

Responses to comments

Dear referee,

We thank you for your supportive comments on our manuscript. Our detailed response in blue text to your comments is attached. Changes to the manuscript text are underlined.

General comments

Overall I think that the manuscript is very close to being ready; I only have a few comments. The new schematic diagram helps but it is not quite clear to me yet, I would add more details. The conclusions and global extrapolations could be worked a little bit more, they basically only cite the Paasche paper. The PIC: POC ratio results were a great addition!

Response: Agreed. We have changed the schematic diagram (Fig. S1), which showed the experimental setup clearly now.

'In addition, these results will improve our understanding on variation in physiological responses of different *E. huxleyi* populations to climate change, and variation in production of different areas in future oceans.' This sentence was added in **lines 359–362**.

PIC : POC ratios of the Azores and Bergen populations declined with rising $p\text{CO}_2$, whereas PIC : POC ratios of the Canary Islands population were rather constant (Figs. S6, S7). This sentence has shown in **lines 366–368**.

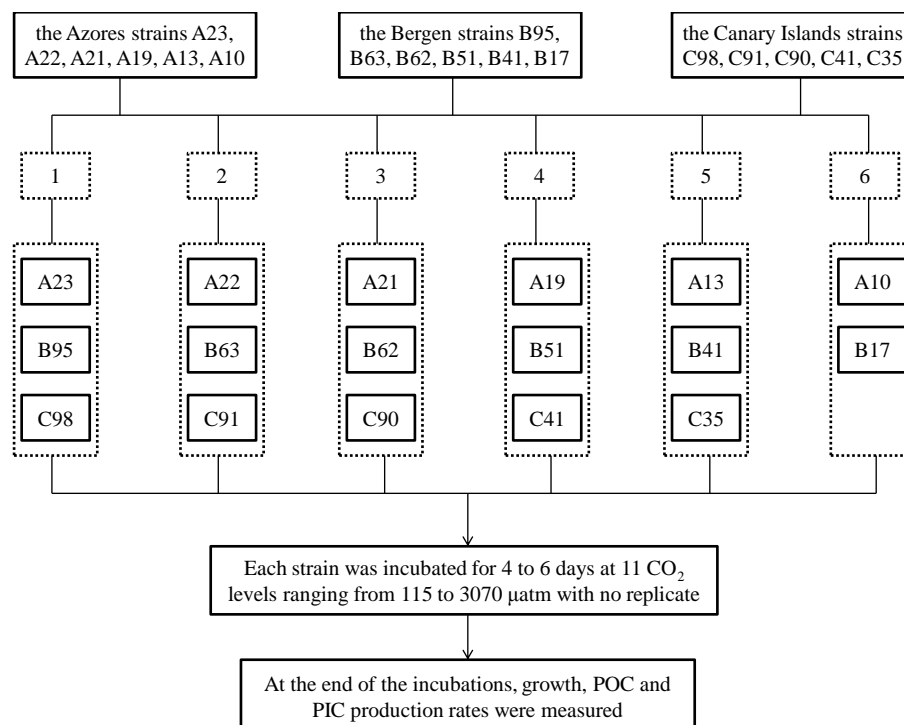


Fig. S1 A flow chat for the experimental processes.

Technical comments

Line 47: suggest deleting or changing this last phrase, not a very good conclusive phrase and this is related to understanding the implications of your findings and global impacts. Perhaps you can say accounting for this variability is important to understand how or whether *Ehux* might adapt to rising CO₂ levels.

Response: Agreed.

‘The existence of distinct responses to changes in carbonate chemistry between and within populations will likely benefit *E. huxleyi* to acclimate and adapt to rising CO₂ levels in the oceans.’ was changed to ‘Accounting for this variability is important to understand how or whether *E. huxleyi* might adapt to rising CO₂ levels.’ **in lines 47–50.**

Line 137 and 147: it does sound like you try to hold your TA constant, I don’t understand why you said you didn’t in your response to my first review....

Response: Yes. In our study, total alkalinity (TA) was constant, which was shown **in line 137**. To response to your first review, we said ‘our CO₂ manipulations are mimicking ongoing ocean acidification where CO₂/pH and DIC changes **at constant TA.**’

Line 237: perhaps I would add “as expected” and add citations

Response: Agreed. We added ‘As expected’ **in line 236.**

Line 258-265: why do you think that there are differences in some sensitivity constants but not in rates?

Response: As shown **in lines 209–210**, ‘*s*, the sensitivity constant, depicts the slope of the decline after optimum CO₂ levels in response to rising H⁺’, which means that sensitivity constant is relevant to rising H⁺. However, growth, POC and PIC production rates are relevant to CO₂ and H⁺ concentrations, and other environmental factors such as temperature and light intensity. So sensitivity constants and rates could show different results.

Line 323: suggest you rephrase to “this illustrates how adaptation to local temperature can ...” I would delete “nicely”, unnecessary.

Response: Agreed. ‘nicely that local temperature adaptation’ was replaced by ‘how adaptation to local temperature’ **in lines 323–324.**

Line 338: add comma after “fashion,”

Response: Agreed. Comma was added after ‘fashion’ **in line 338.**

Line 340-343: your conclusion is based on the assumption that temperature doesn’t change but temperature will increase perhaps even faster than CO₂ and this might have a greater impact. Plus, there could be interactions between temperature and CO₂...

Response: Agreed. Compared to CO₂ concentration, temperature might have a greater impact on growth, POC and PIC production rates of *E. huxleyi*. However, our results cannot show this idea.

Temperature and CO₂ may interactively affect growth, POC and PIC production rates. In this study, incubation temperature (16 °C) may predominantly limit physiological rates of Canary Islands populations. So we said ‘Thus, with rising CO₂, growth, photosynthetic carbon fixation

and calcification rates of the Canary Islands population cannot increase as much as in the Azores and Bergen populations.’ **in lines 340–343.**

Line 363: delete “at increasing temperatures” or change to “and increasing temperatures”

Response: Agreed. We deleted ‘at increasing temperatures’ **in line 366.**

Line 367: please elaborate

Response: ‘As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean.’ was replaced by ‘As changes in PIC : POC ratios of coccolithophore blooms were suggested to impact on biological carbon pump (Rost and Riebesell, 2004), variation in PIC : POC ratios of different populations indicates that different regions might have different changes in marine carbon cycle in the future ocean’. These changes are in **lines 368–374.**

Line 370-72: consider more implications

Response: Agreed. We added ‘and then give a positive feedback to rising atmosphere CO₂ levels’ **in line 379.**

Line 392-95: I disagree, this is just a simple correlation, it doesn’t say anything about dominance of strains. My take is that higher growth rate, means larger population and so, greater production.

Response: Agreed. We changed ‘the dominating strains will also take up or fix dissolved inorganic carbon faster’ to ‘higher growth rate means larger populations and then greater production’. These changes are **in lines 402–404.**

Line 395-396: again, the conclusions need more work

Response: We deleted this sentence ‘When extrapolated to the ocean, *E. huxleyi* blooms may increase the potential of the oceans to absorb CO₂ from the atmosphere and its carbon storage capacity (Blanco-Ameijeiras et al., 2016), which has the potential to mitigate rising CO₂ levels in the atmosphere.’ **in lines 404–407.**

Line 408: I would say something like: in this case, we only studied the effects of rising CO₂ but future studies should take into account simultaneous effects from other interacting factors such as light and temperature variability.

Response: Agreed. We changed these contents ‘, and CO₂ response was modulated by other environmental factors such as temperature and light intensity’ to ‘In this study, we only studied the effects of rising CO₂ but future studies should take into account simultaneous effects from other interacting factors such as light and temperature variability.’. These changes are **in lines 417–421.**

List of changes

1 Lines 47–50: changed ‘The existence of distinct responses to changes in carbonate chemistry between and within populations will likely benefit *E. huxleyi* to acclimate and adapt to rising CO₂ levels in the oceans.’ to ‘Accounting for this variability is important to understand how or whether *E. huxleyi* might adapt to rising CO₂ levels.’

2 Line 236: added ‘As expected’ and changed ‘G’ to ‘g’.

3 Lines 323–324: changed ‘nicely that local temperature adaptation’ to ‘how adaptation to local temperature’.

4 Line 337: deleted ‘De Bodt et al., 2010;’

5 Line 338: added ‘,’.

6 Lines 359–362: added ‘In addition, these results will improve our understanding on variation in physiological responses of different *E. huxleyi* populations to climate change, and variation in production of different areas in future oceans.’.

7 Lines 366: deleted ‘at increasing temperatures’.

8 Line 368: added ‘s’ and ‘, S7’.

9 Lines 368–374: changed ‘As changes in PIC : POC ratios of coccolithophore blooms may impact on the biological carbon pump, different regions might see different changes in the future ocean.’ to ‘As changes in PIC : POC ratios of coccolithophore blooms were suggested to impact on biological carbon pump (Rost and Riebesell, 2004), variation in PIC : POC ratios of different populations indicates that different regions might have different changes in marine carbon cycle in the future ocean.’.

10 Line 379: added ‘and then give a positive feedback to rising atmosphere CO₂ levels’.

11 Lines 402–404: changed ‘the dominating strains will also take up or fix dissolved inorganic carbon faster’ to ‘higher growth rate means larger populations and then greater production’.

12 Lines 404–407: deleted ‘When extrapolated to the ocean, *E. huxleyi* blooms may increase the potential of the oceans to absorb CO₂ from the atmosphere and its carbon storage capacity (Blanco-Ameijeiras et al., 2016), which has the potential to mitigate rising CO₂ levels in the atmosphere.’.

13 Lines 417–421: changed ‘, and CO₂ response was modulated by other environmental factors such as temperature and light intensity.’ to ‘In this study, we only studied the effects of rising CO₂

but future studies should take into account simultaneous effects from other interacting factors such as light and temperature variability.’

14 Line 564: added ‘, doi: 10.1016/j.pocean.2017.10.007, 2017’.

1 **Population-specific responses in physiological rates of *Emiliana huxleyi* to a**
2 **broad CO₂ range**

3

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19 Running head: *population response of Emiliana huxleyi to CO₂*

20

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22 Keywords: CO₂; coccolithophore; physiological rate; population; strain

23 **Abstract**

24 Although coccolithophore physiological responses to CO₂-induced changes in
25 seawater carbonate chemistry have been widely studied in the past, there is limited
26 knowledge on the variability of physiological responses between populations from
27 different areas. In the present study, we investigated the specific responses of growth,
28 particulate organic (POC) and inorganic carbon (PIC) production rates of 3
29 populations of the coccolithophore *Emiliana huxleyi* from three regions in the North
30 Atlantic Ocean (Azores: 6 strains, Canary Islands: 5 strains, and Norwegian coast near
31 Bergen: 6 strains) to a CO₂ partial pressure ($p\text{CO}_2$) range from 120 μatm to 2630
32 μatm . Physiological rates of each population and individual strain increased with
33 rising $p\text{CO}_2$ levels, reached maximum and declined thereafter. Optimal $p\text{CO}_2$ for
34 growth and POC production rates and tolerance to low pH (i.e. high proton
35 concentration) was significantly higher in an *E. huxleyi* population isolated from the
36 Norwegian coast than in those isolated near the Azores and Canary Islands. This may
37 be due to the large environmental variability including large $p\text{CO}_2$ and pH
38 fluctuations in coastal waters off Bergen compared to the rather stable oceanic
39 conditions at the other two sites. Maximum growth and POC production rates of the
40 Azores and Bergen populations were similar and significantly higher than that of the
41 Canary Islands population. This pattern could be driven by temperature-CO₂-
42 interactions where the chosen incubation temperature (16 °C) was slightly below what
43 strains isolated near the Canary Islands normally experience. Our results indicate
44 adaptation of *E. huxleyi* to their local environmental conditions and the existence of

45 distinct *E. huxleyi* populations. Within each population, different growth, POC and
46 PIC production rates at different $p\text{CO}_2$ levels indicated strain-specific phenotypic
47 plasticity. ~~The existence of distinct responses to changes in carbonate chemistry~~
48 ~~between and within populations will likely benefit *E. huxleyi* to acclimate and adapt to~~
49 ~~rising CO_2 levels in the oceans.~~Accounting for this variability is important to
50 understand how or whether *E. huxleyi* might adapt to rising CO_2 levels.

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67 **1 Introduction**

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69 Coccolithophores form a layer of calcium carbonate (CaCO₃) platelets (coccoliths)
70 around their cells. Coccoliths are of biogeochemical importance due to ballasting of
71 organic matter with CaCO₃, a phenomenon which is thought to promote the transport
72 of organic carbon to the deep ocean (Klaas and Archer, 2002; Rost and Riebesell,
73 2004). The coccolithophore *Emiliana huxleyi* forms extensive blooms under
74 favourable light intensity, temperature and nutrient conditions, with different
75 morphotypes in certain regions (Cook et al., 2011; Henderiks et al., 2012; Smith et al.,
76 2012; Balch et al., 2014; Krumhardt et al., 2017).

77 Variable responses of growth, photosynthetic carbon fixation and calcification rates
78 of different *E. huxleyi* strains to rising CO₂ levels have been reported (Langer et al.,
79 2009; Hoppe et al., 2011; Müller et al., 2015; Hattich et al, 2017) and are likely a
80 result of intra-specific variability of genotypes (Langer et al., 2009). Several recent
81 studies observed optimum curve responses in physiological rates of a single *E. huxleyi*
82 strain to a broad *p*CO₂ range from about 20 µatm to 5000 µatm, and linked them to
83 inorganic carbon substrate limitation at low *p*CO₂ and inhibiting H⁺ concentrations at
84 high *p*CO₂ (Bach et al., 2011; 2015; Kottmeier et al., 2016). Until now, studies on the
85 physiological responses of *E. huxleyi* to rising CO₂ are mostly based on a few
86 genotypes and little is known about the potential variability in CO₂ and H⁺ sensitivity
87 between and within populations. Recently, several studies found substantial variations
88 in CO₂ responses for N₂ fixation rates between *Trichodesmium* strains, as well as for

89 growth rates between strains of *Gephyrocapsa oceanica*, *Ostreococcus tauri* and
90 *Fragilariopsis cylindrus* (Hutchins et al., 2013; Schaum et al., 2013; Pancic et al.,
91 2015; Hattich et al., 2017). Hence, multiple strains, ideally from geographically
92 distinct regions should be considered for investigating phytoplankton responses to
93 climate change (Zhang et al., 2014; Blanco-Ameijeiras et al., 2016; Krumhardt et al.,
94 2017).

95 Oceanographic boundaries formed by both ocean currents and environmental
96 factors such as temperature, can limit dispersal of marine phytoplankton, reduce gene
97 flow between geographic populations, and give rise to differentiated populations
98 (Palumbi, 1994). Different populations were found to show different growth rates for
99 *E. huxleyi*, *G. oceanica*, and *Skeletonema marinoi* at the same temperatures, and for
100 *Ditylum brightwellii* at the same light intensities (Brand, 1982; Rynearson and
101 Armbrust, 2004; Kremp et al., 2012; Zhang et al., 2014). Phenotypic plasticity
102 describes the ability of a strain to change its morphology or physiology in response to
103 changing environmental conditions (Bradshaw, 1965). Plasticity can be assessed by
104 analyzing the reaction norm of one trait and a plastic response may allow a strain to
105 acclimate across an environmental gradient and widen its bio-geographical
106 distribution (Reusch, 2014; Levis and Pfennig, 2016).

107 In order to better understand how local adaptation affects the physiological
108 response of *E. huxleyi* to rising CO₂ conditions, we isolated 17 strains from three
109 regions in the Atlantic Ocean, and assessed growth, carbon fixation and calcification
110 responses of the population over a *p*CO₂ range from 120 μatm to 2630 μatm.

111

112 **2 Materials and methods**

113

114 **2.1 Cell isolation sites and experimental setup**

115 *Emiliana huxleyi* strains EHGKL B95, B63, B62, B51, B41 and B17 originated from
116 Raunefjord (Norway 60°18'N, 05°15'E) and were isolated by K. T. Lohbeck in May,
117 2009 (Lohbeck et al., 2012) at ~ 10 °C in-situ water temperature. *E. huxleyi* strains
118 EHGLE A23, A22, A21, A19, A13 and A10 originated from coastal waters near the
119 Azores (38°34'N, 28°42'W) and were isolated by S. L. Eggers in May or June, 2010
120 at ~ 17 °C in-situ water temperature. *E. huxleyi* strains EHGKL C98, C91, C90, C41
121 and C35 originated from coastal waters near Gran Canaria (27°58'N, 15°36'W) and
122 were isolated by K. T. Lohbeck in February, 2014 at ~ 18 °C in-situ water temperature.
123 Seasonal CO₂ concentration in the surface seawater ranges from 240 µatm to 400
124 µatm near Bergen, from 320 µatm to 400 µatm around the Azores and from 320 µatm
125 to 400 µatm around the Canary Islands (Table 1). Monthly surface seawater
126 temperature ranges from 6.0 to 16.0 °C near Bergen, 15.6 to 22.3 °C around the
127 Azores and from 18.0 to 23.5 °C around the Canary Islands (Table S1).

128 All 17 strains belong to morphotype A (determined by scanning electron
129 microscopy) and have been deposited in the Roscoff culture collection (RCC) under
130 the official names as shown above. Genetically different isolates, here called strains,
131 were identified by 5 microsatellite markers (P02E09, P02B12, P02F11, EHMS37,
132 EHMS15) (Table S2). For a description of primer testing, deoxyribonucleic acid

133 (DNA) extraction, DNA concentration measurements, and polymerase chain reaction
134 (PCR) protocols see Zhang et al. (2014). The Azores and Bergen strains had been
135 used earlier by Zhang et al. (2014).

136 The six or five (in case of Canary Islands) strains of each region were used to test
137 the physiological response to varying CO₂ concentrations at constant total alkalinity
138 (TA). The experiment was performed in six consecutive incubations, with one strain
139 from each population (Azores, Bergen, Canary Islands) being cultured at a time (Fig.
140 S1). Monoclonal populations were always grown in sterile-filtered (0.2 µm diameter,
141 Sartobran[®] P 300, Sartorius) artificial seawater medium (ASW) as dilute batch
142 cultures at 200 µmol photons m⁻² s⁻¹ light intensity under a 16/8 h light/dark cycle
143 (light period: 5:00 a.m to 9:00 p.m.) at 16 °C which we consider to be a compromise
144 for the three different origins of the strains. Nutrients were added in excess (with
145 nitrate and phosphate concentrations of 64 µmol kg⁻¹ and 4 µmol kg⁻¹, respectively).
146 For the preparation of ASW and nutrient additions see Zhang et al. (2014). Calculated
147 volumes of Na₂CO₃ and hydrochloric acid were added to the ASW to achieve target
148 CO₂ levels at an average total alkalinity (TA) of 2319 ± 23 µmol kg⁻¹ (Pierrot et al.,
149 2006; Bach et al., 2011). Each strain was grown under 11 CO₂ levels ranging from
150 115 µatm to 3070 µatm without replicate. Mean response variables of all strains with
151 a population were calculated and mean CO₂ levels of all strains within a population
152 ranged from 120 µatm to 2630 µatm. Cells grew in the experimental conditions for at
153 least 7 generations, which corresponded to 4–7 days depending on cell division rates.
154 Cells were cultured for 4 days in 120–925 µatm CO₂, for 5 days in 1080–1380 µatm

155 CO₂, and for 6 or 7 days in 1550–2630 μatm CO₂. Initial cell concentration was 200
156 cells ml⁻¹ (estimated from measured pre-culture concentrations and known dilution)
157 and final cell concentration was lower than 100,000 cells ml⁻¹. Dissolved inorganic
158 carbon (DIC) concentrations and $p\text{CO}_2$ levels changed less than 7% and 11%,
159 respectively, during the experimental growth phase.

160

161 **2.2 pH_T and total alkalinity measurements**

162 At 10:00 a.m. on the last day of incubations (at day 4–7 depending on CO₂
163 concentration), pH_T and TA samples were filtered (0.2 μm diameter, Filtropur S 0.2,
164 Sarstedt) by gentle pressure and stored at 4°C for a maximum of 14 days. The entire
165 sampling lasted less than 2 h. The pH_T sample bottles were filled with considerable
166 overflow and closed tightly with no space. pH_T was measured spectrophotometrically
167 (Cary 100, Agilent) using the indicator dye *m*-cresol purple (Sigma-Aldrich) similar
168 to Carter et al. (2013) with constants of acid dissociation for the protonated and un-
169 protonated forms reported in Clayton and Byrne (1993). TA was measured by open-
170 cell potentiometric titration (862 Compact Titrosampler, Metrohm) according to
171 Dickson et al. (2003). The carbonate system was calculated from measured TA, pH_T,
172 (assuming 4 $\mu\text{mol kg}^{-1}$ of phosphate and 0 $\mu\text{mol kg}^{-1}$ of silicate) using the CO₂
173 System Calculations in MS Excel software (Pierrot et al., 2006) with carbonic acid
174 constants K_1 and K_2 as determined by Roy et al. (1993).

175

176 **2.3 Growth rate measurements**

177 At 1:00 p.m. on the last day of incubation, 25 ml samples were used to measure cell
178 concentration. Cell concentration was determined within two hours using a Z2 Coulter
179 Particle Counter (Beckman). Growth rate (μ) was calculated according to:

$$180 \quad \mu = (\ln N_1 - \ln N_0) / d \quad (1)$$

181 where N_1 is cell concentration on the last day of incubation, N_0 is 200 cells mL⁻¹, and
182 d is the time period for growth of algae in days.

183

184 **2.4 Particulate organic (POC) and inorganic (PIC) carbon measurements**

185 At 3:00 p.m. on the last day of incubation, cells for total particulate (TPC) and total
186 organic (TOC) carbon were filtered onto GF/F filters which were pre-combusted at
187 500 °C for 8 h. Samples of background particulate carbon (BPC) were determined in a
188 similar way but using filtered ASW without algae, which was previously adjusted to
189 target $p\text{CO}_2$ levels, and allowed to age for about 7 days under incubation conditions
190 (*see above*). All samples were placed at -20°C. BPC filters were used as blanks to
191 correct for organic carbon in the medium. TOC and BPC filters were acid fumed.
192 Afterwards, all filters were dried for 8 h at 60 °C. TPC, TOC and BPC were measured
193 using an Elemental Analyzer (EuroEA, Hekatech GmbH). The percentages of BPC in
194 TPC were about 20% at cell densities < 10,000 cells ml⁻¹ and about 10% at cell
195 densities > 40,000 cells ml⁻¹. POC was calculated as the difference between TOC and
196 BPC. PIC was calculated as the difference between TPC and TOC. POC and PIC
197 production rates were calculated as:

$$198 \quad \text{POC production rate} = \mu \text{ (d}^{-1}\text{)} \times (\text{TOC} - \text{BPC}) \text{ (pg C cell}^{-1}\text{)} \quad (2)$$

199 PIC production rate = μ (d^{-1}) \times (TPC – TOC) ($pg\ C\ cell^{-1}$) (3)

200

201 **2.5 Data analysis**

202 In a broad pCO_2 range, physiological rates are expected to initially increase quickly
 203 until reaching an optimum and then decline towards further increasing CO_2 levels (e.g.
 204 Krug et al. 2011). Hence we used the following modified Michaelis-Menten equation
 205 (Bach et al. 2011) which was fitted to measured cellular growth, POC and PIC
 206 production rates and yield theoretical optimum pCO_2 and maximum values for each of
 207 the three populations (combining the data of five or six strains) (Bach et al., 2011).

208
$$y = \frac{X \times pCO_2}{Y + pCO_2} - s \times pCO_2 \quad (4)$$

209 where X and Y are fitted parameters, and s , the sensitivity constant, depicts the slope
 210 of the decline after optimum CO_2 levels in response to rising H^+ . Based on the fitted X ,
 211 Y and s , we calculated pCO_2 optima (K_m) (equation 5) and maximum growth, POC
 212 and PIC production rates following Bach et al., (2011).

213
$$K_m = \sqrt{\frac{X \times Y}{s}} - Y \quad (5)$$

214 The relative values for growth, POC and PIC production rates were calculated as
 215 ratios of growth, POC and PIC production rates at each pCO_2 level to the maximum
 216 (highest) rates. We obtained the relative sensitivity constant by fitting function (4)
 217 based on relative growth, POC and PIC production rates.

218 A one-way ANOVA was then used to test for statistically significant differences in
 219 theoretical optimum pCO_2 , maximum value and relative sensitivity constant between
 220 populations. A Tukey HSD test was conducted to determine the differences between

221 strains from different populations. A Shapiro–Wilk’s analysis was tested to analyze
222 residual normality. Statistical calculations were carried out using *R* and significance
223 was shown by $p < 0.05$.

224

225 **3 Results**

226

227 **3.1 Carbonate chemistry parameters**

228 Carbonate system parameters are shown in Table 2. Average $p\text{CO}_2$ levels of the ASW
229 ranged from 125 μatm to 2490 μatm for the Azores population, from 120 μatm to
230 2280 μatm for the Bergen population, and from 130 μatm to 2630 μatm for the
231 Canary Islands population. Corresponding pH_T values of the ASW ranged from 8.46
232 to 7.33 for the Azores population, from 8.47 to 7.37 for the Bergen population, and
233 from 8.45 to 7.31 for the Canary Islands population.

234

235 **3.2 Measured growth, POC and PIC production rates of each population**

236 As expected, G growth rates, POC and PIC production rates of the three *E. huxleyi*
237 populations increased with rising $p\text{CO}_2$, reached a maximum, and then declined with
238 further $p\text{CO}_2$ increase (Fig. 1). Growth rates of the Azores and Bergen populations
239 were larger than those of the Canary Islands population at all investigated $p\text{CO}_2$ levels
240 (Fig. 1a). With rising $p\text{CO}_2$ levels beyond the $p\text{CO}_2$ optimum, decline in growth rates
241 was more pronounced in the Azores and Canary Islands populations than in the
242 Bergen population (Fig. 1b).

243 Measured POC production rates of the Azores and Bergen populations were larger
244 than those of the Canary Islands population at all $p\text{CO}_2$ levels (Fig. 1c) and decline in
245 POC production rates with increasing $p\text{CO}_2$ levels beyond the $p\text{CO}_2$ optimum was
246 larger in the Azores and Canary Islands populations than in the Bergen population
247 (Fig. 1d).

248 Measured PIC production rates at investigated $p\text{CO}_2$ levels did not show significant
249 differences among the Azores, Bergen and Canary Islands populations (Fig. 1e).
250 Exceptions were that at 365–695 μatm , PIC production rates of the Azores population
251 were larger than those of the Canary Islands population (all $p < 0.05$).

252

253 **3.3 Physiological responses of populations to $p\text{CO}_2$**

254 Calculated optimum $p\text{CO}_2$ for growth, POC and PIC production rates of the Bergen
255 population were significantly larger than those of the Azores and Canary Islands
256 populations (all $p < 0.05$) (Fig. 2a–c). Optimum $p\text{CO}_2$ for these physiological rates
257 between the Azores and Canary Islands population were not different (all $p > 0.1$).

258 Calculated maximum growth rates, POC and PIC production rates were not
259 significantly different between the Azores and the Bergen populations (all $p > 0.1$)
260 (Fig. 2d–f). Maximum growth rate and POC production rate of the Canary Islands
261 population were significantly lower than those of the Azores and Bergen populations
262 (both $p < 0.01$) (Fig. 2d,e). Maximum PIC production rates of the Canary Islands
263 population were significantly lower than that of the Azores population ($p < 0.05$),
264 while there was no difference to the Bergen population ($p > 0.1$) (Fig. 2f).

265 Fitted relative sensitivity constants for growth and POC production rates of the
266 Bergen population were significantly lower than those of the Azores and Canary
267 Islands populations ($p < 0.01$) (Fig. 2g, h). Fitted relative sensitivity constants for
268 growth and POC production rates between the Azores and Canary Islands populations
269 were not significantly different ($p > 0.1$). Fitted relative sensitivity constants for PIC
270 production rates did not show difference among three populations ($p = 0.13$) (Fig. 2i).

271

272 **3.4 Physiological responses of individual strains to $p\text{CO}_2$**

273 Measured growth rates, POC and PIC production rates of 17 *E. huxleyi* strains showed
274 optimum curve response patterns to the broad $p\text{CO}_2$ gradient (Fig. 3). Variations in
275 calculated $p\text{CO}_2$ optima, maximum values and relative sensitivity constants of
276 physiological rates were found between the strains (Table 3).

277 For all strains within each population, optimum $p\text{CO}_2$ of POC production rates
278 were larger than optimum $p\text{CO}_2$ of growth rates or PIC production rates with the
279 exception of optimum $p\text{CO}_2$ of POC and PIC production rates of *E. huxleyi* strain
280 EHGLE A22 (Table 3). Compared to the Azores and Bergen populations, strains
281 isolated near the Canary Islands showed larger variation in optimum $p\text{CO}_2$ of PIC
282 production rates. Within the Azores population, variations in maximum values (V_{max})
283 and relative sensitivity constants (rs) of growth, POC and PIC production rates of all
284 strains were larger than those within the Bergen and Canary Islands populations (Fig.
285 3).

286

287 4 Discussion

288

289 We investigated growth, POC and PIC production rates of 17 *E. huxleyi* strains from
290 three populations to a broad $p\text{CO}_2$ range (120–2630 μatm). The three populations
291 differed significantly in growth and POC production rates at the investigated $p\text{CO}_2$
292 levels. The reaction norms of the individual strains and populations equaled an
293 optimum curve for all physiological rates (Figs. 1 and 3). However, we detected
294 distinct $p\text{CO}_2$ optima for growth, POC and PIC production rates, and different H^+
295 sensitivities for growth and POC production rates among them (Fig. 2). These results
296 indicate the existence of distinct populations in the cosmopolitan coccolithophore *E.*
297 *huxleyi*.

298 In comparison to the Azores and Canary Islands populations, variability in growth
299 rates between strains of the Bergen population was smaller even though they had
300 higher growth rates at all $p\text{CO}_2$ levels (Fig. 3). Furthermore, the Bergen population
301 showed significantly higher $p\text{CO}_2$ optima and lower H^+ sensitivity for growth and
302 POC production rates (Fig. 2). These findings indicate that the Bergen population may
303 be more tolerant to changing carbonate chemistry in terms of its growth and
304 photosynthetic carbon fixation rates. The Bergen strains were isolated from coastal
305 waters, while the Azores and Canary Islands strains were isolated from a more
306 oceanic environment. Seawater carbonate chemistry of coastal waters is usually more
307 dynamic than in the open ocean (Cai, 2011). In fact, previous studies have reported
308 that CO_2 and pH variability of the seawater off Bergen was larger than off the Azores

309 and Canary Islands (Table 1). In addition, due to riverine input, seawater upwelling
310 and metabolic activity of plankton communities, environmental variability in coastal
311 waters are larger than in open-ocean ecosystems (Duarte and Cerbrian, 1996). Doblin
312 and van Sebille (2016) suggested that phytoplankton populations should be constantly
313 under selection when experiencing changing environmental conditions. In this case,
314 the Bergen population, exposed to larger CO₂ or pH fluctuations, may have acquired a
315 higher capacity to acclimate to changing carbonate chemistry resulting in a higher
316 tolerance (or lower sensitivity) to rising CO₂ levels. In contrast, the Azores and
317 Canary Islands populations experience similar, less variable seawater carbonate
318 chemistry conditions in their natural environment, which could explain why they also
319 show similar *p*CO₂ optima and H⁺ sensitivity for physiological rates (Fig. 2).

320 In an earlier study (Zhang et al., 2014), growth rates of the same Azores and
321 Bergen strains as used here were measured at 8–28 °C. While at 26–28 °C the Bergen
322 strains grew slower than the Azores strains, at 8 °C the Azores strains grew slower
323 than the Bergen strains. This illustrates ~~nicely that local temperature adaptation~~ how
324 adaptation to local temperature can significantly affect growth of *E. huxleyi* strains in
325 laboratory experiments. Considering these findings and the temperature ranges of the
326 three isolation locations (Table S1), the incubation temperature of 16 °C used in the
327 present study was lower than the minimum sea surface temperature (SST) commonly
328 recorded at the Canary Islands. In contrast, SSTs of 16 °C and lower have been
329 reported for Azores and Bergen waters (Table S1). When exposed to 16 °C, growth
330 rate of the Canary Islands population might have been already below their optimum

331 and hence significantly reduced in comparison to the other populations (Fig. 2d).

332 Furthermore, compared to the Canary Islands population, the Azores population
333 had higher maximum growth and POC production rates, and similar optimum CO₂ for
334 these physiological rates. Again, this might be related to sub-optimal incubation
335 conditions as temperature has been found to significantly modulate CO₂ responses in
336 coccolithophores in terms of maximum rates, CO₂ optima and half-saturation, and H⁺
337 sensitivity (De Bodt et al., 2010; Sett et al., 2014; Gafar et al., 2018; Gafar and Schulz,
338 2018). In a similar fashion, light can also modulate CO₂ responses, hence different
339 requirements by strains adapted to different light availabilities could also explain our
340 observations (Zhang et al., 2015; Gafar et al., 2018; Gafar and Schulz, 2018). Thus,
341 with rising CO₂, growth, photosynthetic carbon fixation and calcification rates of the
342 Canary Islands population cannot increase as much as in the Azores and Bergen
343 populations. In addition, the Canary Islands population showed smallest variability in
344 optimum pCO₂ and maximum values for growth and POC production rates (Fig. 2).
345 The reason may be that low incubation temperature predominantly limited growth and
346 POC production rates of the Canary Islands population, and decreased the sensitivities
347 of these physiological rates to rising pCO₂.

348 Before we started this experiment, strains isolated from the Azores, Bergen and
349 Canary Islands grew as stock cultures at 15 °C and 400 µatm for 4 years, 5 years and
350 3 months, respectively. Schaum et al. (2015) provide evidence that long-term
351 laboratory incubation affects responses of phytoplankton to different pCO₂ levels.
352 Thus, it is conceivable that the same selection history in the laboratory incubation

353 may contribute to a more similar response of growth, POC and PIC production rates
354 between the Azores and Bergen populations at low $p\text{CO}_2$ levels (Fig. 1).

355 Our results indicate that *E. huxleyi* populations are adapted to the specific
356 environmental conditions of their origin, resulting in different responses to increasing
357 $p\text{CO}_2$ levels. The ability to adapt to diverse environmental conditions is supposed to
358 be one reason for the global distribution of *E. huxleyi* (Paasche, 2002), spanning a
359 temperature range of about 30 °C. In addition, these results will improve our
360 understanding on variation in physiological responses of different *E. huxleyi*
361 populations to climate change, and variation in production of different areas in future
362 oceans. The optimum temperature for growth of the Bergen population was about 22
363 °C and was 5 °C higher than the maximum SST in Bergen waters (Zhang et al. 2014).

364 Furthermore, in comparison to the Azores and Canary Islands populations, larger
365 optimum $p\text{CO}_2$ of growth rate indicates that the Bergen population may benefit more
366 from the rising CO_2 levels ~~at increasing temperatures~~. PIC : POC ratios of the Azores
367 and Bergen populations declined with rising $p\text{CO}_2$, whereas PIC : POC ratios of the

368 Canary Islands population were rather constant (Figs. S6, S7). ~~As changes in PIC :
369 POC ratios of coccolithophore blooms may impact on the biological carbon pump,
370 different regions might see different changes in the future ocean. As changes in PIC :
371 POC ratios of coccolithophore blooms were suggested to impact on biological carbon
372 pump (Rost and Riebesell, 2004), variation in PIC : POC ratios of different
373 populations indicates that different regions might have different changes in marine
374 carbon cycle in the future ocean.~~ In natural seawater, due to ocean currents and gene

375 flow, populations at any given location may get replaced by immigrant genotypes
376 transported there from other locations (Doblin and van Sebille, 2016). In addition, *E.*
377 *huxleyi* is thought to utilize HCO_3^- for calcification which generates protons, and
378 increase in proton concentration may mitigate the potential of the ocean to absorb
379 atmospheric CO_2 and then give a positive feedback to rising atmosphere CO_2 levels
380 (Paasche, 2002).

381 Within a population, individual strains showed different growth, POC and PIC
382 production rates at different $p\text{CO}_2$ levels, indicating phenotypic plasticity of
383 individual strains (Reusch, 2014). Phenotypic plasticity constitutes an advantage for
384 individual strains to acclimate and adapt to elevated $p\text{CO}_2$ by changing fitness-
385 relevant traits and potentially to attenuate the short-term effects of changing
386 environments on fitness-relevant traits (Schaum et al., 2013).

387 The strain-specific CO_2 -response curves revealed considerable physiological
388 diversity in co-occurring strains (Fig. 3). Physiological variability makes a population
389 more resilient, increases its ability to persist in variable environments and potentially
390 forms the basis for selection (Gsell et al., 2012; Hattich et al., 2017). It is clear that
391 other environmental factors such as light intensity, temperature and nutrient
392 concentration affect the responses of physiological rates of individual *E. huxleyi*
393 strains to changing carbonate chemistry, and thus change the physiological variability
394 within populations (Zhang et al., 2015; Feng et al., 2017). However, different
395 sensitivities and requirements of each strain to the variable environments can allow
396 strains to co-exist within a population in the natural environment (Hutchinson, 1961;

397 Reed et al., 2010; Krueger-Hadfield et al., 2014). In a changing ocean, strain
398 succession is likely to occur and shift the population composition (Blanco-Ameijeiras
399 et al., 2016; Hattich et al., 2017). Strains with higher growth rates or other
400 competitive abilities may outcompete others (Schaum et al., 2013). Further, a
401 significant positive correlation between growth and POC production rate or POC
402 quota (Fig. S5) indicates that ~~the dominating strains will also take up or fix dissolved~~
403 ~~inorganic carbon faster~~higher growth rate means larger populations and then greater
404 production. ~~When extrapolated to the ocean, *E. huxleyi* blooms may increase the~~
405 ~~potential of the oceans to absorb CO₂ from the atmosphere and its carbon storage~~
406 ~~capacity (Blanco-Ameijeiras et al., 2016), which has the potential to mitigate rising~~
407 ~~CO₂ levels in the atmosphere.~~

408

409 **5 Conclusions**

410 In the present study, we found population-specific responses in physiological rates of
411 *E. huxleyi* to a broad $p\text{CO}_2$ range, which may have arisen from local adaptation to
412 environmental conditions at their origins. The existence of distinct *E. huxleyi*
413 populations and phenotypic plasticity of individual strains may both be important for
414 *E. huxleyi* when adapting to natural environmental variability and to ongoing climate
415 changes. Our results suggest that when assessing phytoplankton responses to
416 changing environments on a global scale, variability in population and strain
417 responses need to be considered, ~~and CO₂ response was modulated by other~~
418 ~~environmental factors such as temperature and light intensity.~~ In this study, we only

419 studied the effects of rising CO₂ but future studies should take into account
420 simultaneous effects from other interacting factors such as light and temperature
421 variability.

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428 *Author contributions.* YZ, LTB, UR designed the experiment. YZ, LL, RK performed
429 the experiment. YZ prepare the manuscript and all authors analysed the data,
430 reviewed and improved the manuscript.

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433 *Competing interests.* The authors declare that they have no conflict of interest.

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463 **References**

- 464 Bach, L. T., Riebesell, U., Gutowska, M. A., Federwisch, L., and Schulz, K. G.: A
465 unifying concept of coccolithophore sensitivity to changing carbonate chemistry
466 embedded in an ecological framework, *Prog. Oceanogr.*, 135, 125–138, doi:
467 10.1016/j.pocean.2015.04.012, 2015.
- 468 Bach, L. T., Riebesell, U., and Schulz, K. G.: Distinguishing between the effects of
469 ocean acidification and ocean carbonation in the coccolithophore *Emiliana*
470 *huxleyi*, *Limnol. Oceanogr.*, 56, 2040–2050, doi: 10.4319/lo.2011.56.6.2040, 2011.
- 471 Balch, W. M., Drapeau, D. T., Bowler, B. C., Lyczkowski, E. R., Lubelczyk, L. C.,
472 Painter, S. C., and Poulton, A. J.: Surface biological, chemical, and optical
473 properties of the Patagonian Shelf coccolithophore bloom, the brightest waters of
474 the Great Calcite Belt, *Limnol. Oceanogr.*, 59, 1715–1732, doi:
475 10.4319/lo.2014.59.5.1715, 2014.
- 476 Blanco-Ameijeiras, S., Lebrato, M., Stoll, H. M., Iglesias-Rodriguez, D., Müller, M.
477 N., Méndez-Vicente, A., and Oschlies, A: Phenotypic variability in the
478 coccolithophore *Emiliana huxleyi*, *PLoS ONE*, 11, e0157697, doi:
479 10.1371/journal.pone.0157697, 2016.
- 480 Bradshaw, A. D.: Evolutionary significance of phenotypic plasticity in plants, *Adv.*
481 *Genet.*, 13, 115–155, doi: 10.1016/S0065-2660(08)60048-6, 1965.
- 482 Brand, L. E.: Genetic variability and spatial patterns of genetic differentiation in the
483 reproductive rates of the marine coccolithophores *Emiliana huxleyi* and
484 *Gephyrocapsa oceanica*, *Limnol. Oceanogr.*, 27, 236–245, doi:

485 10.4319/lo.1982.27.2.0236, 1982.

486 Cai W. J.: Estuarine and coastal ocean carbon paradox: CO₂ sinks or sites of terrestrial
487 carbon incineration?, *Ann. Rev. Mar. Sci.*, 3, 123–145, doi: 10.1146/annurev-
488 marine-120709-142723, 2011.

489 Carter, B. R., Radich, J. A., Doyle, H. L., and Dickson, A. G.: An automated system
490 for spectrophotometric seawater pH measurements, *Limnol. Oceanogr.: Methods*,
491 11, 16–27, doi: 10.4319/lom.2013.11.16, 2013.

492 Clayton, T. D., and Byrne, R. H.: Spectrophotometric seawater pH measurements–
493 Total hydrogen-ion concentration scale calibration of m-cresol purple and at-sea
494 results, *Deep Sea Res. I*, 40, 2115–2129, doi: 10.1016/0967-0637(93)90048-8,
495 1993.

496 Cook, S. S., Whittock, L., Wright S. W., and Hallegraeff, G. M.: Photosynthetic
497 pigment and genetic differences between two southern ocean morphotypes of
498 *Emiliana huxleyi* (Haptophyta), *J. Phycol.*, 47, 615–626, doi: 10.1111/j.1529-
499 8817.2001.00992.x, 2011.

500 Dickson, A. G., Afghan, J. D., and Anderson, G. C.: Reference materials for oceanic
501 CO₂ analysis: a method for the certification of total alkalinity, *Mar. Chem.*, 80,
502 185–197, doi: 10.1016/S0304-4203(02)00133-0, 2003.

503 Doblin, M. A., and van Sebille, E.: Drift in ocean currents impacts intergenerational
504 microbial exposure to temperature, *Proc. Natl. Acad. Sci. USA.*, 113, 5700–5705,
505 doi: 10.1073/pnas.1521093113, 2016.

506 Duarte, C. M., and Cerbrian, J.: The fate of marine autotrophic production, *Limnol.*

507 Oceanogr., 41, 1758–1766, 1996.

508 Feng Y. Y., Roleda, M. Y., Armstrong, E., Boyd, P. W., and Hurd, C. L.:
509 Environmental controls on the growth, photosynthetic and calcification rates of a
510 Southern Hemisphere strain of the coccolithophore *Emiliana huxleyi*, Limnol.
511 Oceanogr., 62, 519–540, doi: 10.1002/lno.10364, 2017.

512 Gafar, N. A., Eyre, B. D., and Schulz, K. G.: A conceptual model for projecting
513 coccolithophorid growth, calcification and photosynthetic carbon fixation rates
514 in response to global ocean change, Front. Mar. Sci., 4, 433, doi:
515 10.3389/fmars.2017.00433, 2018.

516 Gafar, N. A., and Schulz, K. G.: A niche comparison of *Emiliana huxleyi* and
517 *Gephyrocapsa oceanica* and potential effects of climate change, Biogeosci.
518 Discuss., doi: 10.5194/bg-2018-88.

519 González-Dávila, M., and Santana-Casiano, M.: Seasonal and interannual variability
520 of sea-surface carbon dioxide species at the European Station for Time Series in
521 the Ocean at the Canary Islands (ESTOC) between 1996 and 2000, Glob.
522 Biogeochem. Cycles, 17, 1076, doi: 10.1029/2002GB001993, 2003.

523 Gsell, A. S., de Senerpont-Domis, L. N., Przytulska-Bartosiewicz, A., Mooij, W. M.,
524 van Donk, E, and Ibelings, B. W.: Genotype-by-temperature interactions may help
525 to maintain clonal diversity in *Asterionella formosa* (Bacillariophyceae), J.
526 Phycol., 48, 1197–1208, doi: 10.1111/j.1529-8817.2012.01205.x, 2012.

527 Hattich, G. S. I., Listmann, L., Raab, J., Ozod-Seradj, D., Reusch, T. B. H., and
528 Matthiessen, B.: Inter- and intraspecific phenotypic plasticity of three

529 phytoplankton species in response to ocean acidification, *Biol. Lett.*, 13,
530 20160774, doi: 10.1098/rsbl.2016.0774, 2017.

531 Henderiks, J., Winter, A., Elbrächter, M., Feistel, R., van der Plas, A., Nausch, G.,
532 and Barlow, R.: Environmental controls on *Emiliana huxleyi* morphotypes in the
533 Benguela coastal upwelling system (SE Atlantic), *Mar. Ecol. Prog. Ser.*, 448, 51–
534 66, doi:10.3354/meps09535, 2012.

535 Hoppe, C. J. M., Langer, G., and Rost, B.: *Emiliana huxleyi* shows identical
536 responses to elevated pCO₂ in TA and DIC manipulations, *J. Exp. Mar. Biol.*
537 *Ecol.*, 406, 54–62, doi: 10.1016/j.jembe.2011.06.008, 2011.

538 Hutchins, D. A., Fu, F. X., Webb, E. A., Walworth, N., and Tagliabue, A.: Taxon-
539 specific response of marine nitrogen fixers to elevated carbon dioxide
540 concentrations, *Nat. Geosci.*, 6, 790–795, doi: 10.1038/ngeo1858, 2013.

541 Hutchinson, G. E.: The paradox of the plankton, *Am. Nat.*, 95, 137–145, 1961.

542 Klaas, C., and Archer, D. E.: Association of sinking organic matter with various types
543 of mineral ballast in the deep sea: Implications for the rain ratio, *Glob.*
544 *Biogeochem. Cycles*, 16, 1116, doi: 10.1029/2001GB001765, 2002.

545 Kottmeier, D. M., Rokitta, S. D., and Rost, B.: H⁺-driven increase in CO₂ uptake and
546 decrease in HCO₃⁻ uptake explain coccolithophores' acclimation responses to
547 ocean acidification, *Limnol. Oceanogr.*, 61, 2045–2057, doi: 10.1002/lno.10352,
548 2016.

549 Kremp, A., Godhe, A., Egardt, J., Dupont, S., Suikkanen, S., Casabianca, S., and
550 Penna, A.: Intraspecific variability in the response of bloom-forming marine

551 microalgae to changed climate conditions, *Ecol. Evol.*, 2, 1195–1207, doi:
552 10.1002/ece.3.245, 2012.

553 Krueger-Hadfield, S. A., Balestreri, C., Schroeder, J., Highfield, A., Helaou ä, P.,
554 Allum, J., Moate, R., Lohbeck, K. T., Miller, P. I., Riebesell, U., Reusch, T. B. H.,
555 Rickaby, R. E. M., Young, J., Hallegraeff, G., Brownlee, C., and Schroeder, D. C.:
556 Genotyping an *Emiliania huxleyi* (prymnesiophyceae) bloom event in the North
557 Sea reveals evidence of asexual reproduction, *Biogeosciences*, 11, 5215–5234, doi:
558 10.5194/bg-11-5215-2014, 2014.

559 Krug, S. A., Schulz, K. G., and Riebesell, U.: Effects of changes in carbonate
560 chemistry speciation on *Coccolithus braarudii*: a discussion of coccolithophorid
561 sensitivities, *Biogeosciences*, 8, 771–777, doi: 10.5194/bg-8-771-2011, 2011.

562 Krumhardt, K. M., Lovenduski, N. S., Debora Iglesias-Rodriguez, M., and Kleypas, J.
563 A.: Coccolithophore growth and calcification in a changing ocean, *Prog.*
564 *Oceanogr.*, 159, 276–295, doi: [10.1016/j.pocean.2017.10.007](https://doi.org/10.1016/j.pocean.2017.10.007), 2017.

565 Langer, G., Nehrke, G., Probert, I., Ly, J., and Ziveri, P.: Strain-specific responses of
566 *Emiliania huxleyi* to changing seawater carbonate chemistry, *Biogeosciences*, 6,
567 2637–2646, doi: 10.5194/bg-6-2637-2009, 2009.

568 Levis, N. A., and Pfennig, D. W.: Evaluating ‘plasticity-first’ evolution in nature: key
569 criteria and empirical approaches, *Trends Eco. Evol.*, 31, 563–574, doi:
570 10.1016/j.tree.2016.03.012, 2016.

571 Lohbeck, K. T., Riebesell, U., and Reusch, T. B. H.: Adaptive evolution of a key
572 phytoplankton species to ocean acidification, *Nat. Geosci.*, 5, 346–351, doi:

573 10.1038/ngeo1441, 2012.

574 Müller, M. N., Trull, T. W., and Hallegraeff, G. M.: Differing responses of three
575 Southern Ocean *Emiliana huxleyi* ecotypes to changing seawater carbonate
576 chemistry, *Mar. Ecol. Prog. Ser.*, 531, 81–90, doi: 10.3354/meps11309, 2015.

577 Omar, A. M., Olsen, A., Johannessen, T., Hoppema, M., Thomas, H., and Borges, A.
578 V.: Spatiotemporal variations of $f\text{CO}_2$ in the North Sea, *Ocean Sci.*, 6, 77–89,
579 doi:10.5194/os-6-77-2010, 2010.

580 Paasche, E.: A review of the coccolithophorid *Emiliana huxleyi* (Prymnesiophyceae),
581 with particular reference to growth, coccolith formation, and calcification-
582 photosynthesis interactions, *Phycologia*, 40, 503–529, doi: 10.2216/i0031-8884-
583 40-6-503.1, 2002.

584 Palumbi, S. R.: Genetic divergence, reproductive isolation, and marine speciation.
585 *Ann. Rev. Ecol. Evol. Syst.*, 25, 547–572, doi:
586 10.1146/annurev.es.25.110194.002555, 1994.

587 Pancic, M., Hansen, P. J., Tammilehto, A., and Lundholm, N.: Resilience to
588 temperature and pH changes in a future climate change scenario in six strains of
589 the polar diatom *Fragilariopsis cylindrus*, *Biogeosciences*, 12, 4235–4244, doi: 10.
590 5194/bg-12-4235-2015, 2015.

591 Pierrot, D., Lewis, E., and Wallace, D. W. R.: MS Excel program developed for CO₂
592 system calculations, ORNL/CDIAC-105, Carbon Dioxide Information Analysis
593 Centre, Oak Ridge National Laboratory, U.S., Department of Energy, 2006.

594 Reed, T. E., Waples, R. S., Schindler, D. E., Hard, J. J., and Kinnison, M. T.:

595 Phenotypic plasticity and population viability: the importance of environmental
596 predictability, *Roc. R. Soc. B*, 277, 3391, doi: 10.1098/rspb.2010.0771, 2010.

597 Reusch, T. B. H.: Climate change in the oceans: Evolutionary versus phenotypically
598 plastic responses of marine animals and plants, *Evol. Appl.*, 7, 104–122, doi:
599 10.1111/eva.12109, 2014.

600 R ós, A. F., Pérez, F. F., Álvarez, M., Mintrop, L., González-Dávila, M., Santana-
601 Casiano, J. M., Lefèvre, L., and Watson, A. J.: Seasonal sea-surface carbon
602 dioxide in the Azores area, *Mar. Chem.*, 96, 35–51, doi:
603 10.1016/j.marchem.2004.11.001, 2005.

604 Rost, B., and Riebesell, U.: Coccolithophores and the biological pump: responses to
605 environmental changes, in: *Coccolithophores – From Molecular Biology to*
606 *Global Impact*, edited by: Thierstein, H. R. and Young, J. R., Springer, Berlin, 99–
607 125, 2004.

608 Roy, R. N., Roy, L. N., Lawson, M., Vogel, K. M., Moore, C. P., Davis W., and
609 Millero, F. J.: Thermodynamics of the dissociation of boric acid in seawater at S 5
610 35 from 0 degrees C to 55 degrees C, *Mar. Chem.*, 44, 243–248,
611 doi:10.1016/0304-4203(93)90206-4, 1993.

612 Rynearson, T. A., and Armbrust, E. V.: Genetic differentiation among populations of
613 the planktonic marine diatom *Ditylum Brightwellii* (Bacillariophyceae), *J. Phycol.*,
614 40, 34–43, doi: 10.1046/j.1529-8817.2004.03089.x, 2004.

615 Schaum, E., Rost, B., Millar, A. J., and Collins, S.: Variation in plastic responses of a
616 globally distributed picoplankton species to ocean acidification, *Nat. Clim.*

617 Change, 3, 298–302, doi: 10.1038/nclimate1774, 2013.

618 Schaum, E., Rost, B., Collins, S.: Environmental stability affects phenotypic evolution
619 in a globally distributed marine picoplankton, *The ISME Journal*, 10, 75–84, doi:
620 10.1038/ismej.2015.102, 2015.

621 Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.:
622 Temperature modulates coccolithophorid sensitivity of growth, photosynthesis
623 and calcification to increasing seawater $p\text{CO}_2$, *PLoS ONE*, 9, e88308, doi:
624 10.1371/journal.pone.0088308, 2014.

625 Smith, H. E. K., Tyrrell, T., Charalampopoulou, A., Dumousseaud, C., Legge, O. J.,
626 Birchenough, S., Pettit, L. R., Garley, R., Hartman, S. E., Hartman, M. C., Sagoo,
627 N., Daniels, C. J., Achterberg, E. P., and Hydes, D. J.: Prodominance of heavily
628 calcified coccolithophores at low CaCO_3 saturation during winter in the Bay of
629 Biscay, *Proc. Natl. Acad. Sci. USA*, 109, 8845–8849, doi:
630 10.1073/pnas.1117508109, 2012.

631 Wisshak, M., Form, A., Jakobsen, J., and Freiwald, A.: Temperate carbonate cycling
632 and water mass properties from intertidal to bathyal depths (Azores),
633 *Biogeosciences*, 7, 2379–2396, doi:10.5194/bg-7-2379-2010, 2010.

634 Zhang, Y., Klapper, R., Lohbeck, K. T., Bach, L. T., Schulz, K. G., Reusch, T. B. H.,
635 and Riebesell, U.: Between- and within-population variations in thermal reaction
636 norms of the coccolithophore *Emiliana huxleyi*, *Limnol. Oceanogr.*, 59, 1570–
637 1580, doi: 10.4319/lo.2014.59.5.1570, 2014.

638 Zhang, Y., Bach, L. T., Schulz, K. G., and Riebesell, U.: The modulating effect of

639 light intensity on the response of the coccolithophore *Gephyrocapsa oceanica* to
640 ocean acidification, *Limnol. Oceanogr.*, 60, 2145–2157, doi:10.1002/lno.10161,
641 2015.

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662 **Figure Legends**

663 **Figure 1.** Optimum curve responses of measured and relative growth, particulate
664 organic (POC) and inorganic carbon (PIC) production rates of three *Emiliana huxleyi*
665 populations to a $p\text{CO}_2$ range from 120 μatm to 2630 μatm . Responses of measured (a)
666 and relative (b) growth rates to $p\text{CO}_2$. Responses of measured (c) and relative (d)
667 POC production rates to $p\text{CO}_2$. Responses of measured (e) and relative (f) PIC
668 production rates to $p\text{CO}_2$. Using the nonlinear regression model derived by Bach et al.
669 (2011), the curves were fitted based on average growth, POC and PIC production
670 rates of six strains from the Azores and Bergen, and of five strains from the Canary
671 Islands. Vertical error bars represent standard deviations of six growth, POC and PIC
672 production rates for the Azores and Bergen populations, and five growth, POC and
673 PIC production rates for the Canary Islands population. Horizontal error bars
674 represent standard deviations of six $p\text{CO}_2$ levels for the Azores and Bergen
675 populations and five $p\text{CO}_2$ levels for the Canary Islands populations. At the
676 population levels, 120 μatm and 2630 μatm was the lowest and highest $p\text{CO}_2$ level,
677 respectively.

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679 **Figure 2.** Calculated optimum $p\text{CO}_2$, calculated maximum value and fitted relative
680 sensitivity constant of growth, POC and PIC production rates of each population. (a)
681 optimum $p\text{CO}_2$ of growth rate; (b) optimum $p\text{CO}_2$ of POC production rates; (c)
682 optimum $p\text{CO}_2$ of PIC production rates; (d) maximum growth rate, (e) maximum
683 POC production rate, (f) maximum PIC production rate; (g) relative sensitivity

684 constant of growth rate; **(h)** relative sensitivity constant of POC production rate; **(i)**
685 relative sensitivity constant of PIC production rate. The line in the middle of each box
686 indicates the mean of 6 or 5 optimum $p\text{CO}_2$, 6 or 5 maximum values, and 6 or 5
687 relative sensitivity constants for growth, POC and PIC production rates in each
688 population. Bars indicate the 99% confidence interval. The maximum or minimum
689 data is shown as the small line on the top or bottom of the bar, respectively. Letters in
690 each panel represent statistically significant differences (Tukey HSD, $p < 0.05$).

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692 **Figure 3.** Optimum curve responses of growth, POC and PIC production rates of
693 individual *E. huxleyi* strains in the Azores (left), Bergen (medium) and Canary Islands
694 (right) populations to a CO_2 range from 115 μatm to 3070 μatm . Growth rates of each
695 strain as a function of $p\text{CO}_2$ within the Azores (**a**), Bergen (**b**) and Canary Islands (**c**)
696 populations. POC production rates of each strain as a function of $p\text{CO}_2$ within the
697 Azores (**d**), Bergen (**e**) and Canary Islands (**f**) populations. PIC production rates of
698 each strain as a function of $p\text{CO}_2$ within the Azores (**g**), Bergen (**h**) and Canary
699 Islands (**i**) populations. At the strain levels, 115 μatm and 3070 μatm was the lowest
700 and highest $p\text{CO}_2$ level, respectively.

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706 **Table 1.** Surface seawater CO₂ levels and pH at the Azores, Bergen and Canary
 707 Islands.

	Location	Mean seasonal CO₂ (µatm)	Mean seasonal pH (total scale)	CO₂ variability (µatm)	References
Azores	38°34'N, 28°42'W	320 – 400	8.005 – 8.05	80	R ós et al., 2005 Wisshak et al., 2010
Bergen	60°18'N, 05°15'E	240 – 400	7.98 – 8.22	200	Omar et al., 2010
Canary Islands	27°58'N, 15°36'W	320 – 400	8.005 – 8.05	80	Gonz ález-D ávila et al., 2003

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725 **Table 2.** Carbonate chemistry parameters (mean values for the beginning and end of
726 the incubations) of the artificial seawater for each *Emiliana huxleyi* population. pH
727 and TA samples were collected and measured before and at the end of incubation.
728 Data are expressed as mean values of six strains in the Azores and Bergen population,
729 and five strains in the Canary Islands population.

	$p\text{CO}_2$ (μatm)	pH (total scale)	TA (μmol kg^{-1})	DIC (μmol kg^{-1})	HCO_3^- (μmol kg^{-1})	CO_3^{2-} (μmol kg^{-1})	CO_2 (μmol kg^{-1})	Ω
Azores	125 \pm 3	8.46 \pm 0.01	2358 \pm 12	1844 \pm 11	1485 \pm 13	355 \pm 5	5 \pm 0	8.5 \pm 0.1
	300 \pm 20	8.16 \pm 0.03	2339 \pm 27	2031 \pm 17	1803 \pm 18	218 \pm 13	11 \pm 1	5.2 \pm 0.3
	360 \pm 19	8.09 \pm 0.02	2322 \pm 30	2052 \pm 14	1849 \pm 9	190 \pm 10	13 \pm 1	4.5 \pm 0.3
	500 \pm 26	7.97 \pm 0.02	2301 \pm 23	2100 \pm 16	1933 \pm 14	149 \pm 8	18 \pm 1	3.5 \pm 0.2
	695 \pm 20	7.85 \pm 0.01	2317 \pm 11	2167 \pm 13	2023 \pm 14	118 \pm 2	25 \pm 1	2.8 \pm 0.1
	875 \pm 40	7.76 \pm 0.02	2320 \pm 19	2206 \pm 13	2076 \pm 10	99 \pm 5	32 \pm 1	2.4 \pm 0.1
	1110 \pm 119	7.66 \pm 0.05	2303 \pm 19	2222 \pm 23	2101 \pm 25	80 \pm 8	40 \pm 4	1.9 \pm 0.2
	1315 \pm 104	7.59 \pm 0.03	2308 \pm 18	2251 \pm 26	2133 \pm 26	70 \pm 4	48 \pm 4	1.7 \pm 0.1
	1665 \pm 107	7.50 \pm 0.03	2311 \pm 11	2286 \pm 15	2169 \pm 14	57 \pm 3	60 \pm 4	1.4 \pm 0.1
	1935 \pm 175	7.44 \pm 0.04	2308 \pm 15	2302 \pm 24	2183 \pm 21	50 \pm 4	70 \pm 6	1.2 \pm 0.1
2490 \pm 132	7.33 \pm 0.02	2320 \pm 12	2350 \pm 15	2220 \pm 13	40 \pm 2	90 \pm 5	0.9 \pm 0.1	
Bergen	120 \pm 3	8.47 \pm 0.01	2354 \pm 18	1834 \pm 18	1470 \pm 17	359 \pm 2	4 \pm 0	8.6 \pm 0.1
	290 \pm 16	8.17 \pm 0.02	2337 \pm 21	2024 \pm 12	1793 \pm 14	220 \pm 10	11 \pm 1	5.3 \pm 0.2
	355 \pm 18	8.10 \pm 0.02	2315 \pm 23	2045 \pm 11	1840 \pm 7	192 \pm 10	13 \pm 1	4.6 \pm 0.2
	490 \pm 18	7.98 \pm 0.02	2302 \pm 19	2096 \pm 14	1926 \pm 12	152 \pm 6	18 \pm 1	3.6 \pm 0.1
	670 \pm 22	7.86 \pm 0.01	2317 \pm 11	2162 \pm 10	2016 \pm 10	121 \pm 3	24 \pm 1	2.9 \pm 0.1
	855 \pm 52	7.77 \pm 0.03	2326 \pm 19	2206 \pm 15	2074 \pm 14	101 \pm 6	30 \pm 2	2.4 \pm 0.1
	1080 \pm 53	7.67 \pm 0.02	2316 \pm 26	2232 \pm 20	2110 \pm 18	83 \pm 5	39 \pm 2	2.0 \pm 0.1
	1280 \pm 71	7.60 \pm 0.02	2318 \pm 15	2257 \pm 17	2138 \pm 17	72 \pm 4	46 \pm 3	1.7 \pm 0.1
	1550 \pm 122	7.52 \pm 0.03	2300 \pm 19	2266 \pm 28	2150 \pm 27	60 \pm 4	56 \pm 4	1.4 \pm 0.1
	1800 \pm 235	7.47 \pm 0.05	2301 \pm 19	2286 \pm 33	2168 \pm 30	53 \pm 6	65 \pm 9	1.3 \pm 0.1
2280 \pm 147	7.37 \pm 0.02	2309 \pm 20	2326 \pm 27	2201 \pm 24	42 \pm 2	82 \pm 5	1.0 \pm 0.1	
Canary Islands	130 \pm 3	8.45 \pm 0.01	2344 \pm 38	1842 \pm 32	1491 \pm 26	347 \pm 7	5 \pm 0	8.3 \pm 0.2
	310 \pm 11	8.15 \pm 0.01	2317 \pm 24	2020 \pm 25	1798 \pm 25	210 \pm 4	11 \pm 1	5.0 \pm 0.1
	375 \pm 14	8.07 \pm 0.01	2295 \pm 14	2040 \pm 12	1846 \pm 13	182 \pm 5	14 \pm 1	4.3 \pm 0.1
	505 \pm 32	7.96 \pm 0.02	2297 \pm 19	2097 \pm 20	1930 \pm 23	148 \pm 7	18 \pm 1	3.5 \pm 0.2
	695 \pm 18	7.85 \pm 0.01	2312 \pm 20	2163 \pm 17	2020 \pm 15	118 \pm 3	25 \pm 1	2.8 \pm 0.1
	925 \pm 73	7.74 \pm 0.04	2319 \pm 26	2211 \pm 15	2083 \pm 12	95 \pm 8	33 \pm 3	2.3 \pm 0.1
	1180 \pm 53	7.64 \pm 0.02	2310 \pm 25	2239 \pm 20	2120 \pm 19	76 \pm 4	43 \pm 2	1.8 \pm 0.1
	1380 \pm 104	7.58 \pm 0.03	2323 \pm 5	2271 \pm 10	2154 \pm 11	68 \pm 5	50 \pm 4	1.6 \pm 0.1
	1740 \pm 98	7.48 \pm 0.02	2319 \pm 16	2298 \pm 16	2180 \pm 15	55 \pm 3	63 \pm 4	1.3 \pm 0.1
	2140 \pm 258	7.40 \pm 0.05	2312 \pm 9	2320 \pm 16	2197 \pm 13	46 \pm 5	78 \pm 10	1.1 \pm 0.1
2630 \pm 284	7.31 \pm 0.04	2317 \pm 13	2363 \pm 20	2225 \pm 14	37 \pm 3	98 \pm 8	0.8 \pm 0.1	

731 **Table 3.** Calculated optimum $p\text{CO}_2$, calculated maximum value (V_{max}) and fitted
732 relative sensitivity constant (rs , %) of growth, POC and PIC production rates of each
733 *E. huxleyi* strain.

strain	Growth rate			POC production rate			PIC production rate		
	optimum $p\text{CO}_2$ (μatm)	V_{max} (d^{-1})	rs	optimum $p\text{CO}_2$ (μatm)	V_{max} ($\text{pg C cell}^{-1} \text{d}^{-1}$)	rs	optimum $p\text{CO}_2$ (μatm)	V_{max} ($\text{pg C cell}^{-1} \text{d}^{-1}$)	rs
A23	392	1.21	0.22	673	12.47	0.50	323	13.45	0.38
A22	436	1.27	0.16	591	17.33	0.33	635	12.28	0.40
A21	392	1.25	0.22	707	15.45	0.50	396	16.73	1.11
A19	371	1.26	0.24	512	16.17	0.56	480	18.92	0.67
A13	244	1.08	0.13	756	9.84	0.63	471	11.72	0.57
A10	432	1.32	0.20	549	14.42	0.48	385	11.69	0.24
B95	534	1.26	0.10	762	13.46	0.20	562	9.13	0.33
B63	436	1.26	0.11	633	16.66	0.27	615	12.93	0.45
B62	456	1.29	0.11	945	17.27	0.18	488	14.00	0.43
B51	499	1.29	0.11	660	16.77	0.35	492	11.87	0.48
B41	542	1.25	0.09	984	18.34	0.38	553	9.46	0.37
B17	490	1.32	0.14	761	15.19	0.30	625	12.77	0.47
C98	400	1.03	0.16	644	8.44	0.54	440	6.40	0.31
C91	393	0.97	0.21	413	4.83	0.60	195	10.87	0.33
C90	384	0.97	0.12	546	8.28	0.34	284	8.52	0.50
C41	393	1.01	0.14	609	7.64	0.45	545	11.15	0.30
C35	378	1.05	0.17	596	8.87	0.44	464	12.68	0.34

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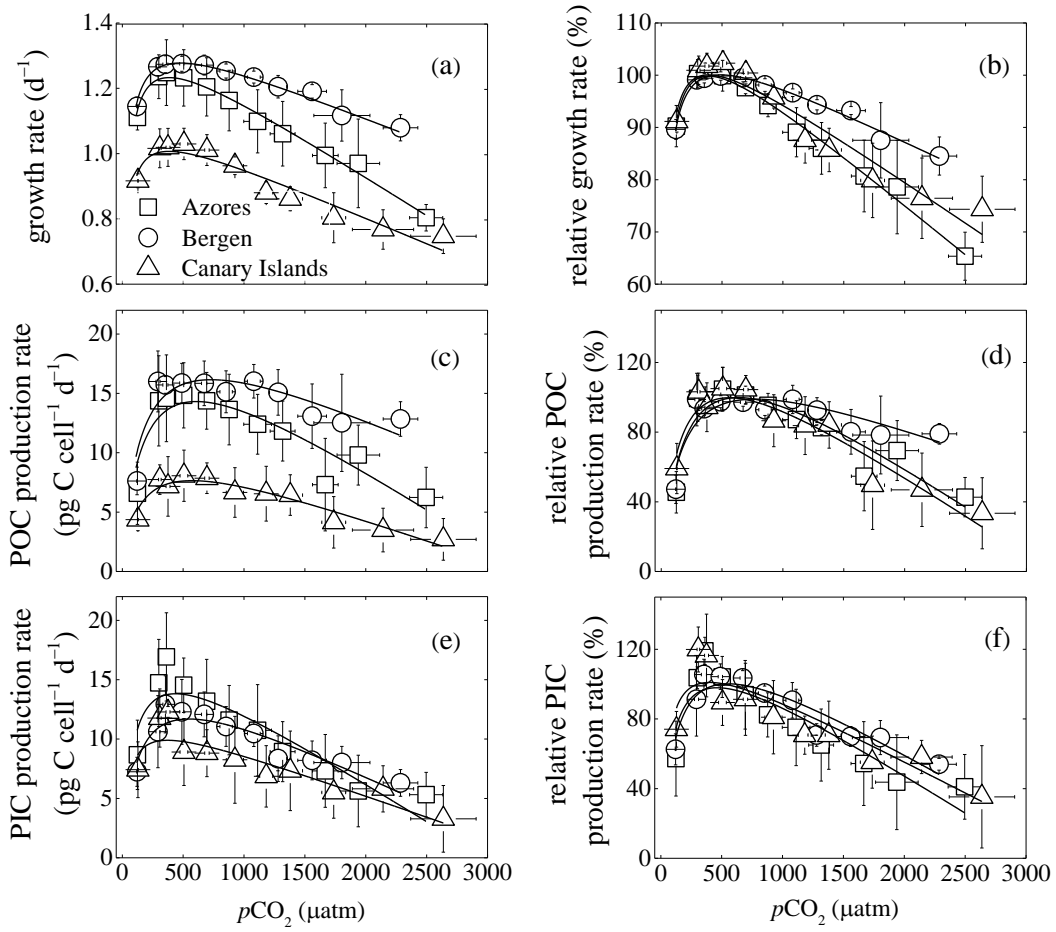
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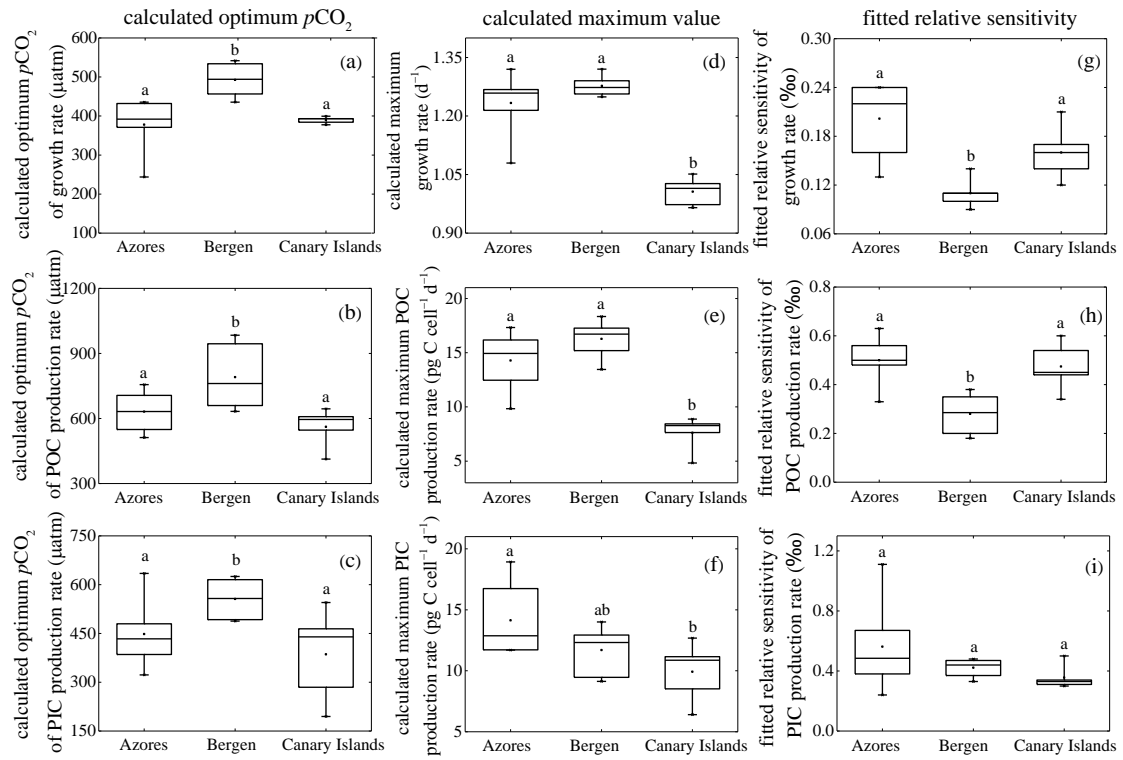
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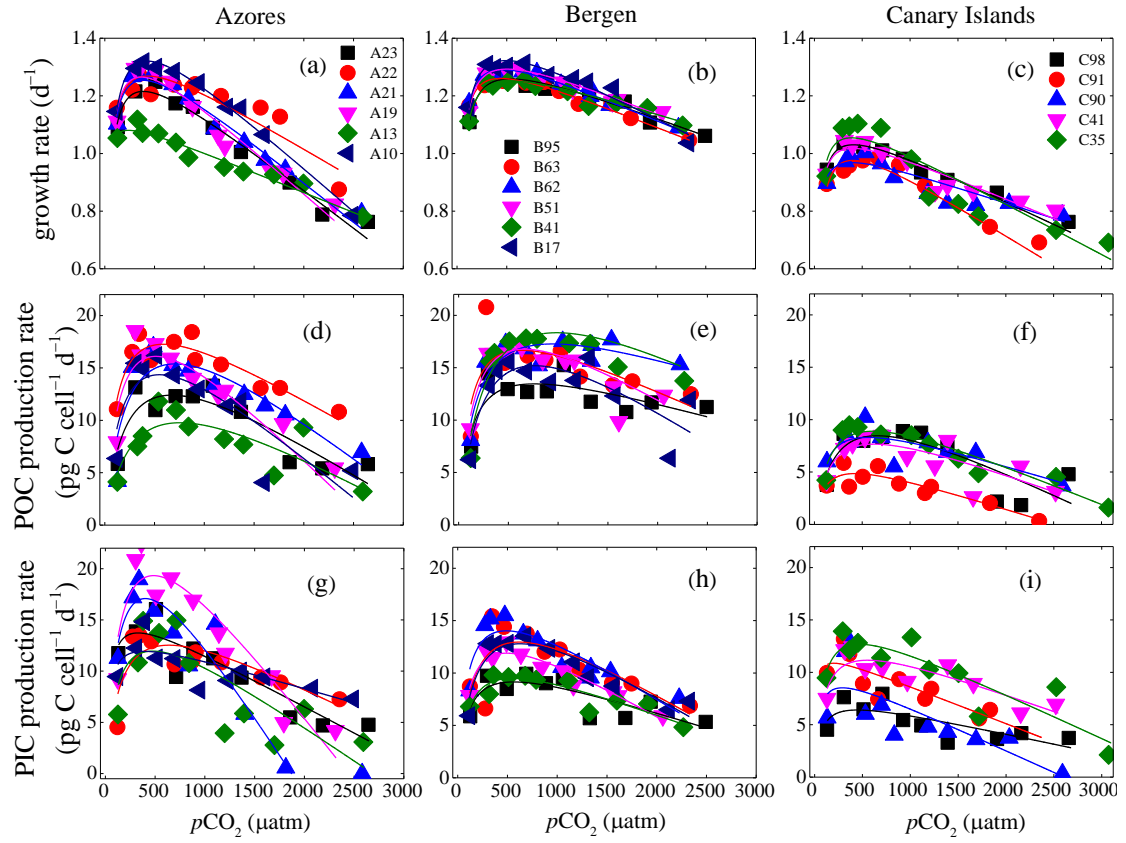
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