

Impact of salinity on element incorporation in two benthic foraminiferal species with contrasting Magnesium contents

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Abstract. Accurate reconstructions of seawater salinity could provide valuable constraints for studying past ocean circulation, the hydrological cycle and sea level change. Controlled growth experiments and field studies have shown the potential of foraminiferal Na/Ca as a direct salinity proxy. Incorporation of minor and trace elements in foraminiferal shell carbonate varies, however, greatly between species and hence extrapolating calibrations to other species needs validation by additional (culturing) studies. Salinity is also known to impact other foraminiferal carbonate-based proxies, such as Mg/Ca for temperature and Sr/Ca for sea water carbonate chemistry. Better constraints on the role of salinity on these proxies will therefore improve their reliability. Using a controlled growth experiment spanning a salinity range of 20 units and analysis of element composition on single chambers using laser ablation-Q-ICP-MS, we here show that Na/Ca correlates positively with salinity in two benthic foraminiferal species (*Ammonia tepida* and *Amphistegina lessonii*). The Na/Ca values differ between the two species, with an approximately 2-fold higher Na/Ca in *A. lessonii* than in *A. tepida*, coinciding with an offset in their Mg content (~35 mmol/mol versus ~2.5 mmol/mol for *A. lessonii* and *A. tepida*, respectively). Despite the offset in average Na/Ca values, the slopes of the Na/Ca-salinity regressions are similar between these two species (0.077 versus 0.064 mmol/mol change per salinity unit). In addition, Mg/Ca and Sr/Ca are positively correlated with salinity in cultured *A. tepida*, but show no correlation with salinity for *A. lessonii*. Electron microprobe mapping of incorporated Na and Mg of the cultured specimens shows that within chamber walls of *A. lessonii*, Na/Ca and Mg/Ca occur in elevated bands in close proximity to the primary organic lining. Between species, Mg-banding is relatively similar, albeit that Mg content is 10 times lower and that variation within the chamber wall is much less pronounced in *A. tepida*. In addition, Na-banding is much less prominent in this species than it is in *A. lessonii*. Inter-species differences in element banding reported here are hypothesized to be caused by differences in biomineralization controls responsible for element uptake.

33 **1. Introduction**

34 Sea water salinity varies over time and space as a function of continental ice volume, evaporation,
35 precipitation and river runoff. Salinity reconstructions could provide important constraints on past
36 ocean circulation, the hydrological cycle and glacial-interglacial sea level changes. Currently, most
37 reconstructions of salinity are indirect and based on the correlation between salinity and $\delta^{18}\text{O}_{\text{water}}$,
38 assuming this relationship to be constant over space and time (Rohling and Bigg, 1998). An
39 independent salinity proxy may reduce the uncertainties inherently associated with such approaches
40 (Rohling and Hilgen, 2007) and should preferably be based on one of the main components of sea
41 water salinity, for instance sodium (Na). Results from a culture study showed that the sodium content
42 of foraminiferal calcite ($\text{Na}/\text{Ca}_{\text{cc}}$) correlates positively and linearly with salinity for the benthic low-
43 Mg, symbiont-barren species *Ammonia tepida*, with a sensitivity of 0.22 mmol/mol for every change of
44 1 salinity unit between salinities 30 and 38.6 (Wit et al., 2013). Various culture studies earlier showed
45 that also Mg/Ca is affected by salinity, but responds more strongly to changes in temperature (Lea et
46 al., 1999; Dissard et al., 2010b; Nürnberg et al., 1996; Hönisch et al., 2013). Although an effect of
47 salinity on foraminiferal Sr/Ca_{cc} has been reported in some studies (Kırsakürek et al., 2008; Dissard et
48 al., 2010b; Wit et al., 2013) other studies did not find a relation between salinity and foraminiferal
49 Sr/Ca (Dueñas-Bohórquez et al., 2009; Diz et al., 2012; Allen et al., 2016), which lead to the
50 hypothesis that foraminiferal Sr/Ca mainly reflects sea water inorganic carbon chemistry (Keul et al.,
51 2017) in addition to its response to temperature (Lea et al., 1999; Raja et al., 2007). Hence, an
52 independent salinity proxy would not only be useful for constraining past (changes in) salinity, but also
53 improve temperature reconstructions based on Mg/Ca_{cc} and reconstructions of past sea water carbonate
54 chemistry based on Sr/Ca.

55 Following the culture-based Na/Ca_{cc}-salinity calibration for *A. tepida* (Wit et al., 2013), a culture study
56 with planktonic symbiont-bearing species also showed a significant linear relationship for
57 *Globigerinoides ruber* (Allen et al., 2016). Although no significant relationship was observed in this
58 study for *G. sacculifer* (Allen et al., 2016), a recent field calibration observed positive linear
59 relationships for both these species (Mezger et al., 2016). Still, the Na/Ca-salinity sensitivities observed
60 between the different species and studies differed considerably (ranging from a change in 0.074 to 0.66
61 mmol/mol in Na/Ca_{cc} for a change in 1 salinity unit). Whereas Wit et al. (2013) suggested an
62 incorporation mechanism similar to that observed in inorganic calcite, field and culture studies also
63 show that different species of foraminifera have varying calcite chemistries, thereby resulting in the
64 need of species-specific calibrations similar to many other foraminiferal trace metal-based proxies (e.g.
65 Elderfield and Ganssen, 2000; Rosenthal et al., 2000; Anand et al., 2003; Bemis et al., 1998; Toyofuku
66 et al., 2011). For example, Mg/Ca_{cc} values are different between groups of low-Mg-, high-Mg hyaline
67 and porcelaneous foraminifera (Toyofuku et al., 2000; Segev and Erez, 2006; Raja et al., 2007), which
68 also seems to be reflected in other co-precipitated cations (De Nooijer et al., 2017). Hence, calibration
69 of Na/Ca_{cc} as a function of salinity for other species is not only necessary to test the applicability of this
70 novel proxy for other groups of foraminifera, but also allows testing whether monovalent cations
71 follow the inter-species trends described for divalent cations (Terakado et al., 2010).

72 Here we calibrated Na-, Mg- and Sr-incorporation in the intermediate-Mg calcite benthic foraminiferal
 73 species *Amphistegina lessonii* and the low-Mg calcite species *Ammonia tepida* over a salinity range of
 74 20 units (from 25 to 45). We thus compare the El/Ca versus salinity trends in a tropical, symbiont-
 75 bearing species (*A. lessonii*) to a temperate intertidal symbiont-barren species (*A. tepida*) and both of
 76 them to existing calibrations. The chemical composition of newly formed calcite was determined by
 77 Laser Ablation Inductively Coupled Plasma Mass Spectrometry (LA-Q-ICP-MS), providing insights in
 78 concentrations and variability therein between specimens and between single chambers. To investigate
 79 intra-specimen variability at the scale of the chamber wall we also performed Electron Probe Micro
 80 Analysis (EPMA), mapping the Ca, Na and Mg distribution throughout the chamber wall for specimens
 81 of both species cultured.

82 2. Methods

83 2.1 Culture media preparation and chemistry

84 In total, 50 L of sea water with a salinity of 50 was prepared by sub-boiling 0.2 μm filtered North
 85 Atlantic sea water for 48 hours at 45 °C. Subsequently, culture media were obtained by diluting this
 86 high-saline sea water with double de-ionized sea water ($\sim 18 \mu\Omega$) in batches of approximately 10L with
 87 salinity increasing from 25 to 45 in steps of 5 units, resulting in 5 unique salinity conditions. Using a
 88 single batch of concentrated sea water to subsequently dilute to the desired salinities ensures constant
 89 element to Ca ratios. Salinity of the media was measured with a salinometer (VWR CO310), based on
 90 conductivity. Culture media were stored in Nalgene containers and kept in the dark at 10 °C. Sea water
 91 pH was determined with a pH meter (pH110, VWR). Subsamples were taken prior to and at the end of
 92 the experiment and analyzed for DIC and element concentrations to monitor the effect of sub-boiling
 93 on the sea water's inorganic carbon chemistry and element composition (Table 1). Subsamples for DIC
 94 were collected in headspace-free vials and conserved with a saturated HgCl_2 solution (10 μl HgCl_2 /10
 95 ml sample). DIC measurements were performed on an autoanalyzer spectrometric system TRAACS
 96 800; Stoll et al. (2001). This analysis requires only a small amount of sample, while yielding high
 97 accuracy ($\pm 2 \mu\text{mol/kg}$) and precision ($\pm 1.5 \mu\text{mol/kg}$). The minor and major elemental composition of
 98 the culture media was measured using a sector field ICP-MS (Element2, Thermo Scientific) by
 99 sampling 1 ml from the culture media and dilution by a factor 300 with 0.14 M HNO_3 (Table 1).

100 **Table 1.** Experiment culture media measurements per salinity condition. Carbonate ion concentrations
 101 and saturation state with respect to calcite (at 25 °C) were calculated using CO2SYS (Van Heuven et
 102 al., 2011) and the equilibrium constants K1 and K2 of Mehrbach et al. (1973), as reformulated by
 103 (Dickson and Millero, 1987).

Experiment	Na/Ca _{sw} mol/mol	Mg/Ca _{sw} mol/mol	Sr/Ca _{sw} mmol/mol	Salinity	DIC $\mu\text{mol/kg}$	pH	[CO ₃ ²⁻] mmol/kgSW	Ω_{calcite}
S25	48.84	5.61	9.37	25.2	1087.3	8.32	164.90	4.28
S30	49.79	5.69	9.45	30.3	1305.3	8.28	205.98	5.15
S35	48.56	5.51	9.04	35.2	1512.0	8.22	258.84	6.22
S40	48.50	5.62	9.19	40.0	1734.4	8.17	267.23	6.16
S45	48.90	5.73	9.21	45.2	1947.4	8.10	284.67	6.23

104

105 2.2 Collection of foraminifera and culture set-up

106 Surface sediment samples containing foraminifera (*A. lessonii*) were collected from the Indo-Pacific
107 Coral Reef aquarium in Burgers' Zoo (Arnhem, The Netherlands; Ernst et al., 2011) and a tidal flat
108 near Den Oever, the Wadden Sea (*A. tepida*, genotype T6; Hayward et al., 2004). Sediment was stored
109 in aerated aquaria at 25°C (*A. lessonii*) and 10°C (*A. tepida*) with a day/night cycle of 12/12 hours,
110 similar to conditions in the coral reef aquarium and Wadden Sea, respectively. From both stocks, living
111 specimens, recognized by chambers that were filled with yellow cytoplasm and pseudopodial activity,
112 were isolated.

113 The culture protocol was the same for both species to facilitate comparison of obtained Element/Ca
114 ratios between species. Since our specimens of *A. tepida* are from a location with a much larger
115 temperature range than where *A. lessonii* is derived from (Ernst et al., 2011; van Aken, 2008; De
116 Nooijer et al., 2014), both species were incubated at 25 °C. Living specimens were placed in groups of
117 25 individuals in Petri dishes with approximately 70 ml of North Atlantic surface sea water (0.2 µm
118 filtered) and fed with fresh cells of the algae *Dunaliella salina*. After reproduction, which occurred in
119 approximately 2/3 of all incubated specimens in both species, 2-3 chambered juveniles were isolated.
120 The use of specimens from reproduction events guarantees that virtually all chambers present at the end
121 of the experiment were produced under the culture conditions (De Nooijer et al., 2014). Strains of
122 specimens of the reproduction events were divided over Petri dishes (resulting in 2-10 individuals per
123 dish) with approximately 10 ml culture medium and stored in a temperature controlled incubator set at
124 25 °C with a day/night cycle of 12/12 hours. The culture media in the Petri dishes were replaced once
125 every week, after which specimens were fed with 1 ml concentrated and freeze-dried *Dunaliella salina*
126 diluted with the culture medium for each salinity condition, to minimize changes in salinity when
127 feeding the foraminifers. The amount of food was adjusted so that it was not depleted after a week, at
128 the same time not resulting in an excess of debris and hence reduce bacterial growth. Petri dishes were
129 sealed with a lid to minimize evaporation. After 6-8 weeks, specimens were harvested and transferred
130 to microvials to clean the specimens' carbonate shells from cell material. Specimens were cleaned with
131 an adapted version of the Barker protocol (Barker et al., 2003), only applying the organic removal /
132 oxidation step, in which NaOH was replaced with NH₄OH, in order to avoid Na-contamination of our
133 samples. Organic matter was removed by adding 1% H₂O₂ buffered with 0.1M NH₄OH at 90 °C and
134 gentle ultrasonication (80kHz, 50% power, in degas mode) for 1 min, which is known not to affect
135 obtained Mg/Ca and Sr/Ca (Barker et al., 2003). Specimens were subsequently rinsed 3 times with
136 double de-ionized water, dried in a laminar flow cabinet, after which their size was determined (i.e. the
137 maximum diameter crossing the centre of the specimen). The specimens were thereafter stored until
138 geochemical analyses (LA-Q-ICP-MS; 2.2.2 and EPMA; 2.4).

139 2.2.2 Foraminiferal calcite chemistry

140 Specimens were fixed on a laser ablation-stub using double sided tape, carefully positioning them to
141 allow ablation of the last chambers (Appendix A). Element concentrations of individual chambers were

142 measured with LA-ICP-MS (Reichart et al., 2003). The last 1-3 chambers of each specimen were
 143 ablated using a circular spot with a diameter of 60 μm (*A. tepida*) and 80 μm (*A. lessonii*)
 144 (NWR193UC, New Wave Research) in a helium environment in a New Wave TV2 dual-volume cell
 145 (cup volume of $\sim 1\text{ cm}^3$) at a repetition rate of 6 Hz and an energy density of approximately 1 J/cm².
 146 The aerosol was transported to a quadrupole ICP-MS (iCap, Thermo Scientific) on a helium flow at a
 147 rate of 0.7 L/min, with 0.4 L/min Argon make-up gas being added before entering the torch. Nitrogen
 148 gas was added at a rate of 5 ml/ minute to enhance sensitivity of the analysis. Before entering the torch,
 149 the aerosol/ Ar/ He mixture passed through an in-house made smoothing device to reduce temporal
 150 variations in signal strength. Monitored masses included ⁷Li, ¹¹B, ²³Na, ²⁴Mg, ²⁵Mg, ²⁷Al, ⁴³Ca, ⁴⁴Ca,
 151 ⁶⁰Ni, ⁶⁶Zn, ⁸⁸Sr, ¹³⁷Ba and ²³⁸U, with one full cycle through the different masses taking 120 ms.
 152 Calibration was performed against a MACS-3 (synthetic calcium carbonate) pressed powder carbonate
 153 standard with ⁴³Ca as an internal standard. Count rates for the different masses were directly translated
 154 into element/Ca_{cc} (El/Ca_{cc}) ratios. Internal precision based on MACS-3 is 4% for Na, 3% for Mg and
 155 4% for Sr. Accuracy and relative analytical errors, based on measuring international standards JCp-1
 156 coral (*Porites* sp.) powder and the NIST (National Institute of Standards and Technology) SRM 610
 157 and SRM 612 (glass) are listed in Table 2. The relatively large offset between the glass standards and
 158 the pressed powders (both MACS-3 and JCp-1) is known not to influence obtained El/Ca_{cc} ratios when
 159 either one is used as calibration standard (Hathorne et al., 2008), but due to the similar matrix, MACS-
 160 3 was chosen as calibration standard.

161 **Table 2.** Accuracies (Ac) and precisions (Pr) for Na, Mg and Sr for the various standards analyzed.

Standard	n	Ac Na (%)	Pr Na (%)	Ac Mg (%)	Pr Mg (%)	Ac Sr (%)	Pr Sr (%)
JCp-1	51	99	6	96	6	96	4
NIST610	32	119	3	104	2	110	3
NIST612	29	119	3	104	2	110	2

162
 163 In total, 675 chambers were measured (336 for *Amphistegina* and 339 for *Ammonia*), resulting in
 164 between 52 to 125 single chamber measurements per salinity condition per species. These
 165 measurements were done on the last three (final or F, penultimate or F-1 and pre-penultimate or F-2)
 166 chambers of these specimens (see Table 3 for number of specimens and average number of spots per
 167 specimen). Element concentrations were calculated from the time (i.e. ablation depth) resolved profiles
 168 using an adapted version of the SILLS (Signal Integration for Laboratory Laser Systems; Guillong et
 169 al., 2008)) package for MATLAB (for details see Van Dijk et al., 2017a), while taking care to exclude
 170 contaminations potentially present on chamber walls (examples of profile selection: Duenas-Bohorquez
 171 et al., 2011; Wit et al., 2013; Mewes et al., 2014; Mezger et al., 2016; Van Dijk et al., 2017b).
 172 Measurements with ablation yields or integrations times <5 s were excluded from further analysis.
 173 The LA measurements were also used to investigate the co-occurrence of elements within specimens.
 174 Since there is variability in Ca counts between the laser ablation measurements, single-spot based
 175 Element/Ca_{cc} ratios may cause spurious correlation due to coupled differences in Ca counts. To test
 176 whether observed correlations between Na/Ca_{cc}, Sr/Ca_{cc} and Mg/Ca_{cc}, based on single-spots, are due to
 177 the use of a common denominator (Ca), we performed a Monte Carlo simulation. In short, the

178 correlation coefficients between randomly drawn single-spot Mg concentration, divided by measured
179 Ca, and measured Na/Ca_{cc} concentrations were compared to the correlation coefficient of measured
180 Na/Ca_{cc} and Mg/Ca_{cc} concentration ratios in our dataset. By using a Kernel fit of the measured data set
181 to draw the random data set and using the measured Ca as a common denominator we effectively
182 simulate the spurious correlation. The Monte Carlo results show that inter-element correlations are not
183 spurious, since the measured correlation coefficient is significantly higher than the distribution of the
184 correlation coefficients between 10,000 randomly drawn El¹ concentrations/measured Ca concentration
185 and measured El²/Ca concentrations (Appendix B).

186 Furthermore, to test whether Sr/Ca_{cc} and Na/Ca_{cc} variability in *A. lessonii* is not caused by variability in
187 Mg content due to a potential closed sum effect (since high amounts of incorporated Mg cations could
188 reduce the Ca content of the shell and hence result in apparently elevated Sr/Ca_{cc} and Na/Ca_{cc}), we
189 calculated maximum variability due to the sole effect of Mg-substitution. For *A. lessonii*, variability
190 (standard deviation) of ±0.09 mmol/mol in Na/Ca_{cc} and ±0.016 mmol/mol in Sr/Ca_{cc} around the mean
191 could be caused by variability in Mg/Ca_{cc} (assuming Mg substitutes for Ca in the calcite lattice, and Mg
192 plus Ca approximates 1 mol per mol calcite). This may have influenced the Sr/Ca_{cc} and Na/Ca_{cc}
193 regression slopes over salinity and also the calculated inter-element correlation coefficients, but only
194 by a maximum of ±1% for both elements, which is considerably lower than the total observed
195 variability within the dataset of 16% and 9%, respectively.

196

197 2.3 Electron Microprobe Mapping

198 To investigate variation of element distribution across the chamber wall, a number of cultured
199 specimens were prepared for Electron Microprobe Analysis (EPMA). From each of the five salinity
200 conditions, six specimens from both species were selected and embedded in resin (Araldite 2020) in an
201 aluminum ring (diameter 1 cm) in a vacuum chamber. Samples were polished with a final polishing
202 step using a diamond emulsion with grains of 0.04 µm. This procedure resulted in exposure of a cross-
203 section of the foraminiferal chamber wall from which areas for EPMA mapping were selected
204 (Appendix A). These areas were selected for being perpendicular to the shell outer surface, resulting in
205 pores completely crossing the exposed chamber wall. Elemental distributions were mapped in
206 chambers prior to F-3 to study the element distribution across the various layers of calcite (lamella)
207 produced with the addition of each new chamber in rotaliid foraminifera (Reiss, 1957; 1960).
208 Elemental distribution in the shell wall was measured using a field emission Electron Probe Micro
209 Analyser (JEOL JXA-8530F HyperProbe) at 7.0kV with a dwell time of 350 ms, using a spot diameter
210 of 80 nm and a step size between 0.1538 µm and 0.4072 µm (130 x 130 pixels).

211 Spatial resolution of the EPMA mapping was determined using the software package CASINO (monte
212 Carlo Simulation of electron trajectory in Solids, v 2.48). With the input parameters identical as used
213 in our analysis (80 nm spot size, beam current 7 KeV, etc.), the simulated surface radius of the
214 backscattered electrons (i.e. the spatial resolution) equals 590 nm. Semi-quantitative El/Ca_{cc} profiles
215 were calculated by averaging the El/Ca_{cc} intensities parallel to the banding direction and applying a
216 constant calibration factor obtained from LA-ICP-MS measurements on the same specimen, similar to

217 the procedure of Eggins et al. (2004). We did not use the depth-resolved laser ablation-profiles for this
218 purpose, but used the average value from the profiles for correlation to the EPMA-derived intensities.

219 3. Results

220 3.1 Foraminiferal calcite element ratios and partitioning coefficients as a function of salinity

221 Per treatment, from lowest to highest salinity, average Na/Ca_{cc} of the newly formed calcite varied
222 between 9.3 and 10.8 mmol/mol for *A. lessonii* and between 4.7 and 6.4 mmol/mol (highest salinity)
223 for *A. tepida* (Fig. 1), with a corresponding partition coefficient (note that partition coefficients are
224 'apparent', not taking into account speciation/activity of Na) ranging from 1.90×10^{-4} to 2.20×10^{-4} and
225 from 0.97×10^{-4} to 1.30×10^{-4} for *Amphistegina* and *Ammonia*, respectively (Table 3). For both
226 species, sets of single-specimen Na/Ca_{cc} show slightly skewed distributions towards higher Na/Ca_{cc} for
227 all salinities (Kolmogorov-Smirnov test, at the 95% confidence level). Combining all specimens (based
228 on the average of single-spot measurements per specimen), Na/Ca_{cc} shows a positive linear relationship
229 with salinity for both *A. lessonii* and *A. tepida* ($\text{Na/Ca}_{\text{cc}} = 0.077 \pm 0.017 \times S + 7.13 \pm 0.60$, $F_{1,186} = 20.9$,
230 $p < 0.001$ for *A. lessonii* and $\text{Na/Ca}_{\text{cc}} = 0.064 \pm 0.013 \times S + 3.29 \pm 0.44$, $F_{1,172} = 25.9$, $p < 0.001$ for *A.*
231 *tepida*; Fig. 1). The observed average relative standard deviation between specimens in Na/Ca_{cc} at each
232 of the 5 salinities is 15% for *A. lessonii* and 20% for *A. tepida*. The variance in Na/Ca_{cc} between
233 individual specimens explained by salinity is $\eta^2 = 0.08$ for *A. lessonii* and $\eta^2 = 0.14$ for *A. tepida*.

234 Specimen's average Mg/Ca_{cc} and Sr/Ca_{cc} correlate positively with salinity in *A. tepida* ($\text{Mg/Ca}_{\text{cc}} =$
235 $0.060 \pm 0.011 \times S + 0.51 \pm 0.38$, $F_{1,172} = 29.9$, $p < 0.001$ and $\text{Sr/Ca}_{\text{cc}} = 0.014 \pm 12 \times 10^{-4} \times S + 1.00 \pm$
236 0.04 , $F_{1,337} = 254$, $p < 0.001$), whereas neither ratio correlates with salinity in *A. lessonii*. Average
237 relative standard deviations for the 5 salinity conditions per element are 27% for Mg/Ca_{cc} and 9% for
238 Sr/Ca_{cc} in *A. lessonii* and 32% in Mg/Ca_{cc} and 7% for Sr/Ca_{cc} for *A. tepida*. In *A. lessonii*, the
239 proportion of variance in Sr/Ca_{cc} explained by salinity is $\eta^2 = 0.04$ ($p < 0.01$) (Mg/Ca_{cc} not significant)
240 and for *A. tepida*, the proportion of variance in Sr/Ca_{cc} explained by salinity is $\eta^2 = 0.44$ and in Mg/Ca_{cc}
241 $\eta^2 = 0.19$ ($p < 0.001$).

242 Single-spot analyses on *Ammonia tepida* show that Na/Ca_{cc} and Mg/Ca_{cc} are significantly correlated
243 within the salinity treatments, except for condition S=30 (Fig. 2). For the individual salinity treatments,
244 single-spot Sr/Ca_{cc} and Mg/Ca_{cc}, as well as Na/Ca_{cc} and Sr/Ca_{cc} are not correlated significantly with
245 each other, except for S=25. Between salinity treatments, distributions in this species shift towards
246 higher Na/Ca_{cc}, Sr/Ca_{cc} and Mg/Ca_{cc} values with increasing salinity, although for the range between 30-
247 40 Na/Ca_{cc} distributions remain rather similar (Fig. 2). For *Amphistegina lessonii*, distributions of
248 Sr/Ca_{cc} and Mg/Ca_{cc} ratios overlap largely between salinities, and only Na/Ca_{cc} distributions shift
249 towards higher values (Fig. 2). Within each salinity condition however, single-spot Na/Ca_{cc}, Mg/Ca_{cc}
250 and Sr/Ca_{cc} in this species are positively correlated amongst each other, whereby the Na/Ca_{cc} intercept
251 of these relationships increases with increasing salinity (Fig. 2 and Appendix C).

252

253 **Table 3.** Average El/Ca_{cc} ratios of the foraminiferal calcite (based on average of average specimens
254 value per salinity (Sal) condition (S25-S45)) \pm standard error and corresponding apparent partitioning

255 coefficients, defined as $D_{El}=(El/Ca_{cc})/(El/Ca_{seawater})$ for *A. lessonii* (A.l.) and *A. tepida* (A.t). 'n/spots'
 256 stands for number of specimens and average number of spots per specimen.

Sal	n/spots	Na/Ca _{cc} mmol/mol	D _{Na}	Mg/Ca _{cc} mmol/mol	D _{Mg}	Sr/Ca _{cc} mmol/mol	D _{Sr}
<i>A.l.</i>							
S25	65/2.6	9.29±0.27	1.90×10 ⁻⁴	33.35±1.20	5.94×10 ⁻³	1.80±0.026	0.199
S30	74/1.9	9.47±0.21	1.90×10 ⁻⁴	32.10±1.20	5.64×10 ⁻³	1.74±0.020	0.189
S35	103/1.9	9.63±0.18	1.98×10 ⁻⁴	32.71±1.07	5.94×10 ⁻³	1.76±0.018	0.191
S40	50/2	10.25±0.31	2.11×10 ⁻⁴	35.22±2.60	6.27×10 ⁻³	1.74±0.034	0.184
S45	44/1.4	10.78±0.30	2.20×10 ⁻⁴	33.80±1.68	5.90×10 ⁻³	1.82±0.036	0.189
<i>A.t.</i>							
S25	109/2.5	4.75±0.11	0.97×10 ⁻⁴	1.90±0.06	3.40×10 ⁻⁴	1.34±0.016	0.148
S30	58/1.8	5.63±0.22	1.13×10 ⁻⁴	2.41±0.09	4.24×10 ⁻⁴	1.44±0.013	0.156
S35	59/1.8	5.58±0.19	1.15×10 ⁻⁴	2.85±0.24	5.17×10 ⁻⁴	1.50±0.012	0.163
S40	93/1.8	5.70±0.16	1.17×10 ⁻⁴	2.73±0.15	4.86×10 ⁻⁴	1.55±0.017	0.164
S45	201.3	6.39±0.37	1.31×10 ⁻⁴	3.27±0.27	5.70×10 ⁻⁴	1.61±0.038	0.168

257

258 3.2 Size and chamber effect on Na/Ca_{cc} and inter-specimen variance

259 Specimens of *A. lessonii* produced most new chambers at salinities of 25, 30 and 35, closest to the
 260 salinity in their “natural” habitat (Burgers Zoo aquarium, salinity (33.9-34.3; Ernst et al., 2011). Size
 261 averages are not significantly different between these salinity treatments, based on a Kruskal-Wallis
 262 test, whereas specimens grown at salinities 40 and 45 were significantly smaller than those from lower
 263 salinities, reflecting lower chamber addition rates over the course of the culturing experiment at higher
 264 salinity (Fig. 3). Combining all specimens, Na/Ca_{cc} is not significantly related to size in *A. lessonii*.
 265 Specimens of *A. tepida* produced less chambers at salinity 45, possibly because such a high salinity is
 266 probably close to its tolerance levels (Murray, 2014), even though this species is adapted to relatively
 267 large salinity shifts in their tidal flat habitat. Specimens in the lower salinity groups (25, 30, 35) grew
 268 larger compared to specimens grown at two the highest salinity groups (Fig. 3). Combining all
 269 specimens, Na/Ca_{cc} is significantly related to size in *A. tepida*, yet with a small slope (-0.003) and just
 270 within the 95% confidence interval (p=0.04).

271 Within each salinity tested, single-chambered Na/Ca_{cc} is slightly positively related to size for the
 272 specimens of *A. lessonii* cultured at salinities 25 (slope = 0.008, R² = 0.32, p < 0.01), 30 (slope = 0.002,
 273 R² = 0.11, p < 0.05 and 35 (slope = 0.005, R² = 0.18, p < 0.001). For the same species, Mg/Ca_{cc} is
 274 positively correlated to size at salinities 25, 30 and 35, with a similar slope of 0.03 (p < 0.05). Sr/Ca_{cc}
 275 also shows a positive relationship to size within salinities 25, 30 and 35 with slopes of 0.0007, 0.0003,
 276 0.0005 (p < 0.001) respectively. For *A. tepida*, there is only a slight negative correlation between size
 277 and Sr/Ca_{cc} for specimens cultured at salinity 25 (slope = 9.9×10⁻⁴, p < 0.001) and no significant
 278 correlation for the other conditions, or between size and Na/Ca_{cc} and Mg/Ca_{cc} in any of the salinities.

279 At the lowest salinity, Na/Ca_{cc} in the F-chamber (newest chamber) show slight (0.9 mmol/mol Na/Ca
 280 higher median) but significant higher values than the F-2 chambers for *A. lessonii* (multicompare test
 281 based on Kruskal-Wallis test, p < 0.05). For specimens of *A. lessonii* cultured at other salinities and for
 282 *A. tepida* at any of the salinities tested, there no significant correlations between Na/Ca_{cc} and chamber
 283 position were observed (note that only chamber positions F to F-2 were taken into account, as for the

284 lower chamber position sample numbers were insufficient). Furthermore, chamber position shows no
285 significant effect on Mg/Ca_{cc} and Sr/Ca_{cc}.

286 To further investigate the variance between and within individuals, a multiway ANOVA was
287 performed to investigate the effect on Na/Ca_{cc} per salinity condition. Inter-individual variance is
288 significant and larger than the variance between chamber groups and intra-individual variance in all
289 salinity groups, with the between individual variability accounting for $\eta^2 = 0.75 \pm 0.11 / 0.84 \pm 0.03$ of
290 the variance ($p < 0.001$) for *A. lessonii* and *A. tepida*, respectively. The variance due to chamber
291 position is not significant and the remaining intra-individual variance accounts for $\eta^2 = 0.09 \pm 0.05 /$
292 0.08 ± 0.05 for *A. lessonii* and *A. tepida*, respectively.

293 3.3 Elemental distributions in the chamber wall

294 EPMA maps of cross-sectioned chamber walls of *A. lessonii* show, within the resolution limits of the
295 technique, that bands of elevated Na/Ca_{cc} intensities overlap with zones of elevated Mg/Ca_{cc} (Fig. 4 and
296 appendix C). Mg bands show higher amplitudes than Na bands, but clearly coincide spatially.
297 Comparing EPMA maps with the backscatter SEM image of the exposed sections shows that the bands
298 with the highest Na/Ca_{cc} and Mg/Ca_{cc} occur in the proximity of the organic linings, which are clearly
299 visible in the backscatter SEM image (Fig. 4), with a number of high Na- and Mg-rich bands with
300 slightly lower maximum intensities occurring towards the outer chamber surface coinciding with
301 subsequent organic linings. For *A. tepida*, one band of elevated Mg/Ca_{cc} band is visible coinciding with
302 the POS with no clear Na/Ca_{cc} banding being detected.

303 4. Discussion

304 4.1 The effect of salinity and DIC on Na/Ca_{cc}, Mg/Ca_{cc} and Sr/Ca_{cc}

305 The single-specimen Na/Ca_{cc} data of the cultured *A. lessonii* and *A. tepida* both correlate positively
306 with salinity (Table 3, Fig. 1). This is in line with previous calibrations (for *Ammonia tepida*, Wit et al.,
307 2013; for cultured *Globigerinoides ruber*, Allen et al., 2016; for field-collected *G. ruber* and *G.*
308 *sacculifer*, Mezger et al., 2016). However, our Na/Ca-salinity calibration for *A. tepida* is somewhat less
309 sensitive than that observed earlier for the same species (Wit et al., 2013). An offset in Na/Ca_{cc} values
310 between calibrations for a single species has been reported previously for the planktonic *G. ruber* and
311 *G. sacculifer* (e.g. Mezger et al., 2016; Allen et al., 2016). Such an apparent discrepancy between
312 studies may be caused by differences between cultures or in situ conditions in one of the conditions not
313 focussed on (e.g. carbon chemistry, light intensity). Alternatively, subtle analytical differences (e.g.
314 differences in cleaning procedures), statistical reasons (for example differences in the number of
315 analyses or sample size) or the effect of genotypic variability on element incorporation (Sadekov et al.,
316 2016) may also play a role. Although the calibration presented here consists of many more data points
317 compared to those in Wit et al. (2013), we do not want to dismiss the latter as several parameters (like
318 cleaning procedures or the source of the seawater used for the culture media) inherently vary

319 (marginally) between studies. As such the difference observed between studies merely illustrates the
320 potential range for this species.

321 Contrasts in sensitivities such as observed for Na/Ca_{cc} between calibrations also apply to Mg/Ca_{cc} and
322 Sr/Ca_{cc}, both of which here show an increase with salinity in *A. tepida* but not in *A. lessonii* (Fig. 1).
323 Previous culturing experiments with *Ammonia tepida*, however, showed a smaller sensitivity of
324 Mg/Ca_{cc} to salinity (0.029-0.0044 mmol/mol change per salinity unit; Dissard et al., 2010) than that
325 reported here (0.06). Still, all these sensitivities are considerably lower than that reported in Kisakürek
326 et al. (2008) for the planktonic *G. ruber* (0.23 when Mg/Ca_{cc} is assumed to increase linearly with
327 salinity), but in the same range as that reported by Nürnberg et al. (1996) for *G. sacculifer* (0.05). The
328 sensitivity of Sr/Ca_{cc} to salinity in *A. tepida* (0.014; Table 3) is comparable to that for *O. universa*
329 (0.008; Lea et al., 2008), *G. ruber* (0.02; Kisakürek et al., 2008) and similar to the significant effect of
330 salinity on Sr incorporation in the same species (0.01-0.02, depending on temperature) found by
331 Dissard et al. (2010).

332 Sea water carbonate chemistry is an additional factor potentially affecting trace metal uptake (e.g. Lea
333 et al., 1999; Keul et al., 2017; Russell et al., 2004). Since salinity and dissolved inorganic carbon
334 concentration in the culture media co-varied in our experiments similar to the natural environment
335 (Table 1), Na/Ca_{cc} in our cultured specimens also correlates positively to sea water [DIC]. However,
336 sodium incorporation has been shown to be independent from changes in carbonate chemistry in
337 cultured *Amphistegina gibbosa* and several other benthic hyaline and porcelaneous species (Van Dijk et
338 al., 2017a). Additionally, Allen et al., (2016) also found no significant effect of carbonate chemistry
339 (i.e. varying [CO₃²⁻]) on Na incorporation in cultured *G. ruber*, suggesting that the variability in
340 Na/Ca_{cc} observed here in *A. lessonii* can be attributed to changes in salinity rather than [DIC].

341 However, future studies should disentangle the impacts of DIC and salinity on Na/Ca, in order to
342 increase proxy confidence in areas where Na/Ca and DIC relationships differ from the global average.

343 Previous studies showed that Sr/Ca_{cc} correlates positively to [DIC] in *A. tepida* (Keul et al., 2017),
344 which may account for part of the correlation between Sr/Ca_{cc} and salinity reported here for this
345 species. The published sensitivity of Sr/Ca_{cc} to [DIC] is approximately 2×10⁻⁵ mmol/mol change in
346 Sr/Ca_{cc} for every 1 μmol/kg change in [DIC], likely representing the maximum potential effect of DIC
347 on Sr partitioning given that others found no significant effect (Dissard et al., 2010a). For a change in
348 ~850 μmol/kg (Table 1), this would amount to an increase in Sr/Ca_{cc} of 0.019 mmol/mol (Keul et al.,
349 2017) over the salinity range studied here, thereby accounting for approximately 7% of the total
350 observed change in Sr/Ca_{cc} (Table 3). Inorganic carbon chemistry is known to affect growth rates and
351 shell weights in benthic foraminifera (Dissard et al., 2010a; Keul et al., 2013), which in turn, may
352 affect incorporation of Sr and Mg, hence providing a mechanistic link between inorganic carbon
353 chemistry and element partitioning.

354 El/Ca ratios of specimens of both species grown within each salinity condition are characterized by a
355 relatively large variability. In the overall data set, salinity only explains around 8% of the variation in
356 Na incorporation for *A. lessonii* and 14%, 19% and 44% of Na, Mg and Sr incorporation in *A. tepida*.
357 Nevertheless, for *A. lessonii*, the Na/Ca mean values (which translates to the values obtained from
358 traditional solution-ICP-MS) fit the regression model relatively well (Fig. 1). However, given the low

359 sensitivity, many specimens are required to reduce the uncertainty (Appendix E). This is reflected by
360 the relatively wide prediction bounds for the Na/Ca_{cc} salinity regressions, indicating an uncertainty
361 associated with a single Na/Ca_{cc} measurement. The relatively large inter-specimen variability in
362 element/Ca_{cc} ratios has been reported and discussed before (e.g. Sadekov et al., 2008; De Nooijer et al.,
363 2014a), but the cause for this variability remains to be identified.

364 4.2 Inter-specimen, inter-species and intra-shell El/Ca_{cc} variability

365 Single-chamber measurements show that Na/Ca_{cc} for both species varies between chambers (i.e.
366 specimens) with a RSD (Relative Standard Deviation) of 15%-20%, despite identical culture conditions
367 (Fig. 1). Since the analytical error on Na/Ca_{cc} accounts for approximately 2% (Table 2), a large portion
368 of the observed variability between specimens must be due to ontogeny and/or inter-specimen
369 differences in biomineralization controls (De Nooijer et al., 2014).

370 Foraminiferal shell size at salinities 40 and 45 are significantly smaller than those cultured at lower
371 salinities. When combining data from all salinities, however, there is no (*A. lessonii*) or only a very
372 small (*A. tepida*) negative correlation between Na/Ca_{cc} and shell size, as opposed to a more substantial
373 negative correlation as observed by Wit et al. (2013). In fact, there appears to be a growth optimum
374 around salinity 30-35, whereas growth at higher salinities might hampered (Fig. 2). This may indicate
375 that the earlier observed negative correlation between size and Na/Ca was the result of indirect co-
376 variation with salinity rather than a causal relationship resulting in lower Na/Ca values in smaller
377 specimens. This is corroborated by our observation that, for individuals grown at a similar salinity, the
378 relationship between Na/Ca_{cc} and size is either slightly positive or absent. Hence, size unlikely affects
379 the observed inter-specimen variability in Na/Ca_{cc}, which is supported by the absence of a correlation
380 between chamber position (and hence ontogenetic stage) and Na/Ca_{cc}. This implies that measuring
381 specimens of different size fractions or measuring different or multiple chambers should not
382 significantly affect the application of the Na/Ca_{cc} salinity proxy. However, sufficient specimens (n>30,
383 for an error margin <5% at the 95% confidence level; Sadekov et al., 2008; De Nooijer et al., 2014a)
384 are required for measurements. As most variability is between individuals rather than between
385 chambers (section 3.3), analyzing more chambers of the same specimen would increase the accuracy of
386 the measurement, but not improve the precision of the salinity estimate, given the large inter-specimen
387 variability. Without a major effect of ontogeny, physiological processes at the organismal level are
388 more likely to cause observed large inter-specimen variability in Na/Ca_{cc}, however these processes
389 remain poorly understood.

390 In *A. lessonii*, single-spot Na/Ca_{cc}, Sr/Ca_{cc} and Mg/Ca_{cc} are correlated amongst each other within each
391 salinity condition (Fig. 3). Correlation coefficients between the three element ratios are similar for the
392 different salinities, with superimposed an increase in the Na/Ca_{cc} relative to that of Mg/Ca_{cc} and Sr/Ca_{cc}
393 with increasing salinity (Appendix C). In contrast, single-spot Sr/Ca_{cc} and Mg/Ca_{cc} in *A. tepida* are not
394 correlated, whereas incorporation of all these elements increases significantly with salinity. Within
395 salinities Mg/Ca_{cc} and Na/Ca_{cc} are significantly correlated in 4 out of the 5 salinities, but with much
396 lower correlation coefficients compared to *A. lessonii* (Fig. 3 and Appendix C). However, between the

397 different salinities these elements are correlated in *A. tepida*, implying that for *A. tepida* salinity is one
398 of the actual parameters controlling element uptake.

399

400 Within conditions, the correlations between both Sr/Ca and Na/Ca with Mg/Ca in *A. lessonii* differs
401 from the correlation of Sr/Ca with Mg/Ca (correlation absent) and Na/Ca with Sr/Ca (weaker
402 correlation) for *A. tepida*. The differences between the correlations likely reflects differences in their
403 calcification pathways (e.g. transport of ions to the site of calcification) and/or might be explained by
404 differences in lattice strain due to the higher Mg-content in *A. lessonii*, whereas this effect is expected
405 to be less prominent in low-Mg species such as *A. tepida* (Evans et al., 2015). Differences in the
406 calcification pathway may also explain why Sr/Ca_{cc} and Mg/Ca_{cc} are correlated to salinity in *A. tepida*,
407 but not in *A. lessonii* (4.1). How elements are transported to the site of calcification and what is the role
408 of sea water-vacuolization, -leakage, trans-membrane-transport of ions, pH regulation and precipitation
409 rate and how this differs between species and specimens, remains to be discovered. The overall element
410 composition of the calcite precipitated by *A. lessonii* suggests that the calcification process of this
411 species may more closely resemble inorganic calcite precipitation, compared to that in *Ammonia tepida*
412 and other low-Mg calcite precipitating species. As a result, more elements (like Mg) are incorporated
413 and crystal lattice strain in intermediate-Mg calcite species is elevated, which may promote
414 incorporation of other elements through stress compensation (Mucci and Morse, 1983; Mewes et al.,
415 2015). This would explain the observed inter-element correlations within salinities. Another difference
416 between the species studied here may be caused by differences in CaCO₃ phase shifts during calcite
417 precipitation (e.g. Bots et al., 2012; De Yoreo et al., 2015). A metastable vaterite pre-cursor phase
418 recently found in two planktonic species may explain the low Mg incorporation relative to inorganic
419 calcite (Jacob et al., 2017). The higher Mg contents of *A. lessonii* could be related to the (partial)
420 absence of a vaterite-calcite transformation in this species. An Amorphous Calcium Carbonate (ACC)
421 pre-cursor phase has been observed in other marine biomineralising organisms (e.g. Weiner et al.,
422 2003; Giuffrè et al., 2015) and often been hypothesised to play a role in foraminiferal calcification
423 (Erez, 2003; De Nooijer et al., 2014b), although it has not yet been directly detected. A higher Mg
424 concentration at the site of calcification could hypothetically result in a phase shift from amorphous
425 calcium carbonate (ACC) directly into to calcite, whereby Mg is stabilizing the ACC, as described by
426 Littlewood et al. (2017). In inorganic calcite, the absence of a vaterite precursor phase also enhances
427 the incorporation of other metals incompatible to calcite, such as Sr (Littlewood et al., 2017) and a
428 similar process could hypothetically contribute to inter-species differences in element partitioning
429 similar to that observed here. Although the strong fractionation against Mg in *A. tepida* could reflect
430 double fractionation through a vaterite-calcite transformation (Jacob et al., 2017) the low-Mg content
431 might as well reflect a more enclosed site of calcification, whereby ions are mainly transported trans-
432 membrane (Nehrke et al., 2013; De Nooijer et al., 2014b). The experiments here do not allow
433 distinguishing between these (and other) potential mechanisms. Trans-membrane transport (TMT) of
434 Ca²⁺ and concomitant leakage of Mg²⁺ and Sr²⁺ might be more sensitive to differences in ionic strength
435 and element concentrations, hence possibly explaining the salinity effect on the incorporation of these
436 elements in *A. tepida* whereas it does not in *A. lessonii*, assuming that TMT relatively contributes more

437 to the supply of ions to the site of calcification in this species compared to *A. lessonii*, which might be
438 relatively more dependent on sea water vacuolisation. However, since there are many, both biotic and
439 abiotic, mechanisms that can (simultaneously) influence (coupled) element partitioning, it is
440 challenging to resolve the exact mechanism responsible for inter-specimen and inter-species
441 differences in El/Ca.

442 In both species, Mg is found to be elevated in bands located close to the primary organic sheet
443 and to other organic layers (Fig. 4), present in rotaliid species due to their lamellar calcification mode
444 (Reiss, 1957, 1960). This is similar to reports of within-chamber wall banding in many elements in
445 other species (Branson et al., 2016; Eggins et al., 2004; Sadekov et al., 2005; Paris et al., 2014; Spero et
446 al., 2015; Fehrenbacher et al., 2017; Kunioka et al., 2006; Steinhardt et al., 2015; Hathorne et al.,
447 2009). In planktonic species element banding has been related to diurnal light-dark cycles rather than
448 the addition of a new lamella with chamber addition (Spero et al., 2015; Fehrenbacher et al., 2017).
449 Whether in the species studied here, chamber addition (and hence element banding) is related to day
450 night cycles remains to be investigated. As in other studies, the Na- and Mg- bands are spatially
451 correlated (Fig. 4). For *Ammonia tepida*, the banding in both elements is less pronounced than for
452 *Amphistegina lessonii*, which is likely related to the (much) lower average Mg/Ca_{cc} and Na/Ca_{cc} ratios
453 in the former species. This inter-species difference observed in the Mg- and Na-maps implies that the
454 concentration of Mg and Na within the high concentration band is lower in *A. tepida* than in *A. lessonii*.
455 Alternatively, as the observations are close to the spatial resolution of the method, the observed pattern
456 could also be due to the band's width being smaller in *A. tepida* compared to *A. lessonii*. When
457 comparing the distribution of the two elements within one specimen, the Mg/Ca_{cc} bands are more
458 pronounced than those of Na/Ca_{cc}, particularly for *A. lessonii*, which is in line with the higher Mg of
459 this species (Fig. 4).

460 The spatial correlation between the intra-shell distributions of Mg and Na, associated to the organic
461 linings, suggests a coupled control on these elements during the calcification process, which is in line
462 with the observed inter-specimen correlations. This suggests that incorporation of these cations is
463 influenced by similar biomineralization mechanisms, related to sea water vacuolization (Erez, 2003;
464 Bentov and Erez, 2006), trans-membrane transport of elements (Nehrke et al., 2013), lattice-strain
465 effect (Evans et al., 2015) and/or metastable precursor phases (Jacob et al., 2017). The relative
466 contributions of these mechanisms might differ between species, resulting in the observed differences
467 in element incorporation and different inter-element correlations between species. Differences in the
468 efficiency of such processes between specimens might cause the observed inter-specimen variability,
469 whereas changes in these processes during the calcification time could be responsible for the observed
470 correlation between elements within the chamber wall.

471 5. Conclusions

472 By extending existing calibrations of the Na/Ca_{cc}-salinity proxy to the intermediate-Mg calcite
473 precipitating benthic foraminifer *Amphistegina lessonii*, we show that the Na/Ca_{cc} increase as a
474 function of salinity is similar to that in previously studied species. The absolute Na/Ca_{cc} for *A. lessonii*

475 is, however, higher than that in *Ammonia tepida*. In *A. tepida*, Mg/Ca_{cc} and Sr/Ca_{cc} are positively
476 correlated to salinity, whereas they are not impacted by salinity in *A. lessonii*. Within each salinity,
477 single chamber-Na/Ca_{cc} and Mg/Ca_{cc} are positively correlated in *A. tepida*, whereas single chamber-
478 Sr/Ca_{cc} is not correlated to either Mg/Ca_{cc} or Na in this species. For *A. lessonii*, all Sr/Ca_{cc}, Mg/Ca_{cc} and
479 Na/Ca_{cc} combinations are positively correlated at the single chamber level. Electron Microprobe
480 Analysis mapping of Na and Mg within chamber walls of cultured specimens shows that in *A. lessonii*,
481 Na/Ca_{cc} and Mg/Ca_{cc} occur in elevated bands in close proximity to the primary organic lining. For
482 specimens of *A. tepida*, Mg-banding appears similar to that in *A. lessonii*, whereas Na-banding is less
483 prominent in this species. These results suggest that biomineralization controls on incorporated
484 elements are different between species.

485 **Acknowledgements**

486 We would like to thank Wim Boer for assistance with LA-ICP-MS measurements, Patrick Laan for sea
487 water measurements and Karel Bakker for DIC measurements. We kindly thank Max Janse (Burgers'
488 Zoo, Arnhem) for providing stock specimens of *A. lessonii* and Kirsten Kooijmans (NIOZ) for
489 providing cultures of *Dunaliella salina*. Sergei Matveev is thanked for assistance with the Electron
490 Microprobe analysis and Leonard Bik for assistance with polishing the samples. This work was carried
491 out under the program of the Netherlands Earth System Science Centre (NESSC), financially supported
492 by the Ministry of Education, Culture and Science (OCW) (Grantnr. 024.002.001) and Darwin Centre
493 for Biogeosciences (program 3020).

494 **References**

- 495 Allen, K. A., Hönisch, B., Eggins, S. M., Haynes, L. L., Rosenthal, Y., and Yu, J.: Trace element
496 proxies for surface ocean conditions: A synthesis of culture calibrations with planktic foraminifera,
497 *Geochimica et Cosmochimica Acta*, 193, 197-221, 2016.
- 498 Anand, P., Elderfield, H., and Conte, M. H.: Calibration of Mg/Ca thermometry in planktonic
499 foraminifera from a sediment trap time series, *Paleoceanography*, 18, 10.1029/2002kpa000846, 2003.
- 500 Barker, S., Greaves, M., and Elderfield, H.: A study of cleaning procedures used for foraminiferal
501 Mg/Ca paleothermometry, *Geochemistry, Geophysics, Geosystems*, 4, 2003.
- 502 Bemis, B. E., Spero, H. J., Bijma, J., and Lea, D. W.: Reevaluation of the oxygen isotopic composition
503 of planktonic foraminifera: Experimental results and revised paleotemperature equations,
504 *Paleoceanography*, 13, 150-160, 10.1029/98pa00070, 1998.
- 505 Bentov, S., and Erez, J.: Impact of biomineralization processes on the Mg content of foraminiferal
506 shells: A biological perspective, *Geochemistry Geophysics Geosystems*, 7, 10.1029/2005gc001015,
507 2006.
- 508 Bots, P., Benning, L. G., Rodriguez-Blanco, J.-D., Roncal-Herrero, T., and Shaw, S.: Mechanistic
509 insights into the crystallization of amorphous calcium carbonate (ACC), *Crystal Growth & Design*, 12,
510 3806-3814, 2012.
- 511 Branson, O., Bonnin, E. A., Perea, D. E., Spero, H. J., Zhu, Z., Winters, M., Hönisch, B., Russell, A.
512 D., Fehrenbacher, J. S., and Gagnon, A. C.: Nanometer-Scale Chemistry of a Calcite Biomineralization
513 Template: Implications for Skeletal Composition and Nucleation, *Proceedings of the National
514 Academy of Sciences*, 201522864, 2016.
- 515 De Nooijer, L. J., Hathorne, E. C., Reichart, G. J., Langer, G., and Bijma, J.: Variability in calcitic
516 Mg/Ca and Sr/Ca ratios in clones of the benthic foraminifer *Ammonia tepida*, *Marine
517 Micropaleontology*, 107, 32-43, 10.1016/j.marmicro.2014.02.002, 2014.

518 De Nooijer, L. J., Brombacher, A., Mewes, A., Langer, G., Nehrke, G., Bijma, J., and Reichart, G.-J.:
519 Ba incorporation in benthic foraminifera, *Biogeosciences Discuss.*, 1-35, 2017.

520 De Yoreo, J. J., Gilbert, P. U., Sommerdijk, N. A., Penn, R. L., Whitlam, S., Joester, D., Zhang, H.,
521 Rimer, J. D., Navrotsky, A., and Banfield, J. F.: Crystallization by particle attachment in synthetic,
522 biogenic, and geologic environments, *Science*, 349, aaa6760, 2015.

523 Dickson, A., and Millero, F.: A comparison of the equilibrium constants for the dissociation of
524 carbonic acid in seawater media, *Deep Sea Research Part A. Oceanographic Research Papers*, 34,
525 1733-1743, 1987.

526 Dissard, D., Nehrke, G., Reichart, G.-J., and Bijma, J.: Impact of seawater pCO₂ on calcification and
527 Mg/Ca and Sr/Ca ratios in benthic foraminifera calcite: results from culturing experiments with
528 *Ammonia tepida*, *Biogeosciences*, 7, 81-93, 2010a.

529 Dissard, D., Nehrke, G., Reichart, G. J., and Bijma, J.: The impact of salinity on the Mg/Ca and Sr/Ca
530 ratio in the benthic foraminifera *Ammonia tepida*: Results from culture experiments, *Geochimica Et*
531 *Cosmochimica Acta*, 74, 928-940, 10.1016/j.gca.2009.10.040, 2010b.

532 Diz, P., Barras, C., Geslin, E., Reichart, G.-J., Metzger, E., Jorissen, F., and Bijma, J.: Incorporation of
533 Mg and Sr and oxygen and carbon stable isotope fractionation in cultured *Ammonia tepida*, *Marine*
534 *Micropaleontology*, 92, 16-28, 2012.

535 Duenas-Bohorquez, A., da Rocha, R. E., Kuroyanagi, A., de Nooijer, L. J., Bijma, J., and Reichart, G.
536 J.: Interindividual variability and ontogenetic effects on Mg and Sr incorporation in the planktonic
537 foraminifer *Globigerinoides sacculifer*, *Geochimica Et Cosmochimica Acta*, 75, 520-532,
538 10.1016/j.gca.2010.10.006, 2011.

539 Dueñas-Bohórquez, A., da Rocha, R. E., Kuroyanagi, A., Bijma, J., and Reichart, G.-J.: Effect of
540 salinity and seawater calcite saturation state on Mg and Sr incorporation in cultured planktonic
541 foraminifera, *Marine Micropaleontology*, 73, 178-189, 2009.

542 Eggins, S. M., Sadekov, A., and De Deckker, P.: Modulation and daily banding of Mg/Ca in *Orbulina*
543 *universa* tests by symbiont photosynthesis and respiration: a complication for seawater thermometry?,
544 *Earth and Planetary Science Letters*, 225, 411-419, 10.1016/j.epsl.2004.06.019, 2004.

545 Elderfield, H., and Ganssen, G.: Past temperature and delta O-18 of surface ocean waters inferred from
546 foraminiferal Mg/Ca ratios, *Nature*, 405, 442-445, 10.1038/35013033, 2000.

547 Erez, J.: The source of ions for biomineralization in foraminifera and their implications for
548 paleoceanographic proxies, *Biomineralization*, 54, 115-149, 10.2113/0540115, 2003.

549 Ernst, S., Janse, M., Renema, W., Kouwenhoven, T., Goudeau, M. L., and Reichart, G. J.: Benthic
550 foraminifera in a large indo-pacific coral reef aquarium, *Journal of Foraminiferal Research*, 41, 101-
551 113, 2011.

552 Fehrenbacher, J. S., Russell, A. D., Davis, C. V., Gagnon, A. C., Spero, H. J., Cliff, J. B., Zhu, Z., and
553 Martin, P.: Link between light-triggered Mg-banding and chamber formation in the planktic
554 foraminifera *Neogloboquadrina dutertrei*, Pacific Northwest National Laboratory (PNNL), Richland,
555 WA (US), Environmental Molecular Sciences Laboratory (EMSL), 2017.

556 Giuffrè, A. J., Gagnon, A. C., De Yoreo, J. J., and Dove, P. M.: Isotopic tracer evidence for the
557 amorphous calcium carbonate to calcite transformation by dissolution–reprecipitation, *Geochimica et*
558 *Cosmochimica Acta*, 165, 407-417, 2015.

559 Guillon, M., Meier, D. L., Allan, M. M., Heinrich, C. A., and Yardley, B. W.: Appendix A6: SILLS:
560 A MATLAB-based program for the reduction of laser ablation ICP-MS data of homogeneous materials
561 and inclusions, *Mineralogical Association of Canada Short Course*, 40, 328-333, 2008.

562 Hathorne, E. C., James, R. H., and Lampitt, R. S.: Environmental versus biomineralization controls on
563 the intratest variation in the trace element composition of the planktonic foraminifera *G. inflata* and *G.*
564 *scitula*, *Paleoceanography*, 24, 10.1029/2009pa001742, 2009.

565 Hönsch, B., Allen, K. A., Lea, D. W., Spero, H. J., Eggins, S. M., Arbuszewski, J., Rosenthal, Y.,
566 Russell, A. D., and Elderfield, H.: The influence of salinity on Mg/Ca in planktic foraminifers–
567 Evidence from cultures, core-top sediments and complementary δ 18 O, *Geochimica et Cosmochimica*
568 *Acta*, 121, 196-213, 2013.

569 Jacob, D., Wirth, R., Agbaje, O., Branson, O., and Eggins, S.: Planktic foraminifera form their shells
570 via metastable carbonate phases, *Nature Communications*, 8, 1265, 2017.

571 Keul, N., Langer, G., de Nooijer, L. J., and Bijma, J.: Effect of ocean acidification on the benthic
572 foraminifera *Ammonia* sp. is caused by a decrease in carbonate ion concentration, *Biogeosciences*, 10,
573 6185-6198, 2013.

574 Keul, N., Langer, G., Thoms, S., de Nooijer, L. J., Reichart, G.-J., and Bijma, J.: Exploring
575 foraminiferal Sr/Ca as a new carbonate system proxy, *Geochimica et Cosmochimica Acta*, 202, 374-
576 386, 2017.

577 Kısakürek, B., Eisenhauer, A., Böhm, F., Garbe-Schönberg, D., and Erez, J.: Controls on shell Mg/Ca
578 and Sr/Ca in cultured planktonic foraminifera, *Globigerinoides ruber* (white), *Earth and Planetary*
579 *Science Letters*, 273, 260-269, 2008.

580 Kunioka, D., Shirai, K., Takahata, N., Sano, Y., Toyofuku, T., and Ujiie, Y.: Microdistribution of
581 Mg/Ca, Sr/Ca, and Ba/Ca ratios in *Pulleniatina obliquiloculata* test by using a NanoSIMS: Implication
582 for the vital effect mechanism, *Geochemistry Geophysics Geosystems*, 7, 10.1029/2006gc001280,
583 2006.

584 Lea, D. W., Mashiotta, T. A., and Spero, H. J.: Controls on magnesium and strontium uptake in
585 planktonic foraminifera determined by live culturing, *Geochimica Et Cosmochimica Acta*, 63, 2369-
586 2379, 10.1016/s0016-7037(99)00197-0, 1999.

587 Littlewood, J. L., Shaw, S., Peacock, C. L., Bots, P., Trivedi, D., and Burke, I. T.: Mechanism of
588 Enhanced Strontium Uptake into Calcite via an Amorphous Calcium Carbonate Crystallization
589 Pathway, *Crystal Growth & Design*, 17, 1214-1223, 2017.

590 Mewes, A., Langer, G., de Nooijer, L. J., Bijma, J., and Reichart, G. J.: Effect of different seawater
591 Mg²⁺ concentrations on calcification in two benthic foraminifers, *Marine Micropaleontology*, 113, 56-
592 64, 10.1016/j.marmicro.2014.09.003, 2014.

593 Mewes, A., Langer, G., Reichart, G.-J., de Nooijer, L. J., Nehrke, G., and Bijma, J.: The impact of Mg
594 contents on Sr partitioning in benthic foraminifers, *Chemical Geology*, 412, 92-98, 2015.

595 Mezger, E., Nooijer, L., Boer, W., Brummer, G., and Reichart, G.: Salinity controls on Na
596 incorporation in Red Sea planktonic foraminifera, *Paleoceanography*, 2016.

597 Mucci, A., and Morse, J. W.: The incorporation of Mg²⁺ and Sr²⁺ into calcite overgrowths: influences
598 of growth rate and solution composition, *Geochimica et Cosmochimica Acta*, 47, 217-233,
599 [http://dx.doi.org/10.1016/0016-7037\(83\)90135-7](http://dx.doi.org/10.1016/0016-7037(83)90135-7), 1983.

600 Murray, J. W.: *Ecology and palaeoecology of benthic foraminifera*, Routledge, 2014.

601 Nehrke, G., Keul, N., Langer, G., de Nooijer, L. J., Bijma, J., and Meibom, A.: A new model for
602 biomineralization and trace-element signatures of Foraminifera tests, *Biogeosciences*, 10, 6759-6767,
603 10.5194/bg-10-6759-2013, 2013.

604 Nürnberg, D., Bijma, J., and Hemleben, C.: Assessing the reliability of magnesium in foraminiferal
605 calcite as a proxy for water mass temperatures, *Geochimica et Cosmochimica Acta*, 60, 803-814, 1996.

606 Paris, G., Fehrenbacher, J. S., Sessions, A. L., Spero, H. J., and Adkins, J. F.: Experimental
607 determination of carbonate-associated sulfate delta S-34 in planktonic foraminifera shells,
608 *Geochemistry Geophysics Geosystems*, 15, 1452-1461, 10.1002/2014gc005295, 2014.

609 Raja, R., Saraswati, P. K., and Iwao, K.: A field-based study on variation in Mg/Ca and Sr/Ca in larger
610 benthic foraminifera, *Geochemistry, Geophysics, Geosystems*, 8, 2007.

611 Reichart, G. J., Jorissen, F., Anschutz, P., and Mason, P. R. D.: Single foraminiferal test chemistry
612 records the marine environment, *Geology*, 31, 355-358, 10.1130/0091-
613 7613(2003)031<0355:sftcrt>2.0.co;2, 2003.

614 Reiss, Z.: The Bilamellidea, nov. superfam., and remarks on Cretaceous globorotaliids, *Contrib*
615 *Cushman Found Foraminifera Res*, 8, 127-145, 1957.

616 Reiss, Z.: Structure of so-called Eponides and some other rotaliiform foraminifera, State of Israel,
617 Ministry of Development, Geological Survey, 1960.

618 Rohling, E., and Hilgen, F.: The eastern Mediterranean climate at times of sapropel formation: a
619 review, *Netherlands Journal of Geosciences/Geologie en Mijnbouw*, 2007.

620 Rosenthal, Y., Lohmann, G. P., Lohmann, K. C., and Sherrell, R. M.: Incorporation and preservation of
621 Mg in *Globigerinoides sacculifer*: Implications for reconstructing the temperature and O-18/O-16 of
622 seawater, *Paleoceanography*, 15, 135-145, 10.1029/1999pa000415, 2000.

623 Russell, A. D., Hönisch, B., Spero, H. J., and Lea, D. W.: Effects of seawater carbonate ion
624 concentration and temperature on shell U, Mg, and Sr in cultured planktonic foraminifera, *Geochimica*
625 *et Cosmochimica Acta*, 68, 4347-4361, 2004.

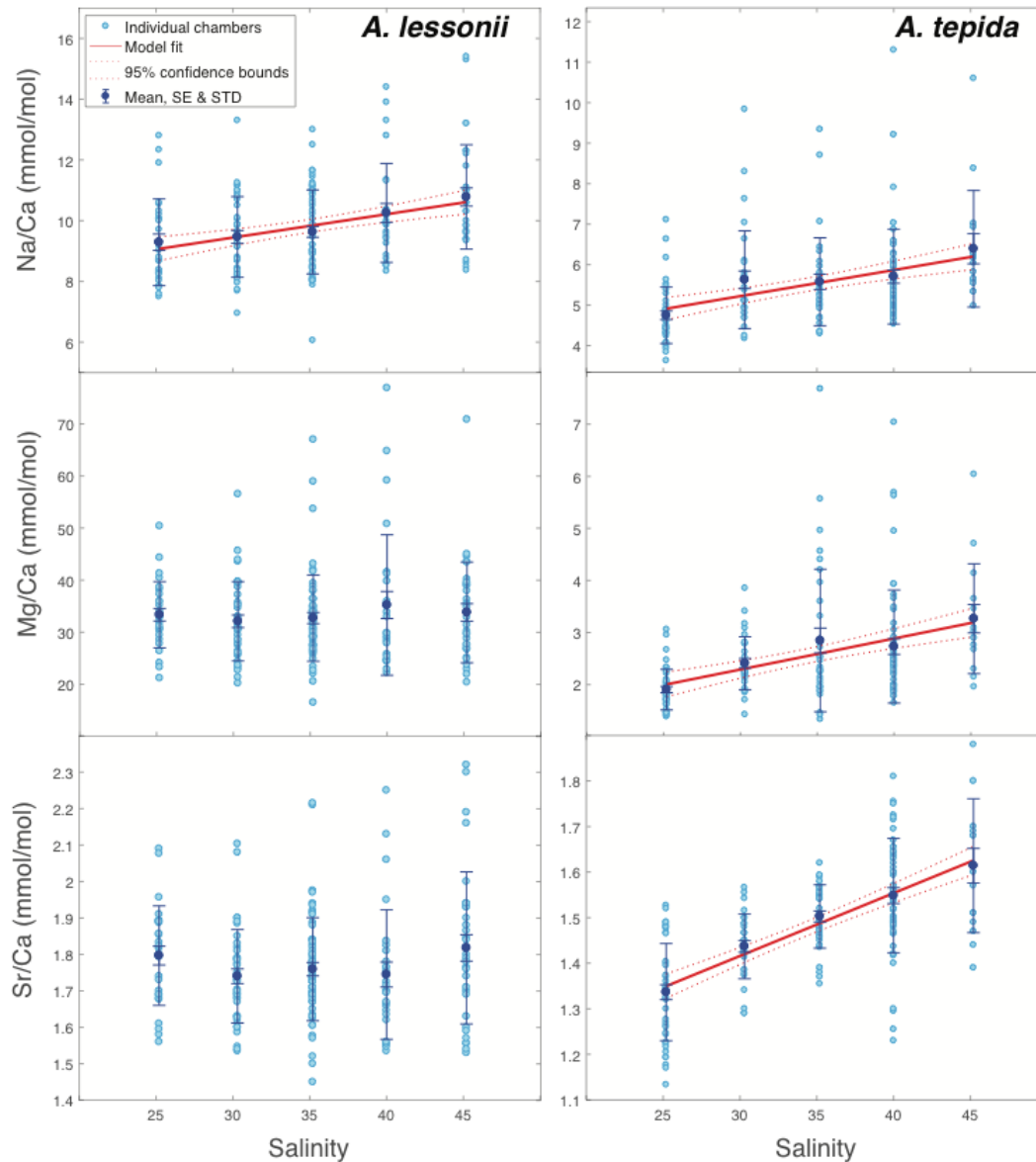
626 Sadekov, A. Y., Eggins, S. M., and De Deckker, P.: Characterization of Mg/Ca distributions in
627 planktonic foraminifera species by electron microprobe mapping, *Geochemistry Geophysics*
628 *Geosystems*, 6, 10.1029/2005gc000973, 2005.

629 Segev, E., and Erez, J.: Effect of Mg/Ca ratio in seawater on shell composition in shallow benthic
630 foraminifera, *Geochemistry Geophysics Geosystems*, 7, 10.1029/2005gc000969, 2006.

631 Spero, H. J., Eggins, S. M., Russell, A. D., Vetter, L., Kilburn, M. R., and Hönisch, B.: Timing and
632 mechanism for intratest Mg/Ca variability in a living planktic foraminifer, *Earth and Planetary Science*
633 *Letters*, 409, 32-42, 10.1016/j.epsl.2014.10.030, 2015.

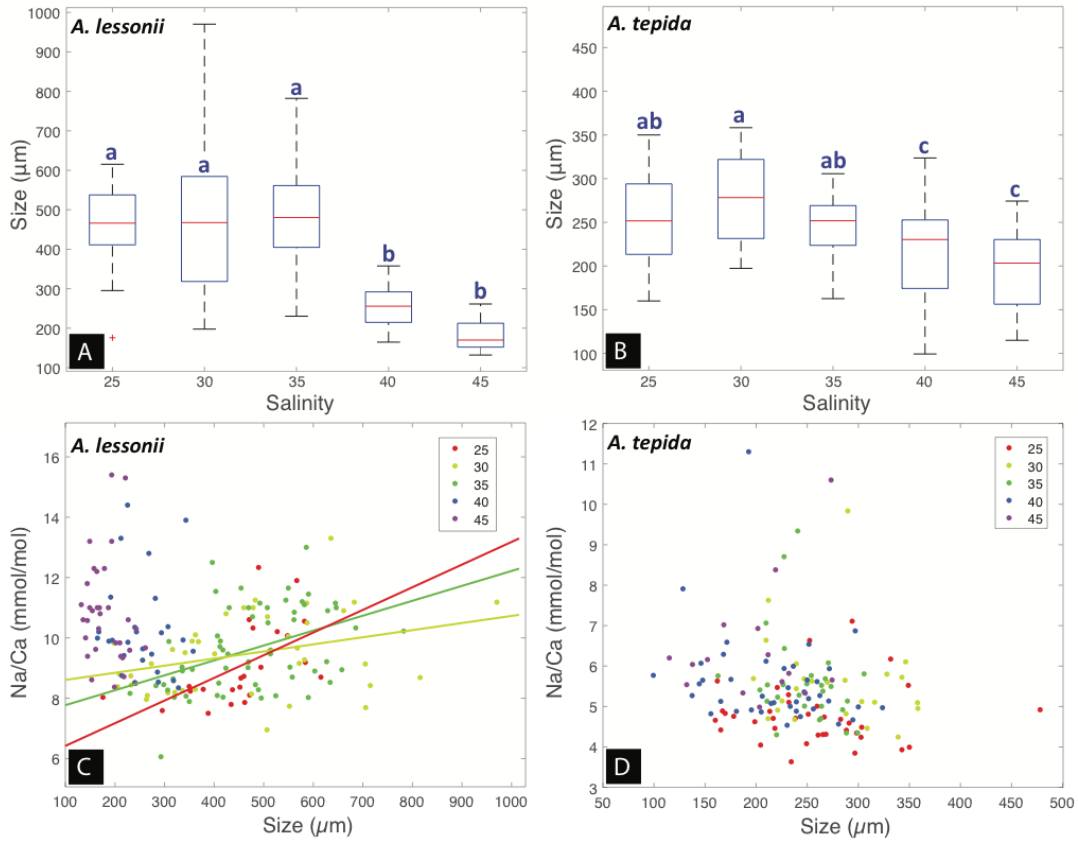
634 Steinhardt, J., de Nooijer, L. L., Brummer, G. J., and Reichart, G. J.: Profiling planktonic foraminiferal
635 crust formation, *Geochemistry, Geophysics, Geosystems*, 16, 2409-2430, 2015.

636 Stoll, M. H. C., Bakker, K., Nobbe, G. H., and Haese, R. R.: Continuous-flow analysis of dissolved
637 inorganic carbon content in seawater, *Analytical Chemistry*, 73, 4111-4116, 10.1021/ac010303r, 2001.
638 Terakado, Y., Ofuka, Y., and Tada, N.: Rare earth elements, Sr, Ba, Fe, and major cation
639 concentrations in some living foraminiferal tests collected from Iriomote Island, Japan: An exploration
640 for trace element behavior during biogenic calcium carbonate formation, *Geochemical Journal*, 44,
641 315-322, 2010.
642 Toyofuku, T., Kitazato, H., Kawahata, H., Tsuchiya, M., and Nohara, M.: Evaluation of Mg/Ca
643 thermometry in foraminifera: Comparison of experimental results and measurements in nature,
644 *Paleoceanography*, 15, 456-464, 10.1029/1999pa000460, 2000.
645 Toyofuku, T., Suzuki, M., Suga, H., Sakai, S., Suzuki, A., Ishikawa, T., de Nooijer, L. J., Schiebel, R.,
646 Kawahata, H., and Kitazato, H.: Mg/Ca and $\delta^{18}\text{O}$ in the brackish shallow-water benthic foraminifer
647 *Ammonia 'beccarii'*, *Marine Micropaleontology*, 78, 113-120, 2011.
648 van Aken, H. M.: Variability of the water temperature in the western Wadden Sea on tidal to centennial
649 time scales, *Journal of Sea Research*, 60, 227-234, 2008.
650 Van Dijk, I., de Nooijer, L. J., and Reichart, G.-J.: Trends in element incorporation in hyaline and
651 porcelaneous foraminifera as a function of pCO₂, *Biogeosciences*, 14, 497, 2017a.
652 Van Dijk, I., de Nooijer, L. J., Wolthers, M., and Reichart, G.-J.: Impacts of pH and [CO₃²⁻] on the
653 incorporation of Zn in foraminiferal calcite, *Geochimica et Cosmochimica Acta*, 197, 263-277, 2017b.
654 Van Heuven, S., Pierrot, D., Rae, J., Lewis, E., and Wallace, D.: MATLAB program developed for
655 CO₂ system calculations, Oak Ridge, Tennessee: Carbon Dioxide Information Analysis Center, Oak
656 Ridge National Laboratory, US Department of Energy. pp. ORNL/CDIAC-105b, 2011.
657 Weiner, S., Levi-Kalisman, Y., Raz, S., and Addadi, L.: Biologically formed amorphous calcium
658 carbonate, *Connective Tissue Research*, 44, 214-218, 2003.
659 Wit, J. C., de Nooijer, L. J., Wolthers, M., and Reichart, G. J.: A novel salinity proxy based on Na
660 incorporation into foraminiferal calcite, *Biogeosciences*, 10, 6375-6387, 10.5194/bg-10-6375-2013,
661 2013.
662



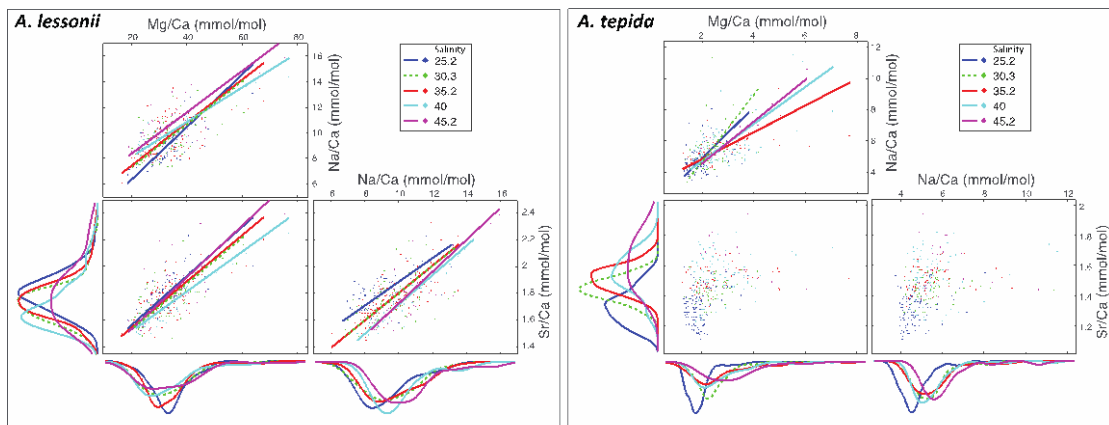
663

664 **Figure 1.** Foraminiferal Na/Ca_{cc}, Mg/Ca_{cc} and Sr/Ca_{cc} versus salinity. Light blue dots represent the average
 665 per specimen (n= 359 for *A. lessonii*, n=339 for *A. tepida*, with 2-3 measured chambers per individual), dark
 666 blue dots indicate the mean, with inner error bars indicating the standard error and outer error bars the
 667 standard deviation for each treatment. The linear regression model (red line) is based on the individuals'
 668 mean, with the 95% confidence interval of the regression in dashed lines.



669

670 **Figure 2.** Boxplots (Panel A and B) showing the size distributions (median, 1st and 3rd quartiles, minimum
 671 and maximum values) for each salinity condition, **n=24, 40, 60, 27, 33** for *A. lessonii* and **n= 38, 24, 28, 41, 15**
 672 **for *A. tepida***. Letters (a, b, c) indicate significant different population means, based on ANOVA ($p < 0.001$).
 673 Panel C and D show the Na/Ca values against size measurements per individual, colour coded per salinity
 674 condition (see legend), for *A. lessonii* and *A. tepida*. Significant linear regression lines are plotted for *A.*
 675 *lessonii*.

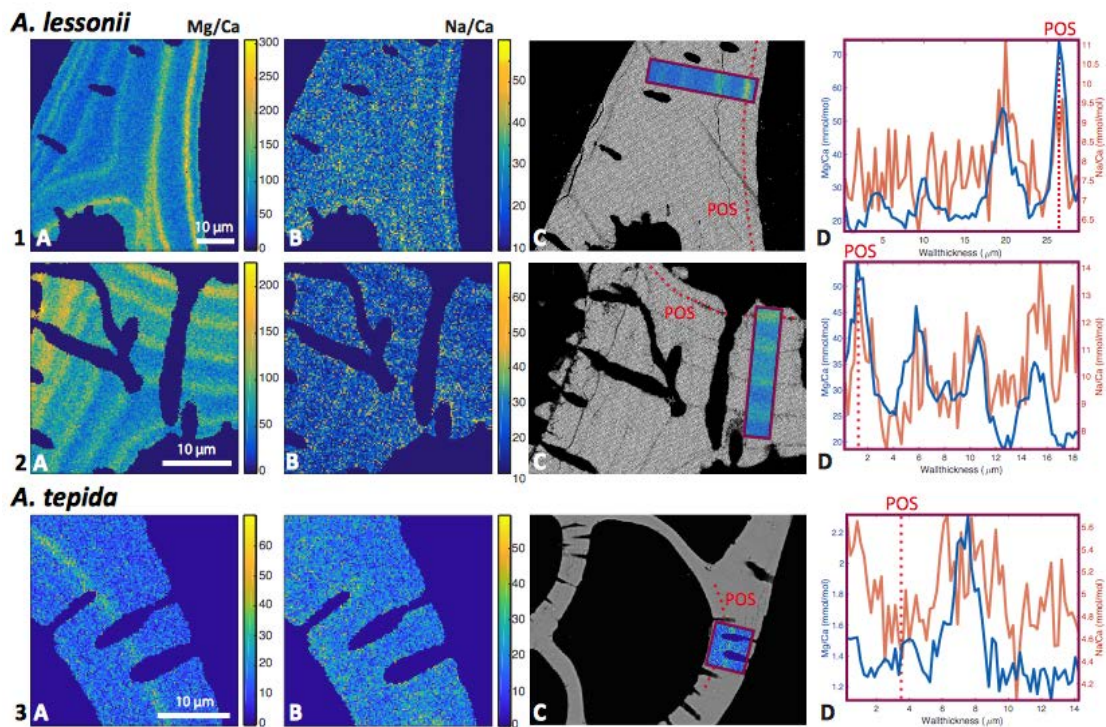


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677 **Figure 3.**

678 **Individual chamber LA-ICP-MS analyses showing correlations between foraminiferal Mg/Ca_{cc}, Sr/Ca_{cc} and**
 679 **Na/Ca_{cc} for *A. tepida* (left) and *A. lessonii* (right) per salinity condition. Significant orthogonal linear**
 680 **regressions** are indicated with a line, colour coded for salinity (see legend). Correlation coefficients, slope
 681 and intercepts of these regressions can be found in Appendix C. In short, within salinity conditions, element

682 ratios are strongly correlated with each other in *A. lessonii*, whereas in *A. tepida*, element ratios do not
 683 (strongly) correlate with each other. When combining all single-spot data in *A. tepida*, element ratios
 684 correlate amongst each other because the incorporation of all three elements increases with salinity, shifting
 685 the distributions to higher values. In *A. lessonii*, only the Na/Ca_{acc} distributions shift towards higher values
 686 with increasing salinity, whereas Mg/Ca_{acc} and Sr/Ca_{acc} distributions are relatively similar between salinity
 687 conditions.



688
 689 **Figure 4. Foraminiferal Mg/Ca_{acc} (A panels; left) and Na/Ca_{acc} (B panels) intensity ratio maps, obtained with**
 690 **EPMA, for two specimens of *A. lessonii* grown at a salinity of 30 (row 1) and 25 (row 2) and one specimen of**
 691 ***A. tepida* (row 3). D panels (right) show profiles for Mg/Ca (blue) and Na/Ca (red), based on averaged**
 692 **EPMA ratios scaled to LA-ICP-MS measurements of the same specimen, of an averaged lateral profile area**
 693 **through the chamber wall perpendicular to the lamella separated by organic linings (purple rectangles C).**
 694 **The transect area is indicated with a purple rectangle, on top of a backscatter SEM image (C), showing that**
 695 **the high EI/Ca bands overlap with the primary organic sheet (POS, marked with dashed red line) and**
 696 **subsequent organic linings. See Appendix D for the results for three more specimens.**

697 **Appendix**

698 **Appendix A.**

699 SEM image of a specimen of *A. lessonii* showing LA-ICP-MS measurement spots (left) and SEM images of
 700 specimens of *A. lessonii* (upper right) and *A. tepida* (lower right) embedded in resin and polished for
 701 Electron Probe Micro Analysis, the mapping area is depicted with a white box.

702 **Appendix B.**

703 Results of the Monte Carlo analysis showing that the measured correlation coefficients for the inter-
704 specimen correlations between the measured El^1/Ca_{cc} and El^2/Ca_{cc} are not caused by a spurious correlation
705 due to the common denominator Ca_{cc} , showing that the measured correlation coefficient is significantly
706 higher than the distribution of the correlation coefficients between 10,000 randomly drawn El^1
707 concentrations/measured Ca concentration and measured El^2/Ca concentrations. This test is based on the
708 concentration results from a single LA-ICP-MS session with specimens of *A. lessonii* cultured at a salinity of
709 35.

710 Appendix C.

711 Results for the orthogonal regressions testing the correlations between single-spot El^1/Ca and El^2/Ca , within
712 each salinity conditions, for *A. lessonii* and *A. tepida*.

713 Appendix D.

714 Foraminiferal Mg/Ca_{cc} and Na/Ca_{cc} (A and B, E and F) intensity ratio maps, obtained with EPMA, for two
715 specimens of *A. lessonii* grown at a salinity of 30 (A-D) and 35 (E-H). Panel D and H show profiles for
716 Mg/Ca (blue) and Na/Ca (red), based on averaged EPMA ratios scaled to LA-ICP-MS measurements in D
717 and on EPMA count ratios in H (no La-ICP-MS data available for this specimen), of an averaged transect
718 area through the chamber wall perpendicular to the POS. The transect areas (purple rectangles) are
719 indicated on top of backscatter SEM images (C and G), showing that the high El/Ca bands overlap with the
720 primary organic sheet (POS, in dashed red line in C, not clear in G)) and subsequent organic linings.

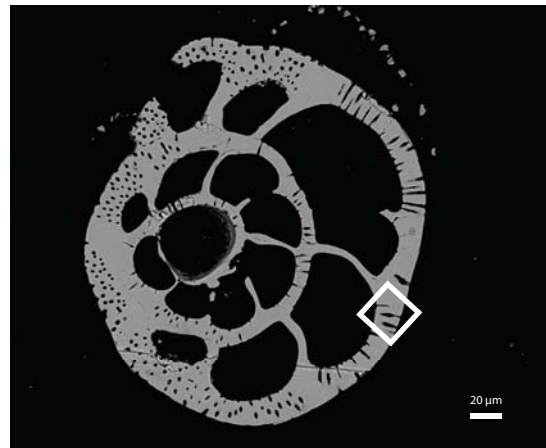
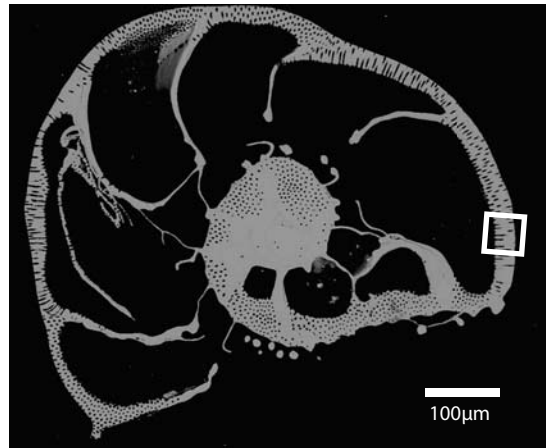
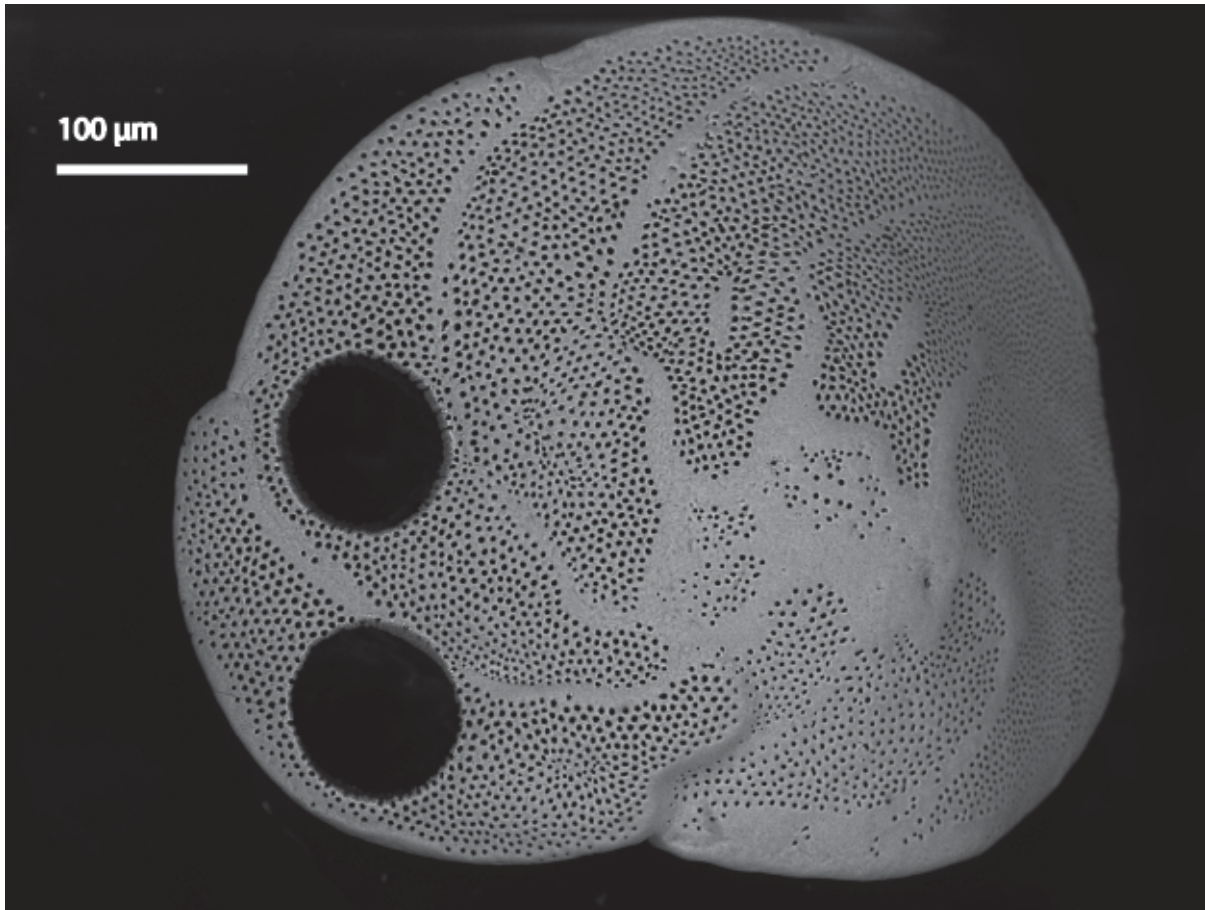
721 Appendix E.

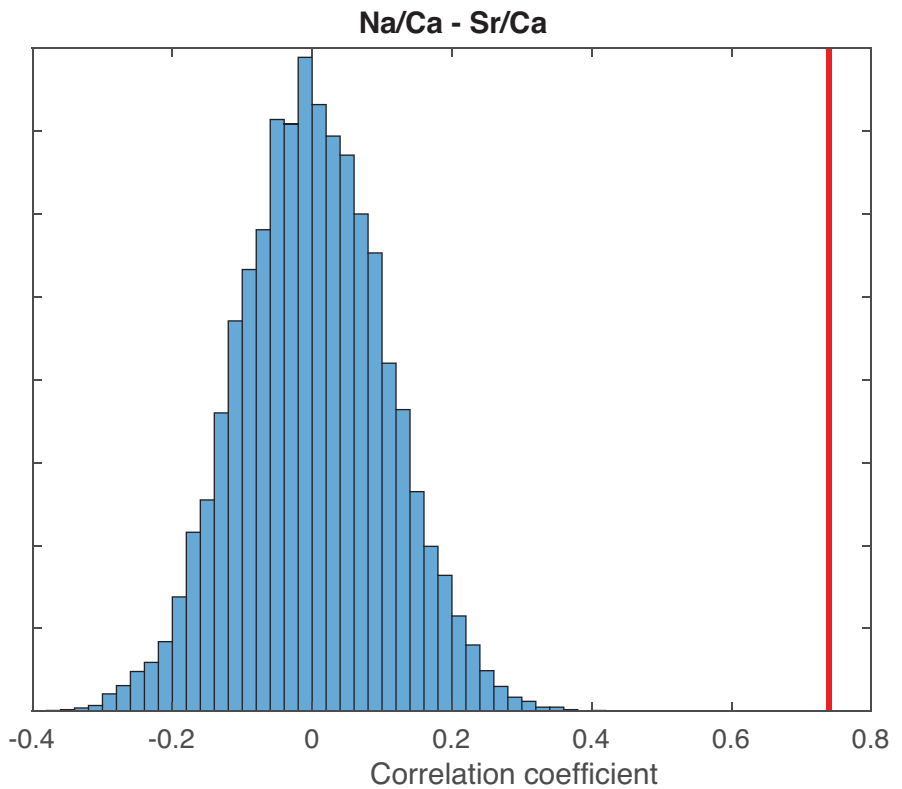
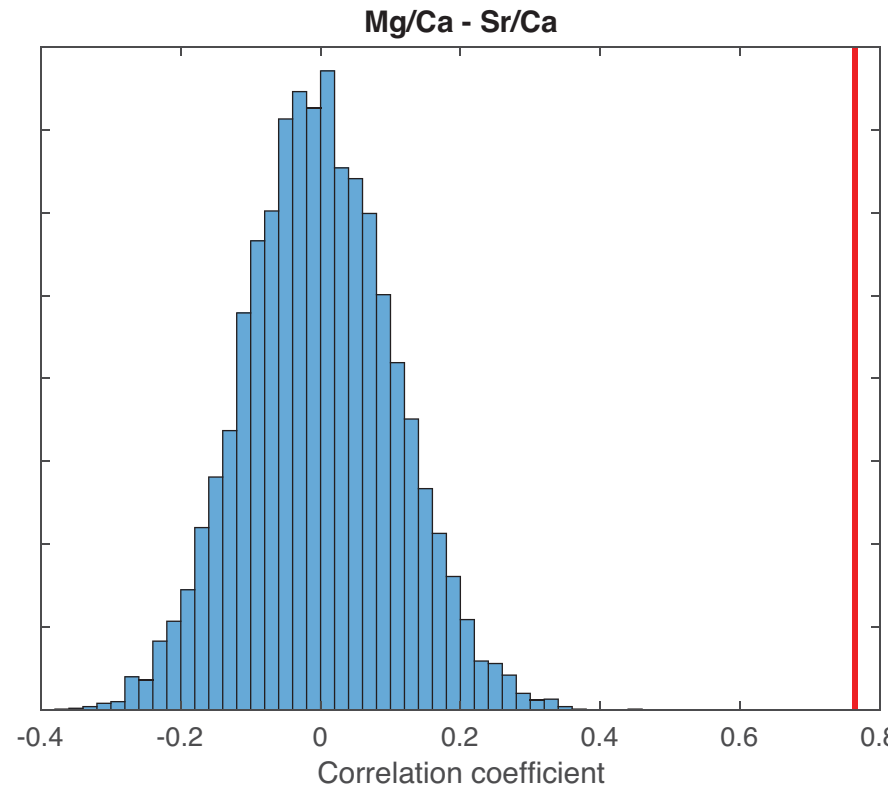
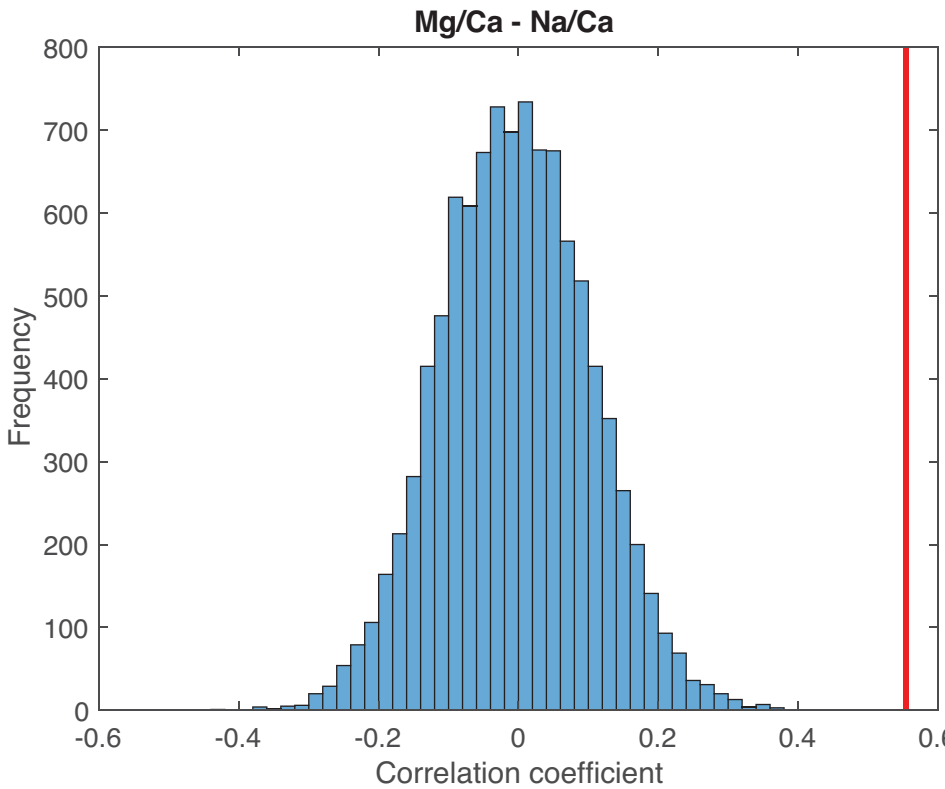
722 Figure showing the relationship between the salinity uncertainty and number of measured specimens for the
723 Na/Ca_{cc} - salinity calibration of *A. lessonii*, calculated following Eq. (1):

$$724 \text{ Salinity uncertainty} = (2 \times \text{RSD} \times \text{Number of specimens}^{-0.5}) / \text{Sensitivity}, \quad (1)$$

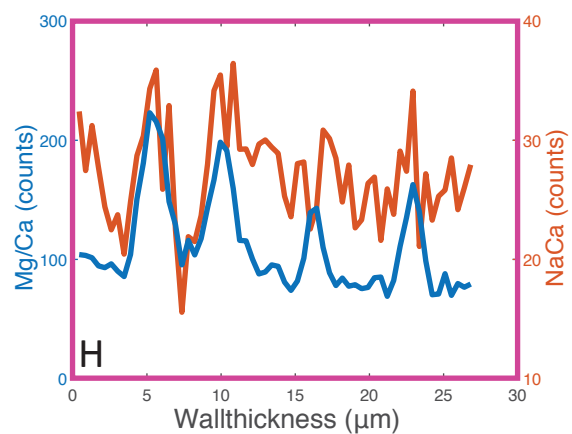
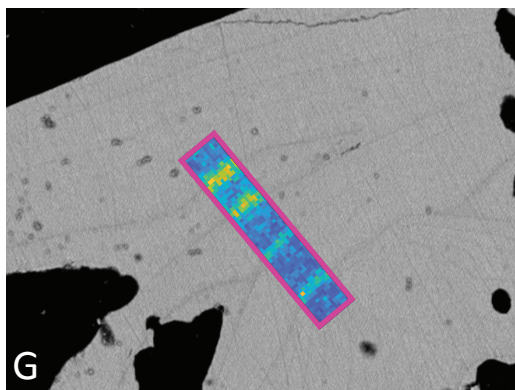
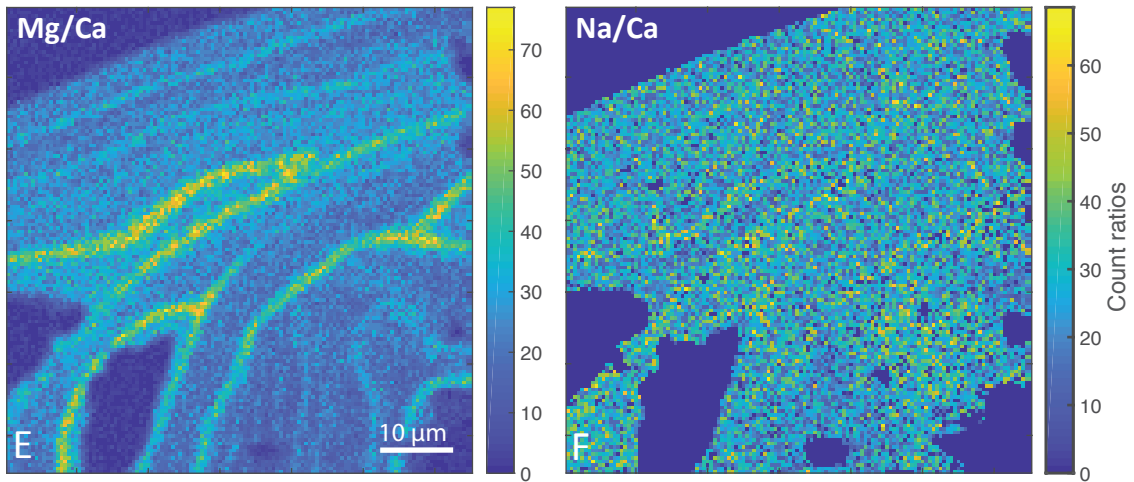
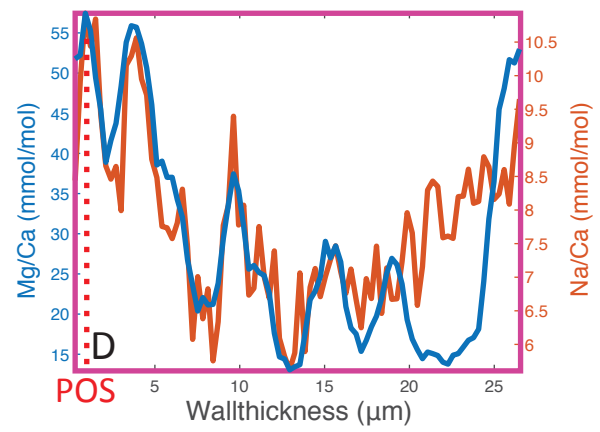
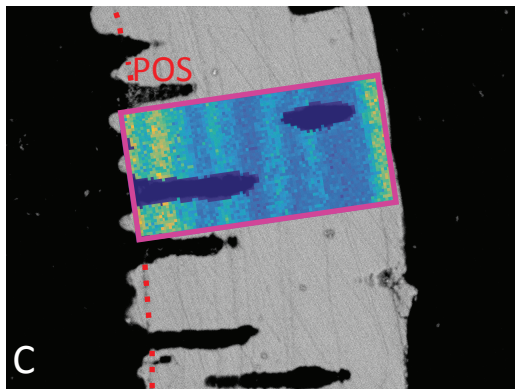
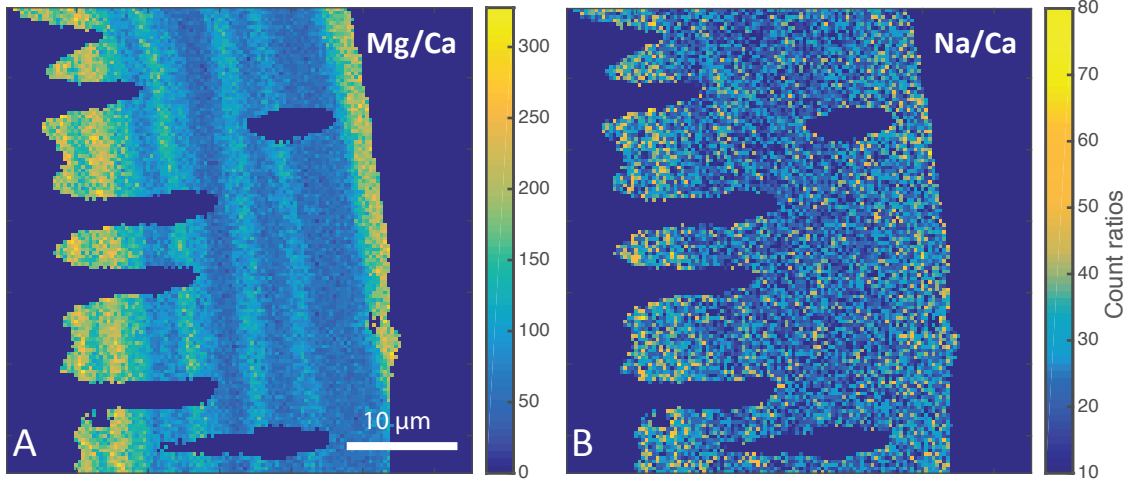
725 whereby sensitivity is the slope of the calibration.

726





A. lessonii



Orthogonal regression results

A. lessonii

p-value (x,y)	r	Slope	y-Intercept
(Mg/Ca, Na/Ca)			
p<0.001	0,71	0,21	2,21
p<0.001	0,52	0,17	3,89
p<0.001	0,57	0,17	4,06
p<0.001	0,89	0,14	5,44
p<0.001	0,69	0,16	5,12
(Mg/Ca, Sr/Ca)			
p<0.001	0,63	0,02	1,19
p<0.001	0,64	0,02	1,18
p<0.001	0,76	0,02	1,19
p<0.001	0,90	0,01	1,22
p<0.001	0,73	0,02	1,14
(Na/Ca, Sr/Ca)			
p<0.001	0,58	0,09	1,00
p<0.001	0,47	0,10	0,78
p<0.001	0,72	0,10	0,77
p<0.001	0,94	0,11	0,62
p<0.001	0,80	0,12	0,53

A. tepida

p-value (x,y)	r	Slope	y-Intercept
(Mg/Ca, Na/Ca)			
p<0.001	0,35	1,64	1,63
p>0.05	0,24	2,14	0,41
p<0.001	0,62	0,86	3,14
p<0.001	0,53	1,19	2,35
p<0.001	0,84	1,34	1,89
(Mg/Ca, Sr/Ca)			
p<0.01	0,28	0,23	0,89
p>0.05	0,20	0,13	1,13
p>0.05	0,18	0,06	1,33
p>0.05	0,17	0,13	1,19
p>0.05	-0,28	-0,16	2,12
(Na/Ca, Sr/Ca)			
p<0.01	0,29	0,14	0,66
p>0.05	0,10	0,06	1,10
p>0.05	0,23	0,07	1,10
p>0.05	0,18	0,11	0,94
p>0.05	-0,32	-0,12	2,34

