# Quantification of lignin oxidation products as vegetation biomarkers in speleothems and cave drip water

Inken Heidke[1], Denis Scholz[2], and Thorsten Hoffmann[1]

[1]Institute of Inorganic Chemistry and Analytical Chemistry, Johannes Gutenberg University of Mainz, Duesbergweg 10-14, 55128 Mainz, Germany
[2]Institute of Geosciences, Johannes Gutenberg University of Mainz, J.-J.-Becher-Weg 21, 55128 Mainz, Germany

**Correspondence:** Thorsten Hoffmann (t.hoffmann@uni-mainz.de)

**Abstract.** Here we present a sensitive method to analyze lignin oxidation products (LOPs) in speleothems and cave drip water to provide a new tool for paleo-vegetation reconstruction. Speleothems are valuable climate archives. However, compared to other terrestrial climate archives, such as lake sediments, speleothems contain very little organic matter. Therefore, very few studies on organic biomarkers in speleothems are available. Our new sensitive method allows to use LOPs as vegetation

5 biomarkers in speleothems.

Our method consists of acid digestion of the speleothem sample followed by solid phase extraction (SPE) of the organic matter. The extracted polymeric lignin is degraded in a microwave assisted alkaline CuO oxidation step to yield monomeric LOPs. The LOPs are extracted via SPE and finally analyzed via ultrahigh-performance liquid chromatography (UHPLC) coupled to electrospray ionization (ESI) and high-resolution orbitrap mass spectrometry (HRMS). The method was applied

10 to stalagmite samples with a sample size of 3–5 g and cave drip water samples with a sample size of 100–200 mL from the Herbstlabyrinth-Advent-Cave in Germany. In addition, fresh plant samples, soil water and powdered lignin samples were analysed for comparison. The concentration of the sum of eight LOPs ($\Sigma 8$) was in the range of 20–84 $\mathrm{ng \cdot g^{-1}}$ for the stalagmite samples and 230–440 $\mathrm{ng \cdot L^{-1}}$ for the cave drip water samples. The limits of quantification for the individual LOPs ranged from 0.3–8.2 ng per sample or 1.5–41.0 $\mathrm{ng \cdot mL^{-1}}$ of the final sample solution.

15 Our method represents a new and powerful analytical tool for paleo-vegetation studies and has great potential to identify the pathways of lignin incorporation into speleothems.

## 1 Introduction

Speleothems are calcareous mineral deposits that form within caves in karstified carbonate rock. The most common types of speleothems are stalagmites, which are formed by water dripping on the ground of the cave, stalactites, which are their coun-

20 terparts on the cave ceiling, and flowstones, which are formed by water films flowing on the cave walls and floor. Speleothems preserve information about climatic and hydrological conditions and the vegetation development above the cave and therefore serve as paleoclimate archives (Fairchild and Baker, 2012; McDermott, 2004). Compared to other paleo-climate archives, such as ice cores and marine or lacustrine sediments, speleothems have certain advantages. They can grow continuously for $10^3$–$10^5$ years, their growth layers are mechanically undisturbed and they do not show a loss of time resolution with increasing

age (Gałuszka et al., 2017; Fairchild et al., 2006). They can be accurately dated up to $500\,000$ years back in time using the $^{230}$Th/U-method (Scholz and Hoffmann, 2008; Richards and Dorale, 2003). Moreover, they occur on all continents except Antarctica and are thus not limited to certain climatic regions.

Most studies of speleothems focus on the analysis of stable isotope ratios ($\delta^{13}$C, $\delta^{18}$O (McDermott, 2004) and inorganic trace elements (Fairchild and Treble, 2009). The organic content of speleothems has so far mostly been analyzed as total organic carbon content or fluorescent organic matter (Quiers et al., 2015). However, in recent years, the interest in molecular organic proxies in climate archives has increased (Giorio et al., 2018; Blyth et al., 2008; Blyth and Watson, 2009; Blyth et al., 2010, 2016). In speleothems, in particular lipid biomarkers, such as fatty acids reflecting changes in vegetational and microbial activities (Xie, 2003; Blyth et al., 2006; Bosle et al., 2014) and membrane lipids (glycerol dialkyl glycerol tetraethers, GDGTs) serving as paleo temperature proxies (Blyth and Schouten, 2013; Baker et al., 2016), have been studied.

Lignin occurs almost exclusively in terrestrial vascular plants and is one of the main constituents of wood and woody plants (Jex et al., 2014). It is a biopolymer that mainly consists of three monomers: sinapyl alcohol, coniferyl alcohol and p-coumaryl alcohol. The proportion of these three monomers varies with the type of plant, such as gymnosperm or angiosperm and woody or non-woody material. Thus, by analyzing the composition of lignin, it is possible to determine the source and type of plant material.

Lignin has been widely used as paleo-vegetation proxy in lake sediment (Tareq et al., 2011) and peat cores (Tareq et al., 2004). In marine sediments (e.g., Zhang et al., 2013) and natural waters (Standley and Kaplan, 1998; Hernes and Benner, 2002), lignin analysis has been used to determine the source of dissolved organic matter. Blyth and Watson (2009) have successfully detected lignin pyrolysis products in speleothems by applying a tetramethylammonium hydroxide (TMAH) thermochemolysis method, but there have been no quantitative studies of lignin in speleothems yet.

Lignin has to be degraded before the molecular composition of its phenolic components can be analyzed. The most common method for degradation of lignin is the alkaline oxidation with cupric oxide (CuO), developed by Hedges and Parker in 1976. This method releases a number of phenolic acids, aldehydes and ketones, which can be divided into four groups: The vanillyl group (V) consisting of vanillic acid, vanillin and acetovanillone, the syringyl group (S) consisting of syringic acid, syringaldehyde and acetosyringone, the cinnamyl group (C) consisting of trans-ferulic acid and p-coumaric acid, and the p-hydroxyl group (P) consisting of p-hydroxybenzoic acid, p-hydroxybenzaldehyde and p-hydroxyacetophenone. Hedges and Mann (1979) analysed fresh plant tissues and showed that the phenols of the syringyl group are only obtained from angiosperm, but not from gymnosperm plant tissues. Likewise, the phenols of the cinnamyl group are only obtained from non-woody and not from woody plant tissues, whereas the phenols of the vanillyl group are found in all kind of vascular plant tissues (angiosperm and gymnosperm, woody and non-woody). These results led to the introduction of the lignin oxidation product (LOP) parameters C/V and S/V, where C, for example, is defined as the sum of all lignin oxidation products of the C-group (Hedges and Mann, 1979). The phenols of the p-hydroxyl group can originate from gymnosperm and non-woody angiosperm plant tissues, but are also oxidation products of protein rich organisms such as bacteria and plankton. Therefore, the P group is not used in the parameters to determine the source of lignin (Jex et al., 2014). The parameter $\Sigma 8$ gives the sum of the eight analytes of the C, S and V-group and is used to estimate the total amount of LOPs in a sample.

The oxidation with CuO has been optimized many times in the past. For example, Goñi and Montgomery (2000) developed a microwave digestion method. Other groups improved the sample clean-up by replacing the formerly used liquid-liquid extraction (LLE) with solid phase extraction (SPE) (Kögel and Bochter, 1985; Kaiser and Benner, 2012). As the CuO oxidation method is broadly used, there are many data sets to compare with. This is certainly an advantage compared to the above mentioned TMAH thermochemolysis method, which is less often used and produces more complex methylated reaction product mixtures (Wysocki et al., 2008). For the detection of the LOPs, gas chromatography coupled to mass spectrometry (GC-MS) is often used, which requires a derivatization step. Liquid chromatography is also used, either in combination with UV detection or coupled to mass spectrometry.

The purpose of this study was to develop and validate a sensitive and selective method for the quantification of LOPs in both speleothem and cave drip water samples using liquid chromatography electrospray ionization mass spectrometry (LC-ESI-MS). This method offers new possibilities for paleo-vegetation reconstruction since it combines the advantages of lignin analysis as a highly specific vegetation biomarker with the above mentioned benefits of speleothems as unique terrestrial climate archives. Lignin as a vegetation biomarker is much more specific for higher plants than for example n-alkanes or fatty acids (Jex et al., 2014), and thus can help to interpret other vegetation markers and stable isotope records. Up to now, lignin analysis for paleo-vegetation reconstruction has only been applied to lake sediments and peat cores, which contain much larger amounts of organic matter than speleothems. Our method allows to analyse the lignin composition of trace amounts of organic matter preserved in speleothems. The stalagmite samples are first acid digested, and the acidic solution is then extracted by SPE. The eluent is then subjected to CuO oxidation in a microwave assisted digestion method. The oxidised sample solutions are again extracted and enriched by SPE, and the LOPs are then separated and detected by ultrahigh-performance liquid chromatography coupled to electrospray ionization high-resolution mass spectrometry (UHPLC-ESI-HRMS).

## 2 Experimental section

### 2.1 Chemicals and materials

Analytical standards of acetosyringone (97%), acetovanillone ($\geq$ 98%), para-coumaric acid ($\geq$ 98%), ethylvanillin (99%), ferulic acid (99%), para-hydroxyacetophenone ($\geq$ 98%), para-hydroxybenzaldehyde ($\geq$ 97.5%), syringaldehyde (98%), syringic acid (> 95%) and cinnamic acid (97%) as well as copper(II) oxide (> 99%) and ammonium iron(II) sulfate (99%) were purchased from Sigma Aldrich. Analytical standards of para-hydroxybenzoic acid (99%) and vanillin (99%) were obtained from Acros Organics, an analytical standard of vanillic acid (98%) was obtained from Alfa Aesar. Sodium hydroxide (pellets, $\geq$ 99%) was purchased from Carl Roth, hydrochloric acid (HCl, suprapure, 30%) from Merck KGaA. Lignin from mainly coniferous wood was obtained from BASF SE. Mixed lignin from wheat straw and various kinds of wood was purchased from Bonding Chemical. Solid phase extraction columns (Oasis HLB, 3 mL tubes, 60 mg packing material) were purchased from Waters. Ultrapure solvents (Optima LC/MS grade) acetonitrile (ACN), water ($H_2O$) and methanol (MeOH) were obtained from Fisher Scientific. Dichloromethane (DCM) ($\geq$ 99.9% (GC)) was obtained from Honeywell Riedel-de Haën. Ultrapure water with 18.2 MΩ resistance was produced using a Milli-Q water system from Merck Millipore (Darmstadt, Germany).
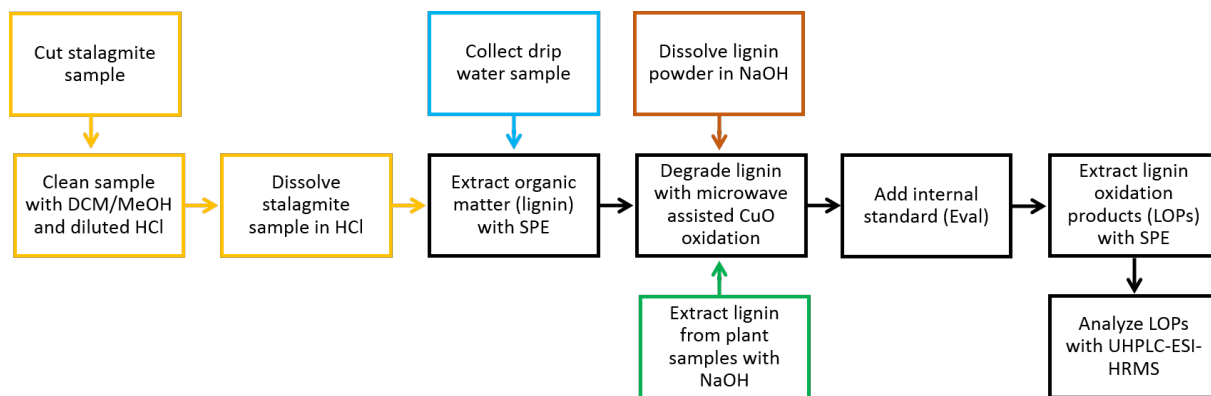
**Figure 1.** Process chart of the overall sample preparation procedure. A detailed description of the individual steps is given in section 2.2.

## 2.2 Methods

The overall sample preparation procedure is shown as process chart in Figure 1. The different steps of the sample preparation will be described in detail in the following paragraphs.

### 2.2.1 Preparation of standards

Stock solutions of all analytical standards were prepared at a concentration of 1 mg·mL$^{-1}$ in ACN. A mixed stock solution of all analytical standards was prepared by dilution of the individual stock solutions to a concentration of 10 µg·mL$^{-1}$ in ACN. The stock solutions were stored at -18 °C. For the external calibration standards, the mixed stock solution was freshly diluted to the appropriate concentrations ranging from 20 ng·mL$^{-1}$ to 2000 ng·mL$^{-1}$ in H2O/ACN 9:1 (v/v). To optimize the SPE procedure for the LOPs, 100 µL of a 1 µg·mL$^{-1}$ mixed standard solution in H$_2$O/ACN 9:1 (v/v) was added to 20 mL of a 2 mol·L$^{-1}$ sodium chloride solution that was acidified to pH 2 with HCl (30%) to simulate the sample solution after the microwave digestion step.

### 2.2.2 Sampling and preparation of stalagmite samples

Stalagmite *NG01* from the Herbstlabyrinth-Advent-Cave, central Germany, was 50 cm long and had a diameter of approximately 15 cm. It was cut along the growth axis using a diamond blade saw. From one of the two halves, a 1 cm thick slab was cut, which was then dated using the $^{230}$Th/U-method (Mischel et al., 2016). This showed that the oldest part of the stalagmite grew at ca. 11 000 years BP, whereas the youngest part stems from recent time. Thus, the stalagmite covers the Holocene. The inner part of the stalagmite slab, close to the growth axis, was already used for stable isotope and trace element (Mischel et al., 2016, 2017) as well as fatty acid analysis (Bosle et al., 2014). Thus, the samples for this study had to be taken from the outer part of one half of the stalagmite slab. Pieces of calcite with approximately 0.5–1.2 cm in width, 2.5–3.7 cm in length and a

weight of 3.0–5.4 g were cut from the slab using a diamond wire saw following the growth lines of the stalagmite. Care was taken to always leave 2 cm space to the outer surface of the stalagmite to avoid contamination and dating problems.

To clean the stalagmite samples, each sample was covered with DCM/MeOH 9:1 (v/v) and cleaned for 10 min at 35 °C in an ultrasonic bath. The solvent was discarded, and the cleaning was repeated a second time. Afterwards, the samples were rinsed with ultrapure water, then each sample was covered with ultrapure water, and 250 μL of HCl (30%) were added to etch away the outermost layer of calcite, which might be contaminated. After 5 min, the samples were rinsed with ultrapure water, dried and weighed. The samples were then placed in clean glass vials and 2.1 mL of HCl (30%) per gram stalagmite were added to dissolve the calcite over night at room temperature. Before extracting the solutions using SPE, they were diluted 1:1 with ultrapure water to prevent clogging of the cartridges.

### 2.2.3   Sampling and preparation of drip water samples

The drip water samples were collected in the framework of a monthly cave monitoring program (Mischel et al., 2016, 2015). All samples presented here were sampled in October 2014 at different drip sites (two fast drip sites, *D1* and *D5*, with a drip rate of 0.3–0.5 drops·s$^{-1}$, one slow drip site, *D2*, with a drip rate of approx. 60 mL·month$^{-1}$, and one sample from a cave pool, *PW*). In addition, soil water (*SW*) was sampled in a meadow above the cave, and rain water (*RW*) was sampled at a weather station above the cave. More information on the sampling techniques can be found in (Mischel et al., 2016, 2015). The samples were collected in pre-cleaned glass vessels. To prevent the growth of microorganisms, 5% (w/w) of acetonitrile were added shortly after sampling. The samples were then stored at 4 °C in the dark for several months. Before extracting the samples using SPE, they were acidified to pH 1–2 with HCl.

### 2.2.4   Preparation of lignin and fresh plant tissue samples

The lignin powder was dissolved in NaOH (2 mol·L$^{-1}$) at a concentration of 1 mg·mL$^{-1}$. 100 μL of this solution were added into the microwave reaction vessel. The plant samples (leaves and branches of Amur maple, and needles and branches of European yew, all collected in Mainz, Germany) were cut in small pieces and dried in an oven at 50 °C for two days. 10 mg·mL$^{-1}$ were soaked in NaOH (2 mol·L$^{-1}$) for several days. 1 mL of this solution was filtered over 1 $\mu$m filters and added into the microwave reaction vessels.

### 2.2.5   Solid phase extraction of organic matter in dissolved stalagmite solution and drip water samples

The SPE cartridges were preconditioned with 3 mL of MeOH followed by 3 mL of ultrapure water, which was acidified to pH 1–2 with HCl. The diluted stalagmite solution or the acidified drip water sample were loaded onto the cartridges using sample reservoirs. The drip rate was always below 1 drop·s$^{-1}$. The cartridges were washed twice with 3 mL of acidified ultrapure water and dried for 20 min by sucking ambient air through the cartridges using a vacuum manifold. The lignin was eluted with 6 portions of 250 μL of MeOH. The solvent was evaporated to almost dryness under a gentle stream of nitrogen at 30 °C. The residue was re-dissolved in 1.5 mL of NaOH (2 mol·L$^{-1}$), the solution was sonicated for 10 min at 45 °C and

added into the microwave reaction vessel. The sample vial was sonicated again with 1.5 mL of NaOH (2 mol·L$^{-1}$) and this solution was added into the microwave reaction vessel, too.

### 2.2.6 Microwave assisted CuO oxidation

The microwave assisted CuO oxidation procedure was performed according to the method described by Goñi and Montgomery
(2000) with slight modifications. An Ethos Plus Microwave Labstation (MLS GmbH, Germany) was used with an HPR-1000/10S high pressure segment rotor, which can hold up to 10 reaction vessels, and an ATC-CE temperature sensor to measure the temperature inside one reaction vessel. 100 mL Teflon vessels were used as reaction vessels. Each vessel was loaded with 250 mg of CuO, 50 mg of (NH$_4$)$_2$Fe(SO$_4$)$_2$·6 H$_2$O and 8 mL of NaOH (2 mol·L$^{-1}$) in total, including the sample solution. The NaOH solution was purged with nitrogen for 30 min before use to remove dissolved oxygen, which could lead to over-
oxidation of the lignin oxidation products. For the same reason, the vessels were purged with an argon flow of 1 mL·min$^{-1}$ for 1 min and then quickly capped to ensure an inert gas atmosphere in the vessels. The vessels were shaken well and then placed in the high-pressure segment rotor of the microwave oven. The temperature was increased to 155 °C in 5 min and then hold at 155 °C for 90 min. Afterwards, the vessels were allowed to cool down to room temperature overnight. Directly after opening the vessels, 50 μL of a 1 μg·mL$^{-1}$ standard solution of ethyl vanillin in H$_2$O/ACN (9:1, v/v) were added as an internal
standard into each vessel except the blank sample. The reaction solutions were transferred to 15 mL centrifuge tubes and the reaction vessels were rinsed twice with 3 mL of NaOH (2 mol·L$^{-1}$). The combined solutions were centrifuged for 10 min at 3000 rpm and the supernatant was decanted into glass vessels. The residue was suspended in 5 mL of NaOH (2 mol·L$^{-1}$) using a vortex mixer, centrifuged again for 10 min at 3000 rpm and the supernatant was combined with the sample solution.

### 2.2.7 Solid phase extraction of LOPs in the oxidized sample solution

The oxidized sample solution was acidified to pH 1–2 with HCl. The conditioning, loading, washing and drying steps of the SPE cartridges were the same as described in section 2.2.5. The LOPs were eluted with four portions of 250 μL of ACN with 2% of ammonia added to reach a basic pH of 8–9. The eluate was evaporated to dryness in a gentle stream of nitrogen at 30 °C and the residue was re-dissolved in 200 μL H$_2$O/ACN (9:1).

### 2.2.8 UHPLC-ESI-HRMS analysis

The analysis of the lignin oxidation products was carried out on a Dionex Ultimate 3000 ultrahigh-performance liquid chromatography system (UHPLC) that was coupled to a heated electrospray ionization source (ESI) and a Q-Exactive Orbitrap high-resolution mass spectrometer (HRMS) (all by Thermo Fisher Scientific). To separate the LOPs, a Hypersil Gold pentafluorophenyl (PFP) column, 50 mm x 2.1 mm with 1.9 μm particle size (also by Thermo Fisher Scientific) was used. The injection volume was 15 μL. A H$_2$O/ACN-gradient program was applied. The gradient started with 10% eluent B (consisting of 98% ACN and 2% H$_2$O) and 90% eluent A (consisting of 98% H$_2$O, 2% ACN and 400 μL·L$^{-1}$ formic acid), which was held for

**Table 1.** Names and abbreviations of the analytes with the respective $m/z$ values of their deprotonated molecular ions [M-H]$^-$ and their specific daughter ions.

| name of analyte | abbreviation | $m/z$ of [M-H]$^-$ | $m/z$ of specific daughter ion (lost neutral fragment) |
|---|---|---|---|
| p-hydroxybenzoic acid | p-Hac | 137.02441 | 93.03455 (-$CO_2$) |
| p-hydroxybenzaldehyde | p-Hal | 121.02943 | 121.02943 (no loss) |
| p-hydroxyacetophenone | p-Hon | 135.04517 | 135.04517 (no loss) |
| vanillic acid | Vac | 167.03498 | 152.01151 (-$CH_3$) |
| vanillin | Val | 151.04007 | 136.01657 (-$CH_3$) |
| acetovanillone | Von | 165.05572 | 150.03220 (-$CH_3$) |
| ethylvanillin (internal standard) | Eval | 165.04518 | 136.01659 (-$CH_2CH_3$) |
| syringic acid | Sac | 197.04555 | 182.02234 (-$CH_3$) |
| syringaldehyde | Sal | 181.05063 | 166.02708 (-$CH_3$) |
| acetosyringone | Son | 195.06628 | 180.04292 (-$CH_3$) |
| trans-ferulic acid | t-Fac | 193.05063 | 134.03734 (-$CH_3$, -$CO_2$) |
| p-coumaric acid | p-Cac | 163.04007 | 119.05024 (-$CO_2$) |
| trans-cinnamic acid (internal standard) | t-Ciac | 147.04520 | 147.04520 (no loss) |

0.5 min. Eluent B was increased to 12% within 2 min, held for 1 min, was further increased to 50% within 1.25 min, held for 0.75 min and increased to 99%. This composition was held for 2 min, then eluent B was decreased to the initial value of 10%.

The ESI source was operated in negative mode, so that deprotonated molecular ions [M-H]$^-$ were formed. The spray voltage was -4.0 kV, the ESI probe was heated to 150 °C to improve the evaporation of the aqueous solvent, the capillary temperature was 350 °C, the sheath gas pressure was 60 psi and the auxiliary gas pressure was 20 psi.

The mass spectrometer was operated in full scan mode with a resolution of 35 000 and a scan range of $m/z$ 80–500. At the respective retention time windows, the full scan mode was alternated with a targeted MS$^2$-mode with a resolution of 17 500 to identify the LOPs by their specific daughter ions, see Table 1. For the MS$^2$-mode (i.e., *parallel reaction monitoring mode* in the software *XCalibur*, provided by Thermo Fisher Scientific), higher-energy collisional dissociation (HCD) was used with 35% normalized collision energy (NCE) for all analytes. The actual collision energy was calculated by the software on basis of mass and charge of the selected precursor ions and was in the range of 10–14 eV.
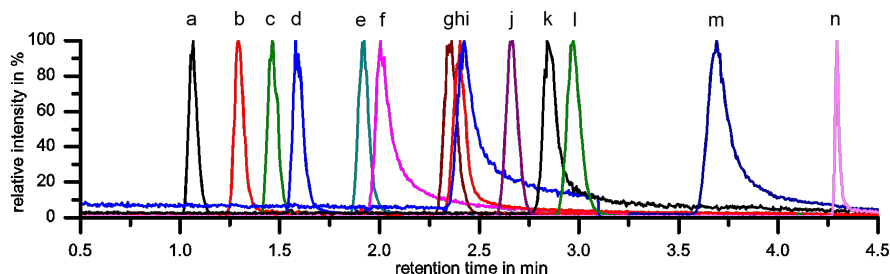
**Figure 2.** Normalized chromatogram of 14 LOP standards on a PFP column. Explanation of the peak numbers (for abbreviations see Table 1): (a) p-Hac, (b) Vac, (c) Sac, (d) p-Hal, (e) p-Hon, (f) Val, (g) p-Cac, (h) Sal, (i) Von, (j) c-Fac, (k) Son, (l) t-Fac, (m) EVal, (n) t-Ciac.

## 3   Results and Discussion

### 3.1   Method development

#### 3.1.1   Separation of LOPs with LC gradient elution and identification of LOPs with MS/MS-experiments

A sufficient separation of the eleven LOPs and two internal standards was achieved within 4.5 min on a PFP column with
H$_2$O/ACN-gradient elution, as can be seen in Fig. 2, which shows the normalized chromatogram of 14 LOP standards. The
analytes were identified via the exact mass of their molecular ion, their retention time compared to standards and their frag-
mentation pattern in the MS$^2$ spectrum. As the chromatograms of the real samples were very complex, all three methods were
indeed required to identify and quantify the analytes. Whenever possible, the quantification was done by integrating the chro-
matographic peak of the molecular ion. However, when the target analyte peak could not be baseline separated from another
signal, the chromatographic peak of a specific daughter ion was used to quantify the analyte.

#### 3.1.2   Optimisation of the solid phase extraction procedure for LOPs

Two different types of SPE-cartridges were tested. The polymer-based Oasis HLB cartridges (hydrophilic lipophilic balanced
polymer, Waters) showed better reproducibility and equal recovery values compared to the silica-based Supelco C18 cartridges
(Sigma Aldrich). The recovery rates could be improved by adding ammonia to the elution solvent, ACN or MeOH, as can be
seen in Figure 2. The basic pH value of the eluent leads to deprotonation of the phenolic hydroxyl group. In this ionic state,
the analytes are better soluble in the polar mobile phase and their adsorption to the stationary phase is weakened. Since we
observed an oxidation of aldehydes and an isomerization of p-coumaric acid and ferulic acid when MeOH was used as elution
solvent – an observation that has been made before (Lima et al., 2007) – ACN with ammonia was used as elution solvent. The
recovery rates ranged from 69% to 101% and are shown in Fig. 3 and Table 2.

Ethyl acetate was tested as elution solvent, too, as used by Kögel and Bochter; however, the recovery rates were lower than
with methanol or acetonitrile. In addition, it was observed that with ethyl acetate, aldehydes were lost in the evaporation step
(Fig. A1 in the supplementary information (SI) available online). The SPE method was tested with spikes of LOP standards of
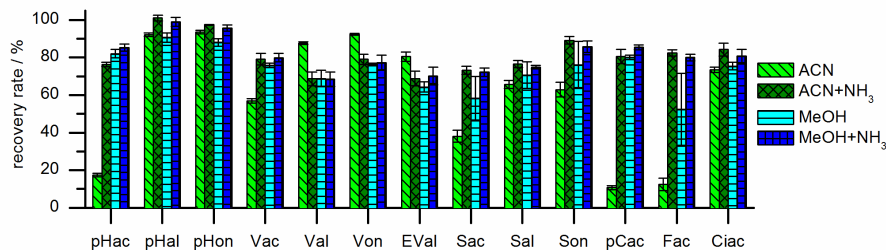
**Figure 3.** Recovery rates of the 13 LOPs on oasis HLB SPE cartridges, eluted with acetonitrile (ACN), acetonitrile with ammonia (ACN+NH$_3$), methanol (MeOH) and methanol with Ammonia (MeOH+NH$_3$). The recovery rates improved significantly if ammonia was added to the elution solvent.

different concentrations reaching from 25 ng to 1000 ng in 20 mL of surrogate sample solution (i.e. $1.25$–$50 \, \mathrm{ng \cdot mL^{-1}}$ in the surrogate sample solution or $125$–$5000 \, \mathrm{ng \cdot mL^{-1}}$ in the final sample solution injected into the LC-MS system). The recovery rate was constant at all concentration levels and the linearity was very good ($R^2 > 0.9990$) for all analytes (Fig. A2 and A3 in the SI).

### 3.1.3   Comparison of different durations and temperatures of the CuO oxidation method

In former studies, the duration of the CuO oxidation method varied between 90 min and 180 min and temperatures of 150 °C or 170 °C have been applied. Therefore, we compared temperatures of 155 °C and 175 °C (the temperature of the microwave program was chosen 5 °C higher than the desired temperature in the Teflon vessels) and durations of 90 min and 180 min, using 100 µg of mixed lignin as standard sample and three subsamples for each constellation. The results are shown in Fig. 4. At a temperature of 175 °C and a duration of 180 min, the concentrations of almost all LOPs were dramatically diminished, probably due to overoxidation. For Val, Von, Sal, pCac and Fac, the highest concentrations were reached with 155 °C and 90 min, every increase in temperature or duration of the oxidation step resulted in a loss of analyte. In consequence, the C/V ratio decreased from 0.037 for 155 °C, 90 min to 0.018 for increased temperature, to 0.014 for increased duration and to 0.009 if both were increased. Similarly, the Vac/Val ratio increased from 0.44 for 155 °C, 90 min to 0.83 for increased temperature and to 0.54 for increased duration. For the Sac/Sal ratio, the increase was from 0.16 for 155 °C, 90 min to 0.37 and 0.28, respectively. These results show, that especially the C-group LOPs, pCac and Fac, as well as the aldehydes Val and Sal and the ketone Von are prone to overoxidation. Therefore, care should be taken to adjust temperature and duration of the CuO oxidation step to avoid overoxidation of the LOPs, otherwise the lignin oxidation parameters C/V, S/V and acid/aldehyde ratios will be distorted. The prevention of overoxidation by the addition of glucose was also tested; however, this did not improve the analysis (see Fig. A4 and Fig. A5 in the SI).
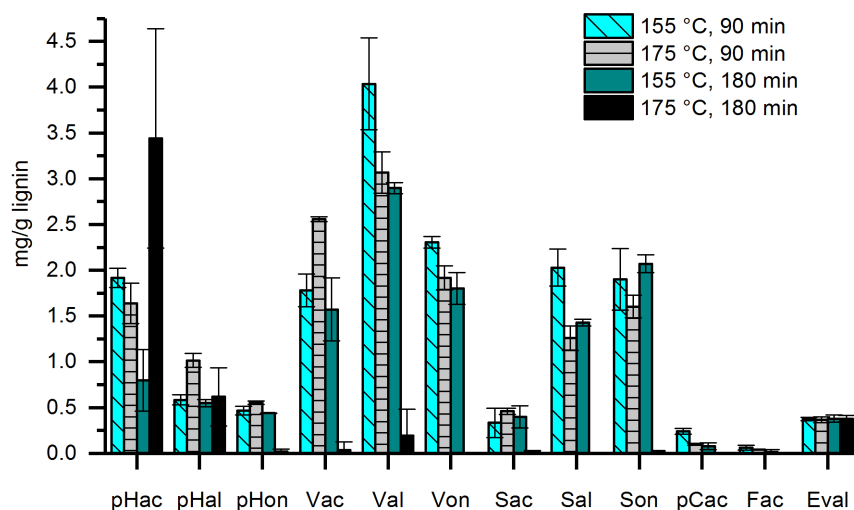
**Figure 4.** Results of a CuO oxidation step at 155 °C and 90 min (light blue bars with diagonal stripes), 175 °C and 90 min (grey bars with horizontal stripes), 155 °C and 180 min (dark cyan bars) and 175 °C and 180 min (black bars). Eval was added after the CuO oxidation step as internal standard

### 3.1.4 Comparison of two sample preparation methods – acid digestion of the stalagmite samples and direct CuO oxidation of stalagmite powder

Obviously, each individual step in the analytical sample preparation method includes the risk of positive or negative artifacts, especially if large amounts of chemicals are added. Therefore, experiments were performed to test whether the HCl dissolving step can be skipped by grinding the stalagmite sample and directly adding the powder into the microwave reaction vessels. 24 g of cleaned stalagmite sample were coarsely crushed and mixed. 12 g of this sample mixture were dissolved in HCl and extracted via SPE as described above. The solution was then divided into three subsamples. The other 12 g were finely ground in a mortar, divided into three subsamples and added directly into the microwave reaction vessels. Figure 5 shows that the LOP concentrations found in the acid digested samples were higher for most analytes than in the ground samples. An explanation for this finding might be that at least a part of the lignin particles is bound in the calcite crystals and is only fully released in the acid digestion method. Blyth et al. (2006) already stated similar findings for lipid biomarkers. Consequently, the acid digestion step is essential for the analysis of the target analytes in speleothems.
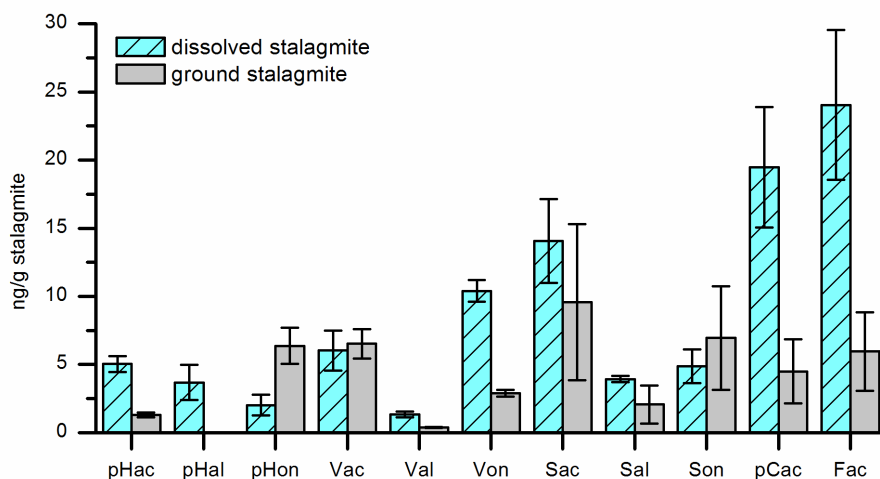
**Figure 5.** Results of dissolved stalagmite samples (light blue bars with diagonal stripes) compared to ground stalagmite samples (grey bars). Dissolving of the samples lead to higher amounts of LOPs.

## 3.2 Method validation and Quality assurance

### 3.2.1 Selectivity

The selectivity of the method was assured by using three parameters for peak identification: the retention time, the exact $m/z$ ratio of the analyte, and the MS$^2$-spectra, as described in section 3.1.1. The variation in the retention time was $\pm$ 0.01 min. To assure that the measured peak area was caused only by the analyte, the corrensponding peak area of the reagent blank measurement was subtracted.

### 3.2.2 Calibration and linearity

External calibration with a standard mixture containing all analytes was performed. The calibration function was obtained using the linear regression method. The parameters of the individual calibration functions are shown in Table A1 in the SI. The concentrations of the standards ranged from 20–500 $\mathrm{ng \cdot mL^{-1}}$ for stalagmite and drip water samples and up to 2000 $\mathrm{ng \cdot mL^{-1}}$ for plant and lignin samples. The calibration was linear in this range.

### 3.2.3 Limits of detection and quantification and reagent blanks

The instrumental limits of detection (LOD) and quantification (LOQ) were calculated by using equations (1) and (2), with $\sigma_0 =$ standard deviation of the peak area of the solvent blank, or, if no signal was detectable for the solvent blank, of the lowest calibration standard, and the slope of the calibration function, $m$. The results are shown in Table A1 in the SI.

5    instrumental limit of detection $$\text{LOD} = \frac{3.3 \cdot \sigma_0}{m} \tag{1}$$

instrumental limit of quantification $$\text{LOQ} = \frac{10 \cdot \sigma_0}{m} \tag{2}$$

To eliminate the influence of possible contamination sources on the results, a reagent blank, which had undergone all sample preparations steps, was analyzed with every batch of samples. The concentrations of LOPs measured in this reagent blank were

10   subtracted from the concentrations measured in the samples. The mean values of six reagent blanks measured on different days are shown in Table 2 (the concentrations refer to the final sample solution injected into the LC-MS system). The values ranged from $1.0 \text{ ng} \cdot \text{mL}^{-1}$ to $680 \text{ ng} \cdot \text{mL}^{-1}$, depending on the analyte (see also 3.2.4). The blank value varied from batch to batch, which is reflected in the standard deviations of the blank values given in Table 2. Therefore, the method detection limit (MDL) and the method quantification limit (MQL) were calculated using only the standard deviation of the peak area of the reagent

15   blank, as shown in equations (3) and (4), with $\sigma_B =$ standard deviation of the peak area of the reagent blank and $m =$ slope of the calibration function. The MDL was below $13.7 \text{ ng} \cdot \text{mL}^{-1}$ for all relevant analytes and the MQL was below $41.5 \text{ng} \cdot \text{mL}^{-1}$ for all relevant analytes.

method detection limit $$\text{MDL} = \frac{3.3 \cdot \sigma_B}{m} \tag{3}$$

20   method quantification limit $$\text{MQL} = \frac{10 \cdot \sigma_B}{m} \tag{4}$$

### 3.2.4 Origin of blank values

The blank values shown in Table 2 reflect the natural occurrence of the different analytes. The highest blank values have been found for the p-hydroxy group, p-coumaric acid, cinnamic acid, vanillin and vanillic acid. The p-hydroxy group is known to originate not only from lignin, but also from protein rich material such as bacteria (Jex et al., 2014). For p-hydroxy acetophe-

25   none, which has a lower blank value than p-hydroxy benzoic acid and p-hydroxy benzaldehyde, it is in discussion whether it originates from lignin or from other sources (Dittmar and Lara, 2001). P-coumaric acid occurs in sporopollenin (Fraser et al.; Montgomery et al., 2016), which is a major component of pollen and fungal spores and also occurs in some form of

**Table 2.** Method detection limit after subtraction of the reagent blank (MDL) in $ng \cdot mL^{-1}$, method quantification limit after subtraction of the reagent blank (MQL) in $ng \cdot mL^{-1}$, mean value of three subsamples of 3.4 g stalagmite after blank subtraction in $ng \cdot mL^{-1}$ and in $ng \cdot g^{-1}$ of the initial stalagmite sample, mean blank value of six reagent blanks measured on different days in $ng \cdot mL^{-1}$, and recovery values of the SPE procedure to extract LOPs (Recov. SPE) in %. All concentrations in $ng \cdot mL^{-1}$ refer to the final sample solution injected into the LC-MS system. The errors stated in this table are standard deviations of $n$ samples. For the methods of calculation used please refer to the text. The abbreviations for the analytes are shown in Table 1.

| analyte | MDL / $ng \cdot mL^{-1}$ | MQL / $ng \cdot mL^{-1}$ | Mean stalagmite / $ng \cdot mL^{-1}$ (n=3) | Mean stalagmite / $ng \cdot g^{-1}$ (n=3) | Mean blank / $ng \cdot mL^{-1}$ (n=6) | Recov. SPE / % (n=3) |
|---|---|---|---|---|---|---|
| p-Hac | 13.8 | 41.9 | $50 \pm 30$ | $2.9 \pm 1.8$ | $155 \pm 130$ | $76 \pm 1$ |
| p-Hal | 25.9 | 78.4 | $25 \pm 85$ | $1.5 \pm 5.0$ | $680 \pm 330$ | $101 \pm 2$ |
| p-Hon | 2.3 | 7.0 | $55 \pm 5$ | $3.2 \pm 0.3$ | $80 \pm 20$ | $97 \pm 0$ |
| Vac | 13.7 | 41.5 | $330 \pm 80$ | $19.4 \pm 4.7$ | $60 \pm 30$ | $79 \pm 3$ |
| Val | 8.2 | 24.8 | $0 \pm 20$ | $0.0 \pm 1.2$ | $65 \pm 30$ | $69 \pm 4$ |
| Von | 3.7 | 11.3 | $1405 \pm 140$ | $82.6 \pm 8.2$ | $20 \pm 10$ | $79 \pm 3$ |
| Sac | 0.3 | 0.8 | $140 \pm 10$ | $8.2 \pm 0.6$ | $6 \pm 3$ | $73 \pm 2$ |
| Sal | 2.3 | 7.1 | $13 \pm 6.5$ | $0.8 \pm 0.4$ | $5 \pm 4$ | $77 \pm 2$ |
| Son | 2.5 | 7.7 | $110 \pm 30$ | $6.5 \pm 1.8$ | $4 \pm 4$ | $89 \pm 2$ |
| t-Fac | 2.0 | 6.2 | $100 \pm 0.5$ | $5.9 \pm 0.0$ | $5 \pm 2$ | $83 \pm 2$ |
| p-Cac | 0.2 | 0.7 | $195 \pm 60$ | $11.5 \pm 3.5$ | $445 \pm 505$ | $81 \pm 4$ |
| Eval (IS) | 0.6 | 1.8 | $147 \pm 4$ | $8.6 \pm 0.2$ | $1 \pm 1$ | $69 \pm 4$ |
| Ciac | 3.8 | 11.6 | $105 \pm 35$ | $6.2 \pm 2.1$ | $100 \pm 20$ | $84 \pm 3$ |

algae (Delwiche et al., 1989). Therefore, para-coumaric acid might be introduced into the sample via the laboratory air or via insufficiently purified water. Vanillin and its oxidised form vanillic acid are frequently used as perfumes and flavourings in food, cosmetics and household cleaning products. Therefore, these compounds might also be introduced into the sample via the laboratory air or via detergents used to clean the lab ware. Cinnamic acid is used as a perfume and flavouring, too, and it also occurs naturally in bacteria, fungi and algae, as it is part of the shikimate pathway (Dewick, 2009). In this study, cinnamic acid was found in the blank and in all samples. Therefore, cinnamic acid is not suitable as internal standard in the analysis of LOPs in natural samples, although it has been used as internal standard in many studies before (Goñi and Montgomery, 2000; Kaiser and Benner, 2012). Ethyl vanillin is much more suitable as internal standard because, as an artificial compound, it has very low blank values and does not occur in natural samples.

### 3.2.5 Repeatability

To determine the repeatability of the sample preparation and analysis method, 10.2 g stalagmite were dissolved, and the solution divided into three subsamples containing 3.4 g stalagmite. The mean values and standard deviations for all analytes are shown

in Table 2. The relative standard deviations ranged from 0.7% to 32% for analytes with more than 2.6 ng (50% for Sal with $2.6 \pm 1.3$ ng). For the p-hydroxy group, the relative standard deviations were higher, but these analytes were not used for the determination of LOP parameters. The LOP parameters calculated from these three subsamples were a C/V ratio of $0.17\pm0.04$ and an S/V ratio of $0.15\pm0.02$. The variability was mainly caused by the CuO oxidation step, which is known to cause relatively

5    high variability even in samples with higher lignin content (for example Hedges and Mann (1979) with standard deviations ranging between 3% and more than 80%). The SPE method used for the extraction of LOPs had standard deviations between 1–6% (Table 2) and therefore did not contribute much to the overall variability of the method.

### 3.2.6    Estimation of uncertainty

According to Konieczka and Namieśnik (2010), the main factors contributing to the uncertainty budget are the uncertainty of

10    the measurement of the weight or volume of the sample, $u_r(\text{sample})$, the repeatability of the sample preparation procedure, $u_r(\text{rep.})$, the recovery determination of the internal standard, $u_r(\text{recov.})$, the calibration step, $u_r(\text{cal.})$, and the uncertainty associated with analyte concentrations close to the limit of detection, $u_r(\text{LOD})$. The combined relative uncertainty $U_r$ is expressed in equation 5.

$$U_r = \sqrt{\left(u_r(\text{sample})\right)^2 + \left(u_r(\text{rep.})\right)^2 + \left(u_r(\text{recov.})\right)^2 + \left(u_r(\text{cal.})\right)^2 + \left(u_r(\text{LOD})\right)^2} \tag{5}$$

15    In our method, $u_r(\text{sample})$ is relatively small with 1 mg or 1 mL, which is usually $< 1\%$. The uncertainty associated with the repeatability of the sample preparation, calculated as the standard deviation of three individually prepared subsamples as explained in section 3.2.5, has the largest influence and can equal 1–30%. The uncertainty of the recovery determination of the internal standard, calculated as the standard deviation of the internal standard, contributes with 1–6%. $u_r(\text{cal.})$, calculated as the standard deviation of the concentration determination of three injections of the same sample into the LC-MS-system, can equal

20    1–15%, but is usually around 3–5%. $u_r(\text{LOD})$, calculated according to equation (6), depends strongly on the concentration $c$ of the analyte.

$$u_r(\text{LOD}) = \frac{\text{LOD}}{c} \tag{6}$$

In the data for stalagmite samples presented in Table 2, $u_r(\text{LOD})$ equals 0.1–5% for most analytes, 17% for Sal and 27–100% for the p-hydroxy group.

25    The errors for all results presented in this work were calculated using the law of propagation of uncertainty. All equations used for calculating concentrations, lignin oxidation parameters and errors are shown in section A4 in the SI.

### 3.3    Application to real samples

### 3.3.1    Analysis of plant and lignin samples

The method was applied to different natural samples from known sources to verify that the C/V and S/V ratios are in accordance

30    with published values. The results are shown in Table 3, and their S/V and C/V ratios are visualized in Figure 5. As expected,

the highest concentrations of LOPs are found in the lignin from conifer wood with a $\Sigma 8$ value of $75.76\,\mathrm{mg \cdot g^{-1}}$ as well as in the lignin from wheat straw and mixed wood with a $\Sigma 8$ value of $14.16\,\mathrm{mg \cdot g^{-1}}$. This means that the CuO oxidation method has a conversion factor of 1.4–7.6% (w/w) if applied to pure lignin, and that the conversion factor also depends on the type of lignin. The plant tissue samples gave LOP concentrations ($\Sigma 8$) of $2.3$–$6.8\,\mathrm{mg \cdot g^{-1}}$ for the wood and bark samples and $1.24$–

5    $1.30\,\mathrm{mg \cdot g^{-1}}$ for the leave and needle samples. These concentrations can be explained by the respective lignin content of the different samples. Figure 5 shows the C/V versus S/V diagram for all samples. The regions for different plant types have been defined by Hedges and Mann in 1979 and are based on the analysis of different plant species. Gymnosperm woody samples contain mainly V-group LOPs. Therefore, they plot close to the origin of the diagram. Angiosperm woody samples contain V- and S-group LOPs, but almost no C-group LOPs. Consequently, they plot close to the S/V-axis. Gymnosperm non-woody

10    samples contain V- and C-group LOPs, but almost no S-group LOPs. Accordingly, they plot close to the C/V-axis. Angiosperm non-woody samples contain all three groups of LOPs and thus show a wide range of C/V and S/V ratios. The analyzed plant samples in our study plot all in or close to the expected regions according to their plant type. Only the maple wood and bark sample and the maple leaves sample plot slightly outside of the regions for angiosperm woody and angiosperm non-woody material, respectively. For the maple wood and bark sample, this could be due to a higher contribution of C-group LOPs in the

15    bark compared to pure woody samples. However, it is important to keep in mind that these regions are just broadly defined and are based on a limited number of analyses and a limited number of different plant species.

**Table 3.** Concentrations of the V, S and C-group LOPs, the sum of all 8 LOPs ($\Sigma 8$) and the ratios C/V and S/V in fresh plant and lignin samples.

| sample | V-group / $\mathrm{mg \cdot g^{-1}}$ | S-group / $\mathrm{mg \cdot g^{-1}}$ | C-group / $\mathrm{mg \cdot g^{-1}}$ | $\Sigma 8$ / $\mathrm{mg \cdot g^{-1}}$ | C/V | S/V |
|---|---|---|---|---|---|---|
| lignin from conifer wood | $75.12 \pm 0.77$ | $0.293 \pm 0.015$ | $0.345 \pm 0.012$ | $75.76 \pm 0.77$ | $0.00 \pm 0.00$ | $0.00 \pm 0.00$ |
| lignin from wheat straw and mixed wood | $7.42 \pm 0.10$ | $6.483 \pm 0.078$ | $0.255 \pm 0.009$ | $14.16 \pm 0.13$ | $0.03 \pm 0.00$ | $0.87 \pm 0.02$ |
| yew wood and bark | $2.25 \pm 0.04$ | $0.024 \pm 0.001$ | $0.083 \pm 0.001$ | $2.35 \pm 0.04$ | $0.04 \pm 0.00$ | $0.01 \pm 0.00$ |
| maple wood and bark | $2.87 \pm 0.04$ | $3.626 \pm 0.089$ | $0.303 \pm 0.002$ | $6.80 \pm 0.10$ | $0.11 \pm 0.00$ | $1.27 \pm 0.04$ |
| yew needles | $0.74 \pm 0.02$ | $0.059 \pm 0.001$ | $0.494 \pm 0.010$ | $1.30 \pm 0.02$ | $0.66 \pm 0.02$ | $0.08 \pm 0.00$ |
| maple leaves | $0.75 \pm 0.02$ | $0.314 \pm 0.005$ | $0.184 \pm 0.003$ | $1.24 \pm 0.02$ | $0.25 \pm 0.01$ | $0.42 \pm 0.01$ |

### 3.3.2   Analysis of stalagmite samples

With a $\Sigma 8$ value of ca. $40$–$110\,\mathrm{ng \cdot (g\,stalagmite)^{-1}}$ (Table 4), the LOP concentration of the stalagmite samples is five orders of magnitude lower than for the vegetation samples and three to four orders of magnitude lower than the typical concentration

20    of sediment samples (e.g., $\Sigma 8$ is $0.15$–$0.75\,\mathrm{mg \cdot (g\,sediment)^{-1}}$ in Tareq et al. (2011)). Because of these low concentrations, 3–5 g stalagmite were required for an analysis to be above the limit of quantification. The C/V ratios of the stalagmite samples
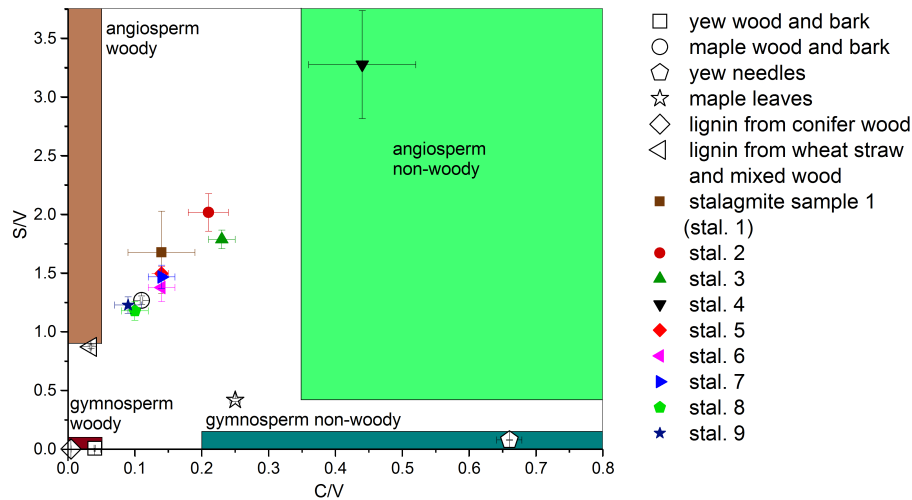
**Figure 6.** Lignin oxidation parameters S/V vs. C/V of different real samples and regions for different sample types defined by Hedges and Mann in 1979.

were all above 0.5, and the S/V ratios were all above 1.0, which suggests a significant contribution of angiosperm woody and angiosperm non-woody vegetation. However, gymnosperm woody and gymnosperm non-woody material might also have contributed to the lignin pool. This suggests a mixed deciduous forest above the cave, and would be in accordance with the results of Litt et al., who analyzed pollen from Holocene lake sediments from the Westeifel Volcanic Field (Litt et al., 2009), which is relatively close to the Herbstlabyrinth.

The nine stalagmite samples were taken at different distances from the top of the stalagmite. This analysis shall serve as a proof of principle for a higher-resolution analysis of the whole stalagmite. In Figure 6, the C/V and S/V ratios are plotted against distance from top (dft). Both ratios show a pronounced peak at 20 cm dft. Furthermore, both ratios show higher values in the top 15 cm and lower values with a decreasing trend between 30 and 50 cm dft. A higher S/V ratio indicates a higher contribution of angiosperm vegetation to the lignin source, and a higher C/V ratio suggests a higher contribution of non-woody vegetation. Therefore, the peak at 20 cm dft could be interpreted as increased input of non-woody, angiosperm vegetation, such as grasses, and less input of wood. The decreasing trend in the lower part of the stalagmite indicates a trend towards more woody, gymnosperm vegetation, such as pine forest. Of course, these presumptions have to be proven by a complete analysis of the stalagmite and a comparison with the other proxy data (Mischel et al., 2016). In addition, a comparison with Holocene pollen records from the area may confirm these preliminary results. Overall, these first results show significant variability of the C/V and S/V ratios and, therefore, the lignin sources. This promising result encourages us to use the analysis of LOPs in stalagmites for paleo-vegetation reconstruction.

**Table 4.** Concentrations of the V-, S- and C-group LOPs and the sum of all 8 LOPs (Σ8) in ng·g$^{-1}$ of the initial stalagmite samples and the ratios C/V and S/V. All samples are from stalagmite *NG01* from the Herbstlabyrinth-Advent-Cave.

| sample | V-group / ng·g$^{-1}$ | S-group / ng·g$^{-1}$ | C-group / ng·g$^{-1}$ | Σ8 / ng·g$^{-1}$ | C/V | S/V |
|---|---|---|---|---|---|---|
| stalagmite sample 1 | 20.9 ± 3.6 | 41.0 ± 3.8 | 3.4 ± 0.7 | 65.3 ± 5.3 | 0.16 ± 0.05 | 1.96 ± 0.39 |
| stalagmite sample 2 | 19.5 ± 1.4 | 45.5 ± 3.2 | 4.7 ± 0.5 | 69.7 ± 3.5 | 0.24 ± 0.03 | 2.33 ± 0.23 |
| stalagmite sample 3 | 17.4 ± 0.8 | 37.1 ± 1.4 | 4.8 ± 0.3 | 59.3 ± 1.6 | 0.27 ± 0.02 | 2.14 ± 0.13 |
| stalagmite sample 4 | 7.6 ± 0.9 | 29.1 ± 2.0 | 3.8 ± 0.3 | 40.6 ± 2.2 | 0.51 ± 0.07 | 3.84 ± 0.51 |
| stalagmite sample 5 | 38.2 ± 2.6 | 67.3 ± 3.4 | 6.4 ± 0.6 | 111.9 ± 4.3 | 0.17 ± 0.02 | 1.76 ± 0.15 |
| stalagmite sample 6 | 24.5 ± 1.5 | 40.0 ± 1.7 | 4.1 ± 0.3 | 68.6 ± 2.3 | 0.17 ± 0.01 | 1.63 ± 0.12 |
| stalagmite sample 7 | 24.4 ± 1.4 | 42.1 ± 2.2 | 3.9 ± 0.2 | 70.4 ± 2.6 | 0.16 ± 0.01 | 1.72 ± 0.14 |
| stalagmite sample 8 | 25.7 ± 1.3 | 35.5 ± 1.7 | 2.9 ± 0.2 | 64.1 ± 2.2 | 0.11 ± 0.01 | 1.38 ± 0.10 |
| stalagmite sample 9 | 39.1 ± 2.8 | 56.2 ± 2.8 | 4.1 ± 0.3 | 99.4 ± 3.9 | 0.10 ± 0.01 | 1.44 ± 0.13 |
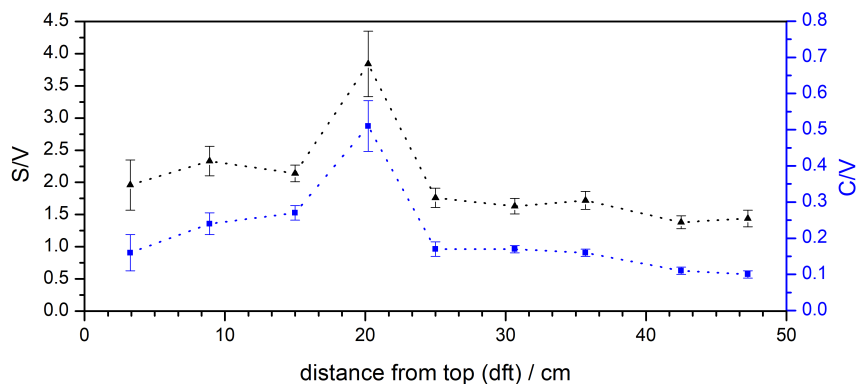
**Figure 7.** S/V (black triangles, left axis) and C/V (blue squares, right axis) ratios of stalagmite *NG01* plotted against the distance from the top of the stalagmite

### 3.3.3 Analysis of cave drip water samples

Very little is known about how lignin is transported from the soil into the cave and how it is incorporated into a stalagmite. To gain further understanding about these processes, it is useful to also analyze lignin in cave drip water. The lignin concentration in cave drip water is even lower than in stalagmite samples, because crystallization of calcite also serves as an enrichment step for the organic components contained in the water. Therefore, a sample volume of 100–200 mL water was used. Here we show the results of the analysis of six different water samples from the Herbstlabyrinth-Advent-Cave, all sampled in October 2014 (Table 5). As expected, the soil water (*SW*) has the largest lignin content with 1.8 μg·L$^{-1}$. The rain water (*RW*) also has a relatively large lignin content of 1.3 μg·L$^{-1}$, which is surprising since this water has not been in contact with soil or

vegetation. The lignin content of the cave drip water samples is much lower, ranging from $0.21\,\mu g \cdot L^{-1}$ for the pool water to $0.36\,\mu g \cdot L^{-1}$ for the fast drip site *D1*. The concentrations of all LOPs decrease from the soil water to the cave drip water, but to a different extent. Whereas V-group LOPs and C-group LOPs decrease by 80–92% and 82–90%, respectively, the concentration of S-group LOPs decreases only by 70–76% (Fig. 8). This is also reflected in higher S/V ratios in the cave drip water than in the soil water, with an increasing trend from the soil water over the two fast drip sites *D1* and *D5* and the slow drip site *D2* to the cave pool water. This could be due to different residence times in the cave and the overlaying karst of the water from the different drip sites. These hypotheses should be proven by a further systematic analysis of cave drip water. This would also enable the study of seasonal variations in the lignin input. The monthly cave monitoring program of Mischel et al. (Mischel et al., 2016, 2015) combined with our new method for the analysis of LOPs even in low-concentration cave drip water could be a valuable tool to further investigate these topics.

**Table 5.** Concentrations of the V-, S- and C-group LOPs, the sum of all 8 LOPs ($\Sigma 8$) and the ratios C/V and S/V in different water samples collected at the Herbstlabyrinth-Advent-Cave in October 2014.

| Sample | Sample volume / L | V-group / $ng \cdot L^{-1}$ | S-group / $ng \cdot L^{-1}$ | C-group / $ng \cdot L^{-1}$ | $\Sigma 8$ / $ng \cdot L^{-1}$ | C/V | S/V |
|---|---|---|---|---|---|---|---|
| RW (rain water) | 0.185 | $918 \pm 69$ | $345 \pm 31$ | $76 \pm 17$ | $1339 \pm 77$ | $0.08 \pm 0.02$ | $0.38 \pm 0.04$ |
| SW (soil water) | 0.076 | $1370 \pm 101$ | $363 \pm 54$ | $42 \pm 38$ | $1775 \pm 121$ | $0.03 \pm 0.03$ | $0.26 \pm 0.04$ |
| D1 (fast dripping) | 0.265 | $271 \pm 21$ | $87 \pm 16$ | $7 \pm 11$ | $365 \pm 29$ | $0.03 \pm 0.04$ | $0.32 \pm 0.07$ |
| D5 (fast dripping) | 0.258 | $175 \pm 20$ | $95 \pm 19$ | $6 \pm 12$ | $275 \pm 30$ | $0.03 \pm 0.07$ | $0.54 \pm 0.13$ |
| D2 (slow dripping) | 0.205 | $157 \pm 15$ | $107 \pm 29$ | $4 \pm 14$ | $269 \pm 35$ | $0.03 \pm 0.09$ | $0.68 \pm 0.19$ |
| PW (pool water) | 0.253 | $114 \pm 23$ | $88 \pm 29$ | $8 \pm 21$ | $210 \pm 42$ | $0.07 \pm 0.18$ | $0.77 \pm 0.29$ |

# 4 Aspects of green analytical chemistry

When developing a new analytical method, it is advantageous to consider how environment-friendly (*green*) the different approaches are. The principles of green analytical chemistry include, among others, to generate as little waste as possible, to eliminate or replace toxic reagents, to miniaturize analytical instruments or to avoid derivatization (Gałuszka et al., 2013; Armenta et al., 2008). In our method, we tried to favour greener approaches over less green approaches whenever possible without sacrificing other qualities like sensitivity. We used solid phase extraction, which consumes considerably less solvent than liquid-liquid extraction, and UHPLC, which is less solvent and time consuming than HPLC. In addition, liquid chromatography does not require a derivatization step, as opposed to gas chromatography. However, the least green step in our method is the CuO oxidation step, as it generates toxic waste and consumes energy. We still chose the CuO oxidation method for our proof of principle analysis because it is the most widely used lignin degradation method for the analysis of LOPs and therefore allows us to compare our results with existing LOP records. In future, however, a greener approach to the degradation of lignin
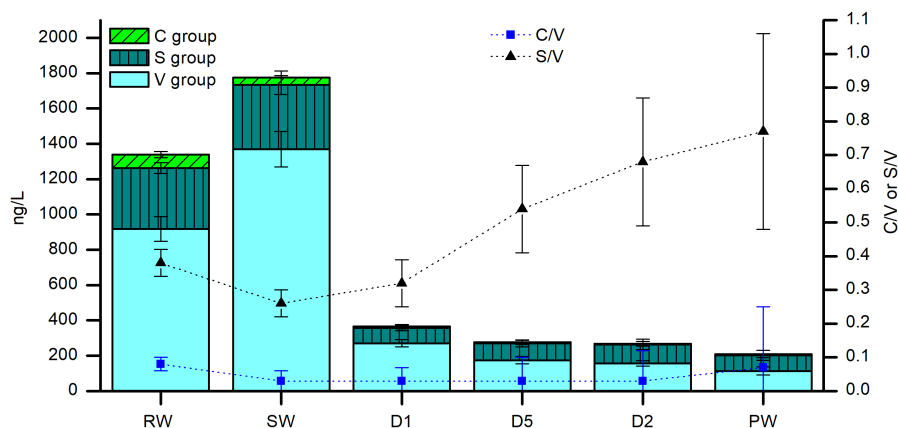
**Figure 8.** LOP concentrations (stacked columns with left axis) and LOP ratios (symbols with right axis) of rain water (*RW*), soil water (*SW*), cave drip water from fast drip sites (*D1* and *D5*), a slow drip site (*D2*) and cave pool water (*PW*). The stacked columns contain the V-group LOPs (light cyan bars), S-group LOPs (dark cyan bars with vertical stripes) and C-group LOPs (green bars with diagonal stripes). Black triangles show the S/V ratio and blue squares show the C/V ratio.

to LOPs should be chosen, which could, for example, be based on electrolysis, preferably in a miniaturized flow cell (Leppla, 2016).

## 5   Conclusions and outlook

We developed a sensitive method for the quantification of LOPs in speleothems and cave drip water and tested it success-
5   fully on samples from the Herbstlabyrinth-Advent-Cave. This is, to our knowledge, the first quantitative analysis of LOPs in speleothems and cave drip water. Our method provides a new and highly specific vegetation proxy for the reconstruction of paleo-vegetation and paleo-climate from speleothem archives. The method was adjusted to the low concentrations of organic matter in speleothems and cave drip water and showed sufficient sensitivity to detect even trace concentrations of lignin. The use of the established CuO oxidation method allows to compare the results to LOP records in other archives. However, as the
10   CuO oxidation step is the main source of variability in our method, an alternative degradation method for lignin with higher re-producibility should be developed. This method could, for example, be based on electrolysis. In addition, LOPs in speleothem samples from other caves in different vegetation and climate zones should be analyzed and compared with stable isotope and trace element records in order to gain more insight into the relation of vegetation, climate and the LOP signal in speleothems. The analysis of cave drip water, sampled monthly within the framework of a cave monitoring program, could elucidate seasonal
15   variations of lignin input as well as possible fractionation processes during its pathway from the soil to the cave.

*Competing interests.* The authors declare that they have no conflict of interest.

*Acknowledgements.* We thank Dr. Simon Mischel for providing stalagmite and cave drip water samples from the Herbstlabyrinth-Advent-Cave. This project has received funding from the European Union's Horizon 2020 research and innovation programme under the Marie Skłodowska-Curie grant agreement No 691037. Denis Scholz acknowledges funding of the German Research Foundation (SCHO 1274/3-1 and SCHO 1274/9-1)

# References

Armenta, S., Garrigues, S., and de La Guardia, M.: Green Analytical Chemistry, TrAC Trends in Analytical Chemistry, 27, 497–511, 2008.

Baker, A., Jex, C. N., Rutlidge, H., Woltering, M., Blyth, A. J., Andersen, M. S., Cuthbert, M. O., Marjo, C. E., Markowska, M., Rau, G. C., and Khan, S. J.: An irrigation experiment to compare soil, water and speleothem tetraether membrane lipid distributions, Organic Geochemistry, 94, 12–20, 2016.

Blyth, A. J. and Schouten, S.: Calibrating the glycerol dialkyl glycerol tetraether temperature signal in speleothems, Geochimica et Cosmochimica Acta, 109, 312–328, 2013.

Blyth, A. J. and Watson, J. S.: Thermochemolysis of organic matter preserved in stalagmites: A preliminary study, Organic Geochemistry, 40, 1029–1031, 2009.

Blyth, A. J., Farrimond, P., and Jones, M.: An optimised method for the extraction and analysis of lipid biomarkers from stalagmites, Organic Geochemistry, 37, 882–890, 2006.

Blyth, A. J., Baker, A., Collins, M. J., Penkman, K. E., Gilmour, M. A., Moss, J. S., Genty, D., and Drysdale, R. N.: Molecular organic matter in speleothems and its potential as an environmental proxy, Quaternary Science Reviews, 27, 905–921, 2008.

Blyth, A. J., Watson, J. S., Woodhead, J., and Hellstrom, J.: Organic compounds preserved in a 2.9million year old stalagmite from the Nullarbor Plain, Australia, Chemical Geology, 279, 101–105, 2010.

Blyth, A. J., Hartland, A., and Baker, A.: Organic proxies in speleothems – New developments, advantages and limitations, Quaternary Science Reviews, 149, 1–17, 2016.

Bosle, J. M., Mischel, S. A., Schulze, A.-L., Scholz, D., and Hoffmann, T.: Quantification of low molecular weight fatty acids in cave drip water and speleothems using HPLC-ESI-IT/MS — development and validation of a selective method, Analytical and bioanalytical chemistry, 406, 3167–3177, 2014.

Delwiche, C. F., Graham, L. E., and Thomson, N.: Lignin-like compounds and sporopollenin coleochaete, an algal model for land plant ancestry, Science (New York, N.Y.), 245, 399–401, 1989.

Dewick, P. M.: The Shikimate Pathway: Aromatic Amino Acids and Phenylpropanoids, chap. 4, pp. 137–186, Wiley-Blackwell, 2009.

Dittmar, T. and Lara, R. J.: Molecular evidence for lignin degradation in sulfate-reducing mangrove sediments (Amazônia, Brazil), Geochimica et Cosmochimica Acta, 65, 1417–1428, 2001.

Fairchild, I. J. and Baker, A.: Speleothem Science, John Wiley & Sons, Ltd, Chichester, UK, 2012.

Fairchild, I. J. and Treble, P. C.: Trace elements in speleothems as recorders of environmental change, Quaternary Science Reviews, 28, 449–468, 2009.

Fairchild, I. J., Smith, C. L., Baker, A., Fuller, L., Spötl, C., Mattey, D., McDermott, F., and E.I.M.F.: Modification and preservation of environmental signals in speleothems, Earth-Science Reviews, 75, 105–153, 2006.

Fraser, W. T., Scott, A. C., Forbes, A. E. S., Glasspool, I. J., Plotnick, R. E., Kenig, F., and Lomax, B. H.: Evolutionary stasis of sporopollenin biochemistry revealed by unaltered Pennsylvanian spores, New Phytologist, 196, 397–401.

Gałuszka, A., Migaszewski, Z., and Namieśnik, J.: The 12 principles of green analytical chemistry and the SIGNIFICANCE mnemonic of green analytical practices, TrAC Trends in Analytical Chemistry, 50, 78–84, 2013.

Gałuszka, A., Migaszewski, Z. M., and Namieśnik, J.: The role of analytical chemistry in the study of the Anthropocene, TrAC Trends in Analytical Chemistry, 97, 146–152, 2017.

Giorio, C., Kehrwald, N., Barbante, C., Kalberer, M., King, A. C., Thomas, E. R., Wolff, E. W., and Zennaro, P.: Prospects for reconstructing paleoenvironmental conditions from organic compounds in polar snow and ice, Quaternary Science Reviews, 183, 1–22, 2018.

Goñi, M. A. and Montgomery, S.: Alkaline CuO Oxidation with a Microwave Digestion System: Lignin Analyses of Geochemical Samples, Analytical chemistry, 72, 3116–3121, 2000.

5   Hedges, J. I. and Mann, D. C.: The characterization of plant tissues by their lignin oxidation products, Geochimica et Cosmochimica Acta, 43, 1803–1807, 1979.

Hedges, J. I. and Parker, P. L.: Land-derived organic matter in surface sediments from the Gulf of Mexico, Geochimica et Cosmochimica Acta, 40, 1019–1029, 1976.

Hernes, P. J. and Benner, R.: Transport and diagenesis of dissolved and particulate terrigenous organic matter in the North Pacific Ocean,

10   Deep Sea Research Part I: Oceanographic Research Papers, 49, 2119–2132, 2002.

Jex, C. N., Pate, G. H., Blyth, A. J., Spencer, R. G., Hernes, P. J., Khan, S. J., and Baker, A.: Lignin biogeochemistry: from modern processes to Quaternary archives, Quaternary Science Reviews, 87, 46–59, 2014.

Kaiser, K. and Benner, R.: Characterization of Lignin by Gas Chromatography and Mass Spectrometry Using a Simplified CuO Oxidation Method, Analytical chemistry, 84, 459–464, 2012.

15   Kögel, I. and Bochter, R.: Characterization of lignin in forest humus layers by high-performance liquid chromatography of cupric oxide oxidation products, Soil Biology and Biochemistry, 17, 637–640, 1985.

Konieczka, P. and Namieśnik, J.: Estimation of uncertainty in analytical procedures based on chromatographic techniques, Journal of chromatography. A, 1217, 881, 2010.

Leppla, D.: Quantification of paleoclimatic proxies using UHPLC-HESI-MS after degradation of lignin – development of an electrochemical

20   method, Masterarbeit, Johannes Gutenberg-Universität, Mainz, 2016.

Lima, D. L., Duarte, A. C., and Esteves, V. I.: Solid-phase extraction and capillary electrophoresis determination of phenols from soil after alkaline CuO oxidation, Chemosphere, 69, 561–568, 2007.

Litt, T., Schölzel, C., Kühl, N., and Brauer, A.: Vegetation and climate history in the Westeifel Volcanic Field (Germany) during the past 11 000 years based on annually laminated lacustrine maar sediments, Boreas, 38, 679–690, 2009.

25   McDermott, F.: Palaeo-climate reconstruction from stable isotope variations in speleothems: a review, Quaternary Science Reviews, 23, 901–918, 2004.

Mischel, S. A., Scholz, D., and Spötl, C.: d18O values of cave drip water: a promising proxy for the reconstruction of the North Atlantic Oscillation?, Climate Dynamics, 45, 3035–3050, 2015.

Mischel, S. A., Scholz, D., Spo tl, C., Jochum, K. P., Schro der-Ritzrau, A., and Fiedler, S.: Holocene climate variability in Central Germany

30   and a potential link to the polar North Atlantic: A replicated record from three coeval speleothems, The Holocene, 2016.

Mischel, S. A., Mertz-Kraus, R., Jochum, K. P., and Scholz, D.: TERMITE: An R script for fast reduction of laser ablation inductively coupled plasma mass spectrometry data and its application to trace element measurements, Rapid communications in mass spectrometry : RCM, 31, 1079–1087, 2017.

Montgomery, W., Potiszil, C., Watson, J. S., and Sephton, M. A.: Sporopollenin, a Natural Copolymer, is Robust under High Hydrostatic

35   Pressure, Macromolecular Chemistry and Physics, 217, 2494–2500, 2016.

Quiers, M., Perrette, Y., Chalmin, E., Fanget, B., and Poulenard, J.: Geochemical mapping of organic carbon in stalagmites using liquid-phase and solid-phase fluorescence, Chemical Geology, 411, 240–247, 2015.

Richards, D. A. and Dorale, J. A.: Uranium-series Chronology and Environmental Applications of Speleothems, Reviews in Mineralogy and Geochemistry, 52, 407–460, 2003.

Scholz, D. and Hoffmann, D.: 230Th/U-dating of fossil corals and speleothems, Quat. Sci. J, 57, 52, 2008.

Standley, L. J. and Kaplan, L. A.: Isolation and analysis of lignin-derived phenols in aquatic humic substances: improvements on the procedures, Organic Geochemistry, 28, 689–697, 1998.

Tareq, S. M., Tanaka, N., and Ohta, K.: Biomarker signature in tropical wetland: lignin phenol vegetation index (LPVI) and its implications for reconstructing the paleoenvironment, The Science of the total environment, 324, 91–103, 2004.

Tareq, S. M., Kitagawa, H., and Ohta, K.: Lignin biomarker and isotopic records of paleovegetation and climate changes from Lake Erhai, southwest China, since 18.5 ka BP, Holocene Lake Records: Patterns, Impacts, Causes And Societal Response Selected Papers from the 3rd LIMPACS Conference, Chandigarh, India, 229, 47–56, 2011.

Wysocki, L. A., Filley, T. R., and Bianchi, T. S.: Comparison of two methods for the analysis of lignin in marine sediments: CuO oxidation versus tetramethylammonium hydroxide (TMAH) thermochemolysis, Organic Geochemistry, 39, 1454–1461, 2008.

Xie, S.: Lipid distribution in a subtropical southern China stalagmite as a record of soil ecosystem response to paleoclimate change, Quaternary Research, 2003.

Zhang, T., Li, X., Sun, S., Lan, H., Du, P., and Wang, M.: Determination of lignin in marine sediment using alkaline cupric oxide oxidation-solid phase extraction-on-column derivatization-gas chromatography, Journal of Ocean University of China, 12, 63–69, 2013.