- 1 Interactive comment on "Effect of elevated pCO_2 on trace gas production during an ocean
- 2 acidification mesocosm experiment" by Sheng-Hui Zhang et al.
- 3 B. Qu (Referee)
- 4 2467327342@qq.com
- 5 Received and published: 27 May 2018
- 6 Increases of anthropogenic emissions of CO₂ since the Industrial Revolution are known to have
- 7 influenced organisms and the delivery of oceanic ecosystem services at a global scale. This is an
- 8 interesting piece of work that shows the effect of elevated pCO_2 on trace gases production
- 9 including DMS and four halocarbon compounds through a mesocosm experiment. The study is
- based on the development of a bloom created by the addition of three different species of cultured
- 11 phytoplankton to nutrient enriched coastal water enclosed in mesocosms. Considering that the
- impact of ocean acidification on DMS and halocarbons remains controversial, it is necessary to
- conductfurther study about this aspect. Overall, this paper is well written and the major points are
- 14 discussed with clarity. I recommend this article to be published in Biogeosciences after
- modification. My major criticism to the manuscript is that the authors point the algae and their
- attached bacteria in the coastal environment were removed through filtration process, have you
- measured the bacterial abundance in the mesocosm before the three different species of algae
- 18 inoculated? In addition, this manuscript lacks the initial concentrations of Phaeodactylum
- 19 tricornuntum, Thalassiosira weissflogii, and Emiliania huxleyi inoculated into the mesocosm.
- There are also some minor thinks that I list below:
- 21 P3, L54 "Further decreases of 0.3–0.4 pH units are predicted by the end of this century (Doneyet
- al., 2009; Orr et al., 2005), which is commonly referred to as ocean acidification (OA)." Please
- 23 update the latest references in this section.
- 24 P3, L61 "DMS is the most important volatile sulfur compound produced from the algal secondary
- 25 metabolite dimethylsulfoniopropionate (DMSP) through complex biological in teractions in
- 26 marine ecosystems (Stefels et al., 2007)." DMSP is not only produced by algae, but also by
- 27 terrestrial plants and marine bacteria. Please re-word this section.
- P4, L75 Replace "attribute" by "attributed".
- 29 P8, L167-L168 What is "LC" and "HC", low CO₂ and high CO₂? Please use the full name for the
- first time in the manuscript.
- P8, L172 The unit of chl a is not unified with Fig. 1, please check.
- 32 P9, L192 Replace "for" by "of"
- P9, L196 delete "growth in"
- 34 P9, L197-198 Replace "increase in Chl a and cell concentrations" by "increase in Chl a
- 35 concentrations and algal cells"
- Response to Reviewer #1:
- 37 Dear Reviewer #1:

- We are grateful to your review of this paper and would like to express our thanks for your helpful
- and constructive comments. We have revised the manuscript and addressed all the comments point
- 40 by point. The main changes we made are as follows:
- 41 Increases of anthropogenic emissions of CO₂ since the Industrial Revolution are known to have
- 42 influenced organisms and the delivery of oceanic ecosystem services at a global scale. This is an
- 43 interesting piece of work that shows the effect of elevated pCO_2 on trace gases production
- 44 including DMS and four halocarbon compounds through a mesocosm experiment. The study is
- based on the development of a bloom created by the addition of three different species of cultured
- 46 phytoplankton to nutrient enriched coastal water enclosed in mesocosms. Considering that the
- 47 impact of ocean acidification on DMS and halocarbons remains controversial, it is necessary to
- 48 conductfurther study about this aspect. Overall, this paper is well written and the major points are
- 49 discussed with clarity. I recommend this article to be published in Biogeosciences after
- modification. My major criticism to the manuscript is that the authors point the algae and their
- attached bacteria in the coastal environment were removed through filtration process, have you
- 52 measured the bacterial abundance in the mesocosm before the three different species of algae
- 53 inoculated? In addition, this manuscript lacks the initial concentrations of *Phaeodactylum*
- 54 tricornuntum, Thalassiosira weissflogii, and Emiliania huxleyi inoculated into the mesocosm.
- Thanks for the reviewer's suggestion and we have added some details about this mesocosm
- experiment in the revised manuscript.
- 57 P6, L125-129 "Emiliania huxleyi (CS-369), Phaeodactylum tricornuntum (CCMA 106), and
- 58 Thalassiosira weissflogii (CCMA 102) were inoculated into the mesocosm bags, with initial
- 59 diatom/coccolithophorid cell ratio was 1:1. The initial concentrations of *Phaeodactylum*
- 60 tricornuntum, Thalassiosira weissflogii, and Emiliania huxleyi inoculated into the mesocosm were
- 61 10, 10, and 20 cells mL⁻¹, respectively."
- 62 P7, L141-142 "Meanwhile, no meaningful numbers of bacteria were counted by flow cytometer in
- the pre-filtered seawater before the inoculations."
- There are also some minor thinks that I list below:
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- al., 2009; Orr et al., 2005), which is commonly referred to as ocean acidification (OA)." Please
- 47 update the latest references in this section.
- Thanks for the reviewer's suggestion and we have updated the latest references in the revised
- 69 manuscript.
- P3, L58-60 "Further decreases of 0.3–0.4 pH units are predicted by the end of this century (Doney
- et al., 2009; Orr et al., 2005; Gattuso et al., 2015), which is commonly referred to as ocean
- acidification (OA)"
- "Gattuso, J. P., Magnan, A., Bille, R., Cheung, W. W. L., Howes, E. L., Joos, F., Allemand, D.,
- Bopp, L., Cooley, S. R., Eakin, C. M., Hoegh-Guldberg, O., Kelly, R. P., Portner, H. O.,

- Rogers, A. D., Baxter, J. M., Laffoley, D., Osborn, D., Rankovic, A., Rochette, J., Sumaila,
- 76 U.R., Treyer, S., Turley, C.: Contrasting futures for ocean and society from different
- anthropogenic CO₂ emissions scenarios. Science, 349 (6243), aac4722, 2015."
- 78 P3, L61 "DMS is the most important volatile sulfur compound produced from the algal secondary
- 79 metabolite dimethylsulfoniopropionate (DMSP) through complex biological in teractions in
- 80 marine ecosystems (Stefels et al., 2007)." DMSP is not only produced by algae, but also by
- 81 terrestrial plants and marine bacteria. Please re-word this section.
- 82 Thanks for the reviewer's suggestion and we have reworded this section in the revised manuscript.
- 83 P3, L67-71 "DMS is the most important volatile sulfur compound produced from
- 84 dimethylsulfoniopropionate (DMSP), which is ubiquitous in marine environments, mainly
- 85 synthesized by marine microalgae (Stefels et al., 2007), a few angiosperms, some corals (Raina et
- 86 al., 2016), and several heterotrophic bacteria (Curson et al., 2017) through complex biological
- interactions in marine ecosystems."
- 88 "Raina, J. B., Tapiolas, D., Motti, C. A., Foret, S., Seemann, T., Tebben, J.: Isolation of an
- antimicrobial compound produced by bacteria associated with reef-building corals. PeerJ, 4,
- 90 e2275, 2016"
- 91 "Curson, A. R., Liu, J., Bermejo Martinez, A., Green, R., Chan, Y., Carrion, O.:
- Dimethylsulfoniopropionate biosynthesis in marine bacteria and identification of the key
- gene in this process. Nat. Microbiol., 2, 17009, 2017."
- 94 P4, L75 Replace "attribute" by "attributed".
- Thanks for the reviewer's suggestion and we have reworded this section in the revised manuscript.
- 96 P4, L80-84 "Several assumptions have been presented to explain these contrasting results and
- 97 attributed the pH-induced variation in DMS-production capability to altered physiology of the
- algae cells or of bacterial DMSP degradation (Vogt et al., 2008; Hopkins et al., 2010, Avgoustidi
- 99 et al., 2012; Archer et al., 2013; Hopkins and Archer, 2014; Webb et al., 2015)"
- P8, L167-L168 What is "LC" and "HC", low CO₂ and high CO₂? Please use the full name for the
- first time in the manuscript.
- Thanks for the reviewer's suggestion and we have used the full name for the first time in the
- revised manuscript.
- P9, L192-195 "The initial chemical parameters of the mesocosm experiment are shown in Table 1.
- The initial mean dissolved nitrate (including NO₃⁻ and NO₂⁻), NH₄⁺, PO₄³⁻ and silicate (SiO₃²⁻)
- concentrations were 54, 20, 2.6 and 41 μ mol L⁻¹ for the low pCO₂ (LC) treatment and 52, 21, 2.4
- and 38 μ mol L⁻¹ for the high pCO_2 (HC) treatment, respectively."
- P8, L172 The unit of chl a is not unified with Fig. 1, please check.
- According to the opinion of reviewer 2#, Fig. 1 was replaced.

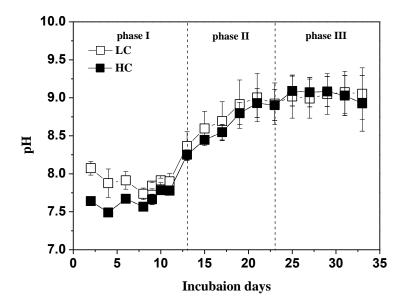


Fig. 1. Temporal changes of pH in the HC (1,000 μ atm, solid squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm standard deviation, n = 3 (triplicate independent mesocosm bags) (Origin 8.0).

114 P9, L192 Replace "for" by "of"

Thanks for the reviewer's suggestion and we have reworded in the revised manuscript according all reviewers' suggestion.

P10, L207-L209 "At the beginning of the experiment, the mean DMS, DMSP and DCB concentrations were all low in both treatments due to the low concentrations of DMS, DMSP and DCB in the original fjord water and possible loss during the filtration procedure (Fig. 2)."

120 P9, L196 delete "growth in"

Thanks for the reviewer's suggestion and we have modified in the revised manuscript.

P10, L217-218 "Compared with DMSP, DMS and DCB concentrations showed similar trends during the mesocosm experiment."

P9, L197-198 Replace "increase in Chl a and cell concentrations" by "increase in Chl a concentrations and algal cells"

Thanks for the reviewer's suggestion and we have modified in the revised manuscript.

P10, L210-212 "The DMSP concentrations in the HC and LC treatments increased significantly along with the increase of Chl *a* concentrations and algal cells, and stayed relatively constant over the following days."

Effect of elevated pCO_2 on trace gas production during an

- ocean acidification mesocosm experiment
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- 135 Shan Gao⁴, Da-Wei Pan³

- 136 ¹ Key Laboratory of Marine Chemistry Theory and Technology, Ministry of Education, Ocean
- University of China, Oingdao 266100, China
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- 155 Author contributions
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Abstract

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phytoplankton, bacteria

A mesocosm experiment was conducted in Wuyuan Bay (Xiamen), China to investigate the effects of elevated pCO₂ on phytoplankton species and production of dimethylsulfide (DMS), dimethylsulfoniopropionate (DMSP) and DMSP-consuming bacteria (DCB) as well as four halocarbon compounds (CHBrCl₂, CH₃Br, CH₂Br₂, and CH₃I). Over a period of 5 weeks, Phaeodactylum tricornuntum outcompeted Thalassiosira weissflogii and Emiliania huxleyi, comprising more than 99% of the final biomass. During the logarithmic growth phase (phase I), DMS concentrations in high pCO₂ mesocosms (HC, 1000 µatm) were 28% lower than those in low pCO₂ mesocosms (LC, 400 µatm). Elevated pCO2 led to a delay in DCB concentrations attached to Thalassiosira weissflogii and Phaeodactylum tricornutum and finally resulted in the delay of DMS concentration in the HC treatment. Unlike DMS, the elevated pCO_2 did not affect DMSP production ability of *Thalassiosira weissflogii* or Phaeodactylum tricornuntum throughout the 5 weeks culture. A positive relationship was detected between CH₃I and Thalassiosira weissflogii and Phaeodactylum tricornuntum during the experiment, and there was a 40% reduction in mean CH₃I concentrations in the HC mesocosms. CHBrCl₂, CH₃Br, and CH₂Br₂ concentrations did not increase with elevated chlorophyll a (Chl a) concentrations compared with DMS(P) and CH3I, and there were no major peaks both in the HC or LC mesocosms. In addition, no effect of elevated pCO_2 was identified for any of the three bromocarbons. Keywords: ocean acidification, dimethylsulfide (DMS), dimethylsulfoniopropionate (DMSP), halocarbons,

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1. Introduction

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Anthropogenic emissions have increased the fugacity of atmospheric carbon dioxide (pCO_2) from the pre-industrial value of 280 µatm to the present-day value of over 400 µatm, and these values will further increase to 800-1000 µatm by the end of this century according to the Intergovernmental Panel on Climate Change (Gattuso et al., 2015). The dissolution of this excess CO₂ into the surface of the ocean directly affects the carbonate system and has lowered the pH by 0.1 units, from 8.21 to 8.10 over the last 250 years. Further decreases of 0.3-0.4 pH units are predicted by the end of this century (Doney et al., 2009; Orr et al., 2005; Gattuso et al., 2015), which is commonly referred to as ocean acidification (OA). The physiological and ecological aspects of the phytoplankton response to this changing environment can potentially alter marine phytoplankton community composition, community biomass, and feedback to biogeochemical cycles (Boyd and Doney, 2002). These changes simultaneously have an impact on some volatile organic compounds produced by marine phytoplankton (Liss et al., 2014; Liu et al., 2017), including the climatically important trace gas dimethylsulfide (DMS) and a number of volatile halocarbon compounds. DMS the most important volatile sulfur compound produced dimethylsulfoniopropionate (DMSP), which is ubiquitous in marine environments, mainly synthesized by marine microalgae (Stefels et al., 2007), a few angiosperms, some corals (Raina et al., 2016), and several heterotrophic bacteria (Curson et al., 2017) through complex biological interactions in marine ecosystems. Although it remains controversial, DMS and its by-products, such as methanesulfonic acid and non-sea-salt sulfate, are suspected to have a prominent part in climate feedback (Charlson et al., 1987; Quinn and Bates, 2011). The conversion of DMSP to

DMS is facilitated by several enzymes, including DMSP-lyase and acyl CoA transferase (Kirkwood et al., 2010; Todd et al., 2007); these enzymes are mainly found in phytoplankton, macroalgae, Symbiodinium, bacteria and fungi (de Souza and Yoch, 1995; Stefels and Dijkhuizen, 1996; Steinke and Kirst, 1996; Bacic and Yoch, 1998; Yost and Mitchelmore, 2009). Several studies have shown a negative impact of decreasing pH on DMS-production capability (Hopkins et al., 2010; Avgoustidi et al., 2012; Archer et al., 2013; Webb et al., 2016), while others have found either no effect or a positive effect (Vogt et al., 2008; Hopkins and Archer, 2014). Several assumptions have been presented to explain these contrasting results and attributed the pHinduced variation in DMS-production capability to altered physiology of the algae cells or of bacterial DMSP degradation (Vogt et al., 2008; Hopkins et al., 2010, Avgoustidi et al., 2012; Archer et al., 2013; Hopkins and Archer, 2014; Webb et al., 2015). Halocarbons also play a significant role in the global climate because they are linked to tropospheric and stratospheric ozone depletion and a synergistic effect of chlorine and bromine species has been reported that they may account for approximately 20% of the polar stratospheric ozone depletion (Roy et al., 2011). In addition, iodocarbons can release atomic iodine (I) quickly through photolysis in the atmospheric boundary layer and I atoms are very efficient in the catalytic removal of O₃ which governs the lifetime of many climate relevant gases including methane and DMS (Jenkins et al., 1991). Compared with DMS, limited attention was received about the effect of OA on halocarbon concentrations. Hopkins et al. (2010) and Webb et al. (2015) measured lower concentrations of several iodocarbons, while bromocarbons were unaffected by elevated pCO₂ through two acidification experiments. In addition, an additional mesocosm study did not elicit significant differences from any halocarbon compounds at up to 1,400 μ atm pCO₂ (Hopkins et al.,

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227 2013).

DMS and halocarbons play a significant role in the global climate and perhaps act a greater extent in the future. Meanwhile, the combined picture arising from existing studies is that the response of communities to OA is not predictable and further studies were required. Based on the controversial results about OA on DMS and halocarbons production, a mesocosm experiment was conducted in Wu Yuan Bay, Xiamen. The aim of this study was to investigate the influence of elevated pCO_2 on diatoms and coccolithophores and to further understand how the productions of DMS and halocarbons respond to OA.

2. Experimental method

2.1 General experimental device

The mesocosm experiments were carried out on a floating platform at the Facility for Ocean Acidification Impacts Study of Xiamen University (FOANIC-XMU, 24.52 N, 117.18 \pm) (for full technical details of the mesocosms, see Liu et al. 2017). Six cylindrical transparent thermoplastic polyurethane bags with domes were deployed along the south side of the platform. The width and depth of each mesocosm bag was 1.5 m and 3 m, respectively. Filtered (0.01 μ m, achieved using an ultrafiltration water purifier, MU801-4T, Midea, Guangdong, China) in situ seawater was pumped into the six bags simultaneously within 24 h. A known amount of NaCl solution was added to each bag to calculate the exact volume of seawater in the bags, according to a comparison of the salinity before and after adding salt (Czerny et al., 2013). The initial in situ pCO_2 was about 650 μ atm. To set the low and high pCO_2 levels, we added Na₂CO₃ solution and CO₂ saturated seawater to the mesocosm bags to alter total alkalinity and dissolved inorganic carbon (Gattuso et al., 2010; Riebesell et al., 2013). Subsequently, during the whole experimental

process, air at the ambient (400 μ atm) and elevated pCO_2 (1000 μ atm) concentrations was continuously bubbled into the mesocosm bags using a CO_2 Enricher (CE-100B, Wuhan Ruihua Instrument & Equipment Ltd., Wuhan, China). Because the seawater in the mesocosm was filtered, the algae in the coastal environment and their attached bacteria were removed and the trace gases produced in the environment did not influence the mesocosm trace gas concentrations after the bags were sealed.

255 2.2 Algal strains

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Emiliania huxleyi (CS-369), Phaeodactylum tricornuntum (CCMA 106), and Thalassiosira weissflogii (CCMA 102) were inoculated into the mesocosm bags, with an initial diatom/coccolithophorid cell ratio of 1:1. The initial concentrations of *Phaeodactylum* tricornuntum, Thalassiosira weissflogii, and Emiliania huxleyi inoculated into the mesocosm were 10, 10, and 20 cells mL⁻¹, respectively. Phaeodactylum tricornuntum and Thalassiosira weissflogii were obtained from the Center for Collections of Marine Bacteria and Phytoplankton of the State Key Laboratory of Marine Environmental Science (Xiamen University). Phaeodactylum tricornuntum was originally isolated from the South China Sea in 2004 and Thalassiosira weissflogii was isolated from Daya Bay in the coastal South China Sea. Emiliania huxleyi was originally isolated in 1992 from the field station of the University of Bergen (Raunefjorden; 60°18'N, 05°15'E). Before being introduced into the mesocosms, the three phytoplankton species were cultured in autoclaved, pre-filtered seawater from Wuyuan Bay at 16 °C (similar to the in situ temperature of Wuyuan Bay) without any addition of nutrients. Cultures were continuously aerated with filtered ambient air containing 400 µatm of CO₂ within plant chambers (HP1000G-D, Wuhan Ruihua Instrument & Equipment, China) at a constant

- bubbling rate of 300 mL min⁻¹. The culture medium was renewed every 24 h to maintain the cells
- of each phytoplankton species in exponential growth. Meanwhile, no meaningful numbers of
- bacteria were counted by flow cytometer in the pre-filtered seawater before the inoculations.
- 2.3 Sampling for DMS(P) and halocarbons
- 275 DMS(P) and halocarbons samples were generally obtained from six mesocosms at 9 a.m., then all
- 276 collected samples were transported into a dark cool box back to the laboratory onshore for
- analysis within 1 h. For DMS analysis, 2 mL sample was gently filtered through a 25 mm GF/F
- 278 (glass fiber) filter and transferred to a purge and trap system linked to a Shimadzu GC-2014 gas
- 279 chromatograph (Tokyo, Japan) equipped with a glass column packed with 10% DEGS on
- 280 Chromosorb W-AW-DMCS (3 m × 3 mm) and a flame photometric detector (FPD) (Zhang et al.,
- 281 2014). For total DMSP analysis, 10 mL water sample was fixed using 50 μL of 50 % H₂SO₄ and
- sealed (Kiene and Slezak, 2006). After > 1 d preservation, DMSP samples were hydrolysed for 24
- 283 h with a pellet of KOH (final pH > 13) to fully convert DMSP to DMS. Then, 2 mL hydrolysed
- sample was carefully transferred to the purge and trap system mentioned above for extraction of
- DMS. For halocarbons, 100 mL sample was purged at 40 °C with pure nitrogen at a flow rate of
- 286 100 mL min⁻¹ for 12 min using another purge and trap system coupled to an Agilent 6890 gas
- chromatograph (Agilent Technologies, Palo Alto, CA, USA) equipped with an electron capture
- detector (ECD) as well as a 60 m DB-624 capillary column (0.53 mm ID; film thickness, 3 μm)
- 289 (Yang et al., 2010). The analytical precision for duplicate measurements of DMS(P) and
- halocarbons was > 10%.
- 291 2.4 Measurements of chlorophyll a
- 292 Chlorophyll a (Chl a) was measured in water samples (200–1,000 mL) collected every 2 d at 9

methanol overnight at 4 °C and centrifuged at 5000 r min⁻¹ for 10 min. The absorbance of the 294 295 supernatant (2.5 mL) was measured from 250 to 800 nm using a scanning spectrophotometer (DU 296 800, Beckman Coulter Inc., Brea, CA, USA). Chl a concentration was calculated according to the 297 equation reported by Porra (2002). 298 2.5 Enumeration of DMSP-consuming bacteria (DCB) 299 The number of DMSP-consuming bacteria (DCB) was estimated using the most probable number 300 (MPN) methodology. The MPN medium consisted of a mixture (1:1 v/v) of sterile artificial sea 301 water (ASW) and mineral medium (Visscher et al., 1991), 3 mL of which was dispensed in 6 mL 302 test tubes, which were closed off by an over-sized cap, allowing gas exchange. Triplicate dilution 303 series were set up. All test tubes contained 1 mmol L⁻¹ DMSP as the sole organic carbon source 304 and were kept at 30 °C in the dark. After 2 weeks, the presence/absence of bacteria in the tubes 305 was verified by DAPI staining (Porter and Feig, 1980). Three tubes containing 3 mL ASW 306 without substrate were used as controls. 307 2.6 Statistical analysis 308 One-way analysis of variance (ANOVA), Tukey's test, and the two-sample t-test were carried out 309 to demonstrate the differences between treatments. A p-value < 0.05 was considered significant. 310 Relationships between DMS(P), halocarbons and a range of other parameters were detected using 311 Pearson's correlation analysis via SPSS 22.0 for Windows (SPSS Inc., Chicago, IL, USA). 312 3. Results and Discussion 313 3.1 Temporal changes in pH, Chl a, Phaeodactylum tricornuntum, Thalassiosira weissflogii, and

a.m. by filtering onto Whatman GF/F filters (25 mm). The filters were placed in 5 mL 100%

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Emiliania huxleyi during the experiment

During the experiment, the seawater in each mesocosm was well combined, and the temperature and salinity were well controlled, with a mean of 16 °C and 29 in all mesocosms, respectively. Meanwhile, we observed significant differences in pH levels between the two CO₂ treatments on days 0–11, but the differences disappeared with subsequent phytoplankton growth (Fig. 1). The phytoplankton growth process was divided into three phases in terms of variations in Chl a concentrations in the mesocosm experiments as described in Liu et al. (2017): i) the logarithmic growth phase (phase I, days 0-13), ii) a plateau phase (phase II, days 13-23, bloom period), and iii) a secondary plateau phase (phase III, days 23-33) attained after a decline in biomass from a maximum in phase II. The initial chemical parameters of the mesocosm experiment are shown in Table 1. The initial mean dissolved nitrate (including NO₃⁻ and NO₂⁻), NH₄⁺, PO₄³⁻ and silicate (SiO_3^{2-}) concentrations were 54, 20, 2.6 and 41 µmol L⁻¹ for the low pCO_2 (LC) treatment and 52, 21, 2.4 and 38 μ mol L⁻¹ for the high pCO_2 (HC) treatment, respectively. The nutrient concentrations (NO₃⁻, NO₂⁻, NH₄⁺ and phosphate) during phase I were consumped rapidly and their concentrations were below or close to the detection limit during phase II (Table 1). In addition, although dissolved inorganic nitrogen (NH₄⁺, NO₃⁻, and NO₂⁻) and phosphate were depleted, Chl a concentration in both treatments (biomass dominated by Phaeodactylum tricornuntum) remained constant over days 12–22, and then declined over subsequent days (Liu et al., 2017). Emiliania huxleyi was only found in phase I and its maximal concentration reached 310 cells mL⁻¹ according to the results of microscopic inspection. *Thalassiosira weissflogii* was found throughout the entire period in each bag, but the maximum concentration was 8,120 cells mL⁻¹, which was far less than the concentration of *Phaeodactylum tricornutum* with a maximum density of about 1.5 million cells mL⁻¹ (Liu et al., 2017).

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3.2 Impact of elevated pCO₂ on DMS and DMSP production

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338 At the beginning of the experiment, the mean DMS, DMSP and DCB concentrations were all low 339 in both treatments due to the low concentrations of DMS, DMSP and DCB in the original fjord 340 water and possible loss during the filtration procedure (Fig. 2). With the growth of phytoplankton, 341 DMS, DMSP and DCB showed slightly different trends during the mesocosm experiment. The 342 DMSP concentrations in the HC and LC treatments increased significantly along with the increase 343 of Chl a concentrations and algal cells, and stayed relatively constant over the following days. A 344 significant positive relationship was observed between DMSP and phytoplankton in the 345 experiment (r = 0.961, p < 0.01 for Phaeodactylum tricornuntum, r = 0.617, p < 0.01 for 346 Thalassiosira weissflogii in the LC treatment, table 2; r = 0.954, p < 0.01 for Phaeodactylum 347 tricornuntum, r = 0.743, p < 0.01 for Thalassiosira weissflogii in the HC treatment, table 3). 348 Compared with DMSP, DMS and DCB concentrations showed similar trends during the 349 mesocosm experiment. DMS concentrations in the LC and HC treatments were 1.03 and 0.74 350 nmol L⁻¹, respectively, while DCB concentrations in the LC and HC treatments were 0.20×10^6 351 and 0.16×10^6 cells mL⁻¹. DMS and DCB concentrations did not increase significantly during 352 phase I, but began to increase rapidly on day 15. DCB concentrations in the LC and HC treatments peaked on days 21 (11.65 \times 10⁶ cells mL⁻¹) and 23 (10.70 \times 10⁶ cells mL⁻¹), while DMS 353 concentrations in the LC and HC treatments peaked on days 25 (112.1 nmol L⁻¹) and 30 (101.9 354 355 nmol L⁻¹). Both DMS and DCB concentrations began to decrease obviously during phase III. 356 Meanwhile, a significant positive relationship was also observed between DMS and 357 Phaeodactylum tricornuntum (r = 0.560, p < 0.05 in the LC treatment; r = 0.635, p < 0.01 in the 358 HC treatment), while no relationship was observed between DMS and Thalassiosira weissflogii

(table 2 and table 3) during the experiment.

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In this study, no difference in mean DMSP concentrations was observed between the two treatments, indicating that elevated pCO₂ had no significant influence on DMSP production in Phaeodactylum tricornuntum and Thalassiosira weissflogii throughout this study. However, a significant 29% reduction in DMS concentrations was detected in the HC treatment compared with the LC treatment (p = 0.016), though no statistical difference for DCB concentrations was found between the LC and HC treatments during phase I. This reduction in DMS concentrations may be attributed to greater consumption of DMS and conversion to DMSO (Webb et al., 2015). In addition, the peak DMS concentration in the HC treatment was delayed 5 days relative to that in the LC treatment during phase II (Fig. 2-A). This result has been observed in previous mesocosm experiments and it was attributed to small scale shifts in community composition and succession that could not be identified with only a once-daily measurement regime (Vogt et al., 2008; Webb et al., 2016). However, this phenomenon can be explained in another straightforward way during this study. Previous studies have showed that marine bacteria play a key role in DMS production and the efficiency of bacteria converting DMSP to DMS may vary from 2 to 100% depending on the nutrient status of the bacteria and the quantity of dissolved organic matter (Simóet al., 2002, 2009; Kiene et al., 1999, 2000). In addition, a significant positive relationship was also observed between DMS and DCB (r = 0.643, p < 0.01 in the LC treatment; r = 0.544, p < 0.01 in the HC treatment) during this experiment. All of these observations point to the importance of bacteria in DMS and DMSP dynamics. During the present mesocosm experiment, DMSP concentrations in the LC treatment decreased slightly on day 23, while the slight decrease appeared on day 29 in the HC treatment (Fig. 2-B). In addition, the time that the DMSP concentration began to decrease was

very close to the time when the highest DMS concentration occurred in both treatments. Moreover, DCB peaked on days 21 (11.65 \times 10⁶ cells mL⁻¹) and 23 (10.70 \times 10⁶ cells mL⁻¹) in the LC and HC treatments, respectively, as shown in Fig. 2-C. Similar to DMS, DCB was also delayed in the HC mesocosm compared to that in the LC mesocosm. Taken together, we inferred that the elevated pCO₂ first delayed growth of DCB in the mesocosm, then the delayed DCB postponed the DMSP degradation process, and eventually delayed the DMS concentration in the HC treatment. In addition, considering that the algae and their attached bacteria were removed through a filtering process before the experiment and the unattached bacteria were maintained in a relatively constant concentration during this mesocosm experiment (Huang et al., 2018), we further concluded that the elevated pCO₂ controlled DMS concentrations mainly by affecting DCB attached to *Thalassiosira weissflogii* and *Phaeodactylum tricornuntum*.

3.3 Impact of elevated pCO₂ on halocarbon compounds

The temporal development in CHBrCl₂, CH₃Br, and CH₂Br₂ concentrations is shown in Fig. 3 (A–C) and the temporal changes of their concentrations were substantially different from those of DMS, DMSP, *Phaeodactylum tricornuntum* and *Thalassiosira weissflogii*. The mean concentrations of CHBrCl₂, CH₃Br and CH₂Br₂ for the entire experiment were 8.58, 7.85, and 5.13 pmol L⁻¹ in the LC treatment and 8.81, 9.73, and 6.27 pmol L⁻¹ in the HC treatment. The concentrations of CHBrCl₂, CH₃Br, and CH₂Br₂ did not increase with the Chl *a* concentration compared with those of DMS and DMSP, and no major peaks were detected in the mesocosms. In addition, no effect of elevated pCO₂ was identified for any of the three bromocarbons, which compared well with previous mesocosm findings (Hopkins et al., 2010, 2013; Webb, et al., 2016). No clear correlation was observed between the three bromocarbons and any of the measured algal

groups (table 2 and table 3), indicating that Phaeodactylum tricornuntum and Thalassiosira weissflogii did not primarily release these three bromocarbons during the mesocosm experiment. Previous studies have reported that large-size cyanobacteria, such as Aphanizomenon flos-aquae, produce bromocarbons (Karlsson et al. 2008) and significant correlations between cyanobacterium abundance and several bromocarbons have been reported in the Arabian Sea (Roy et al., 2011). However, the filtration procedure led to the loss of cyanobacterium in the mesocosms and finally resulted in low bromocarbon concentrations during the experiment, although Phaeodactylum tricornuntum and Thalassiosira weissflogii abundances were high. The temporal dynamics of CH₃I in the HC and LC treatments are shown in Fig. 3-D. The CH₃I concentrations in the LC treatment varied from 0.38 to 12.61 pmol L⁻¹, with a mean of 4.76 pmol L⁻¹. The CH₃I concentrations in the HC treatment ranged between 0.44 and 8.78 pmol L⁻¹, with a mean of 2.88 pmol L⁻¹. The maximum CH₃I concentrations in the HC and LC treatments were both observed on day 23. The range of CH₃I concentrations during this experiment was similar to that measured in the mesocosm experiment (< 1~10 pmol L⁻¹) in Kongsfjorden conducted by Hopkins et al. (2013). In addition, the mean CH₃I concentration in the LC treatment was similar to that measured in the East China Sea, with an average of 5.34 pmol L⁻¹ in winter and 5.74 pmol L⁻¹ in summer (Yuan et al., 2015). Meanwhile, a positive relationship was detected between CH₃I and Chl a, Phaeodactylum tricornuntum and Thalassiosira weissflogii (r = 0.588, p < 0.01 in the LC treatment; r = 0.834, p < 0.01 in the LC treatment for *Phaeodactylum tricornuntum*; r = 0.680 p < 0.010.01 in the LC treatment; r = 0.690, p < 0.01 in the HC treatment for *Thalassiosira weissflogii*; r =0.717, p < 0.01 in the LC treatment; r = 0.741, p < 0.01 in the HC treatment for Chl a). This result agrees with previous mesocosm (Hopkins et al., 2013) and laboratory experiments (Hughes et al.,

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2013; Manley and De La Cuesta, 1997) identifying diatoms as significant producers of CH₃I. Moreover, similar to DMS, the maximum CH₃I concentration also occurred after the maxima of Phaeodactylum tricornuntum and Thalassiosira weissflogii, at about 4 d (Fig. 3-D). This was similar to iodocarbon gases measured in a Norway mesocosm conducted by Hopkins et al. (2010) and chloroiodomethane (CH₂CII) concentrations measured in another Norway mesocosm conducted by Wingenter et al. (2007). Furthermore, the CH₃I concentrations measured in the HC treatment were significantly lower than those measured in the LC treatment during the mesocosm, which is in accord with the discoveries of Hopkins et al. (2010) and Webb et al. (2015) but in contrast to the findings of Hopkins et al. (2013) and Webb et al. (2016). Throughout the mesocosm experiment, there was a 40.2% reduction in the HC mesocosm compared to the LC mesocosm. Considering that the phytoplankton species did not show significant differences in the HC and LC treatments during the experiment, this reduction in the HC treatment was likely not caused by phytoplankton. Apart from direct biological production via methyl transferase enzyme activity by both phytoplankton and bacteria (Amachi et al., 2001), CH₃I is produced from the breakdown of higher molecular weight iodine-containing organic matter (Fenical, 1982) through photochemical reactions between organic matter and light (Richter and Wallace, 2004). Both bacterial methyl transferase enzyme activity and a photochemical reaction may have reduced the CH₃I concentrations in the HC treatment but further experiments are needed to verify this result.

4. Conclusions

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In this study, the effects of increased levels of pCO_2 on marine DMS(P) and halocarbons release were studied in a controlled mesocosm facility. A 28.2% reduction during the logarithmic growth phase and a 5 d delay in DMS concentration was observed in the HC treatment due to the effect of

elevated pCO_2 . Because the seawater in the mesocosm was filtered, the algae in the coastal environment and their attached bacteria were removed and the trace gases produced in the environment did not influence the mesocosm trace gas concentrations after the bags were sealed. Therefore, we attribute this phenomenon to the DMSP-consuming bacteria attached to *Phaeodactylum tricornuntum* and *Thalassiosira weissflogii*. More attention should be paid to the DMSP-consuming bacteria attached to algae under different pH values in future studies. Three bromocarbons compounds were not correlated with a range of biological parameters, as they were affected by the filtration procedure and elevated pCO_2 had no effect on any of the three bromocarbons. The temporal dynamics of CH_3I , combined with strong correlations with biological parameters, indicated biological control of the concentrations of this gas. In addition, the production of CH_3I was sensitive to pCO_2 , with a significant increase in CH_3I concentration at higher pCO_2 . However, without additional empirical measurements, it is unclear whether this decrease was caused by bacterial methyl transferase enzyme activity or by photochemical degradation at higher pCO_2 .

Author contribution: Gui-Peng Yang and Kun-Shan Gao designed the experiments. Sheng-Hui Zhang, Juan Yu and Qiong-Yao Ding carried out the experiments and prepared the manuscript.

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| 608 | Figure captions |
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| 609 | Fig. 1 Temporal changes of pH in the HC (1,000 µatm, solid squares) and LC (400 µatm, white |
| 610 | squares) mesocosms (3,000 L). Data are mean \pm standard deviation, n = 3 (triplicate independent |
| 611 | mesocosm bags) (Origin 8.0). |
| 612 | Fig. 2 Temporal changes in DMS, DMSP and DCB concentrations in the HC (1,000 µatm, black |
| 613 | squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm standard |
| 614 | deviation, $n = 3$ (triplicate independent mesocosm bags). |
| 615 | Fig. 3 Temporal changes in CHBrCl ₂ , CH ₃ Br, CH ₂ Br ₂ and CH ₃ I concentrations in the HC (1,000 |
| 616 | μ atm, black squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm |
| 617 | standard deviation, $n = 3$ (triplicate independent mesocosm bags). |
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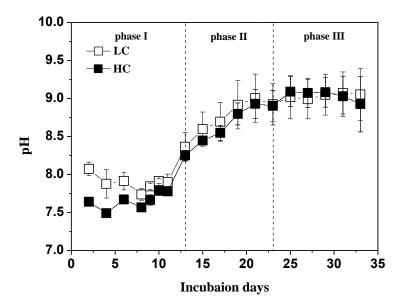
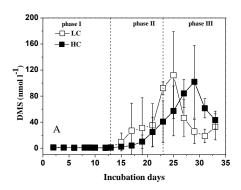
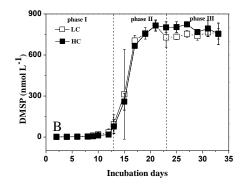


Fig. 1 Temporal changes of pH in the HC (1,000 μ atm, solid squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm standard deviation, n = 3 (triplicate independent mesocosm bags) (Origin 8.0).





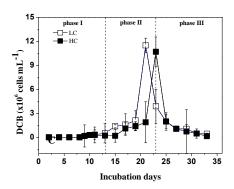


Fig. 2 Temporal changes in DMS (A), DMSP (B), DCB (C) concentrations in the HC (1,000 μ atm, black squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm standard deviation, n = 3 (triplicate independent mesocosm bags) (Origin 8.0).

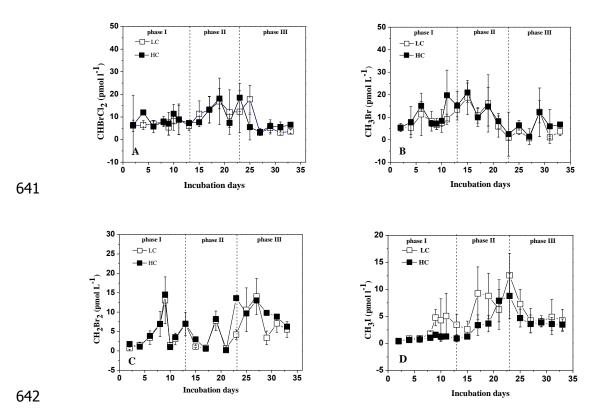


Fig. 3 Temporal changes in CHBrCl₂ (A), CH₃Br (B), CH₂Br₂ (C) and CH₃I (D) concentrations in the HC (1,000 μ atm, black squares) and LC (400 μ atm, white squares) mesocosms (3,000 L). Data are mean \pm standard deviation, n=3 (triplicate independent mesocosm bags) (Origin 8.0).

Table 1. The conditions of DIC, pH_T, pCO₂ and nutrient concentrations in the mesocosm experiments. "-" means that the values were below the detection limit.

| | | pH_T | DIC | pCO_2 | NO ₃ +NO ₂ | $\mathrm{NH_4}^+$ | PO_4^{3-} | SiO ₃ ²⁻ |
|-----------|----|---------|--------------------------|-----------|----------------------------------|----------------------|-------------------------|--------------------------------|
| | | | (µmol kg ⁻¹) | (µatm) | $(\mu mol \ L^{-1})$ | $(\mu mol \ L^{-1})$ | (µmol L ⁻¹) | $(\mu mol\;L^{\text{-}1})$ |
| day 0 | LC | 8.0±0.1 | 2181±29 | 1170~1284 | 52~56 | 19~23 | 2.6±0.2 | 38~40 |
| | НС | 7.5±0.1 | 2333±34 | 340~413 | 51~55 | 19~23 | 2.5±0.2 | 38~39 |
| PhaseI | LC | 7.9~8.4 | 1825~2178 | 373~888 | 15~52 | 1.6~20 | 0.5~2.6 | 31~38 |
| | НС | 7.4~8.2 | 2029~2338 | 1295~1396 | 47~54 | 0.2~21 | 0.7~2.5 | 34~39 |
| Phase II | LC | 8.4~8.5 | 1706~1745 | 46~749 | -~ 15.9 | - | 0.1~0.5 | 10~24 |
| | HC | 8.4~8.6 | 1740~1891 | 59~1164 | 1.1~25 | - | -~0.1 | 29~30 |
| Phase III | LC | 8.5~8.8 | 1673~1706 | 30~43 | - | - | - | 10~16 |
| | HC | 8.6~8.7 | 1616~1740 | 34~110 | - | - | -~0.3 | 24~25 |
| | | | | | | | | |

Table 2. Relationships between DMS, DMSP, Chl *a*, CHBrCl₂, CH₃Br, CH₂Br₂, CH₃I, DCB, *Thalassiosira weissflogii* (*T. weissflogii*) and *Phaeodactylum tricornutum* (*P. tricornutum*) concentrations in the LC treatments.

| | DMS | DMSP | Chl a | CHBrCl ₂ (pmol l ⁻¹) | CH ₃ Br | CH ₂ Br ₂ (pmol l ⁻¹) | CH ₃ I | DCB | T. weissflogii | P. tricornutum |
|---------------------|-------------------------|-------------------------|-----------------------|---|-------------------------|--|-------------------------|---------------------------------------|---------------------------------------|---------------------------|
| | (nmol L ⁻¹) | (nmol L ⁻¹) | (μg L ⁻¹) | (pmor r) | (pmol l ⁻¹) | (pmorr) | (pmol l ⁻¹) | $(\times 10^6 \text{ cells mL}^{-1})$ | $(\times 10^3 \text{ cells mL}^{-1})$ | (cells mL ⁻¹) |
| DMS | 1 | | | | | | | | | |
| DMSP | 0.701** | 1 | | | | | | | | |
| Chl a | 0.597** | 0.792** | 1 | | | | | | | |
| CHBrCl ₂ | 0.526 | 0.280 | 0.559 | 1 | | | | | | |
| CH ₃ Br | -0.413 | -0.230 | 0.196 | 0.313 | 1 | | | | | |
| CH_2Br_2 | 0.310 | 0.180 | 0.001 | -0.136 | -0.308 | 1 | | | | |
| CH ₃ I | 0.694** | 0.654** | 0.717** | 0.596* | -0.151 | 0.129 | 1 | | | |
| DCB | 0.643** | 0.520^{*} | 0.522* | 0.394 | -0.268 | -0.038 | 0.762** | 1 | | |
| T. weissflogii | 0.410 | 0.617** | 0.899** | 0.301 | 0.322 | 0.028 | 0.680** | 0.399 | 1 | |
| P. tricornutum | 0.560* | 0.961** | 0.821** | 0.528 | -0.032 | 0.162 | 0.588** | 0.334 | 0.685** | 1 |

Table 3. Relationships between DMS, DMSP, Chl *a*, CHBrCl₂, CH₃Br, CH₂Br₂, CH₃I, DCB, *Thalassiosira weissflogii* (*T. weissflogii*) and *Phaeodactylum tricornutum* (*P. tricornutum*) concentrations in the HC treatments.

| | DMS | DMSP | Chl a | CHBrCl ₂ (pmol l ⁻¹) | CH ₃ Br | CH ₂ Br ₂ (pmol l ⁻¹) | CH ₃ I | DCB | T. weissflogii | P. tricornutum |
|---------------------|----------------------------|-------------------------|-----------------------|---|-------------------------|--|---------------------------------------|---------------------------------------|---------------------------|----------------|
| | (nmol L ⁻¹) (r | (nmol L ⁻¹) | (μg L ⁻¹) | (pillor r) | (pmol l ⁻¹) | (pmol l ⁻¹) | $(\times 10^6 \text{ cells mL}^{-1})$ | $(\times 10^3 \text{ cells mL}^{-1})$ | (cells mL ⁻¹) | |
| DMS | 1 | | | | | | | | | |
| DMSP | 0.752** | 1 | | | | | | | | |
| Chl a | 0.318* | 0.738** | 1 | | | | | | | |
| CHBrCl ₂ | 0.324 | 0.094 | 0.326 | 1 | | | | | | |
| CH ₃ Br | -0.410 | -0.349 | 0.065 | 0.076 | 1 | | | | | |
| CH_2Br_2 | 0.540* | 0.352 | 0.142 | 0.233 | -0.377 | 1 | | | | |
| CH ₃ I | 0.694** | 0.816** | 0.741** | 0.690* | -0.407 | 0.316 | 1 | | | |
| DCB | 0.544* | 0.522 | 0.549* | 0.532 | -0.311 | 0.368 | 0.851* | 1 | | |
| T. weissflogii | 0.355 | 0.743** | 0.930** | 0.304 | 0.076 | 0.233 | 0.690** | 0.567 | 1 | |
| P. tricornutum | 0.635** | 0.954** | 0.803** | 0.143 | -0.257 | 0.267 | 0.834** | 0.559 | 0.820** | 1 |