

SOIL GREENHOUSE GAS EMISSIONS UNDER DIFFERENT LAND-USE TYPES IN SAVANNA ECOSYSTEMS OF KENYA

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Abstract

~~For effective climate change mitigation strategies, adequate data~~Field measurement data on greenhouse gas (GHG) emissions ~~from a wide range of land use and land cover types area~~ prerequisite. However, GHG field measurement data are still scarce for many land-use types in Africa, causing a high uncertainty in GHG budgets. To address this knowledge gap, we present *in situ* measurements of carbon dioxide (CO₂), nitrous oxide (N₂O), and methane (CH₄) emissions in the ~~lowland part~~lowlands of southern Kenya. We conducted ~~eight chamber measurements~~ measurement campaigns on gas exchange from four dominant land-use types (LUTs) ~~and~~ ~~included~~including (1) cropland, (2) ~~grazed savanna~~, (3) bushland, (3) ~~grazing land~~, and (4) conservation land. ~~Between~~ ~~between~~ 29 November 2017 to 3 November 2018, ~~eight measurement campaigns were conducted~~ accounting for regional seasonality (including wet and dry seasons and transitions periods) ~~in each LUT~~. Mean CO₂ emissions for the whole observation period were significantly higher (p-value<0.05) in the conservation land (75±6 mg CO₂-C m⁻² h⁻¹) compared to the three other sites, which ranged from 45±4 mg CO₂-C m⁻² h⁻¹ (bushland) to 50±5 mg CO₂-C m⁻² h⁻¹ (grazing land). Furthermore, CO₂ emissions varied between seasons, with significantly higher emissions during the wet season than the dry season. ~~In contrast, mean~~Mean N₂O emissions were ~~not different between the four sites, ranging from highest in cropland~~ (2.7±0.6 μg N₂O-N m⁻² h⁻¹) and lowest in bushland (1.2±0.4 μg N₂O-N m⁻² h⁻¹ ~~(in bushland) to~~ 2.7±0.6 μg N₂O-N m⁻² h⁻¹ ~~(in cropland). However~~) but did not vary with season. ~~In fact~~, N₂O

Commented [WS1]: RC1: Ln 25 – the N2O flux was more than double the cropland than bushland, why do you say it was not different between the four sites?

Commented [WS2R1]: We have corrected

emissions were very low both in the wet and dry seasons, with slightly elevated values during the early days of the wet ~~season~~ seasons in all LUTs. On the other hand, CH₄ emissions did not show any significant differences between LUTs and seasons, ~~and most values.~~ Most CH₄ fluxes were below the limit of detection (LOD, $\pm 0.03 \text{ mg-CH}_4\text{-C-m}^{-2}\text{-h}^{-1}$). We attributed the difference in soil CO₂ emissions between the four sites to soil C content, which differed between the sites and was highest in the conservation land. In addition, CO₂ and N₂O emissions positively correlated ~~to~~ with soil moisture, thus an increase in soil moisture led to an increase in emissions. Furthermore, vegetation cover explained the seasonal variation of soil CO₂ emissions as depicted by a strong positive correlation between NDVI and CO₂ emissions, most likely because with more green (active) vegetation cover higher CO₂ emissions occur due to enhanced root respiration compared to drier periods in the year. Soil temperature did not show a clear correlation with either, ~~most CO₂ or N₂O emissions, which is~~ likely due to the low annual variation-variability in soil temperature. ~~We found a strong positive correlation between soil CO₂ and the normalized difference vegetation index (NDVI), but we observed no correlation with~~ between seasons and sites. Based on our results, soil N₂O emissions. We conclude that C, active vegetation cover and soil moisture is a key factor in drivers of soil GHG emissions in ~~these tropical savanna~~ all the tested LUTs. ~~In addition, including vegetation indices in the model greatly improved the results, thus showing the importance of vegetation cover in predicting soil emissions. South Kenya.~~ Our results are within the range of previous GHG flux measurements from soils from various ~~land use types~~ LUTs in other parts of Kenya and contribute to more accurate baseline GHG emission estimates from Africa, which are key to reduce uncertainties in global GHG budgets as well as for informing policymakers when discussing low-emission development strategies.

KEYWORDS: Carbon Dioxide, Nitrous Oxide, Methane, Bushland, Conservation, Grazing land, Cropland.

1. Introduction

Soil is a major source, and in many cases also a sink, of the atmospheric greenhouse gases (GHG) carbon dioxide (CO₂), nitrous oxide (N₂O), and methane (CH₄) (Oertel et al., 2016). The concentrations of these gases have increased since the onset of industrialization in 1970, leading to global warming (IPCC, 2013). GHGs trap the long-wave radiation emitted by the Earth's surface, thus increasing surface temperatures (Arrhenius, 1896). Soil CO₂ emissions originate from root respiration and heterotrophic decomposition of soil organic matter (Oertel et al. 2016). N₂O can be produced from many pathways in the soil nitrogen (N) cycle, but is considered to result primarily

Commented [WS3]: Rephrased based on reviewer one on What is the take home point?

from nitrification and denitrification (Butterbach-Bahl et al. 2013). CH₄ is produced by methanogenesis under anaerobic conditions and consumed by methanotrophic microorganisms under aerobic conditions, with the latter being more important in well-aerated upland soils, which consequently show net CH₄ uptake (i.e. negative flux) (Serrano-silva et al. 2014; Hanson and Hanson 1996). The production and consumption of soil GHGs largely depends on soil physical and chemical properties (Davidson et al., 2006) (e.g. texture, soil organic matter and pH) and are further driven by environmental factors such as soil moisture and soil temperature (Davidson et al., 2006). Thus, soil GHG emissions and uptake along with their controlling factors differ between biomes based on the land use and land-use management.

Land-use changes are reportedly the largest source of anthropogenic GHG emissions in Africa (Valentini et al., 2014). However, *in situ* studies on GHG emissions from various ecosystems in remain scarce, particularly from savanna ecosystems (Castaldi et al., 2006). Savanna is an important land cover type in Africa, covering more than 40 % of its total area (Scholes et al., 1997). In Kenya, savanna and grassland ecosystems cover about 80 % of the total area, comprising various land-use types (LUTs) (GoK, 2013). These ecosystems are subject to accelerating land-use change (Grace et al., 2006) due to population growth (Meyer and Turner, 1992) and land-use management activities (Valentini et al., 2014). Conversion of savanna for small- and large-scale livestock production, crop cultivation, and human settlement is common in Africa (Bombelli et al., 2009). As a consequence, vegetation cover, net primary productivity, allocation of carbon and nutrients in plants and soil (Burke et al., 1998) as well as soil GHG emissions are affected (Abdalla et al., 2018; Carbone et al., 2008).

Overgrazing due to overstocking is a major cause of soil and vegetation degradation in large parts of African savannas (Patton et al., 2007; Abdalla et al., 2018). Factors associated with grazing include animal feeding preferences of certain plant species, thus creating higher pressure for certain species, which decline in numbers over time, leading to species loss and lower pasture nutritive value (Patton et al. 2007). In addition, soil trampling increases soil bulk density and decreases soil water infiltration (Patton et al., 2007). Furthermore, high rates of dung and urine deposition, especially around homesteads and waterholes, create high N concentrations that are toxic for many savanna grass species, affecting vegetation cover and composition (e.g. increase of encroaching species such as *Solanum incanum* L., which is toxic for livestock (van Vegten 1984)). Given that all these factors affect soil properties, soil GHG emissions are most likely similarly affected (Wilsey et al., 2002).

In addition to overgrazing, rapid human population growth leads to more people migrating into savanna ecosystems, which has led to the expansion of cropland (Pellikka et al., 2018; Patton et al.,

Commented [WS4]: RC 1: The introductory paragraph never says what produces and consumes GHGs from the soil?

Commented [WS5R4]: We have included in the introductory

Commented [WS6]: Rephrased

Commented [WS7]: RC: Ln 65. Not all savannah belongs to the ASALs. The humid savannas are relatively wet, with green vegetation almost throughout the year. It is therefore not right to make such a sweeping statement.
AC: We revised to drylands instead of ASALs.

Commented [WS8]: RC 2: Ln 88. Revise the sentence. Overstocking leads to grazing pressure. The way the sentence is written is redundant.

Commented [WS9R8]: Done

Commented [WS10]: Rephrased based on RC2

95 2007). Brink and Eva (2009) found that the area under cropland increased by 57 % between 1975 and
2000 in Africa. In the Horn of Africa, cropland areas increased by 28 % between 1990 and 2010
(Brink et al., 2014), while wooded vegetation in East Africa decreased by 5 % in forests, 16 % in
woodlands, and 19 % in shrublands (Pfeifer et al., 2013). As an additional example, in our study area
Taita Taveta County in Southern Kenya, the area under cropland increased from 30 % in 1987 to 43
100 % in 2011 (Pellikka et al., 2018). However, in the Taita Hills, located in the County, this trend has
slowed down in recent years, while the savanna lowlands are still being cleared to make way for new
cropland (Pellikka et al., 2013).

Croplands in the Kenyan savannas are mostly managed by smallholder farmers with small land sizes
(Waswa and Mburu, 2006). Due to high poverty levels in this region, inputs to improve crop yields,
105 such as the use of fertilizer and herbicide, and mechanized farming are minor (Waswa and Mburu,
2006; CIDP, 2014). Consequently, increase in productivity are mostly generated via cropland
expansion. These smallholder farms are likely to have substantial effects on national GHG emission
budgets (Pelster et al., 2017). Until now, only a few studies have investigated soil GHG emissions
from such agricultural landscapes (Rosenstock et al., 2016), and these studies were mostly carried out
110 in high-potential farming areas such as the Kenyan highlands, which receive >1000 mm rainfall
(FAO, 1996). For example, Rosenstock et al. (2016) showed a large variation of CO₂ and N₂O
emissions both within and between four crop types as affected by environmental conditions and land
management. However, studies measuring GHG emissions from low-productivity croplands in
southern Kenya are to the best of our knowledge still missing. Thus, this study focused on soil GHG
115 emissions from different LUTs relevant for the semi-arid region of Southern Kenya.

Given the vast area covered by savanna, land use and land-cover changes are likely to affect global,
regional, and national C and N cycles, and hence the quantification of their role is vital (Lal, 2004;
Williams et al., 2007). Studies in Kenya have shown large variations of soil GHG emissions in various
savanna ecosystems (Otieno et al., 2010; Oduor et al., 2018), due to land-use (Ondier et al., 2019)
120 and management activities (K'Otuto et al. 2013). Due to the high diversity of these savanna
ecosystems, such studies may not be entirely representative for every region (Ardö et al., 2008).

The lack of reliable soil GHG flux data from natural savanna and cropland limits our understanding
of GHG emissions from African soils (Hickman et al., 2014; Valentini et al., 2014). At the same time,
accurate quantification of GHG emissions from multiple LUTs are essential to allow for reliable
125 estimation of Kenya's national GHG inventory (IPCC, 2019). This is particularly important as Kenya
currently relies on a Tier-1 approach by using default emission factors (EFs) provided in the

Commented [WS11]: RC2: Ln 96-7. ---Croplands are still being cleared from natural vegetation----re-write the sentence, it's not making the intended meaning.
AC: Done. Revised --Natural vegetation is being cleared to make way for the expansion of cropland

Guidelines for Greenhouse Gas Inventories of the UN Intergovernmental Panel on Climate Change (IPCC) to estimate national GHG emission budgets. Following the Paris Climate Agreement (<https://unfccc.int/process-and-meetings/the-paris-agreement/d2hhdC1pcy>), most countries across the globe, including Kenya, have not only agreed to accurately report their GHG emissions at national scales following a Tier-2 approach (i.e. using localized data) but also to mitigate anthropogenic GHG emissions in the upcoming decades, as is communicated via their Nationally Determined Contributions (NDCs). Both can only be achieved with locally derived data.

To address the lack of localized GHG emission data from different LUTs in Kenya, our study aims at: (1) providing crucial baseline data on soil GHG emissions from four dominant land uses, namely conservation land, grazing land, bushland, and cropland, and (2) investigating abiotic and biotic drivers of GHG emissions during different seasons. We hypothesized that GHG emissions in cropland ~~will~~would be higher compared to grazing land, bushland, and conservation land because of larger nutrient inputs (i.e. fertilization) in managed land. Further, we hypothesized that GHG emissions ~~will~~would differ between seasons; more precisely, we expected higher GHG emissions in the wet season than in the dry season ~~because of~~caused by higher soil moisture.

2. Materials and Methods

2.1. Study ~~area~~Area

This study was conducted in the lowlands (800–1000 m a.s.l.) of Taita Taveta County, ~~southern Kenya (Fig. 1), between~~ (latitude 3° 25' S and longitude 38° 20' E-) ~~located in southern Kenya (Fig. 1).~~ Taita Taveta County is one of Kenya's ~~ASAL regions~~dryland areas, with 89 % of the ~~county~~area characterized ~~by~~as arid and semi-arid ~~and arid conditions~~area. The county is divided into three major geographical regions, namely the mountainous zone of the Taita Hills (Dawida, Kasigau, Sagalla), Taita lowlands, and the foot slopes of Mt. Kilimanjaro around Taveta.

In the lowlands, ~~which is our study site,~~vegetation types include woodlands, ~~bushlands,~~ grasslands, and riverine forests/swamps. Tsavo East and Tsavo West National Parks ~~cover~~covers ca. 62 % of the county area ~~for wildlife conservation~~(CIDP, 2014). ~~The parks are open savanna and bush woodland~~bushland that support ~~elephants, buffaloes, lions, antelopes, gazelles, giraffes, zebras, rhinoceroses, large herbivores, predators~~ and a wealth of birdlife. There are 28 ranches designated for livestock production and two wildlife sanctuaries (Taita Hills Wildlife Sanctuary and LUMO Community Wildlife Sanctuary). ~~Livestock in the region is mainly managed in the form of nomadic~~

Commented [WS12]: We harmonize this in the revised manuscript to bushlands as found in Tsavo East and West national park

Commented [WS13]: Rephrased

pastoralism and only limited ranching occurs (CIDP, 2014). Other important land use includes croplands with dryland agriculture in small scale farming operations with low farm inputs (CIDP, 2014), dry thickets and shrublands, and sisal farming (Pellikka et al., 2018). The main soil ~~Other important land uses~~ include cropland under small-scale farming (CIDP, 2014), shrublands, and sisal farming (Pellikka et al., 2018). Soil type is characterized by dark red, very deep, acid sandy clay soil (Ferralsols). Our study sites were located in four of these key land uses in the region, ~~namely including~~ cropland, bushland, wildlife conservation land, and grazing land.

Commented [WS14]: Corrected

~~The lowland area has a bimodal rainfall pattern with two rainy seasons—a long rainy period between March and May and a short rainy period between October and December (CIDP, 2014). The hottest and driest months are January and February, while the dry season from June to October is cooler (Pellikka et al., 2018). Mean annual rainfall is 500 mm and the average annual air temperature is 23 °C, with an average daily minimum temperature of 16.7 °C and a maximum temperature of 28.8 °C (CIDP, 2014).~~

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~~The first site is~~ investigated is a cropland located in Maktau (1070 m a.s.l., Fig. 1). ~~The farm measured approximately 1, Fig. 2a) with a size of about one and a half hectares, cultivated with maize (Zea mays L.) intercropped with beans as the main crops. The farm can be considered as a typical rain-fed smallholder agriculture. Crop farm and crop growing closely follows the rainy seasons, with maize and beans sowed sowing in March, and bean and maize harvesting occurring in June for beans and August, respectively. Other crops on the farm included cowpeas, pigeon peas, cassava, and sweet potatoes. Land preparation was performed by for maize. Animal ploughing using animal traction is done to prepare land before seeding, while and weeding was performed is by hand hoe hoeing. Small quantities of fresh and dry manure (roughly 20 kg, accounting for less than 1 kg of N) were used every month to improve soil fertility by applying approximately 20 kg of mixed dry and fresh animal manure (less than 1 kg of N) every month during the campaign period (Fig. 2a).~~

Commented [WS15]: SC: l. 171: Height a.s.l. is only given for the cropland site and should be given as well for the other sites.
AC: This is done in the revised manuscript for all the sites

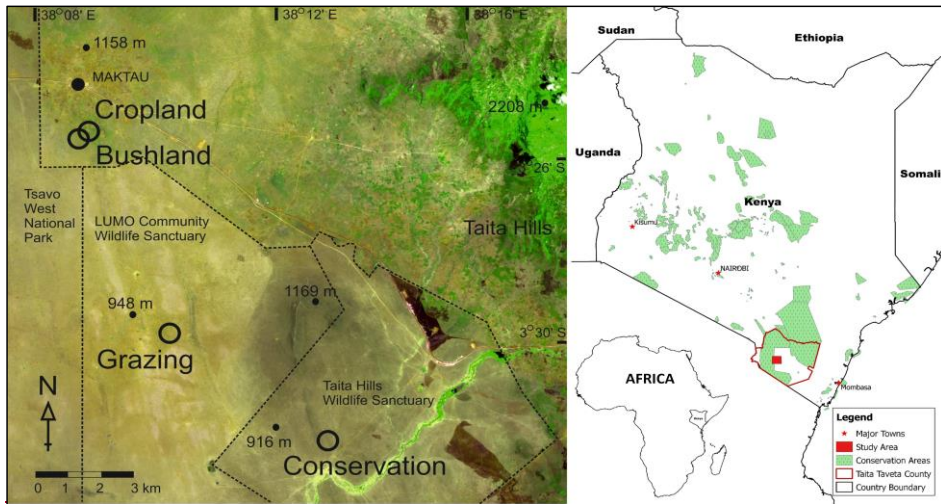


Figure 1. Location of the study sites cropland, bushland, grazing land, and conservation land in the savanna area in the lowlands of Taita Taveta County in southern Kenya.

The second site is located in a private bushland in Maktau next to the cropland (1076 m a.s.l., Fig. 1, Fig. 2e 2b). In Taita Taveta this region, bushland is found both within the conservation areas and under private ownership. In this region, bushland Bushland forms a cover with over 50 % of thorny shrubs and small trees, characterized by *Acacia spp* and *Commiphora spp*. The bushes may vary in height ranging from two to five metres. Herbs and savanna grasses (mostly annual or short-lived perennials) less than one-metre tall form the ground cover. On private Private bushland similar to our study site, individual is used by the farmers usually maintain woodlots of trees and shrubs as part of their farms, where they to generate a small income from forest products such as timber, poles, and firewood, and charcoal to some extent. Additionally, some grazing occurs on the bushland, primarily by livestock owned by the farmer (CIDP, 2014).

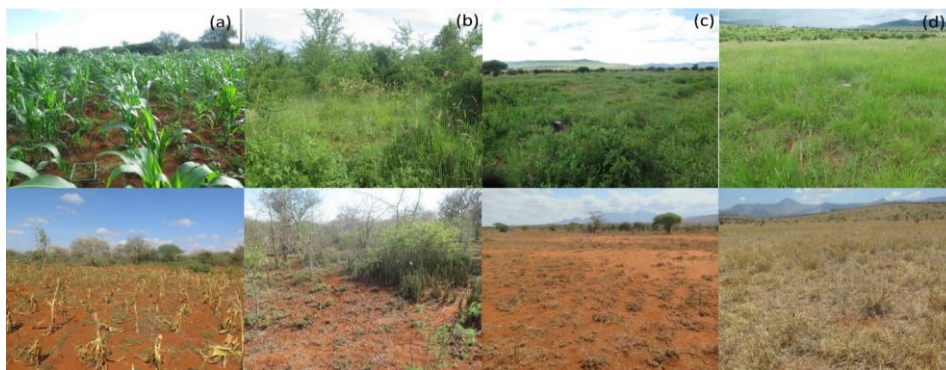
The third site, grazing land (covering approximately 460 km²) is located in the LUMO Community Wildlife Sanctuary (970 m a.s.l. Fig. 1, Fig. 2c) next to Tsavo West National Park and Taita Hills Wildlife Sanctuary and covering approximately 460 km². The sanctuary was formed by merging three ranches, namely the Lualenyi and Mramba communal grazing areas and the Oza group ranch, which were given thus the name "LUMO". This sanctuary is communally owned (GoK, 2013) designated for community livestock grazing by local communities, with, where wildlife is also present, as conservation areas are not necessarily fenced. No individual land ownership occurs, as the land is communally owned (GoK, 2013). Overgrazing However, overgrazing is a major challenge,

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AC: Corrected to spp

especially caused by herders who enter the conservancy illegally, leaving the soil bare for most of the year but especially during the dry season (CIDP, 2014). During this time, livestock (CIDP, 2014).

The fourth site is forced from LUMO into the Taita Hills Sanctuary—the conservation land because of the open boundary it shares with LUMO (Fig. 2d).

The conservation land is located within the Taita Hills Wildlife Sanctuary (928 m a.s.l., Fig. 4). Covering 1, Fig. 2d) covering an area of ca. 110 km² hectares, it. This is a private game sanctuary for wildlife conservation located between LUMO and communal land. The sanctuary is an open savanna grassland dominated by *Schmidtia bulbosa* and *Cenchrus ciliaris* grass species forming an open to closed ground cover, scrublands, shrublands, and scattered woodlands with *Acacia spp.* as main tree species. However, most trees have been damaged by elephants, leaving the landscape more open. The sanctuary is well managed with the application of ecological management tools such as controlled fires. Through these and other conservation efforts, the sanctuary has attracted a higher diversity of large mammals, many of which remain within the unfenced sanctuary throughout the year. Wildlife are the predominant grazers and browsers, although livestock encroachment may be a problem especially during the dry season on the western and eastern borders of the sanctuary (GoK, 2013) (Fig. 2e) (GoK, 2013).



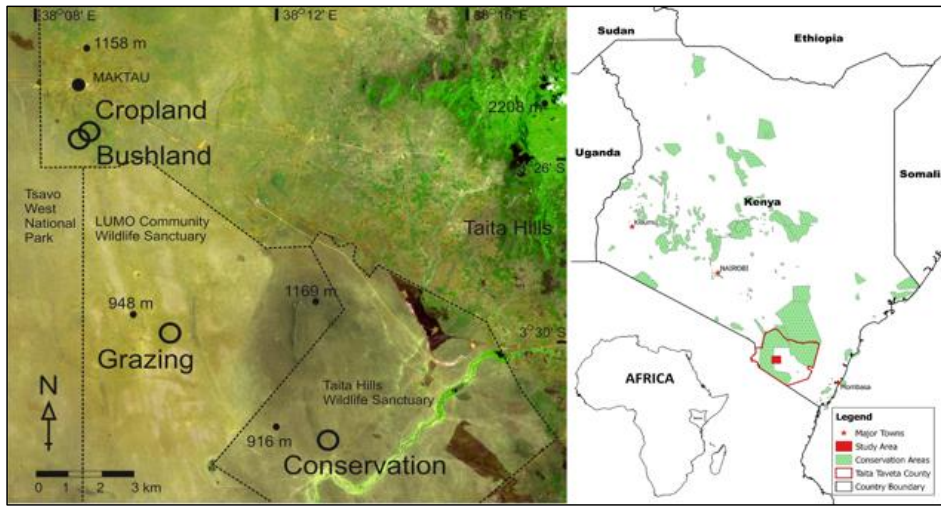
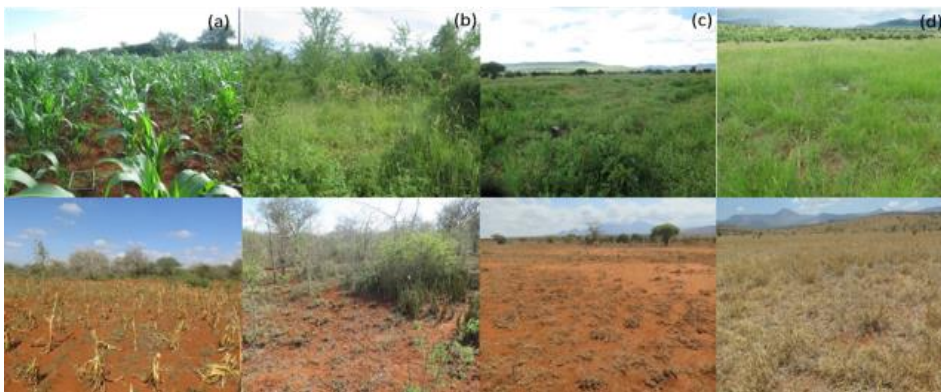


Figure 1. Location of the study sites cropland, bushland, grazing land, and conservation land in the savanna area in the lowlands of Taita Taveta County in southern Kenya.



230 **Figure 2:** The four land-use types: (a) cropland, (b) bushland, (c) grazing land, and (d) conservation
 land. The upper panel shows the land-use types during the wet season, while the ~~the~~ lower panel
 depicts the situation during the dry season. The grey plastic collars visible in upper left photo are
 frames for the GHG flux chambers.

2.2. Defining the ~~Seasons~~seasons

235 We divided ~~theour~~ campaigns into dry and wet seasons, based on ~~anthe~~ agro-climatic concept.
 Therefore, ~~theThe~~ onset ~~date~~ of the wet season was the first wet day of a 3-day wet spell receiving at
 least 20-mm without any 10-day dry spell (< 1-mm) in the ~~followingnext~~ 20-days from 1st of March

for the long wet season and 1st of September for the short wet season (Marteau et al., 2011)(Marteau et al., 2011). Equally, the end of the rainy season was the first of 10 consecutive days during which with no rain occurred. Thus, for this study, the long wet season (LW) was between 2 March to 4 June 2018, and the short wet season (SW) between 23 October and 26 December 2018. The two wet seasons were separated by two dry seasons, the short dry season (SD) from January to February 2018, and the long dry season (LD) from June to September 2018. We had three campaigns during in each of the wet season: during the early days of the wet seasons onset (onset-SW, onset-LW), at the peak of the of the seasons (mid-SW, mid-LW), and towards as well as at the end of the seasons (end-SW, midend-LW).

2.3. Chamber measurements of greenhouse gas emission

Soil atmosphere exchange of CO₂, N₂O, and CH₄ were measured in eight one week campaigns between 29 November 2017 to 3 November 2018 using the static chamber method (Rochette, 2011; Hutchinson et al., 1981). Within each of the four sites, three locations (clusters) were randomly selected and used as replicates for soil GHG concentration measurements. In each cluster, three plastic collars (27 cm × 37.2 cm × 10 cm) were inserted (5–8 cm) into the soil at least 24 hours before the first sample was taken (see Pelster et al., 2017). The collars were left in the ground for the entire measurement period to minimize soil disturbance before the time of measurements (Søe et al., 2004). Only damaged or missing chamber bases (mostly due to livestock or wildlife activity) were replaced, at least 24 hours before the next gas sampling. During each day of a campaign, gas sampling was conducted daily between 7:00 am and 11:00 am, which represent the average flux of the diurnal cycle (Shi et al., 2012; Davidson et al., 1998).

During each gas sampling day, grey opaque PVC lids (27 cm × 37.2 cm × 12 cm) covered with reflective tape were placed onto the collars for 30 mins. Lids were fitted with a fan for gas mixing and a vent to avoid pressure differences between the chamber headspace and outside atmosphere during gas sampling (Pelster et al., 2017). A rubber seal was fitted along the edges of the chamber lid and paper clips were used to hold the lid and collar in place to ensure airtightness. Four gas samples were then collected every 10 minutes (time 0, 10, 20, 30 minutes) after lid deployment (Rochette, 2011). The height of each chamber collar was measured on each sampling date to derive the total chamber volume (total chamber height = height of chamber collar sticking out of the soil + height of the chamber lid). A slightly modified version of the gas pooling method was used to reduce overall sample size while ensuring a good spatial representation of each LUT (see Arias-Navarro et al., 2013). Here, 20 ml of headspace air were collected from each of the three chambers at each time interval

270 with polypropylene syringes (60 ml capacity), resulting in a composite gas sample of 60 ml. The first
275 40 ml were used to flush the vials, and the remaining 20 ml were pushed into 10 ml glass vials, leading
to a slight overpressure to minimize contamination of the gas with ambient air during transportation
(Rochette et al., 2003).

280 Following sample collection, gasSoil-atmosphere exchange of CO₂, N₂O, and CH₄ were measured in
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vials, and the remaining 20 ml overpressured into 10 ml glass vials to minimize contamination of the
gas with ambient air during transportation (Rochette et al., 2003).

300 Gas samples were transported to the laboratory (Mazingira Centre, mazingira.ilri.org) and analysed
using a gas chromatograph (GC, model SRI 8610C ~~gas chromatograph~~). The GC was fitted with
a ⁶³Ni-Electron Capture Detector (ECD) for detecting N₂O concentrations and a Flame Ionization

Detector (FID) fitted with a methanizer for CH₄ and CO₂ analysis. The GC was operated with a Haysep D packed column (3 m, 1/8") at an oven temperature of 70 °C, while ECD and FID detectors were operated at a temperature of 350 °C. Carrier gas (N₂) flow rate was 25 mL min⁻¹ on both FID and ECD lines. In every 40 samples analysed with the GC were eight calibration gases with known CO₂, CH₄, and N₂O concentrations in synthetic air (levels of calibration gases ranged from 400 to 2420 ppm for CO₂, 360 to 2530 ppb for N₂O, and 4.28 to 49.80 ppm for CH₄). Therefore, the gas concentrations of the samples were calculated from peak areas of samples in relation to peak areas of standard gases ~~with known concentrations~~ using a linear model for CO₂ and CH₄ and a power regression for N₂O.

2.4. Greenhouse gas flux calculations

Soil GHG emissions were determined by the rate of change in gas concentration in the chamber headspace over time by linear fitting. The goodness of fit was used to evaluate the linearity of concentration increases/decreases. The dynamics of the CO₂ concentrations over the 30 min deployment period for each gas concentration was assessed to test for chamber leakage due to the typically more robust and continuous flux of CO₂ (Collier et al., 2014; Collier et al., 2014). If the linear model of CO₂ versus deployment time had an R² \geq 0.95 using all four time points (T1, T2, T3, and T4), the measurement was considered valid and four time points were used for analysing the CO₂, N₂O, and CH₄ emissions. However, if R² \leq 0.95 for CO₂ and one data point was a clear outlier, this point was discarded and the three remaining points used for the flux calculation if they showed a strong correlation of CO₂ versus time. Measurements that did not show a clear trend of CO₂ with time were considered faulty, and the entire data point series was discarded. In addition, data points that showed a decrease in CO₂ concentration over time were assumed to indicate leakage and were ~~thus similarly~~ discarded (chambers were opaque, i.e. photosynthesis was inactive during chamber deployment). However, if no leakage was found, negative CH₄ and N₂O emissions were accepted as the uptake of the respective gas by the soil. Emissions were calculated according to Eq. (1):

$$F_{GHG} = \frac{\left(\frac{dc}{dt}\right) \times V_{ch} \times M_w}{A_{ch} \times Mv_{corr}} 60 \times 10^6, \quad (1)$$

where F_{GHG} = soil GHG flux (CO₂, N₂O, or CH₄), $\partial c/\partial t$ = change in chamber headspace gas concentration over time (i.e. slope of the linear regression), V_{ch} = volume of the chamber headspace (m³), M_w = molar weight (g mol⁻¹) of C for CO₂ and CH₄ (12) or N for N₂O (2x N = 28), A_{ch} = area covered by the chamber (m²) and Mv_{corr} = ~~pressure and temperature corrected molar volume~~

(Brümmer et al., 2008) using Eq. (2). With 60 and 10^6 being constants used to convert minutes into hours and microgramme respectively.

Mv_{corr} = Where F_{GHG} = soil GHG flux (CO_2 , N_2O , or CH_4), $\partial c/\partial t$ = change in chamber headspace gas concentration over time (i.e. slope of the linear regression), V_{ch} = volume of the chamber headspace (m^3), M_w = molar weight ($g\ mol^{-1}$) of C for CO_2 and CH_4 (12) or N for N_2O ($2 \times N = 28$), A_{ch} = area covered by the chamber (m^2) and Mv_{corr} = pressure- and temperature-corrected molar volume (Brümmer et al., 2008) using Eq. (2). With 60 and 10^6 being, constants used to convert minutes into hours and micrograms respectively. Temperature in Eq. (2) represent air temperature in the chamber headspace measured during each sampling.

$$Mv_{corr} = 0.02241 \frac{273.15 + Temp(^{\circ}C)}{273.15} \times \frac{Atmospheric\ pressure\ at\ measurement\ (Pa)}{Atmospheric\ pressure\ at\ sea\ level\ (Pa)} \quad (2)$$

The minimum limit of detection (LOD) for each gas was calculated following Parkin et al. (2012) and levels were $\pm 4.9\ mg\ CO_2-C\ m^{-2}\ h^{-1}$ for CO_2 , $\pm 0.04\ \mu g-\mu g\ N_2O-N\ m^{-2}-h^{-1}$ for N_2O , and $\pm 0.03\ mg\ CH_4-C\ m^{-2}\ h^{-1}$ for CH_4 . However, we included all data in the analysis, including those below LOD in line with Croghan and Egeghy (2003), who noted that including such data provides an insight on the distinct measurements, thus giving clarifying the set of environmental observations.

2.5. Auxiliary measurements

During each gas-sampling day, we measured soil moisturewater content (WC) and soil temperature (T) (at a depth of 0–5 cm) adjacent to the collar using a ProCheck handheld data logger with a GS3 sensor (Decagon Devices ProCheck METER Group, Inc. USA). Daily air temperature and precipitation data from November 2017 to November 2018 were obtained from a weather station in Maktau located within the cropland site (Tuure et al., 2019). A soil auger was used to collect soil samples (at a depth of 0–20 cm) during the wet season (22 May 2018) from each land use typesite for soil chemical and physical property analysis. For bulk density, we collected a combination of three samples from each cluster close to each chamber collar at depths of 0–10 cm and 10–20 cm using a soil bulk density ring (Eijkelkamp Agrisearch Equipment, Giesbeek, The Netherlands). Samples were stored in airtight polyethylene bags and kept in a cooler box with ice packs before transportation to the laboratory for further analysis. In the laboratory, samples were stored in a refrigerator (4 °C) and analysed within 10 days.

Field Code Changed

The samples were sieved at <2 mm before analysis. Soil water content was measured by drying soil at 105-°C for 48-h. Soil pH was determined in a 1:2.5 (soil: distilled water) suspension using an electrode pH meter (3540 pH and conductivity Meter, Bibby ~~scientific~~Scientific Ltd, UK). ~~We measured~~ and soil texture using the hydrometer technique (~~Scrimgeour, 2008; Reeuwijk, 2002~~)(Scrimgeour, 2008; Reeuwijk, 2002). Total soil C and N content ~~were analysed using a C/N elemental analyser as follows. A~~ a duplicate of 20 g of fresh sample was oven-dried at 40 °C for 48 hours and ground into a fine powder using a ball mill (Retsch MM400). Approximately 200 mg of the dry sample ~~was~~were measured by elemental analysis (Vario MAX Cube Analyzer Version 05.03.2013).

2.6. Statistical Analysis

~~All statistical~~Statistical analyses were carried out using R ~~Statistical Software (R 3.5.2-~~R Core Team). Spearman correlation coefficients were performed among the variables followed by the Kruskal Wallis test to assess ~~significant differences~~ of soil GHG emissions between the ~~land use types~~LUTs and across seasons. A post-hoc analysis involving pairwise comparisons using the Nemenyi test was performed ~~for gases~~ where significant differences ~~exist~~existed. Significance level was set at $p < 0.05$.

~~We used several functions to assess~~assessed the correlation between soil GHG (~~CO₂ and N₂O~~) with ~~soil temperature and soil water content~~ emissions with T and WC using ~~several functions~~ based on the coefficient of determination (R^2), root-mean-square error (RMSE) and Akaike's information criterion (AIC). There being no difference in the outputs, we present results from the Gaussian function (~~O'Connell, 1990~~)(O'Connell, 1990) for the correlation between soil ~~GHG (CO₂ and N₂O)~~ emissions and ~~soil temperature~~T using Eq. (3), and a quadratic function for correlation with ~~soil water content~~WC using Eq. (4). ~~We also evaluated the combined effect of T and WC on soil GHG emissions by combining Eq. (3) and Eq. (4) into Eq. (5) to assess the effect of these two variables on the emissions.~~

$$Rs = ae^{(bT+cT^2)} \quad (3)$$

$$Rs = a + bWC + cWC^2 \quad (4)$$

~~We also evaluated the combined effect of soil temperature (T) and soil water content (WC) on soil GHG (CO₂ and N₂O) using several functions. Having also found no significant difference in the~~

390 ~~function result outputs, we combined Eq. (3) and Eq. (4) into Eq. (5) to assess the combined effect of these two predictors on soil emissions.~~

$$R_s = e^{(aT + bT^2)} \times (cWC + dWC^2) \quad (5)$$

~~Where R_s is soil GHG (CO_2 and N_2O) emissions, T is soil temperature ($^\circ\text{C}$), and WC is soil water content ($\text{m}^3 \text{m}^{-3}$), while a , b , c , and d represent the model coefficients.~~

395 After no correlation with ~~soil water content and soil temperature was~~ T and a weak correlation with ~~WC were~~ observed, we included ~~information on~~ vegetation cover as a predictor. ~~We~~ For this, we used Normalized Difference Vegetation Index (NDVI) ~~products (MOD13Q1) from MODIS (Moderate Resolution Imaging Spectroradiometer (MODIS) from~~ <https://ladsweb.modaps.eosdis.nasa.gov>. NDVI quantifies vegetation vigour by measuring the difference between reflectance in near-infrared (which green chlorophyll-rich vegetation strongly reflects) and red wavelength areas (which vegetation absorbs) computed using Eq. ~~(6)-(6)~~. ~~MOD13Q1 products from MODIS are NDVI data generated from a 16-day interval at a 250 m spatial resolution as a Level 3 product (Didan, K, 2015).~~

$$NDVI = \frac{NIR + Red}{NIR + Red} \quad (6)$$

405 ~~MODIS NDVI is generated from a 16 day interval at a 250 m spatial resolution as a Level 3 product. To cover our study period, we selected NDVI data that were within the dates of campaign. However, if dates. If no NDVI data fitted within our dates, we used data that were from less than five days before or after the campaign dates, assuming that no significant increase or decrease would occur in the vegetation. The pixels containing the study sites were extracted based on the latitude and longitude of each site. When comparing various functions, linear~~ Linear functions were applied to the seasonal datasets of ~~soil GHG emissions~~ R_s with NDVI to assess the contribution of vegetation ~~indices on soil emissions using Eq. (7) and a combined effect of WC and NDVI on soil GHG~~ CO_2 emissions (R_s) using Eq. ~~(7)-(8)~~.

$$R_s = a + bNDVI \quad (7)$$

415 ~~We also assessed the combined effect of WC and NDVI on soil CO_2 emissions using Eq. (8).~~

$$R_s = a + bNDVI + (cWC + dWC^2) \quad (8)$$

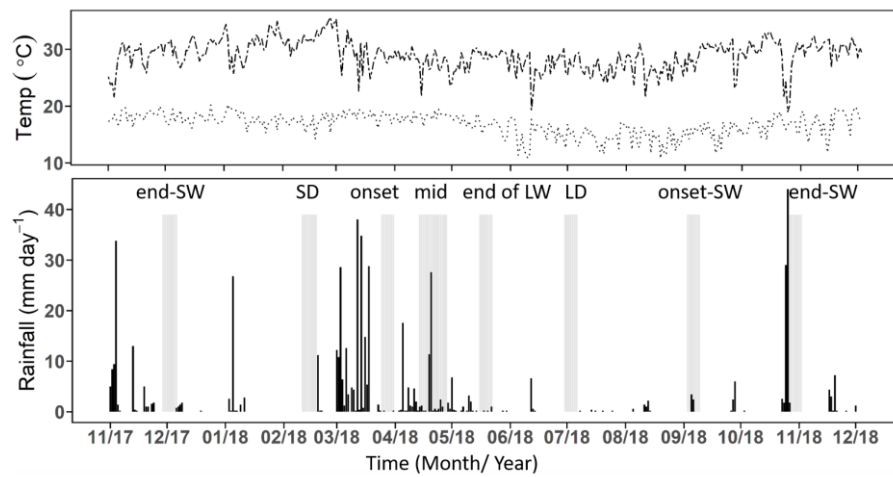
where R_s is soil GHG (CO_2 and N_2O), T is soil temperature, and WC soil volumetric water content ($\text{m}^3 \text{m}^{-3}$) while a , b , c , and d are model coefficients.

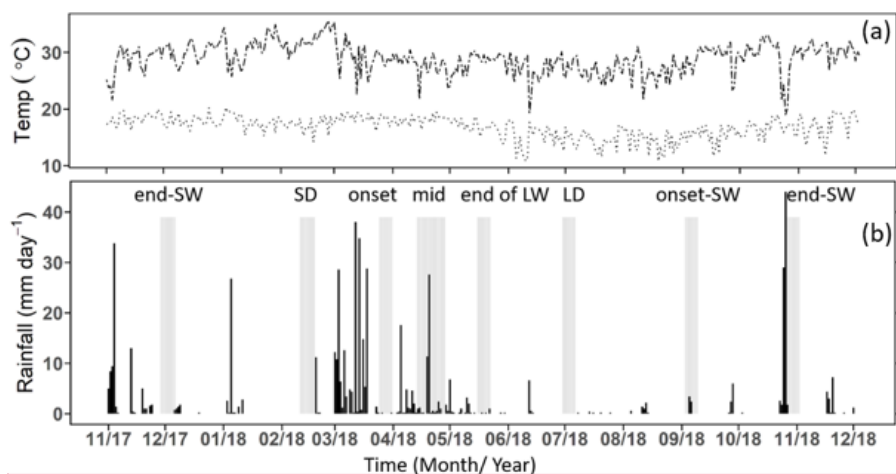
3. Results

420 3.2. Meteorological data

During the 12-month study period, ~~the long rainy periods were observed~~ rains lasted from early March to the end of May, while short ~~rainy periods were observed~~ rains occurred between early September and ~~October~~ December (Fig. 3). The total annual rainfall was 550 mm, which ~~was~~ within the average rainfall ~~quantity~~ expected in the area (CIDP, 2014). ~~Mean~~ ~~The mean~~ annual air temperature was 22.7 °C (min=16.7 °C, max=30.5 °C). January was the hottest month (min=17.4 °C, max=31.9 °C), while June and July (min=14.5± 0.2 °C, max=27± 0.1 °C) were the coolest.

425





Commented [WS17]: RC: Label Fig. 3 as a and b
AC: Done

Figure 3: (a) Daily maximum and minimum air temperature and (b) daily rainfall from lowland in southern Kenya between November 2017 to October 2018 recorded at Maktau weather station. Total annual recorded rainfall was 550 mm. Highlighted grey bars show the days of the sampling campaigns (the season above the grey bars denote SW and LW for the short and long wet season with corresponding onset, mid and end of the wet season, and SD for the short dry season and LD for the long dry season).

3.3. Soil characteristics

Sand was the main component of the soils at all sites (Table 2). Sand content was slightly higher in cropland ($77 \pm 8\%$) compared to the conservation area and bushland (ca. $72 \pm 1\%$). Sand proportion was lower in the grazing land ($64.3 \pm 0.4\%$), while (See Table 2). Grazing land had the highest clay content was higher ($31.7 \pm 0.5\%$) than in the conservation area ($26 \pm 2\%$), bushland ($23.7 \pm 0.7\%$), and while cropland ($19 \pm 2\%$) had the lowest. Soil pH ranged between slightly acidic in the grazing land (6.3 ± 0.3), neutral in the bushland (7.2 ± 0.4), and slightly alkaline in the conservation area and cropland (7.5 ± 0.1 and 7.9 ± 0.2 respectively). Carbon content ranged from 0.93 % in the conservation land to 0.60 % in the cropland. Nitrogen content did not vary significantly between sites (mean = $0.08 \pm 0.01\%$).

Table 1: Soil characteristics of the topsoil (a depth of 0–20 cm) from the four land-use types investigated in this study. Values are given as mean \pm SE.

Commented [WS18]: Rephrased based on RC2

Land Use	Bulk Density (g cm ⁻³)			Soil Texture			
	% N	% C		pH	% Clay	% Sand	% Silt
Bushland	0.08 (0.03)	0.77 (0.5)	1.31 (0.2)	7.2 (0.4)	23.7 (0.7)	71.6 (2.2)	4.7 (2.3)
Conservation land	0.09 (0.02)	0.93 (0.7)	1.27 (0.4)	7.5 (0.1)	26.4 (2.2)	71.6 (0.5)	2.0 (0.0)
Cropland	0.07 (0.04)	0.60 (0.2)	1.26 (0.3)	7.9 (0.2)	19.1 (2.4)	76.9 (8.1)	4.0 (5.1)
Grazing land	0.08 (0.02)	0.83 (0.4)	1.23 (0.5)	6.3 (0.3)	31.7 (0.5)	64.3 (0.4)	4.4 (0.4)

3.4. Soil greenhouse gas emissions

3.4.1. Soil carbon dioxide (CO₂) emissions

450 Mean annual soil CO₂ emissions were ~~significantly higher~~ highest in the conservation land
~~(75±6 mg CO₂-C m⁻² h⁻¹) compared to the other three sites.~~ Concurrently, no significant differences
occurred ~~in CO₂ emissions~~ between grazing land (50±5 mg CO₂-C m⁻² h⁻¹), cropland
~~(47±3 mg CO₂-C m⁻² h⁻¹), and bushland (45±4 mg CO₂-C m⁻² h⁻¹).~~ We observed no ~~significant~~
difference in ~~the~~ CO₂ emissions between the first three seasons, namely SD ~~in February, and~~ onset-
455 LW ~~in March~~ and mid-LW ~~in April~~. However, ~~toward~~ towards the end of the wet season (end-LW) in
May, CO₂ emissions ~~from~~ in the conservation land and grazing land were significantly higher than
~~emissions from~~ cropland and bushland (p<0.05). Through LD, onset-SW, and mid-SW, CO₂
emissions in the conservation land remained significantly higher ~~than emissions from the other three~~
~~sites~~, while ~~emissions those~~ from grazing land dropped during LD and were not different from
bushland or cropland emissions thereafter.

460 Generally, CO₂ emissions were higher ~~during~~ in the wet ~~season~~ seasons than ~~in~~ the dry ~~season~~ seasons
at all sites, ~~showing a bimodal pattern~~ (Fig. 4c). ~~Just after~~ At the onset of the rainy season in early
March, CO₂ emissions increased at all sites by over 200% from SD to LW and dropped during LD by
approximately 70% in grazing land, bushland, and cropland. In the conservation land, the drop from
LW to LD was about ~~20%~~. ~~On average, the highest seasonal mean fluxes were observed during the~~
465 ~~wet season, 20%~~. In the bushland, the highest seasonal mean fluxes were reached in mid-LW
(98±6 mg CO₂-C m⁻² h⁻¹) ~~in early March. However, while~~ in the conservation land
(239±11 ~~mg~~ CO₂-C m⁻² h⁻¹), grazing land (160±16 mg CO₂-C m⁻² h⁻¹), and cropland
(84±12 mg CO₂-C m⁻² h⁻¹), the highest ~~seasonal mean fluxes~~ were observed during end-LW towards
~~the end~~ May. ~~On the other hand, the~~ The lowest seasonal mean CO₂ emissions at all sites were
470 observed during the SD campaign (below 20 mg CO₂-C m⁻² h⁻¹, Fig 4).

When comparing ~~between~~ the two wet seasons (LW and SW), CO₂ emissions were 45-% (bushland),
55-% (conservation land), 56-% (cropland), and 57-% (grazing land) higher ~~during the long wet~~

Commented [WS19]: RC: Different terms are being used, soil respiration; soil CO₂ emissions, ecosystem soil respiration (?) Make this consistent.
AC: We only use the term "soil CO₂ emission".

Commented [WS20]: SC: N₂O fluxes are sometimes given in µg N₂O-N m⁻² h⁻¹ (e.g. l. 295 or 420) and in some graphics they are given in N₂O (µg m⁻² h⁻¹) (e.g. figure 4 or figure 5). This is the same for CO₂. Units should be used consistent, so that you can compare the fluxes.
AC: We made the necessary correction. CO₂ (mg CO₂-C m⁻² h⁻¹), N₂O (µg N₂O-N m⁻² h⁻¹) and CH₄ (mg CH₄-C m⁻² h⁻¹).

Commented [WS21]: RC: Ln 365 – What are the errors on the fluxes? They are so small for soil CO₂ fluxes. Report error and sample size.
AC: The sample size per season is seven daily average values derived from three flux values per day from each land use type. The error bar presented here is the standard error of the three flux values per day.

season in LW than during the short wet season SW (Fig. 5a). For the two dry seasons, CO₂ emissions were also significantly higher in LD than SD across all the sites (in SD all sites recorded emission below 20 mg CO₂-C m⁻² h⁻¹). However, during the LD, CO₂ emissions were 29% (bushland), 38% (cropland), 40% (grazing land), and 77% (conservation) higher than during SD (Fig. 5a). While, as much as CO₂ emissions in cropland, bushland, and grazing land had dropped below to less than 30 mg CO₂-C m⁻² h⁻¹ during LD, CO₂ emissions were still high (118±6 mg CO₂-C m⁻² h⁻¹) in the conservation land, even more than during onset LW (63±9 (118±6 mg CO₂-C m⁻² h⁻¹) and mid-LW (79±5 mg CO₂-C m⁻² h⁻¹) the emissions remained high (Fig 4c).

3.4.2. Soil nitrous oxide (N₂O) emissions

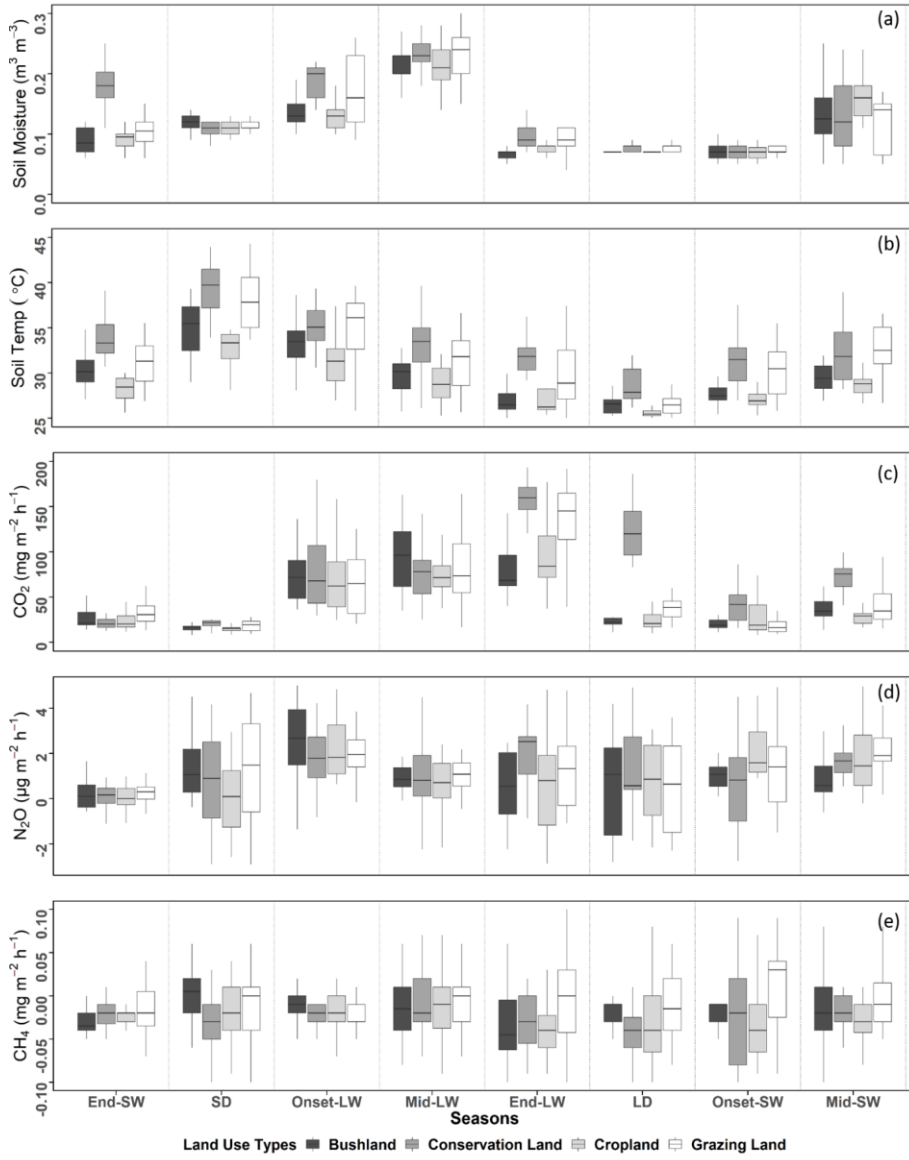
Mean annual N₂O emissions were very low (< 5 μg N₂O-N m⁻² h⁻¹) at all four sites (Fig. 4d). Mean N₂O emissions were higher in the cropland (2.657±0.066 μg N₂O-N m⁻² h⁻¹) recorded the highest mean N₂O emissions than in the conservation land (1.596±0.044 μg N₂O-N m⁻² h⁻¹), grazing land (1.495±0.044 μg N₂O-N m⁻² h⁻¹), and bushland (1.462±0.044 μg N₂O-N m⁻² h⁻¹). N₂O fluxes did not show a clear temporal pattern as observed for CO₂ emissions. Mean seasonal N₂O emissions were very low during both the wet and dry seasons. Within each season, no significant differences in N₂O emissions were observed among the sites. However, at the onset of the rainy season (onset-LW), there were observable increases in N₂O emissions from all the sites. During this period, mean N₂O emissions at all the sites were ca. 2.6±0.4 μg N₂O-N m⁻² h⁻¹. By mid-LW and end-LW periods, N₂O emissions had dropped to nearly zero (<1 μg N₂O-N m⁻² h⁻¹) at all sites. However, in June during LD in June, N₂O emissions in the cropland were significantly higher than at the other three sites (2.35±0.03 μg N₂O-N m⁻² h⁻¹, p <0.05). During this period, the farmer had just harvested his crops.

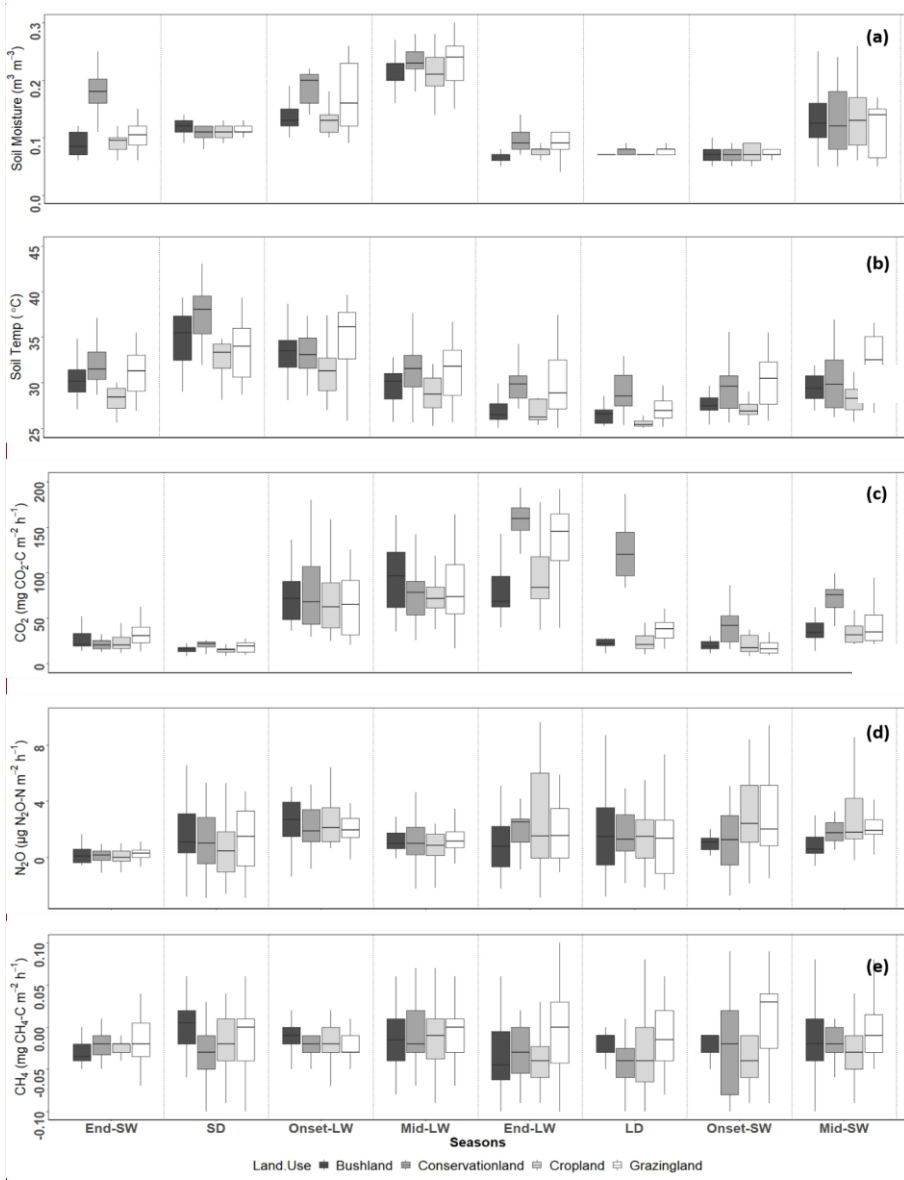
When comparing the two wet seasons, N₂O emissions did not differ between LW and SW at all sites (Fig. 5b). However, short N₂O emission pulses were observed during both seasons. A notable peak of approximately about 70 μg N₂O-N m⁻² h⁻¹ was observed in the cropland on 7 April 2018, which was a week after the application of livestock manure a week before the sampling campaign application. It had also rained the night before the sampling day. At the same site, we also recorded a peak of 55.2 μg N₂O-N m⁻² h⁻¹ on 30 September 2018, likely also due to manure application (personal communication from the farmer Mwadime Mjomba). Other notable peaks were 29.9 μg N₂O-N m⁻² h⁻¹ (in the bushland on 03 September 2018) and 26.59-6 μg N₂O-N m⁻² h⁻¹ (in grazing land on 4 September 2018). These were observed during the SW from chambers with animal

505 droppings ~~within the chambers on these dates. During. For~~ the dry ~~season~~seasons, N₂O emissions did not differ between SD and LD in the bushland, conservation land, and grazing land, while N₂O emissions in the cropland were significantly higher during LD than SD (Fig. 5b).

3.4.3. Soil methane (CH₄) emissions

510 Throughout the study period, CH₄ emissions did not vary significantly among sites and seasons (Fig. 4e and Fig. 5c). The studied sites were mostly CH₄ sinks rather than sources, and CH₄ ~~emissions~~fluxes were very low, ranging from -0.03 to 0.9 mg CH₄-C m⁻² h⁻¹ (Fig. 4e), ~~and were~~ often below the ~~LOD~~limit of detection (0.03 mg CH₄-C m⁻² h⁻¹).

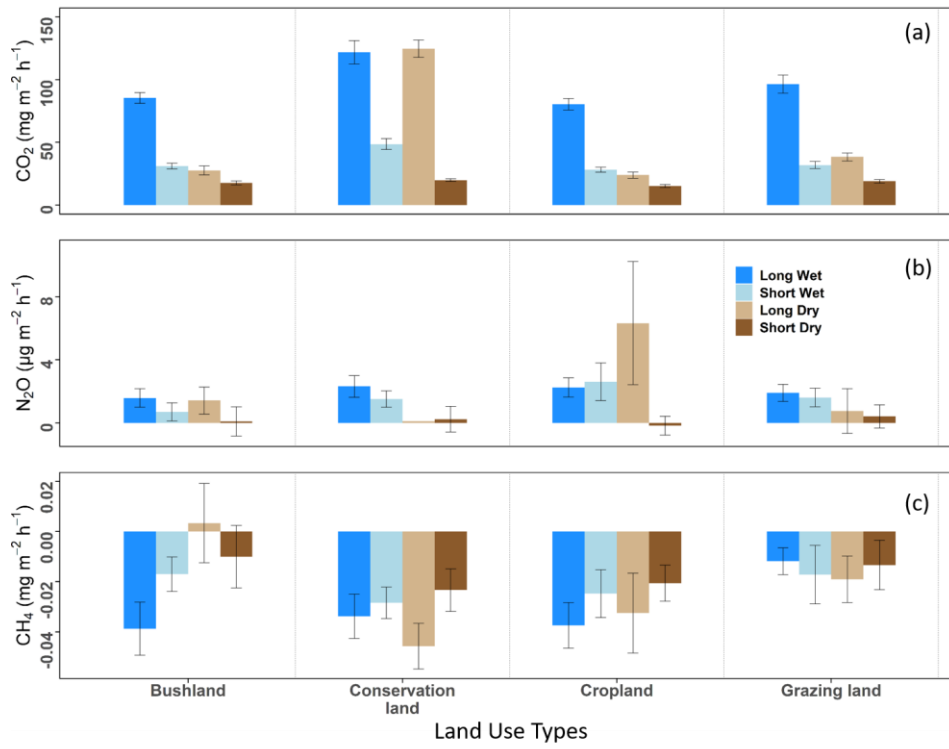




Commented [WS22]: RC: Figures 4 and 5 are good. Keeping the color scheme, the same would be helpful.
AC: The two figures provide different information and having them in the same color scheme might be misleading. Figure 4 shows the difference in emissions between the land uses types while figure 5 gives the differences in emission between the wet and dry season.

515 **Figure 4:** Box plots showing differences in seasonal means for (a) soil moisture, (b) soil temperature, and soil emissions of (c) CO₂, (d) N₂O, and (e) CH₄ for each site from November 2017 to October 2018. Season abbreviations on the x-axis denote SW for the short wet season and LW for

the long wet season with corresponding onset, mid and end of the wet season, along with SD for the short dry season and LD for the long dry season.



520

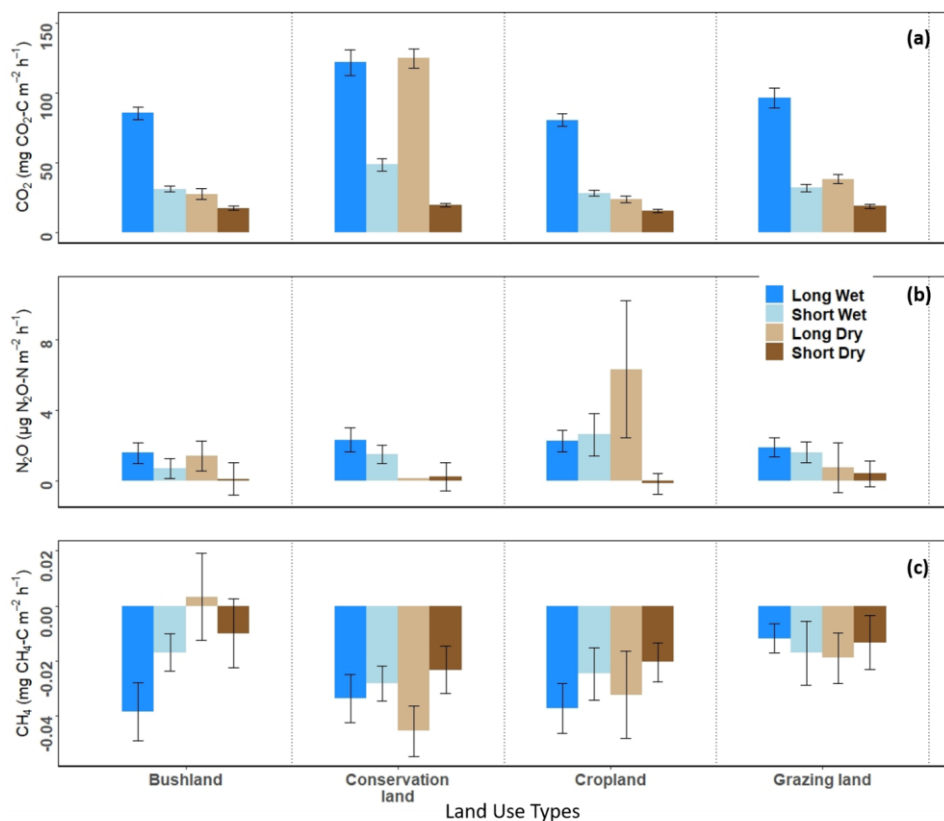


Figure 5: Seasonal differences in mean (a) CO₂, (b) N₂O, and (c) CH₄ emissions between the long and short wet seasons and the long and short dry season for the four land-use types.

3.5. Effects of soil temperature, soil water content, and vegetation indices on soil GHG emissions

525 Soil water content (WC) was highest during the wet season, ranging between (mean 0.13 ± 0.25
 $06 \text{ m}^3 \text{ m}^{-3}$) and lowest during the dry season, ~~ea~~ (mean $0.07 \pm 0.02 \text{ m}^3 \text{ m}^{-3}$) at all sites (see Fig. 4a).
 Soil temperature (T), ~~on the other hand,~~ were ~~higher~~ highest during the SD ($36.7 \pm 2.1 \text{ }^\circ\text{C}$) at all the
 sites and lowest ($24.5 \pm 0.6 \text{ }^\circ\text{C}$) during LD in July (Fig. 4b). Throughout all the campaigns, mean WC
 and mean T were highest in the conservation land, followed by grazing land, bushland, and ~~were~~
 530 lowest in the cropland.

Results from the nonlinear regression analyses Regression results on soil CO₂ and soil N₂O emissions against T and WC are shown in Table 2. The results showed positive correlations between soil CO₂ emissions and WC (p < 0.05). However, the R² was very weak at all sites. Conversely, CO₂ emissions showed no correlation with T (P < 0.05). ~~There was~~ We observed no correlation between N₂O and CH₄ emissions with ~~both~~ either WC or T (p < 0.05). Separating data into the wet and dry season did not improve ~~any of the correlations as we had expected. However, we observed an increase in both CO₂ and N₂O emissions at the onset of the rainy season in March.~~

Table 2: Soil water content (WC) and soil temperature (T) control on carbon dioxide (CO₂) and nitrous oxide (N₂O). ~~Soil CO₂ and N₂O emissions are~~ denoted by *R_s*, while *a*, *b*, and *c* represent the model coefficient

Predictors	Land Use	CO ₂ -C mg m ⁻² h ⁻¹		N ₂ O-N ug m ⁻² h ⁻¹	
<i>R_s = a + bWC + cWC²</i>					
Soil water content (WC)	Bushland	6.12WC + 0.92WC ²	R ² = 0.26***	19.02WC - 64.11WC ²	R ² = 0.008
	Conservation land	135.27WC - 0.57WC ²	R ² = 0.07**	11.63WC - 7.736WC ²	R ² = 0.009
	Cropland	17.83WC + 0.67WC ²	R ² = 0.04***	28.48WC - 66.63WC ²	R ² = 0.005
	Grazing land	15.03WC + 0.79WC ²	R ² = 0.11***	19.81WC - 53.56WC ²	R ² = 0.002
<i>R = ae^(bT+cT²)</i>					
Soil Temperature (T)	Bushland	1.078e ^{0.267T-0.004T²}	R ² = 0.008	360.25e ^{-0.29T-0.004T²}	R ² = 0.008
	Conservation land	0.001e ^{0.817T-0.014T²}	R ² = 0.015**	0.007e ^{0.457T-0.008T²}	R ² = 0.015
	Cropland	4.568e ^{-0.13T+0.002T²}	R ² = 0.008*	0.007e ^{-0.05T+2.42T²}	R ² = 0.008*
	Grazing land	4.136e ^{0.187T-0.003T²}	R ² = 0.015	2.366e ^{0.057T-0.001T²}	R ² = 0.015

***: p<0.0001, **: p<0.001, *: p<0.05

Results from combined ~~soil water content~~ WC and ~~soil temperature~~ T on soil CO₂ and N₂O emissions did not improve the correlation, as shown in Table 3. ~~We concluded that soil emissions could be an overall effect of other factors than these two parameters.~~ Thus, we ~~chose to include~~ included vegetation indices in our model.

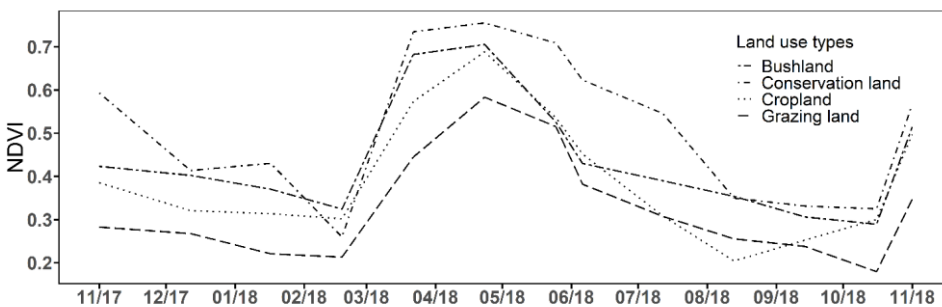
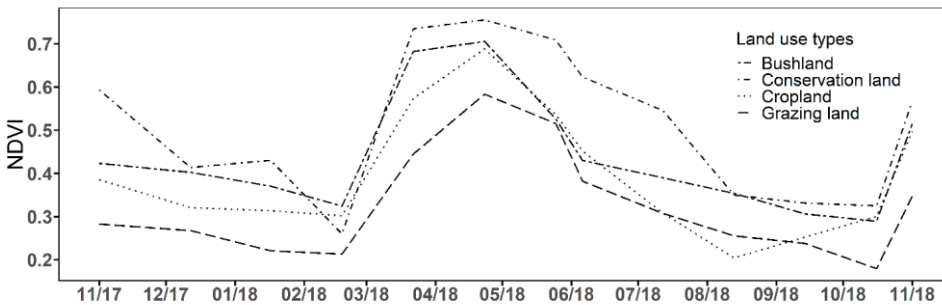
Table 3: Combined effects of soil water content (WC) and soil temperature (T) control on soil CO₂ and N₂O emissions. Soil CO₂ and N₂O emissions denoted by *R_s*, while *a*, *b*, *d*, and *e* represent the model coefficient, (R²) the coefficient of determination, and AIC the Akaike's information criterion

Functions	Land use	a	b	d	e	R ²	AIC
CO₂-C mg m ⁻² h ⁻¹ $R_s = e^{(aT+bT^2)} \times (dWC + eWC^2)$	Bushland	-0.12	0.001	52.774	-0.527	0.31***	1888
	Conservation land	0.90	-0.016	0.0001	0.000	0.10***	2156
	Cropland	-0.39	0.006	3701.901	-84.001	0.08**	1886
	Grazing land	0.14	-0.003	0.842	-0.008	0.12***	2024
N₂O-N ug m ⁻² h ⁻¹ $R_s = e^{(aT+bT^2)} \times (dWC + eWC^2)$	Bushland	-0.50	0.007	2008.345	-58.559	0.009	785
	Conservation land	0.56	-0.010	0.000	0.000	0.003	811
	Cropland	0.67	-0.017	0.003	-0.0001	0.089	911
	Grazing land	0.11	-0.003	0.187	-0.005	0.003	770

***: p<0.0001, **: p<0.001, *: p<0.05

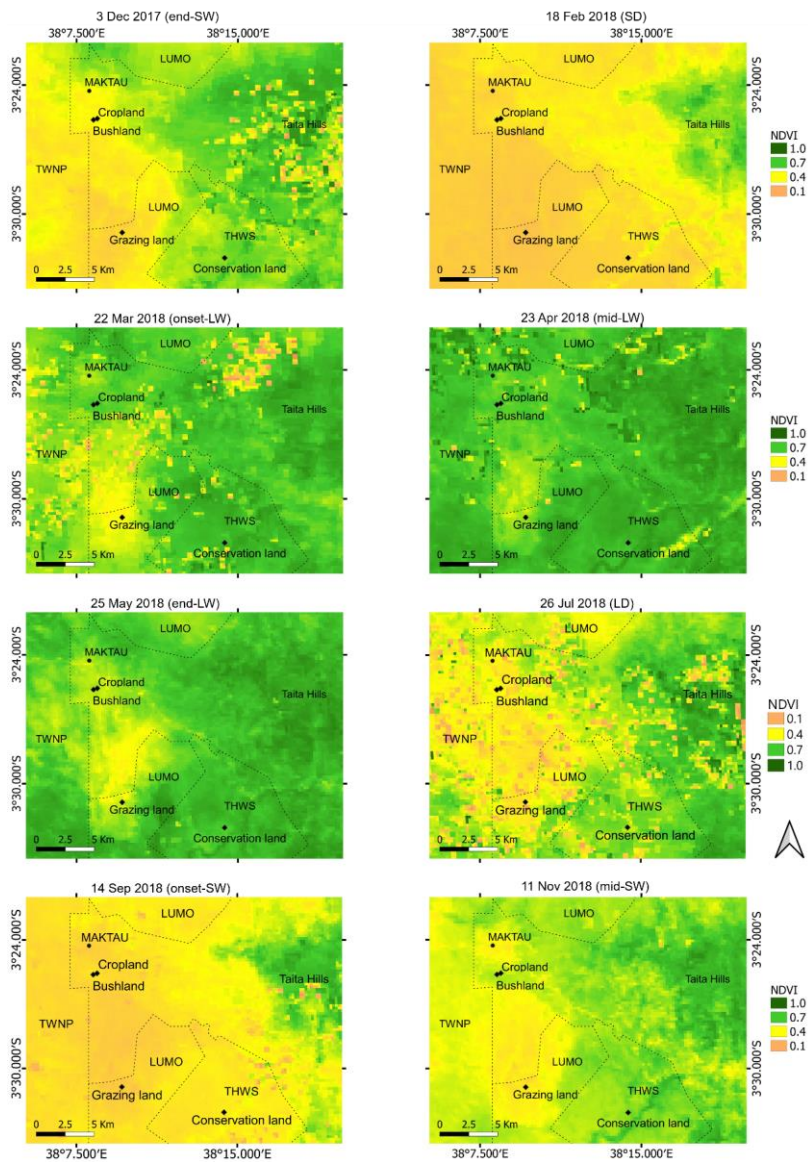
550 The annual change in vegetation cover ~~and during~~ at each ~~campaign in the study areasite~~ are shown
in Fig. (6) ~~and Fig (7) respectively.~~ The highest NDVI values were observed ~~during~~ the LW in April
(ranging from 0.58 to 0.76) and the lowest during the ~~dry season (belowSD (< 0.26)).~~ Vegetation
greenness increased rapidly from mid-March at all sites, ~~which coincided coinciding~~
with the onset of the rainy season and remained high (Fig.6 ~~and Fig. 7).~~ ~~However, at).~~ At end of the rainy season,
555 NDVI gradually dropped. Highest NDVI values occurred in the conservation land (0.51±0.05);
followed by bushland (0.44±0.05), cropland (0.41±0.05), and the lowest ~~values~~ were recorded in the
grazing land (0.33±0.05).

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AC: Done



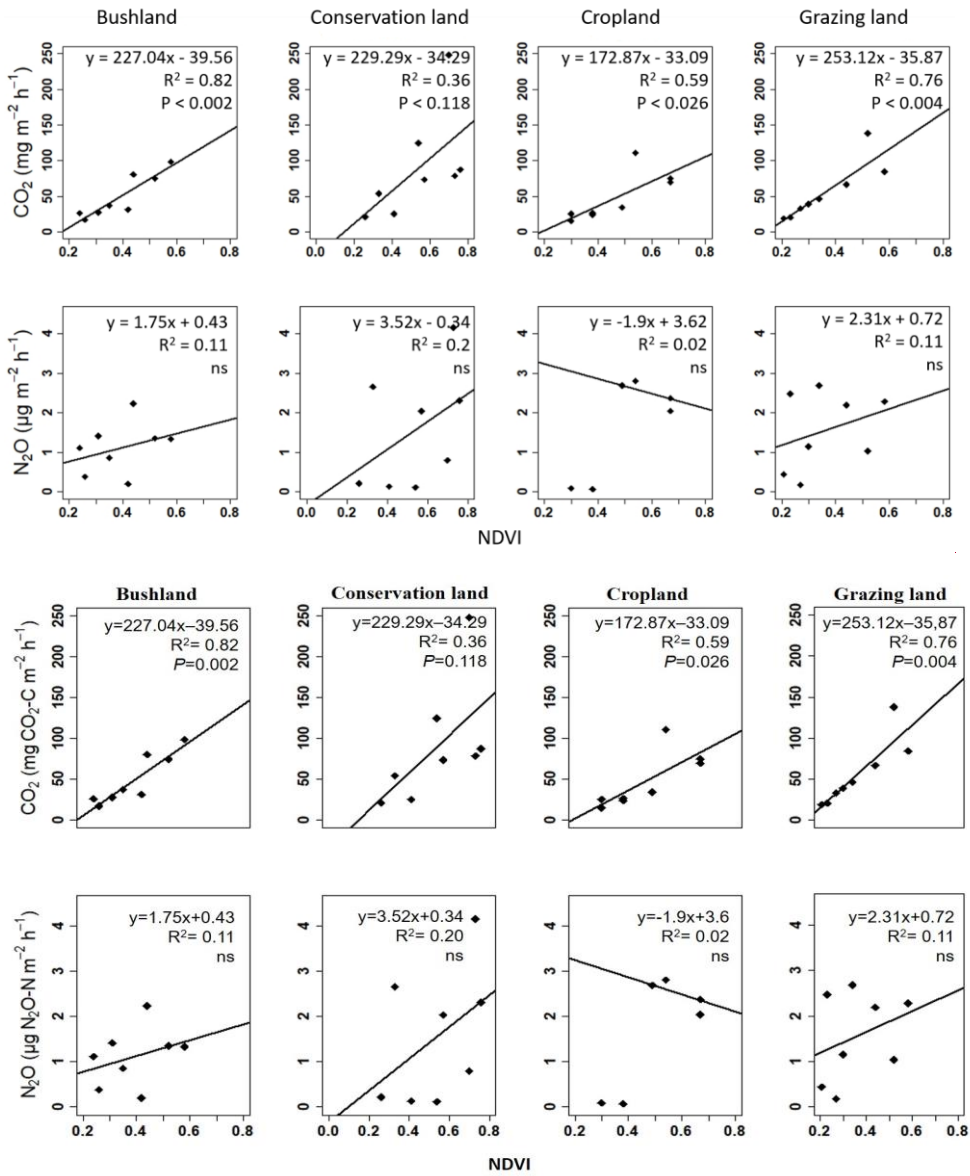
Commented [WS24]: RC1: Figure 6 – put in same color scheme.
AC: This is a relevant point and we tried to have them in the same color scheme of figure 4. However, this then made it very difficult to differentiate between the LUC types.

560 **Figure 6:** Monthly NDVI time series showing the annual trend in vegetation cover from November 2017 to November 2018 for the four land-use types.



565 *Figure 7: Normalized difference vegetation index (NDVI) maps for each campaign (SW and LW denotes short and long wet seasons with corresponding onset, mid and end of the season, and SD and LD for short and long dry season) from November 2017 to November 2018. LUMO is the LUMO community wildlife sanctuary, TWNP stands for the Tsavo west National Park and THWS for Taita Taita hills wildlife sanctuary*

570 ~~Our Regression analysis results show that NDVI was capable of explaining a large degree in the variation of~~ shows a positive correlation between NDVI and seasonal soil-CO₂ emissions ~~in~~ at all the bushland, grazing land, and cropland sites (see Fig. 8). ~~However,~~ 7). Combined WC and NDVI improved the correlation in the conservation land was not even further as shown in Table 4. No significant. ~~We also~~ correlation was observed ~~no significant correlation with~~ between N₂O at all the sites. ~~emissions and NDVI (Fig. 7).~~



ns= not significant

Figure 87: Linear regression analyses of the measured seasonal means for soil CO₂ and soil N₂O emissions during the campaign from November 2017 to November 2018 plotted against NDVI data acquired during each campaign.

Results from combined WC and NDVI improved the correlation outputs further (Table 4). These results highlight the importance of WC and vegetation cover (as represented by NDVI) on soil CO₂ emissions in the study areas.

Table 4: Combined effects of soil water content (WC) and NDVI on soil CO₂ emissions. Soil CO₂ emissions denoted by R_s , while a , b , c , and d represent the model coefficient and (R^2) the coefficient of determination.

	Land use	a	b	c	d	R ²
CO ₂ -C mg m ⁻² h ⁻¹						
$R_s = a + bNDVI + (cWC + dWC^2)$	Bushland	-29.69	196.47	-83.74	781.29	0.86***
	Conservation land	6.48	382.96	-861.98	-256.66	0.82***
	Cropland	26.94	244.54	-1250.22	3269.46	0.79***
	Grazing land	-97.19	396.41	575.60	-3440.80	0.96***

***: p<0.0001, **: p<0.001, *: p<0.05

4. Discussion

4.1. Soil CO₂ emissions

CO₂ emissions from the soil differed significantly between the land use types. Higher CO₂ emissions were particularly observed in the conservation land followed by grazing land and bushland, while lowest levels were recorded from the cropland. We observed the same trend with SOC, and thus we attributed the difference in CO₂ emissions between the land uses to SOC as observed (see Table 1). This is in line with a similar study by La Scala et al. (2000), which recognized SOC as a key driver of CO₂ emissions from the soil, as it is the primary source of energy for soil microorganisms (Lal, 2009).

The differences in CO₂ emissions and SOC among our land use types can be linked to the difference in vegetation cover type, which can alter both biotic and abiotic factors that drives soil CO₂ emissions (Raich and Tufekcioglu, 2000; Pinto et al., 2002) and net carbon assimilation (La Scala et al., 2000). Vegetation types directly influence soil physicochemical properties, which modify soil microbial activities (Raich and Tufekcioglu, 2000). This also affects the quantity of plant carbon allocated belowground (Metealfe et al., 2011). Vegetation additionally affects root respiration by determining

root biomass, and litter quality and quantity (Fanin et al., 2011; Rey et al., 2011). Root respiration and the associated microbial components are important in ecosystem soil respiration. Active roots add directly to soil respiration, while the dead roots and exudates from the roots provides carbon as a source of energy and nutrients for microbial biomass (Tufekcioglu et al., 2001).

610 In the conservation land, thick grasses formed a closed ground cover, especially during the wet seasons (also confirmed by NDVI values). This translated into higher soil respiration due to higher root respiration due to a more active root network and litter production compared to the other sites (Raich and Tufekcioglu, 2000). Besides this, grazing here was low and only occurred by random elephants and other wild mammals. We observed less damage to grass cover in the conservation area
620 compared to grazing land (which was mostly bare due to overgrazing) and bushland. Abdalla et al. (2018) found higher grazing intensity to reduce SOC stocks due to the modification of vegetation cover, which affects litter accumulation and decomposition (Wilsey et al., 2002). This could be the case with the grazing land, which was under heavy grazing by livestock, and also by wildlife. This explains the difference in mean CO₂ emissions between conservation land and grazing land.

625 In the bushland, grazing can be considered moderate when compared to the conservation land and grazing land, as only the farmer's livestock grazed the land. However, the lower CO₂ emissions and lower SOC compared to the grazing land can be a consequence of the higher clay content observed in grazing land (see Table 1). The presence of polyvalent cations in clay forms organo-mineral complexes that protect SOC from microbial and enzymatic decay, which in turn increases SOC storage (Amanuel et al., 2018).

Soil CO₂ emissions and mean SOC were lowest in cropland. Root respiration in cropland depends on periods of live roots in annual crop fields and on the biomass of roots during the initial growing season (Raich et al., 2000). Therefore, the continued removal of crop residues during harvesting and frequent tillage affected both root respiration and SOC. As much as crop residues contribute to carbon stocks
630 through their mineralization (Nandwa, 2001), most maize residues were used as livestock feed and sometimes as fuel, while bean residues were removed completely during the harvest and burned.

Manure inputs provide easily degradable substrates of C and N catalyzing soil emissions (Janssens et al., 2001; Davidson and Janssens, 2006). However, manure input in cropland was very low (approximately 20 kg in a 1.5 hectare farm per month) and thus no measurable effects in CO₂ emissions were detected. This was opposite to our hypothesis. Another reason could be that soil
635 fertility was too low to have a detectable influence on CO₂ emissions (Pelster et al., 2017). Our results

are in the same magnitude with those of Rosenstock et al. (2016); Farai Mapanda et al. (2011), and Pelster et al. (2017), who also did not detect any change in CO₂ emissions after manure application and attributed this to the low input of manure from the maize and sorghum plots.

640 On average, CO₂ emissions were higher during the wet season than during the dry season. At the start of both rainy seasons, CO₂ emissions increased significantly in all land use types. The emission from the conservation land and grazing land are comparable with those in Brümmer et al. (2008), who observed CO₂ emissions ranging between 100 and 250 mg C m⁻² h⁻¹ in a natural savanna in Burkina Faso. Several other studies from similar ecosystem have also documented comparable changes in
645 CO₂ emissions with the onset of the rainy seasons (Castaldi et al., 2006; Livesley et al., 2011; Pinto et al., 2002).

In the cropland, our results during the wet season are similar with those that Rosenstock et al. (2016) measured during the wet season, which ranged between 50 to greater than 200 mg C m⁻² h⁻¹. We attributed the increase in CO₂ emissions during the wet season to the response of soil microbes to soil
650 moisture producing an increase in ecosystem respiration (Livesley et al., 2011; Otieno et al., 2010). Soil moisture allows for soluble substrates and oxygen, both needed by soil microbes, and increasing microbial activity such as decomposition and soil respiration (Davidson et al., 2006; 2009; Grover et al., 2012).

Consequently, higher soil CO₂ emissions during the wet season can be a result of increased root
655 respiration due to more active plant and root growth (Macdonald et al., 2006). Grass sprouts rapidly after the rains, increasing in root network density to maximize the use of available soil moisture on the soil surface (Merbold et al., 2009). This is one explanation for the higher CO₂ emissions in the grassy conservation land, grazing land, and bushland compared to cropland without the grasses. The lower CO₂ emissions in bushland compared to conservation land and grazing land were attributed to
660 the presence of more trees and shrubs, which according to Merbold et al. (2009), respond more slowly than grasses to changes in soil moisture. Therefore, grass production belowground in the conservation land and grazing land was probably more than in the bushland, leading to higher autotrophic respiration and also heterotrophic respiration (Janssens et al., 2001).

To our surprise, the highest seasonal mean CO₂ emissions in conservation land, grazing land, and
665 cropland were observed at the end rather than the at the peak of the wet season. During this time, both documented soil moisture and soil temperature had dropped in all land use types. Thus, we credited the relatively high CO₂ emissions to root respiration that could tap moisture from deeper profiles than

670 microbial activity in the soil (Carbone et al., 2011). According to Carbone et al. (2011), microbial
respiration peaks first with surface soil moisture, whereas root respiration continues to increase
throughout the wet season, controlling emissions as temperatures increase and surface soil moisture
gradually drops. Most microbial activities are found on the soil surface, which are first to wet up and
dry down with rainfall, while roots are located deeper, with access to more water reserves that take
longer to be exhausted (Carbone et al., 2011).

675 **4.1. N₂O emissions**

680 Soil N₂O emissions can vary highly over time, as regulated by factors such as soil moisture,
temperature, aeration, ammonium and nitrate concentrations, pH, and mineralizable carbon
(Butterbach-Bahl et al., 2013). However, we did not document any significant difference in N₂O
emissions between the four land uses. At all the sites, N₂O emissions were very low, and this we
685 attributed to the observed low soil N content (see Table 1). According to Pinto et al. (2002) and
Grover et al. (2012), savanna ecosystems have a very tight N cycling, which transcends to low N
availability. Thus, available N is taken up by vegetation, leaving very little for denitrification
(Castaldi et al., 2006; Mapanda et al., 2011). Our results are consistent with low N₂O emissions
observed in a Brazilian savanna (cerrado) by Wilcke et al. (2005), who also reported low N levels in
their study area. Very low N₂O emissions due to poor nutrient availability have also been observed
in other savanna landscapes (Scholes et al., 1997; Castaldi et al., 2016). Soil N₂O in cropland match
those of Rosenstock et al. (2016), who also attributed the low soil N₂O emissions to poor nutrient
availability in the soil.

690 We did not detect seasonal variations in N₂O emissions. The only exception to otherwise very low
N₂O emissions was after the onset of the rainy season, when N₂O emissions slightly increased at all
sites. Such a pattern has previously been shown by Scholes et al. (1997) in South Africa. Several
other studies have reported a comparable increase in N₂O emissions after the rainy season (Scholes
et al., 1997; Pinto et al., 2002; Castaldi et al., 2006; Livesley et al., 2011). Soil moisture in the savanna
ecosystem controls soil gas diffusion, oxygen (O₂) availability for microbial use and the availability
695 of substrate for microbial communities (Davidson et al., 2000; Butterbach-Bahl et al., 2013).
Therefore, the increase in N₂O emissions at the onset of the wet season is possibly a response of
microbial communities to variation in soil moisture (Rees et al., 2006). In addition, the decomposition
of litter and plant residue facilitated by soil moisture may have further increased N availability.

700 Rosenstock et al. (2016) also recorded low N_2O emissions with average fluxes typically less than $12 \mu g N_2O-N m^{-2} hr^{-1}$ in rainy Kenyan highlands. In cropland, mean N_2O emissions are close to those in Mapanda et al. (2010) in Zimbabwe, who reported an average of $3.3-3.4 \mu g N_2O-N m^{-2} hr^{-1}$. In June and July, the slight increase after the maize and bean harvests could be due to there being no plants to take up the available N and thus some was lost as N_2O .

705 Negative N_2O emissions were detected during the dry season. Such observations could result from the poor N levels observed at all sites. Soil denitrifiers may therefore use N_2O as an N substrate in the absence of NO_2^- and NO_3^- (Rosenkranz et al., 2006). Negative N_2O emissions have also been reported in other tropical savanna soils under similarly dry conditions (Donoso et al., 1993; Castaldi et al., 2006; Livesley et al., 2011).

710 Application of manure in the cropland did not show any significant differences, which is the opposite of what we expected. Several studies have shown both organic and inorganic fertilisers in agricultural land to increase N_2O emissions (Davidson, 2009; Butterbach-Bahl et al., 2013; Hickman et al., 2014). Due to low nutrient levels, as observed, the addition of manure to these soils may not have been sufficient to stimulate high N_2O emissions. However, manure input by the farmer was also very low (less than $12 kg$ of N for the crop growing season). Our conclusion here is that the maize took most of the available N that was added, thus diminishing the pool of N to be lost as N_2O . According to the Taita Development plan, this is a common scenario in the county, which translates to very low yields (CIDP, 2014). Our results are similar to those in Pelster et al. (2017) (generally $< 10 \mu g N_2O-N m^{-2} hr^{-1}$), who also observed no detectable influence of manure application (which they noted was low, between $1-25 kg N ha^{-1}$) on N_2O emissions. Increased nutrient inputs and improved management are required to improve yield and livelihoods, but this may simultaneously lead to increased soil N_2O emissions. Nevertheless, Hickman et al. (2015) suggested that managed agricultural intensification in western Kenya could increase crop yields without immediate increases in N_2O emissions, if application rates remained at or below $100 kg N ha^{-1}$. We believe this could also be the case in our study area.

725 Deposition of faeces and urine by animals while grazing in the grazing land and bushland did not show any statistical difference with cropland and conservation land, as we expected. We found faeces in and close to chambers for most of the sampling days in the grazing land, although we did not see any animal urine in the chambers. In addition, on each sampling day, we also observed animal footprints in our chamber or close by in the same site in the bushland. In the conservation area, as
730 much as animal were grazing, we never observed faeces in or close to our chambers throughout the

campaign period. We only noted the presence of animals close to our chamber from the footprints and the destruction of our chamber from time to time.

1.2. CH₄ emissions

Methane emissions did not vary between the land-use types or with seasons. Most values were below the LOD at all the sites, which is similar to observations made by Rosenstock et al. (2016) and Wanyama et al. (2019).

Soil CO₂ emissions differed significantly between the four LUTs. The highest mean CO₂ emissions were observed in the conservation land followed by grazing land and bushland, and the lowest from cropland. Soil C content, which is the primary source of energy for soil microorganisms that contribute to soil CO₂ emissions (Lal, 2009) also showed the same trend (conservation land > grazing land > bushland > cropland). Therefore, the difference in land use and land-use management activities between our sites played a vital role in modifying both biotic and abiotic factors that drive both soil C content and soil CO₂ emissions (Pinto et al., 2002).

Due to the difference in land use and management, vegetation type and cover differed between our sites. The dense grass network in the conservation land formed an almost closed ground cover, especially in the wet seasons (further confirmed by NDVI values). Being a private sanctuary, only wild mammals (no livestock allowed) grazed and browsed there, and thus we observed less damage on the grass cover throughout all the campaigns as compared to the grazing land (which had large patches of bare soil due to overgrazing) and bushland. This provides a good explanation for the difference in mean CO₂ emissions between these three LUTs, as vegetation is known to affect soil C concentration and root and microbial respiration that directly contribute to soil CO₂ emissions (Fanin et al., 2011; Rey et al., 2011).

With the lowest CO₂ emissions being measured in the cropland, we attribute this observation to the continued tillage and removal of crops and crop residues during land preparation, weeding and harvesting, which affects both root respiration and soil C content (Raich et al., 2000; Nandwa, 2001). In East Africa and especially in smallholder farming systems, most of the crop residues are used as livestock feed and fuel. In addition, manure inputs in cropland are very low (about 20 kg per month on a 1.5 ha farm) and thus no measurable difference in CO₂ emissions was detected before and after manure input, and with the other LUTs. Several other studies observed the same scenario from low

Commented [WS25]: Ln 548. The argument with clays is a bit far-fetched anyway.
AC: We removed this argument in the revised manuscript.

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760 manure input in maize and sorghum plots (Rosenstock et al., 2016; Mapanda et al., 2011, and Pelster et al. 2017).

Commented [WS27]: Rephrased

765 On average, CO₂ emissions were higher during the wet season than during the dry season. At the start of both rainy seasons (SW, LW), CO₂ emissions increased significantly in all LUTs. Emissions from the conservation land and grazing land are comparable to those in Brümmer et al. (2008), who observed CO₂ emissions ranging between 100 and 250 mg CO₂-C m⁻² h⁻¹ in a natural savanna in Burkina Faso. Several other studies from similar ecosystem have also documented comparable changes in CO₂ emissions with the onset of the rainy seasons (Castaldi et al., 2006; Livesley et al., 2011; Pinto et al., 2002). In the cropland, results in the wet season are similar to those measured by Rosenstock et al. (2016), ranging from 50 to > 200 mg m⁻² h⁻¹. We attributed the increase in CO₂ emissions in the wet season to the response of soil microbes and vegetation to soil moisture (Livesley et al., 2011; Otieno et al., 2010). Soil moisture connects microorganisms with soluble substrates (Moyano et al., 2013) and increases microbial activity (Davidson et al., 2006; 2009; Grover et al., 2012) and thereby soil CO₂ emissions.

775 Furthermore, an increase in soil CO₂ emissions during the wet season can also be a result of increased root respiration due to more active plant and root growth (Macdonald et al., 2006). Grasses sprout more rapidly than trees and shrubs with the first rains (Merbold et al., 2009). This provides a possible explanation for the higher CO₂ emissions in the grassy conservation land, grazing land, and bushland compared to cropland during the rainy season. However, grazing land recorded higher CO₂ emissions than bushland (only the farmer's livestock grazed here). The main difference between these two sites – apart from grazing intensity – was that bushland had more trees (*Acacia spp.*) and shrubs (*Commiphora spp.*) and less herbaceous undergrowth than the grazing land, thus providing shade that might have interfered with growth and regrowth of plants below the canopy. Therefore, grass root production in the open conservation land and grazing land was likely higher than in the bushland (Janssens et al., 2001), although we cannot confirm this because root biomass was not determined in this study. In cropland, all grasses and weeds were cleared during regular weeding and therefore did not play a role in root respiration.

790 To our surprise, the highest mean seasonal CO₂ emissions in conservation land, grazing land, and cropland were observed at the end rather than at the peak of the wet season. During this time, both soil moisture and soil temperature had dropped in all LUTs. However, our data was only recorded up to a depth of 5 cm, but roots of perennial grasses, shrubs and trees can tap moisture from greater soil depths (Carbone et al., 2011). According to Carbone et al. (2011), while microbial activity is highest

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AC: deleted and rephrased

and most variable in the upper soil layers, which are first to wet-up and dry-down, roots can access water reserves in deeper soil layers that take longer to be exhausted, and therefore remain active at the end of the wet and into the dry season.

795 **4.2. Soil N₂O emissions**

800 Our results showed very low N₂O emissions from all LUTs, which we attributed to low soil N content observed in all the sites (see Table 1). Savanna ecosystems are characterized by very tight N cycling, which transcends to low N availability (Pinto et al., 2002 and Grover et al., 2012), and most of this N is rapidly taken up by vegetation, leaving very little for denitrification (Castaldi et al., 2006; Mapanda et al., 2011). The N₂O flux results observed from conservation land, grazing land and bushland are consistent with those observed in a Brazilian savanna by Wilcke et al. (2005), and other studies from similar ecosystems reported comparable N₂O flux magnitudes (Scholes et al., 1997; Castaldi et al., 2016; Mapanda et al., 2010). The higher N₂O emissions observed in June and July from our cropland site after the maize and bean harvests likely occurred due to the disturbance and following absence of live plants, which led to higher soil N availability because of less N uptake by plants and increased root decomposition.

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810 In contrast to the patterns observed for CO₂ emissions, we did not detect any seasonal variations in N₂O emissions. The only exception to the otherwise very low N₂O emissions was after the onset of the rainy season, when N₂O emissions slightly increased at all sites. Such patterns have previously been shown by several similar studies (Scholes et al., 1997; Pinto et al., 2002; Castaldi et al., 2006; Livesley et al., 2011). The increase in N₂O flux at the onset of the rains has been attributed to an increase in microbial activity and therefore faster decomposition of litter and plant residue facilitated by an increase in soil moisture, thus increasing N availability (Rees et al., 2006). Furthermore, according to Davidson et al., (2000) and Butterbach-Bahl et al., (2013), soil moisture affects soil gas diffusion, oxygen (O₂) availability, and the movement of substrate necessary for microbial growth and metabolism.

820 Negative N₂O emissions were detected during the dry season. Such observations could result from the low N contents observed at all sites coupled with low soil moisture in the dry season, which facilitates diffusion of atmospheric N₂O into the soil. Soil denitrifiers may, therefore, use N₂O as an N substrate in the absence of NO₂⁻ and NO₃⁻ (Rosenkranz et al., 2006). Negative N₂O emissions have also been reported in other tropical savanna soils under similarly dry conditions (Castaldi et al., 2006; Livesley et al., 2011).

825 Manure application in the cropland was very low (< 12 kg of N in 1.5ha for the crop-growing season), and thus N₂O emissions from cropland were low and not different from the other LUTs, which was
in contrast to what we had hypothesized. Due to low soil N levels in the cropland, the low amount of
manure added was not sufficient to stimulate N₂O emissions, likely because soil N availability was
still limiting for plant and microbial growth (Castaldi et al., 2006). Traditional farming systems in
830 smallholder farms in Africa involve repeated cropping with no or very low N inputs that leads to soil
N mining over time (Chianu et al., 2012). In line with this, in our cropland site maize and beans are
grown during every wet season with no fallow in between years. In addition, the farmer did not use
any chemical fertilizer to increase soil N, and the N input from biological N fixation into the soil was
likely small because beans were harvested for consumption and bean plant residues were used as
livestock feed and not incorporated into the soil. Therefore, the small quantities of manure applied
and legume N fixation may have likely been insufficient to compensate for N loss through leaching
835 and crop harvests. According to the Taita Development plan, this is a common scenario in the county,
which translates to very low crop yields in this region (CIDP, 2014). Another possible explanation
for not detecting the influence of manure on N₂O emissions could be the fact that we did not manage
to sample immediately after manure application and therefore might have missed the instant impact
of manure application on N₂O emissions. However, similar studies by Pelster et al. (2017) and
840 Rosenstock et al. (2016) also did not see any influence of manure application on soil N₂O emissions
and reported N₂O emission values that were generally < 10 µg N₂O-N m⁻² h⁻¹). Equally, the deposition
of dung and urine by animals in the grazing land and bushland did not have any measurable influence
on soil N₂O emissions.

4.3. Soil CH₄ emissions

845 Methane emissions did not vary between the land-use types or with seasons. Most values were below
the LOD at all the sites. Soil water content in our study is clearly the limiting factor for
methanogenesis, which needs anoxic conditions for a certain period until methanogenic archaea are
established (Serrano-silva et al. 2014). Furthermore, soil compaction by animal trampling may have
limited CH₄ diffusion into the soil thus limiting CH₄ consumption by oxidation (Ball et al. 1997). In
850 cropland, continuous tillage interferes with soil structure thus affecting the microenvironment that
favours methanotrophic (Jacinthe et al. 2014). Additionally, low soil C as observed in all the sites
generally leads to low abundance of soil microorganisms and consequently also methane oxidisers
(Serrano-silva et al. 2014). Nevertheless, soils around lakes, waterholes and rivers can be CH₄ sources
in semi-arid savanna ecosystems, but those were not investigated during this study.

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AC: Corrected

4.2.4.4. Effects of soil moisture, soil temperature, and vegetation indices on GHG emissions

Soil moisture and soil temperature are known to be important drivers of soil CO₂ production, and they may change across seasons. Seasonal variations in T were very minor and thus changes in WC were considered to be the main driver of CO₂ emissions in our study, as previously highlighted by Grover et al. (2012). Brümmer et al. (2009) and Livesley et al. (2011) also found that WC controlled CO₂ fluxes from savanna soils, rather than T. Soil moisture determines the rate of soil respiration, including heterotrophic and autotrophic processes that are highly moisture reliant (Ardö et al., 2008; Grover et al., 2012).

Soil temperature and soil moisture are also vital for nitrification and denitrification, as they control the activities of soil microbes and O₂. However, we did not observe a significant relationship between N₂O emissions with both WC and T. Only in the cropland area did we observe a positive correlation with T ($p < 0.05$). As much as previous results have shown a positive relationship between T and N₂O emissions (Castaldi et al., 2010), we did not observe such a relationship in our study. Other studies carried out in savannas by Scholes et al., (1997) and Brümmer et al. (2008) were also unable to link N₂O emissions to variations in T. In fact, N₂O emissions were very low during both the wet and dry seasons, as in (Castaldi et al., 2004), Pelster et al. (2017), and Rosenstock et al. (2016). The most likely clarification for the lack of seasonality would be the low N levels observed at all the sites (Grover et al., 2012).

The vegetation and its status are other important drivers of GHG emissions from soils. NDVI measures the status of vegetation (with a value range of -1 to 1). High NDVI values correspond to high vegetation cover, while low NDVI corresponds to less or no vegetation (Gamon et al., 1995; Butt et al., 2011). Therefore, the drop in NDVI values at the end of the rainy season was a result of the reduction in vegetation cover in the land use types. In the cropland area, this coincided with the harvesting of beans and the drying of the maize plant occurring in June and July. The conservation land showed the highest mean NDVI, mainly due to the dense grassy vegetation, especially during the rainy season. Lowest NDVI values were observed in the grazing land, which was expected because the land is mostly bare caused by overgrazing. Results from the linear regression showed a positive correlation with soil CO₂ emissions with NDVI ($p < 0.05$), explaining between 35 % and 82 % of the variation in soil CO₂ emissions at the four sites. We therefore observed high CO₂ emissions when NDVI was high, which is an indication of more vegetation cover. Several studies have shown that vegetation can affect soil respiration by intercepting radiation, modifying the soil moisture

890 regime, adding litter onto the soil surface, and affecting both plant and root respiration (Myneni et al. 1995; Almagro et al., 2013;). Thus, the inclusion of both NDVI and WC is essential for predicting soil CO₂ emission from savanna soils. Our results confirm the importance of vegetation variability in addition to WC as a key driver of productivity in savannas and is consistent with other studies (Reichstein et al., 2003; Anderson et al., 2008; Lees et al., 2018). Concurrently, the same relationship between NDVI and N₂O emissions could not be proven. Our conclusion was that the low N level played a major impact on overall N₂O emissions in this study.

895 As is common for sub-tropical regions, seasonal variation in soil temperature was small in the study region and therefore soil temperature did not play a big role in modifying soil GHG emissions. Instead, changes in soil moisture were considered to be the main driver of CO₂ emissions in our study, as has previously been highlighted also by other studies (Grover et al. (2012), Brümmer et al. (2009) and Livesley et al. (2011)). However, we did not observe any significant relationship between N₂O emissions with either soil moisture or temperature apart from in the cropland, where we found a positive correlation between N₂O and soil temperature (p < 0.05). As much as previous results have sometimes shown a positive relationship between temperature and N₂O emissions (Castaldi et al., 2010), our results are in line with others (Scholes et al., (1997), Brümmer et al. (2008) who were also unable to link soil N₂O emissions to variations in soil temperature. In fact, N₂O emissions were very low during both the wet and dry seasons, which is similar to the findings of Castaldi et al., (2004). The most likely explanation for the lack of seasonality effects on N₂O emissions would be the low soil N levels observed at all the sites, which was probably the most limiting factor for N₂O emissions and thus overruled all other potential controlling factors (Grover et al., 2012).

910 The vegetation cover as depicted by NDVI represents the status of the vegetation (value range from -1 to 1). High NDVI values correspond to high vegetation cover, while low NDVI correspond to little or no vegetation (Gamon et al., 1995; Butt et al., 2011). Therefore, the increase in NDVI that we observed at the onset of the rainy season indicates sprouting and regrowth of vegetation at that time, while the drop in NDVI values at the end of the rainy season indicates reduction in vegetation cover due to plant senescence and grazing. In the cropland area, low NDVI coincided with the harvesting of beans and the drying of the maize plants in June and July. Highest mean NDVI values were observed in the conservation land, mainly due to the dense grassy vegetation, while the lowest NDVI values were found in the grazing land, which we had expected because this area has large spots without vegetation due to overgrazing. Results from linear regression analysis showed a strong positive correlation of soil CO₂ emissions with NDVI (p < 0.05), explaining between 35 % and 82 %

of the variation in soil CO₂ emissions at the four sites. This means that CO₂ emissions were highest when NDVI (i.e. vegetation cover) was high. Thus, the inclusion of both NDVI and soil moisture measurements is essential for reliably predicting soil CO₂ emission from savanna soils, which is consistent with other studies (Reichstein et al., 2003; Anderson et al., 2008; Lees et al., 2018). Concurrently, the same relationship between NDVI and N₂O emissions could not be proven in our study.

5. Conclusion

The land use management system plays an important role in magnitude and temporal and spatial variability of soil GHG emissions due to changes in vegetation, soil, hydrology, and nutrient availability. However, the effects on GHG emissions change remain uncertain due to limited data being available in most developing countries have large uncertainties due to a lack of data, especially in dry areas and ecosystems facing diverse land-use change. In our study, we quantified soil GHG emissions from four dominant LUTs in the dry lowland/lowlands of southern Kenya, namely bushland, conservation land, cropland, and grazing land. Our results showed significant variation between seasons and the respective land use types-LUTs. CO₂ emissions, in particular, were higher during the wet season than the dry season. The lowest seasonal mean CO₂ emissions were observed in the SD, when soil moisture was very low while soil temperature was very high, compared to the dry season. Most of the variation in CO₂ emissions was could be explained by soil moisture and NDVI and soil moisture, highlighting the fact importance of including proxies for vegetation cover in soil GHG emission studies in savannas. N₂O emissions and CH₄ emissions were of minor importance at all sites. However, we may acknowledge that we might have missed some episodes of the elevated soil N₂O emissions, as these are often episodic, i.e. and of short duration, for examples after fertilization or precipitation events. These can easily be missed unless a continuous measurement and monitoring programme is in place. Following these results, there is still need for more continuous studies to cover spatial and temporal variations in soil emissions from diverse land use types across seasonal and management gradients. Furthermore, continuous measurements allow the detection changes in GHG emission patterns following intensification variation in soil GHG emissions, as well as the inclusion of other LUTs than the ones examined in this study (e.g. wetlands). Nevertheless, we believe that our results are useful for deriving important to reduce uncertainties in GHG emission baselines and to identify reliable and meaningful climate change mitigation interventions by informing the relevant policies.

6. Data availability

The data associated with the manuscript can be obtained from ~~the corresponding author upon request~~ https://figshare.com/articles/Final_data_for_Soil_Greenhouse_Gas_Emissions_under_Different_Land-Use_Types_in_Savanna_Ecosystems_of_Kenya_/11673579

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REVIEWERS COMMENTS

Soil Greenhouse Gas Emissions under Different Land-Use Types in Savannah Ecosystems of Kenya

Authors Response:

1320 The manuscript has been reviewed by 2 reviewers and an additional comment was made in the open discussion. Before, responding to each reviewer and comment individually, we would like to thank for the constructive comments and informative feedback.

The document is structured as follows: each of the reviewer's comment (indicated by RC) is first repeated followed by our response (indicated as AC and in italic). Where relevant we either include a rephrased sentence already or explain on how we intent to implement suggested changes.

1325

Anonymous Referee #1

Overall comments:

1330 **RC:** The manuscript describes a study in four typical land use types in Kenya, Africa. Soil fluxes of CO₂, N₂O, and CH₄ were measured manually eight times over the course of a year. The main strength of the manuscript is that it produces flux estimates of these greenhouse gasses in under-represented ecosystems. Correlations with driving factors of moisture, soil C content, and vegetation activity (NDVI) were explored. The main weakness of the manuscript is the sampling campaign and methods are very limited and coarse, and thus interpretation of the driving factors of the fluxes are much more speculative than could be with greater initial and supporting data. My suggestion would be to reduce the length of the manuscript to focus just on the data collected and acknowledge the weaknesses in the data set. A shorter, more concise, manuscript would be much more effective to get the data out there.

1335

AC: *Once again, we thank the reviewer for pointing out both, the strengths and the limitations in the originally submitted manuscript. Following the concern made on the length of the paper, we will reduce the revised manuscript, focus only on the important issues, and not compromise the quality of the manuscript.*

1340

Abstract

RC: Ln 25 – the N₂O flux was more than double the cropland than bushland, why do you say is was not different between the four sites?

1345 **AC:** *We thank the reviewer for pointing this out. Actually the difference in annual mean N₂O emissions were significantly higher in the cropland ($2.7 \pm 0.6 \mu\text{g m}^{-2} \text{h}^{-1}$) than in the conservation land ($1.6 \pm 0.4 \mu\text{g m}^{-2} \text{h}^{-1}$), grazing land ($1.5 \pm 0.4 \mu\text{g m}^{-2} \text{h}^{-1}$), and bushland ($1.2 \pm 0.4 \mu\text{g m}^{-2} \text{h}^{-1}$) (Kruskal-Willis rank test). In contrast, no difference was observed between the other three sites.*

RC: Ln 31 – Over the course of the measurement period or between sites, CO₂ was correlated with soil moisture?

1350 **AC:** *In all our study sites and across the seasons, soil CO₂ emissions were positively correlated with soil moisture.*

RC: Ln 30-40 - The abstract does not have a clear message. Soil C is important, but soil moisture is driving fluxes, but NDVI is correlated. What is the take home point?

AC: *We will rephrase the abstract as follows in order to provide a clearer message.*

1355 *“Based on our results, soil C and soil moisture are key drivers of soil GHG emissions in all land use types. In addition, vegetation cover explained the seasonal variation of soil CO₂ emissions as depicted by the strong positive correlation of between NDVI and CO₂ emissions. We conclude that with more green (active) vegetation cover higher CO₂ emissions occur due to enhanced root respiration compared to drier periods in the year.*

1360 **Introduction**

RC: The introductory paragraph never says what produces and consumes GHGs from the soil?

1365 **AC:** *We will include the following information into the introduction of the revised manuscript. Each of the GHGs observed is either consumed or produced via biogeochemical processes. For instance, methane is produced by methanogenesis process under anaerobic conditions and consumed by methanotrophic microorganisms under aerobic conditions (Serrano-Silva et al., 2014). CO₂ is in our case (dark chambers) produced in the soil during decomposition of organic matter and root respiration which produces CO₂. Similarly, CO₂ is produced via plant respiration in the case plants were present in the chamber (Oertel et al. 2016). N₂O on the other hand is produced as an intermediate product of denitrification and nitrification processes amongst others (Butterbach-Bahl et al., 2013).*

1370 **RC:** ASALs is an acronym that could be avoided by using drylands, or arid ecosystems. Overall there are many acronyms used that could be avoided.

1375 **AC:** *We thank the reviewer for this suggestion. Overall, we will have a look at the use of abbreviations. In the case of ASAL, we will use drylands in the revised manuscript. A similar point was made by reviewer 2.*

Methods

RC: Ln 187 – ssp

AC: *Corrected to spp*

1380 **RC:** Ln 240 – This is a large assumption. Does the sampling really represent the average flux of the day for your ecosystem? At least one of those references is for a temperate forest where they did measure the 24-hour cycle, which likely has a very different cycle than these ecosystems due to differences and vegetation type and environmental variables.

1385 **AC:** *We agree that this is partially an assumption, and yet one has to compromise, as a more frequent sampling was impossible. Overall, the theory of sampling in the respective morning hours is sound (Parkin and Venterea, 2010), as temperatures increase during the day and thus microbial processes similarly are enhanced. Of course, the fact that moisture and other factors are essential drivers following our results this may be misleading and yet again the drivers suggested are for the whole dataset, i.e. across seasons and not necessarily aiming at explaining diurnal variations. We have done diurnal GHG measurements in other regions in Kenya with a portable laser absorption spectrometer where we found clear temperature dependencies during the course of the day (data not shown here, as this is part of another study, Butterbach-Bahl in prep.).*

1390

RC: Ln 252 – The pooling method reduces the sample # to 3 for each LUT time period instead of 9?

AC: *This is correct, the three gas flux estimates that we have at the end of the day for each LUC type is thanks to the pooling method derived from 9 distinct locations (chambers). With this approach, one can still account for spatial heterogeneity while reducing the overall number of GHG samples to be analysed in the laboratory (Arias-Navarro et al., 2013).*

1395

RC: Ln 290 – was temperature measured in the chamber?

AC: *We recorded the air temperature in the headspace of each chamber at the same interval as gas pooling (recorded at T1, T2, T3 and T4). This temperature is then used to correct the gas flux during flux calculation.*

1400

Results

RC: Ln 385 – What are the errors on the fluxes? They are so small for soil CO₂ fluxes. Report error and sample size.

AC: *The sample size per season is seven daily average values derived from three flux values per day from each land use type. The error bar presented here is the standard error of the three flux values per day.*

1405

RC: Figures 4 and 5 are good. Keeping the color scheme, the same would be helpful.

AC: *The two figures provide different information and having them in the same color scheme might be misleading. Figure 4 shows the difference in emissions between the land uses types while figure 5 gives the differences in emission between the wet and dry season.*

1410

RC: Figure 6 – put in same color scheme.

AC: This is a relevant point and we tried to have them in the same color scheme of figure 4. However, this then made it very difficult to differentiate between the LUC types.

RC: Figure 7 – this is at such a large scale, I don't find it very informative. Fig 6 shows the data used.

1415 **AC:** We agree and we will move Figure 7 into the appendix and keep Figure 6 in the revised manuscript.

Discussion

RC: Different terms are being used, soil respiration; soil CO₂ emissions, ecosystem soil respiration (?) Make this consistent.

1420 **AC:** We thank the reviewer for pointing this out. In the revised manuscript, we will only use the term "soil CO₂ emission".

RC: There is quite a bit of speculation in the discussion. It would be better shortened and more focused on the data collected, not the data lacking that could explain the patterns. This is true for CO₂ and N₂O sections.

1425 **AC:** Following this suggestion and a similar point being raised by Reviewer 2, we will make the discussion shorter and more concise – some examples are:

Example 1

Original - CO₂ emissions from the soil differed significantly between the land-use types. Higher CO₂ emissions were particularly observed in the conservation land followed by grazing land and bushland, while lowest levels were recorded from the cropland. We observed the same trend with SOC, and thus we attributed the difference in CO₂ emissions between the land uses to SOC as observed (see Table 1). This is in line with a similar study by La Scala et al. (2000), which recognized SOC as a key driver of CO₂ emissions from the soil, as it is the primary source of energy for soil microorganisms (Lal, 2009).

1430

Revised - Highest mean CO₂ fluxes were observed from the conservation land followed by grazing land and bushland, and the lowest from cropland. Soil C content also showed the same trend (conservation land > grazing land > bushland > cropland), which is the primary source of energy for soil microorganisms (Lal 2009) and thus affecting CO₂ emissions. Therefore, we attributed this variance to the difference in land use and management activities playing a major role in modifying both biotic and abiotic factors (Pinto et al., 2002).

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Example 2

Original - Soil CO₂ emissions and mean soil C content were lowest in cropland. Root respiration in cropland depends on periods of live roots in annual crop fields and on the biomass of roots during the initial growing season (Raich et al., 2000). Therefore, the continued removal of crop residues

1445 during harvesting and frequent tillage affected both root respiration and SOC. As much as crop residues contribute to carbon stocks through their mineralization (Nandwa, 2001), most maize residues were used as livestock feed and sometimes as fuel, while bean residues were removed completely during the harvest and burned.

1450 **Revised** - With the lowest CO₂ emissions being reported from the cropland, we attribute this observation to the continued tillage and removal of crop and crop residues during land preparation, weeding and harvesting, affecting both root respiration and soil C content (Raich et al., 2000; Nandwa, 2001). In East Africa and in smallholder farming systems, most of the crop residues are used as livestock feed and fuel.

Example 3

1455 **Original** - Manure inputs provide easily degradable substrates of C and N leading to enhanced soil CO₂ emissions (Janssens et al., 2001; Davidson and Janssens, 2006). However, manure input in cropland was very low (approximately 20 kg in a 1.5 ha farm per month) and thus no measurable effects in CO₂ emissions were detected. This was opposite to our hypothesis. Another reason could be that soil fertility was too low to have a detectable influence on CO₂ emissions (Pelster et al., 2017).
1460 Our results are in the same magnitude with those of Rosenstock et al. (2016); Farai Mapanda et al. (2011), and Pelster et al. (2017), who also did not detect any change in CO₂ emissions after manure application and attributed this to the low input of manure to the maize and sorghum plots.

1465 **Revised** - Manure inputs in cropland were very low (about 20 kg per month on a 1.5 ha farm) and thus no measurable difference on CO₂ emissions were detected. Several other studies observed the same scenario from low manure input in maize and sorghum plots (Rosenstock et al., 2016; Farai Mapanda et al., 2011, and Pelster et al. 2017).

Example 4

1470 **Original** - Soil N₂O emissions can vary highly over time, as regulated by factors such as soil moisture, temperature, aeration, ammonium and nitrate concentrations, pH, and mineralizable carbon (Butterbach-Bahl et al., 2013). However, we did not document any significant difference in N₂O emissions between the four land uses. At all the sites, N₂O emissions were very low, and this we attributed to the observed low soil N content (see Table 1). According to Pinto et al. (2002) and Grover et al. (2012), savanna ecosystems have a very tight N cycling, which transcends to low N availability. Thus, available N is taken up by vegetation, leaving very little for denitrification
1475 (Castaldi et al., 2006; Mapanda et al., 2011). Our results are consistent with low N₂O emissions 610 observed in a Brazilian savanna (cerrado) by Wilcke et al. (2005), who also reported low N levels in their study area. Very low N₂O emissions due to poor nutrient availability have also been observed in other savanna landscapes (Scholes et al., 1997; Castaldi et al., 2016). Soil N₂O in cropland match those of Rosenstock et al. (2016), who also attributed the low soil N₂O emissions to poor nutrient
1480 availability in the soil

Revised - Our results showed low N₂O emissions from all the LUTs, which we attributed to the low soil N content observed (see Table 1). Savanna ecosystems are characterized by very tight N cycling,

1485 *which transcends to low N availability (Pinto et al., 2002 and Grover et al.,2012) and most of this N is taken up by vegetation, leaving very little for denitrification (Castaldi et al., 2006; Mapanda et al., 2011). The flux results observed from the conservation land, grazing land and bushland are consistent with those observed in a Brazilian savanna by Wilcke et al. (2005). Many other studies from similar ecosystem reported comparable N₂O flux magnitudes (Scholes et al., 1997; Castaldi et al., 2016; Mapanda et al., 2010). The higher N₂O emissions observed in June and July from our cropland site after the maize and bean harvests are likely occurring due to the absence of plants.*

1490 **RC:** Interesting CH₄ just gets one sentence because it is small: : : but this is important too!

AC: *For the completeness of the paper, methane emissions have to be mentioned. Even though Reviewer 2 suggested to remove everything related to CH₄, we decided to keep this information in the manuscript for 2 reasons: (1) completeness in terms of greenhouse gases, and (2) even if the contribution of CH₄ is low, this is an important results and similar measurements may not have to be repeated.*

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Soil Greenhouse Gas Emissions under Different Land-Use Types in Savannah Ecosystems of Kenya

1500 Authors Response:

Two reviewers have reviewed the manuscript and an additional comment was made in the open discussion. Before, responding to each reviewer and comment individually, we would like to thank for the constructive comments and informative feedback.

1505 The document is structured as follows: each of the reviewer's comment (indicated by RC) is first repeated followed by our response (indicated as AC and in italic). Where relevant we either include a rephrased sentence already or explain on how we intent to implement suggested changes.

RC 2

D. Otieno (Referee)

denotieno@yahoo.com

1510 Received and published: 21 November 2019

1515 **RC:** This is interesting study conducted in semi-arid parts of Kenya, where similar data are quite scarce. The set-up is an area characterized by a series of activities. It is a surprised that there is some form of cultivation/farming in an area that looks more like Tsavo national Park. Nonetheless, the study provides valuable data that extend our knowledge of ecosystem gas fluxes in this part of the world.

AC: *We would like to thank Mr. Otieno for this review and his valuable comments.*

1520 **RC:** The study was conducted in a relatively poor soil. What the authors failed to mention, especially for the cropped and grazed sites was the slope of the field. I tend to imagine that erosion must be playing a critical role in mineralization processes in this place. It looks like the organic/humus, top soil layer is completely gone and what remains is mainly the mineral soils. Unfortunately, the paper is already too long and I will not recommend inclusion of more information on land use history, which would have been helpful in understanding/interpreting these results.

1525 **AC:** *The study area are located is the lowland of the Taita Taveta county, which is very flat. The cropland is at 1070 m a.s.l, bushland at 1076 m a.s.l, grazing land at 970 m a.s.l, and conservation land at 928 m a.s.l. The cropland is received very small quantities of manure and no chemical fertilizer inputs and thus no significant difference in soil C content with the other land use types. In the grazing area, overgrazing was evident as most of the soil was bare especially in dry season. This contributes to soil erosion and compaction of the land by wind and rain and even the livestock while grazing. We will add this information briefly in the revised manuscript.*

1530 **RC:** It's very surprising that temperature and soil moisture had no influence on soil CO₂ fluxes. Could it be the method of data collection, with significant data collection gaps that led to this?

AC: Soil CO₂ emissions were positively correlated to soil moisture. However, variation in soil temperature for the time of measurements during the day in both dry and the wet season were minor, and thus we found no statistically significant effect of soil temperature upon CO₂ emissions for the dataset. Other studies by Brümmer et al. (2009) and Livesley et al. (2011) also found that soil moisture controlled CO₂ emissions from savanna soils, rather than soil temperature. However, if we have had the opportunity to measure more frequently – i.e. following a diurnal course – we are confident that an effect of temperature exists. For instance, we found such diurnal course in GHG emissions in a similar ecosystem in Kenya. This was part of another project and is consequently not shown here.

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RC: For future, the authors need to consider higher frequencies of data collection. In such arid ecosystems, evaporation is quite high and it is likely that critical information is lost by not collecting data more regularly.

AC: For this study, the sampling frequency was based on seasonal variation, thus the campaigns were targeting the wet, transition and dry season and when moisture and/or management practices are likely to impact GHG emissions. Certainly, we would have preferred more frequent measurements, though given the research question asked and the available resource for this project, we had to make a compromise. However, there is a follow-up study in other Land Use Types with measurements that are more frequent.

1545

RC: CH₄ seems to contribute little to this paper, why not exclude it completely? I do not see the two lines of discussion on CH₄ are of major benefit to the readers. The paper is already too long and probably removing all the descriptions on CH₄ could reduce the number of pages.

1550

AC: We agree that the importance of methane emissions is negligible when compared to the other gases. However, our aim was to look at all three GHGs in this study and due to the lack of available GHG emissions data from such land cover types in this region of the world we still think its beneficial to report these here. Certainly, in order to not further lengthen the paper, we decided to keep this information as short as possible.

1555

RC: The word “Soil Organic Carbon SOC” is introduced in the introductory part of the Ms. In the methods, there is total soil carbon and in the results, I met Soil Carbon. In the discussions, SOC becomes the main discussions line. The authors need to be consistent in the use of these terms, otherwise the readers get confused.

1560

AC: We thank the reviewer for pointing this out and we will harmonize in the revised manuscript accordingly.

RC: Ln 65. Not all savannah belongs to the ASALs. The humid savannas are relatively wet, with green vegetation almost throughout the year. It is therefore not right to make such a sweeping statement.

1565

AC: *Noted with thanks and we adjust the revised manuscript accordingly and use drylands instead of ASALs.*

Specific comments

1570 **RC:** Ln 67. Note that shrubs are woody vegetation

AC: *Noted with thanks.*

RC: Ln 88. Revise the sentence. Overstocking leads to grazing pressure. The way the sentence is written is redundant.

AC: *Done*

1575 **RC:** Ln 96-7. ---Croplands are still being cleared from natural vegetation---re-write the sentence, it's not making the intended meaning.

AC: *Done. Revised --Natural vegetation is being cleared to make way for the expansion of cropland*

RC: Ln 104, what's "cropland farming"?

AC: *Here we refer to cropping agriculture in the savanna.*

1580 **RC:** Ln 153. The authors need to be clear on the physiognomic characterization of the vegetation they are studying. Here you have woodlands, bushlands and on line 155 you have wood bushlands, which is which?

AC: *We harmonize this in the revised manuscript to bushland as found in Tsavo East and West national park.*

1585 **RC:** Ln 156 are Lions also grazers?

AC: *Lion are not grazers. On this line, we were mentioning the fauna that the ecosystem supports in general to highlight the importance and the functions of the park.*

RC: Ln 160 –other important land use(s)

AC: *Corrected*

1590 **RC:** Ln 173. Is the farm rain-fed or not? Are there other sources of moisture input apart from rain?

AC: *The farm is totally rain fed.*

RC: Ln 237, how deep was the collar inserted into the soil?

1595 **AC:** *The collars we inserted between 5cm to more than 8cm into the soil. We ensured the collars were inserted so the extend above the surface did not hold water during the rainy season and the collars were less likely to be trampled on and broken by large animals*

Result

RC: Label Fig. 3 as a and b

AC: *Done*

1600 **RC:** Ln 377, Sand proportion was lower than what? In comparative sentences, learn also to use “lowest” or “highest” see ln 417.

AC: *Sand proportion was lowest in the grazing land (64.3±0.4 %) than in the other study three sites. We will adjust the phrasing in the revised manuscript.*

RC: Ln. 456 present data/results according to the chronology of the figures and avoid this back and forth.

1605 **AC:** *Done*

RC: Ln 481. Delete (in) before during.

AC: *Done*

Discussion

1610 **RC:** SOC is only mentioned in the introduction but not in the methodology or results, yet it becomes very prominent in the discussions. Be consistent in the use of terms.

AC: *Noted with thanks.*

RC: Ln 525 is not correct. You cannot attribute the differences only to vegetation. It is definite that land use itself leads to the differences in soil C. Although this is argued correctly in the later sections, this section should be revised.

1615 **AC:** *Corrected*

Ln 548. The argument with clays is a bit far-fetched anyway.

AC: *We removed this argument in the revised manuscript.*

1620 **RC:** Ln 592. ---temperature was measured “down” to 5 cm. I would imagine that 5 cm depth is almost at the surface. What was the deciding factor for installing temp/moisture sensor at this depth? This depth, being close to the surface is associated with very strong temperature fluctuations. It may be one of the reasons why the authors found no temperature correlation with CO₂ efflux. Most grass

roots, cereals included, have roots located within 10 cm, and may extend down to 30 cm. the woody vegetation in such dry places have their roots even deeper. Trying to establish relations with variables measured at 5 cm may not yield positive results.

1625 **AC:** *According to a study by Pavelka et al. (2007), daily dynamic of soil CO₂ fluxes are affected by soil temperature near the soil surface and hence for correlation between soil CO₂ emissions and soil temperature, the measurement of soil temperature at the soil surface, is highly recommended to avoid the inaccuracies. Coupled with this, the ProCheck handheld GS3 sensor (Decagon Devices Inc) for soil moisture and temperature that we were using could only measure up to 5cm. Because, we were*
1630 *taking measurement within or close to the chamber collars, we did not want to cause any soil disturbance. This is also recommended by the GRACENet protocol we were using as our reference protocol.*

RC: Ln 593 check the sentence. How does root respiration tap moisture?

AC: *Noted. Here we mean roots can still tap moisture from deeper profile and thus root respiration can continue even after the surface moisture has dried up.*
1635

RC: Ln 640. Consider soil erosion and volatilization also.

AC: *This is an important point and we thank the reviewer for pointing this out.*

RC: Ln 651. Use “dung” instead of faeces.

AC: *Corrected*

1640 **RC:** Ln 665 what's T? From nowhere, you introduce T.

AC: *T stands for soil temperature.*

1645 **Soil Greenhouse Gas Emissions under Different Land-Use Types in Savannah Ecosystems of Kenya**

Authors Response:

Two reviewers have reviewed the manuscript and an additional comment was made in the open discussion. Before, responding to each reviewer and comment individually, we would like to thank for the constructive comments and informative feedback.

1650 The document is structured as follows: each of the short comment (indicated by SC) is first repeated followed by our response (indicated as AC and in italic). Where relevant we either include a rephrased sentence already or explain on how we intent to implement suggested changes.

Short comments

Cornelius Oertel

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I want so suggest some changes for that publication:

1660 **SC:** N₂O fluxes are sometimes given in $\mu\text{g N}_2\text{O-N m}^{-2} \text{h}^{-1}$ (e.g. l. 295 or 420) and in some graphics they are given in N₂O ($\mu\text{g m}^{-2} \text{h}^{-1}$) (e.g. figure 4 or figure 5). This is the same for CO₂. Units should be used consistent, so that you can compare the fluxes.

AC: *We thank Mr. Oertel for pointing this out and have made the necessary correction. CO₂ ($\text{mg m}^{-2} \text{h}^{-1}$) N₂O ($\mu\text{g m}^{-2} \text{h}^{-1}$) and CH₄ ($\text{mg m}^{-2} \text{h}^{-1}$).*

SC: l. 171: Height a.s.l. is only given for the cropland site and should be given as well for the other sites.

1665 **AC:** *This will be done in the revised manuscript. Cropland is at 1070 m a.s.l, bushland at 1076 m a.s.l, grazing land at 970 m a.s.l, and conservation land at 928 m a.s.l.*

1670