

Response to Editorial Comments

Editor comment: Following the reviewers' suggestions and based to my own reading of your MS and responses to the comments, I believe your paper can be published in Biogeosciences after some significant revisions. The two referees have raised important questions that must carefully addressed in your revised MS. This includes a better description of methodology and the improvement of Nitrogen mass-balance calculations by the different biogeochemical processes you have measured. I also find necessary the addition of "a brief section to the manuscript on the metagenomic analysis of the phytoplankton community" in order to sustain in your MS the discussion on the relationship between N fixation and HABs. I was in general satisfied with most of your answers to the referees comments, and I recommend you incorporate all the necessary information in your revised MS. The two referees also expressed very consistent criticisms on the way you use of two expressions in your MS: the "Nitrogen pendulum" and the "freshwater estuary". I found these comments very relevant and ask you consider them more seriously.

I think the term "pendulum" must be explained more clearly, as (if I well understood) this is the first time the "Nitrogen pendulum" is described. A pendulum oscillates, swings, up and down and right and left... please be more precise explaining the pertinence of the metaphor and what processes make the pendulum move up and down. The use of this expression in the title of your MS is very abrupt and I don't think BG readers will understand immediately what you are referring to in the title. I suggest you remove the expression from the title and include it in the abstract together with a short definition of the concept.

I have much more concern with the use of the term "estuary" or "freshwater estuary", even if the NOAA has defined this last term, the most commonly used definition of an estuary remains the one Cameron and Pritchard (1963): "a semi-enclosed coastal body of water, which has free connection with the open sea, and within which seawater is measurably diluted with freshwater derived from land drainage". More recently, many studies have extended the definition of the estuary to the "tidal river" (and sometimes "freshwater estuarine zone"), as the upper estuarine region without salt intrusion, but influenced by the tide. In Sandusky Bay there is no salt intrusion and there is no tide. So this ecosystem is very far away from the definition of an "estuary". I can understand the need for a specific term to describe the mixing zone between a river and a lake as this is an ecosystem very specific properties. "river-lake mixing zone" is probably the most precise term. I can eventually accept "freshwater estuary" in your MS after a careful definition (this definition is in contradiction with that of Cameron and Pritchard), although I feel this expression must be used sparingly and carefully. In all cases the use of "estuary" alone is inappropriate and your results cannot be extrapolated to "estuarine systems" where the salinity gradient is always a major driving force. For instance in estuaries, salinity stress strongly impacts microbial communities; salinity also generates desorption processes that control the availability of phosphate for phytoplankton. As a consequence, P limitation processes as discussed in your MS will not follow classical trends occurring in estuaries. As an example in your abstract the statement "Estuarine systems such as Sandusky Bay are

mediators of...” is inappropriate because Sandusky Bay is NOT an estuary. In order to avoid confusion, your MS must be consistent with these definitions; Readers must know immediately that your study concerns FRESHWATERS and not ESTUARINE WATERS.

I will be pleased to read a revised version of your MS, together with a short response to my editorial comments.

**With best regards,
Gwenaël Abril, Biogeosciences associate Editor.**

Author response: Thank you for your instructive editorial comments on our manuscript. We appreciate your feedback as well as the reviewers' feedback, both of which have improved the revised manuscript greatly. We have responded to specific reviewer comments below. Regarding the specific points brought up in your editorial comments:

- Methodological revisions are detailed in our response to review #2. This includes clarifications of the calculations involved with the isotope pairing technique, changes to the calculations associated with the N budget, and addition of text to explain rationale for various approaches.
- N budget revisions are made in the methods, results, and discussion. Table 1 has been removed and replaced with a new figure (figure 8). The changes are aimed at improving the ease of comparison among various N sources and sinks, which are now illustrated as a function of discharge patterns. See details on these revisions below in our responses to reviewers #1 and #2.
- A section on metagenomics has been added in section 2.3, the results section, throughout the discussion, and in figure 6. The addition of these data has allowed us to demonstrate (1) that *Planktothrix* dominates the cyanobacterial community throughout the sampling season and (2) N fixers are present as a component of the cyanobacterial assemblage. See details on these revisions below in our responses to reviewers #1 and #2.
- Use of “N pendulum” has been removed from the manuscript. We opted to remove this term completely rather than more explicitly define the term, as multiple internal and external reviewers have expressed confusion over the term. *Importantly, we have changed the title in the manuscript, but we are unsure about how to change the title in the manuscript submission system.*
- Use of “estuary” has been removed in favor of “river-lake mixing zone” and “drowned river mouth.” Usage of “estuary” to describe true estuarine systems has been retained. See details on these revisions below in our responses to reviewers #1 and #2.

We look forward to your evaluation. We much appreciate your time and efforts.

Thank you,
Kateri Salk

Response to Review #1

Reviewer comment: issues in associating chlorophyll concentrations with HABs

Author response: Previous studies have demonstrated a strong correlation between chlorophyll concentrations and phytoplankton biomass in this system, and *Planktothrix* generally dominates phytoplankton biomass (Davis et al. 2015, Conroy et al. 2017). This point is emphasized in our introduction of the study system (page 3, lines 5-7) and in our discussion of the use of chlorophyll concentrations as a proxy for phytoplankton biomass in this system (page 13, lines 5-6). We have added a section to the manuscript on metagenomic analysis of the cyanobacterial community, which demonstrates the dominance of *Planktothrix* as well as the presence of a community of N fixers (page 5, lines 1-16; page 10, lines 20-22; page 14, lines 1-3).

Reviewer comment: use of the term “Great Lakes estuary” or “freshwater estuary” may not be appropriate

Author response: Usage of the term “estuary” when describing systems in the Great Lakes have been changed to “drowned river mouth” (page 1, line 22; page 3, line 4, page 15, line 4), “river-lake mixing zone” (page 3, line 17; page 14, line 29), or “coastal wetland” (page 13, line 2). Instances of the term “estuary” to describe systems fitting the traditional definition of an estuary have been retained (page 11, line 31; page 12, line 13; page 14, line 29, page 15, line 33). When direct comparisons are made between Sandusky Bay and estuarine systems, these comparisons are made with respect to common characteristics including swings in discharge and N loading from the river (e.g., page 15, lines 21-24).

Reviewer comment: p. 11 – denitrification relies on carbon supply as well as nitrate, potential dependence on C:nitrate ratios

Author response: Given the high rates of primary productivity in this system and the high carbon content of the sediments (e.g., Ostrom et al. 2005), the consideration of C limitation on denitrification was not a major concern for this study. Additionally, a recent study showed that sediment C:N ratios were not significant predictors of denitrification rates across the Great Lakes and in Lake Erie (Small et al. 2016). Considerations of C supply are mentioned in the methods along with the assumptions of the IPT method (page 6, lines 25-26).

Ostrom et al. 2005. Evaluation of primary production in Lake Erie by multiple proxies. *Oecologia* 144: 115-124. doi: 10.1007/s00442-005-0032-5

Small et al. 2016. Large differences in potential denitrification and microbial communities across the Laurentian great lakes. *Biogeochemistry* 128: 353. doi: 10.1007/s10533-016-0212-x

Reviewer comment: p. 12 – obvious comments in lines 6-12 on N₂O production as byproduct of denitrification, but logic linking this to last sentence in paragraph 2 is not justified

Author response: The paragraph on N₂O emissions in section 4.2 has been removed. The relevant information from this paragraph has been added to the first paragraph of section 4.2 (page 12, lines 1-2, 5), where we emphasize that denitrification is efficient (very little N₂O released compared to N₂), and that nearshore areas of the Great Lakes have the potential to be influential zones of N removal. This has cleared up the mismatch in logic the reviewer points out.

Reviewer comment: Section 4.3 – Impact of remineralization on N budget needs to be discussed further, particularly the distinction between the gross and net rate of phytoplankton uptake (is uptake a sink?)

Author response: The paragraph on remineralization in the discussion has been edited to include not only the implications of remineralization on rate measurements of NH_4^+ uptake (as was present before) but also the importance of remineralization in eutrophic systems that experience HABs and the significance of a rapidly cycling NH_4^+ pool despite challenges to quantify transient availability and uptake (page 13, lines 21-32). The concept of recycling as a mechanism for generating a supply of N under N limitation is also echoed later in section 4.3 (page 14, lines 16-19).

Reviewer comment: Section 4.3 – These results encourage a thorough investigation of the controls of nitrogen fixation, including a supply of P and Fe

Author response: The reviewer points out an interesting and relevant consideration, which the authors considered as well in the preparation of this manuscript. While Fe limitation of N fixation has precedent in the literature in several systems, we dismissed this mechanism for two reasons: (1) as the Sandusky River discharges into the bay, it brings a high sediment load that is expected to have a sufficient trace metal load to support phytoplankton demands, and (2) the sediments in Sandusky Bay are anoxic, and it is likely that nighttime sediment oxygen demand is high enough that transient anoxia develops at the sediment-water interface, thus enabling P and reduced Fe to be released into the water column and be taken up by phytoplankton. These internal loading events, while outside the scope of this study, have been captured on several occasions over the course of multi-year surveys of the bay. While P limitation was certainly at play during portions of the season in both years, the substantial phytoplankton biomass and high N fixation rates indicated that P was likely not a driving factor for the activity of N fixation. We have added a section to the manuscript on metagenomic analysis of the cyanobacterial community, which demonstrates the presence of a community of N fixers, which lends additional support to the biogeochemical measurements of N fixation (page 5, lines 1-16; page 10, lines 20-22; page 14, lines 1-3).

Reviewer comment: Section 4.4, line 29 – not sure it is straightforward to extrapolate this argument to other coastal areas (likely that water flow is not operating as a pendulum in other coastal systems, where tides and other factors might be stronger controls)

Author response: The final section of the manuscript has been revised with a greater emphasis on Sandusky Bay and the implications of future changes on the role of this system in the Great Lakes system. Extensions to other coastal areas (i.e., Gulf of Mexico) have been retained, as the specific systems referenced have been well-documented as systems that are highly influenced by riverine N loading. We do not suggest that this is the only factor to consider, simply that many systems are highly influenced by hydrology and nutrient loading, so this study could lend insight into other systems (page 15-16, lines 26-4).

Response to Review #2

Reviewer comment: Abstract I suggest referring to dissimilatory nitrate reductive processes as opposed to assimilation

Author response: Assimilatory (phytoplankton uptake) and dissimilatory (microbial N reduction) processes are considered separately, and this language has been revised to more clearly reflect this distinction (page 1, line 15).

Reviewer comment: Use of ‘pendulums’ not really the right term in my opinion. As elaborated on below, I think the key point is that the system is a modulator of nutrient inputs.

Author response: The use of the term “pendulum” has been removed from the manuscript. In cases where swings in hydrology and N dynamics are pointed out, we have chosen to use terms such as “oscillation” and “modulate” as descriptors for how the system functions.

Reviewer comment: Pg 3 line 5 – I don’t think estuary is generally accepted as a term for rivers entering freshwater lakes – I suggest mixing zone.

Author response: Usage of the term “estuary” when describing systems in the Great Lakes have been changed to “drowned river mouth” (page 1, line 22; page 3, line 4, page 15, line 4), “river-lake mixing zone” (page 3, line 17; page 14, line 29), or “coastal wetland” (page 13, line 2). Instances of the term “estuary” to describe systems fitting the traditional definition of an estuary have been retained (page 11, line 31; page 12, line 13; page 14, line 29, page 15, line 33). When direct comparisons are made between Sandusky Bay and estuarine systems, these comparisons are made with respect to common characteristics including swings in discharge and N loading from the river (e.g., page 15, lines 21-24).

Reviewer comment: Methods Isotope analysis delta¹⁵N values are mentioned in the methods, why? The isotope pairing equations use excess ratios of M/Z 29/28 and 30/28 for N₂ and 45/44 and 46/46 for N₂. I suggest deleting all ref to del ¹⁵N and explaining which masses were monitored and how excess ratios were calculated. It also not clear why N₂ was also measured with MIMS or how these data were used.

Author response: The equation for delta (equation 1 in previous version) was removed (as the reviewer points out, we are working not with deltas but with isotope fractions). A description of how P₂₉, P₃₀, P₄₅, and P₄₆ were calculated was added, detailing that the change in isotope fraction of the relevant mass was measured by IRMS and the calculation was corrected for the change in concentration of N₂ or N₂O as measured by MIMS or GC-ECD, respectively (page 6, lines 31-33; page 7, lines 2-4).

Reviewer comment: Phytoplankton N uptake ¹⁵NH₄/NO₃ contamination of ¹⁵N₂. You state that uptake of contamination would have made up less than 5% of measured rates. This depends on the rates. Is this even the case for the lowest measured rates? The main thing that convinced me your data were probably ok, was the fact you could measure low rates in 2016.

Author response: The 5% value for potential contamination is the maximum potential contamination across all measurements, from low to high rates. We calculated the possible overestimation of N fixation given the level of potential contamination reported in Dabundo et

al. (2014), and 5 % was the highest calculated value across the dataset. This statement was revised to clarify this point (page 7, lines 20-22).

Reviewer comment: Budget I don't think converting sediment process rates to volumetric rates is meaningful – these should either be shown as areal rates or total mass for the whole system.

Author response: Sediment process rates have been changed back to areal rates in figure 7, consistent with the units presented in figure 4. Water column rates have been retained as volumetric rates in figures 5 and 7. For calculations of budget, all rates have been converted to kg/d (from either areal or volumetric rates), enabling direct comparisons. The methodology for these conversions is detailed in section 2.6 (page 8, lines 16-19).

Reviewer comment: Line 5 pg 8 – I agree with your point about TKN in the river, but what about in Sandusky Bay? As mentioned below, I think the system is really a modulator that converts NO_x to organic matter and this will be shown clearly in the TKN data if available.

Author response:

TKN concentrations in the Sandusky River and Sandusky Bay were added to the paper in the methods (page 3, line 31-32; page 4, line 8, 22-23; page 8, lines 19-20), results (page 9, line 13-14; page 10, lines 32-33), and figures 3 and 8. This includes an examination of concentrations as well as how TKN loads compare to other aspects of the N budget. The observation of higher TKN concentrations in the bay vs. the river suggests a conversion of DIN into TKN (recycling to organic forms or ammonium), which the reviewer points out is an indication that the system converts nitrate into organic matter and ammonium. We point out the indication of recycling in section 4.3 (page 13, lines 24-26; page 14, lines 17-19).

Reviewer comment: For the nitrate loads, at what time interval were concentration and flow measured? How were these data interpolated to calculate loads?

Author response: Daily nitrate loads were calculated based on daily discharge and nitrate concentration data from the Heidelberg NCWQR. Direct comparisons were made for days when assimilatory and dissimilatory N transformations were measured. These calculations are detailed in section 2.6 (page 8, lines 19-20) and in figure 8.

Reviewer comment: Discussion N removal processes. This was generally good and I agree with the arguments. I felt, however this section could have been a little more quantitative. For example, it is argued that the increase in the N:P was driven in a large part by denitrification. I suggest the authors undertake a back of the envelope calculation to show how the change in the mass of NO₃ in the water column over this period of NO₃-drawdown compares with total denitrification measured over the same period.

Author response: The statement about short hydraulic residence times reducing the capacity for denitrification to remove significant amounts of N has been amended to be more quantitative (page 12, lines 24-26). Given the covarying rates of varying processes over time (N loading, phytoplankton uptake, denitrification) as well as the complex hydrology, we have opted not to make a calculation of the comparison between changing water column NO₃⁻ concentrations and denitrification rates. We feel it is misleading to make a back-of-the-envelope calculation here, as the assumptions we would need to make about hydraulic residence time, N recycling, and

temporal variation in rates would be an oversimplification of this system. Instead, we have focused our efforts on illustrating the N budget as a function of river discharge in figure 8.

Reviewer comment: Discussion N removal processes. Also of relevance here is that Dw (water column driven denitrification) and Dn (water column driven denitrification) are not reported. The breakdown of these is important when considering the drawdown rate of NO_x.

Author response: Given the shallow and well-mixed nature of this system (mean depth 1.6-2.6 m), water column denitrification was not considered as an appreciable denitrification source. This assumption was borne out in the IPT calculations, which enable distinction between Dw and Dn and confirmed that Dw was not active. These detailed results were not included in the manuscript, but a statement describing the assumption of negligible water column denitrification was added (page 4, lines 15-16).

Reviewer comment: Budget I think there was a missed opportunity with the budget to integrate the findings a little more clearly. I suggest that for each period process rates were measured, a budget be undertaken (could be daily or perhaps monthly basis). These budget terms could then graphed to highlight the change from high catchment inputs to high internal inputs via N fixation as flows decreased through to August. This would also highlight the relatively minor importance of denitrification as a sink compared to the inputs. Although the phytoplankton assimilation measurements are a nice part of the paper, I don't think they can be used meaningfully in the budget because they were taken in 2016 when phytoplankton biomass was higher.

Author response: The reviewer brings up a good point that the budget calculations could be more effectively integrated to illustrate the system more clearly. Table 1 has been removed and instead we have added figure 8, which shows all the measured aspects of the N budget as (1) total ranges, (2) individual rate measurements, and (3) as a function of discharge. This figure enables direct comparisons between processes and examinations of how river discharge impacts the system. A description of the budget is detailed in the results (page 10-11, lines 32-4) and throughout the discussion.

Reviewer comment: I think the discussion at line 20 on pg 15 could also talk a little more about the system as a transformer of nitrogen importing DIN and exporting algal biomass as well as N derived from nitrogen fixation. At the moment it is a bit repetitive and not as interesting as it could be. I don't really think the term N pendulum is correct, it really modulates the inputs depending on residence time, with a net export of nitrogen from nitrogen fixation. This finding is consistent with a previous study of a shallow eutrophic lake which often showed net exports of total nitrogen, most likely due to nitrogen fixation. Cook, P.L.M., K.T. Aldridge, S. Lamontagne, and J.D. Brookes. (2010). Retention of nitrogen, phosphorus and silicon in a large semi-arid riverine lake system. Biogeochemistry, 99: 49-63.

Author response: The second paragraph in section 4.4 was revised to include a more detailed discussion of the capacity of the system to transform inorganic forms of N to organic forms, thus modulating the magnitude and composition of N loading downstream to Lake Erie (page 15, lines 16-24). Leading up to this paragraph, supporting points were revised and added (page 13, lines 21-32; page 14, lines 16-19; page 15, lines 9-10). We have cited the suggested study in

section 4.4 as well (page 15, lines 8-9). We have removed the use of the term “pendulum” in favor of terms such as “modulator” and “oscillations.”

Reviewer comment: The last paragraph of the discussion is quite speculative, I suggest remove.

Author response: The final paragraph is meant to put our results in context within the larger Great Lakes system and the potential shifts in hydrology and nutrient regime associated with climate change. We have kept this paragraph in the manuscript, but we have removed two sentences and revised another sentence to limit speculation (page 15-16, lines 26-4).

Reviewer comment: Figure 2, micro symbol now appears as milli.

Author response: There seems to have been a conversion error that made the micro symbol appear as “m.” Thank you for pointing this out, and it is fixed in the revised manuscript.

Reviewer comment: Figure 3 micro symbol as above

Author response: There seems to have been a conversion error that made the micro symbol appear as “m.” Thank you for pointing this out, and it is fixed in the revised manuscript.

Reviewer comment: Figure 4 the letters showing statistically significant groupings are unclear. ^ is carat, not carrot

Author response: The groupings in Figure 5 (formerly figure 4) are the result of an unusual outcome of the Tukey’s post-hoc test of the two-way interaction effects ANOVA, which are valid yet confusing. As this statistical result is not a crucial outcome of the manuscript, the authors have removed the letters from the figure. The misspelling of the ^ symbol is revised.

Reviewer comment: Figure 5a. Why are these rates reported volumetrically? They should be areal as for Fig 3.

Author response: Figure 7a (formerly figure 5a) is now presented in areal units, consistent with rates presented in figure 4.

Nitrogen cycling in Sandusky Bay, Lake Erie: Oscillations between strong and weak export and implications for harmful algal blooms

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Abstract. Recent global water quality crises point to an urgent need for greater understanding of cyanobacterial harmful algal blooms (cHABs) and their drivers. Nearshore areas of Lake Erie such as Sandusky Bay may become seasonally limited by nitrogen (N) and are characterized by distinct cHAB compositions (i.e., *Planktothrix* over *Microcystis*). This study investigated phytoplankton N uptake pathways, determined drivers of N depletion, and characterized the N budget in Sandusky Bay. Nitrate (NO₃⁻) and ammonium (NH₄⁺) uptake, N fixation, and N removal processes were quantified by stable isotopic approaches. Dissimilatory N reduction was a relatively modest N sink, with denitrification, anammox, and N₂O production accounting for 84, 14, and 2 % of sediment N removal, respectively. Phytoplankton assimilation was the dominant N uptake mechanism, and NO₃⁻ uptake rates were higher than NH₄⁺ uptake rates. Riverine N loading was sometimes insufficient to meet assimilatory and dissimilatory demands, but N fixation alleviated this deficit. N fixation made up 23.7-85.4 % of total phytoplankton N acquisition and indirectly supports *Planktothrix* blooms. However, N fixation rates were surprisingly uncorrelated with NO₃⁻ or NH₄⁺ concentrations. Owing to temporal separation in sources and sinks of N to Lake Erie, Sandusky Bay oscillates between a conduit and a filter of downstream N loading to Lake Erie, delivering extensively recycled forms of N during periods of low export. Drowned river mouths such as Sandusky Bay are mediators of downstream N loading, but climate change-induced increases in precipitation and N loading will likely intensify N export from these systems.

1 Introduction

Harmful algal blooms (HABs) are increasing in frequency on a global scale and are stimulated by excessive nutrient loading to aquatic systems (Bricker et al., 2008; Heisler et al., 2008). Lake Erie, in particular, has been subject to increased incidence and expansion of cyanobacterial HABs (cHABs) in recent years (Michalak et al., 2013; Ho and Michalak, 2015; Bullerjahn et al., 2016). These blooms are dominated by cyanobacteria that accumulate the powerful hepatotoxin, microcystin (Carmichael and Boyer, 2016). A more nuanced understanding of the drivers of cHABs, including nutrient cycling, will allow for better prediction and management of blooms in Lake Erie and other ecosystems.

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Phytoplankton biomass and cHABs in Lake Erie have historically been correlated with P loading from river inflows (Kane et al., 2014; Kim et al., 2014). Calls to control eutrophication in Lake Erie have proposed targets for reduced P loading but have largely ignored N (Scavia et al., 2014). However, there is a growing dialogue surrounding the dual management of N and P in lacustrine systems (Gobler et al., 2016; Paerl et al., 2016), particularly as co-limitation of phytoplankton growth by both N and P has been demonstrated in the late summer in Lake Erie (Moon and Carrick, 2007; North et al., 2007; Chaffin et al., 2013; Steffen et al., 2014a). This seasonal N deficiency is consistent with reduced watershed loading of N into the lake's western basin over the past two decades (Stow et al., 2015) combined with active dissimilatory sinks for nitrate (Small et al., 2016). N concentration and speciation also influence toxin production by cHABs (Horst et al., 2014; Monchamp et al., 2014; Davis et al., 2015). Delineating the role of N in controlling cHABs, therefore, requires investigation of spatial and temporal variation in multiple species of N.

Because the Lake Erie catchment is highly agricultural, the majority of bioavailable N in the lake is derived from farming practices and delivered to the lake via river inflows (Robertson and Saad, 2011; Stow et al., 2015). Once in the lake, dissolved inorganic N (DIN) and dissolved organic N (DON) are subject to consumption by competing biological processes. Phytoplankton, including cHAB taxa, commonly take up DIN and DON in the form of ammonium (NH_4^+), nitrate (NO_3^-), and urea (Davis et al., 2015). N fixation, an alternate source of bioavailable N, is expected to occur when DIN is scarce. N fixation activity has been inferred in Lake Erie (MacGregor et al., 2001; Monchamp et al., 2014; Steffen et al., 2014a; Davis et al., 2015) but not quantified for several decades (Howard et al., 1970). Dissimilatory microbial processes may also consume DIN in sediments, representing pathways of permanent N removal whereby N leaves the system as N_2 . Denitrification is expected to be the dominant microbial N removal pathway in freshwaters (Seitzinger et al., 2006), although anammox, a competing pathway, has not been extensively studied in freshwater systems (Yoshinaga et al., 2011; Zhu et al., 2013). Nitrous oxide (N_2O), a potent greenhouse gas, is a byproduct of denitrification and nitrification (Wrage et al., 2001) and is a detrimental consequence of microbial N removal.

Whereas the colonial cyanobacterium *Microcystis* dominates the cHAB community in offshore regions of western Lake Erie, filamentous *Planktothrix* has been shown to persist in N-limited bays and tributaries (Conroy et al., 2007; Kutovaya et al., 2012; Davis et al., 2015). With few exceptions (Pancrace et al., 2017), both cHAB taxa are incapable of N fixation and require dissolved forms of N for growth. *Planktothrix* is a superior scavenger for N (Conroy et al., 2007) and responds strongly to additions of DIN (Donald et al., 2011; 2013). Under N limitation, *Planktothrix* may thus be able to outcompete other phytoplankton for small N inputs. *Planktothrix* is also particularly well-adapted to low irradiances; suspended sediment that rapidly attenuates light is often found in *Planktothrix*-dominated waters (Scheffer et al., 1997; Havens et al., 2003; Oberhaus et al., 2005). Additionally, *Planktothrix* can tolerate the widest temperature range among major bloom-forming cyanobacteria, including *Microcystis*, *Aphanizomenon*, and *Dolichospermum* (*Anabaena*) (Foy et al., 1976; Post et al., 1985). Due to these factors, the persistence of *Planktothrix* in nearshore zones likely operates under a fundamentally different paradigm than

offshore *Microcystis* blooms, and mitigation may require attention to the distinct biogeochemical functioning of these genera in the nearshore vs. offshore.

Sandusky Bay, a [drowned river mouth](#) on the southern shore of Lake Erie, serves as an ideal location in which to examine the relationship between N cycling and cHABs. This system is hypereutrophic (Ostrom et al., 2005), with the cyanobacterium *Planktothrix* dominating phytoplankton biomass from May to October and N fixing phytoplankton making up a small portion of biomass (Davis et al., 2015). Sandusky Bay experiences large fluctuations in NO_3^- concentrations and dissolved N:P ratios throughout the summer (Davis et al., 2015; Conroy et al., 2017), suggesting that the dynamic N cycling may influence cHAB formation in this system. Indeed, growth of *Planktothrix* in Sandusky Bay is stimulated by additions of NO_3^- , NH_4^+ , and urea, indicating that phytoplankton growth is seasonally limited or co-limited by N (Chaffin and Bridgeman, 2014; Davis et al., 2015). Evaluating the mechanisms that promote the persistence of *Planktothrix* in this system will benefit from an examination of N removal processes and inputs from N fixation that directly influence the availability of DIN. A thorough understanding of these processes will also inform the capacity for Sandusky Bay to [mediate nutrient delivery to Lake Erie](#).

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The objectives of this study, conducted in Sandusky Bay, Lake Erie, were to (1) investigate the pathways by which the *Planktothrix*-dominated phytoplankton community acquires N, (2) determine the factors driving N depletion, namely microbial N removal processes and hydrology, and (3) evaluate Sandusky Bay as a [river-lake mixing zone](#) in which N cycling affects cHABs and N loading to Lake Erie. These objectives provide insight into the Grand Challenges of Great Lakes research by addressing variability in ecosystem processes, anthropogenic nutrient forcing and potential for reversibility, and response of this ecosystem to climate change (Sternier et al., 2017).

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2 Methods

2.1 Field Sampling

Sampling took place between May and October in 2015 and 2016 in partnership with the Ohio Department of Natural Resources (ODNR). Water column nutrient and chlorophyll (chl) *a* samples were collected weekly in 2015 and approximately biweekly in 2016. Samples for DIN uptake assays (2016), N fixation assays (2015 and 2016), and microbial N removal assays (2015) were collected monthly. Six stations were sampled: two stations in the inner portion of Sandusky Bay (ODNR4, ODNR6), three stations in the outer portion of Sandusky Bay (ODNR2, ODNR1, and Environment [and Climate Change Canada \[ECCC\]](#) 1163, hereafter 1163), and one station directly outside Sandusky Bay in the western basin of Lake Erie (Bells; Fig. 1). Tributary discharge data from the primary water source to Sandusky Bay, the Sandusky River, were obtained from the USGS stream monitoring station near Fremont, OH (site 04198000; Fig. 1). Hydraulic residence time in Sandusky Bay was estimated by dividing the Bay volume (0.26 km^3 ; Richards and Baker, 1985) by the Sandusky River discharge rate. NO_3^- and total Kjeldahl N (TKN) concentrations in the Sandusky River near the USGS monitoring site were provided by the Heidelberg

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Tributary Loading Program maintained by the National Center for Water Quality Research (NCWQR) at Heidelberg University (Richards et al., 2010).

At each sampling location, water column physical and chemical parameters (pH, conductivity, temperature, dissolved oxygen) were measured using a YSI 600QS sonde (YSI Inc., Yellow Springs, OH, [USA](#)). Water samples were collected by Van Dorn bottle at 1 m depth for analysis of NO_3^- , NH_4^+ , phosphate (PO_4^{3-}), [TKN](#), [chl *a* concentrations](#), [and 16S rRNA metagenomics](#). Samples for dissolved nutrient analysis were filtered immediately upon collection (0.2 μm), kept on ice, and frozen upon return to the lab. [Unfiltered sample water was collected for TKN analysis and also frozen](#). Samples for determination of chl *a* concentrations were collected on 0.2 μm polycarbonate membrane filters and frozen. [Samples for metagenomics analyses were collected using Sterivex cartridge filters \(0.22 \$\mu\text{m}\$; EMD Millipore, Billerica, MA, USA\) which were immediately frozen on dry ice prior to storage at -80 °C](#). Station 1163 has an extensive monitoring history by [ECCC](#) and was chosen for additional water and sediment assays. An additional 20 L carboy was filled with surface water from station 1163 for N uptake assays and sediment incubations. Sediment cores for the evaluation of microbial N removal rates (denitrification, anammox, and N_2O production) by the isotope pairing technique (IPT) were collected at station 1163 using a modified piston corer as described by Smit and Steinman (2015). [Given the shallow and well-oxygenated nature of the bay, microbial N removal in the water column was not considered](#). Intact sediment cores were collected in polycarbonate tubes (7 cm i.d.) to a depth of 25 cm. Water (1 cm) was maintained overlying the sediment during transport to preserve redox gradients.

2.2 Nutrient and Chlorophyll *a* Analyses

Concentrations of $\text{NO}_3^- + \text{NO}_2^-$, NO_2^- , NH_4^+ , and PO_4^{3-} were measured on field-filtered sample water using standard U.S. EPA methods (353.1, 353.2, 350.1, and 365.1, respectively) on a SEAL Analytical QuAAatro continuous segmented flow analyzer (SEAL Analytical Inc., Mequon, WI, [USA](#)). NO_3^- concentration was determined as the difference between $\text{NO}_3^- + \text{NO}_2^-$ and NO_2^- . [TKN was determined on unfiltered water after a sulfuric acid-copper sulfate digestion at 180°C for 60 minutes 380°C for 120 minutes, and then quantified as \$\text{NH}_3\$ on the SEAL analyzer \(U.S. EPA method 351.2\)](#). Seven known concentration standard solutions (including 0) were used for the standard curve ($R^2 > 0.999$), and every-tenth sample was spiked with a known amount of analyte to ensure high accuracy and precision throughout the analysis ($> 95\%$ recovery). Samples with concentrations exceeding the highest standard were diluted and reanalyzed. Values were averaged over two or three replicates. Method detection limits were 0.165, 0.044, 0.558, [0.044, and 1.44 \$\mu\text{mol L}^{-1}\$ for \$\text{NO}_3^- + \text{NO}_2^-\$, \$\text{NO}_2^-\$, \$\text{NH}_4^+\$, \$\text{PO}_4^{3-}\$, and TKN](#), respectively. Extractive chl *a* concentration was measured following Welschmeyer (1994). Filters containing phytoplankton seston were extracted with 90% aqueous acetone overnight at -20 °C followed by measurement of the clarified extract by fluorometry (model TD-700, Turner Designs, Sunnyvale, CA, [USA](#)).

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2.3. 16S rRNA Metagenomic Analysis

DNA was extracted using the PowerWater Sterivex DNA Isolation Kit (MO BIO Laboratories, Inc, Carlsbad, CA, USA) following manufacturer's instructions. Short 16S rRNA Illumina amplicon tag (iTag) sequencing of the V4-V5 hypervariable region of bacterial genomes was completed at the Joint Genome Institute (JGI; Walnut Creek, CA, USA) using an Illumina MiSeq benchtop sequencer (2 × 301 bp reads) according to standard JGI procedures (Tremblay et al., 2015). Primer design for universal amplification of the V4-V5 region of 16S rDNA was based on Parada et al. (2016). Resulting sequences were demultiplexed and contaminating Illumina adaptor sequences were removed using the kmer filter in BBDuk (v37.62) following Singer et al. (2016). Briefly, BBDuk was used to remove reads containing more than 1 'N' base, or with a quality score < 10 across the read, or length < 51 bp or 33% of the full length read. Additional processing using BBDuk (<http://btools.jgi.doe.gov>) mapped reads to masked human, cat, dog and mouse references, discarding hits exceeding 93% identity (Singer et al., 2016).

Processing, clustering and classification of sequenced reads was performed as described previously (Tremblay et al., 2015). Briefly, quality controlled reads were processed by iTagger (v2.2), first by read clustering (97% identity) using algorithms in USEARCH (v9.2; Edgar, 2010), followed by the assignment of operational taxonomic units (OTUs) using the SILVA 16S SSU database (v128; Quast et al., 2013), and finally, analysis of ecological data using QIIME v1.9.1 (Caporaso et al., 2010).

2.4. Sediment Microbial N Removal

Upon return to the lab, water from station 1163 was gently added to cores to a depth of 20 cm. Cores were pre-incubated for 12 h in the dark at *in situ* temperature under gentle aeration to maintain oxic conditions in the overlying water. Following pre-incubation of sediment cores, a sample was taken from the overlying water for DIN concentration analysis (NO_3^- , NO_2^- and NH_4^+), filtered through a precombusted GF/F filter and frozen until analysis. $^{15}\text{N}\text{-NO}_3^-$ (100 $\mu\text{mol L}^{-1}$) was then added to the overlying water in each core. Cores were then capped and statically incubated under gentle stirring throughout the duration of the experiment. An initial equilibration period was employed to allow homogenization of NO_3^- between the overlying water and the NO_3^- reduction zone in the sediment porewater (Dalsgaard et al., 2000). Cores were sacrificed in triplicate or quadruplicate at intervals of 0, 3 or 4, and 6 h, during which time oxic conditions were maintained in the overlying water. Dissolved O_2 in the overlying water was monitored to evaluate the maintenance of oxic conditions throughout the incubation using a YSI 600QS sonde. A final sample for DIN analysis was collected when each core was sacrificed and processed as described above.

Samples for the determination of $\delta^{15}\text{N}_2$ were collected according to Hamilton and Ostrom (2007); briefly, dissolved gases were equilibrated with a He atmosphere, and the headspace was transferred into a pre-evacuated 12 mL Exetainer (Labco Ltd, Lampeter, Ceredigion, UK). Samples for analysis of dissolved N_2 concentrations were siphoned into 12 mL Exetainers to

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overflowing and amended with 200 μL of saturated ZnCl_2 solution to halt biological activity. All Exetainers were stored underwater at room temperature to minimize diffusion of atmospheric N_2 during storage. Samples for analysis of the $\delta^{15}\text{N}_2\text{O}$ and N_2O concentrations were siphoned into 250 and 60 mL serum bottles, respectively, to overflowing and sealed without a headspace with a butyl rubber septum. Biological activity was halted by adding saturated HgCl_2 solution to a final concentration of 0.4 % by volume.

Prior to N_2O concentration analysis, a headspace of 20 mL He was introduced in each 60 mL bottle, maintaining atmospheric pressure with a vent needle. Serum bottles were allowed to equilibrate under gentle shaking for at least 12 h prior to analysis. The headspace was then analyzed by GC-ECD (Shimadzu Greenhouse Gas Analyzer GC-2014, Shimadzu Scientific Instruments, Columbia, MD, USA) for N_2O concentration. The dissolved concentration was calculated based on the headspace equilibrium concentration (Hamilton and Ostrom, 2007).

The isotopic composition of N_2O was analyzed upon introduction of sample water into an enclosed 0.75 L glass vessel that was previously purged of atmospheric air using a gentle flow of He. Dissolved gases were subsequently stripped from the water by sparging the sample with He (Sansone et al., 1997), which carried sample gases into a Trace Gas sample introduction system interfaced to an Isoprime isotope ratio mass spectrometer (IRMS; Elementar Americas, Inc., Mount Laurel, NJ, USA).

Concentrations of dissolved N_2 were analyzed by membrane inlet mass spectrometry (MIMS; Kana et al., 1994). The isotopic composition of N_2 was analyzed by introducing the sample to an evacuated 800 μL sampling loop and then onto a packed molecular sieve (5 Å) column (Alltech, Inc., Deerfield, IL) using He carrier gas within a gas chromatograph (HP-5980, Hewlett Packard, Ramsey, MN) interfaced to an Isoprime IRMS. Analytical reproducibility of standards was 0.3 ‰.

Denitrification, anammox, and N_2O production rates were calculated by the IPT. Calculations were derived from the $\text{IPT}_{\text{anaN}_2\text{O}}$ (Hsu and Kao, 2013), which builds on the R-IPT (Risgaard-Petersen et al., 2003) by enabling quantification of N_2O production simultaneously with denitrification and anammox. This approach relies on the assumption that N removal is limited by N and not by other factors such as C, an assumption that has been supported for Sandusky Bay by Small et al. (2016). Briefly, N_2 production by denitrification ($D_{14-\text{N}_2}$) and anammox (A_{14}) were calculated as:

$$D_{14-\text{N}_2} = (r_{14-\text{N}_2\text{O}} + 1) \times 2r_{14-\text{N}_2\text{O}} \times P_{30} \quad (1)$$

and

$$A_{14} = 2r_{14-\text{N}_2\text{O}} \times (P_{29} - 2r_{14-\text{N}_2\text{O}} \times P_{30}) \quad (2)$$

where $r_{14-\text{N}_2\text{O}}$ is the ratio of ^{14}N to ^{15}N in N_2O , P_{29} and P_{30} the production of $^{29}\text{N}_2$ and $^{30}\text{N}_2$, respectively, were calculated by the increase in isotope fraction of mass 29 or 30 as measured by IRMS and corrected for the change in total N_2 concentration as measured by MIMS. N_2O production ($D_{14-\text{N}_2\text{O}}$) was calculated as:

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$$D_{14-N_2O} = r_{14-N_2O} \times (2P_{46} + P_{45}) \quad (3)$$

where P_{45} and P_{46} are the production of $^{45}\text{N}_2\text{O}$ and $^{46}\text{N}_2\text{O}$, respectively. P_{45} and P_{46} were calculated by the increase in isotope fraction of mass 45 or 46 as measured by IRMS and corrected for the change in total N_2O concentration as measured by GC-ECD.

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5 2.5 Phytoplankton N Uptake

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DIN uptake assays were conducted by adding $^{15}\text{NO}_3^-$ or $^{15}\text{NH}_4^+$ to a serum bottle containing site water to a target concentration of 10 % of ambient NO_3^- or NH_4^+ concentration, respectively. Each assay for a single site was run in triplicate. Bottles were placed in an incubator at *in situ* light intensity, light:dark cycle, and temperature conditions for 24 h.

- 10 N fixation assays were conducted by the dissolution method, which involves the addition of $^{15}\text{N}_2$ -equilibrated water to a water sample rather than a $^{15}\text{N}_2$ bubble (Großkopf et al., 2012). Preparation of $^{15}\text{N}_2$ -equilibrated water involved sparging water in a serum bottle equipped with a butyl rubber septum with He to remove ambient N_2 , followed by injection of $^{15}\text{N}_2$ (98 % atom fraction, Sigma-Aldrich Lot #MBBB0968V) while maintaining atmospheric pressure with a vent needle. Water from station 1163 was transferred into 1.18 L serum bottles and amended with $^{15}\text{N}_2$ -equilibrated water to a final dissolved atom fraction of
- 15 1.14–2.33 %. Each assay for a single site was run in triplicate. Bottles were incubated at *in situ* light and temperature conditions for 24 h.

To evaluate the potential contamination of $^{15}\text{N}_2$ gas with $^{15}\text{NO}_3^-$ or $^{15}\text{NH}_4^+$ (Dabundo et al., 2014), a mass scan of the isotopically enriched gas was performed by IRMS. The mass scan revealed that potential impurities made up < 1 % of the enriched gas. A hypothetical calculation of the incorporation of $^{15}\text{NO}_3^-$ and $^{15}\text{NH}_4^+$, given maximum contamination of the N_2 gas according to Dabundo et al. (2014) and maximum assimilation by phytoplankton, revealed that contamination could have made up between 0 and 5 % of measured N fixation rates.

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- 25 When DIN uptake and N fixation incubations were complete, the samples were vacuum filtered through precombusted GF/F filters. Filters were then dried at 60° C, acidified with 10 % HCl to remove carbonates, and dried again. The concentration and isotopic composition of particulate organic matter (POM) from N fixation assays was analyzed by scraping the contents of dried and acidified filters into tin cups and introducing samples to an elemental analyzer interfaced to an Isoprime IRMS. Analytical reproducibility of standards was 0.2 %.

- 30 DIN uptake rates were calculated according to Dugdale and Wilkerson (1986). The transport rate, or N uptake rate (ρ ; $\mu\text{mol N L}^{-1} \text{ h}^{-1}$), was calculated as:

$$\rho_t = \frac{^{15}\text{N}_{\text{ss}}}{(^{15}\text{N}_{\text{enr}} - ^{15}\text{N}_{\text{na}}) \times T} \times \text{PON}_t$$

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where $^{15}\text{N}_{\text{xs}}$ is the atom percent ^{15}N excess in the POM sample, $^{15}\text{N}_{\text{ent}}$ is the isotope fraction of ^{15}N in the NO_3^- or NH_4^+ pool, and $^{15}\text{N}_{\text{na}}$ is the isotope fraction of ^{15}N in the natural abundance POM. T is the time in hours, and PON_t is the concentration of N in POM ($\mu\text{mol N L}^{-1}$) at the end of the incubation.

- 5 N fixation rates were calculated according to Montoya et al. (1996). The transport rate, or N fixation rate (ρ ; $\mu\text{mol N L}^{-1} \text{ h}^{-1}$), was calculated as:

$$\rho_t = \frac{A_{\text{PNF}} - A_{\text{PN0}}}{(A_{\text{N}_2} - A_{\text{PN0}}) \times T} \times \text{PON}_t \quad (5)$$

where A_{PN0} and A_{PNF} are the isotope fractions of ^{15}N in the POM at the start and end of the incubation, respectively, and A_{N_2} is the isotope fraction of ^{15}N in the labeled N_2 pool. Areal N fixation rates were calculated by scaling volumetric rates over the
10 depth of the photic zone in proportion to light attenuation, under the assumption that water column N fixation is light-dependent (Scott and Grantz, 2013).

2.6 N Budget

Hourly N cycling rates were converted to daily rates by multiplying by the hours of daylight (11.25–15 h) for phytoplankton N uptake and N fixation and by 24 h for sediment N removal processes (denitrification, anammox, and N_2O production). These
15 calculations assumed that phytoplankton N uptake and N fixation were light-dependent but sediment N removal was light-independent. In order to generate a preliminary budget of the magnitude of N sources and sinks in Sandusky Bay, total daily loads of each process for the entire system (kg d^{-1}) were calculated by scaling up volumetric measurements (phytoplankton uptake, N fixation) and areal measurements (denitrification, anammox, N_2O production) to the volume (0.26 km^3) and area (162 km^2) of Sandusky Bay, respectively. The total NO_3^- and TKN load to Sandusky Bay for each sampling date was estimated
20 by multiplying Sandusky River discharge by the NO_3^- and TKN concentrations in the Sandusky River. While point estimates of N cycling processes are not necessarily representative of the entire system, these calculations made it possible to compare the relative magnitudes of N sources, assimilatory uptake processes, and dissimilatory sinks in Sandusky Bay.

2.7 Statistics

Statistical modeling was carried out in R (version 3.2.4). Correlations between (1) nutrient concentrations and discharge and
25 (2) N cycling rates and DIN concentrations were analyzed by linear regression. Pairwise differences between NO_3^- and NH_4^+ uptake rates were analyzed by t -test. Differences in rates of denitrification, anammox, and N_2O production by date were analyzed by one-way ANOVA. Differences in NO_3^- uptake, NH_4^+ uptake, and N fixation by date and station were analyzed by two-way interaction effects ANOVA. In all cases, date was treated as a fixed effect, and temporal autocorrelation was avoided
owing to the spacing of sampling dates approximately one month apart.

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3 Results

Discharge from the Sandusky River was highly episodic in both years (Fig. 2k, 2l) and dominated by significant rain events. Owing to more frequent and intense precipitation in 2015, peak discharge was an order of magnitude higher than in 2016. NO_3^- and PO_4^{3-} concentrations were positively correlated with discharge (NO_3^- : $\text{df} = 135$, $R^2 = 0.26$, $p < 0.0001$; PO_4^{3-} : $\text{df} = 135$, $R^2 = 0.05$, $p < 0.01$), but NH_4^+ concentrations were not correlated with discharge ($\text{df} = 135$, $R^2 = 0.01$, $p = 0.08$). The shortest hydraulic residence time in 2015 was 8 days during peak discharge but in 2016 was 82 days owing to lower peak discharge. In both years, as discharge decreased in the late summer and early fall, hydraulic residence time increased to several months.

NO_3^- concentrations ranged from below detection ($< 0.1 \mu\text{mol L}^{-1}$) to a maximum of $> 650 \mu\text{mol L}^{-1}$ in 2015 and $65 \mu\text{mol L}^{-1}$ in 2016 (Fig. 2a, 2b). Temporal patterns in NO_3^- concentration in the bay followed those in the Sandusky River. In both years, the greatest NO_3^- concentrations were observed in June and July followed by a decline in August that continued through October. The magnitude of these shifts was greater for the inner bay stations (ODNR4, ODNR6) than the outer bay stations (ODNR1, ODNR2, 1163). NO_3^- concentrations were generally higher in the Sandusky River than in Sandusky Bay (Fig. 2a, 2b), whereas TKN concentrations displayed an opposite trend (Fig. 3). NO_3^- concentrations at the nearshore Lake Erie station (Bells) displayed similar temporal patterns to those in the bay in 2015, whereas NO_3^- concentrations at Bells were consistently higher than those in the bay in 2016. NH_4^+ concentrations ranged from below detection ($< 0.5 \mu\text{mol L}^{-1}$) to $17.5 \mu\text{mol L}^{-1}$ across sites during the sampling period and were generally lower in 2016 than 2015 (Fig. 2c, 2d). The greatest NH_4^+ concentrations were observed in the inner bay, and these spikes occurred episodically throughout the sampling season. PO_4^{3-} concentrations ranged from below detection ($< 0.04 \mu\text{mol L}^{-1}$) to $4.25 \mu\text{mol L}^{-1}$ across sites during the sampling period, and the highest concentrations were observed in the inner bay (Fig. 2e, 2f). The molar ratio of DIN to dissolved inorganic P (N:P; $\text{NO}_3^- + \text{NH}_4^+$ to PO_4^{3-}) was highly variable throughout the sampling period in 2015 and 2016, ranging from over 10,000 to below Redfield stoichiometry (16:1) (Fig. 2g, 2h). In general, high N:P values were observed earlier in the season, and low N:P values were observed later in the season as NO_3^- concentrations declined. Chl *a* concentrations ranged from 3.5 to nearly $150 \mu\text{g L}^{-1}$ across sites throughout the sampling period in 2015 and 2016 (Fig. 2i, 2j). Maximum chl *a* concentrations in both years occurred in late August to early September, approximately one month after the peak in NO_3^- and PO_4^{3-} concentrations.

Sediment N removal processes were active across the sampling period at station 1163 in 2015 (Fig. 4). Denitrification rates ranged from 10.02 – $64.81 \mu\text{mol N m}^{-2} \text{h}^{-1}$ over all sampled dates, decreasing over time coincident with declines in water column NO_3^- (ANOVA, $F_{3,24} = 6.53$, $p < 0.01$). Anammox activity was detected on all sampling dates but not in all replicate sediment cores. Anammox rates ranged from 0.52 – $8.10 \mu\text{mol N m}^{-2} \text{h}^{-1}$ across sampled dates, displaying no clear temporal trend (ANOVA, $F_{3,24} = 2.60$, $p = 0.08$). The majority of measured anammox rates were less than $7 \mu\text{mol N m}^{-2} \text{h}^{-1}$, with the exception of two cores on July 27 that displayed elevated anammox rates of 15.15 and $30.75 \mu\text{mol N m}^{-2} \text{h}^{-1}$. N_2O production rates at station 1163 ranged from 0.09 – $2.34 \mu\text{mol N m}^{-2} \text{h}^{-1}$ across sampled dates, decreasing over time coincident with declines in

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water column NO_3^- (ANOVA, $F_{3,24} = 6.85$, $p < 0.01$). Denitrification, anammox, and N_2O production made up an average of 84 %, 14 %, and 2 % of sediment N removal, respectively.

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NO_3^- and NH_4^+ uptake were active in the inner bay, outer bay, and nearshore Lake Erie throughout the sampling period in 2016. NO_3^- uptake rates ranged from 0.01-1.92 $\mu\text{mol N L}^{-1} \text{ h}^{-1}$ (Fig. 5a), and NH_4^+ uptake rates ranged from 0.001-0.11 $\mu\text{mol N L}^{-1} \text{ h}^{-1}$ (Fig. 5b). There was a significant interaction between the effects of date and location on NO_3^- uptake (ANOVA, $F_{6,24} = 133.14$, $p < 0.0001$) and NH_4^+ uptake (ANOVA, $F_{6,24} = 19.43$, $p < 0.0001$). Overall, NO_3^- and NH_4^+ uptake rates were higher within the bay than in the nearshore Lake Erie station. Rates of NO_3^- uptake were significantly higher than rates of NH_4^+ uptake (t -test, $df = 35$, $T = 2.41$, $p = 0.02$). As a result of ambient concentrations being lower than anticipated, there were several instances when the ^{15}N -labeled fraction of NO_3^- or NH_4^+ exceeded 10 % (marked with ^ in Fig. 5a, 5b). However, the elevation of substrate concentration was not associated with unusually high uptake rates in comparison with other dates at the same site, and these observations were retained in the dataset.

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Water column N fixation was active throughout the sampling period in 2015 and 2016, with rates ranging from 0.06-2.16 $\mu\text{mol N L}^{-1} \text{ h}^{-1}$ (Fig. 5c). N fixation rates varied significantly by date (ANOVA, $F_{7,25} = 156.33$, $p < 0.0001$), and rates across dates fell into significant groupings (marked with letters in Fig. 5c). The highest observed rates of N fixation occurred in late July and late August of 2015. The lowest observed rates of N fixation occurred in late July and late August of 2016. Areal rates of N fixation ranged from 309.5-906.9 $\mu\text{mol N m}^{-2} \text{ h}^{-1}$ in 2015 and 0.2-187.8 $\mu\text{mol N m}^{-2} \text{ h}^{-1}$ in 2016.

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Metagenomic analysis revealed that *Planktothrix* spp. dominated the cyanobacterial community (22-99 % of iTag reads) on all but one sampled date at station 1163. Diazotrophs (*Aphanizomenon* sp. and *Dolichospermum* spp.) were present on all sampled dates and made up a minor assemblage of the cyanobacterial community (1-33 % of iTag reads; Fig. 6).

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DIN concentration was positively correlated with daily volumetric rates of denitrification, N_2O production, and NO_3^- uptake, explaining 64, 60, and 75 % of variance in mean rates, respectively (Fig. 7a). Anammox rates were not correlated with DIN concentration due to the high variance observed among dates and replicate sediment cores (Fig. 7b). DIN concentration explained only 12 % of variance in mean NH_4^+ uptake rates, although NH_4^+ concentration alone explained 86 % of variance and was positively correlated with NH_4^+ uptake. N fixation was not significantly correlated with DIN concentration (Fig. 7a). Ranges in daily volumetric rates of assimilatory processes (NO_3^- uptake, NH_4^+ uptake, N fixation) were several orders of magnitude higher than dissimilatory processes (denitrification, anammox, N_2O production), with the exception of NH_4^+ uptake and denitrification, which were within the same magnitude (Fig. 7).

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Daily rates of NO_3^- and TKN loading from the Sandusky River varied by five and two orders of magnitude, respectively, and high loading rates were associated with high discharge (Fig. 8). N fixation often exceeded riverine N loading as a source of N

to Sandusky Bay, particularly during periods of low discharge, NO_3^- uptake was the dominant N uptake process in Sandusky Bay, outpacing NH_4^+ uptake and dissimilatory N removal processes (Fig. 8). On the basis of the total magnitude of ranges, sources and demands of N in this system are tipped in favor of a net source to Lake Erie, ranging from a strong source during periods of high discharge to a weak source during periods of low discharge.

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4 Discussion

4.1 Nutrient Stoichiometry

Sandusky Bay displays considerable seasonal variation in nutrient concentrations, molar dissolved N:P ratios, and chl *a* concentrations, indicative of a system with dynamic changes in hydrology and biogeochemical activity. Maximum chl *a* concentrations in both years ($> 100 \mu\text{g L}^{-1}$) were similar to other hypereutrophic systems (Zhang et al., 2011; Wheeler et al., 2012; Steffen et al., 2014b), as were large swings in NO_3^- concentrations in 2015 (Xu et al., 2010; Steffen et al., 2014b; McCarthy et al., 2016). Elevated nutrient concentrations were associated with high discharge events from the Sandusky River, demonstrating a strong watershed influence on Sandusky Bay. Indeed, the Sandusky River watershed comprises an area 30 times larger than the bay and delivers large nonpoint loads of N and P to its receiving waters (Robertson and Saad, 2011). Between the two study years, discharge from the Sandusky River varied substantially (tenfold higher in 2015), exhibiting large inter-annual variability in hydraulic residence time and nutrient concentrations.

As discharge from the Sandusky River decreased throughout the summer in both years, dissolved N:P ratios fell from a maximum of over 10,000 to below the threshold for N limitation. The decline in N:P ratios is largely driven by decreases in NO_3^- concentration, particularly in 2015, as the range in PO_4^{3-} and NH_4^+ concentration was comparatively narrow. Consumption of NO_3^- could be attributed to both assimilatory and dissimilatory NO_3^- reduction. If phytoplankton were solely responsible for the decline in NO_3^- , nutrients would be expected to be drawn down in molar proportions of approximately 16N:1P (Sterner and Elser, 2002). However, N:P ratios in Sandusky Bay fall sharply throughout the summer, while PO_4^{3-} concentrations are relatively constant by comparison. Although this trend could be influenced by luxury uptake of N by phytoplankton and internal P loading from sediments (Filbrun et al., 2013; McCarthy et al., 2016), the dramatic depletion in DIN compels consideration of microbial N removal processes as a major mechanism for N drawdown in Sandusky Bay.

4.2 N Removal Processes

Marked declines in N:P with time and occurrence of N:P ratios < 16 provide compelling evidence that microbial N removal processes (i.e., denitrification and/or anammox) consume appreciable quantities of NO_3^- in Sandusky Bay. Sediment ^{15}N tracer incubations indicate the primary N removal mechanism in Sandusky Bay is denitrification, which comprised an average of 84 % of sediment N removal across sampling dates. Denitrification rates were positively correlated with DIN concentration, consistent with observations that N supply controls sediment denitrification capacity in estuaries, lakes, and continental shelves

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(Seitzinger et al., 2006). Sandusky Bay exhibited efficient denitrification, with an average of only 2 % of sediment N removal released as N₂O. Denitrification rates in Sandusky Bay (10.02-64.81 μmol N m⁻² h⁻¹) are among the highest reported for the Laurentian Great Lakes. Previous denitrification measurements in offshore zones of the Great Lakes vary over four orders of magnitude, with western and central Lake Erie exhibiting the highest rates at 51 ± 41 μmol N m⁻² h⁻¹ (Small et al., 2014; 2016). Nearshore zones, bays, and river mouths have been observed as areas of enhanced denitrification and N₂O production compared to offshore zones (McCarthy et al., 2007; Small et al., 2014; 2016; Salk et al. 2016), reinforcing that Sandusky Bay and other shallow coastal areas have the potential to act as hotspots of N removal in the Great Lakes system.

Anammox activity in Sandusky Bay was highly variable, even among replicate sediment cores from the same site and date.

High variability was also observed in a study of potential anammox rates in the water column of Sandusky Bay and Lake Erie (Lu et al. 2018). Marked variability may be characteristic of anammox activity in freshwater environments, even across small spatial and temporal scales (Yoshinaga et al., 2011; Zhu et al., 2013; 2015). Anammox made up an average of 14 % of sediment N removal across the sampling period, indicating anammox activity in Sandusky Bay may be typical of shallow estuarine and freshwater systems (Thamdrup and Dalsgaard, 2002; Dalsgaard et al., 2005; Schubert et al., 2006; Dong et al., 2009; 2011; Hsu and Kao, 2013; McCarthy et al., 2016). Measurements of active anammox via isotope tracers are consistent with detection of 16s RNA associated with anammox taxa in the sediment of Lake Erie (Small et al., 2016), demonstrating that anammox has the potential to be an appreciable N removal pathway in this and other nearshore regions within the Great Lakes.

Whereas denitrification and anammox were the primary drivers of dissimilatory N uptake in Sandusky Bay, hydraulic residence time had a marked effect on N removal. In 2015, estimates of water residence time when discharge peaked in June and early July was as low as eight days. By late July, however, hydraulic residence time increased to several months and continued to increase as discharge remained low for the remainder of the summer and early fall. Although N removal rates were greatest when NO₃⁻ concentrations were highest, the capacity for N removal to substantially deplete NO₃⁻ was hindered by the short hydraulic residence time within the Bay. For instance, high N removal rates in late June 2015 coincided with high NO₃⁻ concentrations, but hydraulic residence times of 1-2 weeks would enable only ~ 1/4 of NO₃⁻ to be permanently removed from the system prior to release into Lake Erie. The depletion of N in Sandusky Bay occurred when water residence time lengthened to several months, which provided the opportunity for assimilatory and dissimilatory processes to extensively consume NO₃⁻. Although phytoplankton uptake represents another significant N consumption process, the sharp decline in dissolved N:P ratios during this period (Fig. 2g, 2h) indicates that dissimilatory processes (i.e., denitrification, anammox, N₂O production) were important drivers of NO₃⁻ concentration decline. Thus, Sandusky Bay acts as a conduit for N delivery from the Sandusky River to Lake Erie when hydraulic residence time is short but acts as a filter for NO₃⁻ during periods of long hydraulic residence time. Export of nutrients from Sandusky Bay to Lake Erie during periods of high discharge is illustrated by concomitant NO₃⁻ concentration spikes in Sandusky Bay and at the Lake Erie Bells station in July 2015 (Fig. 2a). Conversely, a signal of NO₃⁻ export from Sandusky Bay into Lake Erie is lost during periods of low discharge, with the Bells

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station displaying higher NO_3^- concentrations than bay stations in 2016 (Fig. 2b). The capacity for Sandusky Bay to act as an alternating conduit and filter for nutrients is consistent with observations in other river mouths and coastal wetlands in the Great Lakes (McCarthy et al., 2007; Larson et al., 2013; Conroy et al., 2017).

4.3 Phytoplankton N Acquisition

Planktothrix-dominated phytoplankton blooms were evident in Sandusky Bay in both 2015 and 2016 (Fig. 6). *Chl a*, a proxy for total phytoplankton biomass (Becker et al., 2009; Millie et al., 2009; Davis et al., 2012), reached maximum levels approximately one month after maximum NO_3^- and PO_4^{3-} concentrations were observed (Fig. 2). This offset in peak nutrient availability and peak phytoplankton biomass has been observed in other years as well (Conroy et al., 2017). The coincidence of peaks in *chl a* with low river discharge is consistent with the idea that long hydraulic residence times create a stable physical environment in which primary producers can flourish (Michalak et al., 2013). Dissolved N:P ratios during the period of highest phytoplankton abundance approached or exceeded the threshold for N limitation, suggesting that *Planktothrix* is successful in acquiring N during periods of scarcity.

Given that *Planktothrix* is an effective competitor for DIN (Conroy et al., 2007), NO_3^- and NH_4^+ were investigated as sources of N for HABs in Sandusky Bay. Indeed, NO_3^- and NH_4^+ uptake were active throughout the summer, demonstrating that low concentrations do not equate to the absence of an actively cycling DIN pool. Transient pools of NH_4^+ generated via water column recycling (Chaffin and Bridgeman, 2014; Donald et al., 2011; 2013; Davis et al., 2015) or sediment regeneration (Paerl et al., 2011; McCarthy et al., 2016) could help to support persistence of *Planktothrix* blooms in the late summer. Uptake rates were proportional to the respective concentrations of substrates, suggesting that although NH_4^+ is a less energetically costly source of N, NO_3^- is utilized by phytoplankton to a greater extent in this system owing to a greater relative abundance.

Remineralization was not specifically measured in NH_4^+ uptake assays, but it is likely that NH_4^+ is rapidly recycled in this system. In comparable systems, remineralization rates are on the same order as uptake rates, and this recycling represents a crucial N supply for cHABs (Hampel et al., 2018). The observation of higher TKN concentrations in Sandusky Bay than in the Sandusky River (Fig. 3) suggests that fixed N and assimilated NO_3^- are recycled into TKN, which is comprised of both organic N forms and NH_4^+ . Given the potential for NH_4^+ remineralization to cause dilution of the ^{15}N -enriched NH_4^+ pool in NH_4^+ uptake assays, a hypothetical dilution due to remineralization was calculated by assuming a remineralization rate equivalent to the NH_4^+ uptake rate. Supposing an immediate dilution rather than a progressive dilution throughout the incubation and an isotopic composition of remineralized NH_4^+ equal to that of particulate N, the calculation represents the maximum possible dilution that could take place during NH_4^+ uptake incubations. The resulting isotope dilution would result in an underestimation of NH_4^+ uptake by $41 \pm 11\%$ (mean \pm SD). Thus, a rapidly recycling pool of NH_4^+ could represent a significant source of N to *Planktothrix* in this system despite the difficulty in quantifying transient availability and uptake.

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N fixation was a major N uptake process in Sandusky Bay, often exceeding rates of NO_3^- and NH_4^+ uptake (Fig. 5). The phytoplankton community responsible for water column N fixation in Sandusky Bay is comprised largely of *Aphanizomenon* and *Dolichospermum* (Fig. 6). Areal N fixation rates ($309.5\text{--}906.9\ \mu\text{mol N m}^{-2}\ \text{h}^{-1}$ in 2015 and $0.2\text{--}187.8\ \mu\text{mol N m}^{-2}\ \text{h}^{-1}$ in 2016) are within the range of those observed in eutrophic lakes ($9.2\text{--}421.2\ \mu\text{mol N m}^{-2}\ \text{h}^{-1}$; Howarth et al., 1988 and references therein), with the exception of extremely high N fixation rates measured in May, July, and August 2015. The occurrence of high N fixation rates under N limitation in eutrophic Sandusky Bay is not surprising, but the observation of N fixation rates exceeding NO_3^- and NH_4^+ uptake rates on occasions when both substrates were readily available is unanticipated.

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On several occasions during periods of low discharge, riverine DIN loading alone could not meet phytoplankton uptake demands (Fig. 8). Consequently, fixation and subsequent recycling of N represent an additional potential source of bioavailable N for cHABs in Sandusky Bay. For dates when comparisons are available, N fixation comprised 23.7–85.4 % of total phytoplankton N uptake (NO_3^- uptake + NH_4^+ uptake + N fixation) in the inner and outer bay. This range is typical of eutrophic lakes, in which N fixation has been observed to comprise 5.5–82.0 % of N inputs (Howarth et al. 1988 and references therein). Given that a high proportion of N in the system is supplied via N fixation, even small transfers of N from diazotrophs could represent an important N source for *Planktothrix*. Specifically, N leaking out of actively N-fixing cells or from decomposing cells has the potential to supplement N supply for non-diazotrophs (Ohlendieck et al., 2000; Beversdorf et al., 2013). Indeed, TKN concentrations were higher in Sandusky Bay than in the Sandusky River during times when NO_3^- concentrations in the bay were very low (Fig. 2, 3), suggesting that fixed N is recycled into the dissolved N pool that can be taken up by non-diazotrophs such as *Planktothrix*. Thus, both riverine N loading and N fixation represent sources of bioavailable N that *Planktothrix* may scavenge during periods of N limitation in Sandusky Bay.

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Rates of N fixation were not correlated with DIN concentrations, a surprising outcome considering that DIN uptake is a less energetically costly process and is predicted to outcompete N fixation when DIN is available (Holl and Montoya, 2005). While this result is unexpected, previous work has shown that the presence of DIN suppresses the synthesis of the nitrogenase complex but not the activity of the existing enzyme (Fogg, 1971; Wolk, 1973; Chang et al., 1980), and high concentrations of NO_3^- only partially suppress heterocyst formation (Ogawa and Carr, 1969; Ohmori and Hattori, 1972). Indeed, N fixation in the presence of DIN and at dissolved N:P ratios greater than 16 have been observed elsewhere (Chen et al., 1996; Spröber et al., 2003; Voss et al., 2004; Moisander et al., 2008; Gao et al., 2014). Given the dynamic nature of hydrology and riverine DIN loading in river-lake mixing zones and estuarine systems, there may be energetic gains from maintaining N fixation machinery that can be utilized quickly following sudden swings in N availability (Moisander et al., 2012). While the precise causes of high rates of N fixation in this hypereutrophic system are not readily apparent, the occurrence of this process indicates that Sandusky Bay not only acts as a conduit for riverine DIN loading to Lake Erie but also a source. Therefore, watershed N management aimed at minimizing N export to Lake Erie may be offset by the introduction of N via N fixation in Sandusky Bay.

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4.4 N Budget

Sandusky Bay displays dynamic swings in hydraulic and nutrient regimes yet consistently develops seasonal N limitation and blooms of *Planktothrix* in the late summer. An examination of the N budget in this system will help to characterize the role of this Great Lakes ~~drowned river mouth~~ in mediating N delivery to Lake Erie (Conroy et al., 2017). On the whole, Sandusky Bay is a source of N to Lake Erie, ~~both as a conduit of watershed N loading and through introductions of fixed N. The magnitude of N delivered downstream is highly dependent on Sandusky River discharge (Fig. 2, 8).~~ During portions of the year when riverine DIN loading is low, N fixation supplements DIN loading to meet assimilatory and dissimilatory N demands, representing a large and crucial balance for the N budget in Sandusky Bay. ~~N fixation has been suggested as a mechanism for balancing the N budget in favor of a net source in other eutrophic lakes (Cook et al., 2010). N fixation and phytoplankton DIN uptake transfer N into organic forms that can be recycled within the system or delivered downstream.~~ The dominance of assimilatory processes suggests that although DIN concentrations are often low in Sandusky Bay, there is an actively cycling N stock within the phytoplankton community that may be utilized by *Planktothrix*. Dissimilatory sinks, although on a smaller magnitude, represent a permanent N sink that may have a greater influence on the development of N limitation than assimilatory ~~processes~~.

The mass balance of N in Sandusky Bay undergoes rapid and dramatic seasonal transitions, shifting the role of the bay from a strong to a weak source of N to Lake Erie. Previous work suggests that oscillations between excess N abundance to N limitation that are consistent from year to year (Conroy et al., 2007; Davis et al., 2015; Conroy et al., 2017). During periods of high discharge and N loading, ~~short hydraulic residence times prevent substantial processing of N and DIN is flushed into Lake Erie. When Sandusky River discharge and N loading are low, sediment N removal and N recycling by phytoplankton result in a smaller pulse of N delivery to Lake Erie that has been extensively processed into organic forms. Sandusky Bay thus oscillates between acting as a conduit and a filter of N. Discharge-driven oscillations in N cycling and downstream delivery may be a common feature in river-lake mixing zones and estuaries,~~ most notably Narragansett Bay (Fulweiler et al., 2007; Fulweiler and Heiss, 2014).

Future projections suggest that climate change will create conditions that are likely to intensify the ~~role of~~ Sandusky Bay and ~~comparable systems as conduits of downstream nutrient delivery.~~ Overall precipitation in the watershed is expected to increase and become dominated by incidences of extreme precipitation (Prein et al., 2017). Moreover, increases in precipitation are predicted to be accompanied of enhanced riverine N loading in the near future (Sinha et al., 2017). ~~Increases in the frequency and intensity of discharge and riverine N loading will likely diminish the capacity for Sandusky Bay to transform and remove N, thus favoring the downstream export of DIN. The increased export of nutrients downstream may support cHABs in the central basin of Lake Erie and exacerbate water quality issues downstream, including hypoxia in the central basin of Lake Erie and the St. Lawrence Estuary (Lehmann et al., 2009; Michalak et al., 2013). Similar climate change-driven shifts in water~~

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An important consideration for the N budget in Sandusky Bay is that N supply and N depletion are temporally separated.

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quality are anticipated in other coastal systems such as the Gulf of Mexico, where enhanced riverine N delivery and predicted increases in magnitude and timing of precipitation will make hypoxia mitigation efforts even more difficult (Turner et al., 2008; 2012). Sandusky Bay may thus serve as a harbinger for what can be expected in many coastal systems that are responding to climate change and increases in N loading.

Data Availability

Data and associated content for this manuscript can be found at https://github.com/KateriSalk/SanduskyBay_NitrogenCycle.git.

Author Contribution

- 10 This study was designed by KS, NO, GB, and RMM. Field sampling was carried out by KS, GB, RMM, JC, and associated research groups. Sample and data analysis were carried out by KS, NO, and JC. KS prepared the manuscript with contributions from NO, GB, RMM, and JC.

Acknowledgements

- 15 We thank Taylor Tuttle, Emily Davenport, Hasand Gandhi, Kristen Slodysko, Erica Fox, Kat Rossos, Callie Nauman, and [Keara Stanislawczyk](#) for their assistance in the field and laboratory. Thank you to Silvia Newell and Mark McCarthy for assistance with MIMS analysis, and Peggy Ostrom, Stephen Hamilton, and Silvia Newell for feedback on the manuscript. This research was supported by the NSF Graduate Research Fellowship [No. 497 DGE1424871], the Michigan State University (MSU) College of Natural Science Hensley Fellowship, the MSU Rose Fellowship in Water Research, the MSU WaterCube program, the Ohio Department of Higher Education's Harmful Algal Bloom Research Initiative [No. R/HAB-2-BOR], and the
- 20 Ohio Sea Grant College Program [No. R/ER-110]. [The work conducted by the U.S. DOE Joint Genome Institute, a DOE Office of Science User Facility, was supported by the Office of Science of the U.S. DOE under Contract No. DE-AC02-05CH11231.](#)

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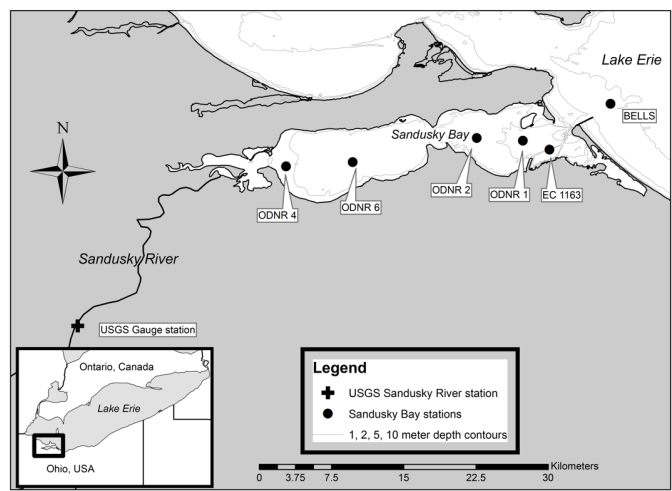
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Table 1. Total range in N sources, assimilatory N uptake processes, and dissimilatory N sinks in Sandusky Bay.
Process ... [1]

Figures



5 Fig. 1: Sampling locations in Sandusky Bay (circles) and Sandusky River monitoring station (USGS monitoring station 04198000, cross).

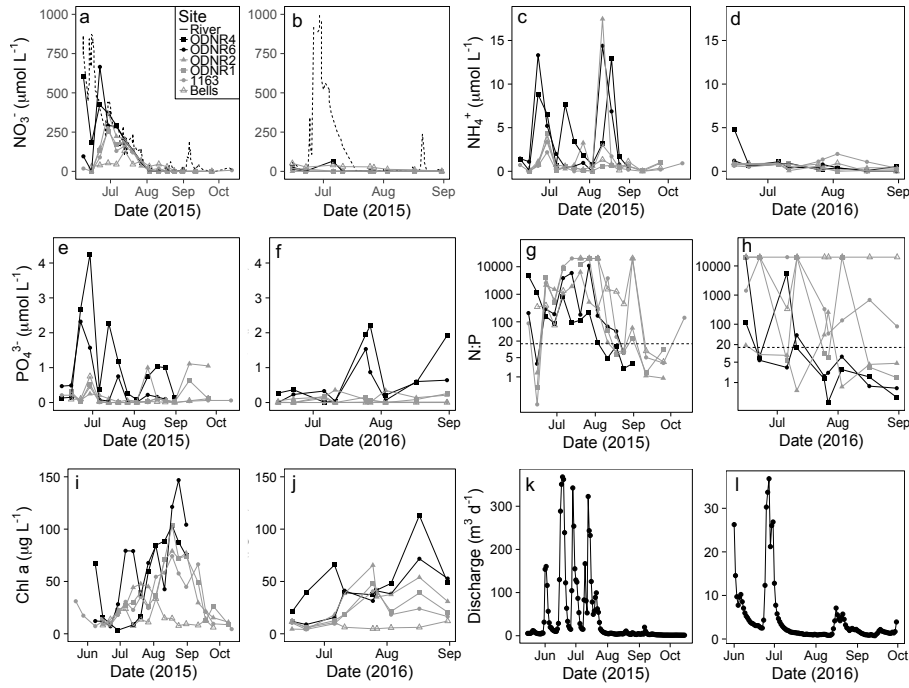


Fig. 2: (a-b) NO_3^- , (c-d) NH_4^+ , (e-f) PO_4^{3-} , (g-h) N:P ($\text{NO}_3^- + \text{NH}_4^+ : \text{PO}_4^{3-}$) ratio, (i-j) chlorophyll a , and (k-l) Sandusky River discharge, in 2015 and 2016, respectively. Six Sandusky Bay sites are presented: inner bay (ODNR and ODNR6, black filled symbols), the outer bay (ODNR2, ODNR1, and 1163, gray filled symbols), and a site outside the bay in the central basin of Lake Erie (Bells, gray open symbols). NO_3^- concentrations in the Sandusky River are presented as a dotted line (a-b). Note the consistent scales between years for single variables (exception: discharge) but differing scales among variables. The dotted line in (g-h) indicates a N:P ratio of 16 (note log scale on y axis).

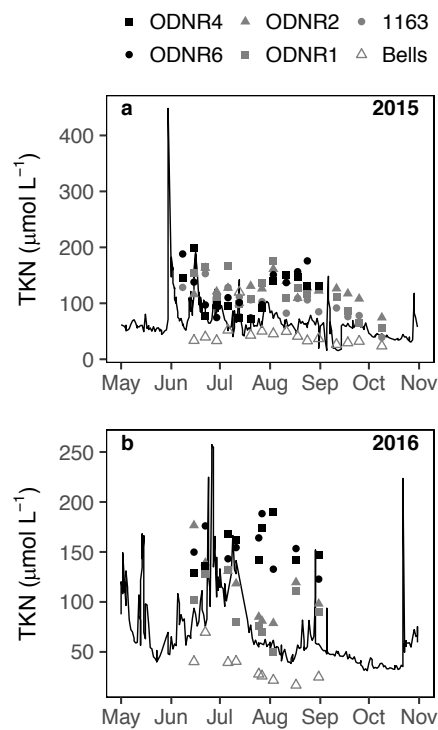


Fig. 3: Total Kjeldahl nitrogen (TKN) concentrations in the Sandusky River (line) and in Sandusky Bay (symbols) in (a) 2015 and (b) 2016. Note the difference in y-axis scales between panels a and b.

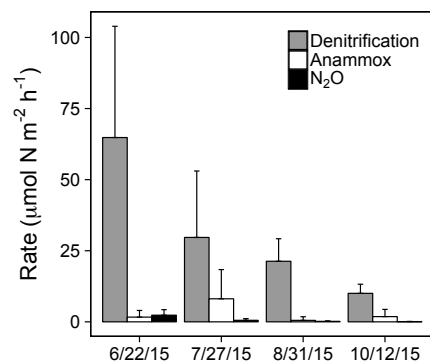


Fig. 4: Denitrification, anammox, and N₂O production rates at station 1163. Error bars represent +1 SD.

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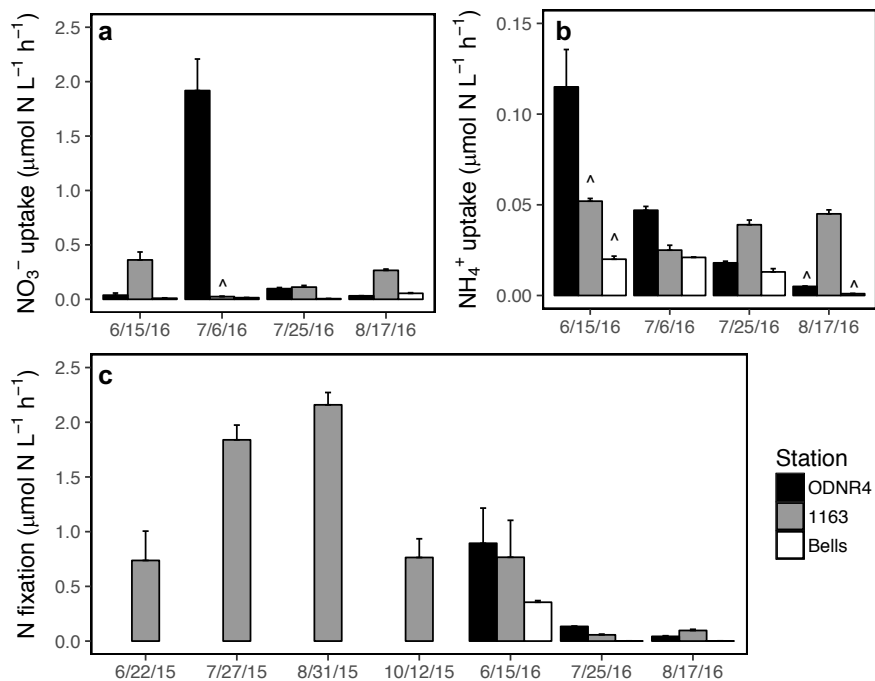


Fig. 5: (a) NO_3^- uptake rates, (b) NH_4^+ uptake rates, and (c) N fixation rates in Sandusky Bay. Rates were measured at stations ODN4 (black), 1163 (gray), and Bells (white). Note the differing y-axis scales among panels. Error bars represent ± 1 SD. Caret symbols (^) indicate incubations in which the ^{15}N addition exceeded 10 % of the ambient NO_3^- or NH_4^+ concentration, owing to unexpectedly low ambient concentrations.

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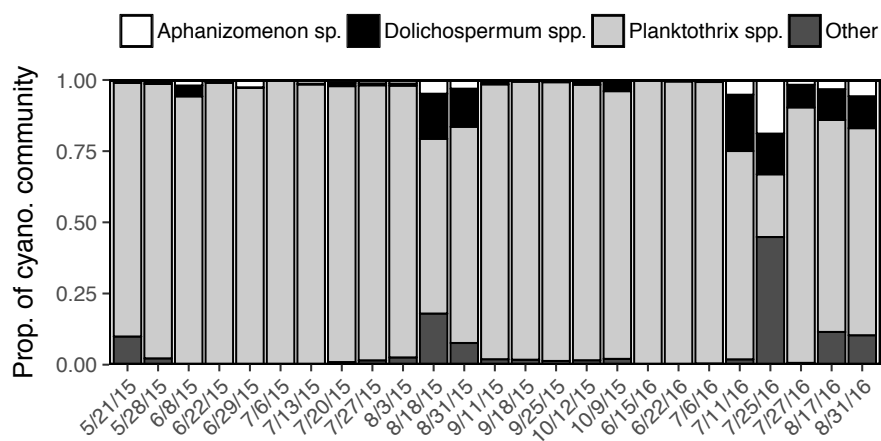


Fig. 6: Proportion of cyanobacterial community derived from iTag reads in 2015 and 2016. A single operational taxonomic unit was detected for *Aphaniizomenon*, and multiple operational taxonomic units were detected for *Dolichospermum* and *Planktothrix*.

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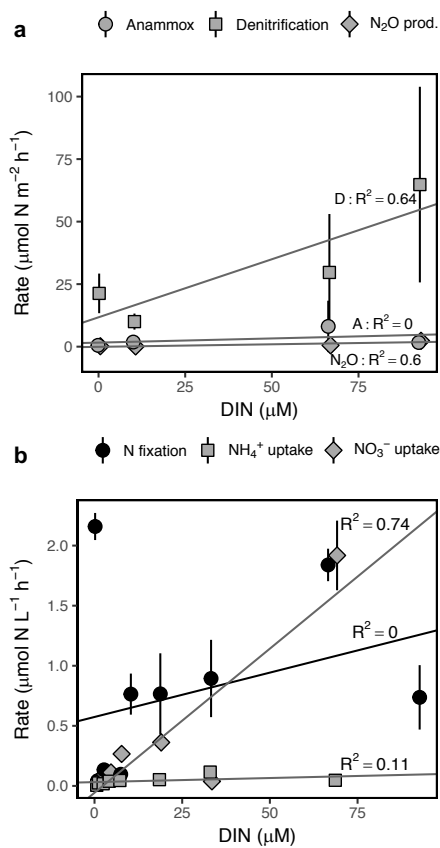
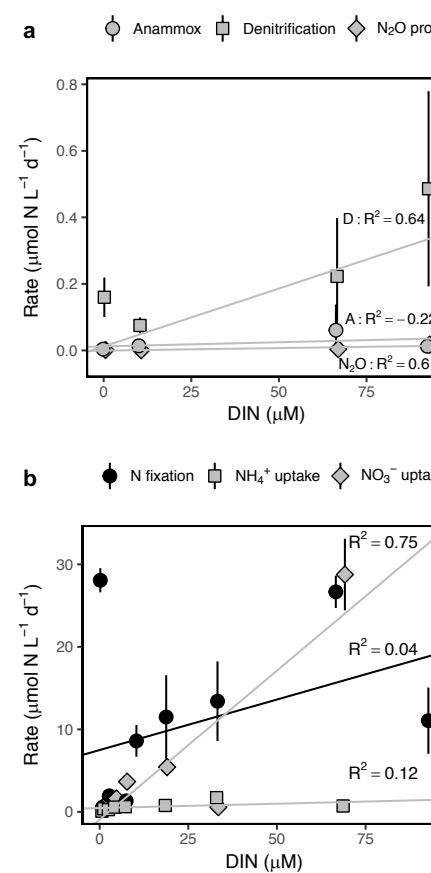


Fig. 7: Rates of (a) dissimilatory and (b) assimilatory N cycling processes as a function of DIN concentration ($\text{NH}_4^+ + \text{NO}_3^- + \text{NO}_2^-$). Note the differing y-axis scales among panels. Error bars represent ± 1 SD. R^2 represents the adjusted R^2 for each linear regression.



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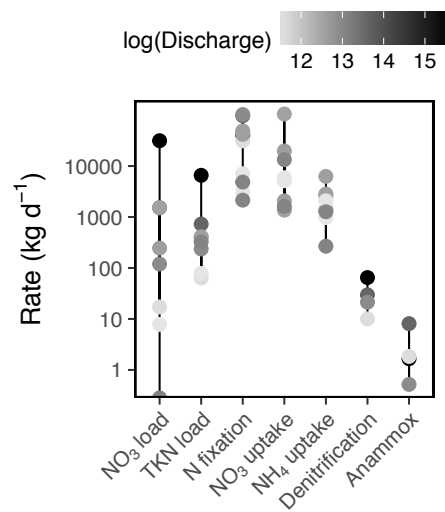


Fig. 8: Rates of Sandusky River N loading and Sandusky Bay N cycling. Note the log scale on the y axis. The color of each point corresponds to the log-transformed Sandusky River discharge on each date ($\text{m}^3 \text{d}^{-1}$). Sandusky River N loads are presented only for the dates on which Sandusky Bay was sampled, but this is representative of the range and distribution of loads throughout the summer.

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Tables

Table 1. Total range in N sources, assimilatory N uptake processes, and dissimilatory N sinks in Sandusky Bay.

Process	Minimum (kg d ⁻¹)	Maximum (kg d ⁻¹)
Sources		
DIN loading	1	355,054
N fixation	2,143	102,179
<i>Total</i>	<i>2,144</i>	<i>457,233</i>
Assimilatory N uptake		
NH ₄ ⁺ uptake	267	6,264
NO ₃ ⁻ uptake	1,360	104,709
<i>Total</i>	<i>1,627</i>	<i>110,973</i>
Dissimilatory N uptake		
Denitrification	274	1,769
Anammox	14	221
N ₂ O production	2	64
<i>Total</i>	<i>290</i>	<i>2,054</i>