

LETTER OF RESPONSE TO REVIEWER'S COMMENTS

Dear Editor

This is a revised version of the MS "Zooplankton diel vertical migration and downward C flux into the Oxygen Minimum Zone in the highly productive upwelling region off Northern Chile" by Tutasi & Escribano.

All comments from both reviewers have been considered in this new version and substantial changes have taken place, but not affecting key conclusions from the study.

Details on changes and corrections are detailed below with the text marked in blue color.

Yours sincerely,

Ruben Escribano

Reviewer 1

The manuscript by Tutasi and Escribano is a quite interesting and valuable report about the vertical distribution of different groups and species of zooplankton in relation to the oxygen minimum zone (OMZ) in the upwelling zone off Chile. They also showed the diel movements of the different zooplankton groups and species and their range of migration in relation to the OMZ. This is a rather important data set of general interest and the results deserve publication. In relation to the requirements made in my previous report, now the authors provide a better explanation about their estimation of migrant biomass and active flux by zooplankton.

However, this reviewer consider their assessment of active flux not properly done. The authors use a value of 0.12 d⁻¹ for respiration at the residence depth of migrants. This is a rather high value of respiration for mesopelagic zooplankton. These rates should be half of the value used by the authors, or even less. As stated in the previous report, I recommended the authors to use the equation given by Ikeda (1985) or even better the equation by Ikeda (2014) considering depth to estimate respiration. The authors assessed the individual biomass of zooplankton which jointly with temperature could provide a rather useful estimate of respiration at depth using these equations. Mortality and egestion should also be much lower than the given values as these parameters are normally lower than respiration rates. Therefore, values of active flux are too high, something that is observed in their comparison to other assessments in Table 4 (excepting Escribano et al., 2009). Although this estimation is carried out in a eutrophic system such as the quite productive upwelling off Chile, these values of active flux are still too large due to the considerations given above. Thus, I consider this a major revision before definitive publication.

R: We thank the reviewer by his(her) comments and appreciate the suggestions to improve the work. We have now re-estimated the vertical flux of C by active migration by re-assessing Respiration (R) of zooplankton using Equations provided by Ikeda (1985) considering body mass of variable taxa and in situ temperature as the key factors affecting R and assuming a minor effect of taxa as demonstrated by Ikeda et al (2014). Since we assessed R by taxa and their biomasses we now provide these estimates as a Supplemental Material, because it is an important information for readers. Regarding Mortality rate, as the reviewer is aware this is a parameter difficult to assess and likely to strongly vary in the field. We are now considering a mortality estimate of 0.04 from published works...although we still think it might be larger at depth upon presence of extremely low oxygen, though certainly requires further work. For C excretion at depth, we have now cited the work of Steinberg et al 2000 (DSR-I) and assumed C excretion may be nearly 30% of C respired by migrants. It is important to note that we now consider C excretion as an important component for C contribution at depth, instead of potential contribution by fecal pellet production.

Because biomass has been re-estimated by using new equations as suggested by second reviewer, then all the estimates of migrant biomass, C flux and C export have substantially changed to much lower values. The vertical flux of C and fraction of C being exported is more comparable to other studies, although still high consistently with a highly productive upwelling zone.

Finally, the authors stated in the material and methods section that migrant biomass was calculated as the difference between day and night biomass in the upper 90 m depth layer. Thus, biomass during night should be larger than during day in this layer, something which is not visually observed in Figure 6. Please, provide some explanation to this.

R: Because of re-assessment of biomasses, the estimates have changed and data are more consistent with normal DVM behaviour, although we noted that some taxa did show reverse patterns of migration as in other studies. However, time-space variation and sampling biases should also be considered. These issues are fully discussed in the Discussion.

Reviewer 2

To start, I can tell that the authors have made an effort to address my and the other reviewer's comments. The paper also has improved in terms of readability, and figures 1, 2 and 3 are much better now.

But there remain a few issues that prevent publication, and therefore I again need to ask for a thorough revision. I still consider this an interesting dataset, and I know how much work went into the image analysis.

The first point is again the biomass estimation. While the figure the authors provided in response to my previous comment looks convincing (then again, it is a log scale), the way they got to their values is absolutely NOT correct, and since a paper like this will be read a lot and

table S3 will be used by others, I absolutely insist that this is corrected. I still think it would be easiest to ditch all the four steps of ellipsoid biovolume calculation, wet weight and dry weight assumptions, and finally carbon conversions, and simply use the image area to carbon conversion regressions (sources for which are already halfway compiled in their own table S3). In table S3, conversion factors from dry weight to carbon are listed. There are several errors in this table. First, citing is sloppy, making it quite difficult to track down the errors (apart from the fact that the full references are missing). For example, Marcolin et al. (2015) is cited a lot, but their values mostly originate from Lehette & Hernandez-Leon 2009. More importantly, THESE VALUES ARE NOT THE FRACTION OF DRY WEIGHT THAT IS MADE UP OF CARBON!!!

Instead, these values are the parameter a (intercept) from the regression $DW = a * Image\ Area^b$. I suspect this mistake somehow happened because many values are near 0.4. However, one could have noted that a 99% carbon content in ostracods would mean they are basically charcoal.

R: We appreciate reviewer's comments and agree that we had miss-interpreted regression parameters. Following the directions from the reviewer, we have now re-estimated all biomass data using the work of Hernandez-Leon et al (2009)... Conversion factors have all been corrected as well according to factors provided by Kiorbe (2013) . Detailed information on used regressions and conversion factors is provided in Tables and Supplemental Material. Revised data of biomasses yield lower estimates of zooplankton C and their migrant biomass. This also substantially affected the estimates of C flux and C export, by reducing their estimates. However, results are consistent with our conclusions and comparable to other studies. New estimates do not affect major findings and C flux by DVM appears as significant in this highly productive upwelling zone, despite presence of an intense OMZ. We also stress the importance of taxonomic structure in determining migrant biomass and downward C export.

Second, the use of a log-axis on a stacked bar chart (figs 4 and 5) is not possible. Please use a linear axis. Actually I don't think the stacked bar chart are really necessary, the taxon-specific values are listed in the tables.

R: We agree and we have now removed the original Fig 4 (data in Table) and original Fig. 5 has been reconstructed in linear scale as suggested.

Third, visualizing the total distribution of biomass (Fig 6) should be kept, but this figure has some issues – first, the x-axis label is a rate. This also does not fit to the figure caption. I still think the most logical way to plot biomass distribution is as bar chart with variable widths on a linear depth axis. This way you can plot biomass concentration ($mg\ C\ m^{-3}$) as the height of the bars, and the bar area gives a good idea on the integral, which makes it nicer to compare depth strata with different thickness (and also relate to the hydrography in fig 2). I made a sketch to illustrate what I mean (it is just a sketch, so it lacks axis tick marks, error bars, and layer names). Left: one panel of Fig 6 now, right: suggestion.

R: Agree.. We have re-made this Figure using a bar with variable width depending on depth strata, using biomass (m³) as to compare strate. We have also added vertical profiles of oxygen in these graphs as to illustrate the actual conditions in terms of oxygenation.

Lastly, the results are still quite lengthy. The words "show" and "exhibit" are used both about fifteen times each. Try to streamline the writing a bit, using mostly active voice ("Copepod abundance was higher here than there" rather than "Copepods exhibited higher abundances here than there").

R: We have fully revised the text avoiding the excess of adjectives.

I do look forward to the final version of the paper.

1 Zooplankton diel vertical migration and downward C flux into the 2 Oxygen Minimum Zone in the highly productive upwelling region off 3 Northern Chile

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10 **Abstract.** The daily vertical movement of zooplankton, known as diel vertical migration (DVM), can enhance the vertical
11 flux of carbon (C) and so contribute to the functioning of the biological pump in the ocean. The magnitude and efficiency of
12 this active transport of C may depend on the size and taxonomic structure of the migrant zooplankton. However, the impact
13 that a variable community structure can have on zooplankton-mediated downward C flux has not been properly addressed.
14 This taxonomic effect may become critically important in highly productive eastern boundary upwelling systems (EBUS),
15 where high levels of zooplankton biomass are found ~~becomes aggregated~~ in the coastal zone and composed by a highly
16 diverse community with variable DVM behavior. In these systems, presence of a subsurface oxygen minimum zone (OMZ)
17 can impose an additional constraint to vertical migration and so influence the downward C export. Here, we address these
18 issues based on a high-resolution zooplankton sampling at three stations off northern Chile (20°S-30°S) during November-
19 December 2015. Automated analysis of zooplankton composition and taxa-structured biomass allowed us to estimate daily
20 migrant biomass by taxa and their amplitude of migration. We found that a higher biomass aggregates above the oxycline,
21 associated with more oxygenated surface waters and this was more evident upon a more intense OMZ. Some taxonomic
22 groups, however, were found closely associated with the OMZ. Most taxa were able to perform DVM in the upwelling zone
23 withstanding severe hypoxia. Also, strong migrants, such as Copepods Eucalanidae and Euphausiids, can exhibit a large
24 migration amplitude (~500 m), remaining either temporarily or permanently during the day or night condition within the core
25 of the OMZ and so contributing to the release of C below the thermocline. Our estimates of DVM-mediated C flux suggested
26 that a mean migrant biomass of ca. 958.34 g mg C m⁻² d⁻¹ may contribute with about 716.36 mg C m⁻² d⁻¹ to the OMZ system
27 through respiration, mortality, and C excretion production of fecal pellets at depth, accounting for ca. 425% of the net
28 primary production, and so implying the existence of an very efficient mechanism to incorporate freshly produced C into the
29 OMZ. This downward C flux mediated by zooplankton is however strongly variable in the space and mostly dependent on
30 the taxonomic structure due to variable migration amplitude and DVM behavior.

31 1 Introduction

32 The Oxygen Minimum Zone (OMZ) in the Southeast Pacific, the fourth largest of the six permanent hypoxic regions in the
33 world oceans (Paulmier *et al.*, 2006), is a key component of the water column and a permanent feature intruding the coastal
34 zone of Ecuador, Peru and Chile (Fuenzalida *et al.*, 2009; Paulmier *et al.*, 2006). In the highly productive upwelling region
35 of Northern Chile, the OMZ is closely linked to wind-driven upwelling in the coastal area and associated to the Equatorial
36 Subsurface Water (ESSW), which is transported southward along the continental shelf by the Peru-Chile Undercurrent
37 (PUC), as far south as 48°S (Fuenzalida *et al.*, 2009; Morales *et al.*, 1996a; Silva and Neshyba, 1979). Off Iquique (20°S) the
38 OMZ is characterized by being thick (500 m), very intense ($< 20 \mu\text{mol kg}^{-1}$) and with O_2 concentrations in the core of OMZ
39 among the lowest found in the global ocean reaching the current detection limit ($< 1 \mu\text{M}$) (Ulloa and Pantoja, 2009), although
40 it becomes thinner at about 30°S (Paulmier *et al.*, 2006).

41 During the last decades, the OMZ systems have attracted much scientific interest because of evidence showing that hypoxic
42 and anoxic conditions in coastal areas are expanding and becoming more intense (Ekau *et al.*, 2010; Stramma *et al.*, 2008).
43 At present, ongoing ocean deoxygenation is widely recognized as linked to global warming and it is rising much concern in
44 modern oceanography (Breitburg *et al.*, 2018).

45 The presence of oxygen-depleted water becomes a critical physiological constraint for pelagic and benthic organisms
46 inhabiting the upwelling zone, impacting their biomass and productivity, the species diversity, distribution, and behaviour
47 and metabolic activity (Wishner *et al.*, 2018; Ekau *et al.*, 2010; Grantham *et al.*, 2004). For instance, diel vertical migration
48 (DVM), a common feature of the various size groups of zooplankton and also one of the most important movements of
49 biomass in the ocean, can also be affected by changes in intensity and distribution of the OMZ (Wishner *et al.*, 2018, 2013;
50 Escribano *et al.*, 2009; Fernández-Álamo and Färber-Lorda, 2006; Hidalgo *et al.*, 2005; Morales *et al.*, 1996; Judkins, 1980).

51 The OMZ can act as an ecological barrier for vertical distribution of many organisms, constraining most zooplankton to a
52 narrow (50 m) upper layer, as shown in the coastal upwelling zone off Chile according to the works of Escribano (2006) and
53 Donoso and Escribano (2014). Zooplankton also become limited to the upper 150 or 300 m in the eastern tropical north
54 Pacific (Wishner *et al.*, 2013). However, the OMZ can also offer refuge for species adapted to live there, creating
55 microhabitats of differing oxygen concentration that are characterized by layers of high zooplankton biomass and abundance,
56 with distinct species zonation (Antezana, 2009; Wishner *et al.*, 2008; Fernández-Álamo and Färber-Lorda, 2006), which, in
57 turn, may have important consequences for carbon (C) cycling and its vertical flux. For example, it is known that
58 zooplankton in the coastal upwelling region off northern Chile may play a significant biogeochemical role by promoting
59 carbon flux into the subsurface OMZ (Escribano *et al.*, 2009). Therefore a significant proportion of the vertical material flux
60 from the euphotic zone to the deep sea (> 200 m) and within the food chain could be determined by DVM of zooplankton
61 (Longhurst and Williams, 1992; Steinberg and Landry, 2017).

62 As important contributors to the functioning of the biological pump, diel zooplankton migrants can actively increase the
63 magnitude of C export by transporting surface-ingested material in their guts to deep waters where it can be metabolized

64 (Steinberg and Landry, 2017). Zooplankton moults or mortality at depth can also contribute to the transportation of
65 assimilated organic biomass into the deep waters (Ducklow *et al.*, 2001). The biological pump process is also thought to be
66 related to the size structure of dominant zooplankton. This because some groups with large body sizes may exhibit a greater
67 range of vertical migration and sometimes higher levels of biomass, and so influencing the biogeochemical fluxes (Dai *et al.*,
68 2016; Ducklow *et al.*, 2001). However, the effect of variable size structure on DVM performance and its consequence for
69 active C transport has not indeed been assessed. Size-structure is certainly related to zooplankton composition which has
70 hardly been properly addressed when examining the role of DVM on C flux. For instance, in areas with hypoxic subsurface
71 layers some species are more active migrants and thus more efficient C-transporters, because they have developed
72 adaptations to low oxygen conditions and can even use the OMZ as their habitat, either temporarily or permanently
73 (Escribano *et al.*, 2009; Gonzalez and Quiñones, 2002; Seibel, 2011). Adaptation may include in some cases reduction of
74 aerobic metabolism by more than 50% during exposure to hypoxia conditions, as a mechanism to facilitate low oxygen
75 tolerance, an so reducing dramatically energy expenditure during daytime within low oxygen waters, and therefore affecting
76 the carbon C flux in areas subjected to low concentrations of oxygen (Seibel *et al.*, 2016).

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78 C-export to depth may also depend on the amount of biomass being produced in the photic zone. Primary production
79 promotes zooplankton feeding and growth and therefore determining C availability for both passive and active transport to
80 depth. In this context, highly productive upwelling zones can be assumed as systems where the C-flux mediated by
81 zooplankton DVM can be enhanced, although certainly depending on the size and taxonomic structure of zooplankton. In
82 these regions, a shallow OMZ might exert a further impact on the C-flux by affecting DVM or zooplankton metabolism at
83 depth. In the present study, based on high resolution sampling and automated analysis of mesozooplankton, we assessed
84 zooplankton vertical migration and downward C to the OMZ in the highly productive upwelling region of Northern Chile.
85 We aimed at understanding the role that taxonomic and size-structure can play in the magnitude and variability of the DVM
86 behaviour interacting with a shallow OMZ, and the implications this interaction can have on the magnitude of the downward
87 C flux in a highly productive coastal upwelling zone

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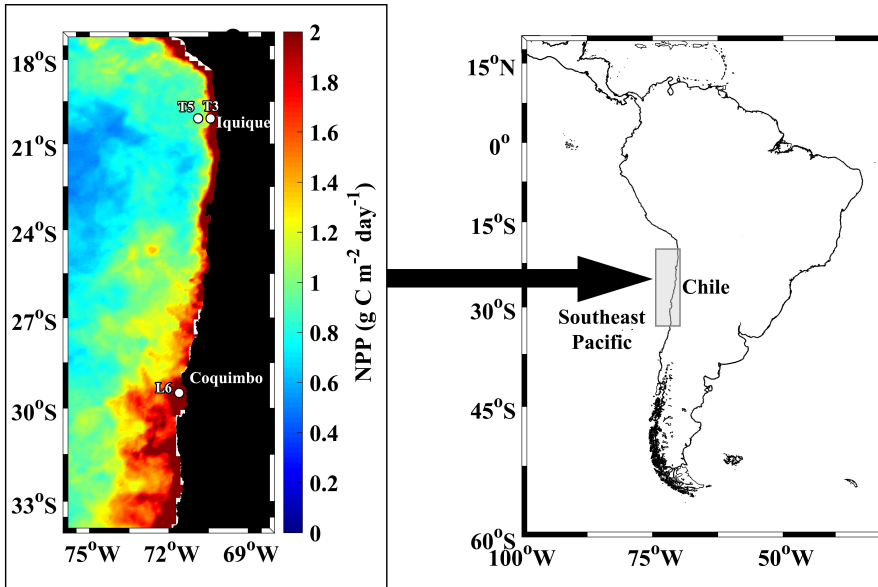
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88 2Methods

89 2.1Study Area

90 The study area was located in the Southeast Pacific Ocean and covered the coastal zone of the northern upwelling region of
91 Chile (21-29°S) (Fig. 1), which is a region known to be subjected to wind-driven upwelling throughout the year and
92 containing an intense and shallow OMZ (Ulloa *et al.*, 2012). The sampling design comprised three stations: Two stations (St.
93 T3 and St. T5) across a zonal section off Iquique (20°S) and a coastal station (St. L6) off Coquimbo (29°S). The study was
94 carried out during the LowpHox cruise conducted in November-December 2015 onboard the R/V *Cabo de Hornos*. At each
95 station, temperature, salinity and dissolved oxygen (DO) were recorded from 1000 m (St.T5 and St.T3) and only 356 m (St.

96 L6) using an oceanographic rosette with a CTD SeaBird 911 (SBE 911 plus) equipped with a SeaBird SBE-43 oxygen sensor
97 and a Seatech Fluorometer. Discrete water samples were also obtained for chemical measurements of oxygen.
98



99
100 **Fig.1.** Study area at the northern upwelling region of Chile, indicating sampling stations during the LowpHox cruise
101 conducted in November-December 2015 onboard the R/V *Cabo de Hornos*. Satellite estimated Net Primary Production
102 (NPP), averaged for November-December 2015 is shown. NPP was obtained from
103 <http://science.oregonstate.edu.ocean.productivity>.

104 2.2 Zooplankton sampling

105 Zooplankton samples were collected during daytime and night-time conditions in two consecutive days at the three stations
106 off northern Chile (T5-T3-L6) (Fig. 1), also as indicated in Table S1 (Supplemental Material). Vertical hauls of
107 zooplankton hauls were performed from 600 to 0 m depth with a Multi Plankton Sampler Hydro-Bios Multinet system with
108 a 0.25 m² opening area and equipped with 200- μ m mesh-size nets. The Multinet towing speed was 1 m s⁻¹ and the flowmeter
109 was in the mouth of the Multinet to estimate the volume of filtered water. Once onboard the collected zooplankton samples
110 were preserved immediately in 5% buffered formalin-seawater solution. At each station, 40 discrete samples were obtained
111 from 600 to 0 m depth. Each sample corresponded to a different depth stratum (30-0, 90-30, 150-90, 400-150 and 600-400m

112 depth). These strata were defined in according to distribution of oxygen concentration and localization of the OMZ (**Fig. 2**).
 113 Then, from the vertical profiles of oxygen and coinciding with the sampled layers of the Multinet, strata were defined as: A
 114 well oxygenated stratum with oxygen approaching air saturation (>250 $\mu\text{mol O}_2 \text{ kg}^{-1}$) was defined as the oxic mixed layer
 115 (OX-ML), an upper O_2 gradient (oxycline), defined by the level where O_2 reaches 4% of the surface O_2 (Paulmier *et al.*,
 116 2006), at whose base is located the upper boundary of the OMZ (45 $\mu\text{mol O}_2$ isoline, OMZ-UB) (Escribano *et al.*, 2009;
 117 Hidalgo *et al.*, 2005; Morales *et al.*, 1999); the OMZ core defined by an upper boundary (OMZ-UC) with the lowest
 118 concentration of O_2 (<20 $\mu\text{mol O}_2 \text{ kg}^{-1}$) and a lower boundary (OMZ-LC) (1 to <20 $\mu\text{mol O}_2 \text{ kg}^{-1}$), and finally a lower O_2
 119 gradient (OMZ-LW). Depth ranges and oxygen levels for these strata are detailed in **Table S2 (Supplemental Material)**.
 120 ~~Once onboard the collected zooplankton samples were preserved immediately in 5% buffered formalin seawater solution.~~
 121

122 2.3 Taxonomic and size measurements

123 Taxonomic identification and enumeration of taxa were carried out by analysis of digitized images obtained with the
 124 Hydroptic ZooScan digital imaging system (Gorsky *et al.*, 2010). Each sample was wet-sieved through a 1000 μm mesh into
 125 two size fractions. Then, each size-fraction was fractionated again separately with a Motoda splitter until the zooplankton
 126 concentration was sufficiently diluted to avoid contact between organisms in the ZooScan scanning frame. Fractioning into
 127 small and large organisms and consequent separate image acquisition of the two size classes prevented underestimates of
 128 large-rare objects which may need less fractioning (Gorsky *et al.*, 2010). 179 subsamples were scanned and digitized at a
 129 resolution of 2400 dpi after manual separation of objects on the scanning tray. After processing the samples with ZooProcess
 130 software, each of the objects was automatically sorted with the help of a learning set, and then the sorting was visually
 131 validated by an expert (for details, see Chang *et al.*, 2012; Gorsky *et al.*, 2010). Organisms making up the ZooScan datasets
 132 were enumerated, measured, biomass-estimated and classified into 27 taxonomic groups, such as Copepods, Chaetognaths,
 133 Euphausiid, Gelatinous and other zooplankton. The abundance (ind. m^{-3}) ~~and stratum-integrated abundance (ind. m^{-2})~~ of
 134 total zooplankton or of each taxonomic group was calculated following Eq. (1):

$$135 \text{Abundance/m}^3 = \frac{N \cdot \text{Subpart}}{\text{Vol} \cdot \text{Townb}} \quad (1)$$

~~$$136 \text{Abundance/m}^3 = \frac{N \cdot \text{Subpart}}{\text{Vol} \cdot \text{Townb}}$$~~

137
 138 were, N is the number of individuals with same prediction (e.g., in last column written “copepod”), Subpart is the splitting
 139 ratio, Vol is net volume and Townb is the number of net tows in a sample. Townb = 1 in our case because of a single tow per
 140 sample. Stratum-integrated abundance (ind. m^{-2}) was obtained after multiplying by width (m) of a given stratum.
 141

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142 **2.4 Biomass estimates and Carbon fluxes**

143 The ZooScan Integrated System also provided zooplankton body size in terms of area (mm²) or volume (mm³) for each
144 organism. We used these organisms' area or volume to estimate dry weight of each individual of different taxonomic groups
145 using published regression equations relating organism size, area or volume to individual weight as detailed in Table S3
146 (Supplemental Material). Mass unit conversions between dry weight (DW) and carbon content (C) were performed using
147 averaged conversion factors obtained for different mixed-zooplankton groups (Kjørboe, 2013) and ichthyoplankton
148 (Childress and Nygaard, 1973) (Table S3 Supplemental Material). Added biomasses (µg C indiv.⁻¹) of individuals within
149 taxonomic categories identified by ZooScan allowed us to estimate total biomass per taxa (mg C m⁻³) for each sample by
150 station, daytime vs night-time condition, and depth strata. Integrated values of biomass per depth strata (mg C m⁻²) and taxa
151 were calculated multiplying by strata width (m).

152 The ellipsoidal biovolume (EBv) was computed instead of equivalent spherical diameter (ESD). Biovolume was used
153 because most of the zooplankton shapes were elongated, and also because ESD overestimates biovolume since spheres have
154 a higher volume: cross-sectional area ratio than other shapes (García-Comas *et al.*, 2014; Herman and Harvey, 2006; Sprules
155 *et al.*, 1998). Ellipsoidal biovolume EBv (mm³) was calculated following Eq. (2):

156
157 _____, (2)

158
159 where, the EBv corresponds to the biovolume of an ellipsoid, Major is the longest axis of the object and Minor is the minor
160 axis of a perfect ellipse of the same area of the measured object, both axes are in mm and were provided by ZooScan. We
161 used EBv (mm³) for estimating biomass (mgC) for all taxonomic groups. Conversions from the EBv to carbon were made in
162 three steps: First EBv was converted into wet weight (WW), assuming 1 mm³ was equivalent to 1 mg WW (Wiebe *et al.*,
163 1975). Then WW was converted to dry weight (DW) assuming that the water content of zooplankton in average was 90%
164 (DM = 0.1 × WW), as suggested by Yamaguchi *et al.* (2014) and Sato *et al.* (2015). Finally, individual biomass was converted
165 to carbon using a factor carbon for each taxonomic group from previous published data Table S3 (Supplemental Material).

166 To calculate the migrant biomass, we integrated biomass in the upper 90 m layer from our two sampled strata 0-30 m and
167 30-60 m. This 0-90 m stratum was considered the approximate above-oxycline layer after examining the vertical profiles of
168 oxygen. Biomass at night was thus subtracted from the corresponding day biomass in this layer to assess daily changes
169 involving migrants as in Putzeys *et al.* (2011). Thus, the negative values of the day-minus-night biomass corresponded to
170 migrant biomass that reached the epipelagic layer at night including organisms inhabiting above and below the oxycline.
171 The proportion of migrant biomass with respect to observed biomass in the upper 90 m of a given taxonomic group was
172 defined at the rate of migration on a daily basis. This rate of migration could thus be used as an index of DMV behaviour for
173 a taxonomic group.

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174 To estimate the active C flux at each sampling station, we considered three processes contributing to C at depth mediated by
 175 migrant zooplankton: respiration (R), excretion (E) and mortality (M) at depth. Respiration at depth (>90 m) was estimated
 176 using the equation provided by Ikeda (1985) which relates individual respiration rate with body mass and temperature,
 177 independently of taxonomic category which may have a minor effect on R in according to a more recent assessment (Ikeda,
 178 2014). Mean body mass ($\mu\text{g C}$) for each taxonomic group from ZooScan estimates and mean in situ temperature were used
 179 to obtain integrated R at each depth strata. Estimates of R for each taxonomic group are shown in Table S6 (Supplemental
 180 Material). Integrated R per station for the 0-600 depth strata along with the corresponding integrated biomass was then used
 181 to estimate the fraction of C being respired at depth by zooplankton. The contribution of C by excretion (E) at depth was
 182 assumed to be 31% of R as suggested by (Steinberg *et al.*, 2000) and daily mortality at depth (M) was considered to be in
 183 the range of 0.3 and 0.5 as suggested by Edvardsen *et al.* (2002), so that a median value of 0.4 as a fraction of migrant
 184 biomass was assumed. Vertical C flux mediated by zooplankton was thus estimated as,

$$186 \text{ C Flux} = \text{MB} \times [(\text{R} + \text{M} + 0.31\text{R})/2] \quad (2)$$

188 used a daily respiration fraction of 0.12. This proportion of biomass being respired was estimated by Hernández-León and
 189 Ikeda (2005) for total zooplankton biomass at a temperature ranging between 13°C and 18°C for mid-latitude areas. We used
 190 this value considering a similar temperature range between surface water and the core of the OMZ (Fig. 2). The daily
 191 contribution of egestion rate to C flux was assumed to be 0.09 of migrant biomass, as suggested by Eseribano *et al.* (2009).
 192 This estimate was derived from a combined biomass of large-sized copepods and euphausiids which are the major groups
 193 contributing to migrant biomass in this upwelling area. This value also considers that most egested faecal pellets may be
 194 produced rapidly in the upper layer upon feeding (Eseribano *et al.*, 2009). The contribution of mortality to C flux was
 195 assumed to be 0.08 d^{-1} of migrant biomass, as a conservative estimate suggested by Ohman and Wood (1996). We thus
 196 estimated total C flux following Eq. (32):

$$198 \text{ C Flux} = \text{MB} \left(\frac{\text{R} + \text{M} + \text{E}}{2} \right) \quad (32)$$

200 where MB is the migrant biomass (mg C m^{-2}), R and M are daily Respiration and Mortality, and E is excretion
 201 expressed as 0.31R. The three processes are divided by 2 assuming a 12 h incursion at depth. We thus estimated total flux as:

$$204 \text{ C Flux} = (0.12 * (\text{MB}) + 0.08 * (\text{MB})) / 2 + 0.09 * (\text{MB}) \quad (43)$$

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205 2.5 Patterns of vertical distribution of migrating zooplankton

206 For the analysis of vertical distribution of organisms, the density estimates of the organisms were standardized to number of
207 individuals per m³ (for each stratum) or per m³ (for integrated values). In order to quantify the presence and extent of DVM
208 of various taxa at each station, we calculated ~~both night: day (N:D) ratios in each stratum and~~ weighted mean depth (WMD)
209 for zooplankton abundance, as a measure of the center of gravity of a population's vertical distribution for each taxon and
210 haul, ~~in~~ according to Andersen *et al.* (2004) following Eq. (354):

$$212 \quad WMD = \frac{\sum (ni \cdot zi \cdot di)}{\sum (ni \cdot zi)}, \quad (345)$$

$$213 \quad WMD = \frac{\sum (ni \cdot zi \cdot di)}{\sum (ni \cdot zi)}$$

214 where d is the mean depth of the strata (m), z the width (m) of the strata and ni the abundance (ind. m⁻³) of a given i
215 taxonomic group.

216 We calculated the amplitude of vertical migration (Δ DVM) as the difference between the WMD of the organisms during the
217 day and the night and therefore this Δ DVM was considered as the criterion to assess the DVM behaviour for each taxonomic
218 group. Positive values indicated normal DVM (pattern of nocturnal ascent by individuals that reside at depth by day) and
219 negative values indicated reverse DVM (pattern of nocturnal descent by individuals that reside near the surface by day). The
220 individuals that occupied the same depth stratum by day and by night, whether near the surface or at depth were considered
221 as non migrant in according to Ohman (1990).

222 2.6 Statistical analysis

223 For statistical analysis, as a criterion for determining if the DVM was significant, we tested for differences in the WMD
224 mean between day and night using a two-tailed t test. We considered the occurrence of DVM when the difference in the
225 WMD mean between day and night was significant ($p < 0.05$). In order to evaluate the similarity/dissimilarity in the
226 abundance and biomass among stations, strata, and day-night conditions, the multivariate grouping techniques were applied
227 ("cluster analysis"), ANOSIM (Two-Way Crossed Analysis) tests and multidimensional scaling (MDS) with the data
228 transformed in PRIMER v 6.1.16 (2013), prior to the application of the Bray-Curtis similarity index (Bray and Curtis, 1957).
229 In general, WMD for taxonomic groups did not exhibit a pronounced bimodal vertical distribution.

230 3 Results

231 3.1 Hydrographic conditions

232 ~~Vertical profiles of temperature, DO, and sigma-t show general oceanographic conditions in all stations (Fig. 2), such as the~~
233 ~~vertical distribution of oxygen and location of the OMZ. From the vertical profiles of oxygen and coinciding with the~~

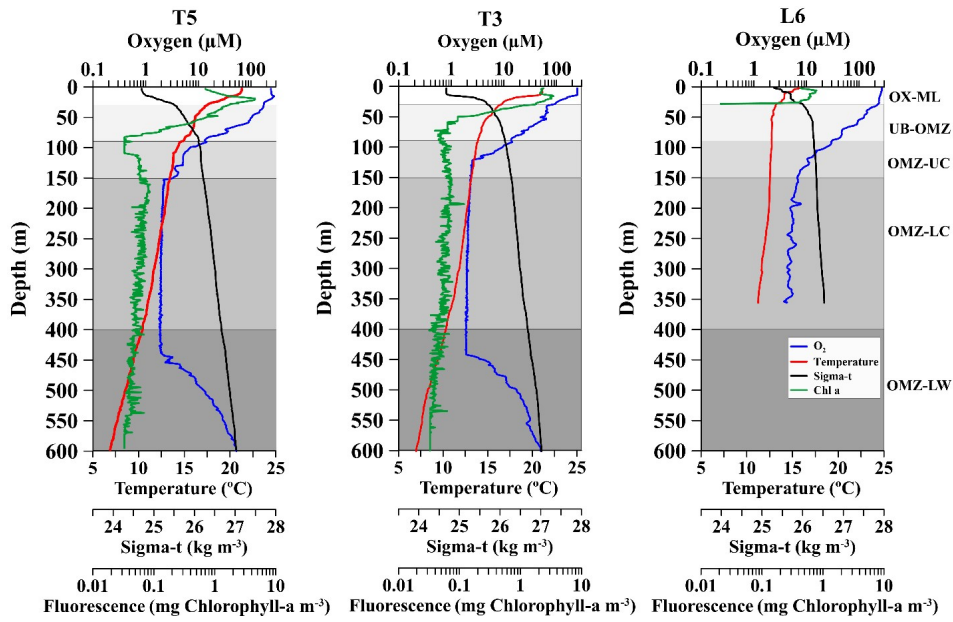
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234 ~~sampled layers of the Multinet different strata were defined: A well oxygenated stratum with oxygen approaching air~~
235 ~~saturation ($>250 \mu\text{mol O}_2 \text{ kg}^{-1}$) defined as oxie mixed layer (OX-ML), an upper O_2 gradient (oxycline), defined by the level~~
236 ~~where O_2 reaches 4% of the surface O_2 (Paulmier *et al.*, 2006), at whose base is located the upper boundary of the OMZ (45~~
237 ~~$\mu\text{mol O}_2$ isoline, OMZ-UB) (Escribano *et al.*, 2009; Hidalgo *et al.*, 2005; Morales *et al.*, 1999); the OMZ core defined by~~
238 ~~an upper boundary (OMZ-UC) with the lowest concentration of O_2 ($<20 \mu\text{mol O}_2 \text{ kg}^{-1}$) and a lower boundary (OMZ-LC) (1~~
239 ~~to $<20 \mu\text{mol O}_2 \text{ kg}^{-1}$), and finally a lower O_2 gradient (OMZ-LW).~~

240 Across the zonal section off Iquique the offshore station (St. T5) and onshore station (St. T3) showed two contrasting
241 hydrographic regimes regarding the OMZ. Station T5 had a less pronounced and thicker OMZ than station T3. At both
242 stations the 5 strata were well defined in the water column (**Fig. 2**). The OX-ML ($>250 \mu\text{mol O}_2 \text{ kg}^{-1}$) was present at 18 m
243 (St. T5) and 15 m (St. T3). The oxycline ~~exhibited a gradual~~ly decrease~~d~~ from oxie ($\sim 250 \mu\text{mol O}_2 \text{ kg}^{-1}$) to suboxic (<20
244 $\mu\text{mol O}_2 \text{ kg}^{-1}$) conditions associated with a strong stratification in the upper 80 m depth. The 45 $\mu\text{mol O}_2$ isoline
245 (OMZ -UB) was ~~present~~ at the base of the oxycline at 70 m (St. T5) and 59 m (St. T3). The OMZ core ($<20 \mu\text{mol O}_2$
246 kg^{-1}) was ~~located~~ below the thermocline and below the 26.5 kg m^{-3} isopycnal following description of Paulmier *et al.* (2006). In
247 the oceanic station (St. T5) the OMZ core was ~~located~~ between 80 to 514 m, while in the coastal station (St. T3) it was
248 between 80 to 507 m with 423 m thickness. The O_2 concentration in the **OMZ** core was ca. 1 $\mu\text{mol O}_2 \text{ kg}^{-1}$. The OMZ-LW at
249 both stations was delimited above the core and below the depth where the O_2 slope changed significantly (slope break >20
250 $\mu\text{mol /m}$) (**Fig. 2**).

251

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252
253

254 **Fig. 2.** Vertical profiles of dissolved oxygen (O_2), temperature, Chlorophyll-a and water density (sigma-t), at 3 stations off
 255 northern Chile (see Fig. 1) during the LowpHox Cruise in the austral spring 2015. Shaded areas represent different layers
 256 defined according to oxygen concentration. OX-M= Oxic mixed layer, OMZ-UB= Upper boundary of the oxygen minimum
 257 zone, OMZ-UC= Upper core of the oxygen minimum zone, OMZ-LC= Lower core of the OMZ and OMZ-LW=lower
 258 oxygen gradient.

259 The structure of the OMZ at the coastal station (St.L6) off Coquimbo ($29^\circ S$) (Fig.1), exhibited was similar characteristics to
 260 St. T3 ($21^\circ S$), but in this area the OMZ was deeper and thinner. The OX-ML was shallower. The OMZ -UB ($45 \mu mol O_2$) in
 261 the base of the oxycline was down to 80 m. The low O_2 concentrations in the core were less intense than at $21^\circ S$ (4 to 20
 262 $\mu mol O_2 kg^{-1}$) and it was located below 100 m (Fig. 2). The OMZ-LW could not be assessed because of lack of CTD data
 263 below 350 m.

264 Additional oceanographic variables showed a surface warming ($>20^\circ C$) and strongly stratified conditions at the three
 265 stations with a sharp thermocline in the upper 100 m, coinciding with the oxycline, whereas Chlorophyll-a maximum ($>5 mg$
 266 Chlorophyll-a m^{-3}) was located in the upper 20 m (Fig. 2).

267 3.2 Zooplankton composition and abundance

268 A total of 27 zooplankton taxa were identified by the ZooScan and ZooProcess (Table S4 Supplemental Material). The
269 number of taxa varied among stations and strata. Across the zonal section off Iquique the number of taxonomic groups
270 fluctuated between 23 (St. T3) and 26 (St. T5), whereas 25 taxa were present off 29°S (St. L6). The most dominant taxa both
271 daytime and night conditions, were: Copepods 87% (in St. T5), 79% (in St. T3) and 69% (in St. L6). This group was
272 constituted by small Copepods, large Copepods, the Copepods Eucalanidae and the Copepods *Acartia* spp; fish eggs
273 constituting 2% (in St. T5), 5% (in St. T3) and 6% (in St. L6); Nauplii being <1% (in St. T5), <1% (in St. T3) and 7% (in St.
274 L6); Appendicularia 5% (in St. T5), 4% (in St. T3) and 3% (in St. L6) (Table S5 Supplemental Material). The remaining
275 19 pooled groups only constituted <6% (in St. T5), 11% (in St. T3) and 15% (in St. L6). The total integrated abundances of
276 zooplankton (0-600 m) by sampling station are shown in Table S4 (Supplemental Material). As based on a Two-Way Crossed
277 Analysis ANOSIM test, this water-column integrated abundance did not show significant differences between day and night
278 samples ($p > 0.05$). However, the abundance of these zooplankton groups regarding stations was significantly
279 different among stations (Two-Way Crossed Analysis ANOSIM $p < 0.05$) were found, so that the stations were treated
280 independently. Off Iquique the abundance showed the lowest values at the onshore station (St. T3 with 18%
281 lower), which was characterized by the strongest and most extensive OMZ in the study area. These values increased at the
282 offshore station (St. T5 with 31% greater), where the OMZ was less pronounced and thicker. Unlike stations T3 and T5, the
283 onshore station off 29°S (St. L6) had a weaker and less extensive OMZ showing the highest zooplankton abundance (51%
284 greater).

285

286 3.3 Diel vertical migration (DVM) and vertical distribution

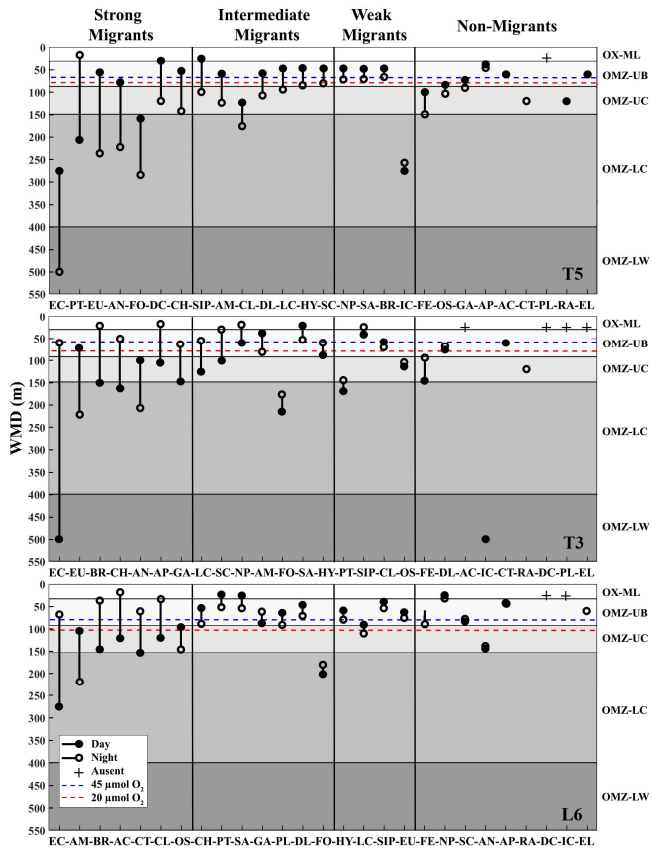
287 3.3.1 Main migrant groups of zooplankton

288 The diel vertical migration of 27 zooplankton taxa in the 0–600 m water column is shown in Fig. 3. These taxa were
289 classified into four groups according to their amplitude of migration (Δ DVM) (Table 1): 1) Strong migrants, represented
290 by taxa with that exhibited a strong DVM and showed a broad range of Δ DVM from 225 to 99 m (in St. T5), 440 to 84 m (in
291 St. T3) and 208 to 87 m (in St. L6). This group constituted 70% of taxa with higher Δ DVM. The composition of taxa in this
292 group was variable at each station (Table 1), but in general was well represented by Eucalanidae Copepods (EC),
293 Euphausiids (EU), *Acartia* Copepods (AC), Ctenophores (CT), Decapods (DC), Annelids (AN), Bryozoan (BR),
294 Pteropods (PT) and Chaetognaths (CH). These taxa were mostly concentrated in the oxic surface stratum (OX-ML), and the
295 OMZ core showing a strong interaction with both the OMZ-UC and the OMZ-LC, and so changing from normoxia to
296 hypoxia condition and vice versa between 0 to 550 m (Fig. 3), 2) Intermediate migrants, represented by taxa with that
297 exhibited a moderate DVM and showed a range of Δ DVM from 73 to 34 m (in St. T5), 70 to 27 m (in St. T3) and 49 to 22 m
298 (in St. L6). This group constituted 23% of taxa with moderate Δ DVM. The composition of taxa in this group was also

299 variable at each station (**Table 1**), but it was mostly represented by small (SC) and large Copepods (LC), Amphipods (AM),
300 Cirripedia Larvae (CL), Gastropods (GA), Siphonophores (SIP) and Appendicularia (AP). These taxa were mostly
301 concentrated in the oxic surface strata (**OX-ML**) and in the **OMZ-UC** showing some interaction with the **OMZ core** and
302 vertically changing from normoxia to hypoxia condition, and vice versa between 0 to 200 m, **3) Weak migrants**, represented
303 by taxa that **with exhibited** a weak DVM and **showed** a range of Δ DVM of 24 to 18 m (in St.T5), 23 to 12 m (in St.T3) and
304 21 to 11 m (in St. L6). This group constituted 5% of taxa of low range of Δ DVM. The composition of taxa in this group was
305 also variable at each station (**Table 1**), but in general it was represented by Hydrozoa (HY), Salps (SA), Platyhelminthes
306 (PT), Decapoda Larvae (DL), Ostracods (OS), Nauplii (NL) and Ichthyoplankton (IC). These taxa were concentrated mainly
307 in the oxic surface strata (**OX-ML**) and in the **OMZ-UP**, but also in the **OMZ-UC** at the onshore stations (Station_T3 and
308 Station L6), showing much less interaction with the **OMZ core**, while spatially moving from normoxia to hypoxia condition
309 and vice versa between 0 to 100 m, and **4) Non-migrants**, represented by taxa which did not exhibit a significant DVM and
310 had a range of Δ DVM from 16 to 0 m (in St. T5), 7 to 0 m (in St. T3) and 6 to 0 m (in St. L6). This group constituted 1% of
311 taxa with not significant Δ DVM. The composition of taxa in this group was also variable at each station (**Table 1**), but in
312 general it was represented by fish eggs (FE), Radiolarian (RA) and Echinoderm larvae (EL).

313

314



315

316 **Fig.3.** Weighed mean depth distribution (WMD) of the zooplankton community interacting with the OMZ off Iquique
 317 (Stations T5 and T3) and off Coquimbo (Station L6) at the northern upwelling area of Chile during the austral spring 2015.
 318 Shaded gray areas represent different layers defined by their oxygen levels (defined in Methods). The taxonomic groups
 319 were classified by automated analysis (ZooScan): EC= Eucalanidae Copepods, AM= amphipods, BR=Bryzoa *larvaen*, AC=
 320 *Acartia* Copepods, CT= Ctenophores, CL=Cirripedia larvae, OS=Ostracods, CH=Chaetognaths, PT= Pteropods, SA=salps,
 321 GA= Gastropods, PL= Platyhelminthes, DL= Decapoda larvae, FO=Foraminifera HY=Hydrozoa, LC=Large Copepods,
 322 SIP=Siphonophores, EU=Euphausiids, FE=fish eggs, NP=Nauplii, SC=Small copepods, AN= Annelid *sae*,
 323 AP=Appendicularian, RA=Radiolarian, DC=Decapods, IC=ichthyoplankton, EL=Echinoderm larvae.

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324

325 **Table 1-** Diel vertical migration indices for 27 taxonomic groups (TAXA) identified and sorted by ZooScan at 3 stations off
 326 northern Chile (see Fig. 1 for acronyms), during the austral spring 2015. Amplitude of migration (Δ DVM) is in meters.
 327 Positive values indicate normal_DVM and negative values indicate reverse DVM (see Methods). ~~N:Dis-Night: day-ratio-of~~
 328 ~~abundane~~. Four groups are defined in according to DVM behaviour. Relative abundances are shown in %.

T5			T3			L6			MIGRANTS
TAXA	Δ DVM	%	TAXA	Δ DVM	%	TAXA	Δ DVM	%	
EC.	225	14	EC.	-440	28	EC.	-208	20	STRONG MIGRANTS > 5%
PT.	-188	12	EU.	149	10	AM.	115	11	
EU.	181	11	BR.	-129	8	BR.	-107	10	
AN.	145	9	CH.	-114	7	AC.	-103	10	
FO.	126	8	AN.	105	7	CT.	-94	9	
DC.	90	6	AP.	-88	6	CL.	-87	8	
CH.	88	6	GA.	-84	5	OS.	49	5	
SIP.	73	5	LC.	-70	5	CH.	35	3	INTERMEDIATE MIGRANTS > 2%
AM.	64	4	SC	-70	5	PT.	29	3	
CL.	51	3	NP.	-42	3	SA.	29	3	
DL.	50	3	AM.	42	3	GA.	-27	3	
LC.	47	3	FO	-38	2	PL.	26	2	
HY.	39	2	SA.	32	2	DL.	22	2	
SC	34	2	HY.	-27	2	FO	-22	2	
NP.	24	2	PT.	-23	1	HY.	21	2	WEAK MIGRANTS > 1%
SA.	23	1	SIP.	-15	1	LC.	19	2	
BR.	18	1	CL.	12	1	SIP.	15	1	
IC.	-18	1	OS.	-9	1	EU.	11	1	
FE	49	3	FE	-52	3	FE	25	2	NON- MIGRANTS <1%
OS.	16	1	DL.	-7	0	NP.	6	1	
GA.	15	1	AC.	-	-	SC	-4	0	
AP.	7	0	IC.	-	-	AN.	-4	0	
AC.	-	-	CT.	-	-	AP.	2	0	
CT.	-	-	RA.	-	-	RA.	-	-	
PL.	-	-	DC.	-	-	DC.	-	-	
RA.	-	-	PL.	-	-	IC.	-	-	
EL.	-	-	EL.	-	-	EL.	-	-	

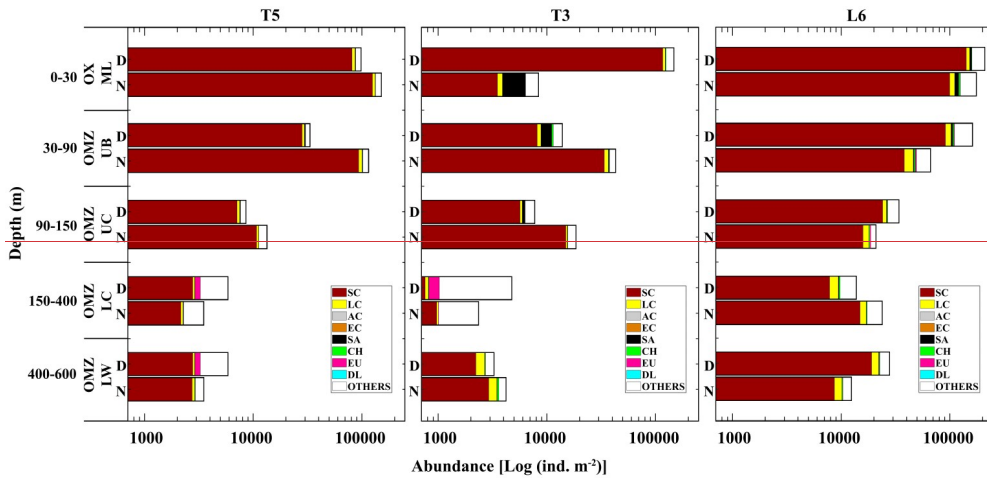
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330 **3.2 Vertical distribution and DVM of dominant groups.**

331 Vertical distributions of zooplankton were assessed for 5 taxonomic groups, which represented 80% of total abundance in
 332 average: Saps (SA-); Copepods represented by small Copepods (SC-), Large Copepods (LC-), Eucalanidae Copepods (EC-)
 333 and Acartia Copepods (AC-); Euphausiids; Decapods Larvae; Chaetognaths and Annelids, Chaetognaths(CH);Euphausiids
 334 (EU) and Decapods Larvae (DL); as well as their patterns of strata-station-abundance relationships are detailed illustrated in
 335 Table S4 (Supplemental Material), Fig. 4. The abundance of these zooplankton groups regarding depth strata was
 336 significantly different (ANOSIM, $p < 0.05$) at each station, and therefore representing distinctive microhabitats characterized
 337 by specific depth and oxygen concentration. In general, the higher abundance (>50%) was found in the shallower strata
 338 and well oxygenated layers (OX-ML and OMZ UB) ($>250 \mu\text{mol O}_2 \text{ kg}^{-1}$), and then it decreased rapidly in the strata
 339 associated with the OMZ core (OMZ-UC and OMZ-LC). Below this stratum a second slight peak in abundance was
 340 present in the OMZ-LW in special at Sts. T5 and L6 all stations, occurring between 400-600 m, both daytime and night
 341 conditions (Fig.4).

342



343

344 **Fig.4.** Vertical distribution of dominant taxonomic groups of zooplankton sampled at 3 stations off northern Chile: off Iquique
 345 at 21° S (Stations T5 and T3) and off Coquimbo at 29° S (Station L6). Abundances are mean values from samples obtained
 346 during 2 consecutive days under night and day conditions in the austral spring 2015. Sampled layers represent different
 347 conditions in terms of oxygen levels (see Methods) and the taxonomic groups and their acronyms are detailed in
 348 Supplemental Material Table S3.

349 As expected, Copepods dominated numerically the zooplankton community both within and outside the OMZ (Fig-4). Small
350 Copepods (SC) were the most abundant (70%), followed by large Copepods (LC) (6%), whereas the Copepods *Acartia* (AC)
351 and Eucalanidae (EC) showed the lowest abundances among Copepods. The largest aggregation of Copepods (pooled data)
352 altogether during the entire study period was observed at the offshore station St. T5 (87%), where abundances reached
353 192088 ind. m⁻². At the onshore station (St. T3) the percentage of contribution of copepods was 79% and 69% at the St L6
354 (Table S5 Supplemental Material). Off Iquique, the highest abundances and biomass were in the shallower strata (OX-
355 ML) at St.T5 (46%) and at St. T3 (47%), and these values were reduced in the core of the OMZ at St. T5 (4 to 1%) and at T3
356 (8 to 1%) between 90-400 m, where oxygen had the lowest concentrations (< 20 μM to 1 μM). At the St. T5 the second peak
357 of abundances showed slight increases was in in the OMZ-LW stratum during daytime condition, where oxygen levels
358 increased after the extremely low levels within the OMZ, while at the onshore station St T3 it was much dimmer and during
359 high condition. At the onshore station off 29° S (St. L6), having a weaker and less extensive OMZ, the vertical distribution of
360 abundance showed was similar vertical distribution. However, the abundance of copepods was lowest in this station (in about
361 69%) in comparison with stations off Iquique, in the core of the OMZ the percentage was between 5 to 3%.

362 DVM of Copepods was pronounced at onshore stations (Stations T3 and L6), but the strength of migration was higher overall
363 at St. T3 off Iquique, as reflected by the migration indices (WMD and ΔDVM) (Table 1). The WMD of these taxa showed
364 had a broad range (17-500 m), which varied significantly among copepods groups and stations, both in day and night
365 samples (p < 0.05) (Fig. 3). During the night, at the offshore station (St. T5) most Copepods exhibited normal DVM and they
366 were concentrated mainly in the oxic surface strata (OX-M) and OMZ-UB (40-60 m) without interacting with the OMZ;
367 except for Eucalanidae which concentrated deeper in the OMZ-LC stratum, associated with the lower core of the OMZ and
368 showing a high ΔDVM (225m). During the day these four groups of Copepods tended to remain deeper in the stratum
369 associated with the lower core of the OMZ (OMZ-UC) and lower O₂ gradient (OMZ-LW), except for small Copepods that
370 remained at the OMZ-UB stratum with and showed a smaller ΔWMD (34 m). At the offshore stations (Stations T3 and L6)
371 the DVM was reverse in most Copepods exhibited reverse DVM, except for large Copepods (LC) which showed slightly
372 normal DVM at St. L6 off 29°S. At night Copepods were tended to concentrate deeper in the stratum associated with the
373 lower core of OMZ (OMZ-UC) and lower O₂ gradient (OMZ-LW), particularly Eucalanidae with which exhibited a strong
374 DVM and with high ΔWMD of 440 m (St. T3) and 208 m (St. L6) and *Acartia* Copepods with 103 m (St. L6) (Table 1),
375 whereas at St. L6 small Copepods (SC) were caught in abundance at the OMZ-UB stratum down to 82-90 m depth,
376 respectively (Fig. 3). During the day, Copepods remained shallower than at night, although they concentrated at different
377 depths. Small Copepods were tended to concentrate in the oxic surface strata OX-ML (St. T3) and remained in the upper
378 boundary of the OMZ (St. L6) without detectable DVM, as judging by the small difference between their day- and night-
379 time distributions (ADVM ca. 4 m). Large Copepods (LC), as expected, showed a normal migration and tended to stay ed
380 inside the OMZ, and concentrated in the OMZ-UC stratum (St. L6) and OMZ-UB (St. T3). Finally, Eucalanidae, with
381 which exhibited a strong DVM tended to distribute in the OMZ-UC (St. T3) and the OMZ UB (St. L6) (Fig. 3).

383 Following Copepods, Salps were the second most abundant taxa of zooplankton, showing the largest aggregation at the
384 onshore station St. L6 (~3%), where abundances made up to $10347 \pm 219 \text{ ind. m}^{-2} \text{ d}^{-1}$. The WMD of Salps showed a range
385 between 54 and 25 m depth (Table 1) and appeared to aggregate deeper at all stations, also present in the OMZ (Fig. 3).
386 Similarly to Salps, most Chaetognaths showed the largest aggregation at the onshore station St. L6 (~2%), where their
387 abundances reached up to $4755 \pm 1038 \text{ ind. m}^{-2} \text{ d}^{-1}$. The abundance and biomass of this group increased in the upper boundary
388 of the OMZ (OMZ-UB) during day and night. No DVM was discernible for this group in this station, because of the slight
389 difference between their day and night time distributions (Fig. 4). By contrast, off Iquique they appeared to perform a strong
390 DVM between the OMZ-UB and the OMZ-UC strata, as indicated by the migration indices (WMD- Δ DVM) (Table
391 1). However at the onshore station (St. T3) they showed a reverse DVM.
392 Unlike copepods, the Euphausiids exhibited were more abundant greater abundances at the onshore station L6 (<1%), where
393 they reached up to $1683 \pm 473 \text{ ind. m}^{-2} \text{ d}^{-1}$. The OMZ-UB stratum was the most abundant in this station, with a
394 peak of abundance during the daytime, however no DVM was detectable, judging by the small difference between their day-
395 and night-time distributions (Fig. 3). Off Iquique, also the highest abundance was in OMZ-UB stratum at night, but with a
396 second peak in OMZ-LC stratum during daytime in both station (Sts. T3 and T5) (Table S4 Supplemental Material). The
397 euphausiids appeared to perform a strong DVM in these stations (Fig. 3), with a vertical range between 236 and 56 m and a
398 mean Δ DVM of 181 m at T5 and at St. T3 between 222 and 73 m with a mean Δ DVM of 149 m (Table 1).
399
400 , although regarding vertical distribution they were mainly concentrated in the OMZ-UB and OMZ-LC in St. T5 and St. T3 at
401 day and at night, respectively. No DVM was detectable in the onshore station (St. L6), judging by the small difference
402 between their day and night time distributions (Fig. 4). However, at the offshore station St. T5 they appeared to perform a
403 strong DVM with a vertical range between 236 and 56 m and a mean Δ DVM of 181 m (Table 1). Decapods larvae were
404 more abundant at St. T5 ($428 \pm 132 \text{ ind. m}^{-2} \text{ d}^{-1}$) and were associated with the OMZ-UB stratum, where they performed a
405 strong normal DVM with a vertical range between 120 and 30 m and a mean Δ DVM of 90 m (Table 1). At the offshore
406 station (St. T3), the surface peak of abundance was in the OX-ML stratum during the day and at in the OMZ-UB cat layer at
407 night in the offshore station (St. T3), where they reached up to $292 \pm 62 \text{ ind. m}^{-2} \text{ d}^{-1}$, with a weak reverse DVM (Δ DVM=7 m).
408 However, at the offshore station St. T5 they appeared to perform a strong normal DVM with a vertical range between 120
409 and 30 m and a mean Δ DVM of 90 m (Table 1). Off Coquimbo (St. L6) they reached up to $400 \pm 88 \text{ ind. m}^{-2} \text{ d}^{-1}$, the OMZ-
410 UB stratum was the most abundant, with a slight second peak in the OMZ-LW stratum during daytime, at this station the
411 vertical range was between 70 and 48 m with a mean Δ DVM of 22 m (Fig. 3).
412
413 The largest aggregation of Chaetognaths was at the onshore station St. L6 (~2%), where their abundances reached up to
414 $4755 \pm 1038 \text{ ind. m}^{-2} \text{ d}^{-1}$. The abundance and biomass of this group increased in the upper boundary of the OMZ (OMZ-UB)
415 during day and night. No DVM was discernible for this group in this station, because of the slight difference between their
416 day- and night-time distributions. By contrast, off Iquique they appeared to perform a strong DVM between the OMZ-UB

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417 and the OMZ-UC strata, as indicated by the migration indices (WMD-ΔDVM) (Table 1). However, at the onshore station
418 (St. T3) they showed a reverse DVM.

419

420 The other main taxon, Annelida was more abundant at the onshore station St. L6, where their abundances reached up to
421 7395±847 ind. m⁻² d⁻¹ (Table S4 Supplemental Material). In the whole area, the highest of abundance was in the OMZ-UB,
422 however a second peak of abundance was in the OMZ-LC during daytime at St T5 and T3 and during the night at St. L6. The
423 DVM of this group was high off Iquique with ΔDVM of 145 m (St. T5) and 105 m (St. T3), while at St. L6 off Coquimbo no
424 DVM was discernible for this group.

425

426 The other main taxa, Appendicularian, showed the largest aggregation at the onshore station T5 (5%), where abundances
427 reached up to 21848 ind. m⁻². The abundance and biomass of this group were mainly found in the oxic surface strata OX-ML
428 at St. T5 and St. L6 both at day and at night. No DVM was detected in these stations, judging by the small difference
429 between their day and night time distributions, while at the onshore station (St. T3) they appeared to have a strong reverse
430 DVM into the OMZ-UC stratum at night as indicated by the migration indices (WMD-ΔDVM) (Table 1). The WMD at this
431 station T3 showed a range between 18 and 106 m depth and ΔDVM of 88 m depth (Table 1).

432 3.3.3 Others groups with vertical distribution associated to OMZ UC

433 The remaining 19 groups that only constituted 112% (in St. T5) 17% (in St. T3) and 276% (in St. L6) in abundance, showed
434 variable. The DVM behaviours was variable at each station, but in general: it was Normal DVM (in St. T5) and reverse
435 DVM (in St. T3-L6.) (Table 1). These groups clearly exhibited different daytime and night depths associated with the
436 OMZ core (OMZ UC-LC). Overall, they tended to reside deeper by day and shallower by night in St. T5 than at the other
437 sites (Fig.3).

438 3.4 Vertical distribution of zooplankton biomass

439 Estimates (mean ± SD) of biomasses of the taxonomic groups integrated by depth strata are summarized in Table 2. These
440 data, averaged from day and night measurements, contrast with the numerical abundances, which were dominated by
441 copepods. In this case, the bulk of zooplankton biomass was dominated by different groups depending on Stations. In terms
442 of biomass, Salps, Copepods, Euphausiids, Decapods L. Chaetognaths, Euphausiids and Decapods Annelids accounted, more
443 or less equally, for >7084% in the whole area (Table 2). At the stations less affected by the effect of OMZ the bulk of
444 biomass was dominated by copepods ~50% at St. T5 and ~40% at St. L6, while Decapods largely dominated the bulk of
445 biomass at Station T3 (~40%), followed by Copepods (19%) and Euphausiids (16 %). (Table 2).

446 The bulk of biomass at Station T5 was dominated by Decapods, Salps, Copepods and Siphonophores >70% in this station,
447 while Salps largely dominated the bulk of biomass at Station T3 (>60%) followed by Chaetognaths (12%). At the onshore
448 Station L6, Copepods and Chaetognaths accounted for about 40% of total biomass (Table 2).

449

450 **Table2.** Mean and Standard Deviation (\pm) of integrated biomass (mg C m⁻²) by taxonomic groups identified and sorted by
 451 ZooScan during daytime/ night-time conditions at three stations (T5, T3 and L6) sampled off northern Chile, during the
 452 austral spring 2015. Mean \pm SD are from n=8 for Stations T5 and T3, and n=4 for Station L6.

TAXA	T5		T3		L6	
	Integrated Biomass		Integrated Biomass		Integrated Biomass	
Salps	652.37	\pm 262.0	9904.93	\pm 1906.6	963.60	\pm 245.6
Copepods						
Large Copepods	552.91	\pm 137.4	231.05	\pm 46.9	838.28	\pm 139.6
Small Copepods	419.12	\pm 96.9	271.77	\pm 48.4	2500.51	\pm 400.3
Eucalanidae Copepods	0.02	\pm 0.0	0.00	\pm 0.0	0.44	\pm 0.2
Acartia Copepods	0.71	\pm 0.3	8.75	\pm 3.2	50.89	\pm 16.3
Chaetognaths	215.71	\pm 72.1	1800.82	\pm 744.1	2261.81	\pm 510.6
Euphausiids	296.39	\pm 69.0	436.24	\pm 115.4	919.43	\pm 250.6
Decapoda Larvae	21.83	\pm 5.4	1074.62	\pm 321.2	549.82	\pm 160.0
Hydrozoa	275.97	\pm 106.9	363.61	\pm 128.2	796.03	\pm 182.7
Annelidae	71.17	\pm 9.8	105.60	\pm 15.4	1044.62	\pm 268.9
Siphonophores-	452.64	\pm 106.5	363.37	\pm 103.3	124.70	\pm 29.9
Decapods	652.79	\pm 279.4	0.00	\pm 0.0	0.00	\pm 0.0
Appendicularian	30.30	\pm 10.1	14.79	\pm 5.9	543.98	\pm 232.7
Ctenophores	0.62	\pm 0.3	119.55	\pm 53.5	102.50	\pm 31.7
Ostracods-	11.92	\pm 1.3	16.51	\pm 1.1	185.56	\pm 35.4
Amphipods	10.62	\pm 2.2	51.17	\pm 15.8	30.73	\pm 6.4
Fish Eggs	17.64	\pm 3.0	10.64	\pm 1.4	47.24	\pm 9.1
Platyhelminthes	0.00	\pm 0.0	0.00	\pm 0.0	62.04	\pm 26.7
Nauplii	3.11	\pm 0.8	1.41	\pm 0.4	25.53	\pm 7.6
Ichthyoplankton	6.98	\pm 2.1	18.83	\pm 8.4	0.00	\pm 0.0
Pteropods	5.58	\pm 2.4	0.89	\pm 0.2	4.27	\pm 1.8
Foraminifera	1.14	\pm 0.4	0.14	\pm 0.0	9.13	\pm 2.5
Gastropods	0.94	\pm 0.2	0.56	\pm 0.1	3.59	\pm 0.4
Cirripedia Larvae	0.25	\pm 0.1	0.14	\pm 0.0	4.10	\pm 1.1
Bryozoan-	0.80	\pm 0.2	0.20	\pm 0.0	1.35	\pm 0.4
Radiolarian	0.04	\pm 0.0	0.08	\pm 0.0	1.03	\pm 0.4
Echinoderm Larvae	0.01	\pm 0.0	0.00	\pm 0.0	0.04	\pm 0.0
Total	3702	\pm 805	14796	\pm 2512	11071	\pm 1635

453

TAXA	T5	T3	L6
	Integrated Biomass	Integrated Biomass	Integrated Biomass

Tabla con formato

Copepods	-	-	-	-	-	-	-
Large C	<u>383.73</u>	±	<u>89.3</u>	<u>286.75</u>	±	<u>50.6</u>	<u>1727.49</u> ± <u>340.8</u>
Small C	<u>467.55</u>	±	<u>115.9</u>	<u>194.51</u>	±	<u>41.4</u>	<u>416.14</u> ± <u>66.6</u>
Eucalanidae C	<u>0.22</u>	±	<u>0.1</u>	<u>2.42</u>	±	<u>0.8</u>	<u>11.78</u> ± <u>4.1</u>
Acartia C	<u>0.02</u>	±	<u>0.0</u>	<u>0.00</u>	±	<u>0.0</u>	<u>0.34</u> ± <u>0.1</u>
Euphausiids	<u>349.48</u>	±	<u>77.6</u>	<u>412.38</u>	±	<u>103.0</u>	<u>1060.58</u> ± <u>305.8</u>
Decapoda L	<u>20.04</u>	±	<u>4.7</u>	<u>1015.61</u>	±	<u>294.6</u>	<u>537.17</u> ± <u>176.8</u>
Chaetognaths	<u>63.97</u>	±	<u>15.7</u>	<u>215.20</u>	±	<u>79.3</u>	<u>641.85</u> ± <u>172.9</u>
Annelida	<u>50.76</u>	±	<u>7.4</u>	<u>79.59</u>	±	<u>11.8</u>	<u>734.67</u> ± <u>196.4</u>
Decapods	<u>442.38</u>	±	<u>187.4</u>	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u> ± <u>0.0</u>
Ostracods	<u>20.83</u>	±	<u>2.9</u>	<u>25.14</u>	±	<u>1.9</u>	<u>171.01</u> ± <u>15.2</u>
Ctenophores	<u>0.45</u>	±	<u>0.2</u>	<u>101.44</u>	±	<u>45.4</u>	<u>100.01</u> ± <u>29.9</u>
Hydrozoan	<u>31.48</u>	±	<u>9.9</u>	<u>28.39</u>	±	<u>8.3</u>	<u>132.34</u> ± <u>39.0</u>
Salps	<u>15.96</u>	±	<u>6.0</u>	<u>129.33</u>	±	<u>33.5</u>	<u>35.14</u> ± <u>9.9</u>
Siphonophores	<u>71.53</u>	±	<u>25.4</u>	<u>55.30</u>	±	<u>16.3</u>	<u>39.17</u> ± <u>9.8</u>
Amphipods	<u>8.78</u>	±	<u>1.7</u>	<u>43.32</u>	±	<u>13.2</u>	<u>28.34</u> ± <u>5.3</u>
Platyhelminthes	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u>	±	<u>0.0</u>	<u>48.83</u> ± <u>21.0</u>
Ichthyoplankton	<u>7.62</u>	±	<u>2.2</u>	<u>28.64</u>	±	<u>12.8</u>	<u>0.00</u> ± <u>0.0</u>
Nauplius L	<u>2.15</u>	±	<u>0.6</u>	<u>1.02</u>	±	<u>0.3</u>	<u>19.78</u> ± <u>5.6</u>
Pteropods	<u>4.20</u>	±	<u>1.8</u>	<u>0.62</u>	±	<u>0.2</u>	<u>5.02</u> ± <u>2.2</u>
Foraminifers	<u>0.45</u>	±	<u>0.1</u>	<u>0.27</u>	±	<u>0.1</u>	<u>7.55</u> ± <u>0.9</u>
Gastropods	<u>0.35</u>	±	<u>0.1</u>	<u>0.21</u>	±	<u>0.0</u>	<u>1.32</u> ± <u>0.2</u>
Cirripedia L	<u>0.17</u>	±	<u>0.0</u>	<u>0.09</u>	±	<u>0.0</u>	<u>1.25</u> ± <u>0.4</u>
Bryozoa L	<u>0.59</u>	±	<u>0.2</u>	<u>0.16</u>	±	<u>0.0</u>	<u>0.70</u> ± <u>0.3</u>
Radiolarian	<u>0.02</u>	±	<u>0.0</u>	<u>0.05</u>	±	<u>0.0</u>	<u>0.60</u> ± <u>0.2</u>
Echinoderm L	<u>0.01</u>	±	<u>0.0</u>	<u>0.00</u>	±	<u>0.0</u>	<u>0.04</u> ± <u>0.0</u>
Appendicularian	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u> ± <u>0.0</u>
Fish Egg	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u>	±	<u>0.0</u>	<u>0.00</u> ± <u>0.0</u>
Total	<u>1943</u>	±	<u>435</u>	<u>2620</u>	±	<u>572</u>	<u>5721</u> ± <u>1019</u>

Tabla con formato

454

455 When assessing the day vs. night vertical distribution of taxonomic groups in terms of their contribution to biomass, different
456 patterns arise compared to numerical abundance. In this case, we used 9 taxonomic categories to examine vertical
457 distribution and DVM in terms of biomass: small Copepods (SC), large Copepods (LC), *Acartia* Copepods (AC),
458 Eucalanidae Copepods (EC), *Salps* (SA), *Chaetognaths* (CH), Euphausiids (EU), Decapod larve (DL), *Chaetognaths* (CH),
459 *Annelids* and all the other taxa (Fig. 45). Contrasting with numerical abundance (Fig. 4), the vertical distribution of biomass
460 was more heterogeneously divided among taxonomic groups and DVM patterns vary strongly between stations. Small
461 Copepods continue to dominate at the St. T5 (24%), with two peak of biomass, a surface peak associated to the upper oxic
462 layer (OX-ML) and OMZ-UB stratum during night condition, and a second peak associated to deeper stratum (OMZ-LW)

463 during daytime. At the onshore Stations T3 and L6 the biomass had a similar vertical distribution but lower (~7%), but at the
464 Station T3 their peak of biomass was in the upper oxic layer (OX-ML) during daytime condition and then it decreases
465 sharply within the OMZ-UB and within the OMZ core (OMZ-UC and OMZ-LC). This abrupt decrease in biomass
466 coincides with the intense OMZ present at this station T3. The second peak of biomass during daytime was in deeper stratum
467 (OMZ-LW), where oxygen conditions seem to be restored. Large Copepods dominate at the onshore St L6(30 %) where
468 their biomass reached up to 1727.49±340.8mg C. m⁻² d⁻¹(Table 2). A surface peak of biomass was associated to OMZ-UB
469 stratum during daytime condition, and a second peak associated to deeper stratum (OMZ-LW) also during daytime. Off
470 Iquique they were the second dominant group, with a surface peak in OX-ML stratum during night at St. T5 and daytime at
471 St. T3, and a second peak in deeper stratum (OMZ-LW), during daytime in both stations (Fig. 4).The biomass of
472 Eucalanidae and Acartia copepods were lower than the other copepods in the whole area, but in general Eucalanidae was
473 associated to the deeper stratum.

474

475 Following Copepods, Euphausiids were the second dominant group in term of biomass in the whole area. In general, their
476 ascent from deep layers to the upper ones at night was also evidenced by increasing proportions of these group in the OMZ-
477 UB stratum at night. The highest biomass was in St L6(19%) where reached up to 1060.58±305.8 mg C. m⁻² d⁻¹(Table 2). A
478 surface peak of biomass was associated to OMZ-UB stratum during night condition, decreasing in the deep strata (Fig. 4).
479 Across of the zonal section off Iquique two peak of biomass was in both stations. A surface peak was in OMZ-UB stratum
480 during night condition followed by lower biomass within the OMZ core, then a second peak was in the OMZ-LC during
481 daytime.

482 Decapod Larvae clearly dominated over copepods in the St. T3 (39%). The high biomass was in OMZ-UB stratum during
483 night condition followed by lower biomass within the OMZ Core. At the St L6 also was a second peak of biomass in the
484 OMZ-LC during night. tended to distribute more homogeneously in the water column, although they also tend to diminish
485 their biomass during the day in the upper layers, while Euphausiids showed their ascent from deep layers to the upper ones at
486 Station T5, but with lower biomass within the OMZ at the onshore stations T3 and L6. Chaetognaths and Annelids were
487 another groups with an exhibiting important vertical movements of biomass between day and night across strata, and
488 like other groups with two peaks of biomass although with not clear night ascent. The high biomass was at St. L6 in both
489 groups (Fig. 45).

490

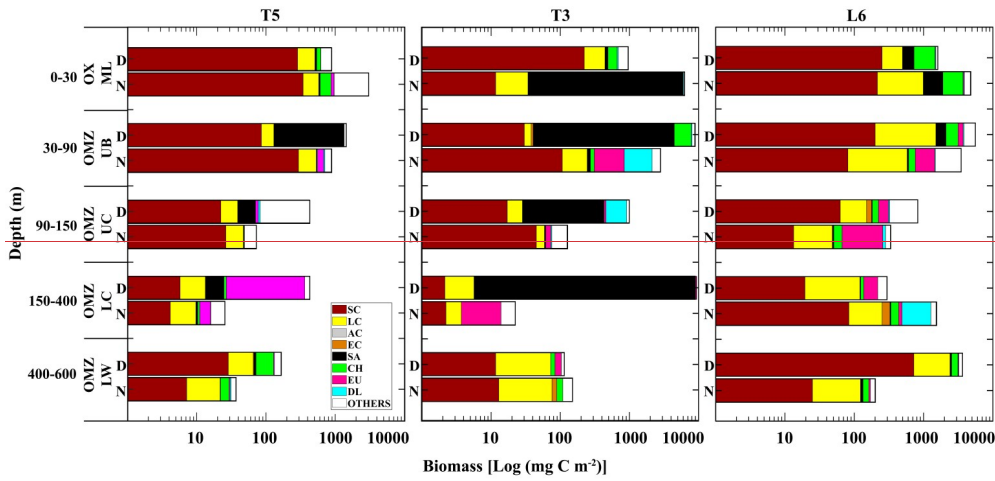
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492 Total added biomass of zooplankton revealed more clearly DVM behaviour of the whole zooplankton community (Fig. 56).
493 The vertical distribution and daytime vs. Night-time variability of zooplankton biomass showed distinctive features
494 associated with the OMZ structure, with significant differences (p<0.05) between strata for both daytime and night-time
495 samplings, as based on the ANOSIM test (p<0.05). In the whole area most of the biomass was concentrated in a narrow band
496 within the OX-ML and OMZ UB strata, associated with more oxygenated surface waters, with reduced values in deeper

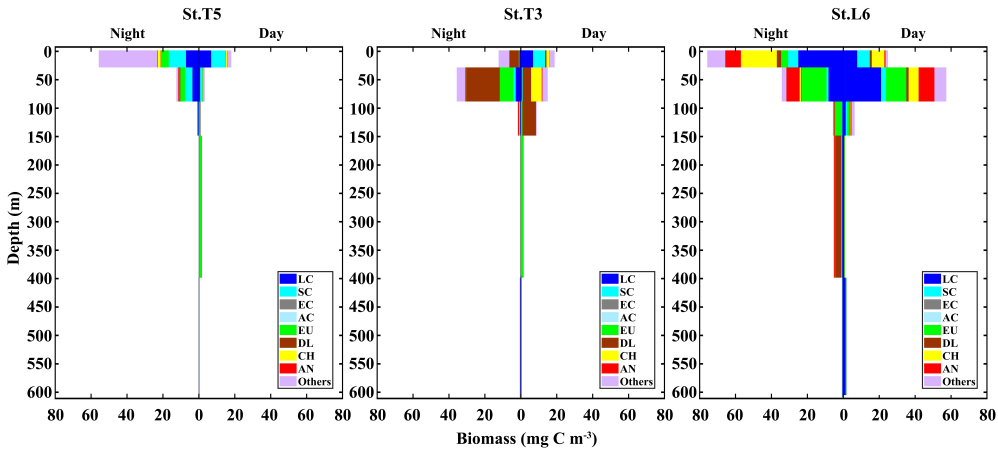
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497 waters associated with the OMZ core, in special at the onshore station off Iquique (St. T3) (Fig.56). Overall, we observed
 498 that highest values of biomass were ~~present~~ during the night at the shallower sampling stratum (Ox-ML) and in the
 499 subsurface during the day. There was also an important increase in biomass at the deepest stratum (OMZ-LW) during the
 500 daytime and night condition.

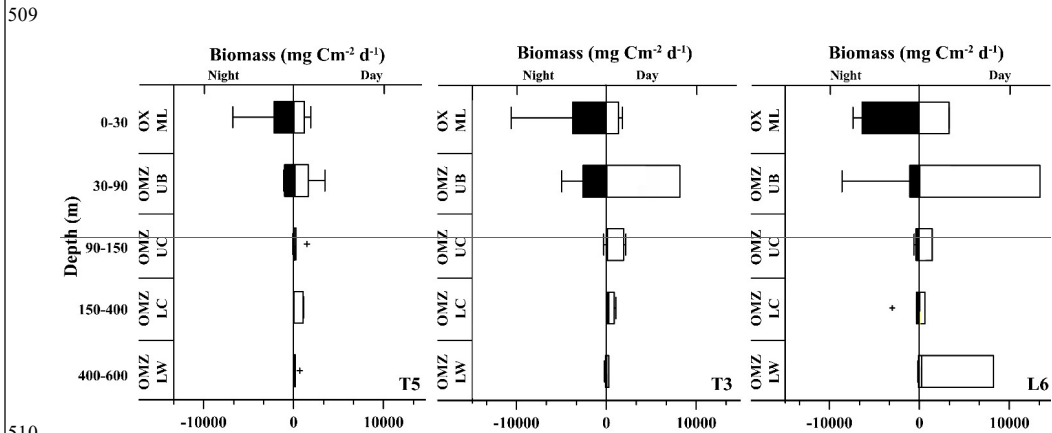
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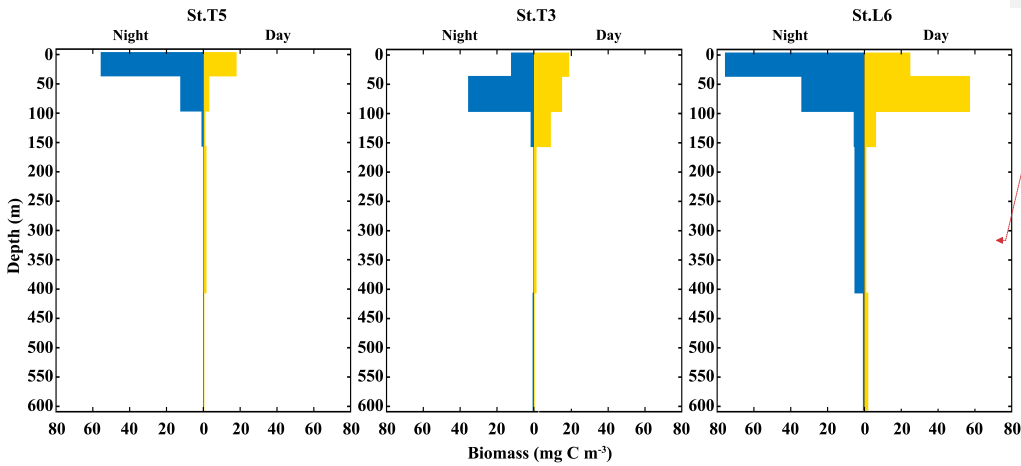
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504
 505 **Fig.45.** Daytime vs. Night-time vertical distribution of biomass of different taxonomic groups at 3 stations off
 506 northern Chile: off Iquique (Stations T5 and T3) and off Coquimbo (Station L6). Data are from night and day replicated
 507 samples during two consecutive days in the austral spring 2015. Values represent means from sampling size n=4 for St.T5
 508 and St. T3, and n=2 for St. L6. *Layers were defined from variable levels of oxygen concentration (see Methods).*



510



Con formato: Centrado

511

512 **Fig. 56.** Vertical distribution of total zooplankton biomass during daytime and night-time conditions at 3 stations off
 513 northern Chile: off Iquique (Stations T5 and T3) and off Coquimbo (Station L6) during two consecutive days in the austral
 514 spring 2015. Some outliers are plotted individually (+). Error bars are standard deviations from sampling size (n=4 for St. T5
 515 and St. T3 and n=2 for St. L6). Layers were defined from variable levels of oxygen concentration (see Methods).

516 **3.4.1 Migrant biomass of the zooplankton taxa**

517 The migrant biomass of the zooplankton taxa and the rate of migration (RM), represented by the proportion of biomass (%)
 518 being vertically moved daily from the upper 90 m are shown in **Table 3**. Most of dominant groups showed a high rate of
 519 migration as reflected in the RM. In terms of migrant biomass, Decapod larvae, Euphausiids, Decapods, Copepods and
 520 Chaetognaths, Euphausiids, Hydrozoa and Decapod larvae accounted for a large proportion of total migrant biomass
 521 (>81%), although high estimates of migrant biomasses were also associated with high standard deviations, indicating
 522 strong variation among replicated samples (**Table 3**). Presence of zero values in **Table 3** represents absence of a given
 523 taxonomic group in the upper 90 m layer or extremely low values of biomass under both conditions' day and night, so that
 524 such groups did not contribute, or they had a non-significant contribution to total migrant biomass.

525
526

527 **Table 3.** Migrant biomass (mg C m⁻²) and rate of migration (RM) (%) for taxonomic groups of zooplankton sampled at
 528 northern Chile at 3 stations: off Iquique (Stations T5 and T3) and off Coquimbo (Station L6) during the austral spring 2015.
 529 RM represents the proportion (%) of migrant biomass with respect to total biomass found at night in the 0-90 m for a given
 530 taxonomic group. SD= Standard deviation of the migrant biomass estimated from n=4 (St. T5 and St. T3) and n=2 (St. L6).

Con formato: Justificado

TAXA	ST-T5				ST-T3				ST-L6		
	MB	SD	RM %		MB	SD	RM %		MB	SD	RM %
<u>Decapoda L</u>	<u>24.50</u>	<u>17.32</u>	<u>1</u>	-	<u>968.95</u>	<u>685.15</u>	<u>40</u>	-	<u>12.19</u>	<u>8.62</u>	<u>1</u>
<u>Euphausiids</u>	<u>269.50</u>	<u>190.56</u>	<u>11</u>	-	<u>455.14</u>	<u>321.83</u>	<u>19</u>	-	<u>247.96</u>	<u>175.34</u>	<u>10</u>
<u>Decapods</u>	<u>876.93</u>	<u>620.08</u>	<u>36</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>
<u>Chaetognaths</u>	<u>20.31</u>	<u>14.36</u>	<u>1</u>	-	<u>390.72</u>	<u>276.28</u>	<u>16</u>	-	<u>105.39</u>	<u>74.52</u>	<u>4</u>
<u>Copepods</u>											
<u>Small C.</u>	<u>223.31</u>	<u>157.90</u>	<u>9</u>	-	<u>139.02</u>	<u>98.30</u>	<u>6</u>	-	<u>128.19</u>	<u>90.64</u>	<u>5</u>
<u>Large C.</u>	<u>183.07</u>	<u>129.45</u>	<u>8</u>	-	<u>43.78</u>	<u>30.96</u>	<u>2</u>	-	<u>245.28</u>	<u>173.44</u>	<u>10</u>
<u>Eucalanidae C.</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>1.10</u>	<u>0.78</u>	<u>0</u>	-	<u>0.12</u>	<u>0.09</u>	<u>0</u>
<u>Acartia C.</u>	<u>0.04</u>	<u>0.03</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.62</u>	<u>0.44</u>	<u>0</u>
<u>Ctenophores</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>202.89</u>	<u>143.46</u>	<u>8</u>	-	<u>81.61</u>	<u>57.71</u>	<u>3</u>
<u>Annelids</u>	<u>32.50</u>	<u>22.98</u>	<u>1</u>	-	<u>4.03</u>	<u>2.85</u>	<u>0</u>	-	<u>156.14</u>	<u>110.41</u>	<u>6</u>
<u>Hydrozoa</u>	<u>35.83</u>	<u>25.34</u>	<u>1</u>	-	<u>7.80</u>	<u>5.52</u>	<u>0</u>	-	<u>79.04</u>	<u>55.89</u>	<u>3</u>
<u>Siphonophores</u>	<u>34.89</u>	<u>24.67</u>	<u>1</u>	-	<u>58.18</u>	<u>41.14</u>	<u>2</u>	-	<u>0.33</u>	<u>0.23</u>	<u>0</u>
<u>Amphipods</u>	<u>5.62</u>	<u>3.98</u>	<u>0</u>	-	<u>32.50</u>	<u>22.98</u>	<u>1</u>	-	<u>32.93</u>	<u>23.29</u>	<u>1</u>
<u>Salps</u>	<u>27.36</u>	<u>19.34</u>	<u>1</u>	-	<u>38.56</u>	<u>27.27</u>	<u>2</u>	-	<u>4.57</u>	<u>3.23</u>	<u>0</u>
<u>Ostracods</u>	<u>17.80</u>	<u>12.59</u>	<u>1</u>	-	<u>12.92</u>	<u>9.14</u>	<u>1</u>	-	<u>36.37</u>	<u>25.72</u>	<u>2</u>
<u>Platyhelminthes</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>63.08</u>	<u>44.61</u>	<u>3</u>
<u>Pteropods</u>	<u>3.97</u>	<u>2.81</u>	<u>0</u>	-	<u>0.07</u>	<u>0.05</u>	<u>0</u>	-	<u>9.85</u>	<u>6.96</u>	<u>0</u>
<u>NaupliusL</u>	<u>3.06</u>	<u>2.17</u>	<u>0</u>	-	<u>0.36</u>	<u>0.25</u>	<u>0</u>	-	<u>2.84</u>	<u>2.01</u>	<u>0</u>
<u>Foraminiphers</u>	<u>0.09</u>	<u>0.07</u>	<u>0</u>	-	<u>0.03</u>	<u>0.02</u>	<u>0</u>	-	<u>3.11</u>	<u>2.20</u>	<u>0</u>
<u>Cirripedia L</u>	<u>0.26</u>	<u>0.18</u>	<u>0</u>	-	<u>0.01</u>	<u>0.01</u>	<u>0</u>	-	<u>1.87</u>	<u>1.32</u>	<u>0</u>
<u>Brvozoa L</u>	<u>1.06</u>	<u>0.75</u>	<u>0</u>	-	<u>0.18</u>	<u>0.13</u>	<u>0</u>	-	<u>0.83</u>	<u>0.59</u>	<u>0</u>
<u>Gastropods</u>	<u>0.22</u>	<u>0.16</u>	<u>0</u>	-	<u>0.12</u>	<u>0.09</u>	<u>0</u>	-	<u>0.11</u>	<u>0.08</u>	<u>0</u>
<u>Echinoderm larvae</u>	<u>0.02</u>	<u>0.01</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.07</u>	<u>0.05</u>	<u>0</u>
<u>Ichthvoplankton</u>	<u>0.06</u>	<u>0.04</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>
<u>Appendicularia</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>
<u>Fish Eggs</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>
<u>Radiolarian</u>	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>	-	<u>0.00</u>	<u>0.00</u>	<u>0</u>

533 **4 Discussion**

534 Studies on zooplankton DVM and the active transport of C₂ mediated by zooplankton have been documented previously for
535 the Pacific Ocean and for other areas of the world's ocean, as summarized in **Table 4**. However, downward C flux due to
536 DVM in highly productive upwelling regions, such as northern Chile, which is also characterized by severe subsurface
537 hypoxic conditions upon presence of a shallow OMZ, is still poorly understood. Some studies have shown that hypoxic

538 conditions can interfere with DMV of many meso- and macrozooplankton species (Wishner *et al.*, 2013; Ekau *et al.*, 2010;
539 Escribano *et al.*, 2009; Apablaza and Palma, 2006; Antezana, 2002; Escribano, 1998). These studies have shown that small
540 differences in oxygen concentration can make a large difference for zooplankton behavior, physiology and adaptation
541 (Wishner *et al.*, 2018; Kiko *et al.*, 2016; Seibel, 2011; Gonzalez and Quiñones, 2002; Escribano and McLaren, 1999).
542 Therefore, it seems that the OMZ can play a very significant role influencing vertical distribution, DVM and ultimately the
543 downward C flux mediated by zooplankton.

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544 Our approach to assess downward C flux into the Oxygen Minimum Zone, based on estimates of the migrant biomass and
545 our proposed migration indices, allowed us on one hand to examine the contribution that different zooplankton can have for
546 vertical flux of C and hence export production, and in the other hand to assess zooplankton responses (e.g. vertical
547 distribution and DVM performance) to changes in environmental conditions over the vertical gradient, such as temperature,
548 water density and the abrupt changes in oxygenation levels. In this subtropical upwelling region, vertical gradients are much
549 stronger than in temperate upwelling zones. For example, the coastal zone in this region is more stratified and with a very
550 shallow OMZ (<50 m) with a weak seasonal signal and moderate upwelling throughout the year (Paulmier and Ruiz-Pino,
551 2009; Fuenzalida *et al.*, 2009; Escribano *et al.*, 2004). This means that zooplankton must cope with hypoxic conditions
552 during their entire life cycle, except for some species that may reside in near surface water (<30 m), such as *C. chilensis* and
553 *C. brachiatus* which have been reported as mostly restricted to the upper layer without performing any substantial DVM
554 (Escribano *et al.*, 2012, 2009; Escribano and Hidalgo, 2000; Escribano, 1998).

Con formato: Español (Chile)

Código de campo cambiado

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555 The vertical distribution and diurnal variability of zooplankton biomass seem to be disturbed by the OMZ, such that high
556 biomass aggregates above the oxycline in a narrow band within the OX-ML and OMZ-UB layers, associated with more
557 oxygenated surface waters, whereas extremely low biomass reside in deeper waters, in particular within the OMZ core. This
558 condition was more evident in the coastal station off Iquique (St. T3), characterized by the most intense OMZ in the whole
559 study area. In the eastern tropical north Pacific, biomass distribution seemed different, exhibiting a secondary peak at depth
560 during the daytime within the upper oxycline or OMZ core (Wishner *et al.*, 2013).

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561 Regarding the estimates of biomass for each of the taxonomic groups, our approaches can certainly introduce variation,
562 depending on selected regressions and conversion factors from highly diverse body shapes and body densities of the
563 zooplankton taxa affecting the estimates of [body area and biovolume](#), dry weight and C content. Various approaches have
564 been adopted for converting sizes to body masses. For example, Lehette and Hernández-León (2009) provided some general
565 regression equations for subtropical and Antarctic zooplankton describing the relationship between scanned area and body
566 mass (C content). These authors also proposed two separate regressions for crustacean and gelatinous zooplankton, because
567 of different body densities. [In our study, we adopted more direct estimates of body masses by converting individual areas or
568 volumes \(from ZooScan\) using published regressions for separate taxonomic groups. Also, in our samples there was a high
569 diversity of taxonomic groups as identified by ZooScan, such that unique regressions for crustacean and gelatinous
570 organisms may lead to strong biases in body mass estimates, because of high variability in C content, which is the key
571 component of body mass needed to estimate C flux. Therefore, \[the use of taxa-specific conversion factors, as those detailed\]\(#\)](#)

572 in our Table S3 (Supplemental Material) is strongly recommended following Yamaguchi *et al.* (2014) and Sato *et al.* (2015);
573 we converted biovolumes into dry weights using a mean body density and mean water content of zooplankton, estimated
574 across taxa by Wiebe (1975), but then taxa-specific Carbon: dry weight ratios collected from published works. In any case,
575 the use of single conversion factors between body volume and mass and wet weight and dry weight among taxa must be
576 considered as source of variability in the estimates of taxonomic biomasses.

577 Despite the apparently hostile oxygen-deficient habitat, associated with the OMZ, we found that most taxa were able to
578 perform DVM in the upwelling zone withstanding severe hypoxia. Even, several zooplankton groups are strong migrants,
579 exhibiting large DVM amplitude (~500 m). Among them, an important migrant group is comprised by the copepods
580 Eucalanidae which have been described as even being able to enter the core of the OMZ, and then migrate downward to the
581 lower limit of the OMZ, which is slightly more oxygenated (Hidalgo *et al.*, 2005). In our study however, their contribution to
582 total migrant biomass was too small (ca. 0.42 ± 0.06 mg C m⁻² d⁻¹), as compared to the estimate made by Hidalgo *et al.* (2005). In
583 fact, the migrant biomass and rate of migration of this group was non-significant when considering DVM between the upper
584 90 m and below, suggesting a little or no contribution to downward flux of C for this group of copepods. However it seems
585 that Eucalanidae remain below the oxycline or nearby the base of the oxycline day and night, as shown by their weighted
586 mean depth (WMD) and therefore suggesting that they may still contribute to vertical flux by feeding at the base of the
587 oxycline at night and then migrating into the OMZ during the day.

588 Other taxa, such as Euphausiids, *Acartia* spp., other copepods, Ctenophores, Decapods, Annelidae, Bryozoa La, Pteropods
589 and Chaetognaths tended to concentrate their populations inside the OMZ core showing a strong link to the OMZ with
590 important movement throughout the water column. Antezana (2010) showed that *E. mucronata*, an endemic and abundant
591 euphausiid in the coastal upwelling zone off Chile, is a well-adapted species to vertically migrate into the core of the OMZ.
592 In fact, the euphausiids studied here showed a large DVM amplitude (~250 m), descending into the core of the OMZ and
593 below 250 m each day. In general, all strong migrants' taxa showed a strong interaction with the core of OMZ, remaining
594 there either temporarily or permanently during the day or night condition, contributing in this way to the release of C below
595 the thermocline, despite presence of hypoxic conditions.

596 Our estimates of DVM-mediated C flux showed that migrant biomass (9583433 ± 778889 mg C m⁻² d⁻¹) and C flux estimates
597 (71674 ± 64291 mg C m⁻² d⁻¹) of the major taxa performing DVM, were greater than those reported for the Pacific Ocean, both
598 in oligotrophic, such as Hawaii, and mesotrophic waters, as the subarctic North Pacific (Steinberg *et al.*, 2008), and even
599 greater than that informed by Yebra *et al.* (2005) within eddies with enhanced biological production. Most of these previous
600 estimates however have not been done in regions with severe hypoxia or anoxia at mid water depths (e.g. Kiko *et al.*, 2016),
601 such as the highly productive upwelling region of the coastal zone off northern of Chile, where the oxygen concentrations
602 may fall below <1 μmol in the core of OMZ (Paulmier and Ruiz-Pino, 2009). Moreover, only few works have considered the
603 whole zooplankton community (Table 4). High productivity and strong aggregation of zooplankton in coastal areas of this
604 region (Escribano *et al.*, 2000; Escribano and Hidalgo, 2000) may promote greater amounts of migrant biomass. This
605 requires however that DVM should not be majorly constrained by presence of the OMZ and that most migrant taxa are

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606 tolerant to low-oxygen. On the other hand, our estimates of downward C flux were substantially lower than previous ones
607 reported off northern Chile by Hidalgo *et al.* (2005) for *Eucalanus inermis* alone (14.1 mg C m⁻² d⁻¹) and for copepods and
608 euphausiids by Escribano *et al.* (2009) (7200 mg C m⁻² d⁻¹) (Table 4). Although, such previous estimates may be too high,
609 considering the level of primary production in the upwelling zone of Chile (~10000 mg C m⁻² d⁻¹, the maximum estimated
610 value) (Daneri *et al.*, 2000), suggesting that previous works may have overestimated active transport of C. Although the
611 work of Escribano *et al.* (2009) was based on samples obtained with a Tucker Trawl net, which can be more efficient in
612 capturing large-sized zooplankton or macrozooplankton (Escribano *et al.*, 2007), as compared with the vertically towed
613 Multinet. This means that our estimates mainly based on mesozooplankton may not include the contribution of some
614 macrozooplankton, and therefore such values may be greater.

615

616 **Table 4.** Comparison of active transport of carbon (AC) (mg C m⁻² d⁻¹) by vertically migratory taxa in Pacific Ocean. Diel
617 vertically migratory taxa (DVM), productivity primary (PP) (mg C m⁻² d⁻¹), migrant biomass (MB) (mg C m⁻²), respiratory
618 loss (R) (mg C m⁻² d⁻¹), faecal pellets production (F) (mg C m⁻² d⁻¹) and mortality (M) (mg C m⁻² d⁻¹). Where provided by
619 authors, estimated passive export (POC) is listed. Fluxes refer to carbon export beneath the epipelagic zone (150–200 m
620 depth, depending on the study) in mg C m⁻²d⁻¹

Location	Taxa	PP	MB	AC	R	F	M	% POC	References
N.Hawaii ALOHA	DVM Zooplankton		108–216	7.1	2.6–4.8			12–18	Al-Mutairi and Landry (2001)
N. Hawaii ALOHA			157.9	3.2–13.6	3.7			18	Steinberg <i>et al.</i> (2008)
N.W. Pac.	DVM Metridia	418	144	9	3	5	1	23.1–61.8	Kobari <i>et al.</i> (2008)
N.W. Pac.	DVM copepods			8				22.3	Takahashi <i>et al.</i> (2009)
N.E. Pac.	Mesopelagic fishes	170			23.9				Davison <i>et al.</i> (2013)
Eastern Equator			96.0± 25.2		4.2± 1.2		2.9±0.8	18.4	Zhang and Dam (1997)
Eastern Equator	DVM zooplankton		154.8± 32.4		7.3± 1.4		5.4±1.1	25.4	Zhang and Dam (1997)
Central Equator (HNLC)			52.9		6		–	4	Rodier and Borgne (1997)
Western Equator			46.9		3		–	6	Rodier and Borgne (1997)
E. Eq. Pac.	DVM Zooplankton		1214	7.1	7.1			204	Rodier and Borgne (1997)
Western Equator	DVM Zooplankton		144–447	23.53–9.97	7.3–19.1		2.6–4.4	13–35	Hidaka <i>et al.</i> (2002)
Equator divergence			2.8–21.8		0.9–1.2			<1–2	Roman <i>et al.</i> (2002)
Oligotrophic area			30.2–33.8		1.3–1.7			4	Roman <i>et al.</i> (2002)

E.S.Pac. N.Chile		5503				Gonzalez <i>et al.</i> (1998)	<i>et</i>
E.S.Pac. N.Chile		10000				Daneri <i>et al.</i> (2000)	
E.S.Pac. N.Chile	DVM Eucalanus		8.0 -34	14.1		Hidalgo <i>et al.</i> , (2005)	
E.S.Pac. N.Chile	DVM zooplankton		37810	7200	670 0	Escribano <i>et al.</i> (2009)	
E.S.Pac. N.Chile	DVM zooplankton	2833±1155	9583433 ±889778	71674±64 294		This study	

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Con formato: Fuente: Cursiva

621 Differences of our estimates with previous works may also be accounted by strong variability of zooplankton abundance in
622 the upwelling zone. In fact, our estimates of migrant biomasses of the different taxonomic groups based on 2 days of
623 sampling and two replicates for each condition (day and night) are strongly variables, as shown by the standard errors in
624 **Table 3** which can be as much as 100% from the mean value. Therefore, comparisons must take caution upon strong time-
625 space variation when assessing zooplankton abundance. Nevertheless, a strong spatial variation in migrant biomass was also
626 evident when comparing the three sampling stations. For instance, St. L6 had more biomass than the other stations, but much
627 less migrant biomass, and thus a very low contribution to vertical flux of C by DVM. At station L6, large copepods,
628 Euphausiids, Annelids and Chaetognaths largely contributed to biomass, although they did not show significant DVM.
629 Therefore, species composition and their DVM behaviour appear as a key factor to determine the downward flux of C
630 mediated by active transport.

631 Even although the OMZ did not greatly prevent DVM migration, zooplankton behaviour appeared disrupted or exhibited
632 reversed patterns, depending on vertical distribution of OMZ and on the taxonomic group being considered. This behaviour
633 was more evident in the onshore stations (Stations T3 and L6), but in particular in the station off Iquique (St. T3) that also
634 showed a higher migration rate (60%). According to Ekau *et al.* (2010), other indirect effects could also be caused by the
635 hypoxia conditions, such as changes in prey availability, prey size or predation risk, as well as changes in species
636 composition, the strength of which depends on the duration and intensity of the hypoxic events. This could explain why
637 individuals within a single population can perform reverse, normal, or non DVM, apparently depending on the more
638 important source of mortality: predation by nocturnal feeding, normally migrating carnivorous zooplankton, or visually
639 hunting planktivorous fish (Ohman, 1990). These kind of DVM behaviors can only be better assessed and understood when
640 looking at the population level, although again time-space variation in zooplankton abundance in a highly heterogeneous
641 upwelling zone should be kept in mind.

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642 Concerning C fluxes, our estimates of active transport of carbon by zooplankton were greater than estimates of passive C
643 sinking obtained off northern Chile at 60 m depth off Antofagasta (23°S) by Gonzalez *et al.*(1998) based on sediment traps
644 (125 to 176mg C m⁻² d⁻¹). Regarding the question on how efficient is active C transport mediated by DVM, we obtained
645 satellite-based (<http://science.oregonstate.edu/ocean.productivity>) estimates of net primary production (monthly means for
646 November-December 2015) for the coastal area (Stations T3 and L6) and the coastal transition zone (Station T5), averaged
647 for the months of November and December 2015. Our estimates of downward C flux represented a mean of ca. 425 % of

648 export of Carbon resulting from net primary production in the upwelling region, estimated in the range of 1500-3500 mg C
649 $\text{m}^{-2} \text{d}^{-1}$ (**Table 5**). If we consider this is accounted only by mesozooplankton, then an important fraction of freshly produced
650 C might be taken downward by zooplankton, and this DVM-mediated C flux ought to be taking into account when and
651 analysing and modelling the C budget in the upwelling zone.

|652

653 **Table5.** Mean net primary production rate and estimates of daily downward C flux due to passive sinking and mediated by
654 diel vertical migration (DVM) of mesozooplankton at three stations (T5, T3 and L6) in the coastal upwelling region off
655 northern Chile during the austral spring 2015. Primary production represents satellite-based estimates of monthly mean
656 (November-December 2015) at the 3 sampling stations. Passive C flux is a mean value estimated from sediment traps by
657 González et al. (2000) off Antofagasta (northern Chile, 23 °S) for January 1997. Total biomass and epipelagic biomass are
658 mean observed values from day-night conditions after 2 consecutive days of sampling.

659

Station	T5	T3	L6	Mean ± SD
Primary Production mg C m ⁻² d ⁻¹	1500	3500	3500	2833 ± 1155
Passive Carbon Flux mg C m ⁻² d ⁻¹				151 ± 36
Integrated Abundance ind. m ⁻² (0–600 m)	221735	127085	371235	240018 ± 123097
Total biomass mg C m ⁻² d ⁻¹ (0–600m)	19433702	262014796	572114071	34289856 ± 20155646
Epipelagic biomass (mg C m ⁻²) (0–90m)	15813126	19889391	42597679	26096732 ± 14433238
Migrant biomass (mg C m ⁻² d ⁻¹)	16862863	10484457	1392977	9583433 ± 778889
Rate of Biomass migration (%)	7051	4226	37	38 ± 3412
Active Carbon Flux (mg C m ⁻² d ⁻¹)	139428	63996	11598	71674 ± 64291
Total Active Carbon exported (%)	29.3	21.8	170.3	3.825 ± 74.8

660

661 5. Conclusions

662 In the coastal upwelling zone off northern Chile the presence of a subsurface oxygen minimum zone (OMZ) can impose an
663 important constraint for diel vertical migration of zooplankton and so influencing the downward C export mediated by
664 zooplankton. We found that most of the zooplankton biomass aggregates above the oxycline, associated with more
665 oxygenated surface waters and this was evident upon presence of a more intense OMZ. Some taxonomic groups however,
666 were found closely associated with the OMZ and several taxa were able to perform DVM in the upwelling zone withstanding
667 severe hypoxia. Also strong migrants, such as large sized copepods and copepod of the group Eucalanidae and Euphausiids,
668 can exhibit a large migration amplitude (~500 m), remaining either temporarily or permanently during the day or night
669 condition within the core of the OMZ, and so contributing to the release of C below the oxycline (and thermocline). Our
670 estimates of DVM-mediated C flux suggested that a mean migrant biomass of 3.4–957.7 mg C m⁻² d⁻¹ may contribute with
671 about 71.10–6g mg C m⁻² d⁻¹ to the OMZ system through respiration, mortality, and C excretion, aproduction of fecal pellets

672 at depth, accounting for ca. 425% of the net primary production, and so implying the existence of a very efficient mechanism
673 to incorporate freshly produced C into the OMZ. This downward C flux mediated by zooplankton DVM is however strongly
674 depending on the taxonomic structure due to variable migration amplitude and DVM behavior. These estimates should also
675 consider the strong temporal-spatial variation in zooplankton abundance in the upwelling zone for comparison purposes.

676

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