



1 **Plant functional traits determined the latitudinal**
2 **variations in soil microbial functions: evidence from a**
3 **forest transect in China**
4

5 Zhiwei Xu^{1,2}, Guirui Yu^{3,4,*}, Xinyu Zhang^{3,4,*}, Ruili Wang⁵, Ning Zhao⁶, Nianpeng He^{3,4}, Qiufeng
6 Wang^{3,4}

7 1. Institute for Peat and Mire Research, College of Geographical Sciences, Northeast Normal
8 University, Changchun, 130024, China

9 2. Jilin Provincial Key Laboratory for Wetland Ecological Processes and Environmental Change in
10 the Changbai Mountains, Changchun, 130024, China

11 3. Key Laboratory of Ecosystem Network Observation and Modeling, Institute of Geographic
12 Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 10010, China.

13 4. College of Resources and Environment, University of Chinese Academy of Sciences, Beijing,
14 100190, China

15 5. College of Forestry, Northwest A&F University, Yangling, 712100, China

16 6. Key Laboratory of Remote Sensing of Gansu Province, Heihe Remote Sensing Experimental
17 Research Station, Cold and Arid Regions Environmental and Engineering Research Institute,
18 Chinese Academy of Sciences, Lanzhou 730000, China

19

20

21 * Corresponding author at: Key Laboratory of Ecosystem Network Observation and Modeling,
22 Institute of Geographic Sciences and Natural Resources Research, Chinese Academy of Sciences,
23 Beijing 100101, China.No. 11A, Datun Road, Chaoyang District, Beijing, 100101, China. Tel.: +86-
24 10-64889268; fax: +86 10 64889432.

25 E-mail: yugr@igsnr.ac.cn (G. Y.), zhangxy@igsnr.ac.cn (X. Z.)



26 **Abstract.** Plant functional traits have increasingly been studied as determinants of ecosystem
27 properties, especially for soil biogeochemical processes. While the relationships between biological
28 community structures and ecological functions are a central issue in ecological theory, these
29 relationships remain poorly understood at the large scale. We selected nine forests along the North–
30 South Transect of Eastern China (NSTEC) to determine how plant functional traits influence the
31 latitudinal pattern of soil microbial functions, and how soil microbial communities and functions are
32 linked at the regional scale. We found that there was considerable variation in the profiles of different
33 substrate use along the NSTEC. Soil microorganisms from temperate forests mainly metabolized
34 high-energy substrates, while those from subtropical forests used all the substrates equally. The soil
35 silt content and plant functional traits together shaped the biogeographical pattern of the soil
36 microbial substrate use. Soil organic matter decomposition rates were significantly higher in
37 temperate forests than in subtropical and tropical forests, which was consistent with the pattern of
38 soil microbial biomass carbon concentrations. Soil organic matter decomposition rates were also
39 significantly and negatively related to soil dissolved organic carbon concentrations, and carboxylic
40 acid, polymer, and miscellaneous substrates. The soil microbial community structures and functions
41 were significantly correlated along the NSTEC. Soil carbohydrate and polymer substrate use were
42 mainly related to soil G⁺ bacterial and actinomycetes biomass, while the use of amine and
43 miscellaneous substrates were related to soil G⁻ bacteria, fungal biomass, and the F/B ratio. The
44 contributions of different groups of microbial biomass to the production of soil enzyme activities
45 differed. The relationship between soil microbial community structure and functions supported that
46 there was functional dissimilarity.



47 **1 Introduction**

48 The catabolic diversity of soil microbial communities is a useful indicator of how microbial functions
49 adapt to environmental stress and can be used to test fundamental questions about soil biological
50 resistance and resilience (Jagadamma et al., 2014; Swallow and Quideau, 2015). We need robust
51 information about functional diversity to understand the role of microbial communities in different
52 environments (Preston-Mafham et al., 2002). Biological community structure and function are
53 intimately linked in ecological processes, and their relationships are a central issue in ecological
54 theory (Talbot et al., 2014). Therefore, a major goal in ecological research is to identify and
55 understand the mechanisms and relationships that control the structure and function of microbial
56 community at large spatial scales.

57 Numerous studies have documented how environmental and anthropogenic perturbations
58 impact on the structure, diversity (Zhou et al., 2016), and enzyme activities (Peng and Wang, 2016;
59 Xu et al., 2017) of soil microbial communities (Tu et al., 2016), and have reported forests in the same
60 climatic zone develop similar microbial communities. Other researchers have examined spatial
61 patterns in soil microbial function at different scales. For example, in their study of Changbai
62 Mountain, China, Tian et al. (2015) found that the soil microbial metabolic activity was moderately
63 spatially dependent, and that the functional diversity was much more spatially dependent. Other
64 researchers have reported differences in soil microbial activities among forest types, with high local
65 variation and significant separation along regional climate gradients (Brockett et al., 2012; Cao et al.,
66 2016). Soil microbes from different climatic zones have different affinities for carbon substrates. For
67 example, microorganisms from boreal pine forest soils used carboxylic acids more efficiently, but
68 decomposed amino acids much less efficiently, than microorganisms from temperate forest soils
69 (Klimek et al., 2016). Despite this, because of limitations in analytical methods, questions still remain
70 about how soil microbial functions vary at the regional scale.

71 The functional diversity of soil microbial communities is regulated by physico-chemical soil
72 properties (Gartzia-Bengoetxea et al., 2016), climate (Cao et al., 2016), and the composition of plant
73 cover (Sherman and Steinberger, 2012). For example, the geographic patterns of soil microbial
74 activity reflect the climate, soil pH, and total phosphorus concentrations over large spatial scales (Cao



75 et al., 2016). Research has shown that substrate-induced respiration rates were higher in soil microbial
76 communities that developed under beech and holm oak forests than under oak and pine forests
77 (Gartzia-Bengoetxea et al., 2016). Plant functional traits have increasingly been studied as
78 determinants of ecosystem properties, especially for soil biogeochemical processes (De Vries et al.,
79 2012; Pei et al., 2016). Soil bacteria phospholipid fatty-acids (PLFAs) were found to be positively
80 correlated with the community-weighted means (CWM) of plant functional traits (leaf nitrogen (N)
81 concentration) (De Vries et al., 2012). The plant leaf dry matter content and the leaf carbon (C) to
82 nitrogen (N) ratio both influence the multivariate soil microbial community structure, and these
83 factors positively promote the abundances of specific microbial functional groups (Pei et al., 2016).
84 Limited soil resources, particularly in tropical forests, mean that soil microorganisms may be more
85 reliant on plants than soil for C and nutrients via rhizosphere exudation or litter production, which
86 varies among plant species (Russell et al., 2007; Raich et al., 2014; Waring et al., 2015). While soil
87 functional diversity has been used as an indicator of microbial metabolic potential, there have been
88 few studies of the integrated effects of climate, vegetation, and soil substrate availability on large-
89 scale soil microbial functional diversity.

90 Although the functional characteristics of soil microorganisms are at least as important as their
91 patterns of community structure in biogeochemical studies, the links between microbial community
92 structure and microbial functions are poorly understood. There are two current hypotheses about how
93 microbes determine ecosystem process rates. In functional redundancy, different microbes perform
94 the same function and so changes in the microbial community structure do not necessarily lead to a
95 change in soil function (Balsler and Firestone, 2005; Strickland et al., 2009). For example, Banerjee
96 et al. (2016) showed that the abundance of different bacterial and fungal groups changed by up to
97 300-fold under straw- and nutrient-amended treatments but that the decomposition rate remained
98 similar, indicating possible functional redundancy. The functional redundancy hypothesis has
99 recently been challenged by a counter-hypothesis, referred to as functional dissimilarity, which
100 suggests that diversity brings stability, and that every species plays a unique role in ecosystem
101 function (Fierer et al., 2007; Waldrop and Firestone, 2006). Soil microbial community composition
102 therefore, combined with environmental variables, may ultimately determine ecosystem process rates.
103 Waldrop and Firestone (2006) showed that G^+ bacteria were mainly responsible for the



104 decomposition of pine needles and soil organic matter, but G^- bacteria were mainly responsible for
105 the decomposition of starch and xylose, which are easy to break down. Philippot et al. (2013), when
106 studying the diversity of denitrifiers, showed that the loss of microbial diversity could result in
107 decreases of between 4- and 5-fold in denitrification activity. In the Mediterranean, losses in the mass
108 of decomposing leaf litter from shrub species accelerated as detritivore assemblages became more
109 functionally dissimilar (Coulis et al., 2015). These studies suggest that the importance of functional
110 redundancy in soil microbial communities has been overstated, so studies of the relationships
111 between soil microbial communities and their functions in natural ecosystems are urgently needed.

112 The North-South Transect of Eastern China (NSTEC) extends from a cold temperate coniferous
113 forest in the north to a tropical rainforest in the south, and includes almost all the forest types found
114 in the Northern Hemisphere (Zhang and Yang, 1995) (Fig. 1 and Table 1). This transect, therefore,
115 provides the optimal environment for investigating geographical patterns in microbial communities
116 and their responses to environmental changes at the large scale. In this study, we examined spatial
117 patterns in soil labile C concentrations, soil organic matter (SOM) decomposition rates, and
118 metabolic activity and functional diversity of microbes in nine forest biomes along the NSTEC. We
119 assessed how abiotic factors, such as climate, soil physical and chemical properties, and biotic factors,
120 in the form of community-weighted means (CWM) of plant functional traits, contributed to soil
121 functional diversity at the regional scale. We also examined the links between soil microbial
122 community structure (PLFAs) and function (SOM decomposition rate, enzyme activities, and
123 microbial substrate use). We tested the following three hypotheses in this study, that (1) the profiles
124 of soil microbial substrate use would vary along a latitudinal gradient, (2) biogeographical patterns
125 of soil microbial substrate use would be constrained by climate and plant functional traits, and (3)
126 the relationships between soil microbial community and functions would demonstrate functional
127 dissimilarity.

128 **2 Material and methods**

129 **2.1 Study area and soil sampling**

130 We selected nine forest ecosystems along the NSTEC, namely Huzhong (HZ), Liangshui (LS),



131 Changbai (CB), Dongling (DL), Taiyue (TY), Shennong (SN), Jiulian (JL), Dinghu (DH), and
132 Jianfeng (JF) (18°44'–51°46'N, 128°53'–108°51'E) (Fig. 1, Table 1). For further information
133 regarding soil characterization and site descriptions see Xu et al. (2017). Forest soils have been
134 classified following the U.S. soil taxonomy and are described in Table 1 (Soil Survey Staff, 2010),
135 where information about the climate and the dominant vegetation at each site is also presented.

136 Soil samples were collected from four random plots in July and August 2013. The information
137 of the sampling process are available in Xu et al. (2017). Briefly, we established four sampling plots
138 measured 30 × 40 m and collected soil samples from a depth of between 0 and 10 cm at between 30
139 and 50 points in each plot along an S-shape. On return to the laboratory, the fresh soil samples were
140 immediately sieved through a 2-mm mesh and subdivided into three subsamples. One subsample was
141 stored briefly at 4°C until analysis for soil enzyme activities and soil pH. Another was stored briefly
142 at –20°C until analysis for PLFAs and Eco-Biolog. The third was air-dried, sieved through a 0.25
143 mm mesh, and analyzed for soil nutrients.

144 2.2 Soil analysis

145 Soil pH was measured at a soil-to-water ratio of 1:2.5. Soil organic carbon (SOC) and total N (TN)
146 concentrations were determined by dry combustion of ground samples (100-mesh) in a C/N analyzer
147 (Elementar, Vario Max CN, Germany). Total phosphorus (TP) was determined with a flow injection
148 auto-analyzer following digestion with H₂SO₄-HClO₄ (Huang et al., 2011). After extraction with
149 distilled water at a soil: distilled water ratio of 1:5, dissolved organic carbon (DOC) concentrations
150 were determined by Liqui TOC II (Elementar, Liqui TOC II, Germany) (Jones and Willett, 2006).
151 Soil microbial biomass carbon (MBC) was measured using the chloroform fumigation and direct
152 extraction technique (Vance et al., 1987). A conversion factor of 2.64 was used to convert extracted
153 C to biomass C. The silt fractions (<53 μm) of the samples were separated by wet-sieving and then
154 were freeze-dried in the laboratory, as described by Six et al. (2000). The soil properties are shown
155 in Table 2. We followed the method described by Bååth et al. (2003) for PLFA analysis and PLFAs
156 are expressed in units of nmol g⁻¹. The four enzymatic activities of β-glucosidase (BG), N-
157 acetylglucosaminidase (NAG), acid phosphatase (AP), and leucine aminopeptidase (LAP)
158 responsible for soil C, N, and phosphorous cycling, were measured following the procedure outlined



159 in Saiya-Cork et al. (2002) and are expressed in units of $\text{nmol h}^{-1} \text{g}^{-1}$. Information about PLFA and
160 enzyme activities are presented in Table S1.

161 2.3 Vegetation data

162 As described by Xu et al. (2018), we collected litter and sun-exposed and mature leaves (leaf blades
163 for trees) from between five and ten individuals of each plant species at each site and determined
164 their TN and TC concentrations. We calculated the specific leaf area (SLA, the one-sided area of a
165 fresh leaf divided by its oven-dried mass, $\text{m}^2 \text{kg}^{-1}$), leaf dry matter content (LDMC, the oven-dried
166 mass of a leaf divided by its water-saturated fresh mass, mg g^{-1}), leaf C concentrations (leaf C, g kg^{-1}),
167 and leaf N concentrations (leaf N, g kg^{-1}) for ten fully expanded leaves of each sampled individual.
168 We also calculated the community-weighted means (CWM), as described by Garnier et al. (2004).
169 The diversity of the tree species and plant functional traits are summarized in Table S2.

170 2.4 Microbial substrate use

171 Microbial functional diversities were determined using a Biolog EcoPlate™ (Biolog Inc., Hayward,
172 California, USA) as described by Garland and Mills (1991). Briefly, approximately 10 g of fresh soil
173 was suspended in 100 ml saline solution (0.85% NaCl) and shaken on an orbital shaker for 30 min at
174 190 rpm. A 150 μl aliquot of supernatant from 1:1 000 dilutions of each soil sample was added to
175 each well. The plates were incubated at 25°C, and the absorbance at 590 nm was measured using a
176 microplate reader (GENios Pro™, Tecan Trading AG, Männedorf, Switzerland) every 24 h up to 240
177 h (0, 24, 48, 72, 96, 120, 144, 168, 192, 216, and 240 h).

178 The Richness (R), Shannon-Weiner diversity index (H'), Shannon evenness index (E), and
179 Simpson dominance index (D) were calculated from the absorption values after EcoPlate™
180 incubation for 96 h (Gomez et al., 2006). Additionally, the 31 C sources were divided into six groups,
181 namely carbohydrates, carboxylic acids, amines, amino acids, polymers, and miscellaneous, as
182 suggested by Zak et al. (1994). The average absorbance of all C sources within each group was
183 computed as the intensity of the single substrate use. The soil microbial metabolic intensities (S) were
184 estimated by the area underneath $AWCD$ vs. t , and were obtained by integrating the equation against
185 time t (Guckert et al., 1996):



186
$$S = \sum [(v_i + v_{i-1})/2 \times (t_i + t_{i-1})]$$

187 where v_i was average optical density of the i th incubation time.

188 2.5 SOM decomposition rate

189 Four replicates from each sampling site with a 60% water-holding capacity were incubated at 20°C.
190 In brief, 40 g of each fresh soil sample were put into a 150-ml incubation bottle, and the samples
191 were then adjusted so that their moisture content corresponded to a water-holding capacity of 60%.
192 During the 4-week incubation period, the soil respiration rates were measured on days 1, 7, 14, 21,
193 and 28 using an automatic system. The SOM decomposition rates were calculated as described in the
194 study of Xu et al. (2015).

195 2.6 Statistical analysis

196 One-way analysis of variance (ANOVA) followed by a post hoc Tukey HSD test were used to test
197 the significance of the differences among the soil properties, C use, functional diversity, and SOM
198 decomposition rates in the different forest ecosystems. We tested the relationships between labile C,
199 soil microbial community structure, microbial function, and the SOM decomposition rates with the
200 Pearson correlation test. Differences were considered significant when $P < 0.05$, with marginal
201 significance set at $P < 0.01$. All ANOVA and regression analyses were performed using SPSS 19.0
202 for Windows. Data are reported as the mean \pm SE.

203 We used redundancy analysis (RDA) to examine the relationship between the environmental
204 variables and soil microbial substrate use. The environmental variables were the same as those
205 described in Xu et al. (2018), including climate, soil properties, litter properties, and plant functional
206 traits. Before RDA, we conducted forward selection of the environmental variables that were
207 significantly correlated with variations in the microbial substrate use profile using stepwise
208 regression and the Monte Carlo Permutation Test. We used CANOCO software 4.5 (Ter Braak and
209 Smilauer 2002) for the RDA and stepwise regression. The environmental properties, which were
210 significantly correlated with the microbial substrate use in the RDA, were stressed in the plots.

211 3 Results



212 3.1 Patterns in the microbial substrate use, soil labile carbon, and SOM decomposition rates

213 Of the forests along the NSTEC, the C metabolic intensity of soil microbes was lowest in HZ and LS;
214 the C metabolic intensity of soil microbes differed significantly between JF and the other forests (Fig.
215 2), which indicates that the color development was significantly higher in the tropical forest soils
216 than in the subtropical and temperate forest soils and is consistent with the variations in the AWCD
217 (Fig.S1). The average values of R , H' , and D were significantly different among the nine forest soils
218 and were highest in JF, SN, and CB (Table 3).

219 Across the nine forests, soil microorganisms used the six substrate groups in the same order; the
220 carboxylic acid substrate was used most, followed by amino acids, carbohydrates, polymers, amines,
221 and miscellaneous substrates (Fig. 3). Microorganisms in the boreal and temperate forests mainly
222 metabolized carbohydrates, amino acids, and carboxylic acids, while those from the subtropical and
223 tropical forests used the substrates in equal proportions.

224 Overall, soil MBC concentrations in the boreal and temperate forests were three to eight times
225 higher than those of the subtropical and tropical forests. In contrast, the average DOC concentration
226 in the tropical and subtropical forest soils ranged from 311 to 458 mg kg⁻¹, which was significantly
227 higher than the average concentration in the temperate and boreal forest soils, where the average
228 concentrations ranged from 204 to 284 mg kg⁻¹ (Table 2). The average SOM decomposition rates in
229 the subtropical forests ranged from 0.64 to 2.42 μg C g⁻¹ d⁻¹, and were significantly lower than the
230 rates in the temperate forests, which ranged from 3.43 to 4.61 μg C g⁻¹ d⁻¹ (Table S3).

231 3.2 Effect of environmental properties on soil microbial substrate use

232 Redundancy analysis showed that the variations in soil microbial substrate use were strongly and
233 positively correlated with the CWM values of LDMC, leaf N, and leaf C, and strongly and negatively
234 correlated with the soil silt content and SMC (Fig. 4). The RDA2 of soil microbial substrate use was
235 strongly positively correlated with TN and SOC, but negatively correlated with mean annual
236 precipitation (MAP) (Fig. 5). RDA1 mainly represented the plant functional traits, soil texture, and
237 micro-meteorological conditions, while RDA2 represented climate and soil nutrients. Overall, the
238 soil silt content and the CWM values of plant functional traits were the main predictors of the



239 latitudinal variation in the soil microbial substrate use along the NSTEC.

240 3.3 Relationships between soil microbial substrate use, enzyme activities, and PLFAs

241 Microbial carbohydrate use was positively related with bacterial biomass and actinomycic biomass
242 (Fig.5). Microbial polymer use was negatively related with bacterial biomass and actinomycic
243 biomass. Microbial amines use was negatively related with G⁻ bacterial and fungal biomass.
244 Miscellaneous substrate use was positively related with fungal biomass and G⁺/G⁻ bacterial biomass
245 (Fig.5).

246 The abundance of G⁻ bacteria was positively associated first with the specific activities of BG,
247 whereas actinomycetes and G⁺ bacteria were positively associated with BG and LAP. Soil fungi were
248 negatively associated with BG (Fig.5).

249 3.4 Relationships between SOM decomposition rate, PLFAs, enzyme activity, and microbial
250 metabolic activities

251 The SOM decomposition rates were significantly and positively related to soil MBC concentrations
252 but significantly and negatively related to soil DOC concentrations (Fig. 6a and b). Except for amino
253 acid and amine substrates, the SOM decomposition rates were significantly and positively related to
254 microbial metabolic activities (AWCD) and carbohydrate substrate use (Fig. 6c and d) and negatively
255 related to carboxylic acid, polymer, and miscellaneous substrate use (Fig. 6e, g, and i).

256 The SOM decomposition rates were significantly and positively correlated with total PLFAs
257 ($r=0.456$, $P=0.005$), bacteria ($r=0.3836$, $P=0.021$), actinomycetes ($r=0.500$, $P=0.002$), and G⁻
258 bacteria PLFAs ($r=0.520$, $P=0.001$) (Fig. 7a, b, d, and f) but were negatively correlated with fungal
259 PLFAs ($r=-0.370$, $P=0.026$), F/B ($r=-0.513$, $P=0.001$), and the G⁺/G⁻ ($r=-0.496$, $P=0.002$) (Fig. 7c,
260 g, and h). Except for LAP activity, soil enzyme activities were significantly and positively correlated
261 with the SOM decomposition rates ($P<0.01$) (Fig. 7i, j, and l).

262 4 Discussion

263 4.1 Response of soil labile C and SOM decomposition rates to variations in forest type



264 Soil organic matter is one of the most important C pools in terrestrial ecosystems. The concentrations
265 of soil DOC in the temperate forests were lower than those in subtropical forests but soil MBC
266 concentrations were higher in temperate forests than in subtropical forests. This reflects the results
267 of previous regional and global studies (Tian et al., 2010; Xu et al., 2013), and shows that the
268 production/consumption ratio of soil DOC was lower, but that microbial C immobilization was higher,
269 in the high latitude forests (Fang et al., 2014). Soil DOC, as a labile SOM fraction with a rapid
270 turnover, is one of the primary energy sources for microorganisms. The higher temperatures and
271 precipitation in subtropical and tropical forests lead to higher turnover rates (Fang et al., 2014), so
272 soil DOC concentrations were highest in subtropical, and MBC concentrations were lowest, in
273 tropical forests. However, in temperate forests, more C is assimilated into microbial biomass, so that
274 less C is lost through chemical and physical processes (Liu et al., 2010). Also, because the
275 decomposition ability of different microbe groups varies, the differences in the soil microbial
276 communities in different forest ecosystems may also be responsible for the spatial variations in the
277 soil DOC and MBC concentrations along the NSTEC (Hagedorn et al., 2008).

278 Heterotrophic soil respiration is sustained by the decomposition of SOM. The SOM
279 decomposition rates along the NSTEC were greater in temperate forests than in subtropical forests,
280 which was consistent with the variations in the soil MBC and SOC concentrations. These results
281 indicate that, as found in other studies, large scale SOM decomposition rates are driven by the
282 amounts of substrate available (Yu et al., 2010). Changes in the availability of C in SOM may affect
283 the microbial resource strategies, which may in turn influence the SOM decomposition rate.

284 4.2 Latitudinal variation in microbial substrate use

285 The AWCD reflects the sole C source use ability of the soil microbial community (Garland and Mills,
286 1991). Of the six groups of C substrates, microbial communities in the temperate forests mainly used
287 carbohydrates, carboxylic acids, and amino acids, which suggests that microorganisms in temperate
288 forests probably use high-energy substrates that degrade easily (Kunito et al., 2009). The latitudinal
289 pattern of soil microbial C substrate use was mainly related to the soil silt contents and the CWMs of
290 LDMC, leaf C, and leaf N concentrations, indicating that the quality of nutrients from plant inputs
291 had a major influence on microbial carbon use efficiency. Plant species with high SLA, high N



292 concentrations in leaves, and low LDMC can result in bacterial-dominated soil microbial
293 communities in grasslands (Orwin et al., 2010). Looking beyond individual traits, related tree species
294 may cultivate microbial communities with similar preference for carbon sources through the
295 coevolution of plants and microbes (Liu et al., 2012; Buscot, 2015).

296 As hypothesized, the soil microbial community composition was explained by the CWMs of
297 plant traits at the regional scale. Carbon substrate use was negatively correlated with leaf N
298 concentrations (Table S2). Bacterially dominated soil microbial communities develop from leaf litter
299 comprised of N-rich leaves from fast growing species (De Vries et al., 2012), while leaves with low
300 N concentrations will promote fungal domination (Orwin et al., 2010; De Vries et al., 2012). In line
301 with this, fungal biomass decreased, and bacterial biomass increased, as the CWM leaf N content
302 increased, and is associated with fast-growing, N-exploitative plants (Xu et al., 2018). Leaf N
303 concentrations are considered as indicators of plant growth and resource uptake (Wright et al., 2004).
304 The results from this study show that, along the NSTEC, high leaf N restrained microbial C substrate
305 use and was a good indicator of the competition between plants for soil N (Pei et al., 2016). Soil
306 microbes and nearby plants may have been competing for N in the soil.

307 We also found that the C substrate use was negatively correlated with the leaf C concentrations
308 (Table S2). High latitude plants may have higher leaf C concentrations than plants at lower latitudes
309 so that they can balance the osmotic pressure of cells and resist freezing (Millard et al., 2007; Hoch
310 and Körner, 2012). The increased C was most likely in the form of an increase in non-structural C,
311 including starch, low molecular weight sugars, and storage lipids that are easy to break down. Plant
312 functional traits play an important role in shaping soil microbial communities (Pei et al., 2016), so
313 soil microorganisms from the temperate forests mainly metabolized high-energy substrates
314 (carbohydrates, carboxylic acids, and amino acids).

315 The LDMC is the ratio of the leaf dry weight to the fresh weight and has been used as a proxy
316 for the ratio of structural compounds to assimilatory tissue (mesophyll and epidermis, Van Arendonk
317 and Poorter, 1994). High values of LDMC indicate large amounts of vascular tissue, cellulose,
318 insoluble sugars, and leaf lignin that are difficult to decompose (Poorter and Bergkotte, 1992); C
319 substrates such as carbohydrates, carboxylic acid, and amino acid are, however, easy to decompose
320 (Myers et al., 2001). In line with this, the use of carbohydrate, carboxylic acid, and amino acid



321 substrates was negatively related to the CWMs of the LDMC (Table S2). Pei et al. (2016) reported
322 that the LDMC was an important driver of multivariate soil microbial community structure and G⁻
323 bacterial abundance.

324 Soil texture regulates soil biological processes and so affects the soil microbial community
325 structure (Sessitsch et al., 2001). In the present study, microbial C substrate use was significantly and
326 positively related to the soil silt content. Soil types and textures varied along the NSTEC. Soil texture
327 influences how microbes use organic matter, and has a strong influence on soil moisture, nutrient
328 availability, and retention (Veen and Kuikman, 1990). Fine-textured soils with a higher silt content
329 are known to be more favorable for bacterial growth than soils with a lower silt content because of
330 their greater water-holding capacity and nutrient availability, and because they are better protected
331 from bacterial grazers (Carson et al., 2010). We found that the microbial C substrate use was higher
332 in LS, CB, SN, and JL than in the other forests, reflecting their fine-grained soils and high silt contents,
333 which ranged from 60% to 80%.

334 4.3 Links between soil microbial community structure and function

335 The quality and changes in the amounts of SOM are influenced by the biomass, vegetation coverage,
336 root distribution, microbial specie (Raich and Schlesinger, 1992). The SOM decomposition rates
337 were higher in temperate forests than in tropical forests and may reflect the higher soil microbial
338 biomass (Wang et al., 2016). In line with this, SOM decomposition rates were positively related with
339 soil MBC concentrations and different groups of PLFAs. The inverse relationships between SOM
340 decomposition rates and DOC, and between SOM decomposition rates and the use of some individual
341 C substrates along the NSTEC, indicate a shift in the soil C turnover from open to closed with
342 increases in the soil labile C concentrations. Further, soil nutrients have a strong influence on the
343 spatial patterns of soil microbial communities. Thus, soil DOC and MBC do not influence SOM
344 decomposition rates directly, but indirectly by regulating microbial properties (Boberg et al., 2014;
345 Wei et al., 2014). Because different communities of microbes have different SOM use efficiencies
346 (Balser and Wixon, 2009; Lipson et al., 2009; Monson et al., 2006), changes in the microbial
347 community structure may influence the decomposition rates of organic matter (Lipson et al., 2009;
348 Keiblinger et al., 2010).



349 Shifts in microbial community composition may influence enzyme production (DeForest et al.,
350 2012; Waldrop et al., 2000; Brockett et al., 2012). Different microbial groups require different
351 amounts of nutrients to construct biomass, or have enzymes that differ in their affinity for nutrients.
352 For example, fungi tend to have higher C/N or C/P ratios while heterotrophic bacteria typically have
353 lower C/N or C/P ratios (Stern and Elser, 2002). We found that the relative abundances of the G⁺
354 bacteria and actinomycetes communities were associated with the specific activities of hydrolytic
355 enzymes involved in C and N acquisition (BG and LAP), whereas the relative abundance of the G⁻
356 bacteria was correlated with soil NAG activities involved in chitin degradation. Waldrop et al. (2000)
357 found that phosphatase activity was significantly correlated with the abundance of various bacterial
358 PLFAs. Soil BG was mainly responsible for cellulose degradation and was involved in breaking
359 down complex organic compounds (cellobiose) into small molecule substrates (glucose) in favor of
360 acquiring C through microbial community growth. Other studies have found that G⁺ bacteria were
361 positively correlated with the cellobiohydrolase that was responsible for degrading complex
362 compounds (Waldrop et al., 2000). Fungi are commonly considered as producers of oxidative
363 enzymes. Therefore, the influence of fungal biomass on variations in enzyme activities was minimal
364 (Kivlin and Treseder, 2014). The linkages between enzyme activity and community composition may
365 provide some insight into the microbial mechanisms that drive the decomposition of macromolecular
366 C compounds.

367 The soil microbial community structure and functions were significantly correlated along the
368 NSTEC. Soil carbohydrate and polymer substrate use were mainly related to soil G⁺ bacterial and
369 actinomycic biomass, but amines and miscellaneous substrates were mainly related to soil G⁻
370 bacterial, fungal biomass, and the F/B ratio. Soil bacteria mainly decomposed simple carbohydrates,
371 organic acids, and amino acids, whereas soil fungi mainly decomposed recalcitrant compounds
372 (Myers et al., 2001; Treonis et al., 2004). Shifts in the microbial community composition may
373 influence enzyme production if microbial groups need nutrients at lower concentrations to construct
374 biomass, or have enzymes that differ in their affinity for nutrients. In agreement with our study,
375 numerous other researchers have reported significant correlations between PLFA profiles and enzyme
376 activities (DeForest et al., 2012; Brockett et al., 2012; Riah-Anglet et al., 2015). Soil BG and AP
377 activities were positively related with bacterial and actinomycic biomass and negatively related with



378 fungal biomass. Soil NAG activities were weakly and positively related with fungal biomass in the
379 present study, and may have been mainly produced by fungal populations (Valášková et al., 2007).
380 These results suggest that that overall ecosystem functioning may suffer if soil microbial groups are
381 lost, which confirms the functional dissimilarity hypothesis. However, to gain an improved
382 understanding of the mechanisms that drive these relationships, we need to carry out further studies
383 with different experimental techniques.

384 **5 Conclusions**

385 In this study we examined the patterns in labile C concentrations, SOM decomposition rates,
386 microbial substrate use, and functional diversity and identified a combination of abiotic and biotic
387 factors that influenced soil microbial functional diversity at the regional scale. The MBC
388 concentration and SOM decomposition rates were significantly lower, and the soil DOC
389 concentrations and microbial metabolic activities were higher, in the subtropical and tropical forests
390 than in the temperate forests. For the first time, we showed that, along with the soil silt content, CWM
391 plant traits explained variations in soil microbial C substrate use at the regional scale. Soil microbial
392 community structure and function were strongly related, which suggest that the loss of soil microbial
393 groups may have consequences for overall ecosystem functioning, which confirms the functional
394 dissimilarity hypothesis.

395 *Data accessibility.* Requests for data and materials should be addressed to N.H. (henp@igsnr.ac.cn) and G.Y.
396 (yugr@igsnr.ac.cn).

397

398 *Author contributions.* Z.W.X., G.R.Y. and X.Y.Z. planned and designed the research. Z.W.X., N.P.H., R.L.W., and
399 N.Z. conducted fieldwork. Z.W.X., G.R.Y., X.Y.Z., and Q.F.W wrote the manuscript. All authors contributed
400 critically to the drafts and gave final approval for publication.

401 *Competing interests.* The authors declare that they have no conflict of interest.

402

403 **Acknowledgements**

404 This research was jointly supported by the National Natural Science Foundation of China (41601084, 41571251), the
405 National Key R&D Program of China (2016YFA0602301), and Science and Technology Research Project of Jilin
406 Province (JJKH20190283KJ).

407 **References**

- 408 van Arendonk, J.J.C.M., Poorter, H. The chemical composition and anatomical structure of leaves of grass species
409 differing in relative growth rate, *Plant Cell Environ.*, 17, 963-970, 1994.
- 410 Bååth, E., Anderson, T.H. Comparison of soil fungal/bacterial ratios in a pH gradient using physiological and PLFA-
411 based techniques, *Soil Biol. Biochem.*, 35, 955-963, 2003.
- 412 Balsler, T.C., Firestone, M.K. Linking microbial community composition and soil processes in a California annual
413 grassland and mixed-conifer forest, *Biogeochemistry*, 73, 395-415, 2005.
- 414 Balsler, T.C., Wixon, D.L. Investigating biological control over soil carbon temperature sensitivity, *Global Change*
415 *Biol.*, 15, 2935-2949, 2009.
- 416 Banerjee, S., Kirkby, C.A., Schmutter, D., Bissett, A., Kirkegaard, J.A., Richardson, A.E. Network analysis reveals
417 functional redundancy and keystone taxa amongst bacterial and fungal communities during organic matter
418 decomposition in an arable soil, *Soil Biol. Biochem.*, 97, 188-198, 2016.
- 419 Boberg, J.B., Finlay, R.D., Stenlid, J., Ekblad, A., Lindahl, B.D. Nitrogen and carbon reallocation in fungal mycelia
420 during decomposition of boreal forest litter, *PLOS ONE*, 9, e92897, 2014.
- 421 Brockert, B.F.T., Prescott, C.E., Grayston, S.J. Soil moisture is the major factor influencing microbial community
422 structure and enzyme activities across seven biogeoclimatic zones in western Canada, *Soil Biol. Biochem.*, 44,
423 9-20, 2012.
- 424 Buscot, F. Implication of evolution and diversity in arbuscular and ectomycorrhizal symbioses, *J. Plant Physiol.*, 172,
425 55-61, 2015.
- 426 Cao, H., Chen, R., Wang, L., Jiang, L., Yang, F., Zheng, S., Wang, G., Lin, X. Soil pH, total phosphorus, climate and
427 distance are the major factors influencing microbial activity at a regional spatial scale, *Sci. Rep-UK*, 6, 25815,
428 2016.
- 429 Carson, J.K., Gonzalez-Quinones, V., Murphy, D.V., Hinz, C., Shaw, J.A., Gleeson, D.B. Low pore connectivity
430 increases bacterial diversity in soil, *Appl. Environ. Microb.*, 76, 3936-3942, 2010.
- 431 Coulis, M., Fromin, N., David, J.-F., Gavinet, J., Clet, A., Devidal, S., Roy, J., Hättenschwiler, S. Functional
432 dissimilarity across trophic levels as a driver of soil processes in a Mediterranean decomposer system exposed
433 to two moisture levels, *Oikos*, 124, 1304-1316, 2015.
- 434 De Vries, F.T., Manning, P., Tallwin, J.R., Mortimer, S.R., Pilgrim, E.S., Harrison, K.A., Hobbs, P.J., Quirk, H.,
435 Shipley, B., Cornelissen, J.H., Kattge, J., Bardgett, R.D. Abiotic drivers and plant traits explain landscape-scale
436 patterns in soil microbial communities, *Ecol. Lett.*, 15, 1230-1239, 2012.
- 437 Deforest, J., Smemo, K., Burke, D., Elliott, H., Becker, J. Soil microbial responses to elevated phosphorus and pH in
438 acidic temperate deciduous forests, *Biogeochemistry*, 109, 189-202, 2012.
- 439 Fang, H., Cheng, S., Wang, Y., Yu, G., Xu, M., Dang, X., Li, L., Wang, L. Changes in soil heterotrophic respiration,
440 carbon availability, and microbial function in seven forests along a climate gradient, *Ecol. Res.*, 29, 1077-1086,
441 2014.
- 442 Fiere, N., Bradford, M. A. Toward an ecological classification of soil bacteria, *Ecology*, 88, 1354-1364, 2007.
- 443 Garland, J.L., Mills, A.L. Classification and characterization of heterotrophic microbial communities on the basis of
444 patterns of community-level sole carbon source utilization, *Appl. Environ. Microb.*, 57, 2351-2359, 1991.
- 445 Garnier, E., Cortez, J., Billes, G., Navas, M.L., Roumet, C., Debussche, M., Gérard, L., Alain, B., David, A., Astrid,
446 B., Cathy, N., Jean-Patrick, T. Plant functional markers capture ecosystem properties during secondary
447 succession, *Ecology*, 85, 2630-2637, 2004.
- 448 Gartzia-Bengoetxea, N., Kandeler, E., Martínez de Arano, I., Arias-González, A. Soil microbial functional activity is
449 governed by a combination of tree species composition and soil properties in temperate forests, *Appl. Soil Ecol.*,
450 100, 57-64, 2016.
- 451 Gomez, E., Ferreras, L., Toresani, S. Soil bacterial functional diversity as influenced by organic amendment
452 application, *Bioresour. Technol.*, 97, 1484-1489, 2006.
- 453 Guckert, J.B., Carr, G.J., Johnson, T.D., et al. Community analysis by Biolog: curve integration for statistical analysis
454 of activated sludge microbial habitats, *J. Microbiol. Meth.*, 27, 183-197, 1996.
- 455 Huang, Z., Clinton, P.W., Baisden, W.T., Davis, M.R. Long-term nitrogen additions increased surface soil carbon
456 concentration in a forest plantation despite elevated decomposition, *Soil Biol. Biochem.*, 43, 302-307, 2011.
- 457 Jagadamma, S., Mayes, M.A., Steinweg, J.M., Schaeffer, S.M. Substrate quality alters the microbial mineralization
458 of added substrate and soil organic carbon, *Biogeosciences*, 11, 4665-4678, 2014.
- 459 Jones, D., Willett, V. Experimental evaluation of methods to quantify dissolved organic nitrogen (DON) and dissolved
460 organic carbon (DOC) in soil, *Soil Biol. Biochem.*, 38, 991-999, 2006.
- 461 Keiblinger, K.M., Hall, E.K., Wanek, W., Szukics, U., Hammerle, I., Ellersdorfer, G., Bock, S., Strauss, J., Sterflinger,
462 K., Richter, A., Zechmeister-Boltenstern, S. The effect of resource quantity and resource stoichiometry on
463 microbial carbon-use-efficiency. *FEMS Microb. Ecol.* 73, 430-440, 2010.
- 464 Kivlin, S.N. & Treseder, K.K. Soil extracellular enzyme activities correspond with abiotic factors more than fungal
465 community composition, *Biogeochemistry*, 117, 23-37, 2014.
- 466 Klimek, B., Chodak, M., Jaźwa, M., Niklińska, M. Functional diversity of soil microbial communities in boreal and
467 temperate Scots pine forests, *Eur. J. Forest Res.*, 135, 731-742, 2016.
- 468 Kunito, T., Akagi, Y., Park, H.-D., Toda, H. Influences of nitrogen and phosphorus addition on polyphenol oxidase
469 activity in a forested Andisol, *Eur. J. Forest Res.*, 128, 361-366, 2009.
- 470 Lipson, D.A., Monson, R.K., Schmidt, S.K., Weintraub, M.N. The trade-off between growth rate and yield in
471 microbial communities and the consequences for under-snow soil respiration in a high elevation coniferous forest,
472 *Biogeochemistry*, 95, 23-35, 2009.



- 473 Liu, X., Liang, M., Etienne, R.S., Wang, Y., Staehelin, C., Yu, S. Experimental evidence for a phylogenetic Janzen-
474 Connell effect in a subtropical forest, *Ecol. Lett.*, 15, 111-118, 2012.
- 475 Liu, Z., Liu, G., Fu, B., Wu, Y., Hu, H., Fu, S. Changes in the soil microbial community with a pine plantation
476 restoration in a dry valley of the upper reaches of the Minjiang River, southwest China, *Ann N Y Acad. Sci.*,
477 1195 Suppl 1, E82-95, 2010.
- 478 Millard, P., Sommerkorn, M., Grelet, G.A. Environmental change and carbon limitation in trees: a biochemical,
479 ecophysiological and ecosystem appraisal, *New Phytol.*, 175, 11-28, 2007.
- 480 Monson, R.K., Lipson, D.L., Burns, S.P., Turnipseed, A.A., Delany, A.C., Williams, M.W., Schmidt, S.K. Winter
481 forest soil respiration controlled by climate and microbial community composition, *Nature*, 439, 711-714, 2006.
- 482 Myers, R.T., Zak, D.R., White, D.C., Peacock, A. Landscape-level patterns of microbial community composition and
483 substrate use in upland forest ecosystems, *Soil Sci. Soc. Am. J.*, 65, 359-367, 2001.
- 484 Orwin, K.H., Buckland, S.M., Johnson, D., Turner, B.L., Smart, S., Oakley, S., Bardgett, R.D. Linkages of plant traits
485 to soil properties and the functioning of temperate grassland, *J. Ecol.*, 98, 1074-1083, 2010.
- 486 Pei, Z., Eichenberg, D., Bruelheide, H., Kröber, W., Kühn, P., Li, Y., von Oheimb, G., Purschke, O., Scholten, T.,
487 Buscot, F., Gutknecht, J.L.M. Soil and tree species traits both shape soil microbial communities during early
488 growth of Chinese subtropical forests, *Soil Biol. Biochem.*, 96, 180-190, 2016.
- 489 Peng, X., Wang, W. Stoichiometry of soil extracellular enzyme activity along a climatic transect in temperate
490 grasslands of northern China, *Soil Biol. Biochem.*, 98, 74-84, 2016.
- 491 Philippot, L., Spor, A., Henault, C., Bru, D., Bizouard, F., Jones, C.M., Sarr, A., Maron, P.A. Loss in microbial
492 diversity affects nitrogen cycling in soil, *ISME J.*, 7, 1609-1619, 2013.
- 493 Poorter, H., Bergkotte, M. Chemical composition of 24 wild species differing in relative growth rate, *Plant Cell*
494 *Environ.*, 15, 221-229, 1992.
- 495 Preston-Mafham, J., Boddy, L., Randerson, P.F. Analysis of microbial community functional diversity using sole-
496 carbon-source utilisation profiles—a critique, *FEMS Microb. Ecol.*, 42, 1-14, 2002.
- 497 Raich, J.W., Clark, D.A., Schwendenmann, L., Wood, T.E. Aboveground tree growth varies with belowground carbon
498 allocation in a tropical rainforest environment, *PLoS one*, 9, e100275, 2014.
- 499 Riah-Anglet, W., Trinsoutrot-Gattin, I., Martin-Laurent, F., Laroche-Ajzenberg, E., Norini, M.-P., Latour, X., Laval,
500 K. Soil microbial community structure and function relationships: A heat stress experiment, *Appl. Soil Ecol.*, 86,
501 121-130, 2015.
- 502 Russell, A.E., Raich, J.W., Valverde-Barrantes, O.J., Fisher, R.F. Tree species effects on soil properties in experimental
503 plantations in tropical moist forest, *Soil Sci. Soc. Am. J.*, 71, 1389, 2007.
- 504 Saiya-Cork, K.R., Sinsabaugh, R.L., Zak, D.R. The effects of long term nitrogen deposition on extracellular enzyme
505 activity in an *Acer saccharum* forest soil, *Soil Biol. Biochem.*, 34, 1309-1315, 2002.
- 506 Sessitsch, A., Weilharter, A., Gerzabek, M.H., Kirchmann, H., Kandeler, E. Microbial Population Structures in Soil
507 Particle Size Fractions of a Long-Term Fertilizer Field Experiment, *Appl. Environ. Microb.*, 67, 4215-4224, 2001.
- 508 Sherman, C., Steinberger, Y. Microbial functional diversity associated with plant litter decomposition along a climatic
509 gradient, *Microb. Ecol.*, 64, 399-415, 2012.
- 510 Six, J., Elliott, E. T., Paustian, K. Soil structure and soil organic matter: II. A normalized stability index and the effect
511 of mineralogy, *Soil Sci. Soc. Am. J.*, 64, 1042-1049, 2000.
- 512 Sterner, R. W. & Elser, J. J. Ecological stoichiometry: the biology of elements from molecules to the biosphere.
513 Princeton, New Jersey, USA: Princeton University Press, 2002.
- 514 Strickland, M.S., Lauber, C., Fierer, N., Bradford, M.A. Testing the functional significance of microbial community
515 composition, *Ecology*, 90, 441-451, 2009.
- 516 Soil Survey Staff Keys to Soil Taxonomy, 11th ed. USDA-Natural Resources Conservation Service, Washington, DC,
517 2010.
- 518 Swallow, M.J.B., Quideau, S.A. A method for determining community level physiological profiles of organic soil
519 horizons, *Soil Sci. Soc. Am. J.*, 79, 536-542, 2015.
- 520 Talbot, J.M., Bruns, T.D., Taylor, J.W., Smith, D.P., Branco, S., Glassman, S.I., Erlandson, S., Vilgalys, R., Liao, H.L.,
521 Smith, M.E., Peay, K.G. Endemism and functional convergence across the North American soil microbiome, *P.*
522 *Natl. Acad. Sci. USA.*, 111, 6341-6346, 2014.
- 523 Ter Braak, C.J.F., Smilauer, P. CANOCO Reference manual and CanoDraw for Windows User's guide: Software for
524 Canonical Community Ordination (Version 4.5). Microcomputer, 2002.
- 525 Tian, H., Chen, G., Zhang, C., Melillo, J.M., Hall, C.A.S., Pattern and variation of C:N:P ratios in China's soils: a
526 synthesis of observational data, *Biogeochemistry*, 98, 139-151 2010.
- 527 Tian, J., McCormack, L., Wang, J., Guo, D., Wang, Q., Zhang, X., Yu, G., Blagodatskaya, E., Kuzyakov, Y. Linkages
528 between the soil organic matter fractions and the microbial metabolic functional diversity within a broad-leaved
529 Korean pine forest, *Eur. J. Soil Biol.*, 66, 57-64, 2015.
- 530 Treonis, A.M., Ostle, N.J., Stott, A.W., Primrose, R., Grayston, S.J., Ineson, P. Identification of groups of
531 metabolically-active rhizosphere microorganisms by stable isotope probing of PLFAs, *Soil Biol. Biochem.*, 36,
532 533-537, 2004.
- 533 Tu, Q., Deng, Y., Yan, Q., Shen, L., Lin, L., He, Z., Wu, L., Van Nostrand, J.D., Buzzard, V., Michaletz, S.T., Enquist,
534 B.J., Weiser, M.D., Kaspari, M., Waide, R.B., Brown, J.H., Zhou, J. Biogeographic patterns of soil diazotrophic
535 communities across six forests in the North America, *Mol. Ecol.*, 25, 2937-2948, 2016.
- 536 Valášková, V., Šnajdr, J., Bittner, B., Cajthaml, T., Merhautová, V., Hofrichter, M., Baldrian, P. Production of
537 lignocellulose-degrading enzymes and degradation of leaf litter by saprotrophic basidiomycetes isolated from a
538 *Quercus petraea* forest, *Soil Biol. Biochem.*, 39, 2651-2660, 2007.
- 539 Vance, E.D., Brookes, P.C., Jenkinson, D.S. An extraction method for measuring soil microbial biomass C, *Soil Biol.*



- 540 Biochem., 19, 703-707, 1987.
- 541 Veen, J.A.V., Kuikman, P.J. Soil structural aspects of decomposition of organic-matter by micro-organism,
542 Biogeochemistry, 11, 213-233, 1990.
- 543 Waldrop, M.P., Firestone, M.K. Response of microbial community composition and function to soil climate change,
544 Microb. Ecol., 52, 716-724, 2006.
- 545 Waldrop, M. P., Balsler, T. C. & Firestone, M. K. Linking microbial community composition to function in a tropical
546 soil, Soil Biol. Biochem., 32, 1837-1846, 2000.
- 547 Wang, Q., He, N., Yu, G., Gao, Y., Wen, X., Wang, R., Koerner, S.E., Yu, Q. Soil microbial respiration rate and
548 temperature sensitivity along a north-south forest transect in eastern China: Patterns and influencing factors, J.
549 Geophys. Res.(Biogeosciences), 121, 399-410, 2016.
- 550 Waring, B.G., Alvarez-Cansino, L., Barry, K.E., Becklund, K.K., Dale, S., Gei, M.G., Keller, A.B., Lopez, O.R.,
551 Markesteijn, L., Mangan, S., Riggs, C.E., Rodriguez-Ronderos, M.E., Segnitz, R.M., Schnitzer, S.A., Powers,
552 J.S. Pervasive and strong effects of plants on soil chemistry: a meta-analysis of individual plant 'Zinke' effects,
553 Proc. R. Soc. B. The Royal Society, 282, 20151001, 2015.
- 554 Wei, H., Guenet, B., Vicca, S., Nunan, N., AbdElgawad, H., Pouteau, V., Shen, W., Janssens, I.A. Thermal acclimation
555 of organic matter decomposition in an artificial forest soil is related to shifts in microbial community structure,
556 Soil Biol. Biochem., 71, 1-12, 2014.
- 557 Wright, I.J., Reich, P.B., Westoby, M., Ackerly, D.D., Baruch, Z., Bongers, F., et al. The worldwide leaf economics
558 spectrum, Nature, 428, 821-827, 2004.
- 559 Xu, X., Thornton, P.E., Post, W.M. A global analysis of soil microbial biomass carbon, nitrogen and phosphorus in
560 terrestrial ecosystems, Global Ecol. Biogeogr., 22, 737-749, 2013.
- 561 Xu, Z., Yu, G., Zhang, X., Ge, J., He, N., Wang, Q., Wang, D. The variations in soil microbial communities, enzyme
562 activities and their relationships with soil organic matter decomposition along the northern slope of Changbai
563 Mountain, Appl. Soil Ecol., 86, 19-29, 2015.
- 564 Xu, Z., Yu, G., Zhang, X., He, N., Wang, Q., Wang, S., Wang, R., Zhao, N., Jia, Y., Wang, C. Soil enzyme activity and
565 stoichiometry in forest ecosystems along the North-South Transect in eastern China (NSTEC), Soil Biol.
566 Biochem., 104, 152-163, 2017.
- 567 Xu, Z., Yu, G., Zhang, X., He, N., Wang, Q., Wang, S., Xu, X., Wang, R., Zhao, N. Biogeographical patterns of soil
568 microbial community as influenced by soil characteristics and climate across Chinese forest biomes, Appl. Soil
569 Ecol., 124, 298-305, 2018.
- 570 Yu, G., Zheng, Z., Wang, Q., Fu, Y., Zhuang, J., Sun, X., Wang, Y. Spatiotemporal pattern of soil respiration of
571 terrestrial ecosystems in China: the development of a geostatistical model and its simulation, Environ. Sci.
572 Technol., 44, 6074-6080, 2010.
- 573 Zak, J.C., Willig, M.R., Moorhead, D.L., Wildman, H.G. Functional diversity of microbial communities: a quantitative
574 approach, Soil Biol. Biochem., 26, 1101-1108, 1994.
- 575 Zhang, X. S., Yang, D. A. Allocation and study on global change transects in China, Quaternary Sci., 1, 43-52, 1995.
576 (In Chinese)
- 577 Zhou, J., Deng, Y., Shen, L., Wen, C., Yan, Q., Ning, D., Qin, Y., Xue, K., Wu, L., He, Z., Voordeckers, J.W., Nostrand,
578 J.D., Buzzard, V., Michaletz, S.T., Enquist, B.J., Weiser, M.D., Kaspary, M., Waide, R., Yang, Y., Brown, J.H.
579 Temperature mediates continental-scale diversity of microbes in forest soils, Nat. Commun., 7, 12083, 2016.



580 **Figures legends**

581 **Fig. 1.** Distribution of typical forest ecosystems along the North-South Transect of eastern China (NSTEC). The
582 abbreviations for the sampling sites from north to south are as follows: HZ, Huzhong; LS, Liangshui; CB, Changbai;
583 DL, Dongling; TY, Taiyue; SN, Shennong; JL, Jiulian; DH, Dinghu; JF, Jiangfeng. These abbreviations are used for
584 the nine forests throughout.

585 **Fig.2.** Variations in soil microbial substrate use during a 240-h incubation for the nine forests. Different colors
586 represent different forest types: Yellow, coniferous forest; Dark yellow, coniferous broad-leaved mixed forest; Purple,
587 deciduous broad-leaved forest; Olive, subtropical evergreen broad-leaved forest; Orange, Tropical monsoon forest.
588 Different lowercase letters indicate significant differences among forests in the same climate zone. The abbreviations
589 of the sampling sites are given in Table 1.

590 **Fig. 3.** Characteristics of microbial use of (a) carbohydrates, (b) carboxylic acids, (c) amino acids, (d) polymers, (e)
591 amines, and (f) miscellaneous along the NSTEC. The representatives of different colors were showed in Figure 2.

592 **Fig.4.** Redundancy analysis (RDA) ordination biplot of soil microbial carbon sources use efficiency and
593 environmental properties. The representatives of different colors were showed in Figure 2. The dotted lines and solid
594 lines represent the environmental variables and lipid signatures and carbon sources. The abbreviations of the variables
595 in this figure are as follows: MAP, mean annual precipitation. The vegetation data: LDMC, leaf dry matter weight;
596 Leaf C, leaf carbon content; Leaf N, leaf nitrogen content; SLA, specific leaf area. Soil properties included SMC, soil
597 moisture content; Silt, soil silt content; TN, soil total nitrogen; SOC, soil organic carbon. The abbreviations of the
598 sampling sites were given in Table 1

599 **Fig.5.** The heatmap of the Pearson's correlation coefficients between the use of individual substrates and microbial
600 PLFAs and soil enzyme activities. Note: The abbreviations of the variables: Actino-, actinomycetes; F/B,
601 fungi/bacteria; G⁺, gram positive bacteria; G⁻, gram negative bacteria; G⁺/G⁻, Gram-positive bacteria/ Gram-negative
602 bacteria. BG, β -1, 4-glucosidase; NAG, β -1,4-N-acetylglucosaminidase; LAP, leucine aminopeptidase; AP, acid
603 phosphatase. ** $P < 0.01$, * $P < 0.05$.

604 **Fig. 6.** Relationships between soil carbon mineralization rates ($\mu\text{g C g}^{-1}\text{d}^{-1}$) and microbial biomass C (MBC), soil
605 dissolved organic C (DOC), average well color development (AWCD), and individual substrate use.

606 **Fig. 7.** Relationships between soil carbon mineralization rates ($\mu\text{g C g}^{-1}\text{d}^{-1}$) and different groups of soil microbial
607 PLFAs (a-h) and enzyme activities (i-l).

608
609

610 **Supporting information**

611 **Table S1** Average values of forest soil enzyme activities and different PLFA groups along the NSTEC.

612 **Table S2** Plant diversity and community weighted means of plant functional traits

613 **Table S3** Soil organic matter (SOM) decomposition rates during the 28 days of incubation time (Mean \pm SE) ($\mu\text{g C}$
614 $\text{g}^{-1}\text{d}^{-1}$)

615 **Fig. S1** Variations in the average well color development (AWCD) values during a 240-h incubation for the nine
616 forests. The abbreviations of the sampling sites are the same as those in Table 1.



617 **Table 1.** The main characteristics of the sampling sites along the North South Transect of East China
 618

Sampling Sites	Longitude (E)	Latitude (N)	Elevation (m)	MAT ^b (°C)	MAP ^b (mm)	Vegetation types	Soil type
HZ ^a	123°01'12"	51°46'48"	850	-3.7	473	Cold temperate coniferous forest	Spodosols
LS	128°53'51"	47°11'06"	401	0.01	648	Temperate conifer broad-leaved mixed forest	Albi-Boric Argosols
CB	128°05'27"	42°24'16"	758	2.8	691	Temperate conifer broad-leaved mixed forest	Albi-Boric Argosols
DL	115°25'24"	39°57'27"	972	6.6	539	Warm temperate deciduous broad-leaved forest	Alfisols
TY	112°04'39"	36°41'43"	1668	6.0	644	Warm temperate deciduous broad-leaved forest	Alfisols
SN	110°29'43"	31°19'15"	1510	8.5	1447	Subtropical deciduous evergreen mixed forest	Inceptisols
JL	114°26'28"	24°35'05"	562	18.2	1770	Subtropical evergreen broad-leaved forest	Ultisols
DH	112°32'14"	23°10'25"	240	21.8	1927	Subtropical monsoon evergreen broad-leaved forest	Ultisols
JF	108°51'26"	18°44'18"	809	23.2	2266	Tropical monsoon forest	Ultisols

619 a: HZ, Huzhong; LS, Liangshui; CB, Changbai; DL, Dongling; TY, Taiyue; SN, Shennong; JL, Jiulian; DH, Dinghui; JF, Jiangfeng.

620 b: MAT, mean annual temperature; MAP, mean annual precipitation.



621 **Table 2.** Soil properties of different sampling sites

Sampling site	pH	ST (°C)	SMC (%)	Silt (%)	SOC (g kg ⁻¹)	MBC (mg kg ⁻¹)	DOC (mg kg ⁻¹)	TN (g kg ⁻¹)	TP (g kg ⁻¹)
HZ	6.79±0.02a	10.3±0.15g	45.3±0.90c	56±1.2c	42.29±0.47b	350±6.0a	240±7.6e	2.90±0.16d	0.87±0.02b
LS	6.17±0.02b	15.9±0.02f	46.9±0.76c	64±0.3b	62.08±7.20a	316±0.7a	204±4.9f	4.59±0.29b	0.59±0.02c
CB	6.37±0.04b	16.0±0.06f	102.8±0.25a	76±0.6a	72.38±2.00a	178±8.8b	314±8.6c	6.05±0.17a	1.67±0.08a
DL	6.87±0.02a	17.8±0.14e	32.4±0.30e	6±2.4e	38.83±0.41c	43±0.8e	284±2.6d	3.17±0.04d	0.56±0.01c
TY	6.85±0.05a	16.0±0.12f	36.0±0.23d	49±1.4d	41.34±2.75c	115±4.0c	226±13.8f	2.43±0.15e	0.52±0.01c
SN	6.93±0.01a	18.4±0.12d	50.5±0.63b	74±0.3a	36.13±1.26c	72±13.1e	311±13.2c	3.76±0.05c	0.81±0.01b
JL	5.57±0.19b	25.3±0.01a	39.0±0.89d	68±0.3b	31.55±1.82c	89±19.7d	387±1.9b	2.28±0.09e	0.36±0.01d
DH	5.43±0.03c	24.4±0.04b	37.8±0.38d	50±1.8d	28.47±0.54d	38±0.1e	334±7.7c	1.77±0.02f	0.20±0.01e
JF	6.32±0.01c	22.5±0.07c	38.6±0.12d	49±0.2d	29.38±0.94d	140±1.3c	458±6.6a	1.99±0.02e	0.15±0.01e

622 Note: ST=temperature of 0–10 cm soil; SMC=soil moisture content; Silt=soil silt content; SOC=soil organic carbon; MBC=microbial biomass carbon; DOC=dissolved organic carbon; TN=soil
 623 total nitrogen; TP=soil total phosphorus. Values were presented as means ± SE (n=4). The abbreviations of the sampling sites were given in the Table 1.

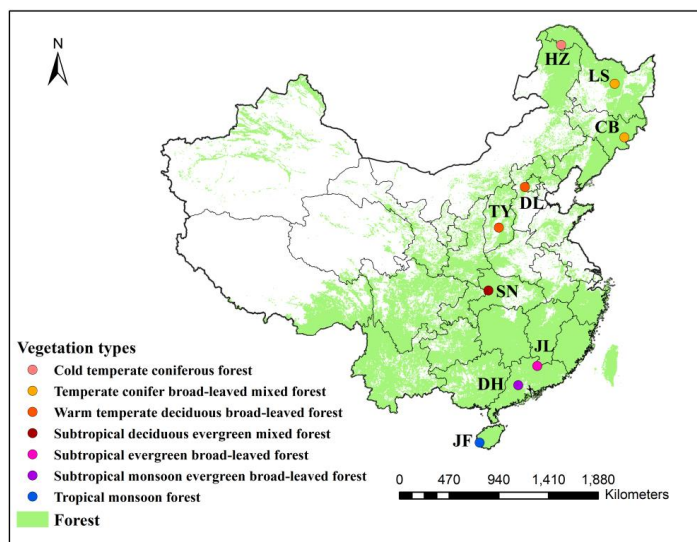
624 **Table 3.** Functional diversity of soil microbial communities in forest ecosystems along the NSTEC

Sampling sites	Richness (<i>R</i>)	Shannon <i>H'</i>	Shannon <i>E</i>	Simpson <i>D</i>
HZ	14.08±0.34d	2.65±0.03d	1.01±0.007b	0.91±0.002c
LS	25.29±0.14b	3.12±0.02b	0.98±0.003c	0.95±0.001a
CB	27.00±0.27a	3.22±0.01a	0.98±0.001c	0.95±0.001a
DL	11.54±0.47e	2.52±0.03e	1.04±0.010a	0.87±0.005d
TY	22.33±0.87c	3.02±0.02c	0.98±0.002c	0.94±0.001a
SN	28.10±0.34a	3.24±0.01a	0.97±0.001c	0.95±0.001a
JL	23.54±0.07c	3.04±0.01c	0.96±0.001c	0.93±0.003b
DH	25.65±0.71b	3.11±0.01b	0.97±0.001c	0.93±0.002b
JF	27.63±0.68a	3.19±0.02a	0.96±0.001c	0.95±0.002a

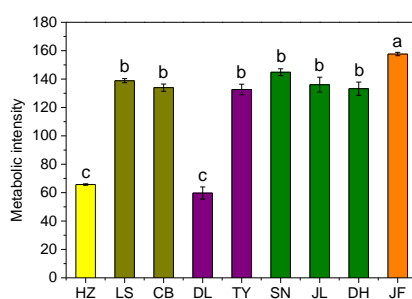
625 Indices were calculated based on the optical density values after incubation for 96 h. Data are expressed as
 626 means±standard errors. Different lowercase letters indicate significant differences among forests. The abbreviations
 627 of the sampling sites are the same as those used in Table 1.



628 **Figures**

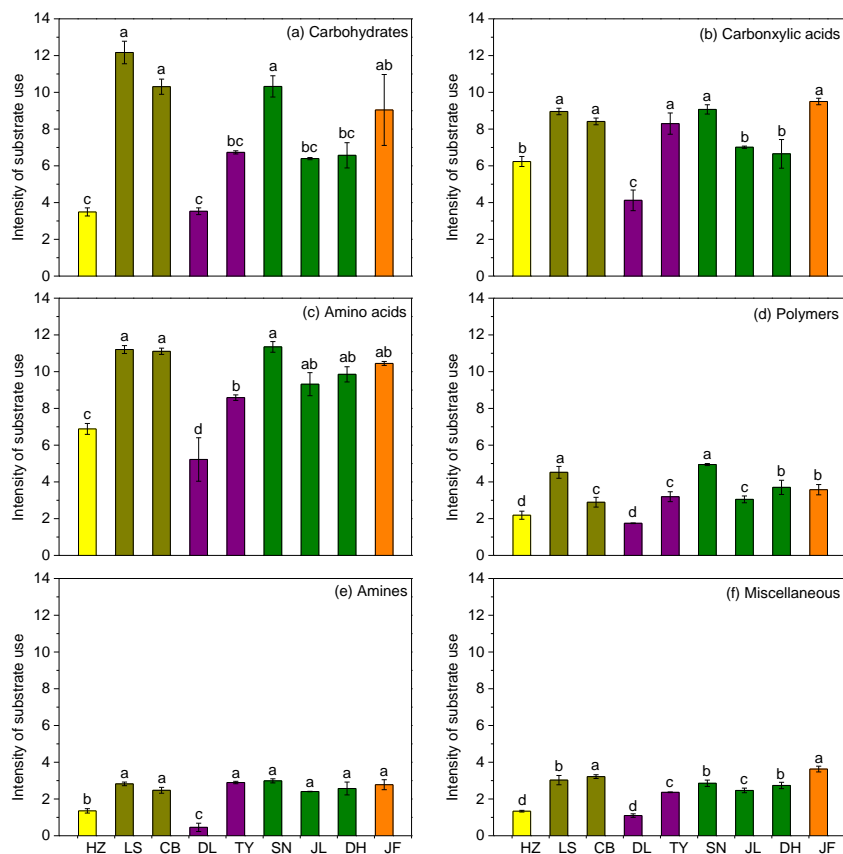


629 **Figure 1.** Distribution of typical forest ecosystems along the North-South Transect of eastern China (NSTEC).
630 The abbreviations of sampling sites from north to south are as follows: HZ, Huzhong; LS, Liangshui; CB,
631 Changbai; DL, Dongling; TY, Taiyue; SN, Shennong; JL, Jiulian; DH, Dinghu; JF, Jiangfeng.
632



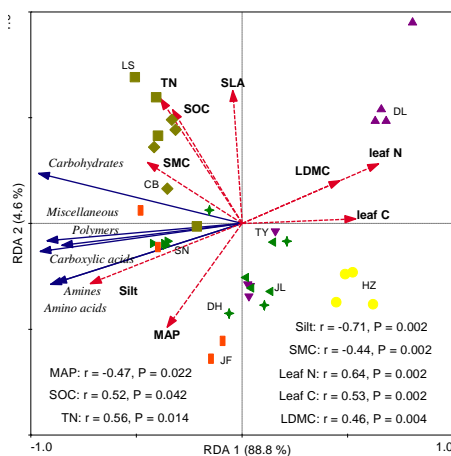
633
634
635
636
637
638

Figure 2. Variations in soil microbial substrate use during a 240-h incubation for the nine forests. Different colors represent different forest types: Yellow, coniferous forest; Dark yellow, coniferous broad-leaved mixed forest; Purple, deciduous broad-leaved forest; Olive, subtropical evergreen broad-leaved forest; Orange, Tropical monsoon forest. Different lowercase letters indicate significant differences among forests in the same climate zone. The abbreviations of the sampling sites are given in Table 1.



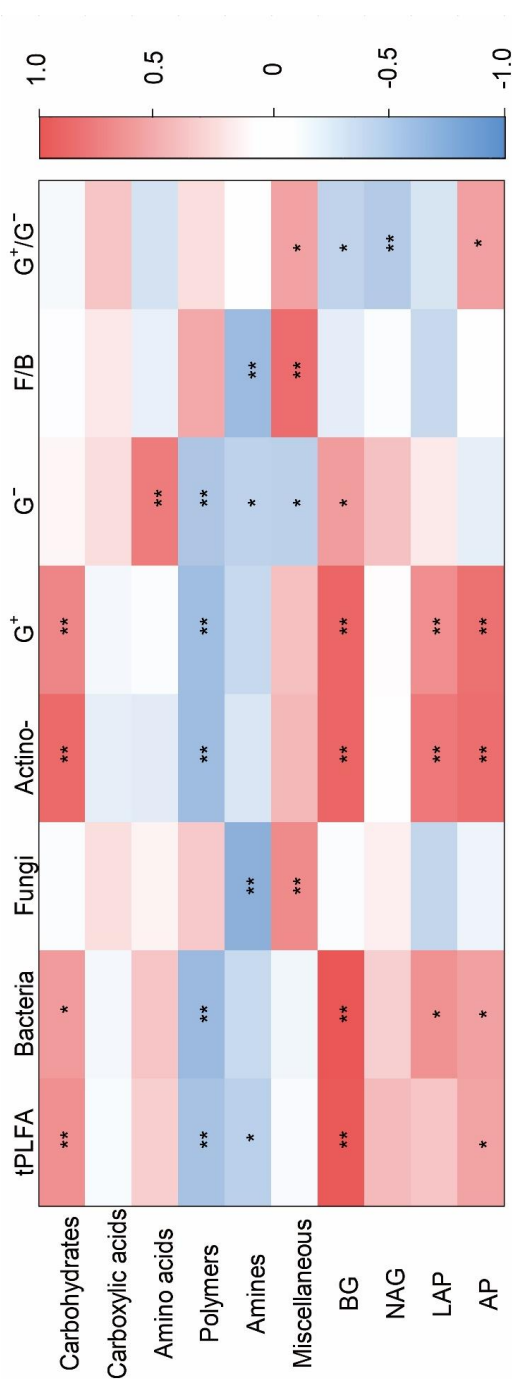
639
640
641
642

Figure 3. Characteristics of microbial use of (a) carbohydrates, (b) carboxylic acids, (c) amino acids, (d) polymers, (e) amines, and (f) miscellaneous along the NSTEC. The representatives of different colors were showed in Figure 2.

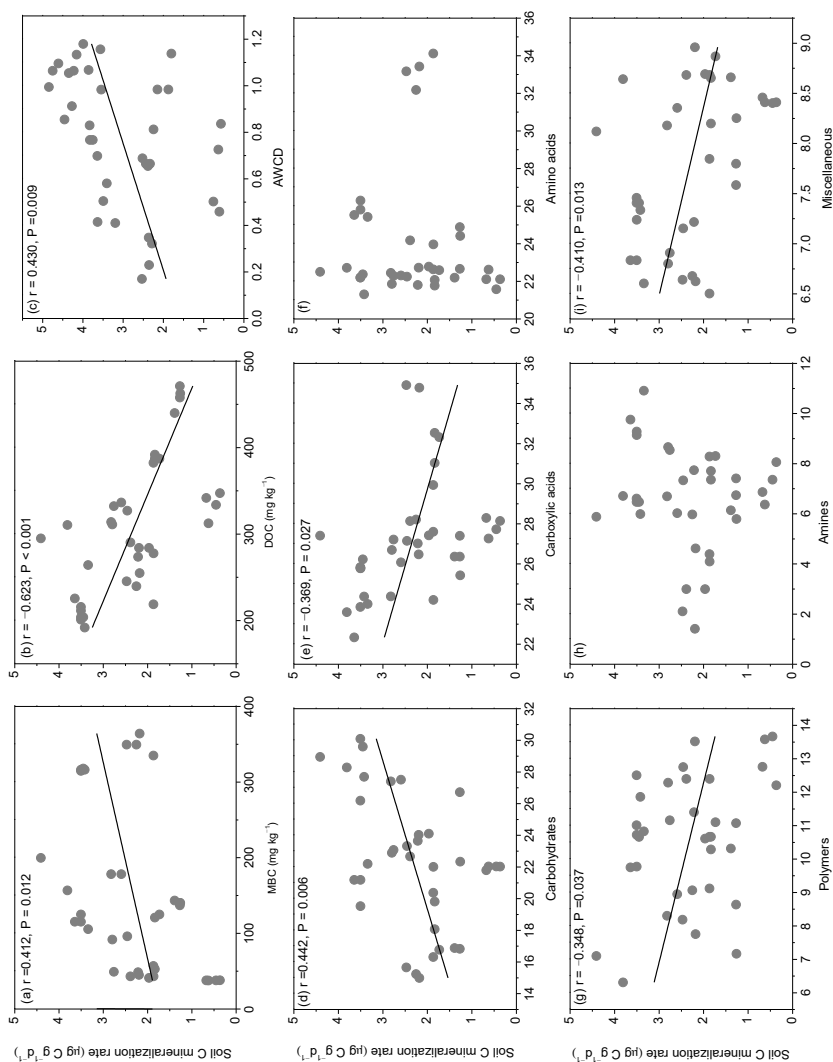


643
644
645
646
647
648
649
650

Figure 4. Redundancy analysis (RDA) ordination biplot of soil microbial carbon sources use efficiency and environmental properties. The representatives of different colors were showed in Figure 2. The dotted lines and solid lines represent the environmental variables and lipid signatures and carbon sources. The abbreviations of the variables in this figure are as follows: MAP, mean annual precipitation. The vegetation data: LDMC, leaf dry matter weight; Leaf C, leaf carbon content; Leaf N, leaf nitrogen content; SLA, specific leaf area. Soil properties included SMC, soil moisture content; Silt, soil silt content; TN, soil total nitrogen; SOC, soil organic carbon. The abbreviations of the sampling sites were given in Table 1.



651
 652 **Figure 5.** The heatmap of the Pearson's correlation coefficients between the use of individual substrates and microbial PLFAs and soil enzyme activities. Note: The abbreviations of the variables:
 653 Actino-, actinomycetes; F/B, fungi/bacteria; G⁺, gram positive bacteria; G⁻, gram negative bacteria; G⁺/G⁻, Gram-positive bacteria/ Gram-negative bacteria. BG, β-1, 4-glucosidase; NAG, β-1,4-
 654 N-acetylglucosaminidase; LAP, leucine aminopeptidase; AP, acid phosphatase. ***P*< 0.01, **P*< 0.05.



655
 656 **Figure 6.** Relationships between soil carbon mineralization rates ($\mu\text{g C g}^{-1} \text{d}^{-1}$) and microbial biomass C (MBC), soil dissolved organic C (DOC), average well color development (AWCD), and
 657 use of individual substrates.

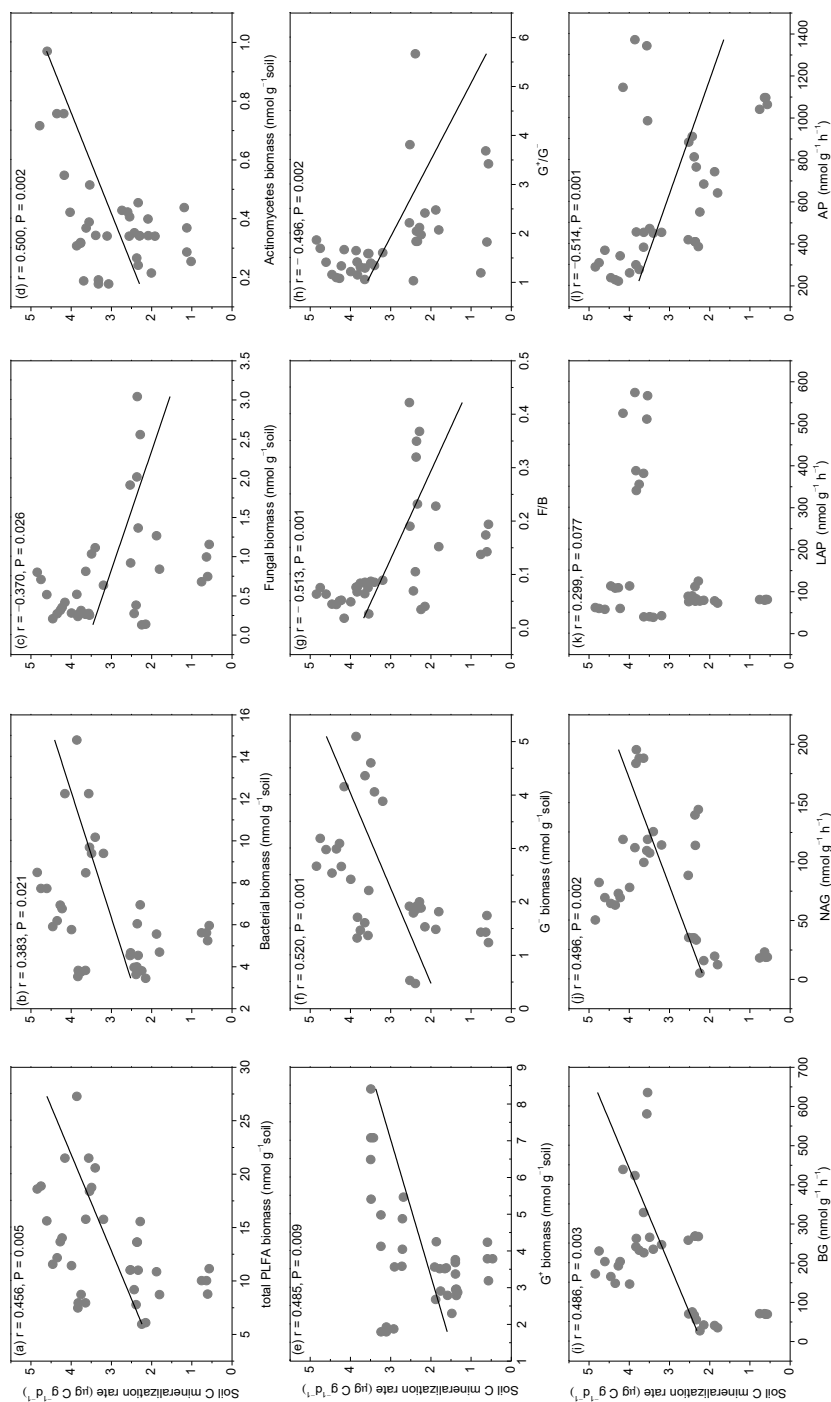


Figure 7. Relationships between soil carbon mineralization rates ($\mu\text{g C g}^{-1} \text{d}^{-1}$) and different groups of soil microbial PLFAs (a-h) and enzyme activities (i-l).