Sources, fluxes, and behaviors of fluorescent dissolved organic matter (FDOM) in the Nakdong-River Estuary, Korea

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Abstract

We monitored seasonal variations of dissolved organic carbon (DOC), stable carbon isotope of DOC (δ^{13}C-DOC), and fluorescent dissolved organic matter (FDOM) in water samples from a fixed station in the Nakdong-River Estuary, Korea. Sampling was performed every hour during spring tide once a month from October 2014 to August 2015. The concentrations of DOC and humic-like FDOM showed significant negative correlations against salinity (r^2=0.42-0.98, $p < 0.0001$), indicating that the river-originated DOM components are the major source and behave conservatively in the estuarine mixing zone. The extrapolated δ^{13}C-DOC values (-27.5 – 24.5‰) in fresh water confirm that both components are mainly of terrestrial origin. The slopes of humic-like FDOM against salinity were 60-80% higher in the summer and fall, due to higher terrestrial production of humic-like FDOM. The slopes of protein-like FDOM against salinity, however, were 70-80% higher in spring, due to higher biological production in river water. Our results suggest that there are large seasonal changes in riverine fluxes of humic and protein-like FDOM to the ocean.

1. Introduction

The global annual flux of dissolved organic carbon (DOC) via rivers is approximately 0.17 - 0.36 × 10^{15} g (Meybeck, 1982; Ludwig et al., 1996; Dai et al., 2012). The DOC delivered from riverine discharges as well as *in situ* production through biological activities significantly affects carbon and biogeochemical cycles in coastal waters (Hedges, 1992; Bianchi et al., 2004; Bauer et al., 2013; Moyer et al., 2015).

Generally, DOC includes fluorescent dissolved organic matter (FDOM), which emits fluorescent light due to its chemical characteristics. As FDOM accounts for 20 - 70% of the DOC in coastal waters (Coble, 2007) and controls the penetration of harmful UV radiation in the euphotic zone, it plays a critical role in carbon cycles as well as biological production. In addition, FDOM is known as a powerful indicator of humic and protein-like substances (Coble, 2007) in coastal waters. River discharge is generally the main source of humic-like FDOM in coastal waters, although it is also produced through *in situ* microbial activity (Romera-Castillo et al., 2011). In contrast, protein-like FDOM is known to be from biological production as well as anthropogenic sources (Baker and Spencer, 2004). Terrestrial humic substances behave conservatively in coastal areas due to their refractory characteristics (Del Castillo et al., 2000), whereas protein substances behave non-conservatively in many estuaries due to their relatively rapid production and degradation (Vignudelli et al., 2004).

The magnitudes of DOC and FDOM fluxes from rivers are generally dependent on rainfall, discharge, and temperatures (Maie et al., 2006; Jaffé et al., 2004; Huang and Chen,

2009). In the estuarine mixing zone, intensive biogeochemical processes occur through photo-oxidation, microbial degradation, or physicochemical transformations (i.e., flocculation, sedimentation) (Bauer and Bianchi, 2011; Moran et al., 1991; Benner and Opsahl, 2001; Raymond and Bauer, 2001). Recent studies have demonstrated large seasonal variations as high as 40%, in DOC export from rivers to the ocean (Burns et al., 2008; Bianchi et al., 2004; Dai et al., 2012). However, the seasonal variations in sources, fluxes, and behaviors of DOC and FDOM in the estuarine mixing zone are still poorly understood.

In this study, we analyzed DOC, δ^{13}C-DOC, and FDOM in estuarine water samples collected monthly from the Nakdong-River estuary. Sampling was conducted at a fixed platform, which has been utilized for monitoring various environmental parameters. This sampling station is advantageous because we can collect water samples for a wide range of salinities throughout tidal fluctuations. Using the data obtained from this unique station, we were able to determine (1) the behaviors of DOM in the estuarine mixing zone, (2) the fluxes of DOM from rivers based on the slopes between salinities and DOM components, and (3) the changes in DOM sources using δ^{13}C-DOC in the estuarine samples. The slope measurement in the mixing zone may represent the endmember of DOM components in rivers better than site-specific measurements in the river, by integrating larger spaces and times.

2. Materials and methods

2.1 Study site

The Nakdong-River Estuary, which is the estuary of the longest river in Korea, is a

major source of water supplying the needs for drinking, agriculture, and industry. The main channel of Nakdong River is approximately 510 km in length with a watershed area of approximately 23,380 km^2. It faces the south-eastern coastal area of the Korean peninsula, passing through Busan which is the second largest city in Korea. The mean annual precipitation is 1150 mm, and most precipitation (60-70%) occurs during the summer monsoon and typhoon seasons (Jeong et al., 2007). To manage water supply and saltwater intrusion, estuary dams were constructed in the mouth of the river in 1987.

2.2 Sampling

Water samples were collected at the sampling site which is located 560 m downstream from the dam (Fig. 1). The sampling period was from October 2014 to August 2015. The 2-L water sampling was conducted every hour for 24 hours during spring tide using an auto-sampler (RoboChem^{TM}Autosampler, Model S3-1224N, Centennial Technology, Korea), with a depth of the water intake 1 m below the surface. After samples were collected in acid-cleaned polyethylene bottles, they were moved to the laboratory within 24 hours. All water samples were filtered using pre-combusted GF/F filters. The FDOM samples were stored in pre-combusted amber glass vials and kept below 4°C in a refrigerator before analysis. The DOC and δ^{13}C-DOC samples were acidified to pH ~2 using 6 M HCl to avoid bacterial activities and stored in pre-combusted glass ampoules. Ampoules were fire-sealed to prevent the samples from any contaminations. The samples were analyzed for DOC and CDOM within a week. Salinity was measured using a YSI Pro Series conductivity probe sensor in the laboratory. The real-time and compulsory discharge volume data from the dam are available at http://www.water.or.kr, provided by K-Water. The monitoring program at this station is

maintained by Korea Environment Management Corporation (KOEM). The water temperature data are recorded automatically at the site. The data are available at https://www.koem.or.kr.

2.3 Analytical methods

The concentrations of DOC were determined by a high-temperature catalytic oxidation (HTCO) method using a TOC-VCPH analyzer (Shimadzu, Japan). Standardization was performed based on the calibration curve of acetanilide in ultra-pure water. The acidified samples were purged with carbon dioxide (CO_2) free carrier gas for 2 min to remove inorganic carbon. The samples were then injected into a combustion column packed with Pt-coated alumina beads and heated to 720°C. The CO_2 evolving from combusted organic carbon was detected by a non-dispersive infrared detector (NDIR). Our DOC method was verified with Deep Seawater Reference (DSR) samples for DOC (44-46 μmol L^{-1}) produced by the University of Miami, USA.

The values of δ^{13}C-DOC were measured using a TOC–IR–MS instrument consisting of an IR–MS instrument coupled with a Vario TOC cube (Isoprime, Elementar, Germany). The TOC instrument uses a common high-temperature catalytic combustion method (Kirkels et al., 2014). The analytical method is fully described in Kim et al. (2015). Briefly, 10 mL of filtered samples was purged with O_2 gas for 20–30 min to completely remove DIC after the samples were acidified to pH ~2. Then, 1 mL of the sample was injected into Pt-impregnated catalyst in a quartz tube. In this tube, the DOC was converted completely to CO_2 at 750°C, which was then fed through a water trap followed by a halogen trap. After DOC was detected by an NDIR

detector, the CO_2 gas was entered the TOC–IR–MS interface by the O_2 carrier gas. In the interface, the CO_2 was transferred to the IR–MS instrument following the removal of any interfering gasses. The δ^{13}C-DOC value of blank was measured using the Low Carbon Water (LCW) from Hansell lab (University of Miami), which contains less than 2 μM DOC. Certified IAEA-CH6 sucrose (International Atomic Energy Agency, −10.45 ± 0.03‰) prepared with the low carbon water was used as a standard solution. A standard sample was analyzed for every sample queue (once before or after ten samples) to check a drifting effect during the measurements. The blank correction was performed using a method previously described in De Troyer et al. (2010) and Panetta et al. (2008). Our measurement result of δ^{13}C-DOC for the DSR (University of Miami) was −21.5 ± 0.1‰, which is consistent with the results reported by Panetta et al. (2008) and Lang et al. (2007). The reproducibility of TOC–IR–MS was ~0.3‰.

FDOM fluorescence was determined in a scan mode using a spectrofluorometer (SCINCO FluoroMate FS-2) within two days after sampling. Emission (Em) spectra were collected from 250 to 600 nm at 2 nm intervals at excitation (Ex) wavelengths from 250 nm to 500 nm at 5 nm intervals. Backgrounds were subtracted for fresh distilled water prepared daily from the sample data to eliminate Raman Scatter peaks (Zepp et al., 2004). All data were obtained in counts per second (cps) and converted to a ppb quinine sulfate standard solution in 0.1 N sulfuric acid at Ex/Em of 350/450 nm. The inner filter effect was negligible for these estuarine water samples using this instrument since the correlation between the uncorrected and corrected values for the inner filter effect was very significant for some samples (r^2=1, n=5). EEMs-PARAFAC analysis was performed using a MATLAB R2013a program with a DOMFluor toolbox.

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148 **3. Results and Discussion**

149 Salinities ranged from 0.1 to 28.5 over the sampling period of a year. Salinities in the
150 sampling location were dependent primarily on the volume of river-water discharge from the
151 dam. The volumes of river discharge were relatively larger in October, April, July, and May.
152 The mean annual surface water temperature was 16°C, with the lowest temperature (avg. 8°C)
153 in December and the highest temperature in August (avg. 26°C).

154

155 *3.1 Behaviors and sources of DOC in the estuarine mixing zone*

156 The concentrations of DOC ranged from 100 to 300 μM, with the highest concentrations in
157 July (avg. 243 μM) and the lowest concentrations in February (avg. 115 μM), consistent with
158 the typical DOC concentration ranges in coastal waters (Wang et al., 2004; Raymond and Bauer,
159 2001). The concentrations of DOC correlated significantly with salinities (r^2 = 0.59-0.92, $p <$
160 0.0001), indicating that DOC behaves conservatively in the mixing zone of this estuary (Fig.
161 2A), which is commonly observed in estuarine mixing zones (Laane, 1980; Mantoura and
162 Woodward, 1983; Del Castillo et al., 2000; Clark et al., 2002; Jaffé et al., 2004).

163

164 If the high salinity periods are excluded, both the slope and y-intercept of DOC
165 concentrations versus salinities were highest in July (Fig. 1), which could be due to a higher
166 terrestrial DOC loading in the summer period, as observed in Horsens Fjord, Denmark
167 (Markager et al., 2011). For this comparison, we excluded the high-salinity periods (>20),
168 including December, January, February, and June, since they showed a narrow and low DOC

concentration range (103-163 μM), resulting in large uncertainties by extrapolating them to the fresh water.

The carbon isotope values in the Nakdong-River Estuary ranged from -28.2 ‰ to -17.6. In order to determine the source of DOC in fresh water, we plotted δ^{13}C-DOC values against salinities (Fig. 2B). The conservative mixing curve of δ^{13}C values can be obtained using the two endmember mixing equation (Spiker, 1980; Raymond and Bauer, 2001):

$$\delta^{13}C_s = \frac{F_r\delta^{13}C_r[DOC]_r+(1-F_r)\times\delta^{13}C_m[DOC]_m}{[DOC]_s} \quad (1)$$

where $\delta^{13}C_s$, $\delta^{13}C_r$ and $\delta^{13}C_m$ are the δ^{13}C-DOC values at a given sample salinity, river endmember salinity, and marine endmember salinity, respectively; F_r is the riverine freshwater fraction calculated from the measured salinities; $[DOC]_s$ and $[DOC]_m$ are the DOC concentrations at a given salinity and marine endmember salinity, respectively; $[DOC]_r$ is the endmember DOC value for the river water (Fig. 1).

The riverine DOC endmember values (S=0‰) ranged from 174 to 284 μM. The marine endmember value (S=29‰) of DOC is 100 μM with the δ^{13}C-DOC value of -19‰. If these values from each month are applied, the δ^{13}C-DOC endmember values for the river water extrapolated to be from -27.5 to -24.5‰ (average: -26.2‰). Overall, the carbon isotope values of our samples are fitted well into the conservative mixing curve of the overall trend, with a slight change using different endmember values for different months (Fig. 2B). In general, δ^{13}C-DOC values range from −22 to −18‰ for marine phytoplankton, from −34‰ to −23‰ for terrestrial C3 plants, and from −16‰ to −10‰ for terrestrial C4 plants (Gearing 1988; Clark

and Fritz, 1997). Carbon isotope values in our study confirm that the main source of DOC in the estuarine mixing zone is dominantly from terrestrial C3 plants over all seasons. However, the value was heavier at lower salinity ranges (S<10) in March and April samples, perhaps in association with the higher biological production in the river.

3.2 Behaviors and sources of FDOM in the estuarine mixing zone

Four components were identified in the water samples from the EEMs dataset. Based on the excitation-emission peak location, Component 1 ($FDOM_H$, Ex/Em = 320/418 nm) is found to be a terrestrial humic-like component (C peak) shown by Coble (2007). Component 2 ($FDOM_P$, Ex/Em = 280/328 nm) is found to be a tryptophan-like component (T peak), which is produced by microbial processes. Component 3 (Ex/Em = 300,325/364 nm) is found to be a marine humic-like component (M peak). Since Component 3 values were significantly correlated with Component 1 (r^2=0.95) values, we simply focused on Component 1 ($FDOM_H$) and Component 2 ($FDOM_P$) for data interpretations.

The concentrations of $FDOM_H$ ranged from 2.4 to 19.7 quinine sulfate unit (QSU), with the highest concentration in July (avg. 17.6 QSU) and the lowest concentration in June (avg. 3.4 QSU) (Fig. 2C). The concentrations of $FDOM_P$ ranged from 0.6 to 22.4 QSU, with the highest concentration in March (avg. 15.1 QSU) followed by October (avg. 13.6 QSU) (Fig. 2D).

212 The concentrations of both FDOM components were significantly correlated with
213 salinities (r^2 = 0.42-0.98, $p<0.0001$ for $FDOM_H$ and r^2 = 0.27-0.96, $p<0.0001$ for $FDOM_P$),
214 indicating that they are conservative in the mixing zone (Fig. 2). The slopes of $FDOM_H$ and
215 $FDOM_P$ for each month ranged from -0.15 to -0.59 and -0.15 to -0.71, respectively. The higher
216 $FDOM_H$ slopes in July and October were similar to the trend of DOC (Fig. 2C), which could
217 be due to higher terrestrial FDOM production. However, the seasons (March and April) in
218 which higher $FDOM_P$ slopes occurred differ from those of DOC and $FDOM_H$, indicating that
219 both FDOM components have different source inputs (Fig. 2D).

220

221 Although there are large differences in scattering of FDOM components against
222 salinities, it is very difficult to compare scatterings for different seasons in order to discuss the
223 different behaviors of DOM since the scattering is generally larger for the narrow salinity
224 ranges. If the winter data are excluded, in March, during the highest biological production
225 period in the river, the correlation coefficient against salinities was the highest for $FDOM_P$ and
226 lowest for $FDOM_H$. In contrast, in June, during the highest fluvial DOM discharge period, the
227 correlation coefficient against salinities was the highest for $FDOM_H$ and lowest for $FDOM_P$.
228 This suggests that the biological production and removal, together with other generally known
229 factors such as photo-degradation and sedimentary inputs, may affect the scattering of these
230 FDOM components in the estuarine mixing zone.

231

232 As such, there was a significant positive correlation between $FDOM_H$ and DOC
233 concentrations throughout all sampling periods (r^2 = 0.93, $p<0.0001$) (Fig. 3A), suggesting that
234 the main source of $FDOM_H$ and DOC is terrestrial based on δ^{13}C-DOC values. Since FDOM

does not usually contribute to a major portion of DOC, a positive correlation between FDOM and DOC has only been observed in specific areas, such as river-estuarine systems (Del Vecchio and Blough, 2004; Coble, 2007). Stedmon et al., (2006) demonstrated that stronger correlations were observed between DOC and FDOM as humic substances derived from terrestrial DOM are more colored than DOM produced *in situ*. In general, terrestrial DOM occurring in rivers originates mainly from plant decomposition and leaf litter in the form of humic substances (Huang and Chen, 2009). As such, Gueguen et al., (2006) showed that humic materials are more effectively extracted from soils during August and September under high temperatures. Thus, higher $FDOM_H$ slopes in August, October, and November, relative to the other periods, could be associated with higher terrestrial inputs of degradation products of soil organic matter (Dowell, 1985; Qualls et al., 1991).

In the study region, $FDOM_P$ concentrations were poorly correlated with DOC concentrations (r^2=0.11) (Fig. 3B). The slopes of $FDOM_P$ concentrations against DOC concentrations varied significantly over different seasons, with steeper gradients in the spring (March and April) and fall (October). In general, $FDOM_P$ is known to be produced efficiently by biological production in water (Coble, 1996; Belzile et al., 2002; Steinberg et al., 2004; Zhao et al., 2017). Thus, higher $FDOM_P$ concentrations, relative to DOC concentrations, in the spring and fall seems to be associated with the spring and fall phytoplankton blooms in river waters (Mayer et al., 1999; Zhang et al., 2009).

3.3 Fluxes of DOC and FDOM in the estuarine mixing zone

257 The fluxes of DOC and FDOM from rivers to the ocean are calculated using the
258 endmember values (C) of these components in rivers multiplied by the river discharge volumes
259 (Q) for each month (Fig. 4). For this estimation, we assumed that (1) the endmember values
260 are the same as the intercepts of the DOC, $FDOM_H$, and $FDOM_P$ versus salinity plots, and (2)
261 the endmember values measured in the spring tides represent the concentrations of these
262 components for each month.

263

264 River discharge was highest in April and July following heavy precipitation, and the
265 largest discharge volume was about five-fold higher than that of winter discharges (Fig. 4A).
266 However, the monthly variations of DOC endmember (y-intercept) values were quite constant,
267 ranging from 174 - 284 μM. This indicates that the concentrations of DOC in the river are
268 independent of river discharge volumes (Fig. 4B). The DOC endmember values were highest
269 in December, followed by July and June (Fig. 4B). The monthly variation trend of $FDOM_H$
270 endmember values was similar to that of DOC, except for the December value. Excluding the
271 December values, the $FDOM_P$ endmember values were highest in March, February, and
272 October. These endmember trends are consistent with the slope variations explained in the
273 previous section. Although there are large uncertainties in fresh water endmember values of
274 DOC and FDOM in winter owing to narrow, high salinity ranges, we used the endmember
275 values for the flux comparisons since the contribution of the uncertainties may be relatively
276 small due to smaller river discharge volumes in winter.

277

278 The riverine DOC flux ranged from 1.6×10^6 mol day^{-1} (February) to 12.3×10^6 mol
279 day^{-1} (July), indicating that there are large variations of DOC fluxes to the ocean. The riverine

flux of $FDOM_H$ and $FDOM_P$ ranged from 1.4×10^9 QSU m^3 day^{-1} (December) to 23.1×10^9 QSU m^3 day^{-1} (July) and from 1.6×10^9 QSU m^3 day^{-1} (June) to 16.4×10^9 QSU m^3 day^{-1} (March), respectively. The seasonal variation trend of $FDOM_H$ was similar to that of DOC. The fluxes of $FDOM_P$ in December and March were twofold higher than those of $FDOM_H$ whereas the flux of $FDOM_H$ in July was 2-3 folds higher than that of $FDOM_P$. This shows that the fluxes of both components of FDOM differ significantly by seasons owing to the different source inputs even though their magnitudes are controlled mainly by river discharges

It is well known that the single sampling event is not enough to capture the full range of natural variability in DOM abundance over all seasons (Stedmon et al., 2006; Huang and Chen, 2009; Markager et al., 2011; Dai et al., 2012; Moyer et al., 2015). Overall, our results show that monthly variations are significant. This implies that our understanding of DOC fluxes from large rivers is largely biased, depending on sampling resolution, methods, and hydrogeological settings of a specific river. For example, if summer data are extrapolated to annual river water discharge, the DOC and $FDOM_H$ fluxes can be overestimated up to three times for the Nakdong River.

4. Conclusions

The concentrations of $FDOM_H$ and DOC showed significant negative correlations against salinities throughout all sampling periods, indicating that they behave conservatively in this estuarine mixing zone. The slopes of both DOC and $FDOM_H$ concentrations versus salinities were highest in July, due to the largest terrestrial DOC loadings. The carbon isotope values showed that the main source of DOC in the estuarine mixing zone is terrestrial C3 plants

over all seasons. The slopes of $FDOM_P$ versus salinity were relatively higher in March and April in association with the spring phytoplankton blooms in river and estuarine waters. The monthly fluxes of DOC, $FDOM_H$, and $FDOM_P$ showed large seasonal variations (5-10 folds), suggesting that the estimation of annual riverine fluxes of DOC, $FDOM_H$, and $FDOM_P$ requires careful considerations of seasonal changes in rivers.

Competing interests

The authors declare that they have no conflict of interest.

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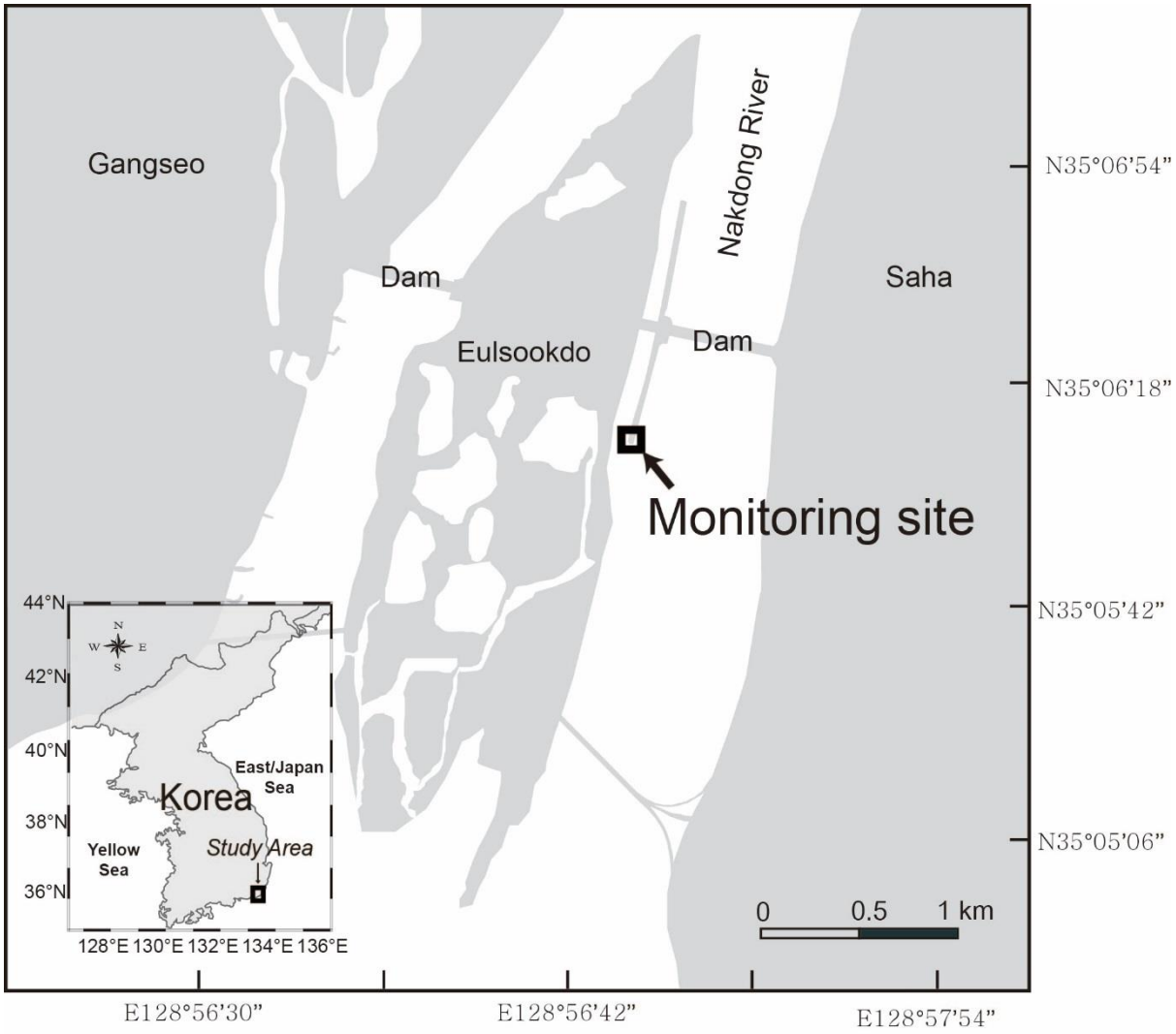

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444 Figure 1. Map of the Nakdong-River Estuary. The square indicates a fixed monitoring site,

445 located 560 m downstream from the dam.

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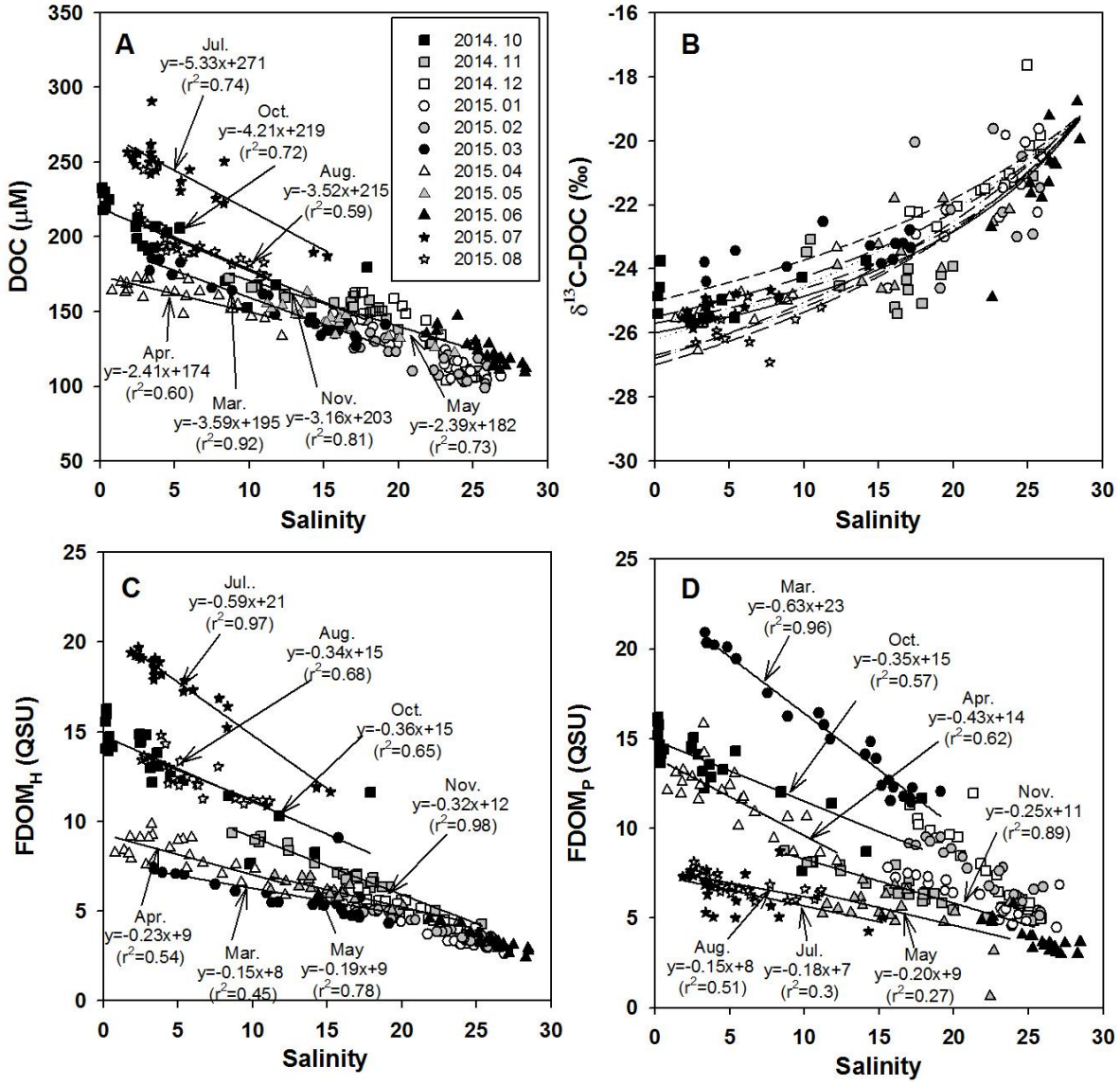
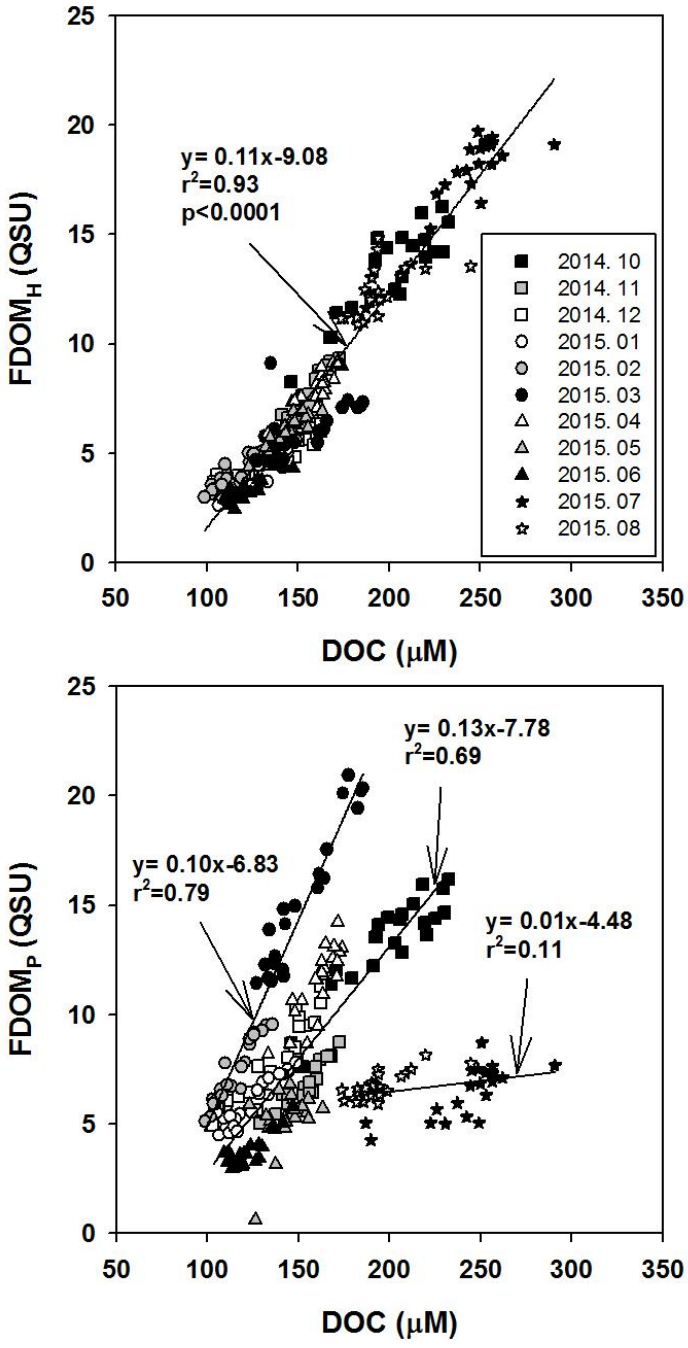


Figure 2. Salinities versus the concentrations of (A) DOC, (B) δ^{13}C-DOC, (C) $FDOM_H$, and (D) $FDOM_P$. The values for the regression lines are excluded for high-salinity periods (>20), including December, January, February, and June, which have large uncertainties in extrapolation. The solid curve (B) is the average conservative mixing line for the two endmember mixing equation. The dotted lines represent the monthly changes in mixing lines for the different monthly endmember values.



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455 Figure 3. The plots of the concentrations of DOC versus the concentrations of (A) $FDOM_H$ and

456 (B) $FDOM_P$.

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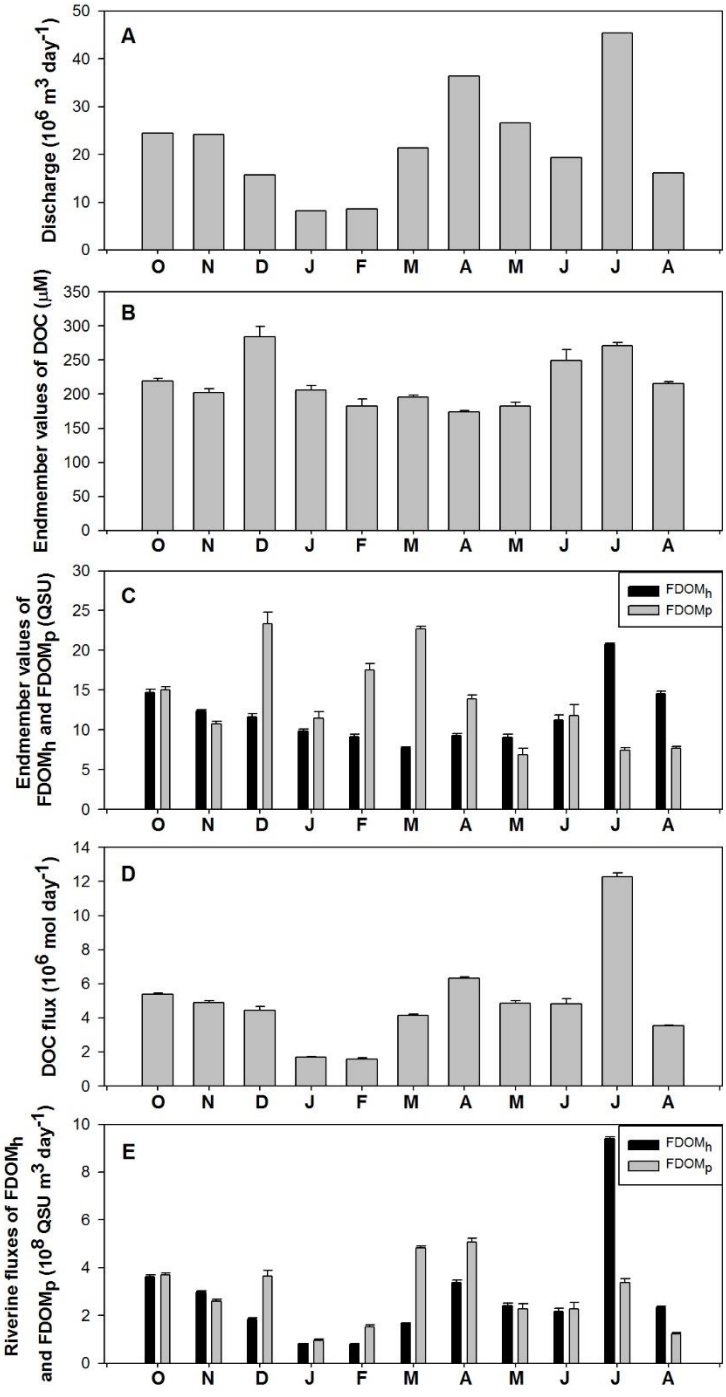

458

459 Figure 4. Temporal variations in discharge volumes, the endmember values of DOC, $FDOM_H$,

460 and $FDOM_P$, and riverine fluxes of DOC, $FDOM_H$, and $FDOM_P$ in the Nakdong-River Estuary

461 from October 2014 to August 2015.