

## Responses to comments

Dear referees,

We thank you for your supportive comments on our manuscript. Our detailed response in blue text to your comments is attached. Changes to the manuscript text are underlined.

### Responses to comments of referee 1

#### Summary

Zhang et al. conducted a series of experiments with multiple strains of *Emiliana huxleyi* isolated from 3 different North Atlantic populations. Each strain was incubated under a broad range of pCO<sub>2</sub> concentrations (about 120-2600 $\mu$ atm) but with constant total alkalinity to discern between effects due to changes in the carbonate systems and changes in CO<sub>2</sub> levels. The physiological responses that Zhang et al measured were growth rates, PIC and POC production rates. They conclude that there were differences among strains and among populations but those differences depended on the physiological rate.

#### General comments

The manuscript is very well written. The ideas, methods and discussion are also clear and well structured, making the manuscript flow very well. This is high quality and thorough work and it deserves to be published. However, my main comment is perhaps related to the novelty of the work and I will make some suggestions as to how this could be addressed. Zhang et al. do a good job citing some of the previous relevant studies but their work would be better served by emphasizing how their work is significantly different and why this is important. We already know from studies like Iglesias-Rodriguez, Bach, Langer, etc., that there are CO<sub>2</sub> effects in coccolithophore's physiological rates and we also know from Langer et al.'s work that these are species-specific and strain-specific responses, so (in my humble opinion) there is not much surprise in finding that there are population-specific differences. Throughout the manuscript the authors hint at the ideas of phenotypic plasticity and environmental variability. This, on the other hand is not so common, and I suggest that the authors elaborate more on this. They already show the pCO<sub>2</sub> and temperature ranges in those 3 sites and it is used to explain the results. Fully accounting for this variability at the original field site is important and they should emphasize that. Acknowledging this variability is usually not done.

Response: We thank this referee for the positive comments. We summarized responses of growth, POC and PIC production rates of different *Emiliana huxleyi* strains to CO<sub>2</sub> and found that most of these studies focused on a few strains or a narrow range of CO<sub>2</sub> level (Table R1). In this study, we used 17 strains and measured growth, POC and PIC production rates at 120  $\mu$ atm to 2630  $\mu$ atm, which are different from previous studies. **These contents were shown in lines 84–87.**

**For lines 333–336:** When exposed to 16 °C, growth rate of the Canary Islands population might have been already below their optimum and hence significantly reduced in comparison to the other populations ~~thus it grew slower than the other populations~~ (Fig. 2d).

**For lines 388–391:** Phenotypic plasticity constitutes an advantage for individual strains to acclimate and adapt to elevated  $p\text{CO}_2$  by changing ~~their~~ fitness-relevant traits and potentially to attenuate the effects of changing environments on fitness-relevant traits (Schaum et al., 2013).

**For lines 395–397:** Physiological variability makes a population more resilient, ~~and~~ increases its ability to persist in variable environments and potentially forms the basis for selection (Gsell et al., 2012; Hattich et al., 2017).

Table R1. Summary of the physiological responses of different *E. huxleyi* genotypes to various  $p\text{CO}_2$  ranges at constant alkalinity condition. Symbols indicate: ↑ increased response, — no response, ↓ decreased response, ∩ optimum response.

<i>E. huxleyi</i> genotype	Isolated site	$p\text{CO}_2$ range (μatm)	Growth rate	POC pro.	PIC pro.	Incubation temp. (°C)	Reference
AC472	South Pacific, New Zealand	400 to 760	↑	—	↑	19	Fiorini et al., (2011)
EHTB 11.15	Trumpeter Bay, Tasmania	375 to 1650	—	∩	↓	14	Müller et al., (2015)
EHSO 5.14	Southern Ocean	300 to 1680	↓	∩	∩	14	Müller et al., (2015)
EHSO 5.11	Southern Ocean	259 to 1255	∩	∩	∩	14	Müller et al., (2015)
NIWA1108	Chatham Rise, New Zealand	80 to 1080	↑	↑	∩	4-25	Feng et al., (2017)
PLY M219 (NZEH)	New Zealand	380 to 750	↓	↓	↓	20	Shi et al., (2009)
PLY M219 (NZEH)	New Zealand	404 to 1066	↓	↑	↓	15	Hoppe et al., (2011)
PML B92/11A	Bergen, Norway	152 to 885	—	↑	↓	15	Riebesell et al., (2000)
PML B92/11A	Bergen, Norway	20 to 6000	∩	∩	∩	15	Bach et al., (2011)
RCC1212	South Atlantic, off South Africa	194 to 1096	↓	∩	↓	20	Langer et al., (2009)
RCC1216	Tasman Sea, off New Zealand	218 to 1201	↓	↑	↓	17	Langer et al., (2009)
RCC1238	North Atlantic, off Japan	206 to 929	↑	∩	—	20	Langer et al., (2009)
RCC1256	North Atlantic, off Iceland	193 to 915	↓	∩	∩	17	Langer et al., (2009)
RCC1256	Iceland	191 to 846	↓	↓	↓	15	Hoppe et al., (2011)
NZEH	New Zealand	280 to 750	↓	↑	↑	19	Iglesias-Rodriguez et al., (2008)

- Bach, L. T., Riebesell, U., and Schulz, K.G.: Distinguishing between the effects of ocean acidification and ocean carbonation in the coccolithophore *Emiliana huxleyi*, *Limnol. Oceanogr.*, 56, 2040–2050, 2011.
- Beardall, J., and Raven, J. A.: Potential effects of global change on microalgal photosynthesis, growth and ecology, *Phycologia*, 43, 26–40, 2004.
- Feng, Y. Y., Roleda, M. Y., Armstrong, E., Boyd, P. W., and Hurd, C. L.: Environmental controls on the growth, photosynthetic and calcification rates of a Southern Hemisphere strain of the coccolithophore *Emiliana huxleyi*, *Limnol. Oceanogr.*, 62, 519–540.
- Fiorini, S., Middelburg, J. J., and Gattuso, J. P.: Testing the effects of elevated pCO<sub>2</sub> on coccolithophores (prymnesiophyceae): comparison between haploid and diploid life stages, *J. Phycol.*, 47, 1281–1291, 2011.
- Hoppe, C. J. M., Langer, G., and Rost, B.: *Emiliana huxleyi* shows identical responses to elevated pCO<sub>2</sub> in TA and DIC manipulations, *J. Exp. Mar. Biol. Ecol.*, 406, 54–62, 2011.
- Iglesias-Rodriguez, M. D., Halloran, P. R., Rickaby, R. E. M., Hall, I. R., Colmenero-Hidalgo, E., Gittins, J. R., Green, D. R. H., Tyrrell, T., Gibbs, S. J., von Dassow, P., Rehm, E., Armbrust, E.V., and Boessenkool, K.P.: Phytoplankton calcification in a high-CO<sub>2</sub> world, *Science*, 320, 336–340, 2008.
- Jones, B. M., Iglesias-Rodriguez, M. D., Skipp, P. J., Edwards, R. J., Greaves, M. J., Young, J. R., Elderfield, H., and O'Connor, C. D.: Responses of the *Emiliana huxleyi* proteome to ocean acidification, *PLoS One* 8(4), e61868, doi: 10.1371/journal.pone.0061868, 2013.
- Langer, G., Nehrke, G., Probert, I., Ly, J., and Ziveri, P.: Strain-specific responses of *Emiliana huxleyi* to changing seawater carbonate chemistry, *Biogeosciences*, 6, 2637–2646, 2009.
- Müller, M. N., Trull, T. W., and Hallegraeff, G. M.: Differing responses of three Southern Ocean *Emiliana huxleyi* ecotypes to changing seawater carbonate chemistry, *Mar. Ecol. Prog. Ser.*, 531, 81–90, 2015.
- Riebesell, U., Zondervan, I., Rost, B., Tortell, P. D., Zeebe, R. E., and Morel, F. M. M.: Reduced calcification of marine plankton in response to increased atmospheric CO<sub>2</sub>, *Nature*, 407, 364–367, 2000.
- Shi, D., Xu, Y., and Morel, F. M. M.: Effects of the pH/pCO<sub>2</sub> control method on medium chemistry and phytoplankton growth, *Biogeosciences*, 6, 1199–1207, 2009.

### Specific comments

While isolating the effect of CO<sub>2</sub> from changes in TA is a great idea, it also poses the question of whether the same experiment should have been repeated letting the TA change with CO<sub>2</sub> concentration. It begs the question of "how would the results look like if TA could change?". After all, this is a more realistic situation and it would contribute to our understanding of *E. huxleyi* responses to a changing World. While I acknowledge that this would be an entire new project, I think it is my role to bring it up. Perhaps acknowledging the caveat would be enough.

Response: We did not 'isolate the effect of CO<sub>2</sub> from changes in TA', and our CO<sub>2</sub> manipulations are mimicking ongoing ocean acidification where CO<sub>2</sub>/pH and DIC changes at constant TA.

As shown in Tables R2 and R3, rising  $p\text{CO}_2$  level dominantly decreased pH at increasing TA conditions. According to studies of Bach et al. (2011), after optimum  $\text{CO}_2$  levels, low pH inhibited growth, POC and PIC production. Thus, we expected that growth, POC and PIC production rates should show optimal curve responses to a broad  $\text{CO}_2$  range at changing TA.

Table R2. Carbonate chemistry parameter at constant  $p\text{CO}_2$  levels.

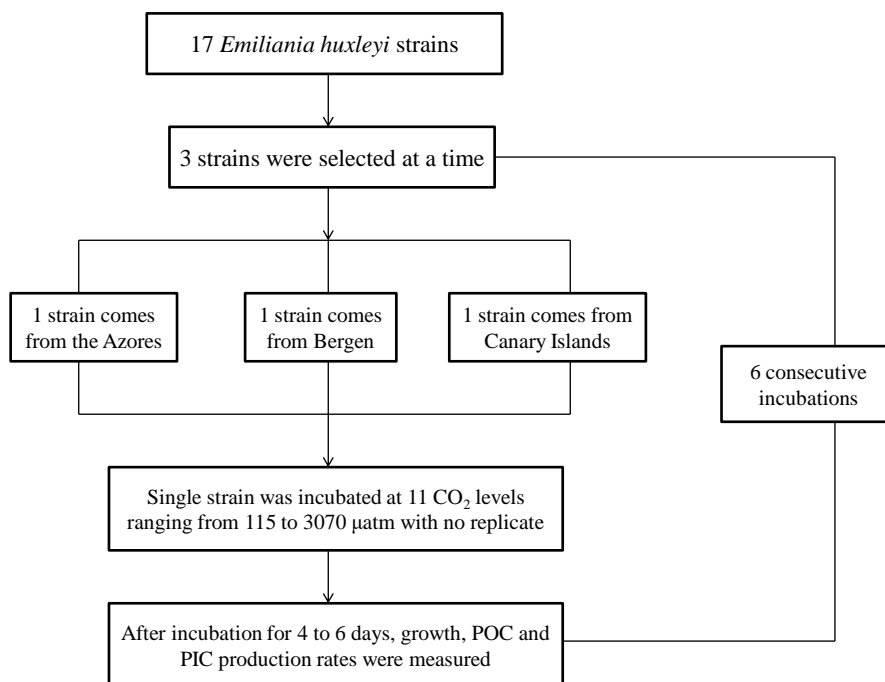
TA ( $\mu\text{mol L}^{-1}$ )	DIC ( $\mu\text{mol kg}^{-1}$ )	pH (total scale)	$p\text{CO}_2$ ( $\mu\text{atm}$ )	$\text{HCO}_3^-$ ( $\mu\text{mol kg}^{-1}$ )	$\text{CO}_3^{2-}$ ( $\mu\text{mol kg}^{-1}$ )	$\text{CO}_2$ ( $\mu\text{mol kg}^{-1}$ )	$\Omega$
1500	1351.6	7.887	400	1245.0	93.7	12.9	2.24
1600	1436.9	7.912	400	1318.8	105.1	12.9	2.51
1700	1521.8	7.935	400	1391.8	117.1	12.9	2.80
1800	1606.3	7.957	400	1463.9	129.5	12.9	3.10
1900	1690.4	7.978	400	1535.1	142.4	12.9	3.41
2000	1774.2	7.997	400	1605.4	155.8	12.9	3.73
2100	1857.5	8.016	400	1675.0	169.6	12.9	4.06
2200	1940.6	8.033	400	1743.8	183.8	12.9	4.40
2300	2023.3	8.050	400	1811.9	198.4	12.9	4.75
2400	2105.6	8.066	400	1879.2	213.5	12.9	5.11
2500	2187.7	8.081	400	1945.9	228.9	12.9	5.47
2600	2269.4	8.095	400	2011.8	244.7	12.9	5.85
2700	2350.8	8.109	400	2077.1	260.8	12.9	6.24
2800	2432.0	8.122	400	2141.8	277.3	12.9	6.63

Table R3. Carbonate chemistry parameter at changing  $p\text{CO}_2$  levels and changing TA conditions.

TA ( $\mu\text{mol L}^{-1}$ )	DIC ( $\mu\text{mol kg}^{-1}$ )	pH (total scale)	$p\text{CO}_2$ ( $\mu\text{atm}$ )	$\text{HCO}_3^-$ ( $\mu\text{mol kg}^{-1}$ )	$\text{CO}_3^{2-}$ ( $\mu\text{mol kg}^{-1}$ )	$\text{CO}_2$ ( $\mu\text{mol kg}^{-1}$ )	$\Omega$
1500	1254.0	8.134	200	1101.0	146.5	6.5	3.51
1600	1436.9	7.912	400	1318.8	105.1	12.9	2.51
1700	1576.7	7.783	600	1470.2	87.1	19.4	2.08
1800	1701.7	7.694	800	1598.6	77.2	25.8	1.85
1900	1819.7	7.628	1000	1716.2	71.2	32.3	1.70
2000	1934.0	7.576	1200	1827.9	67.3	38.8	1.61
2100	2046.0	7.534	1400	1936.0	64.7	45.2	1.55
2200	2156.4	7.500	1600	2041.7	63.0	51.7	1.51
2300	2265.8	7.470	1800	2145.8	61.8	58.1	1.48
2400	2374.4	7.445	2000	2248.7	61.1	64.6	1.46
2500	2482.4	7.422	2200	2350.7	60.7	71.1	1.45
2600	2590.1	7.403	2400	2452.0	60.6	77.5	1.45
2700	2697.4	7.386	2600	2552.8	60.6	84.0	1.45
2800	2804.4	7.370	2800	2653.2	60.8	90.4	1.45

I am a bit confused about how the incubations were done (not saying it is wrong) but perhaps a diagram or flow chart would be helpful. I mention this in the technical comments section as well.

Response: We agree with this referee and present a flow chart which shows the experimental protocol. This flow chart was added in the supplement information as Figure S1.



**Figure R1 (S1).** A flow chart of the experimental protocol.

Also, how realistic are CO<sub>2</sub> levels greater than 1500 uatm?

Response: According to business-as-usual CO<sub>2</sub> emissions (RCP8.5), atmospheric CO<sub>2</sub> level are projected higher than 1500 ppmv after 2200 (Meinshausen et al. 2011).

Meinshausen, M., Smith, S. J., Calvin, K., Daniel, J. S., Kainuma, M. L. T., Lamarque, J., Matsumoto, K., Montzka, S. A., Raper, S. C. B., Riahi, K., Thomson, A., Velders, G. J. M., and van Vuuren, D. P. P.: The RCP greenhouse gas concentrations and their extensions from 1765 to 2300, *Climatic Change*, 109, 213–241, 2011.

It is very interesting that they found almost no differences in PIC production rates among populations, yet growth and POC production rates did show differences at the population level. Why do you think this is? One factor that the authors mention briefly is temperature, I think that temperature-adaptation and temperature-CO<sub>2</sub> interactions might have a greater role in explaining the differences than what the authors attribute to it. In some ways the 3 populations sit along a gradient of temperature and CO<sub>2</sub> and depending on which physiological rate is studied, one parameter might be more important than the other. Zhang et al do mention that growing certain cultures under suboptimal temperatures may have set that

strain or population at a disadvantage from the beginning. Interactions between temperature and CO<sub>2</sub> effects should not be discarded.

Response: We thank the referee for this suggestion.

**For lines 337–348:** These contents ‘One of the reasons may be that compared to the Azores and Bergen populations, 16 °C likely causes lower carbon uptake and carbon-use efficiency of the Canary Islands population (Sett et al., 2014).’ were replaced by ‘Furthermore, compared to the Canary Islands population, the Azores population had higher maximum growth and POC production rates, and similar optimum CO<sub>2</sub> for these physiological rates. Again, this might be related to sub-optimal incubation conditions as temperature has been found to significantly modulate CO<sub>2</sub> responses in coccolithophores in terms of maximum rates, CO<sub>2</sub> optima and half-saturation, and H<sup>+</sup> sensitivity (De Bodt et al., 2010; Sett et al., 2014; Gafar et al., 2018; Gafar and Schulz, 2018). In a similar fashion light can also modulate CO<sub>2</sub> responses, hence different requirements by strains adapted to different light availabilities could also explain our observations (Zhang et al., 2015; Gafar et al., 2018; Gafar and Schulz, 2018).’

**For lines 350–355:** In addition, the Canary Islands population showed smallest variability in optimum pCO<sub>2</sub> and maximum values for growth and POC production rates (Fig. 2). The reason may be that low incubation temperature predominantly limited growth and POC production rates of the Canary Islands population, and decreased the sensitivities of these physiological rates to rising pCO<sub>2</sub>.

De Bodt, C., Van Oostende, N., Harlay, J., Sabbe, K., and Chou, L.: Individual and interacting effects of pCO<sub>2</sub> and temperature on *Emiliania huxleyi* calcification: study of the calcite production, the coccolith morphology and the coccosphere size, *Biogeosciences*, 1401–1412, 2010.

Gafar, N. A., Eyre, B. D., and Schulz, K. G. : A conceptual model for projecting coccolithophorid growth, calcification and photosynthetic carbon fixation rates in response to global ocean change, *Front. Mar. Sci.*, 4, 433, doi: 10.3389/fmars.2017.00433, 2018.

Gafar, N. A., and Schulz, K. G. : A niche comparison of *Emiliania huxleyi* and *Gephyrocapsa oceanica* and potential effects of climate change, *Biogeosci. Discuss.*, doi: 10.5194/bg-2018-88.

Another consideration is that Zhang et al do a great job by showing that there are different ranges of variability in the places where they were isolated from and they use this argument to explain the differences. However, their cultures are maintained at a constant CO<sub>2</sub> concentration (and light pattern and temperature). As the authors suggest in this manuscript, the next generation of experiments should account for variability at its origin and hence variable environmental parameters (within a given range) in experimental designs. Plasticity and adaptation are key parameters to consider in the future.

Response: we agree with this referee.

**For lines 388–391:** Phenotypic plasticity constitutes an advantage for individual strains to acclimate and adapt to elevated pCO<sub>2</sub> by changing fitness-relevant traits and potentially to attenuate the short-term effects of changing environments on fitness-relevant traits (Schaum et al.,

2013).

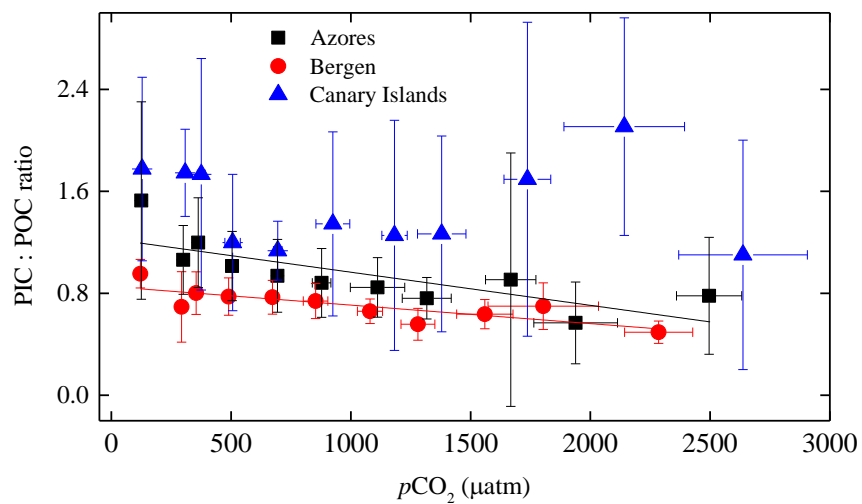
Finally, Zhang et al found some very interesting results, some of which were not fully explored. For example, the optimum  $p\text{CO}_2$  is higher for Bergen than the other 2 regions, but the temperature optimum in Bergen is lower, what are the implications for future projections? Similarly, all strains but one showed that the  $p\text{CO}_2$  optimum for POC is greater than the optimum for PIC and growth rates, how do you think this might affect future PIC: POC ratios? What about the sensitivity constant results? OR Bergen populations experiencing the higher  $\text{CO}_2$  optimum and smallest variability between strains vs. Canary islands showing lowest optimums but highest variabilities in  $\text{CO}_2$  optimums..... These are just some examples of other interesting avenues to explore in the discussion.

Response: Agreed. The optimum temperature for growth of the Bergen population was about 22 °C and was 5 °C higher than the maximum SST in Bergen waters (Zhang et al. 2014). Furthermore, in comparison to the Azores and Canary Islands populations, larger optimum  $p\text{CO}_2$  of growth rate indicates that the Bergen population may benefit more from the rising  $\text{CO}_2$  levels at increasing temperatures. These contents were added in lines 367–372.

As shown in Fig. R2 (or S6 in the supplement), PIC : POC ratios of the Azores and Bergen populations declined with rising  $p\text{CO}_2$ , whereas PIC : POC ratios of the Canary Islands population were rather constant (Fig. S6). As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean. These contents underlined were added in lines 372–376.

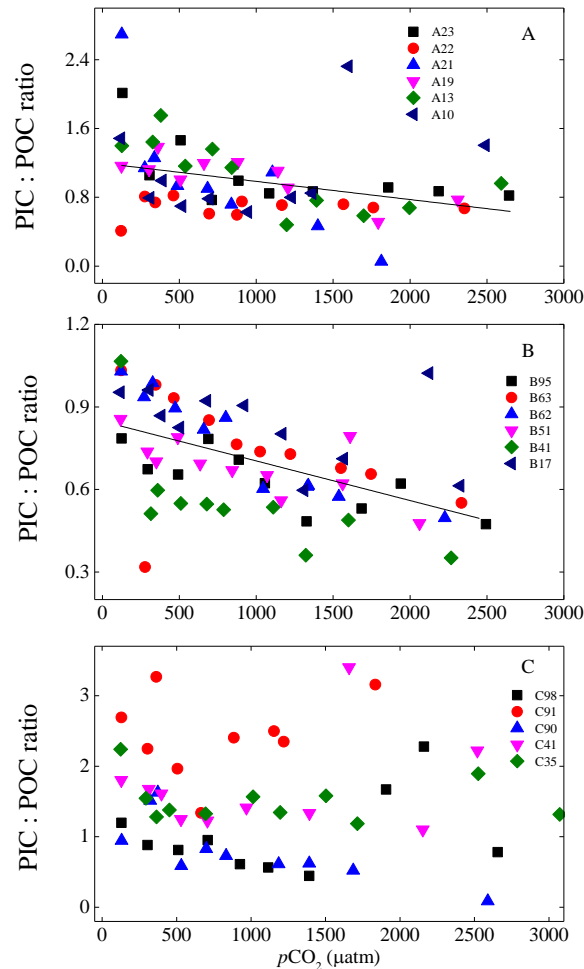
As shown in lines 304–306 in the manuscript or in Fig. 2, low sensitivity constant of growth rate of the Bergen population corresponded to high optimum  $\text{CO}_2$  level.

**For lines 350–355:** In addition, the Canary Islands population showed smallest variability in optimum  $p\text{CO}_2$  and maximum values for growth and POC production rates (Fig. 2). The reason may be that low incubation temperature predominantly limited growth and POC production rates of Canary Islands population, and decreased the sensitivities of these physiological rates to rising  $p\text{CO}_2$ .



**Figure R2 (S6).** Responses of PIC : POC ratio of the Azores (square), Bergen (circular) and Canary Islands (diamond) populations to a  $\text{CO}_2$  range from 120  $\mu\text{atm}$  to 2630  $\mu\text{atm}$ .





**Figure R3 (S7).** Response of PIC : POC ratio of individual *E. huxleyi* strain in the Azores (A), Bergen (B) and Canary Islands (C) populations to a CO<sub>2</sub> range from 115 µatm to 3070 µatm.

Technical comments Line 39: than that of

Response: **For lines 40–42:** Maximum growth and POC production rates of the Azores and Bergen populations were similar and significantly higher than that of the Canary Islands population.

Line 44-45: carbonate chemistry responses? Should it say instead "responses to changes in carbonate chemistry changes"?

Response: **For lines 48–51:** The existence of distinct ~~carbonate chemistry~~ responses to changes in carbonate chemistry between and within populations will likely benefit *E. huxleyi* to acclimate and adapt to rising CO<sub>2</sub> levels in the oceans.

Line 76: I recommend checking this new publication: Krumhardt et al. 2017. Coccolithophore growth and calcification in a changing ocean <https://doi.org/10.1016/j.pocean.2017.10.007>

Response: Krumhardt et al. (2017) developed an empirical coccolithophore model to investigate responses of growth and calcification of coccolithophores to changing environments (temperature,

CO<sub>2</sub>, nutrient concentrations). This paper is now cited on line 76.

**For lines 73–76:** The coccolithophore *Emiliania huxleyi* forms extensive blooms under favourable light intensity, temperature and nutrient conditions, with different morphotypes in certain regions (Cook et al., 2011; Henderiks et al., 2012; Smith et al., 2012; Balch et al., 2014; [Krumhardt et al., 2017](#)).

Krumhardt, K. M., Lovenduski, N. S., Debora Iglesias-Rodriguez, M., and Kleypas, J. A.: Coccolithophore growth and calcification in a changing ocean, *Prog. Oceanogr.*, 159, 276–295.

Line 135: "consecutive incubations" and then in Line 146 "each strain was grown under 11 CO<sub>2</sub> levels: : ." then in line 150 and 158 "at least 7 generations: : :4-7 days depending on CO<sub>2</sub> concentration : : :". can you explain the method in more detail, I am bit confused. Perhaps a supplementary diagram or flow chart figure would help.

Response: As mentioned above, a flow chart showing the experimental protocol was added to the supplement information (Figure S1).

**For lines 138–140:** The experiment was performed in six consecutive incubations, with one strain from each population (Azores, Bergen, Canary Islands) being cultured at a time ([Fig. S1](#)).

Line 202: For Eq 4 and 5, you cited Bach et al 2011, but could you please elaborate on this method. Can you also explain the sensitivity constant a bit more?

Response: **For lines 202–209:** [In a broad  \$p\text{CO}\_2\$  range, physiological rates are expected to initially increase quickly until reaching an optimum and then decline towards further increasing CO<sub>2</sub> levels \(e.g. Krug et al. 2011\). Hence we used the following modified Michaelis-Menten equation \(Bach et al. 2011\) which was fitted to measured cellular growth, POC and PIC production rates and yielding theoretical optimum  \$p\text{CO}\_2\$  and maximum values for each of the three populations \(combining the data of five or six strains\) \(Bach et al., 2011\).](#)

**For lines 211–212:** [s, the sensitivity constant, depicts the slope of the decline after optimum CO<sub>2</sub> levels in response to rising H<sup>+</sup>.](#)

Line 207: Do these refer to figure S3?

Response: **(Lines 218–220)** Relative growth, POC and PIC production rates of each population are shown in Fig. 1b,d,f. Relative POC and PIC quotas of each population were shown in Fig. S2.

Line 295: "These findings indicate that the Bergen population may be more tolerant....." This is a great result! Environmental variability can tell us something about phenotypic plasticity.

Response: **(Lines 315–317)** Large environmental variability usually results in high tolerance of phytoplankton (Doblin and van Sebille, 2016). In this study, we cannot say that large environmental variability result in large or low phenotypic plasticity.

Doblin, M. A., and van Sebille, E.: Drift in ocean currents impacts intergenerational microbial exposure to temperature, *Proc. Natl. Acad. Sci. USA.*, 113, 5700–5705, doi: 10.1073/pnas.1521093113, 2016.

Line 323 "likely causes the lower the carbon: : ." consider moving "the"

**Response: For lines 345–348:** we delete ‘*One of the reasons may be that compared to the Azores and Bergen populations, 16 °C likely causes lower the carbon uptake and carbon use efficiency of the Canary Islands population*’

Line 343: add and "s" to proton

**Response: For lines 379–382:** In addition, *E. huxleyi* is thought to utilize  $\text{HCO}_3^-$  for calcification which generates protons, and increase in proton concentration may mitigate the potential of the ocean to absorb atmospheric  $\text{CO}_2$  (Paasche, 2002).

Line 345: consider adding "and" before "corresponding"

**Response: For lines 382–385:** we deleted this sentence ‘*Thus, due to population-specific growth and PIC production rates or quotas, changes in species composition, corresponding changes in PIC productions, may affect the ability of the ocean to take up  $\text{CO}_2$ .*’

Line 352: this conclusion seems to be out of place and not well justified

**Response: For lines 391–393:** we deleted this sentence ‘*Additionally, our results also suggest that strain-specific PIC quota may be the basis of variation in coccoliths of *E. huxleyi* within the morphotype A (Fig. S4) (Young, 1994; Paasche, 2002).*’

Lines 334-372: some very interesting ideas here but these paragraphs need some tightening.

**Response:** According to suggestions of this referee, we added and deleted some contents **in lines 365–382:** The ability to adapt to diverse environmental conditions is supposed to be one reason for the global distribution of *E. huxleyi* (Paasche, 2002), spanning a temperature range of about 30 °C. The optimum temperature for growth of the Bergen population was about 22 °C and was 5 °C higher than the maximum SST in Bergen waters (Zhang et al. 2014). Furthermore, in comparison to the Azores and Canary Islands populations, larger optimum  $p\text{CO}_2$  of growth rate indicates that the Bergen population may benefit more from the rising  $\text{CO}_2$  levels at increasing temperatures. PIC : POC ratios of the Azores and Bergen populations declined with rising  $p\text{CO}_2$ , whereas PIC : POC ratios of the Canary Islands population were rather constant (Fig. S6). As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean. In natural seawater, due to ocean currents and gene flow, populations at any given location may get replaced by immigrant genotypes transported there from other locations (Doblin and van Sebille, 2016). In addition, *E. huxleyi* is thought to utilize  $\text{HCO}_3^-$  for calcification which generates protons, and increase in proton concentration may mitigate the potential of the ocean to absorb atmospheric  $\text{CO}_2$  (Paasche, 2002).

Line 367-369: do you mean "dominated" or "dominating"? not sure I follow this argument.

**Response: For lines 408–411:** Further, a significant positive correlation between growth and POC

production rate or POC quota (Fig. S5) indicates that the dominating strains will also take up or fix dissolved inorganic carbon faster.

**Responses to comments of referee 2 are shown as following:**

#### GENERAL COMMENTS

The paper by Zhang et al. presents results from a large number of experiments on multiple geographically distinct strains of the coccolithophore *Emiliana huxleyi*. Each strain was exposed to a wide range of pCO<sub>2</sub> concentrations and the authors examined differences in growth rates, photosynthetic rates (POC production) and calcification rates (PIC production). The authors conclude that significant variability exists in population-level sensitivity of physiological rates (most clearly growth and POC production) to pCO<sub>2</sub>. The paper is well written, with the data supporting the conclusions and the authors make some important and insightful conclusions. I have only two minor comments.

The first comment relates to a lack of any discussion or presentation of the variability in PIC:POC ratios and POC (or PIC) production between the different strains. Further information on the level of inter-strain variability in these parameters would strengthen and support the wider implications and conclusions made in the discussion. The second comment relates to the authors consideration of variability and stability in the different environmental conditions of the strain isolation locations – a large factor in these differences is likely to relate to different seasonal cycles and environmental drivers (ice-melt, riverine input, upwelling, etc). However, the authors only hint at the different factors influencing the relative stability of the different locations. Large-scale environmental differences will directly relate to the stability of the environment, as well as differing potential future perturbations for each

of them. Again, making these differences more explicit would support the wider implications of the study.

Response: We cultured 17 *Emiliana huxleyi* strains at 11  $p\text{CO}_2$  levels with no replicate. At each  $p\text{CO}_2$  level, there is no replicate and this is the main reason that we did not discuss variability in physiological rates between strains within the population.

Regarding the variability in the PIC : POC ratio between the populations, we added these contents ‘PIC : POC ratios of the Azores and Bergen populations declined with rising  $p\text{CO}_2$ , whereas PIC : POC ratios of the Canary Islands population were rather constant (Fig. S6). As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean.’ **in lines 372–376.**

**For lines 313–315:** we added these contents: ‘In addition, due to riverine input, seawater upwelling and metabolic activity of plankton communities, environmental variability in coastal water are larger than in open-ocean ecosystems (Duarte and Cerbrian, 1996).’

Duarte, C. M., and Cerbrian, J.: The fate of marine autotrophic production, *Limnol. Oceanogr.*, 41, 1758–1766, 1996.

#### SPECIFIC COMMENTS

Ln 27: Clarity is needed in the abstract on what the authors mean in terms of population-specific responses.

Response: In this study, ‘population-specific responses’ mean that growth, POC and PIC production rates of three *Emiliana huxleyi* populations were different at the same incubation conditions.

**For lines 27–32:** In the present study, we investigated the ~~population~~-specific responses of growth, particulate organic (POC) and inorganic carbon (PIC) production rates of ~~17 strains~~ 3 populations of the coccolithophore *Emiliana huxleyi* from three regions in the North Atlantic Ocean (Azores: 6 strains, Canary Islands: 5 strains, and Norwegian coast near Bergen: 6 strains) to a  $\text{CO}_2$  partial pressure ( $p\text{CO}_2$ ) range from 120  $\mu\text{atm}$  to 2630  $\mu\text{atm}$ .

Ln 28: More information on number of strains per environment would be good in the abstract.

Response: **For lines 27–32: see above.**

Ln 32: ‘expected optimum curve responses’ – may be expected by authors but not clear in the abstract. Some further background would be good.

Response: **For lines 32–34:** Physiological rates of each population and individual strain increased with rising  $p\text{CO}_2$  levels, reached maximum and declined thereafter.

Ln 37: Could the authors elaborate more in terms of the role of seasonality (or lack thereof) in the stability of oceanic conditions.

Response: **For lines 37 to 39:** This may be due to the large environmental variability including large  $p\text{CO}_2$  and pH fluctuations in coastal waters off Bergen compared to the rather stable oceanic

conditions at the other two sites.

In the discussion section, **for lines 313–315:** we added this sentence ‘In addition, due to riverine input, seawater upwelling and metabolic activity of plankton communities, environmental variability in coastal water are larger than in open-ocean ecosystems (Duarte and Cerbrian, 1996).’

Lns 91-92: Would the authors consider adding ‘geographically-distinct’ strains to this line to emphasize both the importance of their own insights and the more general need to consider different strains of other widespread species.

Response: **For lines 91–94:** Hence, multiple strains, ideally from geographically distinct regions should be considered for investigating phytoplankton responses to climate change (Zhang et al., 2014; Blanco-Ameijeiras et al., 2016; Krumhardt et al., 2017).

Krumhardt, K. M., Lovenduski, N. S., Debora Iglesias-Rodriguez, M., and Kleypas, J. A.: Coccolithophore growth and calcification in a changing ocean, *Prog. Oceanogr.*, 159, 276–295.

Lns 103-104: A plastic response also allows a strain to acclimate across an environmental gradient and widen its bio-geographical distribution. Rather than focus on just environmental change, what about environmental variability.

Response: **For lines 103–106:** Plasticity can be assessed by analyzing the reaction norm of one trait and a plastic response may allow a strain to acclimate across an environmental gradient and widen its bio-geographical distribution (Reusch, 2014; Levis and Pfennig, 2016).

Ln 126: How were all strains characterized and confirmed to be morphotype A (i.e. Distal shield length? Central area characteristics?)?

Response: Morphotype A was confirmed by scanning electron microscope.

**For Lines 128–129:** All 17 strains belong to morphotype A (determined by scanning electron microscopy) and have been deposited in the Roscoff culture collection (RCC)

Ln 140-141: Is this statement (‘the best compromise’) appropriate based on the authors end conclusion that the low experiment temperature relative to optimum growth conditions for the Canary Islands strains led to their low growth (and POC production)? It seems to be a compromise that had a definitive influence on the end outcome of the experiments. Is it not simpler to just delete this section (from the point of ‘which ..’ to the end) and come back to this in the discussion?

Reponse: **For lines 140–144:** Monoclonal populations were always grown in sterile-filtered (0.2  $\mu\text{m}$  diameter, Sartobran<sup>®</sup> P 300, Sartorius) artificial seawater medium (ASW) as dilute batch cultures at 200  $\mu\text{mol photons m}^{-2} \text{s}^{-1}$  light intensity under a 16/8 h light/dark cycle (light period: 5:00 a.m to 9:00 p.m.) at 16 °C which we consider to be ~~the best~~ a compromise for the three different origins of the strains.

Our results showed that low incubation temperature led to low growth and POC production rates of the Canary Islands population. In the discussion section, we compared influence of

temperature on physiological rate of three populations. **For lines 324–336:** In an earlier study (Zhang et al., 2014), growth rates of the same Azores and Bergen strains as used here were measured at 8–28 °C. While at 26–28 °C the Bergen strains grew slower than the Azores strains, at 8 °C the Azores strains grew slower than the Bergen strains. This illustrates nicely that local temperature adaptation can significantly affect growth of *E. huxleyi* strains in laboratory experiments. Considering these findings and the temperature ranges of the three isolation locations (Table S1), the incubation temperature of 16 °C used in the present study was lower than the minimum sea surface temperature (SST) commonly recorded at the Canary Islands. In contrast, SSTs of 16 °C and lower have been reported for Azores and Bergen waters (Table S1). When exposed to 16 °C, growth rate of the Canary Islands population might have been already below their optimum and hence significantly reduced in comparison to the other populations (Fig. 2d).

Lns 152-153 (cf Lns 174-175): How were initial cell densities measured/estimated?

Response: **(In line 156)** There was 590 ml seawater in the 500 ml glass bottles. Before cells were inoculated to new seawater, final cell concentrations ( $C_0$ ) were measured, and we calculated the inoculated volumes ( $V$ ) according to  $V = (200 \text{ cell/ml} \times 590 \text{ ml})/C_0$ . By using this method, we think that the initial cell concentration was 200 cell/ml.

For lines **155–157:** Initial cell concentration was 200 cells  $\text{ml}^{-1}$  (estimated from measured pre-culture concentrations and known dilution) and final cell concentration was lower than 100,000 cells  $\text{ml}^{-1}$ .

Lns 289-290: An important result that should be emphasized in the abstract and conclusions.

Response: In the abstract, we added this content **in lines 45–46:** Our results indicate adaptation of *E. huxleyi* to their local environmental conditions and the existence of distinct *E. huxleyi* populations.

In the conclusion: we added this sentence **in lines 420–423:** The existence of distinct *E. huxleyi* populations and phenotypic plasticity of individual strains may both be important for *E. huxleyi* when adapting to natural environmental variability and to ongoing climate changes.

Lns 322-324: Suggest deleting ‘causes’ from this sentence.

Response: **For lines 345–348:** we delete these contents ‘*One of the reasons may be that compared to the Azores and Bergen populations, 16 °C likely causes lower the carbon uptake and carbon-use efficiency of the Canary Islands population (Sett et al., 2014).*’

Ln 351-352: Another potentially important conclusion, especially given the emphasis on determining time-dependent (or space-dependent) variations in coccolith-specific PIC quotas. However, the current paper lacks any details of the strain-specific variability in PIC quota and to what extent the different trends in  $p\text{CO}_2$ -sensitivity (e.g. Fig. 3e) are driven by changes in growth rate and/or cellular (or coccolith) specific PIC quota. Can strain-specific information on PIC quota be added to the supplementary material to support this point with experimental data?

Response: PIC quota of population is shown in figure S2, and PIC quota of individual strain is shown in Figure S4. We measured PIC quota of individual strains at 11  $p\text{CO}_2$  levels **with no replicate**. This is the reason that we did not discuss PIC quota of individual strains.

We deleted this sentence **in lines 391–393**: ‘*Additionally, our results also suggest that strain-specific PIC quota may be the basis of variation in coccoliths of E. huxleyi within the morphotype A (Fig. S3) (Young, 1994; Paasche, 2002).*’

Ln 374: A two line conclusion seems relatively short based on the significant statements made in the conclusions. Either expand or delete?

Response: We added main result **in lines 420–423**: The existence of distinct *E. huxleyi* populations and phenotypic plasticity of individual strains may both be important for *E. huxleyi* when adapting to natural environmental variability and to ongoing climate changes.

**For lines 425–426**: we added ‘, and CO<sub>2</sub> response was modulated by other environmental factors such as temperature and light intensity.’



## List of changes

### Abstract

- 1 Lines 26–27: add ‘from different areas’
- 2 Line 27: delete ‘population’
- 3 Lines 28–29: add ‘3 populations’
- 4 Line 29: delete ‘17 strains’
- 5 Line 30: add ‘: 6 strains’, add ‘: 5 strains’
- 6 Line 31: add ‘: 6 strains’
- 7 Lines 32–33: change ‘displayed the expected optimum curve responses to the  $p\text{CO}_2$  gradient’ to ‘increased with rising  $p\text{CO}_2$  levels, reached maximum and declined thereafter’
- 8 Line 36: change ‘a’ to ‘the’
- 9 Line 37: change ‘fjord’ to ‘coast’
- 10 Line 38: add ‘environmental variability including large’, and delete ‘variability’
- 11 Line 39: add ‘fluctuations’
- 12 Line 41: add ‘that’
- 13 Lines 42–43: change ‘One of the reasons may be that the’ to ‘This pattern could be driven by temperature- $\text{CO}_2$ -interactions where the’
- 14 Line 44: change ‘is’ to ‘was’
- 15 Line 46: add ‘and the existence of distinct *E. huxleyi* populations’
- 16 Lines 48–49: delete ‘carbonate chemistry’
- 17 Line 49: add ‘to changes in carbonate chemistry’
- 18 Line 50: add ‘and adapt’

### Introduction

- 1 Line 76: add ‘; Krumhardt et al., 2017’
- 2 Line 91: change ‘These indicate that’ to ‘Hence,’
- 3 Lines 91–92: add ‘, ideally from geographically distinct regions’
- 4 Line 93: add ‘;’
- 5 Line 94: add ‘Krumhardt et al., 2017’
- 6 Line 105: change ‘to environmental change’ to ‘across an environmental gradient and widen its bio-geographical distribution’

### Materials and methods

- 1 Lines 128–129: add ‘(determined by scanning electron microscopy)’
- 2 Line 129: change ‘at’ to ‘in’
- 3 Lines 139–140: add ‘(Fig. S1)’
- 4 Line 143: change ‘the best’ to ‘a’
- 5 Lines 156–157: add ‘(estimated from measured pre-culture concentrations and known dilution)’
- 6 Lines 202–206: add ‘In a broad  $p\text{CO}_2$  range, physiological rates are expected to initially increase quickly until reaching an optimum and then decline towards further increasing  $\text{CO}_2$  levels’

(e.g. Krug et al. 2011). Hence we used the following modified Michaelis-Menten equation (Bach et al. 2011) which was fitted to measured cellular growth, POC and PIC production rates'

7 Lines 206–207: delete 'The nonlinear regression model (4) was used to fit growth, POC and PIC production rates'

8 Line 207: add 'and'

9 Line 211: delete 'is', add ',' and delete 'which indicates'

10 Line 212: delete 'the effect of', and add 'depicts the slope of the decline after optimum CO<sub>2</sub> levels in response to'

11 Line 213: delete 'the'

12 Line 214: add '(equation 5)', delete 'for physiological rates according to equation (5)', add 'and', and change 'M' to 'm'

13 Line 215: delete 'were calculated by using equation (4) based on  $K_m$ .'

14 Line 216: add 'following Bach et al., (2011).'

## Discussion

1 Lines 313–315: add 'In addition, due to riverine input, seawater upwelling and metabolic activity of plankton communities, environmental variability in coastal waters are larger than in open-ocean ecosystems (Duarte and Cerbrian, 1996).'

2 Line 317: change 'ed with' to 'ing'

3 Line 329: add 'the', and change 'ed' to 'ion'

4 Lines 334–336: change 'thus it grew slower than the other populations' to 'hence significantly reduced in comparison to the other populations'

5 Lines 337–345: add 'Furthermore, compared to the Canary Islands population, the Azores population had higher maximum growth and POC production rates, and similar optimum CO<sub>2</sub> for these physiological rates. Again, this might be related to sub-optimal incubation conditions as temperature has been found to significantly modulate CO<sub>2</sub> responses in coccolithophores in terms of maximum rates, CO<sub>2</sub> optima and half-saturation, and H<sup>+</sup> sensitivity (De Bodt et al., 2010; Sett et al., 2014; Gafar et al., 2018; Gafar and Schulz, 2018). In a similar fashion light can also modulate CO<sub>2</sub> responses, hence different requirements by strains adapted to different light availabilities could also explain our observations (Zhang et al., 2015; Gafar et al., 2018; Gafar and Schulz, 2018).'

6 Lines 345–348: delete 'One of the reasons may be that compared to the Azores and Bergen populations, 16 °C likely causes lower the carbon uptake and carbon-use efficiency of the Canary Islands population (Sett et al., 2014).'

7 Lines 350–355: add 'In addition, the Canary Islands population showed smallest variability in optimum  $p\text{CO}_2$  and maximum values for growth and POC production rates (Fig. 2). The reason may be that low incubation temperature predominantly limited growth and POC production rates of the Canary Islands population, and decreased the sensitivities of these physiological rates to rising  $p\text{CO}_2$ .'

8 Line 365: delete 'reflected in'

9 Line 366: add 'supposed to be one reason for'

10 Lines 367–376: add 'The optimum temperature for growth of the Bergen population was about

22 °C and was 5 °C higher than the maximum SST in Bergen waters (Zhang et al. 2014). Furthermore, in comparison to the Azores and Canary Islands populations, larger optimum  $p\text{CO}_2$  of growth rate indicates that the Bergen population may benefit more from the rising  $\text{CO}_2$  levels at increasing temperatures. PIC : POC ratios of the Azores and Bergen populations declined with rising  $p\text{CO}_2$ , whereas PIC : POC ratios of the Canary Islands population were rather constant (Fig. S6). As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean.'

11 Line 377: change 'populations' to 'immigrant'

12 Line 378: add 'genotypes'

13 Lines 378–379: delete 'when having a higher potential to adapt to a changing environment'

14 Line 380: change 'take up' to 'is thought to utilize', and change 'to calcify and' to 'for calcification which'

15 Line 381: add 's'

16 Lines 382–385: delete 'Thus, due to population-specific growth and PIC production rates or quotas, changes in species composition, corresponding changes in PIC productions, may affect the ability of the ocean to take up  $\text{CO}_2$ .'

17 Line 389: add 'acclimate and', and delete 'their'

18 Lines 390–391: add 'and potentially to attenuate the short-term effects of changing environments on fitness-relevant traits'

19 Lines 391–393: delete 'Additionally, our results also suggest that strain-specific PIC quota may be the basis of variation in coccoliths of *E. huxleyi* within the morphotype A (Fig. S3) (Young, 1994; Paasche, 2002).'

20 Line 396: add ',', delete 'and', and add 'and'

21 Line 397: add 'potentially forms the basis for selection'

22 Line 404: add 'a', and delete 's'

23 Line 406: add 'er', and add 'or other'

24 Line 407: add 'competitive abilities', add 's', and delete 'strains in the oceans'

25 Line 408: change 'S' to 'Further, a s'

26 Line 409: change '4' to '5', change 'suggests' to 'indicates', and change 'ed' to 'ing'

27 Line 410: change 'can' to 'will', add 'or fix', and delete 'from the oceans or'

28 Line 411: delete 'fix carbon faster', and change 'this' to 'When extrapolated to the ocean, *E. huxleyi* blooms'

29 Line 412: change 'or the' to 'and its'

30 Line 413: delete 'of the oceans when large *E. huxleyi* blooms occur'

31 Line 414: change 'will' to 'has the potential to'

## Conclusions

1 Lines 420–423: add 'The existence of distinct *E. huxleyi* populations and phenotypic plasticity of individual strains may both be important for *E. huxleyi* when adapting to natural environmental variability and to ongoing climate changes.'

2 Line 424: change 'or' to 'and'

3 Lines 425–426: add ', and  $\text{CO}_2$  response was modulated by other environmental factors such as

temperature and light intensity.’

## References

- 1 Lines 506–507: add ‘Duarte, C. M., and Cerbrian, J.: The fate of marine autotrophic production, *Limnol. Oceanogr.*, 41, 1758–1766, 1996.’
- 2 Lines 512–515: add ‘Gafar, N. A., Eyre, B. D., and Schulz, K. G. : A conceptual model for projecting coccolithophorid growth, calcification and photosynthetic carbon fixation rates in response to global ocean change, *Front. Mar. Sci.*, 4, 433, doi: 10.3389/fmars.2017.00433, 2018.’
- 3 Lines 516–518: add ‘Gafar, N. A., and Schulz, K. G. : A niche comparison of *Emiliana huxleyi* and *Gephyrocapsa oceanica* and potential effects of climate change, *Biogeosci. Discuss.*, doi: 10.5194/bg-2018-88.’
- 4 Lines 559–561: add ‘Krug, S. A., Schulz, K. G., and Riebesell, U. : Effects of changes in carbonate chemistry speciation on *Coccolithus braarudii*: a discussion of coccolithophorid sensitivities, *Biogeosciences*, 8, 771–777, doi: 10.5194/bg-8-771-2011, 2011.’
- 5 Lines 562–564: add ‘Krumhardt, K. M., Lovenduski, N. S., Debora Iglesias-Rodriguez, M., and Kleypas, J. A.: Coccolithophore growth and calcification in a changing ocean, *Prog. Oceanogr.*, 159, 276–295.’
- 6 Lines 634–635: delete ‘Young, J. R.: Variation in *Emiliana huxleyi* coccolith morphology in samples from the Norwegian EHUX experiment, 1992, *Sarsia*, 79, 417–425, 1994.’

1 **Population-specific responses in physiological rates of *Emiliana huxleyi* to a**  
2 **broad CO<sub>2</sub> range**

3

4 **Yong Zhang,<sup>1,5,\*</sup> Lennart T. Bach,<sup>1</sup> Kai T. Lohbeck,<sup>1,2,6</sup> Kai G. Schulz,<sup>3</sup> Luisa**  
5 **Listmann,<sup>2</sup> Regina Klapper,<sup>4</sup> Ulf Riebesell<sup>1</sup>**

6 <sup>1</sup>Biological Oceanography, GEOMAR Helmholtz-Centre for Ocean Research Kiel,  
7 Kiel, Germany

8 <sup>2</sup>Evolutionary Ecology of Marine Fishes, GEOMAR Helmholtz-Centre for Ocean  
9 Research Kiel, Kiel, Germany

10 <sup>3</sup>Centre for Coastal Biogeochemistry, School of Science, Environment and  
11 Engineering, Southern Cross University, Lismore, NSW, Australia

12 <sup>4</sup>Goethe-University, Institute for Ecology, Evolution and Diversity; Senckenberg  
13 Gesellschaft für Naturforschung, Senckenberg Biodiversity and Climate Research  
14 Centre, Frankfurt am Main, Germany

15 <sup>5</sup>State Key Laboratory of Marine Environmental Science, College of Ocean and Earth  
16 Sciences, Xiamen University (Xiang-An Campus), Xiamen 361102, China

17 <sup>6</sup>Department of Marine Sciences, University of Gothenburg, Gothenburg, Sweden

18

19 Running head: *population response of Emiliana huxleyi to CO<sub>2</sub>*

20

21 \*Correspondence to: Yong Zhang ([zhangyong1983@xmu.edu.cn](mailto:zhangyong1983@xmu.edu.cn))

22 Keywords: CO<sub>2</sub>; coccolithophore; physiological rate; population; strain

23 **Abstract**

24 Although coccolithophore physiological responses to CO<sub>2</sub>-induced changes in  
25 seawater carbonate chemistry have been widely studied in the past, there is limited  
26 knowledge on the variability of physiological responses between populations from  
27 different areas. In the present study, we investigated the ~~population~~-specific responses  
28 of growth, particulate organic (POC) and inorganic carbon (PIC) production rates of 3  
29 populations-17 strains of the coccolithophore *Emiliana huxleyi* from three regions in  
30 the North Atlantic Ocean (Azores: 6 strains, Canary Islands: 5 strains, and Norwegian  
31 coast near Bergen: 6 strains) to a CO<sub>2</sub> partial pressure ( $p\text{CO}_2$ ) range from 120  $\mu\text{atm}$  to  
32 2630  $\mu\text{atm}$ . Physiological rates of each population and individual strain ~~displayed the~~  
33 ~~expected optimum curve responses to the  $p\text{CO}_2$  gradient~~ increased with rising  $p\text{CO}_2$   
34 levels, reached maximum and declined thereafter. Optimal  $p\text{CO}_2$  for growth and POC  
35 production rates and tolerance to low pH (i.e. high proton concentration) was  
36 significantly higher in an *E. huxleyi* population isolated from ~~at~~ the Norwegian  
37 ~~fjord~~coast than in those isolated near the Azores and Canary Islands. This may be due  
38 to the large environmental variability including large  $p\text{CO}_2$  and pH variability  
39 fluctuations in coastal waters off Bergen compared to the rather stable oceanic  
40 conditions at the other two sites. Maximum growth and POC production rates of the  
41 Azores and Bergen populations were similar and significantly higher than that of the  
42 Canary Islands population. ~~One of the reasons may be that the~~ This pattern could be  
43 driven by temperature-CO<sub>2</sub>-interactions where the chosen incubation temperature (16  
44 °C) ~~is~~ was slightly below what strains isolated near the Canary Islands normally

45 experience. Our results indicate adaptation of *E. huxleyi* to their local environmental  
46 conditions and the existence of distinct *E. huxleyi* populations. Within each  
47 population, different growth, POC and PIC production rates at different  $p\text{CO}_2$  levels  
48 indicated strain-specific phenotypic plasticity. The existence of distinct ~~carbonate~~  
49 ~~chemistry~~—responses to changes in carbonate chemistry between and within  
50 populations will likely benefit *E. huxleyi* to acclimate and adapt to rising  $\text{CO}_2$  levels  
51 in the oceans.

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67 **1 Introduction**

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69 Coccolithophores form a layer of calcium carbonate (CaCO<sub>3</sub>) platelets (coccoliths)  
70 around their cells. Coccoliths are of biogeochemical importance due to ballasting of  
71 organic matter with CaCO<sub>3</sub>, a phenomenon which is thought to promote the transport  
72 of organic carbon to the deep ocean (Klaas and Archer, 2002; Rost and Riebesell,  
73 2004). The coccolithophore *Emiliana huxleyi* forms extensive blooms under  
74 favourable light intensity, temperature and nutrient conditions, with different  
75 morphotypes in certain regions (Cook et al., 2011; Henderiks et al., 2012; Smith et al.,  
76 2012; Balch et al., 2014; [Krumhardt et al., 2017](#)).

77 Variable responses of growth, photosynthetic carbon fixation and calcification rates  
78 of different *E. huxleyi* strains to rising CO<sub>2</sub> levels have been reported (Langer et al.,  
79 2009; Hoppe et al., 2011; Müller et al., 2015; Hattich et al, 2017) and are likely a  
80 result of intra-specific variability of genotypes (Langer et al., 2009). Several recent  
81 studies observed optimum curve responses in physiological rates of a single *E. huxleyi*  
82 strain to a broad *p*CO<sub>2</sub> range from about 20 µatm to 5000 µatm, and linked them to  
83 inorganic carbon substrate limitation at low *p*CO<sub>2</sub> and inhibiting H<sup>+</sup> concentrations at  
84 high *p*CO<sub>2</sub> (Bach et al., 2011; 2015; Kottmeier et al., 2016). Until now, studies on the  
85 physiological responses of *E. huxleyi* to rising CO<sub>2</sub> are mostly based on a few  
86 genotypes and little is known about the potential variability in CO<sub>2</sub> and H<sup>+</sup> sensitivity  
87 between and within populations. Recently, several studies found substantial variations  
88 in CO<sub>2</sub> responses for N<sub>2</sub> fixation rates between *Trichodesmium* strains, as well as for



89 growth rates between strains of *Gephyrocapsa oceanica*, *Ostreococcus tauri* and  
90 *Fragilariopsis cylindrus* (Hutchins et al., 2013; Schaum et al., 2013; Pancic et al.,  
91 2015; Hattich et al., 2017). ~~These indicate that~~ Hence, multiple strains, ideally from  
92 geographically distinct regions should be considered for investigating phytoplankton  
93 responses to climate change (Zhang et al., 2014; Blanco-Ameijeiras et al., 2016;  
94 Krumhardt et al., 2017).

95 Oceanographic boundaries formed by both ocean currents and environmental  
96 factors such as temperature, can limit dispersal of marine phytoplankton, reduce gene  
97 flow between geographic populations, and give rise to differentiated populations  
98 (Palumbi, 1994). Different populations were found to show different growth rates for  
99 *E. huxleyi*, *G. oceanica*, and *Skeletonema marinoi* at the same temperatures, and for  
100 *Ditylum brightwellii* at the same light intensities (Brand, 1982; Rynearson and  
101 Armbrust, 2004; Kremp et al., 2012; Zhang et al., 2014). Phenotypic plasticity  
102 describes the ability of a strain to change its morphology or physiology in response to  
103 changing environmental conditions (Bradshaw, 1965). Plasticity can be assessed by  
104 analyzing the reaction norm of one trait and a plastic response may allow a strain to  
105 acclimate ~~to environmental change~~ across an environmental gradient and widen its  
106 bio-geographical distribution (Reusch, 2014; Levis and Pfennig, 2016).

107 In order to better understand how local adaptation affects the physiological  
108 response of *E. huxleyi* to rising CO<sub>2</sub> conditions, we isolated 17 strains from three  
109 regions in the Atlantic Ocean, and assessed growth, carbon fixation and calcification  
110 responses of the population over a pCO<sub>2</sub> range from 120 µatm to 2630 µatm.

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## 112 **2 Materials and methods**

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### 114 **2.1 Cell isolation sites and experimental setup**

115 *Emiliana huxleyi* strains EHGKL B95, B63, B62, B51, B41 and B17 originated from  
116 Raunefjord (Norway 60°18'N, 05°15'E) and were isolated by K. T. Lohbeck in May,  
117 2009 (Lohbeck et al., 2012) at ~ 10 °C in-situ water temperature. *E. huxleyi* strains  
118 EHGLE A23, A22, A21, A19, A13 and A10 originated from coastal waters near the  
119 Azores (38°34'N, 28°42'W) and were isolated by S. L. Eggers in May or June, 2010  
120 at ~ 17 °C in-situ water temperature. *E. huxleyi* strains EHGKL C98, C91, C90, C41  
121 and C35 originated from coastal waters near Gran Canaria (27°58'N, 15°36'W) and  
122 were isolated by K. T. Lohbeck in February, 2014 at ~ 18 °C in-situ water temperature.  
123 Seasonal CO<sub>2</sub> concentration in the surface seawater ranges from 240 µatm to 400  
124 µatm near Bergen, from 320 µatm to 400 µatm around the Azores and from 320 µatm  
125 to 400 µatm around the Canary Islands (Table 1). Monthly surface seawater  
126 temperature ranges from 6.0 to 16.0 °C near Bergen, 15.6 to 22.3 °C around the  
127 Azores and from 18.0 to 23.5 °C around the Canary Islands (Table S1).

128 All 17 strains belong to morphotype A (determined by scanning electron  
129 microscopy) and have been deposited atin the Roscoff culture collection (RCC) under  
130 the official names as shown above. Genetically different isolates, here called strains,  
131 were identified by 5 microsatellite markers (P02E09, P02B12, P02F11, EHMS37,  
132 EHMS15) (Table S2). For a description of primer testing, deoxyribonucleic acid

133 (DNA) extraction, DNA concentration measurements, and polymerase chain reaction  
134 (PCR) protocols see Zhang et al. (2014). The Azores and Bergen strains had been  
135 used earlier by Zhang et al. (2014).

136 The six or five (in case of Canary Islands) strains of each region were used to test  
137 the physiological response to varying CO<sub>2</sub> concentrations at constant total alkalinity  
138 (TA). The experiment was performed in six consecutive incubations, with one strain  
139 from each population (Azores, Bergen, Canary Islands) being cultured at a time (Fig.  
140 S1). Monoclonal populations were always grown in sterile-filtered (0.2 µm diameter,  
141 Sartobran<sup>®</sup> P 300, Sartorius) artificial seawater medium (ASW) as dilute batch  
142 cultures at 200 µmol photons m<sup>-2</sup> s<sup>-1</sup> light intensity under a 16/8 h light/dark cycle  
143 (light period: 5:00 a.m to 9:00 p.m.) at 16 °C which we consider to be ~~the best a~~  
144 compromise for the three different origins of the strains. Nutrients were added in  
145 excess (with nitrate and phosphate concentrations of 64 µmol kg<sup>-1</sup> and 4 µmol kg<sup>-1</sup>,  
146 respectively). For the preparation of ASW and nutrient additions see Zhang et al.  
147 (2014). Calculated volumes of Na<sub>2</sub>CO<sub>3</sub> and hydrochloric acid were added to the ASW  
148 to achieve target CO<sub>2</sub> levels at an average total alkalinity (TA) of 2319 ± 23 µmol kg<sup>-1</sup>  
149 <sup>1</sup> (Pierrot et al., 2006; Bach et al., 2011). Each strain was grown under 11 CO<sub>2</sub> levels  
150 ranging from 115 µatm to 3070 µatm without replicate. Mean response variables of all  
151 strains with a population were calculated and mean CO<sub>2</sub> levels of all strains within a  
152 population ranged from 120 µatm to 2630 µatm. Cells grew in the experimental  
153 conditions for at least 7 generations, which corresponded to 4–7 days depending on  
154 cell division rates. Cells were cultured for 4 days in 120–925 µatm CO<sub>2</sub>, for 5 days in

155 1080–1380  $\mu\text{atm CO}_2$ , and for 6 or 7 days in 1550–2630  $\mu\text{atm CO}_2$ . Initial cell  
156 concentration was 200 cells  $\text{ml}^{-1}$  (estimated from measured pre-culture concentrations  
157 and known dilution) and final cell concentration was lower than 100,000 cells  $\text{ml}^{-1}$ .  
158 Dissolved inorganic carbon (DIC) concentrations and  $p\text{CO}_2$  levels changed less than 7%  
159 and 11%, respectively, during the experimental growth phase.

160

## 161 **2.2 $\text{pH}_T$ and total alkalinity measurements**

162 At 10:00 a.m. on the last day of incubations (at day 4–7 depending on  $\text{CO}_2$   
163 concentration),  $\text{pH}_T$  and TA samples were filtered (0.2  $\mu\text{m}$  diameter, Filtropur S 0.2,  
164 Sarstedt) by gentle pressure and stored at 4°C for a maximum of 14 days. The entire  
165 sampling lasted less than 2 h. The  $\text{pH}_T$  sample bottles were filled with considerable  
166 overflow and closed tightly with no space.  $\text{pH}_T$  was measured spectrophotometrically  
167 (Cary 100, Agilent) using the indicator dye *m*-cresol purple (Sigma-Aldrich) similar  
168 to Carter et al. (2013) with constants of acid dissociation for the protonated and un-  
169 protonated forms reported in Clayton and Byrne (1993). TA was measured by open-  
170 cell potentiometric titration (862 Compact Titrosampler, Metrohm) according to  
171 Dickson et al. (2003). The carbonate system was calculated from measured TA,  $\text{pH}_T$ ,  
172 (assuming 4  $\mu\text{mol kg}^{-1}$  of phosphate and 0  $\mu\text{mol kg}^{-1}$  of silicate) using the  $\text{CO}_2$   
173 System Calculations in MS Excel software (Pierrot et al., 2006) with carbonic acid  
174 constants  $K_1$  and  $K_2$  as determined by Roy et al. (1993).

175

## 176 **2.3 Growth rate measurements**

177 At 1:00 p.m. on the last day of incubation, 25 ml samples were used to measure cell  
178 concentration. Cell concentration was determined within two hours using a Z2 Coulter  
179 Particle Counter (Beckman). Growth rate ( $\mu$ ) was calculated according to:

$$180 \quad \mu = (\ln N_1 - \ln N_0) / d \quad (1)$$

181 where  $N_1$  is cell concentration on the last day of incubation,  $N_0$  is 200 cells mL<sup>-1</sup>, and  
182  $d$  is the time period for growth of algae in days.

183

#### 184 **2.4 Particulate organic (POC) and inorganic (PIC) carbon measurements**

185 At 3:00 p.m. on the last day of incubation, cells for total particulate (TPC) and total  
186 organic (TOC) carbon were filtered onto GF/F filters which were pre-combusted at  
187 500 °C for 8 h. Samples of background particulate carbon (BPC) were determined in a  
188 similar way but using filtered ASW without algae, which was previously adjusted to  
189 target  $p\text{CO}_2$  levels, and allowed to age for about 7 days under incubation conditions  
190 (*see* above). All samples were placed at -20°C. BPC filters were used as blanks to  
191 correct for organic carbon in the medium. TOC and BPC filters were acid fumed.  
192 Afterwards, all filters were dried for 8 h at 60°C. TPC, TOC and BPC were measured  
193 using an Elemental Analyzer (EuroEA, Hekatech GmbH). The percentages of BPC in  
194 TPC were about 20% at cell densities < 10,000 cells ml<sup>-1</sup> and about 10% at cell  
195 densities > 40,000 cells ml<sup>-1</sup>. POC was calculated as the difference between TOC and  
196 BPC. PIC was calculated as the difference between TPC and TOC. POC and PIC  
197 production rates were calculated as:

$$198 \quad \text{POC production rate} = \mu \text{ (d}^{-1}\text{)} \times (\text{TOC} - \text{BPC}) \text{ (pg C cell}^{-1}\text{)} \quad (2)$$

$$\text{PIC production rate} = \mu \text{ (d}^{-1}\text{)} \times (\text{TPC} - \text{TOC}) \text{ (pg C cell}^{-1}\text{)} \quad (3)$$

## 2.5 Data analysis

In a broad  $p\text{CO}_2$  range, physiological rates are expected to initially increase quickly until reaching an optimum and then decline towards further increasing  $\text{CO}_2$  levels (e.g. Krug et al. 2011). Hence we used the following modified Michaelis-Menten equation (Bach et al. 2011) which was fitted to measured cellular growth, POC and PIC production rates. The nonlinear regression model (4) was used to fit growth, POC and PIC production rates and yielding theoretical optimum  $p\text{CO}_2$  and maximum values for each of the three populations (combining the data of five or six strains) (Bach et al., 2011).

$$y = \frac{X \times p\text{CO}_2}{Y + p\text{CO}_2} - s \times p\text{CO}_2 \quad (4)$$

where  $X$  and  $Y$  are fitted parameters, and  $s$  is the sensitivity constant, which indicates the effect of depicts the slope of the decline after optimum  $\text{CO}_2$  levels in response to rising  $\text{H}^+$ . Based on the fitted  $X$ ,  $Y$  and  $s$ , we calculated the  $p\text{CO}_2$  optima ( $K_m$ ) (equation 5) for physiological rates according to equation (5) and  $M$  maximum growth, POC and PIC production rates were calculated by using equation (4) based on  $K_m$  following Bach et al., (2011).

$$K_m = \sqrt{\frac{X \times Y}{s}} - Y \quad (5)$$

The relative values for growth, POC and PIC production rates were calculated as ratios of growth, POC and PIC production rates at each  $p\text{CO}_2$  level to the maximum (highest) rates. We obtained the relative sensitivity constant by fitting function (4)

221 based on relative growth, POC and PIC production rates.

222 A one-way ANOVA was then used to test for statistically significant differences in  
223 theoretical optimum  $p\text{CO}_2$ , maximum value and relative sensitivity constant between  
224 populations. A Tukey HSD test was conducted to determine the differences between  
225 strains from different populations. A Shapiro–Wilk’s analysis was tested to analyze  
226 residual normality. Statistical calculations were carried out using *R* and significance  
227 was shown by  $p < 0.05$ .

228

## 229 **3 Results**

230

### 231 **3.1 Carbonate chemistry parameters**

232 Carbonate system parameters are shown in Table 2. Average  $p\text{CO}_2$  levels of the ASW  
233 ranged from 125  $\mu\text{atm}$  to 2490  $\mu\text{atm}$  for the Azores population, from 120  $\mu\text{atm}$  to  
234 2280  $\mu\text{atm}$  for the Bergen population, and from 130  $\mu\text{atm}$  to 2630  $\mu\text{atm}$  for the  
235 Canary Islands population. Corresponding  $\text{pH}_T$  values of the ASW ranged from 8.46  
236 to 7.33 for the Azores population, from 8.47 to 7.37 for the Bergen population, and  
237 from 8.45 to 7.31 for the Canary Islands population.

238

### 239 **3.2 Measured growth, POC and PIC production rates of each population**

240 Growth rates, POC and PIC production rates of the three *E. huxleyi* populations  
241 increased with rising  $p\text{CO}_2$ , reached a maximum, and then declined with further  $p\text{CO}_2$   
242 increase (Fig. 1). Growth rates of the Azores and Bergen populations were larger than

243 those of the Canary Islands population at all investigated  $p\text{CO}_2$  levels (Fig. 1a). With  
244 rising  $p\text{CO}_2$  levels beyond the  $p\text{CO}_2$  optimum, decline in growth rates was more  
245 pronounced in the Azores and Canary Islands populations than in the Bergen  
246 population (Fig. 1b).

247 Measured POC production rates of the Azores and Bergen populations were larger  
248 than those of the Canary Islands population at all  $p\text{CO}_2$  levels (Fig. 1c) and decline in  
249 POC production rates with increasing  $p\text{CO}_2$  levels beyond the  $p\text{CO}_2$  optimum was  
250 larger in the Azores and Canary Islands populations than in the Bergen population  
251 (Fig. 1d).

252 Measured PIC production rates at investigated  $p\text{CO}_2$  levels did not show significant  
253 differences among the Azores, Bergen and Canary Islands populations (Fig. 1e).  
254 Exceptions were that at 365–695  $\mu\text{atm}$ , PIC production rates of the Azores population  
255 were larger than those of the Canary Islands population (all  $p < 0.05$ ).

256

### 257 **3.3 Physiological responses of populations to $p\text{CO}_2$**

258 Calculated optimum  $p\text{CO}_2$  for growth, POC and PIC production rates of the Bergen  
259 population were significantly larger than those of the Azores and Canary Islands  
260 populations (all  $p < 0.05$ ) (Fig. 2a–c). Optimum  $p\text{CO}_2$  for these physiological rates  
261 between the Azores and Canary Islands population were not different (all  $p > 0.1$ ).

262 Calculated maximum growth rates, POC and PIC production rates were not  
263 significantly different between the Azores and the Bergen populations (all  $p > 0.1$ )  
264 (Fig. 2d–f). Maximum growth rate and POC production rate of the Canary Islands



265 population were significantly lower than those of the Azores and Bergen populations  
266 (both  $p < 0.01$ ) (Fig. 2d,e). Maximum PIC production rates of the Canary Islands  
267 population were significantly lower than that of the Azores population ( $p < 0.05$ ),  
268 while there was no difference to the Bergen population ( $p > 0.1$ ) (Fig. 2f).

269 Fitted relative sensitivity constants for growth and POC production rates of the  
270 Bergen population were significantly lower than those of the Azores and Canary  
271 Islands populations ( $p < 0.01$ ) (Fig. 2g, h). Fitted relative sensitivity constants for  
272 growth and POC production rates between the Azores and Canary Islands populations  
273 were not significantly different ( $p > 0.1$ ). Fitted relative sensitivity constants for PIC  
274 production rates did not show difference among three populations ( $p = 0.13$ ) (Fig. 2i).

275

### 276 **3.4 Physiological responses of individual strains to $p\text{CO}_2$**

277 Measured growth rates, POC and PIC production rates of 17 *E. huxleyi* strains showed  
278 optimum curve response patterns to the broad  $p\text{CO}_2$  gradient (Fig. 3). Variations in  
279 calculated  $p\text{CO}_2$  optima, maximum values and relative sensitivity constants of  
280 physiological rates were found between the strains (Table 3).

281 For all strains within each population, optimum  $p\text{CO}_2$  of POC production rates  
282 were larger than optimum  $p\text{CO}_2$  of growth rates or PIC production rates with the  
283 exception of optimum  $p\text{CO}_2$  of POC and PIC production rates of *E. huxleyi* strain  
284 EHGLE A22 (Table 3). Compared to the Azores and Bergen populations, strains  
285 isolated near the Canary Islands showed larger variation in optimum  $p\text{CO}_2$  of PIC  
286 production rates. Within the Azores population, variations in maximum values ( $V_{\max}$ )

287 and relative sensitivity constants ( $r_s$ ) of growth, POC and PIC production rates of all  
288 strains were larger than those within the Bergen and Canary Islands populations (Fig.  
289 3).

290

#### 291 **4 Discussion**

292

293 We investigated growth, POC and PIC production rates of 17 *E. huxleyi* strains from  
294 three populations to a broad  $p\text{CO}_2$  range (120–2630  $\mu\text{atm}$ ). The three populations  
295 differed significantly in growth and POC production rates at the investigated  $p\text{CO}_2$   
296 levels. The reaction norms of the individual strains and populations equaled an  
297 optimum curve for all physiological rates (Figs. 1 and 3). However, we detected  
298 distinct  $p\text{CO}_2$  optima for growth, POC and PIC production rates, and different  $\text{H}^+$   
299 sensitivities for growth and POC production rates among them (Fig. 2). These results  
300 indicate the existence of distinct populations in the cosmopolitan coccolithophore *E.*  
301 *huxleyi*.

302 In comparison to the Azores and Canary Islands populations, variability in growth  
303 rates between strains of the Bergen population was smaller even though they had  
304 higher growth rates at all  $p\text{CO}_2$  levels (Fig. 3). Furthermore, the Bergen population  
305 showed significantly higher  $p\text{CO}_2$  optima and lower  $\text{H}^+$  sensitivity for growth and  
306 POC production rates (Fig. 2). These findings indicate that the Bergen population may  
307 be more tolerant to changing carbonate chemistry in terms of its growth and  
308 photosynthetic carbon fixation rates. The Bergen strains were isolated from coastal

309 waters, while the Azores and Canary Islands strains were isolated from a more  
310 oceanic environment. Seawater carbonate chemistry of coastal waters is usually more  
311 dynamic than in the open ocean (Cai, 2011). In fact, previous studies have reported  
312 that CO<sub>2</sub> and pH variability of the seawater off Bergen was larger than off the Azores  
313 and Canary Islands (Table 1). In addition, due to riverine input, seawater upwelling  
314 and metabolic activity of plankton communities, environmental variability in coastal  
315 waters are larger than in open-ocean ecosystems (Duarte and Cerbrian, 1996). Doblin  
316 and van Sebille (2016) suggested that phytoplankton populations should be constantly  
317 under selection when experiencinged with changing environmental conditions. In this  
318 case, the Bergen population, exposed to larger CO<sub>2</sub> or pH fluctuations, may have  
319 acquired a higher capacity to acclimate to changing carbonate chemistry resulting in a  
320 higher tolerance (or lower sensitivity) to rising CO<sub>2</sub> levels. In contrast, the Azores and  
321 Canary Islands populations experience similar, less variable seawater carbonate  
322 chemistry conditions in their natural environment, which could explain why they also  
323 show similar *p*CO<sub>2</sub> optima and H<sup>+</sup> sensitivity for physiological rates (Fig. 2).

324 In an earlier study (Zhang et al., 2014), growth rates of the same Azores and  
325 Bergen strains as used here were measured at 8–28 °C. While at 26–28 °C the Bergen  
326 strains grew slower than the Azores strains, at 8 °C the Azores strains grew slower  
327 than the Bergen strains. This illustrates nicely that local temperature adaptation can  
328 significantly affect growth of *E. huxleyi* strains in laboratory experiments.

329 Considering these findings and the temperature ranges of the three isolationed  
330 locations (Table S1), the incubation temperature of 16 °C used in the present study

331 was lower than the minimum sea surface temperature (SST) commonly recorded at  
332 the Canary Islands. In contrast, SSTs of 16 °C and lower have been reported for  
333 Azores and Bergen waters (Table S1). When exposed to 16 °C, growth rate of the  
334 Canary Islands population might have been already below their optimum and hence  
335 significantly reduced in comparison to the other populations ~~thus it grew slower than~~  
336 ~~the other populations~~ (Fig. 2d).

337 Furthermore, compared to the Canary Islands population, the Azores population  
338 had higher maximum growth and POC production rates, and similar optimum CO<sub>2</sub> for  
339 these physiological rates. Again, this might be related to sub-optimal incubation  
340 conditions as temperature has been found to significantly modulate CO<sub>2</sub> responses in  
341 coccolithophores in terms of maximum rates, CO<sub>2</sub> optima and half-saturation, and H<sup>+</sup>  
342 sensitivity (De Bodt et al., 2010; Sett et al., 2014; Gafar et al., 2018; Gafar and Schulz,  
343 2018). In a similar fashion light can also modulate CO<sub>2</sub> responses, hence different  
344 requirements by strains adapted to different light availabilities could also explain our  
345 observations (Zhang et al., 2015; Gafar et al., 2018; Gafar and Schulz, 2018). One of  
346 ~~the reasons may be that compared to the Azores and Bergen populations, 16 °C likely~~  
347 ~~causes lower the carbon uptake and carbon use efficiency of the Canary Islands~~  
348 ~~population (Sett et al., 2014).~~ Thus, with rising CO<sub>2</sub>, growth, photosynthetic carbon  
349 fixation and calcification rates of the Canary Islands population cannot increase as  
350 much as in the Azores and Bergen populations. In addition, the Canary Islands  
351 population showed smallest variability in optimum pCO<sub>2</sub> and maximum values for  
352 growth and POC production rates (Fig. 2). The reason may be that low incubation

353 temperature predominantly limited growth and POC production rates of the Canary  
354 Islands population, and decreased the sensitivities of these physiological rates to  
355 rising  $p\text{CO}_2$ .

356 Before we started this experiment, strains isolated from the Azores, Bergen and  
357 Canary Islands grew as stock cultures at 15 °C and 400  $\mu\text{atm}$  for 4 years, 5 years and  
358 3 months, respectively. Schaum et al. (2015) provide evidence that long-term  
359 laboratory incubation affects responses of phytoplankton to different  $p\text{CO}_2$  levels.  
360 Thus, it is conceivable that the same selection history in the laboratory incubation  
361 may contribute to a more similar response of growth, POC and PIC production rates  
362 between the Azores and Bergen populations at low  $p\text{CO}_2$  levels (Fig. 1).

363 Our results indicate that *E. huxleyi* populations are adapted to the specific  
364 environmental conditions of their origin, resulting in different responses to increasing  
365  $p\text{CO}_2$  levels. The ability to adapt to diverse environmental conditions is ~~reflected in~~  
366 supposed to be one reason for the global distribution of *E. huxleyi* (Paasche, 2002),  
367 spanning a temperature range of about 30 °C. The optimum temperature for growth of  
368 the Bergen population was about 22 °C and was 5 °C higher than the maximum SST  
369 in Bergen waters (Zhang et al. 2014). Furthermore, in comparison to the Azores and  
370 Canary Islands populations, larger optimum  $p\text{CO}_2$  of growth rate indicates that the  
371 Bergen population may benefit more from the rising  $\text{CO}_2$  levels at increasing  
372 temperatures. PIC : POC ratios of the Azores and Bergen populations declined with  
373 rising  $p\text{CO}_2$ , whereas PIC : POC ratios of the Canary Islands population were rather  
374 constant (Fig. S6). As changes in PIC : POC ratios of coccolithophore blooms may

375 impact on the biological carbon pump, different regions might see different changes  
376 in the future ocean. In natural seawater, due to ocean currents and gene flow,  
377 populations at any given location may get replaced by ~~populations~~ immigrant  
378 genotypes transported there from other locations ~~when having a higher potential to~~  
379 ~~adapt to a changing environment~~ (Doblin and van Sebille, 2016). In addition, *E.*  
380 *huxleyi* ~~take up is thought to utilize~~  $\text{HCO}_3^-$  ~~to calcify and for calcification which~~  
381 generates protons, and increase in proton concentration may mitigate the potential of  
382 the ocean to absorb atmospheric  $\text{CO}_2$  (Paasche, 2002). ~~Thus, due to population-~~  
383 ~~specific growth and PIC production rates or quotas, changes in species composition,~~  
384 ~~corresponding changes in PIC productions, may affect the ability of the ocean to take~~  
385 ~~up  $\text{CO}_2$ .~~

386 Within a population, individual strains showed different growth, POC and PIC  
387 production rates at different  $p\text{CO}_2$  levels, indicating phenotypic plasticity of  
388 individual strains (Reusch, 2014). Phenotypic plasticity constitutes an advantage for  
389 individual strains to acclimate and adapt to elevated  $p\text{CO}_2$  by changing ~~their~~ fitness-  
390 relevant traits and potentially to attenuate the short-term effects of changing  
391 environments on fitness-relevant traits (Schaum et al., 2013). ~~Additionally, our results~~  
392 ~~also suggest that strain-specific PIC quota may be the basis of variation in coecoliths~~  
393 ~~of *E. huxleyi* within the morphotype A (Fig. S3) (Young, 1994; Paasche, 2002).~~

394 The strain-specific  $\text{CO}_2$ -response curves revealed considerable physiological  
395 diversity in co-occurring strains (Fig. 3). Physiological variability makes a population  
396 more resilient, ~~and~~ increases its ability to persist in variable environments and

397 potentially forms the basis for selection (Gsell et al., 2012; Hattich et al., 2017). It is  
398 clear that other environmental factors such as light intensity, temperature and nutrient  
399 concentration affect the responses of physiological rates of individual *E. huxleyi*  
400 strains to changing carbonate chemistry, and thus change the physiological variability  
401 within populations (Zhang et al., 2015; Feng et al., 2017). However, different  
402 sensitivities and requirements of each strain to the variable environments can allow  
403 strains to co-exist within a population in the natural environment (Hutchinson, 1961;  
404 Reed et al., 2010; Krueger-Hadfield et al., 2014). In a changing oceans, strain  
405 succession is likely to occur and shift the population composition (Blanco-Ameijeiras  
406 et al., 2016; Hattich et al., 2017). Strains with higher growth rates or other  
407 competitive abilities may outcompete ~~others~~ ~~strains in the oceans~~ (Schaum et al.,  
408 2013). Further, a Ssignificant positive correlation between growth and POC  
409 production rate or POC quota (Fig. 4S5) ~~suggests~~ indicates that the dominated ing  
410 strains ~~can~~ will also take up or fix dissolved inorganic carbon faster ~~from the oceans or~~  
411 ~~fix carbon faster.~~ When extrapolated to the ocean, *E. huxleyi* blooms ~~This~~ may  
412 increase the potential of the oceans to absorb CO<sub>2</sub> from the atmosphere ~~or the~~ and its  
413 carbon storage capacity ~~of the oceans when large *E. huxleyi* blooms occur~~ (Blanco-  
414 Ameijeiras et al., 2016), which ~~will~~ has the potential to mitigate rising CO<sub>2</sub> levels in  
415 the atmosphere.

416

## 417 **5 Conclusions**

418 In the present study, we found population-specific responses in physiological rates of

419 *E. huxleyi* to a broad  $p\text{CO}_2$  range, which may have arisen from local adaptation to  
420 environmental conditions at their origins. The existence of distinct *E. huxleyi*  
421 populations and phenotypic plasticity of individual strains may both be important for  
422 *E. huxleyi* when adapting to natural environmental variability and to ongoing climate  
423 changes. Our results suggest that when assessing phytoplankton responses to  
424 changing environments on a global scale, variability in population ~~or~~and strain  
425 responses need to be considered, and  $\text{CO}_2$  response was modulated by other  
426 environmental factors such as temperature and light intensity.

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442 *Author contributions.* YZ, LTB, UR designed the experiment. YZ, LL, RK performed  
443 the experiment. YZ prepare the manuscript and all authors analysed the data,  
444 reviewed and improved the manuscript.

445

446

447 *Competing interests.* The authors declare that they have no conflict of interest.

448

449

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463 **References**

- 464 Bach, L. T., Riebesell, U., Gutowska, M. A., Federwisch, L., and Schulz, K. G.: A  
465 unifying concept of coccolithophore sensitivity to changing carbonate chemistry  
466 embedded in an ecological framework, *Prog. Oceanogr.*, 135, 125–138, doi:  
467 10.1016/j.pocean.2015.04.012, 2015.
- 468 Bach, L. T., Riebesell, U., and Schulz, K. G.: Distinguishing between the effects of  
469 ocean acidification and ocean carbonation in the coccolithophore *Emiliana*  
470 *huxleyi*, *Limnol. Oceanogr.*, 56, 2040–2050, doi: 10.4319/lo.2011.56.6.2040, 2011.
- 471 Balch, W. M., Drapeau, D. T., Bowler, B. C., Lyczkowski, E. R., Lubelczyk, L. C.,  
472 Painter, S. C., and Poulton, A. J.: Surface biological, chemical, and optical  
473 properties of the Patagonian Shelf coccolithophore bloom, the brightest waters of  
474 the Great Calcite Belt, *Limnol. Oceanogr.*, 59, 1715–1732, doi:  
475 10.4319/lo.2014.59.5.1715, 2014.
- 476 Blanco-Ameijeiras, S., Lebrato, M., Stoll, H. M., Iglesias-Rodriguez, D., Müller, M.  
477 N., Méndez-Vicente, A., and Oschlies, A: Phenotypic variability in the  
478 coccolithophore *Emiliana huxleyi*, *PLoS ONE*, 11, e0157697, doi:  
479 10.1371/journal.pone.0157697, 2016.
- 480 Bradshaw, A. D.: Evolutionary significance of phenotypic plasticity in plants, *Adv.*  
481 *Genet.*, 13, 115–155, doi: 10.1016/S0065-2660(08)60048-6, 1965.
- 482 Brand, L. E.: Genetic variability and spatial patterns of genetic differentiation in the  
483 reproductive rates of the marine coccolithophores *Emiliana huxleyi* and  
484 *Gephyrocapsa oceanica*, *Limnol. Oceanogr.*, 27, 236–245, doi:

485 10.4319/lo.1982.27.2.0236, 1982.

486 Cai W. J.: Estuarine and coastal ocean carbon paradox: CO<sub>2</sub> sinks or sites of terrestrial  
487 carbon incineration?, *Ann. Rev. Mar. Sci.*, 3, 123–145, doi: 10.1146/annurev-  
488 marine-120709-142723, 2011.

489 Carter, B. R., Radich, J. A., Doyle, H. L., and Dickson, A. G.: An automated system  
490 for spectrophotometric seawater pH measurements, *Limnol. Oceanogr.: Methods*,  
491 11, 16–27, doi: 10.4319/lom.2013.11.16, 2013.

492 Clayton, T. D., and Byrne, R. H.: Spectrophotometric seawater pH measurements—  
493 Total hydrogen-ion concentration scale calibration of m-cresol purple and at-sea  
494 results, *Deep Sea Res. I*, 40, 2115–2129, doi: 10.1016/0967-0637(93)90048-8,  
495 1993.

496 Cook, S. S., Whittock, L., Wright S. W., and Hallegraeff, G. M.: Photosynthetic  
497 pigment and genetic differences between two southern ocean morphotypes of  
498 *Emiliana huxleyi* (Haptophyta), *J. Phycol.*, 47, 615–626, doi: 10.1111/j.1529-  
499 8817.2001.00992.x, 2011.

500 Dickson, A. G., Afghan, J. D., and Anderson, G. C.: Reference materials for oceanic  
501 CO<sub>2</sub> analysis: a method for the certification of total alkalinity, *Mar. Chem.*, 80,  
502 185–197, doi: 10.1016/S0304-4203(02)00133-0, 2003.

503 Doblin, M. A., and van Sebille, E.: Drift in ocean currents impacts intergenerational  
504 microbial exposure to temperature, *Proc. Natl. Acad. Sci. USA.*, 113, 5700–5705,  
505 doi: 10.1073/pnas.1521093113, 2016.

506 [Duarte, C. M., and Cerbrian, J.: The fate of marine autotrophic production, \*Limnol.\*](#)

507 [Oceanogr., 41, 1758–1766, 1996.](#)

508 Feng Y. Y., Roleda, M. Y., Armstrong, E., Boyd, P. W., and Hurd, C. L.:

509 Environmental controls on the growth, photosynthetic and calcification rates of a

510 Southern Hemisphere strain of the coccolithophore *Emiliana huxleyi*, Limnol.

511 Oceanogr., 62, 519–540, doi: 10.1002/lno.10364, 2017.

512 [Gafar, N. A., Eyre, B. D., and Schulz, K. G. : A conceptual model for projecting](#)

513 [coccolithophorid growth, calcification and photosynthetic carbon fixation rates](#)

514 [in response to global ocean change, Front. Mar. Sci., 4, 433, doi:](#)

515 [10.3389/fmars.2017.00433, 2018.](#)

516 [Gafar, N. A., and Schulz, K. G. : A niche comparison of \*Emiliana huxleyi\* and](#)

517 [\*Gephyrocapsa oceanica\* and potential effects of climate change, Biogeosci.](#)

518 [Discuss., doi: 10.5194/bg-2018-88.](#)

519 González-Dávila, M., and Santana-Casiano, M.: Seasonal and interannual variability

520 of sea-surface carbon dioxide species at the European Station for Time Series in

521 the Ocean at the Canary Islands (ESTOC) between 1996 and 2000, Glob.

522 Biogeochem. Cycles, 17, 1076, doi: 10.1029/2002GB001993, 2003.

523 Gsell, A. S., de Senerpont-Domis, L. N., Przytulska-Bartosiewicz, A., Mooij, W. M.,

524 van Donk, E, and Ibelings, B. W.: Genotype-by-temperature interactions may help

525 to maintain clonal diversity in *Asterionella formosa* (Bacillariophyceae), J.

526 Phycol., 48, 1197–1208, doi: 10.1111/j.1529-8817.2012.01205.x, 2012.

527 Hattich, G. S. I., Listmann, L., Raab, J., Ozod-Seradj, D., Reusch, T. B. H., and

528 Matthiessen, B.: Inter- and intraspecific phenotypic plasticity of three

529 phytoplankton species in response to ocean acidification, *Biol. Lett.*, 13,  
530 20160774, doi: 10.1098/rsbl.2016.0774, 2017.

531 Henderiks, J., Winter, A., Elbrächter, M., Feistel, R., van der Plas, A., Nausch, G.,  
532 and Barlow, R.: Environmental controls on *Emiliana huxleyi* morphotypes in the  
533 Benguela coastal upwelling system (SE Atlantic), *Mar. Ecol. Prog. Ser.*, 448, 51–  
534 66, doi:10.3354/meps09535, 2012.

535 Hoppe, C. J. M., Langer, G., and Rost, B.: *Emiliana huxleyi* shows identical  
536 responses to elevated pCO<sub>2</sub> in TA and DIC manipulations, *J. Exp. Mar. Biol.*  
537 *Ecol.*, 406, 54–62, doi: 10.1016/j.jembe.2011.06.008, 2011.

538 Hutchins, D. A., Fu, F. X., Webb, E. A., Walworth, N., and Tagliabue, A.: Taxon-  
539 specific response of marine nitrogen fixers to elevated carbon dioxide  
540 concentrations, *Nat. Geosci.*, 6, 790–795, doi: 10.1038/ngeo1858, 2013.

541 Hutchinson, G. E.: The paradox of the plankton, *Am. Nat.*, 95, 137–145, 1961.

542 Klaas, C, and Archer, D. E.: Association of sinking organic matter with various types  
543 of mineral ballast in the deep sea: Implications for the rain ratio, *Glob.*  
544 *Biogeochem. Cycles*, 16, 1116, doi: 10.1029/2001GB001765, 2002.

545 Kottmeier, D. M., Rokitta, S. D., and Rost, B.: H<sup>+</sup>-driven increase in CO<sub>2</sub> uptake and  
546 decrease in HCO<sub>3</sub><sup>-</sup> uptake explain coccolithophores' acclimation responses to  
547 ocean acidification, *Limnol. Oceanogr.*, 61, 2045–2057, doi: 10.1002/lno.10352,  
548 2016.

549 Kremp, A., Godhe, A., Egardt, J., Dupont, S., Suikkanen, S., Casabianca, S., and  
550 Penna, A.: Intraspecific variability in the response of bloom-forming marine

551 microalgae to changed climate conditions, *Ecol. Evol.*, 2, 1195–1207, doi:  
552 10.1002/ece.3.245, 2012.

553 Krueger-Hadfield, S. A., Balestreri, C., Schroeder, J., Highfield, A., Helaou ä, P.,  
554 Allum, J., Moate, R., Lohbeck, K. T., Miller, P. I., Riebesell, U., Reusch, T. B. H.,  
555 Rickaby, R. E. M., Young, J., Hallegraeff, G., Brownlee, C., and Schroeder, D. C.:  
556 Genotyping an *Emiliania huxleyi* (prymnesiophyceae) bloom event in the North  
557 Sea reveals evidence of asexual reproduction, *Biogeosciences*, 11, 5215–5234, doi:  
558 10.5194/bg-11-5215-2014, 2014.

559 [Krug, S. A., Schulz, K. G., and Riebesell, U. : Effects of changes in carbonate](#)  
560 [chemistry speciation on \*Coccolithus braarudii\*: a discussion of coccolithophorid](#)  
561 [sensitivities, \*Biogeosciences\*, 8, 771–777, doi: 10.5194/bg-8-771-2011, 2011.](#)

562 [Krumhardt, K. M., Lovenduski, N. S., Debora Iglesias-Rodriguez, M., and Kleypas, J.](#)  
563 [A.: Coccolithophore growth and calcification in a changing ocean, \*Prog.\*](#)  
564 [Oceanogr.](#), 159, 276–295.

565 Langer, G., Nehrke, G., Probert, I., Ly, J., and Ziveri, P.: Strain-specific responses of  
566 *Emiliania huxleyi* to changing seawater carbonate chemistry, *Biogeosciences*, 6,  
567 2637–2646, doi: 10.5194/bg-6-2637-2009, 2009.

568 Levis, N. A., and Pfennig, D. W.: Evaluating ‘plasticity-first’ evolution in nature: key  
569 criteria and empirical approaches, *Trends Eco. Evol.*, 31, 563–574, doi:  
570 10.1016/j.tree.2016.03.012, 2016.

571 Lohbeck, K. T., Riebesell, U., and Reusch, T. B. H.: Adaptive evolution of a key  
572 phytoplankton species to ocean acidification, *Nat. Geosci.*, 5, 346–351, doi:

573 10.1038/ngeo1441, 2012.

574 Müller, M. N., Trull, T. W., and Hallegraeff, G. M.: Differing responses of three  
575 Southern Ocean *Emiliana huxleyi* ecotypes to changing seawater carbonate  
576 chemistry, *Mar. Ecol. Prog. Ser.*, 531, 81–90, doi: 10.3354/meps11309, 2015.

577 Omar, A. M., Olsen, A., Johannessen, T., Hoppema, M., Thomas, H., and Borges, A.  
578 V.: Spatiotemporal variations of  $f\text{CO}_2$  in the North Sea, *Ocean Sci.*, 6, 77–89,  
579 doi:10.5194/os-6-77-2010, 2010.

580 Paasche, E.: A review of the coccolithophorid *Emiliana huxleyi* (Prymnesiophyceae),  
581 with particular reference to growth, coccolith formation, and calcification-  
582 photosynthesis interactions, *Phycologia*, 40, 503–529, doi: 10.2216/i0031-8884-  
583 40-6-503.1, 2002.

584 Palumbi, S. R.: Genetic divergence, reproductive isolation, and marine speciation.  
585 *Ann. Rev. Ecol. Evol. Syst.*, 25, 547–572, doi:  
586 10.1146/annurev.es.25.110194.002555, 1994.

587 Pancic, M., Hansen, P. J., Tammilehto, A., and Lundholm, N.: Resilience to  
588 temperature and pH changes in a future climate change scenario in six strains of  
589 the polar diatom *Fragilariopsis cylindrus*, *Biogeosciences*, 12, 4235–4244, doi: 10.  
590 5194/bg-12-4235-2015, 2015.

591 Pierrot, D., Lewis, E., and Wallace, D. W. R.: MS Excel program developed for CO<sub>2</sub>  
592 system calculations, ORNL/CDIAC-105, Carbon Dioxide Information Analysis  
593 Centre, Oak Ridge National Laboratory, U.S., Department of Energy, 2006.

594 Reed, T. E., Waples, R. S., Schindler, D. E., Hard, J. J., and Kinnison, M. T.:

595 Phenotypic plasticity and population viability: the importance of environmental  
596 predictability, *Roc. R. Soc. B*, 277, 3391, doi: 10.1098/rspb.2010.0771, 2010.

597 Reusch, T. B. H.: Climate change in the oceans: Evolutionary versus phenotypically  
598 plastic responses of marine animals and plants, *Evol. Appl.*, 7, 104–122, doi:  
599 10.1111/eva.12109, 2014.

600 R ós, A. F., Pérez, F. F., Álvarez, M., Mintrop, L., González-Dávila, M., Santana-  
601 Casiano, J. M., Lefèvre, L., and Watson, A. J.: Seasonal sea-surface carbon  
602 dioxide in the Azores area, *Mar. Chem.*, 96, 35–51, doi:  
603 10.1016/j.marchem.2004.11.001, 2005.

604 Rost, B., and Riebesell, U.: Coccolithophores and the biological pump: responses to  
605 environmental changes, in: *Coccolithophores – From Molecular Biology to*  
606 *Global Impact*, edited by: Thierstein, H. R. and Young, J. R., Springer, Berlin, 99–  
607 125, 2004.

608 Roy, R. N., Roy, L. N., Lawson, M., Vogel, K. M., Moore, C. P., Davis W., and  
609 Millero, F. J.: Thermodynamics of the dissociation of boric acid in seawater at S 5  
610 35 from 0 degrees C to 55 degrees C, *Mar. Chem.*, 44, 243–248,  
611 doi:10.1016/0304-4203(93)90206-4, 1993.

612 Rynearson, T. A., and Armbrust, E. V.: Genetic differentiation among populations of  
613 the planktonic marine diatom *Ditylum Brightwellii* (Bacillariophyceae), *J. Phycol.*,  
614 40, 34–43, doi: 10.1046/j.1529-8817.2004.03089.x, 2004.

615 Schaum, E., Rost, B., Millar, A. J., and Collins, S.: Variation in plastic responses of a  
616 globally distributed picoplankton species to ocean acidification, *Nat. Clim.*



617 Change, 3, 298–302, doi: 10.1038/nclimate1774, 2013.

618 Schaum, E., Rost, B., Collins, S.: Environmental stability affects phenotypic evolution  
619 in a globally distributed marine picoplankton, *The ISME Journal*, 10, 75–84, doi:  
620 10.1038/ismej.2015.102, 2015.

621 Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.:  
622 Temperature modulates coccolithophorid sensitivity of growth, photosynthesis  
623 and calcification to increasing seawater  $p\text{CO}_2$ , *PLoS ONE*, 9, e88308, doi:  
624 10.1371/journal.pone.0088308, 2014.

625 Smith, H. E. K., Tyrrell, T., Charalampopoulou, A., Dumousseaud, C., Legge, O. J.,  
626 Birchenough, S., Pettit, L. R., Garley, R., Hartman, S. E., Hartman, M. C., Sagoo,  
627 N., Daniels, C. J., Achterberg, E. P., and Hydes, D. J.: Prodominance of heavily  
628 calcified coccolithophores at low  $\text{CaCO}_3$  saturation during winter in the Bay of  
629 Biscay, *Proc. Natl. Acad. Sci. USA*, 109, 8845–8849, doi:  
630 10.1073/pnas.1117508109, 2012.

631 Wisshak, M., Form, A., Jakobsen, J., and Freiwald, A.: Temperate carbonate cycling  
632 and water mass properties from intertidal to bathyal depths (Azores),  
633 *Biogeosciences*, 7, 2379–2396, doi:10.5194/bg-7-2379-2010, 2010.

634 ~~Young, J. R.: Variation in *Emiliana huxleyi* coccolith morphology in samples from~~  
635 ~~the Norwegian EHUX experiment, 1992, *Sarsia*, 79, 417–425, 1994.~~

636 Zhang, Y., Klapper, R., Lohbeck, K. T., Bach, L. T., Schulz, K. G., Reusch, T. B. H.,  
637 and Riebesell, U.: Between- and within-population variations in thermal reaction  
638 norms of the coccolithophore *Emiliana huxleyi*, *Limnol. Oceanogr.*, 59, 1570–

639 1580, doi: 10.4319/lo.2014.59.5.1570, 2014.

640 Zhang, Y., Bach, L. T., Schulz, K. G., and Riebesell, U.: The modulating effect of  
641 light intensity on the response of the coccolithophore *Gephyrocapsa oceanica* to  
642 ocean acidification, *Limnol. Oceanogr.*, 60, 2145–2157, doi:10.1002/lno.10161,  
643 2015.

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662 **Figure Legends**

663 **Figure 1.** Optimum curve responses of measured and relative growth, particulate  
664 organic (POC) and inorganic carbon (PIC) production rates of three *Emiliana huxleyi*  
665 populations to a  $p\text{CO}_2$  range from 120  $\mu\text{atm}$  to 2630  $\mu\text{atm}$ . Responses of measured (a)  
666 and relative (b) growth rates to  $p\text{CO}_2$ . Responses of measured (c) and relative (d)  
667 POC production rates to  $p\text{CO}_2$ . Responses of measured (e) and relative (f) PIC  
668 production rates to  $p\text{CO}_2$ . Using the nonlinear regression model derived by Bach et al.  
669 (2011), the curves were fitted based on average growth, POC and PIC production  
670 rates of six strains from the Azores and Bergen, and of five strains from the Canary  
671 Islands. Vertical error bars represent standard deviations of six growth, POC and PIC  
672 production rates for the Azores and Bergen populations, and five growth, POC and  
673 PIC production rates for the Canary Islands population. Horizontal error bars  
674 represent standard deviations of six  $p\text{CO}_2$  levels for the Azores and Bergen  
675 populations and five  $p\text{CO}_2$  levels for the Canary Islands populations. At the  
676 population levels, 120  $\mu\text{atm}$  and 2630  $\mu\text{atm}$  was the lowest and highest  $p\text{CO}_2$  level,  
677 respectively.

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679 **Figure 2.** Calculated optimum  $p\text{CO}_2$ , calculated maximum value and fitted relative  
680 sensitivity constant of growth, POC and PIC production rates of each population. (a)  
681 optimum  $p\text{CO}_2$  of growth rate; (b) optimum  $p\text{CO}_2$  of POC production rates; (c)  
682 optimum  $p\text{CO}_2$  of PIC production rates; (d) maximum growth rate, (e) maximum  
683 POC production rate, (f) maximum PIC production rate; (g) relative sensitivity

684 constant of growth rate; **(h)** relative sensitivity constant of POC production rate; **(i)**  
685 relative sensitivity constant of PIC production rate. The line in the middle of each box  
686 indicates the mean of 6 or 5 optimum  $p\text{CO}_2$ , 6 or 5 maximum values, and 6 or 5  
687 relative sensitivity constants for growth, POC and PIC production rates in each  
688 population. Bars indicate the 99% confidence interval. The maximum or minimum  
689 data is shown as the small line on the top or bottom of the bar, respectively. Letters in  
690 each panel represent statistically significant differences (Tukey HSD,  $p < 0.05$ ).

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692 **Figure 3.** Optimum curve responses of growth, POC and PIC production rates of  
693 individual *E. huxleyi* strains in the Azores (left), Bergen (medium) and Canary Islands  
694 (right) populations to a  $\text{CO}_2$  range from 115  $\mu\text{atm}$  to 3070  $\mu\text{atm}$ . Growth rates of each  
695 strain as a function of  $p\text{CO}_2$  within the Azores **(a)**, Bergen **(b)** and Canary Islands **(c)**  
696 populations. POC production rates of each strain as a function of  $p\text{CO}_2$  within the  
697 Azores **(d)**, Bergen **(e)** and Canary Islands **(f)** populations. PIC production rates of  
698 each strain as a function of  $p\text{CO}_2$  within the Azores **(g)**, Bergen **(h)** and Canary  
699 Islands **(i)** populations. At the strain levels, 115  $\mu\text{atm}$  and 3070  $\mu\text{atm}$  was the lowest  
700 and highest  $p\text{CO}_2$  level, respectively.

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706 **Table 1.** Surface seawater CO<sub>2</sub> levels and pH at the Azores, Bergen and Canary  
 707 Islands.

	<b>Location</b>	<b>Mean seasonal CO<sub>2</sub> (µatm)</b>	<b>Mean seasonal pH (total scale)</b>	<b>CO<sub>2</sub> variability (µatm)</b>	<b>References</b>
Azores	38°34'N, 28°42'W	320 – 400	8.005 – 8.05	80	R ós et al., 2005 Wisshak et al., 2010
Bergen	60°18'N, 05°15'E	240 – 400	7.98 – 8.22	200	Omar et al., 2010
Canary Islands	27°58'N, 15°36'W	320 – 400	8.005 – 8.05	80	Gonz ález-D ávila et al., 2003

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725 **Table 2.** Carbonate chemistry parameters (mean values for the beginning and end of  
726 the incubations) of the artificial seawater for each *Emiliana huxleyi* population. pH  
727 and TA samples were collected and measured before and at the end of incubation.  
728 Data are expressed as mean values of six strains in the Azores and Bergen population,  
729 and five strains in the Canary Islands population.

	$p\text{CO}_2$ ( $\mu\text{atm}$ )	pH (total scale)	TA ( $\mu\text{mol}$ $\text{kg}^{-1}$ )	DIC ( $\mu\text{mol}$ $\text{kg}^{-1}$ )	$\text{HCO}_3^-$ ( $\mu\text{mol}$ $\text{kg}^{-1}$ )	$\text{CO}_3^{2-}$ ( $\mu\text{mol}$ $\text{kg}^{-1}$ )	$\text{CO}_2$ ( $\mu\text{mol}$ $\text{kg}^{-1}$ )	$\Omega$
Azores	125 $\pm$ 3	8.46 $\pm$ 0.01	2358 $\pm$ 12	1844 $\pm$ 11	1485 $\pm$ 13	355 $\pm$ 5	5 $\pm$ 0	8.5 $\pm$ 0.1
	300 $\pm$ 20	8.16 $\pm$ 0.03	2339 $\pm$ 27	2031 $\pm$ 17	1803 $\pm$ 18	218 $\pm$ 13	11 $\pm$ 1	5.2 $\pm$ 0.3
	360 $\pm$ 19	8.09 $\pm$ 0.02	2322 $\pm$ 30	2052 $\pm$ 14	1849 $\pm$ 9	190 $\pm$ 10	13 $\pm$ 1	4.5 $\pm$ 0.3
	500 $\pm$ 26	7.97 $\pm$ 0.02	2301 $\pm$ 23	2100 $\pm$ 16	1933 $\pm$ 14	149 $\pm$ 8	18 $\pm$ 1	3.5 $\pm$ 0.2
	695 $\pm$ 20	7.85 $\pm$ 0.01	2317 $\pm$ 11	2167 $\pm$ 13	2023 $\pm$ 14	118 $\pm$ 2	25 $\pm$ 1	2.8 $\pm$ 0.1
	875 $\pm$ 40	7.76 $\pm$ 0.02	2320 $\pm$ 19	2206 $\pm$ 13	2076 $\pm$ 10	99 $\pm$ 5	32 $\pm$ 1	2.4 $\pm$ 0.1
	1110 $\pm$ 119	7.66 $\pm$ 0.05	2303 $\pm$ 19	2222 $\pm$ 23	2101 $\pm$ 25	80 $\pm$ 8	40 $\pm$ 4	1.9 $\pm$ 0.2
	1315 $\pm$ 104	7.59 $\pm$ 0.03	2308 $\pm$ 18	2251 $\pm$ 26	2133 $\pm$ 26	70 $\pm$ 4	48 $\pm$ 4	1.7 $\pm$ 0.1
	1665 $\pm$ 107	7.50 $\pm$ 0.03	2311 $\pm$ 11	2286 $\pm$ 15	2169 $\pm$ 14	57 $\pm$ 3	60 $\pm$ 4	1.4 $\pm$ 0.1
	1935 $\pm$ 175	7.44 $\pm$ 0.04	2308 $\pm$ 15	2302 $\pm$ 24	2183 $\pm$ 21	50 $\pm$ 4	70 $\pm$ 6	1.2 $\pm$ 0.1
2490 $\pm$ 132	7.33 $\pm$ 0.02	2320 $\pm$ 12	2350 $\pm$ 15	2220 $\pm$ 13	40 $\pm$ 2	90 $\pm$ 5	0.9 $\pm$ 0.1	
Bergen	120 $\pm$ 3	8.47 $\pm$ 0.01	2354 $\pm$ 18	1834 $\pm$ 18	1470 $\pm$ 17	359 $\pm$ 2	4 $\pm$ 0	8.6 $\pm$ 0.1
	290 $\pm$ 16	8.17 $\pm$ 0.02	2337 $\pm$ 21	2024 $\pm$ 12	1793 $\pm$ 14	220 $\pm$ 10	11 $\pm$ 1	5.3 $\pm$ 0.2
	355 $\pm$ 18	8.10 $\pm$ 0.02	2315 $\pm$ 23	2045 $\pm$ 11	1840 $\pm$ 7	192 $\pm$ 10	13 $\pm$ 1	4.6 $\pm$ 0.2
	490 $\pm$ 18	7.98 $\pm$ 0.02	2302 $\pm$ 19	2096 $\pm$ 14	1926 $\pm$ 12	152 $\pm$ 6	18 $\pm$ 1	3.6 $\pm$ 0.1
	670 $\pm$ 22	7.86 $\pm$ 0.01	2317 $\pm$ 11	2162 $\pm$ 10	2016 $\pm$ 10	121 $\pm$ 3	24 $\pm$ 1	2.9 $\pm$ 0.1
	855 $\pm$ 52	7.77 $\pm$ 0.03	2326 $\pm$ 19	2206 $\pm$ 15	2074 $\pm$ 14	101 $\pm$ 6	30 $\pm$ 2	2.4 $\pm$ 0.1
	1080 $\pm$ 53	7.67 $\pm$ 0.02	2316 $\pm$ 26	2232 $\pm$ 20	2110 $\pm$ 18	83 $\pm$ 5	39 $\pm$ 2	2.0 $\pm$ 0.1
	1280 $\pm$ 71	7.60 $\pm$ 0.02	2318 $\pm$ 15	2257 $\pm$ 17	2138 $\pm$ 17	72 $\pm$ 4	46 $\pm$ 3	1.7 $\pm$ 0.1
	1550 $\pm$ 122	7.52 $\pm$ 0.03	2300 $\pm$ 19	2266 $\pm$ 28	2150 $\pm$ 27	60 $\pm$ 4	56 $\pm$ 4	1.4 $\pm$ 0.1
	1800 $\pm$ 235	7.47 $\pm$ 0.05	2301 $\pm$ 19	2286 $\pm$ 33	2168 $\pm$ 30	53 $\pm$ 6	65 $\pm$ 9	1.3 $\pm$ 0.1
2280 $\pm$ 147	7.37 $\pm$ 0.02	2309 $\pm$ 20	2326 $\pm$ 27	2201 $\pm$ 24	42 $\pm$ 2	82 $\pm$ 5	1.0 $\pm$ 0.1	
Canary Islands	130 $\pm$ 3	8.45 $\pm$ 0.01	2344 $\pm$ 38	1842 $\pm$ 32	1491 $\pm$ 26	347 $\pm$ 7	5 $\pm$ 0	8.3 $\pm$ 0.2
	310 $\pm$ 11	8.15 $\pm$ 0.01	2317 $\pm$ 24	2020 $\pm$ 25	1798 $\pm$ 25	210 $\pm$ 4	11 $\pm$ 1	5.0 $\pm$ 0.1
	375 $\pm$ 14	8.07 $\pm$ 0.01	2295 $\pm$ 14	2040 $\pm$ 12	1846 $\pm$ 13	182 $\pm$ 5	14 $\pm$ 1	4.3 $\pm$ 0.1
	505 $\pm$ 32	7.96 $\pm$ 0.02	2297 $\pm$ 19	2097 $\pm$ 20	1930 $\pm$ 23	148 $\pm$ 7	18 $\pm$ 1	3.5 $\pm$ 0.2
	695 $\pm$ 18	7.85 $\pm$ 0.01	2312 $\pm$ 20	2163 $\pm$ 17	2020 $\pm$ 15	118 $\pm$ 3	25 $\pm$ 1	2.8 $\pm$ 0.1
	925 $\pm$ 73	7.74 $\pm$ 0.04	2319 $\pm$ 26	2211 $\pm$ 15	2083 $\pm$ 12	95 $\pm$ 8	33 $\pm$ 3	2.3 $\pm$ 0.1
	1180 $\pm$ 53	7.64 $\pm$ 0.02	2310 $\pm$ 25	2239 $\pm$ 20	2120 $\pm$ 19	76 $\pm$ 4	43 $\pm$ 2	1.8 $\pm$ 0.1
	1380 $\pm$ 104	7.58 $\pm$ 0.03	2323 $\pm$ 5	2271 $\pm$ 10	2154 $\pm$ 11	68 $\pm$ 5	50 $\pm$ 4	1.6 $\pm$ 0.1
	1740 $\pm$ 98	7.48 $\pm$ 0.02	2319 $\pm$ 16	2298 $\pm$ 16	2180 $\pm$ 15	55 $\pm$ 3	63 $\pm$ 4	1.3 $\pm$ 0.1
	2140 $\pm$ 258	7.40 $\pm$ 0.05	2312 $\pm$ 9	2320 $\pm$ 16	2197 $\pm$ 13	46 $\pm$ 5	78 $\pm$ 10	1.1 $\pm$ 0.1
2630 $\pm$ 284	7.31 $\pm$ 0.04	2317 $\pm$ 13	2363 $\pm$ 20	2225 $\pm$ 14	37 $\pm$ 3	98 $\pm$ 8	0.8 $\pm$ 0.1	

731 **Table 3.** Calculated optimum  $p\text{CO}_2$ , calculated maximum value ( $V_{\text{max}}$ ) and fitted  
732 relative sensitivity constant ( $rs$ , %) of growth, POC and PIC production rates of each  
733 *E. huxleyi* strain.

strain	Growth rate			POC production rate			PIC production rate		
	optimum $p\text{CO}_2$ ( $\mu\text{atm}$ )	$V_{\text{max}}$ ( $\text{d}^{-1}$ )	$rs$	optimum $p\text{CO}_2$ ( $\mu\text{atm}$ )	$V_{\text{max}}$ ( $\text{pg C cell}^{-1} \text{d}^{-1}$ )	$rs$	optimum $p\text{CO}_2$ ( $\mu\text{atm}$ )	$V_{\text{max}}$ ( $\text{pg C cell}^{-1} \text{d}^{-1}$ )	$rs$
A23	392	1.21	0.22	673	12.47	0.50	323	13.45	0.38
A22	436	1.27	0.16	591	17.33	0.33	635	12.28	0.40
A21	392	1.25	0.22	707	15.45	0.50	396	16.73	1.11
A19	371	1.26	0.24	512	16.17	0.56	480	18.92	0.67
A13	244	1.08	0.13	756	9.84	0.63	471	11.72	0.57
A10	432	1.32	0.20	549	14.42	0.48	385	11.69	0.24
B95	534	1.26	0.10	762	13.46	0.20	562	9.13	0.33
B63	436	1.26	0.11	633	16.66	0.27	615	12.93	0.45
B62	456	1.29	0.11	945	17.27	0.18	488	14.00	0.43
B51	499	1.29	0.11	660	16.77	0.35	492	11.87	0.48
B41	542	1.25	0.09	984	18.34	0.38	553	9.46	0.37
B17	490	1.32	0.14	761	15.19	0.30	625	12.77	0.47
C98	400	1.03	0.16	644	8.44	0.54	440	6.40	0.31
C91	393	0.97	0.21	413	4.83	0.60	195	10.87	0.33
C90	384	0.97	0.12	546	8.28	0.34	284	8.52	0.50
C41	393	1.01	0.14	609	7.64	0.45	545	11.15	0.30
C35	378	1.05	0.17	596	8.87	0.44	464	12.68	0.34

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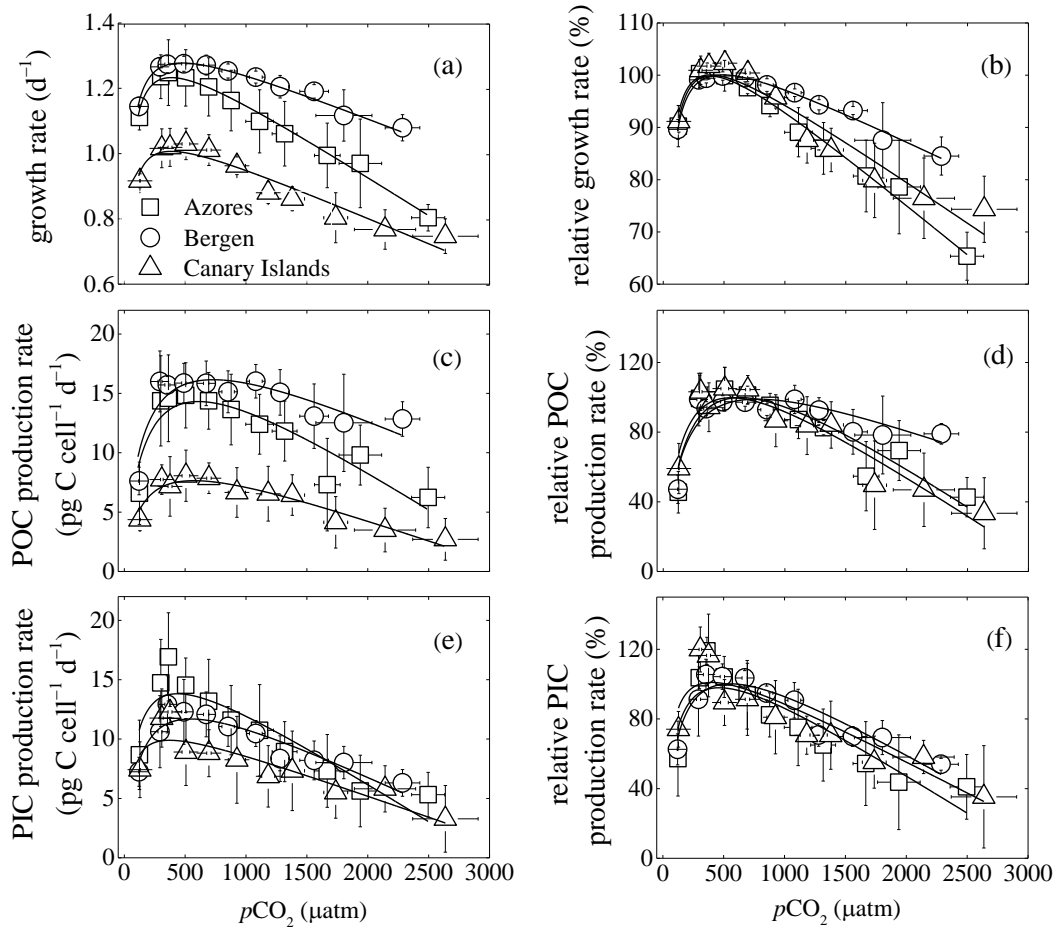
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748 Figure 1

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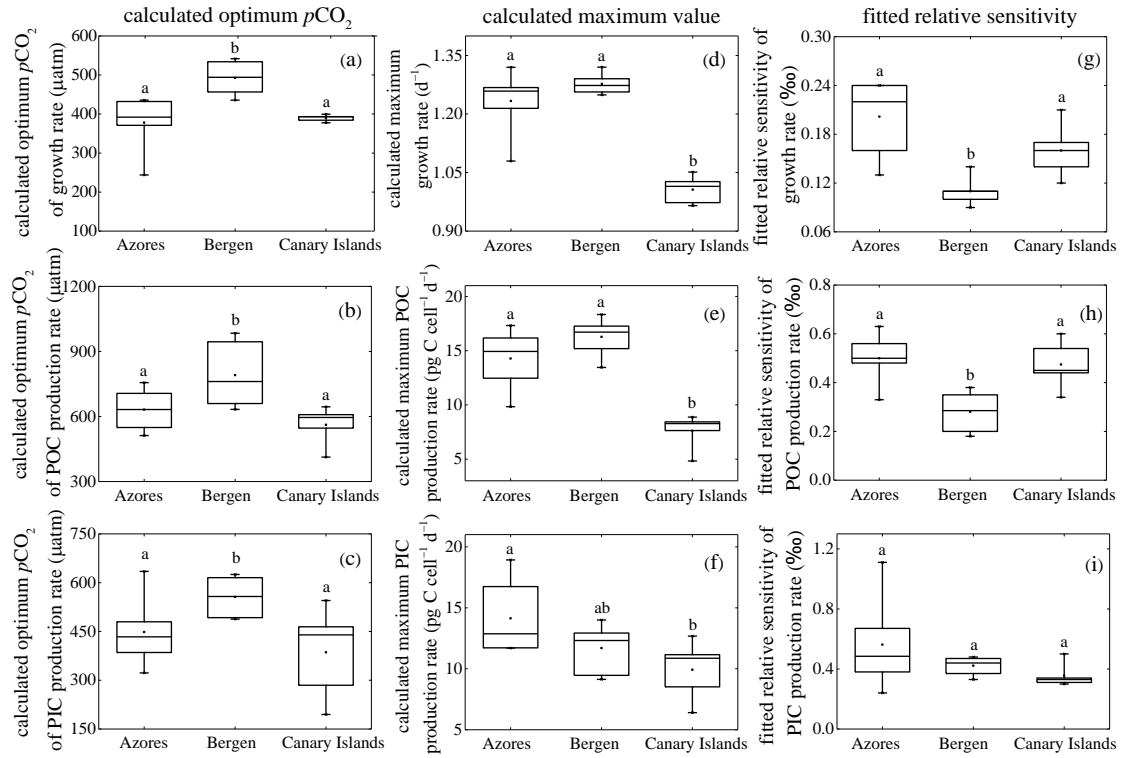
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760 Figure 2

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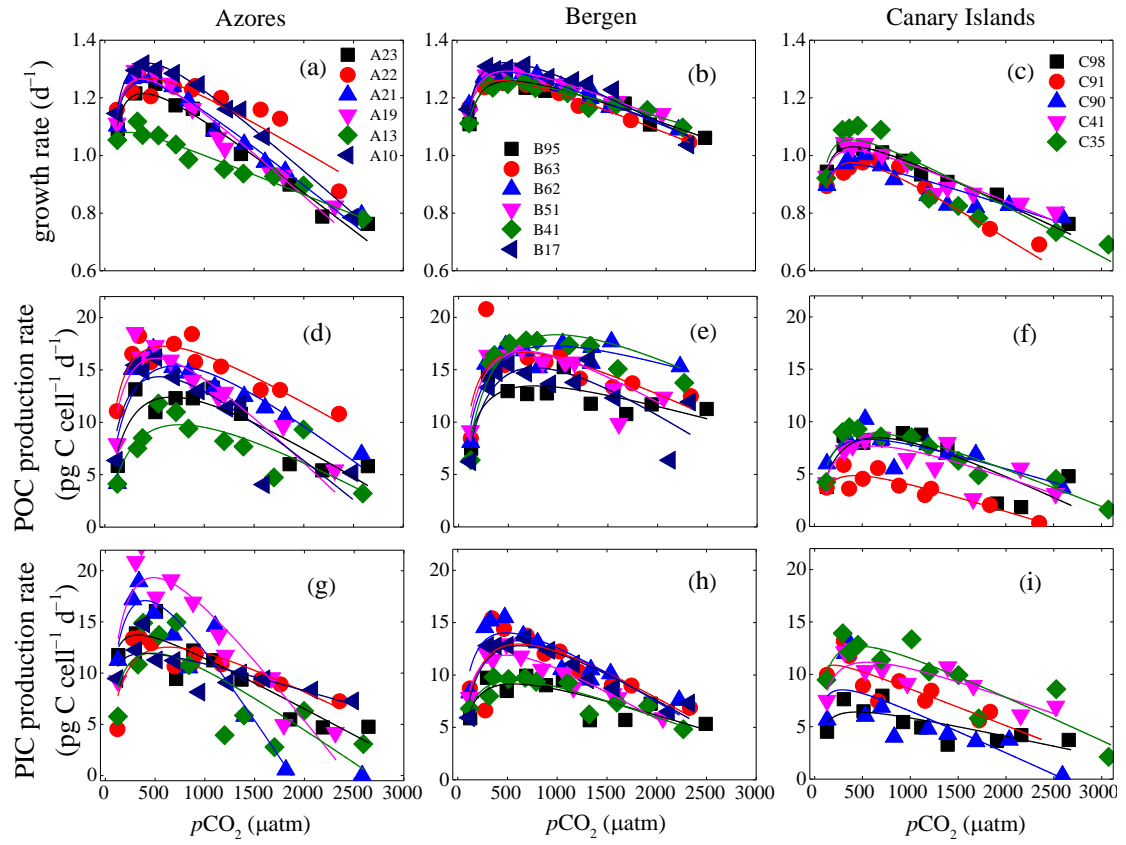
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773 Figure 3

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