

1 **An intercomparison of oceanic methane and nitrous oxide measurements**

2
3 Samuel T. Wilson^{1*}, Hermann W. Bange², Damian L. Arévalo-Martínez², Jonathan Barnes³,
4 Alberto V. Borges⁴, Ian Brown⁵, John L. Bullister⁶, Macarena Burgos^{1,7}, David W. Capelle⁸,
5 Michael Casso⁹, Mercedes de la Paz^{10†}, Laura Farías¹¹, Lindsay Fenwick⁸, Sara Ferrón¹, Gerardo
6 Garcia¹¹, Michael Glockzin¹², David M. Karl¹, Annette Kock², Sarah Laperriere¹³, Cliff S.
7 Law^{14,15}, Cara C. Manning⁸, Andrew Marriner¹⁴, Jukka-Pekka Myllykangas¹⁶, John W.
8 Pohlman⁹, Andrew P. Rees⁵, Alyson E. Santoro¹³, Philippe D. Tortell⁸, Robert C. Upstill-
9 Goddard³, David P. Wisegarver⁶, Guiling L. Zhang¹⁷, Gregor Rehder¹²

10

11 ¹University of Hawai'i at Manoa, Daniel K. Inouye Center for Microbial Oceanography:
12 Research and Education (C-MORE), Honolulu, Hawai'i, USA

13 ²GEOMAR Helmholtz Centre for Ocean Research Kiel, Düsternbrooker Weg 20 24105 Kiel,
14 Germany

15 ³Newcastle University, School of Natural and Environmental Sciences, Newcastle upon Tyne,
16 UK

17 ⁴Université de Liège, Unité d'Océanographie Chimique, Liège, Belgium

18 ⁵Plymouth Marine Laboratory, Plymouth, UK

19 ⁶National Oceanic and Atmospheric Administration, Pacific Marine Environmental Laboratory,
20 Seattle, Washington, USA

21 ⁷Universidad de Cádiz, Instituto de Investigaciones Marinas, Departamento Química-Física
22 Cádiz, Spain

23 ⁸University of British Columbia, Vancouver, Department of Earth, Ocean and Atmospheric
24 Sciences, British Columbia, Canada

25 ⁹U.S. Geological Survey, Woods Hole Coastal and Marine Science Center, Woods Hole, USA

26 ¹⁰Instituto de Investigaciones Marinas, Vigo, Spain

27 ¹¹University of Concepción, Department of Oceanography and Center for climate research and
28 resilience (CR2), Concepción, Chile

29 ¹²Leibniz Institute for Baltic Sea Research Warnemünde, Rostock, Germany

30 ¹³University of California Santa Barbara, Department of Ecology, Evolution, and Marine
31 Biology, Santa Barbara, USA

32 ¹⁴National Institute of Water and Atmospheric Research (NIWA), Wellington, New Zealand

33 ¹⁵Department of Chemistry, University of Otago, Dunedin, New Zealand

34 ¹⁶University of Helsinki, Department of Environmental Sciences, Helsinki, Finland

35 ¹⁷Ocean University of China, Department of Marine Chemistry, Qingdao, China

36

37 †Current address: Instituto Español de Oceanografía, Centro Oceanográfico de A Coruña, A

38 Coruña, Spain

39

40 *corresponding author: stwilson@Hawaii'i.edu

41 **Abstract.** Large scale climatic forcing is impacting oceanic biogeochemical cycles and is
42 expected to influence the water-column distribution of trace gases including methane and nitrous
43 oxide. Our ability as a scientific community to evaluate changes in the water-column inventories
44 of methane and nitrous oxide depends largely on our capacity to obtain robust and accurate
45 concentration measurements which can be validated across different laboratory groups. This
46 study represents the first formal, international, intercomparison of oceanic methane and nitrous
47 oxide measurements whereby participating laboratories received batches of seawater samples
48 from the subtropical Pacific Ocean and the Baltic Sea. Additionally, compressed gas standards
49 from the same calibration scale were distributed to the majority of participating laboratories to
50 improve the analytical accuracy of the gas measurements. The computations used by each
51 laboratory to derive the dissolved gas concentrations were also evaluated for inconsistencies (*e.g.*
52 pressure and temperature corrections, solubility constants). The results from the intercomparison
53 and intercalibration provided invaluable insights into methane and nitrous oxide measurements.
54 It was observed that analyses of seawater samples with the lowest concentrations of methane and
55 nitrous oxide had the lowest precisions. In comparison, while the analytical precision for
56 samples with the highest concentrations of trace gases was better, the variability between the
57 different laboratories was higher; 36% for methane and 27% for nitrous oxide. In addition, the
58 comparison of different batches of seawater samples with methane and nitrous oxide
59 concentrations that ranged over an order of magnitude revealed the ramifications of different
60 calibration procedures for each trace gas. Overall, this study builds upon the intercomparison
61 results to develop a framework for improving oceanic methane and nitrous oxide measurements,
62 with the aim of precluding future analytical discrepancies between laboratories.

63 **1. Introduction**

64 The increasing mole fractions of greenhouse gases in the Earth's atmosphere are causing long-
65 term climate change with unknown future consequences. Two greenhouse gases, methane and
66 nitrous oxide, together contribute approximately 23% of total radiative forcing attributed to well-
67 mixed greenhouse gases (Myhre et al., 2013). It is imperative that the monitoring of methane
68 and nitrous oxide in the Earth's atmosphere is accompanied by measurements at the Earth's
69 surface to better inform the sources and sinks of these climatically important trace gases. This
70 includes measurements of dissolved methane and nitrous oxide in the marine environment,
71 which is an overall source of both gases to the overlying atmosphere (Nevison et al., 1995;
72 Anderson et al., 2010; Naqvi et al., 2010; Freing et al., 2012; Ciais et al., 2014).

73 Oceanic measurements of methane and nitrous oxide are conducted as part of established
74 time-series locations, along hydrographic survey lines, and during disparate oceanographic
75 expeditions. Within low to mid-latitude regions of the open ocean, the surface waters are
76 frequently slightly super-saturated with respect to atmospheric equilibrium for both methane and
77 nitrous oxide. There is typically an order of magnitude range in concentration along a vertical
78 water-column profile at any particular open ocean location (e.g. Wilson et al., 2017). In contrast
79 to the open ocean, near-shore environments, which are subject to river inputs, coastal upwelling,
80 benthic exchange and other processes, have higher concentrations and greater spatial and
81 temporal heterogeneity (e.g. Schmale et al., 2010; Upstill-Goddard and Barnes, 2016).

82 Methods for quantifying dissolved methane and nitrous oxide have evolved and somewhat
83 diverged since the first measurements were made in the 1960s (Craig and Gordon 1963;
84 Atkinson and Richards 1967). Some laboratories employ purge-and-trap methods for extracting
85 and concentrating the gases prior to their analysis (e.g. Zhang et al., 2004; Bullister and
86 Wisegarver, 2008; Capelle et al., 2015; Wilson et al., 2017). Others equilibrate a seawater
87 sample with an overlying headspace gas and inject a fixed volume of the gaseous phase into a
88 gas analyzer (e.g. Upstill-Goddard et al., 1996; Walter et al., 2005; Farias et al., 2009). **The**
89 **purge and trap technique is typically more sensitive by 2-3 orders of magnitude over headspace**
90 **equilibrium. However, the purge and trap technique requires more time for sample analysis and**
91 **it is more difficult to automate the injection of samples into the gas analyzer. Headspace**
92 **equilibrium sampling is most suited for volatile compounds that can be efficiently partitioned**
93 **into the headspace gas volume from the seawater sample. Its limited sensitivity can be**

94 compensated by large volume analysis (*e.g.* Upstill-Goddard et al., 1996). Additional
95 developments for continuous underway surface seawater measurements use equilibrator systems
96 of various designs coupled to a variety of detectors (*e.g.* Weiss et al., 1992; Butler et al., 1989;
97 Gülzow et al., 2011; Arévalo-Martínez et al., 2013). Determining the level of analytical
98 comparability between different laboratories for discrete samples of methane and nitrous oxide is
99 an important step towards improved comprehensive global assessments. Such intercomparison
100 exercises are critical to determining the spatial and temporal variability of methane and nitrous
101 oxide across the world oceans with confidence, since no single laboratory can single-handedly
102 provide all the required measurements at sufficient resolution. Previous comparative exercises
103 have been conducted for other trace gases *e.g.* carbon dioxide, dimethylsulphide, and sulfur
104 hexafluoride (Dickson et al., 2007; Bullister and Tanhua, 2010; Swan et al., 2014) and for trace
105 elements (Cutter et al., 2013). These exercises confirm the value of the intercomparison concept.

106 To instigate this process for methane and nitrous oxide, a series of international
107 intercomparison exercises were conducted between 2013 and 2017, under the auspices of
108 Working Group #143 of the Scientific Committee on Oceanic Research (SCOR) ([www.scor-](http://www.scor-int.org)
109 [int.org](http://www.scor-int.org)). Discrete seawater samples collected from the subtropical Pacific Ocean and the Baltic
110 Sea were distributed to the participating laboratories (Table 1). The samples were selected to
111 cover a representative range of concentrations across marine locations, from the oligotrophic
112 open ocean to highly productive waters, and in some instances sub-oxic, coastal waters. An
113 integral component of the intercomparison exercise was the production and distribution of
114 methane and nitrous oxide gas standards to members of the SCOR Working Group. The
115 intercomparison exercise was conceived and evaluated with the following four questions in
116 mind:

117 Q1. What is the agreement between the SCOR gas standards and the ‘in-house’ gas standards
118 used by each laboratory?

119 Q2. How do measured values of dissolved methane and nitrous oxide compare across
120 laboratories?

121 Q3. Despite the use of different analytical systems, are there general recommendations to reduce
122 uncertainty in the accuracy and precision of methane and nitrous oxide measurements?

123 Q4. What are the implications of inter-laboratory differences for determining the spatial and
124 temporal variability of methane and nitrous oxide in the oceans?

125

126 **2. Methods**

127 **2.1 Calibration of nitrous oxide and methane using compressed gas standards**

128 Laboratory-based measurements of oceanic methane and nitrous oxide require separation of the
129 dissolved gas from the aqueous phase, with the analysis conducted on the gaseous phase.

130 Calibration of the analytical instrumentation used to quantify the concentration of methane and
131 nitrous oxide is nearly always conducted using compressed gas standards, the specifics of which
132 vary between each laboratory. Therefore, the reporting of methane and nitrous oxide datasets
133 ought to be accompanied by a description of the standards used, including their methane and
134 nitrous oxide mole fractions, the declared accuracies, and the composition of their balance or
135 ‘make-up’ gas. For both gases, the highest accuracy commercially available standards have
136 mole fractions close to current day atmospheric values. These standards can be obtained from
137 national agencies including National Oceanic and Atmospheric Administration Global
138 Monitoring Division (NOAA GMD), the National Institute of Metrology China, and the Central
139 Analytical Laboratories of the European Integrated Carbon Observation System Research
140 Infrastructure (ICOS-RI). By comparison, it is more difficult to obtain highly accurate methane
141 and nitrous oxide gas standards with mole fractions exceeding modern-day atmospheric values.
142 This is particularly problematic for nitrous oxide due to the nonlinearity of the widely used
143 Electron Capture Detector (ECD) (Butler and Elkins, 1991).

144 The absence of a widely available high mole fraction, high accuracy nitrous oxide gas
145 standard was noted as a primary concern at the outset of the intercomparison exercise.
146 Therefore, a set of high-pressure primary gas standards was prepared for the SCOR Working
147 Group by John Bullister and David Wisegarver at NOAA Pacific Marine and Environmental
148 Laboratory (PMEL). One batch, referred to as Air Ratio Standard (ARS), had methane and
149 nitrous oxide mole fractions similar to modern air and the other batch, referred to as Water Ratio
150 Standard (WRS) had higher methane and nitrous oxide mole fractions for calibration of high
151 concentration water samples. These SCOR primary standards were checked for stability over a
152 12 month period and assigned mole fractions on the same calibration scale, known as ‘SCOR-
153 2016.’ A comparison was conducted with NOAA standards prepared on the SIO98 calibration
154 scale for nitrous oxide and the NOAA04 calibration scale for methane. Based on the comparison
155 with NOAA standards, the uncertainty of the methane and nitrous oxide mole fractions in the

156 ARS and the uncertainty of the methane mole fraction in the WRS were all estimated at better
157 than 1%. By contrast, the uncertainty of the nitrous oxide mole fraction in the WRS was
158 estimated at 2-3%. The gas standards were distributed to twelve of the laboratories involved in
159 this study (Table 1). The technical details on the production of the gas standards and their
160 assigned absolute mole fractions is included in Bullister et al. (2016).

161

162 **2.2 Collection of discrete samples of nitrous oxide and methane**

163 Dissolved methane and nitrous oxide samples for the intercomparison exercise were collected
164 from the subtropical Pacific Ocean and the Baltic Sea. Pacific samples were obtained on 28
165 November 2013 and 24 February 2017 from the Hawai'i Ocean Time-series (HOT) long-term
166 monitoring site, Station ALOHA, located at 22.75 N, 158.00 W. The November 2013 samples
167 are included in Figure S1 and S2 in the Supplement, but are not discussed in the main Results or
168 Discussion because fewer laboratories were involved in the initial intercomparison, and the
169 results from these samples support the same conclusions obtained with the more recent sample
170 collections. Seawater was collected using Niskin-like bottles designed by John Bullister (NOAA
171 PMEL), which help minimize contamination of trace gases, in particular chlorofluorocarbons
172 and sulfur hexafluoride (Bullister and Wisegarver, 2008). The bottles were attached to a rosette
173 with a conductivity-temperature-depth (CTD) package. Seawater was collected from two depths:
174 700 m and 25 m, where the near-maximum and minimum water-column concentrations for
175 methane and nitrous oxide at this location can be found. **The 25 m samples were always well
176 within the surface mixed layer, which ranged from 100 to 130 m depth during sampling.**

177 Replicate samples were collected from each bottle, with one replicate reserved for analysis at the
178 University of Hawai'i to evaluate variability between sampling bottles. Seawater was dispensed
179 from the Niskin-like bottles using Tygon® tubing into the bottom of borosilicate glass bottles,
180 allowing overflow of at least two sample volumes and ensuring the absence of bubbles. Most
181 sample bottles were 240 mL in size and were sealed with no headspace using butyl-rubber
182 stoppers and aluminum crimp-seals. A few laboratory groups requested smaller crimp-sealed
183 glass bottles ranging from 20-120 mL in volume and two laboratories used 1 L glass bottles
184 which were closed with a glass stopper and sealed with Apiezon® grease. Seawater samples
185 were collected in quadruplicate for each laboratory. All samples were preserved using saturated
186 mercuric chloride solution (100 µL of saturated mercuric chloride solution per 100 mL of

187 seawater sample) and stored in the dark at room temperature until shipment. The choice of
188 mercuric chloride as the preservative for dissolved methane and nitrous oxide was due to its long
189 history of usage. It is recognized that other preservatives have been proposed (e.g. Magen et al.,
190 2014, Bussmann et al., 2015), however pending a community-wide evaluation of their
191 effectiveness over a range of microbial assemblages and environmental conditions for both
192 methane and nitrous oxide, it is not evident that they are a superior alternative to mercuric
193 chloride.

194 Samples from the western Baltic Sea were collected during 15-21 October 2016, onboard the
195 R/V *Elisabeth Mann Borgese* (Table 2). Since the Baltic Sea consists of different basins with
196 varying concentrations of oxygen beneath permanent haloclines (Schmale et al., 2010), a larger
197 range of water-column methane and nitrous oxide concentrations were accessible for inter-
198 laboratory comparison compared to Station ALOHA. For all seven Baltic Sea stations, the
199 water-column was sampled into an on-deck 1,000 L water tank that was subsequently
200 subsampled into discrete sample bottles. At three stations (BAL1, BAL3, and BAL6), the water
201 tank was filled from the shipboard high-throughput underway seawater system. For deeper
202 water-column sampling at the stations BAL2, BAL4, and BAL5, the water tank was filled using
203 a pumping CTD system (Strady et al., 2008) with a flow rate of 6 L min⁻¹ and a total pumping
204 time of approximately 3 h. For the final deep water-column station, BAL7, the pump that
205 supplied the shipboard underway system was lowered to a depth of 21 m to facilitate a shorter
206 pumping time of approximately 20 mins. Subsampling the water tank for all samples took
207 approximately 1 h in total and the total sampling volume was less than 100 L. To verify the
208 homogeneity of the seawater during the sampling process, the first and last samples collected
209 from the water tank were analyzed by Newcastle University onboard the research vessel. In
210 contrast to the Pacific Ocean sampling, which predominantly used 240 mL glass vials, each
211 laboratory provided their own preferred vials and stoppers for the Baltic Sea samples. Seawater
212 samples were collected in triplicate for each laboratory. All samples were preserved with 100
213 µL of saturated mercuric chloride solution per 100 ml of seawater sample, with the exception of
214 samples collected by U.S. Geological Survey, who analyzed unpreserved samples onboard the
215 research vessel.

216

217 **2.3. Sample analysis**

218 Each laboratory measured dissolved methane and nitrous oxide slightly differently. A full
219 description of each laboratory's method can be found in Table S6 and Table S7 in the
220 Supplement for methane and nitrous oxide, respectively.

221 The majority of laboratories measured methane and nitrous oxide by equilibrating the
222 seawater sample with an overlying headspace and subsequently injecting a portion of the gaseous
223 phase into the gas analyzer. This method has been conducted since the 1960s when gas
224 chromatography was first used to quantify dissolved hydrocarbons (McAuliffe, 1963). The
225 headspace was created using helium, nitrogen, or high-purity air to displace a portion of the
226 seawater sample within the sample bottle. Alternatively, a subsample of the seawater was
227 transferred to a gas-tight syringe and the headspace gas subsequently added. The volume of the
228 vessel used to conduct the headspace equilibration ranged from 20 ml borosilicate glass vials to 1
229 L glass vials and syringes used by Newcastle University and U.S. Geological Survey,
230 respectively. The dissolved gases equilibrated with the overlying headspace at a controlled
231 temperature for a set period of time that ranged from 20 min to 24 h. The equilibration process
232 was typically enhanced by some initial period of physical agitation. After equilibration, an
233 aliquot of the headspace was transferred into the gas analyzer (GA) by either physical injection,
234 displacement using a brine solution, or injection using a switching valve. Some laboratories
235 incorporated a drying agent and a carbon dioxide scrubber prior to analysis. The gas sample
236 passed through a multi-port injection valve containing a sample loop of known volume, which
237 transferred the gas sample directly onto the analytical column within the oven of the GA.
238 Calibration of the instrument was achieved by passing the gas standards through the injection
239 valve.

240 The final gas concentrations using the headspace equilibration method was calculated by:

241

$$242 \quad [1] \quad C_{gas} [\text{nmol L}^{-1}] = \left(\beta x P V_{wp} + \frac{xP}{RT} V_{hs} \right) / V_{wp}$$

243

244 where β is the Bunsen solubility of nitrous oxide (Weiss and Price, 1980) or methane
245 (Wiesenburg and Guinasso, 1979) in $\text{nmol L}^{-1} \text{ atm}^{-1}$, x is the dry gas mole fraction (ppb)
246 measured in the headspace, P is the atmospheric pressure (atm), V_{wp} is the volume of water
247 sample (mL), V_{hs} is the volume (mL) of the created headspace, R is the gas constant (0.08205746

248 L atm K⁻¹mol⁻¹), and T is equilibration temperature in Kelvin (K). An example calculation is
249 provided in Table S8 in the Supplement.

250 In contrast to the headspace equilibrium method, five laboratories used a purge-and-trap
251 system for methane and/or nitrous oxide analysis (Table S6 and Table S7 in the Supplement).
252 These systems were directly coupled to a Flame Ionization Detector (FID) or ECD, with the
253 exception of University of British Columbia, where a quadrupole mass spectrometer with an
254 electron impact ion source and Faraday cup detector were used (Capelle et al., 2015). The
255 purge-and-trap systems were broadly similar, each transferring the seawater sample to a sparging
256 chamber. Sparging times typically ranged from 5-10 min and the sparge gas was either high
257 purity helium or high purity nitrogen. **In addition to commercially available gas scrubbers,**
258 **purification of the sparge gas was achieved by passing it through stainless steel tubing packed**
259 **with Poropak Q and immersed in liquid nitrogen. This is a recommended precaution to**
260 **consistently achieve a low blank signal of methane.** The elutant gas was dried using Nafion or
261 Drierite, and subsequently cryotrapped on a sample loop packed with Poropak Q to aid retention
262 of methane and nitrous oxide. **Cryotrapping was achieved for methane using liquid nitrogen (-**
263 **195°C) and either liquid nitrogen or cooled ethanol (-70°C) for nitrous oxide.** Subsequently, the
264 valve was switched to inject mode and the sample loop was rapidly heated to transfer its contents
265 onto the analytical column. Calibration was achieved by injecting standards via sample loops
266 using multi-port injection valves. Injection of standards upstream of the sparge chamber allowed
267 for calibration of the purge-and-trap gas handling system, in addition to the GA. Calculation of
268 the gas concentrations using the purge-and-trap method was achieved by application of the ideal
269 gas law to the standard gas measurements:

270 [2] $PV = nRT$

271 where P , R , and T are the same as Equation 1, V represents the volume of gas injected (L),
272 and n represents moles of gas injected. Rearranging Equation 2 yields the number of moles of
273 methane or nitrous oxide gas for each sample loop injection of compressed gas standards. These
274 values were used to determine a calibration curve based on the measured peak areas of the
275 injected standards, and thereafter derive the number of moles measured for each unknown
276 sample. To calculate concentrations of methane or nitrous oxide in a water sample, the number
277 of moles measured were divided by the volume (L) of seawater sample analyzed. An example
278 calculation is provided in Table S8 in the Supplement.

279

280 **2.4 Data analysis**

281 The final concentrations of methane and nitrous oxide are reported in nmol kg^{-1} . The analytical
282 precision for each batch of samples obtained by each of the individual laboratories was estimated
283 from the analysis of replicate seawater samples and reported as the coefficient of variation (%).
284 The values reported by each laboratory for all the batches of seawater samples are shown in
285 Tables S1 to S4 in the Supplement. Due to the observed inter-laboratory variability, it is likely
286 that the median value of methane and nitrous oxide for each batch of samples does not represent
287 the absolute *in situ* concentration. As this complicates the analytical accuracy for each
288 laboratory, we instead calculated the percentage difference between the median concentration
289 determined for each set of samples and the mean value reported by an individual laboratory. The
290 presence of outliers was established using the Interquartile Range (IQR) and by comparing with
291 one standard deviation applied to the overall median value.

292

293 **3. Results**

294 **3.1 Comparison of methane and nitrous oxide gas standards**

295 Six laboratories compared their existing ‘in-house’ standards of methane with the SCOR
296 standards. This was done by calibrating in-house standards and deriving a mixing ratio for the
297 SCOR standards which were treated as unknowns. Four laboratories reported methane values for
298 either the ARS or WRS within 3% of their absolute concentration, whereas two laboratories
299 reported an offset of 6% and 10% between their in-house standards and the SCOR standards
300 (Table S6 in the Supplement). For those laboratories who measured the SCOR standards to
301 within 3% or better accuracy, observed offsets in methane concentrations from the overall
302 median cannot be due to the calibration gas.

303 Seven laboratories compared their own in-house standards of nitrous oxide with the prepared
304 SCOR standards. Six laboratories reported values of nitrous oxide for the ARS which were
305 within 3% of the absolute concentration, with the remaining laboratory reporting an offset of
306 10% (Table S7 in the Supplement). The majority of these laboratories (five out of six groups)
307 compared the SCOR ARS with NOAA GMD standards, which have a balance gas of air instead
308 of nitrogen. Some laboratories with analytical systems that incorporated fixed sample loops (*e.g.*
309 1 or 2 ml loops housed in a 6-port or 10-port injection valve) had difficulty analyzing the WRS,

310 as the peak areas created by the high mole fraction of the standard exceeded the signal typically
311 measured from in-house standards or acquired by sample analysis, by an order of magnitude.
312 The high mole fraction of the WRS was not an issue when multiple sample loops of varying
313 sizes were incorporated into the analytical system, which was the case for purge-and-trap based
314 designs. **For the two laboratories with an in-house standard of comparable mole fraction to the**
315 **WRS, an offset of 3% and a >20% offset was reported.**

316

317 **3.2 Methane concentrations in the intercomparison samples**

318 Overall, median methane concentrations in seawater samples collected from the Pacific Ocean
319 and the Baltic Sea ranged from 0.9 to 60.3 nmol kg⁻¹ (Table 2). Out of 101 reported values, 3
320 outliers were identified using the IQR criterion and were not included in further analysis. The
321 methane data values for each batch of samples analyzed by each laboratory, including the mean
322 and standard deviation, the number of samples analyzed, and the % offset from the overall
323 median value are reported in Table S1 and Table S2 in the Supplement. **Analysis conducted by**
324 **the University of Hawai'i of methane and nitrous oxide from each Niskin-like bottle used in the**
325 **Pacific Ocean sampling did not reveal any bottle-to-bottle differences. Furthermore, analysis by**
326 **Newcastle University showed there was no difference between the first and the last set of**
327 **samples collected from the 1000 L collection used in the Baltic Sea sampling.**

328 The two Pacific Ocean sampling sites had the lowest water-column concentrations of
329 methane (Fig. 1a and 1b). The PAC1 samples collected from within the mesopelagic zone,
330 where methane concentrations have been reported to be less than 1 nmol kg⁻¹ (Reeburgh, 2007;
331 Wilson et al., 2017), showed a distribution of reported concentrations skewed towards the higher
332 values. For the PAC1 samples, seven out of twelve laboratories reported values ≤ 1 nmol kg⁻¹
333 and the mean coefficient of variation for all laboratories was 11% (Table 2). In contrast to the
334 mesopelagic samples, the methane concentrations for the near-surface seawater samples (PAC2)
335 were close to atmospheric equilibrium (Fig. 1b). Measured concentrations of methane for PAC2
336 samples ranged from 1.9 to 3.8 nmol kg⁻¹ and the mean coefficient of variation for all
337 laboratories was 7%. Similar to the PAC1 samples, PAC2 also had a distribution of data skewed
338 towards the higher concentrations.

339 Three Baltic Sea sampling sites (BAL1, BAL3, and BAL6) had median methane
340 concentrations that ranged from 4.1 to 5.7 nmol kg⁻¹ (Fig. 1c). The BAL1 samples also showed a

341 skewed distribution of reported values towards higher concentrations, as seen in PAC1 and
342 PAC2 samples. However, this was not evident in BAL3 or BAL6, which had the closest
343 agreement between the reported methane concentrations. For these three sets of Baltic Sea
344 samples, the mean coefficient of variation for all laboratories ranged from 4% (BAL3) to 9%
345 (BAL1). The next three Baltic Sea samples (BAL4, BAL5, and BAL7) had methane
346 concentrations that ranged from 18.8 to 35.4 nmol kg⁻¹ (Fig. 1d). These three sets of samples
347 had a normal distribution of data and the closest agreement between the reported concentrations
348 for all of the Pacific Ocean and Baltic Sea samples. Furthermore, for these three sets of samples,
349 the mean coefficient of variation for all laboratories was 4% (Table 2). The final Baltic Sea
350 sample (BAL2) had the highest concentrations of methane, with a median reported value of 60.3
351 nmol kg⁻¹, and a large range of values (45.2 to 67.2 nmol kg⁻¹; Fig. 1e). The BAL2 samples had
352 the lowest overall mean coefficient of variation for all laboratories; 2% (Table 2).

353 Further analysis of the data was conducted to better comprehend the factors that caused the
354 observed inter-laboratory variability in methane measurements. The deviation from median
355 values was calculated for each sample collected from the Baltic Sea (Fig. 2). The Pacific Ocean
356 samples (PAC1 and PAC2) were not included in this analysis due to the skewed distribution of
357 data. There were also some instances in the Baltic Sea samples, where the median concentration
358 might not have realistically represented the absolute *in situ* methane concentration. This was
359 most likely to have occurred at low concentrations due to the skewed distribution of reported
360 concentrations (*e.g.* BAL1) or at high concentrations where there was a large range in reported
361 values (*e.g.* BAL2). The results revealed that a few laboratories (Datasets D, F, and G) were
362 consistently within or close to 5% of the median value for all batches of seawater samples (Fig.
363 2). Some laboratories (*e.g.* Datasets B, C, and H) had a higher deviation from the median value
364 at higher methane concentrations. Two laboratories (Datasets J and K) had a higher deviation
365 from the median value at lower methane concentrations. Finally, in some cases it was not
366 possible to determine a trend (Datasets A and E), due to the variability.

367 The reasons behind the trends for each dataset became more apparent when considering the
368 effect of the inclusion or exclusion of low standards in the calibration curve on the resulting
369 derived concentrations (Fig. 3). The FID has a linear response to methane at nanomolar values
370 and therefore a high level of accuracy across a relatively wide range of *in situ* methane
371 concentrations can be obtained with the correct slope and intercept. To demonstrate this,

372 calibration curves for methane were provided by the University of Hawai'i. These revealed
373 minimal variation in the slope value when calibration points were increased from low mole
374 fractions (Fig. 3a) to higher mole fractions (Fig. 3b). However, the intercept value was sensitive
375 to the range of calibration values used, and this effect was further exacerbated when only the
376 higher calibration points were included (*i.e.* Fig. 3c). **The relevance to final methane**
377 **concentrations is demonstrated by considering the values reported by the University of Hawai'i**
378 **for PAC2 samples (Fig. 1b). An almost 30% increase in final methane concentration occurs**
379 **from the use of the calibration equation in Figure 3c, compared to Figure 3a. This derives from a**
380 **measured peak area for methane of 62 for a sample with a volume of 0.076 L and a seawater**
381 **density of 1024 kg m⁻³, yielding a final methane concentration of 2.1 and 2.8 nmol kg⁻¹ using the**
382 **equations from Figure 3a and 3c, respectively. With this understanding on the effect of FID**
383 **calibration, we consider it likely that the increased deviation from median values at high methane**
384 **concentrations (Datasets B, C, and H) results from differences in calibration slope between each**
385 **laboratory. In contrast, the datasets with a higher offset at low methane concentrations (Datasets**
386 **J and K) could be due to erroneous low standard values causing a skewed intercept. In addition,**
387 **there may be other factors including sample contamination, discussed in Section 3.4.**

388

389 **3.3 Nitrous oxide concentrations in the intercomparison samples**

390 Overall, median nitrous oxide concentrations in seawater samples collected from the Pacific
391 Ocean and the Baltic Sea ranged from 3.4 to 42.4 nmol kg⁻¹ (Table 2). Of the 113 reported
392 values, ten outliers were identified using the IQR criterion and were not included in further
393 analysis. The nitrous oxide data values for each batch of samples analyzed by each laboratory,
394 including the mean and standard deviation, the number of samples analyzed, and the % offset
395 from the overall median value are reported in Table S3 and Table S4 in the Supplement.

396 For six sets of seawater samples, BAL1, BAL2, BAL3, BAL6, BAL7, and PAC2, the
397 concentrations of nitrous oxide were close to atmospheric equilibrium. The reported values
398 ranged from 7.7 to 12.7 nmol kg⁻¹ in the Baltic Sea (Fig. 4a) and from 5.9 to 7.6 nmol kg⁻¹ in the
399 Pacific Ocean (Fig. 4b). For the Pacific Ocean near-surface (**mixed layer**) sampling site (PAC2),
400 the theoretical value of nitrous oxide concentration in equilibrium with the overlying atmosphere
401 is also shown (Fig. 4b). For these six samples with concentrations close to atmospheric

402 equilibrium, the mean coefficient of variation for all laboratories ranged from 3% (BAL3 and
403 PAC2) to 5% (BAL1) (Table 2).

404 For the three other sets of samples (BAL4, BAL5, and PAC1), the nitrous oxide
405 concentrations deviated significantly from atmospheric equilibrium (Fig. 4c, 4d, and 4e). At one
406 sampling site, BAL4 (Fig. 4c), nitrous oxide was under-saturated with respect to atmospheric
407 equilibrium and reported concentrations ranged from 2.1–5.5 nmol kg⁻¹. As observed in the low
408 concentration Pacific Ocean methane samples, there was a skewed distribution of the data
409 towards the higher nitrous oxide concentrations. The BAL4 samples also had the highest
410 variability (*i.e.* lowest precision), with a mean coefficient of variation of 8% (Table 2). The two
411 remaining samples (PAC1 and BAL5) had much higher concentrations of nitrous oxide, as
412 expected for low-oxygen regions of the water-column. In contrast to the samples with near
413 atmospheric equilibrium concentrations of nitrous oxide, there was a low overall agreement
414 between the independent laboratories for PAC1 and BAL5 nitrous oxide concentrations (Fig. 4d,
415 4e). At PAC1 and BAL5, nitrous oxide concentrations ranged from 34.3–45.8 nmol kg⁻¹ (Fig.
416 4d) and 30.1–45.9 nmol kg⁻¹, respectively (Fig. 4e). The mean coefficient of variation for all
417 laboratories was 4% for BAL5 samples compared to 3% for PAC1 samples.

418 The deviation of individual nitrous oxide concentrations from the median value provides
419 insight into the variability associated with their measurements (Fig. 5). The BAL1 dataset was
420 not included in this analysis due to its skewed data distribution and the high inter-laboratory
421 variability for BAL5 indicated that the median value may differ from the absolute nitrous oxide
422 concentration for this sample. For the low nitrous oxide Baltic Sea and Pacific Ocean samples
423 (Fig. 5a), the majority of data points were within 5% of the median values. Furthermore, for the
424 majority of laboratories, the data points for separate seawater samples clustered together
425 indicating some consistency to the extent they varied from the overall median value. Exceptions
426 to this observation include Datasets E, C, L, and K (Fig. 5a) which demonstrated varying
427 precision and accuracy. At high nitrous oxide concentrations (Fig. 5b), there are fewer data
428 points within 5% of the median value compared to low nitrous oxide concentrations (Fig. 5a).
429 Therefore, for PAC1 and BAL5 samples, 6 and 7 data points fall within 5% of the median value,
430 respectively. Furthermore, only three laboratories (Datasets F, G, and K) had data for both
431 Pacific Ocean and Baltic Sea samples within 5% of the median value. This could have been

432 caused by inconsistent analysis between different batches of samples or by variable sample
433 collection and transportation.

434 The likely factors that caused these offsets in nitrous oxide concentrations among
435 laboratories include sample analysis and calibration of the gas analyzers. Calibration of the ECD
436 is nontrivial and at least two prior publications have discussed nitrous oxide calibration issues
437 (Butler and Elkins, 1991; Bange et al., 2001). The laboratories participating in the nitrous oxide
438 intercomparison employed different calibration procedures (Fig. 6). Some used a linear fit and
439 kept their analytical peak areas within a narrow range (Fig. 6a), while others used a step-wise
440 linear fit and therefore used different slopes for low and high nitrous oxide mole fractions (Fig.
441 6b). Finally, some applied a polynomial curve (Fig. 6c) and sometimes two different polynomial
442 fits, for low and high concentrations. The difficulty in calibrating the ECD was evidenced by the
443 deviation from median values as multiple datasets show good precision but consistent offsets at
444 the lowest (Fig. 5a) and highest (Fig. 5b) final concentrations of nitrous oxide.

445

446 **3.4 Sample storage and sample bottle size**

447 Because prolonged storage of samples can influence dissolved gas concentrations, including
448 methane and nitrous oxide, the intercomparison dataset was analyzed for sample storage effects
449 (Table S5 in the Supplement). It should, however, be noted that assessing the effect of storage
450 time on sample integrity was not a formal goal of the intercomparison exercise and replicate
451 samples were not analyzed at repeated intervals by independent laboratories, as would normally
452 be required for a thorough analysis. Nonetheless our results did provide some insights into
453 potential storage-related problems. Most notably, there were indications that an increase in
454 storage time caused increased concentrations and increased variability for methane samples with
455 low concentrations, *i.e.* PAC1 and PAC2 samples which had median methane concentrations of
456 0.9 and 2.3 nmol kg⁻¹, respectively (Fig. 7). In comparison, for samples of nitrous oxide with
457 low concentrations there was no trend of increasing values as observed for samples with low
458 methane concentrations.

459 Another variable which differed between laboratories for the intercomparison exercise was
460 the size of samples bottle, which ranged from 25 ml to 1 liter for the different laboratories.
461 There was no observed difference between the methane and nitrous oxide values obtained from
462 the various sampling bottles and it was concluded that sampling bottles were not a controlling

463 factor for the observed differences between laboratories. We note, however, the potential for
464 greater air bubble contamination in smaller bottles.

465

466 **4. Discussion**

467 The marine methane and nitrous oxide analytical community is growing. This is reflected in the
468 increasing number of corresponding scientific publications and the resulting development of a
469 global database for methane and nitrous oxide (Bange et al., 2009). Like all Earth observation
470 measurements, there is a need for intercomparison exercises of the type reported here, for data
471 quality assurance, and for appropriate reporting practices (National Research Council, 1993). To
472 the best of our knowledge, the work presented here is the first formal intercomparison of
473 dissolved methane and nitrous oxide measurements. Based on our results, we discuss the lessons
474 learned and our recommendations moving forward, by addressing the four questions that were
475 posed in the Introduction.

476

477 **4.1 What is the agreement between the SCOR gas standards and the ‘in-house’ gas 478 standards used by each laboratory?**

479 It is typical for laboratories to source some, or all, of their compressed gas standards from
480 commercial suppliers. National agencies, such as NOAA GMD or National Institute of
481 Metrology China, also provide standards to the scientific community. The national agencies
482 typically offer a lower range in concentrations than commercial suppliers, but their standards
483 tend to have a higher level of accuracy. Of the twelve laboratories participating in the
484 intercomparison, eight reported using national agency standards, with seven of them using gases
485 sourced from NOAA GMD. Since the methane and nitrous oxide mole fractions of these
486 national agency standards are equivalent to modern-day atmospheric mixing ratios, they are
487 similar to the SCOR ARS distributed to the majority of laboratories in this study. Laboratories
488 in receipt of the SCOR standards were asked to predict their mole fractions based on those of
489 their own in-house standards. For the majority that conducted this exercise, there was good
490 agreement (<3% difference) between the NOAA GMD and the SCOR ARS for both methane
491 and nitrous oxide. For three laboratories, a larger offset was observed between the NOAA GMD
492 and the SCOR ARS. There was also a good prediction for the higher methane content SCOR
493 WRS, facilitated by the linear response of the FID (Fig. 3). In contrast, the nitrous oxide mole

494 fraction in the SCOR WRS exceeded the typical working range for several laboratories and it
495 was difficult for them to cross-compare with their in-house standards. This reflects an analytical
496 set-up that involves on-column injection via a 6-port or 10-port valve with one or two sample
497 loops, respectively. The sample loops have a fixed volume and their inaccessibility makes it
498 difficult to replace them by a smaller loop size. Therefore either dilution of the standard is
499 required, or smaller loops need to be incorporated into the calibration protocol. The two
500 laboratories that compared their in-house standards with the SCOR WRS reported an offset of
501 3% and >20%. This indicates that variability between standards can be an issue for obtaining
502 accurate dissolved concentrations and provides support for the production of a widely available
503 high concentration nitrous oxide standard. We strongly recommend that all commercially
504 obtained standards are cross-checked against primary standards, such as the SCOR ARS and
505 WRS. This should be conducted at least at the beginning and end of their use to detect any drift
506 that may have occurred during their lifetime. With due diligence and care, the SCOR standards
507 provide the capability for cross-checking personal standards for years to decades (Bullister et al.,
508 2016).

509

510 **4.2 How do measured values of methane and nitrous oxide compare across laboratories?**

511 **Methane:** The methane intercomparison highlighted the variability that exists between
512 measurements conducted by independent laboratories. At low methane concentrations, a skewed
513 distribution of methane data was observed, which was particularly evident in PAC1 (Fig. 1a).
514 Potential causes include calibration procedures (Section 3.2) and/or sample contamination which
515 is more prevalent at low concentrations (Section 3.4). For some laboratories, the low methane
516 concentrations are close to their detection limit, which is determined by the relatively low
517 sensitivity of the FID and the small number of moles of methane in an introduced headspace
518 **equilibrated** with seawater. An approximate working detection limit for methane analysis via
519 headspace equilibration is 1 nmol kg^{-1} , although some laboratories improve upon this by having
520 a large aqueous: gaseous phase ratio during the equilibration process (*e.g.* Upstill-Goddard et al.,
521 1996). Depending upon the volume of sample analyzed, purge-and-trap analysis can have a
522 detection limit much lower than 1 nmol kg^{-1} (*e.g.* Wilson et al., 2017). Methane measurements
523 in aquatic habitats with methane concentrations near the limit of analytical detection include
524 mesopelagic and high latitude environments distal from coastal or benthic inputs (*e.g.* Rehder et

525 al., 1999; Kitidis et al., 2010; Fenwick et al., 2017). Of additional concern is that the skewed
526 distribution of methane concentrations also occurs in samples collected both from the surface
527 ocean (PAC2; Fig. 1b) and coastal environments (BAL1; Fig. 1c). Methane concentrations
528 between 2–6 nmol kg⁻¹ are within the detection limit of all participating laboratories. To address
529 this we recommend that laboratories restrict sample storage to the minimum time required to
530 analyze the samples and incorporate internal controls into their sample analysis (Section 4.4).

531 There was an improvement in the overall agreement between the laboratories for samples
532 with higher methane concentrations. However, some of the highest variability between the
533 laboratories was observed at the highest concentrations of methane analyzed (BAL2; Fig. 1e).
534 This high degree of variability resulted in significant uncertainty in the absolute *in situ*
535 concentration. Methane concentrations of this magnitude and higher are found in coastal
536 environments (Zhang et al., 2004; Jakobs et al., 2014; Borges et al., 2017) and in the water-
537 column associated with seafloor emissions (*e.g.* Pohlman et al., 2011). These environments are
538 considered vulnerable to climate induced changes and eutrophication, and therefore it is
539 necessary that independent measurements are conducted to the highest possible accuracy to
540 allow for inter-laboratory and inter-habitat comparisons. To address this we recommend that
541 reference material be produced and distributed between laboratories.

542
543 **Nitrous oxide:** Some of the trends discussed for methane were also evident in the nitrous oxide
544 data. For the samples with the lowest nitrous oxide concentrations a skewed data distribution
545 was observed, as found for methane (Fig. 4c). Such low nitrous oxide concentrations are typical
546 of low-oxygen water-column environments (<10 μmol kg⁻¹). Therefore, the analytical bias
547 towards measuring values higher than the absolute *in situ* concentrations is particularly pertinent
548 to oceanographers measuring nitrous oxide in oxygen minimum zones and other low-oxygen
549 environments (Naqvi et al., 2010; Farías et al., 2015; Ji et al., 2015). The low concentrations of
550 nitrous oxide still exceed detection limits by at least an order of magnitude for even the less-
551 sensitive headspace method due to the high sensitivity of the ECD. Therefore, the bias towards
552 reporting elevated values for low concentrations of nitrous oxide is related less to analytical
553 sensitivity and is more a consequence of calibration issues. During the intercomparison exercise
554 ECD calibration was identified as a nontrivial issue for all participating laboratories and it
555 deserves continuing attention. In particular, the nonlinearity of the ECD means that low and

556 high nitrous oxide concentrations are more vulnerable to error since the values fall outside of the
557 most frequented part of the calibration curve. This is particularly true if a linear fit is used to
558 calibrate the ECD (Fig. 6a). To circumvent this problem, one laboratory used a step-wise linear
559 function while other laboratories used a quadratic function. The usefulness of multiple
560 calibration curves for low and high nitrous oxide concentrations was highlighted during the
561 intercomparison exercise, although this necessitates some consideration of the threshold for
562 switching **between** different calibration curves.

563 The majority of seawater samples analyzed had nitrous oxide concentrations ranging from 7–
564 11 nmol kg⁻¹ (Fig. 4a, 4b), which are close to atmospheric equilibrium values, as shown for the
565 Pacific Ocean (Fig. 4b). Collective analysis of these samples gives insight into the precision and
566 accuracy associated with surface water nitrous oxide analysis (Fig 5a). This is discussed further
567 in the context of implementing internal controls for methane and nitrous oxide (Section 4.4). For
568 samples with the highest nitrous oxide concentrations, *i.e.* exceeding 30 nmol kg⁻¹, there was
569 high variability between the concentrations reported by the independent laboratories. This was
570 most evident for the BAL5 samples (Fig. 4e) and similar to the variability observed at the highest
571 methane concentrations analyzed (Fig. 1e). It is difficult to assess how much of this variability
572 was specifically due to the differences in calibration practices between the laboratories and the
573 differences in gas standards with high nitrous oxide mole fractions, but at least some of it can be
574 attributed to this. These results form the basis for a proposed production of reference material
575 for both trace gases.

576

577 **4.3 Are there general recommendations to reduce uncertainty in the accuracy and** 578 **precision of methane and nitrous oxide measurements?**

579 There are several analytical recommendations resulting from this study. The use of highly
580 accurate standards and the appropriate calibration fit is an essential requirement for both
581 headspace equilibration and the purge-and-trap technique. It was shown that both analytical
582 approaches can yield comparable values for methane and nitrous oxide, with the main
583 differences observed at low methane concentrations. At sub-nanomolar methane concentrations,
584 four out of the six laboratories that reported methane concentrations <1 nmol kg⁻¹ used a purge-
585 and-trap analysis.

586 This study also revealed that sample storage time can be an important factor. Specifically,
587 the results from this study corroborate the findings of Magen et al. (2014) who showed that
588 samples with low concentrations of methane and more susceptible to increased values as a result
589 of contamination. The contamination was most likely due to the release of methane and other
590 hydrocarbons from the septa (Niemann et al., 2015). Since the release of hydrocarbons occurs
591 over a period of time, it is recommended to keep storage time to a minimum and to store samples
592 in the dark. It should be noted that sample integrity can also be compromised due to other
593 factors including inadequate preservation, outgassing, and adsorption of gases onto septa. For all
594 these reasons, it is recommended to conduct an evaluation of sample storage time for the
595 environment that is being sampled.

596 One useful item that was not included as part of the intercomparison exercise but can help
597 decrease uncertainty in the accuracy and precision of methane and nitrous oxide measurements
598 are internal control measurements. Internal controls represent a self-assessment quality control
599 check to validate the analytical method and quantify the magnitude of uncertainty. Appropriate
600 internal controls for methane and nitrous oxide consist of air-equilibrated seawater samples.
601 Their purpose is to provide checks for methane concentrations ranging from 2–3 nmol kg⁻¹ and
602 for nitrous oxide concentrations from 5–9 nmol kg⁻¹. The air used in the equilibration process
603 could be sourced from the ambient environment if sufficiently stable or from a compressed gas
604 cylinder after cross-checking the concentration with the appropriate gas standard. Air-
605 equilibrated samples provide reassurance that the analytical system is providing values within the
606 correct range. Air-equilibrated samples also indicate the certainty associated with calculating the
607 saturation state of the ocean with respect to atmospheric equilibrium. This is particularly
608 relevant when the seawater being sampled is within a few percent of saturation. Finally, these
609 air-equilibrated samples provide an estimate of analytical accuracy, which is infrequently
610 reported for methane or nitrous oxide. At present, only a few studies report the analysis of air-
611 equilibrated seawater alongside water-column samples (Bullister and Wisegarver, 2008; Capelle
612 et al., 2015; Bourbonnais et al., 2017; Wilson et al., 2017). It is likely that wider implementation
613 would facilitate internal assessment of the analytical system. Since the main equipment required
614 is a water-bath and an overhead stirrer, the production is not cost-prohibitive. A
615 recommendation of this intercomparison exercise is that laboratories routinely use air-
616 equilibrated seawater samples to provide an estimate of analytical accuracy.

617 In addition to the self-assessments provided by the analysis of air-equilibrated seawater, this
618 study revealed the need for reference seawater to help assess the accuracy of high concentration
619 methane and nitrous oxide measurements. Reference seawater in this instance refers to batches
620 of dissolved methane and nitrous oxide samples prepared in the laboratory using an equilibrator
621 set-up, as used for dissolved inorganic carbon (Dickson et al., 2007). In the absence of plans for
622 additional intercomparison exercises, the provision of reference seawater will allow laboratories
623 to continue evaluating their own measurements. **Finally, the lessons learned during the**
624 **intercomparison exercises will be the basis for a forthcoming Good Practice Guide for dissolved**
625 **methane and nitrous oxide.**

626

627 **4.4 What are the implications of interlaboratory differences for determining the spatial and** 628 **temporal variability of methane and nitrous oxide in the oceans?**

629 The key outcome of this study was the identification of differences in methane and nitrous oxide
630 concentrations for the same batch of seawater samples measured by several independent
631 laboratories. Emergent from this is the distinct possibility that any given laboratory will
632 incorrectly report data, thereby increasing uncertainty over the saturation states of both gases.
633 The tendency to over-estimate methane concentrations close to atmospheric equilibrium means
634 that marine emissions of methane to the overlying atmosphere will be also overestimated (Bange
635 et al., 1994; Upstill-Goddard and Barnes, 2016). In contrast, for nitrous oxide there does not
636 appear to be either an under-estimation or over-estimation of concentrations. Consequently, there
637 is generally a lower inherent uncertainty in its surface ocean saturation state, as previously
638 proposed (Law and Ling, 2001; Forster et al., 2009).

639 The inter-laboratory differences highlighted by this study should be viewed in the context of
640 numerous individual efforts to assess temporal and/or spatial trends in methane and nitrous oxide
641 by way of time-series observations (Bange et al., 2010; Farías et al., 2015; Wilson et al., 2017;
642 Fenwick and Tortell, 2018), repeat hydrographic survey lines (de la Paz et al., 2017), and single
643 expeditions. While the value of these in integrating the behaviour of methane and nitrous oxide
644 into the hydrography and biogeochemistry of local-regional ecosystems is beyond question, their
645 value would be enhanced by the rigorous cross-validation of analytical protocols. Without this,
646 perceived small temporal and/or spatial changes in water-column concentrations in any given
647 region are difficult to verify unless the data all originate from a single laboratory. In addition,

648 the value of a global methane and nitrous oxide database (*e.g* Bange et al., 2009) would to some
649 extent be compromised by the uncertainty. Taking due account of the analytical variability
650 between laboratories will clearly be vital to any future assessment of the changing methane and
651 nitrous oxide budgets of the oceans.

652

653 **5. Conclusions**

654 Overall, the intercomparison exercise was invaluable to the growing community of ocean
655 scientists interested in understanding the dynamics of dissolved methane and nitrous oxide in the
656 water-column. The level of agreement between independent measurements of dissolved
657 concentrations was evaluated in the context of several contributing factors, including sample
658 analysis, standards, calibration procedures, and sample storage time. Importantly, the
659 intercomparison represents a concerted effort from the scientists involved to critically assess the
660 quality of their data, and to initiate the steps required for further improvements.

661 Recommendations arising from the intercomparison include routine cross-calibration of working
662 gas standards against primary standards, minimizing sample storage time, incorporating internal
663 controls (air-equilibrated seawater) alongside routine sample analysis, and the future production
664 of reference seawater for methane and nitrous oxide measurements. These efforts will help
665 resolve temporal and spatial variability, which is necessary for constraining methane and nitrous
666 oxide emissions from aquatic ecosystems and for evaluating the processes that govern their
667 production and consumption in the water-column.

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850 **Table 1.** List of laboratories that participated in the intercomparison. All laboratories measured
 851 both methane and nitrous oxide except U.S. Geological Survey (methane only), U.C. Santa
 852 Barbara (nitrous oxide only), and NOAA PMEL (nitrous oxide from the Pacific Ocean). Also
 853 indicated are the twelve laboratories that received the SCOR gas standards of methane and
 854 nitrous oxide.

Institution	Lead Scientist	SCOR Standards
University of Hawai'i, USA	Samuel Wilson	Yes
GEOMAR, Germany	Hermann Bange	Yes
Newcastle University, UK	Robert Upstill-Goddard	Yes
Université de Liège, Belgium	Alberto Vieira Borges	No
Plymouth Marine Laboratory, UK	Andrew Rees	Yes
NOAA PMEL, USA	John Bullister	Yes
IIM-CSIC, Spain	Mercedes de la Paz	Yes
CACYTMAR, Spain	Macarena Burgos	No
University of Concepción, Chile	Laura Farías	Yes
IOW, Germany	Gregor Rehder	Yes
University of California Santa Barbara, USA	Alyson Santoro	Yes
National Institute of Water and Atmospheric Research, NZ	Cliff Law	Yes
University British Columbia, Canada	Philippe Tortell	Yes
U.S. Geological Survey, USA	John Pohlman	No
Ocean University of China, China	Guiling Zhang	Yes

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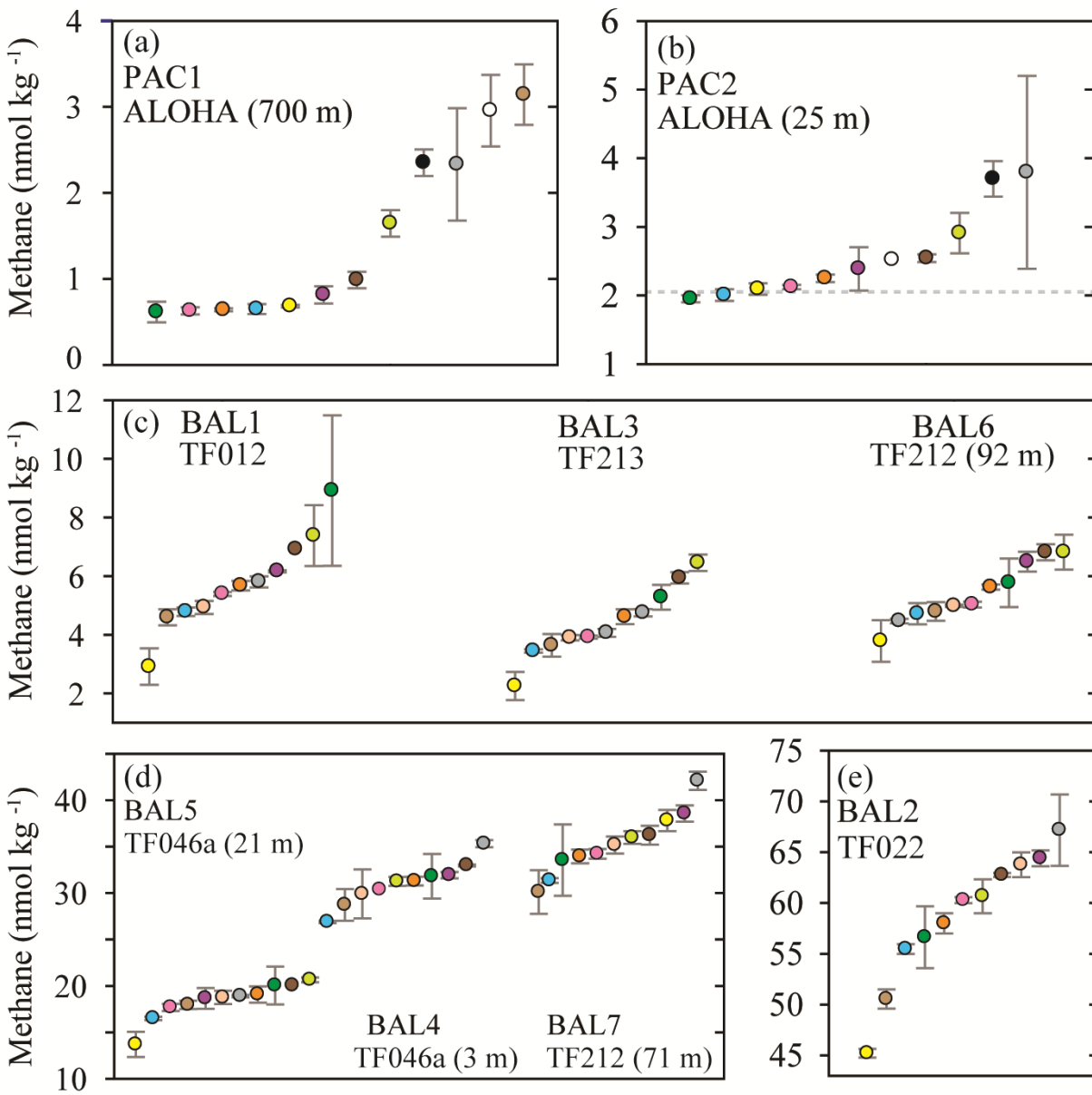
857 **Table 2.** Pertinent information for each batch of methane and nitrous oxide samples. This
 858 includes contextual hydrographic information, median and mean concentrations of methane and
 859 nitrous oxide, range, number of outliers, and the overall average coefficient of variation (%).
 860

Sampling parameters									
Sample ID	PAC1	PAC 2	BAL1	BAL2	BAL3	BAL4	BAL5	BAL6	BAL7
Location	22.75N 158.00W	22.75N 158.00W	54.32N 11.55E	54.11N 11.18E	55.25N 15.98E	55.30N 15.80E	55.30N 15.80E	54.47N 12.21E	54.47N 12.21E
Location name	Station ALOHA	Station ALOHA	TF012	TF022	TF213	TF212	TF212	TF046a	TF046a
Sampling date	24.2.17	24.2.17	16.10.16	17.10.16	18.10.16	19.10.16	20.10.16	21.10.16	21.10.16
Sampling depth (m)	25	700	3	22	3	92	71	3	21
Seawater temperature (°C)	23.6	5.1	12.0	13.6	12.2	6.6	6.7	11.8	13.4
Salinity	34.97	34.23	13.85	17.37	7.87	18.40	18.08	8.81	17.65
Density (kg m ⁻³)	1024	1027	1010	1013	1006	1014	1014	1006	1013
Nitrous oxide									
Number of datasets	13	13	12	13	12	13	12	13	12
Outliers	0	1	2	1	1	0	1	2	2
Median N ₂ O conc. (nmol kg ⁻¹)	42.4	7.0	11.0	9.4	11.1	3.4	40.2	11.0	9.6
Mean N ₂ O conc. (nmol kg ⁻¹)	41.3	7.0	11.1	9.2	11.0	3.4	39.0	10.8	9.5
Range	34.3-45.8	5.9-7.6	10.1-12.7	7.7-11.0	9.6-11.6	2.1-5.5	30.1-45.9	9.5-11.5	8.0-10.4
Average coeff. variation (%)	2.8	4.4	4.5	4.2	2.7	7.5	4.0	2.6	4.4
Methane									
Number of datasets	12	12	11	11	11	11	11	11	11
Outliers	0	1	0	0	0	1	1	0	0
Median CH ₄ conc. (nmol kg ⁻¹)	0.9	2.3	5.7	60.3	4.1	31.3	18.8	5.0	35.2
Mean CH ₄ conc. (nmol kg ⁻¹)	1.8	2.6	5.8	58.6	4.4	31.1	18.8	5.4	35.4
Range	0.6-3.1	1.9-3.8	2.9-8.9	45.2-67.2	2.5-6.5	26.9-35.3	16.5-20.7	3.8-6.8	30.1-42.1
Average coeff. variation (%)	10.9	7.2	8.6	2.1	4.3	3.5	4.2	6.5	3.5

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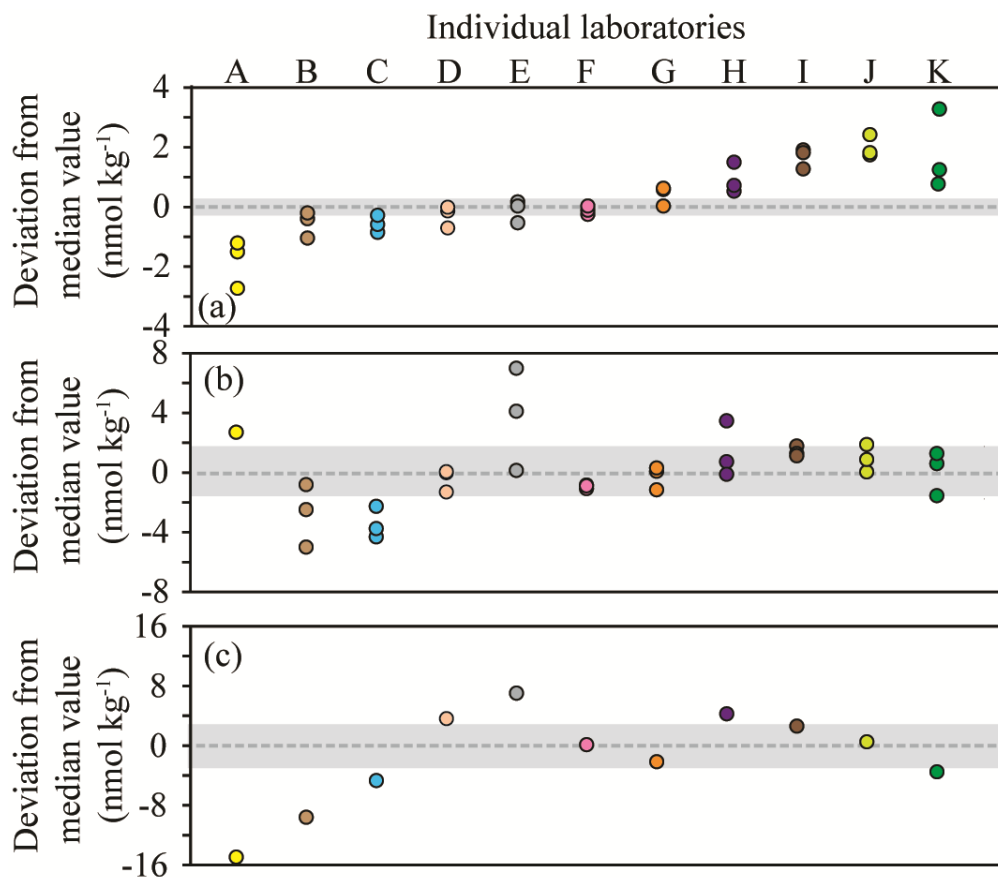
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863 **Figures**



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 865
 866 Figure 1. Concentrations of methane measured in nine separate seawater samples collected from
 867 the Pacific Ocean (Fig. 1a, 1b) and the Baltic Sea (Fig. 1c, 1d, 1e). The dashed grey line
 868 represents the value of methane at atmospheric equilibrium (Fig. 1b.) Individual data points are
 869 plotted sequentially by increasing value. The same color symbol is used for each laboratory in
 870 all plots.

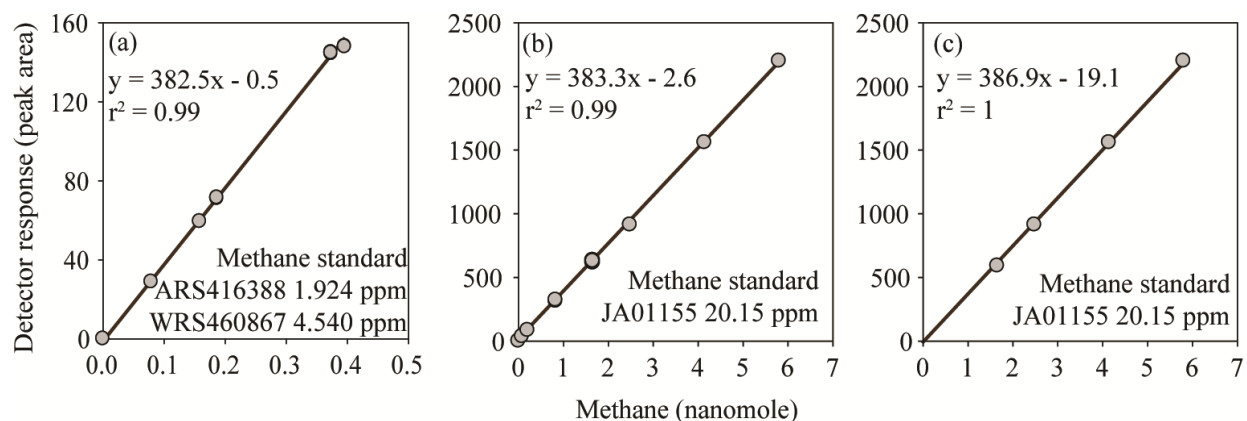
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 874 Figure 2. Deviation from the median methane concentration (reported as absolute values in nmol
 875 kg^{-1}) for the seven Baltic Sea samples. The batches of seawater samples include BAL1, BAL3,
 876 and BAL6 (Fig. 2a), BAL4, BAL5, and BAL7 (Fig. 2b), and BAL2 (Fig. 2c). The shaded grey
 877 area indicates values $\leq 5\%$ of the median concentration. The color scheme for each laboratory
 878 dataset is identical to that used in Figure 1 and the letters allocated to each dataset are to facilitate
 879 cross-referencing in the text. Note that the y-axis scale varies between the Figures.

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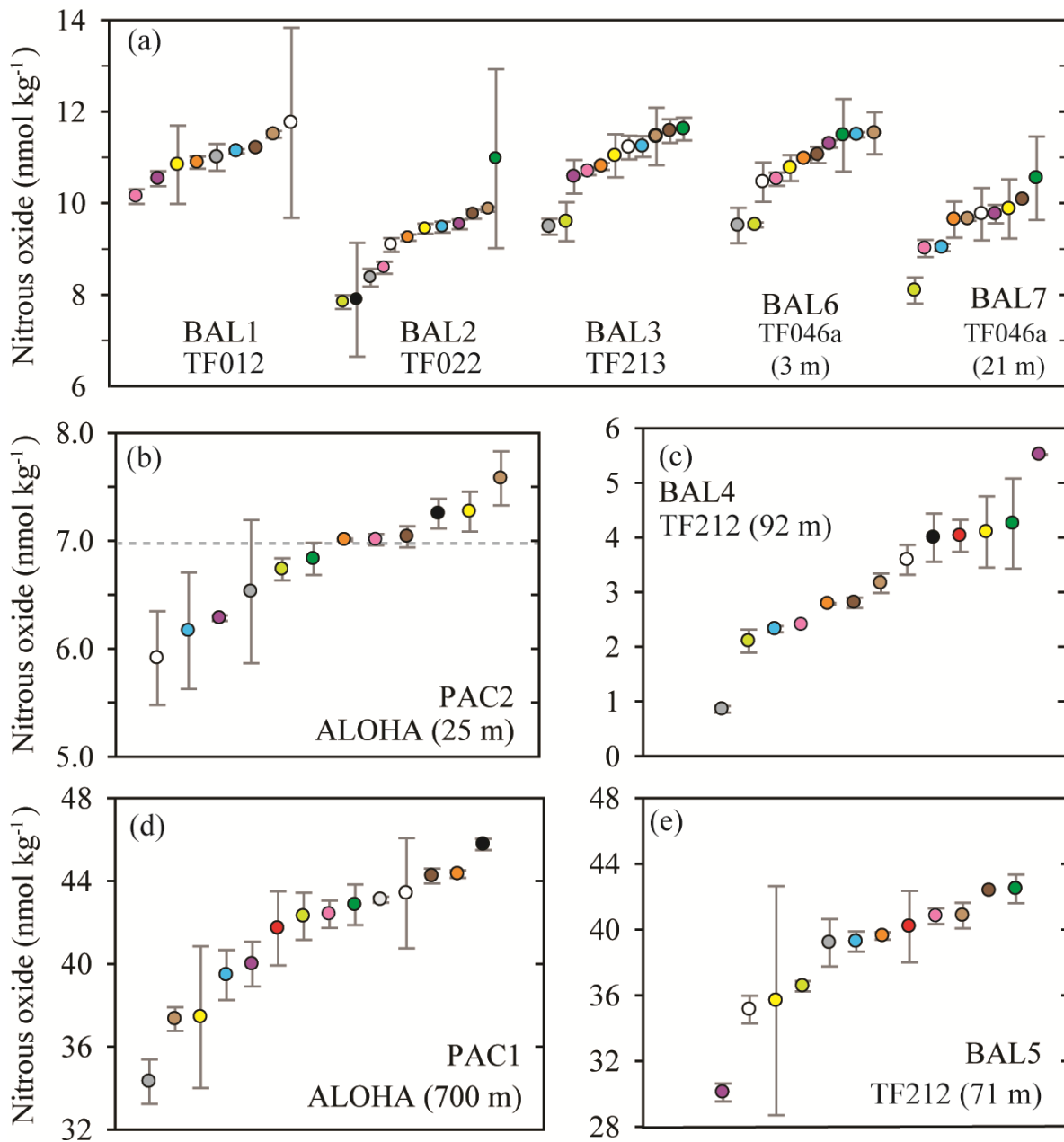
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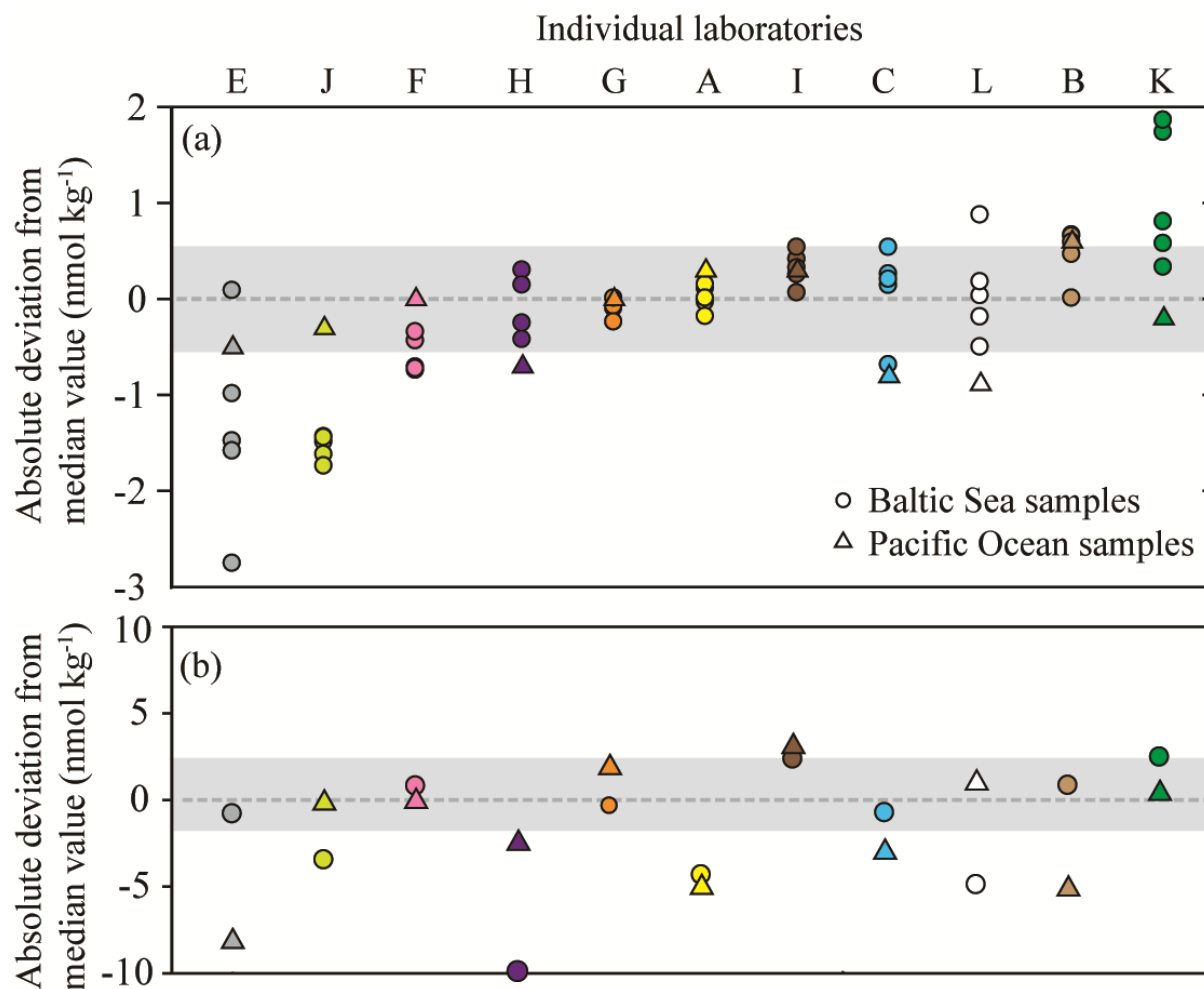
883 Figure 3. FID response to methane, fitted with a linear regression calibration. The inclusion
884 (Fig. 3a and Fig. 3b) or exclusion (Fig. 3c) of low methane values cause the calibration slope and
885 intercept to vary. However, the observed variation in the calibration slope does not have a
886 significant effect on the final calculated concentrations of methane. In contrast, variation in the
887 intercept does have an effect on the final concentrations of methane.

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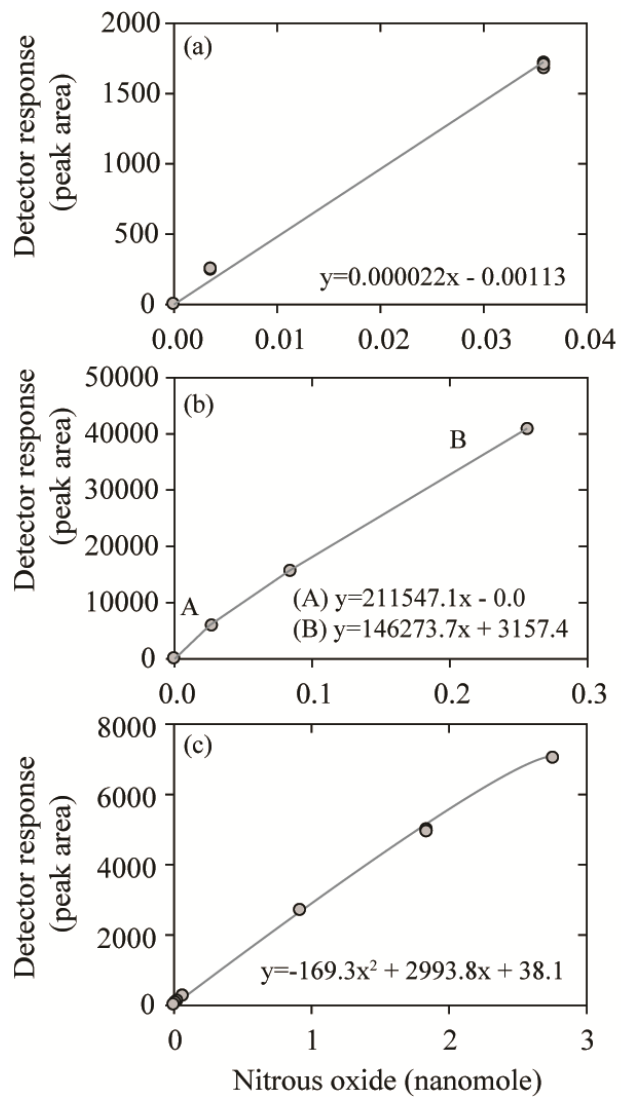
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Figure 4. Concentrations of nitrous oxide measured in nine separate samples from the Baltic Sea and the Pacific Ocean. The dashed grey line represents the value of nitrous oxide at atmospheric equilibrium (Fig. 4b). Individual data points are plotted sequentially by increasing value. The same color symbol is used for each laboratory in all plots.



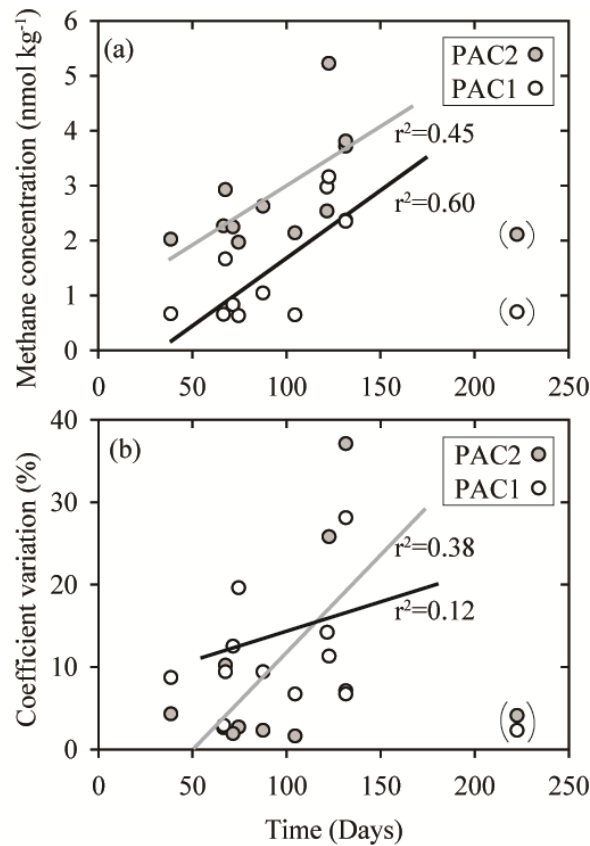
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898 Figure 5. Deviation from the median value (reported in absolute units) for nitrous oxide datasets.
 899 The batches of samples include BAL1,2,3,6,7 (Fig. 5a) and PAC2 and BAL5 (Fig. 5b). The
 900 Baltic Sea samples are represented by circles and the Pacific Ocean samples are represented by
 901 triangles. The shaded area indicates a deviation $\leq 5\%$ from the median value, based on a water-
 902 column concentration of 11 nmol kg^{-1} and 42 nmol kg^{-1} for Fig. 5a and 5b, respectively. The
 903 color scheme for each laboratory dataset is identical to that used in Figure 4 and the letters
 904 allocated to each dataset are to facilitate cross-referencing in the text. Note the y-axis for Fig 5a
 905 and 5b are plotted on a different scale.



906

907 Figure 6. Three calibration curves for nitrous oxide measurements using an ECD including linear
 908 (Fig. 6a), multilinear (Fig. 6b), and quadratic (Fig. 6c) fits.



910

911 Figure 7. Comparison of sample storage times with measured concentrations of methane (Fig.
 912 7a) and coefficient variation (Fig. 7b) for two sets of seawater samples (PAC1 and PAC2)
 913 **collected in February 2017**. These two sets of seawater samples had the lowest methane
 914 concentrations and appear to be influenced by the duration of storage time. The data points
 915 enclosed in parentheses were not included in the regression analysis. The PAC1 regression line
 916 is black and the PAC2 regression line is grey. All of the storage times are included in the
 917 Supplementary Material.