



- 1 Elemental composition of invertebrates shells composed of different CaCO₃ polymorphs at
- 2 different ontogenetic stages: a case study from the brackish Gulf of Gdansk (the Baltic Sea)
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- 19 Abstract
- 20 In this study, the concentrations of 12 metals: Ca, Na, Sr, Mg, Ba, Mn, Cu, Pb, V, Y, U and
- 21 Cd in shells of bivalve molluscs (aragonitic: Cerastoderma glaucum, Mya arenaria and
- 22 Limecola balthica and bimineralic: Mytilus trossulus) and arthropods (calcitic: Amphibalanus
- *improvisus*) were obtained. The main goal was to determine the incorporation patterns of
- shells built with different calcium carbonate polymorphs. The role of potential biological
- 25 control on the shell chemistry was assessed by comparing the concentrations of trace elements
- 26 between younger and older individuals (different size classes). The potential impact of
- 27 environmental factors on the observed elemental concentrations in the studied shells is
- 28 discussed. Specimens were collected from brackish waters of the Baltic Sea (the Gulf of
- 29 Gdansk). For every species, 40 individuals (ten in each size class) were selected. Pre-cleaned
- 30 shells were analysed by ICP-OES and ICP-MS to determine the concentrations of metals.
- 31 The distributions of elements both differ between species and exhibit high
- 32 intraspecific variability. Calcitic shells preferentially incorporated Mg > Sr > Na, aragonitic
- shells incorporated Na > Sr > Mg, and bimineralic shells accumulated Na approximately two





34	times more intensively, than Mg and Sr which remained at similar levels. Among all species,
35	the calcitic shells of A. improvisus most effectively concentrated the majority of the studied
36	elements, especially Mg > Mn > Ba, which was contrary to the shells of aragonitic molluscs
37	that contained the lowest levels of trace elements. The size-dependent distributions of
38	elements in shells did not exhibit a consistent pattern. The highest significant differences were
39	found for the bimineralic shells of <i>M. trossulus</i> , while the smallest were found for aragonitic
40	shells; if any variability occurred, it was observed in heavy metals (Pb, Cd).
41	Our results indicate that elemental variability, especially that of Mg and Sr, is
42	dominated by the properties of the crystal lattice. The inconsistent variability of trace element
43	concentrations between species and within single populations supports the important role of
44	species-specific biological control of the biomineralization process and indicates that
45	environmental factors have a significant influence on the incorporation of trace elements into
46	the shells.
47	
48	Key words: calcite, aragonite, trace metals, ontogeny, shell, ICP-MS, ICP-OES
49	
50	1. Introduction
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68 of ocean history (Freitas et al., 2006; Gillikin et al., 2006; Khim et al., 2003; Ponnurangam et al., 2016; Vander Putten et al., 2000). However, as recent studies have indicated (Dove, 69 70 2010), the influence of environmental parameters on shell precipitation could be very complex. The mineralogy and chemistry of shells are likely to be not only linked to 71 environmental conditions but are also simultaneously controlled by the organism itself. Even 72 a single population or closely related species within the same habitat may exhibit different 73 74 accumulation strategies (Rainbow et al. 2000). Calcite and aragonite differ in structure, symmetry and properties. The calcium ion has 75 9-fold coordination in aragonite and 6-fold coordination in calcite, and its coordination creates 76 77 hexagonal and cubic packing (Putnis, 1992). Due to the spatial structure, larger cations, such as Sr, fit better in aragonite, while smaller cations, such as Mg, are energetically favoured in 78 calcite. In natural systems, calcite commonly incorporates Mg (Morse et al., 2007; Reeder, 79 80 1983; Wang and Xu, 2001), and increasing Mg concentrations are known to increase the 81 calcite solubility (Kuklinski and Taylor, 2009; Smith et al., 2006). At the temperatures and pressures of the Earth's surface, low-Mg calcite is the most stable form of CaCO₃ (de Boer, 82 83 1977). Nevertheless, in many marine organisms, aragonite and high-Mg calcite are the dominant phases precipitated from seawater (Dickson, 2004). The solution chemistry (Cusack 84 85 and Freer, 2008), temperature (Balthasar and Cusack, 2015), pressure (Allison et al., 2001), CaCO₃ saturation state (Watson et al., 2012), pCO₂ (Lee and Morse, 2010) and phylogenesis 86 (Kuklinski and Taylor, 2009; Smith et al., 1998; Smith and Girvan, 2010) are known to 87 influence shell mineralogy. The main driving force controlling the mineralogy of precipitated 88 CaCO₃ is the ratio of Mg to Ca ions in seawater (Cusack and Freer, 2008; Morse et al., 2007)... 89 A Mg/Ca > 2 favours the precipitation of aragonite and high-Mg calcite. At high Mg/Ca 90 91 ratios, such as those in modern seawater (Mg/Ca = 5.2), calcitic structures incorporate Mg, 92 which is observed to inhibit calcite nucleation and growth, whereas aragonite nucleation is not affected by magnesium in solution (De ChoudensSanchez and Gonzalez, 2009; Morse et al., 93 2007). Most low-Mg calcite-producing taxa are known to actively control the amount of 94 95 magnesium incorporated. Since the Mg/Ca ratio of the calcifying fluid is dissimilar to that of ambient seawater, organisms might be unaffected by changes in the environmental Mg/Ca 96 ratio (Bentov and Erez, 2005; De Nooijer et al., 2014). Several studies of organisms that 97 98 secrete calcareous skeletons have shown that lower seawater temperatures are correlated with 99 the secretion of calcite skeletons with low Mg contents, rather than more soluble high-Mg calcite or aragonite skeletons (Taylor and Reid, 1990). The mineralogy of many calcareous 100 structures changes with latitude, likely as a result of the temperature gradient from the poles 101





102	to the equator (Kuklinski and Taylor, 2009; Loxton et al., 2014; Taylor et al., 2014).
103	However, the secretion of calcite by marine organisms inhabiting low-temperature areas and
104	aragonite in warmer waters is not an absolute rule, as many polar molluses form fully
105	aragonitic shells (Cairns and Macintyre, 1992). Thermodynamics predicts that aragonite is the
106	stable phase at pressures higher than 5000 hPa (roughly 40 m depth), and calcite is the stable
107	phase at lower pressures. However, aragonite is the major constituent of shells or pearls,
108	indicating its metastable formation in shallow waters (Sunagawa et al., 2007). The
109	incorporation of Sr was suggested to play a significant role in the biomineralogical
110	precipitation of aragonite (Allison et al., 2001). Many studies have demonstrated a clear
111	correlation between the concentration of Sr in the hard parts and precipitation of the aragonite
112	layer (Iglikowska et al., 2016; Reeder, 1983). The ionic radius of Sr is larger than that of Ca;
113	thus, Sr is more likely to form 9-fold coordination, which triggers metastable aragonite
114	nucleation (Sunagawa et al., 2007).
115	Although many studies have demonstrated that the biological control of shell
116	composition is often more important than the environmental control (Carré et al., 2006;
117	Freitas et al., 2005, 2006; Gillikin et al., 2005), there is evidence to the opposite trend (Klein
118	et al., 1996; Pearce and Mann, 2006). In biologically controlled mineralization, the organism
119	drives the process of nucleation and growth of the minerals independently from the
120	environmental conditions. Basically, endogenous factors manifest through co-regulation of all
121	the structures and functions of the organism, including its sex, growth rate, metabolism, and
122	feeding strategy (Lowenstam and Weiner, 1989). Throughout the lifespan, the biogenic
123	system experiences ontogenetic trends and seasonal variations in physiology, determining
124	metabolic expenses on the basis of life's needs. Biological effects have been repeatedly used
125	to explain any shifts of elemental concentrations in shells from a theoretical equilibrium
126	(Davis et al., 2000; Watson et al., 1995). Ontogenetic fluctuations of the growth rate and
127	metabolic activity affect the intensity of the metal uptake (Lee et al., 1998). Vander Putten et
128	al. (2000) concluded that the seasonality of the accumulation of Mg, Sr and Pb in Mytilus
129	edulis shells shows significant similarity across individuals, with a maximum during spring
130	and early summer, and profiles cannot be explained by seasonal variations in the seawater
131	composition. (Carré et al., 2006) developed a model of ion transport in bivalve shells that
132	shows that Ca channels are less ion-selective when Ca ²⁺ fluxes are higher. Other studies have
133	found that trace metal uptake rate increases as mussel filtration rate increases (Janssen and
134	Scholtz, 1979).
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136 composition of the carbonate shell and how much this process is affected by the environment (Casella et al., 2017; Schöne and Krause, 2016; Weiner et al., 2001). Although calcareous 137 138 parts serve as a powerful tool for the interpretation of the environmental conditions, detailed insight into the different factors controlling the composition of biological CaCO₃ is 139 fundamental for understanding the chemical profiles. Tracking potential sources of variation 140 within calcifying organisms of many species with diverse mineralogy should provide valuable 141 insight into the patterns driving the biomineralization process (Smith et al., 1998). 142 The aim of this study is, by examining the trace element concentrations in calcitic, 143 aragonitic and bimineralic shells. to assess the potential differences in free ion binding in two 144 crystalline polymorphs in the low-salinity environmental system using the example of the 145 southern Baltic Sea. Brackish waters influence the activities of metal ions, enhancing their 146 bioavailability (Fritioff et al., 2005) and contributing to lower Ca²⁺ concentrations, CaCO₃ 147 saturation states, alkalinity and buffering capacity compared to those of more saline waters 148 149 (Beldowski et al., 2010; Cai et al., 2010; Findlay et al., 2007). The seasonal changes (mainly the surface temperature variability, periodic lowering of salinity, irregular inflow of pollutants 150 151 and intensity of biological processes) were found to affect both the metal concentration and the physiological processes of living organisms (Urey et al., 1951). For this study we selected 152 153 dominant species of the calcifying benthic fauna in this region. This includes the bivalve molluscs: aragonitic Cerastoderma glaucum, Mya arenaria, and Limecola balthica and 154 bimineralic Mytilus trossulus, as well as an arthopod, Amphibalanus improvisus, with a fully 155 calcitic shell. The investigated species are deposit and filter feeders and have sessile or semi-156 sessile lifestyles. Their lifespan is long enough to include various seasons and years covering 157 158 therefore a range of environmental conditions during their existence. The potential influence 159 of local environmental conditions on the observed trace element concentrations in the shells of the studied organisms is discussed. A comparison of concentration levels between size 160 classes is also performed, to demonstrate any potential biological control of the shell 161 chemistry (assuming that the larger specimens were older than the smaller specimens of the 162 163 same species).

164 2. Material and methods

165 2.1. Study area

166 The study area is located in the Gulf of Gdansk in the southern region of the Baltic Sea (Fig.

167 1). The Gulf is partially sheltered from the northwest by the Hel Peninsula and from the west

and south by the Polish coastline (Kruk-Dowgiałło and Szaniawska, 2008; Rainbow et al.,





169	2004). This location makes the seawater the most turbulent in January and the calmest in
170	June, with weak bottom currents and minimal tidal amplitudes. The hydrophysical parameters
171	of the Gulf are mostly driven by the temperate climate and the following seasonal changes.
172	Differences in air temperature and water mixing cause seasonal fluctuations of the surface
173	water temperature, ranging from approximately 4 to 22°C (Uścinowicz, 2011). The Gulf of
174	Gdansk is a low-salinity system under the influence of brackish water from the open southern
175	Baltic Sea and fresh waters from rivers, mainly the Vistula River; the Vistula is the largest
176	river in Poland and has an average annual inflow in the estuary of 1080 m ³ s ⁻¹ , which varies
177	seasonally from 250 to 8000 m ^{3} s ⁻¹ and has a maximum in spring (Cyberski et al., 2006).
178	Thus, the average water salinity in the gulf is 7, varying from approximately 5.5 in summer to
179	8.4 in winter (Bulnheim and Gosling, 1988; Szefer et al., 2002).
180	The Gulf is known as an area highly exposed to human impact. This is due to intensive
181	usage of its resources and to anthropogenic emissions originating from various coastal
182	sources, river inflows and atmospheric deposition. The most significant input of industrial and
183	municipal pollution into the Gulf of Gdansk is derived from the Vistula River, which
184	transports pollutants from a catchment area of 194,000 km ² (Pruszak et al., 2005). Both the
185	water discharge and sediment load into the Gulf are strongly seasonally dependent. The
186	maximum river flow typically occurs in spring, with the minimum in autumn and winter,
187	mainly due to periods of snow and ice melting (Cyberski et al., 2006; Pruszak et al., 2005;
188	Szefer et al., 1996). Winter storms cause the re-deposition of sediments further into the Gulf
189	of Gdansk (Damrat et al., 2013). Because of the natural conditions, mainly, the limited water
190	exchange, contaminants brought by the river remain in the ecosystem for decades,
191	accumulating in the sediments and in living organisms (Glasby et al. 2004; Szumiło-Pilarska
192	et al. 2016).
193	The benthic community of the Gulf of Gdansk is characterized by the dominance of
194	calcifying filter feeders and deposit feeders (Kruk-Dowgiałło and Dubrawski, 1998). The low
195	salinity (~7) and alkalinity, which, in most regions of the Baltic Sea, are lower than those in
196	oceanic surface waters, cause reduced CO_3^{2-} and Ca^{2+} concentrations in seawater, thus
197	lowering the CaCO ₃ saturation state (Ω). Furthermore, due to the seasonality of the
198	temperature and water chemistry, the amplitude of Ω in the Baltic Sea is high in comparison
199	to that in the open ocean. It alternates between approximately 1 to 5 for calcite and 0.5 to 2.5
200	for aragonite (Findlay et al., 2007). Seasonal changes also affect the feeding behaviour. The
201	capacity for food processing is evolutionarily adapted to the concentrations of suspended
202	matter, primarily phytoplankton, that prevail in the biotope during the productive seasons of





- 203 the year. During these periods, organisms exploit their growth potential (Pierscieniak et al.,
- 204 2010; Staniszewska et al., 2016).
- 205

206 2.2. Species

The barnacle Amphibalanus improvisus (Darwin, 1854) (Arthopoda, Maxillopoda), 207 commonly named the bay barnacle, is a small sessile crustacean that typically exists in 208 shallow coastal zones that are less than 10 m deep. It is widespread around the Atlantic and 209 has been dispersed by shipping to many parts of the world, now having a worldwide 210 distribution. It is a euryhaline and eurythermal species that is absent only from the Arctic and 211 Antarctic seas (Kerckhof, 2002). A. improvisus is a filter feeder that inhabits hard substrates. 212 In the Baltic Sea, the reproduction of barnacles starts in spring with temperatures over 10°C 213 and ends in autumn. The species grows to approximately 10 mm in diameter with a maximum 214 215 height of approximately 6 mm, and generally, it has longevity of one year (Żmudziński, 1990). It has a conical shell composed of six fused calcite plates (Weidema, 2000). 216 The clam Cerastoderma glaucum (Bruguière, 1789) (Mollusca, Bivalvia), commonly 217 218 known as the lagoon cockle, is a saltwater clam found along the coasts of Europe and North Africa, including in the Mediterranean and Black Seas, the Caspian Lake, and the low-salinity 219 Baltic Sea. It is a euryhaline species living in salinities of 4-84. C. glaucum can tolerate 220 habitats with wide range of temperatures, from periodically freezing to above 30°C. It is a 221 filter feeder that burrows 2-4 cm below the surface in soft sediments at shallow depths of 222 secluded bays. It actively lives near the sediment surface, acting as a biodiffuser (Urban-223 Malinga et al., 2013). The clam is surrounded by a ribbed aragonite shell, which is externally 224 vellowish to greenish brown (Jelnes et al., 1971). In the brackish environment of the Gulf of 225 Gdansk, C. glaucum spawns in May-July and typically lives up to 4 years, achieving a height 226 227 of 27 mm (Żmudziński, 1990). The soft-shell clam Mya arenaria (Linnaeus, 1758) (Mollusca, Bivalvia) is a marine 228 229 invasive species introduced into European waters from the Atlantic coasts of North America 230 (Behrends et al., 2005). It has a wide global distribution, mainly due to its adaptability to varying environments with salinities between 4 and 35 and temperatures between -2 and 231 28°C (Gofas, 2004; Strasser et al., 1999). M. arenaria burrows into the sediment up to 20-30 232 cm below the surface of the sea floor. *M. arenaria* is a filter feeder, filtering organic particles 233 234 and microinvertabrates using long fused siphons, and a deposit feeder. In the Gulf of Gdansk,

235 *M. arenaria* is a common inhabitant of shallow waters up to a depth of 30 m. It spawns once





236 or twice a year in spring or summer, at temperatures of $10 - 15^{\circ}$ C. Individuals live 10 - 12years. They have aragonitic shells and grow up to 70 mm (Żmudziński, 1990). 237 238 The clam Limecola balthica (Linnaeus, 1758) (Mollusca, Bivalvia) lives in the northern parts of the Atlantic and Pacific oceans, in Subarctic and European waters from 239 southern France to the White Sea and Pechora Sea, including the Baltic Sea (Strelkov et al., 240 2007). It is a euryhaline clam capable of living in a wide range of water salinities from 3 to 40 241 and at temperatures from -2 to more than 30°C (Sartori and Gofas, 2016). L. balthica is a 242 filter feeder and deposit feeder that lives buried a few centimetres below the surface. It has a 243 semi-sessile lifestyle, with the ability to undertake periodic migrations (Hiddink et al., 2002). 244 In the Baltic Sea, L. balthica lives at depths of up to 40 m and grows to 24 mm. Adults 245 reproduce in spring when the water temperature reaches 10°C and live 12 years (Żmudziński, 246 1990). They have aragonitic shells varying in colour between individuals and locations, 247 mainly exhibiting white, pink, yellow and orange (Sartori and Gofas, 2016). 248 249 The mussel Mytilus trossulus (Gould, 1850) (Mollusca, Bivalvia) is one of three closely related taxa in the *Mytilus edulis* complex of blue mussels, which, collectively, are 250 251 widely distributed in the temperate and cold-water coasts of the Northern Hemisphere and are often dominant organisms on hard substrates of shallow nearshore habitats (Rainbow et al., 252 253 1999; Wenne et al., 2016). Generally, *M. trossulus* has a life span of approximately 12 years and grows to 100 mm, yet in the estuarine environment of the Gulf of Gdansk, it reaches a 254 maximum length of approximately 50 mm (Gofas, 2004; Żmudziński, 1990). Mussels are 255 sessile filter feeders, mainly depending on phytoplankton. They reproduce from late spring to 256 early autumn, depending on the temperature and food abundance (Larsson et al., 2017; 257 Lauringson et al., 2014; Rainbow et al., 2004). The shell of *M. trossulus* is bimineralic and 258 259 consists of two calcium carbonate layers: an outer calcite and inner aragonite layer in variable proportions between individuals (Dalbeck, 2008; Piwoni-Piórewicz et al., 2017). 260 261 2.3. Sample collection and preparation 262 263 Samples of shells were gathered by a Van Veen grab sampler from four stations, GN, MA, M2 and MW, located in the Gulf of Gdansk (Table 1, Fig. 1). No single location that would 264

be a source of several species presenting different forms of $CaCO_3$ was found. Three species

- were found at MA, M2 and MW (one at each station), while two species were collected at
- station GN (see Table 1). To ensure that the samples were not contaminated or modified by
- solutions of preservatives, the collected material was transported alive in tanks filled with

seawater to the laboratory, where sample preparation took place.





270	By measuring the shell heights of clams and shell diameters of barnacles using a
271	calliper with an accuracy of ± 1 mm, the shells were classified into four size classes. The
272	division into size classes was performed on the basis of the size reached by each species in the
273	southern Baltic Sea environment (Table 2). Forty shells (ten in each class) were selected for
274	A. improvisus, C. glaucum, M. arenaria and L. balthica. For M. trossulus, 20 shells were
275	selected, with five in each class, while the results for the rest were obtained from Piwoni-
276	Piorewicz et al. (2017). After the removal of soft tissues, each shell was viewed under a
277	stereoscopic microscope to examine the epibiotic flora and fauna, which could contaminate
278	the sample and bias the chemical analysis. If any organisms were present on the shell, they
279	were carefully removed. To remove the biofilm, the periostracum was scraped with a scalpel,
280	and pre-cleaned shells were placed in an ultrasonic bath (InterSonic IS-7S) in ultra-pure water
281	for 30 minutes and then dried at 70°C for 24 hours. The shells were then crushed and ground
282	into a fine powder with an agate mortar and pestle. Aliquots (2.8 – 849 mg, mean = 132 mg)
283	of the powdered samples were weighed using a 5-digit analytical balance, placed into a 15 ml
284	plastic tube (Sarstedt [®]) and dissolved in a mixture of 1.5 ml concentrated nitric acid (HNO ₃ ,
285	Sigma Aldrich [®] , Trace SELECT for trace analysis), 1.5 ml ultra-pure water and 0.3 ml 30%
286	hydrogen peroxide (H ₂ O ₂ , Merck [®] Suprapure grade). After 24 hours at 70°C, the liquid
287	samples were diluted to c. 15 ml by weight with ultra-pure water.

288

2.4. Elemental analysis

289 Concentrations of chemical elements in the digested samples were determined at the Natural 290 History Museum, London, using a Thermo iCap 6500 Duo inductively coupled plasma optical 291 emission spectrometer (ICP-OES) for Ca, Mg, Na and Sr and an Agilent 7700x inductively 292 coupled plasma mass spectrometer (ICP-MS) for Ba, Cd, Cu, Pb, U, V, Mn and Y. 293 294 Calibration of the ICP-OES analysis was performed using solutions that were matrix-matched to the high calcium concentrations in the samples. Multiple wavelengths for each element 295 were recorded, and line selection was performed by accounting for the suitability of the 296 297 wavelength to the concentrations in the samples. The accuracy and reproducibility of the analyses were checked using two calcium-rich certified reference materials (CRMs), JLs-1 298 Limestone and JDo-1 Dolomite (both from the Geological Survey of Japan), which were 299 dissolved by the total digestion method using hydrofluoric acid. The reference materials were 300 301 diluted to match the concentrations of Ca in the sample solutions (Imai et al., 1996). The limits of quantification (LOQ) of the ICP-MS analysis were generally determined 302

as the concentration corresponding to ten times the standard deviation of the signal obtained 303





- by analysing 5% HNO₃ solution (6–7 times) in each individual run. ICP-MS was run in
 helium (He) mode (5 ml min⁻¹ He, 99.9995% purity) for most trace elements to minimize the
 molecular interference from plasma and solution components and Ca from the samples. The
 accuracy and reproducibility was checked by periodic analyses of JLs-1 and JDo-1 (Imai et
- 308 al., 1996).
- 309
- 310 2.5. Statistical analyses
- 311 To evaluate the effect of the shell size (ontogenesis stage) on the concentrations of trace
- 312 elements in calcareous parts of A. improvisus, C. glaucum, M. arenaria L. balthica and M.
- 313 trossulus, the concentrations of Ca, Na, Sr, Mg, Ba, Mn, Cu, Pb, V, Y, U and Cd were
- examined in the four size classes separately for each species. The data were not normally
- 315 distributed (Shapiro-Wilk test); therefore, significant differences between the mean
- 316 concentrations of the selected elements in the size classes were identified by one-way
- 317 Kruskal-Wallis nonparametric ANOVA (p-value = 0.05) and post-hoc Dunn's tests for
- 318 multiple independent groups. Statistical computing and graphical visualizations were
- 319 performed in RStudio.
- 320
- 321 3. Results
- 322 The species exhibited similar accumulation orders since they had the highest concentration of Na, Sr and Mg and the lowest concentrations of U and Cd (Table 3, Fig. 2). However, the 323 concentrations of metals were different between the calcitic shells of A. improvisus, the 324 325 aragonitic shells of C. glaucum, M. arenaria, L. balthica and the bimineralic shells of M. trossulus, showing high variability (Fig. 3). Manganese was the most variable and increased 326 from $1.4 \pm 1.0 \text{ mg kg}^{-1}$ in aragonitic *M. arenaria* to $620 \pm 155 \text{ mg kg}^{-1}$ in calcitic *A*. 327 328 *improvisus*. Trace elements (V, Cu, Y, Cd, Ba, Pb, U) were generally present at higher concentrations in calcitic shells than in aragonitic ones. Shells of M. arenaria had the lowest 329 330 concentrations of impurities. The smallest variability between species was found for Sr, Na and Ca (Fig. 3). 331 The results of the Kruskal-Wallis nonparametric ANOVA test, which was used to 332 333 compare the shell chemistry between the four size classes in each population, revealed the
- lack of ontogenetic changes within aragonitic *M. arenaria*. The smallest variability was found
- in aragonitic *L. balthica*, for which only the concentration of Na (p = 0.014) decreased with
- shell growth. The third of the aragonite species, *C. glaucum*, showed high variability of the
- 337 four elements. Throughout the lifespan, shells were built with increasing concentrations of Sr





338	(p = 0.002), which was contrary to the decreasing concentration of Na $(p = 0.015)$, Mn $(p = 0.015)$
339	0.014) and Cd ($p = 0.009$). Similarly, the calcitic shells of <i>A. improvisus</i> also showed
340	variability of four elements between size classes, namely, Mg ($p = 0.007$), V ($p = 0.011$), Cu
341	(p = 0.027) and Pb $(p = 0.004)$. However, in this case, the sequences of the ontogenetic
342	changes were not straightforward but, rather, had a tendency to fluctuate between the
343	youngest and the oldest individuals. The highest variability was found within the bimineralic
344	shells of <i>M. trossulus</i> . The size classes differed in terms of five elements. The incorporation
345	of V ($p < 0.001$), Cu ($p < 0.001$), Y ($p = 0.013$), Cd ($p = 0.002$) and U ($p < 0.001$) into the
346	crystal lattice decreased in larger mussels (Fig. 3, Fig. 4).
347	Detailed analyses of the significant differences in the studied elements between the
348	size classes based on post-hoc Dunn's tests for multiple independent groups indicated that the
349	variations were not linear (Fig. 5). In L. balthica aragonitic shells, sodium concentration
350	decreased with ageing, showing differences between the size classes I and III and I and IV. In
351	the shells of C. glaucum, Sr concentration increased gradually, reaching a peak in the size
352	class IV, while statistically significant differences were observed between the size classes I
353	and IV and the III and IV. An inverse pattern was observed for sodium, yet its concentration
354	differed only between the youngest and oldest individuals. The aragonitic shells were
355	characterized by a common trend of the Mn and Cd concentrations, which decreased from I
356	size class and later reached a plateau. In the calcitic shells of A. improvisus, the concentrations
357	of Mg, V, Cu and Pb decreased in larger individuals. The levels of magnesium in shells
358	statically differed between the size classes III and IV; likewise, shells from the size class III
359	had the highest concentrations. Vanadium and lead occurred at the highest concentrations in
360	the shells of the youngest individuals, and later, they oscillated around a similar average. Co
361	decreased in growing shells of barnacles, reaching the minimum in the size class III. There
362	were no statistically significant differences between the classes III and IV, and the oldest half
363	of the group maintained the downward trend. However, it is worth noting that Cu
364	concentrations in many of the oldest shells are several times higher. The bimineralic shells of
365	M. trossulus were characterized by the highest variability of trace element concentrations
366	between size classes. The trend of decreasing concentrations was clearly marked for V, Co,
367	Cd and U. The concentration of Y decreased in larger shells, but the oldest shells showed an
368	upward trend; therefore, significant differences were found between the size classes I and III
369	(Fig. 5).
370	

371 4. Discussion





372 Magnesium and strontium are some of the most commonly studied elements in the skeletons of marine calcifiers due to their important role in the biological precipitation of CaCO3 373 374 (Allison et al., 2001; De Choudens Sanchez and Gonzalez, 2009; Morse et al., 2007). We found statistically significant differences in the concentrations of Mg and Sr between the 375 calcitic, aragonitic and bimineralic species. This confirms that the crystal structure of the shell 376 plays a key role in the incorporation of major elements. In the calcitic A. improvisus, Mg was 377 the dominant impurity, while aragonitic C. glaucum, M. arenaria and L. balthica had 378 concentrations of Sr over 15 times higher than those of Mg. The shell of *M. trossulus* contains 379 layers of calcite and aragonite, thus favouring both Mg and Sr. Not surprisingly, in this study, 380 mussels were characterized by similar concentrations of Mg and Sr, reaching 0.106 ± 0.017 381 mg kg⁻¹ and 0.117 ± 0.010 mg kg⁻¹, respectively (Table 3, Fig. 3). These observations are in 382 agreement with results obtained by other authors for different calcareous species (Iglikowska 383 et al., 2016; Morse et al., 1997; Reeder, 1983; Wang and Xu, 2001). Such results confirm the 384 385 importance of the Mg and Sr ionic radius in terms of the aragonite and calcite lattice properties (England, 2005). The studied calcitic A. improvisus, without a doubt, has the 386 highest Mg concentration in its shells $(0.394 \pm 0.032 \text{ mg kg}^{-1})$; however, the Sr concentration 387 was also higher in barnacles $(0.225 \pm 0.019 \text{ mg kg}^{-1})$ compared to that in aragonitic clams 388 $(0.170 - 0.216 \text{ mg kg}^{-1})$ (Table 3, Fig. 3). Kinetic and biological effects (Urey et al., 1951) 389 influence the partitioning of Sr between the shell and seawater, and Sr in shells is known to 390 391 significantly exceed its concentration expected at the thermodynamic equilibrium (Schöne et al., 2010). An important consideration for Sr concentrations in calcite is that Mg incorporation 392 into the shells may increase the Sr affinity through changes in the calcite crystal morphology 393 394 (Carpenter and Arbor, 1992). Greater inclusion of minor ions into the shell structure may 395 result in crystal defects, contributing to a higher incorporation rate of less compatible ions into the crystal lattice (Dalbeck, 2008; Davis et al., 2000). Furthermore, it was found that 396 barnacles differ from other calcitic species, having an uncommonly high Sr level (Carpenter 397 and Arbor, 1992; Ullmann et al., 2018), and the highest concentration in A. improvisus in the 398 399 Gulf of Gdansk could act as species-specific adaptation (Fig. 3). 400 Sodium belongs to the group of proxies used in the robust reconstruction of salinity, and this relationship seems to be related mainly to the increase of Na activity in seawater with 401 increasing salinity (Findlater et al., 2014; Wit et al., 2013). Although we observe that for 402 403 invertebrates collected from stations with similar salinities between 6.9 and 7.3, the aragonitic species incorporated more Na than did the calcitic and bimineralic species (Figs. 3 and 4). The 404

405 concentration of Na ranged, on average, from 0.2 wt% in *M. trossulus* to 0.32 wt% in *C.*





406 glaucum (Table 3). This finding could mean that the polymorph orientation affects Na concentration in shells. Following the orientation of the crystal lattice, Na, as a cation larger 407 408 than Ca, performs better in the aragonite structure (Davis et al., 2000). Dalbeck (2008) compared the minor element variations between the polymorphs of Mytilus edulis bimineralic 409 shells, and faound that Na occurred in higher concentrations in the aragonitic layer. Although 410 the Na/Ca ratio is frequently discussed as a climate proxy, the degree to which Na 411 concentration in biogenic carbonate represents the environmental salinity varies from one 412 species to another (Vander Putten et al., 2000; Schöne and Krause Jr, 2016). This study shows 413 lack of pattern in the distribution of Na between individuals from one population and between 414 species. 415

The concentrations of manganese and barium in the calcitic shells of A. improvisus 416 were several orders of magnitude larger than those in other species, reaching 625 ± 160 mg 417 kg^{-1} and 73.2 ± 19.3 mg kg⁻¹, respectively (Table 3). The bimineralic shells of the mussel M. 418 trossulus collected from the same location as A. improvisus (GN) contained 56.6 ± 20.0 mg 419 kg^{-1} of Mn and $17.0 \pm 6.7 mg kg^{-1}$ of Ba, while the aragonitic shells of clams incorporated 420 from $1.419 \pm 1.01 \text{ mg kg}^{-1}$ to $11.933 \pm 10.47 \text{ mg kg}^{-1}$ of Mn and from $2.63 \pm 1.58 \text{ mg kg}^{-1}$ to 421 15.12 ± 10.15 mg kg⁻¹ of Ba (Table 4, Fig. 3). This trend does not seem to be strongly linked 422 to the crystal lattice orientation because smaller Mn cations are energetically favoured in 423 calcitic species, while larger cations, such as Ba, better fit into aragonite (Dalbeck, 2008; 424 Davis et al., 2000; Wang and Xu, 2001). It was expected that aragonite shells would 425 incorporate Ba more intensively than would calcitic shells (Findlater et al., 2014; Gillikin et 426 al., 2006), yet this trend was not observed (Figs. 3 and 4). The organisms derive metal ions 427 from both the surrounding water and food (Freitas et al. 2006; Gillikin et al. 2005; Poulain et 428 429 al. 2015). It is highly probable that intra-species variabilities reflect the seasonal variability of Mn and Ba concentrations in seawater, while differences between species partly depend on 430 feeding strategies. Findlater et al. (2014), based on shells of fossil freshwater bivalves and 431 gastropods from the same location, concluded that a clear biochemical vital effect was evident 432 433 for Mn and Ba. Dissolved Mn is intensively delivered to the oceans by riverine input, and in 434 the Gulf of Gdansk, the geographical range of the Vistula River inflow plays a very important role in the Mn distribution (Uścinowicz, 2011). However, much lower concentrations of Mn 435 436 were found in the shells of molluses than in barnacles from the same location (Figs. 3 and 4). 437 A similar relationship was found in the soft tissues of M. trossulus and A. improvisus collected from different locations in the Gulf of Gdansk in May 1998 (Rainbow et al., 2000). 438 The range of Mn in the soft tissues of *M. trossulus* varied from 19.0 to 41.0, while that in *A.* 439





improvisus ranged from 187 to 307 mg kg⁻¹. Therefore, A. *improvisus* is a very effective 440 accumulator of Mn. The disproportionality of Mn and Ba concentrations between molluscs 441 442 and barnacles could reflect the ability of molluscs to limit vital processes and to reduce the calcification rate during stressful conditions (Berge et al., 2006; Hiebenthal et al., 2012). A 443 limited food base, low temperatures, saturation states during colder months and even the high 444 river inflow in early spring (Cyberski et al., 2006; Pruszak et al., 2005; Szefer et al., 1996) can 445 reduce the precipitation rates of C. glaucum, M. arenaria, L. balthica and M. trossulus. In 446 contrast, the calcified shells of barnacles grow more or less continuously (Bourget and Crisp, 447 1975) and represent yearly calcification. The Mn and Ba incorporation from seawater into 448 calcite and aragonite has been shown to correlate with changes in primary production and 449 phytoplankton blooms (Freitas et al., 2006; Lazareth et al., 2003; Vander Putten et al., 2000). 450 As many clams and barnacles are filter-feeding organisms, Mn- or Ba-rich particles are 451 ingested with food and from the water column. Schoemann et al. (1998) recorded a seasonal 452 453 increase of dissolved Mn in the water column after peaks in phytoplankton blooms. This partly explains why *M. arenaria* had the smallest concentrations of Mn and Ba. This clam 454 455 represents a deposit-feeding mode and thus may have limited access to suspended particles. Uptake and accumulation in filter-feeders correlates with metal concentrations in seawater, 456 457 whereas accumulation in deposit-feeders most likely reflects the metal concentrations in the sediments (Newman and McIntosh, 1982). Luoma and Jenne (1976) concluded that the degree 458 of metal accumulation by deposit-feeders was directly related to the characteristic rate of 459 metal desorption from a particular type of sediment. 460

Shells of A. improvisus, characterized by the highest concentrations of Mg, Sr, Mn and 461 Ba, were simultaneously the richest in most other trace metals, which were likewise more 462 463 concentrated in bimineralic *M. trossulus* than in aragonitic clams (Table 3, Fig. 2). It was mentioned previously that impurities arise in deformed crystal lattices (Dalbeck, 2008; Davis 464 et al., 2000), and these lattices could be changed through the intensive incorporation of Ca 465 substitutions. However, it is also important to mention that the ionic radii of smaller cations, 466 467 such as V, Cd, Y, Pb, U and Cu, are more energetically favoured in calcite than in aragonite (Morse et al., 1997; Reeder, 1983; Wang and Xu, 2001). 468 Trace metal concentrations in shells are usually linked to local environmental 469

470 conditions. If heavy metals in the shells of calcifying invertebrates reflect the environmental

471 levels of bioavailable ions, the highest concentrations of most trace metals in the shells of *A*.

472 *improvisus* and *M. trossulus* from GN station (Table 4, Fig. 3) are probably related to local

473 contamination. Rainbow et al. (2004) compared the differences between the bioavailabilities





of seven metals (Cu, Zn, Fe, Pb, Ni, Cd and Mn) to *A. improvisus* and *M. trossulus* across five
localities in the Gulf of Gdansk, including GN station, which was often the richest metal
source for marine organisms. Among the stations selected in this study, GN is closer than
MA, MW and M2 to the mouth of the Vistula River, which is a major source of metal
contaminants in the marine system (Fig. 1).

In this study, individuals were collected over a wide range of sizes from each station 479 480 (Table 2), representing different ages and various periods of time, living under the influence of seasonal changes. The southern Baltic Sea is driven by cyclical environmental dynamics, 481 which evoke physiological stress, reduce the food base and determine the biogeochemical 482 cycle (Elder and Collins, 1991). It is certain that the biological processes of an animal 483 precipitating carbonate complicate the calcification course (Vander Putten et al., 2000; 484 Zacherl et al., 2003). However, the physiology of an organism is likely dependent on the 485 environmental conditions it experiences (Gillikin et al., 2005). The recorded concentrations of 486 487 metals in all populations exhibited marked variability among individuals within a single location (Table 3, Fig. 3), an attribute which was previously noticed by several researchers 488 489 (Gillikin et al., 2005; Vander Putten et al., 2000). The wide range of inter-species variability suggests that the uptake of metals, and the subsequent bioavailability, are highly dependent on 490 491 biological factors such as the growth rate, feeding strategy, genotype and phylogeny. Within species, organisms from juveniles to adults experience morphological and functional changes 492 related to sex, metabolic rate or reproductive stage, which affect the biomineralization process 493 (Carré et al., 2006; Freitas et al., 2006; Gillikin et al., 2005; Schöne et al., 2010, 2011; Warter 494 et al., 2018). In shells of the studied A. improvisus, C. glaucum, M. arenaria, L. balthica and 495 496 *M. trossulus*, differences in elemental concentrations over the lifespan did not produce a 497 consistent pattern. From the chemical profiles, it is evident that the size-dependent distributions of elements in the shells were only marked in a few cases (Figs. 5 and 6). The 498 studied A. improvisus and M. trossulus showed the greatest chemical variability between size 499 classes (Fig. 4). However, in this case, the varied elements were not uniform for A. improvisus 500 501 and *M. trossulus*, even though the organisms came from one location (Table 1). No size-502 related trend was observed for Y, U, and Cd in barnacles or for Pb in molluscs (Fig. 3). It was previously found that different species in the same habitat might show different patterns of 503 504 metal accumulation, which are caused by specific routes of uptake or differences in the 505 physiological pathway of metals (Rainbow, 2002, 1995). Rainbow et al. (1999) tested A. improvisus and M. trossulus from the Gulf of Gdansk as environmental biomonitors by 506 measuring the concentrations of Co, Zn, Fe, Cd, Pb, Mn and Ni in soft tissues. They revealed 507





that the mussels and barnacles gathered at one location did not show the same variation inmetal bioavailabilities.

510 In this study, generally, when statistical differences between size classes were recorded, the concentrations of metals decreased with the size of the shell. The reverse trend 511 was found only for Sr in the shells of C. glaucum, in which the concentration of Sr increased 512 with size (Fig. 4). It is possible that the younger specimens had a greater calcium carbonate 513 precipitation rate and uptake of trace elements due to a higher metabolism in the earlier stages 514 of the animal life (Dalbeck, 2008). Rosenberg and Hughes (1991) suggested that areas of 515 higher shell curvature, such as the umbo, require greater metabolic expenditure, resulting in 516 an increase of metals uptake. This supports the idea of a higher metabolic rate and metal 517 concentrations in younger specimens (Szefer et al., 2002). This study also indicates that the 518 smallest individuals seem to be the most efficient accumulators of trace elements. This is 519 520 mostly expressed in the cases of a few trace elements, in which the concentration decreased across the four size classes in the shells of A. improvisus, M. trossulus and C. glaucum (Fig. 521 4). There are several studies confirming a negative relationship between metal concentrations 522 523 and shell size (Martincic et al., 1992; Piwoni-Piórewicz et al., 2017; Ritz et al., 1982). Catsiki and Gialamas (1994) suggested that apart from metabolic processes, an active detoxification 524 mechanism is responsible for this trend, and its efficiency is higher in older and larger 525 individuals. 526 Nevertheless, many of the elements studied herein showed a lack of statistically 527 significant relationships between the shell sizes. Few trace metals in the shells of A. 528 improvisus (Y, U, Cd), C. glaucum (V, Cu, Y, Pb, U) and M. trussulus (Pb) and all trace 529 metals in the shells of L. balthica and M. arenaria (V, Cu, Y, Pb, U, Cd) showed no 530 531 significant variability related with the growth of organisms (Fig. 3). This is not unusual in 532 biomineralization studies. Saavedra et al. (2004) observed no differences between Cd, Pb, Cr, Ni, As, Cu and Zn concentrations for different shell lengths of the raft Mytilus 533 galloprovincialis separated into four size classes. Protasowicki, Dural, & Jaremek (2008) 534 535 similarly found that the concentrations of Hg, Pb, Cd, Cu, Zn, Cr, Ni, Fe, Mn, V, Li and Al in the shells of the mussel Mytilus edulis from the Polish coast of the Baltic Sea did not vary 536 between shell sizes. This inconsistent pattern proves that biological effects modify the way in 537 538 which environmental data are recorded in the biogenic hard parts. 539 540 5. Summary





541	The shells of the calcitic A. improvisus, the aragonitic C. glaucum, L. balthica, M. arenaria
542	and the bimineralic <i>M. trossulus</i> from the Gulf of Gdansk are accumulators of a wide
543	spectrum of metals from the surrounding environment. The selected taxa differed in terms of
544	the chemical compositions of shells, both between taxa and within a single population.
545	Elemental variability is strongly associated with the properties of the crystal lattice, which are
546	different between calcite and aragonite. Remarkably high variability of elemental
547	concentrations between species, especially for manganese, supports the importance of
548	species-specific biological control of the biomineralization process. The high variability of
549	elemental values within a single population indicates that both physiological and
550	environmental factors have an influence on the incorporation of these trace elements into the
551	shell. Given that the species were obtained from various sites, an impact from local
552	environmental factors cannot be excluded.
553	This study has shown that interpreting the chemical profiles of calcium carbonate
554	shells to determine the marine environment conditions is recommended to investigate how
555	metal concentrations vary among large numbers of species from contrasting habitats.
556	Nevertheless, we recommend using shells of similar sizes to reduce the ontogenetic
557	variability.
557 558	variability.
557 558 559	Author contribution
557 558 559 560	Author contribution Anna Piwoni-Piórewicz and Piotr Kukliński designed the study and carried them out.
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- 917
- 918 Tables and Figures
- 919 Table 1 Details of the research stations utilized for the collection of organisms. The
- 920 environmental parameters T (temperature) and salinity were measured near the bottom during
- 921 sample collection

Station	Latitude N	Longitude E	Collected species	Depth (m)	T (°C)	Salinity	Date
GN	54°32.0'	18°48.1'	Amphibalanus improvisus	36	3.1	7.3	May 2013
GN	54°32.0'	18°48.1'	Mytilus trossulus	36	3.1	7.3	May 2013
M2	54°38.9'	18°33.0'	Cerastoderma glaucum	10	16.9	6.9	June 2014
MA	54°37.2'	18°32.8'	Mya arenaria	10	19.8	6.9	June 2014
MW	54°37.4'	18°37.5'	Limecola balthica	31	4.7	7.0	June 2014

922

923 Table 2 The range of shell sizes for each studied species within a given size class

Success	Minoral trma	Size cla	asses (mi	m)	Maximum size in the	
species	wineral type	1	2	3	4	Gulf of Gdansk (mm)*
Amphibalanus improvisus	Calcite	3-4	5-6	7-8	9-10	10
Mytilus trossulus	Bimineralic	6-5	16-25	26-35	36–44	50
Cerastoderma glaucum	Aragonite	4-8	9-12	12-16	16-20	27
Mya arenaria	Aragonite	10-20	20-30	30-40	40–49	70
Limecola balhtica	Aragonite	4—7	8-11	12-15	15-18	24

*Based on the literature data (Zmudziński, 1990)

925

926 Table 3 The concentrations of the studied elements in shells of the collected organisms

Amphibala	nus improvisus						
Elements	Mineral type	Unit	Ν	Mean	±1 SD	Min	Max
Ca			37	31.6	2.60	27.6	39.0
Mg		%	37	0.394	0.032	0.337	0.472
Na		wt	37	0.285	0.047	0.222	0.508
Sr			37	0.225	0.019	0.198	0.283
V			36	0.466	0.156	0.155	0.78
Mn	cite		37	625	160	414	1161
Cu	Cal	_	37	3.90	2.20	1.20	8.30
Y		ι çû	37	0.21	0.12	0.070	0.78
Cd		ng l	37	0.062	0.020	0.029	0.125
Ва		T	37	73.0	20.1	59.0	145
Pb			37	1.34	0.79	0.456	3.86
U			37	0.108	0.043	0.035	0.220
Cerastode	rma glaucum						
Elements	Mineral type	Unit	Ν	Mean	±1 SD	Min	Max
Ca	t n og a	t	40	34.8	1.60	30.4	38.1





	l .						
Mg			40	0.008	0.003	0.005	0.023
Na			40	0.322	0.021	0.277	0.361
Sr			40	0.170	0.014	0.143	0.199
V			37	0.069	0.061	0.003	0.239
Mn			40	11.9	10.5	0.614	48.8
Cu		_	40	4.08	6.29	0.188	30.6
Y		[_] g [_]	40	0.070	0.074	0.006	0.401
Cd		ng	37	0.011	0.009	0.001	0.043
Ва		1	40	15.0	10.0	1.22	35.4
Pb			34	0.14	0.10	0.006	0.298
U			40	0.096	0.084	0.003	0.339
Limecola l	palthica						
Elements	Mineral type	Unit	Ν	Mean	±1 SD	Min	Max
Ca			40	34.5	1.70	29.0	37.5
Mg		t %	40	0.008	0.003	0.003	0.018
Na		×	40	0.31	0.03	0.25	0.367
Sr			40	0.177	0.021	0.135	0.244
V	te		35	0.035	0.031	0.002	0.102
Mn	goni		40	7.90	7.10	0.264	28.5
Cu	Arag	÷	40	3.00	3.40	0.047	11.5
Y	ł	kg'	40	0.08	0.11	0.002	0.421
Cd		mg	37	0.007	0.008	0.0004	0.031
Ba			40	13.0	10.0	0.638	27.6
Pb			39	0.14	0.16	0.007	0.585
U			40	0.034	0.024	0.004	0.088
				0.051	0.024	0.001	0.000
Mya arena	ria			0.051	0.024	0.001	0.000
Mya arena Elements	<i>ria</i> Mineral type	Unit	N	Mean	±1 SD	Min	Max
Mya arena Elements Ca	<i>ria</i> Mineral type	Unit	N 40	Mean 36.3	±1 SD 1.8	Min 30.5	Max 39.1
Mya arena Elements Ca Mg	<i>ria</i> Mineral type	Unit % t	N 40 40	Mean 36.3 0.015	±1 SD 1.8 0.002	Min 30.5 0.011	Max 39.1 0.021
Mya arena Elements Ca Mg Na	<i>ria</i> Mineral type	Unit % tx	N 40 40 40	Mean 36.3 0.015 0.308	±1 SD 1.8 0.002 0.021	Min 30.5 0.011 0.26	Max 39.1 0.021 0.345
Mya arena Elements Ca Mg Na Sr	ria Mineral type	Unit % ₩	N 40 40 40 40	Mean 36.3 0.015 0.308 0.216	±1 SD 1.8 0.002 0.021 0.028	Min 30.5 0.011 0.26 0.168	Max 39.1 0.021 0.345 0.295
Mya arena Elements Ca Mg Na Sr V	ria Mineral type	Unit % tw	N 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016	±1 SD 1.8 0.002 0.021 0.028 0.014	Min 30.5 0.011 0.26 0.168 0.001	Max 39.1 0.021 0.345 0.295 0.062
Mya arena Elements Ca Mg Na Sr V Wn	ria Mineral type	Unit %	N 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00	Min 30.5 0.011 0.26 0.168 0.001 0.276	Max 39.1 0.021 0.345 0.295 0.062 5.32
Mya arena Elements Ca Mg Na Sr V Mn Cu	ria Mineral type	Unit %t	N 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114
Mya arena Elements Ca Mg Na Sr V Mn Cu Y	ria Mineral type	kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd	ria Mineral type atiuogery	mg kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 37	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba	ria Mineral type Piuogeny	mg kg_1 wt %	N 40 40 40 40 40 40 40 40 40 37 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.116
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb	ria Mineral type atuo8 atuo8 M	mg kg ⁻¹ mt %	N 40 40 40 40 40 40 40 40 40 37 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U	ria Mineral type Pragounite	mg kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.003	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tree	ria Mineral type atuo Burger Mineral type	Unit % tw %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.001 0.574 0.002 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tra Elements	nia Mineral type atuo Besulus Mineral type	Unit % tw Sygmu Unit	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005 ±1 SD 2.8	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 20.7
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tro Elements Ca Mg	ria Mineral type otiuoBerry Mineral type	Unit % tw Unit Unit	N 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106	±1 SD ±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 2.8 0.017	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.001 Min 25.2 0.061	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tra Elements Ca Mg	ria Mineral type officiency Mineral type	wt % mg kg wt %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.021	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.201
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tree Elements Ca Mg Na Sr	ria Mineral type ajuogery ssulus Mineral type	mg kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.201	$\begin{array}{c} \pm 1 \text{ SD} \\ 1.8 \\ 0.002 \\ 0.021 \\ 0.028 \\ 0.014 \\ 1.00 \\ 0.024 \\ 0.018 \\ 0.003 \\ 1.60 \\ 0.014 \\ 0.005 \\ \hline \pm 1 \text{ SD} \\ 2.8 \\ 0.017 \\ 0.031 \\ 0.01 \\ \end{array}$	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.574 0.002 0.574 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.142
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tro Elements Ca Mg Na Sr	ria Mineral type augustus Mineral type	Unit % tw Unit Unit	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.117	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.031 0.01	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.001 25.2 0.061 0.147 0.101	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.143
Mya arena Elements Ca Mg Na Sr V Mn Cu Y Cd Ba Pb U U Mytilus tro Elements Ca Mg Na Sr V V	ria Mineral type augustus Mineral type	Mt % mt % mt %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.117 0.78 54.0	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.031 0.01 0.52	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.001	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.143 0.291 0.143
Mya arenaElementsCaMgNaSrVMnCuYCdBaPbUMytilus tracElementsCaMgNaSrVMnCr	ria Mineral type Puragouite Mineral type	tw wt % mg kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.002 2.60 0.015 0.0002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.117 0.78 54.0	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.031 0.01 0.021 0.024 0.017 0.031 0.01 0.025 15.0	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.001 25.2 0.061 0.147 0.101 0.206 31.8 2.26	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.143 2.30 92.6 5.9 2.6
Mya arenaElementsCaMgNaSrVMnCuYCdBaPbUMytilus tracElementsCaMgNaSrVMnCuVMnCuV	ria Mineral type Puragouite Mineral type	ng kg ⁻¹ wt % mg kg ⁻¹ wt %	N 40 40 40 40 40 40 40 40 40 40 40 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.117 0.78 54.0 14.0	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.031 0.01 0.52 15.0 10.0	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.574 0.002 0.574 0.002 0.574 0.001 Min 25.2 0.061 0.147 0.101 0.206 31.8 3.36 0.012	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.143 2.30 92.6 58.3 4.97
Mya arenaElementsCaMgNaSrVMnCuYCdBaPbUMytilus tracElementsCaMgNaSrVMnCuYCaMgNaSrVMnCuYCuYCuY	ria Mineral type Provide type Mineral type	mg kg ⁻¹ wt % mg kg ⁻¹ wt %	N 40 21 41 40 35 40 40 35 40 40	Mean 36.3 0.015 0.308 0.216 0.016 1.40 0.031 0.020 0.002 2.60 0.015 0.006 Mean 34.4 0.106 0.201 0.117 0.78 54.0 14.0 0.70	±1 SD 1.8 0.002 0.021 0.028 0.014 1.00 0.024 0.018 0.003 1.60 0.014 0.005 ±1 SD 2.8 0.017 0.031 0.01 0.52 15.0 10.0 1.20 2.1	Min 30.5 0.011 0.26 0.168 0.001 0.276 0.006 0.003 0.0002 0.574 0.002 0.574 0.002 0.574 0.001 Min 25.2 0.061 0.147 0.101 0.206 31.8 3.36 0.040	Max 39.1 0.021 0.345 0.295 0.062 5.32 0.114 0.089 0.016 6.26 0.077 0.022 Max 39.7 0.14 0.291 0.143 2.30 92.6 58.3 4.86 0.620 0.202







931 Fig. 1 The study area, with sampling stations indicated by black dots (stations details see

932 Table 1)







934

Fig. 2 The distribution of logarithmically transformed mean metal concentrations in shells of

936 Amphibalanus improvisus, Cerastoderma glaucum, Limecola balthica, Mya arenaria and

937 Mytilus trossulus







Fig. 3 Metal concentrations in analysed shells of *Amphibalanus improvisus, Cerastoderma*glaucum, Limecola balthica, Mya arenaria and Mytilus trossulus. The figure shows the raw
data as grey dots, with boxplots representing a ±1 standard deviation around the means and
whiskers indicating the minimum and maximum concentrations. The grey areas show
elements with statistically significant differences between size classes





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946

947 Fig. 4 Concentrations of elements with statistically significant differences between the four

size classes (for size class details see Table 2) in the shells of *Amphibalanus improvisus*,

949 Cerastoderma glaucum, Limecola balthica and Mytilus trossulus. The figure shows the raw

950 data as black dots, with the light grey areas representing the standard deviations around the

951 means (middle line)







953

954 Fig. 5 Pairwise comparisons of metal concentrations between size classes (for size class

955 details see Table 2) in the shells of Amphibalanus improvisus, Cerastoderma glaucum,

956 Limecola balthica and Mytilus trossulus. The broken line separates statistically significant

957 dependence (black points) from statistically non-significant dependence (empty points)