

1 **Author response to Associate Editor**

2 Associate Editor Decision: Publish subject to minor revisions (review by editor) (02
3 May 2018) by Emilio Marañón

4 Comments to the Author:

5 Dear authors,

6 Thank you for submitting a revised version of your manuscript. You have addressed
7 adequately all the points that had been raised. However, there seems to be a problem
8 with the reported values of the C:N ratio (Fig. 3F). Given the values of POC and PON
9 concentration, the molar C:N ratio should range roughly between 10-16. Please correct
10 both the figure and the text describing it (lines 326-331). Also, the units of C:N could be
11 simplified to the more commonly used molC:molN.

12 Best regards,

13 Emilio Marañón

14 **Response:** the POC:PON data has been corrected in **Fig. 3** and in the manuscript text
15 from line 326 with the correct units (molC:molN)

16

17 **Effects of elevated CO₂ and temperature on phytoplankton community**
18 **biomass, species composition and photosynthesis during an**
19 **experimentally induced autumn bloom in the Western English**
20 **Channel**

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25

26 **Abstract**

27 The combined effects of elevated pCO₂ and temperature were investigated during an
28 experimentally induced autumn phytoplankton bloom *in vitro* sampled from the Western
29 English Channel (WEC). A full factorial 36-day microcosm experiment was conducted under

30 year 2100 predicted temperature (+ 4.5 °C) and pCO₂ levels (800 μatm). Over the experimental
31 period total phytoplankton biomass was significantly influenced by elevated pCO₂. At the end of
32 the experiment, biomass increased 6.5-fold under elevated pCO₂ and 4.6-fold under elevated
33 temperature relative to the ambient control. By contrast, the combined influence of elevated
34 pCO₂ and temperature had little effect on biomass relative to the control. Throughout the
35 experiment in all treatments and in the control, the phytoplankton community structure shifted
36 from dinoflagellates to nanophytoplankton . At the end of the experiment, under elevated pCO₂
37 nanophytoplankton contributed 90% of community biomass and was dominated by *Phaeocystis*
38 spp.. Under elevated temperature, nanophytoplankton comprised 85% of the community
39 biomass and was dominated by smaller nano-flagellates. In the control, larger nano-flagellates
40 dominated whilst the smallest nanophytoplankton contribution was observed under combined
41 elevated pCO₂ and temperature (~40 %). Under elevated pCO₂, temperature and in the control,
42 there was a significant decrease in dinoflagellate biomass. Under the combined effects of
43 elevated pCO₂ and temperature, dinoflagellate biomass increased and was dominated by the
44 harmful algal bloom (HAB) species, *Prorocentrum cordatum*. At the end of experiment,
45 Chlorophyll a (Chl *a*) normalised maximum photosynthetic rates (P^{B_m}) increased > 6-fold under
46 elevated pCO₂ and > 3-fold under elevated temperature while no effect on P^{B_m} was observed
47 when pCO₂ and temperature were elevated simultaneously. The results suggest that future
48 increases in temperature and pCO₂ simultaneously do not appear to influence coastal
49 phytoplankton productivity but significantly influence community composition during autumn
50 in the WEC.

51 **1. Introduction**

52 Oceanic concentration of CO₂ has increased by ~42% over pre-industrial levels, with a
53 continuing annual increase of ~0.4%. Current CO₂ level has reached ~400 μatm and has been
54 predicted to rise to >700 μatm by the end of this century(IPCC, 2013), with estimates exceeding
55 1000 μatm (Matear and Lenton, 2018; Raupach et al., 2007; Raven et al., 2005). With increasing
56 atmospheric CO₂, the oceans continue to absorb CO₂ from the atmosphere, which results in a
57 shift in oceanic carbonate chemistry resulting in a decrease in seawater pH or 'Ocean
58 Acidification' (OA). The projected increase in atmospheric CO₂ and corresponding increase in
59 ocean uptake, is predicted to result in a decrease in global mean surface seawater pH of 0.3
60 units below the present value of 8.1 to 7.8 (Wolf-gladrow et al., 1999). Under this scenario, the
61 shift in dissolved inorganic carbon (DIC) equilibria has wide ranging implications for
62 phytoplankton photosynthetic carbon fixation rates and growth (Riebesell, 2004).

63 Concurrent with OA, elevated atmospheric CO₂ and other climate active gases have warmed the
64 planet by ~0.6 °C over the past 100 years (IPCC, 2007). Atmospheric temperature has been

65 predicted to rise by a further 1.8 to 4 °C by the end of this century (Alley et al., 2007).
66 Phytoplankton metabolic activity may be accelerated by increased temperature (Eppley, 1972),
67 which can vary depending on the phytoplankton species and their physiological
68 requirements (Beardall et al., 2009; Boyd et al., 2013). Long-term data sets already suggest that
69 ongoing changes in coastal phytoplankton communities are likely due to climate shifts and other
70 anthropogenic influences (Edwards et al., 2006; Smetacek and Cloern, 2008; Widdicombe et al.,
71 2010). The response to OA and temperature can potentially alter the community composition,
72 community biomass and photo-physiology. Understanding how these two factors may interact,
73 synergistically or antagonistically, is critical to our understanding and for predicting future
74 primary productivity (Boyd and Doney, 2002; Dunne, 2014).

75 Laboratory studies of phytoplankton species in culture and studies on natural populations in
76 the field have shown that most species exhibit sensitivity, in terms of growth and
77 photosynthetic rates, to elevated pCO₂ and temperature individually. To date, only a few studies
78 have investigated the interactive effects of these two parameters on natural populations (e.g.
79 Coello-Camba et al., 2014; Feng et al., 2009; Gao et al., 2017; Hare et al., 2007). Most laboratory
80 studies demonstrate variable results with species-specific responses. In the diatom
81 *Thalassiosira weissflogii* for example, pCO₂ elevated to 1000 µatm and + 5 °C temperature
82 synergistically enhanced growth, while the same conditions resulted in a reduction in growth
83 for the diatom *Dactyliosolen fragilissimus* (Taucher et al., 2015). Although there have been fewer
84 studies on dinoflagellates, variable responses have also been reported (Errera et al., 2014; Fu et
85 al., 2008). In natural populations, elevated pCO₂ has stimulated the growth of pico- and
86 nanophytoplankton (Boras et al., 2016; Engel et al., 2008) while increased temperature has
87 reduced their biomass (Moustaka-Gouni et al., 2016; Peter and Sommer, 2012). In a recent field
88 study on natural phytoplankton communities, elevated temperature (+ 3°C above ambient)
89 enhanced community biomass but the combined influence of elevated temperature and pCO₂
90 reduced the biomass (Gao et al., 2017).

91 Phytoplankton species composition, abundance and biomass has been measured since 1992 at
92 the time-series station L4 in the western English Channel (WEC), to evaluate how global
93 changes could drive future shifts in phytoplankton community structure and carbon
94 biogeochemistry. At this station, sea surface temperature and pCO₂ reach maximum values
95 during late summer and start to decline in autumn. During October, mean seawater
96 temperatures at 10 m decrease from 15.39 °C (± 0.49 sd) to 14.37 °C (± 0.62 sd). Following a
97 period of CO₂ oversaturation in late summer, pCO₂ returns to near-equilibrium at station L4 in
98 October when mean pCO₂ values decrease from 455.32 µatm (± 63.92 sd) to 404.06 µatm (±
99 38.55 sd) (Kitidis et al., 2012).

100 From a biological perspective, the autumn period at station L4 is characterised by the decline of
101 the late summer diatom and dinoflagellate blooms (Widdicombe et al., 2010) when their
102 biomass approaches values close to the time series minima (diatom biomass range: 6.01 (\pm 6.88
103 sd) – 2.85 (\pm 3.28 sd) mg C m⁻³; dinoflagellate biomass range: 1.75 (\pm 3.28 sd) – 0.66 (\pm 1.08 sd)
104 mg C m⁻³). Typically, over this period nanophytoplankton becomes numerically dominant and
105 biomass ranges from 20.94 (\pm 33.25 sd) – 9.38 (\pm 3.31 sd) mg C m⁻³, though there is
106 considerable variability in this biomass.

107 Based on the existing literature, the working hypotheses of this study are that: (1) community
108 biomass will increase differentially under individual treatments of elevated temperature and
109 pCO₂; (2) elevated pCO₂ will lead to taxonomic shifts due to differences in species-specific CO₂
110 concentrating mechanisms and/or RuBisCO specificity; (3) photosynthetic carbon fixation rates
111 will increase differentially under individual treatments of elevated temperature and pCO₂; (4)
112 elevated temperature will lead to taxonomic shifts due to species-specific thermal optima; (5)
113 temperature and pCO₂ elevated simultaneously will have synergistic effects.

114 The objective of the study was therefore to investigate the combined effects of elevated pCO₂
115 and temperature on phytoplankton community structure, biomass and photosynthetic carbon
116 fixation rates during the autumn transition from diatoms and dinoflagellates to
117 nanophytoplankton at station L4 in the WEC.

118 **2. Materials and methods**

119 **2.1 Perturbation experiment, sampling and experimental set-up**

120 Experimental seawater containing a natural phytoplankton community was sampled at station
121 L4 (50 ° 15' N, 4 ° 13' W) on 7th October 2015 from 10 m depth (40 L). The experimental
122 seawater was gently pre-filtered through a 200 µm Nitex mesh to remove mesozooplankton
123 grazers, into two 20 L acid-cleaned carboys. While grazers play an important role in regulating
124 phytoplankton community structure (e.g. Strom, 2002), our experimental goals considered only
125 the effects of elevated temperature and pCO₂, though the mesh size used does not remove
126 microzooplankton. In addition, 320 L of seawater was collected into sixteen 20 L acid-cleaned
127 carboys from the same depth for use as experimental media. Immediately upon return to the
128 laboratory the media seawater was filtered through an in-line 0.2 and 0.1 µm filter (Acropak™,
129 Pall Life Sciences) then stored in the dark at 14 °C until use. The experimental seawater was
130 gently and thoroughly mixed and transferred in equal parts from each carboy (to ensure
131 homogeneity) to sixteen 2.5 L borosilicate incubation bottles (4 sets of 4 replicates). The
132 remaining experimental seawater was sampled for initial (T₀) concentrations of nutrients, Chl
133 *a*, total alkalinity, dissolved inorganic carbon, particulate organic carbon (POC) and nitrogen

134 (PON) and was also used to characterise the starting experimental phytoplankton community.
135 The incubation bottles were placed in an outdoor simulated in-situ incubation culture system
136 and each set of replicates was linked to one of four 22 L reservoirs filled with the filtered
137 seawater media. Neutral density spectrally corrected blue filters (Lee Filter no. 061) were
138 placed between polycarbonate sheets and mounted to the top, sides and ends of the incubation
139 system to provide ~50 % irradiance, approximating PAR measured at 10 m depth at station L4
140 on the day of sampling prior to starting experimental incubations (see **Fig. S1**, supplementary
141 material for time course of PAR levels during the experiment). The media was aerated with CO₂
142 free air and 5 % CO₂ in air precisely mixed using a mass flow controller (Bronkhorst UK
143 Limited) and used for the microcosm dilutions as per the following experimental design: (1)
144 control (390 µatm pCO₂, 14.5 °C matching station L4 in-situ values), (2) high temperature (390
145 µatm pCO₂, 18.5 °C), (3) high pCO₂ (800 µatm pCO₂, 14.5 °C) and (4) combination (800 µatm
146 pCO₂, 18.5 °C).

147 Initial nutrient concentrations (0.24 µM nitrate + nitrite, 0.086 µM phosphate and 2.14 µM
148 silicate on 7th October 2015) were amended to 8 µM nitrate+nitrite and 0.5 µM phosphate.
149 Pulses of nutrient inputs frequently occur at station L4 from August to December following
150 heavy rainfall events and subsequent riverine inputs to the system (e.g. Barnes et al., 2015). Our
151 nutrient amendments simulated these in situ conditions and were held constant to maintain
152 phytoplankton growth. Previous pilot studies highlighted that if these concentrations were not
153 maintained, the phytoplankton population crashes (Keys, 2017). As the phytoplankton
154 community was sampled over the transitional phase from diatoms and dinoflagellates to
155 nanophytoplankton, the in situ silicate concentration was maintained to reproduce the silicate
156 concentrations typical of this time of year (Smyth et al., 2010). Nutrient concentrations were
157 measured at time point T0 only.

158 Media transfer and sample acquisition was driven by peristaltic pumps. Following 48 hrs
159 acclimation in batch culture, semi-continuous daily dilution rates were maintained at between
160 10-13 % of the incubation bottle volume throughout the experiment. CO₂ enriched seawater
161 was added to the high CO₂ treatment replicates every 24 hrs, acclimating the natural
162 phytoplankton population to increments of elevated pCO₂ from ambient to ~800 µatm over 8
163 days followed by maintenance at ~800 µatm as per the method described by Schulz *et al*,
164 (2009). Adding CO₂ enriched seawater is the preferred protocol, since some phytoplankton
165 species are inhibited by the mechanical effects of direct bubbling (Riebesell et al., 2010; Shi et
166 al., 2009) which causes a reduction in growth rates and the formation of aggregates (Love et al.,
167 2016). pH was monitored daily to adjust the pCO₂ of the experimental media (+/-) prior to
168 dilutions to maintain target pCO₂ levels in the incubation bottles. The seasonality in pH and total

169 alkalinity (TA) are fairly stable at station L4 with high pH and low dissolved inorganic carbon
170 (DIC) during early summer, and low pH, high DIC throughout autumn and winter (Kitidis et al.,
171 2012). By maintaining the carbonate chemistry over the duration of the experiment, we aimed
172 to simulate natural events at the study site.

173 To provide sufficient time for changes in the phytoplankton community to occur and to achieve
174 an ecologically relevant data set, the incubation period was extended well beyond short-term
175 acclimation. Previous pilot studies using the same experimental protocols highlighted that after
176 ~20 days of incubation, significant changes in community structure and biomass were observed
177 (Keys, 2017). These results were used to inform a more relevant incubation period of 30+ days.

178 **2.2 Analytical methods, experimental seawater**

179 **2.2.1 Chlorophyll *a***

180 Chl *a* was measured in each incubation bottle. 100 mL triplicate samples from each replicate
181 were filtered onto 25 mm GF/F filters (nominal pore size 0.7 μm), extracted in 90 % acetone
182 overnight at -20 °C and Chl *a* concentration was measured on a Turner Trilogy™ fluorometer
183 using the non-acidified method of Welschmeyer (1994). The fluorometer was calibrated against
184 a stock Chl *a* standard (*Anacystis nidulans*, Sigma Aldrich, UK), the concentration of which was
185 determined with a Perkin Elmer™ spectrophotometer at wavelengths 663.89 and 750.11 nm.
186 Samples for Chl *a* analysis were taken every 2-3 days.

187 **2.2.2 Carbonate system**

188 70 mL samples for total alkalinity (TA) and dissolved inorganic carbon (DIC) analysis were
189 collected from each experimental replicate, stored in amber borosilicate bottles with no head
190 space and fixed with 40 μL of super-saturated Hg_2Cl_2 solution for later determination (Apollo
191 SciTech™ Alkalinity Titrator AS-ALK2; Apollo SciTech™ AS-C3 DIC analyser, with analytical
192 precision of 3 $\mu\text{mol kg}^{-1}$). Duplicate measurements were made for TA and triplicate
193 measurements for DIC. Carbonate system parameter values for media and treatment samples
194 were calculated from TA and DIC measurements using the programme CO₂sys (Pierrot et al.,
195 2006) with dissociation constants of carbonic acid of Mehrbach *et al.*, (1973) refitted by Dickson
196 and Millero (Dickson and Millero, 1987). Samples for TA and DIC were taken for analysis every
197 2-3 days throughout the experiment.

198 **2.2.3 Phytoplankton community analysis**

199 Phytoplankton community analysis was performed by flow cytometry (Becton Dickinson Accuri
200™ C6) for the 0.2 to 18 μm size fraction following Tarran *et al.*, (2006) and inverted light
201 microscopy was used to enumerate cells > 18 μm (BS EN 15204,2006). For flow cytometry, 2

202 mL samples fixed with glutaraldehyde to a final concentration of 2 % were flash frozen in liquid
203 nitrogen and stored at -80 °C for subsequent analysis. Phytoplankton data acquisition was
204 triggered on both chlorophyll fluorescence and forward light scatter (FSC) using prior
205 knowledge of the position of *Synechococcus* sp. to set the lower limit of analysis. Density plots of
206 FSC vs. CHL fluorescence, phycoerythrin fluorescence vs. CHL fluorescence and side scatter
207 (SSC) vs. CHL fluorescence were used to discriminate *Synechococcus* sp., picoeukaryote
208 phytoplankton (approx. 0.5–3 µm), coccolithophores, cryptophytes, *Phaeocystis* sp. single cells
209 and nanophytoplankton (eukaryotes >3 µm, excluding the coccolithophores, cryptophytes and
210 *Phaeocystis* sp. single cells), (for further information on flow cytometer calibration for
211 phytoplankton size measurements, see supplementary material). For inverted light microscopy,
212 140 mL samples were fixed with 2 % (final concentration) acid Lugol's iodine solution and
213 analysed by inverted light microscopy (Olympus™ IMT-2) using the Utermöhl counting
214 technique (Utermöhl, 1958; Widdicombe *et al.*, 2010). Phytoplankton community samples were
215 taken at T0, T10, T17, T24 and T36.

216 **2.2.4 Phytoplankton community biomass**

217 The smaller size fraction identified and enumerated through flow cytometry;
218 picophytoplankton, nanophytoplankton, *Synechococcus*, coccolithophores and cryptophytes
219 were converted to carbon biomass (mg C m⁻³) using a spherical model to calculate mean cell
220 volume:

$$221 \left(\frac{4}{3} * \pi * r^3\right) \quad \text{Equation 1.}$$

222 and a conversion factor of 0.22 pg C µm⁻³ (Booth, 1988). A conversion factor of 0.285 pg C µm⁻³
223 was used for coccolithophores (Tarran *et al.*, 2006) and cell a volume of 113 µm³ and carbon
224 cell⁻¹ value of 18 pg applied for *Phaeocystis* spp. (Widdicombe *et al.*, 2010). *Phaeocystis* spp.
225 were identified and enumerated by flow cytometry separately to the nanophytoplankton class
226 due to high observed abundance in in the high pCO₂ treatment. Mean cell measurements of
227 individual species/taxa were used to calculate cell bio-volume for the 18 µm + size fraction
228 according to Kovala and Larrance (1966) and converted to biomass according to the equations
229 of Menden-Deuer & Lessard, (2000).

230 **2.2.5 POC and PON**

231 Samples for particulate organic carbon (POC) and particulate organic nitrogen (PON) were
232 taken at T0, T15 and T36. 150 mL samples were taken from each replicate and filtered under
233 gentle vacuum pressure onto pre-ashed 25mm glass fibre filters (GF/F, nominal pore size 0.7
234 µm). Filters were stored in acid washed petri-slides at -20 °C until further processing. Sample

235 analysis was conducted using a Thermoquest Elemental Analyser (Flash 1112). Acetanilide
236 standards (Sigma Aldrich, UK) were used to calibrate measurements of carbon and nitrogen and
237 also used during the analysis to account for possible drift in measured concentrations.

238 **2.2.6 Chl fluorescence-based photophysiology**

239 Photosystem II (PSII) variable chlorophyll fluorescence parameters were measured using a fast
240 repetition rate fluorometer (FRRf) (FastOcean sensor in combination with an Act2Run
241 laboratory system, Chelsea Technologies, West Molesey, UK). The excitation wavelengths of the
242 FRRf's light emitting diodes (LEDs) were 450, 530 and 624 nm. The instrument was used in
243 single turnover mode with a saturation phase comprising 100 flashlets on a 2 μ s pitch and a
244 relaxation phase comprising 40 flashlets on a 50 μ s pitch. Measurements were conducted in a
245 temperature-controlled chamber at 15 °C. The minimum (F_o) and maximum (F_m) Chl
246 fluorescence were estimated according to Kolber et al., (1998). Maximum quantum yields of PSII
247 were calculated as:

$$248 \quad F_v / F_m = (F_m - F_o) / F_m \quad \text{Equation 2.}$$

249 PSII electron flux was calculated on a volume basis (JV_{PSII} ; $\text{mol e}^- \text{m}^{-3} \text{d}^{-1}$) using the absorption
250 algorithm (Oxborough et al., 2012) following spectral correction by normalising the FRRf LED
251 emission to the white spectra using Fast^{PRO} 8 software. This step required inputting the
252 experimental phytoplankton community fluorescence excitation spectra values (FES). Since we
253 did not measure the FES of our experimental samples, we used mean literature values for each
254 phytoplankton group calculated proportionally (based on percentage contribution to total
255 estimated biomass per phytoplankton group) as representative values for our experimental
256 samples. The JV_{PSII} rates were converted to chlorophyll specific carbon fixation rates (mg C (mg
257 $\text{Chl } a)^{-1} \text{m}^{-3} \text{h}^{-1}$), calculated as:

$$258 \quad JV_{PSII} \times \varphi_{E:C} \times MW_C / \text{Chl } a \quad \text{Equation 3}$$

259 where $\varphi_{E:C}$ is the electron requirement for carbon uptake ($\text{molecule CO}_2 \text{ (mol electrons)}^{-1}$), MW_C
260 is the molecular weight of carbon and $\text{Chl } a$ is the $\text{Chl } a$ measurement specific to each sample.
261 $\text{Chl } a$ specific JV_{PSII} based photosynthesis-irradiance curves were conducted in replicate batches
262 between 10:00 – 16:00 to account for variability over the photo-period at between 8 - 14
263 irradiance intensities. The maximum intensity applied was adjusted according to ambient
264 natural irradiance on the day of sampling. Maximum photosynthetic rates of carbon fixation
265 (P^B_m), the light limited slope (α^B) and the light saturation point of photosynthesis (I_k) were
266 estimated by fitting the data to the model of Webb et al., (1974):

$$267 \quad P^B = (1 - e^{-\alpha \times I / P^B_m}) \quad \text{Equation 4}$$

268 Due to instrument failure during the experiment, samples for FRRf fluorescence-based light
269 curves were taken at T36 only.

270 **2.3 Statistical analysis**

271 To test for effects of temperature, pCO₂ and possible time dependence of the measured response
272 variables (Chl *a*, total biomass, POC, PON, photosynthetic parameters and biomass of individual
273 species), generalized linear mixed models with the factors pCO₂, temperature and time (and all
274 interactions) were applied to the data between T0 and T36. Analyses were conducted using the
275 lme4 package in R (R Core Team (2014). R Foundation for Statistical Computing, Vienna,
276 Austria).

277 **3. Results**

278 Chl *a* concentration in the WEC at station L4 from 30 September - 6th October 2015 (when sea
279 water was collected for the experiment) varied between 0.02-5 mg m⁻³, with a mean
280 concentration of ~1.6 mg m⁻³ (**Fig. 1 A**). Over the period leading up to phytoplankton
281 community sampling, increasing nitrate and silicate concentrations coincided with a Chl *a* peak
282 on 23rd September (**Fig. 1 B**). Routine net trawl (20 µm) sample observations indicated a
283 phytoplankton community dominated by the diatoms *Leptocylindrus danicus* and *L. minimus*
284 with a lower presence of the dinoflagellates *Prorocentrum cordatum*, *Heterocapsa* spp. and
285 *Oxytoxum gracile*. Following decreasing nitrate concentrations, there was a *P. cordatum* bloom
286 on 29th September, during the week before the experiment started (data not shown).

287 **3.1 Experimental carbonate system**

288 Equilibration to the target high pCO₂ values (800 µatm) within the high pCO₂ and combination
289 treatments was achieved at T10 (**Fig. 2 A & B**). These treatments were slowly acclimated to
290 increasing levels of pCO₂ over 7 days (from the initial dilution at T3) while the control and high
291 temperature treatments were acclimated at the same ambient carbonate system values as those
292 measured at station L4 on the day of sampling. Following equilibration, the mean pCO₂ values
293 within the control and high temperature treatments were 394.9 (± 4.3 sd) and 393.2 (± 4.8 sd)
294 µatm respectively, while in the high pCO₂ and combination treatments mean pCO₂ values were
295 822.6 (± 9.4) and 836.5 (± 15.6 sd) µatm, respectively. Carbonate system values remained stable
296 throughout the experiment (For full carbonate system measured and calculated parameters, see
297 **Table S1** in supplementary material).

298 **3.2 Experimental temperature treatments**

299 Mean temperatures in the control and high pCO₂ treatments were 14.1 (± 0.35 sd) °C and in the
300 high temperature and combination treatments the mean temperatures were 18.6 (± 0.42 sd) °C,

301 with a mean temperature difference between the ambient and high temperature treatments of
302 4.46 (\pm 0.42 sd) °C (Supplementary material, **Fig. S2 A & B**).

303

304 **3.3 Chlorophyll *a***

305 Mean Chl *a* in the experimental seawater at T0 was 1.64 (\pm 0.02 sd) mg m⁻³ (**Fig. 3 A**). This
306 decreased in all treatments between T0 to T7, to ~0.1 (\pm 0.09, 0.035 and 0.035 sd) mg m⁻³ in the
307 control, high pCO₂ and combination treatments, while in the high temperature treatment at T7
308 Chl *a* was 0.46 mg m⁻³ (\pm 0.29 sd) ($z = 2.176$, $p < 0.05$). From T7 to T12 Chl *a* increased in all
309 treatments which was highest in the combination (4.99 mg m⁻³ \pm 0.69 sd) and high pCO₂
310 treatments (3.83 mg m⁻³ \pm 0.43 sd). Overall, Chl *a* was significantly influenced by experimental
311 time, independent of experimental treatments (**Table 1**). At T36 Chl *a* concentration in the
312 combination treatment was higher (6.87 (\pm 0.58 sd) mg m⁻³) than all other treatments while the
313 high temperature treatment concentration was higher (4.77 (\pm 0.44 sd) mg m⁻³) than the control
314 and high pCO₂ treatment. Mean concentrations for the control and high pCO₂ treatment at T36
315 were not significantly different at 3.30 (\pm 0.22 sd) and 3.46 (\pm 0.35 sd) mg m⁻³ respectively
316 (pairwise comparison $t = 0.78$, $p = 0.858$).

317 **3.4 Phytoplankton biomass**

318 The starting biomass in all treatments was 110.2 (\pm 5.7 sd) mg C m⁻³ (**Fig. 3 B**). The biomass was
319 dominated by dinoflagellates (~50%) with smaller contributions from nanophytoplankton
320 (~13%), cryptophytes (~11%), diatoms (~9%), coccolithophores (~8%), *Synechococcus* (~6%)
321 and picophytoplankton (~3%). Total biomass was significantly influenced in all treatments over
322 time (**Table 1**) and at T10, it was significantly higher in the high temperature treatment when
323 biomass reached 752 (\pm 106 sd) mg C m⁻³ ($z = 2.769$, $p < 0.01$). Biomass was significantly higher
324 in the elevated pCO₂ treatment (interaction of time x high pCO₂) (**Table 1**), reaching 2481 (\pm
325 182.68 sd) mg C m⁻³ at T36, ~6.5-fold higher than the control ($z = 3.657$, $p < 0.001$). Total
326 biomass in the high temperature treatment at T36 was significantly higher than the
327 combination treatment and ambient control ($z = 2.744$, $p < 0.001$), which were 525 (\pm 28.02 sd)
328 mg C m⁻³ and 378 (\pm 33.95 sd) mg C m⁻³, respectively. Reaching 1735 (\pm 169.24 sd) mg C m⁻³,
329 biomass in the high temperature treatment was ~4.6-fold higher than the control.

330 POC followed the same trends in all treatments between T0 and T36 (**Fig. 3 C**) and was in close
331 range of the estimated biomass ($R^2 = 0.914$, **Fig. 3 D**). POC was significantly influenced by the
332 interaction of time x high pCO₂ and time x high temperature (**Table 1**). At T36 POC was
333 significantly higher in the high pCO₂ treatment (2086 \pm 155.19 sd mg m⁻³) followed by the high

334 temperature treatment (1594 ± 162.24 sd mg m^{-3}), ~5.4-fold and 4-fold higher than the control,
335 respectively, whereas a decline in POC was observed in the control and combination treatment.
336 PON followed the same trend as POC over the course of the experiment, though it was only
337 significantly influenced by the interaction between time x high pCO_2 (**Fig. 3 E, Table 1**). At T36
338 concentrations were $147 (\pm 12.99$ sd) and $133 (\pm 15.59$ sd) mg m^{-3} in the high pCO_2 and high
339 temperature treatments respectively, while PON was $57.75 (\pm 13.07$ sd) mg m^{-3} in the
340 combination treatment and $47.18 (\pm 9.32$ sd) mg m^{-3} in the control. POC:PON ratios were
341 significantly influenced by the interaction of time x high pCO_2 and time x high temperature
342 (**Table 1**). The largest increase, from 3.028×10^{-5} to $1.632 \times 10^{-4} \mu\text{M C}:\mu\text{M N}$ ($\pm 1.299 \times 10^{-5}$ sd)
343 was in the high pCO_2 treatment (4.5-fold higher than the control at T36), followed by an
344 increase to 1.232×10^{-4} ($\pm 1.404 \times 10^{-5}$ sd) $\mu\text{M C}:\mu\text{M N}$ in the high temperature treatment (3-
345 fold higher than the control at T36). POC:PON in the combination treatment also increased over
346 time and was 45% higher than the control at T36 ($4.200 \times 10^{-5} \pm 5.550 \times 10^{-6}$ sd) $\mu\text{M C}:\mu\text{M N}$
347 (**Fig. 3 F**). The largest increase of 33 %, from 10.72 to 14.26 (± 1.73 sd) $\text{molC}:\text{molN}$ was in the
348 high pCO_2 treatment (73% higher than the control), followed by an increase of 32 % to 9.83 (\pm
349 1.82 sd) $\text{molC}:\text{molN}$ in the combination treatment (19% higher than the control), and an
350 increase of 17 % to 12.09 (± 2.14 sd) $\text{molC}:\text{molN}$ in the high temperature treatment (46%
351 higher than the control). In contrast, the POC:PON ratio in the control declined by 20 % from T0
352 to T36, from 10.33 to 8.26 (± 0.50 sd) $\text{molC}:\text{molN}$ (**Fig. 3 F**).

353

354 3.5 Community composition

355 From T0 to T24 the community shifted away from dominance of dinoflagellates in all
356 treatments, followed by further regime shifts between T24 and T36 in the control and
357 combination treatments. At T36 diatoms dominated the phytoplankton community biomass in
358 the ambient control (**Fig. 4 A**), while the high temperature and high pCO_2 treatments exhibited
359 near mono-specific dominance of nanophytoplankton (**Figs. 4 B & C**). The most diverse
360 community was in the combination treatment where dinoflagellates and *Synechococcus* became
361 more prominent (**Fig. 4 D**).

362 Between T10 and T24 the community shifted to nanophytoplankton in all experimental
363 treatments. This dominance was maintained to T36 in the high temperature and high pCO_2
364 treatments whereas in the ambient control and combination treatment, the community shifted
365 away from nanophytoplankton (**Fig. 5 A**). Nanophytoplankton biomass was significantly higher
366 in the high pCO_2 treatment (**Table 2**) with biomass reaching $2216 (\pm 189.67$ sd) mg C m^{-3} at
367 T36. This biomass was also high (though not significantly throughout the experiment until T36)

368 in the high temperature treatment (T36: 1489 (\pm 170.32 sd) mg C m⁻³, $z = 1.695$, $p = 0.09$)
369 compared to the control and combination treatments. In the combination treatment
370 nanophytoplankton biomass was 238 (\pm 14.16 sd) mg C m⁻³ at T36 which was higher than the
371 control, though not significantly (162 \pm 20.02 sd mg C m⁻³). In addition to significant differences
372 in nanophytoplankton biomass amongst the experimental treatments, treatment-specific
373 differences in cell size were also observed. Larger nano-flagellates dominated the control (mean
374 cell diameter of 6.34 μ m), smaller nano-flagellates dominated the high temperature and
375 combination treatments (mean cell diameters of 3.61 μ m and 4.28 μ m) whereas *Phaeocystis* spp.
376 dominated the high pCO₂ treatment (mean cell diameter 5.04 μ m) and was not observed in any
377 other treatment (Supplementary material, **Fig. S3 A-D**).

378 At T0, diatom biomass was low and dominated by *Coscinodiscus wailessi* (48 %; 4.99 mg C m⁻³),
379 *Pleurosigma* (25 %; 2.56 mg C m⁻³) and *Thalassiosira subtilis* (19 %; 1.94 mg C m⁻³). Small
380 biomass contributions were made by *Navicula distans*, undetermined pennate diatoms and
381 *Cylindrotheca closterium*. Biomass in the diatom group remained low from T0 to T24 but
382 increased significantly through time in all treatments (**Table 2**), with the highest biomass in the
383 high pCO₂ treatment (235 \pm 21.41 sd mg C m⁻³, **Fig. 5 B**). The highest diatom contribution to
384 total community biomass at T36 was in the ambient control (52 % of biomass; 198 \pm 17.28 sd
385 mg C m⁻³). In both the high temperature and combination treatments diatom biomass was lower
386 at T36 (151 \pm 10.94 sd and 124 \pm 19.16 sd mg C m⁻³, respectively). In all treatments, diatom
387 biomass shifted from the larger *C. wailessi* to the smaller *C. closterium*, *N. distans*, *T. subtilis* and
388 *Tropidoneis* spp., the relative contributions of which were treatment-specific. Overall *N. distans*
389 dominated diatom biomass in all treatments at T36 (ambient control: 112 \pm 24.86 sd mg C m⁻³,
390 56 % of biomass; high temperature: 106 \pm 17.75 sd mg C m⁻³, 70 % of biomass; high pCO₂: 152 \pm
391 19.09 sd mg C m⁻³, 61 % of biomass; and combination: 111 \pm 20.97 sd mg C m⁻³, 89 % of
392 biomass; Supplementary material, **Fig. S4 A-D**).

393 The starting dinoflagellate community was dominated by *Gyrodinium spirale* (91 %; 49 mg C m⁻³)
394 with smaller contributions from *Katodinium glaucum* (5 %; 2.76 mg C m⁻³), *Prorocentrum*
395 *cordatum* (3 %; 1.78 mg C m⁻³) and undetermined *Gymnodiniales* (1 %; 0.49 mg C m⁻³). At T36
396 Dinoflagellate biomass was significantly higher in the combination treatment (90 \pm 16.98 sd mg
397 C m⁻³, **Fig. 5 C, Table 2**) followed by the high temperature treatment (57 \pm 6.87 sd mg C m⁻³,
398 **Table 2**). There was no significant difference in dinoflagellate biomass between the high pCO₂
399 treatment and ambient control at T36 when biomass was low. In the combination treatment, the
400 dinoflagellate biomass became dominated by *P. cordatum* which contributed 59 \pm 12.95 sd mg C
401 m⁻³ (66 % of biomass in this group).

402 *Synechococcus* biomass was significantly higher in the combination treatment (reaching $59.9 \pm$
403 4.30 sd mg C m⁻³ at T36, **Fig. 5 D, Table 2**) followed by the high temperature treatment ($30 \pm$
404 5.98 sd mg C m⁻³, **Table 2**). In both the high pCO₂ treatment and control *Synechococcus* biomass
405 was low (~ 7 mg C m⁻³ in both treatments at T36), though an initial significant response to high
406 pCO₂ was observed between T0 – T10 (**Table 2**). In all treatments and throughout the
407 experiment, relative to the other phytoplankton groups, biomass of picophytoplankton (**Fig. 5**
408 **E**), cryptophytes (**Fig. 5 F**) and coccolithophores (**Fig. 5 G**) remained low, though there was a
409 slight increase in picophytoplankton in the combination treatment (11.26 ± 0.79 sd mg C m⁻³;
410 **Table 2**).

411 Microzooplankton was dominated by *Strombolidium* spp. in all treatments throughout the
412 experiment, though biomass was low relative to the phytoplankton community (**Fig. 6**).
413 Following a decline from T0 to T10, microzooplankton biomass increased in all but the high CO₂
414 treatment until T17 when biomass diverged. The biomass trajectory maintained an increase in
415 the control when at T36 it was highest at ~ 1.6 mg C m⁻³, 90% higher than the high temperature
416 treatment (0.83 mg C m⁻³). Microzooplankton biomass was significantly lower in the high CO₂
417 treatment at T36 ($z = -2.100$, $p = 0.036$) and undetected in the combination treatment at this
418 time point (**Table 2**).

419

420 **3.6 Chl *a* fluorescence-based photophysiology**

421 At T36, FRRf photosynthesis-irradiance (PE) parameters were strongly influenced by the
422 experimental treatments. P^{B_m} was significantly higher in the high pCO₂ treatment (18.93 mg C
423 (mg Chl *a*)⁻¹ m⁻³ h⁻¹), followed by the high temperature treatment (9.58 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹;
424 **Fig. 7, Tables 3 & 4**). There was no significant difference in P^{B_m} between the control and
425 combination treatments (2.77 and 3.02 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹). Light limited photosynthetic
426 efficiency (α^B) also followed the same trend and was significantly higher in the high pCO₂
427 treatment (0.13 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹ ($\mu\text{mol photon m}^{-2} \text{s}^{-1}$)⁻¹) followed by the high
428 temperature treatment (0.09 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹ ($\mu\text{mol photon m}^{-2} \text{s}^{-1}$)⁻¹; **Tables 3 & 4**). α^B
429 was low in both the control and combination treatment (0.03 and 0.04 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹
430 ($\mu\text{mol photon m}^{-2} \text{s}^{-1}$)⁻¹, respectively). The light saturation point of photosynthesis (E_k) was
431 significantly higher in the high pCO₂ treatment relative to all treatments (144.13 $\mu\text{mol photon}$
432 $\text{m}^{-2} \text{s}^{-1}$), though significantly lower in the combination treatment relative to both the high pCO₂
433 and high temperature treatments (**Tables 3 & 4**).

434 **4. Discussion**

435 Individually, elevated temperature and pCO₂ resulted in the highest biomass and maximum
436 photosynthetic rates (P^{B_m}) at T36, when nanophytoplankton dominated. The interaction of
437 these two factors had little effect on total biomass with values close to the ambient control, and
438 no effect on P^{B_m}. The combination treatment, however, exhibited the greatest diversity of
439 phytoplankton functional groups with dinoflagellates and *Synechococcus* becoming dominant
440 over time.

441 Elevated pCO₂ has been shown to enhance the growth and photosynthesis of some
442 phytoplankton species which have active uptake systems for inorganic carbon (Giordano et al.,
443 2005; Reinfelder, 2011). Elevated pCO₂ may therefore lead to lowered energetic costs of carbon
444 assimilation in some species and a redistribution of the cellular energy budget to other
445 processes (Tortell et al., 2002). In this study, under elevated pCO₂ where the dominant group
446 was nanophytoplankton, the most abundant species was the haptophyte *Phaeocystis* spp.
447 Photosynthetic carbon fixation in *Phaeocystis* spp. is presently near saturation with respect to
448 current levels of pCO₂ (Rost et al., 2003). Dominance of this spp. under elevated pCO₂ may be
449 due to lowered grazing pressure since microzooplankton biomass was lowest in the high CO₂
450 treatment throughout the experiment. The increased biomass and photosynthetic carbon
451 fixation in this experimental community under elevated pCO₂ is due to the community shift to
452 *Phaeocystis* spp.. The increased biomass in the high temperature treatment (where
453 microzooplankton biomass remained stable between T17 to T36, though lower than the
454 control) may be attributed to enhanced enzymatic activities, since algal growth commonly
455 increases with temperature until after an optimal range (Boyd et al., 2013; Goldman and
456 Carpenter, 1974; Savage et al., 2004). Optimum growth temperatures for marine phytoplankton
457 are often several degrees higher than environmental temperatures (Eppley, 1972; Thomas et al.,
458 2012). Nanophytoplankton also dominated in this treatment and while *Phaeocystis* spp. was not
459 discriminated, no further classification was made at a group/species level. Reduced biomass in
460 the control from T24 onwards may be due to increased grazing pressure given the highest
461 concentrations of microzooplankton biomass were observed in the control. Conversely,
462 microzooplankton biomass declined significantly from T17 in the combination treatment,
463 indicating reduced grazing pressure while phytoplankton biomass also declined from this time
464 point. Nutrient concentrations were not measured beyond T0 and we cannot therefore exclude
465 the possibility that differences in nutrient availability may have contributed to observed
466 differences between control and high temperature and high CO₂ treatments.

467 **4.1 Chl *a***

468 Biomass in the control peaked at T25 followed by a decline to T36. Correlated with this, Chl *a*
469 also peaked at T25 in the control and declined to 3.3 mg m⁻³ by T27, remaining close to this

470 value until T36. Biomass in the combination treatment peaked at T20 followed by decline to
471 T36 whereas Chl *a* in this treatment declined from T20 to T25 followed by an increase at T27
472 before further decline similar to the biomass. Chl *a* peaked in this treatment again at T36 (6.8
473 mg m⁻³). We attribute the increase in Chl *a* between T25 – T27 (coincident with an overall
474 biomass decrease) to lower species specific carbon:Chl *a* ratios as a result of the increase in
475 dinoflagellates, *Synechococcus* and picophytoplankton biomass from T25. We speculate that the
476 decline in biomass under nutrient replete conditions in the combination treatment was
477 probably due to slower species-specific growth rates when diatoms and dinoflagellates became
478 more prominent in this treatment. Carbon:Chl *a* in diatoms and dinoflagellates have previously
479 been demonstrated to be lower than nano- and picophytoplankton (Sathyendranath et al.,
480 2009) This contrasts the results reported in comparable studies as Chl *a* is generally highly
481 correlated with biomass, (e.g. Feng et al., 2009). Similar results were reported however by Hare
482 et al., (2007) which indicates that Chl *a* may not always be a reliable proxy for biomass in mixed
483 communities.

484 **4.2 Biomass**

485 This study shows that the phytoplankton community response to elevated temperature and
486 pCO₂ is highly variable. pCO₂ elevated to ~800 µatm induced higher community biomass, similar
487 to the findings of Kim et al., (2006), whereas in other natural community studies no CO₂ effect
488 on biomass was observed (Delille et al., 2005; Maugendre et al., 2017; Paul et al., 2015). A ~4.5
489 °C increase in temperature also resulted in higher biomass at T36 in this study, similar to the
490 findings of Feng et al., (2009) and Hare et al., (2007) though elevated temperature has
491 previously reduced biomass of natural nanophytoplankton communities in the Western Baltic
492 Sea and Arctic Ocean (Coello-Camba et al., 2014; Moustaka-Gouni et al., 2016). When elevated
493 temperature and pCO₂ were combined, community biomass exhibited little response, similar to
494 the findings of Gao et al., (2017), though an increase in biomass has also been reported (Calbet
495 et al., 2014; Feng et al., 2009). Geographic location and season also play an important role in
496 structuring the community and its response in terms of biomass to elevated temperature and
497 pCO₂. (Li et al., 2009; Morán et al., 2010). This may explain part of the variability in responses
498 observed from studies on phytoplankton during different seasons and provinces.

499 **4.3 Carbon:Nitrogen**

500 In agreement with others, the results of this experiment showed highest increases in C:N under
501 elevated pCO₂ alone (Riebesell et al., 2007). C:N also increased under high temperature,
502 consistent with the findings of Lomas and Glibert, (1999) and Taucher et al., (2015). It also
503 increased when pCO₂ and temperature were elevated, albeit to a lesser degree , which was also

504 observed by Calbet et al., (2014), but contrasts other studies that have observed C:N being
505 unaffected by the combined influence of elevated pCO₂ and temperature (Deppeler and
506 Davidson, 2017; Kim et al., 2006; C. Paul et al., 2015). C:N is a strong indicator of cellular protein
507 content (Woods and Harrison, 2003) and increases under elevated pCO₂ and warming may lead
508 to lowered nutritional value of phytoplankton which has implications for zooplankton
509 reproduction and the biogeochemical cycling of nutrients.

510 **4.4 Photosynthetic carbon fixation rates**

511 At T36, under elevated pCO₂ P^B_m was > 6 times higher than in the control, but only one time
512 point was measured so we are not able to make decisive conclusions. Riebesell et al., (2007) and
513 Tortell et al., (2008) also reported an increase in P^B_m under elevated pCO₂. By contrast other
514 observations on natural populations under elevated pCO₂ reported a reduction in P^B_m (Feng et
515 al., 2009; Hare et al., 2007). Studies on laboratory cultures have shown that increases in
516 temperature cause an increase photosynthetic rates (Feng et al., 2008; Fu et al., 2007; Hutchins
517 et al., 2007), similar to what we observed in this study. In the combined pCO₂ and temperature
518 treatment, we found no effect on P^B_m, which has also been observed in experiments on natural
519 populations (Coello-Camba and Agustí, 2016; Gao et al., 2017). This contrasts the findings of
520 Feng et al., (2009) and Hare et al., (2007) who observed the highest P^B_m when temperature and
521 pCO₂ were elevated simultaneously. In this study, increases in α^B and E_k under elevated pCO₂,
522 and a decrease in these parameters when elevated pCO₂ and temperature were combined also
523 contrasts the trends reported by Feng et al., (2009). We should stress however, that while our
524 photophysiological measurements support our observed trends in community biomass, they
525 were made on a single occasion at the end of the experiment. Future experiments should focus
526 on acquiring photophysiological measurements throughout.

527 Species specific photosynthetic rates have been demonstrated to decrease beyond their thermal
528 optimum (Raven and Geider, 1988) which can be modified through photoprotective rather than
529 photosynthetic pigments (Kiefer and Mitchell, 1983). This may explain the difference in P^B_m
530 between the high pCO₂ and high temperature treatments (in addition to differences in
531 nanophytoplankton community composition in relation to *Phaeocystis* spp. discussed above), as
532 the experimental high temperature treatment in this study was ~4.5 ° C higher than the control.

533 There was no significant effect of combined elevated pCO₂ and temperature on P^B_m, which was
534 strongly influenced by taxonomic differences between the experimental treatments. Warming
535 has been shown to lead to smaller cell sizes in nanophytoplankton (Atkinson et al., 2003; Peter
536 and Sommer, 2012), which was observed in the combined treatment together with decreased
537 nanophytoplankton biomass. Diatoms also shifted to smaller species with reduced biomass,

538 while dinoflagellate and *Synechococcus* biomass increased at T36. Dinoflagellates are the only
539 photoautotrophs with form II RuBisCO (Morse et al., 1995) which has the lowest
540 carboxylation:oxygation specificity factor among eukaryotic phytoplankton (Badger et al.,
541 1998), which may give dinoflagellates a disadvantage in carbon fixation under present ambient
542 pCO₂ levels. Phytoplankton growth rates are generally slower in surface waters with high pH
543 (≥9) resulting from photosynthetic removal of CO₂ by previous blooms and the associated
544 nutrient depletion (Hansen, 2002; Hinga, 2002). Though growth under high pH provides
545 indirect evidence that dinoflagellates possess CCMs, direct evidence is limited and points to the
546 efficiency of CCMs in dinoflagellates as moderate in comparison to diatoms and some
547 haptophytes (Reinfelder, 2011 and references therein). Given that dinoflagellates accounted for
548 just ~20% of biomass in the combination treatment, exerting a minor influence on community
549 photosynthetic rates, further work is required to explain the lower P^B_m under the combined
550 influence of elevated pCO₂ and temperature compared to the individual treatment influences.
551 We applied the same electron requirement parameter for carbon uptake across all treatments,
552 though in nature and between species, there can be considerable variation in this parameter
553 (e.g. 1.15 to 54.2 mol e⁻ (mol C)⁻¹; Lawrenz et al., 2013) which can co-vary with temperature,
554 nutrients, Chl *a*, irradiance and community structure. Better measurement techniques at
555 quantifying this variability are necessary in the future.

556 **4.5 Community composition**

557 Phytoplankton community structure changes were observed, with a shift from dinoflagellates to
558 nanophytoplankton which was most pronounced under single treatments of elevated
559 temperature and pCO₂. Amongst the nanophytoplankton, a distinct size shift to smaller cells was
560 observed in the high temperature and combination treatments, while in the high pCO₂
561 treatment *Phaeocystis* spp. dominated. Under combined pCO₂ and temperature from T24
562 onwards however, dinoflagellate and *Synechococcus* biomass increased and nanophytoplankton
563 biomass decreased. An increase in pico- and nanophytoplankton has previously been reported
564 in natural communities under elevated pCO₂ (Bermúdez et al., 2016; Boras et al., 2016;
565 Brussaard et al., 2013; Engel et al., 2008) while no effect on these size classes has been observed
566 in other studies (Calbet et al., 2014; Paulino et al., 2007). Moustaka-Gouni et al., (2016) also
567 found no CO₂ effect on natural nanophytoplankton communities but increased temperature
568 reduced the biomass of this group. Kim et al., (2006) observed a shift from nanophytoplankton
569 to diatoms under elevated pCO₂ alone while a shift from diatoms to nanophytoplankton under
570 combined elevated pCO₂ and temperature has been reported (Hare et al., 2007). A variable
571 response in *Phaeocystis* spp. to elevated pCO₂ has also been reported with increased growth
572 (Chen et al., 2014; Keys et al., 2017), no effect (Thoisen et al., 2015) and decreased growth

573 (Hoogstraten et al., 2012) observed. *Phaeocystis* spp. can outcompete other phytoplankton and
574 form massive blooms (up to 10 g C m⁻³) with impacts on food webs, global biogeochemical
575 cycles and climate regulation (Schoemann et al., 2005). While not a toxic algal species,
576 *Phaeocystis* spp. are considered a harmful algal bloom (HAB) species when biomass reaches
577 sufficient concentrations to cause anoxia through the production of mucus foam which can clog
578 the feeding apparatus of zooplankton and fish (Eilertsen & Raa, 1995).

579 Recently published studies on the response of diatoms to elevated pCO₂ and temperature vary
580 greatly. For example, Taucher et al., (2015) showed that *Thalassiosira weissflogii* incubated at
581 1000 µatm pCO₂ increased growth by 8 % while for *Dactyliosolen fragilissimus*, growth
582 increased by 39 %; temperature elevated by + 5°C also had a stimulating effect on *T. weissflogii*
583 but inhibited the growth rate of *D. fragilissimus*; and when the treatments were combined
584 growth was enhanced in *T. weissflogii* but reduced in *D. fragilissimus*. In our study, elevated pCO₂
585 increased biomass in diatoms (time dependent), but elevated temperature and the combination
586 of these factors reduced the signal of this response. A distinct size-shift in diatom species was
587 observed in all treatments, from the larger *Coscinodiscus* spp., *Pleurosigma* and *Thalassiosira*
588 *subtilis* to the smaller *Navicula distans*. This was most pronounced in the combination treatment
589 where *N. distans* formed 89 % of diatom biomass. *Navicula* spp. previously exhibited a
590 differential response to both elevated temperature and pCO₂. At + 4.5 °C and 960 ppm CO₂
591 Torstensson et al., (2012) observed no synergistic effects on the benthic *Navicula directa*.
592 Elevated temperature increased growth rates by 43 % while a reduction of 5 % was observed
593 under elevated CO₂. No effects on growth were detected at pH ranging from 8 – 7.4 units in
594 *Navicula* spp. (Thoisen et al., 2015), while there was a significant increase in growth in *N.*
595 *distans* along a CO₂ gradient at a shallow cold-water vent system (Baragi et al., 2015).

596 *Synechococcus* grown under pCO₂ elevated to 750 ppm and temperature elevated by 4 °C
597 resulted in increased growth and a 4-fold increase in P_{Bm} (Fu et al., 2007) which is similar to the
598 results of the present study.

599 The combination of elevated temperature and pCO₂ significantly increased dinoflagellate
600 biomass to 17 % of total biomass. This was due to *P. cordatum* which increased biomass by
601 more than 30-fold from T0 to T30 (66 % of dinoflagellate biomass in this treatment). Despite
602 the global increase in the frequency of HABs few studies have focussed on the response of
603 dinoflagellates to elevated pCO₂ and temperature. In laboratory studies at 1000 ppm CO₂,
604 growth rates of the HAB species *Karenia brevis* increased by 46 %, at 1000 ppm CO₂ and + 5 °C
605 temperature it's growth increased by 30 % but was reduced under elevated temperature alone
606 (Errera et al., 2014). A combined increase in pCO₂ and temperature enhanced both the growth

607 and P^B_m in the dinoflagellate *Heterosigma akashiwo*, whereas in contrast to the present findings,
608 only pCO₂ alone enhanced these parameters in *P. cordatum* (Fu et al., 2008).

609 5. Implications

610 Increased biomass, P^B_m and a community shift to nanophytoplankton under individual increases
611 in temperature and pCO₂ suggests a potential negative feedback on atmospheric CO₂, whereby
612 more CO₂ is removed from the ocean, and hence from the atmosphere through an increase in
613 photosynthesis. The selection of *Phaeocystis* spp. under elevated pCO₂ indicates the potential for
614 negative impacts on ecosystem function and food web structure due to the formation of hypoxic
615 zones which can occur under eutrophication, inhibitory feeding effects and lowered fecundity in
616 many copepods associated with this species (Schoemann et al., 2005; Verity et al., 2007). While
617 more CO₂ is fixed, selection for nanophytoplankton in both of these treatments however, may
618 result in reduced carbon sequestration due to slower sinking rates of the smaller phytoplankton
619 cells (Bopp et al., 2001; Laws et al., 2000). When temperature and pCO₂ were elevated
620 simultaneously, community biomass showed little response and no effects on P^B_m were
621 observed. This suggests no change on feedback to atmospheric CO₂ and climate warming in
622 future warmer high CO₂ oceans. Additionally, combined elevated pCO₂ and temperature
623 significantly modified taxonomic composition, by reducing diatom biomass relative to the
624 control with an increase in dinoflagellate biomass dominated by the HAB species, *P. cordatum*.
625 This has implications for fisheries, ecosystem function and human health.

626 6. Conclusion

627 These experimental results provide new evidence that increases in pCO₂ coupled with rising sea
628 temperatures may have antagonistic effects on the autumn phytoplankton community in the
629 WEC. Under future global change scenarios, the size range and biomass of diatoms may be
630 reduced with increased dinoflagellate biomass and the selection of HAB species. The
631 experimental simulations of year 2100 temperature and pCO₂ demonstrate that the effects of
632 warming can be offset by elevated pCO₂, maintaining current levels of coastal phytoplankton
633 productivity while significantly altering the community structure, and in turn these shifts will
634 have consequences on carbon biogeochemical cycling in the WEC.

635 **Data availability:** Experimental data used for analysis will be made available (DOI will be
636 created)

637 **Author contributions:** Matthew Keys collected, measured, processed and analysed the data and
638 prepared the figures. Drs Gavin Tilstone and Helen Findlay conceived, directed and sought the

639 necessary funds to support the research. Matthew Keys and Dr Gavin Tilstone wrote the paper
640 with input from Claire Widdicombe and Professor Tracy Lawson. Claire Widdicombe supervised
641 and advised on phytoplankton taxonomic classifications.

642 **Competing interests:** The authors declare that they have no conflict of interest.

643 **Acknowledgements:** G.H.T, H.S.F. and C.E.W were supported by the UK Natural Environment
644 Research Council's (NERC) National Capability – The Western English Channel Observatory
645 (WCO). C.E.W was also partly funded by the NERC and Department for Environment, Food and
646 Rural Affairs, Marine Ecosystems Research Program (Grant no. NE/L003279/1). M.K. was
647 supported by a NERC PhD studentship (grant No. NE/L50189X/1). We thank Glen Tarran for his
648 training, help and assistance with flow cytometry, The National Earth Observation Data Archive
649 and Analysis Service UK (NEODAAS) for providing the MODIS image used in Fig 1. and the crew
650 of RV Plymouth Quest for their helpful assistance during field sampling.

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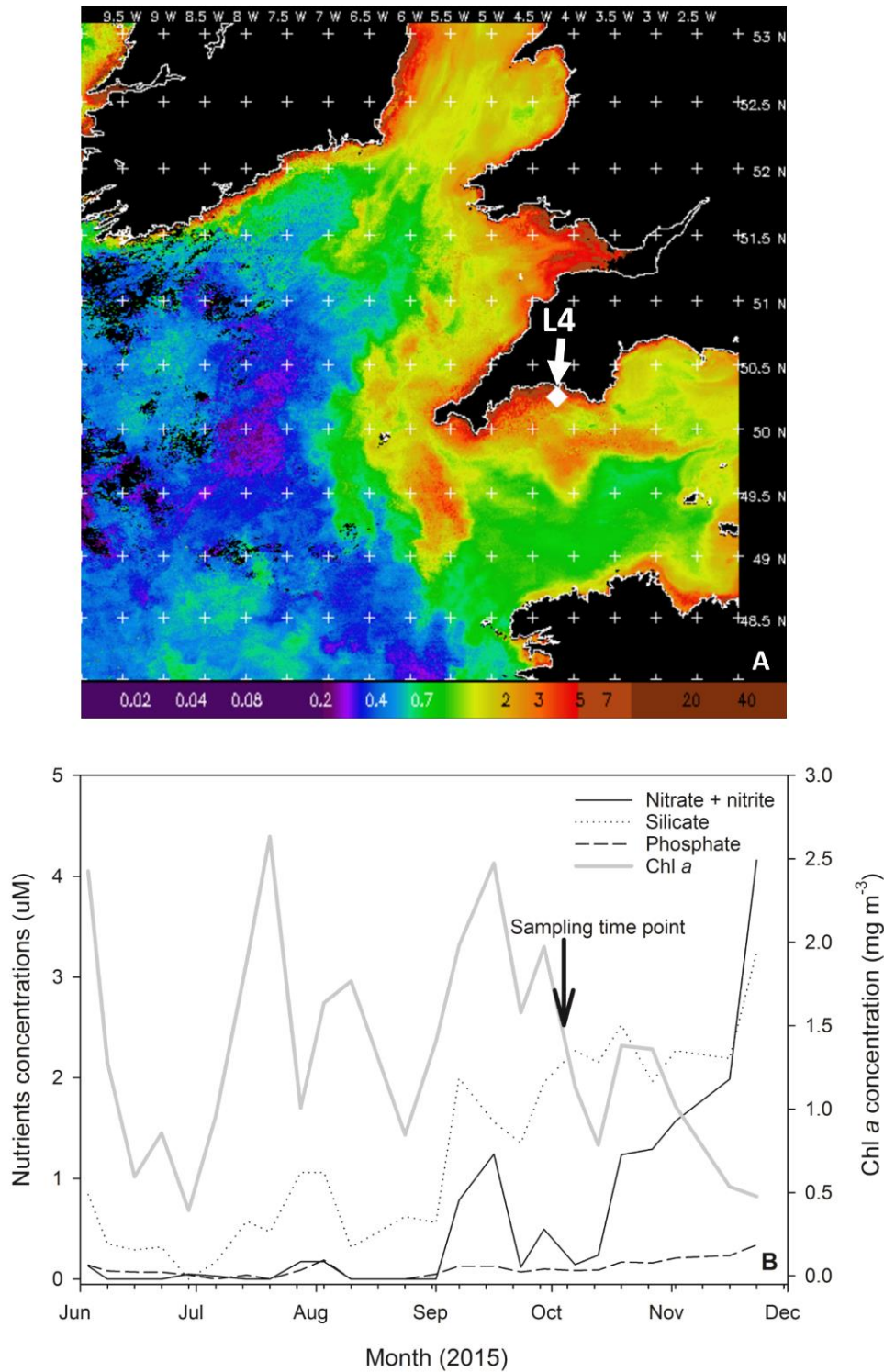


Fig. 1. (A). MODIS weekly composite chl *a* image of the western English Channel covering the period 30th September – 6th October 2015 (coincident with the week of phytoplankton community sampling for the present study), processing courtesy of NEODAAS. The position of coastal station L4 is marked with a white diamond. (B). Profiles of weekly nutrient and chl *a* concentrations from station L4 at a depth of 10 m over the second half of 2015 in the months prior to phytoplankton community sampling (indicated by black arrow and text).

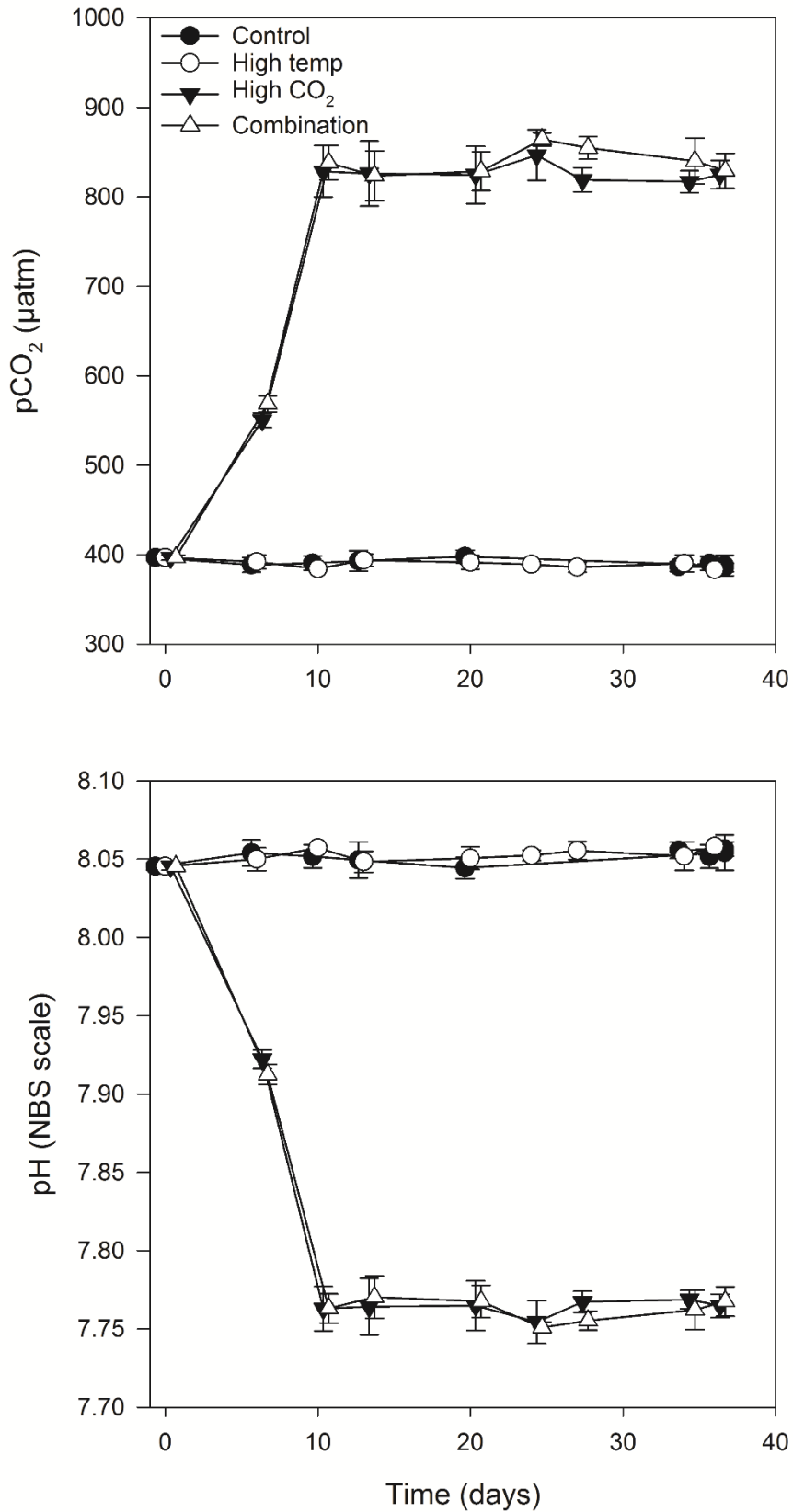


Fig. 2. Calculated values of partial pressure of CO₂ in seawater (pCO₂) (A) and pH (B) from direct measurements of total alkalinity and dissolved inorganic carbon. (For full carbonate system values see **Table S1.**, supplementary material)

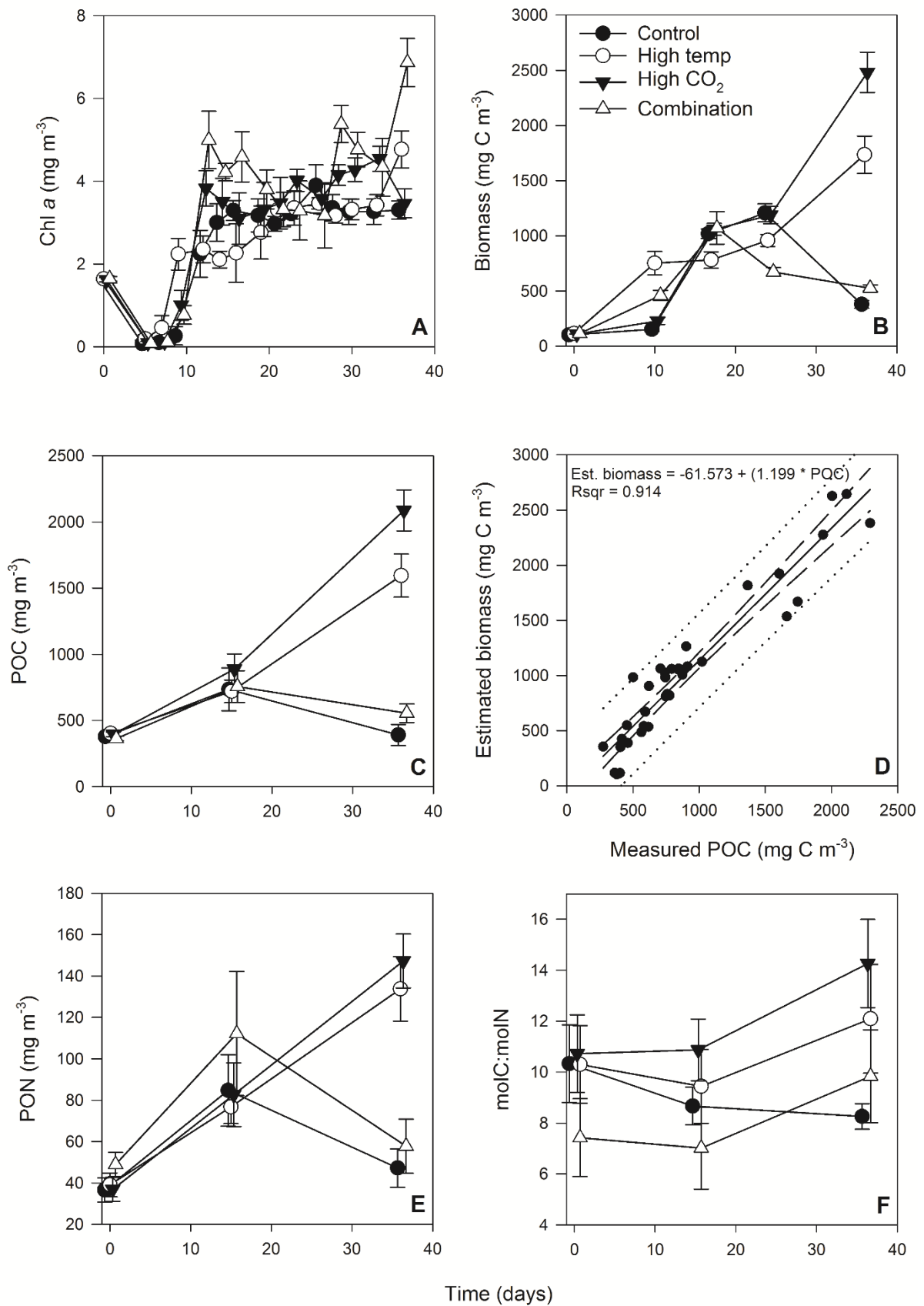


Fig. 3. Time course of chl *a* (A), estimated phytoplankton biomass (B), POC (C), regression of estimated phytoplankton carbon vs measured POC (D), PON (E) and POC:PON (F).

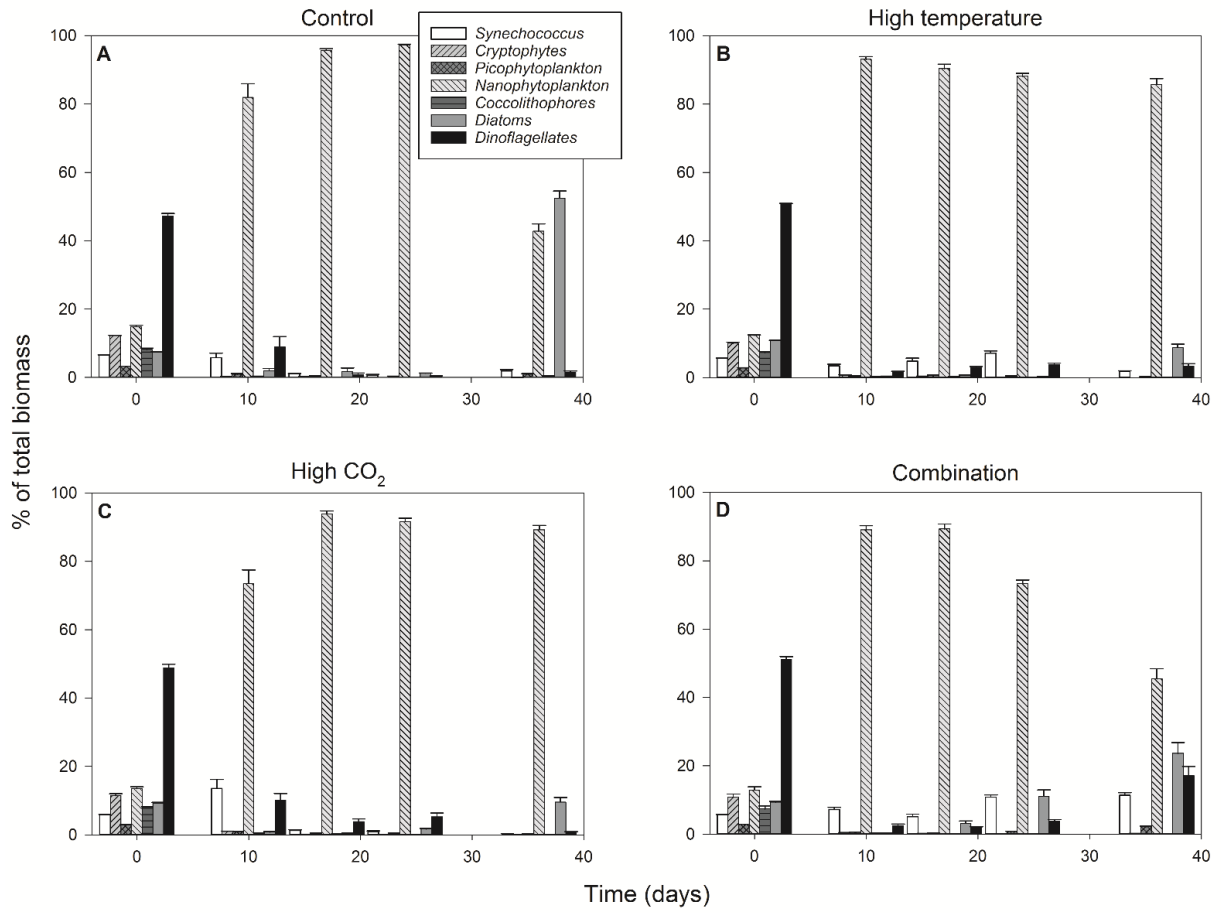


Fig. 4. Percentage contribution to community biomass by phytoplankton groups/species throughout the experiment in the control (A), high temperature (B), high CO₂ (C) and combination treatments (D).

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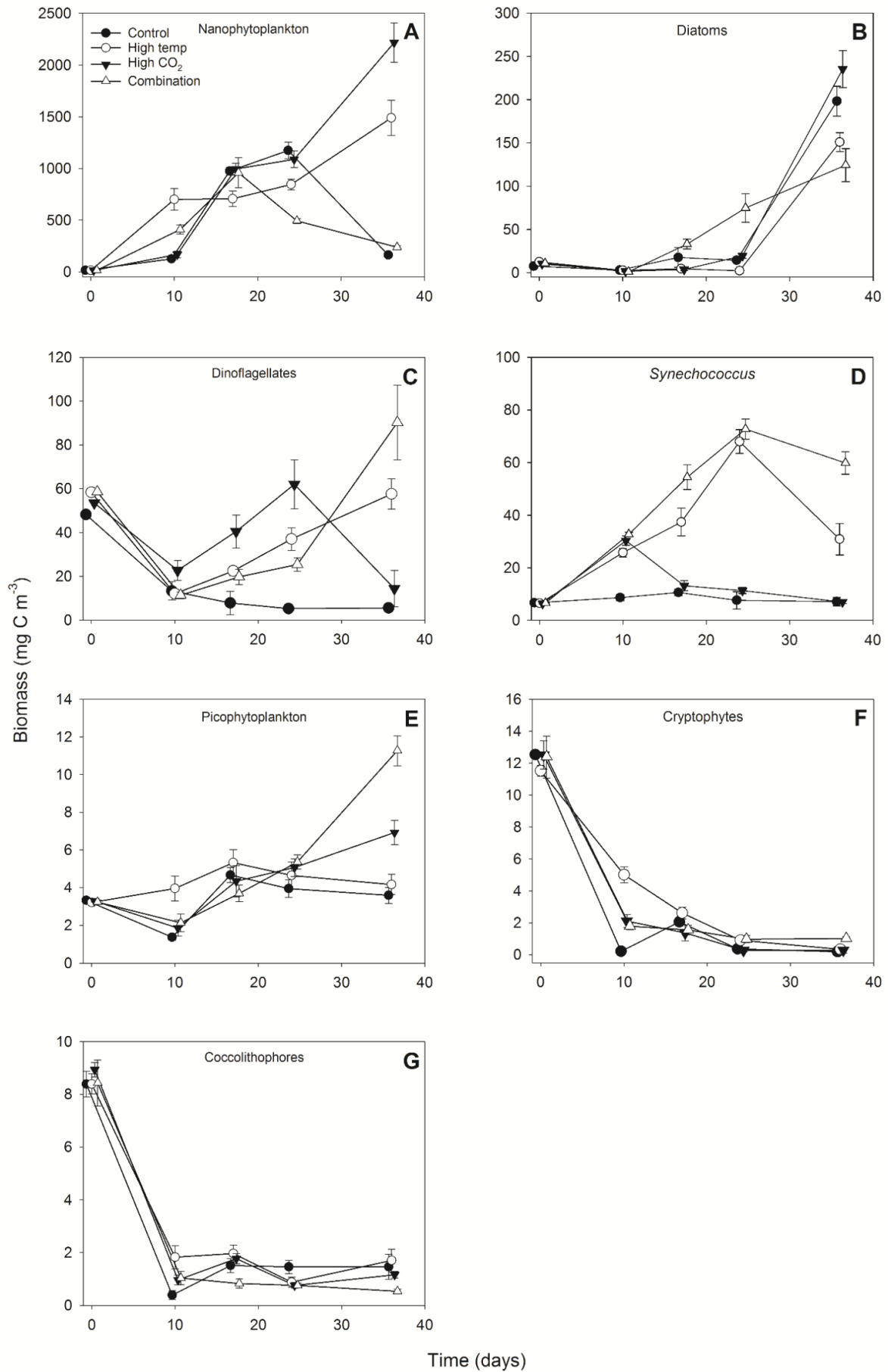


Fig. 5. Response of individual phytoplankton groups to experimental treatments.

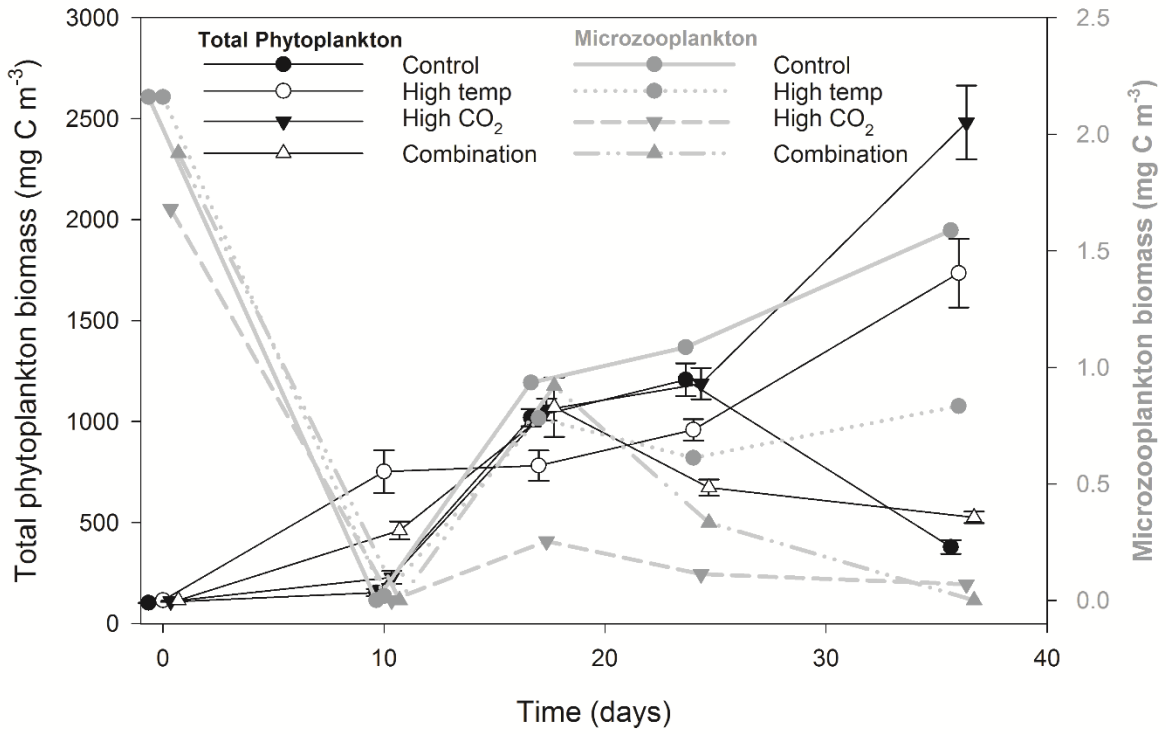


Fig. 6. Microzooplankton biomass (dominated by *Strombolidium* sp.) relative to total phytoplankton biomass.

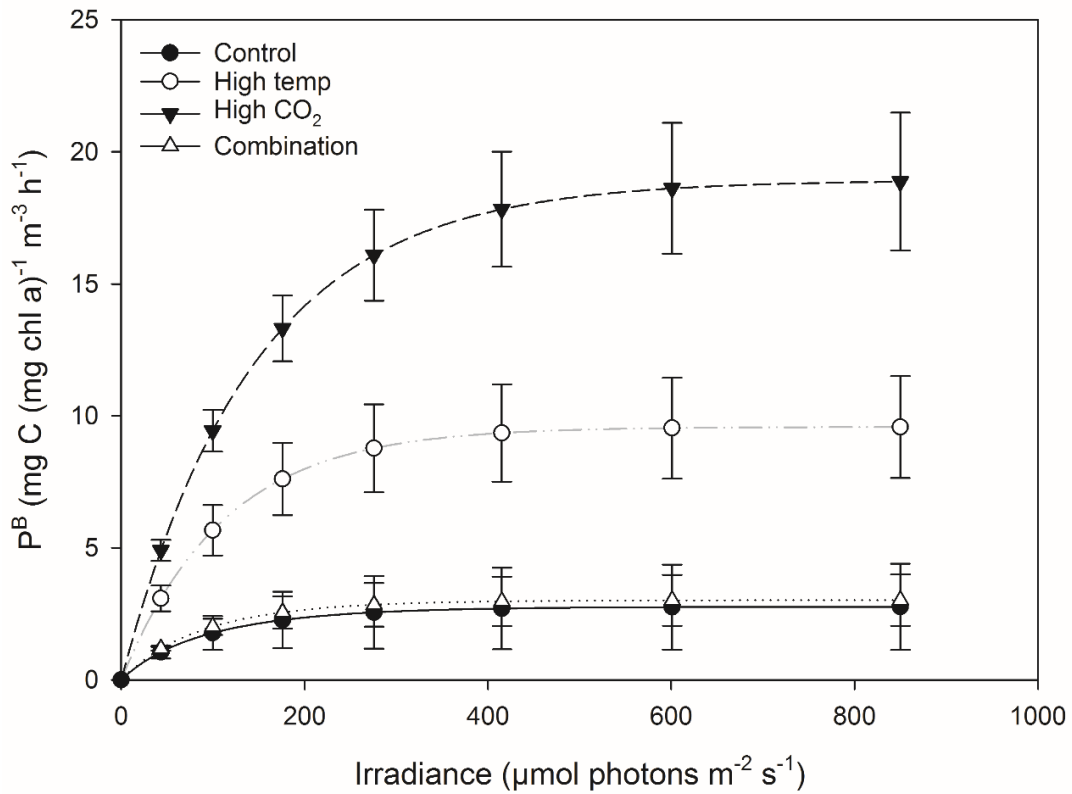


Fig. 7. Fitted parameters of FRRf-based photosynthesis-irradiance curves for the experimental treatments on the final experimental day (T36)

Table 1. Results of generalized linear mixed model testing for effects of time, temperature, pCO₂ and all interactions on chl *a*, phytoplankton biomass and particulate organic carbon and nitrogen. Significant results are in bold; * p < 0.05, ** p < 0.01, *** p < 0.001.

| Response variable | n | df | z-value | p | sig |
|---|----------|-----------|----------------|------------------|------------|
| <u>Chla (mg m⁻³)</u> | | | | | |
| High temp | 516 | 507 | 0.412 | 0.680 | |
| High pCO ₂ | 516 | 507 | 0.664 | 0.507 | |
| Time | 516 | 507 | 3.815 | <0.001 | *** |
| High temp x high pCO ₂ | 516 | 507 | 1.100 | 0.271 | |
| Time x high temp | 516 | 507 | -0.213 | 0.831 | |
| Time x high CO ₂ | 516 | 507 | -0.011 | 0.991 | |
| Time x high temp x high CO ₂ | 516 | 507 | 0.340 | 0.734 | |
| <u>Estimated biomass (mg C m⁻³)</u> | | | | | |
| High temp | 80 | 71 | 0.092 | 0.927 | |
| High pCO ₂ | 80 | 71 | 2.102 | 0.036 | * |
| Time | 80 | 71 | 2.524 | 0.012 | * |
| High temp x high pCO ₂ | 80 | 71 | 1.253 | 0.210 | |
| Time x high temp | 80 | 71 | 1.866 | 0.062 | |
| Time x high CO ₂ | 80 | 71 | 4.414 | <0.001 | *** |
| Time x high temp x high CO ₂ | 80 | 71 | -1.050 | 0.294 | |
| <u>POC (mg m⁻³)</u> | | | | | |
| High temp | 48 | 38 | -0.977 | 0.328 | |
| High pCO ₂ | 48 | 38 | -0.866 | 0.386 | |
| Time | 48 | 38 | -0.203 | 0.839 | |
| High temp x high pCO ₂ | 48 | 38 | -0.29 | 0.772 | |
| Time x high temp | 48 | 38 | 3.648 | <0.001 | *** |
| Time x high CO ₂ | 48 | 38 | 4.333 | <0.001 | *** |
| Time x high temp x high CO ₂ | 48 | 38 | 0.913 | 0.361 | |
| <u>PON (mg m⁻³)</u> | | | | | |
| High temp | 48 | 38 | -0.640 | 0.522 | |
| High pCO ₂ | 48 | 38 | -0.479 | 0.632 | |
| Time | 48 | 38 | 0.202 | 0.84 | |
| High temp x high pCO ₂ | 48 | 38 | 0.667 | 0.505 | |
| Time x high temp | 48 | 38 | 1.674 | 0.094 | |
| Time x high CO ₂ | 48 | 38 | 2.037 | < 0.05 | * |
| Time x high temp x high CO ₂ | 48 | 38 | -0.141 | 0.730 | |
| <u>POC:PON molC:mol N</u> | | | | | |
| High temp | 48 | 38 | 0.222 | 0.824 | |
| High pCO ₂ | 48 | 38 | 0.029 | 0.977 | |
| Time | 48 | 38 | 0.184 | 0.854 | |
| High temp x high pCO ₂ | 48 | 38 | 0.990 | 0.322 | |
| Time x high temp | 48 | 38 | 2.377 | 0.017 | * |
| Time x high CO ₂ | 48 | 38 | 2.748 | 0.005 | ** |
| Time x high temp x high CO ₂ | 48 | 38 | -0.215 | 0.829 | |

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Table 2. Results of generalized linear mixed model testing for significant effects of time, temperature, pCO₂ and all interactions on phytoplankton species biomass. Significant results are in bold;

* p < 0.05, ** p < 0.01, *** p < 0.001.

| Response variable | n | df | z-value | p | sig |
|---|----------|-----------|----------------|------------------|------------|
| Diatoms (mg C m⁻³) | | | | | |
| High temp | 80 | 70 | -0.216 | 0.829 | |
| High pCO ₂ | 80 | 70 | -0.895 | 0.371 | |
| Time | 80 | 70 | 2.951 | 0.003 | ** |
| High temp x high pCO ₂ | 80 | 70 | 1.063 | 0.288 | |
| Time x high temp | 80 | 70 | -1.151 | 0.250 | |
| Time x high CO ₂ | 80 | 70 | 0.560 | 0.576 | |
| Time x high temp x high CO ₂ | 80 | 70 | 0.368 | 0.713 | |
| Dinoflagellates (mg C m⁻³) | | | | | |
| High temp | 80 | 70 | -0.018 | 0.986 | |
| High pCO ₂ | 80 | 70 | 0.487 | 0.627 | |
| Time | 80 | 70 | -2.347 | 0.019 | * |
| High temp x high pCO ₂ | 80 | 70 | -0.166 | 0.868 | |
| Time x high temp | 80 | 70 | 1.857 | 0.063 | |
| Time x high CO ₂ | 80 | 70 | 1.009 | 0.313 | |
| Time x high temp x high CO ₂ | 80 | 70 | 2.207 | 0.027 | * |
| Nanophytoplankton (mg m⁻³) | | | | | |
| High temp | 80 | 70 | -0.371 | 0.710 | |
| High pCO ₂ | 80 | 70 | -2.108 | 0.035 | * |
| Time | 80 | 70 | 2.162 | 0.031 | * |
| High temp x high pCO ₂ | 80 | 70 | 0.79 | 0.430 | |
| Time x high temp | 80 | 70 | 1.695 | 0.090 | |
| Time x high CO ₂ | 80 | 70 | 3.563 | <0.001 | *** |
| Time x high temp x high CO ₂ | 80 | 70 | -0.806 | 0.420 | |
| Synechococcus (mg m⁻³) | | | | | |
| High temp | 80 | 70 | 3.333 | <0.001 | *** |
| High pCO ₂ | 80 | 70 | 2.231 | 0.026 | * |
| Time | 80 | 70 | 0.049 | 0.961 | |
| High temp x high pCO ₂ | 80 | 70 | 2.391 | 0.017 | * |
| Time x high temp | 80 | 70 | 4.076 | <0.001 | *** |
| Time x high CO ₂ | 80 | 70 | -1.553 | 0.1204 | |
| Time x high temp x high CO ₂ | 80 | 70 | 5.382 | <0.001 | *** |
| Picophytoplankton (mg m⁻³) | | | | | |
| High temp | 80 | 70 | 0.951 | 0.342 | |
| High pCO ₂ | 80 | 70 | -0.472 | 0.637 | |
| Time | 80 | 70 | 0.897 | 0.370 | |
| High temp x high pCO ₂ | 80 | 70 | -1.188 | 0.235 | |
| Time x high temp | 80 | 70 | -0.219 | 0.827 | |
| Time x high CO ₂ | 80 | 70 | 1.411 | 0.158 | |
| Time x high temp x high CO ₂ | 80 | 70 | 2.792 | 0.005 | ** |
| Coccolithophores (mg C m⁻³) | | | | | |
| High temp | 80 | 70 | -0.408 | 0.683 | |
| High pCO ₂ | 80 | 70 | -0.308 | 0.758 | |
| Time | 80 | 70 | 0.211 | 0.833 | |
| High temp x high pCO ₂ | 80 | 70 | -0.319 | 0.750 | |

Table 2 cont'd

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|---|----|----|--------|------------------|------------|
| Time x high temp | 80 | 70 | 0.269 | 0.788 | |
| Time x high CO ₂ | 80 | 70 | 0.295 | 0.768 | |
| Time x high temp x high CO ₂ | 80 | 70 | 0.502 | 0.615 | |
| Cryptophytes (mg C m⁻³) | | | | | |
| High temp | 80 | 70 | 0.207 | 0.836 | |
| High pCO ₂ | 80 | 70 | 0.256 | 0.798 | |
| Time | 80 | 70 | -5.289 | <0.001 | *** |
| High temp x high pCO ₂ | 80 | 70 | -0.349 | 0.727 | |
| Time x high temp | 80 | 70 | 1.885 | 0.059 | |
| Time x high CO ₂ | 80 | 70 | 0.167 | 0.867 | |
| Time x high temp x high CO ₂ | 80 | 70 | 1.694 | 0.090 | |
| Microzooplankton (mg C m⁻³) | | | | | |
| High temp | 80 | 70 | 0.138 | 0.890 | |
| High pCO ₂ | 80 | 70 | -0.142 | 0.887 | |
| Time | 80 | 70 | 0.418 | 0.676 | |
| High temp x high pCO ₂ | 80 | 70 | 0.314 | 0.753 | |
| Time x high temp | 80 | 70 | -0.930 | 0.352 | |
| Time x high CO ₂ | 80 | 70 | -2.100 | 0.036 | * |
| Time x high temp x high CO ₂ | 80 | 70 | -1.996 | 0.046 | * |

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Table 3. FRRf-based photosynthesis-irradiance curve parameters for the experimental treatments on the final day (T36).

| Parameter | Control | sd | High temp | sd | High CO ₂ | sd | Combination | sd |
|----------------------------------|---------|-------|-----------|------|----------------------|-------|-------------|-------|
| P^B_m | 2.77 | 1.63 | 9.58 | 1.94 | 18.93 | 2.65 | 3.02 | 0.97 |
| α | 0.03 | 0.01 | 0.09 | 0.01 | 0.13 | 0.01 | 0.04 | 0.00 |
| I_k | 85.33 | 45.47 | 110.93 | 6.09 | 144.13 | 17.91 | 86.38 | 33.06 |

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989 **Table 4.** Results of generalised linear model testing for significant effects of temperature, CO₂ and temperature
 990 x CO₂ on phytoplankton photophysiology at T36; P^B_m (maximum photosynthetic rates), α (light limited slope)
 991 and I_k (light saturated photosynthesis). Significant results are in bold; * p < 0.05, ** p < 0.001, *** p < 0.0001.

| Response variable | n | df | t-value | p | sig |
|---|----------|-----------|----------------|--------------------|------------|
| <u>P^B_m</u> | | | | | |
| High temp | 12 | 8 | 7.353 | < 0.0001 | *** |
| High pCO ₂ | 12 | 8 | 8.735 | < 0.0001 | *** |
| High temp x high pCO ₂ | 12 | 8 | -8.519 | < 0.0001 | *** |
| <u>α</u> | | | | | |
| High temp | 12 | 8 | 13.03 | < 0.0001 | *** |
| High pCO ₂ | 12 | 8 | 15.15 | < 0.0001 | *** |
| High temp x high pCO ₂ | 12 | 8 | -14.82 | < 0.0001 | *** |
| <u>I_k</u> | | | | | |
| High temp | 12 | 8 | 2.018 | 0.0783 | |
| High pCO ₂ | 12 | 8 | 2.541 | 0.0347 | * |
| High temp x high pCO ₂ | 12 | 8 | -2.441 | 0.0405 | * |

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