1	Saltwater reduces potential CO ₂ and CH ₄ production in peat soils from a coastal freshwater
2	forested wetland
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21	Keywords: extracellular enzyme activity, sea-level rise, methanogenesis, microbial biomass
22	carbon, carbon isotopes, ghost forest
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24 Abstract A major concern for coastal freshwater wetland function and health are the effects of saltwater intrusion on greenhouse gas production from peat soils. Coastal freshwater forested 25 wetlands are likely to experience increased hydroperiod with rising sea level, as well as saltwater 26 intrusion. These potential changes to wetland hydrology may also alter forested wetland structure 27 and lead to a transition from forest to shrub/marsh wetland ecosystems. Loss of forested 28 29 wetlands is already evident by dying trees and dead standing trees ("ghost" forests) along the Atlantic Coast of the US, which will result in significant alterations to plant carbon (C) inputs, 30 particularly that of coarse woody debris, to soils. We investigated the effects of salinity and 31 32 wood C inputs on soils collected from a coastal freshwater forested wetland in North Carolina, USA, and incubated in the laboratory with either freshwater or saltwater (2.5 or 5.0 ppt) and with 33 or without the additions of wood. Saltwater additions at 2.5 ppt and 5.0 ppt reduced CO₂ 34 production by 41 and 37 %, respectively, compared to freshwater. Methane production was 35 reduced by 98 % (wood-free incubations) and by 75-87 % (wood-amended incubations) in 36 37 saltwater treatments compared to the freshwater plus wood treatment. Additions of wood also resulted in lower CH₄ production from the freshwater treatment and higher CH₄ production from 38 saltwater treatments compared to wood-free incubations. The δ^{13} CH₄-C isotopic signature 39 40 suggested that in wood-free incubations, CH₄ produced from the freshwater treatment originated primarily from the acetoclastic pathway, while CH₄ produced from the saltwater treatments 41 originated primarily from the hydrogenotrophic pathway. These results suggest that saltwater 42 43 intrusion into coastal freshwater forested wetlands will reduce CH₄ production, but long-term changes in C dynamics will likely depend on how changes in wetland vegetation and microbial 44 45 function influence C cycling in peat soils.

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47 **1 Introduction**

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Sea-level rise (SLR) threatens coastal regions around the world. Significantly, the rate of 49 SLR is not uniform around the globe, with the highest rate occurring along the Atlantic coast of 50 North America between Cape Hatteras and Cape Cod due to factors including local currents, 51 52 tides, and glacial isostatic rebound (Karegar et al., 2017; Sallenger et al., 2012). Along with economic and cultural impacts, health of coastal forested ecosystems are expected to be impacted 53 by SLR (Langston et al., 2017; Kirwan and Gedan 2019). For instance, salinization of coastal 54 55 freshwater wetlands will likely impact vegetation community dynamics and regeneration in low lying (< 1m) wetlands (Langston et al., 2017). Understanding how coastal wetland ecosystems 56 respond to extreme events, long-term climate change and a rapidly rising sea is essential to 57 developing the tools needed for sustainable management of natural resources, and the building of 58 resilient communities and strong economies. Because it has more than 5,180 km² of coastal 59 60 ecosystems and urban areas below 1 m elevation, the state of North Carolina is highly vulnerable to climate change and SLR and therefore saltwater intrusion (Riggs and Ames, 2008, Titus and 61 Richman, 2001). 62

As sea level changes, coastal plant communities move accordingly up and down the continental shelf. In recent geologic time, sea level has risen about 3 m over the past ~2,500 years from sea level reconstructions adjacent to our study site (Kemp et al., 2011). The rate of SLR has varied greatly over that time, with periods of stability and change, and a geologically unprecedented acceleration in recent decades. The current distribution of coastal forested wetlands reflects the hydrologic equilibrium of the recent past climate, but the widespread mortality of such forests suggests that the rate of SLR is in a time of rapid change at a rate

potentially faster than the forest's capacity to move upslope, resulting in widespread death of
coastal freshwater forested wetlands (Kirwan and Gedan 2019). Furthermore, dying coastal
forests will alter the quantity and quality of organic matter inputs to the soil as vegetation shifts
occur, as well as introduce a large pulse of woody debris into soils. This has the potential to alter
C cycling processes responsible for storage of C in peat soils or loss of C as CO₂ and CH₄
(Winfrey and Zeikus, 1977).

76 Wetlands store more than 25% of global terrestrial soil C in deep soil organic matter deposits due to their unique hydrology and biogeochemistry (Batjes, 1996; Bridgham et al., 77 78 2006). Carbon storage capacity is especially high in forested wetlands characterized by abundant woody biomass, forest floors of *Spaghnum* spp., and deep organic soils. Across the US 79 Southeast, soil organic C (SOC) in soils increases with proximity to the coast and is greatest in 80 coastal wetlands (Johnson and Kern, 2003). Carbon densities are even higher in the formations 81 of organic soils (Histosols) that occur across the region, typically ranging from 687 to 940 t ha⁻¹, 82 but can be as high as 1,447 t ha⁻¹ (Johnson and Kern, 2003). As noted, forested wetlands, which 83 historically have contributed to terrestrial C sequestration, are in serious decline and processes 84 leading to destabilization of accumulated soil C are not represented in broad-scale ecosystem and 85 86 land-surface models. The extent of changes in soil C cycling processes attributable to altered hydroperiod, saltwater intrusion, and structural changes in vegetation in these ecosystems 87 remains unclear. 88

Saltwater intrusion, a direct result of SLR, into freshwater wetlands alters soil C cycling
processes (Ardón et al., 2016; Ardón et al., 2018), particularly that of methanogenesis (Baldwin
et al., 2006; Chambers et al., 2011; Dang et al., 2018; Marton et al., 2012), and microbial activity
(e.g., extracellular enzyme activity, Morrissey et al., 2014; Neubauer et al., 2013). Saltwater

contains high concentrations of ions, particularly sulfate (SO_4^{2-}) , which support high rates of 93 SO₄²⁻ reduction compared to freshwater wetlands (Weston et al., 2011). Sulfate acts as a terminal 94 electron acceptor in anaerobic respiration of SOC, and SO₄²⁻ reducers will typically increase in 95 96 abundance in response to saltwater intrusion and out-compete other anaerobic microorganisms, particularly methanogens, for C (Bridgham et al. 2013; Dang et al., 2019; Winfrey and Zeikus, 97 1977). The effect of SO_4^{2-} on soil C cycling and competitive interactions with other anaerobic 98 99 microbial processes also appears dependent on the concentration of the ion (Chambers et al., 2011). Even within freshwater forested wetlands, hydrology and microtopography interact to 100 influence the amount of SO_4^{2-} within soils experiencing different levels of saturation, and 101 therefore rates of SO_4^{2-} reduction (Minick et al., 2019a). A majority of saltwater intrusion 102 studies on soil C dynamics though have focused on tidal freshwater wetlands, whereas non-tidal 103 104 freshwater wetlands have received relatively little attention, partially due to their more confined distribution across the landscape. Nonetheless, they occupy critical zones within the coastal 105 wetland ecosystem distribution and will be influenced by SLR differently than that of tidal 106 wetlands. Tidal wetlands may experience short-term pulses of saltwater with tidal movement of 107 water, while SLR effects on saltwater intrusion into non-tidal freshwater wetlands may result in 108 109 more long-term saltwater inundation. This difference in saltwater inundation period may influence rates of soil CO₂, CH₄ production, and microbial activity (Neubauer et al., 2013); and 110 therefore should be considered in light of the hydrologic properties of non-tidal wetlands. 111 112 Saltwater intrusion into freshwater systems may also influence the CH₄ production pathways (Dang et al., 2019; Weston et al., 2011), as a result of saltwater-induced shifts in 113 methanogenic microbial communities (Baldwin et al., 2006; Chambers et al., 2011; Dang et al., 114 115 2019). Stable isotope analysis of CO₂ and CH₄ indicate that acetoclastic methanogenesis is the

116 major CH_4 producing pathway in freshwater wetlands (Angle et al., 2016), but the influence of saltwater on the pathway of CH₄ formation in non-tidal freshwater forested wetlands has rarely 117 been studied, particularly through the lens of CO_2 and CH_4 stable C isotope analysis. As ¹³C 118 isotopic analysis of CH₄ is non-destructive and is long-proven as a reliable indicator of the CH₄ 119 production pathway (Whiticar et al., 1986), utilization of this analysis provides easily attainable 120 121 information on the effects of freshwater compared to saltwater on CH₄ production dynamics in coastal wetland ecosystems experiencing SLR-induced changes in hydrology and vegetation. 122 Our goal in this study was to test whether saltwater additions alter the production of CO₂, 123 124 CH₄, and microbial activity from organic soils of a non-tidal temperate freshwater forested wetland in coastal North Carolina, US, and whether effects differ in response to additions of 125 wood. Although many studies have focused on salinity pulses in tidal freshwater wetlands, less 126 127 attention has been given to the effects of sustained saltwater intrusion on soil C dynamics. We expect saltwater intrusion due to SLR will be more persistent in non-tidal wetlands. Therefore, 128 129 we investigated the effects of sustained saltwater inundation, using a laboratory microcosm experiment, on greenhouse gas production and microbial activity (e.g., microbial biomass C and 130 extracellular enzyme activity). Wood additions to microcosms were utilized to mimic the 131 132 potential large pulses of wood to peat soils as forest dieback occurs along the aquatic-terrestrial fringes of the Atlantic Coast and these wetlands transition to shrub/marsh ecosystems (Kirwan 133 and Gedan 2019); thereby providing a large and widespread pulse of coarse woody debris to 134 135 wetland soils and potentially altering soil C cycling.

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137 2 Methods

139 2.1 Field Site Description

141	The field site was located in the Alligator River National Wildlife Refuge (ARNWR) in
142	Dare County, North Carolina (35°47'N, 75°54'W) (Figure 1). The ARNWR was established in
143	1984 and is characterized by a diverse assemblage of non-tidal pocosin wetland types (Allen et
144	al., 2011). ARNWR has a network of roads and canals, but in general contains vast expanses of
145	minimally disturbed forested- and shrub-wetlands. Thirteen plots were established in a 4 km ²
146	area in the middle of a bottomland hardwood forest surrounding a 35-meter eddy covariance flux
147	tower (US-NC4 in the AmeriFlux database; Minick et al., 2019a). Of the 13 plots (7 m radius),
148	four central plots were utilized for this study which have been more intensively measured for
149	plant and soil properties and processes (Miao et al. 2013, Miao et al., 2017, Minick et al 2019a,
150	2019b, Mitra et al. 2019). Over-story plant species composition was predominantly composed of
151	black gum (Nyssa sylvatica), swamp tupelo (Nyssa biflora), bald cypress (Taxodium distichum),
152	with occasional red maple (Acer rubrum), sweet gum (Liquidambar styraciflua), white cedar
153	(Chamaecyparis thyoides), and loblolly pine (Pinus taeda). The understory was predominantly
154	fetterbush (Lyonia lucida), bitter gallberry (Ilex albra), red bay (Persea borbonia), and sweet bay
155	(Magnolia virginiana). Mean air temperature and precipitation from climate records of an
156	adjacent meteorological station (Manteo AP, NC, 35°55'N, 75°42'W, National Climatic Data
157	Center) for the period of 2008 – 2018 was 17.0 \pm 0.30 °C and 932 \pm 38 mm, respectively. These
158	wetlands are characterized by a hydroperiod that responds over short time scales and is driven
159	primarily by variable precipitation patterns. Soils are classified as a Pungo series (very poorly
160	drained dystic thermic typic Haplosaprist) with a deep, highly decomposed muck layer overlain
161	by a shallow, less decomposed peat layer and underlain by highly reduced mineral sediments of

162	Pleistocene origin (Riggs, 1996). Soils from the surface of hummocks have a pH of 4.2 ± 0.1 , C
163	concentration of 49 \pm 1.3 %, and a $\delta^{13}C$ value of -29.1 \pm 0.29 ‰ (Minick et al. 2019b). Ground
164	elevation is below < 1 m above sea level. Sea-level rise models of coastal NC show that
165	ARNWR will experience almost complete inundation by 2100, with attendant shifts in
166	ecosystem composition (DOD, 2010).
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168	2.2 Sample Collection
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170	Soil samples were collected on February 6, 2018, from surface organic soils by removing
171	seven 10 x 10 cm ⁻² monoliths from hummocks to the depth of the root mat (approximately 6.3
172	cm) using a saw and a 10 x 10 cm ⁻² PVC square. The seven soil samples were composited by
173	plot and stored on ice for transport back to the laboratory. In the laboratory, roots and large
174	organic matter were removed by hand and gently homogenized. Soils samples were then stored
175	in the dark at 4°C for seven weeks before initiating the laboratory incubation.
176	Freshwater and saltwater for the experiment was collected from water bodies surrounding
177	the ARNWR on March 7, 2018 (Figure 1). Freshwater was collected from Milltail Creek, which
178	runs Northwest from the center of ARNWR to Alligator River and drains our forested wetland
179	study site. Freshwater salt concentration was 0 ppt. Saltwater was collected from Roanoke
180	Sound to the east of ARNWR and had a salt concentration of 19 ppt (Figure 1). Freshwater and
181	saltwater were mixed together to get the desired salt concentration for the saltwater treatments
182	(2.5 and 5.0 ppt). These concentrations of saltwater were chosen due to the salinity levels in the
183	Croatan and Pamlico Sounds, which are adjacent to ARNWR (Figure 1). Salinity in these waters
184	range from approximately 1 to 5 ppt (unpublished data). Prior to mixing, freshwater and

185 saltwater was filtered through a Whatman #2 filter (8 µm). Neither saltwater nor freshwater were sterile filtered, therefore microbial communities from each water source were mixed 186 together and added to the incubations. This could influence the response of soil microbes to the 187 various treatments, but also represents what would occur under future projections of SLR in this 188 region and the resulting mixing of freshwater and saltwater within the wetland. Four water 189 samples of each freshwater and saltwater mixture were sent to the NCSU Environmental and 190 Agricultural Testing Service laboratory for analysis of total organic C (TOC), ammonium 191 (NH_4^+) , nitrate (NO_3^-) , phosphate (PO_4^-) , SO_4^- , calcium (Ca^{2+}) , magnesium (Mg^{2+}) , sodium 192 (Na⁺), potassium (K⁺), and chlorine (Cl⁻). Analysis of TOC was made using a TOC analyzer 193 (Schimadzu Scientific Instruments, Durham, NC). Analysis of NH₄⁺, NO₃⁻, and PO₄⁻, was made 194 using Latchat Quikchem 8500 flow injection analysis system (Lachat Insturments, Milwaukee, 195 WI). Sulfate and Cl⁻ were measured on a Dionex ion chromatograph (Thermo Fisher Scientific, 196 Waltham, MA). Finally, a Perkin Elmer 8000 inductively-coupled plasma-optical emission 197 spectrometer (Perkin Elmer, Waltham, MA) was used to analyze water samples for Ca^{2+} , Mg^{2+} , 198 Na^+ , K^+ , and Cl^- . 199

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201 **2.3 Incubation Setup**

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Incubation water treatments included: 1) soils incubated at 65 % water holding capacity (WHC) (Dry); 2) soils incubated at 100% WHC with freshwater (0 ppt); 3) soils incubated at 100% WHC with a saltwater concentration of 2.5 ppt (2.5 ppt); and 4) soils incubated at 100% WHC with a saltwater concentration of 5.0 ppt (5.0 ppt). A subsample of each fresh soil (soils stored at 4 °C) was dried at 105°C to constant mass to determine gravimetric soil water content.

208 Approximately 150 - 200 g fresh soil (20 - 25 g dry weight) collected from each plot was weighed into 1 L canning jars. For water addition estimates, WHC was calculated by placing a 209 subsample of fresh soil (approximately 2 g fresh weight) in a funnel with a Whatman #1 filter 210 and saturating with deionized H_2O (d H_2O). The saturated sample was allowed to drain into a 211 conical flask for 2 h. After 2 h, the saturated soil was weighed, dried at 105 °C to constant mass, 212 and weighed again to determine WHC. It is important to note that the 100% WHC moisture 213 level resulted in soils being completely flooded (with either freshwater or saltwater) with water 214 covering the surface of the incubated soils, thereby allowing for the development of CH₄ 215 216 producing conditions similar to that observed in the field for surface soils. After soil and water additions, the remaining headspace was estimated for each individual incubation vessel 217 (approximately 750 mL) and used in the calculation of gas production rates. Following wood-218 219 additions (see below), incubation vessels from each of the eight treatments were incubated in the dark in the laboratory for 98 d at 20 - 23 °C. 220

Two sets of incubations were set up with the above mentioned water treatments. We 221 222 added ¹³C-depleted American sweetgum (*Liquidamber styraciflua*) wood to half the incubation 223 vessels (0.22 g wood per g soil) (wood-amended), while the other half were incubated without 224 wood (wood-free). Trees were grown at the Duke FACE site under elevated CO₂ concentrations (200 ppm CO_2 above ambient) using natural gas derived CO_2 with a depleted ¹³C signature 225 compared to that of the atmosphere (Feng et al., 2010; Schlesinger et al., 2006). The site was 226 227 established in 1983 after clear cut and burn (Kim et al., 2016). Trees were grown under elevated CO_2 from 1994 to 2010 at which point they were harvested (Kim et al., 2016). Cookies were 228 removed from harvested trees, dried to a constant moisture level and stored at -20 °C until use. 229 230 The bark layer was removed and the outer six tree rings of multiple cookies were removed with a

231 chisel. Wood was then finely ground in a Wiley Mill (Thomas Scientific, Swedesboro, NJ, USA) and analyzed for C content and ¹³C signature on a Picarro G2201-i Isotopic CO₂/CH₄ 232 Analyzer outfitted with a Costech combustion module for solid sample analysis (Picarro Inc., 233 Sunnyvale, CA USA). For δ^{13} C analysis of solids (e.g., wood, microbial biomass extracts, soils), 234 certified solid standards were used to develop a standard curve from the expected and measured 235 δ^{13} C values (R² > 0.999). These standards included USGS 40 (L-glutamic acid) (δ^{13} C = -26.39 236 ‰; USGS Reston Stable Isotope Laboratory, Reston, VA, USA), protein ($\delta^{13}C = -26.98$ ‰; 237 Elemental Microanalysis Ltd, Okehampton, UK), urea ($\delta^{13}C = -48.63$ ‰; Elemental 238 Microanalysis Ltd, Okehampton, UK), atropine ($\delta^{13}C = -18.96$ %; Costech Analytical 239 Technologies, Inc, Valencia, CA, USA), and acetanilide ($\delta^{13}C = -28.10$ ‰; Costech Analytical 240 Technologies, Inc, Valencia, CA, USA). For C concentration, atropine standards were weighed 241 out over a range of C concentrations that encompassed the expected C concentrations of the 242 unknown samples and within the measurement range of the instrument. A standard curve for C 243 concentration was also developed from the expected and measured C concentration of the 244 atropine standards ($R^2 > 0.99$). All unknown sample's C concentration and δ^{13} C value were 245 adjusted using the linear equations derived from the appropriate standard curve. The δ^{13} C values 246 were reported in parts per thousand (‰) relative to the Vienna Pee Dee Belemnite (VPDB) 247 standard. Wood had a C content of 45.6 ± 0.21 % and δ^{13} C value of -40.7 ± 0.06 ‰, which was 248 within the range of -42 to -39 ‰ measured on fresh pine needles and fine roots (Schlesinger et 249 al., 2006), and more depleted in ¹³C compared to that measured in hummock surface soils from 250 our site (-29.1 \pm 0.29 ‰; Minick et al. 2019b). 251

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253 2.4 CO₂ and CH₄ Sample Collection and Analysis

255	Headspace gas samples were collected from incubation vessels 15 times over the course
256	of the 98 d incubation (days 1, 4, 8, 11, 15, 19, 25, 29, 29, 47, 56, 63, 70, 84, 98). Incubation lids
257	were loosened between measurements to allow for gas exchange with the ambient atmosphere.
258	Four blank incubations (empty jars; no soil, water, or wood) were set up and treated in the exact
259	same manner as incubations containing soil, water, and wood. Blanks were used to measure soil-
260	free CO ₂ and CH ₄ concentrations in incubations, which were always well below the detection
261	limit of the gas analyzer (described below). Prior to each measurement, incubation vessels were
262	removed from the dark, sealed tightly, and flushed at 20 psi for three minutes with CO ₂ /CH ₄ free
263	zero air (Airgas, Radnor, PA, USA). Following flushing, incubation vessels were immediately
264	placed back in the dark (2-6 h over the first 39 days and 12-18 h over the remainder of the
265	incubation) before taking a gas sample for analysis. Approximately 300 mL of headspace gas
266	was removed using a 50 mL gas-tight syringe and transferred to an evacuated 0.5 L Tedlar gas
267	sampling bag (Restek, Bellefonte, PA, USA). Simultaneous analysis of CO2 and CH4
268	concentrations and $\delta^{13}C$ isotopic signature were conducted on a Picarro G2201-i Isotopic
269	CO_2/CH_4 Analyzer (Picarro Inc., Sunnyvale, CA USA). For $\delta^{13}C$ analysis of gases (e.g., CO_2
270	and CH ₄), certified gas standards were used to develop a standard curve from the expected and
271	measured δ^{13} C values (R ² > 0.99). The gas standards for ¹³ CO ₂ analysis included gas tanks
272	containing: 1) 372 ppm CO ₂ with a δ^{13} C value of -11.0 ± 0.25 ‰ (Airgas, Inc., Radnor, PA); 2)
273	420 ppm CO ₂ with a δ^{13} C value of -10.3 ± 0.18 ‰ (Airgas, Inc., Radnor, PA); 3) 768 ppm CO ₂
274	with a $\delta^{13}C$ value of -29.5 \pm 0.14 ‰ (Airgas, Inc., Radnor, PA); and 4) 3000 ppm CO_2 with a
275	δ^{13} C value of -34.4 ± 0.3 ‰ (Airgas, Inc., Radnor, PA). The gas standards for 13 CH ₄ analysis
276	included gas tanks containing: 1) 1.75 ppm CH ₄ with a δ^{13} C value of -43.2 ± 0.07 ‰ (Airgas,

Inc., Radnor, PA); 2) 2.00 ppm CH₄ with a δ^{13} C value of -42.7 ± 0.20 ‰ (Airgas, Inc., Radnor, 277 PA); 3) 10.00 ppm CH₄ with a δ^{13} C value of -68.6 ± 1.00 ‰ (Airgas, Inc., Radnor, PA); and 4) 278 15.08 ppm CH₄ with a δ^{13} C value of -29.5 ± 0.14 ‰ (Airgas, Inc., Radnor, PA). For CO₂ and 279 280 CH₄ concentration, a concentrated gas standard (gas mix containing 4043 ppm CO₂ and CH₄) (Airgas, Inc., Radnor, PA) was diluted with zero air gas, providing a range of CO₂ and CH₄ 281 concentrations that encompassed the expected gas concentrations of the unknown samples. A 282 standard curve for gas concentration was developed from the expected and measured gas 283 concentration of the diluted gas standards ($R^2 > 0.99$). All unknown gas sample CO₂ and CH₄ 284 concentrations and δ^{13} C values were adjusted using the linear equations derived from the 285 appropriate standard curve. The δ^{13} C values were reported in parts per thousand (‰) relative to 286 the Vienna Pee Dee Belemnite (VPDB) standard. Production rates of CO₂-C and CH₄-C were 287 288 calculated as well as daily cumulative CO₂-C and CH₄-C production summed over the course of the 98 d incubation. Small subsamples (approximately 1.0 g dry weight) of soil were removed 289 periodically from each incubation vessel for extracellular enzyme analysis (see below). Removal 290 of soil was accounted for in subsequent calculations of gas production rates. Incubation vessel 291 292 water levels (mass basis) were checked and adjusted three times per week using either freshwater or saltwater. 293

The proportion of wood-derived CO_2 at each sampling date was calculated using ${}^{13}CO_2$ data and the ${}^{13}C$ of depleted wood (-40.07) in a two pool flux model (Fry 2006), with the depleted wood signature as one end-point and the ${}^{13}CO_2$ of wood-free incubations as the other endpoint.

299 % C = ((
$$\delta^{13}CO_{2wood + soil} - \delta^{13}CO_{2wood - free soil}$$
) / ($\delta^{13}C_{wood} - \delta^{13}CO_{2wood - free soil}$)) *100

Where $\delta^{13}CO_{2wood+ soil}$ is the $\delta^{13}C$ value of CO₂ produced from soils incubated with the addition of ¹³C-depleted wood, $\delta^{13}C_{wood-free soil}$ is the $\delta^{13}C$ value of CO₂ produced from soils incubated without the addition of ¹³C-depleted wood, and $\delta^{13}C_{wood}$ is the average $\delta^{13}C$ value of the ¹³C-depleted wood. Total wood-derived CO₂ was calculated using cumulative CO₂ produced over the 98 d incubation and the average ¹³CO₂ across the whole incubation.

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307 2.5 Soil Characteristics

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Soil organic C concentration and δ^{13} C were analyzed on the four replicate soil samples 309 prior to the start of the incubation (initial soil samples) and on soils from each of the thirty 310 311 incubations following the 98 d incubation period. The initial C analysis was performed on samples removed prior to incubation. Soils were finely ground in a Wiley Mill (Thomas 312 Scientific, Swedesboro, NJ, USA) prior to analysis on a Picarro G2201-i Isotopic CO₂/CH₄ 313 Analyzer outfitted with a Costech combustion module for solid sample analysis (Picarro Inc., 314 Sunnyvale, CA USA). Carbon concentration and ¹³C calibration standards were the same as 315 those described for the analysis of the ¹³C-depleted wood. 316

Soil pH and redox potential (Eh = mV) were measured in each incubation within one hour following sampling of headspace gas. Soil pH was measured on the four replicate soil samples immediately prior to the start of the incubation with a glass electrode in a 1:2 mixture (by mass) of soil and distilled water (dH₂O). Soil redox potential (Eh = mV) was measured using a Martini ORP 57 ORP/°C/°F meter (Milwaukee Instruments, Inc., Rocky Mount, NC, USA).

324 2.6 Microbial Biomass Carbon and $\delta^{13}C$ Isotopic Signature

326	Microbial biomass C (MBC) was estimated on soils collected from incubations on day 1
327	(after 24 hour post-treatment incubation) and day 98 (following the end of the incubation). The
328	chloroform fumigation extraction (CFE) method was adapted from Vance et al. (1987) in order
329	to estimate MBC and δ^{13} C. Briefly, one subsample of soil (approximately 0.5 g dry weight each)
330	was placed in a 50 mL beaker in a vacuum desiccator to be fumigated. Another subsample was
331	placed into an extraction bottle for immediate extraction in 0.5 M K_2SO_4 by shaking for 1 hr and
332	subsequently filtering through Whatman #2 filter paper to remove soil particles. The samples in
333	the desiccator were fumigated with ethanol-free chloroform $(CHCl_3)$ and incubated under
334	vacuum for 3 d. After the 3 d fumigation, samples were extracted similar to that of non-
335	fumigated samples. Filtered 0.5 M K_2SO_4 extracts were dried at 60 °C in a ventilated drying
336	oven and then ground to a fine powder with mortar and pestle before analysis of C concentration
337	and $\delta^{13}C$ on a Picarro G2201-i Isotopic CO ₂ /CH ₄ Analyzer outfitted with a Costech combustion
338	module for solid sample analysis (Picarro Inc., Sunnyvale, CA USA). Carbon concentration and
339	¹³ C calibration standards were the same as those described for the analysis of the ¹³ C-depleted
340	wood. Microbial C biomass was determined using the following equation:
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$342 \qquad \qquad MBC = EC / k_{EC}$

344 where the chloroform-labile pool (EC) is the difference between C in the fumigated and non-fumigated extracts, and k_{EC} (extractable portion of MBC after fumigation) is soil-specific 345 and estimated as 0.45 (Joergensen, 1996). 346 The δ^{13} C of MBC was estimated as the δ^{13} C of the C extracted from the funigated soil 347 sample in excess of that extracted from the non-fumigated soil sample using the following 348 equation: 349 350 $\delta^{13}C_{MBC}$ (‰) = (($\delta^{13}C_f \ge C_f$) – ($\delta^{13}C_{nf} \ge C_{nf}$)/($C_f - C_{nf}$) 351 352 where C_f and C_{nf} is the concentration (mg kg⁻¹ soil) of C extracted from the fumigated 353 and non-fumigated soil samples, respectively, and $\delta^{13}C_f$ and $\delta^{13}C_{nf}$ is the ¹³C natural abundance 354 355 (‰) of the fumigated and non-fumigated soil samples, respectively. 356 2.5 Extracellular Enzyme Analysis 357 358 The potential activity of five extracellular enzymes was quantified on soil samples and on 359 360 days 1, 8, 35, and 98 of the soil incubation. The enzymes chosen for this experiment represent a range of compounds in which they degrade, including fast and slow cycling C compounds, as 361 well as ones that target nitrogen (N), phosphorus (P), and sulfate (S) containing compounds. The 362 363 Enzyme Commission number (EC) is stated in parenthesis after each enzyme, which classifies them by the chemical reaction catalyzed by each enzyme. The specific enzymes measured were: 364 β-glucosidase (BG; EC: 3.2.1.21), xylosidase (XYL; EC 3.2.1.37), peroxidase (PER; EC: 365

366 1.11.1.7), β -glucosaminidase (NAGase; EC: 3.2.1.30), alkaline phosphatase (AP; EC: 3.1.3.1),

and arylsulfatase (AS; EC: 3.1.6.1). Carbon-degrading enzymes BG, XYL, and PER degrade
sugar, hemicellulose, and lignin, respectively, while the N-degrading enzyme, NAGase, degrades
chitin. Enzymes AP and AS degrade phosphorus and SO₄²⁻ containing compounds, respectively.
Substrates for all enzyme assays were dissolved in 50 mM, pH 5.0 acetate buffer solution for a
final concentration of 5 mM substrate.

Hydrolytic enzymes (BG, XYL, NAGase, AP, and AS) were measured using techniques 372 outlined in Sinsabaugh et al. (1993). Approximately 0.8 g dry weight of soil sample was 373 suspended in 50 mL of a 50 mM, pH 5.0 acetate buffer solution and homogenized in a blender 374 375 for 1 min. In a 2 mL centrifuge tube, a 0.9 mL aliquot of the soil-buffer suspension was combined with 0.9 mL of the appropriate 5 mM p-nitrophenyl substrate solution for a total of 376 three analytical replicates. Additionally, duplicate background controls consisting of 0.9 mL 377 aliquot of soil-buffer suspension plus 0.9 mL of acetate buffer were analyzed, as well asfour 378 substrate controls consisting of 0.9 mL substrate solution plus 0.9 mL buffer. The samples were 379 agitated for 2-5 hr. Samples were then centrifuged at 8,160 g for 3 min. Supernatant (1.5 mL) 380 was transferred to a 15 mL centrifuge tube containing 150 µL 1.0 M NaOH, followed by the 381 addition of 8.35 mL dH₂O. The resulting mixture was vortexed and a subsample transferred to a 382 383 cuvette and the optical density at 410 nm was measured on a spectrophotometer (Beckman Coulter DU 800 Spectrophotometer, Brea, CA, USA). 384

The oxidative enzyme (PER) was measured using techniques outlined in Sinsabaugh et al. (1992). PER is primarily involved in oxidation of phenolic compounds and depolymerization of lignin. The same general procedure for hydrolytic enzymes was followed utilizing a 5 mM L-3,4-Dihydroxyphenylalanine (L-DOPA) (Sigma-Aldrich Co. LLC, St. Louis, MO, USA) solution as the substrate plus the addition of 0.2 mL of 0.3% H₂O₂ to all sample replicates and substrate

controls. After set up of analytical replicates and substrate and background controls, the samples
were agitated for 2-3 hr. Samples were then centrifuged at 8,160 g for 3 min. The resulting
supernatant turns an intense indigo color. Supernatant (1.4 mL) was transferred directly to a
cuvette and the optical density at 460 nm was measured on a spectrophotometer.

For all enzymes, the mean absorbance of two background controls and four substrate controls was subtracted from that of three analytical replicates and divided by the molar efficiency (1.66/µmol), length of incubation (h), and soil dry weight. Enzyme activity was expressed as µmol substrate converted per g dry soil mass per hour (µmol g⁻¹ h⁻¹). Daily cumulative enzyme activity was calculated and summed over the course of the 98 d incubation.

399

400 **2.6 Statistical Analysis**

401

Water chemistry, cumulative CO₂ production, cumulative CH₄ production, cumulative 402 enzyme activity, post-incubation SOC concentration and δ^{13} C, and wood-derived and wood-403 associated SOC, CO₂, and MBC were analyzed using a one-way ANOVA (PROC GLM 404 package). Microbial biomass C, MBC ¹³C, pH, Eh, δ^{13} CO₂, and δ^{13} CH₄ were analyzed using 405 406 repeated-measures ANOVA (PROC MIXED package) with time (Time) as the repeated measure and the incubation treatment as the fixed effect. All data for wood-free and wood-amended soils 407 were analyzed separately. Raw data were natural log-transformed where necessary to establish 408 409 homogeneity of variance. If significant main effects or interactions were identified in the oneway or repeated-measures ANOVA (P < 0.05), then post-hoc comparison of least-squares means 410 was performed. All statistical analyses were performed using SAS 9.4 software (SAS Institute, 411 412 Cary, NC, USA).

414	3 Results
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415

416 **3.1 Water and Soil Properties**

417

Freshwater had higher concentrations of TOC compared to the saltwater treatments
(Table 1). Concentration of SO₄²⁻, Cl⁻, Na⁺, Ca²⁺, Mg²⁺, and K⁺ were higher in saltwater
treatments compared to freshwater and were approximately twice as high in the 5.0 ppt saltwater
treatment compared to 2.5 ppt saltwater (Table 1).

Initial (pre-incubation) SOC concentration was 490 ± 27 g kg⁻¹ with a δ^{13} C value of -28.5 ± 0.32 ‰. After 98 d of incubation, SOC concentration in wood-free incubations was lower in the 5.0 ppt saltwater treatment, although no difference in soil δ^{13} C was found between treatments (Table 2). For wood-amended incubations, post-incubation SOC concentration was lower in the 5.0 ppt saltwater treatment compared to the dry and freshwater treatment (Table 2). Overall, the δ^{13} C of wood-free (-29.5 ± 0.08 ‰) and wood-amended soils (-30.5 ± 0.12 ‰) after 98 days of incubation were significantly different (F = 49.6; *P* < 0.0001).

Soil pH was significantly lower in the saltwater treatments in both wood-free and woodamended soils compared to the dry and freshwater treatments (Table 3; Figure 2A-B). After an initial drop of pH in saltwater treatments (wood-free and wood-amended) to between 3.2 and 3.4 pH, pH steadily climbed back up to between 3.8 and 4.2 pH (Figure 2A-B). In wood-free soils, differences in soil Eh between treatments was variable over time, with both the 5.0 ppt saltwater treatment and the freshwater treatment having the lowest redox potential at different time points throughout the incubation (Table 3; Figure 2C), but fell below -124 mV on average. In wood-

436	amended soils, Eh dropped quickly to between -200 and -400 mV over the first 30 days for
437	saltwater incubated soils (Table 3; Figure 2D), before rising to between -100 to 0 mV for the rest
438	of the incubation period. In freshwater incubated soils, Eh rose quickly back to between -50 to 50
439	mV by day 15 and remained at this level for the rest of the incubation period, while saltwater
440	treatments had significantly lower Eh between days 8 and 25.
441	
442	3.2 CO ₂ , CH ₄ , δ ¹³ CO ₂ -C, and δ ¹³ CH ₄ -C
443	
444	In wood-free incubations, cumulative CO ₂ production was not different between the dry
445	and freshwater treatments, but was higher than that produced from saltwater treatments (Table 4;
446	Figure 3A). Cumulative CO ₂ produced from wood-amended soils was highest in the dry
446 447	Figure 3A). Cumulative CO ₂ produced from wood-amended soils was highest in the dry treatment compared to all other treatments (Table 4; Figure 3B). Wood-derived CO ₂ (calculated
446 447 448	Figure 3A). Cumulative CO ₂ produced from wood-amended soils was highest in the dry treatment compared to all other treatments (Table 4; Figure 3B). Wood-derived CO ₂ (calculated as the difference between cumulative CO ₂ produced from wood-amended and wood-free
446 447 448 449	Figure 3A). Cumulative CO ₂ produced from wood-amended soils was highest in the dry treatment compared to all other treatments (Table 4; Figure 3B). Wood-derived CO ₂ (calculated as the difference between cumulative CO ₂ produced from wood-amended and wood-free incubations) was highest in the dry treatment (Table 4; Figure 3C). This finding was also

450 confirmed by calculating cumulative wood-derived C using the ¹³C two-pool mixing model, with 451 the highest proportion found in the dry treatment (54 \pm 4.6 %) compared to soils incubated with

452 freshwater (42 \pm 1.7 %), 2.5 ppt saltwater (37 \pm 1.0 %), and 5.0 ppt saltwater (38 \pm 1.5 %) (F =

453 10.1; P = 0.001).

454 Cumulative CH₄ production was highest in the freshwater treatment compared to the
455 saltwater treatments in both wood-free and wood-amended incubations (Table 4; Figure 3D-E).
456 The difference between cumulative CH₄ produced from wood-amended and wood-free
457 incubations was lower (and exhibited a negative response to wood additions) in the freshwater

treatment compared to both saltwater treatments (Table 3; Figure 3F), which both had a slightpositive response to wood additions.

460	The CO ₂ :CH ₄ ratio, in wood-free incubations, was calculated only for soils incubated
461	under saturated conditions with freshwater or saltwater. The CO _{2:} CH ₄ ratio, in wood-free
462	incubations, was highest in freshwater (6 \pm 3.4), compared to the 2.5 ppt saltwater (136 \pm 33.9)
463	and 5.0 ppt saltwater (102 \pm 30.3) (F = 24.8; P = 0.0002). The CO ₂ :CH ₄ ratio, in wood-amended
464	incubations, was highest in freshwater (9 \pm 0.8), compared to the 2.5 ppt saltwater (53 \pm 20.3)
465	and 5.0 ppt saltwater (107 \pm 37.7) (F = 9.2; P = 0.007).
466	The δ^{13} CO ₂ -C and wood-derived CO ₂ (estimated by 13 C two-pool mixing model)
467	exhibited a time by treatment interaction for both wood-free and wood-amended incubations
468	(Table 3; Figure 4A-B). In general, $\delta^{13}CO_2$ -C in wood-free and wood-amended incubations was
469	depleted in the dry treatment (and remained steady throughout the incubation period) compared
470	to all other treatments, especially after day 15. The proportion of wood-derived CO ₂ was
471	initially higher in freshwater and saltwater treatments (after day 1) but gradually dropped over
472	the course of the incubation, while the proportion of wood-derived CO ₂ from the dry treatment
473	dropped quickly after the first sampling date (day 1) and remained steady (approximately 50-60
474	%) for the remainder of the incubation period (Figure 4C).

The δ^{13} CH₄-C (Table 3; Figure 5) exhibited a treatment and time effect (Table 3; Figure 5A-B), but only for wood-free incubations. For wood-free incubations, average ¹³CH₄-C across the course of the incubation was enriched in the freshwater treatment (-67.8 ± 2.4 ‰) compared to the 2.5 ppt (-80.1 ± 2.4 ‰) and 5.0 ppt (-82.3 ± 2.0 ‰) saltwater treatments (Figure 5C). No difference in the δ^{13} CH₄-C was found in wood-amended incubations (Figure 4b, d), which ranged from between -78 to -75 ‰ for all treatments.

3.3 Microbial Biomass Carbon and Extracellular Enzyme Activity

484	Initially, MBC was lowest in the dry treatment of wood-free incubations and in the 5 ppt
485	treatment of wood-amended incubations (Table 3; Table 5). Following the 98 day incubation,
486	MBC was highest in the dry treatment of wood-free incubations, with no differences between the
487	other treatments. In wood-amended incubations, final MBC was also highest in the dry
488	treatment compared to both saltwater treatments. Initial $\delta^{13}C$ of MBC did not differ between
489	treatments in either the wood-free or wood amended soils (Table 3; Table 5). After the 98 day
490	incubation, ¹³ C of MBC in the wood-free treatments was depleted in the freshwater treatment
491	and enriched in the 5.0 ppt saltwater treatment. In wood-amended incubations, ¹³ C of MBC was
492	depleted in the dry treatment and enriched in the freshwater and 5.0 ppt saltwater treatments.
493	Furthermore, the proportion of wood-derived MBC (as estimated by ¹³ C mixing model
494	calculations) was highest in the dry treatment (31 %) and the 2.5 ppt saltwater treatment (21%)
495	compared to the freshwater treatment (4%) (Table 5).
496	In wood-free incubations, activity of BG and NAGase was higher, while PER was lower,
497	in the dry treatment compared to the saltwater treatments (Table 4; Table 5). Activity of AS was
498	higher in the dry and freshwater treatments compared to saltwater treatments, in both wood-free
499	and wood-amended incubations. In wood-amended incubations, BG and NAGase were highest
500	in the dry treatment compared to the saltwater treatments. In the freshwater treatment, wood
501	addition reduced activity of BG and NAGase compared to wood-free incubations (Figure 6A-B),
502	but enhanced PER activity (Figure 6C). Wood addition also reduced AS and P activity across all

treatments compared to wood-free incubations (Figure 6D-E).

505 4 Discussion

506

As forests within the lower coastal plain physiographic region of the southeastern US 507 continue to experience increasing stresses from SLR, changes in microbial C cycling processes 508 509 should be expected. Our results, combined with other field and lab experiments, confirm that saltwater intrusion into coastal freshwater forested wetlands can result in reductions in CO2 and 510 CH₄ production (Ardón et al., 2016; Ardón et al., 2018), but this may be balanced by long- and 511 512 short-term effects of saltwater intrusion on these C cycling processes (Weston et al., 2011), as well as changes in C inputs due to forest-to-marsh transition. Further, wood additions to these 513 wetland soils may reduce CH₄ production under freshwater conditions compared to the absence 514 wood additions (Figure 3C and 3F), but slightly enhance CH₄ production under saltwater 515 conditions. Our results also demonstrate that substantial quantities of CH₄ can be produced from 516 517 freshwater wetland soils with redox potential between -100 to 100 mV, which may be related to the specific pathway of CH₄ production (acetoclastic versus hydrogenotrophic) (Angle et al., 518 2016), and challenges the widespread assumption that methanogenesis only occurs at very low 519 520 redox potentials. Changes in the water table depth at the ARNWR is driven primarily by precipitation patterns (Minick et al., 2019a), resulting in the influx of oxygenated waters. 521 522 Periodic *in situ* measurements of redox potential at the ARNWR indicate that standing water is 523 relatively aerated (Eh = 175 - 260 mV), while surface soils of hummocks when not submerged are more aerated (Eh = 320 mV) than submerged hollow surface soils (Eh = 100 to 150 mV) and 524 525 deeper organic soils (20 - 40 cm depth; Eh = 50 to 90 mV) (unpublished data). Furthermore, our 526 results indicate that additions of new C to soils as wood may result in short-term reductions in

redox potential as anaerobic processes are enhanced due to the added C substrate and terminal electron acceptors are quickly reduced. As SLR continues to rise over the next century, more persistent saltwater intrusion may occur as rising brackish waters mix with non-tidal freshwater systems having important implications for both above- and below-ground C cycling dynamics. Although our study only looked at these effects in a controlled laboratory experiment, these data provide a baseline understanding of potential changes in C cycling dynamics in these wetlands due to SLR.

Saltwater additions decreased CO₂ production compared to freshwater in the wood-free 534 535 soils, although post-incubation MBC and extracellular enzyme activity (e.g., BG, NAGase, and AP) were not different between these treatments. This has been found in other pocosin wetland 536 soils on the coast of North Carolina (Ardón et al. 2018). Variable effects of salinity (and/or SO₄²⁻ 537 additions) have been found on soil respiration, with some studies showing an increase (Marton et 538 al., 2012; Weston et al., 2011), a decrease (Lozanovska et al. 2016; Servais et al. 2019), or no 539 540 change (Baldwin et al., 2006). Krauss et al. (2012) found that permanently flooded saltwater treatments (expected in non-tidal wetlands) in a simulated coastal swamp mesocosm reduced soil 541 respiration, whereas saltwater pulses (expected in tidal wetlands) had a variable effect on soil 542 543 respiration. Alternatively, CO_2 production was not reduced in the saltwater compared to freshwater treatments in wood-amended soils, while post-incubation MBC was lower in the 544 saltwater compared to freshwater, which suggests a shift in microbial carbon use efficiency. 545 546 Methane production was higher in the freshwater compared to saltwater treatments in both wood-amended and wood-free incubations. Numerous others studies have found that 547 548 saltwater reduces CH₄ fluxes compared to freshwater, both within the field and laboratory. 549 Reduced CH₄ production from saltwater treated soils primarily results from the availability of

550	more energetically favorable terminal electron acceptors (primarily SO_4^{2-}), which leads to the
551	competitive suppression of methanogenic microbial communities by SO4 ²⁻ reducing communities
552	(Bridgham et al., 2013; Chambers et al., 2011; Winfrey and Zeikus, 1977), as methanogens and
553	SO_4^{2-} reducers compete for acetate and electrons (Le Mer and Roger, 2001). Dang et al. (2019)
554	did find partial recovery over time of the methanogenic community following saltwater
555	inundation to freshwater soil cores, but interestingly this community resembled that of microbes
556	performing hydrogenotrophic methanogenesis and not acetoclastic methanogenesis. Activity of
557	arylsulfatase was also lower in saltwater amended soils. This also indicates a functional change
558	in the microbial community, as microbes in the saltwater treatment are utilizing the readily
559	available SO_4^{2-} pool, while microbes in the freshwater and dry treatments are still actively
560	producing SO ₄ ²⁻ -liberating enzymes to support their metabolic activities. Findings by Baldwin et
561	al. (2006) support the effects of saltwater on changing the microbial community structure as
562	well, in which reductions in CH ₄ production in NaCl treated freshwater sediments were
563	accompanied by a reduction in archaeal (methanogens) microbial population, establishing a link
564	between shifting microbial populations and changing CH4 flux rates due to saltwater intrusion.
565	Changes in the CH ₄ production due to saltwater additions appear to be related to the
566	dominant CH ₄ producing pathway. The ¹³ CH ₄ isotopic signature in wood-free freshwater
567	incubated soils indicated that acetoclastic methanogenesis was the dominant CH4 producing
568	pathway, while hydrogenotrophic methanogenesis dominated in the saltwater treatments.
569	Acetoclastic methanogenesis produces isotopically enriched CH ₄ compared to that of the
570	hydrogenotrophic methanogenesis (Chasar et al., 2000; Conrad et al. 2010; Krohn et al. 2017;
571	Sugimoto and Wada, 1993; Whiticar et al., 1986; Whiticar 1999). The differences in C
572	discrimination between the two pathways is greater for the hydrogenotrophic compared to the

573 acetoclastic pathway, resulting in more depleted (-110 to -60 ‰) and more enriched (-60 ‰ to -50 ‰) ¹³CH₄, respectively. This has been confirmed in field and laboratory experiments (Conrad 574 et al. 2010; Krohn et al. 2017; Krzycki et al., 1987; Sugimoto and Wada, 1993; Whiticar et al., 575 1986; Whiticar, 1999). Baldwin et al. (2006) also found that saltwater additions promoted the 576 hydrogenotrophic methanogenic pathway. Further, recent studies have found that saltwater 577 578 additions to soils result in a shift in the relative abundance of hydrogenotrophic methanogens (Chambers et al. 2011; Dang et al 2019), supporting the idea that saltwater may alter not only the 579 production of CH₄ but also the pathway of methane production. 580

581 Changes in freshwater and saltwater hydrology due to rising seas is leading to dramatic shifts in the dominant plant communities within the ARNWR and across the southeastern US 582 583 (Connor et al., 1997; DOD, 2010; Langston et al., 2017; Kirwan and Gedan 2019). This has the potential to alter the soil C balance due to introduction of large amounts of coarse woody debris 584 as trees die. In our laboratory experiment, additions of wood resulted in changes in both CO₂ 585 586 and CH₄ production, but the direction of change depended on if soils were incubated with freshwater or saltwater. Wood additions increased CO₂ production compared to wood-free soils, 587 except in the freshwater treatment. This was particularly evident in the dry treatment where 588 589 wood additions increased CO₂ production by approximately 32 %. For the dry treatment, woodamended soils had the highest MBC and NAGase activity as microbes were likely immobilizing 590 more N to support metabolic activities in the presence of added C (Fisk et al., 2015; Minick et 591 592 al., 2017). Higher respiration with wood additions in the saltwater treatments likely resulted from enhanced metabolic activity of SO_4^{2-} reducing microbes in the presence of an added C source. 593 594 On the other hand, wood additions resulted in a decline in CH₄ production from the freshwater 595 treatment, while slightly enhancing CH₄ production from the saltwater treatments. Wood

596 additions also resulted in much lower redox potential, particularly in the saltwater treatments, and coupled with ¹³CH₄ stable isotope composition may have driven the higher levels of CH₄ 597 production (via hydrogenotrophic methanogenesis) in the wood plus saltwater treatments. The 598 suppression of CH₄ production by wood additions in the freshwater treatment was somewhat 599 surprising given the positive effects of C additions on CH₄ production recently found in 600 601 freshwater sediments (West et al. 2012), but likely resulted from enhancement of other, more energetically favorable redox reactions with the addition of a C source (e.g., wood). Furthermore, 602 wood additions to freshwater incubations resulted in a decrease in MBC and activity of BG and 603 604 NAGase enzymes compared to wood-free incubations and an increase in PER activity. This suggests that the microbial communities have altered their functional capacity in response to 605 wood additions when exposed to freshwater. The CO₂:CH₄ ratio further indicated that, in 606 freshwater, CH₄ production was quite high in relation to CO₂ production. This ratio was 607 significantly higher for saltwater treatments as CH₄ production dropped drastically compared to 608 609 freshwater. In wood-free incubations, the CO₂:CH₄ trend between freshwater and saltwater treatments was parabolic, but was linear upward in wood-amended soils. This suggests that 610 interactions between saltwater concentration and coarse woody debris (in the form of dead and 611 612 dying trees; Kirwan and Gedan 2019) may be important to understand when determining effects 613 of saltwater intrusion on greenhouse gas production in freshwater forested wetlands. 614 Findings from this study indicate that substantial changes in the greenhouse gas 615 production and microbial activity are possible due to saltwater intrusion into freshwater wetland ecosystems but that the availability of C in the form of dead wood (as forests transition to marsh) 616

618 wetlands, saltwater intrusion may reduce both CO_2 and CH_4 emissions from soils to the

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may alter the magnitude of this effect. At ARNWR and similar coastal freshwater forested

619	atmosphere. Sea-level rise will likely lead to dramatic and visually striking changes in
620	vegetation, particularly transitioning forested wetlands into shrub or marsh wetlands (Kirwan and
621	Gedan 2019), which has resulted in the widespread occurrence of "ghost" forests along the
622	Atlantic coast (Kirwan and Gedan 2019). As forested wetlands are lost, dead trees could provide
623	a significant source of C to already C-rich peat soils, with the potential to alter CO ₂ and CH ₄
624	production. The long-term effect of forest-to-marsh transition on ecosystem C storage will likely
625	depend on the balance between dead wood inputs and effects of SLR and vegetation change on
626	future C inputs and soil microbial C cycling processes. Future work should include investigation
627	of these C cycling and microbial processes at the field-scale and expand to a wider range of non-
628	tidal wetlands within the southeastern US region.
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630	Author contribution
631	
632	All authors contributed to the conception and design of the study. KM wrote the first draft of the
633	manuscript. KM collected the samples from the field and performed laboratory analysis. All
634	authors contributed to manuscript revision and approved the submitted version.
635	
636	Competing Interest
637	
638	The authors declare that they have no conflict of interest.
639	
640	Acknowledgements
641	

642	We thank numerous undergraduate researchers for their invaluable help collecting samples from
643	the field and analyzing samples in the laboratory. We also thank two reviewers for their
644	comments, which significantly improved the manuscript. Primary support was provided by
645	USDA NIFA (Multi-agency A.5 Carbon Cycle Science Program) award 2014-67003-
646	22068. Additional support was provided by DOE NICCR award 08-SC-NICCR-1072, the
647	USDA Forest Service Eastern Forest Environmental Threat Assessment Center award 13-JV-
648	11330110-081, and Carolinas Integrated Sciences and Assessments award 2013-0190/13-
649	2322. The USFWS Alligator River National Wildlife Refuge provided helpful scientific
650	discussions, the forested wetland research site, and valuable in-kind support.
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Tables and Figures

Table 1. Total organic C (TOC) and ion concentrations (mg L⁻¹) in freshwater (0 ppt), 2.5 ppt saltwater, and 5.0 ppt saltwater. Standard errors of the mean are in parenthesis (n=4). Values with different superscript lowercase letters are significantly different (P < P

0.05).

Treatment	TOC	SO_4^{2-}	Cl-	Na^+	$\mathrm{NH_4^+}$	NO_3^-	PO_4^{3-}	Ca ²⁺	Mg^{2+}	\mathbf{K}^+
0 ppt 2.5 ppt 5.0 ppt	44 (0.3) ^a 40 (0.7) ^b 38 (0.1) ^b	$1 (0.1)^{a}$ 162 (1.3) ^b 319 (6.5) ^c	17 (0.2) ^a 1391 (42.8) ^b 2695 (22.6) ^c	8 (0.1) ^a 538 (19.2) ^b 1039 (15.9) ^c	0.00 (0.000) ^a 0.06 (0.004) ^b 0.07 (0.004) ^b	$\begin{array}{c} 0.00 \ (0.000)^{\mathbf{a}} \\ 0.06 \ (0.000)^{\mathbf{a}} \\ 0.07 \ (0.004)^{\mathbf{a}} \end{array}$	0.00 (0.000) ^a 0.01 (0.000) ^a 0.01 (0.000) ^b	$1 (0.0)^{a}$ 23 (0.3) ^b 44 (1.0) ^c	$1 (0.0)^{a}$ 64 (2.6) ^b 125 (2.1) ^c	0.2 (0.0) ^a 19 (0.3) ^b 36 (0.4) ^c
841	00(011)	017 (010)	2000 (22.0)	1003 (1013)				(110)	120 (211)	
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Table 2. Post-incubation soil organic C (SOC) concentration (g kg⁻¹), SOC δ^{13} C (‰), and wood-derived SOC (%) (estimated from ¹³C

two pool mixing model) for soil samples collected from the field and incubated for 98 d in the laboratory under dry conditions (Dry)

or fully saturated with freshwater (0 ppt) or saltwater (2.5 and 5.0 ppt) and with (+ Wood) or without addition of 13 C-depleted wood.

855 Standard errors of the mean are in parenthesis (n=4). Data from wood-free and wood-amended soils were analyzed separately. Values

followed by different superscript lowercase letters are significantly different between the four treatments of the wood-free or wood-

857 amended soils (P < 0.05).

Treatment	Post-SOC Concentration (g kg ⁻¹)	Post-SOC δ ¹³ C (‰)	Wood-derived SOC (%)
Dry	495 (1.5) ^b	-29.5 (0.20) ^a	•
0 ppt	493 (3.3) ^b	-29.5 (0.18) ^a	
2.5 ppt	488 (4.9) ^b	-29.5 (0.20) ^a	
5.0 ppt	460 (8.6) ^a	-29.5 (0.16) ^a	
Dry + Wood	491 (4.7) ^{ab}	-30.4 (0.30) ^a	8 (2.5)
0 ppt + Wood	502 (4.6) ^a	-30.7 (0.22) ^a	12 (0.4)
2.5 ppt + Wood	477 (4.9) ^{bc}	-30.6 (0.35) ^a	10 (1.4)
5.0 ppt + Wood	470 (4.6) ^c	-30.4 (0.14) ^a	10 (2.0)

Table 3. Results (F-values and significance) from the repeated measures ANOVA of pH, Eh, microbial biomass C (MBC), δ^{13} C

isotopic signature of MBC, δ^{13} CO₂, and δ^{13} CH₄ measured in soils collected from a coastal freshwater forested wetland and incubated

in the laboratory for 98 d under fully saturated conditions with either freshwater or saltwater (2.5 ppt and 5.0 ppt). Data from wood-

874 free and wood-amended soils were analyzed separately.

Source	pН	Eh	MBC	MBC ¹³ C	$\delta^{13}CO_2$	δ ¹³ CH
Wood-Free						
Treatment	26.6***	4.5*	3.7*	3.2*	351.7***	60.5**
Time	4.4***	40.7***	40.9***	15.8**	24.2***	8.3***
Treatment x Treatment	1.22	3.7***	27.3***	3.3*	6.4***	1.1
Wood-Amended						
Treatment	29.0***	13.6***	39.9***	2.6	129.8***	0.3
Time	18.3***	30.1***	111.0***	3.7	34.8***	1.4
Treatment x Treatment	1.4	3.4***	24.2***	5.5**	8 3***	1.0
k	*P < 0.05,	**P<0.01,	***P<0.00	001	0.0	1.0
*	*P < 0.05,	**P < 0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
3	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
2	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P<0.01,	***P<0.00	001	0.5	1.0
8	*P < 0.05,	**P < 0.01,	***P<0.00	001	0.5	1.0

Table 4. Results (F-values and significance) from the one-way ANOVA of cumulative gas production and extracellular enzyme
activity (BG: β-glucosidase; PER: peroxidase; NAGase: glucosaminidase; AP: alkaline phosphatase; and AS: arylsulfatase) from soils
collected from a coastal freshwater forested wetland and incubated in the laboratory for 98 d under dry conditions or fully saturated
conditions with either freshwater or saltwater (2.5 ppt and 5.0 ppt). Data from wood-free and wood-amended soils were analyzed
separately.

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	Source	CO_2	CH ₄	BG	PER	NAGase	AP	AS
	Wood-Free							
	Treatment	20.4***	15.6***	7.2**	11.9**	9.5**	0.9	15.8**
	Wood-Amended							
	Treatment	13.3**	36.7***	16.6**	2.5	32.0***	2.3	31.2***
897	*P < 0.05, **H	P < 0.01, *** I	P < 0.0001					
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Table 5. Initial (1 d) and final (98 d) microbial biomass C (MBC) (mg kg⁻¹), MBC δ^{13} C (‰), wood-derived MBC (%) (estimated

916 using ¹³C two pool mixing model), and cumulative extracellular enzyme activity (μ mol g⁻¹) (BG: β -glucosidase; PER: peroxidase;

917 NAGase: glucosaminidase; AP: alkaline phosphatase; and AS: arylsulfatase) for soils incubated under dry conditions (Dry) or

saturated conditions with freshwater (0 ppt) or saltwater (2.5 and 5.0 ppt) and with (+ Wood) or without addition of 13 C-depleted

919 wood. Standard errors of the mean are in parenthesis (n=4). Values followed by different superscript lowercase letters are

920 significantly different between the four treatments for the wood-free or wood-amended soils (P < 0.05).

Treatment	Initial MBC Concentration (mg kg ⁻¹)	Final MBC Concentration (mg kg ⁻¹)	Initial MBC δ ¹³ C (‰)	Final MBC δ^{13} C (‰)	Wood- derived MBC (%)	BG	PER	NAGase	AP	AS
Dry	2238 (400) ^c	4077 (387) ^a	-27.0 (0.43) ^a	-28.4 (0.28) ^{ab}		547 (37) ^a	176 (14) ^a	240 (20) ^a	7599 (1038) ^a	47 (2) ^a
0 ppt	3982 (196) ^{ab}	2657 (344) ^b	-27.3 (0.19) ^a	-28.9 (0.16) ^a		479 (18) ^{ab}	197 (38) ^a	194 (11) ^{ab}	6308 (517) ^a	47 (8) ^a
2.5 ppt	7334 (1177) ^a	2495 (195) ^b	-27.8 (0.51) ^a	-27.9 (0.03) ^{ab}		389 (33) ^b	412 (75) ^b	159 (9) ^b	6539 (183) ^a	19 (3) ^b
5.0 ppt	6483 (104) ^{ab}	2114 (135) ^b	-27.0 (0.30) ^a	-27.4 (0.15) ^b		379 (27) ^b	490 (30) ^b	154 (8) ^b	6387 (529) ^a	15 (2) ^b
Dry + Wood	4444 (579) ^a	5174 (249) ^a	-29.3 (0.40) ^a	-32.1 (0.44) ^a	31 (4.9) ^a	554 (37) ^a	243 (22) ^a	275 (17) ^a	7247 (887) ^a	40 (2) ^a
0 ppt + Wood	5376 (330) ^a	1832 (102) ^b	-29.8 (0.37) ^a	-29.4 (0.15 ^b	4 (1.1) ^b	349 (24) ^b	275 (44) ^a	153 (11) ^b	4965 (459) ^a	36 (3) ^a
2.5 ppt + Wood	5173 (405) ^a	748 (124) ^c	-30.1 (0.25) ^a	-30.4 (0.95) ^{ab}	21 (7.8) ^a	368 (12) ^b	365 (30) ^a	150 (6) ^b	5548 (653) ^a	14 (3) ^b
5.0 ppt + Wood	2123 (400) ^b	790 (87) ^c	-29.9 (0.43) ^a	-29.7 (0.37) ^b	18 (1.9) ^{ab}	369 (13) ^b	326 (38) ^a	150 (6) ^b	5893 (495) ^a	13 (2) ^b
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- 923 Figure 1. Location of the Alligator River National Wildlife Refuge (ARNWR) in eastern North
- 924 Carolina (NC) and surrounding water bodies. The enlarged map shows surrounding freshwater
- 925 (Alligator River and Albermarle Sound) and saltwater (Pamlico Sound, Croatan Sound, and
- 926 Roanoke Sound) bodies. The star represents the approximate location of soil and freshwater
- 927 (from Milltail Creek) sampling locations within the freshwater forested wetlands of ARNWR.
- 928 The black circle represents the approximate location of saltwater sampling (at the Melvin
- Daniels Bridge, Roanoke Sound) from the Roanoke Sound. The saltwater was sampled
- 930 approximately 20 miles east of the soil and freshwater samples.



Figure 2. pH for wood-free soils (A) and wood-amended soils (B) and redox potential for woodfree soils (C) and wood-amended soils (D) measured over the course of the 98 d laboratory incubation. Symbols represent mean with standard error (n=4). Treatment means with different lowercase letters are significantly different within a sampling time point (P < 0.05).



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Figure 3. Cumulative CO₂ production from wood-free soils (A), wood-amended soils (B), and the wood-associated CO₂ production (C); and cumulative CH₄ production for wood-free soils (D), wood amended soils (E), and the wood-associated CH₄ production (F). Panels C and F refer to the difference between wood-amended and wood-free soils. Bars represent mean with standard error (n=4). Bars with different uppercase letters are significantly different (P < 0.05).

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Figure 4. The δ^{13} CO₂ values measured over the course of the 98 d laboratory incubation for wood-free soils (A), wood-amended soils (B), and the proportion of wood-derived CO₂ (C). Bars represent mean with standard error (n=4). Treatment means with different lowercase letters are significantly different within a sampling time point (*P* < 0.05).



990Figure 5. The δ^{13} CH₄ values measured over the course of the 98 d laboratory incubation for991wood-free soils (A) and wood-amended soils (B) and the average δ^{13} CH₄ across the entire992incubation for wood-free soils (C) and wood-amended soils (D). Symbols or bars represent993mean with standard error (n=4). Treatment means with different lowercase letters are994significantly different within a sampling time point (P < 0.05).</td>





1014Figure 6. Wood-associated (wood-amended – wood-free) enzyme activity. Bars represent mean1015with standard error (n=4). Treatment means with different upper letters are significantly1016different (P < 0.05).

