



1 **C₅ glycolipids of heterocystous cyanobacteria track symbiont abundance in the diatom *Hemiaulus hauckii***
2 **across the tropical north Atlantic**

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11 **Abstract.** Diatom-diazotroph associations (DDAs) include marine heterocystous cyanobacteria found as exo-
12 and endosymbionts in multiple diatom species. Heterocysts are the site of N₂ fixation and have a thickened cell
13 walls containing unique heterocyst glycolipids which maintain a low oxygen environment within the heterocyst.
14 The endosymbiotic cyanobacteria *Richelia intracellularis* found in species of the diatom genus *Hemiaulus* and
15 *Rhizosolenia* makes heterocyst glycolipids (HGs) containing pentose (C₅) moieties that are distinct from limnetic
16 cyanobacterial HGs with C₆ moieties. Here we applied a method for analysis of intact polar lipids (IPLs) to the
17 study of HGs in suspended particulate matter (SPM) and surface sediment from across the tropical North
18 Atlantic. The study focused on the Amazon plume region where DDAs are documented to form extensive
19 surface blooms in order to examine the utility of C₅ HGs as markers for DDAs as well as their transportation to
20 underlying sediments. C₅ HGs were detected in both marine SPM and surface sediments. We found a significant
21 correlation between the water column concentration of C₅ HGs and DDA symbiont counts. In particular, the
22 concentrations of both the C₅ HGs (1-(O-ribose)-3,27,29-triacontanetriol (C₅ HG₃₀ triol) and 1-(O-ribose)-
23 3,29,31-dotriacontanetriol (C₅ HG₃₂ triol)) in SPM exhibited a significant correlation with the number of
24 *Hemiaulus hauckii* symbionts. This result strengthens the idea that C₅ HGs can be applied as biomarkers for
25 marine endosymbiotic heterocystous cyanobacteria. The presence of the C₅ HGs in surface sediment provides
26 evidence that they are effectively transported to the sediment and hence have potential as biomarkers for studies
27 of the contribution of DDAs to the paleo-marine N-cycle.

28 **1 Introduction**

29 Cyanobacteria are cosmopolitan oxygenic photoautotrophs that play an important role in the global carbon and
30 nitrogen cycles. Marine cyanobacteria are the major fixers of dinitrogen (N₂) in modern tropical and subtropical
31 oligotrophic oceans (Karl et al., 1997; Lee et al., 2002). Because N₂ fixation is sensitive to oxygen, cyanobacteria
32 have evolved a range of different strategies in order to combine the incompatible processes of oxygenic
33 photosynthesis and N₂ fixation. One strategy, found only in filamentous cyanobacteria, is to fix N₂ in
34 differentiated cells known as heterocysts (Wolk, 1973; Rippka et al., 1979). Free-living heterocystous
35 cyanobacteria are rare in the open ocean (Staal et al., 2003); however, heterocystous taxa are abundant as both
36 exo- and endosymbionts in diatoms (Foster et al., 2011; Gómez et al., 2005; Luo et al., 2012; Villareal, 1991;



37 Villareal et al., 2011, 2012). These diatom-diazotroph associations (DDAs) can fully support the nitrogen (N)
38 requirements of both host and symbiont (Foster et al., 2011; Villareal, 1990) which explains the presence of
39 these symbioses in oligotrophic offshore environments such as the North Pacific gyre (Venrick, 1974). In the
40 western tropical north Atlantic Ocean, these symbiotic associations produce nearly 70% of total N demand in the
41 surface waters (Carpenter et al., 1999) as non-symbiotic diatom blooms deplete N in the Amazon River plume
42 and create N-poor conditions with residual P and Si (Subramaniam et al., 2008; Weber et al., 2017).

43 In all non-symbiotic cyanobacteria studied to date, the heterocyst cell walls contain heterocyst
44 glycolipids (HGs) (Abreu-Grobois et al., 1977; Bauersachs et al., 2009a, 2014; Gambacorta et al., 1995; Nichols
45 and Wood, 1968). These HGs comprise a hexose head group (C₆) glycosidically bound to long chain diols, triols,
46 or hydroxyketones (cf. Fig. 1) (Bauersachs et al., 2009b, 2011; Bryce et al., 1972; Gambacorta et al., 1998). In
47 contrast, the endosymbiotic heterocystous cyanobacterium *Richelia intracellularis* (found within the marine
48 diatoms *Hemiaulus hauckii* and *H. membranaceus*; (Villareal, 1991) contained HGs with a pentose sugar head
49 group (C₅) rather than a C₆ sugar (Fig. 1) (Schouten et al., 2013). The structural difference in the glycolipids of
50 marine endosymbiotic heterocystous cyanobacteria compared to the free-living counterparts was hypothesized to
51 be an adaptation to the high intracellular O₂ concentrations within the host diatom (Schouten et al., 2013).

52 In the first study of the C₅ HGs in the natural environment, these compounds were found in suspended
53 particulate material (SPM) and surface sediment from the Amazon plume but not in lake sediments or river SPM
54 (Bale et al., 2015). However, HGs with a C₅ sugar moiety comprising a shorter C₂₆ carbon chain were tentatively
55 identified in a culture of freshwater cyanobacterium *Aphanizomenon ovalisporum* UAM 290 and in suspended
56 particulate matter from three freshwater environments in Spain (Wörmer et al., 2012). Thus, it remains to be
57 demonstrated whether distinctive C₅ HGs are unambiguously associated with DDAs in the marine environment.
58 In addition, the genera *Rhizosolenia*, *Guinardia* and *Hemiaulus* all contain species harboring heterocystous
59 cyanobacteria. DDA taxonomic relationships and host-symbiont specificity are only partially defined (Hilton et
60 al 2014, Foster and Zehr 2006, Janson et al 1999), suggesting additional clarification of how diverse HGs are
61 distributed within DDAs is required.

62 In this study, we applied a novel Ultra High Pressure Liquid Chromatography- High Resolution Mass
63 Spectrometry (UHPLC-HRMS) method to analyze the concentration of HG lipids in SPM from the oligotrophic
64 open Atlantic Ocean to the region affected by the Amazon River plume. We compared lipid concentrations with
65 the number of diazotrophic symbionts to examine the applicability of HGs to trace these organisms. Furthermore,
66 we also analyzed HG lipids in the surface sediment along the transect to examine the transport of these
67 compounds to the geological record and potential for use as a molecular tracer for DDA N₂ fixation.

68 2 Methods

69 2.1 Cruise track and physiochemical parameters

70 Sampling was carried out during a 4 week research cruise (64PE393) onboard the R/V *Pelagia* from 26th August
71 – 21st September 2014. The cruise followed a >5000 km transect and sampling occurred at 23 stations, starting
72 at Cape Verde and finishing at the island of Barbados (Fig. 2). The cruise track began close to the Cape Verde
73 EEZ boundary and proceeded approximately south-westerly across the Atlantic (Fig. 2). Aquarius sea-surface
74 salinity (SSS) satellite data (30 day composite, centered on 01-Sept-14) clearly indicated the influence of the



75 freshwater Amazon discharge in the region, i.e. surface salinity < 33 (Fig. 2). Discrete CTD measurements of
76 salinity (contour lines Fig. 3a, Table 1) generally agreed with the satellite data as to the geographical spread of
77 the Amazon River plume. However, the region was highly dynamic with the plume location shifting hundreds of
78 km over the course of the cruise as noted in the sequential 7-day Aquarius SSS composites (Fig. S1 -
79 Supplemental material).

80 Temperature and salinity were measured using a Sea-Bird SBE911+ conductivity–temperature–depth
81 (CTD) system equipped with a 24 × 12 L Niskin bottles rosette sampler. Fluorescence was measured with a
82 Chelsea Aquatracka MKIII fluorometer. Chlorophyll fluorescence was not calibrated against discrete chlorophyll
83 and is reported as relative fluorescence units (RFU). Seawater samples for dissolved inorganic nutrient analysis
84 were taken from the Niskin bottles in 60 ml high-density polyethylene syringes with a three way valve and
85 filtered over Acrodisc PF syringe filters (0.8/0.2 µm Supor Membrane, PALL Corporation) into pre-rinsed 5 mL
86 polyethylene vial. Dissolved orthophosphate (PO₄) and nitrogen (NO₃, NO₂ and NH₄) were stored in dark at 4°C
87 until analysis onboard (within 18 h) using a QuAatro autoanalyzer (Grasshoff, 1983; Murphy and Riley, 1962).
88 Samples for dissolved reactive silicate (Si) analysis (Strickland and Parsons, 1968) were stored dark at 4°C until
89 analysis using the same system as above upon return to NIOZ. The detection limits were calculated as: PO₄
90 0.004 µmol L⁻¹, NH₄ 0.030 µmol L⁻¹, NO₃+NO₂ 0.005 µmol L⁻¹ and NO₂ 0.002 µmol L⁻¹.

91 2.2 Phytoplankton pigment composition and enumeration of diazotrophs

92 Samples for diazotroph enumeration were collected in polycarbonate bottles of which 500-1170 ml was filtered
93 under gentle vacuum (< 5 psi) through a 10 µm pore-size polycarbonate filter (47 mm diameter). Filters were
94 placed onto 75 X 50 mm glass slides (Corning 2947) and 2-3 drops of non-fluorescent immersion oil (Cargille
95 type DF) placed on the slide. A glass cover slip (45 x 50 mm; Fisherbrand 12-545-14) was placed on the filter
96 sample and allowed to sit while the immersion oil cleared the filter. The sample was subsequently viewed under
97 transmitted light and epi-fluorescence illumination light filter (530-560 nm excitation, 572-648 nm emission;
98 Olympus BX51) for counting/identifying trichomes and host cells as well as photomicrography (Olympus
99 DP70).

100 For phytoplankton pigment analysis, seawater was filtered through 0.7 µm glass fiber GF/F filters (Pall
101 Corporation, Washington). The filters were extracted in 4 mL 100% methanol buffered with 0.5 mol L⁻¹
102 ammonium acetate, homogenized for 15 s, and analyzed by high performance liquid chromatography (HPLC).
103 The relative abundances of the different taxonomic groups were determined using CHEMTAX (Mackey et al.,
104 1996; Riegman and Kraay, 2001).

105 2.3 SPM and surface sediment collection

106 Three McLane *in situ* pumps (McLane Laboratories Inc., Falmouth) were used to collect suspended particulate
107 matter (SPM) from the water column for lipid analysis. They were generally deployed at three depths: the
108 surface (3 - 5 m), the bottom wind mixed layer (BWML) and the deep chlorophyll-*a* maximum (DCM), with
109 some additional sampling at 200 m (Table 1). They pumped between 90 and 380 L with a cut-off at a pre-
110 programmed pressure threshold and the SPM was collected on pre-ashed 0.7 µm, 142 mm, GF/F filters (Pall
111 Corporation, Washington) and immediately frozen at -80°C. At Sta., as part of a different study (Besseling et al



112 in prep), 12 additional sampling points were carried out to produce a high resolution depth profile (Table 2)
113 where the SPM was collected on pre-ashed 0.3 μm GF75 filters (Avantec, Japan).

114 Sediment was collected at each station in 10 cm diameter, 60 cm length multicores. For sediment
115 sampling site, triplicate cores were collected, always from a single multicore deployment (with a maximum of 60
116 cm between core centers). The cores were sliced into 1 cm slices using a hydraulic slicer and each slice was
117 stored separately in a geochemical bag and immediately frozen at -80°C . For this study we analyzed the 0–1 cm
118 (surface sediment) slice. For TOC analysis, sediment was freeze dried and analysis was carried out using a Flash
119 2000 series Elemental Analyzer (Thermo Scientific) equipped with a TCD detector.

120 **2.4 Lipid extraction**

121 The extraction of lipids from freeze dried filtered seawater or sediment samples was carried out using a modified
122 Bligh-Dyer extraction (Bale et al., 2013). The samples were extracted in an ultrasonic bath for 10 min with 5 –
123 20 ml of single-phase solvent mixture of methanol (MeOH): dichloromethane (DCM): phosphate buffer (2:1:0.8,
124 v:v:v). After centrifugation ($1000 \times g$ for 5 min, room temperature, Froilabo Firlabo SW12 with swing out rotor)
125 to separate the solvent extract and residue, the solvent mixture was collected in a separate flask. This was
126 repeated three times before DCM and phosphate buffer were added to the single-phase extract to induce phase
127 separation, producing a new ratio of MeOH:DCM:phosphate buffer (1:1:0.9 v:v:v). After centrifugation ($1000 \times$
128 g for 5 min), the DCM phase was collected in a glass round-bottom flask and the remaining MeOH:phosphate
129 buffer phase was washed two additional times with DCM. Rotary evaporation was used to reduce the combined
130 DCM phase before it was evaporated to dryness under a stream of N_2 .

131 **2.5 Analysis of intact polar lipids**

132 Whereas previous studies of heterocyst glycolipids have applied high performance liquid chromatography
133 multiple reaction monitoring (MRM) mass spectrometry (HPLC– MS^2) method (e.g., Bale et al. (2015)), in this
134 study we used an Ultra High Pressure Liquid Chromatography-High Resolution Mass Spectrometry (UHPLC-
135 HRMS) method, designed for the analysis of a wide range of intact polar lipids (Moore et al., 2013). The
136 UHPLC-HRMS method was adapted by replacement of hexane with heptane as the non-polar solvent in the
137 eluent, to reduce the toxic nature of hexane relative to heptane in terms of a work place health hazard (Buddrick
138 et al., 2013; Carelli et al., 2007; Daughtrey et al., 1999). Our UHPLC-HRMS method was as follows: we used an
139 Ultimate 3000 RS UHPLC, equipped with thermostatted auto-injector and column oven, coupled to a Q Exactive
140 Orbitrap MS with Ion Max source with heated electrospray ionization (HESI) probe (Thermo Fisher Scientific,
141 Waltham, MA). Separation was achieved on an Acquity UPLC BEH HILIC column (150 x 2.0 mm, 2.1 μm
142 particles, pore size 12 nm; Waters, Milford, MA) maintained at 30 $^{\circ}\text{C}$. Elution was achieved with (A) heptane-
143 propanol-formic acid-14.8 mol L^{-1} aqueous NH_3 (79:20:0.12:0.04, v/v/v/v) and (B) propanol water-formic acid-
144 14.8 mol L^{-1} aqueous NH_3 (88:10:0.12:0.04, v/v/v/v) starting at 100% A, followed by a linear increase to 30% B
145 at 20 min, followed by a 15 min hold, and a further increase to 60% B at 50 min. Flow rate was 0.2 ml min^{-1} ,
146 total run time was 70 min, followed by a 20 min re-equilibration period. Positive ion ESI settings were: capillary
147 temperature, 275 $^{\circ}\text{C}$; sheath gas (N_2) pressure, 35 arbitrary units (AU); auxiliary gas (N_2) pressure, 10 AU; spray
148 voltage, 4.0 kV; probe heater temperature, 275 $^{\circ}\text{C}$; S-lens 50 V. Target lipids were analyzed with a mass range of
149 m/z 350–2000 (resolution 70,000 ppm), followed by data-dependent tandem MS^2 (resolution 17,500 ppm), in



150 which the ten most abundant masses in the mass spectrum were fragmented successively (normalized collision
151 energy, 35; isolation width, 1.0 m/z). The Q Exactive was calibrated within a mass accuracy range of 1 ppm
152 using the Thermo Scientific Pierce LTQ Velos ESI Positive Ion Calibration Solution. During analysis dynamic
153 exclusion was used to temporarily exclude masses (for 6 s) in order to allow selection of less abundant ions for
154 MS². In addition, an inclusion list (within 3 ppm) was used, containing all known HGs, in order to obtain
155 confirmatory fragment spectra.

156 Before analysis, the extracts were re-dissolved in a mixture of heptane, isopropanol and water (72:27:1,
157 v:v:v) which contained two internal standards (IS), a platelet-activating factor (PAF) standard (1-O-hexadecyl-2-
158 acetyl-sn-glycero-3-phosphocholine, 5 ng on column) and a short-chain glycolipid standard, n-dodecyl- β -D-
159 glucopyranoside ($\geq 98\%$ Sigma-Aldrich, 20 ng on column; cf. Bale et al. (2017)). The samples were then filtered
160 through 0.45 μm mesh True Regenerated Cellulose syringe filters (4 mm diameter; Grace Alltech). The injection
161 volume was each sample was 10 μl . For quantification the relative response factor (RRF) between the n-dodecyl-
162 β -D-glucopyranoside IS and an isolated C₆ HG (1-(O-hexose)-3,25-hexacosanediol (Bale et al., 2017) was
163 determined to be 6.63.

164 The 12 samples collected at Sta. 10 (0.3 μm GF75 filters, Table 2) were analyzed on the same UHPLC-
165 HRMS system, but with hexane instead of heptane in the mobile phase. Also, the n-dodecyl- β -D-
166 glucopyranoside IS was not added, so quantification was based the PAF IS and correcting for the RRF between
167 the n-dodecyl- β -D-glucopyranoside IS and the PAF IS.

168 2.6 Statistical analysis

169 T-tests and Pearson correlations were determined using Sigmaplot software (version 13.0).

170 3. Results

171 3.1 Physicochemical conditions and phytoplankton assemblage

172 Stations 1-6, 12 and 22 correspond to oceanic stations (SSS > 35, following the convention of Subramaniam et
173 al. (2008)), with Sta. 7-11 in the intermediate salinity range (30 – 35). Originally termed mesohaline
174 (Subramaniam et al., 2008), we use ‘intermediate salinity’ to avoid confusion with the older use of mesohaline in
175 coastal systems to refer to 5-18 waters (Elliott and McLusky, 2002). Only Sta. 11, with a SSS of 29.2, was in the
176 low salinity range defined by Subramaniam et al. (2008). Temperature was uniformly high across the cruise track
177 (>27°C in the euphotic zone) with the 25° C isotherm deepening along the cruise track (Fig. 3a). Stations 1-6
178 exhibited typical tropical open-ocean conditions with an average 36.3 ± 0.2 and depleted surface (3 – 5 m)
179 inorganic nutrient concentrations (on average $0.01 \pm 0.01 \mu\text{mol L}^{-1} \text{PO}_4$, $0.02 \pm 0.20 \mu\text{mol L}^{-1} \text{NO}_3, \text{NO}_2$ and 0.86
180 $\pm 0.09 \mu\text{mol L}^{-1} \text{Si}$ (Fig. 3c,d). Surface NO_3, NO_2 and PO_4 concentration remained low at Sta. 7 – 11 (Fig. 3c),
181 although PO_4 increased slightly to $0.07 \mu\text{mol L}^{-1}$. Si concentrations increased > 10-fold at Sta. 7-11 and was on
182 average of $12.1 \pm 4.4 \mu\text{mol L}^{-1}$ (Fig. 3d). Sta. 12 was shallow and close to the coast but was just north of the
183 point of plume retroreflection (cf. Fig. 2), as evidenced by an increased SSS (35.4) and relatively lower Si
184 concentration ($3.14 \mu\text{mol L}^{-1}$). The deep chlorophyll (Chl) maximum (DCM; cf. maxima in chl fluorescence
185 (Fig. 3b) was associated with the nutricline (cf. Fig. 3c) over most of the transect, with the highest DCM
186 fluorescence at the oceanic stations Sta. 1 and 2. There was a secondary surface Chl fluorescence maximum at



187 Sta. 11 which was the most nearshore, lowest salinity station. Just north of the plume, Station 12 displayed a
188 more mixed water column profile with uniformly elevated Chl fluorescence to ~75 m (Fig. 3b).

189 From the coastal shelf of French Guiana (Sta. 11 and 12), the cruise progressed in a northerly direction
190 towards the Caribbean. The Amazon River influence was again evident after Sta. 13, but decreased with
191 distance, with SSS ranging from 32.8 at Sta. 13 to a maximum of 35.6 at Sta. 22. Surface $\text{NO}_{3+\text{NO}_2}$ remained
192 low through Sta. 13 – 23 (on average $0.01 \pm 0.00 \mu\text{mol L}^{-1}$), while PO_4 was variable but generally decreased to
193 open ocean levels (from $0.01 \mu\text{mol L}^{-1}$ at Sta. 13 to below the limit of detection at Sta. 23). Si dropped from 10.4
194 $\mu\text{mol L}^{-1}$ at Sta. 13 to $3.94 \mu\text{mol L}^{-1}$ at Sta. 23.

195 The phytoplankton pigment composition analysis at Sta. 1-6 was dominated by the cyanobacteria
196 *Prochlorococcus* which made up around 50% of total Chl a in the surface waters (Table S1). At Sta. 7-10 and
197 18-23, the phycoerythrin-containing cyanobacteria (e.g. *Synechococcus*) dominated the phytoplankton
198 community. In general, at Sta. 7-10 the share of Chrysophytes and Prymnesiophyceae pigments was relatively
199 larger. The share of Chrysophyceae was particularly large at the DCM, even dominating the phytoplankton
200 community biomass at Sta. 15-23 (Table S1). Diatoms (Bacillariophyceae) contributed substantially in the
201 surface waters of Sta. 8, up to 21% of total Chl a.

202 3.2 Diazotroph enumeration

203 The diazotroph cyanobacteria were divided into 5 categories: three of them are symbionts, i.e. with the diatoms
204 *Rhizosolenia cf. clevei*, *Hemiaulus hauckii*, and *Guinardia cylindrus* DDAs, and two are non-symbionts, i.e.
205 *Trichodesmium* colonies (>10 trichomes organized into a coherent structure), and free *Trichodesmium* trichomes
206 (Fig. 4a-e, Table S3). Total DDA abundance was low (0-21 combined DDA *Richelia* trichomes L^{-1}) at Sta. 1-6.
207 *Hemiaulus* DDA abundance was greatest at Sta. 8 (ca. 4.0×10^3 trichomes L^{-1}) with a secondary maximum at Sta.
208 17 (0.8×10^3 trichomes L^{-1}), both in the surface (<5 m) waters. *Rhizosolenia* DDA abundance was lower than
209 *Hemiaulus* DDA abundance at Sta. 7 (*Rhizosolenia* DDA ca. 60 trichomes L^{-1}) and at Sta. 15 and 16
210 (*Rhizosolenia* DDA, ca. 80 trichomes L^{-1}). *Rhizosolenia* DDAs were not observed below 31.6 salinity (Fig. 4a).
211 *Hemiaulus* DDA were observed down to 27.1-27.6 salinity at ~80-100 trichomes L^{-1} (Fig. 4b). Free
212 *Trichodesmium* trichomes were broadly distributed (Fig. 4d). Free *Trichodesmium* trichomes often occurred
213 across a wide depth range, down to 75 m at Sta. 17. *Trichodesmium* colonies were seen sporadically and with
214 distributions dominated by two sampling points (Sta. 6, 32 m and Sta. 21, 60 m) where colony abundance > 25
215 colonies L^{-1} . A single observation of colonies at depth under the low salinity plume generated contour lines
216 suggesting a generalized presence at depth. However, removal of this observation (Sta. 14, 61 m) removed this
217 trend and resulted in distinct separation of the colony distributions, i.e. two areas of increased biomass associated
218 with salinity gradients at the edge of the river plume.

219 3.3 Heterocyst glycolipids in suspended particulate matter

220 We analyzed heterocyst glycolipids (HGs) in SPM from along the cruise transect collected at the surface,
221 bottom wind mixed layer (BWML) and the DCM. Two C_5 HGs were detected in the SPM, i.e. 1-(O-ribose)-
222 3,27,29-triacontanetriol and 1-(O-ribose)-3,29,31-dotriacontanetriol (C_5 HG₃₀ and C_5 HG₃₂ triol respectively,
223 Fig. 1). C_5 HG₃₂ triol represented on average 98 ± 4 of the combined concentration of the two HGs. Previous
224 studies of C_5 HGs have identified 1-(O-ribose)-3,29-triacontanediol (C_5 HG₃₀ diol, Fig. 1) in both cultures and



225 environmental samples (Bale et al., 2015; Schouten et al., 2013), but these were not seen in the SPM or surface
226 sediment analyzed in this study.

227

228 The concentrations of the two C₅ HGs were highest in the surface waters of Sta. 8 and showed a second local
229 maxima at Sta. 16 (Table 1 and Fig. 4f). The surface concentration of the dominant HG, i.e. C₅ HG₃₀ triol, ranged
230 between 0 and 4800 pg L⁻¹. The range in concentration was 50-fold lower at the DCM (0-200 pg L⁻¹, Table S2).
231 The three samples from 200 m depth showed lowest concentrations, ranging between 20.6 and 127 pg L⁻¹.
232 Overall, the C₅ HG₃₀ triol was consistently present in the higher concentration of the two (Fig. 4f). The minor
233 HG, i.e. C₅ HG₃₂ triol, ranged between 0 – 10% of their combined concentration at the surface and BWML, was
234 between 0 – 5% at the DCM and 0 – 17% at 200 m (cf. Fig. 4f contour lines and Table S2).

235 HGs with a C₆ sugar head group were not detected in any SPM samples with the exception of one
236 sample, taken at Sta. 20a from the DCM (65 m). 1-(O-hexose)-3,25-hexacosanediol (C₆ HG₂₆ diol, Fig. 1) and 1-
237 (O-hexose)-3-keto-25-hexacosanol (C₆ HG₂₆ keto-ol) were confidently identified from their [M+H]⁺ accurate
238 mass (*m/z* 577.4674 and 575.4517 respectively) and their fragmentation patterns, which followed published
239 reports (Bauersachs et al., 2009b). C₆ HG₂₆ diol and C₆ HG₂₆ keto-ol were present at concentrations of 0.3 and
240 0.4 ng L⁻¹ respectively (data not shown), both ~10 times higher than the concentration of the C₅ HG₃₀ triol in this
241 sample (Table 1).

242 At Sta. 10, besides SPM samples collected on 0.7 μm GF/Fs, SPM samples were also collected at
243 depths down to 3000 m using 0.3 μm GF75 filters (Table 2). As with the 0.7 μm SPM samples at station 10, C₅
244 HG₃₀ triol was consistently present in higher concentration than C₅ HG₃₂ triol (which represented on average
245 only 1.4 % ± 0.7 of their combined concentration). The concentrations and depth trends (to 200 m) of the two C₅
246 HGs did not differ between the 0.3 μm and 0.7 μm filter SPM samples (Fig. 5). For both the 0.3 μm samples and
247 the 0.7 μm samples, the combined concentration of C₅ HG₃₀ triol and C₅ HG₃₂ triol was highest at 200 m, 108 pg
248 L⁻¹. In the 0.3 μm samples, both concentrations decreased below 200 m, although both C₅ HGs remained
249 detectable at 3000 m depth.

250 3.4 Heterocyst glycolipids and bulk properties in surface sediment

251 As with the SPM, C₅ HG₃₀ triol and C₅ HG₃₂ triol were detected in the surface sediment of seventeen stations
252 (Table 1). C₅ HG₃₀ triol was also here consistently present in the higher concentration of the two (C₅ HG₃₂ triol
253 represented on average 9.4 % ± 3.0 of their combined concentration). The C₅ HG₃₀ diol was not detected in any
254 surface sediment, alike the SPM samples. HGs with a C₆ sugar head group were also not detected in any surface
255 sediment. In the sediment underlying the high-salinity open ocean stations (1, 3, 5) the combined concentration
256 of the two C₅ HGs was low (2.0 – 3.7 ng g⁻¹, Table 1). It was high at Sta. 7 and 8 (10.6 and 16.3 ng g⁻¹), while
257 Sta. 9 – 17 contained mid-range concentrations (5.2 – 14.8 ng g⁻¹), with the exception of the two coastal-shelf
258 stations (11 and 12) where the concentration was at its lowest (0.2 and 0.3 ng g⁻¹). At the final 4 stations (20a,
259 21a, 22 and 23) the combined concentration returned to high levels (11.2 – 19.0 ng g⁻¹). For context, the TOC
260 was relatively stable between Sta. 1 and 10 (av. 0.6 ± 0.1 %, n=7) then low at Sta. 11 and 12 (av. 0.2 ± 0.1 %).
261 Sta. 13 exhibited the highest TOC of all the stations (1.2 ± 0.0 %), and TOC decreased steadily at all stations
262 thereafter, and was 0.6 ± 0.0 % at Sta. 23.

263 **4. Discussion**264 **4.1 Heterocyst glycolipids and DDAs in the water column**

265 The Amazon plume has been extensively documented to support high numbers of the diatom-diazotroph
266 associations (DDA) such as *Hemiaulus hauckii*-*Richelia intracellularis* and *Rhizosolenia clevei*-*Richelia*
267 *intracellularis* (Carpenter et al., 1999; Foster et al., 2007; Goes et al., 2014; Subramaniam et al., 2008; Weber et
268 al., 2017). Our study took place outside the high Amazon flow period and the Chl concentrations and DDA
269 counts encountered on this cruise did not reach the values seen in ‘bloom conditions’ described during previous
270 studies in the region (Carpenter et al., 1999; Subramaniam et al., 2008). However, the DDA counts in certain
271 stations were up to 3 orders of magnitude higher than surrounding waters and comparable to the open ocean
272 DDA blooms seen in the North Pacific gyre (Villareal et al., 2011, 2012). These strong gradients permitted to
273 investigate relationships between DDA and HG distributions.

274 The concentration of the C₅ HGs were correlated with the cell counts of different diazotrophs. The
275 concentrations of both the C₅ HGs (1-(O-ribose)-3,29,31-dotriacontanetriol (C₅ HG₃₀ triol, Fig. 1) and 1-(O-
276 ribose)-3,27,29-triacontanetriol (C₅ HG₃₂ triol) exhibited the most significant positive Pearson correlation with
277 the number of *Hemiaulus* symbionts ($p \leq 0.001$, $r = 0.79$ and 0.78 respectively, $n=54$). While C₅ heterocyst
278 glycolipids (HG) have been found in cultures of DDAs (Bale et al., 2015; Schouten et al., 2013), our study of
279 the tropical north Atlantic provides to our knowledge for the first time, environmental evidence that C₅ HGs
280 track the abundance and distribution of DDAs.

281 Interestingly, there was no significant correlation found between the number of *Rhizosolenia* symbionts
282 and the concentration of the C₅ HGs (C₅ HG₃₀ triol: $p = 0.07$, $r = 0.23$ and C₅ HG₃₂ triol: $p = 0.14$, $r = 0.19$),
283 except when the surface and BWML of Sta. 8 were excluded from the analysis (C₅ HG₃₀ triol: $p \leq 0.001$, $r =$
284 0.88 ; and C₅ HG₃₂ triol: $p \leq 0.001$, $r = 0.83$). This difference may in part be due to the lower number of
285 *Rhizosolenia*/*Guinardia* symbionts relative to *Hemiaulus* symbionts (on average *Rhizosolenia* symbionts in this
286 study represented 24 ± 34 % of the sum of *Rhizosolenia* and *Hemiaulus* symbionts), similar to previous findings
287 that *Hemiaulus* dominated over *Rhizosolenia* in the Amazon plume (Foster et al., 2007) and Caribbean region
288 (Villareal, 1994). Furthermore, culture studies have shown that *Rhizosolenia* symbionts contain only trace
289 amounts of C₅ HG₃₀ triol, (Bale et al., 2015), whereas this is a dominant HG in *Hemiaulus* symbionts (Schouten
290 et al., 2013). In this study, *Rhizosolenia* DDAs were not observed below salinities of 31.5, while *Hemiaulus*
291 DDAs were observed at salinities of 27.5 (Fig. 6). The plume is highly dynamic and it is unclear whether this is a
292 significant niche separation between the taxa or simply mixing and loss of the less abundant *Rhizosolenia* DDA
293 at the water volumes being counted.

294 Unfortunately, a unique biomarker for *Rhizosolenia* and *Guinardia* symbionts has not been identified to
295 date (Bale et al., 2015; Schouten et al., 2013). There was also a significant correlation between C₅ HG₃₂ triol,
296 (but not C₅ HG₃₀ triol) and the counts of *Guinardia cylindrus* (formerly *Rhizosolenia cylindrus*) ($p \leq 0.03$, $r =$
297 0.49 , $n=21$). This was the only species for which there was a correlation with C₅ HG₃₂ triol but not C₅ HG₃₀ triol.
298 This DDA has not been cultured and nothing is known about the heterocyst lipid composition of this species.
299 These results suggest C₅ HG₃₀ triol may be synthesized by this species.

300 At approximately half of the sampling points, glycolipids could be detected in SPM where no DDAs
301 were observed by microscopy. These sampling points generally contained low combined concentration of the
302 two C₅ HG lipids ($0 - 62.9$ pg L⁻¹, $n=22$), compared to the sampling points where DDAs were detected ($18 -$



303 5300 pg L⁻¹, n=32). These two groups were significantly different from each other (as determined by t-test, $p =$
304 <0.001). This difference may be result of the difference in total sampling volumes between the two methods.
305 Microscopic examinations were carried out using 0.5 - 1.2 L per sample. Although for lipid analysis 90 - 400 L
306 were filtered, the individual analyses on the UHPLC-HRMS system each represented between 0.5 and 5 L of
307 seawater. However, the far greater initial sample volume leads to a higher probability that the lipid samples
308 would contain symbiont chains than the microscopy samples. In addition, microscopic examinations may have
309 missed free heterocysts and heterocysts that were incorporated into unrecognizable masses in aggregates whereas
310 UHPLC-HRMS may have still detected the associated HGs. Indeed, copepod grazing in the plume (Conroy et
311 al., 2016) will repackage *Richelia* trichomes, and little is known of the effects of gut passage on heterocyst and
312 HG integrity. It should also be noted that because sampling for diazotroph enumeration and for lipid analysis
313 occurred via different methods, there was a time offset of ≤ 5 h and a depth offset of ≤ 20 m between the two
314 sampling events representing the same water column phenomena (surface, BWML and DCM)

315 Unexpectedly, a significant correlation was also found for C₅ HG₃₀ triol and C₅ HG₃₂ triol and the
316 number of *Trichodesmium* colonies ($p \leq 0.001$, $r = 0.68$ and 0.67 , $n=54$), and for C₅ HG₃₀ triol and the number of
317 *Trichodesmium* filaments ($p \leq 0.05$, $r = 0.30$, $n=54$). These correlations could be coincidental as C₅ HG
318 producing organisms have not been described in association with *Trichodesmium* nor would *Trichodesmium* be
319 expected to produce HGs itself as it does not use heterocysts to fix nitrogen. A recent study in the North Pacific
320 Subtropical Gyre found that *Trichodesmium* colonies were harboring an endobiontic heterocystous cyanobacteria
321 of the genus *Calothrix* (Momper et al., 2015). However, analyses of the HG content of both freshwater and
322 marine *Calothrix* cultures have to date only revealed the presence of C₆ HGs, not C₅ HGs (Bauersachs et al.,
323 2009a; Schouten et al., 2013; Wörmer et al., 2012). Furthermore, no heterocystous cyanobacteria were observed
324 in *Trichodesmium* from the Caribbean (Borstad, 1978) or southwest Sargasso Sea (Siddiqui et al., 1992).
325 *Trichodesmium* is reported to have a physiological differentiated cell (diazocyte) that permits N₂-fixation in an
326 oxygenated colony or trichome, and which lacks the thickened cell envelope of heterocysts where HGs are
327 localized (Sandh et al., 2012).

328 While elevated HGs were statistically more associated with the DDA blooms than either free or
329 colonial *Trichodesmium*, there was frequently a co-occurrence of *Trichodesmium* with the DDA taxa (Fig. 4)
330 which could also contribute to the unexpected correlation. The *Trichodesmium* distribution appears to contrast
331 with the findings of Foster et al. (2007), Goes et al. (2014) and Subramaniam et al. (2008), who all concluded
332 that changing nutrient availability as reflected in the salinity gradient along the Amazon River plume led to
333 zonation of the diazotroph community. However, their data were examining more pronounced DDA cell
334 abundance concentrations under much higher Amazon plume flow conditions. The broader features of our
335 observations, i.e. a low salinity region with higher nutrient concentrations and few diazotrophs transitioning to
336 strong diazotroph gradients in the salinity gradient to oceanic conditions, are in concordance with their
337 observations.

338 Two C₆ HGs, generally associated with free-living heterocyst forming cyanobacteria from freshwater or
339 brackish environments (Bale et al., 2015, 2016, Bauersachs et al., 2009b, 2010, 2011; Bühring et al., 2014;
340 Wörmer et al., 2012) were identified only in the DCM of Sta. 20a (C₆ HG₂₆ diol and C₆ HG₂₆ keto-ol). Whereas
341 in this study the two C₆ HGs were found at a similar concentration to each other, previous studies have reported
342 that C₆ HG₂₆ keto-ol was detected a minor component relative to the more abundant C₆ HG₂₆ diol (Bale et al.,



343 2015, 2016, Bauersachs et al., 2009a, 2009b, 2011; Schouten et al., 2013; Wörmer et al., 2012). An earlier study
344 executed nearer to the mouth of the Amazon river detected trace levels of C₆ HG₂₆ diol (but not C₆ HG₂₆ keto-ol)
345 in surface sediments (Bale et al., 2015). In contrast, both C₆ HG₂₆ diol and C₆ HG₂₆ keto-ol were recorded in
346 freshwater Amazon River water and floodplain lake sediment.

347 There are reports of cyanobacterial species in cohabitation with other vegetal such as the floating
348 macroalgae *Sargassum* (Carpenter, 1972; Hanson, 1977; Phlips et al., 1986) and *Trichodesmium* (Momper et al.,
349 2015). While the HG content of the cyanobacteria in these co-habitations has not been investigated, these
350 cyanobacteria are in the same families as known C₆ HG producers (Bauersachs et al., 2009a; Schouten et al.,
351 2013; Wörmer et al., 2012). *Trichodesmium* was not detected by microscopy at this sampling point, however as
352 stated above, there is an apparent difference the limit of detection between counting by microscopy and lipid
353 analysis by UHPLC-HRMS. Floating ‘fields’ of *Sargassum* were regularly encountered during the research
354 cruise, with the maximum observations occurring around Sta. 16 (pers. obs.). Further work on the HG
355 composition of the cyanobacteria found in these cohabitations would be necessary to draw conclusions as to
356 whether they contributed to the source of the two C₆ HGs detected at this sampling point.

357 4.2 C₅ Heterocyst glycolipids below the DCM

358 While the concentration of the C₅ HGs was generally highest within the mixed layer (ML, cf. Fig. 4f), Sta. 10
359 exhibited an increase in C₅ HG concentration with depth with C₅ HGs in both the 0.3 μm and 0.7 μm samples
360 increasing with depth to a maximum at 200 m (Fig. 5). The two size fraction profiles were carried out
361 approximately 12 hours apart and suggests that the HG maxima at 200 m was a feature for at least this period of
362 time. Sta. 9 was the only other station where the C₅ HG concentration (0.7 μm) at the DCM was higher than in
363 the ML (cf. Table S1). Foster et al. (2007) reported that DDAs are high in the ML but can increase below the ML
364 down to at least 100 m. Sediment trap studies in the North Pacific and tropical North Atlantic ocean have found
365 significant contributions by DDAs to the vertically exported particulate organic carbon (Karl et al., 2012;
366 Scharek et al., 1999; Subramaniam et al., 2008). While our study did not utilize sediment traps to collect sinking
367 particles, a proportion of the matter collected by in situ filtration is probably sinking rather than suspended
368 (Abramson et al., 2010). C₅ HGs have been found in surface sediment at depths up to 3000 m underlying our
369 water column sampling points (this study and Bale et al. (2015)), supporting the hypothesis that DDAs are
370 effectively transported in this environment from the water column to the sediment. These sinking particles could
371 be due to bloom-termination and aggregation or sinking of zooplankton fecal pellets.

372 4.3 C₅ Heterocyst glycolipids in surface sediment

373 As was found in a previous study concentrating on a smaller area close to the mouth of the Amazon (Bale et al.,
374 2015), the presence of a similar distribution of C₅ HGs in SPM and surface sediment indicates that HG producers
375 sink, probably enhanced by the mineral ballast as well as matrix protection provided by the association with
376 diatom silica skeletons. The total C₅ HG concentration in surface sediments was more spatially homogenous than
377 the distribution in the SPM (Table 1). Other than the two stations very close to the coast (where currents were
378 high and % TOC was at its lowest), the HGs were detected in comparably high levels from Sta. 7 onwards. This
379 reflects the wide spatial range of the HG-producers through an ‘integrated’ multi-decadal record of their
380 deposition. Each year between June and January, the Amazon plume is retroflected offshore, across the Atlantic



381 towards Africa due to the actions of the North Brazil Current and the North Equatorial Countercurrent, which
382 may account for the presence of the C₅ HGs in the surface sediments of Sta. 1 - 10. The rest of the year the
383 Amazon water flows northwestward towards the Caribbean Sea as the countercurrent and the retroflexion
384 weaken or vanish (Muller-Karger et al., 1988), in turn accounting for the C₅ HGs in the surface sediments of Sta.
385 13 - 23.

386 5. Conclusions

387 C₅ HGs were detected in the water column of the tropical North Atlantic and their concentrations correlated
388 strongly with DDAs. Furthermore, the HGs tracked the movement of the DDAs to the surface sediments in areas
389 known to be impacted by high seasonal DDA input (under the Amazon plume) whereas the HG concentration in
390 sediment farther away from plume was low. We conclude that C₅ HGs provide a robust, reliable method for
391 detecting DDAs in the marine environment. The apparent stability and specificity of C₅ HGs mean that they have
392 high potential for use in future work examining the presence and N-cycling role of DDAs in the past.

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580 **Figure legends**

581 **Figure 1.** Structures of the heterocyst glycolipids detected in this study C₆ glycolipids: 1-(O-hexose)-3,25-
582 hexacosanediol (C₆ HG₂₆ diol), 1-(O-hexose)-3-keto-25-hexacosanol (C₆ HG₂₆ keto-ol). C₅ glycolipids: 1-(O-
583 ribose)-3,29-triacontanediol (C₅ HG₃₀ diol), 1-(O-ribose)-3,27,29-triacontanetriol (C₅ HG₃₀ triol), 1-(O-ribose)-
584 3,29,32-dotriacontanetriol (C₅ HG₃₂ triol). Grey box indicates glycolipids associated with DDAs

585

586 **Figure 2.** Map of tropical North Atlantic showing the study site. Location of the stations indicated. Aquarius sea-
587 surface salinity (SSS) satellite data from ERDAPP (30 day composite, centered on 01-Sept-14,
588 <https://coastwatch.pfeg.noaa.gov/erddap/index.html>).

589

590 **Figure 3.** Water column characteristics along the cruise track. Color scales show a) temperature, b) chlorophyll
591 fluorescence (from fluorometer on CTD), c) PO₄ (color scale) and d) Si. Contour lines show salinity (a, b, d) and
592 NO₃ + NO₂ (c). Station numbers noted above plots, distance along transect from the Cape Verde Islands below.

593

594 **Figure 4.** Diazotroph abundance along the cruise track. Color scales show a) *Rhizosolenia* symbionts (trichomes
595 L⁻¹), b) *Hemiaulus* symbionts (trichomes L⁻¹), c) *Guinardia* symbionts (trichomes L⁻¹), d) *Trichodesmium* (free
596 trichomes L⁻¹) and e) *Trichodesmium* (colonies L⁻¹) while contour lines show salinity (a – e). f) Color scale
597 shows concentration of C₅ HG₃₀ triol (pg L⁻¹) while contour lines show C₅ HG₃₂ triol % (of C₅ total sum). Station
598 numbers above plots, distance along transect from the Cape Verde Islands below. Dots in Fig. 4a-c indicate
599 sampling depth for the salinity contours. Fig. 4d-e indicate sampling depth for HG lipids (Fig. 4f.). See
600 comments in text regarding *Trichodesmium* colony contouring artifacts.

601

602 **Figure 5.** Station 10, down column profile of C₅ HG sum (C₅ HG₃₀ triol + C₅ HG₃₂ triol, pg L⁻¹) from 0.7 μm
603 GF/F filters (grey broken line) and 0.3 μm GF75 filters (solid black line).

604

605 **Figure 6.** Plots of cell numbers and HG concentrations with a color scale showing salinity. a) *Rhizosolenia*
606 symbionts (trichomes L⁻¹), b) *Hemiaulus* symbionts (trichomes L⁻¹), c) C₅ HG₃₀ triol (pg L⁻¹), d) C₅ HG₃₂ triol
607 (pg L⁻¹) . Station numbers on x axis.

608

609

610



611 Table 1. Glycolipid concentrations from sea surface (3 – 5m) and surface sediment for all stations. For
612 concentrations at BWML and DCM see Table S3. † = No sediment collected, ns = not sampled.

Station	Lat	Long	Date	Water depth (m)	Salinity	Sea surface		Surface sediment		TOC (%)
						C ₅ HG ₃₀ triol (pg L ⁻¹)	C ₅ HG ₃₂ triol (pg L ⁻¹)	C ₅ HG ₃₀ triol (ng g ⁻¹)	C ₅ HG ₃₂ triol (ng g ⁻¹)	
1	15.02	-30.56	29/08/14	5500	36.4	18.0	0.00	1.7 ± 0.4	0.2 ± 0.1	0.6 ± 0.0
2 †	14.35	-32.58	30/08/14	6300	36.5	ns	ns	†	†	†
3	13.16	-36.21	31/08/14	5190	36.4	24.6	0.00	3.3 ± 0.8	0.4 ± 0.1	0.6 ± 0.0
4 †	12.41	-38.50	01/09/14	4810	36.2	40.9	0.00	†	†	†
5	10.83	-40.47	02/09/14	4620	36.0	8.9	0.00	2.3 ± 0.9	0.3 ± 0.1	0.5 ± 0.1
6 †	9.41	-42.10	03/09/14	3610	36.1	27.3	0.00	†	†	†
7	7.52	-44.28	04/09/14	4650	33.5	773	66.2	14.6 ± 6.8	1.7 ± 1.0	0.7 ± 0.0
8	6.49	-45.45	05/09/14	4250	31.9	4837	469	9.7 ± 1.7	1.0 ± 0.1	0.6 ± 0.0
9 *	5.60	-46.40	06/09/14	3770	32.2	24.6	0.00	4.8 ± 0.6	0.4 ± 0.0	0.5 ± 0.0
10 *	6.68	-47.49	07/09/14	4080	31.3	13.3	0.00	6.8 ± 4.4	0.7 ± 0.2	0.7 ± 0.0
11	5.53	-51.50	10/09/14	80	29.2	0.00	0.00	0.2 ± 0.1	0.01 ± 0.01	0.1 ± 0.0
12	6.07	-52.46	10/09/14	70	35.4	3.01	0.00	0.3 ± 0.1	0.01 ± 0.01	0.3 ± 0.1
13	7.60	-53.02	11/09/14	1000	32.8	31.1	0.00	7.4 ± 3.0	0.9 ± 0.4	1.2 ± 0.0
14 *	9.53	-51.32	12/09/14	4840	31.4	31.6	6.2	13.5 ± 1.4	1.3 ± 0.2	0.9 ± 0.0
15 †	8.95	-49.98	13/09/14	4660	32.7	565	24.4	†	†	†
16	10.22	-51.88	14/09/14	4940	33.9	391	27.3	13.0 ± 6.1	1.6 ± 0.5	1.0 ± 0.1
17	9.90	-53.27	15/09/14	4750	31.6	379	15.5	9.4 ± 3.0	0.9 ± 0.3	0.9 ± 0.1
18 †	9.37	-55.20	16/09/14	3590	33.2	611	67.2	†	†	†
19 †	10.52	-55.48	16/09/14	4180	32.8	390	34.5	†	†	†
20a	11.27	-54.16	17/09/14	4790	33.9	2.3	0.0	17.6 ± 7.0	1.4 ± 1.2	0.8 ± 0.0
20b †	11.47	-54.21	17/09/14	4830	34.2	67.7	0.0	†	†	†
21a	13.02	-54.67	18/09/14	5040	33.8	196	9.8	12.9 ± 1.7	1.6 ± 0.1	0.6 ± 0.0
21b †	13.20	-54.72	18/09/14	5170	34.8	249	6.5	†	†	†
22	14.80	-55.18	19/09/14	5500	35.6	48.6	0.4	13.4 ± 4.5	2.4 ± 1.1	0.7 ± 0.1
23	15.79	-57.05	20/09/14	5320	34.0	106	5.9	9.8 ± 3.5	1.4 ± 0.5	0.6 ± 0.0



613

615 Table 2. The additional SPM samples collected for high resolution depth profile at Sta. 10 (0.3 μm GF/F). * = Deep chlorophyll maximum.

Sampling depth (m)	Salinity	Temperature ($^{\circ}\text{C}$)	$\text{C}_5 \text{HG}_{30}$ triol ($\mu\text{g L}^{-1}$)	$\text{C}_5 \text{HG}_{32}$ triol ($\mu\text{g L}^{-1}$)	Sum ($\mu\text{g L}^{-1}$)
20	35.3	28.6	5.6	0.0	5.6
50*	36.4	27.3	9.6	0.0	9.6
200	35.2	11.4	108	0.3	108
400	34.7	7.4	24.4	0.1	24.5
600	34.6	6.3	29.0	0.1	29.1
800	34.6	5.1	22.5	0.4	22.9
1000	34.7	4.7	11.9	0.2	12.2
1200	34.8	4.8	12.6	0.2	12.8
1500	35.0	4.5	16.2	0.3	16.5
2000	35.0	3.4	18.3	0.3	18.6
2500	34.9	2.8	20.7	0.4	21.0
3000	34.9	2.4	22.3	0.3	22.6

Supplement

Supplementary Figures

620 **Figure S1.** Aquarius sea-surface salinity (SSS) satellite data (7 day composites), centered on (DD/MM/YY) a) 27/08/14, b) 03/09/14, c) 10/09/14, d) 17/09/14 and e) 24/09/14 showing highly dynamic plume location. Approximate location of R/V *Pelagia* indicated with purple circle.

Supplementary Tables

625 **Table S1.** Phytoplankton composition from Chemtax software based on pigment analysis. Numbers represent fraction of total Chl a. Fractions greater than 0.5 are highlighted in red and fractions between 0.1 and 0.2 are highlighted in purple.

Table S2. Diazotroph enumeration data. 5 categories: three are symbionts (syms), with the diatoms *Rhizosolenia clevei*, *Hemiaulus hauckii*, and *Guinardia cylindrus*, and two non-symbionts, *Trichodesmium* colonies and free *Trichodesmium* trichomes. Units are trichomes L^{-1} .

630 **Table S3.** Glycolipid concentration data. Concentration of $\text{C}_5 \text{HG}_{30}$ triol and $\text{C}_5 \text{HG}_{32}$ triol ($\mu\text{g L}^{-1}$) along with concentration of Chl a (ng L^{-1}) as measured by HPLC.