



1 **Effects of elevated CO₂ and temperature on phytoplankton community**
2 **biomass, species composition and photosynthesis during an autumn**
3 **bloom in the Western English Channel**

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8

9 **Abstract**

10 The combined effects of elevated pCO₂ and temperature were investigated during an autumn
11 phytoplankton bloom in the Western English Channel (WEC). A full factorial 36-day microcosm
12 experiment was conducted under year 2100 predicted temperature (+ 4.5 °C) and pCO₂ levels
13 (800 µatm). The starting phytoplankton community biomass was 110.2 (± 5.7 sd) mg carbon (C)
14 m⁻³ and was dominated by dinoflagellates (~50 %) with smaller contributions from
15 nanophytoplankton (~13 %), cryptophytes (~11 %) and diatoms (~9 %). Over the experimental
16 period total biomass was significantly increased by elevated pCO₂ (20-fold increase) and
17 elevated temperature (15-fold increase). In contrast, the combined influence of these two
18 factors had little effect on biomass relative to the ambient control. The phytoplankton
19 community structure shifted from dinoflagellates to nanophytoplankton at the end of the
20 experiment in all treatments. Under elevated pCO₂ nanophytoplankton contributed 90% of
21 community biomass and was dominated by *Phaeocystis* spp., while under elevated temperature
22 nanophytoplankton contributed 85% of the community biomass and was dominated by smaller
23 nano-flagellates. Under ambient conditions larger nano-flagellates dominated while the smallest
24 nanophytoplankton contribution was observed under combined elevated pCO₂ and temperature
25 (~40 %). Dinoflagellate biomass declined significantly under the individual influences of
26 elevated pCO₂, temperature and ambient conditions. Under the combined effects of elevated
27 pCO₂ and temperature, dinoflagellate biomass almost doubled from the starting biomass and
28 there was a 30-fold increase in the harmful algal bloom (HAB) species, *Prorocentrum cordatum*.
29 Chlorophyll a normalised maximum photosynthetic rates (P^{B_m}) increased > 6-fold under
30 elevated pCO₂ and > 3-fold under elevated temperature while no effect on P^{B_m} was observed
31 when pCO₂ and temperature were elevated simultaneously. The results suggest that future
32 increases in temperature and pCO₂ do not appear to influence coastal phytoplankton



33 productivity during autumn in the WEC which would have a negative feedback on atmospheric
34 CO₂.

35 1. Introduction

36 Oceanic uptake of atmospheric CO₂ has increased by ~42% over pre-industrial levels, with an
37 on-going annual increase of ~0.4%. Current CO₂ level has reached ~400 μatm and has been
38 predicted to rise to >700 μatm by the end of this century (Alley *et al.*, 2007), with estimates
39 exceeding 1000 μatm (Raupach *et al.*, 2007; Raven *et al.*, 2005). The oceans are absorbing CO₂
40 from the atmosphere, which results in a shift in oceanic carbonate chemistry resulting in a
41 decrease in seawater pH or ‘Ocean Acidification’ (OA). The projected increase in atmospheric
42 CO₂ and corresponding increase in ocean uptake, is predicted to result in a decrease in global
43 mean seawater pH of 0.3 units below the present value of 8.1 to 7.8 (Wolf-gladrow *et al.*, 1999).
44 Under this scenario, the shift in dissolved inorganic carbon (DIC) equilibria has wide ranging
45 implications for phytoplankton photosynthetic carbon fixation rates and growth (Riebesell,
46 2004).

47 Concurrent with OA, elevated atmospheric CO₂ and other climate active gases have warmed the
48 planet by ~0.6 °C over the past 100 years (IPCC, 2007). Atmospheric temperature has been
49 predicted to rise by a further 1.8 to 4 °C by the end of this century (Alley *et al.*, 2007).
50 Phytoplankton metabolic activity may be accelerated by increased temperature (Eppley, 1972),
51 which can vary depending on the phytoplankton species and their physiological requirements
52 (Beardall and Stojkovic, 2006). Long-term data sets already suggest that ongoing changes in
53 coastal phytoplankton communities are likely due to climate shifts and other anthropogenic
54 influences (Edwards *et al.*, 2006; Smetacek and Cloern, 2008; Widdicombe *et al.*, 2010). The
55 response to OA and temperature can potentially alter the community composition, community
56 biomass and photo-physiology. Understanding how these two factors may interact
57 (synergistically or antagonistically) is critical to our understanding and for predicting future
58 primary productivity (Boyd and Doney, 2002).

59 Laboratory studies of phytoplankton species in culture and studies on natural populations in
60 the field have shown that most species exhibit sensitivity, in terms of growth and
61 photosynthetic rates, to elevated pCO₂ and temperature individually. To date, only a few studies
62 have investigated the interactive effects of these two stressors on natural populations (e.g.
63 Coello-Camba *et al.*, 2014; Feng *et al.*, 2009; Gao *et al.*, 2017; Hare *et al.*, 2007). Most laboratory
64 studies have varied results with species-specific responses, for example, with the diatom
65 *Thalassiosira weissflogii*, pCO₂ elevated to 1000 μatm and + 5 °C temperature increase
66 synergistically enhanced growth, while the same conditions resulted in a reduction in growth



67 for the diatom *Dactyliosolen fragilissimus* (Taucher et al., 2015). Although there have been fewer
68 studies on dinoflagellates, similar variability in responses has been observed, e.g. (Errera et al.,
69 2014; Fu et al., 2008). In natural populations, elevated pCO₂ has stimulated growth in pico- and
70 nanophytoplankton communities (Engel et al., 2008) while increased temperature has reduced
71 biomass of these groups (Moustaka-Gouni et al., 2016). In a recent field study on natural
72 phytoplankton communities, elevated temperature (+ 3°C above ambient) enhanced community
73 biomass in natural populations but the combined influence of elevated temperature and pCO₂
74 caused a reduction in biomass (Gao et al., 2017).

75 Phytoplankton species composition, abundance and biomass has been measured at the time-
76 series station L4 in the western English Channel (WEC) since 1992, to evaluate how global
77 changes could drive future shifts in phytoplankton community structure and carbon
78 biogeochemistry. To compliment the biological time series, key environmental parameters for
79 monitoring the health and state of the WEC are measured weekly including depth profiles of
80 seawater temperature. Dissolved inorganic carbon (DIC) and total alkalinity (TA) has been
81 sampled at station L4 since 2008. Over the past 50 years a 0.5 °C warming has been observed in
82 the WEC (Smyth et al., 2010). The DIC and TA time series is relatively short and as such there is
83 no significant trend in the calculated pCO₂, although it has shown an increase.

84 Based on the existing literature, the working hypotheses of this study are that: (1) community
85 biomass will increase differentially under individual treatments of elevated temperature and
86 pCO₂; (2) elevated pCO₂ will lead to taxonomic shifts due to differences in species-specific CO₂
87 concentrating mechanisms and/or RuBisCO specificity; (3) photosynthetic carbon fixation rates
88 will increase differentially under individual treatments of elevated temperature and pCO₂; (4)
89 elevated temperature will lead to taxonomic shifts due to species-specific thermal optima; (5)
90 temperature and pCO₂ elevated simultaneously will have synergistic effects.

91 The objectives of the study were to investigate: 1) the combined effects of elevated pCO₂ and
92 temperature on phytoplankton community structure, biomass and photosynthetic carbon
93 fixation rates during the autumn transition from diatoms and dinoflagellates to
94 nanophytoplankton at station L4; 2) assess the natural variability in phytoplankton community
95 structure and the carbon biomass of the dominant species observed in the experimental
96 community relative to long-term observations at station L4 over two decades (1993-2014); and
97 3) assess the distribution of biomass of the dominant species observed at the end of the
98 experiment relative to the in-situ gradients of temperature and pCO₂ observed at station L4. The
99 effects of elevated pCO₂ and temperature on phytoplankton succession in autumn is presently
100 unknown.



101 2. Materials and methods

102 2.1 Time series, phytoplankton community composition

103 Station L4 (50° 15'N, 4° 13'W) is located 13 km SSW of Plymouth in a water depth of ~54 m
104 (Harris, 2010) and is regarded as one of Europe's principal coastal time series sites. Sampling is
105 conducted on a weekly basis (weather permitting) and has been on-going since 1992
106 (<http://www.westernchannelobservatory.org>). Phytoplankton taxonomic composition was
107 enumerated from seawater samples collected from 10 m depth, fixed with 2 % (final
108 concentration) acid Lugol's iodine solution and analysed by inverted light microscopy using the
109 Utermöhl counting technique (Utermöhl, 1958; Widdicombe *et al.*, 2010). For phytoplankton
110 carbon biomass values; taxa-specific mean cell bio-volumes were calculated following Kovala &
111 Larrance, (1966) and converted to carbon using the equations of Menden-Deuer & Lessard,
112 (2000).

113 2.2 Perturbation experiment, sampling and experimental set-up

114 Experimental seawater containing a natural phytoplankton community was sampled at station
115 L4 on 7th October 2015 from 10 m depth (40 L). The experimental seawater was gently pre-
116 filtered through a 200 µm Nitex mesh to remove zooplankton grazers, into two 20 L acid-
117 cleaned carboys. In addition, 320 L of seawater was collected into sixteen 20 L acid-cleaned
118 carboys from the same depth for use as experimental media. Immediately upon return to the
119 laboratory the media seawater was filtered through an in-line 0.2 and 0.1 µm filter (Acropak™,
120 Pall Life Sciences) then stored in the dark at 14 °C until use. The experimental seawater was
121 gently and thoroughly mixed and transferred in equal parts from each carboy (to ensure
122 homogeneity) to sixteen 2.5 L borosilicate incubation bottles (4 sets of 4 replicates). The
123 remaining experimental seawater was sampled for initial (T0) concentrations of nutrients,
124 chlorophyll *a*, total alkalinity, dissolved inorganic carbon, particulate organic carbon (POC) and
125 nitrogen (PON) and was also used to characterise the starting experimental phytoplankton
126 community. The incubation bottles were placed in an outdoor simulated in-situ incubation
127 culture system and each set of replicates were linked to one of four 22 L reservoirs filled with
128 the filtered seawater media. Neutral density spectrally corrected blue filters (Lee Filter no. 301)
129 were placed between polycarbonate sheets and mounted to the top, sides and ends of the
130 incubation system to provide ~50 % irradiance, approximating PAR measured at 10 m depth at
131 station L4 on the day of sampling prior to starting experimental incubations. The media was
132 aerated with CO₂ free air and 5 % CO₂ in air precisely mixed using a mass flow controller
133 (Bronkhorst UK Limited) and used for the microcosm dilutions as per the following
134 experimental design: (1) control (390 µatm pCO₂, 14.5 °C matching station L4 in-situ values),



135 (2) high temperature (390 $\mu\text{atm pCO}_2$, 18.5 °C), (3) high pCO_2 (800 $\mu\text{atm pCO}_2$, 14.5 °C) and (4)
136 combination (800 $\mu\text{atm pCO}_2$, 18.5 °C).

137 Initial nutrient concentrations (measured at 0.24 μM nitrate + nitrite, 0.086 μM phosphate and
138 2.14 μM silicate on 7th October 2015) were amended to 8 μM nitrate+nitrite and 0.5 μM
139 phosphate to provide favourable growth conditions. As the phytoplankton community was
140 sampled over the transitional phase from diatoms and dinoflagellates to nanophytoplankton,
141 the in-situ silicate concentration was maintained to reproduce the silicate concentrations
142 typical of this time of year (Smyth et al., 2010). Media transfer and sample acquisition was
143 facilitated by peristaltic pumps and semi-continuous daily dilution rates were set between 10-
144 13 % of the incubation bottle volume following 48 hrs acclimation in batch culture. CO_2
145 enriched seawater was added to the high CO_2 treatment replicates every 24 hrs, acclimating the
146 natural phytoplankton population to increments of elevated pCO_2 from ambient to $\sim 800 \mu\text{atm}$
147 over 8 days followed by maintenance at $\sim 800 \mu\text{atm}$ as per the method described by Schulz *et al.*,
148 (2009). This protocol was preferred since some phytoplankton species are inhibited by the
149 mechanical effects of direct bubbling (Riebesell et al., 2010; Shi et al., 2009) which can cause a
150 reduction in growth rates and the formation of aggregates (Love et al., 2016). pH was monitored
151 daily to adjust the pCO_2 of the experimental media (+/-) prior to dilutions to maintain target
152 pCO_2 levels in the incubation bottles.

153 2.3 Analytical methods, experimental seawater

154 2.3.1 Chlorophyll *a*

155 Chlorophyll *a* (Chl *a*) was measured in each incubation bottle. 100 mL triplicate samples from
156 each replicate were filtered onto 25 mm GF/F filters (nominal pore size 0.7 μm), extracted in 90
157 % acetone overnight at -20 °C and chl *a* was estimated on a Turner Trilogy™ fluorometer using
158 the non-acidified method of Welschmeyer (1994). The fluorometer was calibrated against a
159 stock Chl *a* standard (*Anacystis nidulans*, Sigma Aldrich, UK), the concentration of which was
160 determined with a Perkin Elmer™ spectrophotometer at wavelengths 663.89 and 750.11 nm.
161 Samples for Chl *a* were taken every 2-3 days.

162 2.3.2 Carbonate system

163 70 mL samples for total alkalinity (TA) and dissolved inorganic carbon (DIC) analysis were
164 collected from each experimental replicate, stored in amber borosilicate bottles with no head
165 space and fixed with 40 μL of super-saturated Hg_2Cl_2 solution for later determination (Apollo
166 SciTech™ Alkalinity Titrator AS-ALK2; Apollo SciTech™ AS-C3 DIC analyser, with analytical
167 precision of 3 $\mu\text{mol kg}^{-1}$). Duplicate measurements were made for TA and triplicate



168 measurements for DIC. Carbonate system parameter values for media and treatment samples
169 were calculated from TA and DIC measurements using the programme CO₂sys (Pierrot et al.,
170 2006) with dissociation constants of carbonic acid of Mehrbach *et al.*, (1973) refitted by Dickson
171 and Millero (Dickson and Millero, 1987). Samples for TA and DIC were taken every 2-3 days.

172 2.3.3 Phytoplankton community analysis

173 Phytoplankton community analysis was performed by flow cytometry (Becton Dickinson Accuri
174 [™] C6) for the 0.2 to 18 µm size fraction following Tarran *et al.*, (2006) and inverted light
175 microscopy was used to enumerate cells > 18 µm (BS EN 15204,2006). For flow cytometry, 2
176 mL samples fixed with glutaraldehyde to a final concentration of 2 % were flash frozen in liquid
177 nitrogen and stored at -80 °C for later analysis. For inverted light microscopy, 140 mL samples
178 were fixed with 2 % (final concentration) acid Lugol's iodine solution and analysed by inverted
179 light microscopy (Olympus[™] IMT-2) using the Utermöhl counting technique (Utermöhl, 1958;
180 Widdicombe *et al.*, 2010). Phytoplankton community samples were taken at T0, T10, T17, T24
181 and T36.

182 2.3.4 Phytoplankton community biomass

183 The smaller size fraction identified and enumerated through flow cytometry;
184 picophytoplankton, nanophytoplankton, *Synechococcus*, coccolithophores and cryptophytes
185 were converted to carbon biomass (mg C m⁻³) using a spherical model to calculate mean cell
186 volume:

$$187 \left(\frac{4}{3} * \pi * r^3\right)$$

188 and a conversion factor of 0.22 pg C µm⁻³ (Booth, 1988). A conversion factor of 0.285 pg C µm⁻³
189 was used for coccolithophores (Tarran et al., 2006) and cell a volume of 113 µm³ and carbon
190 cell⁻¹ value of 18 pg applied for *Phaeocystis* spp. (Widdicombe *et al.*, 2010). *Phaeocystis* spp.
191 were identified and enumerated by flow cytometry separately to the nanophytoplankton class
192 due to high observed abundance in in the high pCO₂ treatment. Mean cell measurements of
193 individual species/taxa were used to calculate cell bio-volume for the 18 µm + size fraction
194 according to Kovala and Larrance (1966) and converted to biomass according to the equations
195 of Menden-Deuer & Lessard, (2000).

196 2.3.5 POC and PON

197 Samples for particulate organic carbon (POC) and particulate organic nitrogen (PON) were
198 taken at T0, T15 and T36.150 mL samples were taken from each replicate and filtered under
199 gentle vacuum onto pre-ashed 25mm glass fibre filters (GF/F, nominal pore size 0.7 µm). Filters



200 were stored in acid washed petri-slides at -20 °C until further processing. Sample analysis was
201 conducted using a Thermoquest Elemental Analyser (Flash 1112). Acetanilide standards (Sigma
202 Aldrich, UK) were used to calibrate measurements of carbon and nitrogen and also used during
203 the analysis to account for any drift in measured values.

204 2.3.6 Chl fluorescence-based photophysiology

205 Photosystem II (PSII) variable Chl fluorescence parameters were measured using a fast
206 repetition rate fluorometer (FRRf) (FastOcean sensor in combination with an Act2Run
207 laboratory system, Chelsea Technologies, West Molesey, UK). The excitation wavelengths of the
208 FRRf's light emitting diodes (LEDs) were 450, 530 and 624 nm. The instrument was used in
209 single turnover mode with a saturation phase comprising 100 flashlets on a 2 µs pitch and a
210 relaxation phase comprising 40 flashlets on a 50 µs pitch. Measurements were conducted in a
211 temperature controlled chamber at 15 °C. The minimum (F_o) and maximum (F_m) Chl
212 fluorescences were estimated according to Kolber et al., (1998). Maximum quantum yields of
213 PSII were calculated as:

$$214 F_v / F_m = (F_m - F_o) / F_m$$

215 PSII electron flux was calculated on a volume basis (JV_{PSII} ; mol e⁻ m⁻³ d⁻¹) using the absorption
216 algorithm (Oxborough et al., 2012) following spectral correction by normalising the FRRf LED
217 emission to the white spectra using Fast^{PRO} 8 software. This step required inputting the
218 experimental phytoplankton community fluorescence excitation spectra values (FES). Since we
219 did not measure the FES of our experimental samples, we used mean literature values for each
220 phytoplankton group calculated proportionally (based on percentage contribution to total
221 estimated biomass per phytoplankton group) as representative values for our experimental
222 samples. The JV_{PSII} rates were converted to Chl specific carbon fixation rates (mg C (mg Chl *a*)⁻¹
223 m⁻³ h⁻¹), calculated as:

$$224 JV_{PSII} \times \varphi_{E:C} \times MW_C / \text{Chl } a$$

225 where $\varphi_{E:C}$ is the electron requirement for carbon uptake (molecule CO₂ (mol electrons)⁻¹), MW_C
226 is the molecular weight of carbon and Chl *a* is the Chl *a* measurement specific to each sample.
227 Chl specific JV_{PSII} based photosynthesis-irradiance curves were conducted in replicate batches
228 between 10:00 – 16:00 to account for variability over the photo-period at between 8 - 14
229 irradiance intensities. The maximum intensity applied was adjusted according to ambient
230 natural irradiance on the day of sampling. Maximum photosynthetic rates of carbon fixation
231 (P_m^B), the light limited slope (α^B) and the light saturation point of photosynthesis (I_k) were
232 estimated by fitting the data to the model by Webb et al., (1974):



233 $P^B = (1 - e^{-\alpha \times I / P^B_m})$

234 Samples for FRRf fluorescence-based light curves were taken at T36.

235 2.4 Statistical analysis

236 To test for effects of high pCO₂, high temperature and high pCO₂ x high temperature on the
237 measured response variables (Chl *a*, total community biomass, POC, PON, photosynthetic
238 parameters and biomass of individual species), generalised least squares models with the
239 factors pCO₂, temperature and time (and all interactions) were applied to the data between T0
240 and T36 incorporating an auto-regressive correlation structure of the order (1) to account for
241 auto correlation. To test for significant differences between experimental treatments at T36 in
242 all measured parameters, generalized linear models were applied to the data. Where main
243 effects were established, pairwise comparisons were performed using the method of Herberich
244 *et al.*, (2010) for data with non-normality and/or heteroscedasticity. Weekly biomass values
245 from the L4 time-series were averaged over years to elucidate the variability and seasonal
246 cycles of the dominant species observed in the experimental community at T36, relative to the
247 time-series observations. The distribution of these species biomass at station L4 was also
248 analysed relative to the in-situ gradients of temperature (1993-2014) and pCO₂ (2008-2014)
249 using frequency histograms. Analyses were conducted using the R statistical package (R Core
250 Team (2014). R: A language and environment for statistical computing. R Foundation for
251 Statistical Computing, Vienna, Austria).

252 3. Results

253 Chl *a* concentration in the WEC ranged between 0.02~5 mg m⁻³ from 30 September - 6th
254 October 2015, with a concentration of ~1.6 mg m⁻³ at station L4 (**Fig. 1 A**). Over the period
255 leading up to phytoplankton community sampling, increasing nitrate and silicate concentrations
256 coincided with a Chl *a* peak on 23rd September (**Fig. 1 B**). Routine net trawl (20 μm) sample
257 observations indicated a phytoplankton community dominated by the diatoms *Leptocylindrus*
258 *danicus* and *L. minimus* with a lower presence of the dinoflagellates *Prorocentrum cordatum*,
259 *Heterocapsa* spp. and *Oxytoxum gracile*. Following decreasing nitrate concentrations, this
260 community transitioned to a *P. cordatum* bloom on 29th September, the week before
261 experimental community sampling (data not shown).

262 3.1 Experimental carbonate system

263 Equilibration to the target high pCO₂ values (800 μatm) within the high pCO₂ and combination
264 treatments was achieved at T10 (**Fig. 2 A**). These treatments were slowly acclimated to
265 increasing levels of pCO₂ over 7 days (from the initial dilution at T3) while the control and high



266 temperature treatments were acclimated at the same ambient carbonate system values as that
267 from station L4 on the day of sampling. Following equilibration, the mean $p\text{CO}_2$ values within
268 the control and high temperature treatments were $394.9 (\pm 4.3 \text{ sd})$ and $393.2 (\pm 4.8 \text{ sd}) \mu\text{atm}$
269 respectively, while in the high $p\text{CO}_2$ and combination treatments mean $p\text{CO}_2$ values were 822.6
270 (± 9.4) and $836.5 (\pm 15.6 \text{ sd}) \mu\text{atm}$, respectively. Carbonate system values remained stable
271 throughout the experiment (**Fig. 2 B-D**).

272 3.2 Experimental temperature treatments

273 Mean temperatures in the control and high $p\text{CO}_2$ treatments were $14.1 (\pm 0.35 \text{ sd}) ^\circ\text{C}$ and in the
274 high temperature and combination treatments the mean temperatures were $18.6 (\pm 0.42 \text{ sd}) ^\circ\text{C}$.
275 There was a mean temperature difference between the ambient and high temperature
276 treatments of $4.46 (\pm 0.42 \text{ sd}) ^\circ\text{C}$ (Supporting information, Fig. S1 A & B).

277

278 3.3 Chlorophyll *a*

279 Mean Chl *a* in the experimental seawater at T0 was $1.64 (\pm 0.02 \text{ sd}) \text{ mg m}^{-3}$ (**Fig. 3 A**). This
280 decreased in all treatments between T0 to T7, to $\sim 0.1 (\pm 0.09, 0.035 \text{ and } 0.035 \text{ sd}) \text{ mg m}^{-3}$ in the
281 control, high $p\text{CO}_2$ and combination treatments, while in the high temperature treatment at T7
282 Chl *a* was $0.46 \text{ mg m}^{-3} (\pm 0.29 \text{ sd})$. From T7 to T12 there was an increase in Chl *a* in all
283 treatments which was highest in the combination ($4.99 \text{ mg m}^{-3} \pm 0.69 \text{ sd}$) and high $p\text{CO}_2$
284 treatments ($3.83 \text{ mg m}^{-3} \pm 0.43 \text{ sd}$) (**Table 1**). At T36 Chl *a* concentration in the combination
285 treatment was significantly higher than all other treatments at $6.87 (\pm 0.58 \text{ sd}) \text{ mg m}^{-3}$ (**Table**
286 **2**) while the high temperature treatment concentration was significantly higher than the control
287 and high $p\text{CO}_2$ treatments at $4.77 (\pm 0.44 \text{ sd}) \text{ mg m}^{-3}$ (**Table 2**). Mean concentrations for the
288 control and high $p\text{CO}_2$ treatments at T36 were not significantly different at $3.30 (\pm 0.22 \text{ sd})$ and
289 $3.46 (\pm 0.35 \text{ sd}) \text{ mg m}^{-3}$ respectively (pairwise comparison $t = 0.78, p = 0.858$).

290

291 3.4 Phytoplankton biomass

292 The starting biomass in all treatments was $110.2 (\pm 5.7 \text{ sd}) \text{ mg C m}^{-3}$ (**Fig. 3 B**) and the
293 community biomass was dominated by dinoflagellates ($\sim 50\%$) with smaller contributions from
294 nanophytoplankton ($\sim 13\%$), cryptophytes ($\sim 11\%$), diatoms ($\sim 9\%$), coccolithophores ($\sim 8\%$),
295 *Synechococcus* ($\sim 6\%$) and picophytoplankton ($\sim 3\%$). Total biomass increased significantly in
296 all treatments over time (**Table 1**) and at T10, it was significantly higher in the high
297 temperature treatment when the biomass reached $752 (\pm 106 \text{ sd}) \text{ mg C m}^{-3}$. At T36 however,
298 total biomass was significantly higher in the high $p\text{CO}_2$ treatment (**Table 1**) and reached 2481



299 (± 182.68 sd) mg C m⁻³, which increased more than 20-fold from T0. Total biomass in the high
300 temperature treatment increased more than 15-fold to 1735 (± 169.24 sd) mg C m⁻³ at T36 and
301 was significantly higher than the combination treatment and ambient control, which were 525
302 (± 28.02 sd) mg C m⁻³ and 378 (± 33.95 sd) mg C m⁻³, respectively (**Table 2**).

303 Measured POC followed the same trends as estimated biomass in all treatments between T0 and
304 T36 (**Fig. 3 C**) and despite some variability between the two measures, POC was within the
305 range of estimates ($R^2 = 0.914$, **Fig. 3 D**). At T36, POC was significantly greater in the high pCO₂
306 treatment (2086 \pm 155.19 sd mg m⁻³) followed by the high temperature treatment (1594 \pm
307 162.24 sd mg m⁻³), which were significantly greater than the control and combination treatment
308 (**Table 1**). PON followed the same trends as POC over the course of the experiment (**Fig. 3 E**,
309 **Table 1**): at T36 concentrations were 147 (± 12.99 sd) and 133 (± 15.59 sd) mg m⁻³ in the high
310 pCO₂ and high temperature treatments respectively, while PON was 57.75 (± 13.07 sd) mg m⁻³
311 in the combination treatment and 47.18 (± 9.32 sd) mg m⁻³ in the control (**Table 1**). POC:PON
312 ratios increased significantly over time in all treatments except for the control. The largest
313 increase of 33 %, from 10.72 to 14.26 mg m⁻³ (± 1.73 sd) was in the high pCO₂ treatment,
314 followed by an increase of 32 % to 9.83 (± 1.82 sd) mg m⁻³ in the combination treatment, and an
315 increase of 17 % to 12.09 (± 2.14 sd) mg m⁻³ in the high temperature treatment. In contrast, the
316 POC:PON ratio in the control declined by 20 % from T0 to T36, from 10.33 to 8.26 (± 0.50 sd)
317 mg m⁻³ (**Fig. 3 F, Table 1**).

318 3.5 Community composition

319 At T36 diatoms dominated the phytoplankton community biomass in the ambient control with a
320 substantial contribution from nanophytoplankton (**Fig. 4 A**), while the high temperature and
321 high pCO₂ treatments exhibited near mono-specific dominance of nanophytoplankton (**Figs. 4 B**
322 **& C**). The most diverse community was in the combination treatment where dinoflagellates and
323 *Synechococcus* became more prominent (**Fig. 4 D**).

324 Between T10 and T24 the community shifted to nanophytoplankton in all experimental
325 treatments. This dominance was maintained through to T36 in the high temperature and high
326 pCO₂ treatments whereas in the ambient control and combination treatment, the community
327 shifted away from nanophytoplankton (**Fig. 5 A**). At T36 nanophytoplankton biomass was
328 significantly higher in the high pCO₂ treatment followed by the high temperature treatment
329 (**Table 2**) when biomass attained 2216 (± 189.67 sd) mg C m⁻³ and 1489 (± 170.32 sd) mg C m⁻³,
330 respectively. In the combination treatment nanophytoplankton biomass was 238 (± 14.16 sd)
331 mg C m⁻³ at T36 which was significantly higher compared to the ambient control (162 \pm 20.02 sd
332 mg C m⁻³; **Table 2**). In addition to significant differences in nanophytoplankton biomass



333 amongst the experimental treatments, treatment-specific differences in cell size was observed.
334 Larger nano-flagellates dominated the control (mean cell diameter of 6.34 μm), smaller nano-
335 flagellates dominated the high temperature and combination treatments (mean cell diameters
336 of 3.61 μm and 4.28 μm) whereas *Phaeocystis* spp. dominated the high pCO₂ treatment (mean
337 cell diameter 5.04 μm) and was not observed in any other treatment (Supporting Information,
338 Fig. S2 A-D).

339 Low starting biomass of diatoms at T0 was dominated by *Coscinodiscus wailessi* (48 %; 4.99 mg
340 C m⁻³), *Pleurosigma* (25 %; 2.56 mg C m⁻³) and *Thalassiosira subtilis* (19 %; 1.94 mg C m⁻³). Small
341 biomass contributions were made by *Navicula distans*, undetermined pennate diatoms and
342 *Cylindrotheca closterium*. Biomass in the diatom group remained low from T0 to T24 but
343 increased at T36 in all treatments, with significantly higher biomass in the high pCO₂ treatment
344 (235 ± 21.41 sd mg C m⁻³, **Fig. 5 B, Table 2**). The highest diatom contribution to total
345 community biomass at T36 was in the ambient control (52 % of biomass; 198 ± 17.28 sd mg C
346 m⁻³). In both the high temperature and combination treatments diatom biomass was
347 significantly lower at T36 (151 ± 10.94 sd and 124 ± 19.16 sd mg C m⁻³, respectively). In all
348 treatments at T36, diatom biomass shifted away from dominance of the larger *C. Wailessii* to the
349 comparatively smaller *C. closterium*, *N. distans*, *T. subtilis* and *Tropidoneis* spp., the relative
350 contributions of which were treatment-specific. Overall *N. distans* dominated diatom biomass in
351 all treatments at T36 (ambient control: 112 ± 24.86 sd mg C m⁻³, 56 % of biomass; high
352 temperature: 106 ± 17.75 sd mg C m⁻³, 70 % of biomass; high pCO₂: 152 ± 19.09 sd mg C m⁻³, 61
353 % of biomass; and combination: 111 ± 20.97 sd mg C m⁻³, 89 % of biomass; Supporting
354 Information, Fig. S3 A-D).

355 The starting dinoflagellate community was dominated by *Gyrodinium spirale* (91 %; 49 mg C m⁻³
356 ³), with smaller contributions from *Katodinium glaucum* (5 %; 2.76 mg C m⁻³), *Prorocentrum*
357 *cordatum* (3 %; 1.78 mg C m⁻³) and undetermined *Gymnodiniales* (1 %; 0.49 mg C m⁻³). At T36
358 dinoflagellate biomass was significantly higher in the combination treatment (90 ± 16.98 sd mg
359 C m⁻³, **Fig. 5 C, Table 2**) followed by the high temperature treatment (57 ± 6.87 sd mg C m⁻³,
360 **Table 2**). There was no significant difference in dinoflagellate biomass between the high pCO₂
361 treatment and ambient control at T36 when biomass was low. In the combination treatment,
362 dinoflagellate biomass shifted away from the larger *G. spirale* and was dominated by *P.*
363 *cordatum* which contributed 59 ± 12.95 sd mg C m⁻³ (66 % of biomass in this group).

364 *Synechococcus* biomass was significantly higher at T36 in the combination treatment (59.9 ±
365 4.30 sd mg C m⁻³, **Fig. 5 D, Table 2**) followed by the high temperature treatment (30 ± 5.98 sd
366 mg C m⁻³, **Table 2**). In both the high pCO₂ treatment and ambient control at T36 *Synechococcus*
367 biomass was low (~7 mg C m⁻³ in both treatments). Relative to the other phytoplankton groups,



368 biomass of picophytoplankton (**Fig. 5 E**), cryptophytes (**Fig. 5 F**) and coccolithophores (**Fig. 5**
369 **G**) remained low in all treatments throughout the experiment. Though picophytoplankton
370 responded positively to the high pCO₂ and combination treatments at T36 (high pCO₂: 6.93 ±
371 0.63 sd mg C m⁻³; combination: 11.26 ± 0.79 sd mg C m⁻³; **Table 2**).

372 **3.6 Chl fluorescence-based photophysiology**

373 At T36, FRRf PI parameters were strongly influenced by the experimental treatments. P^{B_m} was
374 significantly higher in the high pCO₂ treatment (18.93 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹), followed by the
375 high temperature treatment (9.58 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹; **Fig. 6, Tables 3 & 4**). There was no
376 significant difference in P^{B_m} between the ambient control and combination treatment (2.77 and
377 3.02 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹). Light limited photosynthetic efficiency (α^B) also followed the
378 same trend and was significantly higher in the high pCO₂ treatment (0.13 mg C (mg Chl *a*)⁻¹ m⁻³
379 h⁻¹ (μmol photon m⁻² s⁻¹)⁻¹) followed by the high temperature treatment (0.09 mg C (mg Chl *a*)⁻¹
380 m⁻³ h⁻¹ (μmol photon m⁻² s⁻¹)⁻¹) (**Tables 3 & 4**). α^B was low in both control and combination
381 treatments (0.03 and 0.04 mg C (mg Chl *a*)⁻¹ m⁻³ h⁻¹ (μmol photon m⁻² s⁻¹)⁻¹, respectively). The
382 light saturation point of photosynthesis (*I_k*) was significantly higher in the high pCO₂ treatment
383 relative to all treatments where *I_k* was 144.13 μmol photon m⁻² s⁻¹, though significantly lower in
384 the combination treatment relative to both the high pCO₂ and high temperature treatments
385 (**Tables 3 & 4**).

386 **3.7 Natural variability of biomass in the WEC, station L4 time series.**

387 Nanophytoplankton is a critical component of the station L4 carbon budget. The mean annual
388 total nanophytoplankton biomass over the time series (1993-2014) was 586 (± 16.54 sd) mg C
389 m⁻³ with maximum annual biomass of 1182 mg C m⁻³ in 2002 (62 % of total annual
390 phytoplankton biomass) and minimum annual biomass of 262 mg C m⁻³ in 2008 (23 % of total
391 annual phytoplankton biomass). In 8 of the 21 years of time series observations,
392 nanophytoplankton contributed more than 40 % of the station L4 carbon budget. Consistently
393 over the seasonal cycle at L4, mean nanophytoplankton biomass > 10 mg C m⁻³ occurred from
394 early April until the end of October, exhibiting sustained long-term seasonality relative to other
395 phytoplankton groups, though maximal biomass was constrained between April and the 3rd
396 week of June with one exception (**Fig. 7 A**).

397 *N. distans* dominated diatom biomass in the experimental communities though was found to be
398 a very minor component of the diatom carbon budget at station L4 (0.04 % of total annual
399 diatom biomass). Weekly *N. distans* biomass averaged over the time series was very low and
400 ranged from below the limit of detection to ~0.2 mg C m⁻³ with maximum total annual biomass



401 of $\sim 0.5 \text{ mg C m}^{-3}$ in 2005 (**Fig. 7 B**). Seasonality of maximal *N. distans* biomass was constrained
402 to September-October when the mean maximal biomass was 0.03 mg C m^{-3} .

403 *P. cordatum* dominated the dinoflagellate biomass in the experimental communities, and made a
404 significant contribution to total biomass in the combination treatment. Weekly *P. cordatum*
405 biomass averaged over the time series at L4 ranged from 0.004 to 107 mg C m^{-3} and exhibited
406 strong seasonality. Mean total annual *P. cordatum* biomass was 25.5 mg C m^{-3} with maximum
407 annual biomass of 233 mg C m^{-3} in 2006 (minimum annual biomass of $0.004 \text{ mg C m}^{-3}$ in 1994).
408 The bloom peak (taken as an increase in biomass $> 1.0 \text{ mg C m}^{-3}$) usually occurred in September
409 although as early as mid-June (2001 and 2013) in some cases (**Fig. 7 C**). Mean maximal biomass
410 was 12.7 mg C m^{-3} with positive anomalies occurring in 5 out of 21 years throughout the time-
411 series (ranging from 15 to 107 mg C m^{-3}). *P. cordatum* contributed on average, 9.2 % of the total
412 annual dinoflagellate biomass with a maximum contribution of ~ 55 % in 2006 and 12 % and 63
413 % when averaged over the bloom period from mid-June to end-September. *P. cordatum*
414 contributed 3.4 % of total phytoplankton biomass during the bloom period and ~ 32 % of total
415 phytoplankton biomass in 2006 during an unprecedented bloom when biomass attained 107
416 mg C m^{-3} in 2006 (**Fig. 7 D**).

417 Group and species-specific optimal temperature ranges were found when considering how the
418 dominant experimental species were temporally distributed relative to in-situ temperatures:
419 nanophytoplankton exhibited a bi-modal distribution with 31 % biomass between $9\text{-}11^\circ\text{C}$ and
420 30 % between $15\text{-}16.5^\circ\text{C}$, with 6 % above 16.5°C (**Fig. 8 A**). ~ 60 % of *N. distans* biomass
421 occurred between $14\text{-}16^\circ\text{C}$ and 2 % biomass above 16°C (**Fig. 8 B**). In contrast, 66 % of *P.*
422 *cordatum* biomass occurred between $14\text{-}16^\circ\text{C}$ and 24 % above 16°C (**Fig. 8 C**). Biomass
423 distribution relative to station L4 in-situ pCO_2 levels (2008-2014) also demonstrated
424 group/species-specific optimal ranges. 71 % nanophytoplankton biomass occurred at a pCO_2
425 range of $245\text{-}410 \mu\text{atm}$, 24 % $410\text{-}515 \mu\text{atm}$ and 5 % between $515\text{-}680 \mu\text{atm}$ (**Fig. 8 D**). *N.*
426 *distans* followed a similar trend with 72% biomass between a pCO_2 range of $245\text{-}410 \mu\text{atm}$, 26%
427 between $410\text{-}515 \mu\text{atm}$ and $< 2\%$ beyond $515 \mu\text{atm}$ (**Fig. 8 E**). By contrast, 97% of *P. cordatum*
428 biomass occurred between $245\text{-}410 \mu\text{atm}$ with 3 % biomass occurring beyond $410 \mu\text{atm}$ (**Fig. 8**
429 **F**).

430 4. Discussion

431 Individually, elevated temperature and pCO_2 resulted in the highest biomass and maximum
432 photosynthetic rates (P^{B_m}), when nanophytoplankton dominated. The interaction of these two
433 factors had little effect on total biomass with values close to the ambient control, and no effect



434 on P^B_m . The combination treatment, however, exhibited the greatest diversity of phytoplankton
435 functional groups with dinoflagellates and *Synechococcus* becoming more prominent.

436 Elevated pCO_2 has been shown to enhance the growth and photosynthesis of some
437 phytoplankton species which have active uptake systems for inorganic carbon (Giordano et al.,
438 2005; Reinfelder, 2011). Elevated pCO_2 may therefore lead to lowered energetic costs of carbon
439 assimilation in some species and a redistribution of the cellular energy budget to other
440 processes (Tortell et al., 2002). In the present study, under elevated pCO_2 where the dominant
441 group was nanophytoplankton, the community was dominated by the bloom-forming
442 haptophyte *Phaeocystis* spp. Photosynthetic carbon fixation in *Phaeocystis* spp. is presently near
443 saturation with respect to current levels of pCO_2 (Rost et al., 2003). Inorganic carbon acquisition
444 in this species has been shown to be equal to, or more efficient than that of diatoms. Indeed, in
445 *Phaeocystis* spp, extracellular carbonic anhydrase is regulated by CO_2 (aq), and HCO_3^- is utilized
446 as a carbon source in photosynthesis, indicating more efficient use of CO_2 (Elzenga et al., 2000;
447 Rost et al., 2003), and thus providing an advantage to *Phaeocystis* spp. when more CO_2 is
448 present. Therefore, the increased biomass and photosynthetic carbon fixation seen here under
449 elevated pCO_2 can be attributed to the community shift to *Phaeocystis* spp.. The increased
450 biomass seen in the high temperature treatment in this study, may be attributed to enhanced
451 enzymatic activities, since algal growth commonly increases with temperature until after the
452 optimal range (Boyd et al., 2013; Goldman and Carpenter, 1974; Savage et al., 2004) and
453 optimum growth temperatures for marine phytoplankton are often several degrees higher than
454 environmental temperatures (Eppley, 1972; Thomas et al., 2012).

455 **4.1 Chl *a***

456 Chl *a* concentration was significantly higher in the combination treatment at T36 when total
457 biomass was lower, but Chl *a* was significantly lower in the high pCO_2 treatment when biomass
458 was significantly higher than all other treatments. This contrasts the results reported in
459 comparable studies as Chl *a* is generally highly correlated with biomass, (e.g. Feng et al., 2009).
460 Similar results were reported however by Hare et al., (2007) which indicates that Chl *a* may not
461 always be a reliable proxy for biomass in mixed communities. Differences in Chl *a* may therefore
462 be attributed to taxonomic differences in community composition.

463 **4.2 Biomass**

464 This study shows that the phytoplankton community response to elevated temperature and
465 pCO_2 is highly variable. pCO_2 elevated to $\sim 800 \mu atm$ induced higher community biomass in
466 agreement with Kim et al., (2006) and Riebesell et al., (2007), whereas in other natural
467 community studies no CO_2 effect on biomass was observed (Delille et al., 2005; Maudgendre et al.,



468 2017; Paul et al., 2015). A ~ 4.5 °C increase in temperature also resulted in higher biomass in
469 this study, similar to the findings of Feng et al., (2009) and Hare et al., (2007) though elevated
470 temperature has previously reduced biomass of natural nanophytoplankton communities in the
471 Western Baltic Sea and Arctic Ocean (Coello-Camba et al., 2014; Moustaka-Gouni et al., 2016).
472 When elevated temperature and $p\text{CO}_2$ were combined, community biomass exhibited little
473 response, similar to the findings of Gao et al., (2017), though an increase in biomass has also
474 been reported (Calbet et al., 2014; Feng et al., 2009). Geographic location and season also play
475 an important role in structuring the community and its response in terms of biomass to elevated
476 temperature and $p\text{CO}_2$, e.g. (Li et al., 2009; Morán et al., 2010).

477 4.3 Carbon:Nitrogen

478 In agreement with others, the results of this experiment showed highest increases in C:N under
479 elevated $p\text{CO}_2$ alone (Riebesell et al., 2007). C:N also increased under high temperature,
480 consistent with the findings of Lomas and Glibert, (1999) and Taucher et al., (2015) and was
481 stimulated to a lesser degree when $p\text{CO}_2$ and temperature were elevated simultaneously, which
482 was also observed in the study of Calbet et al., (2014), but contrasts other studies that have
483 observed C:N to be unaffected by the combined influence of elevated $p\text{CO}_2$ and temperature, e.g.
484 (Deppeler and Davidson, 2017; Kim et al., 2006; C. Paul et al., 2015). C:N is a strong indicator of
485 cellular protein content (Woods and Harrison, 2003) and increases under elevated $p\text{CO}_2$ and
486 warming may likely lead to lowered nutritional value of phytoplankton with consequences for
487 zooplankton reproduction and biogeochemical cycles.

488 4.4 Photosynthetic carbon fixation rates

489 At T36, under elevated $p\text{CO}_2$ P^{B}_{m} was > 6 times higher than the ambient control, which has also
490 been reported in elevated $p\text{CO}_2$ perturbation experiments by Riebesell et al., (2007) and Tortell
491 et al., (2008), but contrasts other observations on natural populations where the effect of
492 elevated $p\text{CO}_2$ alone was found to reduce P^{B}_{m} (Feng et al., 2009; Hare et al., 2007). Studies on
493 laboratory cultures have shown that increases in temperature increase photosynthetic rates
494 (Feng et al., 2008; Fu et al., 2007; Hutchins et al., 2007), similar to our findings. We found that
495 there was no effect on P^{B}_{m} under the combined treatment which has also been observed in
496 experiments on natural populations (Coello-Camba and Agustí, 2016; Gao et al., 2017). This
497 strongly contrasts the findings of Feng et al., (2009) and Hare et al., (2007) who observed the
498 highest P^{B}_{m} when temperature and $p\text{CO}_2$ were elevated simultaneously. Increases in α^{B} and I_k
499 under elevated $p\text{CO}_2$, and a decrease in these parameters when elevated $p\text{CO}_2$ and temperature
500 were combined is opposite to the trends reported by Feng et al., (2009).



501 Photosynthetic rates have been demonstrated to decrease beyond a temperature of 20 °C
502 (Raven and Geider, 1988) which can be modified through photoprotective rather than
503 photosynthetic pigments (Kiefer and Mitchell, 1983). This may explain the difference in P^B_m
504 between the high pCO_2 and high temperature treatments (in addition to differences in
505 nanophytoplankton community composition in relation to *Phaeocystis* spp. discussed above), as
506 the experimental high temperature treatment in the present study was ~ 4.5 °C higher than
507 ambient.

508 There was no significant effect of combined elevated pCO_2 and temperature on P^B_m , which was
509 strongly influenced by taxonomic differences between the experimental treatments. Warming
510 has been shown to lead to smaller cell sizes in nanophytoplankton (Atkinson et al., 2003; Peter
511 and Sommer, 2012), which was observed in the combined treatment together with decreased
512 nanophytoplankton biomass. Diatoms also shifted to smaller species with reduced biomass,
513 while dinoflagellate and *Synechococcus* biomass increased at T36. Dinoflagellates are the only
514 photoautotrophs with form II RuBisCO (Morse et al., 1995) which has the lowest
515 carboxylation:oxygenation specificity factor among eukaryotic phytoplankton (Badger et al.,
516 1998), giving dinoflagellates a disadvantage in carbon fixation under present ambient pCO_2
517 levels. Dinoflagellates generally grow at slower rates in surface waters with high pH (≥ 9)
518 resulting from photosynthetic removal of CO_2 by previous blooms (Hansen, 2002; Hinga, 2002).
519 Though growth under high pH provides indirect evidence that dinoflagellates possess CCMs,
520 direct evidence is limited and points to the efficiency of CCMs in dinoflagellates as moderate in
521 comparison to diatoms and some haptophytes (Reinfelder, 2011 and references therein). This
522 may explain the lower P^B_m in the combined treatment compared to elevated pCO_2 and
523 temperature individually.

524 4.5 Community composition

525 Phytoplankton community structure changes were observed, with a shift from dinoflagellates to
526 nanophytoplankton which was most pronounced under single treatments of elevated
527 temperature and pCO_2 . Amongst the nanophytoplankton, a distinct size shift to smaller cells was
528 observed in the high temperature and combination treatments, while in the high pCO_2
529 treatment *Phaeocystis* spp. dominated. Under combined pCO_2 and temperature at T36 however,
530 dinoflagellate and *Synechococcus* biomass increased at the expense of nanophytoplankton.

531 An increase in pico- and nanophytoplankton has previously been reported in natural
532 communities under elevated pCO_2 (Bermúdez et al., 2016; Boras et al., 2016; Brussaard et al.,
533 2013; Engel et al., 2008) while no effect on these size classes has been observed in other studies
534 (Calbet et al., 2014; Paulino et al., 2007). Moustaka-Gouni et al., (2016) also found no CO_2 effect



535 on natural nanophytoplankton communities but increased temperature reduced the biomass of
536 this group. Kim et al., (2006) observed a shift from nanophytoplankton to diatoms under
537 elevated pCO₂ alone while a shift from diatoms to nanophytoplankton under combined elevated
538 pCO₂ and temperature has been reported (Hare et al., 2007). A variable response in *Phaeocystis*
539 spp. to elevated pCO₂ has also been reported with increased growth (Chen et al., 2014; Keys et
540 al., 2017), no effect (Thoisen et al., 2015) and decreased growth (Hoogstraten et al., 2012)
541 observed. *Phaeocystis* spp. can outcompete other phytoplankton and form massive blooms (up
542 to 10 g C m⁻³) with impacts on food webs, global biogeochemical cycles and climate regulation
543 (Schoemann et al., 2005). While not a highly toxic algal species, *Phaeocystis* spp. are considered
544 a harmful algal bloom (HAB) species when biomass reaches sufficient concentrations to cause
545 anoxia through the production of mucus foam which can clog the feeding apparatus of
546 zooplankton and fish (Eilertsen & Raa, 1995).

547 The response of diatoms to elevated pCO₂ and temperature has been variable. For example, A
548 study by Taucher et al., (2015) showed that *Thalassiosira weissflogii* incubated at 1000 µatm
549 pCO₂ increased growth by 8 % while for *Dactyliosolen fragilissimus*, growth increased by 39 %;
550 temperature elevated by + 5°C also had a stimulating effect on *T. weissflogii* but inhibited the
551 growth rate of *D. fragilissimus*; and when the treatments were combined growth was enhanced
552 in *T. weissflogii* but reduced in *D. fragilissimus*. In partial agreement, the results of the present
553 experiment show that elevated pCO₂ increased biomass in diatoms but elevated temperature
554 and the combination of these factors reduced biomass. A distinct size-shift in diatom species
555 was observed in all treatments, from the larger *Coscinodiscus* spp., *Pleurosigma* and
556 *Thalassiosira subtilis* to the smaller *Navicula distans*. This was most pronounced in the
557 combination treatment where *N. distans* contributed 89 % of diatom biomass. *Navicula* spp.
558 previously exhibited a differential response to both elevated temperature and pCO₂. At + 4.5 °C
559 and 960 ppm CO₂ Torstensson et al., (2012) observed no synergistic effects on the benthic
560 *Navicula directa*. Elevated temperature increased growth rates by 43 % while a reduction of 5 %
561 was observed under elevated CO₂. No effects on growth were detected at pH ranging from 8 –
562 7.4 units on *Navicula* spp. (Thoisen et al., 2015), while growth in *N. distans* was significantly
563 stimulated along a CO₂ gradient at a shallow cold-water vent system (Baragi et al., 2015).

564 *Synechococcus* grown under pCO₂ elevated to 750 ppm and temperature elevated by 4 °C
565 resulted in increased growth and a 4-fold increase in P^{B_m} (Fu et al., 2007) which is similar to the
566 results of the present study.

567 The combination of elevated temperature and pCO₂ significantly increased dinoflagellate
568 biomass which almost doubled, accounting for 17 % of total biomass. This was due to *P.*
569 *cordatum* which increased biomass by more than 30-fold between T0 and T30 (66 % of



570 dinoflagellate biomass in this treatment). Despite the global increase in the frequency of HABs
571 few studies have focussed on the response of dinoflagellates to elevated pCO₂ and temperature.
572 In laboratory studies at 1000 ppm CO₂, growth rates of the HAB species *Karenia brevis* increased
573 by 46 %, at 1000 ppm CO₂ and + 5 °C temperature it's growth increased by 30 % but was
574 reduced under elevated temperature alone (Errera et al., 2014). A combined increase in pCO₂
575 and temperature enhanced both the growth and P^B_m in the dinoflagellate *Heterosigma akashiwo*,
576 whereas in contrast to the present findings, only pCO₂ alone enhanced these parameters in *P.*
577 *cordatum* (Fu et al., 2008).

578 Among HAB species, *P. cordatum* is widely distributed geographically in temperate and
579 subtropical waters, has detrimental effects at the organismal and environmental levels and is
580 potentially harmful to humans via shellfish poisoning (Heil et al., 2005). Recent increases in the
581 frequency, magnitude and distribution of harmful phytoplankton species has focussed attention
582 on the unique physiological, ecological and toxicological aspects of the species involved
583 (Andeson et al., 2002; Hallegraeff, 1993). Ocean acidification combined with warming could
584 potentially affect the growth and toxicity of HAB species (Fu et al., 2012). Recent studies on
585 several diatom and dinoflagellate species suggest that ocean acidification combined with
586 elevated temperature may dramatically increase the toxicity of some harmful groups (e.g.
587 Flores-Moya et al., 2012; Fu et al., 2010; Sun et al., 2011; Tatters et al., 2012). The ecology and
588 bloom dynamics of *P. cordatum* have been well documented in selected environments (e.g.
589 Chesapeake Bay, Baltic Sea). The spread of this species to previously unreported areas through
590 either ballast water transport, aquaculture development, or increasing eutrophication, has been
591 reported (Heil et al., 2005). In Chesapeake Bay *P. cordatum* 'mahogany tides' have been
592 associated with anoxic/hypoxic events, fish kills, aquaculture shellfish kills and the loss of
593 aquatic vegetation (Tango et al., 2005). Over the last two decades *P. cordatum* has established
594 itself as a dominant summer phytoplankton species in the Baltic Sea but so far there are no
595 reports of toxic blooms (Hajdu et al., 2000). However, for the first time a *P. cordatum* bloom was
596 recorded in February 2002 at Bolinao, Northern Philippines and was coincident with a mass
597 aquaculture fish kill resulting in losses estimated at US\$120,000 (Azanza et al., 2005). Several
598 clones of *P. cordatum* were found to produce a water-soluble neuro-toxin during bloom decline
599 in culture studies (Grzebyk et al., 1997). More recently, a series of positive bioassays for
600 tetrodotoxins (TTXs) was observed in mussels (Vlamiš et al., 2015) which coincided with the
601 simultaneous presence of a *P. cordatum* bloom. Data analysis from previous years (2006 –
602 2012) identified multiple sample cases for toxins in aquaculture production areas coinciding
603 with *P. cordatum* blooms.



604 In addition to strong links to toxic algal events, mixotrophy has also been reported in *P.*
605 *cordatum*. In a study by Stoecker et al., (1997) up to 50% of *P. cordatum* sampled from
606 Chesapeake Bay in the summer contained cryptophyte material. The authors concluded that *P.*
607 *cordatum* feeding is a mechanism for supplementing carbon nutrition and this may explain why
608 the ratio of nanophytoplankton:dinoflagellates was significantly lower in our combination
609 experimental community compared to the other treatments.

610 **4.6 Natural variability of biomass in the WEC, station L4 time series.**

611 During autumn in the WEC, sea surface temperature and pCO₂ start to decline following their
612 respective time series maximal values at station L4. During October, mean seawater
613 temperatures at 10 m decrease from 15.39 °C (± 0.49 sd) to 14.37 °C (± 0.62 sd). Following a
614 period of CO₂ oversaturation in late summer, pCO₂ returns to near-equilibrium at station L4 in
615 October when mean pCO₂ values decrease during this month from 455.32 µatm (± 63.92 sd) to
616 404.06 µatm (± 38.55 sd) (Kitidis et al., 2012). As is the case with seawater warming, predicted
617 future ocean acidification is likely to impact coastal phytoplankton communities in autumn
618 when the present upper limit of the pCO₂ threshold increases during this period of surface
619 ocean-atmosphere equilibrium (Riebesell, 2004).

620 From a biological perspective, the autumn period at station L4 is characterised by the decline of
621 the late summer diatom and dinoflagellate blooms (Widdicombe et al., 2010) when biomass of
622 these two groups approaches values close to the time series minima (diatom biomass range:
623 6.01 (± 6.88 sd) – 2.85 (± 3.28 sd) mg C m⁻³; dinoflagellate biomass range: 1.75 (± 3.28 sd) – 0.66
624 (± 1.08 sd) mg C m⁻³). Typically, over this period nanophytoplankton becomes numerically
625 dominant when biomass of this group ranges from 20.94 (± 33.25 sd) – 9.38 (± 3.31 sd), though
626 the time series shows high variability in this biomass.

627 Comparative analyses of the WEC time series and the dominant species from the experimental
628 treatments showed that nanophytoplankton contributes significantly to the station L4 carbon
629 budget. The in-situ bimodal distribution of nanophytoplankton biomass at cold and warm
630 temperature ranges indicates a potential tolerance to temperature increase. Most
631 nanophytoplankton biomass occurred at an in-situ pCO₂ range between 245-410 µatm but
632 almost 10 % of biomass was distributed between 515-680 µatm, indicating some tolerance to
633 elevated pCO₂ during periods of CO₂ oversaturation at station L4. The dominant diatom species
634 in the experimental communities, *N. distans*, was a very minor contributor to diatom biomass
635 over the time series with most biomass constrained to very narrow in-situ temperature and
636 pCO₂ ranges of 14-16 °C and 245-410 µatm. *P. cordatum* dominated dinoflagellate biomass in
637 the combination treatment but was generally a low biomass contributor to dinoflagellate



638 biomass over the time series, with the exception of one unprecedented bloom in 2006. *P.*
639 *cordatum* biomass exhibited higher thermal in-situ optima with most biomass observed
640 between 14–16 °C and almost a quarter of biomass above 16 °C, indicating tolerance to
641 temperature increase, though the majority of biomass at station L4 occurred at times of low in-
642 situ pCO₂ (245–350 μatm) with just 3% beyond 410 μatm. These trends suggest conditions of
643 warming may favour nanophytoplankton and *P. cordatum*, elevated pCO₂ may favour
644 nanophytoplankton and both factors combined may favour both species. These observations are
645 consistent with the experimental results.

646 5. Implications

647 Increased biomass, P^{B_m} and a community shift to nanophytoplankton under individual increases
648 in temperature and pCO₂ suggests a potential positive feedback on atmospheric CO₂, whereby
649 more CO₂ is removed from the ocean, and hence from the atmosphere by photosynthetic
650 activity. The selection of *Phaeocystis* spp. under elevated pCO₂ indicates the potential for
651 negative impacts on ecosystem function and food web structure associated with this species
652 (Schoemann et al., 2005; Verity et al., 2007). However, while more CO₂ is photosynthesised,
653 selection for nanophytoplankton in both of these treatments may actually result in reduced
654 carbon sequestration due to slower sinking rates of these smaller phytoplankton cells (Bopp et
655 al., 2001; Laws et al., 2000). When temperature and pCO₂ were elevated simultaneously,
656 community biomass showed little response and no effects on P^{B_m} were observed, suggesting a
657 negative feedback on atmospheric CO₂ and climate warming in future warmer high CO₂ oceans.
658 Additionally, combined elevated pCO₂ and temperature significantly modified taxonomic
659 composition, by reducing diatom biomass relative to the ambient control with increasing
660 dinoflagellate biomass dominated by the HAB species, *P. cordatum*. This has implications for
661 fisheries, ecosystem function and human health.

662 6. Conclusion

663 These experimental results provide new evidence that increases in pCO₂ coupled with rising sea
664 temperatures may have antagonistic effects on the autumn phytoplankton community in the
665 WEC. Under future global change scenarios, the size range and biomass of diatoms may be
666 reduced with increased dinoflagellate biomass and the selection of HAB species. The
667 experimental simulations of year 2100 temperature and pCO₂ demonstrate that the effects of
668 warming can be offset by elevated pCO₂ potentially reducing coastal phytoplankton productivity
669 and significantly altering the community structure, and in turn these shifts will have
670 consequences on carbon biogeochemical cycling in the WEC.



671 **Data availability:** Experimental data used for analysis will be made available (DOI will be
672 created)

673 **Author contributions:** Matthew Keys collected, measured, processed and analysed the data and
674 prepared the figures. Drs Gavin Tilstone and Helen Findlay conceived, directed and sought the
675 necessary funds to support the research. Matthew Keys and Dr Gavin Tilstone wrote the paper
676 with input from Claire Widdicombe and Professor Tracy Lawson. Claire Widdicombe supervised
677 and advised on phytoplankton taxonomic classifications.

678 **Competing interests:** The authors declare that they have no conflict of interest.

679 **Acknowledgements:** G.H.T, H.S.F. and C.E.W were supported by the UK Natural Environment
680 Research Council's (NERC) National Capability – The Western English Channel Observatory
681 (WCO). C.E.W was also partly funded by the NERC and Department for Environment, Food and
682 Rural Affairs, Marine Ecosystems Research Program (Grant no. NE/L003279/1). M.K. was
683 supported by a NERC PhD studentship (grant No. NE/L50189X/1). We thank Glen Tarran for his
684 training, help and assistance with flow cytometry, The National Earth Observation Data Archive
685 and Analysis Service UK (NEODAAS) for their help in providing the MODIS image used in Fig 1.
686 and the crew of RV Plymouth Quest for their helpful assistance during field sampling.

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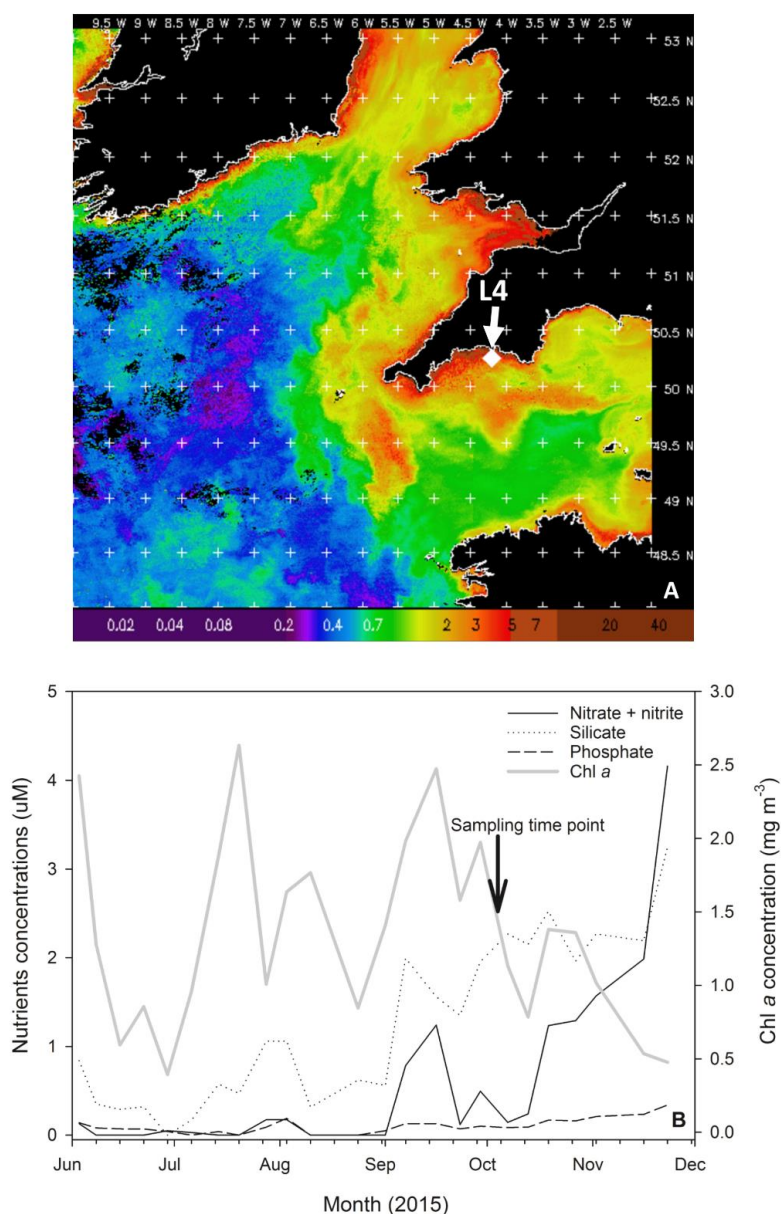


Fig. 1. (A). MODIS weekly composite chl *a* image of the western English Channel covering the period 30th September – 6th October 2015 (coincident with the week of phytoplankton community sampling for the present study), processing courtesy of NEODAAS. The position of coastal station L4 is marked with a white diamond. (B). Profiles of weekly nutrient and chl *a* concentrations from station L4 at a depth of 10 m over the second half of 2015 in the months prior to phytoplankton community sampling (indicated by black arrow and text).

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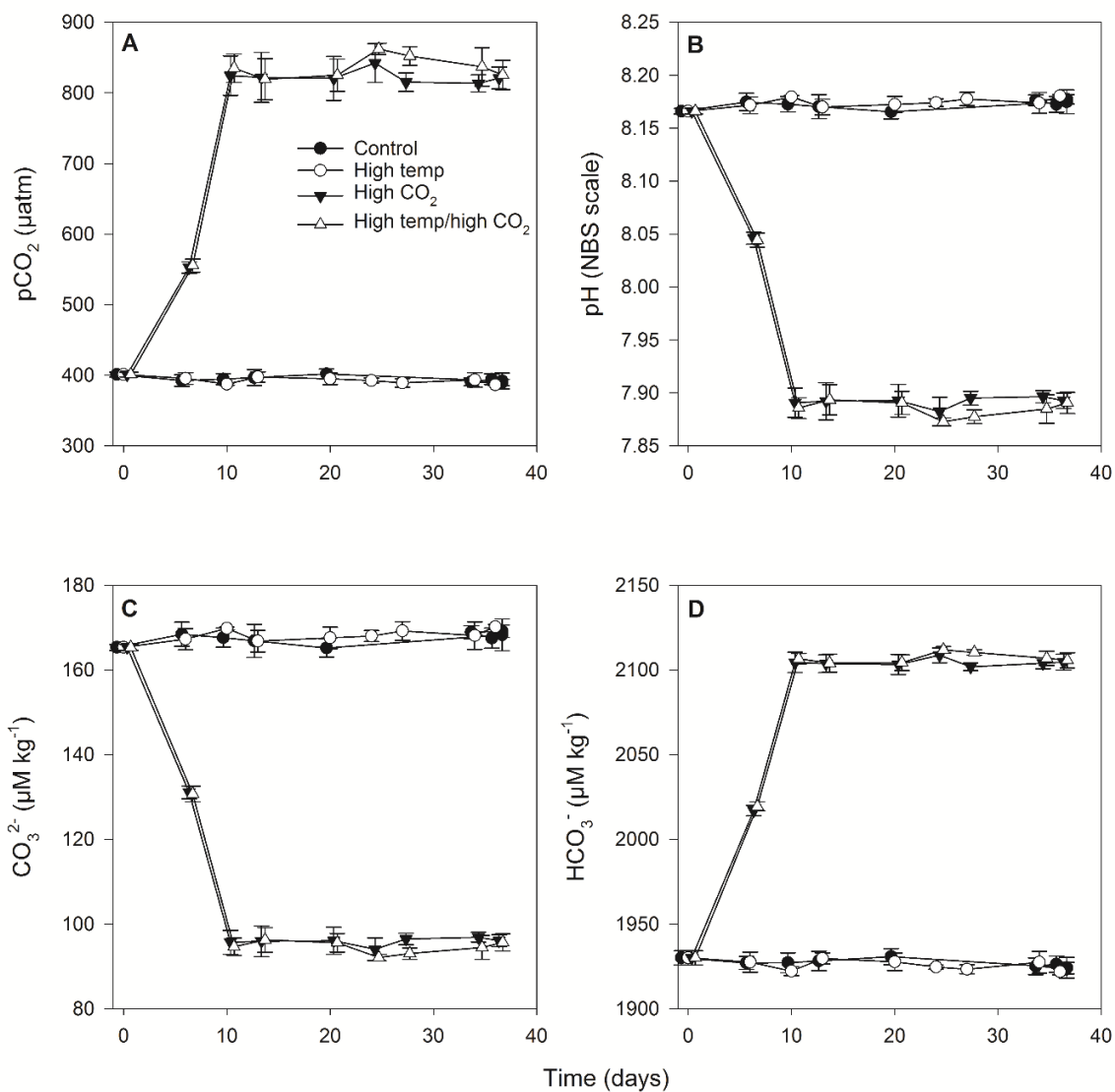


Fig. 2. Carbonate system values of the experimental phytoplankton incubations. **(A)**, partial pressure of CO_2 in seawater ($p\text{CO}_2$), **(B)**, pH on the NBS scale, **(C)**, carbonate concentration (CO_3^{2-}) and **(D)**, bicarbonate concentration (HCO_3^-) were estimated from direct measurements of total alkalinity and dissolved inorganic carbon.

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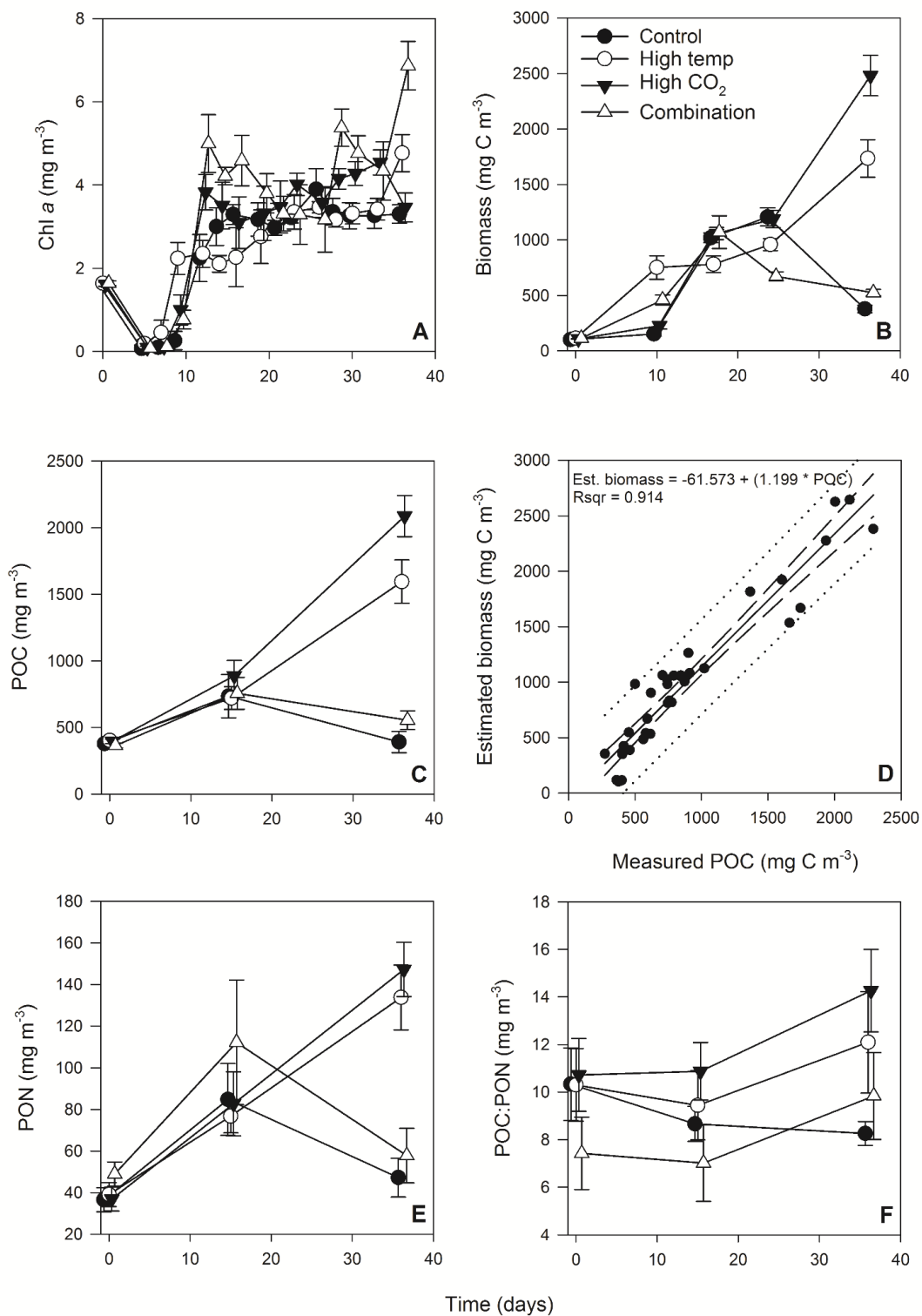


Fig. 3. Time course of chl *a* (A), estimated phytoplankton biomass (B), POC (C), regression of estimated phytoplankton carbon vs measured POC (D), PON (E) and POC:PON (F).



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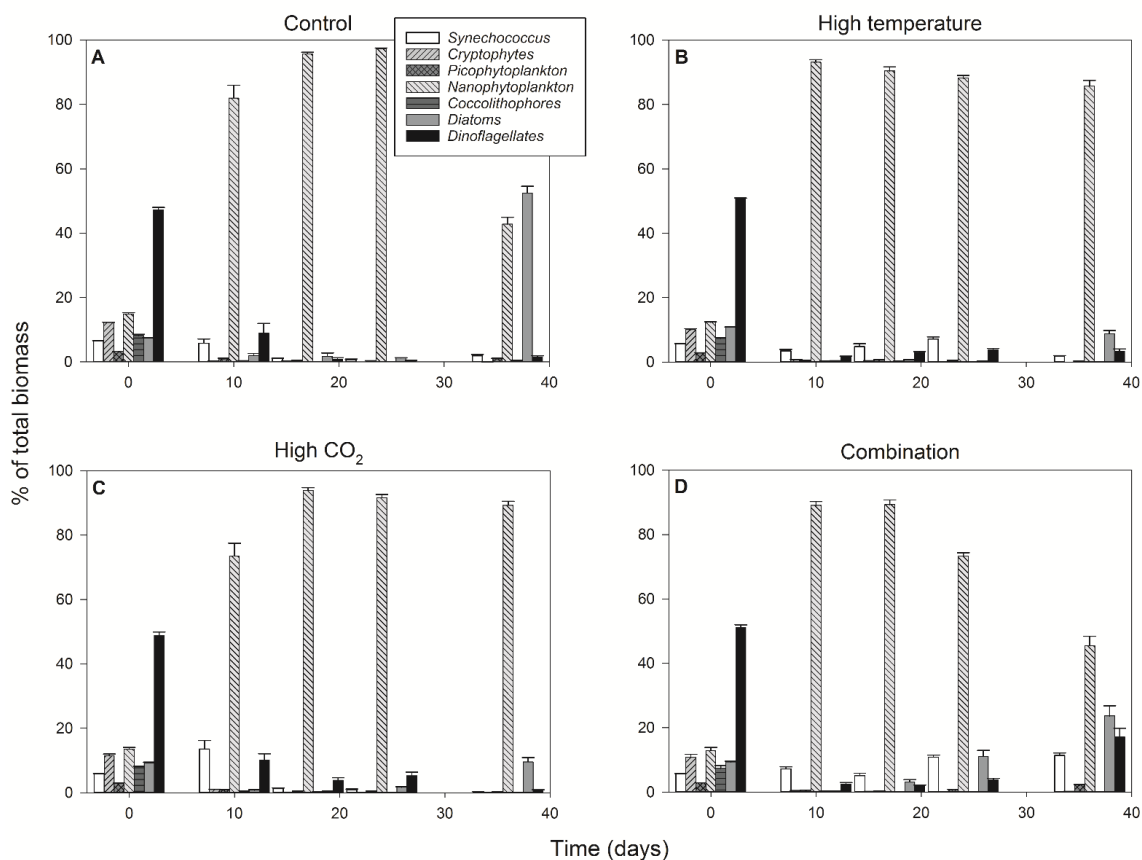


Fig. 4. Percentage contribution to community biomass by phytoplankton groups/species throughout the experiment in the control (A), high temperature (B), high CO₂ (C) and combination treatments (D).

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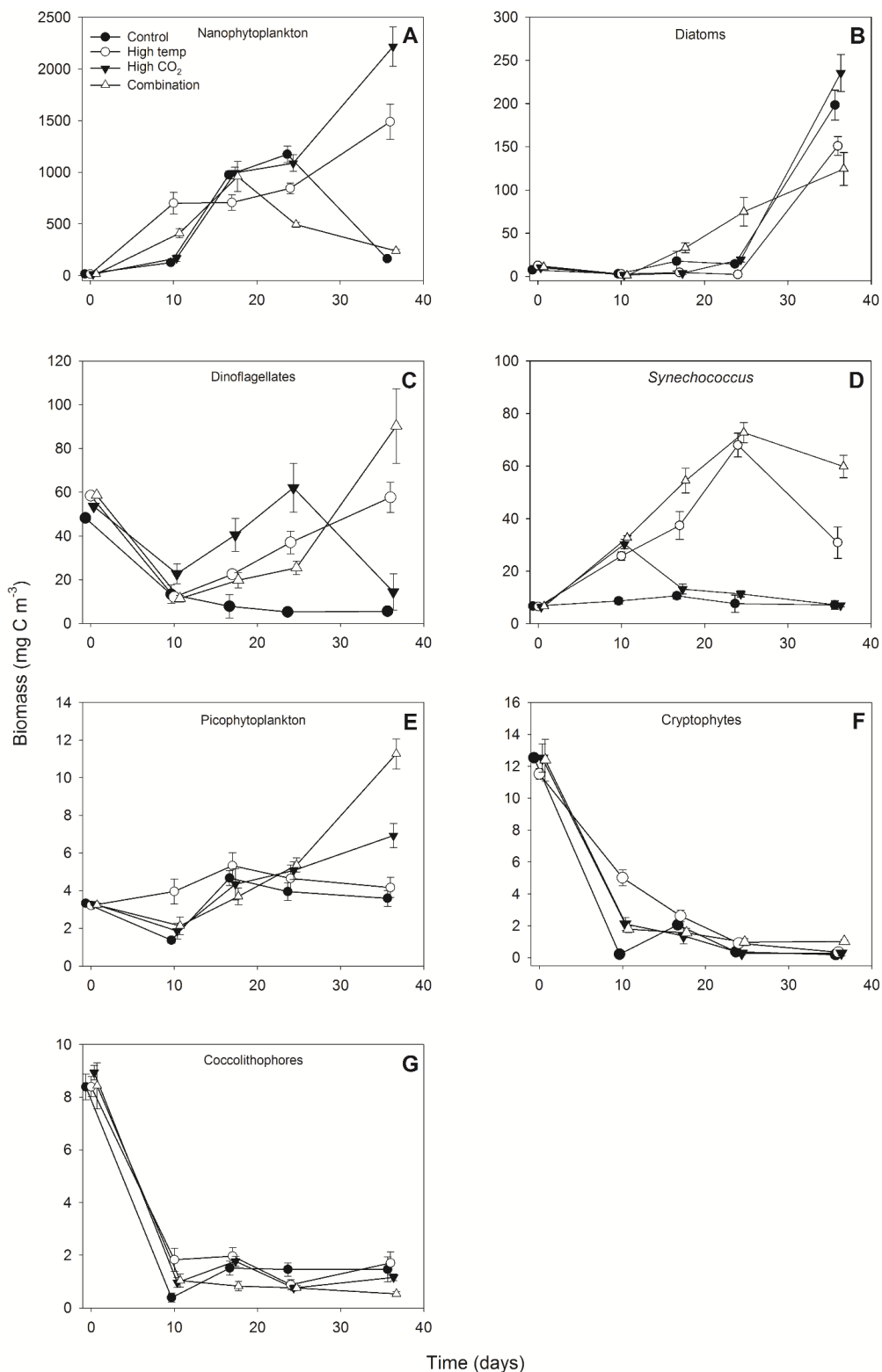


Fig. 5. Response of individual phytoplankton groups to experimental treatments.

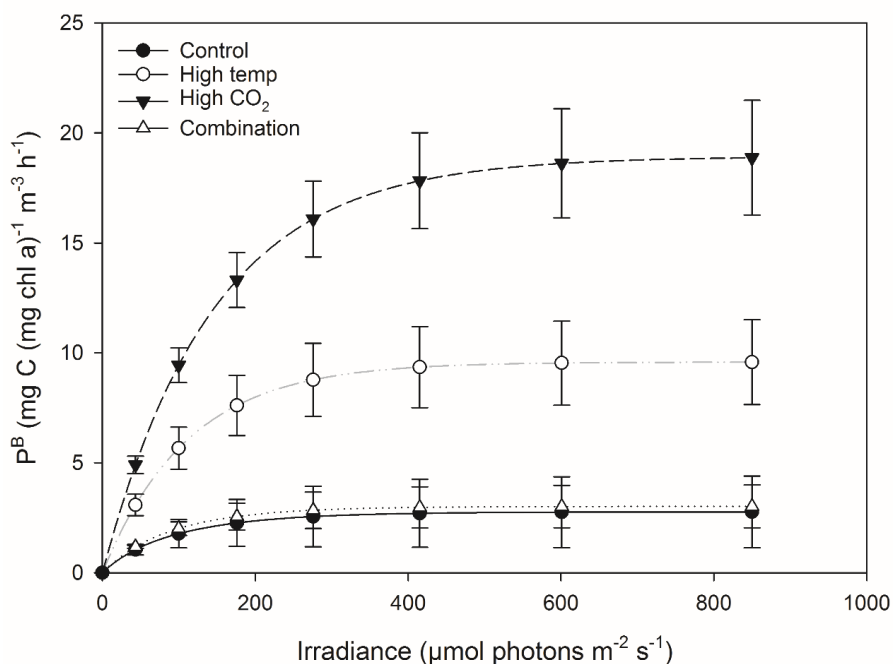


Fig. 6. Fitted parameters of FRRf-based photosynthesis-irradiance curves for the experimental treatments on the final experimental day (T36)

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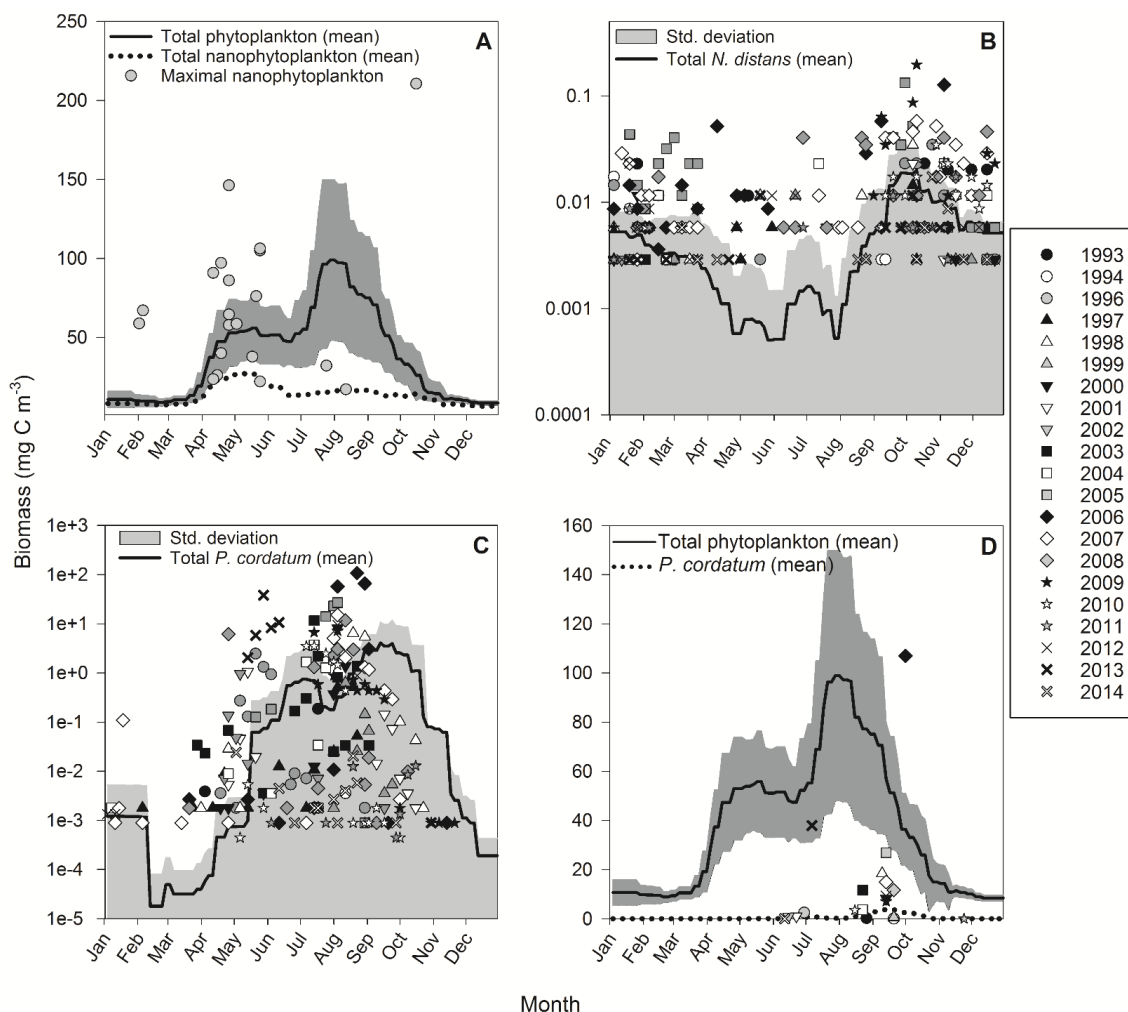


Fig. 7. (A) Temporal weekly profile of total phytoplankton carbon biomass at station L4 between 1993-2014. Black line is smoothed running average of total phytoplankton, grey area is standard deviation, dotted line is smoothed running average of total nanophytoplankton biomass and grey circles are maximal nanophytoplankton biomass from weekly observations (mean maxima of 70 mg C m^{-3}). (B) Seasonal profiles of *Navicula distans* (common log scale) between 1993-2014. Black line is smoothed running average over the time series, grey area is standard deviation and all symbols are observed data values by year. (C) Seasonal profiles of *Prorocentrum cordatum* (common log scale) between 1993-2014. Black line is smoothed running average over the time series, grey area is standard deviation and all symbols are observed data values by year. (D) Maximal *P. cordatum* biomass values relative to total phytoplankton biomass. Black line is smoothed running average of total phytoplankton biomass, grey area is standard deviation, dotted line is mean *P. cordatum* biomass and symbols are maximal *P. cordatum* biomass from weekly observations by year (as per figure legend for B & C.).

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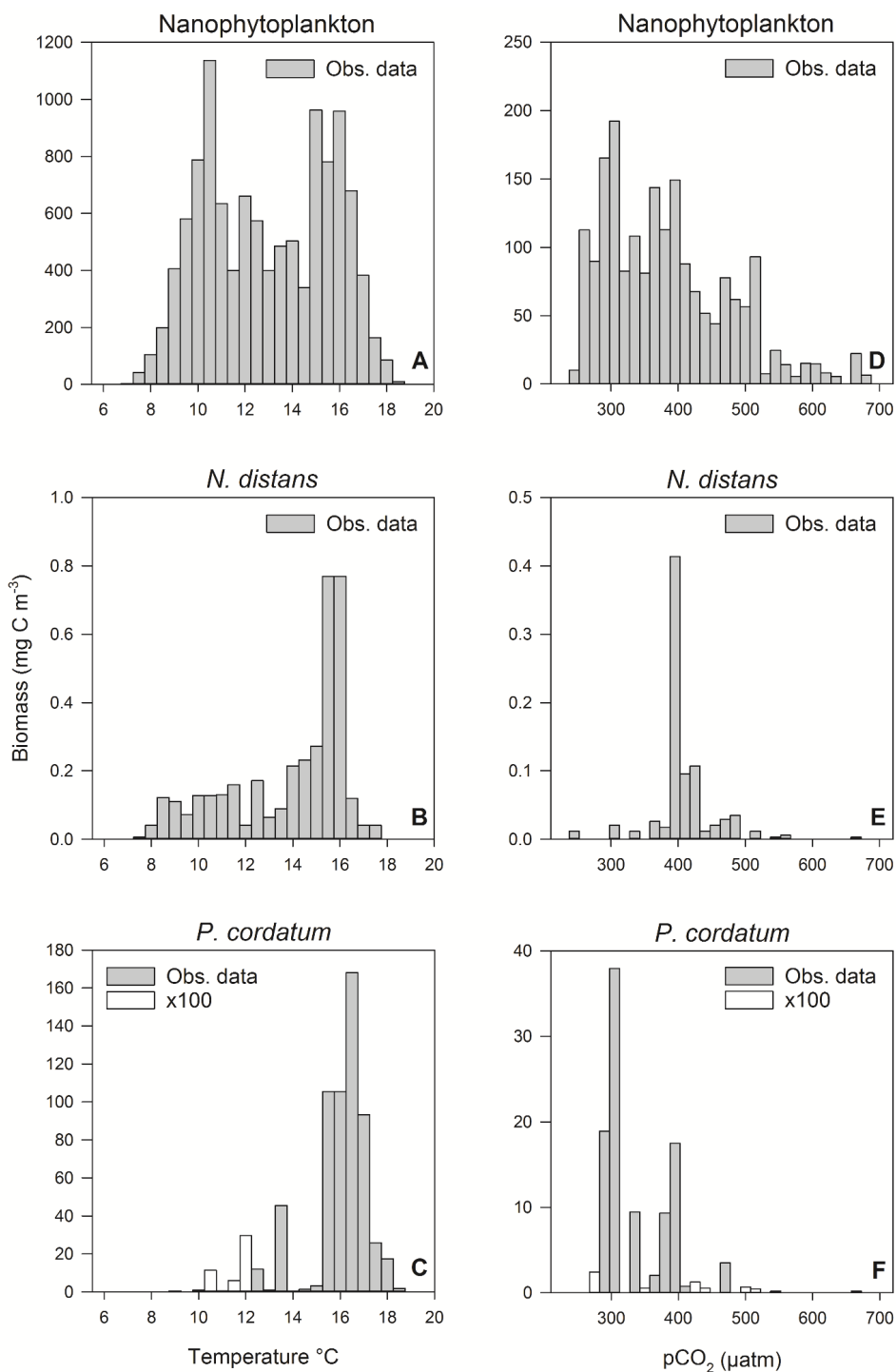


Fig. 8. Frequency distribution of biomass at station L4 along the *in-situ* gradients of temperature (1993-2014) and pCO₂ (2008-2014) for nanophytoplankton (A & D) *N. distans* (B & E) and *P. cordatum* (C & F).



.045 **Table 1.** Results of generalised least-squares model testing for main effects of time, high temperature, high CO₂
 .046 and all interactions on chl *a*, phytoplankton biomass and particulate organic nitrogen. Significant results are in
 .047 bold; * $p < 0.05$, ** $p < 0.001$, *** $p < 0.0001$.

Response variable	n	df	t-value	p	sig
Chla (mg m⁻³)					
Time	80	72	3.782211	0.0003	**
High temp	80	72	0.688339	0.4935	
High CO ₂	80	72	0.765811	0.4463	
Time x high temp	80	72	0.330431	0.742	
Time x high CO ₂	80	72	-0.596962	0.5524	
High temp x high CO ₂	80	72	0.338096	0.7363	
Time x high temp x high CO ₂	80	72	1.302498	0.1969	
Estimated biomass (mg C m⁻³)					
Time	80	72	3.339498	0.0013	*
High temp	80	72	-0.144359	0.8856	
High CO ₂	80	72	-1.008942	0.3164	
Time x high temp	80	72	3.189888	0.0021	*
Time x high CO ₂	80	72	4.751901	0.0000	***
High temp x high CO ₂	80	72	0.341905	0.7334	
Time x high temp x high CO ₂	80	72	0.449075	0.6547	
POC (mg m⁻³)					
Time	48	40	-0.27037	0.7883	
High temp	48	40	-1.2607	0.2147	
High CO ₂	48	40	-1.13796	0.2619	
Time x high temp	48	40	5.31006	0.0000	***
Time x high CO ₂	48	40	6.24182	0.0000	***
High temp x high CO ₂	48	40	-0.38194	0.7045	
Time x high temp x high CO ₂	48	40	1.21692	0.2308	
PON (mg m⁻³)					
Time	48	40	0.276438	0.7836	
High temp	48	40	-1.447791	0.1555	
High CO ₂	48	40	-1.571726	0.1239	
Time x high temp	48	40	4.78625	0.0000	***
Time x high CO ₂	48	40	5.493647	0.0000	***
High temp x high CO ₂	48	40	0.95334	0.3461	
Time x high temp x high CO ₂	48	40	-0.126291	0.9001	
POC:PON (mg m⁻³)					
Time	48	40	-3.248155	0.0024	*
High temp	48	40	-0.206777	0.8372	
High CO ₂	48	40	-0.055976	0.9556	
Time x high temp	48	40	2.433457	0.0195	*
Time x high CO ₂	48	40	3.838128	0.0004	***
High temp x high CO ₂	48	40	-2.932253	0.0055	*
Time x high temp x high CO ₂	48	40	2.40294	0.021	*



.049 **Table 2.** Results of generalised linear model testing for significant effects of temperature, CO₂ and temperature
 .050 x CO₂ on chl *a* and phytoplankton biomass at the experiment end (T36). Significant results are in bold;
 .051 * $p < 0.05$, ** $p < 0.001$, *** $p < 0.0001$.

Response variable	n	df	z-value	p	sig
<u>Chl <i>a</i> mg m⁻³</u>					
High temp	16	12	7.413	< 0.0001	***
High CO ₂	16	12	0.804	0.437	
High temp x high CO ₂	16	12	18.043	< 0.0001	***
<u>Total biomass (mg C m⁻³)</u>					
High temp	16	12	28.953	< 0.0001	***
High CO ₂	16	12	36.042	< 0.0001	***
High temp x high CO ₂	16	12	5.899	< 0.0001	***
<u>Diatoms (mg C m⁻³)</u>					
High temp	16	12	-4.43	< 0.0001	***
High CO ₂	16	12	3.036	0.0024	**
High temp x high CO ₂	16	12	-7.243	< 0.0001	***
<u>Dinoflagellates (mg C m⁻³)</u>					
High temp	16	12	9.848	< 0.0001	***
High CO ₂	16	12	1.805	0.2927	
High temp x high CO ₂	16	12	11.902	< 0.0001	***
<u>Nanophytoplankton (mg m⁻³)</u>					
High temp	16	12	32.9	< 0.0001	***
High CO ₂	16	12	39.04	< 0.0001	***
High temp x high CO ₂	16	12	5.22	< 0.0001	***
<u>Synechococcus (mg m⁻³)</u>					
High temp	16	12	7.045	< 0.0001	***
High CO ₂	16	12	-0.091	0.928	
High temp x high CO ₂	16	12	10.739	< 0.0001	***
<u>Picophytoplankton (mg m⁻³)</u>					
High temp	16	12	0.413	0.679486	
High CO ₂	16	12	2.02	0.043435	*
High temp x high CO ₂	16	12	3.773	< 0.0001	***
<u>Coccolithophores (mg C m⁻³)</u>					
High temp	16	12	0.276	0.782	
High CO ₂	16	12	-0.368	0.713	
High temp x high CO ₂	16	12	-1.265	0.206	
<u>Cryptophytes (mg C m⁻³)</u>					
High temp	16	12	0.404	0.686	
High CO ₂	16	12	0.273	0.785	
High temp x high CO ₂	16	12	1.341	0.18	

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.056 **Table 3.** FRRf-based photosynthesis-irradiance curve parameters for the experimental treatments on the final
 .057 day (T36).

Parameter	Control	sd	High temp	sd	High CO ₂	sd	Combination	sd
P^B_m	2.77	1.63	9.58	1.94	18.93	2.65	3.02	0.97
α	0.03	0.01	0.09	0.01	0.13	0.01	0.04	0.00
I_k	85.33	45.47	110.93	6.09	144.13	17.91	86.38	33.06

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.060 **Table 4.** Results of generalised linear model testing for significant effects of temperature, CO₂ and temperature
 .061 x CO₂ on phytoplankton photophysiology; P^B_m (maximum photosynthetic rates), α (light limited slope) and I_k
 .062 (light saturated photosynthesis). Significant results are in bold; * $p < 0.05$, ** $p < 0.001$, *** $p < 0.0001$.

Response variable	n	df	t-value	p	sig
P^B_m					
High temp	12	8	7.353	< 0.0001	***
High pCO ₂	12	8	8.735	< 0.0001	***
High temp x high pCO ₂	12	8	-8.519	< 0.0001	***
α					
High temp	12	8	13.03	< 0.0001	***
High pCO ₂	12	8	15.15	< 0.0001	***
High temp x high pCO ₂	12	8	-14.82	< 0.0001	***
I_k					
High temp	12	8	2.018	0.0783	
High pCO ₂	12	8	2.541	0.0347	*
High temp x high pCO ₂	12	8	-2.441	0.0405	*

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