Reviews and syntheses: Insights into deep-sea food webs and global environmental gradients revealed by stable isotope ( $\delta^{15}N$ ,  $\delta^{13}C$ ) and fatty acid trophic biomarkers

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Abstract. Biochemical markers developed initially for food-web studies of terrestrial and shallow-1 2 water environments have only recently been applied to deep-sea ecosystems (i.e. in the early 3 2000s). For the first time since their implementation, this review took a close look at the existing 4 literature in the field of deep-sea trophic ecology to synthesize current knowledge. Furthermore, it 5 provided an opportunity for a preliminary analysis of global geographic (i.e. latitudinal, along a depth gradient) trends in the isotopic ( $\delta^{15}N$ ,  $\delta^{13}C$ ) and fatty acid composition of deep-sea macro-6 7 and megafauna from heterotrophic systems. Results revealed significant relationships along the 8 latitudinal and bathymetric gradients. Deep-sea animals sampled at temperate and polar latitudes 9 displayed lower isotopic ratios and greater proportions of essential  $\omega$ 3 long-chain polyunsaturated fatty acids (LC-PUFA) than did tropical counterparts. Furthermore,  $\delta^{15}N$  and  $\delta^{13}C$  ratios as well as 10 11 proportions of arachidonic acid increased with increasing depth. Since similar latitudinal trends in the isotopic and fatty acid composition were found in surface water phytoplankton and particulate 12 13 organic matter, these results highlight the link across latitudes between surface primary production and deep-water communities. Because global climate change may affect quantity and quality (e.g. 14 15 levels of essential w3 PUFA) of surface primary productivity, and by extension those of its downward flux, the dietary intake of deep-sea organisms may likely be altered. In addition, because 16 17 essential  $\omega$ 3 PUFA play a major role in the response to temperature variations, climate change may 18 interfere with the ability of deep-sea species to cope with potential temperature shifts. Importantly, 19 methodological disparities were highlighted that prevented in-depth analyses, indicating that further 20 studies should be conducted using standardized methods in order to generate more reliable global predictions. 21

# 22 **1 Introduction**

## 23 **1.1 Historical background of biochemical biomarkers in deep-sea food-web studies**

While the use of biochemical biomarkers in marine food-web studies has a long and successful 24 25 tradition in shallow-water ecosystems, starting from the 1970s with the use of stable isotopes (McConnaughey and McRoy, 1979) and lipids (Lee et al., 1971), their application in deep-water 26 environments is relatively new (e.g. lken et al., 2001; Polunin et al., 2001; Howell et al., 2003). 27 28 Undoubtedly, technological advances made over the past few decades have allowed the 29 exploration of ever deeper ecosystems with more refined techniques. Iken et al. (2001) were among the first to provide a comprehensive analysis of a deep-sea food web, which was sampled 30 at a depth of ~4840 m at the Porcupine Abyssal Plain (PAP, Northeast Atlantic), by using bulk 31 32 stable N and C isotope ratios ( $\delta^{15}$ N and  $\delta^{13}$ C respectively) as trophic markers. In the same year, Polunin et al. (2001) used the same approach to study the trophic relationships of a slope 33 megafaunal assemblage collected off the Balearic Islands (western Mediterranean). Since these 34 35 first two investigations, several others have been carried out across different oceanic regions and 36 climes, such as the Canadian Arctic (Iken et al., 2005), the Arabian Sea (Jeffreys et al., 2009), and the Sea of Japan (Kharlamenko et al., 2013). Furthermore, over the past decade, it has become 37 evident that the simultaneous use of different trophic markers (e.g.  $\delta^{15}N$ ,  $\delta^{13}C$ , and fatty acids, FA) 38 and techniques (e.g. bulk or compound specific isotope analysis, as well as FA, gut content and 39 morphometric analyses) provides a more complete picture of trophic structure and dynamics. 40 Indeed, while the first investigations relied on a single method (Iken et al., 2001; Polunin et al., 41 2001: Howell et al., 2003), the latest trend in deep-sea food-web studies favours an integrative 42 approach, which maximizes the efficiency of each technique, while increasing the resolution of the 43 investigation (e.g. Stowasser et al., 2009; Parzanini et al., 2017). 44

For the first time since the implementation of trophic markers in studies of deep-sea food
webs, this review synthesizes current knowledge in this growing field of research, mainly focusing

on heterotrophic ecosystems (i.e. relying on photosynthetic primary production). In addition, it
provides a preliminary overview of large-scale geographic trends from the analysis of isotopic and
FA data for macro- and megafauna, along with guidance for future investigations. In particular, the
present contribution i) briefly defines various trophic biomarkers and their respective advantages; ii)
describes deep-sea food webs, based on examples from the literature; iii) lists the sources of
variation among the different studies to highlight pitfalls and gaps; and iv) provides a preliminary
guantitative analysis across studies by using relevant datasets.

## 54 **1.2 Comparison of major trophic markers**

The analysis of gut contents was among the first techniques (together with in situ observation of 55 56 feeding behaviors) applied in trophic ecology and food-web studies in aquatic systems (Gartner et al., 1997; Michener and Kaufman, 2007). Subsequently, other methods were developed as 57 58 alternative or supplementary means of studying diet and feeding behaviors within the same ecosystems. Among them, the use of biochemical markers as trophic tracers rapidly grew in 59 60 popularity in food-web ecology, since it is relatively simple and should overcome many of the issues ascribed to gut content analysis (Michener and Kaufman, 2007). In this regard, Table 1 lists 61 strengths and drawbacks of gut content analysis and of the two most popular biochemical 62 63 techniques, i.e. bulk stable isotope and FA analyses. For instance, bulk stable isotope and FA 64 analyses may, theoretically, be performed on any species, regardless of feeding mode and food 65 sources, whereas gut content analysis can only be applied to those organisms characterized by a sufficiently large and full stomach. Except in cases where individuals are too small and have to be 66 analyzed whole, biochemical analyses are typically conducted on target tissues (e.g. muscle) that 67 68 provide long-term dietary data and reduce intra-individual variability (Table 1). In addition, the use 69 of biochemical tracers requires shorter processing times than gut content analysis. Thanks to this integrative approach and faster output, the application of food-web tracers has been particularly 70 helpful in deep-sea studies, which are often plaqued by financial and logistical constraints. 71 72 Furthermore, due its relative ease of use, it has favoured the analysis of wider sets of taxa/feeding

guilds, primary producers included, rather than focusing on one or a few focal groups. However, the
interpretation of isotopic and FA data is complex, and both techniques require dedicated and
sophisticated instrumentation (e.g. gas chromatograph, mass spectrometer) and knowledge of
intrinsic sources of variations (see Sect. 1.4). Although each method needs a sufficient sample
size, only gut content analysis may provide direct and clear taxonomic evidence of the diet (Table
1). Therefore, as stated above, the latest trend in trophic ecology advocates a multifaceted
approach, on the understanding that each technique may offer unique and valuable data.

The principle behind the use of food-web tracers is that the biochemical signature of 80 consumers reflects that of their diet. Among them,  $\delta^{15}N$  and  $\delta^{13}C$  are the most popular. While the 81 82 former is used to study trophic positions and dietary sources, with an enrichment factor of 2-4‰ between a consumer and its food (Minagawa and Wada, 1984); the latter undergoes little 83 84 fractionation (<1‰) and, therefore, is used to distinguish primary food sources (McConnaughey 85 and McRoy, 1979). For further details, refer to Sulzman (2007) and Michener and Kaufman (2007) who have provided extensive reviews on the chemistry behind stable isotopes and their use as 86 87 food-web tracers, respectively. In addition, sterols, FA and amino acids, which are important constituents of lipids (for the former two) and proteins (for the latter), have successfully been used 88 to study trophic relationships and dietary sources in deep-water systems (Howell et al., 2003; 89 90 Drazen et al. 2008a, 2008b). Their use is based on the principle that certain FA and amino acids 91 are considered essential for animals, being required for optimal fitness. However, most species 92 cannot synthesize these essential compounds *de novo* and, therefore, they must gain them through their diet. Indeed, only primary producers and a few consumers possess the enzymatic apparatus 93 94 to synthesize essential FA and amino acids *de novo*. Conversely, a few taxa are unable to synthesize sterols de novo, which are critical for them; therefore, they have to acquire these 95 96 essential sterols through diet (Martin-Creuzburg and Von Elert, 2009). Because sterols, FA, and amino acids undergo little or no alteration when consumed, it is possible to detect dietary sources 97 within the consumers' tissues (Parrish et al., 2000). The isotopic signature of amino acids can also 98 be used to study trophic position through compound specific analysis ( $\delta^{15}N$ ), as some of these 99

acids show trophic enrichment (Bradley et al., 2015). Detailed information about FA analysis was
outside the scope of this study, and is provided by Parrish (2009) and Iverson (2009); whereas the
use of sterols as food-web tracers was outlined in Martin-Creuzburg and Von Elert (2009) and
Parrish et al. (2000). McClelland and Montoya (2002) and Larsen et al. (2009), conversely, discuss
the use of amino acids as trophic biomarkers.

# 105 **1.3 Understanding deep-sea food webs through biochemical markers**

106 As there is no photosynthetically-derived primary production in the deep sea, deep-water ecosystems are mostly heterotrophic (Gage, 2003), and may hence largely rely on particulate 107 108 organic matter (POM) that passively sinks from the surface waters as a primary source of nutrients 109 (Hudson et al., 2004). Nonetheless, food can also be actively transported down by those animals that carry out vertical diel migrations through the water column (Trueman et al., 2014); it can also 110 111 be provided by the occasional fall of large animal carcasses (Smith and Baco, 2003); and/or by lateral inputs, from inland and shelf areas towards abyssal offshore regions (Pfannkuche, 2005). 112 113 Although most of the deep-water ecosystems are heterotrophic, a few, such as hydrothermal vents and cold seeps, are fuelled by chemical energy (e.g. methane, hydrogen sulfide) and rely on 114 chemosynthetic microorganisms for the production of organic matter. Each of these primary food 115 116 sources has a specific isotopic composition and biochemical signature, resulting from a 117 combination of chemical and physical processes reflective of its origin. By knowing the composition of the food source(s) that fuel(s) a given food web, it is possible to re-construct its trophic structure 118 and dynamics. Conversely, by measuring the signatures of the food-web components, it is possible 119 120 to assess food sources on which they rely. For instance, lken et al. (2001) showed that 121 phytodetritus was the primary energy input of the deep-sea benthic community at PAP, and also defined two different trophic pathways: a pelagic and isotopically lighter one in which sinking POM 122 and small pelagic prey constituted the main food sources; and a benthic and more isotopically 123 enriched trophic pathway, fuelled by degraded sedimented POM. In fact, once POM settles on the 124 125 seafloor, it undergoes continuous degradation by microbes and is reworked through bioturbation

126 and feeding activities, thus leading to a more isotopically enriched material relative to the sinking one (Iken et al., 2001). Depending on the primary food source they relied on, benthic organisms at 127 PAP were thus characterized by either lower or higher values of  $\delta^{15}$ N. Similar scenarios of dual 128 129 trophic pathways characterizing benthic systems were also found by Iken et al. (2005) in the Canadian Arctic; Drazen et al. (2008b) in the North Pacific; Reid et al. (2012) within the benthic 130 community sampled on the mid-Atlantic Ridge; Valls et al. (2014) in the western Mediterranean; 131 and Parzanini et al. (2017) in the Northwest Atlantic. Moreover, Kharlamenko et al. (2013) used 132 both stable isotopes and FA to study the dietary sources of benthic invertebrates collected along 133 the continental slope (500-1600 m depth) in the Sea of Japan. The authors recognized different 134 135 trophic pathways (i.e. planktonic, benthic, microbial) and dietary sources by using biochemical tracers; and they proposed a strong link with the primary production of the surface waters, as the 136 137 FA composition of the deep-sea echinoderms and mollusks was similar to that of the shallow-water 138 counterparts.

As POM sinks through the water column, its  $\delta^{15}N$  increases, reflecting the preferential 139 140 assimilation of the lighter isotope, <sup>14</sup>N by microbes; in particular, a gradient in POM  $\delta^{15}$ N has been 141 detected with depth, where POM at greater depths is more enriched (Altabet et al., 1999). For this 142 reason, Mintenbeck et al. (2007) carried out a study in the high-Antarctic Weddell Sea to assess whether this gradient was reflected in the isotopic signature of POM consumers sampled at 50-143 1600 m. In this regard, only those organisms feeding directly on sinking POM (e.g. suspension 144 145 feeders) showed increasing values of  $\delta^{15}N$  with depth, whereas the increase was less evident for the deposit feeders (Mintenbeck et al., 2007). Similar results for suspension feeders were obtained 146 by Bergmann et al. (2009) who analyzed a benthic food web sampled at the deep-water 147 observatory HAUSGARTEN, west of Svalbard (Arctic), between 1300 and 5600 m depth. 148 149 Conversely, deposit feeders exhibited a negative trend along the bathymetric gradient in terms of 150  $\delta^{15}$ N, and predator/scavengers were not affected. In another study, Sherwood et al. (2008) did not 151 detect any relationships with depth in the  $\delta^{15}$ N values measured from cold-water corals collected on 152 a slope environment in the Northwest Atlantic. Among the explanations suggested for these

153 inconsistencies and differences among feeding groups, Mintenbeck et al. (2007) and Sherwood et 154 al. (2008) included feeding preferences with respect to the size and sinking velocity of POM. According to these authors, only those organisms feeding on small particles of sinking POM should 155 156 reflect a bathymetric gradient in  $\delta^{15}$ N. In fact, small-sized particles sink at a lower velocity and, therefore, experience high rates of degradation, with more evident changes in  $\delta^{15}N$  (Mintenbeck et 157 al., 2007). Based on these findings, depth-stratified sampling should ideally be conducted when 158 studying a system characterized by a bathymetric gradient, as it would prevent biases in the 159 interpretation of the isotopic data. 160

Deep-water systems are generally characterized by a limited food supply, as the quantity of 161 162 food being transferred from the surface to the bottom diminishes with increasing depth (Gage, 163 2003). In addition, in temperate areas, food arrives as intermittent pulses, following the spring and 164 late summer blooms of primary (and secondary) productivity. For this reason, deep-water benthic 165 communities can only rely on fresh, high-quality phytodetritus within short temporal windows 166 following algal blooms; whereas reworked and resuspended POM fuels these communities for the 167 rest of the year (Lampitt, 1985). Deep-sea benthic organisms have hence developed adaptations 168 and strategies to increase their feeding success and minimize competition for food, including 169 trophic niche expansion and specialization. In this regard, certain benthic taxa (e.g. pennatulacean corals, hexactinellid sponges) and/or feeding groups (e.g. suspension and deposit feeders) at PAP 170 showed vertical extension of their trophic niches (i.e. omnivory) which, according to Iken et al. 171 172 (2001), was most likely driven by a strong competition for food. In other words, some species belonging to the same taxon or feeding guild shared similar food sources (i.e. exhibiting similar 173 174  $\delta^{13}$ C values), but they were located at different trophic levels (i.e. exhibiting a wide range of  $\delta^{15}$ N). Similarly, Jeffreys et al. (2009) reported trophic niche expansion among and within feeding guilds 175 176 sampled between 140 and 1400 m depth, at the Pakistan margin (Arabian Sea). Pennatulacean corals and other sestonivorous cnidarians, for example, displayed the greatest niche expansion; 177 178 they fed not only on POM, but also on small invertebrates (e.g. zooplankton). Moreover, ophiuroids, 179 which are typically selective deposit feeders, switched to an omnivorous diet under food-limited

180 conditions (Jeffreys et al., 2009). Apart from trophic niche expansion, Iken et al. (2001) proposed 181 that specialization on certain food items represented another adaptation developed by benthic organisms at PAP to mitigate competition for food. Holothuroid echinoderms, for instance, were 182 thought to accomplish food specialization through a combination of different factors involving 183 changes in morphology, mobility, and digestive abilities (Iken et al., 2001). Further examples of 184 trophic niche segregation and food partitioning, as strategies to minimize competition, were also 185 reported for deep-sea demersal fishes in the Northwest Mediterranean Sea (Papiol et al., 2013) 186 and for asteroid echinoderms in the Northwest Atlantic (Gale et al., 2013). Howell et al. (2003) 187 detected trophic niche expansion across different species of deep-sea asteroids (1053-4840 m) by 188 189 analyzing their FA composition. In particular, multivariate analysis of FA proportions discriminated 190 three different feeding guilds among the asteroids analysed, including mud ingesters, 191 predators/scavengers, and suspension feeders.

#### 192 **1.4 Sources of variation across studies**

When comparing studies relying on biochemical analysis, there are numerous sources of variation, which may influence results and findings, and also prevent the detection of similarities and general trends. However, their importance may depend on the scale of the investigation (i.e. local, regional, or global). In this section, the main sources of variation are illustrated and explained by type (Table 2).

## 198 **1.4.1 Biological sources**

Age, size, and sex, whether related to diet, determine natural intraspecific variability in the isotopic and FA compositions of organisms, which may affect data interpretation of small spatial scale investigations. At a basic level, sessile and sedentary taxa typically experience a transition from a pelagic to a benthic lifestyle between the larval and the juvenile stage (Rieger, 1994). Research has also shown that certain deep-sea fish experience changes in diet with age, typically with younger individuals preying upon benthic organisms and adults feeding on prey that are larger and of

205 benthopelagic origin (Mauchline and Gordon, 1984; Eliassen and Jobling, 1985). Stowasser et al. 206 (2009) combined stable isotope analysis (SIA) and FA analysis to detect ontogenetic shifts in the diet of the fish Coryphaenoides armatus and Antimora rostrata, collected at depths between 785 207 208 and 4814 m at PAP (Northeast Atlantic). By looking at their biochemical composition, the two species switched from active predation to scavenging with increasing size. Similar results are 209 reported in Drazen et al. (2008c) for macrourid fish species from the eastern North Pacific. 210 Conversely, although Reid et al. (2013) detected size-related trends in the  $\delta^{13}$ C of deep-water fish 211 collected from the Mid-Atlantic Ridge at 2400-2750 m depth, the authors were not able to 212 distinguish whether these results were due to ontogenetic changes in diet or merely to an effect of 213 214 increasing size, within the size-range sampled. Moreover,  $\delta^{15}N$  and trophic position may increase 215 with body size in adult shallow-water fish, as larger predatory fish ingest larger, more isotopically 216 enriched prey (Badalamenti et al., 2002; Galván et al., 2010).

217 The potential influence of sex as a source of variation in biomarker studies has not received as much attention and remains ambiguous. Nonetheless, Boyle et al. (2012) studied whether diet 218 219 and trophic position varied between sexes in deep-sea fish species collected at 55 -1280 m depth 220 in the eastern North Pacific using gut content and stable isotope analysis of muscle tissue. The 221 authors did not detect any difference between sexes, but variations in trophic position were encountered when analyzing fish of different sizes (Boyle et al., 2012). An investigation of the 222 oceanic squid Todarodes filippovae sampled within a depth range of 13-380 m in the southwestern 223 224 Indian Ocean by Cherel et al. (2009), revealed that females had higher values of  $\delta^{15}$ N, and thus occupied a higher trophic position. However, because T. filippovae exhibits sexual dimorphism in 225 body size, this difference was ultimately shown to be driven by size, i.e. no  $\delta^{15}$ N-variations were 226 detected when females and males of similar sizes were compared (Cherel et al., 2009). Sex may 227 228 constitute a source of variation in relation to diet in those species that exhibit extreme cases of sexual dimorphism, as in deep-sea anglerfish (Shine, 1989). However, investigation of the role of 229 230 sex on intraspecific variability will need to be carried out across a broader taxonomic scope before 231 drawing generalizations.

#### 232 **1.4.2 Environmental sources**

233 Larger-scale (e.g. regional, global) comparative studies among deep-sea habitats are complicated 234 by the wide bathymetric ranges they may occupy, anywhere between 200 and ~11 000 m depth. Depth may constitute a major driver of variation of  $\delta^{15}N$  and  $\delta^{13}C$  in deep-sea organisms for two 235 main reasons. First, as mentioned earlier, biodegradation processes occurring within the water 236 column may favour the enrichment of POM as it sinks, thus influencing the stable isotope 237 238 composition of those organisms that directly feed on it (Mintenbeck et al., 2007; Bergmann et al., 239 2009). Second, size-based trends and shifts in diet, hence in the isotopic composition, with depth have been reported for deep-sea demersal fish (Collins et al., 2005; Mindel et al., 2016a, 2016b). 240 241 Likewise, deep-sea species may exhibit different lipid and FA compositions along a bathymetric 242 gradient, reflecting physiological adaptations to changing temperature and pressure with depth (Parzanini et al., 2018b). 243

244 Geographic location (e.g. latitude) and season, linked to level/type of surface primary production, nitrogen supply dynamics, as well as temperature, are also important factors to 245 consider when comparing studies, as large-scale temporal and spatial differences may be detected 246 in the organisms' isotopic composition. Stowasser et al. (2009), for instance, combined stable 247 248 isotope and FA acid analyses to study seasonal variations in the diet of 5 species of demersal fish collected between 785 and 4814 m in the Northeast Atlantic. The authors found overall that stable 249 isotope and FA composition of fish varied temporally, and that these differences most likely 250 251 reflected timing and strength of food inputs sinking from surface waters. However, not all the species (e.g. Coryphaenoides armatus) exhibited a strong seasonality in their biochemical 252 253 composition, probably due to the high trophic position of the species and the length of the food web 254 analyzed obscuring the effects of the seasonal POM inputs (Stowasser et al., 2009). Colombo et al. (2016) detected a latitudinal gradient in the FA composition of marine species, with higher levels of 255 256 ω3-polyunsaturated fatty acids in organisms collected at polar and temperate regions in comparison to tropical ones. Large-scale geographic effects will be further explored below, in the 257

exploratory analytical section; however, Fig. 1 shows where food-web studies accomplished via
biochemical tracers have been carried out in heterotrophic ecosystems, highlighting important
geographic heterogeneity, especially the limited number of investigations in the southern
hemisphere.

## 262 **1.4.3 Analytical sources**

263 Several aspects of the SIA methodology can generate variability among studies, including type(s) 264 of tissue chosen for analysis, as well as sample treatment and storage, thus influencing interpretation of small-scale investigations. For instance, lipids have lower <sup>13</sup>C in comparison to 265 proteins and carbohydrates (DeNiro and Epstein, 1977), lipid-rich tissues hence display lower  $\delta^{13}$ C 266 values. In addition, there are tissues, such as liver in fish and gonads in other taxa, which are 267 268 characterized by higher turnover rates of lipids than others (e.g. white muscle), and hence incorporate information only on the recent diet. To avoid biases caused by the presence of lipids in 269 270 tissues, several approaches may be used. Stowasser et al. (2009) and Boyle et al. (2012), for 271 example, opted to extract lipid from the tissues prior to analysis, whereas Sherwood et al. (2008), 272 Fanelli et al. (2011a, 2011b) and Papiol et al. (2013) applied a mathematical correction to their  $\delta^{13}$ C data, based on the elemental C to N ratio (C:N) characterizing the samples. Other authors, such as 273 Polunin et al. (2001) and Carlier et al. (2009), did not apply any treatment. In the case of 274 mathematical corrections, two equations are currently used for deep-sea organisms, those 275 276 proposed by Post et al. (2007) and Hoffman and Sutton (2010). Since lipid extraction increases values of δ<sup>15</sup>N in deep-sea fish muscle tissue (Hoffman and Sutton, 2010), this practice is not 277 recommended. Conversely, mathematical corrections seem to be preferable when dealing with 278 279 lipids, and they have already been applied in several studies, including those mentioned above. 280 Some marine organisms, such as corals and echinoderms, contain carbonate skeletal elements. Since inorganic carbonate has higher  $\delta^{13}$ C values than other fractions (Pinnegar and 281 Polunin, 1999), it is a widespread practice to acidify these types of samples. Variations occur when 282 283 acidification is executed on samples that are simultaneously run for  $\delta^{15}N$  and  $\delta^{13}C$ , as the treatment

may affect δ<sup>15</sup>N data (Bunn et al., 1995). Whenever feasible, depending on both financial
constraints and the sizes of the organisms, processing samples separately for each isotope would
therefore be advisable, as in Carlier et al. (2009), Sherwood et al. (2008), and Papiol et al. (2013).

The tissues of elasmobranchs (e.g. sharks, rays) contain urea and trimethylamine oxide, 287 which are both <sup>15</sup>N-depleted; therefore, their presence may affect stable isotope data (Hussey et 288 al., 2012; Kim and Koch, 2012; Churchill et al., 2015). As for the inorganic carbonate issue, there is 289 no agreement among studies. Nonetheless, the removal of urea prior to analysis or the use of 290 arithmetic corrections are among the most common solutions applied to deal with the presence of 291 292 these compounds. In addition, the former seems to be the more commonly recommended and 293 performed, as the application of mathematical corrections requires the calculation of species-294 specific discrimination factors, which is not always feasible (Hussey et al., 2012).

295 Sample storage is also crucial to obtain reliable data, since non-optimal preservation 296 methods may compromise the outcome of the investigation. Regarding the storage temperature, 297 while biological samples for gut content and stable isotope analysis are commonly frozen at -20°C, 298 if not processed soon after their collection; those for lipid analysis are either stored at -80°C 299 (recommended) or at -20°C prior to further processing in the lab. Since storage at -20°C might not 300 completely prevent lipid degradation, especially if samples are analyzed after several years, rapid initial processing of samples and vacuum packing may reduce potential issues when freezing at -301 80°C is not logistically feasible. In addition, freezing is highly recommended over chemical storage 302 303 for stable isotope analysis, as there is evidence that formalin/ethanol considerably alters the 304 isotopic ratios in biological tissues (Arrington and Winemiller, 2002; Syväranta et al., 2011; Xu et al., 2011). 305

306 2 Preliminary comparative analysis

The study of large-scale trends in biological variables (e.g. distribution, biochemical composition, biodiversity) may not only help understand general functioning and structure of ecosystems, but it may also allow us to make predictions and support conservation initiatives. While several studies already exist on large-scale distribution and biodiversity patterns of deep-sea species (Rex et al.,
1993; Stuart et al., 2003; Ramirez-Llodra et al., 2010), a similar approach has yet to be applied to
trophodynamics. This preliminary analysis detected global spatial trends (i.e. along latitudinal and
depth gradients) in the isotopic and FA composition of deep-water animals for the first time since
the application of biochemical tracers to the study of trophic ecology in the deep sea.

Latitudinal gradients have been detected in  $\delta^{13}$ C of plankton and POM collected from 315 surface waters in both the southern and northern hemispheres, with decreasing values towards the 316 polar regions (Sackett et al., 1965; Rau et al. 1982; Francois et al., 1993). Both environmental (e.g. 317 temperature, nutrient supply) and biological (e.g. plankton metabolism) factors have been proposed 318 319 to explain such trends (Rau et al., 1982; Francois et al., 1993). The stable N isotope signature of 320 surface primary production may also vary regionally, depending on the nutrient (mainly N) supply to 321 the phytoplankton, as well as its community structure and cell size (Choy et al., 2015; Hetherington 322 et al., 2017). Oligotrophic areas, characterized by marked oxygen minimum zones and by high denitrification rates, such as the eastern tropical Pacific Ocean, typically have higher  $\delta^{15}$ N values 323 (Hetherington et al., 2017). In addition, latitudinal trends have been detected in the FA composition 324 of marine organisms, which tend to have higher levels of essential  $\omega$ 3 long-chain polyunsaturated 325 fatty acids (LC-PUFA) in the polar and temperate regions in comparison to the tropical ones 326 (Colombo et al., 2016). As POM is the main food source of most deep-sea food webs (Gage, 2003; 327 Hudson et al., 2004), we hypothesized that a) similar latitudinal gradients exist in the isotopic and 328 329 essential PUFA composition of deep-water organisms; and that b) the strength of these trends varied among organisms from different habitats, i.e. pelagic, demersal, and benthic, as diversely 330 331 dependant on POM. Furthermore, as both isotopic and lipid composition of POM and as deep-sea taxa varied along a depth gradient in the deep North Pacific (Lewis, 1967; Altabet et al., 1999), 332 333 North Atlantic (Polunin et al., 2001; Parzanini et al., 2018a, 2018b, 2017) and Arctic Ocean (Bergmann et al., 2008), we hypothesized that similar trends could be extended to the global scale. 334

## 335 2.1 Materials and methods

#### 336 2.1.1 Data set

This analysis focused on studies that used either bulk stable isotope or FA analysis, or a 337 combination of them, to infer trophic relationships of deep-water macro- and megafauna, as well as 338 to study deep-sea food webs, from heterotrophic ecosystems. Experimental studies, as well as 339 340 investigations on chemosynthetic habitats (e.g. hydrothermal vents) were excluded a priori to avoid possible biases. In fact, these habitats are fuelled by primary dietary sources, e.g. methane, whose 341 isotopic and FA composition is substantially different than that of POM (Rau and Hedges, 1979; 342 343 Saito and Osako, 2007). Table 3 outlines the full data set collated for the present analysis, which 344 includes 52 different studies. The literature search was carried out through Scopus and Google Scholar portals using the following key words: stable isotopes, fatty acids, food webs, deep sea, 345 346 trophic ecology, and trophic relationships. Additional sources provided by an anonymous referee were also included. These studies were used to analyze global trends in  $\delta^{15}N$ ,  $\delta^{13}C$ , and the 347 348 essential arachidonic (ARA, 20:4 $\omega$ 6), eicosapentaenoic (EPA; 20:5 $\omega$ 3) and docosahexaenoic (DHA, 22:6ω3) acids across deep-water communities. ARA, EPA, and DHA are the most important 349 nutrients in aquatic ecosystems, required by organisms for optimal health (Parrish 2009), as well as 350 excellent trophic biomarkers. In fact, whereas EPA and DHA are typically used as biomarkers in 351 352 diatoms and dinoflagellates respectively (Parrish, 2013), in the deep sea, ARA is associated with microorganisms from the sediment (Howell et al. 2003). Our study focused on these three FA since 353 they are present in all the organisms under analysis. 354

#### 355 2.1.2 Variables considered

Each species from each investigation was sorted by latitude (i.e. tropical, 0 - 30°; temperate, 30 -60°; and polar, 60 - 90°), habitat (i.e. pelagic, demersal, and benthic) depth at collection (i.e. mesopelagic, 200 – 1000 m; bathypelagic, 1000 – 4000 m; and abyssopelagic, >4000 m, for pelagic species; bathyal 200 - 4000 m; abyssal, 4000 - 6000 m; and hadal, > 6000 m, for benthic

360 species), and phylum (i.e. Annelida, Arthropoda, Brachiopoda, Bryozoa, Chaetognatha, Chordata, 361 Cnidaria, Hemichordata, Echinodermata, Mollusca, Nematoda, Nemertea, Porifera, and Sipuncula). Information about species habitat was either obtained through WoRMS and FishBase online 362 363 databases or was already included in the source paper. In addition, species were labelled as 364 "meso-bathypelagic" and "bathyal-abyssal", if the depth at collection was not specified further, but the whole set of samples for a study was collected within those zones. In the current analysis, 365 tissue type, acidification treatment, sampling season, sex, and age were not considered as 366 variables, because i) they were assumed to not play a major role in global-scale investigations 367 and/or ii) this information was not always provided. In addition, tests were performed on lipid-368 corrected and uncorrected  $\delta^{13}C$  data pooled together. For analyses regarding stable isotope 369 370 composition ( $\delta^{15}N$ ,  $\delta^{13}C$ ), data were obtained from lken et al. (2005), Mincks et al. (2008), 371 Bergmann et al. (2009), Quiroga et al. (2014), and van Oevelen et al. (2018), for polar regions; Iken et al. (2001), Madurell et al. (2008), Sherwood et al. (2008), Carlier et al. (2009), Fanelli et al. 372 (2009), Stowasser et al. (2009), Fanelli et al. (2011a, 2011b), Boyle et al. (2012), Reid et al. (2012), 373 374 Fanelli et al. (2013), Gale et al. (2013), Kharlamenko et al. (2013), Papiol et al. (2013), Reid et al. 375 (2013), Tecchio et al. (2013), Kiyashko et al. (2014), Trueman et al. (2014), Valls et al. (2014a, 376 2014b), Kopp et al. (2018), Parzanini et al. (2017), Preciado et al. (2017) and Parzanini et al. (2018a), for temperate latitudes; and Jeffreys et al. (2009), Churchill et al. (2015), Shipley et al. 377 (2017), and Richards et al. (2019), for tropical regions (Table S1). FA composition (ARA, EPA, and 378 379 DHA) data were collected from Pétursdóttir et al. (2008a, 2008b), and Würzberg et al. (2011a, 2011b, 2011c), for polar areas; Lewis (1967), Howell et al. (2003), Hudson et al. (2004), Økland et 380 al. (2005), Drazen et al. (2008a, 2008b), Stowasser et al. (2009), Murdukhovich et al. (2018), 381 Parzanini et al. (2018a), Salvo et al. (2018), van Oevelen et al. (2018), and Kharlamenko et al. 383 382 (2018), for temperate regions; and Jeffreys et al. (2009) and Shi et al. (2018), for tropical regions 384 (Table S2).

## 385 **2.2 Statistical analysis**

Comparisons among multiple groups of deep-sea organisms were run through t-tests and oneway 386 analysis of variance (ANOVA). In particular, isotopic (i.e.  $\delta^{15}N$ ,  $\delta^{13}C$ ) and FA (i.e. ARA, EPA and 387 388 DHA) data were compared across organisms from different latitudes (i.e. tropical, temperate and 389 polar), habitats (i.e. pelagic, demersal, benthic), and collection depths (i.e. mesopelagic, bathypelagic, meso-bathypelagic, abyssopelagic, bathyal, bathyal-abyssal, abyssal, and hadal) to 390 detect any significant differences. When the normality assumption was violated, Mann-Whitney 391 rank sum test, Kruskal-Wallis oneway ANOVA on ranks, and Dunn's method pairwise comparisons 392 393 were performed instead. In addition, multivariate statistics, i.e. principal coordinate analysis (PCO) and permutational MANOVA (PERMANOVA) were used to study the variability in the isotopic and 394 395 FA composition of deep-water organisms across different latitudes, habitats, collection depths, and 396 phyla. In addition, a distance based linear model (DistLM) was run to assess which of these four 397 factors contributed the most to such a variability. PCO, PERMANOVA, and DistLM were run on 398 resemblance matrices, based on Euclidean distance for the isotopic data, and Bray-Curtis for the 399 FA data. Data were not normalized or transformed prior to analysis. Univariate statistics was conducted using Sigmaplot 12.5, while PCO, PERMANOVA and DistLM were run through Primer 400 401 7.0 with the add-on package PERMANOVA+ (Clarke and Gorley, 2006).

## 402 **2.3 Results**

Analyses revealed both latitudinal and depth-related trends for isotopic and essential FA composition. In particular, mean values ( $\pm$ SD) of  $\delta^{15}$ N and  $\delta^{13}$ C were significantly lower in deep-sea fauna sampled at high latitudes than in that collected at low latitudes ( $\delta^{15}$ N, ANOVA on Ranks, *H* = 35.6, *p* ≤ 0.001;  $\delta^{13}$ C, ANOVA on Ranks, *H* = 277.9, *p* ≤ 0.001; Fig. 2). Conversely, no difference was detected across latitudes in terms of ARA, but mean proportions ( $\pm$ SD) of EPA and DHA were significantly greater at polar latitudes than at temperate and tropical areas (EPA, ANOVA on Ranks, *H* = 11.4, *p* = 0.003; DHA, ANOVA on Ranks, *H* = 63.6, *p* ≤ 0.001; Fig. 3). Similarly, PERMANOVA

410 detected significant differences across latitudes in terms of both stable isotopes [Pseudo-F = 81.4,

411 p(perm) = 0.0001] and essential FA [Pseudo-F = 11.0, p(perm) = 0.0001].

When deep-water species were analyzed separately according to their habitat, the same 412 latitudinal trend in the isotopic composition were shown for deep-water benthic species ( $\delta^{15}$ N, 413 414 ANOVA on Ranks, H = 40.5,  $p \le 0.001$ ;  $\delta^{13}$ C, ANOVA on Ranks, H = 171.2,  $p \le 0.001$ ); whereas, for demersal and pelagic species, only the  $\delta^{13}$ C ratios were significantly lower at higher latitudes 415 (ANOVA on Ranks, H = 105.7,  $p \le 0.001$ , for demersal species; ANOVA on Ranks, H = 11.5, p =416 0.003, for pelagic species). PERMANOVA showed that the isotopic composition of deep-sea 417 animals was indeed statistically different across the three habitats [Pseudo-F = 112.6, p(perm) = 418 419 0.0001], and benthic and demersal species had higher stable N and C isotope ratios than the 420 pelagic counterparts (p < 0.05). Conversely, only benthic and pelagic species revealed a latitudinal 421 gradient in their essential FA composition (EPA, ANOVA on Ranks, H = 12.1, p = 0.002; DHA, ANOVA on Ranks, H = 43.6,  $p \le 0.001$ , for benthic species; EPA, ANOVA, H = 6.4, p = 0.011, for 422 pelagic taxa). In this regard, pelagic, demersal, and benthic taxa had a different essential FA 423 424 composition (ARA, ANOVA on Ranks, H = 39.7,  $p \le 0.001$ ; EPA, ANOVA on Ranks, H = 12.5, p =425 0.002; DHA, ANOVA on Ranks, H = 76.9,  $p \le 0.001$ ; Pseudo-F = 19.7, p(perm) = 0.0001). Benthic 426 species had the highest proportions of ARA and EPA (p < 0.05); while demersal species had the highest levels of DHA, although similar to those of pelagic species. 427

While mean values of both stable N and C isotope ratios significantly increased with depth 428 429 for benthic and demersal species ( $\delta^{15}$ N, ANOVA on Ranks, H = 63.9,  $p \le 0.001$ ;  $\delta^{13}$ C, ANOVA on Ranks, H = 126.2,  $p \le 0.001$ ), only  $\delta^{13}$ C ratios showed the same trend in pelagic taxa (ANOVA on 430 431 Ranks, H = 125.5,  $p \le 0.001$ ). Proportions of EPA significantly decreased along the bathymetric gradient for pelagic taxa (ANOVA on Ranks, H = 12.3, p = 0.002), and levels of ARA were 432 433 significantly higher at abyssal depths for benthic and demersal species (ANOVA on Ranks, H =39.7,  $p \le 0.001$ ). In addition, for benthic and demersal fauna, levels of  $\delta^{15}N$ ,  $\delta^{13}C$ , and ARA 434 increased for benthic and demersal organisms with increasing depth ( $\delta^{15}N$ , ANOVA on Ranks, H =435 84.7,  $p \le 0.001$ ; δ<sup>13</sup>C, ANOVA on Ranks, H = 105.0,  $p \le 0.001$ ; ARA, ANOVA on Ranks, H = 22.8, 436

437  $p \le 0.001$ ). PERMANOVA revealed significant differences in the isotopic [Pseudo-F = 74.6, p(perm)438 = 0.0001] and essential FA composition [Pseudo-F = 8.6, p(perm) = 0.0001] across collection 439 depths.

Among the four variables considered (i.e. latitude, habitat, collection depth, and phylum), analyses revealed that 'habitat' and 'phylum' were the most important factors influencing the variability of the stable isotope (respectively 12 and 9%; DistLM, *adjusted*  $R^2 = 0.4$ ) and FA (respectively 8 and 11%; DistLM, *adjusted*  $R^2 = 0.3$ ) composition of deep-water organisms (Fig. 4).

#### 444 **2.4 Discussion**

445 The present analysis shows for the first time, the existence of a) latitudinal trends in both stable isotope and essential FA composition of deep-sea organisms, with decreasing  $\delta^{13}$ C ratios and 446 447 increasing ω3 LC-PUFA towards the poles; b) global bathymetric trends in the isotopic composition 448 of deep-water fauna for which mean levels of  $\delta^{15}N$ ,  $\delta^{13}C$ , and ARA increased with increasing depth. In addition, it provides further evidence of the link, across latitudes and depth, between surface 449 450 primary production of the surface waters and the deep-water consumers. The present findings generally align with reports of decreasing values of  $\delta^{13}$ C in surface-waters plankton and POM 451 452 towards the polar regions, in both the southern and northern hemisphere (Sackett et al., 1965; Rau 453 et al., 1982; Francois et al., 1993), as well as of increasing POM isotopic ratios along a bathymetric gradient (Altabet et al., 1999). They also agree with Colombo et al. (2016) who noticed that 454 proportions of  $\omega$ 3 LC-PUFA were higher in marine organisms from polar and temperate regions in 455 456 comparison to tropical regions, and with Parzanini et al. (2018a) who detected increasing 457 proportions of ARA along a slope area in the deep Northwest Atlantic. Water temperature, in combination with other abiotic (e.g. oceanographic and 458 biogeochemical processes, nutrient supply) and biological factors (e.g. species metabolism, 459 460 taxonomic composition of deep-water communities, microbial remineralization processes) seems to play a role in these trends (Rau et al., 1982; Francois et al., 1993; Altabet et al., 1999; Colombo et 461 462 al., 2016). In particular, water temperature influences isotopic fractionation processes and, typically,

463 higher fractionation is associated with lower temperatures (Sackett et al., 1965). High fractionation rates are also linked to the pronounced denitrification activities characterizing oligotrophic areas 464 465 such as observed in some areas of the tropics (Hetherington et al., 2017). This may explain the higher  $\delta^{15}$ N ratios of the deep-sea organisms from the tropical latitudes analyzed in this study. 466 Furthermore, water temperature affects membrane fluidity, and lower temperatures decrease the 467 fluidity of cell membrane (Parrish, 2013; Colombo et al., 2016). Thus, in order to maintain normal 468 membrane function and condition, i.e. health, ectotherms may counteract variations in water 469 temperature by readjusting their FA composition (Cossins and Lee, 1985; Parrish, 2013). For 470 example, larger proportions of long chain unsaturated FA (e.g. ARA, EPA) within the lipid bilayer 471 472 help increase membrane fluidity (Parrish 2013), as these molecules are characterized by a higher 473 flexibility (DeLong and Yayanos, 1985; Colombo et al., 2016).

474 Trends in the isotopic and FA composition of deep-sea organisms were also seen along a 475 depth gradient. As a proxy for water temperature as well as nutrient supply, depth may influence 476 biochemical composition of marine consumers (Parzanini et al., 2018a, 2018b). POM becomes 477 more isotopically enriched while sinking to deeper depth due to microbial degradation (Altabet et al., 1999). Thus, the isotopic composition of deep-water organisms which feed on POM may vary 478 479 accordingly (Mintenbeck et al., 2007). In the present analysis, levels of ARA were globally higher at deeper depths, similar to the study by Parzanini et al. (2018a), which may be due to i) a higher 480 reliance of deeper-dwelling organisms on the benthic-detrital trophic pathway; and/or ii) the need to 481 482 maintain membrane fluidity at low temperatures via increasing the unsaturation levels of membrane phospholipids. 483

Finding latitudinal trends in the biochemical composition of deep-water organisms that mirror results from shallow depths provides further evidence of the link between the two systems, in that deep-sea benthic communities rely on POM sinking from the surface water as a primary food source (Gage, 2003; Hudson et al., 2004). Close dependence of deep-sea food webs on nearsurface processes raises important concerns. According to the latest climate estimates, both air and water temperatures have been rising, and continue to increase; and seawater pH has already

490 dropped by 0.1 units due to large  $CO_2$  emissions, and is expected to decrease further (IPCC, 491 2017). Furthermore, models predict that increasing surface water temperature will favor 492 stratification, while reducing vertical mixing as well as enhancing variability in the transport of 493 primary production and energy (i.e. carbon) transport to the deep sea (Smith et al., 2009; Jones et 494 al., 2014; Sweetman et al., 2017). At the same time, deep-water benthic biomass is expected to decrease due to the increasing variability in the food supply, which may in turn affect health and 495 functioning of benthic ecosystems, as well as global biogeochemical cycles (Jones et al., 2014). 496 Hixson and Arts (2016) showed that the FA composition of the six most common fresh- and salt-497 water phytoplankton species responded to temperature and, specifically, that their  $\omega$ 3 PUFA levels 498 499 decreased with increasing temperature. Not only do  $\omega$ 3 PUFA, such as EPA and DHA, play an 500 important role in the response to temperature variations in aquatic systems, but they are also 501 essential nutrients and are highly required by aquatic organisms for optimal growth and health 502 (Parrish, 2009). A case in point, Rossoll et al. (2012) showed experimentally that growth and reproduction of the copepod Acartia tonsa were severely compromised by the alteration of FA 503 content and composition of its primary food source, the diatom Thalassiosira pseudonana, exposed 504 505 to high  $CO_2$  levels. The present investigation, therefore, suggests that changes in amounts and 506 composition of surface production could also result in changes in essential nutrients and biomarkers in deep-sea benthic organisms that feed on it, with possible cascading effects 507 throughout deep-water food webs. Such variations may alter nutrient intake of deep-sea benthic 508 509 organisms, as well as trophodynamics; and they may also influence species' abilities to cope with deep cold waters. 510

# 511 **3 Conclusions**

This investigation provides a first summary of the information available on deep-sea food webs inferred by bulk stable isotope and FA analyses, providing guidance for future studies and a glimpse at global-scale patterns in the biochemical composition of deep-water organisms from heterotrophic ecosystems. Food-web tracers represent a powerful tool that can help elucidate the

516 structure and dynamics of food webs from shallow to deeper waters, and support management 517 initiatives. However, this tool is even more effective when combined with other techniques (e.g. gut 518 content analysis), as each method provides uniquely valuable data. When comparing studies, it 519 emerges that there are multiple sources of variations, whether biological, environmental, and/or analytical. Depending on the scale of the investigation, these differences are more or less 520 521 susceptible to biases, suggesting that they have to be considered and acknowledged when attempting cross-comparisons even though they may be contextually acceptable. The preliminary 522 analysis conducted here detected latitudinal and bathymetric trends in the isotopic and FA 523 composition of deep-sea species. In light of global climate change and the link between surface 524 525 production and deep-sea communities, changes in amounts and composition of surface production 526 may influence the essential nutrient intake (e.g.  $\omega$ 3 PUFA) of deep-water organisms. Because  $\omega$ 3 527 PUFA are involved in the response to temperature variations in ectotherms, climate change may 528 also affect the ability of these species to cope with potential temperature shifts. However, more 529 studies are required to help detect global trends, especially in those areas that are still poorly 530 understood (most deep-sea areas) or not yet investigated (e.g. in the southern hemisphere). In addition, it is necessary to standardize analytical methods to limit their influence and help 531 compensate for natural variability. 532

# 533 Data availability

All data used for analysis can be found as supplementary material, in Table S1 and S2.

Table S1. Dataset applied to analyze trends in the isotopic composition of deep-sea
animals.

Table S2. Dataset applied to analyze trends in the essential FA composition of deep sea animals.

# 539 Author contribution

- 540 All the authors contributed to the manuscript conceptualization and methodology. CP was
- responsible of data curation, formal analysis, investigation, and in writing the original draft of the
- 542 manuscript. CCP, JH, and AM reviewed and edited the draft. Lastly, CCP and AM provided
- 543 supervision, as well as funds to this project.

# 544 **Competing interests**

545 The authors declare that they have no conflict of interest.

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# **Tables**

**Table 1** Comparison outlining the major strengths and drawbacks of gut content, stable isotope, and FA analysis.

Gut content analysis	Stable isotope analysis	FA analysis		
Direct evidence of diet	Indirect evidence of diet (assumption validation required)	Indirect evidence of diet (assumption validation required)		
Snap shot of the most recent meal	Integrative over time	Integrative over time		
Small sample sizes may lower representativity of diet	Small sample sizes may lower representativity of diet	Small sample sizes may lower representativity of diet		
Inter-individual variability can only be accounted for with appropriate sample size	Inter-individual variability minimized due to integrative nature	Inter-individual variability likely but minimized due to integrative nature		
Temporal variability can only be accounted for with appropriate sample size	Temporal variability minimized due to integrative nature	Temporal variability minimized due to integrative nature		
Partly dependent on sex in cases where there are dietary differences between sexes	Partly dependent on sex in cases where there are dietary differences between sexes	Partly dependent on sex in cases where there are dietary differences between sexes		
May be sensitive to body size (e.g. onthogenetic dietary changes)	May be sensitive to body size, whether or not size influences diet	Dependent body size if size affects diet		
Species with large stomachs and slow digestion rates are easier to study	Applies to all species, but requires enough material (see below)	Applies to all species, but requires enough materia (see below)		
The analysis cannot be carried out with empty stomachs	Independent of stomach fullness	Independent of stomach fullness		
Digestion rates may bias contents recovered	Independent of digestion process	Independent of digestion process		
Small specimens with small stomachs are more difficult to study	Small specimens may have to be pooled, guts included	Small specimens may have to be pooled, guts included		
Only gut content is analyzed	Typically applied to target tissues	Typically applied to target tissues		
Interpretation is relatively easy, unless food is highly digested, and the evidence obtained cannot be misinterpreted, taxonomically speaking	Data interpretation is complex (post-analysis mathematical corrections are often applied)	Data interpretation is complex (linked to FA biomarkers as food tracers)		
Long processing time	Relatively short processing time	Relatively short processing time		
Little instrumentation, low cost (unless high resolution scopes are used)	Medium technology, medium/high cost	Medium technology, medium/high cost		

Biological	Analytical	Environmental	
Taxonomy	Sample gear	Depth	
Sex	Sample storage	Season	
Age	Sample treatment (e.g. Acidification of organisms containing carbonatic anatomical elements; Lipid removal; urea removal)	Primary productivity levels at surface Latitude	
Size	Mathematical correction (i.e. whether applied and which one)		
Feeding habits	Tissue type	Temperature	
General physiological condition		Ocean region	
		Geological feature (e.g. shelf, slope, canyon, plain, trench)	

**Table 2** Sources of variations across studies, distinguished by type (i.e. biological, environmental, analytical).

**Table 3** List of trophic ecology studies in deep-sea heterotrophic systems, carried out using stable isotopes (bulk) and lipids (including FA) as food-web tracers. Experimental studies were excluded *a priori*. Reference, method(s) applied, latitude, sampling depth, ocean region, and taxa analyzed are reported for each study. Polar latitudes include investigations between 60 - 90° N/S, whereas temperate and tropical latitudes represent studies carried out within 0 - 30° N and 30 - 60° N, respectively. References are ordered according to sampling depth(s).

References	Method(s)	Latitude	Depth	Ocean region	Taxa analyzed
	wethod(s)		(m)	Ocean region	
Mintenbeck et al. 2007	Stable isotopes	Polar	50-1600	Weddell Sea (Antarctic)	Benthic bryozoans, cnidarians, crustaceans, echinoderms, echiurans, mollusks, sponges, sipuncules, and tunicates
Quiroga et al. 2014	Stable isotopes	Polar	250-322	Weddell Sea	Benthic annelids, crustaceans, bryozoans, tunicates, cnidarians, echinoderms, molluscs, nemertea worms, sponges and sipunculans.
van Oevelen et al. 2018	Stable isotopes, Lipids	Polar/Temperate	270-850	Trænadjupet Trough (Norwegian continental shelf), Belgica Mounds (Porcupine Seabight)	Cold-water coral communities
Mincks et al. 2008	Stable isotopes	Polar	550-650	Bellinghausen Sea	Benthic annelids, cnidarians, echinoderms, molluscs, sponges, and demersal fish
Würzberg et al. 2011a	Lipids	Polar	600-5337	Weddell Sea (Antarctic)	Shelf and deep-sea peracarid crustaceans + foraminiferans
Würzberg et al. 2011b	Lipids, Gut contents	Polar	600-2150	Weddell Sea (Antarctic)	Demersal fish
Würzberg et al. 2011c	Lipids	Polar	600-5337	Weddell Sea (Antarctic)	Shelf and deep-sea polychaetes
lken et al. 2005	Stable isotopes	Polar	800-2082	High Arctic Canadian Basin	Benthic cnidarians, crustaceans, echinoderms, echiurans, mollusks, and polychaetes; pelagic crustaceans
Pétursdóttir et al. 2008a	Stable isotopes, Lipids	Polar	1000-2000	Reykjanes Ridge (North Atlantic)	Mesopelagic crustaceans and fish
Pétursdóttir et al. 2008b	Stable isotopes, Lipids	Polar	1000-2001	Reykjanes Ridge (North Atlantic)	Mesopelagic crustaceans and fish
Bergmann et al. 2009	Stable isotopes	Polar	1300-5600	HAUSGARTEN observatory, west Svalbard (Arctic)	Benthic cnidarians, crustaceans, echiurans, echinoderms, mollusks, nemertean worms, polychaetes, priapulids, sponges, and tunicates; Demersal fish

Valls et al. 2014a	Stable isotopes	Temperate	40-400	Balearic Basin (western Mediterranean)	Mesopelagic fish and zooplankton
Sherwood et al. 2008	Stable isotopes	Temperate	47-1433	Northwest Atlantic	Cold-water corals
Hamoutene et al. 2008*	Lipids	Temperate	50-1500	Cape Chidley, and southern Grand Bank (Northwest Atlantic)	Cold-water corals
Boyle et al. 2012	Stable isotopes, Gut contents	Temperate	55-1280	eastern North Pacific	Benthic cnidarians, crustaceans, echinoderms, and mollusks; polychaetes; demersal fish
Polunin et al. 2001	Stable isotopes	Temperate	200-1800	Balearic Basin (western Mediterranean)	Demersal fish
Valls et al. 2014b	Stable isotopes	Temperate	250-850	Balearic Basin (western Mediterranean)	Hyperbenthic echinoderms and hyperbenthic/pelagic crustaceans, elasmobranchs and mollusks
Gale et al. 2013	Stable isotopes, Gut contents	Temperate	258-1418	Northwest Atlantic	Echinoderms
Carlier et al. 2009	Stable isotopes	Temperate	300-1100	Ionian Sea (central Mediterranean)	Cold-water coral community
Parzanini et al. 2018a	Stable isotopes, Lipids, Elemental	Temperate	310-1413	Northwest Atlantic	Slope cnidarians, crustaceans, echinoderms, fish, mollusks, sponges and tunicates
Parzanini et al. 2018b	Lipids Stable	Temperate	310-1413	Northwest Atlantic	Slope cnidarians, crustaceans, echinoderms, fish, mollusks, sponges and tunicates
Parzanini et al. 2017	isotopes, Gut contents, Morphometrics	Temperate	310-1413	Northwest Atlantic	Pelagic and demersal fish
Madurell et al. 2008	Stable isotopes	Temperate	350-780	Balearic Basin (western Mediterranean)	Suprabenthic crustaceans and fish
Kopp et al. 2018	Stable isotopes	Temperate	415-516	Celtic Sea (Northeast Atlantic)	Epifaunal crustaceans, mollusks, and fish
Papiol et al. 2013	Stable isotopes	Temperate	423-1175	Balearic Basin (western Mediterranean)	Benthopelagic crustaceans
Fanelli et al. 2013	Stable isotopes	Temperate	445-2198	Balearic Basin (western Mediterranean)	Slope crustaceans and mollusks
Økland et al. 2004	Lipids	Temperate	500-1600	Porcupine Bank and western continental slope (Northeast Atlantic)	Demersal fish
Trueman et al. 2014	Stable isotopes	Temperate	500-1500	Hatton Bank (Northeast Atlantic)	Demersal fish
Kharlamenko et al. 2013	Stable isotopes, Lipids	Temperate	500-1600	Sea of Japan	Echinoderms and mollusks
Preciado et al. 2017	Stable isotopes, Gut contents	Temperate	625-1800	Galicia Bank (Northeast Atlantic)	Demersal fish and pelagic/demersal crustaceans
Fanelli et al. 2009	Stable isotopes	Temperate	650-780	Algerian Basin (western Mediterranean)	Mesopelagic crustaceans and fish; benthic crustaceans
Fanelli et al. 2011a	Stable isotopes, Gut contents	Temperate	650-800	Balearic Basin (western Mediterranean)	Zooplankton and micronekton
Fanelli et al. 2011b	Stable isotopes	Temperate	650-1000	Balearic Basin (western Mediterranean)	Epibenthic/infaunal nemertin worms, polychaetes, sipuncules, mollusks, crustaceans, echinoderms

Salvo et al. 2017	Lipids	Temperate	770-1370	Northwest Atlantic	Cold water corals
Stowasser et al. 2009	Stable isotopes, Lipids, Gut contents	Temperate	785-4814	Porcupine Seabight and Abyssal Plain (Northeast Atlantic)	Moridae and Macrouridae fish
Hudson et al. 2004	Lipids	Temperate	800-4850	Porcupine Seabight and Abyssal Plain (Northeast Atlantic)	Holoturoids
Howell et al. 2003	Lipids	Temperate	1053-4840	Porcupine Abyssal Plain (Northeast Atlantic)	Asteroids
Tecchio et al. 2013	Stable isotopes	Temperate	1200-3000	Mediterranean Sea (western + central + eastern)	Zooplankton
Reid et al. 2012	Stable isotopes	Temperate	2400-2750	Mid-Atlantic Ridge (North Atlantic)	Benthic cnidarians, crustaceans, echinoderms,
Reid et al. 2013	Stable isotopes	Temperate	2404-2718	Mid-Atlantic Ridge (North Atlantic)	Deep-sea fish
Kiyashko et al. 2014	Stable isotopes	Temperate	2481-3666	Sea of Japan	Benthic annelids, crustaceans, ascidians, cnidarians, echinoderms, molluscs and sponges
Mordukhovich et al. 2018	Lipids	Temperate	3352-4722	Sea of Okhotsk and Pacific Ocean	Deep-sea macro-benthic nematodes
Kharlamenko et al. 2018	Lipids	temperate	>4000	Sea of Okhotsk	Benthic annelids, echinoderms, molluscs, and sipunculans
Drazen et al. 2008a	Lipids	Temperate	4100	eastern North Pacific	Ophiuroids and holoturoids
Drazen et al. 2008b	Lipids	Temperate	4100	eastern North Pacific	Cnidarians, polychaetes and crustaceans, demersal and pelagic crustaceans and fish
Drazen et al. 2008c*	Stable isotopes, Gut contents	Temperate	4100	eastern North Pacific	Macrourid fish
Drazen et al. 2009	Lipids	Temperate	4100	eastern North Pacific	Macrourid fish and cephalopods
lken et al. 2001	Stable isotopes	Temperate	4840	Porcupine Abyssal Plain (Northeast Atlantic)	Demersal/Benthic cnidarians, crustaceans, echinoderms, echiurans, fish, mollusks, nematodes, polychaetas, sipuncules, and tunicates
Lewis, 1967	Lipids	Tropical	0-4000	Off San Diego and Baja California (eastern Pacific)	Demersal and pelagic crustaceans and fish
Jeffreys et al. 2009	Stable isotopes, Lipids	Tropical	140-1400	Arabian Sea	Crustaceans, cnidarians, and echinoderms
Churchill et al. 2015	Stable isotopes, Gut contents	Tropical	250-1200	south-central Gulf of Mexico, off Florida to Louisiana (western Atlantic)	Elasmobranchs
Shipley et al. 2017	Stable isotopes	Tropical/Polar	472-1024	Exuma Sound (The Bahamas), Lancaster Sound (Canadian Arctic)	Elasmobranchs
Richards et al. 2019	Stable isotopes	Tropical	1000-3000	Gulf of Mexico	Meso-bathypelagic fish
Shi et al. 2018	Lipids	Tropical	>6000 m	Pacific Ocean	Benthic amphipods

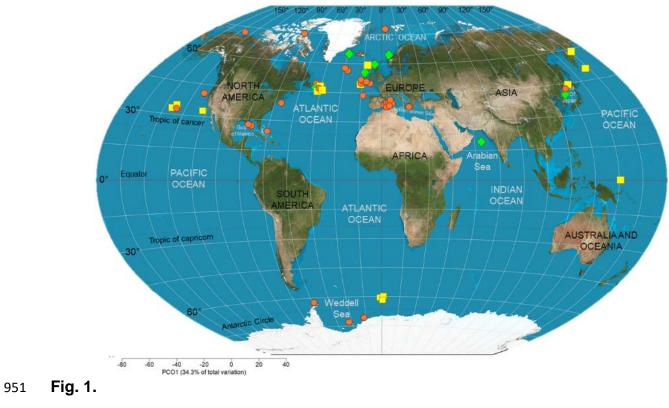
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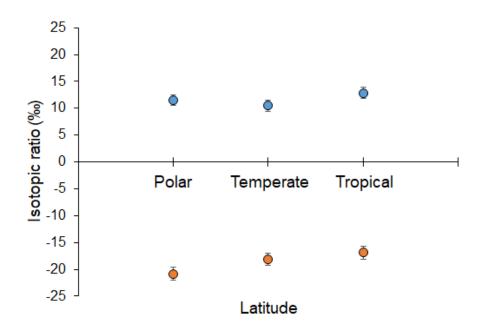
**Fig. 1.** Deep-sea biomarker studies in the world ocean. Symbols indicate where the studies listed in Table 2 have been carried out. In detail, red circles represent those investigations that have used stable isotopes as food web tracers; whereas yellow squares and green diamonds indicate those which used lipids and a combination of SIA and FA analysis, respectively.

**Fig. 2.** Stable N and C isotopic composition of deep-sea animals across latitudes. Mean values of  $\delta^{15}$ N (blue circles above) and  $\delta^{13}$ C (orange circles below) (‰) measured in deep-sea organisms across polar, temperate, and tropical latitudes. Bars represent standard deviation (polar, n = 235; temperate, n = 1469; tropical, n = 41).

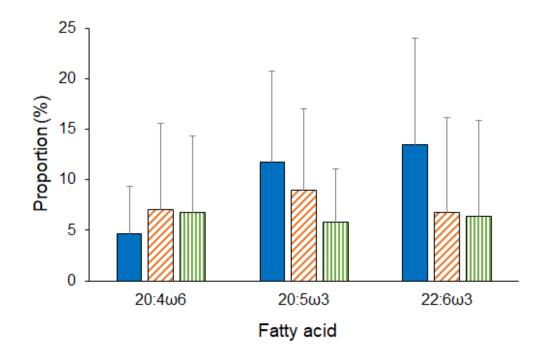
Fig. 3. Essential FA composition of deep-sea animals across latitudes. Mean proportions of
essential FA measured in the tissues of deep-sea animals from polar (blue bars), temperate
(orange diagonal striped bars), and tropical (green vertical striped bars) latitudes. Bars represent
standard deviation (polar, n = 176; temperate, n = 227; tropical, n = 11).

Fig. 4. Differences in terms of biochemical compositions among deep-sea animals from various habitats. Principal coordinate analysis plots representing differences in terms of isotopic (above) and essential FA composition (below) of deep-water species. In both cases, the variable 'habitat' resulted one of the most important factors, contributing 12 and 8% respectively to the variability in the biochemical composition of the deep-sea species.

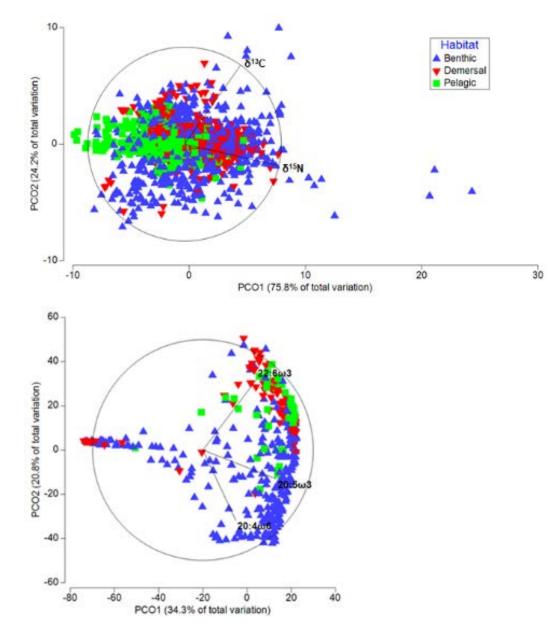




**Fig. 2.** 



**Fig. 3.** 



**Fig. 4**.