The stable carbon isotope signature of methane produced by saprotrophic fungi

Moritz Schroll[1]*, Frank Keppler[1,2], Markus Greule[1], Christian Eckhardt[3], Holger Zorn[4], Katharina Lenhart[5,6]*

[1]Institute of Earth Sciences, Heidelberg University, Im Neuenheimer Feld 236, 69120 Heidelberg, Germany
[2]Heidelberg Center for the Environment (HCE), Heidelberg University, 69120 Heidelberg, Germany
[3]Departement of Plant Ecology, Justus Liebig University Giessen, IFZ, 26 - 32, 35392 Giessen, Germany
[4]Institute of Food Chemistry & Food Biotechnology, Justus Liebig University Giessen, IFZ, 58, 35392 Giessen, Germany
[5]Centre for Organismal Studies (COS), Im Neuenheimer Feld 230, 69120 Heidelberg, Germany
[6]Bingen University of Applied Sciences, Berlinstraße 109, Bingen 55411, Germany

Correspondence to: Moritz Schroll (Moritz.Schroll@geow.uni-heidelberg.de) and Katharina Lenhart (k.lenhart@th-bingen.de)

Abstract. Methane (CH_4) is the most abundant organic compound in the atmosphere with emissions from many biotic and abiotic sources. Recent studies have shown that CH_4 production occurs under aerobic conditions in eukaryotes, such as plants, animals, algae and saprotrophic fungi. Saprotrophic fungi play an important role in nutrient recycling in terrestrial ecosystems via decomposition of plant litter. Although the CH_4 production by saprotrophic fungi has been reported, so far, no data for stable carbon isotope values of the emitted CH_4 ($\delta^{13}C$-CH_4 values) is available. In this study, we measured the $\delta^{13}C$ values of CH_4 and carbon dioxide ($\delta^{13}C$-CO_2 values) emitted by the two saprotrophic fungi *Pleurotus sapidus* (oyster mushroom) and *Laetiporus sulphureus* (sulphur shelf) cultivated, on three different substrates pine wood (*Pinus sylvestris*), grass (mixture of *Lolium perenne*, *Poa pratensis*, *Festuca rubra*) and corn (*Zea mays*), reflecting both C_3 and C_4 plants with distinguished bulk $\delta^{13}C$ values. Applying Keeling plots, we found that the $\delta^{13}C$ source values of CH_4 emitted from fungi cover a wide range from -40 mUr to -69 mUr depending on the growth substrate and fungal species. Whilst little apparent carbon isotopic fractionation (in the range of -0.3 mUr to 4.6 mUr) was calculated for $\delta^{13}C$ values of CO_2 released from *P. sapidus* and *L. sulphureus* relative to the bulk $\delta^{13}C$ values of the growth substrates, much larger carbon isotopic fractionations (ranging from -22 mUr to -42 mUr) were observed for the formation of CH_4. Although the two fungal species showed similar $\delta^{13}CH_4$ source values when grown on pine wood, $\delta^{13}CH_4$ source values differed substantially between the two fungal species when grown on grass or corn. We found that the source values of $\delta^{13}CH_4$ emitted by saprotrophic fungi are highly dependent on the fungal species and the metabolized substrate. The source values of $\delta^{13}CH_4$ cover a broad range and overlap with values reported for methanogenic archaea, thermogenic degradation of organic matter and other eukaryotes.

1 Introduction

Methane (CH$_4$) is an important greenhouse gas that is emitted by several abiotic sources (e.g. fossil fuel, biomass burning, geological processes) and biotic sources (e.g. wetlands, agriculture and waste, fresh waters) to the atmosphere (Kirschke et al., 2013; Saunois et al., 2016, 2019). In the past, biotic CH$_4$ production has been attributed exclusively to strictly anaerobic microorganisms, such as methanogens that are ubiquitous in wetlands, rice paddies, landfills and the intestines of termites and ruminants (Kirschke et al., 2013). The discovery of CH$_4$ emissions from dead and living plants under oxic conditions (Keppler et al., 2006, 2009) paved the way for the search of new biogenic CH$_4$ sources. Since then, several previously unknown CH$_4$ sources were discovered, including endothelial cells of rat liver (Boros and Keppler, 2019; Ghyczy et al., 2008), plant cell cultures (Wishkerman et al., 2011), marine algae (Klintzsch et al., 2019; Lenhart et al., 2016), marine and terrestrial cyanobacteria (Bižić et al., 2020), humans (Keppler et al., 2016) and saprotrophic fungi (Lenhart et al., 2012).

Fungi play a central role in ecosystems by decomposing organic matter and thereby recycling formerly bound carbon and nutrients (Grinhut et al., 2007). This process is especially important in forests where fungi are essential for wood decay and have a great impact on the carbon and nitrogen cycles in these environments (Ralph and Catcheside, 2002). White rot fungi (e.g. *Trametes versicolor* or *Pleurotus ostreatus*) are able to decompose the chemically complex structural component lignin, whereas brown rot fungi (e.g. *Serpula lacrymans* or *Gloeophyllum trabeum*) mainly metabolize cellulose and hemicellulose (Ten Have and Teunissen, 2001; Leonowicz et al., 1999; Valášková and Baldrian, 2006). Fungi have already been determined to be involved in the synthesis of CH$_4$ during wood decay (Beckmann et al., 2011; Mukhin and Voronin, 2007, 2008) by breakdown of large macromolecules to smaller molecules, thereby providing bacteria and archaea with essential substrate. Elevated levels of CH$_4$ were found in fungus-infected wood stems with oxygen concentrations ranging from 1 to 14 % (Hietala et al., 2015). Here, CH$_4$ production was associated with anoxic microsites in the xylem, indicating that at least part of the CH$_4$ was produced by methanogenic archaea. Nevertheless, Lenhart et al., 2012 demonstrated that basidiomycetes are able to produce CH$_4$ under aerobic conditions without the presence of methanogenic archaea. Therefore, fungi might be an underestimated source of CH$_4$ in the global CH$_4$ cycle.

Stable carbon isotopes (expressed as δ^{13}C values) have often been used to investigate sources and sinks of CH$_4$ on the global scale (Whiticar, 1993). As different CH$_4$ sources have characteristic δ^{13}C values, δ^{13}C-CH$_4$ values might be used to quantify the individual contributions of various sources regionally and/or globally (Dlugokencky et al., 2011; Hein et al., 1997; Nisbet et al., 2016; Quay et al., 1999; Tyler, 1986; Whiticar, 1999). The short lifetime of CH$_4$ in the atmosphere (range from 9.7 $\pm$ 1.5 to 11.2 $\pm$ 1.3 years) (Naik et al., 2013; Prather et al., 2012; Voulgarakis et al., 2013) assures that global isotopic δ^{13}C-CH$_4$ patterns represent the average of recent inputs by various sources and allows the quantification of respective source strengths (Mikaloff Fletcher et al., 2004b, 2004a).

Additionally, stable isotopes provide information about the formation processes of CH$_4$. Traditionally, three formation categories of δ^{13}C-CH$_4$ values have been identified: biogenic, with typical δ^{13}C-CH$_4$ values ranging from ~-55 mUr to -70 mUr, thermogenic (ranging from ~-25 mUr to -55 mUr) and pyrogenic (ranging from ~-13 mUr to -25 mUr) (Kirschke et al.,

65 2013). However, stable isotope values of recently identified CH_4 sources, i.e. human CH_4 emissions (-56 mUr to -95 mUr) (Keppler et al., 2016), plant-derived CH_4 (-52 mUr to -69 mUr) (Keppler et al., 2006), and abiotic UV-induced CH_4 formation by plants (-52 mUr to -67 mUr) (Vigano et al., 2009) also need to be considered.

In this study, we investigated the stable carbon isotope source signatures of CH_4 and CO_2 released by the two basidiomycetes *Pleurotus sapidus* (white rot fungus) and *Laetiporus sulphureus* (brown rot fungus). Both fungi were cultivated under sterile
70 conditions on three different substrates (pine wood, grass, and corn) with varying bulk $\delta^{13}C$ values. We examined the influence of fungal species and growth substrate on $\delta^{13}C\text{-}CH_4$ and $\delta^{13}C\text{-}CO_2$ values and compared the $\delta^{13}C\text{-}CH_4$ values from the two fungal species with those of other known sources reported from the literature.

2 Material and Methods

2.1 Selected fungi

75 *P. sapidus* (Pleurotaceae, DSMZ 8266) and *L. sulphureus* (Polyporacaeae, DSMZ 1014) were chosen for this experiment because of their capability to emit CH_4 (Lenhart et al., 2012), their ecological and physiological characteristics (*P. sapidus* is a white rot fungus and *L. sulphureus* is a brown rot fungus) and well-established practical handling under laboratory conditions.

2.2 Cultivation of fungi and incubation experiments

Pine wood (*Pinus sylvestris*), grass (mixture of *Lolium perenne*, *Poa pratensis*, *Festuca rubra*) and corn (*Zea mays*) were
80 selected as growth substrates. Pine wood was chosen to investigate if white rot and brown rot fungi differ in $\delta^{13}C\text{-}CH_4$ and $\delta^{13}C\text{-}CO_2$ values released during wood decay. Therefore, dead pine wood branches were collected from the forest floor and shredded to small wood chips with a length of about 5 cm (Natura 1800L; Glora, Witten, Germany). The wood chips were dried at 60°C for 48h and stored in a flask (Weck, Hanau, Germany). Grass (C_3 plant) and corn (C_4 plant) were selected because of their different stable isotope values. As the metabolic pathway for carbon fixation is biochemically different in C_3 and C_4
85 plants, plant biomass differs in $\delta^{13}C$ values, which in turn might lead to different $\delta^{13}C$ values of CH_4 and CO_2 released by fungi. Therefore, typical garden lawn was manually cut, dried at 70 °C, and stored in a flask. The corn substrate consisted of conventional corn flour.

The substrates were autoclaved and filled into 2.7 l flasks (Weck, Hanau, Germany) and inoculated with pure fungal submerged cultures under sterile conditions according to Lenhart et al., 2012. After addition of the fungi, the flasks were closed with lids
90 and a rubber band sealing. To allow gas exchange during the growth time of the fungi (about two weeks), a hole in the centre of every lid was fitted with a cotton stopper. Before the start of the incubation experiments, the flasks were aerated under sterile conditions in order to start the incubation at atmospheric CH_4 mixing ratios. Additionally, to seal the flasks airtight the cotton stoppers were replaced by sterile silicone stoppers (Saint-Gobain Performance Plastics, Charny, France).

For the incubation experiments, *P. sapidus* und *L. sulphureus* were incubated on the three substrates, while substrates were
95 incubated as control treatments. Before the incubation experiments, the substrates were sterilized by autoclaving at 121 °C and

2 bar pressure for 20 minutes. The incubation experiments were conducted as three replicates per treatment. The duration of the incubation accounted for up to 40 h. All incubations were conducted at room temperature (23 ± 1.5 °C). At every sampling point, 40 ml air was taken from the flasks for gas concentration measurements and an additional 40 ml were taken for δ^{13}C-CH_4 stable isotope ratio mass spectrometry (IRMS) analysis. The gas samples were taken with airtight 60 ml PE syringes (Plastipak, BD, Franklin Lakes, USA) and transferred into 12 ml evacuated Exetainers (Labco, High Wycombe, UK). Subsequently a volume of atmospheric air equivalent to the volume of the removed sample was added into each flask directly after sampling. Mixing ratios and stable isotope values of CH_4 were corrected according to the dilution.

When calculating the fungal CH_4 and CO_2 production rates, we subtracted substrate derived CH_4 and CO_2 production rates (determined in the control treatments) from the respective fungi containing samples. Additionally, for the calculation of the fungal production rates only sample points showing a linear increase in CH_4 and CO_2 were taken into account.

To account for differences in the metabolic activity of the fungi, we additionally measured respiration rates, assuming that metabolic activity correlates with respiration and therefore CO_2 emissions of the fungi. Hence, we related fungal derived CH_4 emissions to respiration by calculating the $CH_4 : CO_2$ emission ratio.

2.3 Analysis of CH_4 and CO_2 via gas-chromatography

Samples were analysed using a gas chromatograph (GC, Bruker Greenhouse Gas Analyser 450-GC) equipped with a flame ionization detector (FID) and an electron capture detector (ECD) for the detection of CH_4 and CO_2, respectively. The detector temperatures were set at 300 °C (FID) and 350 °C (ECD). Five reference gases (Deuste Steininger GmbH) were used for calibrating the GC-system. The reference gases were in the range of 1 parts per million by volume (ppmv) to 21 ppmv and 304 ppmv to 40,000 ppmv for CH_4 and CO_2, respectively. Gas peaks were integrated using Galaxie software (Varian Inc., Palo Alto, CA, USA).

2.4 Definition of δ values and isotope apparent fractionation

In this paper, all stable carbon isotope ratios are expressed in the conventional 'delta' δ notation, meaning the relative difference of the isotope ratio of a substance compared to the standard substance Vienna Peedee Belemnite (V-PDB) (Eq. (1)).

$$\delta^{13}C = \frac{\left(\frac{^{13}C}{^{12}C}\right)_{sample}}{\left(\frac{^{13}C}{^{12}C}\right)_{V\text{-}PDB}} - 1 \tag{1}$$

The apparent fractionation (ε_{app}) between fungal δ^{13}C-CH_4 or δ^{13}C-CO_2 values and the δ^{13}C values of the substrates was calculated according to Eq. (2).

$$\varepsilon_{app\ CH4\ or\ CO2} = \frac{(\delta^{13}C + 1)_{fungal\ CH_4\ or\ CO_2}}{(\delta^{13}C + 1)_{substrate}} - 1 \tag{2}$$

We follow the proposal of Brand and Coplen, 2012 and use the term 'urey' (Ur) as the isotope delta unit, in order to conform
with the guidelines for the International System of Units (SI). Hence, isotope delta values that were formerly given as -70 ‰,
are expressed as -70 mUr.

2.5 Measurements of $\delta^{13}CH_4$ and $\delta^{13}CO_2$ values

Stable carbon isotope values of CH_4 and CO_2 were measured using a continuous flow isotope mass spectrometry system (CF-
IRMS). A HP 6890N GC (Agilent, Santa Clara, USA) was linked to a preconcentration unit for CH_4 measurements and an
autosampler A200S (CTC Analytics, Zwingen, Switzerland) for CO_2 analysis. The GC was equipped with a CP-PoraPLOT Q
capillary column (Varian, Palo Alto, USA) (27,5 m x 0.25 mm i.d., film thickness 8 µm). The GC was operated with an injector
temperature of 200°C, isothermal oven temperature of 30°C, split injection (10:1) and a constant carrier gas flow of 1.8 ml
min^{-1} (methane-free helium). The GC was coupled to a Delta^{PLUS}XL isotope ratio mass spectrometer (ThermoQuest Finnigan,
Bremen, Germany) via an oxidation reactor and a GC Combustion III Interface (ThermoQuest Finnigan, Bremen, Germany).
The oxidation reactor was employed with the following properties: ceramic tube (Al_2O_3), length 320 mm, 1.0 mm i.d., with
Ni/Pt wires inside activated by oxygen, reactor temperature 960 °C.

For CH_4 measurements with the preconcentration unit, headspace gas samples were transferred to an evacuated 40 ml sample
loop. Methane was trapped on Hayesep D, separated from other compounds by the GC and then introduced into the IRMS
system via an open split. The monitor gas was carbon dioxide of high purity (carbon dioxide 4.5, Messer Griesheim, Frankfurt,
Germany) with a known $\delta^{13}C$ value of -23.6 mUr (calibrated at MPI for Biogeochemistry in Jena, Germany). All $\delta^{13}C$ values
were corrected using two CH_4 reference standards (Isometric instruments, Victoria, Canada) with $\delta^{13}C$ values of -23.9 ± 0.2
mUr and -54.5 ± 0.2 mUr that were calibrated against IAEA and NIST reference substances. The normalization of the sample
values was done according to Paul et al., 2007.

2.6 Bulk isotope analysis of fungal substrates

Stable carbon isotope values of the bulk substrate were measured using an Elemental Analyzer Flash EA 11112 (Thermo
Fischer Scientific, Germany) coupled to a Delta V IRMS (Thermo Fischer Scientific, Germany). Therefore, 0.06 mg of the
substrate was put into a tin cup and combusted in the Elemental Analyzer. The resulting gases were separated in a GC by a
CP-PoraPLOT Q capillary column (Varian, Palo Alto, USA) (27,5 m x 0.25 mm i.d., film thickness 8 µm) and then reached
the Delta V IRMS via a Conflo IV Universal Continuous Flow Interface (Thermo Fischer Scientific, Germany). Isotope values
were corrected using USGS 40 and USGS 41 standards.

2.7 Determination of isotopic source signature of CH_4 and CO_2 applying Keeling plots

For the determination of $\delta^{13}C$ source values of CH_4 and CO_2 the Keeling plot method was used (Keeling, 1958; Pataki et al.,
2003) (Eq. (3)):

$$\delta^{13}C_a = c_b(\delta^{13}C_b - \delta^{13}C_s)\left(\frac{1}{c_a}\right) + \delta^{13}C_s \tag{3}$$

where c_a is the mixing ratio of CH_4/CO_2 in the headspace, $\delta^{13}C_a$ is the $\delta^{13}C$ value of CH_4/CO_2 in the headspace, c_b is the mixing ratio of background CH_4/CO_2, $\delta^{13}C_b$ is the $\delta^{13}C$ value of background CH_4/CO_2 and $\delta^{13}C_s$ the $\delta^{13}C$ source value of the CH_4/CO_2. For a more detailed description of the application of Keeling plots for determination of CH_4 source signature we refer to the study by Keppler et al., 2016.

$\delta^{13}C$-CH_4 source signatures were calculated after the Keeling plot method for each flask. Results of the Keeling plots are then given as the arithmetic mean of the three individual flasks per treatment with standard deviations (n=3).

$\delta^{13}C$-CH_4 source signatures of each flask of *P. sapidus* and *L. sulphureus* grown on pine were corrected for CH_4 emissions and $\delta^{13}C$-CH_4 values of the "pine" control samples using the following mass balance approach (Eq. (4)).

$$\delta^{13}C_{\text{fungi corrected}} = \frac{\left(P(CH_4)_{\text{fungi}} * \delta^{13}C_{\text{fungi}}\right) - \left(P(CH_4)_{\text{pine}} * \delta^{13}C_{\text{pine}}\right)}{\left(P(CH_4)_{\text{fungi}} - P(CH_4)_{\text{pine}}\right)} \tag{4}$$

, where $P(CH_4)_{\text{fungi/pine wood}}$ is the CH_4 emitted by the fungi or pine wood and $\delta^{13}C_{\text{fungi/pinewood}}$ is the $\delta^{13}C$-CH_4 source signature of the fungi or pine wood derived from Keeling plots. Corrected $\delta^{13}C$-CH_4 source values for *P. sapidus* and *L. sulphureus* are given as the arithmetic mean of the three individual flasks per treatment with standard deviations (n=3).

The determination coefficient (R^2) of the Keeling plots showed values higher than 0.93, except for *P. sapidus* grown on grass ($R^2=0.51$). The lower R^2 value for *P. sapidus* grown on grass is probably a result of the marginal changes of $\delta^{13}C$-CH_4 values due to the small increase of the CH_4 mixing ratio compared to the background CH_4 mixing ratio. Therefore, the low R^2 does not necessarily indicate a weaker relationship between CH_4 mixing ratio and $\delta^{13}C$-CH_4.

2.8 Statistics

Mixing ratios and production rates of CH_4, CO_2, $\delta^{13}C$-CH_4 and $\delta^{13}C$-CO_2 values and $\delta^{13}C$ source values are presented as arithmetic mean of three independent replicates with standard deviations (SD; n = 3). Linear regression analysis, arithmetic means and SDs were calculated using Excel (Microsoft Excel for Office 365 MSO). Two-way analysis of variance (ANOVA) (SigmaPlot 12.2.0.45, USA) were carried out to test for "species" and "substrate" related effects on $\delta^{13}C$-CH_4 and $\delta^{13}C$-CO_2 source values for each treatment. Differences at the $p < 0.05$ level were referred to as significant.

3 Results and Discussion

In this section, we firstly present the results of CH_4 and CO_2 production from the two fungal species grown on the three different substrates. This includes emission rates of CH_4 and CO_2 from the control treatments of pine wood, grass and corn as well as the molar ratio of CH_4 and CO_2. Secondly, we then present the respective stable isotope values measured for CH_4 and CO_2 during the incubation experiments and calculate the stable isotope source values of CH_4 and CO_2 released by the fungi applying Keeling plots. We then compare these values with stable carbon isotope values of the bulk organic matter by

calculating the apparent fractionation. Finally, we compare $\delta^{13}C$ source values of fungal derived CH_4 with values known for other CH_4 sources from the literature.

3.1 Release of CH_4 and CO_2 from *P. sapidus* and *L. sulphureus*
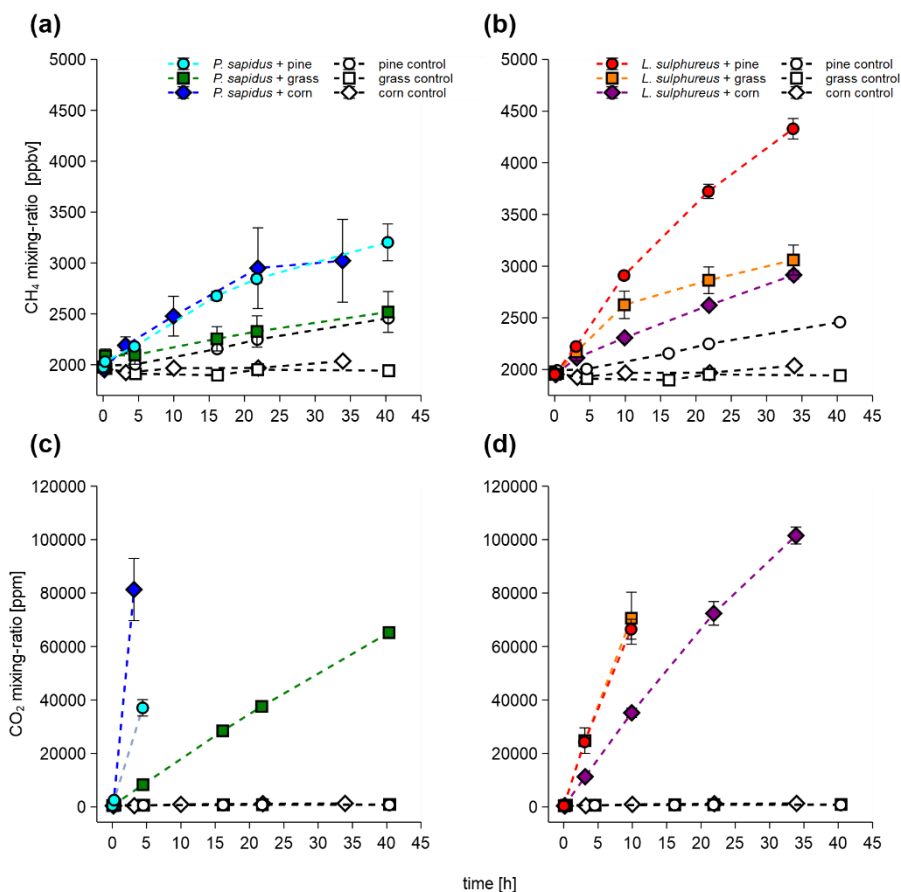
Figure 1: Mixing ratios of CH_4 and CO_2 of *P. sapidus* (a, c) and *L. sulphureus* (b, d) grown on pine wood, grass, and corn. Mixing ratios are presented as mean values with standard deviation SD (n=3).

All incubation experiments in which fungi were grown on different substrates showed a significant increase in CH_4 compared to the respective substrate control (Fig. 1 a, c). Calculated emission rates for CH_4 and CO_2 are presented in Table 1. *L. sulphureus* grown on grass (7.5 ± 1.3 nmol h^{-1}) showed the highest emission rate of CH_4, followed by *L. sulphureus* grown on pine (6.2 ± 0.3 nmol h^{-1}), *P. sapidus* grown on corn (4.4 ± 1.9 nmol h^{-1}), *L. sulphureus* grown on corn (2.6 ± 0.1 nmol h^{-1}), *P. sapidus* grown on pine (2.5 ± 0.2 nmol h^{-1}) and *P. sapidus* grown on grass (1.4 ± 0.5 nmol h^{-1}). Please note that CH_4 and CO_2 emission rates are not related to fungal biomass. Therefore, differences in the emission rates might be due to varying fungal biomass of the subsamples. Instead, CH_4 production was related to CO_2 production by determining the molar emission ratio

between CH_4 and CO_2 (µmol CH_4 : mol CO_2). CO_2 production thereby reflects the amount of fungal biomass and is also an indicator for the metabolic activity of the fungi.

The control flasks did not show significant changes in their CH_4 and CO_2 mixing ratios over time, except for CH_4 in pine wood controls (1.3 ± 0.1 nmol h^{-1}). However, in the control flasks of corn small CH_4 emission rates of 0.25 ± 0.01 nmol h^{-1} were observed and in the control 'grass' the CH_4 mixing ratio slightly decreased over time (-0.05 ± 0.04 nmol h^{-1}). Whilst the pine wood and corn control flasks showed a small increase in the CH_4 mixing ratio, they did not show an increase in CO_2 mixing ratios. These data rule out a contamination by microbial heterotrophs, as this would cause a measurable CO_2 increase within the flasks. The CH_4 increase in the substrate controls might be attributed to CH_4 release by dead plant material as it was already shown by Keppler et al., 2006 and Vigano et al., 2009. Within the scope of these experiments, no analytic test for microbial contamination was conducted. Nevertheless, Lenhart et al., 2012 clearly showed that with the performed method of cultivation of fungi and incubation experiments no methanogenic archaea were present, using three different methods (Fluorescence in situ hybridization (FISH), confocal laser scanning microscopy (CLSM) and quantitative real time PCR). Furthermore, CH_4 and CO_2 release and the CH_4 : CO_2 emission ratios in our incubations are similar to the experiments of Lenhart et al., 2012 and do not indicate microbial contamination. Therefore, we assume that in our investigations no contamination with bacteria or methanogenic archaea was present.

For *P. sapidus* grown on corn and *L. sulphureus* grown on grass, no further linear increase in CH_4 was observed after 22 h and 10 h, respectively. This might be due to a reduced decay of organic matter and slower fungal metabolism because of higher CO_2 and lower O_2 mixing ratios.

A drastic increase in CO_2 mixing ratios relative to the controls was observed in all flasks containing fungi (Fig. 1 b, d). The CO_2 emission rates are shown in Table 1. CO_2 production rates ranged from 176 ± 4 µmol h^{-1} to 2910 ± 410 µmol h^{-1} for *P. sapidus* grown on grass and *P. sapidus* grown on corn, respectively. These highly variable CO_2 production rates might reflect different fungal biomass and metabolic activity (mineralisation of organic matter). In the control treatments, tiny increases in the CO_2 mixing ratio were detected ranging from 0.64 ± 0.12 µmol h^{-1} to 0.91 ± 0.14 µmol h^{-1}. Only one flask (corn control) showed a somewhat higher increase in CO_2 (7.76 µmol h^{-1}), which is most likely caused by microbial contamination of the flask. However, no increase in the CH_4 mixing ratio was detected (see supplementary material). Therefore, this control flask was excluded from further calculations.

Mean CH_4 and CO_2 emission rates and CH_4 : CO_2 emission ratios of all treatments are presented in Table 1. Higher ratios indicate a higher CH_4 production during decay of the substrates. Thereby, both fungal species and substrate affect the CH_4 : CO_2 emission ratio ($p<0.001$). For *P. sapidus,* CH_4 : CO_2 emission ratios are more variable (1.4 to 8.0 µmol CH_4/mol CO_2) compared to *L. sulphureus* (6.7 – 9.6 µmol CH_4/mol CO_2). This variation might be due to differences in the fungi's enzyme sets required for organic matter decay, as *P. sapidus* is a white rot fungus and *L. sulphureus* is a brown rot fungus. At present the biochemical pathways that lead to CH_4 are still unknown, although compounds such as the sulphur-bound methyl-group of methionine and glucose have been identified to act as carbon precursors of fungal-derived CH_4 (Lenhart et al., 2012).

Lenhart et al., 2012 found $CH_4 : CO_2$ ratios of fungi that ranged between 8 µmol CH_4/mol CO_2 and 17 µmol CH_4/mol CO_2, which is in the same order of magnitude as the $CH_4 : CO_2$ ratios determined in this study. It should be noted that $CH_4 : CO_2$ ratios of Lenhart et al., 2012 were given in ppbv CH_4 : % CO_2 and for better comparability $CH_4 : CO_2$ ratios were converted to fit the units used in this study (µmol CH_4 : mol CO_2).

235

Table 1: CH_4 and CO_2 production rates and molar $CH_4 : CO_2$ emission ratios of the fungi incubated on different substrates. Values are presented as mean values of three independent replicates with SD (n = 3), except for the control "corn" (n=2).

Fungi	Substrate	CH_4 production rate [nmol h^{-1}]	CO_2 production rate [µmol h^{-1}]	$CH_4 : CO_2$ ratio [µmol/mol]
P. sapidus	pine	2.5 ± 0.2	901 ± 79	2.8 ± 0.4
	grass	1.4 ± 0.5	176 ± 4	8.0 ± 2.8
	corn	4.4 ± 1.9	2910 ± 419	1.4 ± 0.5
L. sulphureus	pine	6.2 ± 0.3	724 ± 42	8.6 ± 1.0
	grass	7.5 ± 1.3	771 ± 103	9.6 ± 0.5
	corn	2.6 ± 0.1	385 ± 20	6.7 ± 0.4
control	pine	1.3 ± 0.1	0.64 ± 0.12	-
	grass	-0.05 ± 0.04	0.91 ± 0.14	-
	corn	0.25	0.66	-

3.2 Stable carbon isotope values of CH₄ and CO₂
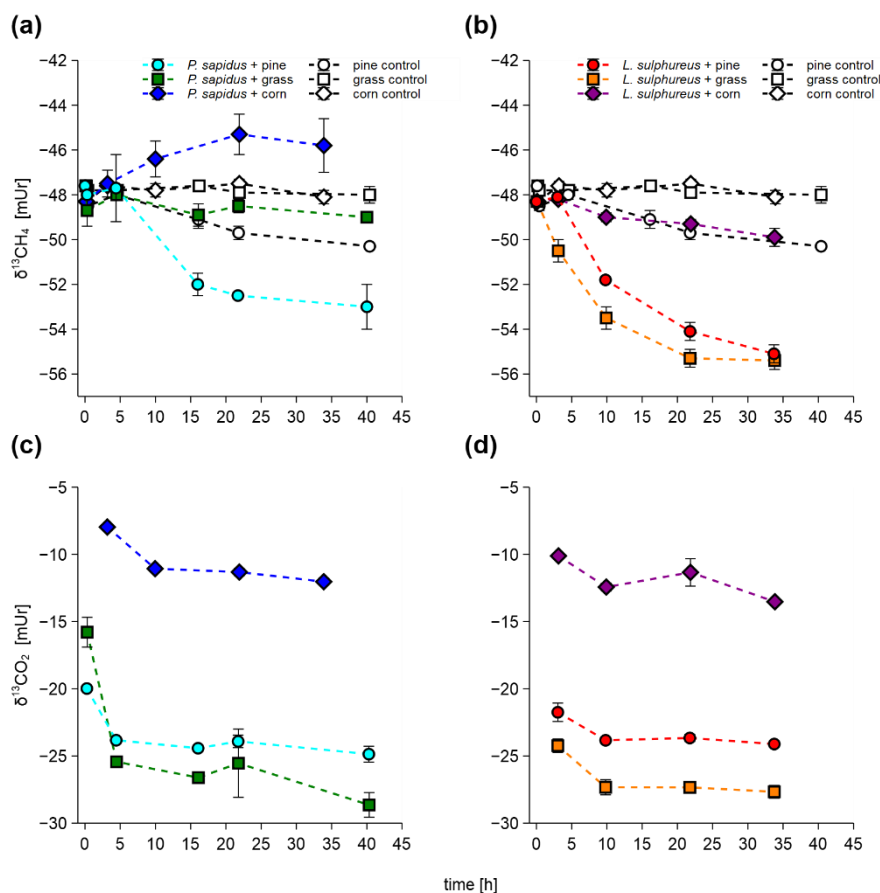
Figure 2: Stable carbon isotope values of CH_4 and CO_2 of *P. sapidus* (a, c) and *L. sulphureus* (b, d) grown on pine, grass, and corn. Values are presented as mean values with SD (n=3), except for $\delta^{13}CO_2$ values of *L. sulphureus* grown on corn (n=2).

Stable carbon isotope values of CH_4 and CO_2 measured from the incubation experiments are presented in Fig. 2. All incubations show a trend towards more negative $\delta^{13}C$-CH_4 values (less ^{13}C) with time except for *P. sapidus* grown on corn, where a tendency towards more positive $\delta^{13}C$-CH_4 values was observed (Fig. 2 a, b). During the incubation, $\delta^{13}C$-CH_4 values changed from -47.7 ± 0.1 mUr (for incubation of *P. sapidus* grown on pine/grass) and -48.2 ± 0.1 mUr (for incubation of *P. sapidus* grown on corn and *L. sulphureus* grown on pine/grass/corn) to -53.0 ± 0.7 mUr (*P. sapidus* grown on pine), -48.7 ± 0.3 mUr (*P. sapidus* grown on grass), -45.8 ± 1.2 mUr (*P. sapidus* grown on corn), -55.1 ± 0.4 mUr (*L. sulphureus* grown on pine), -55.4 ± 0.4 mUr (*L. sulphureus* grown on grass) and -49.9 ± 0.4 mUr (*L. sulphureus* grown on corn). The controls showed no significant shift in $\delta^{13}C$-CH_4 values except for the control "pine", where an increase in the CH_4 mixing ratio along with more

negative values of δ^{13}C-CH$_4$ values occurred over time. This was accounted for when calculating the δ^{13}C-CH$_4$ source signatures for *P. sapidus* grown on pine and *L. sulphureus* grown on pine (see materials and methods 2.7).

The δ^{13}C-CO$_2$ values showed a trend towards more negative values within the first three to four hours of incubation (Fig. 2 c, d). After this time only minor changes of the δ^{13}C-CO$_2$ values occurred. Final δ^{13}C-CO$_2$ values of the incubation were -24.9 ± 0.6 mUr (*P. sapidus* grown on pine), -28.6 ± 0.9 mUr (*P. sapidus* grown on grass), -12.0 ± 0.3 mUr (*P. sapidus* grown on corn), -24.1 ± 0.1 mUr (*L. sulphureus* grown on pine), -27.7 ± 0.5 mUr (*L. sulphureus* grown on grass) and -13.0 ± 0.5 mUr (*L. sulphureus* grown on corn).

Table 2: Calculated δ^{13}C-CH$_4$ and δ^{13}C-CO$_2$ source signatures, δ^{13}C values of the substrates, and $\varepsilon_{app\,CH4}$ and $\varepsilon_{app\,CO2}$. Values are presented as mean values with the SD (n=3).

Fungi	Substrate	δ^{13}C-CH$_4$ source [mUr]	δ^{13}C-CO$_2$ source [mUr]	δ^{13}C substrate [mUr]	$\varepsilon_{app\,CH4}$ [mUr]	$\varepsilon_{app\,CO2}$ [mUr]
P. sapidus	pine	-65.3 ± 1.1	-24.1 ± 0.1		-38.4 ± 1.2	4.0 ± 0.1
	grass	-52.9 ± 1.6	-27.4 ± 1.3		-21.8 ± 1.7	4.6 ± 1.3
	corn	-39.8 ± 2.0	-12.0 ± 0.3		-28.5 ± 2.0	-0.3 ± 0.3
L. sulphureus	pine	-61.4 ± 0.5	-25.0 ± 0.5		-34.4 ± 0.6	3.0 ± 0.4
	grass	-69.2 ± 1.9	-29.0 ± 0.5		-38.6 ± 2.0	2.9 ± 0.5
	corn	-53.4 ± 1.1	-12.8 ± 0.3		-42.2 ± 1.1	-1.1 ± 0.3
control	pine			-28.0 ± 0.5		
	grass			-31.5 ± 0.6		
	corn			-11.7 ± 0.1		

3.3 δ^{13}C-CH$_4$ and δ^{13}C-CO$_2$ source signatures of fungi
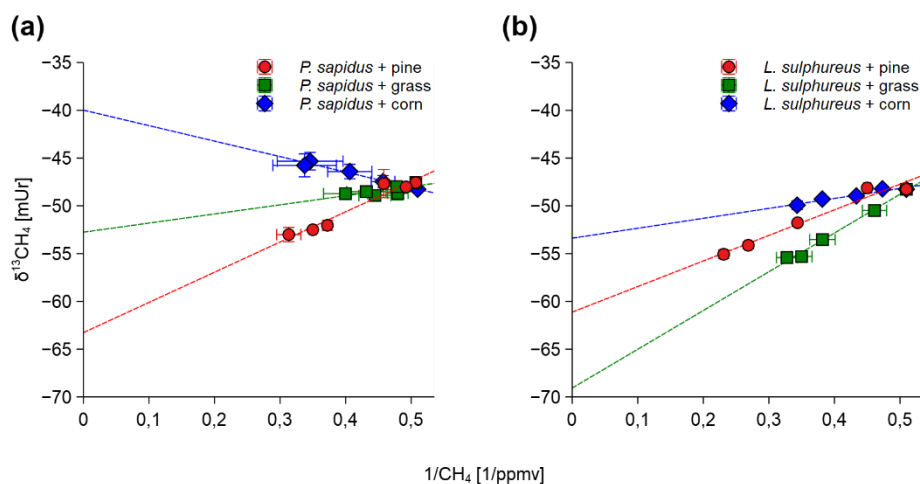
Figure 3: Keeling plots shown for *P. sapidus* (a) and *L. sulphureus* (b) grown on three substrates. Sample points in the graphs are given as the arithmetic mean of δ^{13}C-CH$_4$ or δ^{13}C-CO$_2$ values with SD (n=3) on the y-axis and the arithmetic mean of the
265   inverted mixing ratio of CH$_4$ or CO$_2$ with SD (n=3) on the x-axis.

The δ^{13}C-CH$_4$ source signatures determined via a Keeling plot analysis (Fig. 3) that range from -69.2 ± 1.9 mUr (*L. sulphureus* grown on grass) to -39.8 ± 2.0 mUr (*P. sapidus* grown on corn) are presented in Table 2. Average δ^{13}C-CH$_4$ source signatures for each fungal species, considering all three substrates, are -52.6 mUr for *P. sapidus* and -61.3 mUr for *L. sulphureus*. These
270   results suggest that the fungal species significantly influence the isotopic values of the emitted CH$_4$ (*p*<0.001). A possible explanation for this observation could be the different enzyme sets of both fungi decomposing different components of the growth substrates, as *P. sapidus* belongs to white rot fungi and *L. sulphureus* is a brown rot fungus. However, detailed investigations of the metabolic pathways leading to CH$_4$ formation were beyond the scope of this study.

Furthermore, a significant effect of the growth substrate on δ^{13}C-CH$_4$ source signatures was observed (*p*<0.001). δ^{13}C-CH$_4$
275   source signatures by *P. sapidus* were more positive compared to those of *L. sulphureus* when grown on grass (Δ=16.3 mUr) and corn (Δ=13.6 mUr) (Fig. 4). When grown on pine wood, δ^{13}C-CH$_4$ source signatures were similar with *P. sapidus* showing slightly more negative values (Δ=-3.9 mUr). Methane emitted by both fungi grown on corn was generally more enriched in ^{13}C (less negative δ^{13}C-CH$_4$ source values) compared to the fungi grown on pine wood and grass. This might be easily explained by the δ^{13}C values of the growth substrates corn (-11.7 mUr, typical for C$_4$-plants) being roughly 20 mUr less
280   negative in their δ^{13}C values compared to the C$_3$-plants pine wood (-28.0 mUr) and grass (-31.5 mUr).
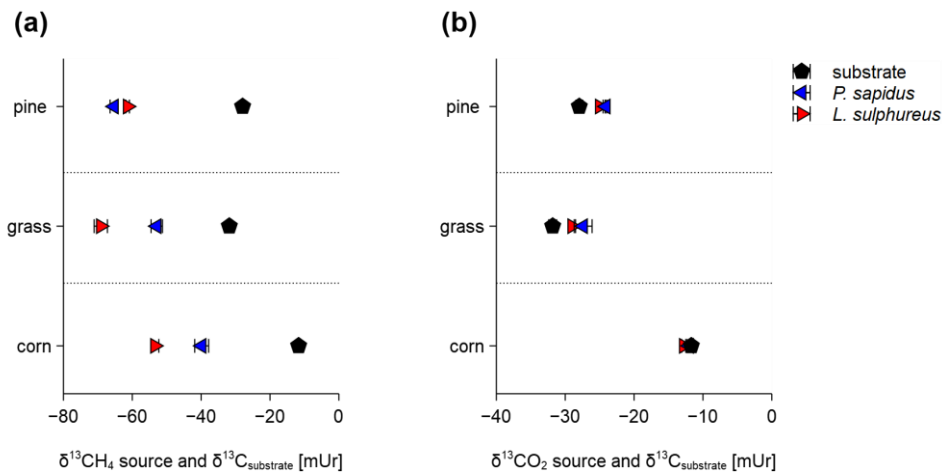
Figure 4: Calculated source signatures of $\delta^{13}C$-CH_4 values (a) and $\delta^{3}C$-CO_2 values (b) by *P. sapidus*, *L. sulphureus* and the $\delta^{13}C$ values of the substrate. The data points represent the mean values of the individual Keeling plots with SD (n=3).

285   Comparison of calculated $\delta^{13}C$-CH_4 source signatures with measured bulk $\delta^{13}C$ values of the substrates shows that CH_4 emitted by both fungi is generally depleted in ^{13}C compared to the respective substrates (Fig. 4a). Based on this data we further calculated the apparent fractionation ($\varepsilon_{app\,CH4}$) between the $\delta^{13}C$-CH_4 source signatures and the bulk $\delta^{13}C$ values of the growth substrates. The apparent fractionation was calculated as up to the present no metabolic pathway for the formation of CH_4 in fungi is known and therefore currently only the initial $\delta^{13}C$ signatures of the substrates and the calculated $\delta^{13}C$-CH_4 source

290   signatures of the fungi can be compared. The values of $\varepsilon_{app\,CH4}$ that range from -21.8 mUr (*P. sapidus* grown on grass) to -42.2 mUr (*L. sulphureus* grown on corn) are presented in Table 2. When grown on pine wood $\varepsilon_{app\,CH4}$ values are similar for *P. sapidus* (-38.4 ± 1.2 mUr) and *L. sulphureus* (-34.4 ± 0.6 mUr). The differences in $\varepsilon_{app\,CH4}$ values between both fungal species are distinct when grown on grass (*P. sapidus*: -21.8 ± 1.7 mUr, *L. sulphureus:* -38,6 ± 2.0 mUr) and corn (*P. sapidus*: -28,5 ± 2.0 mUr, *L. sulphureus*: -42.2 ± 1.1 mUr).

295   The calculated $\delta^{13}C$-CO_2 source signatures of both fungi (Table 2) range from -29.0 ± 0.5 mUr (*L. sulphureus* grown on grass) to -12.0 ± 0.3 mUr (*P. sapidus* grown on corn*)*. $\delta^{13}C$-CO_2 source signatures are in a similar range for both fungi for all three substrates. However, CO_2 emitted by *L. sulphureus* is slightly more depleted in ^{13}C for all three substrates compared to *P. sapidus*. Hence, an effect of fungal species on the stable carbon isotope values of CO_2 is significant (p=0.008). Also, the used substrates were found to influence $\delta^{13}C$-CO_2 values significantly (p<0.001).

300   The $\delta^{13}C$-CO_2 source signatures of the fungi show only small deviations from the bulk $\delta^{13}C$ values of the respective substrates (Fig. 4b). However, for both fungi grown on pine wood and grass, $\delta^{13}C$-CO_2 values are slightly less negative (a few mUr) compared to the bulk substrate. This observation is rather unexpected, as $\delta^{13}C$-CO_2 values are usually more negative with respect to $\delta^{13}C$ values of growth substrates due to fractionation during the metabolism (Bowling et al., 2008). However, when

grown on corn δ^{13}C-CO$_2$ source signatures by both fungi are more negative compared to the substrate and calculated $\varepsilon_{app\ CO2}$

305   values (Table 2) are -1.1 ± 0.3 mUr and +4.6 ± 1.3 mUr for *L. sulphureus* grown on corn and *P. sapidus* grown on grass, respectively.

The results of the incubation experiments show that there are distinct differences in the δ^{13}C-CH$_4$ and δ^{13}C-CO$_2$ values released by both fungi. While the δ^{13}C-CO$_2$ source signatures are similar to the δ^{13}C values of the substrate (with $\varepsilon_{app\ CO2}$ values up to 4.6 mUr), the δ^{13}C-CH$_4$ source signatures deviate strongly from the respective substrate, with $\varepsilon_{app\ CH4}$ values of up to -42.2

310   mUr. This either indicates that metabolic pathways leading to the formation of CH$_4$ and CO$_2$ have different fractionation and/or that fungal CH$_4$ and CO$_2$ are derived from different precursor compounds of the respective substrate. The growth substrates used for this study (pine wood, grass, corn) contain distinct amounts of cellulose, hemicellulose, lignin and other compounds at different proportions (in contrast to only using pure glucose or cellulose as growth substrate). Hence, the δ^{13}C-CH$_4$ and δ^{13}C-CO$_2$ source signatures depend on the specific metabolic pathways used by the fungal species as well as the chemical

315   composition of the growth substrate. The selected fungi and used growth substrates provide a first solid basis for the potential range of δ^{13}C-CH$_4$ values that might occur in nature.

3.4 Fungal δ^{13}C-CH$_4$ values compared with known CH$_4$ sources

Figure 5 compares the δ^{13}C-CH$_4$ values emitted by fungi in relation to other known CH$_4$ sources in the environment that have been reported from the literature. The red bars indicate typical biogenic (formerly only considered to be produced by archaea)

320   CH$_4$ sources with emissions from wetlands, ruminants, landfills and rice paddies where δ^{13}C-CH$_4$ values are usually ranging from -85 mUr to -40 mUr. Abiotic CH$_4$ sources (including thermogenic or pyrolytic processes) stemming from natural gas, coal mining and biomass burning are characterized by less negative δ^{13}C values usually ranging from -55 mUr to -20 mUr. In addition gas hydrates which might be formed by both microbial and abiotic processes cover a wide range of δ^{13}C values (-29 mUr to -73 mUr), depending on its forming mechanisms (Kvenvolden, 1995). The δ^{13}C source signatures of plant derived CH$_4$

325   have been reported to be in the range of -72 mUr to -45 mUr (Keppler et al., 2006; Vigano et al., 2009) depending on the photosynthetic pathways (C$_3$, C$_4$ or CAM). Furthermore, there is a tendency towards more negative δ^{13}C-CH$_4$ values when the respective plant was treated with UV radiation (Vigano et al., 2009). δ^{13}C-CH$_4$ source signatures of humans which might include both formation by microbes in the gut but also from cellular processes show a rather wide range with values between -95 mUr and -56 mUr (Keppler et al., 2016). The results of our experiments conducted with two fungal species and three

330   different growth substrates provide a range of δ^{13}C-CH$_4$ source values from -69 mUr to -40 mUr. This range overlaps with other eukaryotic sources, most microbial CH$_4$ sources and even some abiotic CH$_4$ sources such as natural gas or emissions from coal mining.
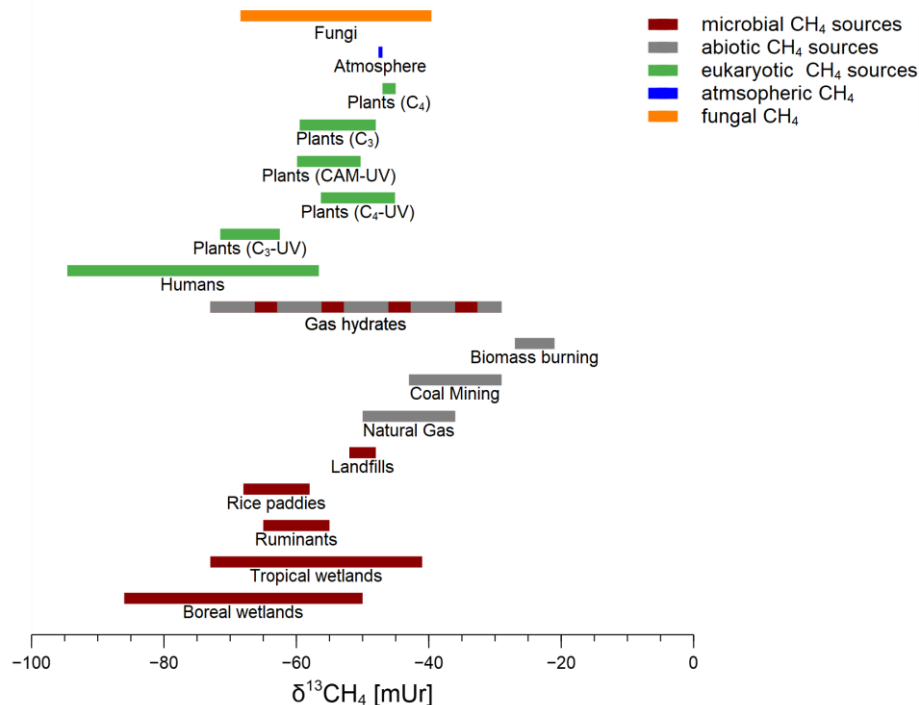
Figure 5: Range of δ ^{13}C-CH$_4$ values of microbial CH$_4$ sources (red), abiotic CH$_4$ sources (grey), eukaryotic CH$_4$ sources (green), atmospheric CH$_4$ (blue) and fungal CH$_4$ from this study (orange). The red and grey dashed bar indicates a mixture of microbial and abiotic CH$_4$ formation processes for gas hydrates (Kvenvolden, 1995). Data taken from (Brownlow et al., 2017; Keppler et al., 2006, 2016; Kvenvolden, 1995; Nisbet et al., 2016; Quay et al., 1999; Vigano et al., 2009).

4 Conclusion

This study provided the first analysis of stable carbon isotope values of CH$_4$ emitted by two saprotrophic fungi that were grown on three different substrates. δ^{13}C-CH$_4$ and δ^{13}C-CO$_2$ source values were found to be dependent on the fungal species, as well as the substrates decomposed by the fungi. δ^{13}C-CH$_4$ source values of the fungi were found to be in the range of -69 mUr to -40 mUr and therefore overlap with δ^{13}C-CH$_4$ values reported for other CH$_4$ sources such as methanogenic archaea, eukaryotes and from abiotic CH$_4$ sources (e.g. natural gas, coal mining). Stable carbon isotope values of CH$_4$ in combination with flux measurements are often applied for a better understanding of regional and global CH$_4$ cycling. However, in recent years it has become clear that many biogenic CH$_4$ sources include complex CH$_4$ formation processes, resulting in different isotopic fractionation patterns depending on several biochemical and abiotic factors. Thus, studying ecosystems in which more than one major CH$_4$ source has to be expected (e.g. methanogenic archaea, fungi, cyanobacteria or plants) gets increasingly complicated as distinguishing between each individual source solely by stable carbon isotope values might be highly

challenging. Therefore, additional tools are needed to better identify the sources but also to disentangle sources and sinks. In
350 future research, stable hydrogen isotopic values of CH_4 (δ^2H-CH_4 values) or even applications of clumped isotopes might prove
suitable tools for better distinction between different CH_4 sources and thus to better constrain the global CH_4 budget.

Data availability. We provide the data in heiDATA, which is an institutional repository for research data of Heidelberg
355 University (https://doi.org/10.11588/data/DQYPMC)

Author contributions. MS, KL, and FK conceived the study and designed the experiments; HZ provided fungal cultures, MS
performed the experiments under the supervision of FK and KL; CE helped with gas measurements; MG measured stable
isotope values of greenhouse gases; MS, FK, and KL analysed the data; MS, FK, HZ, MG, and KL, discussed the results, and
360 MS, KL and FK wrote the paper.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgments. We thank Anette Giesemann for analytical measurements of stable carbon isotope values of the bulk
365 substrates. We are grateful to Bianka Daubertshäuser for technical support with the cultivation of the fungi and to Lukas Kohl
for encouraging us to perform this study. We acknowledge financial support by the Deutsche Forschungsgemeinschaft.

Financial support. This research has been supported by the German Research Foundation (DFG Grant Numbers KE 884/8-2,
KE 884/16-2) and (LE3381/1-1).

370

375

380

References

Beckmann, S., Krüger, M., Engelen, B., Gorbushina, A. A. and Cypionka, H.: Role of bacteria, archaea and fungi involved in methane release in abandoned coal mines, Geomicrobiol. J., 28(4), 347–358, doi:10.1080/01490451.2010.503258, 2011.

385

Bižić, M., Klintzsch, T., Ionescu, D., Hindiyeh, M. Y., Günthel, M., Muro-Pastor, A. M., Eckert, W., Urich, T., Keppler, F. and Grossart, H. P.: Aquatic and terrestrial cyanobacteria produce methane, Sci. Adv., 6(3), eaax5343, doi:10.1126/sciadv.aax5343, 2020.

390   Boros, M. and Keppler, F.: Methane production and bioactivity-A link to oxido-reductive stress, Front. Physiol., 10(SEP), 1244, doi:10.3389/fphys.2019.01244, 2019.

Bowling, D. R., Pataki, D. E. and Randerson, J. T.: Carbon isotopes in terrestrial ecosystem pools and CO_2 fluxes, New Phytol., 178(1), 24–40, doi:10.1111/j.1469-8137.2007.02342.x, 2008.

395

Brand, W. A. and Coplen, T. B.: Stable isotope deltas: Tiny, yet robust signatures in nature, Isotopes Environ. Health Stud., 48(3), 393–409, doi:10.1080/10256016.2012.666977, 2012.

Brownlow, R., Lowry, D., Fisher, R. E., France, J. L., Lanoisellé, M., White, B., Wooster, M. J., Zhang, T. and Nisbet, E. G.: 400   Isotopic Ratios of Tropical Methane Emissions by Atmospheric Measurement, Global Biogeochem. Cy, 31(9), 1408–1419, doi:10.1002/2017GB005689, 2017.

Dlugokencky, E. J., Nisbet, E. G., Fisher, R. and Lowry, D.: Global atmospheric methane: Budget, changes and dangers, Philos. Trans. R. Soc. A Math. Phys. Eng. Sci., 369(1943), 2058–2072, doi:10.1098/rsta.2010.0341, 2011.

405

Ghyczy, M., Torday, C., Kaszaki, J., Szabó, A., Czóbel, M. and Boros, M.: Hypoxia-induced generation of methane in mitochondria and eukaryotic cells - An alternative approach to methanogenesis, Cell. Physiol. Biochem., 21(1–3), 251–258, doi:10.1159/000113766, 2008.

410   Grinhut, T., Hadar, Y. and Chen, Y.: Degradation and transformation of humic substances by saprotrophic fungi: processes and mechanisms, Fungal Biol. Rev., 21(4), 179–189, doi:10.1016/j.fbr.2007.09.003, 2007.

Ten Have, R. and Teunissen, P. J. M.: Oxidative mechanisms involved in lignin degradation by white-rot fungi, Chem. Rev., 101(11), 3397–3413, doi:10.1021/cr000115l, 2001.

415 Hein, R., Crutzen, P. J. and Heimann, M.: An inverse modeling approach to investigate the global atmospheric methane cycle, Global Biogeochem. Cycles, 11(1), 43–76, doi:10.1029/96GB03043, 1997.

Hietala, A. M., Dörsch, P., Kvaalen, H. and Solheim, H.: Carbon dioxide and methane formation in norway spruce stems infected by white-rot fungi, Forests, 6(9), 3304–3325, doi:10.3390/f6093304, 2015.

420

Keppler, F., Hamilton, J. T. G., Braß, M. and Röckmann, T.: Methane emissions from terrestrial plants under aerobic conditions, Nature, 439(7073), 187–191, doi:10.1038/nature04420, 2006.

Keppler, F., Boros, M., Frankenberg, C., Lelieveld, J., McLeod, A., Pirttilä, A. M., Röckmann, T. and Schnitzler, J. P.: Methane

425 formation in aerobic environments, Environ. Chem., 6, 459–465, doi:10.1071/EN09137, 2009.

Keppler, F., Schiller, A., Ehehalt, R., Greule, M., Hartmann, J. and Polag, D.: Stable isotope and high precision concentration measurements confirm that all humans produce and exhale methane, J. Breath Res., 10(1), doi:10.1088/1752-7155/10/1/016003, 2016.

430

Kirschke, S., Bousquet, P., Ciais, P., Saunois, M., Canadell, J. G., Dlugokencky, E. J., Bergamaschi, P., Bergmann, D., Blake, D. R., Bruhwiler, L., Cameron-Smith, P., Castaldi, S., Chevallier, F., Feng, L., Fraser, A., Heimann, M., Hodson, E. L., Houweling, S., Josse, B., Fraser, P. J., Krummel, P. B., Lamarque, J. F., Langenfelds, R. L., Le Quéré, C., Naik, V., O'doherty, S., Palmer, P. I., Pison, I., Plummer, D., Poulter, B., Prinn, R. G., Rigby, M., Ringeval, B., Santini, M., Schmidt, M., Shindell,

435 D. T., Simpson, I. J., Spahni, R., Steele, L. P., Strode, S. A., Sudo, K., Szopa, S., Van Der Werf, G. R., Voulgarakis, A., Van Weele, M., Weiss, R. F., Williams, J. E. and Zeng, G.: Three decades of global methane sources and sinks, Nat. Geosci., 6(10), 813–823, doi:10.1038/ngeo1955, 2013.

Klintzsch, T., Langer, G., Nehrke, G., Wieland, A., Lenhart, K. and Keppler, F.: Methane production by three widespread

440 marine phytoplankton species: Release rates, precursor compounds, and potential relevance for the environment, Biogeosciences, 16(20), 4129–4144, doi:10.5194/bg-16-4129-2019, 2019.

Kvenvolden, K. A.: A review of the geochemistry of methane in natural gas hydrate, Org. Geochem., 23(11–12), 997–1008, doi:10.1016/0146-6380(96)00002-2, 1995.

445

Lenhart, K., Bunge, M., Ratering, S., Neu, T. R., Schüttmann, I., Greule, M., Kammann, C., Schnell, S., Müller, C., Zorn, H. and Keppler, F.: Evidence for methane production by saprotrophic fungi, Nat. Commun., 3, doi:10.1038/ncomms2049, 2012.

Lenhart, K., Klintzsch, T., Langer, G., Nehrke, G., Bunge, M., Schnell, S. and Keppler, F.: Evidence for methane production by the marine algae *Emiliania huxleyi*, Biogeosciences, 13(10), 3163–3174, doi:10.5194/bg-13-3163-2016, 2016.

Leonowicz, A., Matuszewska, A., Luterek, J., Ziegenhagen, D., Wojtaś-Wasilewska, M., Cho, N. S., Hofrichter, M. and Rogalski, J.: Biodegradation of lignin by white rot fungi, Fungal Genet. Biol., 27(2–3), 175–185, doi:10.1006/fgbi.1999.1150, 1999.

Mikaloff Fletcher, S. E., Tans, P. P., Bruhwiler, L. M., Miller, J. B. and Heimann, M.: CH_4 sources estimated from atmospheric observations of CH_4 and its $^{13}C/^{12}C$ isotopic ratios: 1. Inverse modeling of source processes, Global Biogeochem. Cycles, 18(4), 1–17, doi:10.1029/2004GB002223, 2004a.

Mikaloff Fletcher, S. E., Tans, P. P., Bruhwiler, L. M., Miller, J. B. and Heimann, M.: CH_4 sources estimated from atmospheric observations of CH_4 and its $^{13}C/^{12}C$ isotopic ratios: 2. Inverse modeling of CH_4 fluxes from geographical regions, Global Biogeochem. Cy., 18(4), 1–15, doi:10.1029/2004GB002224, 2004b.

Mukhin, V. A. and Voronin, P. Y.: Methane emission during wood fungal decomposition, Dokl. Biol. Sci., 413(1), 159–160, doi:10.1134/S0012496607020202, 2007.

Mukhin, V. A. and Voronin, P. Y.: A new source of methane in boreal forests, Appl. Biochem. Microbiol., 44(3), 297–299, doi:10.1134/S0003683808030125, 2008.

Naik, V., Voulgarakis, A., Fiore, A. M., Horowitz, L. W., Lamarque, J. F., Lin, M., Prather, M. J., Young, P. J., Bergmann, D., Cameron-Smith, P. J., Cionni, I., Collins, W. J., Dalsøren, S. B., Doherty, R., Eyring, V., Faluvegi, G., Folberth, G. A., Josse, B., Lee, Y. H., MacKenzie, I. A., Nagashima, T., Van Noije, T. P. C., Plummer, D. A., Righi, M., Rumbold, S. T., Skeie, R., Shindell, D. T., Stevenson, D. S., Strode, S., Sudo, K., Szopa, S. and Zeng, G.: Preindustrial to present-day changes in tropospheric hydroxyl radical and methane lifetime from the Atmospheric Chemistry and Climate Model Intercomparison Project (ACCMIP), Atmos. Chem. Phys., 13(10), 5277–5298, doi:10.5194/acp-13-5277-2013, 2013.

Nisbet, E. G., Dlugokencky, E. J., Manning, M. R., Lowry, D., Fisher, R. E., France, J. L., Michel, S. E., Miller, J. B., White, J. W. C., Vaughn, B., Bousquet, P., Pyle, J. A., Warwick, N. J., Cain, M., Brownlow, R., Zazzeri, G., Lanoisellé, M., Manning, A. C., Gloor, E., Worthy, D. E. J., Brunke, E. G., Labuschagne, C., Wolff, E. W. and Ganesan, A. L.: Rising atmospheric methane: 2007–2014 growth and isotopic shift, Global Biogeochem. Cycles, 30(9), 1356–1370, doi:10.1002/2016GB005406, 2016.

Paul, D., Skrzypek, G. and Fórizs, I.: Normalization of measured stable isotopic compositions to isotope reference scales – a review, Rapid Commun. Mass Spectrom., 21(18), 3006–3014, doi:10.1002/rcm.3185, 2007.

485

Prather, M. J., Holmes, C. D. and Hsu, J.: Reactive greenhouse gas scenarios: Systematic exploration of uncertainties and the role of atmospheric chemistry, Geophys. Res. Lett., 39(9), n/a-n/a, doi:10.1029/2012GL051440, 2012.

Quay, P., Stutsman, J., Wilbur, D., Snover, A., Dlugokencky, E. and Brown, T.: The isotopic composition of atmospheric

490   methane, Global Biogeochem. Cycles, 13(2), 445–461, doi:10.1029/1998GB900006, 1999.

Ralph, J. P. and Catcheside, D. E. A.: Biodegradation by White-Rot Fungi, in Industrial Applications, pp. 303–326, Springer Berlin Heidelberg., 2002.

495   Saunois, M., Bousquet, P., Poulter, B., Peregon, A., Ciais, P., Canadell, J. G., Dlugokencky, E. J., Etiope, G., Bastviken, D., Houweling, S., Janssens-Maenhout, G., Tubiello, F. N., Castaldi, S., Jackson, R. B., Alexe, M., Arora, V. K., Beerling, D. J., Bergamaschi, P., Blake, D. R., Brailsford, G., Brovkin, V., Bruhwiler, L., Crevoisier, C., Crill, P., Covey, K., Curry, C., Frankenberg, C., Gedney, N., Höglund-Isaksson, L., Ishizawa, M., Ito, A., Joos, F., Kim, H. S., Kleinen, T., Krummel, P., Lamarque, J. F., Langenfelds, R., Locatelli, R., Machida, T., Maksyutov, S., McDonald, K. C., Marshall, J., Melton, J. R.,

500   Morino, I., Naik, V., O'Doherty, S., Parmentier, F. J. W., Patra, P. K., Peng, C., Peng, S., Peters, G. P., Pison, I., Prigent, C., Prinn, R., Ramonet, M., Riley, W. J., Saito, M., Santini, M., Schroeder, R., Simpson, I. J., Spahni, R., Steele, P., Takizawa, A., Thornton, B. F., Tian, H., Tohjima, Y., Viovy, N., Voulgarakis, A., Van Weele, M., Van Der Werf, G. R., Weiss, R., Wiedinmyer, C., Wilton, D. J., Wiltshire, A., Worthy, D., Wunch, D., Xu, X., Yoshida, Y., Zhang, B., Zhang, Z. and Zhu, Q.: The global methane budget 2000-2012, Earth Syst. Sci. Data, 8(2), 697–751, doi:10.5194/essd-8-697-2016, 2016.

505

Saunois, M., Stavert, A. R., Poulter, B., Bousquet, P., Canadell, J. G., Jackson, R. B., Raymond, P. A., Dlugokencky, E. J., Houweling, S., Patra, P. K., Ciais, P., Arora, V. K., Bastviken, D., Bergamaschi, P., Blake, D. R., Brailsford, G., Bruhwiler, L., Carlson, K. M., Carrol, M., Castaldi, S., Chandra, N., Crevoisier, C., Crill, P. M., Covey, K., Curry, C. L., Etiope, G., Frankenberg, C., Gedney, N., Hegglin, M. I., Höglund-Isakson, L., Hugelius, G., Ishizawa, M., Ito, A., Janssens-Maenhout,

510   G., Jensen, K. M., Joos, F., Kleinen, T., Krummel, P. B., Langenfelds, R. L., Laruelle, G. G., Liu, L., Machida, T., Maksyutov, S., McDonald, K. C., McNorton, J., Miller, P. A., Melton, J. R., Morino, I., Müller, J., Murgia-Flores, F., Naik, V., Niwa, Y., Noce, S., O'Doherty, S., Parker, R. J., Peng, C., Peng, S., Peters, G. P., Prigent, C., Prinn, R., Ramonet, M., Regnier, P., Riley, W. J., Rosentreter, J. A., Segers, A., Simpson, I. J., Shi, H., Smith, S. J., Steele, P. L., Thornton, B. F., Tian, H., Tohjima, Y., Tubiello, F. N., Tsuruta, A., Viovy, N., Voulgarakis, A., Weber, T. S., van Weele, M., van der Werf, G. R.,

515   Weiss, R. F., Worthy, D., Wunch, D., Yin, Y., Yoshida, Y., Zhang, W., Zhang, Z., Zhao, Y., Zheng, B., Zhu, Q., Zhu, Q. and Zhuang, Q.: The Global Methane Budget 2000-2017, Earth Syst. Sci. Data Discuss., 1–138, doi:10.5194/essd-2019-128, 2019.

Tyler, S. C.: Stable carbon isotope ratios in atmospheric methane and some of its sources, J. Geophys. Res., 91(D12), 13232, doi:10.1029/jd091id12p13232, 1986.

520

Valášková, V. and Baldrian, P.: Degradation of cellulose and hemicelluloses by the brown rot fungus *Piptoporus betulinus* - Production of extracellular enzymes and characterization of the major cellulases, Microbiology, 152(12), 3613–3622, doi:10.1099/mic.0.29149-0, 2006.

525 Vigano, I., Röckmann, T., Holzinger, R., van Dijk, A., Keppler, F., Greule, M., Brand, W. A., Geilmann, H. and van Weelden, H.: The stable isotope signature of methane emitted from plant material under UV irradiation, Atmos. Environ., 43(35), 5637–5646, doi:10.1016/j.atmosenv.2009.07.046, 2009.

Voulgarakis, A., Naik, V., Lamarque, J. F., Shindell, D. T., Young, P. J., Prather, M. J., Wild, O., Field, R. D., Bergmann, D., 530 Cameron-Smith, P., Cionni, I., Collins, W. J., Dalsøren, S. B., Doherty, R. M., Eyring, V., Faluvegi, G., Folberth, G. A., Horowitz, L. W., Josse, B., MacKenzie, I. A., Nagashima, T., Plummer, D. A., Righi, M., Rumbold, S. T., Stevenson, D. S., Strode, S. A., Sudo, K., Szopa, S. and Zeng, G.: Analysis of present day and future OH and methane lifetime in the ACCMIP simulations, Atmos. Chem. Phys., 13(5), 2563–2587, doi:10.5194/acp-13-2563-2013, 2013.

535 Whiticar, M. J.: Stable Isotopes and Global Budgets, in Atmospheric Methane: Sources, Sinks, and Role in Global Change, pp. 138–167, Springer Berlin Heidelberg., 1993.

Whiticar, M. J.: Carbon and hydrogen isotope systematics of bacterial formation and oxidation of methane, Chem. Geol., 161(1), 291–314, doi:10.1016/S0009-2541(99)00092-3, 1999.

540

Wishkerman, A., Greiner, S., Ghyczy, M., Boros, M., Rausch, T., Lenhart, K. and Keppler, F.: Enhanced formation of methane in plant cell cultures by inhibition of cytochrome c oxidase, Plant Cell Environ., 34(3), 457–464, doi:10.1111/j.1365-3040.2010.02255.x, 2011.