



1 **A quest for the biological sources of the ubiquitous**
2 **long chain alkyl diols in the marine realm.**

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4 **Sergio Balzano^{a*}, Julie Lattaud^a, Laura Villanueva^a, Sebastiaan Rampen^a, Corina**
5 **P.D. Brussaard^a, Judith van Bleijswijk^a, Nicole Bale^a, Jaap S. Sinninghe Damsté^{a,b},**
6 **Stefan Schouten^{a,b}**

7
8 ^a Department of Marine Microbiology and Biogeochemistry (MMB) and Utrecht University,
9 NIOZ Royal Netherlands Institute for Sea Research, PO Box 59, 1790 AB Den Burg, The
10 Netherlands

11 ^b Department of Earth Sciences, Faculty of Geosciences, Utrecht University, Budapestlaan 4,
12 3584 CD Utrecht, The Netherlands

13

14 * Correspondence to Sergio Balzano. E-mail address: Sergio.balzano@nioz.nl (S. Balzano).

15



16 **Abstract**

17 Long chain alkyl diols (LCDs) are widespread in the marine water column and
18 sediments but their biological sources are mostly unknown. Here we combine lipid analyses
19 with 18S rRNA gene amplicon sequencing on suspended particulate matter (SPM) collected in
20 the photic zone of the tropical North Atlantic at 24 stations to infer relationships between LCDs
21 and potential LCD-producers. The C₃₀ 1,15-diol was detected in all SPM samples and accounted
22 for >95 % of the total LCDs, while minor proportions of C₂₈ and C₃₀ 1,13-diols, C₂₈ and C₃₀
23 1,14-diols as well as C₃₂ 1,15-diol were found. The concentration of the C₃₀ and C₃₂ diols was
24 higher in the mixed layer of the water column compared to the deep chlorophyll maximum
25 (DCM), whereas concentrations of C₂₈ diols were comparable. Sequencing analyses revealed
26 extremely low contributions (\approx 0.1 % of the 18S rRNA gene reads) of known LCD-producers
27 but the contributions from two taxonomic classes to which known producers are affiliated, i.e.
28 Dictyochophyceae and Chrysophyceae, followed a trend similar to that of the concentrations of
29 C₃₀ and C₃₂ diols. Statistical analyses indicated that the abundance of 4 operational taxonomic
30 units (OTUs) of the Chrysophyceae and Dictyochophyceae, along with 23 OTUs falling in other
31 phylogenetic groups, were significantly correlated with C₃₀ diol concentrations. However, it is
32 not clear whether some of these OTUs might indeed correspond to LCD-producers or whether
33 these correlations are just indirect. Furthermore, based on the average LCD-content measured
34 in cultivated LCD-producing algae, the detected concentrations of LCDs in SPM are too high
35 to be explained by the abundances of the suspected LCD-producing OTUs. This is likely
36 explained by the slower degradation of LCDs compared to DNA in the oxic water column and
37 suggests that some of the LCDs found here were likely to be associated to suspended debris,
38 while the DNA from the related LCD-producers had been already fully degraded. This suggests
39 that care should be taken in constraining biological sources of relatively stable biomarker lipids
40 by quantitative comparisons of DNA and lipid abundances.

41

42 **Keywords:** long chain alkyl diols, HCC, tropical North Atlantic, 18S rRNA gene amplicon
43 sequencing, Eustigmatophyceae, Chrysophyceae, Dictyochophyceae



44 1. INTRODUCTION

45 Long chain alkyl diols (LCDs) are lipids that consist of a linear alkyl chain with 22–38
46 carbons which is hydroxylated at the terminal carbon atom and at an intermediate position,
47 and are usually saturated or monounsaturated. LCDs were identified for the first time in Black
48 Sea sediments (de Leeuw et al., 1981) and have subsequently been found to occur widespread
49 in both suspended particulate matter (SPM) and sediments from both coastal and off-shore
50 sites throughout the World Ocean (Jiang et al., 1994; Versteegh et al., 1997; Rampen et al.,
51 2014b). LCDs can be preserved in marine sediments for long periods of time and their
52 distribution can reflect the environmental conditions at the time they were produced. Several
53 indices, based on ratios between the different diols, have been proposed for the reconstruction
54 of past environmental conditions. The Diol Index, reflecting the proportion of C₂₈ and C₃₀
55 1,14-diols over the sum of C₂₈ and C₃₀ 1,14-diols and C₃₀ 1,15-diol, has been proposed to
56 track ancient upwelling conditions since the 1,14-diols are believed to be mostly produced by
57 upwelling diatoms of the genus *Proboscia* (Rampen et al., 2008). Another index, the long
58 chain diol index (LDI), which is based on the proportion of the C₃₀ 1,15-diol over the C₂₈ and
59 C₃₀ 1,13-diols, shows a strong correlation with sea surface temperature (SST) and is used to
60 determine past SST (Plancq et al., 2014; Rampen et al., 2012; Rodrigo-Gámiz et al., 2015). In
61 addition, since the C₃₂ 1,15-diol is the major component of the LCDs of freshwater
62 Eustigmatophyceae (Rampen et al., 2014a; Volkman et al., 1992), the fractional abundance of
63 C₃₂ 1,15-diol has been suggested to be a marker of riverine input in seawater (Lattaud et al.,
64 2017b; de Bar et al., 2016; Lattaud et al., 2017a). Other markers for riverine inputs in
65 seawater are the C₃₀₋₃₆ 1,ω20-diols which are produced by the freshwater fern *Azolla*
66 (Speelman et al., 2009; Mao et al., 2017). However, application of these proxies in the marine
67 realm remains uncertain as long as the sources of the major marine diols are unknown.

68 The most abundant LCDs in seawater are the saturated C₂₈ and C₃₀ 1,13-diols, C₂₈ and C₃₀
69 1,14-diols, and C₃₀ and C₃₂ 1,15-diols (Rampen et al., 2014b), which are all likely produced by
70 phytoplankton. However, the marine biological sources of LCDs are still not fully clear
71 because, in contrast with the widespread occurrence of LCDs in the sediment, few marine
72 taxa have been shown to contain these lipids. Eustigmatophyceae contain C₃₀ 1,13-, C₃₀ 1,15-,
73 and C₃₂ 1,15-diols (Rampen et al., 2014a; Volkman et al., 1992) but they comprise mostly
74 freshwater species and only a few rare marine representatives from the genus
75 *Nannochloropsis* are known (Fawley and Fawley, 2007; Andersen et al., 1998). Furthermore,
76 the distribution of LCDs in the marine environment does not match that of LCDs of marine
77 Eustigmatophytes (Rampen et al., 2012). Species of the diatom genera *Proboscia* and the



78 dictyocophycean *Apedinella radians* contain C₂₈₋₃₂ 1,14-diols (Sinninghe Damsté et al., 2003;
79 Rampen et al., 2009; Rampen et al., 2011), with the former accounting for significant
80 proportions of marine biomass mostly in upwelling regions (Moita et al., 2003; Lassiter et al.,
81 2006), whereas the latter has been occasionally observed in estuarine environments (Bergesch
82 et al., 2008; Seoane et al., 2005). Few other marine species from classes genetically related to
83 diatoms and Eustigmatophyceae have been recently shown to produce LCDs (Table S1). All
84 the known LCD-producing phytoplankters belong to the eukaryotic supergroup
85 Heterokontophyta, a division which includes, among others, diatoms and brown seaweeds.
86 The widespread occurrence of LCDs in the marine environment, in spite of the restricted
87 abundance and distribution of marine LCD-producers, suggests that these compounds may be
88 produced by unknown phytoplankton species.

89 A way of assessing the sources of biomarker lipids is to compare the abundance of lipids in
90 environmental samples with the composition of the microbial community, as determined by
91 genetic methods. For example, Villanueva et al. (2014) analysed both LCDs and
92 eustigmatophycean 18S rRNA gene sequences in a tropical freshwater lake and found five
93 clades of uncultured Eustigmatophyceae in the top 25 m of the water column of the lake,
94 where LCDs were also abundant. Abundance determination by quantitative polymerase chain
95 reaction (qPCR) highlighted that the number of eustigmatophycean 18S rRNA gene copies
96 peaked at the same depth as the LCDs, suggesting that Eustigmatophyceae are a primary
97 source for LCDs in freshwater (Villanueva et al., 2014). However, one of the limitations of
98 this approach is that it relies on specific eustigmatophycean primers designed based on the
99 sequences available in the genetic databases, which could be biased and not target all the
100 existing LCD biological sources. To compensate for this limitation high throughput amplicon
101 sequencing of the 18S rRNA gene allows the exploration of the total marine microbial
102 communities in deep details (de Vargas et al., 2015; Balzano et al., 2015; Christaki et al.,
103 2014; Logares et al., 2012; Massana et al., 2015; Stoeck et al., 2009) and the combination of
104 this analyses with lipid composition may potentially assist in identifying the main LCD
105 producers in marine settings.

106 In the present study, we quantitatively analysed the composition and abundance of LCDs
107 in suspended particulate matter (SPM) collected along the tropical North Atlantic (Fig. 1A) at
108 different depths in the photic zone (surface, deep chlorophyll maximum and bottom of the
109 wind mixed layer; see also Bale et al., 2018). The 18S rRNA gene abundance and
110 composition of the SPM was also analysed by quantitative PCR (qPCR) and high throughput
111 amplicon sequencing to infer the taxonomic composition and to compare the abundance of the



112 different taxa with that of the LCDs in order to identify the potential marine biological
113 sources of LCDs.



114 2. MATERIAL AND METHODS

115

116 2.1 Cruise transect, ancillary data, and SPM collection

117 Samples were taken during the HCC cruise (64PE393), which took place from 24th August
118 to 21st September 2014 along a transect on the tropical North Atlantic Ocean (see Bale et al.
119 (2018) for details). The transect was from Mindelo (Cape Verde) to a location about 500 km
120 from the Amazon River mouth and then westwards along the coast towards Barbados (Fig.
121 1A). Temperature, salinity and nutrient data have previously been reported in Bale et al.
122 (2018).

123 Seawater was collected from two or three depths at each station to measure the
124 concentration of chlorophyll *a* (Chl-*a*) and the abundances of photosynthetic pico and
125 nanoeukaryotes. Seawater was collected during the up cast using Niskin bottles mounted on a
126 CTD frame. The sampling depths were determined based on the evaluation of the vertical
127 profiles of temperature, salinity, and Chl-*a* fluorescence after the down cast of the CTD
128 deployment. The depth of the bottom wind mixed layer (BWML) and the deep chlorophyll
129 maximum (DCM) were determined based on the lowest position of the mixed layer and the
130 depth at which the highest values of chlorophyll fluorescence were observed. For Chl-*a*
131 determination seawater was collected from the Niskin bottles and filtered through 0.7 µm
132 pore-size glass-fiber (Whatman GF/F) filters, followed by frozen storage. Chl-*a* was extracted
133 with methanol buffered with 0.5 M ammonium acetate, homogenized for 15 s and analysed by
134 high performance liquid chromatography.

135 Photosynthetic pico- and nanoeukaryotes were enumerated by flow cytometry according to
136 the protocol by Marie et al. (2005). In short, 1 mL samples were counted fresh using a
137 Becton-Dickinson FACSCalibur (Erembodegem, Belgium) flow cytometer equipped with an
138 air-cooled Argon laser (488 nm, 15 mW). Phytoplankton were discriminated based on their
139 chlorophyll *a* autofluorescence and scatter signature. Cyanobacteria, i.e. *Synechococcus* and
140 *Prochlorococcus*, were not taken into account for the current study. Size fractionation was
141 performed by gravity filtration with >3 µm average cell diameter phytoplankton groups
142 classified as nanoeukaryotic and those <3 µm average cell diameter as picoeukaryotic
143 phytoplankton.

144 Three McLane *in situ* pumps (McLane Laboratories Inc., Falmouth) were used to collect
145 SPM from the water column for the analysis of both lipids and microbial communities.
146 Similarly to the seawater collected with Niskin bottles for Chl-*a* and flow cytometry analyses,



147 the *in situ* pumps were deployed at the surface (3 - 5 m depth), the BWML and the DCM
148 (Table S2). Between 100 and 400 L seawater were pumped and the SPM was collected on
149 pre-combusted 0.7 μm GF/F filters (Pall Corporation, Washington) and immediately frozen at
150 -80°C . For the determination of the organic carbon concentrations, SPM was freeze dried and
151 analysis was carried out using a Flash 2000 series Elemental Analyzer (Thermo Scientific)
152 equipped with a thermal conductivity detector.

153

154 2.2 Lipid extraction and LCD analyses

155 Lipids were extracted from the GF/F filters as described previously (Lattaud et al.,
156 2017b). Briefly, $\frac{1}{4}$ of the filters were dried using a LyoQuest (Telstart, Life Sciences) freeze-
157 dryer and lipids were extracted using base and acid hydrolysis. The base hydrolysis was
158 achieved with 12 mL of a 1 N KOH in methanol (MeOH) solution by refluxing for 1 h.
159 Subsequently, the pH was adjusted to 4 with 2 M HCl: MeOH (1:1, v/v) and the extract was
160 transferred into a separatory funnel. The residues were further extracted once with MeOH:
161 H₂O (1 : 1, v/v), twice with MeOH, and three times with dichloromethane (DCM). The
162 extracts were combined in the separatory funnel and bidistilled water (6 mL) was added. The
163 combined solutions were mixed, shaken and separated into a MeOH : H₂O and a DCM phase,
164 the DCM phase was removed and collected into a centrifuge tube. The aqueous layer was re-
165 extracted twice with 3 mL DCM. The pooled DCM layers were dried over a Na₂SO₄ column
166 and the DCM was evaporated under a stream of nitrogen. The extract was then acid
167 hydrolyzed with 2 mL of 1.5 N HCl in MeOH solution under reflux for 2 h. The pH was
168 adjusted to 4 by adding 2 N KOH : MeOH. 2 mL of DCM and 2 mL of bidistilled water were
169 added to the hydrolyzed extract, mixed and shaken and, after phase separation, the DCM layer
170 was transferred into another centrifuge tube. The remaining aqueous layer was washed twice
171 with 2 mL DCM. The combined DCM layers were dried over a Na₂SO₄ column, the DCM
172 was evaporated under a stream of nitrogen and a C₂₂ 5,17-diol was added to the extract as
173 internal standard. The extract was separated on an activated aluminium oxide column into
174 three fractions using the following solvents: hexane: DCM (9: 1, v/v), hexane: DCM (1: 1,
175 v/v) and DCM: MeOH (1: 1, v/v). The latter (polar) fraction containing the diols was dried
176 under a gentle nitrogen stream. Diols were derivatized by silylating an aliquot of the polar
177 fraction with 10 μL N,O-Bis(trimethylsilyl) trifluoroacetamide (BSTFA) and 10 μL pyridine,
178 heating for 30 min at 60 $^{\circ}\text{C}$ and adding 30 μL of ethyl acetate. The analysis of diols was
179 performed by gas chromatography-mass spectrometry (GC-MS) using an Agilent 7990B GC
180 gas chromatograph, equipped with a fused silica capillary column (2.5 m x 320 μm) coated



181 with CP Sil-5 (film thickness 0.12 μm), coupled to an Agilent 5977A MSD mass
182 spectrometer. The temperature regime for the oven was the same as that used by Lattaud et al.
183 (2017b). The diols were identified and quantified via SIM (Single Ion Monitoring) of the
184 $m/z=299.3$ (C_{28} 1,14-diol), 313.3 (C_{28} 1,13-diol, C_{30} 1,15-diol), 327.3 (C_{30} 1,14-diol) and
185 341.3 (C_{30} 1,13-diol, C_{32} 1,15-diol) ions (Versteegh et al., 1997; Rampen et al., 2012).
186 Absolute concentrations were calculated using the peak area of the internal standard as a
187 reference.

188

189 **2.3 DNA extraction, PCR, qPCR, and 18S rRNA gene sequencing**

190 On ice a small portion of the GF/F filters, corresponding to 1/16 of their initial size,
191 hence containing SPM from ca. 25 L of seawater, was cut in many small pieces using sterile
192 scissors and tweezers. Filter pieces were then transferred into 2 mL microtubes and the DNA
193 was extracted using a MOBIO powersoil DNA isolation kit (Qiagen) following manufacturer
194 instructions. We amplified the hypervariable V4 region of the 18S rRNA which is considered
195 the best genetic marker for the identification of microbial eukaryotes (Logares et al., 2012;
196 Massana et al., 2015). The V4 is located in a central region (565-584 bp to 964-981 bp for
197 *Saccharomyces cerevisiae*) of the 18S rRNA and it was amplified from the genomic DNA by
198 PCR using the universal eukaryotic primers TAREuk454FWD1 (5'-
199 CCAGCASCYGC GGTAATTCC-3') and TAREuk454REV3 (5'-
200 ACTTTCGTTCTTGATYRA-3') (Stoeck et al., 2010). Primers were modified for multiplex
201 sequencing on a Roche 454 GS FLX system: a 454-adaptor A
202 (CCATCTCATCCCTGCGTGTCTCCGACTCAG), a key (TCAG), and a 10 bp sample-
203 specific Multiple Identifier (MID, Table S3) were bound to the 5' end of the forward primer,
204 whereas a 454-adaptor 2 (CCTATCCCCTGTGTGCCTTGGCAGTCTCAG) and a unique
205 MID (CGTGTC A) were bound to the 5' end of the reverse primer for all the samples. The
206 PCR mixture included 25 μL Phusion Flash High-Fidelity PCR Master Mix (ThermoFisher
207 Scientific) 19.1 μL deionised water, 1.5 μL Dimethyl sulfoxide, 1.7 μL from each primer and
208 25 ng genomic DNA and the V4 region was amplified using the same thermal cycling as
209 described by Logares et al. (2012). Amplicons were visualised on a 1% agarose gel and V4
210 bands were excised and subsequently purified using a QIAquick Gel Extraction Kit (Qiagen)
211 and DNA concentration was measured by Qubit Fluorometric Quantitation (ThermoFisher
212 Scientific). For each sequencing run, 20 samples were pooled in equimolar amount and
213 sequenced using a 454 GS-FLX Plus (Macrogen Korea). Some samples yielded a low



214 number of reads and were re-sequenced; overall 77 samples were sequenced in 5 sequencing
215 runs.

216 To determine the concentration of total 18S rRNA genes within the seawater sampled we
217 carried out qPCR using the same primers and the same cycling conditions as described above.
218 qPCR analysis was performed on a Biorad CFX96™ Real-Time System/C1000 Thermal
219 cycler equipped with CFX Manager™ Software. Abundance of 18S rRNA gene sequences
220 was determined with the same primer pair (TAREuk454FWD1/ TAREuk454REV3) used for
221 the 18S rRNA gene diversity analysis. Each reaction contained 12.5 µL MasterMix phusion,
222 8.25 µL deionised nuclease-free water, 0.75 µL DMSO, 1 µL from each primer and 0.5 µL
223 Sybr green and 1 µL of DNA template. Reactions were performed in iCycler iQ™ 96-well
224 plates (Bio-Rad). A mixture of V4 18S rRNA gene amplicons obtained as described above
225 was used to prepare standard solutions. All qPCR reactions were performed in triplicate with
226 standard curves from 6.4×10^3 to 6.4×10^9 V4 molecules per microliter. Specificity of the
227 qPCR was verified with melting curve analyses (50 °C to 95 °C).

228

229 **2.4 Bioinformatic analyses**

230 Bioinformatic analyses were carried out using the python based bioinformatic pipeline
231 quantitative insight in microbial ecology (QIIME) (Caporaso et al., 2010). Overall we
232 obtained 372 107 raw sequences; reads with a length comprised between 250 and 500 bp, less
233 than 8 homopolymers, and a phred quality ≥ 25 over 50 bp sliding windows were kept for
234 downstream analyses. Chimeric sequences were then identified by comparison with the
235 Protist Ribosomal Database 2 (PR2) (Guillou et al., 2013) using the Uchime algorithm (Edgar
236 et al., 2011) and removed from the dataset along with singletons (i.e. reads not sharing 100 %
237 identity with at least another read).

238 A total of 238 564 reads remaining after quality filtering were clustered into 2457
239 Operational Taxonomic Units (OTUs) based on 95 % sequence identity using Uclust (Edgar,
240 2010). Samples containing less than 1000 sequencing reads were removed from the dataset.
241 The taxonomic affiliation of the OTUs was then inferred by comparison with the PR2
242 (Guillou et al., 2013) using BLAST (Altschul et al., 1990) within the QIIME pipeline. Reads
243 from metazoa and multicellular fungi were removed from the dataset which finally contained
244 1871 OTUs and 184 279 reads. The abundance of the different taxa in each sample were
245 estimated by multiplying the percentage of reads with the concentration of V4 copies
246 measured by qPCR. Taxa containing LCD-producers were extracted from the dataset and
247 plotted using Ocean Data View (Schlitzer, 2017).



248

249 **2.5 Statistical analyses**

250 Linear regression analyses between the concentrations of the different LCDs were
251 performed to assess whether some of the LCDs were likely to derive from a common source.
252 To investigate relationships between LCDs and environmental conditions we calculated the
253 Spearman rank correlation coefficient (ρ) using the R package *vegan* (Dixon, 2003). The
254 environmental data used were temperature, salinity, TOC, nutrients (nitrate, nitrite,
255 ammonium, phosphate, and silica), as well the concentration of Chl-*a* and the abundance of
256 photosynthetic pico and nanoeukaryotes. Samples containing missing data and outliers were
257 removed from the dataset before the calculations. Both correlation coefficients and p-values
258 were calculated and the latter were corrected for False Discovery Rates (Benjamini and
259 Hochberg, 1995). Correlations were considered significant for p-values <0.01 .

260 To investigate the relationships between lipids and microbial taxa we also calculated the
261 Spearman's rank correlation coefficient between the LCD concentrations and the abundance
262 of the different taxa at both OTU and class levels. To this end, taxonomic data were
263 normalized based on the number of V4 copies in the different samples measured by qPCR.

264 Comparisons at class level provide the advantage of pooling distribution data from several
265 closely-related OTUs reducing the number of zeros (samples where a given OTU is absent),
266 as the latter complicate statistical analyses of biological distributions (Legendre and
267 Gallagher, 2001). However pooling OTUs at higher taxonomic levels likely leads to
268 combining the LCD-producers with species falling in the same taxonomic level but which are
269 unable to produce LCDs. We thus removed OTUs observed in fewer than 19 samples (25 %)
270 and compared the resulting OTU table with the LCD concentrations. These analyses were
271 performed using the qiime script *observation_metadata_correlation.py* (Caporaso et al.,
272 2010) and the p-values were corrected for false discovery rates (Benjamini and Hochberg,
273 1995).



274 3. RESULTS

275

276 3.1 Ancillary data

277 The HCC cruise sailed across tropical Atlantic waters (Fig. 1A) in late summer and was
278 targeted at SPM from the photic zone collected at the surface, the bottom water mixed layer
279 (BWML) and the deep chlorophyll maximum (DCM). The extent of the photic zone as well as
280 the depths of both BWML and DCM at each station were assessed based on the vertical
281 profiles of temperature, salinity and chlorophyll fluorescence. Temperature of photic zone
282 waters ranged between 15-29 °C (Fig. 1B), the BWML depth ranged between 9-40 m,
283 whereas the depth of the DCM ranged between 45-105 m. Temperatures varied at the DCM
284 increasing westwards, whereas they were relatively constant at surface and BWML. Salinity
285 varied between 29 and 36.5 PSU (Fig. 1C) at the surface, whereas it was fairly constant in the
286 DCM (36 to 37). The concentration of Chl-*a* varied from 34 to 470 ng L⁻¹ (Fig. 1D), with the
287 lowest values measured at the surface of the easternmost (1 to 6) and westernmost (21 to 23)
288 stations and the relatively higher concentrations in surface waters of the shallowest stations
289 (11 to 13) located above the continental shelf and about 500 km off the Amazon River mouth
290 (Fig. 1A). The POC concentration ranged between 0.6 and 13 mg L⁻¹ and also peaked at
291 surface for the shallowest stations (Fig. 1E).

292 Photosynthetic picoeukaryotes, quantified by flow cytometry, were more abundant at the
293 DCM compared to surface and BWML (Fig. 1F). Their abundance peaked at the DCM of
294 Stations 1 and 2 (>1.5 x 10⁷ cell L⁻¹), whereas for surface waters the highest values were
295 measured at Stations 11 to 13. In contrast, photosynthetic nanoeukaryotes did not vary
296 substantially through the water column and their abundance peaked at the surface of Stn. 17
297 reaching a density of 1.4 x 10⁵ cell L⁻¹ (Fig. 1G).

298

299 3.2 Long chain alkyl diols

300 Six long chain diols were detected, the C₂₈ and C₃₀ 1,13-diols, C₂₈ and C₃₀ 1,14-diols and
301 the C₃₀ and C₃₂ 1,15-diols (Fig. 2). The C₃₀ 1,15-diol dominated all samples, accounting for
302 >95 % of the total LCDs, and its concentration ranged from 100 to 1600 pg L⁻¹. The
303 concentration of the C₂₈ 1,13-diol ranged from 0 (i.e. undetectable) to 55 pg L⁻¹, whereas the
304 highest concentration measured for the C₂₈ 1,14-diol was 64 pg L⁻¹. The other minor diols
305 were usually more abundant than the C₂₈ diols, reaching concentrations up to 190 pg L⁻¹ for
306 the C₃₀ 1,13-diol, 240 pg L⁻¹ for the C₃₀ 1,14-diol, and 480 pg L⁻¹ for the C₃₂ 1,15-diol (Fig. 2).



307 The concentration of the C₂₈ 1,13-diol peaked in the surface waters of Station 10 but it was
308 below the detection limit in 19 samples from different depths and stations (Fig. 2A). The C₂₈
309 1,14-diol reached its highest concentrations at the DCM of Station 12 (64 pg L⁻¹) and at the
310 surface of Station 13 (45 pg L⁻¹) and tended to be more abundant in the waters of the eastern
311 stations (Fig. 2B). The concentrations of both C₂₈ 1,13- and C₂₈ 1,14-diols did not vary
312 significantly with depth (t-test, p-value >0.1), while those of the C₃₀ 1,13-, C₃₀ 1,14-, and C₃₀
313 1,15-diols were higher in the mixed layer (surface and BWML) compared to the DCM (p-
314 value <0.01).

315 The concentration of the C₃₀ 1,13-diol peaked at the surface of Stations 10 and 14 (Fig.
316 2C), while that of the C₃₀ 1,14-diol reached its maximum at the BWML of Stations 7 and 8
317 (Fig. 2D). The highest concentration of the C₃₀ 1,15-diol was measured at the surface of
318 Station 17 (16 ng L⁻¹, Fig. 2E). The concentration of the C₃₂ 1,15-diol peaked in the surface
319 waters of Station 10 and 14 and at the DCM of Station 7 (Fig. 2F) and its concentration did
320 not vary significantly with depth. The concentrations of both the C₃₀ and C₃₂ diols peaked in
321 the mixed layer of Stations 7-10 and 14-17, which are located in close proximity to the
322 Amazon Shelf (Figs. 2C-F).

323

324 3.3 Eukaryotic 18S rRNA gene diversity analysis

325 Sequencing of the hypervariable V4 region of the 18S rRNA gene of 68 SPM samples
326 resulted in 238 564 reads with an average of 4 987 reads per sample (Table S2). Reads were
327 clustered based on 95% sequence identity and, after removal of reads of metazoa and
328 multicellular fungi, we obtained 1871 operational taxonomic units (OTUs). Rarefaction
329 analyses indicate that >90% of the genetic diversity was captured (Figure S1), suggesting that
330 no sample was undersequenced. Most (>90%) reads sequenced here were assigned to
331 Dinophyceae, Syndiniales, Metazoa, Haptophyta, and Radiolaria (Fig. 3). Samples were
332 grouped according to the depth layer (surface, BWML, and DCM) and analysis of similarity
333 (anosim) revealed that the average variance between samples from different groups was
334 higher than the average variance between samples from the same group (p-value ≈ 0.001),
335 indicating that the eukaryotic community was mostly influenced by the water depth rather
336 than the geographic location. The proportion of reads from Dinophyceae, Syndiniales, and
337 Haptophyta was slightly higher in the mixed layer compared to the DCM, whereas Radiolaria
338 and Pelagophyceae tended to be slightly more abundant in deeper waters (Fig. 3). All samples
339 except surface waters from Station 12, the BWML from Station 11 and the DCM from Station
340 22 exhibited high contributions (>50%) from Dinophyceae and Syndiniales (Supplementary



341 Figure S2). Radiolaria dominated the DCM at Station 22, diatoms were relatively abundant (\approx
342 10-20 %) at the surface of Stations 12-14 and the BWML of Station 12 while the contribution
343 of diatom reads was <5 % for all the other samples.

344 18S rRNA gene reads of only four taxa containing known LCD-producers were detected
345 within our dataset: *Proboscia* spp., Florenciellales, *Heterosigma* spp., and Eustigmatophyceae
346 (Table 1). In 33 out of 68 SPM samples we did not detect any 18S rRNA gene read from
347 known LCD-producers, whereas reads from these taxa accounted for <0.1 % of the total 18S
348 rRNA reads in 24 samples, 0.1 to 0.5 % in 8 samples, 0.5 to 1 % in 2 samples and 1.5 % in
349 one sample (Stn. 20, BWML). The 18S rRNA gene reads from putative LCD-producers were
350 mostly recovered from the mixed layer (Table 1). Florenciellales was the most abundant taxon
351 among the known LCD-producers since it exhibited the highest number of reads (99) and was
352 present in 28 out of 68 samples. The other taxa of putative LCD-producers were detected only
353 in 8 (Eustigmatophyceae) or 2 (*Proboscia* sp. and *Heterosigma akashiwo*) samples (Table 1)
354 accounting from 3 (*Proboscia*) to 45 (Eustigmatophyceae) reads. The Eustigmatophyceae
355 (mostly affiliated to *Nannochloropsis oculata*) were found at surface for the Stations 11, 12,
356 and 13, as well as at the DCM of Station 20 (Fig. 4A).

357 Since species genetically related to cultivated microalgae known to produce LCD may
358 also contain LCDs, we expanded our community composition analyses to groups at a higher
359 taxonomic level, and focused on those classes or divisions that contain LCD-producers (Table
360 S1). Specifically we investigated the distribution of Eustigmatophyceae, since they are the
361 most well-known class of LCD-producers, Pelagophyceae and Chrysophyceae, which include
362 the LCD-producers *Sarcinochrysis marina* and *Chrysosphaera parvula*, respectively (Table
363 S1), Dictyochophyceae, which includes *Apedinella radians* (Rampen et al., 2011), and
364 Raphidophyceae, which include two LCD-producers, *H. akashiwo* and *Haramonas dimorpha*
365 (Table S1). We did not detect any representative of Pinguiphyceae, a class which include the
366 LCD producer *Phaeomonas parva* (Table S1). Reads associated to Pelagophyceae, and mostly
367 (97 %) affiliated to *Pelagomonas calceolata*, were recovered more frequently as they were
368 present in 55 samples with an average abundance of 85 reads (2 % of total reads) per sample
369 and a maximum value of 935 reads (12 % of total) in the DCM of Stn. 23 (Fig. 4B).
370 Pelagophyceae reads were mostly detected in the DCM and were particularly abundant at the
371 3 westernmost stations investigated, where they comprised 8 % of total reads (Fig. 4B).

372 Chrysophyceae and Dictyochophyceae were also detected in most samples (54 and 57
373 samples, respectively) and their reads were recovered more frequently at the surface and
374 BWML of the westernmost part of the transect (Stns. 20-23) and at the surface of Stations 3-4



375 (Fig. 4C and D). Their 18S rRNA gene reads reached abundances of up to 55 and 41 reads
376 (0.4 and 0.6 % of total, respectively), for Chrysophyceae and Dictyochophyceae respectively,
377 in the BWML of Station 20 (Table S4). Raphidophyceae were present only in three samples
378 from Stations 11, 12, and 13 (Fig. 4F).

379

380 4. DISCUSSION

381 4.1. Comparison of diol distributions

382 In general, it is thought that 1,13- and 1,15-diols derive from a different source than 1,14-
383 diols in the marine realm (Sinninghe Damsté et al., 2003; Rampen et al., 2007; Rampen et al.,
384 2011). Indeed, linear regressions showed that the concentration of C₃₀ 1,15-diol is
385 significantly correlated with those of the C₃₀ 1,13- and C₃₂ 1,15-diols (Figs. 5A-B). We did
386 not observe any significant correlation between the concentrations of the C₂₈ 1,13- and the C₃₀
387 1,13- or C₃₀ 1,15-diol (Figs. 5C-D), which might be due to the fact that C₂₈ 1,13-diol was only
388 present in low abundance and below detection limit in 19 out of 71 samples. This low
389 abundance of C₂₈ 1,13-diol is consistent with the relatively high temperatures observed for the
390 tropical Atlantic ocean (Fig. 1B), since the LCD core top calibration study has revealed that
391 the fractional abundance of the C₃₀ 1,15-diol is high and that of the C₂₈ 1,13-diol is low when
392 SST is relatively high (Rampen et al., 2012).

393 The concentration of the C₂₈ 1,14-diol was not correlated to that of the C₃₀ 1,14-diol (Fig.
394 5E), potentially suggesting a different origin for the C₂₈ and the C₃₀ 1,14-diols. However, the
395 concentration of C₃₀ 1,14-diol was significantly correlated to the C₃₀ 1,15-diol (Fig. 5F). This
396 is quite surprising as the 1,14-diols in seawater have been suggested to derive from *Proboscia*
397 spp. (Sinninghe Damsté et al., 2003; Rampen et al., 2009), and to a lesser extent from *A.*
398 *radians* (Rampen et al., 2011) whereas the 1,13- and 1,15-diols are thought to be associated
399 with Eustigmatophyceae (Rampen et al., 2014 and references cited therein). Previous studies
400 highlighted indeed good correlations in the fluxes of C₂₈ and C₃₀ 1,14-diols in the water
401 column of the Arabian Sea (Rampen et al., 2007) and the northwestern Indian Ocean
402 (Rampen et al., 2008). *Proboscia* spp. contain also unsaturated 1,14-diols which were not
403 found here; specifically the warm water species *Proboscia indica* is dominated by C_{28:1} and
404 C_{30:1} 1,14-diols (Rampen et al., 2007) suggesting that the 1,14-diols found here do not derive
405 from *Proboscia* spp.. This is confirmed by the absence or very low proportions of 18S rRNA
406 gene reads from the major producers of C₂₈₋₃₀ 1,14 diols, that are *Proboscia* spp. and *A.*
407 *radians* (Table 1). This suggests two different sources for the C₂₈ and the C₃₀ 1,14-diols.



408 Since the C₃₀ 1,15-diol accounted for >95 % of the LCDs, it is possible that the C₃₀ 1,14-diol
409 was biosynthesised in low amounts, along with C₃₀ 1,13-diol, by the producers of C₃₀ 1,15-
410 diol. This is supported by the fact that Eustigmatophyceae can contain small amounts of 1,14-
411 diols along with large quantities of 1,15 diols (Rampen et al., 2014a): specifically the C₂₈ 1,14
412 diol, accounts for up to the 15% of the total LCDs in *Pseudostarastrum enorme*, and lower
413 proportions (1-5%) of C₃₀ 1,14 diols were previously found in *Vischeria punctata* and
414 *Eustigmatos vischeri* (Rampen et al., 2014a).

415 It has been found that the distributions of LCDs can be affected by riverine input, which
416 is reflected by elevated amounts of the C₃₂ 1,15-diol (>10%, de Bar et al., 2016; Lattaud et al.
417 2017b). However, the fractional abundance of the C₃₂ 1,15-diol in the SPM is low (0 to 4 %,
418 data not shown), far lower than the values typically measured in river-influenced ecosystems
419 such as the Iberian Atlantic Margin (de Bar et al., 2016), the Kara Sea (Lattaud et al., 2017b)
420 or the Congo River plume (Versteegh et al., 2000). The HCC cruise took place in a period of
421 the year (August/September) when the water discharge from the Amazon River is typically
422 low (Molleri et al., 2010), thus leading to low inputs of riverine organic matter into the sea.
423 The LCDs in the sampled SPM are thus likely not impacted by terrestrial input of LCDs.

424 Beyond Heterokontophyta, LCDs may also be produced by lower (Speelman et al., 2009)
425 and higher (Wen and Jetter, 2007; Racovita and Jetter, 2016) plants. However, only 4 reads
426 from our dataset were associated with a plant species, i.e. *Panax ginseng* (Table S4), which is
427 not known to contain LCDs. The near absence of 18S rRNA gene reads from higher plants
428 confirms the low riverine input of organic matter in the SPM of the tropical North Atlantic
429 waters analysed here.

430 We explored the variations in the concentrations of LCDs with respect to environmental
431 data. The C₂₈ 1,13- and 1,14-diols, both occurring in low abundance, did not exhibit
432 significant correlations with any of the environmental data measured here (Table 2). In
433 contrast the concentrations of C₃₀ 1,13-, 1,14- and 1,15-diols exhibited significant but weak
434 positive correlations with temperature and dissolved silica and weak negative correlations
435 with salinity and nitrite. The concentration of the C₃₂ 1,15-diol revealed a correlation with the
436 same environmental variables as the C₃₀ diols except for dissolved silica and nitrite and
437 exhibited a weak negative correlation with the concentration of nitrate. The correlations found
438 here are likely simply due to different water masses: the mixed layer, where the highest
439 proportions of LCDs were measured, exhibited indeed higher temperatures and lower
440 salinities compared to the DCM. We repeated the analyses after excluding DCM samples and
441 did not find strong positive or negative correlations between LCDs and environmental



442 variables (data not shown). Thus, there does not seem to be a major control of environmental
443 conditions on the concentrations of LCDs.

444

445 **4.2 Comparison with eukaryotic abundance and diversity**

446 Although LCDs are likely produced by phytoplankton, the variability in LCD abundance
447 is not correlated to that of Chl-*a* concentration, photosynthetic pico- and nanoeukaryotes
448 (Table 2). This lack of correlation suggests that the LCD-producers found here accounted for
449 only a small proportion of phytoplankton. The high proportion of Dinophyceae, Syndiniales,
450 and Radiolaria found within our genetic libraries agree with previous studies on marine
451 microbial communities based on 18S rRNA gene sequencing in different environments
452 (Christaki et al., 2014; Comeau et al., 2011; de Vargas et al., 2015). However, these taxa do
453 not necessarily dominate marine microbial communities and so our results are likely due to a
454 relatively high number of rRNA gene copies per cell (Zhu et al., 2005). Larger-sized
455 dinoflagellates such as *Prorocentrum minimum* and *Amphidinium carterae* can contain up to
456 1000 gene copies per cell compared to <10 of rRNA gene copies for smaller sized (<3 µm)
457 species of Chlorophyta, Pelagophyceae, and Haptophyta (Zhu et al., 2005).

458

459 *4.2.1 LCD-producers*

460 Although the primers used in this study have a perfect match with the 18S sequences of
461 most eukaryotes (including all the classes containing LCD-producers), and the rarefaction
462 curves indicate that we sampled an appropriate (i.e. >90%) proportion of the eukaryotic
463 community, the large number (100-1000) of rRNA gene copies per cell present within
464 dinoflagellates and Radiolaria might have somehow affected the detection of LCD producers.
465 In particular *Nannochloropsis salina* has been shown to possess only 1-2 copies of 18S rRNA
466 gene (Zhu et al., 2005), and similarly, the other marine *Nannochloropsis* species which do not
467 differ greatly in size from *N. salina* (Fawley and Fawley, 2007) are also likely to have a low
468 number of 18S rRNA gene copies. Known species of LCD-producers were present in only 51
469 % of our samples as revealed by sequencing data (Table 1), whereas the major LCD, the C₃₀
470 1,15-diol, was present in all the samples. This suggests that the LCDs found here were (1)
471 either produced by other species which were not detected using the current methodology or
472 (2) that the LCD-producers were undersampled because of their low number of rRNA gene
473 copies per cell, or (3) that the DNA of the LCD-producers was no longer present in the SPM
474 at the moment of sampling. Specifically, marine Eustigmatophyceae were represented by only
475 two OTUs (denovo2075, *Nannochloropsis oculata*, and denovo229, uncultured



476 Eustigmatophyceae, Table S4) detected in only 8 samples, confirming the hypothesis of
477 Volkman et al. (1999) and Rampen et al (2012) that they are not the major producers of LCDs
478 in the marine environment. Even if we expand our analyses of LCD-related species to a
479 higher taxonomic level, we do not find large proportions of 18S rRNA reads (generally <0.9
480 % of total reads) except for the class Pelagophyceae, which accounts for up to 12 % of total
481 reads (Fig. 2A-E). However, Pelagophyceae are unlikely to be the source of any of the LCDs
482 found here because their vertical distribution does not correspond well to that of LCDs, which
483 were either more abundant in the upper layers (C₃₀ 1,13-, 1,14-, and 1,15-diols and C₃₂ 1,15-
484 diol) or did not vary greatly with depth (C₂₈ diols, Fig. 2). Chrysophyceae and
485 Dictyochophyceae were instead more abundant in the upper layers (Fig. 4B-C) and although
486 none of the three known LCD-producers from these taxa (Table S1) seems to contain the C₃₀
487 1,15-diol, which was dominant here, species within the Chrysophyceae and
488 Dictyochophyceae may possibly be a source for the C₃₀ diols.

489 The C₂₈ diols exhibited higher concentrations at the BWML of Station 12 and at surface
490 in Stn. 13 (Fig. 2A and B), and higher proportions of 18S rRNA gene reads were recovered
491 from Pelagophyceae (2.4 %), and Eustigmatophyceae (0.5 %), at the surface of Stations 11-12
492 (Fig. 4D-F). The scattered occurrences of these groups and the mismatches in distributions
493 when compared to the LCDs suggest that the LCDs in the tropical North Atlantic Ocean are
494 unlikely to derive from Pelagophyceae, radial centric diatoms, Raphidophyceae, and/or
495 Eustigmatophyceae.

496 Overall the abundance of known LCD producers is low and scattered and does not match
497 the observed abundance patterns observed for the LCDs, suggesting that most of the LCDs
498 measured here were not produced by any of these species.

499

500 4.2.2 Correlations between OTUs and LCDs

501 Since LCDs have been shown to be present within two genetically distant eukaryotic
502 supergroups, the Heterokontophyta and the Archaeplastida, the latter including plants as well
503 as green and red algae, the genetic and enzymatic machinery required for the biosynthesis of
504 LCDs might be present in other genera and classes, including uncultured species. We,
505 therefore, also compared the concentration of LCDs with the composition of the entire
506 eukaryotic microbial community, normalised with respect to the 18S rRNA gene abundance,
507 at both class and OTU levels to identify co-occurrence patterns. No significant correlation
508 was found at class level (data not shown), whereas the correlations at OTU level were weak
509 (≤ 0.5) but significant (p-value <0.01) for 27 OTUs affiliated to 11 different classes (Table 2).



510 A reason behind the lack of correlation between taxonomic classes and LCDs can be that
511 pooling OTUs at higher taxonomic levels likely leads to combining the LCD-producers with
512 species which are unable to produce LCDs but are falling in the same taxonomic level. The
513 ability from microorganisms to biosynthesise LCDs can indeed vary even between genetically
514 related species, since it has been shown that some genera include both LCD-producers and
515 species which do not contain LCDs (Table S1).

516 The C₃₀ 1,15 diol exhibited significant correlations ($p < 0.01$) with 23 OTUs and overall,
517 27 OTUs were significantly correlated with C₃₀ or to a lesser extent, C₃₂ diols (Table 3). Of
518 the 27 OTUs, 4 OTUs were affiliated to classes containing known LCD-producers
519 (Chrysophyceae and Dictyochophyceae, Table 3). The abundance of the two chrysophycean
520 OTUs (denovo465 and denovo1680, Table 3) exhibited significant correlations with the
521 concentrations of both C₃₀ 1,13 and C₃₀ 1,15 and accounted for 52 % of the total reads from
522 this class and the only known LCD-producer from this class (*Chryso-sphaera parvula*) was
523 found to contain C₃₂ 1,15 diol (Rampen, unpublished results). The 2 OTUs affiliated to
524 Dictyochophyceae (denovo873 and denovo958) and exhibiting positive correlation with
525 LCDs, cluster within Pedinellales and Florenciellales families, respectively, and are thus
526 closely related to two known LCD-producers, *Florenciella parvula* and *Apedinella radians*.
527 However, *F. parvula* contains C₂₄ 1,13-, C₂₄ 1,14-, and C₂₄ 1,15-diols (Rampen, unpublished
528 results) and *A. radians* produces C₂₈, C₃₀, and C₃₂ 1,14-diols (Rampen et al., 2011), whereas
529 the two dictyochophycean OTUs denovo873 and denovo958 exhibited positive correlation
530 with the C₃₀ 1,15-diol (Table 3).

531 The correlation values found here are nearly all low ($\rho \approx 0.4-0.5$), raising the question on
532 whether these relationships reflect the ability of these species to produce LCDs or that these
533 co-occurrences are simply driven by other environmental conditions leading to similar spatial
534 distributions of OTUs and LCDs. Furthermore, other taxa which showed significant
535 correlations with C₃₀ diols occur rarely in the marine environment. For example,
536 Centroheliozoa are mostly known as freshwater predators (Slapeta et al., 2005) and, in
537 seawater, have been sporadically detected in anoxic environments only (Stock et al., 2009;
538 Stoeck et al., 2009), suggesting that the centroheliozoan reads found here are unlikely to
539 derive from active microorganisms. In contrast, other taxa include marine representatives
540 commonly found in the photic zone of seawater and thus the reads found here might derive
541 from living organisms: Syndiniales are intracellular parasites of other marine protists, and the
542 genetic clades found here (Group I Clade 4, Group II Clades 2, 7,8,17, and 23) are commonly
543 detected in the upper 100 m of the water column (Guillou et al., 2008). Spirotrichea instead



544 include several heterotrophic and mixotrophic marine planktonic ciliates (Santoferrara et al.,
545 2017; Agatha et al., 2004), whereas *Phaeocystis* is a widespread primary producer. The
546 uncultured classes exhibiting significant positive correlations with LCDs (Prasino Clade IX
547 and the HAP-3 clade) are also commonly observed in the photic zone (Lopes dos Santos et
548 al., 2016; Shi et al., 2009; Egge et al., 2015). However, cultivated representatives will be
549 required in order to confirm whether species within these taxa are capable of LCD synthesis.
550

551 **4.3 Can 18S rRNA gene-based community composition analysis be used to determine** 552 **LCD biological sources?**

553 The lack of correlations of C₂₈ and C₃₂ diols with any OTUs as well as the low degree of
554 correlation between OTUs and C₃₀ diols and the trace abundance or near absence of known
555 LCD producers suggest that the 18S rRNA genes from the microorganisms which produced
556 these LCDs were either absent, or present below detection level in the seawater sampled. The
557 fact that we sampled >90 % of the OTUs potentially present (Figure S1) and the use of
558 universal eukaryotic primers suggests that LCD-producers have been unlikely to escape
559 detection. However, the relatively low number of rRNA gene copies found for *N. oculata*
560 (Zhu et al., 2005) and likewise also in other smaller-sized marine Eustigmatophyceae, suggest
561 that LCD-producers might have been undersampled with respect to larger-sized species which
562 can contain up to 1000 rRNA copies per cell (Zhu et al., 2005).

563 It should be considered that both the LCDs and DNA in the SPM might derive not only
564 from active or senescent cells, but also from debris (Not et al., 2009). In addition, LCDs can
565 persist in seawater for likely much longer periods than the DNA of the related LCD-
566 producers. Although the biological function of LCDs is unclear for most species, they have
567 been shown to be the building blocks of cell wall polymers in Eustigmatophyceae. In
568 *Nannochloropsis* cell wall, LCDs and long chain alkenols are likely to be bound together
569 through ester and ether bonds to form highly refractory polymers known as algaenans (Gelin
570 et al., 1997; Scholz et al., 2014). These biopolymers are thought to be quite persistent and
571 accumulate in ancient sediments for millions of years (Derenne and Largeau, 2001; de Leeuw
572 et al., 2006; Tegelaar et al., 1989). Indeed, LCDs are ubiquitous in recent surface sediments
573 (Rampen et al., 2012) and ancient sediments of up to 65 million years old (Yamamoto et al.,
574 1996) showing their recalcitrant nature. Recent laboratory experiments highlighted that LCDs
575 from dead biomass of *Nannochloropsis oculata* can persist in seawater for longer than 250
576 days under anoxic (Grossi et al., 2001) and oxic conditions (Reiche et al., unpublished
577 results). In contrast, much shorter turnover times (6 h to 2 months) are typically reported for



578 extracellular DNA in the oxic water column (Nielsen et al., 2007). This suggests that the
579 DNA from LCD-producers likely reflects the living eukaryotic community (recently) present
580 when seawater was sampled, while the LCDs probably represent an accumulation that
581 occurred over longer periods of time (weeks to months or even years).

582 Because of this large difference in turnover rates between LCDs and the DNA from the
583 LCD-producers, 18S rRNA gene analysis of environmental samples may be unsuccessful for
584 identifying LCD-producers. This is seemingly in contrast to a previous study that showed that
585 the LCD concentration in the upper 25 m of the freshwater lake Challa (Tanzania) was related
586 to the number of eustigmatophycean 18S rRNA gene copies (Villanueva et al., 2014).
587 However, Villanueva et al. (2014) used Eustigmatophyceae-biased primers and since this was
588 a lake system, Eustigmatophyceae are likely to be the major source of LCDs in freshwater
589 ecosystems. Importantly, they found a mismatch for the uppermost part of the water column
590 (0–5 m), where high LCD abundance (38–46 ng L⁻¹) coincided with little or no
591 Eustigmatophyceae 18S rRNA gene copies. This pattern was explained by them to be caused
592 by wind-driven and convective mixing of preserved LCDs, while phytoplankton adjusted its
593 buoyancy at greater depth (Villanueva et al., 2014).

594 Laboratory experiments carried out under different conditions of temperature, light
595 irradiance, salinity and nitrate concentrations revealed average cellular LCD content of about
596 23 fg cell⁻¹ (Balzano et al., 2017) for *Nannochloropsis oceanica*. The average LCD
597 concentration in the SPM investigated was ca. 2.6 ng L⁻¹, which would correspond to ca. 1.1 x
598 10⁶ pico/nano algal cells L⁻¹. We detected average phytoplankton abundances of 3.3 x 10⁶ cell
599 L⁻¹ for picoeukaryotes and 3.6 x 10⁴ cell L⁻¹ for nanoeukaryotes. Although nanoplanktonic
600 Eustigmatophyceae might produce larger amounts of LCDs than those measured in our
601 previous study (Balzano et al., 2017), because of their larger cell size, the nanoplankton
602 abundances measured here are two orders of magnitude lower than the levels required to
603 source the LCDs (1.1 x 10⁶ cell L⁻¹). Therefore, if the LCDs measured here were
604 biosynthesised by intact microorganisms in the water column, nanoplankton alone would not
605 be able to source all the LCDs measured, and therefore in addition at least one-third of the
606 picophytoplankton should be able to produce LCDs, which is unrealistic. This supports the
607 idea that most of the LCDs detected here are of fossil nature and not contained in living cells.
608 The higher concentrations of LCDs found in the SPM from the mixed layer compared to the
609 DCM suggest that LCDs were originally produced at a higher frequency in the mixed layer.
610 Moreover their possible fossil nature indicates that LCDs were likely to persist in the mixed
611 layer for long periods, eventually associated with suspended particulate matter.



612

613 The combination of lipid and DNA analyses is often complicated by different turnover
614 rates, especially for refractory compounds such as LCDs. Studies focused on more labile
615 biomarker lipids such as fatty acids or intact polar lipids can be more successful, e.g. with
616 short branched fatty acids (Balzano et al., 2011) or archaeal phospholipids (Buckles et al.,
617 2013; Pitcher et al., 2011). Therefore, care has to be taken in inferring sources of biomarker
618 lipids by the quantitative comparison of DNA abundance with biomarker lipid concentrations.

619



620

621 **5. Conclusions**

622 The combination of lipid analyses and 18S rRNA gene amplicon sequencing revealed
623 some weak correlations between 23 OTUs and C₃₀ diols. Four of these OTUs are affiliated to
624 classes that include few LCD-producing species (i.e. Chrysophyceae and Dictyo-
625 chophyceae), whereas the remaining 19 OTUs belong to taxa in which the presence of LCDs
626 has not been shown. In both cases it remains unclear whether the correlation between these
627 27 OTUs and the C₃₀ diols reflects novel LCD-producers or is driven by other environmental
628 conditions.

629 The abundances of photosynthetic pico and nanoeukaryotes measured here suggest
630 that these microbial populations are highly unlikely to source all the LCDs found. Some of
631 the LCDs found here might be associated with suspended debris rather than intact cells, with
632 the DNA from their producers being already degraded at the time of sampling. DNA
633 degradation rates in the oxygenated water column are indeed faster than those of most lipids,
634 including LCDs. The freshness of the organic matter and the turnover rates of both lipids
635 and DNA in a given environment should thus be considered when identifying the biological
636 sources of a specific class of lipids through DNA sequencing. In particular the 18S rRNA
637 gene amplicon sequencing can be suitable to track LCD sources (1) for simple ecosystems
638 or laboratory/*in situ* mesocosms with high proportions of fresh organic matter and (2) for
639 low oxygen/anoxic environments where extracellular DNA can persist for longer periods.

640

641

642

643 **Acknowledgments**

644 We thank the captains and the crew of the R/V Pelagia for their support during the cruise.
645 We thank Sharyn Ossebaar for help in sample collection, H. Witte, E. Panoto, and S.
646 Vreugdenhil for support in molecular biology, H. Malschaert for the bioinformatics and M.
647 Besseling for helpful discussions. This research was funded by the European Research
648 Council (ERC) under the European Union's Seventh Framework Program (FP7/2007-2013)
649 ERC grant agreement [339206]. S.S. and J.S.S.D. receive financial support from the
650 Netherlands Earth System Science Centre (NESSC).



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886 **Table 1.** Distribution of the 18S rRNA gene reads associated with known LCD-producers
 887

Taxon	Florenciellales	<i>Heterosigma</i>	Eustigmatophyceae	<i>Proboscia</i>	Total ^a
No of samples	28	2	8	2	35
Surface	12	1	2	0	12
BWML ^b	11	0	2	2	13
DCM ^c	5	1	4	0	10
No of reads	99	10	45	3	157
% total	0.04	0.004	0.02	0.001	0.06
Surface	48	4	25	0	77
BWML	41	0	9	3	53
DCM	10	6	11	0	27

888

889 ^a Total number of samples where 18S rRNA gene reads from LCD-producers were found or
 890 total number of reads associated with LCD-producers. Overall 68 samples were screened for the
 891 presence of 18S rRNA genes affiliated to LCD-producers.

892 ^b Bottom water mixed layer

893 ^c Deep chlorophyll maximum

894



895 **Table 2.** Spearman rank correlation coefficients between LCDs and environmental
 896 variables^a.

897

898		C ₂₈ 1,13	C ₂₈ 1,14	C ₃₀ 1,13	C ₃₀ 1,14	C ₃₀ 1,15	C ₃₂ 1,15
900	POC ^b	0.3	0.2	0.2	0.3	0.3	0.3
901	Salinity	-0.2	0.0	-0.5	-0.7	-0.6	-0.6
902	Temperature	0.2	-0.1	0.5	0.5	0.5	0.5
903	Phosphate	0.0	0.2	-0.3	-0.2	-0.3	-0.2
904	Ammonium	0.0	0.1	-0.3	-0.4	-0.4	-0.2
905	Nitrite	-0.2	0.0	-0.6	-0.5	-0.6	-0.4
906	Nitrate	0.0	0.2	-0.4	-0.3	-0.3	-0.5
907	Silica	0.1	0.0	0.4	0.5	0.5	0.4
908	Chl- <i>a</i>	-0.1	0.0	-0.2	-0.2	-0.3	-0.1
909	Picoeukaryotes	-0.1	-0.1	-0.4	-0.3	-0.4	-0.2
910	Nanoeukaryotes	0.0	-0.1	0.1	0.2	0.2	0.2

911

912

913

914

^a Significant (p-value < 0.01) correlation values are in bold.

915

^c Particulate Organic Carbon

916

Table 3. List of Operational Taxonomic Units (OTUs), representing 95 % of sequence identity, which were correlated^a with LCDs in the HCC cruise.

OTU ID	Taxon	Class	C ₃₀ 1,13	C ₃₀ 1,14	C ₃₀ 1,15	C ₃₂ 1,15	Total 1,13	Total 1,14	Total 1,15
denovo2033	Choreotrichia	Spiritricha							0.40
denovo2137	<i>Climacocylis scalaria</i>		0.45		0.49		0.45		0.49
denovo940	<i>Laboea strobila</i>		0.53	0.46	0.60		0.56	0.45	0.59
denovo685	Oligotrichia				0.41				0.40
denovo1804	<i>Pseudotonia</i>		0.56	0.47	0.56	0.47	0.53	0.41	0.57
denovo492	<i>Blastodinium spinulosum</i>	Dinophyceae	0.43	0.44	0.46				0.45
denovo720	<i>Ceratocorys horrida</i>				0.46		0.44		0.45
denovo1682	<i>Neoceratium fusus</i>							0.47	
denovo526	<i>Protodinium simplex</i>			0.43	0.44			0.48	0.43
denovo267	<i>Pyrophacus steinii</i>				0.43				0.42
denovo732	Dino Group I Clade 4	Syndiniales			0.40		0.46		0.41
denovo555	Dino Group II Clade 2				0.49		0.41		0.48
denovo1077	Dino Group II Clade 7				0.44		0.42		0.43
denovo1834	Dino Group II Clade 8				0.44		0.45		0.45
denovo1145	Dino Group II Clade 17		0.50		0.49		0.53	0.42	0.48
denovo2080	Dino Group II Clade 23				0.40		0.43		0.40
denovo725	Prasino Clade 9B	Prasino Clade IX			0.42		0.41		0.41
denovo1066	Pterocystida	Centroheliozoa			0.46				0.46
denovo400	HAP3	Haptophyta		0.47	0.49			0.47	0.48
denovo2132	<i>Phaeocystis</i>						0.46		
denovo972	<i>Haptolina</i>						0.44		
denovo465	Chrysophyceae Clade G	Chrysophyceae	0.44		0.43		0.45		0.42
denovo1680	Chrysophyceae Clade H		0.44		0.42		0.48		0.42
denovo1988	Raphid pennate	diatoms			0.41				0.41
denovo873	Pedinellales	Dictyochophyceae	0.56	0.45	0.55	0.52	0.55		0.56





	denovo958	Florentiales	Dictyochophyceae				
921	denovo2433	Unidentified picozoan	Picozoa	0.49	0.50	0.43	0.45
922							0.47
923							0.46
924							0.44
925							0.55

^a Only significant (p-value < 0.01 after FDR correction) correlations are shown. OTUs closely related to known LCD-producers are in bold.



926 **Figure Legend**

927

928 **Figure 1.** HCC cruise track in the western tropical Atlantic Ocean, physical seawater
929 properties, and biological parameters. (A) Map of the sampling stations. Spatial distribution
930 of (B) temperature, (C) salinity, the concentration of (D) chlorophyll-*a*, (E) organic carbon
931 concentrations, and the abundance of photosynthetic (F) picoeukaryotes and (G)
932 nanoeukaryotes. Data were plotted using ocean data view (ODV) software using kriging for
933 interpolation between datapoints (Schlitzer, 2017).

934

935 **Figure 2.** Spatial distribution of the concentration of LCDs: (A) C₂₈ 1,13, (B) C₂₈ 1,14,
936 (C) C₃₀ 1,13, (D) C₃₀ 1,14, (E) C₃₀ 1,15, and (F) C₃₂ 1,15. Data were plotted using ocean data
937 view (ODV) software using kriging for interpolation between datapoints (Schlitzer, 2017).

938

939 **Figure 3.** Average fractional abundance of the reads obtained by 18S rRNA gene
940 sequencing of SPM from the western tropical Atlantic Ocean over the various classes of
941 eukaryotes. The V4 fragment of the 18S rRNA gene was sequenced using universal
942 eukaryotic primers. Samples were pooled according to depth and the average contribution
943 from each group at the different depth is shown. Error bars represent the standard deviation in
944 the data from the various stations.

945

946 **Figure 4.** Spatial distribution of the 18S rRNA gene fragments related to taxa containing
947 LCD-producers at different stations and depth. (A) Pelagophyceae, (B) Chrysophyceae, (C)
948 Dictyochophyceae, (D) radial centric diatoms, (E) Eustigmatophyceae, and (F)
949 Raphidophyceae. Data were plotted using ocean data view (ODV) software using kriging for
950 interpolation between datapoints (Schlitzer, 2017).

951

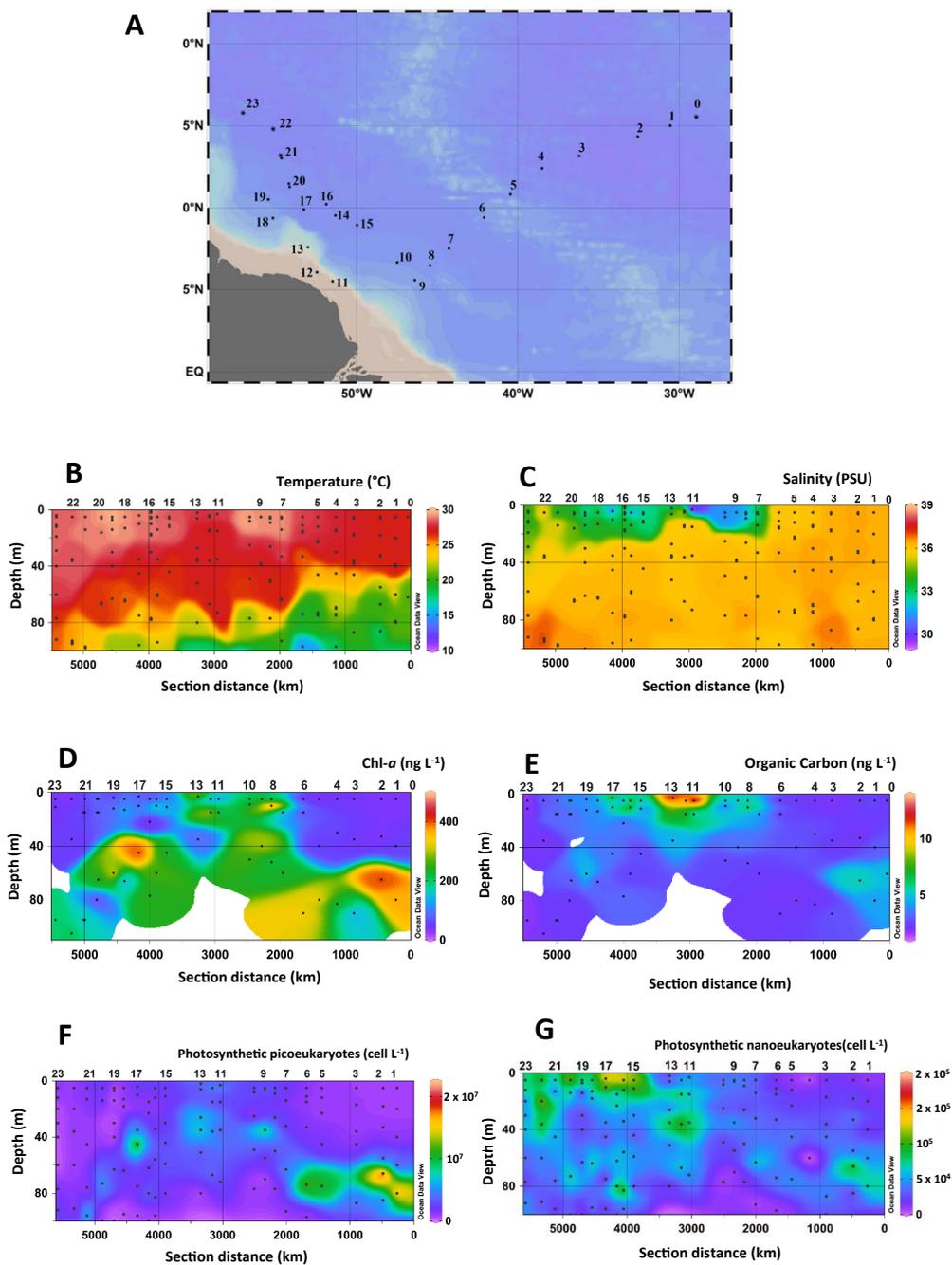


952 **Figure 5.** Scatter plots of the concentrations of the different LCDs in the western tropical
953 Atlantic Ocean. **(A)** C₃₀ 1,13 diol vs C₃₀ 1,15 diol, **(B)** C₃₂ 1,15 diol vs C₃₀ 1,15 diol, **(C)** C₂₈
954 1,13 diol vs C₃₀ 1,13 diol, **(D)** C₂₈ 1,13 diol vs C₃₀ 1,15 diol, **(E)** C₃₀ 1,14 diol vs C₂₈ 1,14
955 diol, **(F)** C₃₀ 1,14 diol vs C₃₀ 1,15 diol.

956



Figure 1



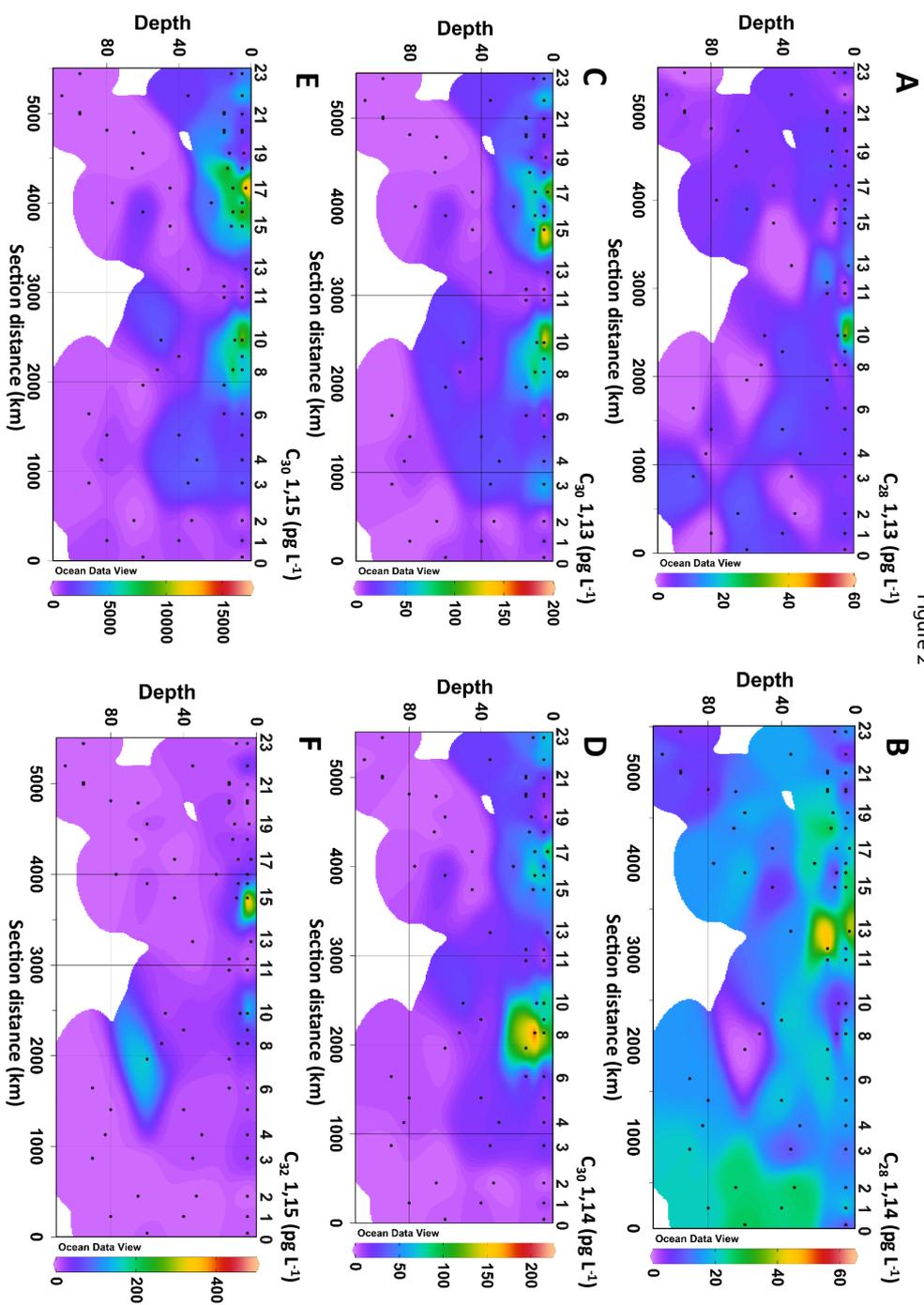
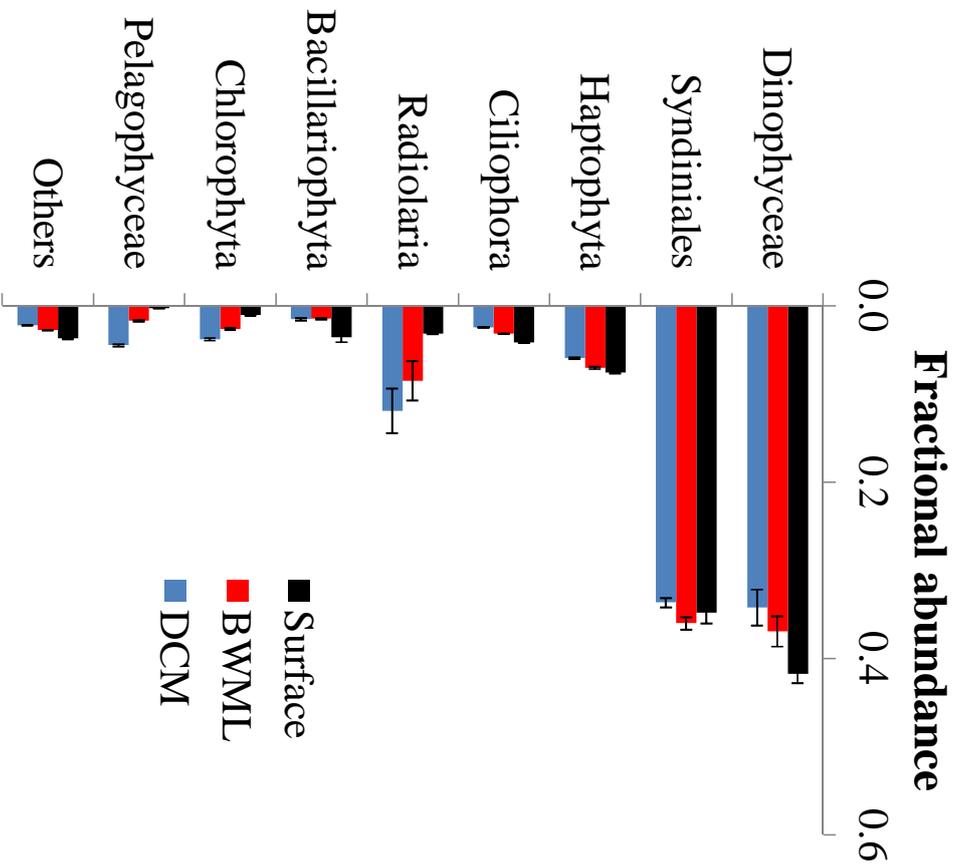


Figure 2

Figure 3



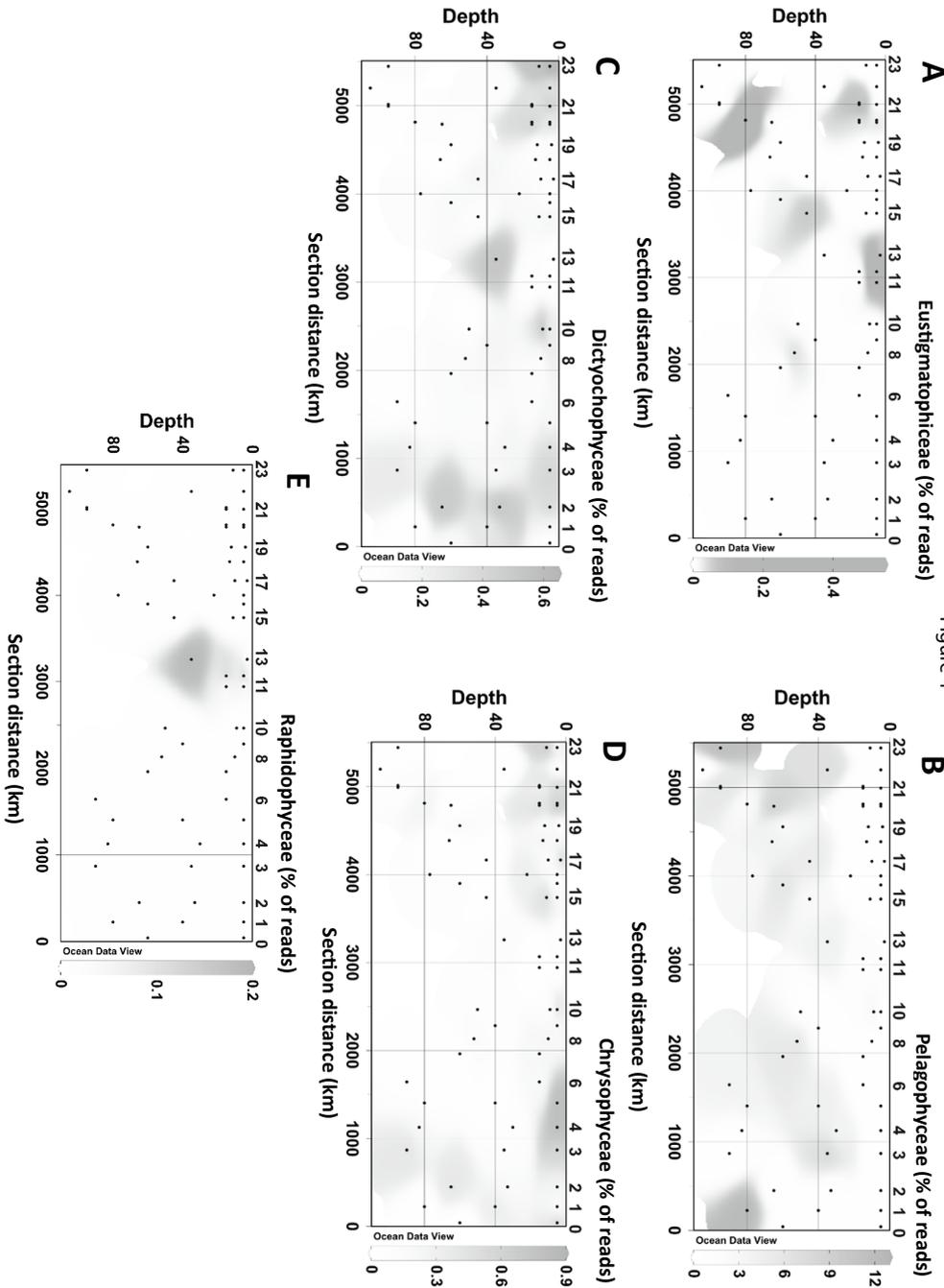


Figure 4





Figure 5

