1.	Dear	editor,

Please find attached a new version of our manuscript entitled "Automatic high-frequency
measurements of full soil greenhouse gas fluxes in a tropical forest" for consideration as
publication in Biogeosciences,

7 We have revised the manuscript based on the feedback of the five reviewers and you can find below
8 a rebuttal letters where we give our detailed responses to each point raised by the reviewers.

- 10 You can also find below the a marked-up manuscript version.
- 12 On behalf of the authors,

14 Elodie Courtois & Clement Sthal

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17 **Referees' comments:**

18

19 **RC1, Referee #1 (Remarks to the Author):**

20

This manuscript presents a protocol for measuring three greenhouse gases (CO₂,CH₄ and N₂O) at high temporal frequency in a tropical forest soil using a combination of commercially available systems. This is a very timely and relevant manuscript, particularly for measurements of CH₄ and N₂O. Continuous high frequency measurements with clear sampling protocols will help researchers capture and model, transient changes in CH₄ and N₂O fluxes often missed by more infrequent sampling strategies. This looks like a nice, efficient system for measuring these important greenhouse gases.

28

29 *Response:* We thank the reviewer for the positive comments and constructive inputs.

30

The technical write-up is comprehensive and easy to follow however I do have technical questions regarding their methodology that should be addressed.

33

1. How did the authors keep moisture from affecting sampling, either within the tubing lines orwhen moving from IRGA to Picarro? Was there any moisture buildup in the tubing lines?

36

37 **Response:** We also worried about moisture problems prior to our experiment, but such problems 38 never occurred. No significant condensation occurred in our system, we assume because 39 temperature variations were typically small below the canopy, varying from 22°C in the night to 40 28°C during the day (see Figure attached). We regularly checked that no liquid water 41 accumulation occurred in the sampling tubes. Moreover, both analysers measure water vapor and 42 account for its effect on concentration of CO₂, CH₄ and N₂O.



44 Hours
45 Figure 1: Daily temperature variation at 2m height under the canopy cover during the study period
46 (June-September 2016).

48

49 Changes in the manuscript: Temperature variations are typically small below the canopy due to 50 the shadowing by dense canopy crown and microclimatic conditions. During the study period, 51 temperature at 2m height varies during the day from 22 °C in the night to 28 °C during the day. 52 The presence of water condensation inside the tubing lines was carefully checked every week and 53 never occurred during the study period.

54

Were the instruments kept within operating temperature ranges: specifically the upper end of
 operating ranges? Where there diel changes in Licor and Picarro instrument temperatures?

Response: In our study site, all the systems were operating below dense understory vegetation and
canopy cover, which naturally create a buffer maintaining air temperature relatively constant over
the year (i.e. at daily and seasonally time scale). There were no diel changes in Licor and Picarro
inside operating temperature: Picarro temperature was monitored and remained between 44.99
°C and 45.01 °C and Licor temperature at 51.7 °C.

63

65

64 3. How often were the instruments calibrated?

66 **Response:** All the systems used at this time, i.e. from June to September 2016, were new and

- 67 received from the manufacturers. We therefore did not re-calibrate them during the study period.
- 68 We are confident that the data of the gas concentrations recorded by our analysers were robust.

70 4. The authors subsampled from the flow downstream of the Li-8100A-IRGA into the Picarro 71 G2308.What was the flow rate through the Picarro analyzer- was it 2.8L/min or some lowerrate?

72 I don't see the flow controller in Figure 1a diagram- is it built into the external pump? 73

74 **Response:** The flow rate of 2.8L/min corresponds to the flow between the chambers and the 75 multiplexer which cannot be adjusted. This high rate inside the chamber allows to achieve a 76 sufficient air mixing in the chamber headspace during the measurements. Flow rates in the 77 subsampling lines (Li8100 and Picarro) were lower and set between 1.5 and 1.7L/min as 78 recommended by the manufacturers.

79

80 **Changes in the manuscript**: Flow rates in the subsampling lines (Li8100 and Picarro) were lower and set between 1.5 and 1.7 L min-1 as recommended by the companies. 81

82 83

84 5. What is the sampling volume inside the Picarro G2308? My concern is that if there is 85 subsampling from a high flow rate at the T piece subsampling loop (Figure 1 a) through a secondary instrument, and then flow is re-merged downstream of the external pump and returned to the closed 86 chamber there may be a dilution effect. This might not impact a 2 minute sampling time period 87 but may have a greater impact on a 25 minute time period. 88 89

- **Response:** The volume of the PICARRO analyser together with tubing from the subsampling loop 90 91 represents 130 cm³. It represents only 2% of the total volume (6088 cm³) and the dilution effect 92 was therefore limited.
- 93

94 6. Did the authors also use their Picarro G2308 to measure CO_2 as well and if so how does that 95 compare to the LI-8150 analyzer? This would provide confidence in running the two systems 96 inline.

97

98 **Response:** One step of our data quality check / quality control consisted in comparing the soil CO_2 effluxes measured by the PICARRO with the soil CO₂ effluxes measured by the Li-8100A. This 99 method was also used to control that there were no leaks. Overall, both systems agreed very well. 100 101 However, while the PICARRO G2308 analyser automatically reports dry mol fraction of CH₄ and 102 N_2O , it only reports CO_2 uncorrected by H_2O concentration. As the precision of measurement was 103 better for CO₂ using the Li8100 and as it also automatically reports dry mol fraction for this gas,

- 104 we decided to use Picarro estimation for CH_4 and N_2O and Li8100 estimation for CO_2 .
- 105
- 106 Specific questions:

107 Pg 2 line 26: are the reference [17,18] in the correct format? This is the only location that lists 108 reference numbers as opposed to first author.

- 109
- 110 **Response:** This has been corrected
- 111
- 112 Pg 3 line 23: the authors used an external pump, however the LI-8150 has an internal diaphragm
- pump- was this turned off or was it used in line with the external pump? 113
- 114

115 **Response:** The system was always operating with the two pumps, the internal diaphragm pump of

- the Li-8150 and the external pump of the PICARRO, turned on, which limited the risk of water
 condensation inside the tubing line.
- 118

119 Pg 3 line 28: the CRDS is not the "only" method that can detect low concentrations of N_2O . Can 120 you change to "one of the only" as opposed to "only".

- 122 *Response:* This has been corrected.
- 123

127

121

Pg 4 line 14: the flow rate of 2.8L/min is very high. Was this flow rate tested to determine if pressure within the closed chambers was altered? I assume that the Licor 8100 chamber tops had their patented pressure relief value installed?

Response: The flow rate of 2.8L/min corresponds to the flow between the chambers and the multiplexer which cannot be adjusted. This high rate inside the chamber allows to achieve a sufficient air mixing in the chamber headspace during the measurements. Flow rates in the subsampling lines (Li8100 and Picarro) were lower and set between 1.5 and 1.7L/min as recommended by the manufacturers.

133

Figure1b) are all the tubing lines from the 16 chambers to the multiplexer 15m (as in the text)inlength? A small point but they look like different lengths in the diagram.

136

137 *Response:* Yes the length of tubing was 15 m for all chambers but as chambers are installed in a
138 grid around the instruments, they are not all 15m away from the instruments. The length of the
139 tubing lines (15 m) is cited in the caption.

140 141

142 **RC2, Referee #2 (Remarks to the Author):**

143

The authors measured in situ soil fluxes of CO2, CH4 and N2O continuously with a commercially 144 available automated chamber system coupled with a CRDS analyser in a tropical forest for four 145 146 months. The manuscript is focused entirely on the methodological aspect of these measurements, stressing the importance of adjusting chamber closure times for the different gases for reliable flux 147 calculations. The effect of closure times on flux calculation results were studied by trying different 148 149 chamber closure times in the field and by adjusting the number of data points used for the actual flux calculation. This is a well-designed study, and overall, the manuscript is well written and 150 structured. Also choosing appropriate chamber closure times and the operation of automated 151 chamber systems are important topics for the soil flux community. However, I recommend 152 publication of this manuscript after major revision because I have some general concerns with this 153 154 manuscript.

- 159 1. Soil Flux Pro provides for each chamber measurement linear and non-linear flux calculations.
- 160 Why did you choose to use only the linear flux calculation results?

Response: We express our deep thanks to reviewer for his positive comments about our manuscript
 and constructive remarks. We have addressed, see below, our answers to each comment / remark.

162 **Response:** We firstly decided to use linear flux calculations only because we thought that the 163 saturation effects characterised by a plateau after a certain time would be low. However, following 164 your comment and comment from other reviewers, we changed flux calculations in the new version 165 of the manuscript to use exponential estimations.

166

167 The underestimation of fluxes with linear regression due to saturation effects is well-known. That's 168 why numerous non-linear calculation schemes have been developed. Could you have significantly 169 reduced the chamber closure time for the N₂O flux calculation if you had used non-linear flux 170 calculation?

170 171

172 *Response:* See previous comment.173

The selection of the flux calculation scheme can change the MDF at least for chamber measurements with only few gas samples over time. Does this effect disappear with highfrequency analysers, i.e. selection of the flux calculation scheme becomes less crucial in that regard? Would there still be a significant difference between the SHORT and LONG flux calculation for the different gases when using non-linear flux estimates?

179

Response: The standard error approach that we used (Nickerson, 2016) is a first order approximation for the MDF from high-frequency measurements and the "true" MDF is a function of the chamber time-series fit type as well (i.e. Linear, exponential, quadratic). Nonetheless, while the use of linear regression resulted in systematically smaller fluxes as compared to exponential regression. It is therefore recommended to initially calculate fluxes with linear regression to

determine the threshold for "low" fluxes and to recalculate them using exponential regression
(Korkiakoski et al., 2017).

187

Korkiakoski, M., Tuovinen, J.-P., Aurela, M., Koskinen, M., Minkkinen, K., Ojanen, P., Penttilä,
T., Rainne, J., Laurila, T. and Lohila, A.: Methane exchange at the peatland forest floor-automatic

190 *chamber system exposes the dynamics of small fluxes, 2017.*

Nickerson, N.: Evaluating gas emission measurements using Minimum Detectable Flux (MDF),
Eosense Inc Dartm. N. S. Can., 2016.

193 2. You write about high-frequency measurements only as sampling measurement plots more 194 frequently over time. However, you could also address the high-frequency sampling during a 195 chamber closure since you use high-frequency gas analysers and work with MDF in your data 196 analysis. There are several automated chamber systems which do not employ high frequency 197 analysers, but still collect discrete gas samples which have to be analysed with a GC. Especially 198 for N₂O it is very interesting to see what fluxes we can capture with CRDS in comparison to GC 199 analysis.

200

201 **Response:** In a previous study in the same environment (Courtois et al., 2018), we estimated that

202 the minimum detectable fluxes using Gas Chromatography analysis of four discrete gas samples

203 over 30 minutes for N2O was $\pm 8.3 \mu g N m - 2 h - 1$. MDF estimated in the present study using high

frequency measurement was 0.002 nmol m-2 s-1 or 0.2 μ g N m-2 h-1 for N₂O which is therefore ~

205 40 times lower. We added a sentence in the manuscript to highlight this interesting fact.

- 207 Courtois, E. A., Stahl, C., Van den Berge, J., Bréchet, L., Van Langenhove, L., Richter, A., Urbina,
 208 I., Soong, J. L., Peñuelas, J. and Janssens, I. A.: Spatial Variation of Soil CO₂, CH₄ and N₂O
 209 Fluxes Across Topographical Positions in Tropical Forests of the Guiana Shield, Ecosystems, 1–
 210 14, 2018.
- 3. You only write how the SHORT and LONG measurements affected the flux estimates. But how
 did they affect the uncertainty of the single flux estimates? How large/small were the error bars
 for the flux estimates?
- 214

Response: Comparison of standard error of single flux estimates using 2 minutes or 25 minutes estimations for two weeks (from August 2nd for August 9th and from August 16th for August 25th) shows that standard errors are always higher for 2 minutes than for 25 minutes estimation for all three gases (Figure 2 below). Nonetheless, we decided not to integrate this figure in the manuscript because it does not add much to the study.

220



Figure 2: Comparison of standard error of single flux estimates using 2 minutes or 25 minutes
 estimations for two weeks (from August 2nd for August 9th and from August 16th for August 25th)
 224

4. Could you have just used one, namely the LONG, closure time for all chambers and only choose
for the flux calculation between SHORT and LONG calculation times? This would be more
practical than rotating closure times between chambers.

- 230 **Response:** Setting all chambers as LONG measurements would have led to a maximum of ~ 3 231 measurements only per chamber and per day. Mixing LONG and SHORT measurements allows to 232 maximise the number of measurements per chamber and per days while ensuring a reliable 233 estimation of the low N₂O fluxes and to capture transient peaks of CH₄ and N₂O.
- 234

221

225

235 236	5. In section 3, the results are clearly presented, but the discussion part is very limited.
237	Response: The main aim of our study was not to identify controls and mechanisms of the soil GHG
238	fluxes but rather to test novel soil GHG systems for continuous high-frequency measurements. We
239	think that this manuscript could be used as technical support to set up new soil systems and
240	contribute to record comparable soil GHG data in other regions around the world. Nonetheless,
241	the discussion has been revised in the new version of the manuscript to integrate discussion on
242	spatio-temporal variability of fluxes based on our study.
243	The second s
244	
245	Specific comments:
246	Page 2, line 26: numbers instead of author names for references
247	
248	Response: This has been corrected.
249	
250	You are not always consisted in how you write company names (capital versus small letters). Also
251	often vou write 'minute' when vou could just use 'min'.
252	
253	Response: We corrected this in the new version of the manuscript.
254	
255	Page 3, line 23: What are the pump specifications? Was it the pump supplied by Picarro with the
256	instrument or did you use another pump?
257	
258	Response: We have included more information in the manuscript about the external pump
259	provided by PICARRO: recirculation pump A0702
260	
261	Page 4, lines 9.: The soil temperature and soil moisture probes, were those the ones which can be
262	directly attached to the chambers?
263	
264	Response: The soil temperature and soil moisture probes were those provided by Li-COR, which
265	are directly attached to the chambers. The probes measured soil temperature and soil moisture
266	around the PVC collars. We have added more details in the text.
267	
268	Page 5, line 5.: Did you use the analytical accuracy specified on the technical data sheets of the
269	analysers or did you perform measurements yourself?
270	
271	Response: We used the analytical accuracy specified on the technical data sheets of the analysers.
272	This is now specified in the new version of the manuscript.
273	
274	Where there significant air pressure and ambient temperature changes at your site over the four
275	months? If yes, did you test how different temperature and pressure values could change the MDF
276	estimate? Is nincl. or excl. the deadband?
277	
278	Response: Please, see our response to reviewer 1 above (2.). In our study site, because all the
279	systems were operating under dense understory vegetation and canopy cover, air temperature

- remained relatively constant over the year (i.e. at daily and seasonally time scale) near the soil
 surface. This is also true for air pressure.
- 282 St
- Page 6, line 6: What CO2 concentrations were reached during LONG closure times (and forCH4)?
- 285 *Response:* CO₂ concentration can reach 2000 ppm and CH₄ concentration 4 ppm.
- Page 6, line 7: I find that confusing in comparison to section 2.5. So considering the deadband, the
 chambers were closed for 3 and 26 minutes, respectively?
- 289

Response: No, the chamber stayed close for 2 minutes and 25 minutes and the first minute was not used for the flux estimation. As we have a sampling frequency of 1 second, it still represents 60 points for curve fitting. Nonetheless, we agree that this could be considered as a too short period for CH_4 and CO_2 estimation using the SHORT (2 minutes) closure time. We therefore compared

the CO₂ and CH₄ estimations with a deadband of 60 seconds (fluxes estimation with 60 seconds)

with a deadband of 30 seconds (fluxes estimation with 90 seconds) for the week from August 16th

296 to August 25th. These two estimations were very well correlated (see figure 3 below) so we decided

- 297 to keep our 60 s deadband results.
- 298



299

Figure 3: Comparison of CO_2 and CH_4 fluxes with a 30 seconds and a 60 seconds deadband for the week from August 16th to August 25th. The red dotted lines represents the 1:1 line.

302

Page 6, line 20: Why did you consider these fluxes as unreliable when the chamber quality check using the R^2 for CO₂ was passed? Are you not unnecessarily filtering out fluxes which are not significantly different from zero, and thus introducing a bias in your data? Because this often happens when using R^2 as a filter criterion for low fluxes.

308 309	Response: Because of the high soil respiration activity, low soil CO_2 fluxes do not really occur in this tropical rainforest, not even during the dry season. When the R^2 criterion for CO_2 was not
310	passed, it always corresponded to situations of imperfect closure of the chamber, due to leaves or
311	small branches lying on the soil collars (381 measurements over 17796, i.e. 2.1%). In these cases,
312	it was therefore necessary to remove flux data for the three gases.
313	
314 315	Page 6, first paragraph of section 3: You had no problems with humidity and the automated chamber system at your site?
316	
317	Response: Please, see our response to reviewer I above (1.); at the given flow rate and the small
318	diel cooling, we had no problems with water condensation inside the tubing lines of our system.
319	
320 321 322	for your small chambers. Except for the necessary descriptions in the method section, you completely disregard the role of chamber volume for choosing the right chamber closure time.
323 224	Besponse. We added a sentence to precise that our result is valid for small chambers only
324 325	Kesponse: we daded a sentence to precise that our result is valid for small chambers only.
326	Page 7, lines 20/21: That sentence does not make any sense to me. 85.6 % of the fluxes were above
327	or below?
328	
329	Response: This sentence has been considerably modified.
330	
331	Page 8, line 4: You didn't show the diurnal variation in your data. This is more a point for the
332	discussion than a conclusion from your presented data.
333	
334	Response: We agree with the reviewer; however, ecological interpretation of our data will require
335	more long-term data and will be published in a future paper. Here, we wanted to provide a
336	technical report with information on how to get robust results on GHG flux estimations rather
337	than on how these fluxes are produced and vary.
338	
339	The references are not well formatted.
340	
341	Response: References were reformatted
342	
343	Table 1: Use superscript for the units.
344	
345	Response: This has been corrected
346	
347	Table 2: Include n for each chamber.
348	
349	Response: This information has been added.
350	
351	Check how the units are written on the y-axis of the figures.
352	
353	Response: This has been checked

- 354
- 355

356 **RC3, Referee #3 (Remarks to the Author):**

357

358 In this manuscript, the authors detail a field-deployed and field-tested system for measuring soil 359 greenhouse gas (GHG) emissions (CO2, CH4 and N2O) from a tropical wet forest; the system 360 leverages a commercially-available automated flux chamber system with a commercially-361 available CRDS analyzer. More specifically, the authors (a) outline the technical protocol for implementing such a system, (b) report the mean fluxes and variability of CO2, CH4 and N2O 362 363 observed over the four-month deployment period and (c) test two chamber closure lengths to determine the most effective experimental design for capturing fluxes above minimum detectable 364 levels. Successfully implementing such a system in the tropics is both difficult and has only been 365 done rarely, so a technical note detailing how to do so is absolutely a contribution to the literature. 366 367 I have two general/more broad comments about the manuscript, which I detail below, and also include several more specific comments at the end of this review 368

369

370 *Response:* We thank the reviewer for the positive comments and constructive inputs.

371

1. Concerns regarding flux calculation protocols

In these automated, high-frequency GHG systems, one important set of experimental design and protocol decisions govern how to calculate flux rates and screen for acceptable data points. The authors lay out fairly transparent information about how they calculated their flux rates for each gas, but I wonder if more discussion of the implications of their calculation choices is warranted. I have a few specific questions. Am I correct in understanding that the authors only calculated flux rates for all three gases based on a linear model (Pg 4, Ln 27)? If so, I wonder why they didn't consider also fitting exponential models to the CH, and N=O fluxes, if not the COs fluxes

379 consider also fitting exponential models to the CH_4 and N_2O fluxes, if not the CO_2 fluxes.

380

Response: We first decided to use linear flux calculations only because we thought that the
 saturation effects characterised by a plateau after a certain time would be low. However, following
 your comment and comment from other reviewers, we changed flux calculations in the new version
 of the manuscript to use exponential estimations.

385

The authors themselves note on Pg 5, Ln 17 that using certain chamber closure times (which, of course, this paper is very interested in) in combination with a linear flux fit can lead to flux underestimation. Couldn't the "optimal" chamber closure time that the authors attempt to find also include some experimental designs with different closure lengths but non-linear flux fits?

390

Response: The standard error approach that we used (Nickerson, 2016) is a first order approximation for the MDF from high-frequency measurements and the "true" MDF is a function of the chamber timeseries fit type as well (i.e. Linear, exponential, quadratic). Nonetheless, while the use of linear regression resulted in systematically smaller fluxes as compared to exponential regression. It is therefore recommended to initially calculate fluxes with linear regression to determine the threshold for "low" fluxes and to recalculate them using exponential regression (Korkiakoski et al., 2017).

Korkiakoski, M., Tuovinen, J.-P., Aurela, M., Koskinen, M., Minkkinen, K., Ojanen, P., Penttilä,
T., Rainne, J., Laurila, T. and Lohila, A.: Methane exchange at the peatland forest floor-automatic
chamber system exposes the dynamics of small fluxes, 2017.

402

403 Nickerson, N.: Evaluating gas emission measurements using Minimum Detectable Flux (MDF),
404 Eosense Inc Dartm. N. S. Can., 2016.

405

406 Additionally, the Picarro G2308 records numerous diagnostic variables alongside GHG 407 concentrations, including measures of moisture, temperature, pressure, etc. From methods section 408 2.6, I am under the impression that fluxes were only struck from the dataset if they were (a) below 409 the MDF, (b) had an R² for CO₂ < 0.9, or (c) only struck N₂O data if the SHORT R² < 0.8. First, 410 the authors might consider including a supplemental figure that justifies their decision not to have 411 data quality rules around CH₄, as they do for N₂O.

412

413 **Response:** We decided to use the same quality check for CH₄ than for N2O and to only struck CH₄ 414 data if the SHORT $R^2 < 0.8$. As you can see in the CH₄ figure, the main emission or consumption 415 peak during the 2 minutes measurements are still present.

416

417 Second, like the other reviewers, I was curious as to how humidity was dealt with, since it appears 418 that the moisture-related diagnostics weren't used to evaluate data quality. Was there a water trap 419 that isn't marked in the instrument set-up diagram?

420

421 *Response:* We regularly checked that no liquid water accumulation occurred in the sampling
422 tubes. Moreover, both analyzers measure also water vapor and its effect on concentration of CO₂,
423 CH₄ and N₂O was accounted for. We also monitored Licor and Picarro inside operating
424 temperature: Picarro temperature remained between 44.99 and 45.01 C and Licor temperature at
425 51.7 C.

426

427 More broadly, as this paper aims to outline best practices for setting up this kind of experiment in 428 the tropics, fleshing out the data management aspect of things would improve the paper, in my 429 opinion.

430

431 *Response:* We think that we already included many aspect of quality check and data management

and more information can be found in the new version of the manuscript. All the subsequent figure
and analysis were done using *R* software and we would be happy to share the codes upon request.

- 433 434
- 435 2. Technical note vs. data exploration paper

The aspects of this paper that serve as a technical note are novel and helpful. That said, the results and discussion section, in which the observed GHG fluxes are analyzed, is perfunctory and is relatively focused on a methodological question: what chamber closure time should be used in this

439 system, and how can others determine what chamber closure time to use in their analogous system?

440 I found myself wishing that there was a more robust analysis of the GHG data itself and the

- 441 ecological implications of their various findings.
- 442 See also my comment below about Table 2.
- 443

444 **Response:** See RC2 (5)

446 <u>Specific comments:</u>

* Pg 7, Ln 30: The authors ultimately recommend a sampling protocol that rotates between short
and long closure times. What is preventing them from recommending always doing LONG
chamber closures and only using the first two minutes of chamber closure time to calculate the
CO₂ flux, thus decreasing the amount of human labor needed to swap out the system program once
a week?

452

453 Response: Setting all chambers as LONG measurements would have led to a maximum of ~ 3 454 measurements only per chamber and per day. Mixing LONG and SHORT measurements allows to 455 maximise the number of measurements per chamber and per days while ensuring a reliable 456 estimation of the low N_2O fluxes and to capture transient peaks of CH₄ and N_2O . Moreover, under 457 tropical conditions, a visit once a week is absolutely necessary to ensure a proper maintenance of 458 the setting (removing fallen leaves, branches, etc...).

459

* Pg 7, Ln 23 / Figure A2: This figure is used to justify why LONG chamber readings weren't
reliable for N₂O flux estimation, but these data don't indicate that the variable fluxes are unreliable,
only that they are variable and considerably larger in magnitude than the SHORT N₂O flux
estimates. Can an additional supplemental figure be added showing the R² values for the LONG
vs. SHORT N₂O fluxes? Or some similar figure that shows why the LONG fluxes are considered
unacceptable?

466

467 **Response:** If we understood well your question, you are asking why we considered SHORT (not 468 LONG) N_2O fluxes as unacceptable. In order to compare the fluxes from SHORT and LONG 469 closure time, we have added a third graph to the Figure 2 (previously supplementary Figure 1). It 470 shows that N_2O fluxes estimated using the SHORT closure time are not correlated with fluxes 471 estimated with the LONG closure time (R^2 of 0.02) and should not be considered.

472

473 * Pg 7, Ln 31: "Our unique system..." and Pg 8, Ln 12: "this is the first time that this experimental set up is described and tested under tropical field conditions." I believe an analogous system was 474 (O'Connell 475 described in Puerto Rico et 2018. Nature Communications, al 476 https://doi.org/10.1038/s41467-018-03352-3), though not as a technical or methods paper.

477

478 **Response:** We realised that, although the suggested reference was in the list of references, it was 479 not mentioned in the text and we now cite this reference in the introduction. In this reference, the 480 authors are measuring CO_2 and CH_4 but not N_2O fluxes. Moreover, the description of the 481 experimental design of the automated soil fluxes measurement system in this paper is very short 482 and does not give the specification of the multiplexer not the exact model of the CRDS Picarro 483 analyser that they used. We therefore still believe that our study is the first one describing in detail 484 the simultaneous measurement of the three GHGs under tropical field conditions.

485

488

489 *Response:* In the revised manuscript, this figure is now presented in the main text (Figure 2) and 490 also includes the N_2O .

^{486 *} Appendix Figure A1: Might be worth including the N₂O comparison just as the CO₂ comparison
487 is included even though the authors discard the LONG CO₂ flux estimates

492 * Table 2: A number of the authors' chambers reported mean N_2O fluxes below 0. This seems 493 worth mentioning in the results and/or discussion.

494

495 **Response:** We agree with the reviewer and, as suggested, we have added some comments about 496 the respective parts of the N_2O and also CH_4 emissions and absorptions. We have estimated that 497 28% of our fluxes indicated a sink for N_2O and 72% a source for N_2O and that 59% of our fluxes 498 indicated a sink for CH_4 and 41% a source for CH_4 .

499

500 SC1, Referee #4 (Remarks to the Author):

501

In the recent years, several studies highlighted the need for continuous measurements of soil GHG 502 503 other than CO2, which has been technically challenging for long time. However, combination of 504 different new instrumentation allows addressing this challenge nowadays. I think this manuscript 505 a timely technical note addressing one of the most important issues regarding continuous 506 measurements: which is the balance between frequency and reliability of measurements? Despite some of the points discussed in the paper are instrument specific considerations (Li8100 and 507 G2308), I think that most of them apply for high-frequency studies using other instrumentation. In 508 509 my opinion, two points could be covered more in depth in order to make the manuscript more 510 strong and inspiring for the community:

511

(i) suitability of linear or exponential fits for estimating GHG fluxes, especially under highemissions and long chamber closure time

514

515 Response: We first decided to use linear flux calculations only because we thought that the 516 saturation effects characterised by a plateau after a certain time would be low. However, following 517 your comment and comment from other reviewers, we changed flux calculations in the new version 518 of the manuscript to use exponential estimations.

519

(ii) which threshold criterions do we have to apply for low rate fluxes and which are the
 consequences of using different criterions on temporal patterns (both short and long term scales)
 and on accumulated emission estimates.

523

Response: Some information regarding this issue were added in the new version of the manuscript.
525

Finally, I want to recognize the challenge of running this complex instrument setup in a tropical
forest. Dealing with high moisture when using IRGAS and CRDS is not easy, but the authors
succeeded.

529

530 *Response:* Thank you for this positive comment.

531

532 I am looking forward to see the data in the full experiment context with their ecological 533 implications. Here you could find some specific comments, suggestions and open discussion 534 points:

- 535 Pg3 L19-28. Li-8100 can detect really small fluxes of CO_2 as well. I guess that the main reason
- 536 for using both Li-8100 and Picarro G2308 is that one instrument controls the chambers and the

537 other measures the three gases. Additionally, measuring simultaneously CO₂ with two independent 538 systems is a good control to validate the proper performance of the instruments. I wonder which

- 539 was the agreement in CO_2 between Li8100 and G2308.
- 540

541 **Response:** In our system, the automated chambers were controlled by the Li-8150, which was 542 controlled by the Li-8100A. The gas analysers were Li-8100A for the CO₂ and PICARRO G2308 543 for the CH₄ and N₂O. We recommend using the Li-8100A to determine the soil CO₂ effluxes but 544 recognise that CO₂ information provided by the PICARRO can still be used to check for potential 545 leaks in the analysers / tubing.

546

P4 L24-28. I don't know if I understand this statement, but SoilFluxPro (Li-COR software) allows
to directly upload hundreds of Picarro files simultaneously (up to 2 months). You can choose to
open all the files in one single file, and directly merge it with the Li-COR data.

550

551 **Response**: When we first used this import function in SoilFluxPro, we realised that there were 552 problems when importing Picarro files (that are split in one file per hour) by suing the function 553 IMPORT. When a measurement was overlapping two distinct hours (for example a flux estimation from 8:50 am to 9:15 am, the function RECOMPUTE in SoilFluxPro only take into account the 554 first Picarro file. We contacted Licor for this issue and they agree that this was a weakness of the 555 556 software. They are currently working on it to implement this in a new version of the software. In the meantime, the use of the R function that we developed and that can be found in the 557 supplementary material to merge all Picarro files in one file allows to overcome this issue. 558 559 P4 L28-30. One of the best things of using SoilFluxPro is that calculates the fluxes using both 560

561 linear and exponential fits, which could result in substantial differences in fluxes (see the attached 562 example from my own data). My experience is that exponential equations usually fits better than 563 linear ones (in terms of \mathbb{R}^2), especially for high flux rates under long chamber closure times.

565 **Response:** See responses to previous comments, exponential fits were now used for all flux 566 computations.

- 568 P4 L29. Which was the actual length of each measurement without including the deadband?
- 569

567

564

570 *Response:* See comment from RC2

571

575

572 P5 L5. Why are you not using CO₂ measured with Picarro?

573
574 *Response:* Please, see our response to reviewer 1 above (6.).

P5 L7-8. I guess that it has to be the volume of the system (chamber, Li8100, Picarro, multiplexer
and tubing). This is really important since the volume of the system is a parameter controlling the
minimum detectable flux, so Table 1 might substantially change depending on this "detail".

580 *Response:* Yes, it is the volume of the whole system.

581

579

582 P5 L16-17. Again, this can be solved using exponential fits.

584 **Response:** See responses below

586 P6 L7, L13 and L17. I guess these are not the correct figures.

587

585

588 **Response:** This has been corrected

589

590 P6 L14-20. As far as I understood, you kept values higher than MDF (for emissions) and lower 591 than -MDF (for sinks), but what happen with values close in between (-MDF < x < MDF)? What 592 did you do with values close to 0 flux? And what happen if a flux was higher than MDF but had low R^2 ? The same applies for N₂O. Which criterion we have to use when measuring gas emissions 593 594 at low rates? Is it a 0 flux, NA, should we keep the calculated flux regardless of the R²? Choosing 595 one or other criterion might have several implications in order to estimate cumulative or mean 596 fluxes, especially if the data does not have normal distribution and it's not 0 centered. In L11 you describe an R² criterion for considering stable micrometeorological and chamber conditions based 597 598 on CO₂. Then, why we should apply other criterion for the other gases if the conditions are stable? I understand that for this might not be super relevant for a technical note, but this is a key question 599 if you want to quantify emissions in natural conditions. In my opinion, this is the core of the study, 600 and one of the most challenging issues we need to address when measuring CH₄, N₂O and other 601 trace gases. When we have high fluxes, everything is clear, but when we have low fluxes, it turns 602 more complicated. We were discussing this issue in Petrakis et al. 2017, but I still don't have the 603 604 answers. I guess there is not a silver bullet.

605

606 **Response:** We agree with the reviewer that this is one of the main challenge of CH₄ and N₂O soil 607 fluxes estimation. For CO₂, fluxes from tropical soils are always high so the R^2 criterion allows to 608 easily detect measurement issues such as imperfect chamber closure. In this case, it is logical to 609 remove fluxes estimation for the three gases. For CH₄ and N₂O, we decided (1) to consider fluxes 610 below MDF as null fluxes (i.e. fluxes so small that they are below detection limit) (2) to consider 611 fluxes above MDF but with a low R^2 were considered as NA (Not available, fluxes estimation 612 impeded by unknown measurement issues)

- 612 *impeded by unknown measurement issues).*
- 613

Figure 1A. There is something in this panel it's not completely clear. As far as I understand, the air goes from the chamber to the multiplexer, to Li8100, to G2308, to the external pump, to the multiplexer and again into the chamber. However, in the schematic view there is a black circuit (T piece sub-sampling loop) that connects the multiplexer, Li8100, G2308 and the external pump. Since the air composition does not change between these four elements, why the subsampling tub was not inserted in serial at one point of the circuit?

620

Response: Inserting the subsampling loop in parallel rather in in serial was a proposition from
the manufacturer that we followed here.

623

Table 2. I wonder which closure time did you use in this table (2 or 25min). It would be interesting a comparison between 2 and 25 closure times. I'm not sure you will find differences in the means. This would suggest that short closure times might not affect the annual balance but deviation of

- 627 the data (as we can see in Ap Figure 2).
- 628

629 **Response:** In this table, we used all fluxes estimation available after quality checking. Your 630 proposition is interesting but as 2 and 25 minutes estimation were not made on the same weeks, it 631 is difficult to compare them. We therefore propose to keep this table as it is but we stated in the 632 table caption that this estimation was made using all the data available and we also added the number of data points that were used for this estimation. 633 634 635 Appendix Figure A1. In my opinion, this is one of the most interesting figures in the manuscript 636 and I think it should be place in the main manuscript. 637 638 **Response:** The figure was placed in the main text 639 Some suggestions: 640 641 a) Regressions will have better fit if you use exponential equations for estimating the flux. For each flux you can choose linear or exponential depending on the R^2 . 642 643 644 **Response:** Exponential fits were now used for all fluxes estimation. 645 b) Could you display R² and the coefficients of the regressions between 2 and 25min? Regression 646 B shows a good fit, but it seems that 2 min fluxes tends to overestimate fluxes compared to 25min 647 estimates. Again, this could be an artifact of using linear regressions and not exponential. 648 649 **Response:** This figure has been modified by using exponential fits and the R^2 were added 650 651 c) It would be interesting plotting the regression for N₂O including all values (without removing 652 data using R^2 or MDF criterions)? This is related to my comment on Table 2. 653 Apendix Figure A2. Please, edit the figure caption. Petrakis, S., Barba, J., Bond-Lamberty, B. and 654 Vargas, R.: Using greenhouse gas fluxes to define soil functional types, Plant Soil, 1–10, 2017. 655 656 SC2, Referee #5 (Remarks to the Author): 657 658 659 This manuscript focused on a very important topic about soil CO2/CH4/N2O fluxes 660 in tropical rainforest. The experiment was well designed. Particularly, this may be the world's first report about in situ and simultaneously measurement of soil CO2/CH4/N2O 661 fluxes at low latitude (between 10°N and 10°S). I would like to give the authors my 662 663 comments. 664 665 **Response:** Thank you for this positive comment. 666 667 1. Important references: To date, through the "Web of Science", I could not find any publication about continuous 668 measurement of soil CO2 efflux (Rs) using the automated chambers in the low latitude tropical 669 forests that between 10 N and 10 S. Though two campaign studies in very humid forests (ï'C 670 s3500 mm of annual precipitation) using automated chambers each in northeastern Australia 671 (17 S) (Kiese and ButterbachBahl, 2002) and northeastern Puerto Rico (18 N) (Wood et al., 672 2013) were conducted only less than 6-month period, they observed similar phenomenon with Rs 673 was higher during the dry season but lower during the wet season. Kiese and ButterbachBahl 674

675 (2002) also measured N2O flux. Conversely, a 4-year continuous measurement of Rs in a 676 seasonal dry (1,250 mm of annual precipitation) tropical forest in western Thailand (14 N)

- 677 showed higher Rs in wet season than that of dry season (Hanpattanakit et al., 2015).
- 678

679 **Response:** In fact, a very recent paper reported continuous monitoring of Rs during three years in the tropical forest of Panama (Rubio and Detto, 2017). Moreover, a previous study conducted 680 681 at the same site as ours (Paracou site, near the Guyaflux tower) also reported 577 days of Rs 682 measurement (Rowland et al., 2014). Both references highlighted a significant effect of soil moisture on seasonal and diurnal cycles of Rs. Together with the two other references from 683 684 tropical that you cited, there provide evidences that Rs in tropical forest soils are typically higher in the wet than in the dry season. The other study that you cited (Hanpattanakit et al., 685 2015) was conducted in a seasonally dry forest which are apparently reacting differently than 686 687 typical tropical wet forest (precipitations > 2000m/year). Nonetheless, the results that we are 688 presenting in our study were conducted from June to September 2016 which corresponds in our 689 site to the end of the wet season and the onset of the dry season. With these data, we cannot 690 discuss seasonal effects, at least one full year, or more, of measurements would be necessary for 691 this. 692 693 Rowland, L., Hill, T. C., Stahl, C., Siebicke, L., Burban, B., Zaragoza-Castells, J., Ponton, S., 694 Bonal, D., Meir, P. and Williams, M.: Evidence for strong seasonality in the carbon storage and 695 carbon use efficiency of an Amazonian forest, Glob. Change Biol., 20(3), 979–991, 2014. 696 697 Rubio, V. E. and Detto, M.: Spatiotemporal variability of soil respiration in a seasonal tropical 698 forest, Ecol. Evol., 7(17), 7104–7116, 2017. 699 700 2. CO2 flux: 701 Empirically, also see the above references, CO2 flux is largely controlled by soil moisture (rain 702 events) at tropical forests. However, based on Fig 3, during 4-month experiment (June-703 September 2016), most of the chambers did not show temporal variation in CO2 flux. Thus, the 704 authors are suggested to add soil moisture (and temperature) data to Fig 3 and provide some 705 discussion about the (lack of) relationships between Rs and soil moisture and temperature. 706 707 **Response:** As discussed in above, a four months period is limited to go deep into such relationships, especially in tropical forest where temporal and spatial variability of fluxes are 708 709 high. You can find below a figure that can now be found in the supplementary material of the 710 manuscript displaying the relationship of the three gases with soil moisture. Nonetheless, going

- deeper in the discussion of the effect of rain event, soil moisture and the relative importance of 711
- 712 spatial, seasonal and diurnal variability of fluxes cannot be done with these dataset that was
- 713 specifically constructed to demonstrate the feasibility of running the system under tropical
- 714 conditions.
- 715
- 716 3. CH4 flux:
- Generally speaking, upland forest soil is a CH4 sink, even lowland tropical forest soil. Compared 717
- to Rs, however, CH4 flux is more complex and generally has large spatial variation, because the 718
- 719 termite activity can emit CH4 thus offset a partial of the soil CH4 sink. I am confused with Table
- 720 2, because ten of the sixteen chambers showed CH4 source. Li-Cor soil chamber (8100-104) can

721 be considered to block most activity of the termite, because the chamber base (collar; 7 cm in

- 722 height) was inserted 7 cm into the soil and left another 4 cm above the soil; in addition, the
- 723 chamber has relative additional big metal base surround the collar. On the other hand, inserted
- 724 chamber base (collar) into the tropical (clay) soil can (sometimes) cause waterlogging inside the
- Li-Cor soil chamber (8100-104), which might convert the CH4 sink to CH4 source. Same with 725
- 726 CO2 flux, temporal variations in CH4 fluxes also could not be detected in Fig. 4. Also, 727
- megascopically, the chambers did not show the common pattern of temporal variation in CH4 728 fluxes (Fig 4). Sure, this forest has plentiful precipitation (about 3000 mm) and very low
- 729 elevation, both of these abiotic factors may cause the site as CH4 source. Thus, the authors are
- 730 suggested to provide some more discussion about (the lack of) spatio-temporal variation in CH4
- 731 flux.
- 732
- 733 **Response:** Again here, this result can be easily explained by the time frame of the study. Tropical
- 734 soils are generally considered as sink at a yearly basis but much study show that there are
- seasonal variation in CH₄ fluxes and that tropical soils tend to shift from a sink in the dry season 735
- 736 to a sources during the wet season. Here, a four months period is limited to go deep into such
- 737 relationships, especially in tropical forest where temporal and spatial variability of fluxes are 738
- high. You can find below a figure that can now be found in the supplementary material of the 739
- manuscript displaying the relationship of the three gases with soil moisture. Nonetheless, going
- deeper in the discussion of the effect of rain event, soil moisture and the relative importance of 740 spatial, seasonal and diurnal variability of fluxes cannot be done with these dataset that was 741
- specifically constructed to demonstrate the feasibility of running the system under tropical
- 742 743 conditions.
- 744
- 745 4. Appendix Figure A1:
- 746 This figure shows a very general (basic) chamber-problem for measurement of soil GHGs fluxes.
- Long closure time will cause higher GHGs concentration (if the soil is GHGs source) or lower 747
- GHGs concentration (if the soil is GHGs sink) inside the chamber, which will induce 748
- 749 underestimation of GHGs flux (saturation effect). Saturation effect is generally positively
- associated with both flux rate and ratio of the effective chamber volume to the measured soil 750 surface area. Empirically, I believe the 2-mintute 751
- 752 closure time is enough for measurement of both CO2 and CH4 flux in tropical forests, even for
- most temperate and boreal forests. For Li-Cor soil chamber (8100-104), the ratio is (0.0040761/ 753
- 754 0.03178=0.12826 m) = 12.3 cm. However, for many of the custommade soil chambers, the ratio
- 755 is generally higher than 12.3 cm, thus this is might be the specific problem (issue) only for Li-
- 756 Cor soil chamber (8100-104). I suggest the authors feedback this problem to Li-Cor and suggest
- 757 Li-Cor to draw this problem to their instrument user manual.
- 758
- 759 **Response:** Thank you for this feedback. Following comments from the other reviewers, we used
- exponential fit for estimating all fluxes which improved this saturation issue. Also, as stated in 760
- 761 the manuscript, we always used 2 minutes estimation for CO_2 fluxes to overcome this issue.
- 762
- 763 5. Also for Appendix Figure A1:
- 764 The authors are suggested to re-draw the Appendix Figure A1 indicating different symbols
- 765 (or color) for each of the four chambers.
- 766

Response: Following comments from the other reviewers, this figure has been moved to the main
text and now also include N₂O. We decided to use different colours (black and grey) for the two
distinct weeks that were used for this comparison instead that different colours for the different
chambers because it allows a better view of the fact that these two weeks are covering almost the

- 771 whole range of fluxes that can be encountered in the site.
- 772
- 773 6. Closure time:

When compared Table 1 with Table 2, the closure time of 10 minutes for measurement of N2Oflux was enough. Thus, the Table 1 is suggested to be deleted.

776

Response: We disagree with this comment. A closure time of 10 minutes would have led to a
MDF of 0.009 instead of 0.002. In this case, only 82% instead of 96% of the fluxes would have
been considered of reliable. We therefore decided to maintain Table 1 in the manuscript as it
allows to show that a MDF of 0.002 can only be achieved with a 25 minutes closure time.

- 781
- 782 7. Additional suggestion 1:
- 783 To prove the data quality or measurement precision, the authors are suggested to
- add a plot showing changes in CO2, CH4 and N2O concentrations in the chambers.
- 785 Following is a sample plot (Sample Fig).
- 786
- *Response:* We think that this information is not adding much to the comprehension of the study
 and it was not included in the new version of the manuscript.
- 789
- 790 8. Additional suggestion 2:
- As I mentioned in the above, this may be the world's first report about in situ and simultaneously

measurement of soil CO2/CH4/N2O fluxes at low latitude (between 10_N and 10_S). I believe

- this paper will be a potential high citation rate if the authors can give some more discussion
- about spatio-temporal variation in CO2/CH4/N2O fluxes and their control factors. For example,
- the coefficient of variation (CV) was used to represent the spatial variation. CV of Rs can be calculated by $CV = (SD/(mean Rs))_{100}$.
- 797
- **Response:** Mean and SD per chambers are available in Table 2 and we added a figure with mean value of each chamber per days for the three gases allowing to visualize the spatio-temporal variability of fluence.
- 800 *variability of fluxes.*
- 801

802 803	Automatic high-frequency measurements of full soil greenhouse gas fluxes in a tropical forest
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823	

824	Abstract. Measuring <i>in situ</i> soil fluxes of carbon dioxide (CO ₂), methane (CH ₄), and nitrous oxide
825	(N ₂ O) continuously at high frequency requires appropriate technology. We tested the combination
826	of a commercial automated soil CO $_2$ flux chamber system (LI-8100A) with a CH $_4$ and N $_2O$
827	analyzer (Picarro G2308) in a tropical rainforest for 4 months. A chamber closure time of 2 minutes
828	was sufficient for a reliable estimation of CO_2 and CH_4 fluxes (100% and 98.5% of fluxes were
829	above Minimum Detectable Flux - MDF, respectively). This closure time was generally not
830	suitable for a reliable estimation of the low N_2O fluxes in this ecosystem but was sufficient for
831	detecting rare major peak events. A closure time of 25 minutes was more appropriate for reliable
832	estimation of most N ₂ O fluxes (85.6% of measured fluxes are above MDF \pm 0.002 nmol m ⁻² s ⁻¹).
833	Our study highlights the importance of adjusted closure time for each gas.
834	
835	
836	Keywords: Soil fluxes, French Guiana, methods, optimizing, Minimum Detectable Fluxes
837	

1. Introduction

840 Carbon dioxide (CO₂), methane (CH₄) and nitrous oxide (N₂O) are the three main 841 greenhouse gases (GHGs) in terms of radiative forcing. Increases in these GHG concentrations in 842 the atmosphere is driving anthropogenic global warming. Understanding the magnitude of GHG fluxes in natural ecosystems has recently become a priority in the study of GHG balances (Merbold 843 844 et al., 2015). Tropical intact forests cover 1392 Mha globally and represent about 70% of the total 845 tropical forest area (1949 Mha), which accounts for the largest area of global forest biomes (~50%). Very few reliable long term datasets on full GHG balances are available from tropical 846 847 ecosystems, despite their known importance for the global cycles of these three GHGs (Dutaur and Verchot, 2007). This is in part due to the challenges of designing and operating continuous, multi-848 849 gas flux analysis systems in tropical forests. Soil processes in particular are responsible for an 850 important part of GHGs that are produced or consumed in tropical ecosystems (Oertel et al., 2016). Soil physical, chemical, and biological characteristics are linked to variation in GHG emissions 851 852 from soils, which in turn can display very high spatial and temporal variability (Arias-Navarro et al., 2017; Silver et al., 1999). 853

854 Historically, soil GHG fluxes (emission or consumption) have been measured using the 855 static chamber method. This involves closing chambers manually for a known period of time, usually 30-60 minutes, and repeated collection of air samples for further analysis via gas 856 chromatography (Verchot et al., 1999, 2000). Fluxes are then computed from the increase (or 857 858 decrease) in gas concentration per unit time, per surface area enclosed by the chamber, and 859 corrected by the volume of the chamber. While these labor-intensive and time-consuming manual 860 measurements are well adapted to capture high spatial flux variability (Arias-Navarro et al., 2017; Pumpanen et al., 2004), they do not capture high temporal variation, which is necessary for the 861

accurate estimation of annual GHG budgets. Moreover, short term, transient spikes in the emission
or consumption of these GHGs likely remains undetected with static chamber methods, imposing
a lost opportunity to fully understand the production or consumption processes of GHGs and their
response to rapidly changing environmental conditions. One of the key challenges of contemporary
GHG flux research is to close these knowledge gaps in order to improve the quantitative prediction
of GHG fluxes (Merbold et al., 2015).

The use of automatic chambers is one approach to obtain continuous estimation of soil 868 GHG flux data at high temporal frequency (several measurements per days) at various sampling 869 870 points. Since the 1970s (Denmead, 1979), a variety of technical solutions for automated flux sampling have been developed (Ambus et al., 2010; Breuer et al., 2000; Görres et al., 2016; 871 Kostyanovsky et al., 2018; O'Connell et al., 2018; Petrakis et al., 2017a; Savage et al., 872 873 2014)(Ambus et al., 2010; Breuer et al., 2000; Görres et al., 2016; Kostyanovsky et al., 2018; 874 Petrakis et al., 2017a; Savage et al., 2014), particularly for soil CO₂ fluxes. However, accurate detection of CH₄ and N₂O fluxes from soils via flow through systems is more difficult than CO₂ 875 876 due to significantly lower background concentrations and lower flux rates (Kostyanovsky et al., 877 2018). The budgetary requirements for large infrastructure and intensive maintenance as compared 878 to manual chamber measurements have prevented the widespread application of automated 879 systems. The use of automated and continuous methods to estimate full GHG budgets in situ remains scarce, especially in complex biomes with extreme climate such as tropical forests. 880 881 Therefore, only a few studies actually address the difficulties and challenges associated with operating these systems under field conditions (Görres et al., 2016; Koskinen et al., 2014). 882

Recent technological advances have now made more automated chamber systems
commercially available, and an increasing number of custom-made systems are being designed

and deployed for soil GHG flux measurements <u>(De Klein and Harvey, 2012) [17,18]</u>. Here, we present a detailed field deployment of a custom built, automated soil GHG flux system – the LI-8100A Soil CO₂ Flux System (LI-COR Biosciences Inc., Lincoln, NE, USA) running in line with a Picarro G2308 (Picarro Inc., Santa Clara, CA, USA).—Using a 4 month dataset of continuous measurements of CO₂, CH₄, and N₂O fluxes simultaneously under tropical forest conditions, we present an optimized sampling protocol for the estimation of the full GHG budget in this ecosystem.

892

893 **2. Methods**

894

2.1. Measurement site

This study was conducted in the Paracou research station (5°15'N, 52°55'W), located in the coastal area of French Guiana, South America. The automated soil GHG flux system was deployed in the footprint of the Guyaflux site, which holds a 55 m-tall tower upon which canopy CO_2 , H₂O and energy fluxes have been monitored since 2004 using the eddy covariance technique (Aguilos et al., 2018; Bonal et al., 2008). The site is covered with tropical pristine forest and located in the northernmost part of the Guiana shield. It is characterized by a succession of small, elliptical hills rising to 10–40 m a.s.l., sometimes associated with plateaus of similar altitude.

The soils are mostly nutrient-poor acrisols (FAO-ISRICISSS, 1998) with pockets of sandy ultisols developed over a Precambrian metamorphic formation called the 'Bonidoro series', and composed of schist and sandstone, sporadically traversed by veins of pegmatite, aplite and quartz (Bonal et al., 2008). The forest around the tower is characteristic of a tropical pristine forest with both high tree density (~ 620 trees with a dbh>10 cm ha⁻¹) and species richness (~ 140 species ha⁻¹). The climate is highly seasonal due to the north/south movement of the Inter-Tropical Convergence Zone. The wet season, characterized by heavy rain events, lasts for 8 months (December–July) and alternates with a 4 month dry period (August–November) during which precipitation is typically lower than 100 mm per month. For the period 2004-2015, annual rainfall quantities were on average 3103 mm year⁻¹, relative extractable water (an index of soil water availability (Wagner et al., 2011)) varied from 0.93 in the wet season to 0.46 in the dry season and soil temperature was on average 25.1 with little seasonal nor diurnal variation (Aguilos et al., 2018).

915

2.2. Automated sampling system

916 A schematic view of the automatic sampling system is shown in Figure 1(A). The system consisted 917 of four main components: sixteen automated long-term chambers (8100-104, LI-COR 918 Biosciences), a multiplexer to link one chamber at a time to the gas analyzers (LI-8150, LI-COR 919 Biosciences), an infrared gas analyzer (IRGA) to measure CO₂ concentrations (LI-8100A, LI-COR 920 Biosciences), and a cavity ring down spectroscopy (CRDS) instrument to measure CH₄ and N₂O 921 concentrations (G2308, Picarro) that was fitted with an external recirculation pump (A0702, 922 Picarro). Both the IRGA and CRDS systems were necessary to measure all three GHG 923 concentrations due to the different abundances and flux rates of CO₂, CH₄ and N₂O. The IRGA 924 methodology is accurate and precise enough to detect small CO_2 concentration changes at high 925 background concentrations (approximately 400 ppmv; parts per million in volume units). However, the detection of small changes in CH₄ and N₂O concentrations, even at their low 926 927 background atmospheric concentrations in the order of 2000 ppbv (ppbv=parts per billion in 928 volume units) and 300 ppby, respectively, requires higher accuracy and precision levels that can 929 only be detected with the CRDS.



932 Figure 1: Experimental Design: (A) Schematic view of the installation composed of four main components: sixteen automated long-term chambers (8100-104, LI-COR Biosciences), a 933 934 multiplexer to link one of these chambers to the gas analyzers (LI-8150, LI-COR Biosciences), an 935 infrared gas analyzer (IRGA) to measure CO₂ concentrations (LI-8100A, LI-COR Biosciences), 936 and a cavity ring down spectroscopy (CRDS) instrument to measure CH4 and N2O concentrations (G2308, Picarro) that was fitted with an external pump. (B) Schematic representation of the grid 937 938 with the shelter housing the equipment in the middle and the 16 chambers (grey dots) linked to the Li-8150 multiplexer with 15 meters cables (black lines). (C) Picture of the instruments in the field. 939 940

941 Power supply was delivered through a 12 kVa generator (Perkins STORM15) fitted with 942 batteries located 400 m away from the instruments. Both the CO₂ analyzer control unit and the 943 multiplexer (LI-COR) had their own weather-proof casing, requiring no additional protection in 944 the field. Nonetheless, in consideration of the high precipitation at the site, these devices were placed under a wooden shelter for added protection. The CH₄ and N₂O analyzer (Picarro), its 945 external pump and a computer monitor were housed in a waterproof shelter that was specifically 946 947 designed to host them (Figure 1(C)). The Li-8100 and the G2308 computers were connected through ethernet connection to ensure time synchronization. The sixteen automated soil chambers 948 (8100-104, LI-COR Biosciences) were installed in a grid in the forest (Figure 1(B)) covering in 949 total an area of approximately 300 m² (15 m x 20 m). Each chamber was only closed during 950 individual chamber measurement periods, and was fully open when not sampling. The PVC collars 951 952 that were provided with the 8100-104 automatic chambers were inserted in the soil one month 953 prior to the first measurement (20.3 cm inner diameter/21.3 cm outer diameter; enclosed soil area \sim 318 cm²; insertion depth \sim 7cm; offset \sim 4cm; green PVC). When the chambers close, they are 954 955 automatically lowered so that they cover each soil collar and ensure a fully sealed chamber. The 956 chamber lid does not directly rest on the collar rim, but on a metal plate surrounding the collar, 957 leaving the collar undisturbed and minimizing lateral leaks (Hupp et al., 2009).

The 16 chambers were connected via 15 m Bev-a-line tubing (8 mm inner diameter) with the multiplexer (LI-8150), which allows for switching between each of the 16 chambers in any given sequence. Soil temperature (at a depth of 10 cm) was monitored with 8100-201 Ω thermistor probes (Omega Engineering Inc., Stamford, CT, USA), and soil volumetric water content (0-10 cm) was monitored with 8100-202 ECH₂O Model EC-5 soil moisture sensors (Decagon Devices 963 Inc., Pullman, WA, USA). Soil temperature and soil volumetric water content were directly 964 connected to the chambers and recorded by the Licor system using the same time step.

Each chamber was purged for 15 sec prior to each measurement and 45 sec after each 965 966 measurement in order to flush the lines and restore background gas levels in the system. The flow rate during the purging and the measurements was $\sim 2.8 \text{ Lmin}^{-1}$ between the Li-8150 and the 967 chambers, which ensures sufficient air mixing in the chamber headspace during the measurements 968 969 (Görres et al., 2016). Flow rates in the subsampling lines (Li8100 and Picarro) were lower and set between 1.5 and 1.7 L min⁻¹ as recommended by the manufacturers. The LI-8100 software 970 provided the rate of CO₂ concentration increase in the chamber which was used to quantify the 971 flux of CO₂ from the soil surface into the atmosphere (taking into account the enclosed soil surface 972 area and the total system volume). A subsampling loop was inserted after the analyzer (LI-8100A) 973 974 and before the multiplexer (LI-8150), to pull the air sample through the Picarro G2308 CRDS analyzer for the determination of CH₄ and N₂O concentrations and flux estimations, before going 975 976 back to the chamber (Figure 1(A)). All three gas concentrations were recorded every second over 977 the sampling periods.

978

2.3. Flux calculations

979 All fluxes estimation were done by using commercially available Soil Flux pro software (LI-COR 980 Biosciences). An R script (Supplementary file 1) was created to merge all the Picarro files from a given week in order to import them into the Soil flux Pro software. The Picarro creates one file per 981 982 hour and when Picarro files are not merged, Soil flux Pro software is not able to deal with 983 measurements overlapping between two distinct Picarro files (e.g. when a single measurement is done from 9:50 am to 10:15 am) leading to incorrect estimation of CH₄ and N₂O fluxes. CO₂, CH₄ 984 985 and N₂O fluxes were measured as linear changes inexponential fit of gas concentration with time

- 986 using Soil flux Pro software and include a 60 sec dead band to account for soil surface pressure
- 987 disturbances due to the closing of the chamber.

988 **Table 1:** Minimum Detectable Fluxes (MDF) for each gas and for closure times from 2 to 30 989 minutes. The two closure times that were used in this study (2 minutes and 25 minutes) are 990 highlighted in bold.

Closure time (minutes)	$N_2O \text{ (nmol } m^{-2} \text{ s}^{-1}\text{)}$	$CH_4 (nmol m^{-2} s^{-1})$	$CO_2 (nmol m^{-2} s^{-1})$
2	0.100	0.040	2.393
5	0.025	0.010	0.605
10	0.009	0.004	0.214
15	0.005	0.002	0.117
20	0.003	0.001	0.076
25	0.002	0.001	0.054
30	0.002	0.001	0.041

992 **2.4. Minimum Detectable Fluxes**

The minimum detectable flux (MDF) for each gas was estimated by using a metric originally developed by Christiansen et al. (2015), which was modified by Nickerson (2016) to make it more suitable for high-frequency measurements (Christiansen et al., 2015; Nickerson, 2016):

996
$$MDF = \left(\frac{A_a}{t_c \sqrt{n}}\right) \left(\frac{VP}{SRT}\right)$$

Where Aa is the analytical accuracy of the analyzer (25 ppb for N₂O and 10 ppb for CH₄ with the Picarro G2308 and 600 ppb for CO₂ with the Li8100, recorded from the technical data sheets of the analyzers), t_c is the closure time of the chamber in seconds, n is the number of points that are available to compute the flux (i.e. t_c divided by the sampling periodicity, every 1 second in this study), V is the chamber volume (0.0040761 m³), P is the atmospheric pressure (101325 Pa), S is the chamber surface area (0.03178 m²), R is the ideal gas constant (8.314 m³ Pa K⁻¹ mol⁻¹) and T is the ambient temperature (298.15 K). We computed the MDF of each gas for closure times from
2 minutes to 30 minutes in order to select the optimal chamber closure time for each gas in our
integrated system (Table 1).

1006

2.5. Closure time

1007 Selecting the best length of time for soil GHG measurements and accurate flux calculation in an 1008 integrated CO₂, CH₄ and N₂O automated measurement system requires careful consideration. At 1009 low fluxes, longer measurement periods are needed to reach reliable measurements of real 1010 concentration changes, while at high fluxes possible storage and saturation effects in the chamber 1011 headspace might result in non-linear concentration increases and thereby underestimated fluxes. 1012 In order to maximize the detectable percentage of fluxes for N₂O and CH₄ without impeding spatial 1013 coverage and temporal resolution, we built a combined program with two different closure times. 1014 Each week, four out of sixteen chambers were programmed to stay closed for a longer 1015 measurement period to ensure a reliable estimation of low fluxes while the other twelve chambers 1016 were programmed to stay closed for a shorter period to capture diel variation and detect high 1017 fluxes. For the short closure time (SHORT hereafter), we used a 2 minute measurement period 1018 because (1) this is a standard closure time for soil CO_2 flux calculations (Janssens et al., 2000), 1019 even in tropical forests (Epron et al., 2006; Sayer et al., 2007) of this region where MDF for CO₂ 1020 flux is typically low (Bonal et al., 2008; Bréchet et al., 2009; Courtois et al., 2018), (2) corresponding MDFs of CH₄ (0.04 nmol $m^{-2} s^{-1}$ or) and N₂O (0.1 nmol $m^{-2} s^{-1}$) are compatible 1021 1022 with the detection of emission or consumption peaks of these two gases in this region (Courtois et 1023 al., 2018; Petitjean et al., 2015). For the long closure time (LONG hereafter), we decided to use a 1024 25 minute measurement period in order to optimize the trade-off between a reliable estimation of 1025 low N₂O fluxes (Table 1) and a program length that allows for a sufficient number of flux
1026 measurements per chamber and per day.

We therefore programmed the multiplexer for 2.5-h cycles (9-10 measurements per chamber per day), which included four chambers with LONG measurements and twelve chambers with SHORT measurements. Each week, the program was modified manually so that the four LONG measurements were rotated across the chambers. Each chamber was therefore measured with the LONG closure time for one 7 consecutive day period per month (4 weeks).

1032 **2.6. System maintenance and data processing**

The automated sampling system was installed on June 1st 2016 and operated until September 29th 1033 1034 2016 (4 months), totaling 17652-17592 individual measurements for each gas (4326-4098 with 1035 LONG closure time and 13326-13494 with SHORT closure time). Coarse wood debris was were 1036 removed weekly but small litter, such as leaves, fruits, and twigs, was left in the collar area. Every 1037 week, living plants growing inside the collars, and the dead leaves on the chambers, were carefully 1038 removed by hand. The R² value of the exponential increase of -CO₂ over 2 minutes was used as an 1039 indicator that the system was functioning correctly and not impeded by debris (Görres et al., 2016; 1040 Savage et al., 2014). When the R^2 of the regression between time and CO_2 concentration was lower than 0.9, we considered this as an indication that there may have been an issue with the chamber 1041 1042 closing and sealing correctly and removed the flux measurement for all three gases from our 1043 analysis.

For CO₂, we observed a strong concentration saturation effect when using the LONG closure time (25 minutes), leading to an underestimation of fluxes (Supplementary Figure 1Figure 2). All CO₂ flux estimates were therefore based on 2 minute regressions only, using either full concentration measurements of the SHORT closure time or the 2 first minutes of the LONG

closure time. The R² value of the linearly increasing CO₂ over 2 minutes was used as an indicator 1048 1049 that the system was functioning correctly and not impeded by debris (Görres et al., 2016; Savage et al., 2014). When the R^2 of the regression between time and CO_2 concentration was lower than 1050 1051 0.9, we considered this as an indication that there may have been an issue with the chamber closing and sealing correctly and removed the flux measurement for all three gases from our analysis. 1052 Following recommendations (Rubio and Detto, 2017), we removed anomalous values, i.e. CO₂ 1053 fluxes estimation with a difference greater than 5 μ mol m⁻² s⁻¹ between with previous or following 1054 measurements or lower than 0 μ mol m⁻² s⁻¹. For CH₄, we observed only a slight saturation effect 1055 when using the LONG closure time (Supplementary Figure 24). Variation in the flux calculations 1056 did not differ between the SHORT and LONG chamber closure measurements. All fluxes above 1057 or below (for negative fluxes) MDF were considered as reliable and were calculated using the full 1058 1059 data available (2 minutes measurement for the SHORT closure time and 25 minutes for the LONG closure time). N₂O flux calculations were much more variable when measuring with the SHORT 1060 closure time compared to the LONG closure time- (Supplementary Figure 22). Even if fluxes were 1061 1062 above the detection limit, the low fluxes estimated with the SHORT closure time were not reliable as shown by the low correlation in Figure 2. For both CH₄ and N₂O, wWe therefore decided to 1063 apply the following quality check procedure (1) to keep all All the fluxes estimated with the LONG 1064 elosure time that were not complying with MDF criterion even with low R^2 for the regression 1065 1066 between time and N₂O concentrations and were discarded. (2) <u>All to consider all</u> fluxes estimated with the SHORT closure time with a R^2 lower than 0.8 as unreliable were discarded (Savage et al., 1067 2014). (3) We applied the same procedure than for CO_2 regarding anomalous values (difference 1068 greater than 5 nmol $m^{-2} s^{-1}$ between consecutive measurements). 1069



Figure 2: Comparison of 2 minutes and 25 minutes estimations for (A) CO₂ (B) CH₄ and (C) N₂O fluxes. For this, we used measurements made over 25 minutes and recomputed the flux with the two firsts minutes for two weeks (from August 2nd for August 9th in black and from August 16th for August 25th in grey) covering the whole range of fluxes during the study period. All fluxes were computed using exponential fit. The dashed line represent the 1:1 line while the solid grey line represents the linear regression between 2 minutes and 25 minutes estimations (R² of these regressions are indicated on each panel).

1078 **3. Results and discussions**

A cleaning frequency of once a week was necessary and sufficient to remove falling leaves and 1079 1080 branches from the automatic chamber system, prevent leaks and generate a continuous dataset of 1081 soil GHG fluxes from this tropical forest. Temperature variations are typically small below the canopy due to the shadowing by dense canopy crown and microclimatic conditions. During the 1082 1083 study period, temperature at 2m height varies during the day from 22 °C in the night to 28 °C 1084 during the day. The presence of water condensation inside the tubing lines was carefully checked 1085 every week and never occurred during the study period. The automatic chamber system worked well most of the time, but some data gaps did exist. Over the 17724-17592 individual flux 1086 estimations, $\frac{276-343}{1.95}$ %) had to be discarded because of (1) problems in the connection 1087

1088 between the chamber and the multiplexer ($\frac{196-154}{154}$ measurements, $\frac{0.9}{1}$ % of data points); (2)

1089 imperfect chamber closing, which was detected by an insufficient increase of CO_2 (103–189)

1090 measurements, 10.6% of data points).

1091

Table 2: Mean, standard deviation (SD), minimum (Min) and maximum (Max) values of each gas
 and each chamber over the study period. <u>These values are computed using all fluxes estimation</u>
 (either with SHORT or LONG closure time) remaining after quality check. The number (N) of
 fluxes that were used is also indicated for each chamber. The last line of the table is the mean of
 all fluxes by chambers by gas and the min and max for all chambers by gas.

	$CO_2 \ (\mu mol \ m^{-2} \ s^{-1})$					$CH_4 \text{ (nmol m}^{-2} \text{ s}^{-1}\text{)}$			$N_2O \ (nmol \ m^{-2} \ s^{-1})$						
	Mean	Sd	Min	Max	Ν	Mean	Sd	Min	Max	Ν	Mean	Sd	Min	Max	Ν
Chamber 1	7.19	0.93	2.14	10.81	940	10.97	7.73	-2.08	28.79	840	0.10	0.12	-0.48	0.70	284
Chamber 2	7.60	1.11	4.00	12.21	1166	-1.62	1.75	-4.09	11.68	899	0.00	0.14	-1.03	0.75	285
Chamber 3	5.58	0.99	2.11	11.12	1135	0.35	2.95	-2.48	22.94	745	0.03	0.23	-0.61	2.85	208
Chamber 4	7.94	1.37	4.36	12.13	1154	-1.85	1.23	-3.63	6.09	1105	0.04	0.10	-0.66	0.60	224
Chamber 5	4.14	0.92	0.53	10.05	1139	1.37	3.26	-2.20	12.61	752	0.15	0.33	-1.04	3.23	382
Chamber 6	8.87	1.70	3.36	17.68	1070	-1.38	1.78	-3.20	8.04	801	-0.02	0.12	-1.04	0.63	272
Chamber 7	13.47	2.78	0.89	22.12	988	1.37	3.60	-2.63	19.56	749	0.64	1.37	-0.85	7.93	216
Chamber 8	7.44	1.19	2.03	11.02	1099	0.03	2.96	-3.37	18.47	785	0.02	0.15	-1.36	0.84	202
Chamber 9	4.25	1.20	0.44	11.37	1002	2.06	3.13	-2.14	11.53	879	0.02	0.11	-0.62	0.58	332
Chamber 10	5.60	1.30	0.69	13.13	1037	1.21	2.46	-1.91	10.34	657	0.04	0.13	-0.64	0.77	252
Chamber 11	11.97	2.19	6.84	18.78	1004	6.72	7.61	-1.06	41.49	855	0.03	0.17	-1.01	1.04	199
Chamber 12	9.42	2.70	3.45	21.54	968	1.40	6.68	-3.29	41.94	891	0.02	0.09	-0.75	0.30	204
Chamber 13	5.85	1.34	0.42	8.49	944	5.29	5.92	-4.60	26.64	654	0.10	0.19	-0.84	1.71	335
Chamber 14	5.66	1.15	0.72	10.72	987	2.78	6.22	-2.48	35.15	691	0.09	0.17	-0.63	0.93	231
Chamber 15	16.63	3.27	9.42	29.64	850	-0.46	2.05	-3.25	8.26	839	-0.02	0.16	-0.96	0.72	185
Chamber 16	7.35	1.13	3.98	11.37	994	-1.34	1.48	-3.60	6.11	843	0.00	0.11	-1.00	0.83	187
	8.06	1.58	0.42	29.64	16477	1.68	3.80	-4.60	41.94	12985	0.08	0.23	-1.36	7.93	3998



1099

Figure <u>32</u>: Distribution of fluxes: Histogram of (A) CO₂, (B) CH₄ and (C) N₂O fluxes over the study period. For (B) and (C), the dotted line represents null fluxes.

1103 **3.1. CO₂ fluxes**

1104 Additionally to the 343 fluxes than were removed after the firsts steps of quality check procedure, 758 CO₂ fluxes estimations were also considered as anomalous, either because the difference with 1105 previous or following measurements where greater than 5 μ mol m⁻² s⁻¹ (758 measurements, i.e. 1106 4.3%) or because they were lower than 0 μ mol m⁻² s⁻¹ (14 measurements). In total, 16477 CO₂ 1107 fluxes over 17592 (93.6%) can be used over the four month period. CO_2 fluxes were on average 1108 $87.10634 \pm 13.621 \mu$ mol m⁻² s⁻¹-(Table 2) which would correspond to a mean annual soil CO₂ 1109 efflux of 3050 gC m⁻² year⁻¹ which falls into the upper range of the extensive review of mean 1110 annual soil CO₂ effluxes estimations in tropical forest provided recently by Rubio and Detto 1111 (2017). Nonetheless, our study period (June-September) only covered the end of the wet season 1112 and more data are needed to precise this estimation. with a high variability among chambers (Table 1113 2). The minimum flux measured during the study period was 0.19 µmol m⁻² s⁻¹ (Table 2) and the 1114

majority of fluxes were between 3 to 10 μ mol m⁻² s⁻¹ (Figure 2). All two-minute measurements of 1115 1116 CO_2 fluxes from the four-month study period were therefore above the MDF of 2.39 nmol m⁻² s⁻¹ 1117 for the LI8100 analyzer (Table 1). No saturation effect was detected using the SHORT closure 1118 time and estimation of CO_2 over a shorter time period is not recommended (Davidson et al., 2002). 1119 CO₂ fluxes using the LONG closure time would be underestimated due to the buildup of high CO₂ 1120 concentrations due to large fluxes over this long time period (Supplementary Figure 1), 1121 and are not recommended. For small chambers as the one that were used in this study, wWe 1122 therefore conclude that a 2 minute sampling time should be used for CO₂ flux calculations since 1123 the MDF of this short measurement period allowed for the retention of 100% of the data. When the chambers stay closed longer for accurate detection of N₂O fluxes, only the first two minutes of 1124 1125 data should be used for CO₂ flux calculations.

The use of 16 automated flux chambers allowed for the capture of spatial and temporal variability 1126 1127 of soil respiration. Over this four month period, corresponding to the end of the wet season in French Guiana, -temporal variability remained low (Figure 4). This dataset is therefore not long 1128 1129 enough to detect seasonal variation of soil respiration that were highlighted in previous study (Rowland et al., 2014; Rubio and Detto, 2017). We did found that soil respiration tended to 1130 decrease in very humid soils (Supplementary Figure 1) as highlighted previously at the same site 1131 1132 (Rowland et al., 2014) but more data are needed to disentangle precisely the importance of seasonal and diurnal variability from the responses to environmental triggers on soil respiration. 1133 1134 Nonetheless, even during this relatively short period, our data clearly demonstrated a strong spatial 1135 variability of soil respiration, even at a low spatial scale (Figure 5, Table 2), some local spots clearly displaying stronger values of soil respiration during the study period. which is needed to 1136 1137 constrain ecosystem carbon budgets (Figure 3).



Figure 43: CO₂ fluxes through time: CO₂ fluxes for each chamber (1 to 16) over the study period1140with fluxes estimated with SHORT (2 minutes) closure time in black and fluxes estimated with

1141 the 2 first minutes of the LONG (25 minutes) closure time in grey. All panels have the same limits

1142 on the y axis (from 0 to 25 μ mol m⁻² s⁻¹)





145Figure 5: Mean values per days for (A) CO2, (B) CH4 and (C) N2O fluxes over the study period.146Each chamber is represented by a distinct color.

1148 **3.2.CH4 fluxes**

1149 Additionally to the 343 fluxes than were removed after the firsts steps of quality check procedure, CH₄ fluxes estimations were also discarded because of (1) problems with Picarro files (12 1150 1151 measurements), (2) application of the MDF criterion (137 measurements), (3) application of the 1152 R^2 criterion for SHORT closure time (3751 measurements, i.e. 28% of the SHORT measurements) 1153 (4) detection of anomalous values (364 measurements). In total, 12985 CH₄ fluxes over 17592 1154 (73.8%) can be used over the four month period. No saturation effect was detected using the LONG 1155 closure time and fluxes estimated with the SHORT closure time were very well correlated to fluxes 1156 using the LONG closure time, even for small fluxes (Figure 2).- CH4-fluxes were on average 1.06 1157 ± 4.52 nmol m⁻² s⁻¹ with a high variability among chambers (Table 2). Minimum detectable fluxes for CH₄ were \pm 0.04 nmol m⁻²s⁻¹ using the SHORT closure time and \pm 0.001 nmol m⁻²s⁻¹ using the 1158 1159 LONG closure time (Table 1). 98.568.4 % and 9998.29% of fluxes measured with the SHORT and 1160 LONG closure times, respectively, were retained in our quality control data processing over the 1161 four-month study period. These measurement periods, therefore, allowed for the retention of a 1162 large majority of CH₄ emission or consumption fluxes in our data analysis. -CH₄ fluxes were on average 1.706 ± 34.852 nmol m⁻² s⁻¹ with a high variability among chambers 1163 1164 (Table 2) but <u>T</u>the frequency of negative CH₄ fluxes (consumption, 59% of fluxes) was greater

than positive fluxes (emission, <u>41% of fluxes</u>) during this period (Figure <u>32</u>). Most of the time, soils were either consuming or emitting small amounts of CH₄, but transient, large emission peaks were periodically detected at individual chamber locations during the study period (Figure <u>64</u>).
<u>Tropical soils are generally considered as sink at a yearly basis (Dutaur and Verchot, 2007) but it</u> is known that these soils can shift from a source in the wet to a sink in the dry season (Courtois et 169).





1177

1178Figure 64: CH4 fluxes through time: CH4 fluxes for each chamber (1 to 16) over the study period1179with fluxes estimated with SHORT (2 minutes) closure time in black and fluxes estimated with1180LONG (25 minutes) closure time in grey. The dotted line displays the zero flux line. All panels1181have the same limits on the y axis (from -5 to 30 nmol m⁻² s⁻¹)1182

1183 **3.3.** N₂O fluxes

Additionally to the 343 fluxes than were removed after the firsts steps of quality check procedure,
 N₂O fluxes estimations were also discarded because of (1) problems with Picarro files (12
 measurements), (2) application of the MDF criterion (1594 measurements), (3) application of the

R² criterion for SHORT closure time (11643 measurements, i.e. 28% of the SHORT 1187 1188 measurements) (4) detection of anomalous values (364 measurements). In total, 3998 N₂O fluxes over 17592 (22.7%, 140 measurements with the SHORT and 3858 measurements with the LONG 1189 1190 closure time) can be used over the four month period. 94.1% of fluxes measured with the LONG 1191 closure times were retained after our quality control data processing over the four-month study period. N₂O fluxes were on average 0.038 ± 0.537 nmol m⁻² s⁻¹ with a high variability among 1192 1193 chambers (Table 2). Most N₂O fluxes (85.6%) with the LONG closure time were above or below MDF (± 0.002 nmol m⁻²s⁻¹, Table 1) and varied between -2 to 2 nmol m⁻²s⁻¹ (Figure 2). When 1194 measured over 25 minutes, N₂O fluxes in our site could therefore be considered as reliable. Using 1195 1196 the SHORT closure time, most flux estimations had to be discarded because they led to unreliable 1197 flux estimations (Supplementary FFigure 2). Only 7.9% of the measurements using the SHORT 1198 elosure time were retained after our quality control checks for the N2O flux data. Nonetheless, the 1199 high frequency of the SHORT closure times still allowed the detection of one high N₂O emission or consumption events (up to 15 nmol m⁻²s⁻¹) that wereas detected during the study period (Figure 1200 1201 5, chamber 75 and 7). N₂O fluxes were on average 0.1 ± 0.2 nmol m⁻² s⁻¹ with a high variability among chambers (Table 1202 2). At the same chamber, N2O flux can shift from consumption to emission with 28% of fluxes 1203 1204 indicating a sink and 72% a source for N₂O (Figure 3). The high variability in N₂O fluxes that we detected over four months with our automated system are in agreement with the typical high 1205 1206 variability in N₂O fluxes measured from tropical soils over space and time using static chambers 1207 (Arias-Navarro et al., 2017; Courtois et al., 2018). Moreover, N₂O fluxes didn't show any 1208 relationship with surface soil humidity (Supplementary Figure 1), which underline the complexity

1209 of the biological process underlying these fluxes. In a previous study in the same environment





Figure 57: N₂O fluxes through time: N₂O fluxes for each chamber (1 to 16) over the study period1218with fluxes estimated with the SHORT (2 minutes) closure time in black and fluxes estimated with1219the LONG (25 minutes) closure time in grey. The dotted line displays the zero flux line. Due to1220the high differences among chambers, each panel has specific limit on the y axis.1221

4. Conclusions

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Our unique system coupled a Li8100 CO_2 analyzer and multiplexor with a Picarro G2308 CH_4 and N<sub>2</sub>O analyzer to sample 16 automated soil flux chambers with a rotation of SHORT and LONG closure times for the accurate monitoring of three GHG fluxes over four months with high spatial and temporal resolution. The sampling system of SHORT and LONG closure times with a weekly
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1227 rotation presented here has three major advantages, which ultimately can provide high confidence 1228 in the estimation of annual the full GHG budgets of tropical soils: (1) the LONG closure time 1229 allows a reliable estimation of the low N_2O fluxes in this ecosystem, which was clearly not 1230 achieved using a shorter closure time, (2) the number of data points per day are sufficiently high 1231 (9 to 10 measurements per day) to capture potential diurnal variation (Nicolini et al., 2013; Rubio 1232 and Detto, 2017) of the three gases with good spatial replication (16 chambers), (3) periodic 1233 extreme events of high N₂O fluxes can still be detected with the SHORT closure time period, 1234 which occurs at higher frequency than the LONG closure measurements. Our study underlines the 1235 importance of appropriate closure time for each GHG gas for accurate estimation of GHG budgets. 1236 We demonstrate here that the combination of a commercial soil GHG chamber system – the LI-8100A Automated Soil CO₂ Flux System – running in line with a Picarro G2308, enables the 1237 1238 continuous, long-term measurement of CO₂, CH₄, and N₂O simultaneously under tropical 1239 conditions. Similar configurations have been recently implemented in temperate climate (Petrakis et al., 2017b, 2017a), but to our knowledge, this is the first time that this experimental set up is 1240 1241 fully described and tested under tropical field conditions for the measurement of the three soil 1242 GHG fluxes simultaneously. Additionally, our study determined the optimal chamber closure time 1243 for each GHG. This information is crucial for the calculation of accurate soil fluxes at diurnal 1244 timesteps and for the estimation of annual GHG budgets. This combination of automated closed dynamic chambers and advanced GHG analyzers allows for, (1) accounting of short-term 1245 1246 variability in GHG fluxes while taking into account spatial variability, (2) estimating annual GHG 1247 budgets at these locations, (3) tracking the variability in GHG fluxes along hours, days, seasons 1248 and years, and (4) studying the impact of climatic change on soil GHG budgets.

1250 Author contribution. JVB and NA designed the experiments and EAC, CS, BB and DB carried

1251 them out. EAC and CS prepared the manuscript with contributions from all co-authors.

1252 **Competing interests.** The authors declare that they have no conflict of interest.

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1266

1268	Supplementary File 1: R code for merging Picarro files to include them in Soil Flux pro
1269	
1270	## to list all the days in a given directory (Picarro makes one directory per day)
1271	ListDay<-list.files()
1272	Pfile<-list()
1273	## to concatenate all the hourly file in one file per day
1274	for (j in 1:length(ListDay))
1275	{
1276	print(j)
1277	ListFilesPicarro<-list.files(ListDay[j])
1278	Data<-read.table(paste(ListDay[j],"/",ListFilesPicarro[1],sep=""))
1279	for (i in 2:length(ListFilesPicarro))
1280	{
1281	temp<-read.table(paste(ListDay[j],"/",ListFilesPicarro[i],sep=""))
1282	Data<-rbind(Data, temp)
1283	print(i)
1284	
1285	Pfile[[j]]<-Data
1286	
1287	## to concatenante all days and make just one file will all data
1288	MasterData<-Pfile[[1]]
1289	for (k in 2:length(Pfile))
1290	
1291	MasterData<-rbind(MasterData,Pfile[[k]])
1292	print(k)
1293	
1294	## to write the table in a way that SFP can read it
1295	write.table(MasterData, "MasterData.dat", quote=F)
1296	
1297	

1299 Supplementary Figure 1: Relationship between soil surface humidity and (A) CO₂, (B) CH₄ and
 1300 (C) N₂O fluxes over the study period.



1305 **REFERENCES**

Aguilos, M., Hérault, B., Burban, B., Wagner, F. and Bonal, D.: What drives long-term variations
in carbon flux and balance in a tropical rainforest in French Guiana?, Agric. For. Meteorol., 253,
114–123, 2018.

Ambus, P., Skiba, U., Drewer, J., Jones, S., Carter, M. S., Albert, K. R. and Sutton, M.:
Development of an accumulation-based system for cost-effective chamber measurements of inert
trace gas fluxes, Eur. J. Soil Sci., 61(5), 785–792, 2010.

- 1312 Arias-Navarro, C., Díaz-Pinés, E., Klatt, S., Brandt, P., Rufino, M. C., Butterbach-Bahl, K. and
- Verchot, L.: Spatial variability of soil N₂O and CO₂ fluxes in different topographic positions in a
 tropical montane forest in Kenya, J. Geophys. Res. Biogeosciences, 122(3), 514–527, 2017.
- 1514 upped montane forest in Kenya, J. Geophys. Res. Diogeosetences, 122(5), $514^{-}527$, 2017.
- Bonal, D., Bosc, A., Ponton, S., Goret, J.Y., Burban, B., Gross, P., Bonnefond, J., Elbers, J., Longdoz, B. and Epron, D.: Impact of severe dry season on net ecosystem exchange in the
- 1317 Neotropical rainforest of French Guiana, Glob. Change Biol., 14(8), 1917–1933, 2008.
- 1318 Bréchet, L., Ponton, S., Roy, J., Freycon, V., Coûteaux, M.-M., Bonal, D. and Epron, D.: Do tree
- 1319 species characteristics influence soil respiration in tropical forests? A test based on 16 tree species
- 1320 planted in monospecific plots, Plant Soil, 319(1–2), 235–246, 2009.
- Breuer, L., Papen, H. and Butterbach-Bahl, K.: N₂O emission from tropical forest soils of Australia, J. Geophys. Res. Atmospheres, 105(D21), 26353–26367, 2000.
- 1323 Christiansen, J. R., Outhwaite, J. and Smukler, S. M.: Comparison of CO₂, CH₄ and N₂O soil-1324 atmosphere exchange measured in static chambers with cavity ring-down spectroscopy and gas 1325 chromatography, Agric. For. Meteorol., 211, 48–57, 2015.
- Courtois, E. A., Stahl, C., Van den Berge, J., Bréchet, L., Van Langenhove, L., Richter, A., Urbina,
 I., Soong, J. L., Peñuelas, J. and Janssens, I. A.: Spatial Variation of Soil CO₂, CH₄ and N₂O Fluxes
 Across Topographical Positions in Tropical Forests of the Guiana Shield, Ecosystems, 21(7),
 1445–1458, 2018.
- 1330 Davidson, E., Savage, K., Verchot, L. and Navarro, R.: Minimizing artifacts and biases in 1331 chamber-based measurements of soil respiration, Agric. For. Meteorol., 113(1), 21–37, 2002.
- Davidson, E. A., Nepstad, D. C., Ishida, F. Y. and Brando, P. M.: Effects of an experimental
 drought and recovery on soil emissions of carbon dioxide, methane, nitrous oxide, and nitric oxide
- in a moist tropical forest, Glob. Change Biol., 14(11), 2582–2590, 2008.
- De Klein, C. and Harvey, M.: Nitrous oxide chamber methodology guidelines, Glob. Res. Alliance
 Agric. Greenh. Gases Minist. Prim. Ind. Wellingt. N. Z., 2012.
- Denmead, O.: Chamber systems for measuring nitrous oxide emission from soils in the field, Soil
 Sci. Soc. Am. J., 43(1), 89–95, 1979.

- Dutaur, L. and Verchot, L. V.: A global inventory of the soil CH₄ sink, Glob. Biogeochem. Cycles,
 21(4), 2007.
- Epron, D., Bosc, A., Bonal, D. and Freycon, V.: Spatial variation of soil respiration across a
 topographic gradient in a tropical rain forest in French Guiana, J. Trop. Ecol., 22(05), 565–574,
 2006.
- Görres, C.-M., Kammann, C. and Ceulemans, R.: Automation of soil flux chamber measurements:
 potentials and pitfalls, Biogeosciences, 13(6), 1949–1966, 2016.
- Hupp, J. R., Garcia, R. L., Madsen, R. and McDermitt, D. K.: Measurement of CO₂ Evolution in
 a Multiplexed Flask System, vol. 44, pp. 1143–1143, Amer. Soc. Horticultural Science, 2009.
- Janssens, I. A., Kowalski, A. S., Longdoz, B. and Ceulemans, R.: Assessing forest soil CO₂ efflux:
 an in situ comparison of four techniques, Tree Physiol., 20(1), 23–32, 2000.
- Koskinen, M., Minkkinen, K., Ojanen, P., Kämäräinen, M., Laurila, T. and Lohila, A.:
 Measurements of CO₂ exchange with an automated chamber system throughout the year:
 challenges in measuring night-time respiration on porous peat soil, Biogeosciences, 11(2), 347,
 2014.
- Kostyanovsky, K., Huggins, D., Stockle, C., Waldo, S. and Lamb, B.: Developing a flow through
 chamber system for automated measurements of soil N₂O and CO₂ emissions, Measurement, 113,
 172–180, 2018.
- Merbold, L., Wohlfahrt, G., Butterbach-Bahl, K., Pilegaard, K., DelSontro, T., Stoy, P. and Zona,
 D.: Preface: Towards a full greenhouse gas balance of the biosphere, Biogeosciences, 12(2), 453–
 456, 2015.
- Nickerson, N.: Evaluating gas emission measurements using Minimum Detectable Flux (MDF),
 Eosense Inc Dartm. N. S. Can., 2016.
- 1362 Nicolini, G., Castaldi, S., Fratini, G. and Valentini, R.: A literature overview of 1363 micrometeorological CH_4 and N_2O flux measurements in terrestrial ecosystems, Atmos. Environ., 1364 81, 311-319, 2013.
- O'Connell, C. S., Ruan, L. and Silver, W. L.: Drought drives rapid shifts in tropical rainforest soil
 biogeochemistry and greenhouse gas emissions, Nat. Commun., 9(1), 1348, 2018.
- Oertel, C., Matschullat, J., Zurba, K., Zimmermann, F. and Erasmi, S.: Greenhouse gas emissions
 from soils—A review, Chem. Erde-Geochem., 76(3), 327–352, 2016.
- 1369 Petitjean, C., Hénault, C., Perrin, A.-S., Pontet, C., Metay, A., Bernoux, M., Jehanno, T., Viard,
- A. and Roggy, J.-C.: Soil N 2 O emissions in French Guiana after the conversion of tropical forest
 to agriculture with the chop-and-mulch method, Agric. Ecosyst. Environ., 208, 64–74, 2015.
- Petrakis, S., Seyfferth, A., Kan, J., Inamdar, S. and Vargas, R.: Influence of experimental extreme water pulses on greenhouse gas emissions from soils, Biogeochemistry, 133(2), 147–164, 2017a.

Petrakis, S., Barba, J., Bond-Lamberty, B. and Vargas, R.: Using greenhouse gas fluxes to define
soil functional types, Plant Soil, 1–10, 2017b.

Pumpanen, J., Kolari, P., Ilvesniemi, H., Minkkinen, K., Vesala, T., Niinistö, S., Lohila, A.,
Larmola, T., Morero, M. and Pihlatie, M.: Comparison of different chamber techniques for
measuring soil CO₂ efflux, Agric. For. Meteorol., 123(3), 159–176, 2004.

- Rowland, L., Hill, T. C., Stahl, C., Siebicke, L., Burban, B., Zaragoza-Castells, J., Ponton, S.,
 Bonal, D., Meir, P. and Williams, M.: Evidence for strong seasonality in the carbon storage and
 carbon use efficiency of an Amazonian forest, Glob. Change Biol., 20(3), 979–991, 2014.
- Rubio, V. E. and Detto, M.: Spatiotemporal variability of soil respiration in a seasonal tropical forest, Ecol. Evol., 7(17), 7104–7116, 2017.
- Savage, K., Phillips, R. and Davidson, E.: High temporal frequency measurements of greenhouse
 gas emissions from soils, Biogeosciences, 11(10), 2709, 2014.
- Sayer, E. J., Powers, J. S. and Tanner, E. V.: Increased litterfall in tropical forests boosts the
 transfer of soil CO2 to the atmosphere, PLoS One, 2(12), e1299, 2007.
- Silver, W. L., Lugo, A. and Keller, M.: Soil oxygen availability and biogeochemistry along rainfall
 and topographic gradients in upland wet tropical forest soils, Biogeochemistry, 44(3), 301–328,
 1999.
- 1391 Teh, Y., Diem, T., Jones, S., Huaraca Quispe, L. P., Baggs, E., Morley, N., Richards, M., Smith,
- 1392 P. and Meir, P.: Methane and nitrous oxide fluxes across an elevation gradient in the tropical
- 1393 Peruvian Andes, Biogeosciences, 2014.
- Verchot, L. V., Davidson, E. A., Cattânio, H., Ackerman, I. L., Erickson, H. E. and Keller, M.:
 Land use change and biogeochemical controls of nitrogen oxide emissions from soils in eastern
 Amazonia, Glob. Biogeochem. Cycles, 13(1), 31–46, 1999.
- Verchot, L. V., Davidson, E. A., Cattânio, J. H. and Ackerman, I. L.: Land-use change and
 biogeochemical controls of methane fluxes in soils of eastern Amazonia, Ecosystems, 3(1), 41–
 56, 2000.
- Wagner, F., Hérault, B., Stahl, C., Bonal, D. and Rossi, V.: Modeling water availability for trees
 in tropical forests, Agric. For. Meteorol., 151(9), 1202–1213, 2011.
- 1402
- 1403