heijden et al., 2015), we found no significant correlation for both Mg and K between roots and leaves (Table S3). However, negative relationships of stem Mg vs. leaf Mg and stem K vs. leaf K suggest that the plant internal pools of base cations could act as sources of base cation supply for leaves (Weatherall et al., 2006).

Translocation of base cations within plant tissues is one of the main physiological
405 mechanisms buffering low nutrient availabilities in soils (van der Heijden et al., 2015). For instance, supplementation of Mg is a critical process to maintain photosynthesis in forests growing on acid and cation poor soils (Verbruggen and Hermans, 2013). In support of this, we found significantly positive correlations between Mg and soluble sugar concentrations (one of the main photosynthates) in leaves across the three sites
410 (Fig. S4a), while relationships were more pronounced at wet-temperate Mt. (Fig.S4b) where soil pH and exchangeable Mg was the lowest (Fig. 1a, 2b).

Unlike Ca, Mg, Mn and Zn, the concentrations of K and Fe in plant leaves decoupled with their availabilities in the soil (Table S3), which may suggest that not only availability of these nutrients in soils affect their leaf concentrations, but that also
415 other environmental factors (*e.g.*, temperature) played more important roles in affecting plant nutrition (van den Driessche, 1974). We do not know why this decoupling only occurred for K and Fe, but possibly factors such as temperature constrained soil microbial activity and plant metabolism (Körner and Paulsen, 2004)

and subsequently uptake of these nutrients by plants. On the other hand, plants often
420 increase nutrient uptake to compensate for decreased metabolism at low temperature
(Reich and Oleksyn, 2004). Thus, these opposite effects of temperature on K and Fe
concentrations in plant tissues may have obscured their relationships with K and Fe
availability in the soil along the elevational gradients. While plant nutrient
concentrations were mainly influenced by nutrient availabilities in the soil and by
425 plant-internal translocation processes, we found no consistent evidence that plants
accumulate more base cations and micronutrients in their tissues to better adapt to
cold environments at higher elevation.

5 Conclusions

430 We did not find consistent elevational patterns of base cations and micronutrients in
plant-soil systems along three different elevation transects up to the alpine treeline
and shrubline in different regions of China. Rather, our results highlight the essential
roles of specific edaphic properties of soil pH, SOC and extractable nitrate and
ammonium in regulating soil base cation and micronutrient availabilities across
435 climate zones. Soil available base cations and micronutrients were mostly positively
correlated with concentrations of base cations and micronutrients in plant tissues,
except for K and Fe. Our results suggest that base cation and micronutrient
concentrations in plants (trees and shrubs) growing at their upper limits are largely
controlled by their availabilities in the soil rather than by plant adaptations to cold
440 environment at higher elevations.

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Data Availability

Data sets for this paper can be obtained via personal communication.

450 Author contributions

Mai-He Li designed the study; Xue Wang, Heyong Liu, Jinfei Yin and Zhan Shi did the plant and soil measurements; Ruzhen Wang and Xue Wang analyzed the data; Ruzhen Wang wrote the manuscript; Mai-He Li, Feike A. Dijkstra and Artemi Cerdà revised the manuscript; Mai-He Li and Yong Jiang provided funding and laboratory facilities for this study.

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Tables

Table 1 Effects (*F* values) of plant life form (L, tree or shrub), elevation position (E), and their interactions on soil pH, soil organic carbon (SOC), base cations and micronutrients in soils (exchangeable/available form) and plants (total) across sampling sites.

	soil pH	SOC	soil Ca	soil Mg	Soil K	leaf Ca	root Ca	shoot Ca	stem Ca	leaf Mg	root Mg	shoot Mg	stem Mg	leaf K	root K	shoot K	stem K	soil Fe	soil Mn	soil Zn	leaf Fe	root Fe	shoot Fe	stem Fe	leaf Mn	root Mn	shoot Mn	stem Mn	leaf Zn	root Zn	shoot Zn	stem Zn
L	19.6**	0.96	4.72*	5.85*	35.8**	3.05	14.7**	20.20**	21.9**	16.5**	0.09	17.1**	42.6**	5.98*	1.42	26.4**	28.9**	2.65	0.10	4.71*	15.8**	2.84	4.44*	5.41*	0.11	2.41	7.45**	18.0**	3.03	5.10*	2.63	14.5**
E	5.37**	12.0**	0.25	0.61	21.4**	9.10**	6.98**	14.5**	0.00	1.66	0.49	6.82**	3.86*	8.08**	0.67	2.16	3.36*	3.22*	0.15	2.07	0.2	5.08**	3.46*	0.66	4.45*	0.91	0.14	1.72	1.47	0.16	1.31	0.40
L×E	15.6**	3.33*	2.40	0.55	8.61**	2.24	8.55**	6.61**	1.11	1.82	7.20**	5.23**	0.37	0.95	1.01	2.20	2.84	0.05	1.69	4.42*	0.36	3.40*	3.72*	0.61	0.15	2.81	1.22	2.37	0.90	1.93	1.34	2.12

650

* and ** indicate significant level at $P < 0.05$ and 0.01 , respectively.

Table 2 Regression statistics relating soil base cations and micronutrients to other soil physicochemical parameters under all trees and shrubs across three sampling sites.

The TN, C:N, TIN, and Olsen P represent soil total nitrogen, SOC to TN ratio, total inorganic nitrogen, and Olsen phosphorus.

655

	Ca	Mg	K	Fe	Mn	Zn
pH	0.94**	0.87**	-0.06	-0.64**	-0.36**	-0.73**
SOC	0.12	0.24*	0.18	0.30**	0.07	0.37**
TN	0.39**	0.51**	0.18	-0.01	0.03	0.11
C:N	-0.45**	-0.44**	-0.11	0.60**	0.02	0.44**
NO ₃ ⁻	-0.24*	-0.29**	0.28**	-0.05	0.02	0.08
NH ₄ ⁺	-0.38**	-0.33**	-0.10	0.46**	0.15	0.63**
TIN	-0.38**	-0.39**	0.19*	0.17	0.08	0.36**
Olsen P	0.10	0.18	0.49**	-0.05	0.04	0.03

* and ** indicate significant level at $P < 0.05$ and 0.01 , respectively.

Figures

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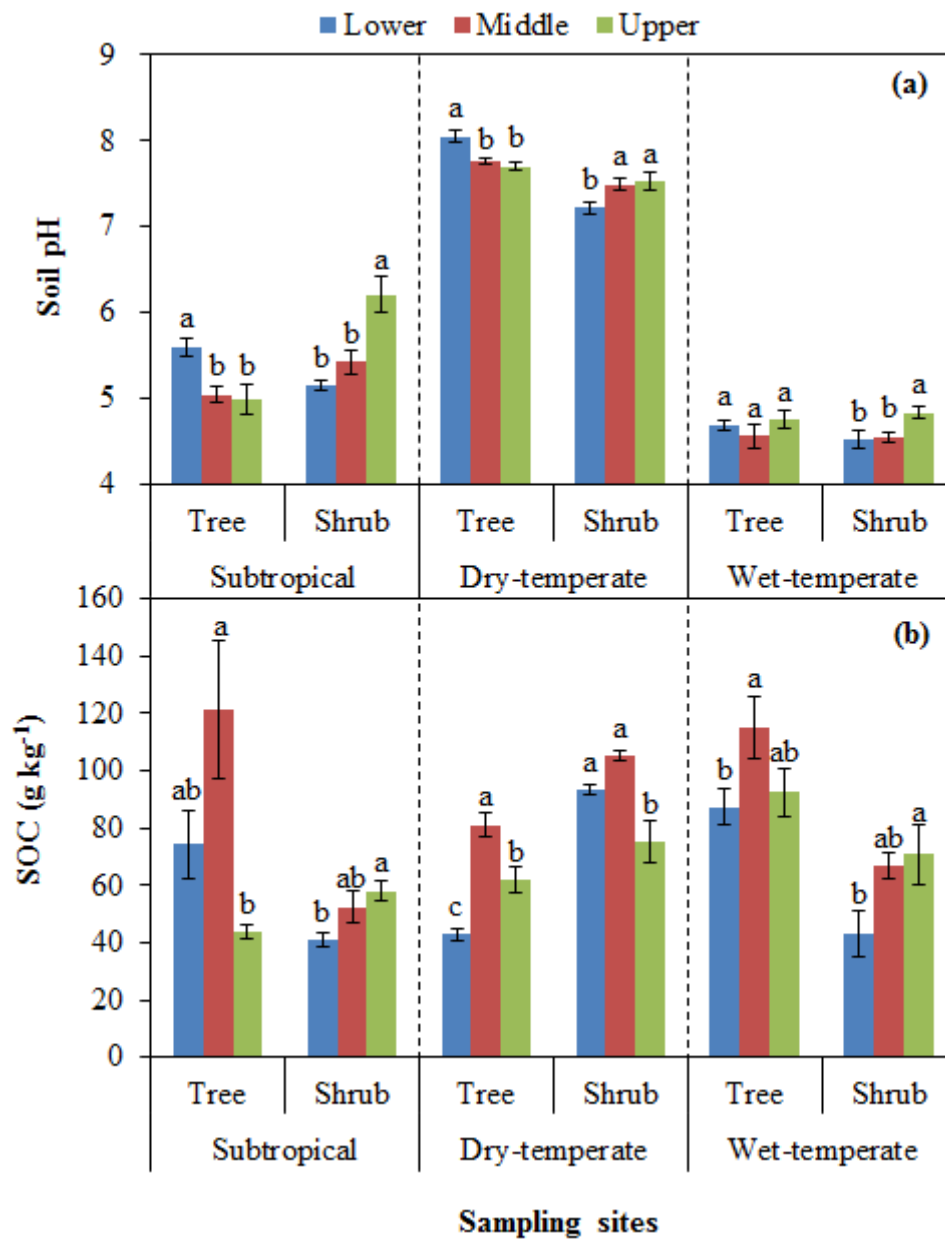


Fig. 1 Soil pH values (a) and soil organic carbon (SOC) concentration (b) at lower and middle elevations as well as at the upper limit of treelines or shrublines for each of the three sites. Different letters indicate significant differences among three elevations within treeline or shrubline for each site.

665

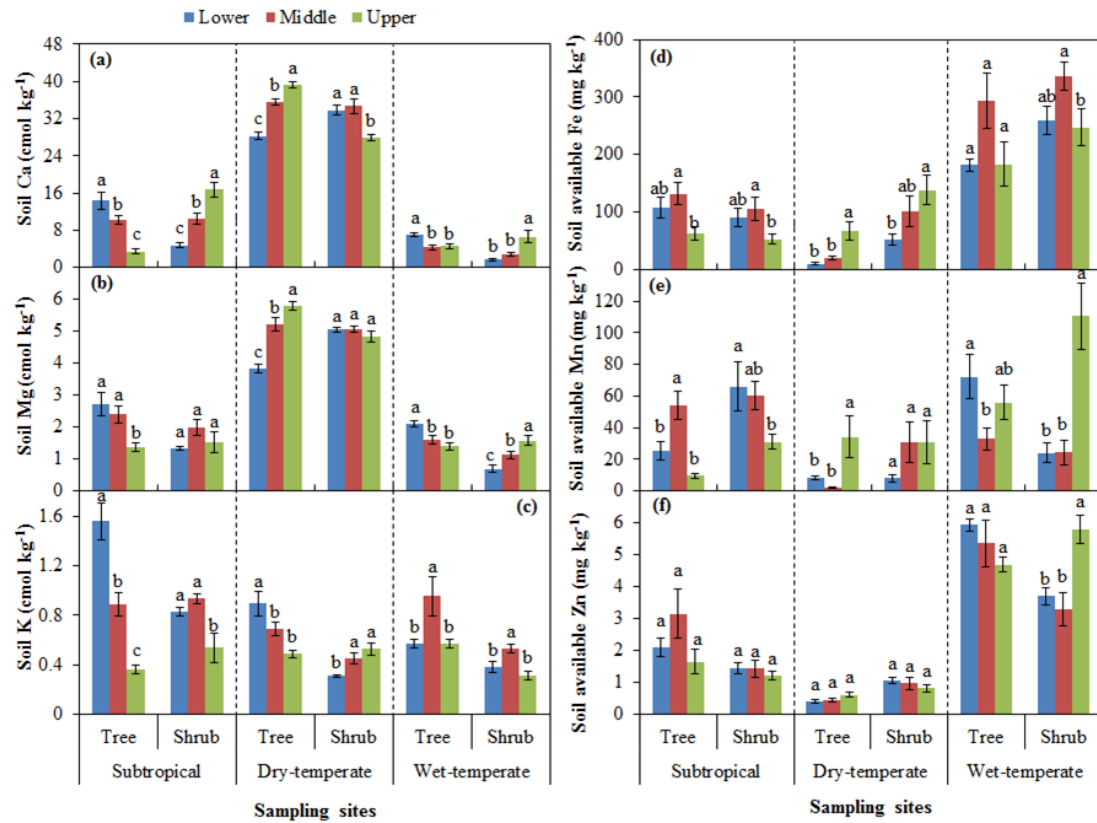
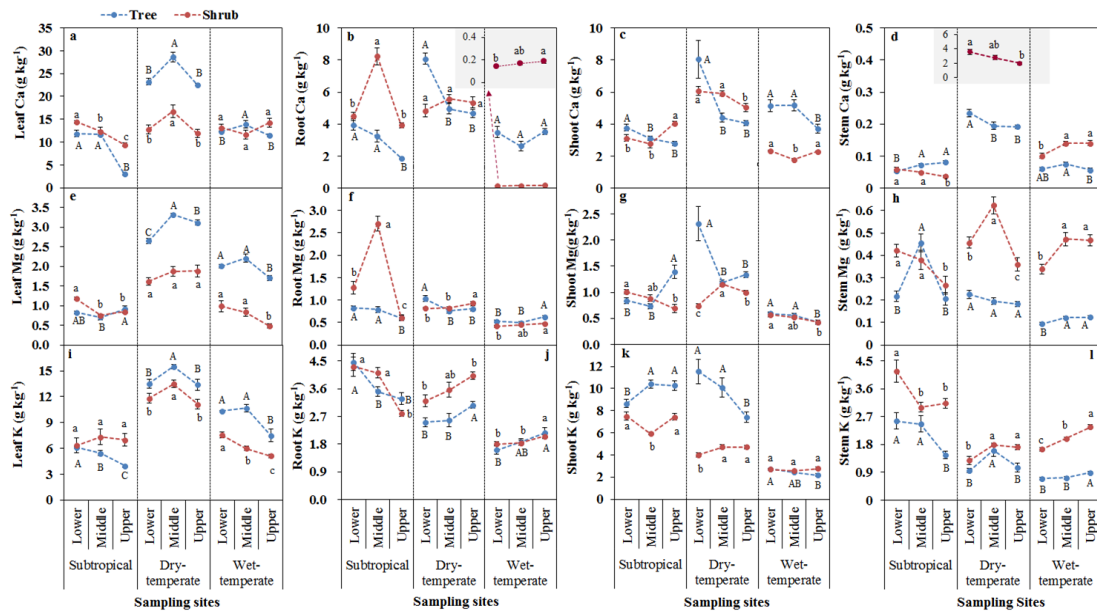


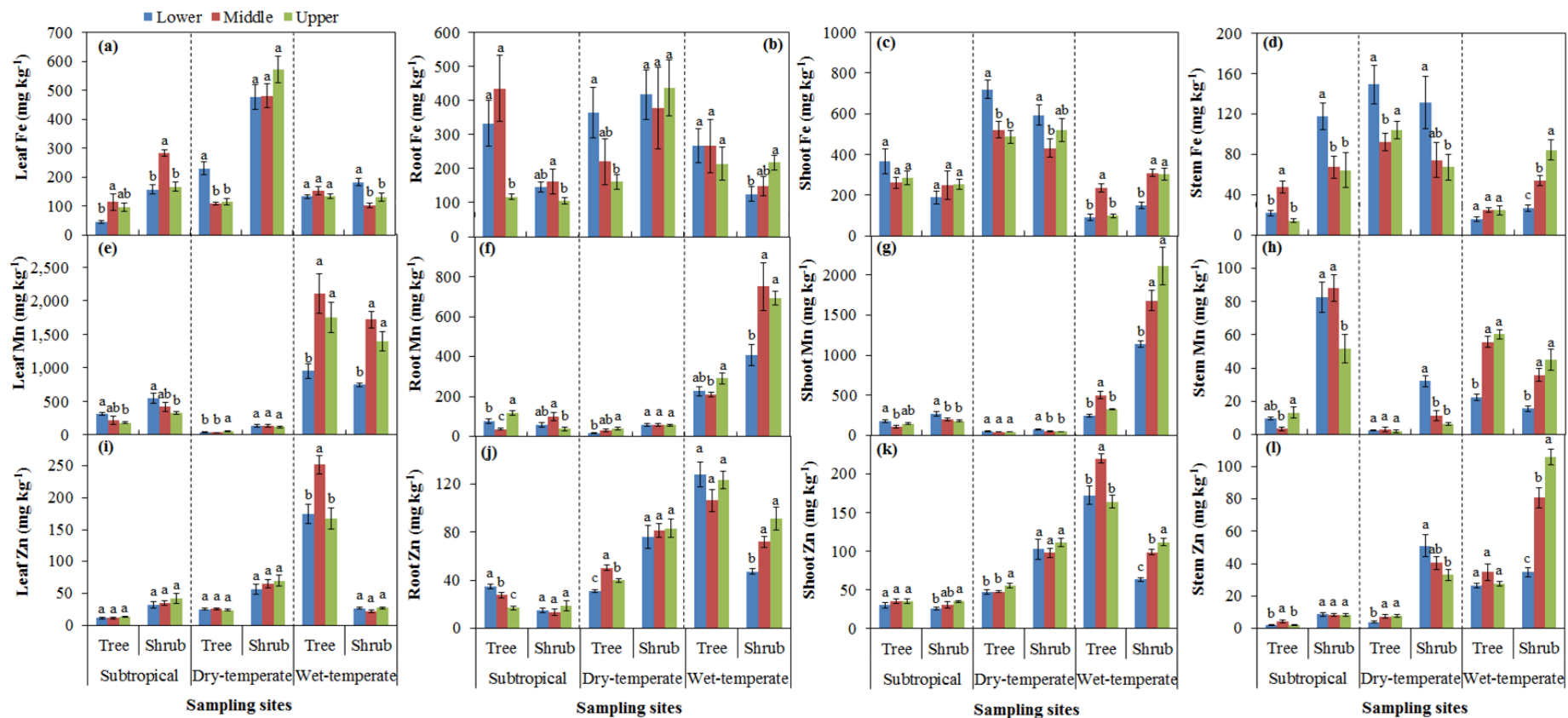
Fig. 2 Concentrations of soil exchangeable base cations of Ca (a), Mg (b) and K (c) and available micronutrients of Fe (d), Mn (e) and Zn (f) at lower and middle

670 elevations as well as at the upper limit of trees or shrubs for each of the three sites.

Different letters indicate significant differences among three elevations within treeline or shrubline for each site.

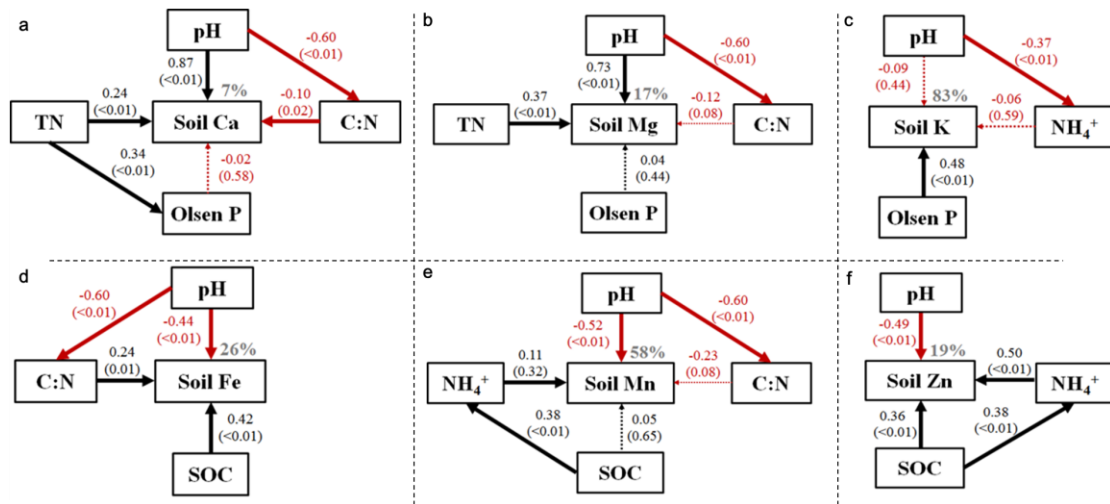


675 **Fig. 3** Base cation concentrations of Ca (a, b, c, d), Mg (e, f, g, h) and K (I, j, k, l) in plant tissues of leaf, root, shoot and stem sapwood at lower and middle elevations as well as at the upper limit of trees or shrubs for each of the three sites. Different letters indicate significant differences among three elevations within treeline or shrubline for each site.



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Fig. 4 Micronutrient concentrations of Fe (a, b, c, d), Mn (e, f, g, h) and Zn (I, j, k, l) in plant tissues of leaf, root, shoot and stem sapwood at lower and middle elevations as well as at the upper limit of trees or shrubs for each of the three sites. Different letters indicate significant differences among three elevations within treeline or shrubline for each site.



685 **Fig. 5** The results of structural equation modelling of the effect of soil parameters on soil exchangeable base cations of Ca, Mg and K, and available micronutrients of Fe, Mn and Cu. Arrows indicate positive (black) and negative (red) effects. Solid and dotted lines represent significant and non-significant relationships, respectively. The number adjacent to each arrow is the standardized path coefficient with corresponding

690 *p* value between brackets.