We appreciate the comments from two anonymous reviews. We answer below in blue text to each comment listed in black adding the explicit action taken in a revised manuscript.

Reviewer #1

This paper develops an analysis of the global land C fluxes for three decades. The aim of the paper is to analyse the performance of a data assimilation system, over a historical period. The paper misses clear science questions that would broaden the interest of the study. I lack a clear understanding of the novelty of this work. The conclusions focus on the result that a single decade of data leads to similar results to 3 decades of assimilation, but I am not clear how robust this result is, and how significant it is. I suggest the authors focus more clearly on the novelty of their experimental results. It would be helpful to focus on science questions that are more broadly relevant to the C cycle science community.

We appreciate the reviewer's concerns and are grateful for the comments, which have helped us to formulate the manuscript more clearly. We present three decades of land C flux reanalysis with the aim to better understand the ability of MPI-CCDAS to make decadal projections of the land C cycle. Our approach was to assimilate different time-periods of data to understand the effect of data selection, but also the effect of withholding more recent information. In this way, we can assess the projection of carbon fluxes for the more recent period not simply against the withheld observations, but also against a model run, which has been constrained by this data, in order to disentangle model limitations (i.e. in reproducing an observation that the model was informed about through data assimilation) from prognostic uncertainty (i.e. in failing to reproduce an observation that it can reproduce when given this information). To our knowledge, this is a novel design and focus. Previous studies looking at the prognostic capacity of a CCDAS only looked at short timescales after the assimilation (Scholze et al. 2007; Schürmann et al. 2016), or in the case of Rayner et al. 2011, the authors used a much simpler model ignoring the interacting effects of water, energy and phenology on the carbon cycle predictions. In the revised manuscript, we will be more explicit and clear in the scientific contribution of our results that are relevant to the carbon cycle scientific community.

Our results demonstrate that it is not necessary to ingest more than one decade of observations to improve important features of the global carbon cycle over the following 1-2 decades, in particular the long-term trend and seasonal amplitude of atmospheric CO₂ concentrations at station level, as well as in the long-term trend, phenological seasonality and interannual variability of FAPAR. These results provide an insight into the amount of data that is necessary in data assimilation systems to improve the representation of the global carbon cycle components. This information might be decisive for opening the possibility of including newly measured data of other global indicators, such as SIF, that only exist for periods of less than a decade.

It might be helpful to add experiments assimilating just FAPAR or just CO2, and test the outputs against the other (withheld) observational dataset. This experiment might indicate whether the model effectively couples canopy processes with atmospheric concentrations.

We appreciate the suggestion of the reviewer regarding the experiment of only assimilating one of the observational data sets at a time. However, we did not want to repeat this experiment because this has been previously done with the same system in

Schürmann et al. 2016. In that work, 5 years of observational data (2005 – 2009) were assimilated in three independent experiments: only FAPAR, only atmospheric CO₂ concentrations and the two data sets simultaneously. The results showed that with this time period it was sufficient to demonstrate, in a nutshell, that when assimilating only FAPAR, the average growing season and vegetation seasonality (indicated by FAPAR) was considerably improved in the northern boreal areas. When atmospheric CO₂ concentrations were the only assimilated observations, the global gross and net carbon fluxes were overall improved, but the GPP in the tropics was significantly reduced when compared to the GPP after the FAPAR-only assimilation.

While we agree with the reviewer that in principle it would have been interesting to repeat these experiments to analyze their effect on long-term trends, this would require too much computational time to be included in this study and it was not the central focus of this work.

The abstract is too long, it needs to be reduced to 300 words and focused on the key outcome

We will significantly shorten the abstract and remain under 300 words including only the key outcome regarding using different periods in the observational data during the assimilation.

Writing style is clunky, missing words – for instance the opening sentence is poorly comprehensible: "The observed contemporary in atmospheric CO2 is driven by anthropogenic emissions from fossil fuels and land-use change". I suggest the authors work together to improve the English, reduce sentence lengths, and shorten the more bloated paragraphs.

We realized that we missed the word "increase" in the opening sentence and appreciate the reviewer for pointing this out. The sentence should read: "The observed contemporary **increase** in atmospheric CO_2 ...", we apologize for this. Besides this correction, the revised manuscript will undergo a thorough revision for the English language with the aim of shortening sentences and paragraphs.

Methods: The global grid is very coarse (8x10) – what are the implications?

A previous work with the dynamic global vegetation model LPG-DGVM (Müller and Lucht, 2007), demonstrated that the effect of increasing the spatial resolution did not have a strong effect on the simulated regional and global carbon fluxes, whereas temporal dynamics are unaffected. To our knowledge, the effects on the simulated C fluxes after assimilation due to changes in the horizontal spatial resolution of the model grid cells solely, has not yet been assessed with current data assimilation systems. One reason for this is that the computational cost of such an approach is prohibitive. A study using ORCHIDAS (Peylin et al., 2016) suggested that the level of complexity of the ecosystem model in a data assimilation system, including its spatial resolution, does not guarantee improvements in the optimization of C fluxes. Such a conclusion is likely valid for the assimilation of atmospheric observations, but does decrease the ability of the CCDAS to adequately constrain phenology parameters given the mixed phenology in these large pixels. We would like to point out, however, that increasing the spatial resolution to a degree that would allow for a clean identification of PFT-wise parameters, exceeds the resolution of many state-of-the-art forward terrestrial biosphere models. We will add this discussion in the revised manuscript.

Why mix such a large spatial grid with such fine temporal scales? Is this valid or required? Why not have a grid that matches TM3?

We are unsure on the meaning of the question by the reviewer, because it combines both spatial and temporal scales. In the CCDAS, we integrate the simulated daily net CO_2 fluxes to monthly scale and transport them with the Jacobian representation of TM3. This approach allows us to account for the non-linear impact of weather anomalies on the surface fluxes, but removes the impact of synoptic atmospheric transport variability on the simulated seasonal and long-term dynamics of atmospheric CO_2 at the monitoring stations. We will improve this explanation in the revised manuscript.

We would have liked to use MPI-CCDAS on the same grid as TM3; however, the land-surface model is currently numerically too expensive to allow these higher resolutions.

Explain what is meant by 'each iteration' (L. 139).

The term "each iteration" refers to every cycle when the model re-calculates the cost function for the difference between the model parameters and the observational constraint. We will add this explicitly to the revised ms for clarification.

It is unclear how FAPAR data are generated from equation 2, which seems to generate NDVI estimates. We require details on the FAPAR observation operator – I could not find any.

We apologize for the error in Eq. 2, which should have read:

"We therefore used as FAPAR proxy the Global Inventory Monitoring and Modeling System (GIMMS) NDVI product for the period 1982 to 2006 (Tucker et al., 2005), and merged it with the TIP-FAPAR product to provide a longer record of vegetation greenness. The maximum and minimum NDVI values were rescaled at the pixel level to coincide with those from the TIP-FAPAR for the overlapping time periods (i.e. 2003 to 2006) following:

$$FAPAR_{mod} = \frac{{}_{NDVI-NDVI_{min,x}}}{{}_{NDVI_{max,x}-NDVI_{min,x}}} \times \left(TIP_{max,x} - TIP_{min,x}\right) + TIP_{min,x}$$
(2)

where x is the period 2003 to 2006 for each data set, $FAPAR_{min}$ and $FAPAR_{max}$ are from the NDVI product, NDVI is the full product from 1982 to 2006, and TIP_{max} and TIP_{min} are from the TIP-FAPAR product. In this way, the merged product maintains the maximum and minimum values from TIP-FAPAR while preserving the temporal dynamics of NDVI. Due to a technical failure in the CCDAS, the final FAPAR product spans only from 1982 to 2006 and the last four years from TIP-FAPAR product (2007-2010) were not included."

Why are there no FAPAR data after 2006? (1. 278)

Unfortunately, this is an error during the assimilation procedure. Due to a technical fault, the CCDAS did not consider the remaining four years of data from the original TIP-FAPAR time-series as planned. We discovered this issue only during the post-processing phase, and we wrote it in the final sentence of the paragraph above. However, we believe that our results are still valid, because the main information gain of the CCDAS in terms of phenology stems from the seasonal cycle, with little effect on the overall trends between the three assimilation experiments with different time

periods. This comment will be added also in the revised manuscript.

What are the implications of not including fire emissions? Why were fire products such as GFED not used to provide this input?

Omitting fire fluxes may impair the ability of the MPI-CCDAS to correctly infer the atmospheric growth rate of CO₂ in years with strong contribution of a fire flux, such as the 1998 El Niño. It was not possible to use GFED-like products in this particular assimilation experiment, because for example GFED4 data do not exist prior to 1997, whereas our inversion started in 1982. Adding them only starting in 1997 would have biased the assimilation procedure. In the discussion manuscript we added already the lack of this process as part of the limitations of the MPI-CCDAS in the discussion section.

I would like to know more about the process to determine which parameters were selected for optimisation? These seems to mostly phenological variables which will link to FAPAR. I would expect to see other parameters related to C turnover, e.g. mortality rates, decomposition rates.

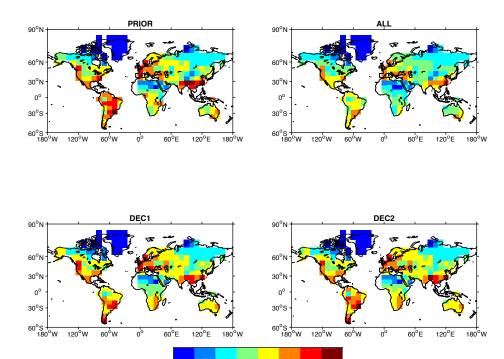
As described in Schürmann et al. 2016, the choices of parameters were done by an extensive parameter sensitivity study with a large set of MPI-CCDAS model parameters for a wide range of biomes. We retained parameters that had a strong effect of the simulated carbon and water fluxes as well as phenology. In Table 1, we listed the parameters selected for the optimization. While the majority are indeed linked to phenology, we considered also parameters linked to photosynthesis and global parameters that control the land carbon turnover. Those are the last four parameters listed in Table 1: for the heterotrophic respiration the temperature sensitivity to respiration (Q_{10}) and a multiplier for initial slow pool ($f_{\rm slow}$) to account for non-steady state conditions at the beginning of the assimilation; for the autotrophic respiration, the leaf fraction of maintenance respiration, and finally an initial atmospheric carbon concentration.

Results:

What are the substantial changes in tropical LAI (l. 306) – how do these match in situ measurements?

As shown in Table 1, the maximum LAI value is one of the optimization parameters and it was prescribed for each PFT. In the case of the tropical evergreen and deciduous trees, this equals to 7 m² m⁻². To provide a numerical context to the LAI changes, we summarize in the table below the mean and maximum LAI values per regions of Fig. 1 for each experiment, and we also show in the figure below, the average maximum LAI for each experiment for the period 1980-2010.

Region	PRIOR	ALL	DEC1	DEC2
	(mean / max)	(mean / max)	(mean / max)	(mean / max)
	$(m^2 m^{-2})$	$(m^2 m^{-2})$	$(m^2 m^{-2})$	$(m^2 m^{-2})$
BE	0.61 / 2.29	0.60 / 1.94	0.70 / 2.42	0.69 / 2.42
BW	0.31 / 1.62	0.30 / 1.44	0.35 / 2.01	0.35 / 2.02
TNE	1.28 / 4.28	1.17 / 3.33	1.31 / 3.49	1.32 / 3.79
TNW	1.26 / 3.11	1.15 / 2.84	1.30 / 3.23	1.30 / 3.21
TE	1.62 / 3.27	1.30 / 2.43	1.63 / 3.20	1.67 / 3.33
TW	2.21 / 3.17	1.68 / 2.27	2.00 / 2.89	2.08 / 3.00
TSE	1.54 / 2.72	1.43 / 2.51	1.86 / 2.77	1.83 / 2.68
TSW	2.42 / 3.69	2.04 / 2.71	2.38 / 3.47	2.43 / 3.66



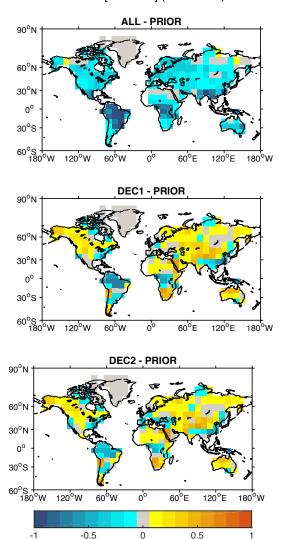
When we compare the LAI mean values between the experiments and the PRIOR results (Fig. 3 in the discussion manuscript), we observe that the largest change in LAI values was in the tropical west area (TW) comprising Brazil, with a decrease in LAI values of up to 24 % in the ALL experiment with respect to the PRIOR, as a response of the maximum LAI decay in the tropical evergreen PFT (visible in the figure above). We argue that this decrease is a response of a global compensating effect to heterotrophic respiration, leading to the lower GPP tropical value.

Ground based observations in the tropical Amazon-Savanna transition forest have been reported with an annual mean LAI value for the total canopy between 2005 and 2008 of 7.4±0.6 m² m⁻², and for the seasonal flooded forest a value of 3.4±0.1 m² m⁻². For the remote sensing data from MODIS, the reported values are 6.2±0.2 m² m⁻² and 5.8±0.3 m² m⁻², respectively (Biudes et al., 2014).

In the eastern Amazon forest, the remote sensing-based LAI has been reported as 6.2 m² m⁻² from LiDAR, and 4.8 m² m⁻² with a low end of 2.0 m² m⁻² from MODIS (Qu et al., 2011). The maximum LAI values from our model results before and after the assimilation (see table above) fall within the values from MODIS and LiDAR. However, this comparison is robust because of the spatial resolution of the different methods: a coarse model grid cell resolution vs. the resolution of ground-based measurements and the resolution of the remote sensing pixels (50x50 m for ground-based and LiDAR data, and 463 x 463 m for MODIS).

As mentioned below, Figure 3 will be moved to the supplement for the sake of keeping the revised manuscript more focused on the main aim, and because the reference to the LAI results were done only sporadically through the manuscript. We will also add (at the editors discretion) another figure in the supplement to show the differences between experiments for the maximum LAI of the mean values.

Max LAI [m^2/m^2] (1980-2010)



What is the increase in R2 (1, 314) – please report in the text.

-0.5

The R² values between the FAPAR observations and model results are: 0.1638 for PRIOR, 0.1984 for ALL, 0.3412 for DEC1 and 0.3402 for DEC2. The values are given in Lines 318-323. We will improve these paragraphs to make clearer that the values are given in those lines by moving the values to the lines above.

The key result seems to be shifts in the timing of CO2 exchanges – it would be interesting to focus more on the shifts in model process representation (parameters) required to allow these changes. I do not find Figure A2 very helpful in this regard. We appreciate the comment from the reviewer, however we believe that an

interpretation of the results at the level of changes in parameters only adds an incomplete picture to the analysis and it is difficult to conclude the effect of the changes. For this reason, we only focused on the relative changes summarized in Fig. A2. We will add a section in the Appendix regarding the assimilation performance where we discuss in more detail the response of each parameter.

The large drop in GPP in the posteriors is significant – what process is this traced to in the parameter adjustment?

This results primarily from a reduction in tropical leaf area index (see Figure 3 in discussion manuscript that will be moved to the supplement for the revised version, and also in the comment above), as well as the drop in the photosynthetic capacity (see larger change in Fig. A2 for parameter f_{photos} in the tropical evergreen and deciduous PFTs) after the assimilation. In our results, we will explicitly add these observations in the appendix part of the assimilation performance mentioned above.

I did not find fig 4 helpful in regard to identifying improvements in IAV modelling – it would help to have some statistics to support the statements here (l. 461). In the revised manuscript, we will present summary statistics (RMSE, R²) in the text, and will move the Figure to the supplementary material.

Discussion

"the mismatch between observations and model output is small, and thus of little concern". This statement needs to be more rigorous – how is 'small' determined? What is the threshold for concern? The lack of tropical IAV suggests too weak an ENSO response – were any relevant parameters included in the CCDAS that would have allowed identification of drought response? Some of the discussion seems circular – that using FAPAR data in the assimilation improves modelling of FAPAR seasonality.

We will clarify this statement a) to note that this corresponds to the IAV of FAPAR, b) that the observed signal is small compared to seasonal variations, and c) the retrieval error or the FAPAR product, which as a global average corresponds to ± 0.2088 (relative units, Schürmann et al. 2016). The assimilation procedure allows changes in the phenology response to water stress (τ_w). However, the assimilation tended to decrease rather than increase the drought sensitivity of tropical phenology given the entire spatially explicit FAPAR time series, and therefore did not allow to capture these excursions (assuming that they are actually driven by drought related changes in LAI).

It would be good to clarify the text to the explicit calibration and validation results to strengthen this section. It seems this result has already been identified in an earlier CCDAS publication, so it is not clear what is novel here.

In the revised manuscript, we will clarify the text and reduce redundancy to the earlier publication. However, since we are using new data and run a different set-up, we believe that it is important to establish the baseline performance of the CCDAS before looking into the novel results of long-term trends. We will condense this section to the absolute necessary to give more space to the novel results (see also our response to the selection of Figures below).

The authors identify problems "results from the structural dependence of the MPICCDAS on few, globally applicable PFT-level parameters, and challenges in using the spatial mixed signal at the model resolution to infer PFT-specific parameters." It would help to develop these ideas some more – do we expect these issues to specifically affect the current analysis and in what ways?

In the revised manuscript, we will expand this discussion. Firstly, this relates to the problem mentioned by the reviewer above with respect to the impact of coarse spatial resolution. We believe that aggregating the remote sensing data into PFT-specific

classes per pixel from a high-resolution grid would allow reducing the problems in the identification of phenological parameters. Secondly, although some of the phenological parameters adapt to mean growing season temperature, some of the thresholds are globally applicable, which causes mainly a problem for temperate grasslands, which cover a wide climatological range. Finding an appropriate means to cluster grasslands into more spatially refined classes would further reduce the errors of the MPI-CCDAS to simulating boreal, temperate and tropical phenology. Finally, some of the global parameters (such as $f_{\text{aut_leaf}}$ and f_{slow}) imply that improvements of modeled fluxes in the boreal regions directly affect fluxes in the tropics, inducing parameter changes to compensate for the altered C fluxes. Such dependence is typical in biosphere models, but may not be ecologically and eco-physiologically correct. Defining these parameters per PFT would reduce such a problem. However, any of these changes would inflate the inverse problem to be solved, therefore increasing computational costs and would not necessarily reduce overall uncertainty (equifinality).

Table 3. The posteriors suggest a Ra:GPP ratio of _65% - it would be useful to discuss this value which seems high for a global estimate.

We believe that this features is a result of the fact that net primary production itself is not well constrained from the atmospheric record. We suspect that two factors contribute to the low NPP:GPP ratio: i) the observed fast coupling between GPP and both autotrophic and heterotrophic respiration, which cannot be reproduced by a state-of-the-art first-order-decay soil carbon turnover model. Since autotrophic respiration in MPI-CCDAS is directly coupled to GPP, increasing the fraction of GPP partitioned to it increases the seasonal cycle of ecosystem respiration; ii) Increasing Ra reduces the net land carbon uptake, and may mask changes in vegetation carbon turnover, which were excluded from the analysis, because their effect on carbon storage was much lower than that of changing $f_{\text{aut_leaf}}$. We note that accounting for the vegetation carbon turnover parameters without any further constraint on NPP would likely not have increased the confidence in the CCDAS outcome because of equifinality.

The very large reduction in soil C stocks from the prior needs further discussion – JSBACH was spun up to steady state for the prior, so I am not clear how the experiments generated 50% drops in this value. How far is the model from steady state with such a reduction in soil C?

The model was spun-up initially until the soil carbon pools reached equilibrium considering pre-industrial forcing. However this new "initial state" for the model is not on steady state when considering climate variability, hence to compensate this, the CCDAS creates an artificial sink of C, leading to a reduction in the soil C stocks, in order to reduce the respiration. Unfortunately, this is unavoidable and is rather a model effect to compensate by contemporary climate changes.

We hear again that a result here repeats an earlier CCDAs result (l. 660), which reduces the novelty of the analysis.

To investigate the mechanisms that influence the patterns observed in the simulated global GPP or NEP after the assimilation is out of the scope of the presented manuscript. The reference to the work of Schürmann et al. 2016 in this line, as well as in previous others throughout the ms, serves as a point of comparison to previous results with the same model but obtained under a different experimental design (5 years only of assimilation), which also contributes to set the preceding performance of

the current set up (see comment above).

Other comments:

Figures – there are too many figures, some of which are of low value, and this distracts from the key message of the paper. Some of the figures in the appendix are referred to several times, so why are they not in the main text (replacing those referenced less).

With the aim of shortening the manuscript and avoid distraction to the main findings, we will remove some figures for the revised version. Specifically we will remove those figures that are less referred to in the main manuscript such as: Fig. 2, on the experimental design and mentioned only once in the main text, it will be moved to appendix. Figures 3 and 4 will be moved to the supplement: Fig. 3, showing the spatial distribution of mean LAI before and after the assimilation, but the manuscript does not focus on the specific changes on LAI after the assimilation, instead R² values are given in Table 2. Fig. 4 shows the interannual variability of FAPAR for the different sub-regions is only mentioned briefly in the results and will be moved to the supplement.

As from figures from the appendix Figures A3 and A4 were mentioned more frequently hence they would be moved to the main text.

L. 59. Citation needed for this 5.6% value

The reference for this value is LeQueré et al., 2018 which is cited in the lines below.

L. 297. The zone between 20-60 is not well described as "sub-tropical"

We will refer to this range of latitudes as north and south temperate zones (TNW and TNE for west and east northern hemisphere, and TSW and STE for west and east southern hemisphere) in the revised manuscript.

Table 1: Add row numbers

Ok, in a first column we will add row numbers.

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simulated land surface carbon cycle, Geoscientific Model Development, 9, 3321-3346, 2016, 10.5194/gmd-9-3321-2016.

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