

1 **The Arctic picoeukaryote *Micromonas pusilla* benefits**
2 **synergistically from warming and ocean acidification**

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14

15 **Abstract**

16 In the Arctic Ocean, climate change effects such as warming and ocean acidification (OA) are
17 manifesting faster than in other regions. Yet, we are lacking a mechanistic understanding of the
18 interactive effects of these drivers on Arctic primary producers. In the current study, one of the
19 most abundant species of the Arctic Ocean, the prasinophyte *Micromonas pusilla*, was exposed
20 to a range of different pCO₂ levels at two temperatures representing realistic current and future
21 scenarios for nutrient-replete conditions. We observed that warming and OA synergistically
22 increased growth rates at intermediate to high pCO₂ levels. Furthermore, elevated temperatures
23 shifted the pCO₂-optimum of biomass production to higher levels. Based on changes in cellular
24 composition and photophysiology, we hypothesise that the observed synergies can be explained
25 by beneficial effects of warming on carbon fixation in combination with facilitated carbon
26 acquisition under OA. Our findings help to understand the higher abundances of picoeukaryotes
27 such as *M. pusilla* under OA, as has been observed in many mesocosm studies.

28 **1 Introduction**

29 With the progress in using molecular tools to describe marine biodiversity in the past decades,
30 the scientific community has become increasingly aware of the underestimated importance of
31 picoeukaryotes, for both primary and export production of the world's oceans (Richardson and
32 Jackson, 2007; Worden and Not, 2008). Larger phytoplankton such as diatoms are efficient
33 vectors for carbon export due to aggregate formation and ingestion by large zooplankton
34 leading to the production of fast-settling faecal pellets (Sherr et al., 2003). In contrast,
35 picoeukaryotes are mainly grazed by smaller heterotrophic protists such as ciliates, which have
36 a low carbon retention, excrete relatively more dissolved material, and thus fuel recycled
37 production (Sherr and Sherr, 2002). Hence, changes in the relative abundance of pico- and
38 nanoeukaryotes can have large implications for food webs and biogeochemistry (Worden et al.,
39 2015).

40 Picoeukaryotes tend to dominate low nutrient environments, which is often attributed to
41 their high surface:volume ratios and mixotrophic capacities (Raven, 1998; McKie-Krisberg and
42 Sanders, 2014). The low nutrient concentrations in the Arctic surface ocean, for example, cause
43 picoeukaryotes to be particularly successful in this region. In fact, the globally occurring
44 prasinophyte *Micromonas pusilla* is considered the most abundant species in the Arctic ocean
45 (Šlapeta et al., 2006; Lovejoy et al., 2007; Marquardt et al., 2016). In this environment, strong
46 stratification causes low nutrient concentrations throughout the summer and autumn months
47 (Tremblay et al., 2015), and the occurrence of the polar night requires organisms to either form
48 resting stages or to have heterotrophic capacities (Tremblay et al., 2009; Lovejoy, 2014; Berge
49 et al., 2015; Vader et al., 2015).

50 Climate change effects manifest faster in the Arctic than anywhere else on the planet
51 (Stocker, 2014). In this region, for example, temperatures are rising more than twice as fast as
52 at the rest of the globe (Miller et al., 2010). The concurrent rapid reduction in ice cover allows
53 for more light penetration and longer growing seasons, while increased stratification due to ice
54 melt and warming constrain nutrient supply to surface waters, both of which will change the
55 dynamics of primary production (Arrigo et al., 2008; Wassmann and Reigstad, 2011). Ocean
56 acidification (OA) is also especially pronounced in the Arctic Ocean, because low temperatures
57 and alkalinity make the system sensitive to anthropogenic CO₂ loading (AMAP, 2013; Qi et
58 al., 2017). Picoeukaryotes such as *M. pusilla* may benefit from these changes and are considered
59 potential winners of climate change. In the Canadian Arctic, for example, picoeukaryote
60 abundances are increasing as surface waters get warmer, fresher and more oligotrophic (Li et
61 al., 2009). Regarding OA effects, the majority of studies on natural phytoplankton assemblages

62 have shown picoeukaryotes, particularly *M. pusilla*, to increase in relative abundance with
63 increasing pCO₂ levels (Engel et al., 2008; Meakin and Wyman, 2011; Newbold et al., 2012;
64 Brussaard et al., 2013; Hussherr et al. 2017; Schulz et al., 2017). Despite the evident sensitivity
65 of *M. pusilla* to changes in pCO₂ levels, a detailed assessment of the OA effects, their
66 interaction with warming as well as the underlying mechanisms in this important species is still
67 missing.

68 Like all photosynthetic organisms, cells of *M. pusilla* need to maintain a balance
69 between energy sources (i.e. light harvesting by the photosynthetic apparatus) and sinks (most
70 importantly carbon fixation in the Calvin cycle) to prevent harmful levels of excitation pressure
71 on the photosynthetic electron transport chain (Behrenfeld et al., 2008). Light harvesting and
72 electron transport in the photosystems are largely independent of changes in temperature and
73 pCO₂ (Mock and Hoch, 2005; Hoppe et al., 2015), but the impact of these drivers on energy
74 sinks can potentially affect the energy balance of the cell: The beneficial effects of elevated
75 pCO₂ observed in phytoplankton are thought to be caused by increased diffusive CO₂ supply,
76 reduced CO₂ leakage, or by lowered costs to operate their CO₂ concentrating mechanisms (Rost
77 et al., 2008; Bach et al., 2013). Elevated temperatures, on the other hand, can change enzyme
78 kinetics including those involved in the Calvin cycle, thus leading to a larger sink of excitation
79 energy (Maxwell et al., 1994; Toseland et al., 2013). Hence, both ocean warming and
80 acidification potentially increase the efficiency of photosynthesis and biomass production, at
81 least up to the organisms' respective optimum levels. Above these levels, temperatures and
82 proton concentrations start to disrupt enzymatic processes, increase the investment into pH
83 homeostasis, and impair the delicate regulation of cellular processes (Levitt, 1980; Taylor et
84 al., 2001; Flynn et al., 2012). Thus, the complex balance between beneficial and detrimental
85 effects will determine whether the combination of warming and OA will synergistically
86 promote or deteriorate phytoplankton growth and biomass build-up.

87 In the current study, we aim to investigate the responses of an Arctic *M. pusilla* strain
88 to warming and OA. To this end, *M. pusilla* was grown at four pCO₂ levels ranging from
89 preindustrial to future scenarios (180-1400 μatm) under 2°C and 6°C, which represent the
90 magnitude of the projected future temperature increase in this region (Collins et al., 2013), but
91 also the current spring and summer temperatures in the environment where the strain was
92 isolated (Hegseth et al., in press).

93 **2 Material & Methods**

94

95 **2.1 Culture conditions**

96 Monoclonal cultures of the picoeukaryote *Micromonas pusilla* (Butcher) I. Manton & M. Parke
97 (isolated in 2014 by K. Wolf in Kongsfjorden, Svalbard, 79°N; local temperature range -1.5 to
98 8°C) taxonomic identification confirmed by rDNA sequencing of SSU, LSU and ITS
99 sequences) were grown in 1-L glass bottles in semi-continuous dilute-batch cultures (max
100 129,000 cells mL⁻¹; diluted every 3-4 days) under constant irradiances of 150 ± 26 μmol
101 photons m⁻² s⁻¹. Media consisted of 0.2 μm sterile-filtered Arctic seawater with a salinity of
102 32.7 enriched with macronutrients, trace metals and vitamins according to F/2_R medium
103 (Guillard and Ryther, 1962). Light intensities were provided by daylight lamps (Philips Master
104 TL-D 18W; emission peaks at wavelength of 440, 560 and 635 nm), adjusted by neutral density
105 screens and monitored using a LI-1400 data logger (Li-Cor) equipped with a 4 π -sensor (Walz).
106 Cells were growing at four different CO₂ partial pressures (pCO₂; 180, 380, 1000, and 1400
107 μatm) and two temperatures ($2.2 \pm 0.3^\circ\text{C}$ and $6.3 \pm 0.2^\circ\text{C}$). Cultures were acclimated to these
108 conditions for at least 7 generations prior to sampling.

109 Different pCO₂ conditions were achieved by aeration of the incubation bottles with air
110 of the respective pCO₂ levels delivered through sterile 0.2- μm air-filters (Midisart 2000,
111 Sartorius stedim) for 24 h prior to inoculation. Gas mixtures were generated using a gas flow
112 controller (CGM 2000 MCZ Umwelttechnik), in which CO₂-free air (<1 ppmv CO₂; Dominick
113 Hunter) was mixed with pure CO₂ (Air Liquide Deutschland). The pCO₂ levels in the gas
114 mixtures were regularly monitored with a non-dispersive infrared analyzer system (LI6252, LI-
115 COR Biosciences), calibrated with CO₂-free air and purchased gas mixtures of 150 ± 10 and
116 1000 ± 20 ppmv CO₂ (Air Liquide Deutschland).

117

118 **2.2 Carbonate chemistry**

119 Samples for total alkalinity (A_T) were filtered through 0.7- μm glass fibre filters (GF/F,
120 Whatman) and stored in borosilicate bottles at 3°C. A_T was estimated from duplicate
121 potentiometric titration (Brewer et al., 1986) using a TitroLine alpha plus (Schott Instruments).
122 A_T values were corrected for systematic errors based on measurements of certified reference
123 materials (CRMs provided by Prof. A. Dickson, Scripps, USA; batch #111; reproducibility ± 5
124 $\mu\text{mol kg}^{-1}$). Total dissolved inorganic carbon (C_T) samples were filtered through 0.2- μm
125 cellulose-acetate filters (Sartorius stedim) and stored in gas-tight borosilicate bottles at 3°C. C_T
126 was measured colorimetrically in triplicates with a QuAAtro autoanalyzer (Seal; Stoll et al.

127 2001). The analyser was calibrated with NaHCO_3 solutions (with a salinity of 35, achieved by
128 addition of NaCl) to achieve concentrations ranging from 1800 to 2300 $\mu\text{mol C}_T \text{ kg}^{-1}$. CRMs
129 were used for corrections of errors in instrument performance such as baseline drifts
130 (reproducibility $\pm 8 \mu\text{mol kg}^{-1}$). Seawater pH_{total} was measured potentiometrically with a two-
131 point calibrated glass reference electrode (IOline, Schott Instruments). An internal TRIS-based
132 reference standard (Dickson et al., 2007) was used to correct for variability on electrode
133 performance (reproducibility ± 0.015 pH units). Following recommendations by Hoppe et al.
134 (2012), seawater carbonate chemistry including pCO_2 was calculated from A_T and pH using
135 CO_2SYS (Pierrot et al., 2006). The dissociation constants of carbonic acid of Mehrbach et al.
136 (1973), as refitted by Dickson and Millero (1987), were used for calculations. Dissociation
137 constants for KHSO_4 were taken from Dickson (1990).

138

139 **2.3 Growth, elemental composition and production rates**

140 Samples for cell counts were fixed with glutaraldehyde (0.5% final concentration). After gentle
141 mixing, samples were stored at room temperature in the dark for 15 min, and subsequently
142 frozen in liquid nitrogen and stored at -80°C . Prior to analysis, samples were thawed on ice and
143 mixed thoroughly. After addition of 10 μL SybrGreen working solution (dissolved in DMSO)
144 and 10 μL YG beads working solution (1 μm -Flouresbrite calibration beads grade YG,
145 Polyscience), samples were counted on an Accuri C6 flow cytometer (BD Biosciences)
146 equipped with a blue solid-state laser (488 nm excitation wavelength) run on medium fluidics
147 settings (35 $\mu\text{L min}^{-1}$; 16 μm core size) with a limit of 50,000 events or 250 μL . Analysis was
148 performed based on red (FL3 channel, >670 nm) and green (FL1 channel, 533 ± 30 nm)
149 fluorescence, as well as sideward and forward light scattering. Specific growth rates constants
150 (μ) were determined from exponential fits of cell counts over 4 consecutive days.

151 Particulate organic carbon (POC) and nitrogen (PON) were measured after filtration
152 onto precombusted (15h, 500°C) GF/F filters (Whatman) and stored at -20°C . and dried for
153 at least 12 h at 60°C prior to sample preparation. Analysis was performed using a CHNS-O
154 elemental analyser (Euro EA 3000, HEKAtech). Contents of POC and PON were corrected for
155 blank measurements and normalised to filtered volume and cell densities to yield cellular
156 quotas. Production rates of POC were calculated by multiplying the cellular quota with the
157 division rate constant k of the respective incubation. Samples for determination of chlorophyll
158 a (Chl a) were filtered onto GF/F filters (Whatman), immediately placed into liquid nitrogen
159 and stored at -80°C until analysis. Chl a was subsequently extracted in 8 mL 90% acetone at

160 4°C over night. Chl *a* concentrations were determined on a fluorometer (TD-700, Turner
161 Designs), using an acidification step (1M HCl) to determine phaeopigments (Knap et al., 1996).

162

163 **2.4 Variable Chl *a* fluorescence**

164 Photophysiological characteristics, based on photosystem II (PSII) variable Chl *a* fluorescence,
165 were measured using a fast repetition rate fluorometer (FRRf; FastOcean PTX, Chelsea
166 Technologies) in combination with a FastAct Laboratory system (Chelsea Technologies). The
167 excitation wavelength of the fluorometer's light-emitting diodes (LEDs) was 450 nm, and the
168 applied light intensity was 21587 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$. The FRRf was used in single turnover
169 mode, with a saturation phase comprising 100 flashlets on a 2 μs pitch and a relaxation phase
170 comprising 40 flashlets on a 50 μs pitch. Measurements from all replicates (n=3) were
171 conducted in a temperature-controlled chamber ($\pm 0.2^\circ\text{C}$) at the respective treatment
172 temperature.

173 After subtraction of a blank value, the minimum (F_0 and F_0' for light-and dark-
174 acclimated measurements, respectively) and maximum Chl *a* fluorescence (F_m , and F_m' for
175 light-and dark-acclimated measurements, respectively) were estimated from iterative
176 algorithms for induction (Kolber et al., 1998) and relaxation phase (Oxborough, 2012) after 15
177 min of dark acclimation, which was sufficient to achieve a dark-acclimated state (data not
178 shown). All fluorescence parameters were calculated by standard equations (Genty et al., 1989;
179 Maxwell and Johnson, 2000). Maximum quantum yields of PSII (apparent PSII photochemical
180 quantum efficiency; F_v/F_m) were calculated as

$$181 \quad F_v/F_m = (F_m - F_0)/F_m \quad (1)$$

182 Fluorescence based photosynthesis-irradiance curves (PI) were conducted at six irradiances (I)
183 between 33 and 672 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$, with an acclimation time of 10 min per light step.
184 Electron transfer rate through PSII (ETR [$\text{mol e}^- (\text{mol RCII})^{-1} \text{s}^{-1}$]) for each light step was
185 calculated as:

$$186 \quad \text{ETR} = ((F_m' - F_0')/F_m') * I \quad (2)$$

187 Following the suggestion by Silsbe and Kromkamp (2012), the light-use efficiency (α [mol e^-
188 $\text{m}^2 (\text{mol RCII})^{-1} (\text{mol photons})^{-1}$]) and the maximum electron transfer rates per RCII (ETR_{max}
189 [$\text{mol e}^- (\text{mol RCII})^{-1} \text{s}^{-1}$]) were estimated by fitting the data to the model by (Webb et al., 1974):

$$190 \quad \text{ETR} = \text{ETR}_{\text{max}} * [1 - e^{-(\alpha * I)/\text{ETR}_{\text{max}}}] \quad (3)$$

191 The light saturation index (E_k [$\mu\text{mol photons m}^{-2} \text{s}^{-1}$]) was then calculated as $\text{ETR}_{\text{max}}/\alpha$.
192 Maximum non-photochemical quenching of Chl *a* fluorescence (NPQ) at irradiances of 672

193 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$ (i.e. the highest irradiance step of the PI curve) were calculated using the
194 normalized Stern-Volmer coefficient, also termed NSV, as described in McKew et al. (2013):

$$195 \quad (F_q'/F_v')-1 = F_0'/F_v' \quad (4)$$

196 where F_q' is the differences between measured and maximal fluorescence (Suggett et al., 2010).

197 F_0' was measured after each light step (with a duration of 90 s).

198

199 **2.5 Statistics**

200 All data is given as the mean of three biological replicates with \pm one standard deviation. To

201 test for significant differences between the treatments, two-way analyses of variance (ANOVA)

202 with additional normality (Kolmogorov-Smirnov) and Post Hoc (Holm-Sidak) tests were

203 performed. The significance level was set to 0.05. Statistical analyses were performed with the

204 program SigmaPlot (SysStat Software Inc, Version 12.5).

205

206 **3 Results**

207

208 **3.1 Carbonate Chemistry**

209 Regular dilution of cultures with pre-aerated seawater medium kept carbonate chemistry stable
210 over the course of the experiment. More specifically, in each bottle the drift in A_T and C_T
211 compared to initial values was $\leq 3\%$ and $\leq 4\%$, respectively (data not shown). Final carbonate
212 chemistry in the 2°C treatments yielded $p\text{CO}_2$ levels of 197 ± 3 , 323 ± 12 , 959 ± 22 and 1380
213 $\pm 53 \mu\text{atm}$ (Table 1). In the 6°C treatments, $p\text{CO}_2$ levels were 198 ± 6 , 394 ± 10 , 1036 ± 31 and
214 $1449 \pm 18 \mu\text{atm}$. Please note that the same $p\text{CO}_2$ level translates into differing dissolved CO_2
215 concentrations at different temperatures due to the temperature dependency of the carbonate
216 system. Specifically, the treatment $p\text{CO}_2$ values translated into up to 13% lower dissolved CO_2
217 concentrations in the 6°C compared to the 2°C treatment (Table 1; cf. Figure SII).
218 Concurrently, the $p\text{CO}_2$ levels at 2°C corresponded to pH_{total} values of 8.30 ± 0.01 , 8.11 ± 0.01 ,
219 7.68 ± 0.01 and 7.52 ± 0.02 , respectively. In the 6°C treatment, pH_{total} values of the four $p\text{CO}_2$
220 treatments were 8.30 ± 0.01 , 8.04 ± 0.01 , 7.65 ± 0.01 and 7.52 ± 0.01 , respectively.

221

222 **3.2 Growth and biomass build-up**

223 Growth rates constants of exponentially growing *M. pusilla* cultures were significantly affected
224 by the applied treatments (Figure 1, Table 2, SI1). Depending on the $p\text{CO}_2$ level, temperature
225 increased growth by 20 to 60% with an average of 0.80 d^{-1} under low and 1.10 d^{-1} under high
226 temperature conditions (two-way ANOVA, $F = 328$, $p < 0.001$). Overall, there was also a
227 positive $p\text{CO}_2$ effect on growth (two-way ANOVA, $F = 9$, $p = 0.001$), even though no linear
228 trends with either $p\text{CO}_2$ or $[\text{CO}_2]$ were observed (Figure 1, SI1). The observed $p\text{CO}_2$ responses
229 also differed between temperature levels, indicating a significant interaction between both
230 drivers (two-way ANOVA, $F = 12$, $p < 0.001$): Under low temperature, growth increased
231 significantly from 180 to 380 μatm $p\text{CO}_2$ (post-hoc, $t = 3.1$, $p = 0.04$), while there was a
232 declining, yet insignificant trend in growth with further increases in $p\text{CO}_2$. Under high
233 temperature, growth was significantly higher under 1000 compared to lower (180 μatm ; post-
234 hoc, $t = 5.6$, $p < 0.001$) and higher $p\text{CO}_2$ levels (1400 μatm ; post-hoc, $t = 5.9$, $p < 0.001$). Thus,
235 warming shifted the optimum range for growth to higher $p\text{CO}_2$ levels (Figure 1A).

236 This trend was also observed in terms of POC production rates (Figure 1B, Table 2),
237 with significant effects of temperature (Table SI1; two-way ANOVA, $F = 356$, $p < 0.001$), $p\text{CO}_2$
238 (two-way ANOVA, $F = 7$, $p = 0.003$), and their interaction (two-way ANOVA, $F = 29$, p
239 < 0.001). At low temperatures, higher production rates were observed at 180 and 380 μatm

240 compared to those at 1000 and 1400 μatm pCO_2 (post-hoc tests, $t = 3.5$, $p = 0.016$ and $t = 3.0$,
241 $p = 0.046$, respectively). At high temperatures, POC production rates were significantly higher
242 at 1000 μatm than at all other pCO_2 levels (post-hoc tests, e.g. $t = 9.1$, $p < 0.001$ for 380 vs.
243 1000 μatm and $t = 7.4$, $p < 0.001$ for 1000 vs. 1400 μatm), again indicating an upward shift in
244 the pCO_2 optimum with warming.

245

246 **3.3 Cellular composition**

247 Overall, POC quota (Figure 2 a, Table 2, Table SII) were significantly higher under elevated
248 compared to low temperature (two-way ANOVA, $F = 24$, $p < 0.001$), but no overarching trend
249 with pCO_2 was observed. Under low temperature, cells had significantly higher POC quota at
250 low pCO_2 levels (180 and 380 μatm) compared to high pCO_2 levels (1000 and 1400 μatm ; all
251 four post-hoc tests significant, e.g. 380 vs 1000 μatm : $t = 2.8$, $p = 0.033$). This trend reversed
252 under high temperature, where POC quota were highest under 1000 and 1400 μatm (post-hoc
253 test, $t = 3.5$, $p = 0.024$). Thus, temperature and pCO_2 levels exhibited a significant interactive
254 effect on POC quota (two-way ANOVA, $F = 10$, $p < 0.001$).

255 Similar trends were observed in terms of cellular PON quota (Figure 2 b, Table 2, Table
256 SII), where temperature (two-way ANOVA, $F = 5$, $p = 0.045$) and its interaction with pCO_2
257 (two-way ANOVA, $F = 10$, $p < 0.001$) significantly affected the results. Here, opposing pCO_2
258 effects under different temperatures were more subtle, with PON quota under low temperatures
259 only being significantly decreased between 380 and 1400 μatm (post-hoc test, $t = 3.3$, $p =$
260 0.027), while under high temperature PON quota significantly increased from 180 and 380 to
261 1000 μatm pCO_2 (post-hoc tests, $t = 3.7$, $p = 0.012$ and $t = 2.8$, $p = 0.028$, respectively).

262 Regarding cellular Chl *a* quota, there were no significant effects of temperature or pCO_2
263 alone (Figure 2 c, Table 2, Table SII), but a significant interaction between the two drivers
264 (two-way ANOVA, $F = 18$, $p < 0.001$): Under low temperature, Chl *a* quota decreased from low
265 (180 μatm) to high pCO_2 levels (1000 and 1400 μatm ; post-hoc tests, $t = 5.0$, $p < 0.001$ and $t =$
266 3.9 , $p = 0.006$, respectively). Under high temperature, the opposite trend was observed, where
267 Chl *a* quota increased from low (180 and 380 μatm) to high pCO_2 levels (1000 and 1400 μatm ;
268 all four post-hoc tests significant, e.g. 380 vs 1000 μatm : $t = 3.0$, $p = 0.027$).

269 Molar C:N ratios of the biomass (Table 2, Table SII) increased with temperature (two-
270 way ANOVA, $F = 14$, $p = 0.002$), yet this overall difference was mainly driven by results at
271 low pCO_2 levels (180 and 380 μatm ; post-hoc tests, $t = 2.7$, $p = 0.017$ and $t = 3.5$, $p = 0.003$,
272 respectively). By itself, pCO_2 did not significantly affect C:N ratios.

273 The ratios of C:Chl *a* (Figure 2 d, Table 2, Table SII) were elevated under high

274 compared to low temperature conditions (two-way ANOVA, $F = 14$, $p = 0.002$), an effect that
275 was most pronounced at $p\text{CO}_2$ levels of $180 \mu\text{atm}$ (post-hoc test, $t = 5.5$, $p < 0.001$). While there
276 was no effect of $p\text{CO}_2$ on C:Chl *a* at low temperature, C:Chl *a* decreased with increasing $p\text{CO}_2$
277 at high temperature (two-way ANOVA, interaction term, $F = 6$, $p = 0.007$; 180 vs. $1400 \mu\text{atm}$
278 at 6°C post-hoc test, $t = 3.9$, $p = 0.008$).

279

280 **3.4 Chl *a* fluorescence-based photophysiology**

281 The effects of the applied treatments on photophysiology were studied by means of FRRf,
282 which investigates photochemistry at photosystem II (PSII). No effects of the applied
283 treatments were observed in most parameters investigated (Table 3, SI1). This was true for the
284 dark-acclimated quantum yield efficiency of PSII (F_v/F_m), which was similar in all treatments
285 with values of 0.45 ± 0.06 , as well as for absorption cross section of PSII light harvesting (σ_{PSII}).

286 Furthermore, the fitted parameters of FRRf-based PI curves (α , ETR_{max} and E_K) were
287 independent of the experimental treatments (Table 3, SI1). In contrast, the rate constant of the
288 reopening of PSII reaction centres (τ_{ES} ; Table 3, SI1) was slightly yet significantly smaller
289 under high temperatures (two-way ANOVA, $F = 6$, $p = 0.029$), even though this overall
290 response also depended on the applied $p\text{CO}_2$ levels (two-way ANOVA, interaction term, $F = 4$,
291 $p = 0.033$).

292 Maximum non-photochemical quenching (NPQ_{max} ; Table 3, SI1) increased
293 significantly with $p\text{CO}_2$ (Table SI1; two-way ANOVA, $F = 0$, $p = 0.002$) while temperature had
294 no effect. Post-hoc tests revealed that this response was mainly driven by high NPQ_{max} values
295 at $1000 \mu\text{atm}$, which were significantly higher than in any other $p\text{CO}_2$ treatment (e.g. $t = 4.1$, p
296 $= 0.006$ for 380 vs. $1000 \mu\text{atm}$ and $t = 3.1$, $p = 0.030$ for 1000 vs. $1400 \mu\text{atm}$).

297 **4 Discussion**

298

299 **4.1 *Micromonas pusilla* benefits from warming**

300 We observed a strong stimulation of growth rates and biomass build-up with increasing
301 temperature (Figure 1, Table 2). Even though the isolate stems from 1.8°C water temperature,
302 the beneficial effects of warming from 2°C to 6°C are not surprising: *M. pusilla* is known to
303 dominate Arctic phytoplankton assemblages in the summer and autumn situations (Lovejoy et
304 al., 2007; Marquardt et al., 2016) when surface temperatures of 6°C or more can be reached
305 (Hegseth et al., in press). Our results are also in line with mesocosm experiments that indicate
306 stimulatory effects of warming on picoplankton abundances (Daufresne et al., 2009; Sommer
307 et al., 2015) as well as with the temperature optimum of 6-8°C observed for another Arctic
308 strain of *M. pusilla* (Lovejoy et al., 2007).

309 Below the temperature optimum of a cell, warming causes an acceleration of the entire
310 metabolism, as enzymatic reactions run faster under these conditions (Eppley, 1972; Brown et
311 al., 2004). In this study, warming caused higher growth rates, POC quotas and biomass
312 production (Figure 2, Tables 2, SI1), indicating that particularly the fixation and storage of
313 carbon was facilitated by increasing temperature. Electron transport processes, on the other
314 hand, were largely independent of temperature (Tables 3, SI1). Thus, temperature affected the
315 balance between electron transport ('light reaction') and carbon fixation in the Calvin cycle
316 ('dark reactions'). Especially under relatively low temperatures, as investigated here, warming
317 can decrease the excitation pressure on the electron transport chain of the photosystems by
318 increasing the temperature-limited turnover rates of enzyme reaction such as RuBisCO (Mock
319 and Hoch, 2005). Thus, cells grown under low temperature need to invest relatively more
320 energy into biosynthesis than into photochemistry compared to cells grown under high
321 temperature (Toseland et al., 2013). While it has been shown that Antarctic diatoms can
322 compensate for slow RuBisCO kinetics by increasing the expression of this enzyme (Young et
323 al., 2014), it is unknown whether such acclimation responses also occur in prasinophytes.
324 Regarding the C:Chl *a* ratio, this can be taken as an indicator on how much resources the cell
325 retains as carbon biomass (e.g. structural and storage compounds) relative to how much is
326 invested into its light harvesting capacities (Halsey and Jones, 2015). In this study, the strong
327 temperature-dependent increase in C:Chl *a* (Figure 2, Table SI1) under potentially limiting
328 pCO₂ levels of 180 µatm suggests that under warming, the balance between light harvesting
329 and carbon fixation was indeed more beneficial for biomass build-up. Furthermore, elevated
330 temperature significantly decreased τ_{ES} (Table 3, SI1), which can serve as a proxy of the rate

331 at which down-stream processes can remove electrons from PSII (Kolber et al., 1998). Thus,
332 our results indicate that the drainage of electrons into carbon fixation was faster under warmer
333 conditions, explaining the higher growth and biomass production under these conditions.

334

335 **4.2 Warming shifts CO₂ optima towards higher pCO₂ levels**

336 Under 6°C and pCO₂ levels expected to be reached by the end of this century, OA had a
337 significantly positive effect on growth and biomass build-up (Figure 1). This finding is in line
338 with previous studies, which have shown that picoeukaryotes can benefit strongly from OA in
339 both laboratory and mesocosm studies (Meakin and Wyman, 2011; Newbold et al., 2012;
340 Schaum et al., 2012; Brussaard et al., 2013; Maat et al., 2014; Schulz et al., 2017). Such positive
341 response to OA could indicate that picoeukaryotes such as *M. pusilla* are mainly dependent on
342 diffusive CO₂ supply and thus directly benefit from higher CO₂ concentrations (Brussaard et
343 al., 2013; Schulz et al., 2013; Schulz et al., 2017). While this could be related to the large surface
344 to volume ratio of small cells, opposite trends within the group of diatoms (i.e. higher sensitivity
345 of larger compared to smaller diatoms; Wu et al.; 2014; Sett et al. 2018) suggest more complex
346 underlying mechanisms at play.

347 Despite this overall effect, growth rates of *M. pusilla* tended to follow a non-linear
348 response curve over the tested range of glacial to elevated future pCO₂ levels (i.e. 180 to 1400
349 µatm), i.e. growth increased with increasing pCO₂ from low to intermediate, but decreased
350 again under the highest pCO₂ levels (Figure 1). Such an optimum behaviour can be expected
351 for most environmental drivers (Harley et al., 2017) and has previously been observed in
352 response to OA (Sett et al., 2014; Wolf et al., 2018; Hoppe et al. 2018). The response patterns
353 in these studies were attributed to a combination of beneficial effects of rising pCO₂ under
354 potentially carbon-limiting conditions for photosynthesis, and negative effects of declining pH
355 on cellular homeostasis and enzyme performance, which manifest mainly at high pCO₂ (Bach
356 et al., 2013). This non-linearity in the observed pCO₂ effects emphasises the importance of
357 experiments with more than two pCO₂ levels in order to properly describe OA-response patterns
358 of organisms.

359 On a more general level, apparent discrepancies between OA studies can be attributed
360 to actual differences in the environmental settings and their interactive effects with pCO₂
361 (Riebesell and Gattuso, 2015). When comparing the two most commonly applied pCO₂ levels,
362 i.e. the present-day and the anticipated end-of-century situation, the effects of OA on most of
363 the investigated physiological parameters are reversed under 6°C compared to 2°C (Figure 3).
364 This illustrates how difficult it is to infer responses to OA from experiments applying only one

365 set of environmental conditions. It is also noteworthy that the combination of OA and warming
366 led to more densely packed cells (no change in cell size based on flow cytometric
367 measurements; data not shown) with similar stoichiometry compared to the control treatment
368 (Table 2). This indicates that cells managed to cope well with the experienced future conditions.
369 Furthermore, warming altered the OA-dependent change in most of the investigated parameters
370 in a direction that indicates higher fitness compared to low temperatures (e.g. higher growth
371 rates and higher elemental quota; Figure 3). Thus, the increase in growth under future compared
372 to ambient conditions was larger than what would be expected by the respective responses to
373 warming and OA in isolation, indicating synergistic beneficial effects of both drivers.

374

375 **4.3 Potential mechanism underlying the interaction between warming and OA**

376 The observed synergistic effects could be explained by their specific impacts on carbon
377 acquisition and fixation. As outlined in the introduction, light and dark reaction of
378 photosynthesis need to be balanced to achieve high biomass production while avoiding
379 photodamage (Behrenfeld et al., 2008). According to our data, this balance is shifted towards
380 higher biomass production rates under warming and OA.

381 At higher temperatures, seawater CO₂ concentrations were lower than under colder
382 conditions (Table 1; Zeebe and Wolf-Gladrow, 2001). At the same time, warming from 2°C to
383 6°C caused up to 60% higher growth and 110% higher biomass build-up rates (Figure 1, Table
384 2). Furthermore, the decrease in τ_{ES} indicates a faster transfer of photochemical energy into
385 downstream processes such as RuBisCO activity (Table 3). Increased carbon demand in concert
386 with lower carbon supply at higher temperatures thus increases the risk of CO₂ shortage in the
387 cell, which in turn causes OA to have larger effects than at colder temperatures. Moreover,
388 warming changes the kinetics of carbon fixation, with RuBisCO increasing its maximum
389 turnover rates but decreasing its affinity for CO₂ (Young et al., 2014). At higher temperature,
390 cells thus have the potential for higher carboxylation rates provided sufficient CO₂ is available
391 (Kranz et al., 2015). Under elevated pCO₂ levels, diffusive CO₂ supply increases and/or costs
392 for active carbon acquisition decrease. Consequently, the positive effect of increasing
393 temperature on the carbon fixation rate can develop its full potential under OA.

394 In conclusion, elevated catabolic activity under warmer conditions can explain the
395 observed upward shift in the CO₂-optimum of growth with increasing temperature (Figure 1),
396 as the corresponding higher carbon demand causes CO₂ fixation to saturate under higher pCO₂
397 levels. In combination with a faster and more efficient machinery for pH homeostasis at
398 elevated temperatures (Morgan-Kiss et al., 2006), this could explain why declining growth rates

399 were only observed at relatively higher $p\text{CO}_2$ levels compared to those under low temperature
400 conditions (Figure 2).

401 **4.4 Implications for the current and future Arctic pelagic ecosystem**

402 Picoeukaryotes such as *M. pusilla* are considered to be potential winners of climate change:
403 They are not only thriving in warmer, more stratified environments, which are predicted to
404 further expand in the future, but also seem to benefit from OA (Li et al., 2009; Schulz et al.,
405 2017). Our results for *M. pusilla* confirm beneficial effects of warming and OA on growth and
406 biomass production under nutrient-replete conditions (Figure 1, Table 2). Hence, in warmer
407 spring conditions of the future, this species may experience growth stimulation under OA,
408 potentially increasing its importance early in the growing season. Currently, *M. pusilla* already
409 dominates Arctic phytoplankton assemblages in the nutrient-limited summer and autumn
410 situations, which were not investigated in this study. Regarding the importance of nutrient
411 availability, laboratory experiments found beneficial OA effects on *M. pusilla* primary
412 production to persist also under P limitation (Maat et al., 2014), while in a mesocosm
413 community, OA-dependent increases in *M. pusilla* abundances disappeared when the system
414 ran into P and N co-limitation (Engel et al., 2008). Thus, it remains to be seen how the combined
415 effects of warming and OA manifest under low nutrient conditions as well as how the responses
416 may depend on sources and types of nutrients (e.g. mixing-delivered nitrate vs. regenerated
417 ammonium).

418 A species' success in the environment does not only depend on individual performance, but
419 also on how it compares to that of competing species. When we compare our results with the
420 responses of the Arctic diatom *Thalassiosira hyalina*, isolated from the same location and
421 exposed to the same experimental conditions (Wolf et al., 2018), the diatom had higher growth
422 rates than the picoeukaryote under most treatment conditions, as can be expected for nutrient-
423 replete conditions (Sarhou et al., 2005). The relative increase in growth rates from ambient (2
424 or 3°C and 380 $\mu\text{atm pCO}_2$) to future conditions (6°C and 1000 $\mu\text{atm pCO}_2$) was, however,
425 much higher for *M. pusilla* than for *T. hyalina*. The fact that our experiments were conducted
426 under nutrient-replete conditions, which typically favour diatoms over picoeukaryotes, may
427 indicate an even stronger increase in fitness (Collins et al., 2014) and could mean that *M. pusilla*
428 gains another competitive advantage over phytoplankton like diatoms in the future, in addition
429 to those resulting from changes in stratification (Li et al., 2009). Thus, our findings suggest
430 higher picoplankton contribution to future Arctic phytoplankton assemblages under non-
431 limiting conditions, e.g. early in the growing season when picoeukaryotes can already
432 contribute quite substantially to the phytoplankton standing stocks (Marquardt et al, 2017,
433 Paulsen et al. 2015). How such competition between diatoms and picoeukaryotes would

434 manifest under nutrient-depleted conditions that strongly favour *M. pusilla* is currently
435 unknown.

436 Even though picoeukaryotes seem to contribute more to the downward export of organic
437 matter than previously assumed (Waite et al., 2000; Richardson and Jackson, 2007), in
438 comparison to e.g. diatoms, they are less efficient vectors for carbon export to depth and have
439 a lower energy transfer along trophic levels (Sherr et al., 2003). Consequently, Arctic food webs
440 dominated by picoeukaryotes would look very different from those fuelled by diatom
441 production (Sherr et al., 2003; Paulsen et al., 2015). Due to its motility and capability to grow
442 mixotrophically, *M. pusilla* is characterized by an exceptionally high cellular C:N ratio
443 compared to other Arctic phytoplankton (Table 2; Halsey et al., 2014; McKie-Krisberg and
444 Sanders, 2014). An increased importance of this species would thus not only affect the food
445 web due to its small size and concurrent grazer preferences, but also in terms of food quality
446 (van de Waal and Boersma, 2012). The expected higher growth rates and thus abundances of
447 this species may thus strengthen the Arctic microbial food web. Together with a concurrent
448 weakening of the classical diatom-fuelled food web, this could have severe implications for the
449 flow of energy and nutrients through future marine Arctic ecosystems (Post, 2016).

450

451 **4.5 Conclusions**

452 This study is the first to show synergistic effects of warming and OA on *M. pusilla*, one of the
453 most abundant species of the worlds' oceans. Individually, both warming and OA cause more
454 efficient biomass build-up under nutrient-replete conditions. Beneficial effects manifest,
455 however, even more strongly in combination, when facilitated carbon acquisition (e.g. due to
456 higher diffusive CO₂ supply) co-occurs with higher fixation rates (e.g. due to higher turnover-
457 rates of RuBisCO). Our results provide an explanation for the observations of previous
458 mesocosm studies, which indicated beneficial effects of OA and warming on *M. pusilla* and
459 other picoeukaryotes. Characterising the responses of this Arctic key species to warming and
460 OA will help to develop mechanistic phytoplankton functional types and more realistic model
461 representation of phytoplankton assemblages as well as their responses to multiple drivers.
462 Future studies are needed to elucidate further multifactorial environmental changes, addressing
463 both abiotic (e.g. changes in light and nutrients) as well as biotic (e.g. heterotrophy,
464 competition, grazers, viruses) interactions.

465 **Author Contributions**

466 C.J.M.H. and B.R. designed the study. C.J.M.H. and C.F. conducted the experiment. C.J.M.H.
467 analysed the data and prepared the manuscript with contributions from B.R. and C.F.

468

469 The authors declare that they have no conflict of interest.

470

471

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Table 1: Seawater carbonate chemistry at the end of the experiments (n=3; mean \pm 1 s.d.). CO₂ partial pressure (pCO₂) and dissolved CO₂ concentrations were calculated from total alkalinity (A_T) and pH_{total} at 2 or 6°C, a salinity of 32.7 using CO₂SYS (Pierrot et al., 2006), and phosphate and silicate concentrations of 10 and 100 $\mu\text{mol kg}^{-1}$, respectively. n.a. indicates that values are not available for this specific treatment.

Temperature [°C]	pCO ₂ level [μatm]	pH total scale	A _T [$\mu\text{mol kg}^{-1}$]	C _T [$\mu\text{mol kg}^{-1}$]	dissolved CO ₂ [$\mu\text{mol kg}^{-1}$]	pCO ₂ [μatm]
2	180	8.3 \pm 0.01	2264 \pm 9	2024 \pm 6	11.6 \pm 0.2	197 \pm 3
	380	8.11 \pm 0.01	2244 \pm 30	2124 \pm 11	19.0 \pm 0.7	323 \pm 12
	1000	7.68 \pm 0.01	2255 \pm 45	2215 \pm 23	56.4 \pm 1.3	959 \pm 22
	1400	7.52 \pm 0.02	2243 \pm 5	n.a.	81.1 \pm 3.1	1380 \pm 53
6	180	8.3 \pm 0.01	2243 \pm 28	1969 \pm 10	10.0 \pm 0.3	198 \pm 6
	380	8.04 \pm 0.01	2256 \pm 21	2058 \pm 7	20.0 \pm 0.5	394 \pm 10
	1000	7.65 \pm 0.01	2262 \pm 22	2178 \pm 14	52.6 \pm 1.6	1036 \pm 31
	1400	7.52 \pm 0.01	2265 \pm 5	n.a.	73.6 \pm 0.9	1449 \pm 18

Table 2: Growth rate constants μ , division rate constants k , POC production rates and cellular quota of Chl a , POC and PON as well as their ratios of *M. pusilla* at the end of the experiment under the different treatment conditions (n=3; mean \pm 1 s.d.). Results from statistical analysis can be found in Table SI2.

Temperature [°C]	pCO ₂ [μ atm]	Growth rate constant μ [d ⁻¹]	Division rate constant k [d ⁻¹]	POC production [fmol cell ⁻¹ d ⁻¹]	POC quota [fmol cell ⁻¹]	PON quota [fmol cell ⁻¹]	Chl a quota [fg cell ⁻¹]	POC:PON [mol mol ⁻¹]	POC:Chl a [g g ⁻¹]
2	180	0.75 \pm 0.04	1.08 \pm 0.05	256 \pm 11	239 \pm 20	28.7 \pm 2.6	28.6 \pm 2.1	8.3 \pm 0.1	100 \pm 7
	380	0.85 \pm 0.03	1.23 \pm 0.04	290 \pm 16	237 \pm 22	30.9 \pm 1.8	24.7 \pm 1.4	7.7 \pm 0.3	115 \pm 10
	1000	0.79 \pm 0.05	1.15 \pm 0.07	224 \pm 23	196 \pm 25	24.9 \pm 4.9	20.0 \pm 2.9	8.0 \pm 0.6	118 \pm 5
	1400	0.82 \pm 0.05	1.18 \pm 0.07	235 \pm 13	199 \pm 16	23.9 \pm 2.3	22.0 \pm 1.5	8.4 \pm 0.1	109 \pm 4
6	180	1.06 \pm 0.03	1.53 \pm 0.05	376 \pm 15	245 \pm 3	26.9 \pm 0.6	21.1 \pm 0.8	9.1 \pm 0.1	140 \pm 7
	380	1.05 \pm 0.03	1.52 \pm 0.04	342 \pm 39	226 \pm 26	25.9 \pm 2.5	22.1 \pm 2.6	8.7 \pm 0.3	123 \pm 11
	1000	1.25 \pm 0.05	1.80 \pm 0.07	497 \pm 12	275 \pm 6	33.7 \pm 2.0	27.2 \pm 0.9	8.2 \pm 0.6	122 \pm 6
	1400	1.05 \pm 0.04	1.52 \pm 0.06	400 \pm 14	263 \pm 7	31.1 \pm 1.7	28.6 \pm 3.2	8.5 \pm 0.3	111 \pm 16

Table 3: : FRR-fluorimetric PSII photochemistry measurements – PSII quantum yield efficiency Fv/Fm [dimensionless], functional absorption cross section (σ_{PSII}) [$\text{nm}^{-2} \text{PSII}^{-1}$], rate of PSII re-opening (τ_{ES} [ms]), maximum non-photochemical quenching at 672 $\mu\text{mol photons m}^{-2} \text{s}^{-1}$ (NPQ_{max} [dimensionless]), maximum light-use efficiency (initial slope α [$\text{mol e}^{-} \text{m}^2 (\text{mol RCII})^{-1} (\text{mol photons})^{-1}$]), maximal absolute electron transfer rates through PSII (ETR_{max} [$\text{mol e}^{-} (\text{mol RCII})^{-1} \text{s}^{-1}$]), and the light saturation index (E_{K} [$\mu\text{mol photons m}^{-2} \text{s}^{-1}$]) under the different temperature and pCO₂ treatments (n=3; mean \pm 1 s.d.). Results from statistical analysis can be found in Table SI2.

Temp	pCO ₂	Fv/Fm	σ_{PSII}	τ_{ES}	NPQ_{max}	α	ETR_{max}	E_{K}
2	180	0.50 \pm 0.01	8.66 \pm 0.35	439 \pm 8	2.26 \pm 0.18	0.42 \pm 0.05	33 \pm 2	81 \pm 13
	380	0.43 \pm 0.09	8.93 \pm 0.26	425 \pm 4	3.51 \pm 0.55	0.32 \pm 0.15	25 \pm 5	91 \pm 44
	1000	0.45 \pm 0.08	8.55 \pm 0.07	448 \pm 1	3.96 \pm 0.71	0.42 \pm 0.03	31 \pm 2	75 \pm 10
	1400	0.47 \pm 0.10	9.06 \pm 0.05	422 \pm 14	2.45 \pm 0.44	0.43 \pm 0.08	31 \pm 7	75 \pm 31
6	180	0.49 \pm 0.01	9.22 \pm 0.22	412 \pm 6	2.51 \pm 0.37	0.49 \pm 0.08	28 \pm 10	59 \pm 28
	380	0.43 \pm 0.12	8.83 \pm 0.17	427 \pm 6	2.83 \pm 0.59	0.38 \pm 0.09	35 \pm 14	90 \pm 17
	1000	0.41 \pm 0.07	8.91 \pm 0.22	422 \pm 11	4.94 \pm 1.46	0.33 \pm 0.09	32 \pm 5	100 \pm 21
	1400	0.45 \pm 0.04	8.71 \pm 0.50	428 \pm 19	2.93 \pm 0.50	0.38 \pm 0.04	40 \pm 5	104 \pm 6

Figure 1: Specific growth rate constant μ (A) and POC production (B) of *M. pusilla* under low (filled symbols) and high temperatures (open symbols) as a function of pCO₂ (n=3; mean \pm 1 s.d.). Results from statistical analysis can be found in Table SI2.

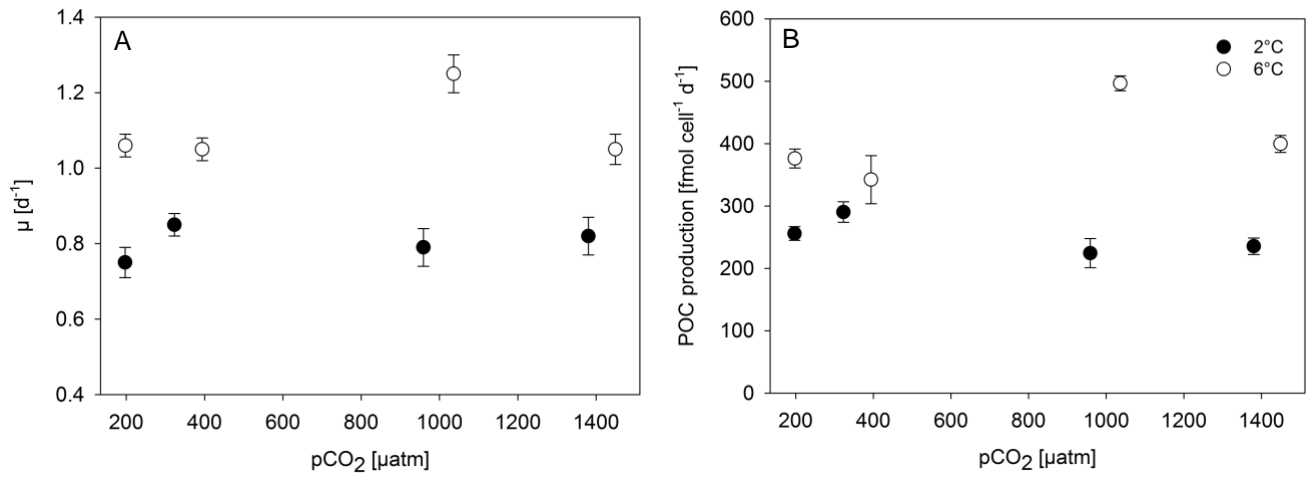


Figure 2: Cellular composition, i.e. POC (A), PON (B) and Chl *a* quota (C) as well as as C:Chl *a* ratios (D), of *M. pusilla* under low (filled symbols) and high temperatures (open symbols) as a function of pCO₂ (n=3; mean ±1 s.d.). Results from statistical analysis can be found in Table SI2.

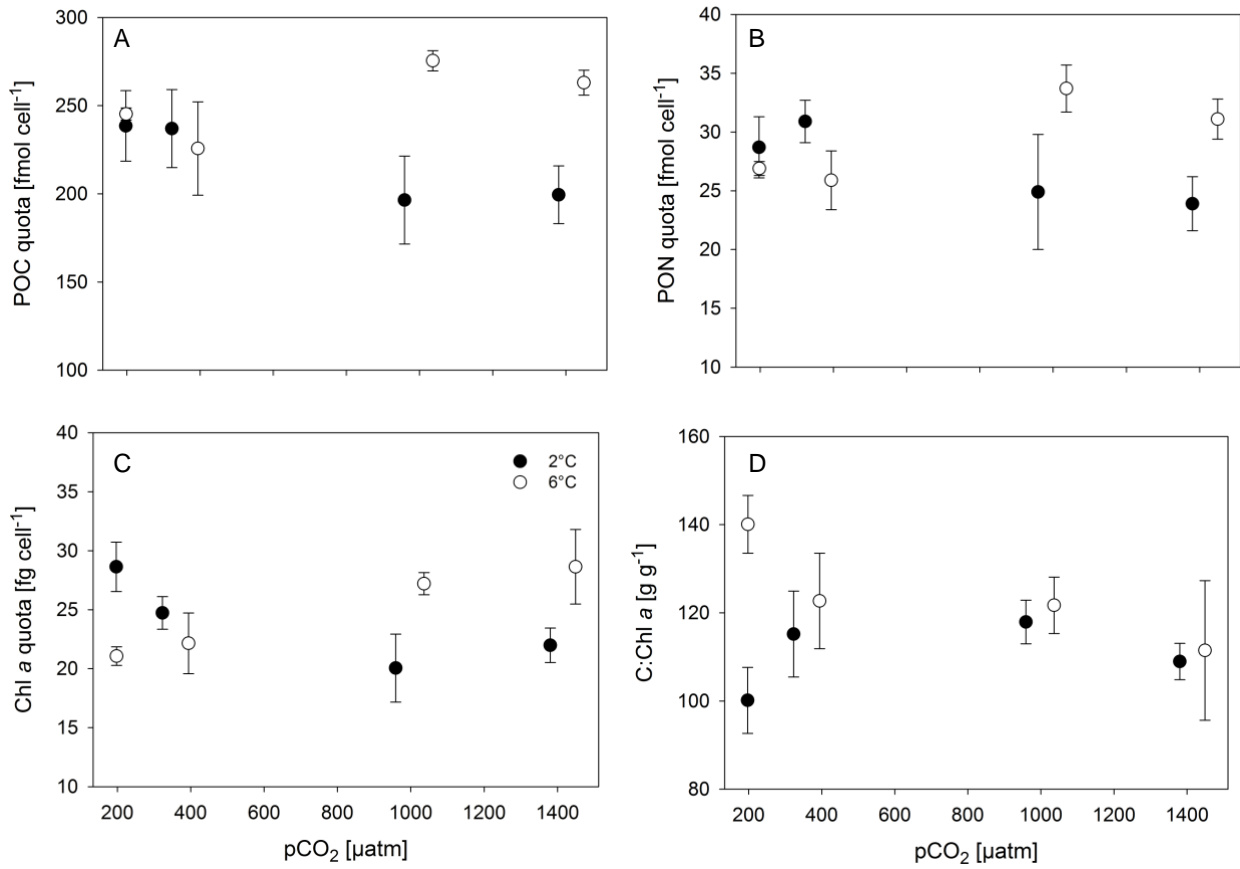
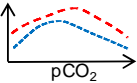
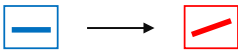
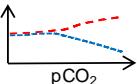
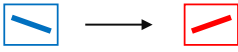
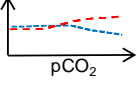
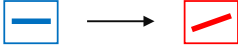
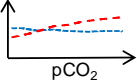
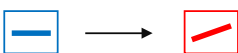
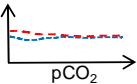
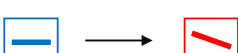


Figure 3: Schematic illustration of results for both temperatures over the entire range of pCO₂ levels as well as focusing on the responses between 380 and 1000 μatm (as the representation for commonly used OA treatments) and their modulation by temperature.

Parameter	Response curves	380 - 1000 μatm effect 2°C → 6°C
Growth		
POC quota		
PON quota		
Chl a quota		
C:N		
C:Chla	