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Dear Dr. Still

Please find enclosed our revised manuscript, mss# bg-2018-322, title, "Quantifying energy use efficiency via entropy production: A case study from longleaf pine ecosystems." We have completed all revisions per the Reviewer's comments and suggestion. We would like to thank them for their constructive reviews which have help to improve the manuscript. Below are the Reviewer's comments in normal black font and our responses to the comments in blue. We have include in our online submission both a marked and clean copy of the revised manuscript. Please contact me if you have any questions regarding the manuscript.

Sincerely,

A handwritten signature in black ink, appearing to read 'Gregory Starr'.

Gregory Starr, Ph.D.  
Professor of Global Change Biology  
University of Alabama

## Reviewer 1: Alex Kleidon

**Reviewer's comment:** *First, the entropy balance is used in Eq. 9, stating that the “overall change in entropy production ( $S$ ) over time ( $t$ ) in  $\text{kJ m}^{-2} \text{K}^{-1}$  of the ecosystem [is estimated] by adding entropy flux and entropy production”. This is incorrect. What Eq. 9 formulates is the entropy balance. It balances the change in entropy on the left hand side of the equation ( $dS/dt$ ) with the sum of all entropy exchange fluxes ( $J$ ) and all entropy production terms ( $\sigma$ ). This balance is typically assumed to be zero in a steady state, i.e.,  $dS/dt = 0$ , which then allows one to diagnose entropy production from the difference in entropy exchange fluxes. This is in fact what the authors do to diagnose entropy production in Eqs. 3.6 and 3.7 to diagnose entropy production by absorption of radiation. Yet, the authors later use  $dS/dt$  in Eq. 4.8 to derive an efficiency. This efficiency should be zero, otherwise they did not do the balancing correctly. So there is a major inconsistency in the methodology that needs to be resolved.*

**Authors' response:** We have changed the calculation for  $dS/dt$  and now focused on the entropy outputs and inputs and internal entropy production, to quantify the change in entropy ( $dS/dt$ ). Please see section 2.5 (Eq. 4.9) and the results section 3.6 in the revisions.

**Reviewer's comment:** *Second, entropy production by absorption of longwave radiation is estimated using net longwave radiation at the surface (Eq. 3.7). What is the justification for using net long- wave radiation, rather than gross fluxes? After all, the downwelling longwave radiation of the surface adds an entropy flux of  $R_{\text{down}}/T_{\text{sky}}$ , while the emission of radiation from the surface exports entropy at the rate of  $R_{\text{up}}/T_{\text{sr}}$ . Using the difference of these two fluxes (assuming that  $dS/dt=0$ ) yields an entropy production of  $\sigma = R_{\text{up}}/T_{\text{sr}} - R_{\text{down}}/T_{\text{sky}}$ , which is not the same as  $(R_{\text{up}} - R_{\text{down}}) * (1/T_{\text{sr}} - 1/T_{\text{sky}})$ . The authors should correct this, or explain why their expression is justified. The same reasoning applies to the application of net ecosystem exchange, where I think that also gross fluxes should be used, not net fluxes.*

**Authors' response:** Thank you for pointing out the mistake. We have adjusted our calculations following the Brunsell et al. (2011) approach using incoming longwave radiation to calculate entropy production as follows:  $R_{\text{L,in}} \times (1/T_{\text{sr}} - 1/T_{\text{sky}})$ . We acknowledge that calculating the  $R_{\text{L,up}}/T_{\text{sr}}$  and  $R_{\text{L,down}}/T_{\text{sky}}$  will estimate the incoming and outgoing entropy transfer associated with longwave radiation, but not the entropy produced due to absorption of longwave radiation and conversion to heat during this process (as shown in Brunsell et al. 2011). Please see section 2.5 (Eq. 4.7) in the manuscript.

We have also eliminated the efficiency ratio calculations of metabolic activity and now quantify the overall change in metabolic energy using solely NEE. We believe that using direct fluxes is superior, as it avoids any influence of model bias, as  $R_{\text{eco}}$  is estimated using temperature data. However, we have changed the analysis to comparing metabolic energy and entropy changes in the systems (see section 2.5 and the results section 3.5), rather than using a ratio to quantify metabolic efficiency. We have also changed our analysis to using daily average half-hourly fluxes for all variables in the manuscript following your comment.

**Reviewer's comment:** *Additional insights gained from entropy fluxes and entropy production*  
The authors link their entropy-based analysis to rather general concepts such as resilience and energy use efficiency. Yet, I do not see the additional insights gained by using entropy production, rather than an analysis based on the entropy, water, and carbon balance. Why does the entropy-based analysis provide more or novel insights that cannot be obtained by just an interpretation based on fluxes? The authors do not really answer this question within the manuscript and do not use the results to show this, as they only focus on an entropy-based analysis.

*In terms of interpreting the observations, I think that there is a critical step missing that relates the observed differences to an interpretation of processes, and this cannot be gained by just looking at entropy. For instance, temperature changes result from changes in the energy balance, as temperature is a measure for heat content. Yet, the energy balance is not even shown or discussed. Likewise, to understand changes in evaporation, I would expect a water balance being discussed. Instead, this study directly diagnoses entropy fluxes and thereby skips this process-based level of interpretation. It does not show and interpret the fluxes of the energy, water, and carbon balances separately, and does not demonstrate that something else can be learned by looking at entropy.*

*By lumping all aspects of the land surface into entropy production, I think that this neglects those aspects that are relevant for ecosystems from those that are irrelevant. The relevant flux for ecosystems is primarily the uptake of carbon, as this provides the chemical energy for terrestrial ecosystems. Plants live from the energy they fix during carbon assimilation, and, quite frankly, care little about the entropy production of other processes.*

*For this manuscript to provide more solid insights, I think it needs a more process based interpretation using the available data, it needs to be more specific regarding those terms that are really relevant to ecosystems, and it needs to at least discuss why there is more to be gained by looking at entropy-based diagnostics.*

**Authors' response:** We have added an analysis and discussion of energy fluxes and the sites' energy balances to show the novelty of the entropy approach and to highlight specifically that the inclusion of entropy production gives more insights about the energy efficiencies and ecosystem function. Please see sections 3.3 and 3.4 for the results. To estimate the entropy budget of ecosystems, it is of particular importance to quantify entropy production based on the absorption of radiation, as this term is of similar magnitude as the entropy fluxes of LE and H together at our sites. We have also added a more thorough introduction and discussion of the topic.

We have also included soil moisture content and rainfall in our analysis to quantify changes in entropy fluxes and entropy production, but an analysis of the whole water budget was beyond the scope of this research project.

We kindly disagree with the reviewer's comment that the relevant flux for ecosystems is solely the carbon flux. For ecosystems (encompassing not only plant organisms), the partitioning of heat fluxes plays a significant role in their function, because the physical and biological processes are interconnected. LE in particular plays a large role in the maintenance of the surface temperature in ecosystems and is one of the largest contributors to entropy export in our ecosystem.

**Minor comments:**

*Abstract: “Our study provides foundational evidence of how MEP can be used to determine resiliency across ecosystems globally” - I am not at all convinced and doubt this conclusion. The authors provide no discussion why a diagnosis based on entropy fluxes yields more or better insights than the diagnosis of energy, water, and carbon balances. I see this as a critical missing bit in this manuscript.*

**Authors' response:** We have adjusted the discussion and methodology to show that entropy metrics can give further insights about differences in ecosystem function at the three longleaf pine sites, in addition to using energy fluxes. Our revisions focus on the entropy import and export, as well as the internal entropy production, to quantify how close these ecosystems are to a thermodynamic steady state.

*Introduction, page 2, line 16: MEP is referred to as a principle in the text. At best, it is a “proposed” principle, or better hypothesis, as it is not generally being accepted.*

**Authors' response:** We have adjusted the sentence accordingly.

*page 2, line 24: How can agricultural systems exceed MEP if MEP already describes the maximum? This does not make sense. What I can imagine is that agricultural systems maintain a different state because of nutrient inputs, but then, the boundary conditions are changed because there are additional exchange fluxes across the system boundary. Also, why would this excessive entropy production be unsustainable? As long as the nutrient input can be maintained, I see no reason why it should be unsustainable.*

**Authors' response:** We have altered our introduction to focus more on the importance of entropy exchanges and entropy production in ecosystems. The section of MEP and MEP in agricultural systems has therefore been eliminated.

*page 3, line 9: What are entropy efficiency ratios? In thermodynamics, efficiency is used to describe the conversion efficiency of one form of energy into another, and this involves entropy (like the well-known Carnot limit). But to speak of efficiency for entropy does not make sense to me.*

**Authors' response:** We have adjusted our revisions to avoid the use of “entropy efficiency”.

*page 6, line 4: How can two unknowns (GEE and Reco) be estimated from one equation? I think there is some information missing here.*

**Authors' response:** We have added a more detailed description of how these fluxes were obtained (see section 2.4).

*page 6, line 8: The authors convert the units from  $W\ m^{-2}\ K^{-1}$  to  $kJ\ m^{-2}\ K^{-1}$ . The unit should be  $kJ\ m^{-2}\ K^{-1}\ month^{-1}$  (i.e., the time is missing, throughout the whole manuscript), since entropy production refers to a rate, and not to an amount. But I do not understand the motivation for not keeping the units*

**Authors' response:** We have adjusted the units accordingly and are now using daily averages of half-hourly energy and entropy fluxes in  $W\ m^{-2}$  and  $W\ m^{-2}\ K^{-1}$ , respectively.

*page 6, line 14: Radiative entropy production actually includes a factor of 4/3, as it does not deal with heat, but with radiation (the additional contribution of 1/3 is due to radiation pressure). I think it needs a brief explanation why this factor was omitted.*

**Authors' response:** We avoided this factor assuming that the incoming and outgoing radiation does not assert radiation pressure (see Ozawa et al. 2003; Kleidon and Lorenz, 2005; Fraedrich and Lunkeit, 2008; Kleidon, 2009; Pascale et al., 2012). Please see section 2.5.

*page 7, line 2: What do you mean by “to calculate the change in entropy of the metabolic system”. Do you refer to entropy production? If you want to estimate entropy production, this would relate to dissipation of carbohydrates, which in turn relates to respiration. So I do not understand why NEE is being used.*

**Authors' response:** As noted above, we have changed the calculation of metabolic energy and entropy solely using NEE data without quantifying an efficiency ratio. Instead we are now comparing results for NEE energy and entropy.

*page 7, line 14: Why is net longwave radiation being used to calculate entropy production? The entropy fluxes of longwave radiation are  $Rl,down/Tsky$  and  $Rl,up/Tsrf$  as the authors write earlier in the manuscript. But this is not the same as  $Rl,net * (1/Tsrf - 1/Tsky)$ . (See major comment above)*

**Authors' response:** We have altered this calculation method in section 2.5. Please see our response above.

*page 7, line 20:  $dS/dt$  refers to the change in entropy with time, not change in entropy production. It should be zero in steady state, otherwise one cannot calculate entropy production from entropy fluxes. (See major comment above)*

**Authors' response:** You are correct. As noted above, we have fixed this error.

*page 7, line 29/30: Why are these expressions referred to as MEP? I see no connection to MEP. They just formulate radiative entropy production. Also, what's the difference to Eq. 3.6 and 3.7?*

**Authors' response:** We have revised this section to make it clearer that we are talking about an assumption. If this assumption does not necessarily reflect reality, it still gives us a means to compare different ecosystems or sites with respect to how they reflect, absorb and emit radiation.

*page 8, line 3: "an ecosystem maximizes its entropy production when it converts all incoming  $R_s$  and  $R_l$  into work". This is not correct. First, work is something different than entropy production. Second, it is impossible to convert all incoming radiation into work, as it would imply that there is no energy left to maintain a temperature that is greater than  $T = 0K$ .*

**Authors' response:** As noted above, we have changed the  $dS/dt$  section, and have excluded this ratio analysis.

*page 8, line 3: "... MEP.. is often negative or 0". No! Entropy production must always be greater or equal to zero, otherwise there is something wrong in the formulations! Spontaneous reductions in entropy are only possible at the microscopic scale during extremely short time periods but are practically irrelevant at the scale of ecosystems.*

**Authors' response:** We have adjusted the sentence accordingly in section 2.6.

*page 8, line 7: "maximum entropy of metabolism". What do you mean by this?*

**Authors' response:** We apologize for the confusion with this statement. We have changed our analysis to using net fluxes and are now looking at metabolic activity, rather than efficiency at the site by comparing energy and entropy fluxes of NEE at the sites.

*page 8, line 13: You express the efficiency as the ratio of the entropy flux associated with net ecosystem exchange to the energy flux of GEE. Should this not compare gross energy fluxes, rather than net exchange to gross exchange*

**Authors' response:** We have excluded the analysis of metabolic efficiency in the revised paper.

*page 8, line 16: This expression merely describes a radiative entropy flux, but not entropy production, or a maximum in entropy production.*

**Authors' response:** We have altered the section including the whole ecosystem entropy budget; this section was omitted.

*page 8, line 18: This expression does not give an efficiency, because in steady state (a condition needed to estimate entropy production from fluxes),  $dS/dt = 0$  so this expression is zero as well.*

**Authors' response:** As noted above, we have altered the calculation.

*I stop here with commenting, because I think that the methodology has a number of flaws that I wonder how much these impact the results. In addition, as expressed earlier, I think that the overall motivation for this entropy-based analysis needs to be improved.*

**Authors' response:** We have substantially altered the introduction and discussion in the revised manuscript to improve clarity about the methodology and as to why entropy metrics are useful in quantifying differences in ecosystem function.

## **Response to reviewer 2**

### **Major Comments**

**Reviewer's comment:** *The authors need to provide a more detailed overview of the concept of entropy. Are these concepts definable for biological systems? What are the caveats? How do they fit in with the second law of thermodynamics (and concepts of disorder and free energy)?*

**Author's response:** We have added more background information on the concept of entropy, specifically tailored towards biological systems.

**Reviewer's comment:** *The framework presented in this study is built on Stoy et al., 2014, which in turn used a formulation by Holdway et al., 2010. These essentially simplify the concept of entropy to temperature normalization of fluxes of energy, carbon and water exchange. While a temperature normalized index for these quantities is likely to be highly useful in itself, does it warrant invoking entropy? Moreover, there are several inconsistencies, and not adequate explanation for how entropy for different fluxes is estimated. For instance, eq 4.6. which the authors define as the entropy efficiency of metabolism, is essentially a ratio of NEE:GPP. This has been previously identified as carbon use efficiency and extensively studied (for. e.g. see DeLucia et al., 2007 and references therein). In many instances, it is unclear how energy and entropy are related. It would be useful to present side-by-side comparisons.*

**Author's response:** You are correct; the concept of thermodynamic entropy applied in our study is essentially a normalization of energy fluxes to temperature, as the magnitude of entropy fluxes and entropy production is a function of the temperature from which flux originated. This can be helpful in determining differences in energy use efficiency in ecosystems, specifically for the sites as these differed in sky, air, surface and soil temperatures. We are now using half-hourly fluxes for all energy and entropy in our calculations. Furthermore, we have omitted the section on metabolic efficiency ratios and are now focusing on metabolic activity in terms of energy and entropy (see section 2.5 and 3.5).

### **Three examples**

**Reviewer's comment:** *1. Page 3, line 3: how does the entropy dissipation through sensible heat relate to energy dissipation? These concepts need to be clarified. 2. Fig. 4. Why look at  $J_{LE}$  instead of  $LE$  fluxes? What is additionally learned from this? 3. Page 10, line 31.  $J_{NEE}$  not being related to soil moisture. This claim (I say claim since data is not shown) would be highly interesting if it is contrasted with the NEE response to soil moisture. There are more rigorous formulations (e.g. Wu et al., 2017) as well as critical discussions (e.g. Volk and Paulus, 2010).*

**Author's response:** An analysis of entropy fluxes is preferable in ecosystems which are exposed to different environmental variables, such as differences in surface and air temperatures, which affect the magnitude of entropy fluxes and entropy production. For example, two systems could have similar magnitudes of  $LE$ , but differ in  $J_{LE}$  due to differences in air or surface temperatures. For the ecosystem which maintains a higher surface/air temperature, the entropy flux would be lower, suggesting that it is less efficient in exporting entropy across its boundaries. By calculating the difference between entropy outputs and inputs, as well as internal entropy production, one can estimate how close an ecosystem is to a thermodynamic "steady state" and

therefore how organized it is. This cannot be accomplished by studying the energy balance alone. We have added a more thorough introduction and discussion of why entropy metrics can be more useful in describing energy use efficiency.

**Reviewer's comment:** *Another cause for concern is that that inferences are not quantitatively supposed. There are several instances where analysis is restricted to 'eyeballing' relationships between different curves, and correlation coefficients are not presented. In some occasions this leads to the authors making inferences that are not backed up by the data that is presented.*

**Author's response:** We have added tables of Type III effect summaries for all models in the supplementary materials.

**Reviewer's comment:** *The writing is overly descriptive, and often disconnected with the conclusions. Is this study describing entropy fluxes and efficiency ratios and how these vary with different environmental conditions, or is it trying to use these variables to understand site differences? The result is an unclear combination of the two. I would recommend the Authors' to stick to a storyline that is supported by the data.*

**Author's response:** Thank you for this valuable comment. We used environmental variables to understand changes and differences in entropy production and fluxes and thus changes in energy efficiency at our three sites. We have added an explanation of the objective of this study.

**Reviewer's comment:** *Finally, there are several instances where the authors discuss the effect of soil moisture and rainfall on various fluxes/processes in the text (e.g. lines 13, 19, 31 on page 10, line 25 on page 11) but do not choose to show these data. In my opinion these data are critical and need to be discussed (since it is a drought recovery study).*

**Author's response:** We have added figures showing all significant effects included in our models.

**Reviewer's comment:** *In light of these observations, I would not recommend this manuscript in its current form for publication in Biogeosciences. I think the authors provide very valuable observations, but should consider either re-framing the study or provide a more critical discussion on the concept of MEP, as well as consider extensive revisions on the writing as well as presentation of data.*

**Author's response:** We have altered the introduction and discussion to reflect your comment.

## Figures

**Reviewer's comment:** *There are several instances where curves are classified as significantly different, but do not appear significantly different from each other at all (Fig. 1d, for instance). The authors need to expand figure captions, since in the current form it is hard to infer what is being shown. E.g. Figure 4 has three time series (one for each site in most panels) but only one for sub panels b and e. It is unclear what data are presented. There are similar issues with Figs. 5-7.*

**Author's response:** We have added supplementary tables with all type 3 effects for all models included in this manuscript to show where there are significant differences among the independent variables. However, we note in figure captions when interactive effects were not significant, thus only showing a single black line (for example see fig. 6 for panels (g), (h) and (o)).

**Reviewer's comment:** *I also feel that the authors rely on too much on summarizing data and do not explain how or why this is done (again, eg. Fig 4b and d). What are the data that are presented in these analyses?*

**Author's response:** We have changed our analysis to estimating entropy from mean half hourly energy fluxes to daily time-steps ( $\text{W m}^{-2} \text{K}^{-1}$ ).

**Reviewer's comment:** *The authors need to include sub panels in the text (Fig. 4a, b etc.).*

**Author's response:** Thank you for this suggestion; we have added sub-panel information to the text.

**Reviewer's comment:** *Figure 1 has inconsistent units for temperature. For instance, subpanels c and e are plotted in units of Kelvin but d and f are in deg. C. Also, VPD is plotted in Figure 1 but not discussed at all amongst other discussions of Fig. 1 (Sec. 3.1).*

**Author's response:** We have added Figure 2 for sky, air, surface and soil temperatures with consistent units. We have also added more text describing differences in VPD (see section 3.1).

**Reviewer's comment:** *Fig 2. Why are monthly means shown here, while the rest of the paper annual means are presented?*

**Author's response:** We are now using daily average half hourly estimates in our models. All figures will show these values on the same timestep, with the exception of SI Figure S1, as it seemed more appropriate to show monthly sums of rainfall, as differences among the sites and years became more apparent this way.

**Reviewer's comment:** *Table 1: Please provide LAI estimates (if available) and also disturbance history, since this is a key component of your overall conclusions.*

**Author's response:** We have added LAI data for the mesic and xeric site. Unfortunately that information was not available for the intermediate site. We have also added fire disturbance history for all sites to Table 1.

## Minor comments

**Reviewer's comment:** *Page 2 Line 1-2: Turbulent exchange of... specify (for e.g. momentum, heat, gases). Line 3: Maybe just use examples related to terrestrial ecosystems? Are these examples of the butterfly effect in terrestrial ecosystems?*

**Author's response:** We have adjusted the sentence accordingly.

**Reviewer's comment:** *Page 5 Lines 5-9: This assumes energy balance closure. Please describe why you closed the energy balance.*

**Author's response:** To accurately describe the entropy balance for ecosystems, we are required to have a closed energy balance. Following your and reviewer 1's comments, we have added a more detailed description of the energy balance and energy partitioning at the three sites (see section 2.2)

**Reviewer's comment:** *Page 6 Eq. 2: Describe briefly how NEE was partitioned into source and sink terms.*

**Author's response:** We have added a more detailed description on how we partitioned NEE into GEE and  $R_{\text{eco}}$  following Whelan et al. (2013) and Starr et al. (2016).

**Reviewer's comment:** *Page 7 eq. 3.6. and 3.7: Unclear why net fluxes are used. Line 23: Are periods of rainfall excluded from the analyses? Where is this described? eq 4.1 and 4.2: Why is 4.1. formulated using incoming radiation whereas as 4.2 using net fluxes?*

**Author's response:** We have revised this section and are estimating entropy production as follows:

$$\boxed{\sigma R_s = R_{s,\text{in}} - R_{s,\text{out}} \left( \frac{1}{T_{\text{sr}}f} - \frac{1}{T_{\text{sun}}} \right)}$$
$$\boxed{\sigma R_l = R_{l,\text{in}} \left( \frac{1}{T_{\text{sr}}f} - \frac{1}{T_{\text{sun}}} \right)}$$

For the actual entropy production calculation only shortwave radiation that was absorbed by the ecosystem would be converted into heat, whereas for the MEP calculation we quantified the maximum entropy production, assuming that an efficient ecosystem would absorb all  $R_{s,\text{in}}$ . In contrast, for  $R_{l,\text{in}}$  energy is absorbed and then reemitted by the ecosystem rather than reflected.

**Reviewer's comment:** *Page 8 eq. 4.8 is essentially carbon use efficiency (see major comment above).*

**Author's response:** We have omitted the section on metabolic ratios and focus on energy and entropy of the metabolic system using NEE.

**Reviewer's comment:** *Page 9 Line 11. Subpanels missing. Lines 21-24: temperatures differences do not appear to be significantly different across sites in Fig. 1.*

**Author's response:** We kindly note that even though these figures appear to not show significant differences in temperature across the sites, our statistical results indicate that there were in fact significant differences in temperature. We have added supplementary type-3 results for all models to show this.

**Reviewer's comment:** *Page 10. Sec. 3.2. Methods for this analysis are not presented. I think this section should be merged with Sec. 2.1. (site description), as it doesn't appear to be a result of this study (unless methods are presented). Line 14: Soil moisture data seems important here (and in other places). Line 15: VPD effects are discussed first but EVI figure shown first in Fig. 4. Line 23: This is not correct according to Fig. 4. Line 23: See major comment above.*

**Author's response:** We have added a description of the methods that were used to estimate understory biomass for the three sites. In addition we have added graphs for all significant effects in the models. The description of all variables has been ordered to be consistent with the text.

**Reviewer's comment:** *Page. 13 Line 1: What does 'preservation' on LE mean? Again, these are hard to interpret in the absence of absolute fluxes (see major comment above). Line 8: Ecosystems do not 'experience' LE (or JLE), but rather the interactions between the ecosystem and the overlying atmosphere determines the LE flux. Line 13: Clarify what this means.*

**Author's response:** As noted above, we have added an analysis of the energy balance to show absolute fluxes. We also have changed the wording and made sure to not "personalize" ecosystems throughout this manuscript.

**Reviewer's comment:** *Page 14 Line 8: should read "at the more biodiverse site (i.e. mesic)" Line 11: What was the contribution of the C4 understory photosynthesis to overall ecosystem photosynthesis? Did you measure this? Lines 25-30: This is incorrect. Annual (and monthly) changes in EVI do not reflect changes in biomass. Biomass includes the carbon stored in the trunks, branches and stems of trees (among other pools), which do not fluctuate in forests at these timescales. Instead, at these timescales EVI is a measure of canopy greenness that is related to net photosynthesis (see Sims et al., 2008).*

**Author's response:** We have changed the sentence accordingly. Unfortunately, we have no estimates of variation in ecosystem fluxes from differences in understory composition at our sites. However, in another study we showed that the understory contributes about 50% to  $R_{\text{eco}}$ , using soil respiration data. We have also corrected the definition of EVI in the text.

# Quantifying energy use efficiency via entropy production: A case study from longleaf pine ecosystems

**Deleted:** maximum

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**Abstract.** Ecosystems are open systems that exchange matter and energy with their environment. They differ in their efficiency in doing so as a result of their location on Earth, structure, and disturbance, including anthropogenic legacy. Entropy has been proposed to be an effective metric to describe these differences as it relates energy use efficiencies of ecosystems to their thermodynamic environment (i.e. temperature) but has rarely been studied to understand how ecosystems with different disturbance legacies respond when confronted with environmental variability. We studied three sites in a longleaf pine ecosystem with varying levels of anthropogenic legacy and plant functional diversity, all of which were exposed to extreme drought. We quantified radiative ( $eff_{rad}$ ), metabolic and overall entropy changes – as well as changes in exported to imported entropy ( $eff_{flux}$ ) in response to drought disturbance, and environmental variability using 24 total years of eddy covariance data (8 years per site). We show that structural and functional characteristics contribute to differences in energy use efficiencies at the three study sites. Our results demonstrate that ecosystem function during drought is modulated by decreased absorbed solar energy and variation in the partitioning of energy and entropy exports owing to differences in site enhanced vegetation index and/or soil water content. Low  $eff_{rad}$  and slow adjustment of  $eff_{flux}$  at the anthropogenically altered site prolonged its recovery from drought by approximately one year. In contrast, stands with greater plant functional diversity (i.e., the ones that included both C3 and C4 species) adjusted their entropy exports when faced with drought, which accelerated their recovery. Our study provides a path forward for using entropy to determine ecosystem function across different global ecosystems.

## 1 Introduction

25 Ecosystems utilize resources, such as solar radiation, nutrients and water, to maintain a state far from thermodynamic equilibrium (Amthor, 2010; Beer et al., 2009; Finzi et al., 2007; Thomas et al., 2016). Understanding ecosystem resource use efficiency is crucial, as anthropogenic and climate induced changes around the globe continue to alter ecosystem structure and function (Haddeland et al., 2014; Porter et al., 2012; Reinmann and Hutyra, 2016; Thom et al., 2017).

30 Ecosystems are open and dynamic systems that exchange matter and energy with their surroundings as described by the ecosystem energy balance:

$$R_n = R_{s,in} - R_{s,out} + R_{l,in} - R_{l,out} = LE + H + G + M \quad (1)$$

**Deleted:** Global ecosystems vary

**Deleted:** function, and therefore resilience to disturbance,

**Deleted:** Resilience can therefore

**Deleted:** difficult to describe solely based on energy partitioning, as it fails to effectively describe how ecosystems use available resources, such as soil moisture. Maximum entropy production (MEP) has been shown to be a better

**Deleted:** the availability of resources.

**Deleted:** biodiversity

**Deleted:** their resilience from

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**Deleted:** MEP ratios. Sites with anthropogenic legacy had ~10% lower overall...

**Deleted:** metabolic

**Deleted:** efficiency compared to more biodiverse

**Deleted:** This resulted in lower resilience and a delay in

**Deleted:** ~1

**Deleted:** Additionally, a set of entropy ratios

**Deleted:** metabolic and overall energy use efficiency explained more clearly site-specific

**Deleted:** , whereas the radiative entropy budget gave more insights about structural complexities at the sites. Our study provides foundational evidence of how MEP can be used to determine resiliency...

**Deleted:** globally

**Deleted:** Energy use efficiency in ecosystems comprises the efficient utilization and conversion of

**Deleted:** the elements affecting energy

**Deleted:** efficiencies

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**Deleted:** resources

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where  $R_n$  is net radiation,  $R_{s,in}$  and  $R_{s,out}$  are incident and upwelling shortwave radiation, and  $R_{l,in}$  and  $R_{l,out}$  are incoming and upwelling longwave radiation, respectively. The terms LE, H and G represent energy exports through latent heat, sensible heat and ground heat fluxes, respectively and M is an energy storage term comprised of changes in biomass accumulation through metabolic processes (Holdaway et al. 2010). M is often neglected due to the assumption of a steady state over longer periods and because M is much smaller in magnitude compared to other fluxes, but it imposes a control on energy fluxes, like  $R_n$ , LE and H, through changes in leaf area and reflective properties, as well as through active biotic control in response to changes in environmental variables (i.e., stomata opening and closing due to water availability (Hammerle et al., 2013).

From equation 1, ecosystem energy exchange is a function of its thermodynamic environment - the heat transfer of a system with its surroundings - which differs based on the different mechanisms by which heat is transported: conduction, convection, radiation. Complicating our understanding of ecosystem energy dynamics is the fact that more frequent fluctuations in environmental variables are expected as a result of global climate change, including extreme events like droughts, which will alter the resource efficiency of ecosystems across the globe and with it their resilience (Franklin et al., 2016; Woodward et al., 2010). It is hypothesized that ecosystems aim to optimize their energy use and thus maximize their balance of entropy production and entropy exports to avoid thermodynamic equilibrium (Schneider and Kay, 1994; Schymanski et al., 2010). The magnitude of entropy production and entropy fluxes in ecosystems depends on thermodynamic gradients (i.e., thermal gradients, chemical gradients, etc.) between organisms and their surroundings (Kleidon, 2010). Ecosystems invest energy to build more complex structures (i.e., self-sustainability; Müller and Kroll, 2011; Virgo and Harvey, 2007), which can enhance their entropy export and therefore keep the ecosystem far from thermodynamic equilibrium (Odum, 1988; Schneider and Kay, 1994; Holdaway et al., 2010; Skene, 2015). For example, forest stands with more vertical structure were found to be more efficient in harvesting available light, which consequently increased their productivity (Bohn and Huth, 2017; Hardiman et al., 2011). Productive sites with greater leaf area can maintain higher latent energy (LE) fluxes, which increases their entropy export (Meysman and Bruers, 2010; Brunsell et al., 2011); LE fluxes also maintain lower ecosystem surface temperatures and thereby greater entropy production. On the contrary, large values of H caused by surface temperatures that are greater than air temperatures, result in lower entropy production (LeMone et al., 2007). This has been shown in deforested landscapes (Bonan, 2008; Khanna et al. 2017), as well as comparative studies of different vegetation types and in ecosystems with heterogeneity in their vegetation distribution (Holdaway et al., 2010; Brunsell et al., 2011; Kuricheva et al., 2017). Here, we evaluate how efficiently ecosystems use energy by assessing ecosystem entropy production as well as by quantifying the ratios in entropy imports and exports ( $eff_{flux}$  and  $dS/dt$ ) in three study ecosystems that represent an edaphic and management gradient. We do so by measuring their structural complexity over an eight-year period via the enhanced vegetation index (EVI) and variation in annual understory biomass, and in relation to the energy and entropy partitioning of incoming energy from solar radiation. We build upon the techniques proposed by Holdaway et al. (2010), Brunsell et al. (2011), and Stoy et al. (2014), by calculating entropy production and entropy fluxes within longleaf pine (*Pinus palustris* Mill.) ecosystems. The sites differed in ecosystem structure (i.e., basal area, Table 1) and plant functional diversity due in part to differences in soil water holding capacity, as well as different levels of anthropogenic legacy. The sites were exposed to a severe drought in the beginning of

**Deleted:** Ecosystem energy use efficiency is often described based on the ability of the system to partition energy into sensible and latent heat fluxes (Hurnya and Pokorný, 2016; Ripl, 1995; Wilson, 2002). However, energy fluxes differ substantially within and among ecosystems, due to differences in structure and function (Kang et al., 2015), which regulate their base rate of energy use (Brunsell et al., 2011). For example, ecosystems exhibit small-scale changes in environmental variables based on their structural properties (Perry, 2002). This spatial heterogeneity can drive differences in microclimate which lead to changes in the turbulent exchange between the land and atmosphere (Perry, 2002). Additionally, differences in initial conditions during the development of ecosystems can have large impacts on their trajectory (e.g., "butterfly-effect"; Dantas-Torres, 2015; Hastings, 1993). The combination of these factors leads to different baselines of energy use efficiency and may even cause instabilities in ecosystems as they further develop (Mori, 2011).

**Deleted:** This has the potential to further intensify instabilities, which may shift or disrupt the dynamic equilibriums of ecosystems (Mori, 2011; Siteur et al., 2016).

**Deleted:** The use of maximum entropy production (MEP) has been proposed as a metric to address ecosystem energy use efficiency (Kleidon et al., 2010; Stoy et al., 2014; Swenson, 1989) based on the notion that the second law of thermodynamics can be applied to whole ecosystems to describe their strategy for resource utilization (e.g. carbon and water fluxes; Dewar, 2005; Whitfield, 2005; 2007).

**Deleted:** MEP is therefore thought to be the most probable outcome of the different ways elements in non-equilibrium can reorganize themselves (Jaynes, 1957b; 1957a), corresponding to a dynamic stable state of an ecosystem with its surroundings (Dyke and Kleidon, 2010; Schneider and Kay, 1994; Svirzhev, 2000). Hence, the principle of MEP allows us to make the least unlikely projection of a system responding to any kind of change (Endres, 2017; England, 2015).<sup>1</sup>

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**Deleted:** ability to build thermodynamic gradients. For example, agricultural systems were shown to exceed MEP, due to the application of fertilizer and periodic irrigation schedules (i.e., artificial energy; Patzek, 2008; Steinborn and Svirzhev, 2000). This excessive entropy production is unsustainable and decreases the ability of ecosystems to adapt to perturbations (Cochran et al., 2016). This is because agricultural practices compensate for natural fluctuations in resource availability, which would normally ... [1]

**Deleted:** Here, we evaluate how efficiently ecosystems use incoming energy by assessing individual components of an ecosystem's entropy production in comparison to the empirical MEP of the system (Stoy et al., 2014). We do so by measuring the ... [2]

**Deleted:** Holdaway et al. (2010) and Brunsell et al. (2011) by calculating entropy ratios to compare and contrast three sites (mesic, intermediate and xeric) within a longleaf pine (*Pinus palustris* Mill.) ecosystem. The three longleaf pine sites differed in ecosystem ... [3]

this study, which we used to quantify entropy exchanges in response to the disturbance. First, we compare and contrast differences in ecosystem energy fluxes (i.e.,  $R_n$ , LE, H, G and the net ecosystem exchange of carbon dioxide, NEE) and entropy fluxes ( $J_{LE}$ ,  $J_H$ ,  $J_G$ ,  $J_{NEE}$  and radiative entropy production) in response to changes in structural and environmental variables (EVI, SWC, VPD, and precipitation). Next, we quantify how entropy exports and entropy production at the different sites 5 adjust to changes in incoming entropy when exposed to drought. We do so by estimating radiative efficiency ( $eff_{rad}$ ), the ratio of entropy production to an empirical maximum entropy production (MEP), and ratios of daily imported and exported entropy fluxes ( $eff_{flux}$ ), as well as through the overall change in entropy ( $dS/dt$ ) at the sites. We hypothesize that: (1) the xeric site will have a higher entropy flux from  $J_H$  and  $J_G$ , but lower  $J_{NEE}$  due to its lower EVI and lower basal area, which will result in more 10 variable  $dS/dt$  compared to the other sites; (2) the mesic site will maintain higher  $eff_{rad}$  due to its greater structural complexity (i.e., plant functional diversity and basal area) and thus greater absorptive capacity for solar radiation compared to the other sites; (3) the intermediate site will have lower  $eff_{rad}$  and  $eff_{flux}$  compared to the mesic and xeric sites, as a result of its lower plant functional diversity (i.e. low abundance of C4 species) and structural complexity, causing lower absorption of solar radiation and export of entropy through LE.

## 2 Materials and Methods

### 15 2.1 Site description

This study was conducted at the Joseph W. Jones Ecological Research Center in southwestern Georgia, USA ( $31.2201^\circ N$ ,  $84.4792^\circ W$ ) from January 2009 to December 2016. The three sites are maintained by frequent low intensity fire on a two-year return interval and were last burned in 2015 (Starr et al., 2016). The climate is humid subtropical with a mean annual precipitation of 1310 mm (Kirkman et al., 2001). Mean temperature extremes range from  $3^\circ C$  to  $16^\circ C$  in winter and  $22^\circ C$  to 20  $33^\circ C$  in summer (NCDC, 2011).

The three sites differ based on soil moisture availability, as a result of differences in soil drainage. The mesic site lies on somewhat poorly drained sandy loam over sandy clay loam and clay textured soils (Goebel et al., 1997; 2001). Soils at the intermediate site are well drained and have a depth to the argillic horizon of  $\sim 165$  cm (Goebel et al., 1997). The xeric site lies on well-drained deep sandy soils with no argillic horizon (Goebel et al., 1997). All sites are situated within 10 km of each other 25 and have average elevations of 165, 155, and 160 m for the mesic, intermediate, and xeric sites, respectively.

95-year-old longleaf pine trees (*Pinus palustris* Mill.) dominate the overstory of all sites, and overall basal area (BA) and diameter at breast height (DBH) varied by site (Table 1). The overstories of each site also contain a small proportion of oak trees; The xeric site has the highest proportion, with 22 %, versus 8 % and 7.7 % at the mesic and intermediate sites, respectively. The understory at the mesic and xeric sites is largely covered with perennial C4 grass species, such as wiregrass 30 (*Aristida beyrichiana* [Trin.]), whereas woody species dominate the intermediate site. composition and abundance of other plant species varies by site (Kirkman et al., 2001; 2016). Soil perturbation at the intermediate site affected species richness, so that wiregrass is almost absent.

We acquired EVI for 2009 through 2016 for all three sites from the online data pool at [lpdaac.usgs.gov](http://lpdaac.usgs.gov) via the NASA Land Processes Distributed Active Archive Center (LP DAAC) and the USGS Earth Resources Observation and Science Center

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**Deleted:** We hypothesize that (1) the xeric site will have a higher entropy flux from sensible heat ( $J_H$ ) due to its lower vegetation coverage and lower basal area, and therefore higher partitioning into H rather than LE, which will lower its overall entropy efficiency compared to the mesic site, (2) the mesic site will be able to maintain higher entropy efficiencies ( $eff_{rad}$ ,  $eff_I$  and  $eff_M$ ) due to its structural complexity (i.e., species diversity and basal area), and (3) the intermediate site will have lower  $eff_{rad}$ ,  $eff_I$  and  $eff_M$  compared to the mesic and xeric sites, as a result of its lower biodiversity and structural complexity, a consequence of its anthropogenic legacy.<sup>1</sup> We show that a combination of different entropy efficiency ratios revealed how structural and functional characteristics in this ecosystem contribute to energy use efficiencies at the three sites. Our results demonstrate that resilience to disturbance is a function of entropy efficiency, as low efficiencies at anthropogenically altered sites prolonged the recovery from drought by approximately a year. In contrast, more biodiverse stands (in overstory and understory) were able to maintain high entropy/energy use efficiencies, which increased their resilience to and recovery from drought.<sup>2</sup>

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(EROS), using MODIS Aqua and Terra data products (MYD13Q1 and MOD13Q1; DAAC, 2008) to quantify changes in ecosystem structure from disturbance. EVI products for the sites were available on an eight day basis and linearly interpolated to obtain daily estimates. We also acquired Palmer Drought Severity Indices (PDSI) for Southwest Georgia from the National Oceanic and Atmospheric Administration data archive for 2009 to 2016 to identify the months of drought disturbance (Dai et al., 2004).

Understory composition and biomass was estimated annually from 2009 through 2013. Thereafter, the collection frequency became biannual, so that 2014 and 2016 were missing in the data collection. Understory biomass was estimated using 0.75 m<sup>2</sup> clip plots, which were randomly located by tossing a plot frame from pre-installed litter trap positions (n = 20 per site; Wiesner et al. 2018). All live and dead vegetation, smaller than 1 m in height was clipped and analyzed in our laboratory. Vegetation was classified by plant family (here, forbs, ferns, legumes, wiregrass, other grasses, and woody plants), and each sample was dried to constant weight.

Net ecosystem exchange of CO<sub>2</sub> measurements. Net ecosystem exchange (NEE) was measured continuously at 10 Hz at all three sites from January 2009 to December 2016 using open-path eddy covariance (EC) techniques (Whelan et al., 2013). Data were stored on CR-5000 dataloggers (Campbell Scientific, Logan, UT). CO<sub>2</sub> and water vapor concentration were measured

15 with an open path infrared gas analyzer (IRGA, LI-7500, LI-COR Inc., Lincoln, NE) and wind velocity and sonic temperature were measured with a three-dimensional sonic anemometer (CSAT3, Campbell Scientific, Logan, UT). These sensors were installed ~4 m above mean canopy height at each site (34.5, 37.5, and 34.9 m for the mesic, intermediate and xeric sites, respectively), ~0.2 m apart to minimize flow distortion between the two instruments and vertically aligned to match the sampling volume of both instruments.

## 20 2.2 Sensible and latent heat flux measurements

Net energy fluxes of LE and H were estimated in W m<sup>-2</sup> using temperature and wind velocity measurements from the sonic anemometer, as well as water vapor density measurements from the IRGA:

$$LE = \lambda \rho_a \bar{w}' \bar{q}' \quad (2.1)$$

$$H = \rho_a c_p (\bar{w}' \bar{T}'_s - 0.000321 \bar{T}_s \bar{w}' \bar{q}') \quad (2.2)$$

25 where  $\lambda$  is the latent heat of vaporization (J kg<sup>-1</sup>),  $\rho_a$  is the density of air (kg m<sup>-3</sup>),  $c_p$  is the specific heat of air (kJ kg<sup>-1</sup> K<sup>-1</sup>),  $\bar{w}'$  is the instantaneous deviation of vertical wind speed (w, m s<sup>-1</sup>) from the mean, and  $\bar{q}'$  and  $\bar{T}'_s$  are the instantaneous deviations of water vapor concentration (kg kg<sup>-1</sup>) and sonic temperature (Kaimal and Gaynor, 1991) from their respective means. The overbars in Eqs. 1.1 and 1.2 signify the time-averaged covariance. Missing H and LE were gap-filled on a monthly basis using simple linear models as a function of R<sub>n</sub>.

30 In cases where energy balance closure was not achieved, energy fluxes of H and LE were corrected using the Bowen method following Twine et al. (2000), where fluxes are adjusted using residual energy, and the estimated Bowen ratio ( $\beta = H/LE$ ), which assumes that  $\beta$  was correctly measured by the EC system.

$$LE = \frac{1}{1+\beta} (R_n - G) \quad (2.3)$$

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 $H = \rho_a c_p (\bar{w}' \bar{T}'_s - 0.000321 \bar{T}_s \bar{w}' \bar{q}')$  (1.2)

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$$H = \beta \times LE$$

(24)

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Closing the energy balance is important to quantify differences in energy and entropy fluxes by site, as according to the First law of Thermodynamics energy is always conserved. To quantify differences in environmental drivers and site variation between energy and entropy fluxes, we established models of average daily energy fluxes (described in section 2.7)

### 5 2.3 Meteorological instrumentation

Meteorological data above the canopy were also collected and stored on the CR-5000 dataloggers (Campbell Scientific, Logan, UT). Meteorological data measured on the towers included: photosynthetically active radiation (PAR; LI-190, LI-COR Inc., Lincoln, NE), global radiation (LI-200SZ, LI-COR Inc., Lincoln, NE), incident and outgoing shortwave and longwave radiation to calculate  $R_n$  (NR01, Hukseflux, thermal sensors, Delft, The Netherlands), precipitation (TE525 Tipping Bucket Rain Gauge, Texas Electronics, Dallas, TX), wind direction and velocity (Model 05103-5, R.M. Young, Traverse City, MI), air temperature ( $T_{air}$ ) and relative humidity (RH; HMP45C, Campbell Scientific, Logan, UT), and barometric pressure (PTB110, Vaisala, Helsinki, Finland).

Soil temperature ( $T_{soil}$ ), volumetric water content of the soil (SWC) and soil heat flux (G) were measured in one location near the base of each tower at each site every 15 seconds and averaged every 30 minutes on an independently powered CR10X datalogger.  $T_{soil}$  was measured at depths of 4 and 8 cm with insulated thermocouples (Type-T, Omega Engineering, INC., Stamford, CT), and G was measured at a depth of 10 cm with soil heat flux plates (HFP01, Hukseflux, Delft, The Netherlands). SWC was measured within the top 20 cm of the soil surface using a water content reflectrometer probe (CS616, Campbell Scientific, Logan, UT).

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### 2.4 Data processing

20 Raw EC data were processed using EdiRe (v.1.4.3.1184; Clement, 1999), which carried out a two-dimensional, coordinate rotation of the horizontal wind velocities to obtain turbulence statistics perpendicular to the local streamline. Fluxes were calculated for half-hour intervals and then corrected for mass transfer resulting from changes in density not accounted for by the IRGA. Barometric pressure data were used to correct fluxes to standard atmospheric pressure. Flux data screening was applied to eliminate 30-min fluxes of NEE, H and LE, resulting from systematic errors as described in Whelan et al. (2013)

25 and Starr et al. (2016). Such errors encompassed (amongst other things): rain, poor coupling of the canopy and the atmosphere (defined by the friction velocity,  $ustar$ ), and excessive variation from half-hourly means.

Gross ecosystem exchange (GEE) and ecosystem respiration ( $R_{eco}$ ) were estimated from eddy covariance measurements of net ecosystem exchange of  $CO_2$  (NEE;  $\mu mol m^{-2} s^{-1}$ ) at a time resolution of half an hour, from which GEE and  $R_{eco}$  can be estimated as follows:

30  $GEE = -NEE + R_{eco}$  (3)  
Missing half hourly data were gap-filled as described in Whelan et al. (2013) and Starr et al. (2016). Daytime and nighttime data were estimated utilizing a Michaelis-Menten approach for ( $PAR > 10 \mu mol m^{-2} s^{-1}$ ) and a modification of the Lloyd and Taylor (1994) model ( $PAR \leq 10 \mu mol m^{-2} s^{-1}$ ), respectively. Monthly equations were used to gap-fill data; however, where too

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few observations were available to produce stable and biologically reasonable parameter estimates, annual equations were used. NEE partitioning to estimate daytime  $R_{\text{eco}}$  was performed by using the nighttime gap-filling equation, and then utilizing equation (3) to estimate GEE. Nighttime GEE was assumed to be zero.

## 2.5 Entropy production calculations

5 **Half-hourly**, Net ecosystem exchange of  $\text{CO}_2$  was converted to  $\text{W m}^{-2}$  ( $\text{NEE}_e$ ), using the assumption that one micromole of  $\text{CO}_2$  stores approximately 0.506 J, where  $1 \text{ J m}^{-2} \text{ s}^{-1}$  equals  $1 \text{ W m}^{-2}$  (Nikolov et al., 1995).

For entropy production and fluxes of shortwave ( $R_s$ ) and longwave radiation ( $R_l$ ) we followed established approaches of Brunsell *et al.* (2011), Holdaway *et al.* (2010), and Stoy *et al.* (2014). The **half-hourly** entropy flux produced through **absorption** of  $R_s$  **emitted by** the surface of the sun ( $J_{Rs}$ ,  $\text{W m}^{-2} \text{ K}^{-1}$ ) was calculated as:

$$10 \quad J_{Rs} = \frac{R_{s,\text{net}}}{T_{\text{sun}}} \quad (4.1)$$

where sun surface temperature ( $T_{\text{sun}}$ ) was assumed to be 5780 K, with  $R_{s,\text{net}}$  defined as **the difference of incident and upwelling**  $R_s$ . **The** entropy flux of  $R_l$  ( $J_{Rl}$ ,  $\text{W m}^{-2} \text{ K}^{-1}$ ) **was** calculated as:

$$11 \quad J_{Rl} = \left( \frac{R_{l,\text{in}}}{T_{\text{sky}}} - \frac{R_{l,\text{out}}}{T_{\text{srf}}} \right) \quad (4.2)$$

where  $R_{l,\text{in}}/T_{\text{sky}}$  is **the** entropy flux of  $R_{l,\text{in}}$  as incoming  $R_l$  ( $J_{Rl,\text{in}}$ ), and  $R_{l,\text{out}}/T_{\text{srf}}$  is **the** entropy flux of  $R_{l,\text{out}}$  as outgoing  $R_l$  ( $J_{Rl,\text{out}}$ ).

15 surface temperature ( $T_{\text{srf}}$ , K) was calculated from upwelling  $R_l$  ( $J_{Rl,\text{out}}$ ):

$$16 \quad T_{\text{srf}} = \left( \frac{R_{l,\text{out}}}{(A \times e_{\text{srf}} \times k_B)} \right)^{1/4} \quad (4.2.1)$$

with emissivity of the surface calculated as  $e_{\text{srf}} = 0.99 - 0.16\alpha$  (Juang *et al.*, 2007), the **view factor**,  $A$  was assumed to be **unity**, and the Stefan-Boltzmann constant  $k_B = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$ . The shortwave albedo ( $\alpha$ ) was calculated as the **daily**, average of noontime outgoing  $R_s$  ( $R_{s,\text{out}}$ ) divided by  $R_{s,\text{in}}$ . The sky temperature,  $T_{\text{sky}}$  (K), was calculated from  $R_{l,\text{in}}$  using the Stefan-

20 Boltzmann equation:

$$17 \quad T_{\text{sky}} = \left( \frac{R_{l,\text{in}}}{(A \times e_{\text{atm}} \times k_B)} \right)^{1/4} \quad (4.2.2)$$

where the emissivity of the atmosphere ( $e_{\text{atm}}$ ) was assumed to be 0.85, following Campbell and Norman (1998).

All other ecosystem entropy fluxes  $J_{LE}$ ,  $J_H$ ,  $J_G$ , and  $J_{NEE}$  ( $\text{W m}^{-2} \text{ K}^{-1}$ ) were calculated by dividing the energy fluxes by temperature as:

$$25 \quad J_x = \frac{x}{T_y} \quad (4.3)$$

where  $x = LE, H, G$  and  $NEE_e$ , and  $T_y$  **was** assumed to be  $T_{\text{air}}$  (for  $J_{LE}$ ,  $J_H$  and  $J_{NEE}$ , K) or  $T_{\text{soil}}$  (for  $J_G$ , in K). We used the energy of  $NEE$  ( $NEE_e$ ) directly measured with the EC towers, to calculate **the** change in entropy of the metabolic system.

We also calculated entropy produced from evaporation associated with mixing of saturated air from the canopy with the fraction of air in the atmosphere that has RH below 100 % ( $JLE_{\text{mix}}$ ), following Holdaway *et al.* (2010):

$$30 \quad JLE_{\text{mix}} = ET \times R_v \times \ln(RH) \quad (4.4)$$

**Deleted:** LE, H, and G fluxes were converted from  $\text{W m}^{-2}$  to  $\text{kJ m}^{-2}$  and summed to monthly estimates.

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where the evapotranspiration rate,  $ET = LE \times \rho_a \lambda$ ,  $\lambda$  is the latent heat of vaporization of water, and  $R_v$  is the gas constant of water vapor (0.461  $\text{kJ kg}^{-1} \text{K}^{-1}$  for moist air).

The sum of entropy of ecosystem fluxes ( $J$ ,  $\text{W m}^{-2} \text{K}^{-1}$ ) for each half-hour was then calculated by adding all entropy fluxes between the surface and atmosphere:

$$5 \quad J = J_{RI} + J_{RS} + J_{LE} + J_H + J_G + J_{NEE} + JLE_{mix} \quad (4.5)$$

The conversion of low entropy  $R_s$  and  $R_i$  to high entropy heat at the surface through absorption of  $R_s$  and  $R_i$ , respectively, was calculated as:

$$\sigma_{RS} = R_{s,net} \left( \frac{1}{T_{srf}} - \frac{1}{T_{sun}} \right) \quad (4.6)$$

$$\sigma_{RI} = R_{i,in} \left( \frac{1}{T_{srf}} - \frac{1}{T_{sky}} \right) \quad (4.7)$$

10 where  $T_{srf}$  is the radiometric surface temperature (Eq. 3.2.1) and  $\sigma_{RS}$  and  $\sigma_{RI}$  are in  $\text{W m}^{-2} \text{K}^{-1}$ .

The overall half-hourly entropy production ( $\sigma$ ,  $\text{W m}^{-2} \text{K}^{-1}$ ) was then calculated as the sum of the entropy productions of  $R_s$  and  $R_i$ :

$$\sigma = \sigma_{RI} + \sigma_{RS}. \quad (4.8)$$

We excluded the factor 4/3, which is associated with the transfer of momentum exerted by electromagnetic radiation on a 15 surface (Wu et al., 2008), in our calculations of  $\sigma$  and  $J$  for entropy production and entropy fluxes because we assumed that radiation pressure at the sites would be negligible (see Ozawa et al. 2003; Kleidon and Lorenz, 2005; Fraedrich and Lunkeit, 2008; Kleidon, 2009; Pascale et al., 2012). Finally, we estimated half-hourly change in entropy production ( $S$ ) over time ( $t$ ) in  $\text{W m}^{-2} \text{K}^{-1}$  of the ecosystem by adding entropy flux of imports ( $J_{RS,net}$ ,  $R_{RI,in}$ ) and exports (i.e.,  $J_{LE}$ ,  $J_H$ ,  $J_G$ ,  $J_{NEE}$ ,  $J_{RI,up}$ ,  $JLE_{mix}$ ) and entropy production:

$$20 \quad \frac{dS}{dt} = J + \sigma. \quad (4.9)$$

Note that this approach does not account for entropy production due to frictional dissipation of entropy from rainfall or subsurface water flow, as these would be of even smaller magnitude than entropy production from metabolic activity of the ecosystem (Brunsell et al., 2011). Here negative  $dS/dt$  represents the export of entropy of the ecosystem to its surroundings.

## 2.6 Ecosystem entropy models for radiation and ecosystem fluxes

25 We estimated half-hourly MEP of the radiation budget ( $MEP_{rad}$ ) in  $\text{W m}^{-2} \text{K}^{-1}$ , to compare site differences in radiation energy use and entropy dissipation.

Empirical MEP ( $MEP_{rad}$ ) was determined following Stoy et al. (2014), by estimating the MEP of half-hourly  $R_s$  ( $MEP_{RS}$ ) and  $R_i$  ( $MEP_{RI}$ ):

$$MEP_{RS} = R_{s,in} \left( \frac{1}{T_{air}} - \frac{1}{T_{sun}} \right) \quad (5.1)$$

$$30 \quad MEP_{RI} = R_{i,net} \left( \frac{1}{T_{srf}} - \frac{1}{T_{air}} \right) \quad (5.2)$$

$$MEP_{rad} = MEP_{RS} + MEP_{RI}. \quad (5.3)$$

**Deleted:**  $JLE_{mix} = ET \times R_v \times \ln(RH)$  (3.4)

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**Deleted:**  $J = J_{RI} + J_{RS} + J_{LE} + J_H + J_G + J_{NEE} + JLE_{mix}$  (3.5)

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We estimated MEP of the radiation budget ( $MEP_{rad}$ ), metabolic

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We assume that under ideal conditions, an ecosystem maximizes its entropy production when it converts all incoming  $R_s$  and  $R_i$  into work (Stoy et al., 2014). While this assumption does not necessarily reflect reality in natural ecosystems, this method gives us a means to compare different sites with respect to their reflective and absorptive capacities *versus* a reference ecosystem that absorbs and dissipates all incident solar energy. Note that  $MEP_{RI}$  is often of lower magnitude than  $MEP_{RS}$ ,

5 because here we assume that an efficient ecosystem would dissipate less energy through sensible heat, such that  $T_{srf}$  would approach  $T_{air}$ .

The half-hourly entropy ratio of radiation is then calculated using  $\sigma$  from Eq. 4.8 as follows:

$$eff_{rad} = \frac{\sigma}{MEP_{rad}} \quad (5.4)$$

We then estimated the variable  $eff_{flux}$  as the ratio of incoming radiation entropy ( $J_{Rs}$  and  $J_{RI,in}$ ) and the sum of exported entropy 10 fluxes ( $J_{LE}$ ,  $J_H$ ,  $J_G$ ,  $J_{NEE}$ , and  $J_{RI,out}$ ) to assess how entropy was partitioned into entropy production and entropy fluxes over the different study years.

## 2.7 Statistical analyses

We estimated average daily values for all response variables to decrease autocorrelation for statistical analysis. We first tested for significant differences in environmental and structural variables among the three sites prior to the entropy analysis. We 15 estimated simple general linear mixed models (GLMM) using the R package *nlme* to look at differences among sites for rain, SWC, vapor pressure deficit (VPD), EVI,  $T_{srf}$ ,  $T_{air}$ ,  $T_{sky}$  and  $T_{soil}$ , as well as  $R_{s,in}$ ,  $R_{s,out}$ ,  $R_{i,in}$  and  $R_{i,out}$ . All response variables were daily means for rainfall we calculated monthly sums to estimate differences among the sites. We included a random effect for day of measurement to account for repeated measurements, as well as an AR(1) structure to account for temporal autocorrelation among measurements. The model of rainfall only included year and site as independent variables and no 20 random effects. Independent variables for the other models were month, year and site, as well as their interactions.

Subsequently, we estimated GLMMs of daily energy ( $R_{n}$ , LE, H, G and  $NEE_s$ ) and entropy fluxes ( $J_{LE}$ ,  $J_H$ ,  $J_G$ , and  $J_{NEE}$ ), entropy production ( $\sigma$ ), entropy ratios ( $eff_{rad}$  and  $eff_{flux}$ ) and overall entropy ( $dS/dt$ ) to quantify their differences by environmental and structural variables by site. For all models we included random effects and an AR(1) autoregressive 25 correlation structure to account for repeated daily measurements. All models initially included independent variables for site, year and month, mean EVI, SWC, VPD and daily rainfall sums. We also included interactions of environmental variables with site, site with year and site with month, to determine changes in the energy efficiency over the study period among sites. Independent variables and their interactions were deemed significant when  $p < 0.05$ . We used a Tukey adjustment to test for significant differences among sites. GLMM Analyses were performed via the R packages *nlme*, *lsmeans*, and *car* (Fox and Weisberg, 2011; Lenth, 2016; Pinheiro et al., 2014).

**Deleted:**  $MEP_{RS} = R_{s,in} \left( \frac{1}{T_{air}} - \frac{1}{T_{sun}} \right)$  (4.1)

$MEP_{RI} = R_{i,in} \left( \frac{1}{T_{srf}} - \frac{1}{T_{air}} \right)$  (4.2)

$MEP_{rad} = MEP_{RS} + MEP_{RI}$  (4.3)

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$$MEP_M = \frac{E_{in}}{T_{air}} \quad (4.5)$$

The entropy efficiency ratio of metabolism was then calculated from metabolic... ... [6]

**Deleted:** flux of actual available energy  $J_{NEE}$ , estimated as  $NEE_s/T_{air}$  (in K), divided by  $MEP_M$ :

$$eff_M = \frac{J_{NEE}}{MEP_M} \quad (4.6)$$

Maximum entropy of all energy fluxes ( $MEP_s$ ) was assumed as **Deleted:** maximum efficiency of an ecosystem, if all incoming energy from  $R_s$  and  $R_i$  was converted to work, which would only depend on  $T_{srf}$ :

$$MEP_s = \frac{R_{s,in} + R_{i,in}}{T_{srf}} \quad (4.7)$$

We then estimated overall site entropy efficiency as:

$$eff_s = \frac{dS/dt}{MEP_s} \quad (4.8)$$

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### 3 Results

#### 3.1 Differences in environmental, radiative and temperature variables among sites

All three sites experienced severe drought from mid-2010 through mid-2012 (Fig. S1, Supplementary Information). There was no significant difference between the mesic and xeric sites in rainfall sums, but the intermediate site had lower rainfall sums (~20 mm per month) compared to the other sites (Table S1). SWC was significantly lower at the xeric (<19 %) compared to mesic and intermediate sites (~20 %) for all years of this study (Fig. 1a and b, Table S2). SWC and EVI decreased during the drought at all sites, but only significantly so at the mesic site. VPD significantly increased at all sites during the drought. For all years, EVI was significantly lower (0.02-0.04) at the xeric site compared to the other two sites (Fig. 1e and f), while the intermediate site had significantly higher EVI compared to the mesic site, except in 2010.

Daily  $T_{srf}$  at the mesic site was significantly higher than the xeric site for all years except 2012, 2014, and 2016 (Fig. 2a). From 2012 to 2016 the intermediate site had higher  $T_{srf}$  compared to the other two sites.  $T_{air}$  was significantly lower at the mesic site compared to the intermediate and xeric sites for all years, except in 2014, and in 2012, when the xeric site had higher  $T_{air}$  compared to the intermediate (Fig. 2a).  $T_{soil}$  was significantly lower at the mesic site compared to the other sites, except in 2013, when there was no significant difference between the mesic and xeric sites. For all years, daily  $T_{soil}$  was significantly higher at the xeric site compared to the intermediate site except for 2011 and 2012, when the intermediate site was significantly higher.

$R_{s,out}$  was significantly higher at the xeric site compared to the other sites, except for 2014, where we found no significant difference between the intermediate and xeric sites. Daily  $R_{s,out}$  was also significantly lower at the mesic site, compared to the intermediate site, except in 2009. Average daily  $R_{l,out}$  was significantly lower at the mesic site compared to the intermediate site during all years, except for 2011 and 2012, and compared to the xeric site for all years, except for 2011. The intermediate site had significantly higher  $R_{l,out}$  compared to the xeric site during 2013, 2014 and 2016. As a consequence of these component fluxes,  $R_n$  was significantly higher at the xeric site compared to the intermediate site for all years except 2009 and 2014 (SI Fig. S2a, Table S3). Average  $R_n$  was significantly lower at the mesic site compared to the xeric site in 2013 and 2016, and was significantly higher compared to the xeric site from 2009 to 2011. Average daily  $R_n$  significantly increased at the intermediate and xeric sites but showed no change at the mesic site with an increase in EVI (SI Fig. S3a).

Environmental, radiative and temperature variables also tended to be significantly different among months within site, and in many instances among sites by month. Differences followed seasonal patterns, as noted in SI Fig. S2 and SI Table S2.

#### 3.2 Understory wiregrass and woody abundance at the sites

Wiregrass was virtually absent at the intermediate site for all years of this study (Fig. 4a), whereas woody species were more abundant compared to the others. The mesic and xeric sites both had higher proportions of wiregrass in the understory (~25 % versus 5 % at the intermediate site), which slightly decreased during 2011 (Fig. 4a). In addition, woody biomass increased to ~75 g m<sup>-2</sup> at the xeric site during 2011, but not at the mesic site. In 2012, woody biomass decreased to ~40 g m<sup>-2</sup> at the xeric

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and intermediate sites and remained low during the following years at the xeric site, but increased at the intermediate site ( $>100$  g m $^{-2}$ , Fig. 4b).

### 3.3 Energy fluxes of H, LE, and G

LE was significantly lower at the intermediate site compared to the mesic site for all years, except 2011, and compared to the xeric site for all years, except for 2015. We found no significant difference between the mesic and xeric sites in 2009, 2010, 2014 and 2016, but for the other years of this study the xeric site had significantly higher LE. LE significantly increased at all sites with higher EVI, with a greater increase at the intermediate and a smaller increase at the xeric site, compared to the mesic site (SI Fig. S3g). LE significantly increased at all sites with an increase in SWC and VPD (SI Fig. S3e and f). LE at the intermediate site was significantly lower compared to the other sites for all levels of VPD (SI Fig. S3g). LE was significantly lower with higher rainfall, with no significant differences among sites (SI Fig. S3h).

There was no significant difference in H between the mesic and intermediate sites, except in 2011 and 2013, when the mesic site was higher than the intermediate, and in 2015 and 2016, when the reverse occurred. H was significantly lower at the xeric site compared to the mesic site for all years except for 2014 and 2016, and compared to the intermediate site for all years except 2011 and 2013. Average H was significantly higher at the mesic site compared to the xeric site during the months of May through October (SI Fig. S2b). The intermediate site had significantly lower H compared to the other two sites for the months of January through March and the xeric site had significantly lower H for June through October. Compared to the other two sites, average H was significantly lower at the intermediate site when EVI was greater than 0.4, and significantly higher at the xeric site for EVI  $> 0.5$  (SI Fig. S3i). Average H significantly decreased at all sites with an increase in SWC (SI Fig. S3j). Average daily H significantly increased at all sites with an increase in VPD, with a lower decrease at the intermediate site (SI Fig. S3k).

G was significantly lower at the intermediate site during 2016 (negative), compared to 2009 through 2011 and 2014. Average daily G was positive during summer months, and negative during winter months (October through March) at all sites (SI Fig. S2b). Average daily G significantly decreased with an increase in EVI at the mesic and intermediate site, but had no significant change at the xeric site (SI Fig. S3m). G was significantly less positive at the xeric site compared to the other sites for EVI  $< 0.3$ , but was significantly more negative at the intermediate site compared to the mesic and xeric sites when EVI was above 0.4. Average G significantly decreased (to negative) with an increase in SWC (SI Fig. S3n), and significantly increased (to positive) with an increase in VPD, but only at the intermediate and xeric sites (SI Fig. S3o). Daily rainfall did not significantly alter G at the sites, but the intermediate had significantly more negative G compared to the other two sites (2-10 W m $^{-2}$ ) when daily rainfall was positive (SI Fig. S3p).

### 3.4 entropy fluxes of $J_H$ , $J_{LE}$ , and $J_{NEE}$ and $J_G$ and entropy production

For all years, average daily  $\sigma$  was significantly higher at the mesic site compared to the intermediate site (by  $> 0.01 - 0.036$  W m $^{-2}$  K $^{-1}$ ; Fig. 5a, Table S4), while  $\sigma$  was not significantly different between the mesic and xeric sites for almost all years (Fig 5a). Average daily  $\sigma$  significantly increased with EVI, independent of site (Fig. 6a), and also significantly increased with

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SWC and VPD, with a greater slope at the xeric site (Fig. 6b and c). Average daily  $\sigma$  significantly decreased at all sites with an increase in rainfall (noting that entropy production from rainfall itself is not considered here and assumed to be approximately equal amongst ecosystems), and  $\sigma$  was significantly lower at the intermediate site during rainy periods compared to the other two sites (Fig. 6d). There was no significant difference in  $\sigma$  at the mesic and xeric sites for all levels of rain.

5 The xeric site had significantly higher average daily  $J_{LE}$ , ranging from  $\sim 0.22$  to  $0.28 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ , versus the intermediate site with  $\sim 0.18$ – $0.25 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$  (Fig. 5a, Table S4) for all years, except 2015.  $J_{LE}$  at the xeric site was also higher than the mesic site in 2011 through 2013 and in 2015, ranging from  $0.2$  to  $0.26 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ . The mesic site had  $\sim 0.01$ – $0.06 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$  higher  $J_{LE}$  compared to the intermediate site, except in 2011.  $J_{LE}$  significantly increased with greater EVI and SWC (Fig. 6e and f).  $J_{LE}$  was significantly higher at the xeric site compared to the other sites for EVI < 0.4.  $J_{LE}$  was significantly higher at the xeric site compared to the other sites when SWC was above 19%, similar to the model of LE.  $J_{LE}$  significantly increased with VPD, and significantly decreased with rainfall (Fig. 6g and h). Unlike the model results for LE, the effects of VPD were not significantly different by site.

10 Models of H and  $J_H$  were similar, except that  $J_H$  in the mesic and xeric sites were not significantly different in 2015 (Fig. 5a, Table S4). Average daily  $J_H$  was significantly higher at the mesic site in 2011 and 2012 ( $\sim 0.2$ – $0.24 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ ) compared to

15 the intermediate ( $\sim 0.19 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ ; Fig. 5a) and xeric sites ( $\sim 0.16$ – $0.20 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ ). In 2009, 2010 and 2012, the xeric site had significantly lower  $J_H$  compared to the other sites (by  $\sim 0.02 \text{ W}_\text{m}^{-2} \text{ K}^{-1}$ ).  $J_H$  decreased only at the mesic and intermediate sites with increasing EVI (Fig. 6i) such that the intermediate site had significantly lower  $J_H$  compared to the other sites when EVI was above 0.4.  $J_H$  decreased with increased SWC at all sites, and the xeric site had significantly lower  $J_H$  compared to the other sites when SWC was above 19% (Fig. 6j). VPD significantly increased  $J_H$  at all three sites, with a greater increase at the xeric site (Fig. 6k).  $J_H$  significantly decreased at all sites with increased rainfall, where the intermediate site had significantly lower  $J_H$  compared to the mesic and xeric sites when rainfall was greater than 40 mm per day (Fig. 6l).

20 Average daily  $J_G$  was not significantly different among the years 2009–2014 and 2016 at the mesic site, but significantly increased during 2015 (Fig. 5a, Table S4), similar to the model results for G. Similarly,  $J_G$  was significantly lower at the intermediate site during 2016 (negative).  $J_G$  at the xeric site was not significantly different by year. Average daily  $J_G$  was

25 positive during summer months, and negative during winter months at all sites (Fig. 5b). Average daily  $J_G$  significantly decreased from positive to negative at the mesic and intermediate sites with an increase in EVI, with no significant change at the xeric site (Fig. 6m), similar to the model of G.  $J_G$  was significantly more negative at the intermediate site compared to the other sites for  $\text{EVI} > 0.4$ . Average  $J_G$  only significantly decreased at the intermediate and xeric sites (to negative), such that  $J_G$  was significantly more negative at the two sites when SWC was above 18% (Fig. 6n).  $J_G$  significantly increased with greater VPD, independent of site (Fig. 6o). Similar to the model of G, daily rainfall did not significantly alter the magnitude of  $J_G$  at the sites. However, the intermediate had significantly more negative  $J_G$  compared to the other two sites when daily rainfall increased (Fig. 6p).

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### **3.5 Metabolic energy and entropy**

Metabolic energy was consistently more negative at the mesic site, compared to the other sites for all years in this study (Fig 7a, Table S5). The intermediate and xeric sites exported metabolic energy from 2009 through 2011, which was greater at the intermediate site for 2010. NEE<sub>e</sub> significantly increased to more negative at all sites during May and significantly decreased

5 during August through October, which resulted in positive NEE<sub>e</sub> at the intermediate site (Fig. 7b). NEE<sub>e</sub> significantly decreased at all sites with an increase in EVI, which was greater at the xeric site (Fig. 7c). An increase in SWC resulted in decreasing NEE<sub>e</sub>, independent of site (Fig. 7d). An increase in VPD significantly decreased NEE<sub>e</sub> to more negative at all sites, with a greater decrease at the intermediate site (Fig. 7e). Increases in rainfall significantly increased NEE<sub>e</sub> to positive at all sites, where the intermediate site had a greater increase compared to the other sites (Fig. 7f).

10 In contrast to the model of NEE<sub>e</sub>, results of the model of J<sub>NEE</sub> indicated that the mesic site had significantly more negative J<sub>NEE</sub> compared to the other sites during most years; but during 2011 J<sub>NEE</sub> was significantly less negative at the site compared to the intermediate and xeric sites (Fig. 7g). The intermediate site had consistently more negative J<sub>NEE</sub> compared to the xeric site, except for 2014 where J<sub>NEE</sub> significantly decreased at the site. J<sub>NEE</sub> was more negative during summer months at all sites with no significant differences between the mesic and xeric sites (Fig. 7h, Table S5). Values of J<sub>NEE</sub> significantly decreased with an

15 increase in EVI, independent of site (Fig. 7i), different from the model of NEE<sub>e</sub>. SWC significantly decreased J<sub>NEE</sub> at all sites, with a greater slope at the mesic site (Fig. 7j). Higher VPD significantly increased J<sub>NEE</sub> similar to the model of NEE<sub>e</sub>; however slopes were more similar among the sites (Fig. 7k). Rainfall significantly decreased J<sub>NEE</sub> to less negative with a greater slope at the intermediate site, similar to the model of NEE<sub>e</sub> (Fig. 7l).

### **3.6 Entropy models**

20 From 2011 through 2016, eff<sub>rad</sub> was significantly higher at the mesic site (0.89-0.93), compared to the intermediate (0.88-0.91) and xeric (0.88-0.92) sites, which were not significantly different (Fig. 8a). Average eff<sub>rad</sub> did not significantly change with EVI, but significantly increased with higher SWC (Fig. 8c), independent of site. Higher VPD significantly decreased values of eff<sub>rad</sub> at all sites (Fig. 8d). The mesic site had significantly higher values of eff<sub>rad</sub> compared to the other two sites for all levels of VPD (Fig. 8d). Rainfall significantly increased values of eff<sub>rad</sub> at all sites, with a greater increase at the intermediate

25 site (Fig. 8e, Table S6). Daily average eff<sub>flux</sub> was significantly greater at the mesic site for most of the measurement period (Fig. 9a, Table S6). eff<sub>flux</sub> was significantly higher at the xeric site compared to the intermediate site for the years 2009, 2011, and 2013 through 2015. For 2012 and 2016 the intermediate site had significantly greater eff<sub>flux</sub> compared to the xeric site. Greater EVI only significantly increased eff<sub>flux</sub> at the mesic site, which had higher eff<sub>flux</sub> compared to the other sites for all levels of EVI (Fig. 30 9c). The intermediate site had significantly lower eff<sub>flux</sub> compared to the xeric site when EVI was above 0.3. An increase in SWC significantly decreased values of eff<sub>flux</sub> only at the intermediate and xeric sites, with a greater decrease at the xeric site (Fig. 9d). Higher VPD significantly decreased eff<sub>flux</sub> at all sites, with a greater decrease at the intermediate site (Fig. 9e). Rainfall significantly increased eff<sub>flux</sub> at all sites, where the intermediate site showed the highest increase (Fig. 9f).

There was no significant difference in  $dS/dt$  among sites for all years and months, except in 2014, where the intermediate site had significantly higher  $dS/dt$  compared to the other sites (Fig. 10a, Table S6). In addition, the xeric site accumulated  $dS/dt$  during 2012 such that it was significantly different from the other sites. An increase in EVI and VPD resulted in a significant increase in  $dS/dt$ , independent of site (Fig. 10c and d). SWC and rainfall were not significant in the model of  $dS/dt$ . The diurnal variation in  $dS/dt$  was greater at the mesic and xeric sites during the drought years 2010, 2011 and 2012, compared to the intermediate site, specifically during nighttime (SI Fig. S4). At the intermediate site  $dS/dt$  varied more during the years 2014 and 2016, as seen by greater entropy accumulation during nighttime hours and greater export during daytime hours for the year 2014.

#### 4 Discussion

10 Here we describe differences in energy use efficiencies of sites with varying structural complexities (i.e., understory composition, basal area, DBH) using metrics of energy and entropy. Different from our expectations, environmental and structural effects on energy and entropy fluxes were not different with the exception of  $NEE_e$  and  $J_{NEE}$ . These results suggest that differences in the thermodynamic environment among sites (i.e., air and surface temperatures) did not contribute to changes in entropy export in response to environmental variables. Nevertheless, metabolic entropy ( $J_{NEE}$ ) decreased during the 15 drought at all sites (Fig. 7), whereas  $NEE_e$  showed no significant change at the intermediate and xeric sites. The different results were a function of  $T_{air}$ , which increased during the summer of 2011, especially at the intermediate and xeric sites, thus lowering the flux of  $J_{NEE}$  (Fig. 7). The decrease in  $J_{NEE}$  suggests that metabolic activity at all sites was similarly affected by low rainfall, increasing VPD, and changes in temperature, demonstrating a decrease in physiological activity of plant species during drought (Barron-Gafford et al., 2013). This decrease in metabolic efficiency supports a previous study at the mesic and 20 xeric sites, which found lower electron transport and carboxylation capacity during drought (Wright et al., 2012). Differences in the underlying reflective capacities at the sites significantly altered their entropy production and resulted in variations in entropy exchanges (Stoy et al., 2014). The more structurally complex mesic site had more negative metabolic entropy ( $J_{NEE}$ ), which translates to greater energy accumulation, in addition to greater radiation entropy and export efficiencies ( $eff_{rad}$ ,  $eff_{flux}$ ) compared to the intermediate site, which had greater land use legacy and was structurally less complex. Although 25 the radiation entropy ratio ( $eff_{rad}$ ) indicated that both the intermediate and xeric sites were equally energy efficient,  $eff_{flux}$  showed prolonged recovery of energy efficiency from drought by one year. Entropy change over time ( $dS/dt$ ) did not significantly vary at the mesic site, but was more variable at the xeric and intermediate sites following the drought. We hypothesized that the xeric site would have higher  $H$  and  $J_H$ , due to its open canopy and sandy soils and therefore lower 30 volumetric heat capacity. In contrast to our first hypothesis, the mesic and intermediate sites and not the xeric site had a more pronounced increase in  $H$  and  $J_H$  when EVI decreased during drought (Fig. 1). Lower  $H$  and  $J_H$  at the xeric site was a consequence of greater energy partitioning into LE, enabled by greater transpiration rates of plant functional types present at the site (deciduous and evergreen oaks in the understory, mid- and overstory; Klein et al., 2013; Renninger et al., 2015; Stoy et al., 2006). This result was confirmed, as  $J_H$  fluxes did not significantly change with an increase in EVI, whereas  $J_{LE}$  increased.

**Deleted:**  $J_{NEE}$  was always negative at the mesic site throughout this study, indicating more energy storage which decreased entropy production locally (Fig. 4). The mesic site had significantly lower  $J_{NEE}$  compared to the intermediate site for all years, except 2011, and 2015. The mesic site had significantly more negative  $J_{NEE}$  compared to the xeric site in 2009 and 2012, whereas the xeric site had significantly more negative  $J_{NEE}$  compared to the intermediate site for the years 2010, and 2012-2014. The mesic site had significantly more negative  $J_{NEE}$  compared to the other two sites when EVI was below 0.35 (Fig. 4). In addition, the xeric site had less positive  $J_{NEE}$  (by  $\sim 1 \text{ kJ m}^{-2} \text{ K}^{-1}$ ) compared to the intermediate site when EVI was  $<0.44$ . For  $T_{air}$  between 17-23 °C, the mesic site had significantly more negative  $J_{NEE}$  compared to the xeric site. When  $T_{air}$  was above 17 °C,  $J_{NEE}$  at the intermediate site was significantly more positive compared to the other sites (by  $> 2 \text{ kJ m}^{-2} \text{ K}^{-1}$ ; Fig. 4).  $J_{NEE}$  became more negative (from  $-1$  to  $-4 \text{ kJ m}^{-2} \text{ K}^{-1}$ ) with increasing VPD, independent of site, and was not affected by SWC or rainfall (data not shown). 

$J_O$  was significantly higher at the xeric site, compared to the intermediate site in 2012. In 2016,  $J_O$  at the intermediate site ( $-35 \text{ kJ m}^{-2} \text{ K}^{-1}$ ) was significantly more negative compared to the mesic and xeric sites ( $0-10 \text{ kJ m}^{-2} \text{ K}^{-1}$ ; Fig. 5). In contrast,  $J_O$  was significantly more positive at the xeric site, compared to other sites when SWC was  $>22 \%$ .  $J_O$  increased from negative to positive with an increase in  $T_{soil}$ , which was not significantly different by site (Fig. 5).  $J_O$  was not affected by VPD, rainfall or EVI. 

In 2010, 2012 and 2013,  $\sigma$  was significantly higher at the mesic site compared to the xeric site (by  $> 50 \text{ kJ m}^{-2} \text{ K}^{-1}$ ; Fig. 5). In 2009, 2011 and 2012  $\sigma$  was significantly higher at the intermediate site (by  $> 50 \text{ kJ m}^{-2} \text{ K}^{-1}$ ) compared to the xeric, but not the mesic site (Fig. 5). Entropy production was unaffected by SWC,  $T_{air}$ , and EVI, but significantly increased with higher VPD (from  $<1400$  to  $1800 \text{ kJ m}^{-2} \text{ K}^{-1}$ ; Fig. 5). 

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... [21]

suggesting that evapotranspiration and the cooling of leaf and soil surfaces had greater influence on the partitioning of available energy. In contrast,  $J_H$  increased more at the mesic and xeric sites with increasing VPD, suggesting that drier air increased the sensible heat flux from the surface to the atmosphere (Massmann et al., 2018). Similarly, as VPD increased so did  $\sigma$  at all sites. This response was also observed in Kuricheva et al. (2017), where drier summers resulted in greater entropy production, likely because an increase in VPD correlated with greater absorption of solar radiation and partitioning to  $H$  (Fig. 3a). Even though plant abundance was lower at the xeric site, its species composition was better adapted to drought conditions, which allowed for higher  $J_{LE}$  compared to the other sites (Roman et al., 2015). Furthermore, an increase in EVI during summer months at the xeric site increased  $J_{LE}$ , demonstrating that greater leaf area enhanced ecosystem function (Peng et al., 2017; Zhu et al., 2016). Interestingly,  $J_{LE}$  did not vary significantly by site with changes in VPD, which supported the findings of Whelan et al. (2013) that all sites had similar stomatal regulation to increases in VPD. Overall, the xeric site had higher  $J_{LE}$  compared to the other sites for  $EVI < 0.5$ , even though the site basal area was almost half that of the mesic and intermediate sites (Table 1). An overstory composed of more oak species at the xeric site (~20 %) along with the  $C_4$  understory resulted in higher transpiration during spring and summer, compared to stands containing just pine trees (Klein et al., 2013; Renninger et al., 2015; Stoy et al., 2006). Additionally,  $C_4$  grasses and oak species at the xeric site were better adapted to drought (i.e., anisohydric response; Osborne and Sack, 2012; Roman et al., 2015), which may enable higher entropy production and lower variability in the structural integrity (i.e., lower decreases in EVI; Fig. 1e). This suggests that the understory plays a crucial role in the structure and function of more open canopy ecosystems (Aoki, 2012; Lin, 2015), in addition to more productive overstory trees during summer. This led to similar entropy export efficiencies at all sites as evidenced by all sites having comparable  $dS/dt$ . Nevertheless, as  $\sigma$  increased with greater absorption of radiation due to an increase in EVI,  $J_H$  decreased as a result of higher SWC, resulting in temporary entropy accumulation at the xeric site during the end of 2012, (SI Fig. 4), which may have contributed to higher  $T_{air}$  compared to the other sites (Fig. 2). In contrast, the mesic site was affected by the interaction of biological and radiative forces, as  $J_{LE}$ ,  $J_{NEE}$  and  $eff_{rad}$  decreased more severely with decreasing plant leaf area compared to the xeric site (lower EVI; Fig. 1e). As a consequence of lower LE and  $J_{LE}$  during the drought, more energy was partitioned into  $H$  in 2011 (Fig. 6), as air, soil and surface temperatures increased due to lower leaf area (Figs. 1 and 2), indicating a shift of ecosystem function (Ban-Weiss et al., 2011) towards lower quality energy degradation (Kuricheva et al., 2017). This initially depleted soil moisture storage at the mesic site (Fig. 1) and further decreased LE and  $J_{LE}$  (Kim and Wang, 2012; Lauri et al., 2014). Nevertheless, the shift in energy partitioning at the mesic site allowed for the maintenance of  $dS/dt$  during drought, by export of entropy which had accumulated during nighttime hours (SI Fig. S4), demonstrating an adaptation of the site to changes in resource availability (Basu et al., 2016; Brodrribb et al., 2014). In contrast, the xeric and intermediate sites showed greater variability in annual  $dS/dt$  following the drought when rainfall returned to pre-drought levels and SWC increased (Fig. 10a). Nevertheless, the rapid increase in  $J_{LE}$  in 2012 at the mesic and xeric sites indicated an increase in ecosystem function through greater evapotranspiration. This provides evidence of recovery following the drought, because  $J_{LE}$  is of higher quality entropy dissipation (Kuricheva et al., 2017), coupling both mass and heat dynamics (Brunsell et al., 2011), whereas  $J_H$  is a function of the thermal gradient (Kleidon, 2010; LeMone et al., 2007).

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**Deleted:** In contrast, plant species at the mesic were adapted to higher soil water conditions, such that  $eff_s$  decreased, as the quality of entropy dissipation through other fluxes leads to lower  $J_{LE}$  (Kuricheva et al., 2017). This is because  $J_{LE}$  couples both mass and heat dynamics (Brunsell et al., 2011), whereas  $J_H$  is only a function of the thermal gradient (Kleidon, 2010; LeMone et al., 2007). Nevertheless,  $J_H$  increased at all sites during the first year of drought and then remained higher compared to pre-drought conditions, indicating a shift of ecological

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In generally, plant species at the mesic site were better adapted to higher soil water conditions, as entropy and energy fluxes did not change as drastically with increasing SWC compared to the other sites. This recovery of EVI following drought also allowed for greater  $eff_{rad}$  at the sites. But  $eff_{rad}$  was higher at the mesic site despite lower EVI compared to the intermediate site. This finding supports our second hypothesis, that sites with greater plant functional diversity maintain greater radiative entropy efficiency, the mesic site efficiently used available energy from incoming solar radiation (Fig. 2) through lower reflection of  $R_s$  and by emitting less longwave radiation (Lin, 2015).  $Eff_{rad}$  decreased during the initial drought year because all sites reflected more  $R_s$ , likely a consequence of a change in EVI, as well as leaf angle from a decrease in SWC and altered plant hydraulics. Higher  $eff_{rad}$  and  $eff_{flux}$  at the mesic site are consistent with enhanced function due to greater plant diversity in the understory (Fig. 4a). For example, wiregrass, a C<sub>4</sub> species, can maintain photosynthetic rates under high temperatures (Osborne and Sack, 2012; Ward et al., 1999), which allows for greater energy storage during unfavorable environmental conditions (Brunsell et al., 2011). Despite higher wiregrass biomass in the understory, the xeric site was less efficient in using available radiation energy, indicated by high  $R_{s,out}$  and  $R_{l,out}$  (Brunsell et al., 2011). Structural limitations of the canopy (i.e., lower basal area) impeded the efficient absorption of available radiation, therefore lowering  $eff_{rad}$  (Norris et al., 2011). Furthermore, larger proportions of deciduous oak trees at the xeric site (Table 1), which typically shed their leaves during the winter, lowered the capacity of the system to acquire radiation (Baldocchi et al., 2004; Fig. 8b). This inefficiency was also confirmed by model results for  $J_{NEE}$ , which, in contrast to  $NEE_e$  revealed lower metabolic efficiency relative to the intermediate site, reflecting differences in basal area and site EVI. These results demonstrate that the mesic site was better adapted to changes in resource availability by way of altering its reflective properties, where energy partitioning adjusted to maintain steady entropy exports relative to incoming entropy (Gunawardena et al., 2017; Otto et al., 2014; Taha et al., 1988). Nevertheless, metabolic activity at all sites was a source for entropy ( $J_{NEE} > 1$ ) during rainy periods, demonstrating an inefficiency in maintaining optimal function when environmental pressure was imposed on the system. greater metabolic efficiency at the mesic site resulted in more rapid increases in the structural complexity as indicated by a decrease in  $R_{s,out}$  following the drought when compared to the intermediate site (Brunsell et al., 2011; Holdaway et al., 2010). Metabolic activity (in energy terms) at the intermediate site was largely dependent on EVI (i.e., leaf area), demonstrating lower biological control of individual plant species (i.e., stomatal control; Urban et al. 2016), but a strong influence of total leaf area on metabolic function and the export of entropy (Brunsell et al., 2011; Fig. 4 and 6). This was further illustrated at the intermediate site through less negative metabolic energy ( $NEE_e$ ) when EVI was  $\sim 0.25$  (Fig. 7c). Even though EVI in 2012 was greater at the intermediate site this did not correspond to higher  $J_{LE}$  (Fig. 5a), which was also shown by a lack of significant change in entropy exports with changes in EVI ( $eff_{flux}$ , Fig. 9c). The result of lower metabolic function at the intermediate site is intriguing as the mesic and intermediate sites were structurally similar, based on similar BA, mean DBH and overstory tree composition (Table 1). The inefficiency appears to be a consequence of anthropogenic modification, which homogenized the ecosystem, leading to a decrease in plant functional types (Table. 1; Fig. 3), thereby reducing values of  $eff_{rad}$  and  $eff_{flux}$ . This result provides evidence that the intermediate site was less efficient in absorbing energy and dissipating entropy compared to the mesic site, resulting in slower adaptation to drought. Similar results were shown in Lin et al. (2015), where disturbed sites had

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predominantly lower entropy production rates, as well as in Lin et al. (2018) where greater surface temperature led to decreased  $\sigma$ , which we also observed at the intermediate site. Our third hypothesis was therefore supported, as the intermediate site had lower  $eff_{flux}$  relative to the mesic and xeric sites. Lower plant functional diversity, specifically the lack of wiregrass, due to soil perturbations that took place prior to stand establishment (>95 years ago), likely lowered metabolic function, which in turn 5 affected entropy exports at the intermediate site and its recovery from drought. For example, a negative  $J_G$  at the intermediate site was observed with increasing SWC suggesting poor soil water drainage, which is also likely a consequence of agricultural legacy (Kozlowski, 1999). A prolonged increase in  $eff_{flux}$  compared to the other sites showed that the intermediate site did not adapt its entropy exports, in addition to greater reflection of  $R_s$  during drought recovery. This result indicates that differences 10 in soil conditions and lower plant functional diversity at the intermediate site reduced entropy exports compared to the other sites (Meysman and Bruers, 2010), such that plant functional types present at the site could not rescue the ecosystem's function during disturbance (Elmqvist et al., 2003). Furthermore, while the intermediate site showed no change in  $dS/dt$  during the drought, following the drought the export of entropy significantly increased, resulting in more unstable conditions (Fig. 10a). The increase in entropy export corresponded to high annual rainfall and soil moisture conditions (Figs. 1 and S1), once more 15 suggesting that soil characteristics were altered due to its agricultural legacy. The lower ability to adapt to changes in resource availability at the intermediate site could induce its degradation if environmental fluctuations become more frequent and severe with climate change (Mori, 2011; Siteur et al., 2016). This could further exacerbate instabilities for nearby sites, as changes in the reflective properties of degraded sites can alter microclimate and weather patterns across whole ecosystems (Norris et al., 2011).

We conclude that the analysis of entropy dynamics, in relation to structural and environmental variables gives valuable insights 20 into the functional complexity of ecosystems and their ability to adapt to drought. A combination of entropy fluxes and entropy ratios revealed how differences in structural and/or functional characteristics affect energy efficiencies in longleaf pine ecosystems. Our results show that all sites demonstrated adaptive capacity to extreme drought, as indicated by a lack of significant change in  $dS/dt$ , except for greater variations at the xeric and intermediate sites following the drought. We show 25 that overall low entropy exports at the site with greater land use legacy had the potential to decrease ecosystem function (Meysman and Bruers, 2010), especially during high rainfall events. Changes in climate and natural and human induced disturbances are becoming more frequent and severe (IPCC, 2014), demanding more predictive power about how changes in ecosystem structure and function will alter resilience to disturbances. Future policy, conservation or restoration applications depend on reliable measures, such as the metrics presented here, to monitor ecosystem function following disturbances 30 (Haddeland et al., 2014; Porter et al., 2012; Reinmann and Hutyra, 2016; Thorn et al., 2017). This is especially critical for anthropogenically modified systems, as their land use history can affect changes in energy use efficiency and thus alter their ability to recover from disturbances (Bürgi et al., 2016; Foster et al., 2003). The application of entropy metrics could improve our understanding of the interaction of structure, function and legacy on energy use efficiency across a variety of global ecosystems.

#### Author contribution

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G.S. and L.B. designed and acquired funding for the research. S.W. and C.S. analyzed the data. P.S. aided S.W. with the theories of ~~entropy~~ and energy density. All authors contributed to writing of the manuscript.

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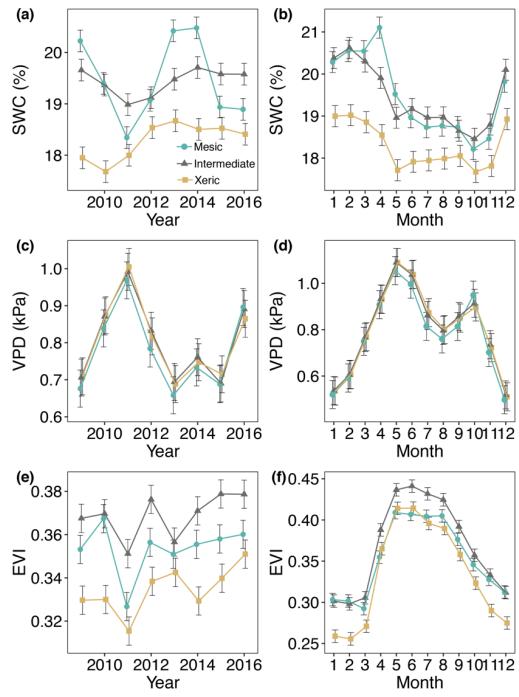
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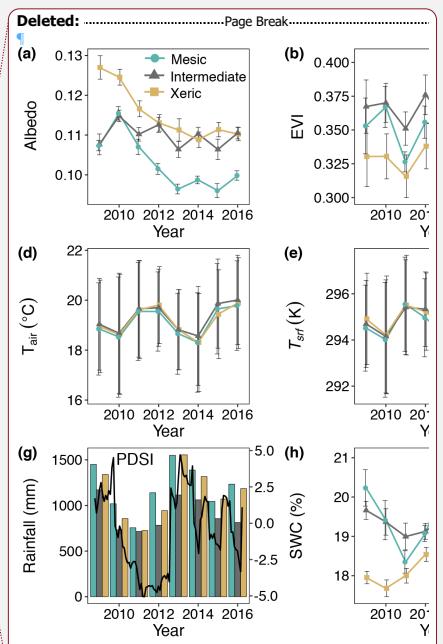
Characteristic	Mesic	Intermediate	Xeric
Mean DBH (cm)	25.9	42.5	22.5
B <sub>A</sub> <i>P. palustris</i> (m <sup>2</sup> ha <sup>-1</sup> )	17.7	14.6	8.9
B <sub>A</sub> all tree spp. (m <sup>2</sup> ha <sup>-1</sup> )	19.0	15.7	11.0
Proportion of oak overstory trees (%)	6.8	7.0	19.1
<u>LAI (m<sup>-2</sup> m<sup>-2</sup>)</u>	<u>1.0<sup>a</sup></u>	<u>unknown</u>	<u>0.69<sup>a</sup></u>
Wiregrass in the understory (%)	28	5	24
Woody species in the understory (%)	12	15	10
Prescribed fire	<u>Early spring of 2009, 2011, 2013, 2015</u>	<u>Early spring of 2009, 2011, 2013, 2015</u>	<u>Early spring of 2009, 2011, 2013, 2015</u>

<sup>a</sup> Wright et al. 2012



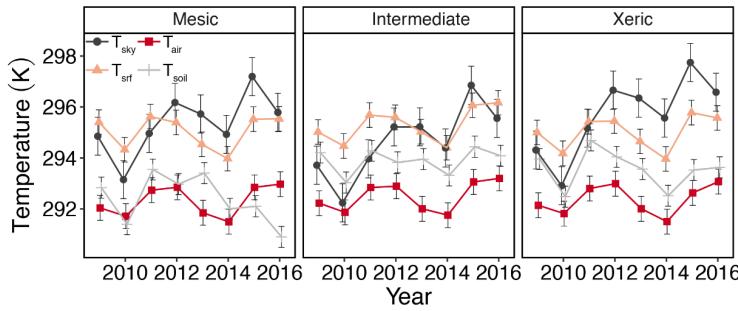
**Figure 1:** Least square mean predicted values from the mixed model results for Environmental and structural variables for the years 2009-2016 at the mesic, intermediate and xeric sites, with average annual (a, c, and e) and monthly (b, d, and f) means of (a and b) soil water content (SWC), (c and d) vapor pressure deficit (VPD), and (e and f) enhanced vegetation index (EVI). Error bars represent standard errors (SE).

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**Figure 1:**

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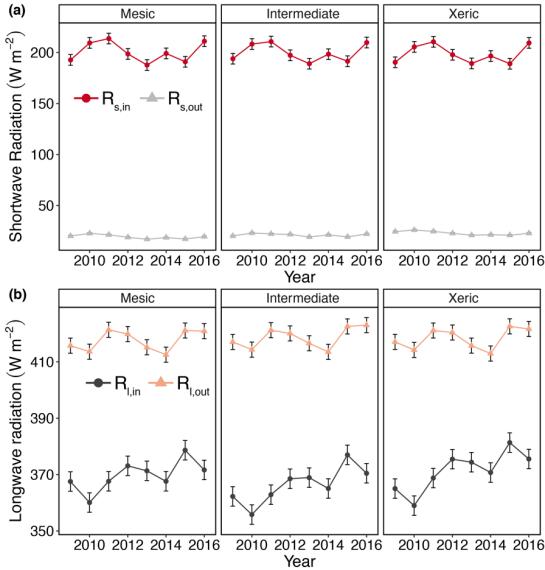


**Figure 2:** Least square mean predicted values from the mixed model results for annual sky temperature ( $T_{sky}$ ), air temperature ( $T_{air}$ ), surface temperature ( $T_{srf}$ ), and soil temperature ( $T_{soil}$ ) at the mesic, intermediate and xeric sites. Error bars represent SE.

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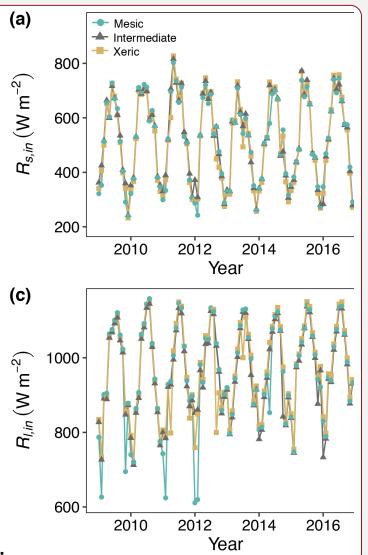
**Deleted:** (d) air temperature ( $T_{air}$ ), (e) sky temperature ( $T_{sky}$ ), (f)

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**Figure 3:** Least square mean predicted values from the mixed model results of annual average radiation at the mesic, intermediate and xeric sites for the years 2009-2016. (a) annual incoming and outgoing shortwave radiation ( $R_{s,in}$  and  $R_{s,out}$ ), and (b) annual incoming and outgoing longwave radiation ( $R_{l,in}$  and  $R_{l,out}$ ). Error bars represent SE.

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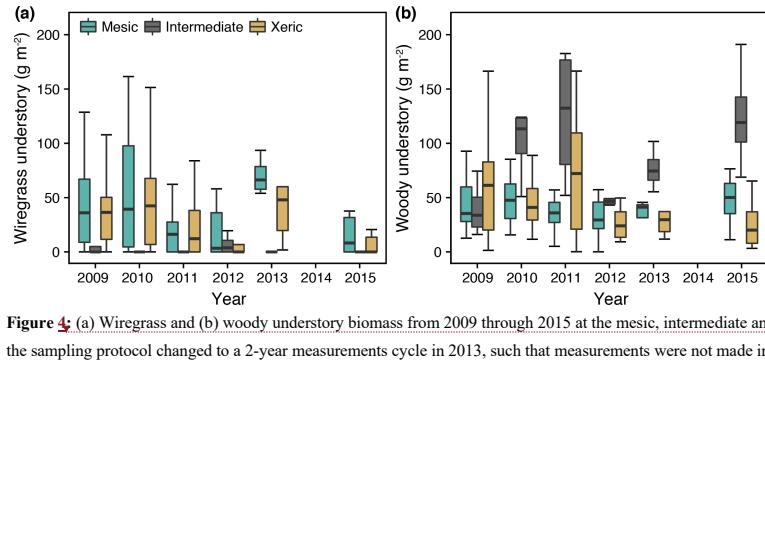
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**Deleted:** Figure 3: (a) Wiregrass and (b) woody understory biomass from 2009 through 2015 at the mesic, intermediate and xeric sites. Note that the sampling protocol changed to a 2-year measurements cycle in 2013, such that measurements were not made in 2014 and 2016.



**Figure 4.** (a) Wiregrass and (b) woody understory biomass from 2009 through 2015 at the mesic, intermediate and xeric sites. Note that the sampling protocol changed to a 2-year measurements cycle in 2013, such that measurements were not made in 2014 and 2016.

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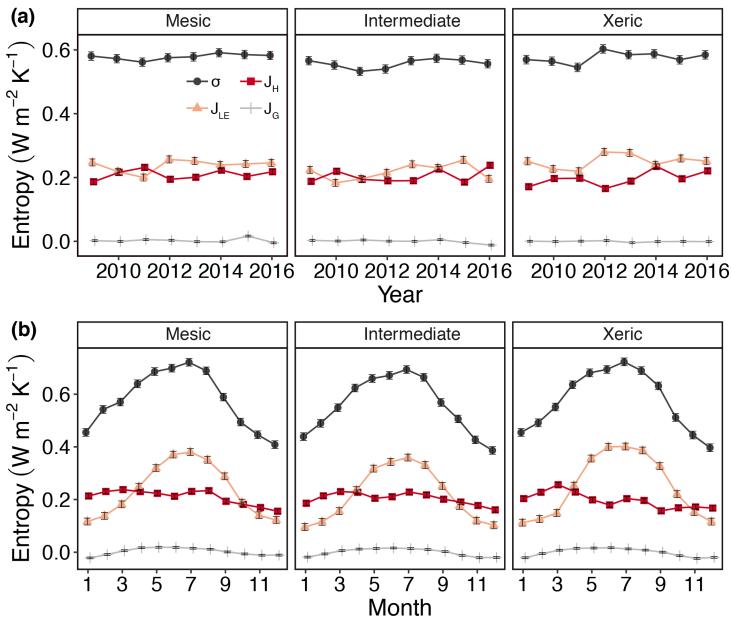
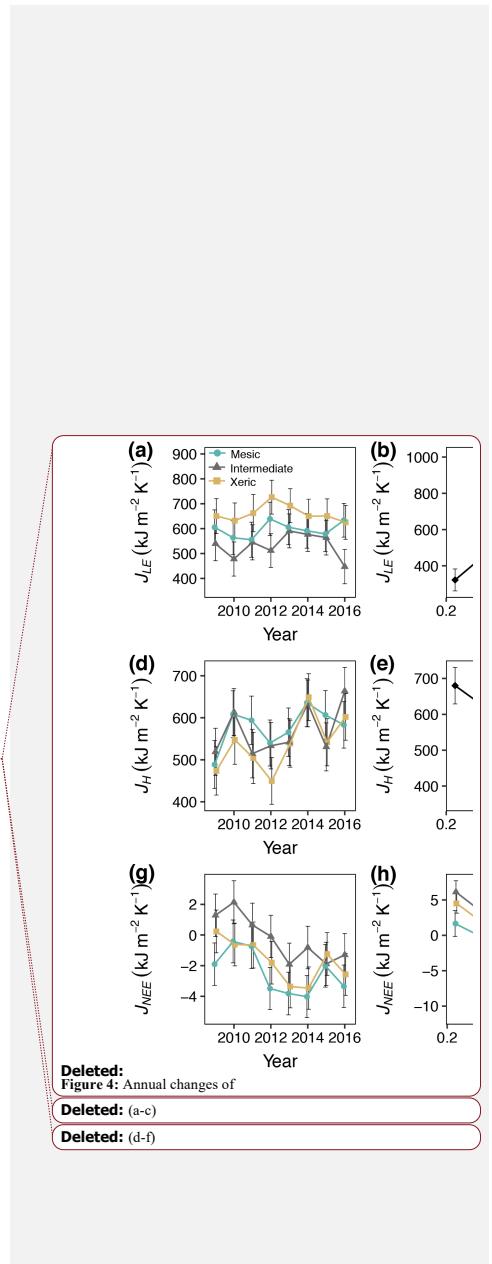
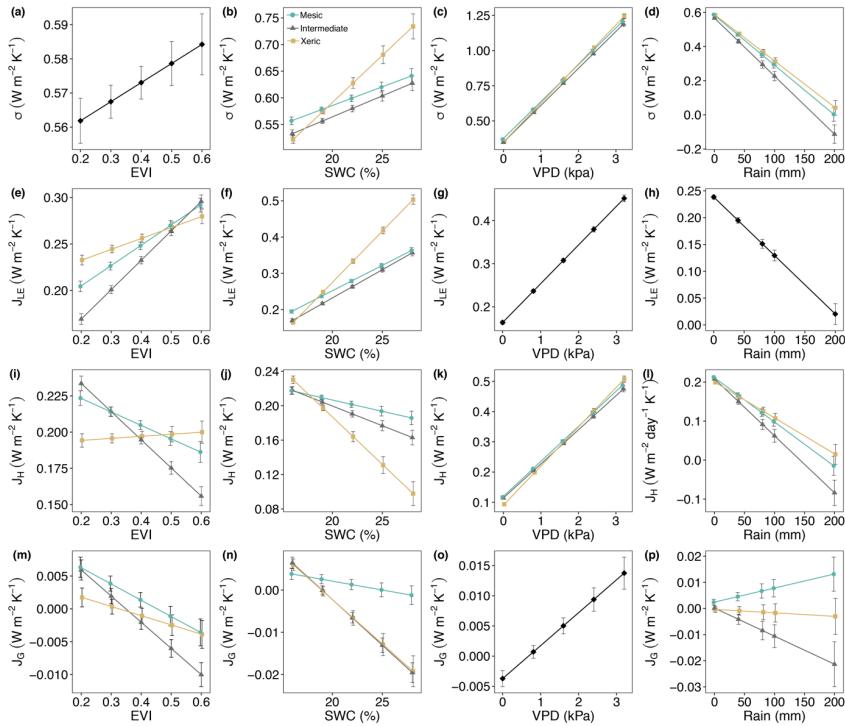


Figure 5: Least square mean predicted values from the mixed model results of annual (a) and monthly (b) average entropy production ( $\sigma$ ) and entropy fluxes of latent energy ( $J_{LE}$ ), sensible heat ( $J_H$ ), and ground heat ( $J_G$ ) at the mesic, intermediate and xeric sites. Error bars represent SE.

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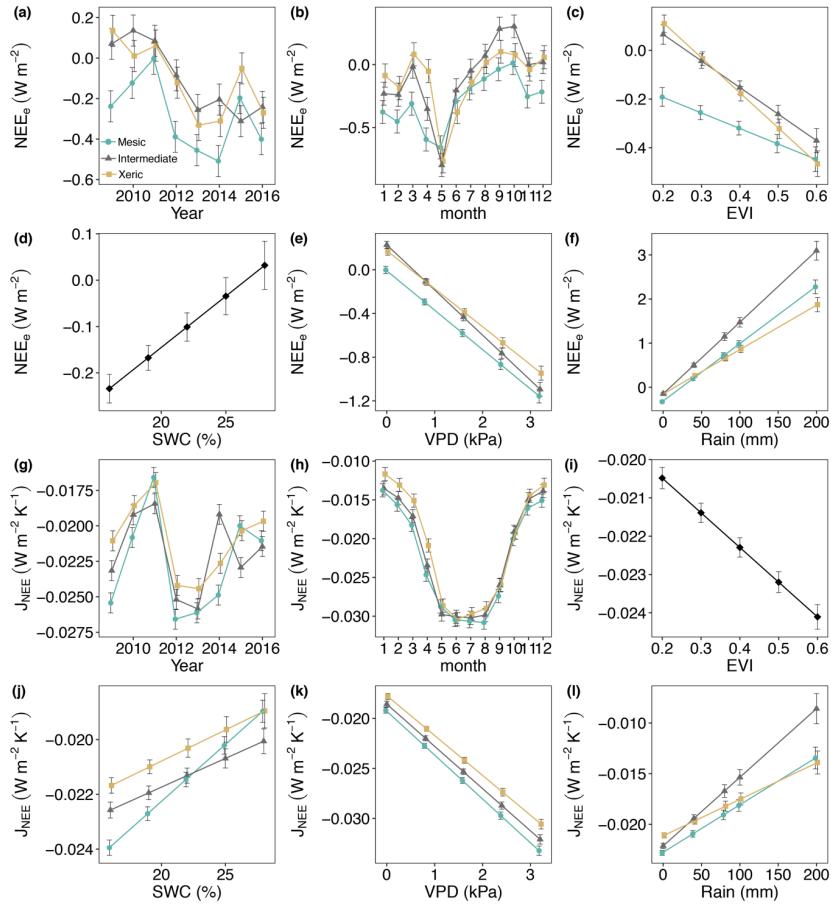




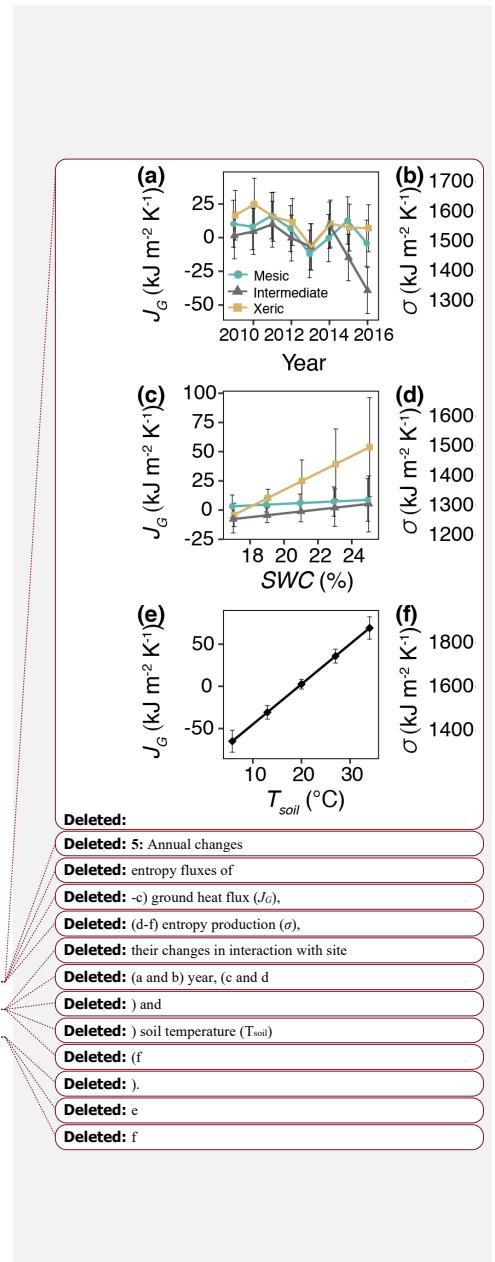
**Figure 6:** Least square mean predicted values from the mixed model results for interactive effects of (a, e, i, m) Enhanced Vegetation Index (EVI), (b, f, j, n) soil water content (SWC), (c, g, k, o) vapor pressure deficit (VPD), and (d, h, l, p) rain on (a-d) entropy production ( $\sigma$ ) and entropy fluxes of (e-h) latent energy ( $J_{LE}$ ), (i-l) sensible heat ( $J_H$ ), and (m-p) ground heat ( $J_G$ ). For (g), (h) and (o) the interaction

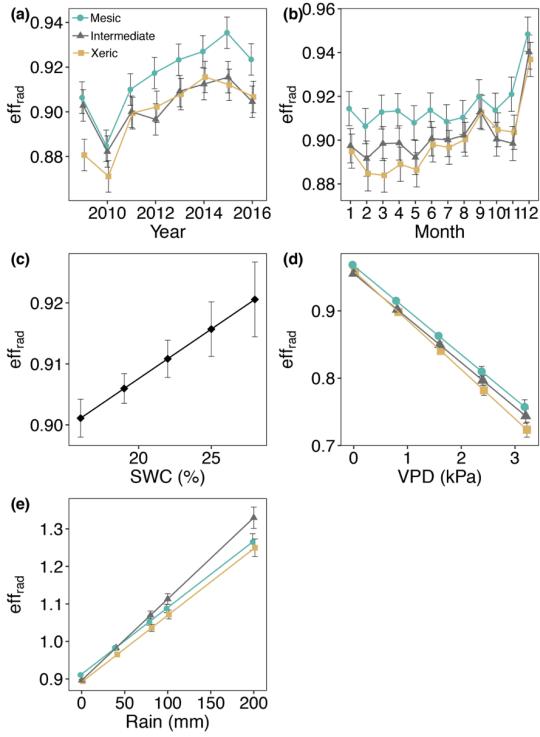
5 with site was not significant, as signified by a single black line. Error bars represent SE.

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- Deleted:** ( ) metabolic entropy production ( $J_{NEE}$ ), and their changes in interaction with site and (a, d, g) year, (b, e, h)
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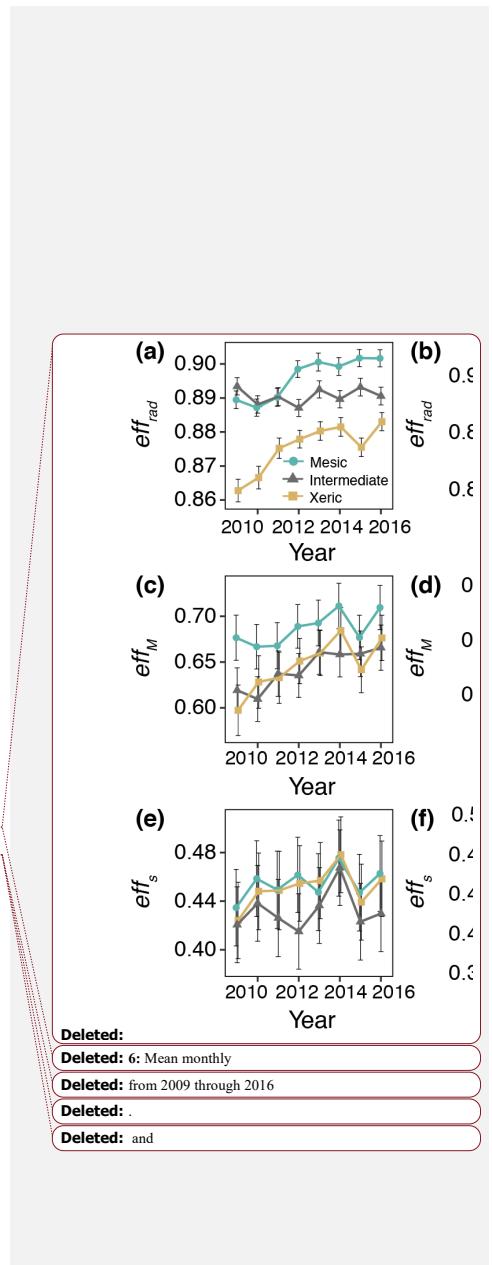
**Figure 7:** Least square mean predictive values from the mixed model results for interactive effects of site with (a and g) year, (b, c, h, and i) enhanced vegetation index (EVI), (d and j) soil water content (SWC), (e and k) vapor pressure deficit (VPD), and (f and l) rain on (a-f) the metabolic energy flux (NEE<sub>e</sub>) and (g and l) metabolic entropy fluxes of (J<sub>NEE</sub>). For (d) and (g) the interaction with site was not significant, as indicated by a single solid black line. Error bars represent SE.

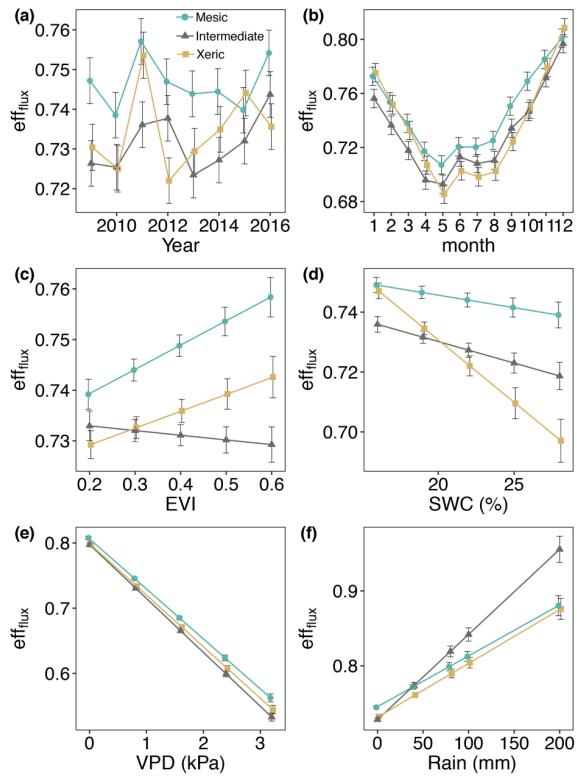




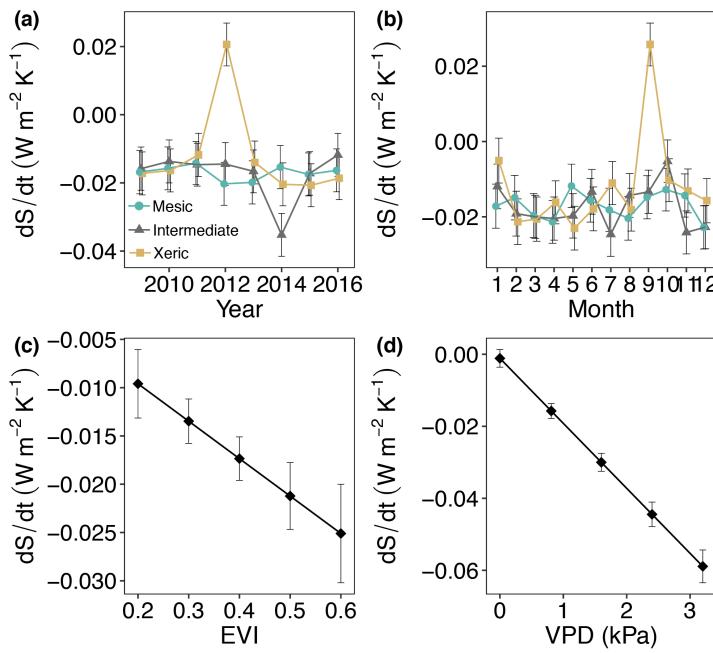
**Figure 8:** Least square mean predicted values from the mixed model results for average daily half-hourly radiative entropy efficiencies ( $eff_{rad}$ ) at the mesic, intermediate and xeric sites for the interactions of site with (a) year, (b) month, (c) soil water content (SWC), for which the interaction with site was not significant, (d) vapor pressure deficit (VPD), and (e) rain. The enhanced vegetation index was not significant in the model. Error bars represent SE.

5





**Figure 9:** Least square mean predicted values from the mixed model results for average daily half-hourly flux entropy efficiencies ( $\text{eff}_{\text{flux}}$ ) at the mesic, intermediate and xeric sites for the interactions of site with (a) year, (b) month, (c) enhanced vegetation index (EVI), (d) soil water content (SWC), (e) vapor pressure deficit (VPD), and (f) rain. Error bars represent SE.



**Figure 10:** Least square mean predicted values from the mixed model results for average daily entropy for the mesic, intermediate and xeric sites by (a) year (a) month, (c) enhanced vegetation index (EVI), and (d) vapor pressure deficit (VPD). Soil water content and rain, as well as the interactions with site were not significant in the model. Error bars represent SE.

*Supplementary Information for bg-2018-322*

“Quantifying energy use efficiency via entropy production: A case study from longleaf pine ecosystems”

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**Table S1:** Type 3 effects for model of rain

<u>Effect</u>	<u>Sum Sq.</u>	<u>Df</u>	<u>F value</u>	<u>Pr(&gt;F)</u>
Site	<u>30607</u>	<u>2</u>	<u>4.2059</u>	<u>0.0159</u>
Year	<u>95938</u>	<u>7</u>	<u>3.7667</u>	<u>&lt; 0.001</u>

**Table S2:** Type 3 effects for the models of environmental variables and radiation

Model	Effect	Chisq	Df	p-value
<b>SWC</b>	Site	6561.692	2	< 0.001
	Year	23.764	7	0.0013
	Month	94.089	11	< 0.001
	Site:Year	2629.617	14	< 0.001
	Site:Month	1398.986	22	< 0.001
<b>VPD</b>	Site	245.268	2	< 0.001
	Year	33.981	7	< 0.001
	Month	100.044	11	< 0.001
	Site:Year	214.101	14	< 0.001
	Site:Month	232.327	22	< 0.001
<b>EVI</b>	Site	2510.727	2	< 0.001
	Year	15.868	7	0.0264
	Month	597.701	11	< 0.001
	Site:Year	294.805	14	< 0.001
	Site:Month	791.727	22	< 0.001
<b>T<sub>sky</sub></b>	Site	2202.369	2	< 0.001
	Year	23.089	7	0.0017
	Month	912.141	11	< 0.001
	Site:Year	440.318	14	< 0.001
	Site:Month	63.082	22	< 0.001
<b>T<sub>srf</sub></b>	Site	438.625	2	< 0.001
	Year	12.844	7	0.076
	Month	1423.846	11	< 0.001
	Site:Year	435.639	14	< 0.001
	Site:Month	778.064	22	< 0.001
<b>T<sub>air</sub></b>	Site	1419.775	2	< 0.001
	Year	9.954	7	0.1912
	Month	1231.11	11	< 0.001
	Site:Year	1311.82	14	< 0.001
	Site:Month	336.866	22	< 0.001
<b>T<sub>soil</sub></b>	Site	5110.24	2	< 0.001
	Year	16.817	7	0.0186
	Month	1901.818	11	< 0.001
	Site:Year	1922.717	14	< 0.001
	Site:Month	5270.008	22	< 0.001
<b>R<sub>s,in</sub></b>	Site	0.9664	2	0.6168
	Year	16.3199	7	0.0224
	Month	763.0665	11	< 0.001
	Site:Year	121.9389	14	< 0.001
	Site:Month	170.75	22	< 0.001
<b>R<sub>s,out</sub></b>	Site	4161.151	2	< 0.001
	Year	48.782	7	< 0.001
	Month	682.874	11	< 0.001
	Site:Year	816.733	14	< 0.001
	Site:Month	1780.397	22	< 0.001
<b>R<sub>lin</sub></b>	Site	2479.339	2	< 0.001
	Year	22.578	7	0.0020
	Month	1005.462	11	< 0.001
	Site:Year	482.99	14	< 0.001
	Site:Month	72.965	22	< 0.001
<b>R<sub>lout</sub></b>	Site	226.43	2	< 0.001
	Year	13.07	7	0.0704
	Month	1433.87	11	< 0.001
	Site:Year	137.39	14	< 0.001
	Site:Month	980.18	22	< 0.001

**Table S3:** Type 3 effects for the models of energy

Model	Effect	Chisq.	Df	Pr(>Chisq)
<b>R<sub>n</sub></b>	Year	<u>20.6658</u>	<u>7</u>	<u>0.0042975</u>
	Month	<u>1927.222</u>	<u>11</u>	<u>&lt; 0.001</u>
	SWC	<u>58.6889</u>	<u>1</u>	<u>&lt; 0.001</u>
	Site	<u>650.5143</u>	<u>2</u>	<u>&lt; 0.001</u>
	EVI	<u>12.2151</u>	<u>1</u>	<u>0.0005</u>
	Rain	<u>140.9816</u>	<u>1</u>	<u>&lt; 0.001</u>
	VPD	<u>1756.8922</u>	<u>1</u>	<u>&lt; 0.001</u>
	Month:Site	<u>120.9114</u>	<u>22</u>	<u>&lt; 0.001</u>
	SWC:Site	<u>24.2945</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:EVI	<u>7.3321</u>	<u>2</u>	<u>0.0256</u>
<b>LE</b>	Site:VPD	<u>16.6743</u>	<u>2</u>	<u>0.0002</u>
	Year:Site	<u>263.8642</u>	<u>14</u>	<u>&lt; 0.001</u>
	Year	<u>20.7768</u>	<u>7</u>	<u>0.0041</u>
	Month	<u>754.2793</u>	<u>11</u>	<u>&lt; 0.001</u>
	SWC	<u>455.4372</u>	<u>1</u>	<u>&lt; 0.001</u>
	Site	<u>476.4295</u>	<u>2</u>	<u>&lt; 0.001</u>
	EVI	<u>149.9341</u>	<u>1</u>	<u>&lt; 0.001</u>
	Rain	<u>116.5615</u>	<u>1</u>	<u>&lt; 0.001</u>
	VPD	<u>1043.0314</u>	<u>1</u>	<u>&lt; 0.001</u>
	Month:Site	<u>369.8495</u>	<u>22</u>	<u>&lt; 0.001</u>
<b>H</b>	SWC:Site	<u>130.9093</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:EVI	<u>43.0759</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:VPD	<u>5.3897</u>	<u>2</u>	<u>0.0676</u>
	Year:Site	<u>564.6937</u>	<u>14</u>	<u>&lt; 0.001</u>
	Year	<u>39.525</u>	<u>7</u>	<u>&lt; 0.001</u>
	Month	<u>108.742</u>	<u>11</u>	<u>&lt; 0.001</u>
	SWC	<u>29.086</u>	<u>1</u>	<u>&lt; 0.001</u>
	Site	<u>90.131</u>	<u>2</u>	<u>&lt; 0.001</u>
	EVI	<u>25.974</u>	<u>1</u>	<u>&lt; 0.001</u>
	Rain	<u>95.918</u>	<u>1</u>	<u>&lt; 0.001</u>
<b>G</b>	VPD	<u>1320.893</u>	<u>1</u>	<u>&lt; 0.001</u>
	Month:Site	<u>301.757</u>	<u>22</u>	<u>&lt; 0.001</u>
	SWC:Site	<u>35.234</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:EVI	<u>41.862</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:VPD	<u>29.24</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:Rain	<u>16.416</u>	<u>2</u>	<u>0.0003</u>
	Year:Site	<u>351.685</u>	<u>14</u>	<u>&lt; 0.001</u>
	Year	<u>9.1742</u>	<u>7</u>	<u>0.2404</u>
	Month	<u>180.4785</u>	<u>11</u>	<u>&lt; 0.001</u>
	SWC	<u>37.8658</u>	<u>1</u>	<u>&lt; 0.001</u>
	Site	<u>200.7208</u>	<u>2</u>	<u>&lt; 0.001</u>
	EVI	<u>33.4003</u>	<u>1</u>	<u>&lt; 0.001</u>
	Rain	<u>0.1512</u>	<u>1</u>	<u>0.6974</u>
	VPD	<u>36.7781</u>	<u>1</u>	<u>&lt; 0.001</u>
	Month:Site	<u>375.8069</u>	<u>22</u>	<u>&lt; 0.001</u>
	SWC:Site	<u>38.7949</u>	<u>2</u>	<u>&lt; 0.001</u>
	Site:EVI	<u>8.2576</u>	<u>2</u>	<u>0.0161</u>
	Site:Rain	<u>14.6424</u>	<u>2</u>	<u>0.0007</u>
	Site:VPD	<u>6.4624</u>	<u>2</u>	<u>0.0395</u>
	Year:Site	<u>990.9702</u>	<u>14</u>	<u>&lt; 0.001</u>

**Table S4:** Type 3 effects for models of entropy

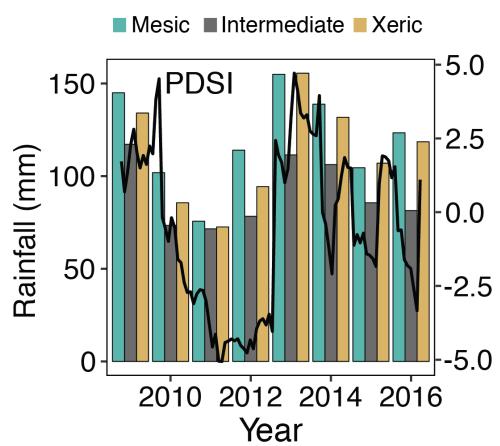
Model	Effect	Chisq	Df	Pr(>Chisq)
<b>σ</b>	Year	4.9661	1	0.664
	Month	606.5841	11	< 0.001
	SWC	39.387	1	< 0.001
	Site	152.0689	2	< 0.001
	EVI	3.2469	1	0.0716
	Rain	250.7508	1	< 0.001
	VPD	2189.9512	1	< 0.001
	Month:Site	156.1348	22	< 0.001
	SWC:Site	23.5738	2	< 0.001
	Site:VPD	10.0017	2	0.0067
<b>JLE</b>	Site:Rain	8.8673	2	0.0119
	Year:Site	122.0744	14	< 0.001
	Year	21.216	1	0.0035
	Month	726.81	11	< 0.001
	SWC	456.76	1	< 0.001
	Site	493.661	2	< 0.001
	EVI	148.839	1	< 0.001
	Rain	127.775	1	< 0.001
	VPD	1011.278	1	< 0.001
	Month:Site	367.42	22	< 0.001
<b>JH</b>	SWC:Site	162.581	2	< 0.001
	Site:EVI	42.076	2	< 0.001
	Year:Site	560.321	14	< 0.001
	Year	38.625	1	< 0.001
	Month	101.071	11	< 0.001
	SWC	25.483	1	< 0.001
	Site	93.504	2	< 0.001
	EVI	25.804	1	< 0.001
	Rain	94.524	1	< 0.001
	VPD	1208.397	1	< 0.001
<b>JG</b>	Month:Site	315.446	22	< 0.001
	SWC:Site	39.127	2	< 0.001
	Site:EVI	44.953	2	< 0.001
	Site:VPD	30.372	2	< 0.001
	Site:Rain	14.251	2	0.0008
	Year:Site	370.91	14	< 0.001
	Year	7.6197	1	0.3673
	Month	180.1628	11	< 0.001
	SWC	35.1066	1	< 0.001
	Site	234.691	2	< 0.001
	EVI	31.1994	1	< 0.001
	Rain	0.8563	1	0.3548
	VPD	29.1953	1	< 0.001
	Month:Site	299.2461	22	< 0.001
	SWC:Site	56.2234	2	< 0.001
	Site:EVI	11.0306	2	0.004
	Site:Rain	22.1752	2	< 0.001
	Year:Site	1082.405	14	< 0.001

**Table S5:** Type 3 effects for models of metabolic energy (NEE<sub>s</sub>) and entropy (J<sub>NEE</sub>)**Deleted:** budget (*eff<sub>rel</sub>*), (c and d) the

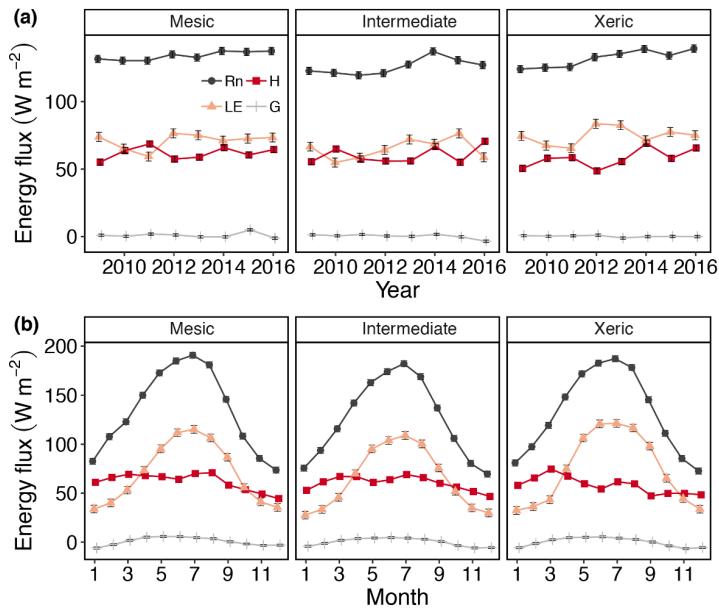
Effect	Chisq	Df	Pr(>Chisq)	Effect
<b>NEE<sub>s</sub></b>	<u>Year</u>	<u>29.646</u>	<u>7</u>	<u>0.0001102</u>
	<u>Month</u>	<u>74.127</u>	<u>11</u>	<u>&lt; 0.001</u>
	<u>SWC</u>	<u>19.826</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Site</u>	<u>779.838</u>	<u>2</u>	<u>&lt; 0.001</u>
	<u>EVI</u>	<u>75.114</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Rain</u>	<u>300.884</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>VPD</u>	<u>327.07</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Month:Site</u>	<u>742.229</u>	<u>22</u>	<u>&lt; 0.001</u>
	<u>Site:EVI</u>	<u>14.519</u>	<u>2</u>	<u>0.0007</u>
	<u>Site:VPD</u>	<u>11.067</u>	<u>2</u>	<u>0.0034</u>
<b>J<sub>NEE</sub></b>	<u>Site:Rain</u>	<u>42.48</u>	<u>2</u>	<u>&lt; 0.001</u>
	<u>Year:Site</u>	<u>520.107</u>	<u>14</u>	<u>&lt; 0.001</u>
	<u>Year</u>	<u>100.0912</u>	<u>7</u>	<u>&lt; 0.001</u>
	<u>Month</u>	<u>734.3098</u>	<u>11</u>	<u>&lt; 0.001</u>
	<u>SWC</u>	<u>102.5001</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Site</u>	<u>472.4768</u>	<u>2</u>	<u>&lt; 0.001</u>
	<u>EVI</u>	<u>123.4161</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Rain</u>	<u>85.0485</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>VPD</u>	<u>839.139</u>	<u>1</u>	<u>&lt; 0.001</u>
	<u>Month:Site</u>	<u>675.38</u>	<u>22</u>	<u>&lt; 0.001</u>

**Table S6:** Type 3 effects for models of entropy efficiency

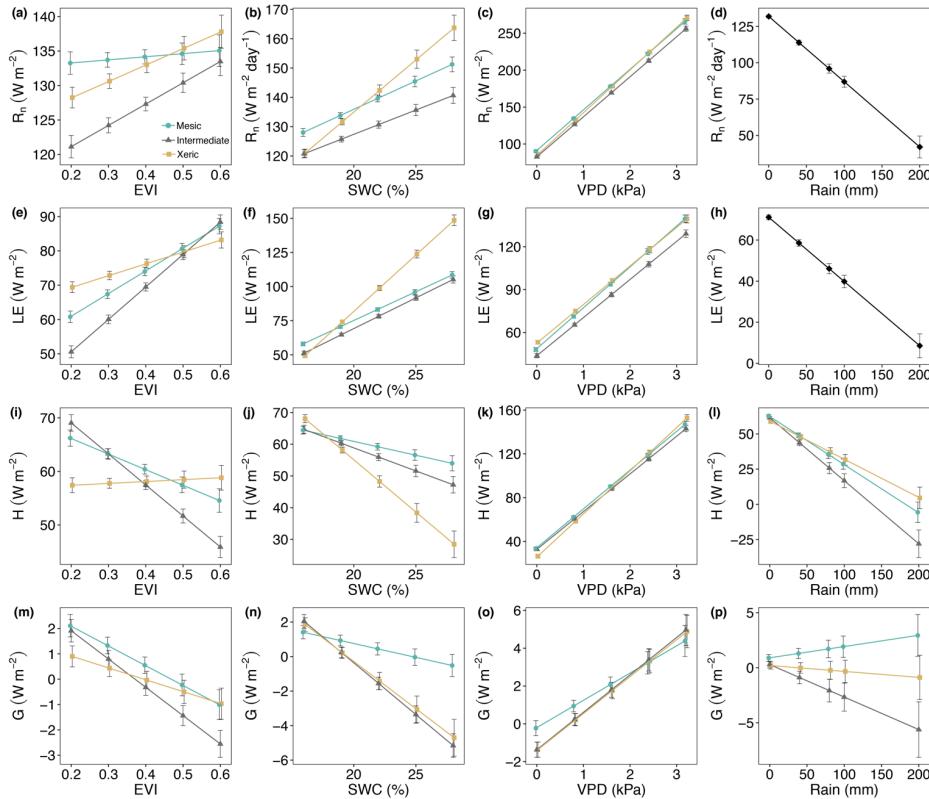
Model	Effect	Chisq	Df	Pr(>Chisq)
<u>eff<sub>rad</sub></u>	Site	<u>321.3179</u>	<u>2</u>	<u>&lt;0.001</u>
	Year	<u>28.9597</u>	<u>7</u>	<u>0.0002</u>
	Month	<u>31.1969</u>	<u>11</u>	<u>0.001</u>
	VPD	<u>269.8436</u>	<u>1</u>	<u>&lt;0.001</u>
	Rain	<u>295.7158</u>	<u>1</u>	<u>&lt;0.001</u>
	SWC	<u>6.6371</u>	<u>1</u>	<u>0.001</u>
	Site:Month	<u>78.91</u>	<u>22</u>	<u>&lt;0.001</u>
	Site:VPD	<u>10.683</u>	<u>2</u>	<u>0.0048</u>
	Site:Rain	<u>17.7766</u>	<u>2</u>	<u>0.0001</u>
	Site:Year	<u>165.2804</u>	<u>14</u>	<u>&lt;0.001</u>
	Mite	<u>938.8639</u>	<u>2</u>	<u>&lt;0.001</u>
	Year	<u>9.2791</u>	<u>7</u>	<u>0.2332</u>
<u>eff<sub>flux</sub></u>	Month	<u>251.1215</u>	<u>11</u>	<u>&lt;0.001</u>
	VPD	<u>1204.1726</u>	<u>1</u>	<u>&lt;0.001</u>
	EVI	<u>5.4535</u>	<u>1</u>	<u>0.0195</u>
	Rain	<u>122.5276</u>	<u>1</u>	<u>&lt;0.001</u>
	SWC	<u>8.9111</u>	<u>1</u>	<u>0.0028</u>
	Site:Month	<u>307.582</u>	<u>22</u>	<u>&lt;0.001</u>
	Site:SWC	<u>25.8864</u>	<u>2</u>	<u>&lt;0.001</u>
	Site:VPD	<u>17.4305</u>	<u>2</u>	<u>0.0002</u>
	Site:Rain	<u>51.4031</u>	<u>2</u>	<u>&lt;0.001</u>
	Site:EVI	<u>15.1919</u>	<u>2</u>	<u>0.0005</u>
	Site:Year	<u>517.3889</u>	<u>14</u>	<u>&lt;0.001</u>
	Site	<u>12.945</u>	<u>2</u>	<u>0.0016</u>
<u>dS/dt</u>	Year	<u>5.9043</u>	<u>7</u>	<u>0.551</u>
	Month	<u>16.8799</u>	<u>11</u>	<u>0.1115</u>
	VPD	<u>114.1762</u>	<u>1</u>	<u>&lt;0.001</u>
	EVI	<u>4.207</u>	<u>1</u>	<u>0.0403</u>
	Site:Month	<u>103.0141</u>	<u>22</u>	<u>&lt;0.001</u>
	Site:Year	<u>135.7525</u>	<u>14</u>	<u>&lt;0.001</u>



**Figure S1:** Monthly rainfall sums and Palmer Drought Severity Index (PDSI) for the mesic, intermediate and xeric sites from 2009 through 2016.



**Figure S2:** Least square mean predicted values from the mixed model results for annual (a) and monthly (b) changes of the energy fluxes of net radiation ( $R_n$ ), latent energy (LE), sensible heat (H), and ground heat (G) at the mesic, intermediate and xeric sites. Error bars represent standard errors (SE).



**Figure S3:** Least square mean predicted values from the mixed model results for interactive effects of (a, e, i, m) enhanced vegetation index (EVI), (b, f, j, n) soil water content (SWC), (c, g, k, o) vapor pressure deficit (VPD) and (d, h, l, p) rain on the energy fluxes of (a-d) net radiation ( $R_n$ ), (e-h) latent energy (LE), (i-l) sensible heat (H), and (m-p) ground heat (G). For (d) and (h) the interaction with site was not significant, as indicated by a single solid black line. Error bars represent SE.

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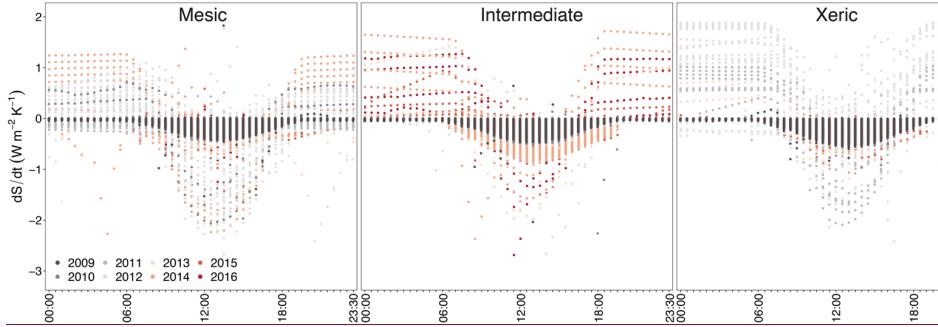


Figure S4: Diurnal changes in entropy ( $dS/dt$ ) at the mesic, intermediate and xeric sites for the years 2011 through 2016.

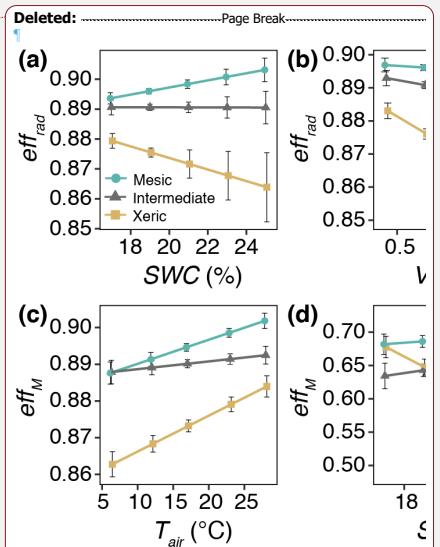


Figure 7: Differences in the entropy efficiencies for radiation entropy efficiency ( $eff_{rad}$ ) in interaction with site and (a) soil water content (SWC), (b) vapor pressure deficit (VPD), and (c) air temperature ( $T_{air}$ ), as well as (d) metabolic entropy efficiency ( $eff_M$ ) in interaction with site and SWC.

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