



The stable carbon isotope signature of methane produced by saprotrophic fungi

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15 **Abstract.** Methane (CH₄) is the most abundant organic compound in the atmosphere with emissions from many biotic and
abiotic sources. Recent studies have shown that CH₄ production occurs under aerobic conditions in eukaryotes such as plants,
animals, algae and saprotrophic fungi. Saprotrophic fungi play an important role in nutrient recycling in terrestrial ecosystems
by their ability to decompose plant litter. Even though the CH₄ production by saprotrophic fungi has been reported, so far, no
data for stable carbon isotope values of the emitted CH₄ ($\delta^{13}\text{C}\text{-CH}_4$ values) is available. In this study we measured the $\delta^{13}\text{C}$
20 values of CH₄ and carbon dioxide ($\delta^{13}\text{C}\text{-CO}_2$ values) emitted by the two saprotrophic fungi *Pleurotus sapidus* and *Laetiporus*
sulphureus cultivated on three different substrates pine wood, grass and corn, reflecting both C₃ and C₄ plants with
distinguished bulk $\delta^{13}\text{C}$ values. Applying keeling plots, we found that the $\delta^{13}\text{C}$ source values of CH₄ emitted from fungi cover
a wide range from -40 mUr to -69 mUr depending on the growth substrate and fungal species. Whilst little apparent carbon
isotopic fractionation (in the range of -0.3 mUr to 4.6 mUr) was calculated for $\delta^{13}\text{C}$ values of CO₂ released from *P. sapidus*
25 and *L. sulphureus* relative to the bulk $\delta^{13}\text{C}$ values of the growth substrates, much larger carbon isotopic fractionations (ranging
from -22 mUr to -42 mUr) were observed for the formation of CH₄. Whilst the two fungal species showed similar $\delta^{13}\text{C}\text{CH}_4$
source values when grown on pine wood, $\delta^{13}\text{C}\text{CH}_4$ source values differed substantially between the two fungal species when
grown on grass or corn. We found that $\delta^{13}\text{C}\text{CH}_4$ source values emitted by saprotrophic fungi are highly dependent on the fungal
species and the metabolized substrate. They cover a broad range of $\delta^{13}\text{C}\text{CH}_4$ values and overlap with values reported for
30 methanogenic archaea, thermogenic degradation of organic matter and other eukaryotes.



1 Introduction

Methane (CH₄) is an important greenhouse gas that is emitted by several abiotic sources (e.g. fossil fuel burning and use, biomass burning, geological processes) and biotic sources (e.g. wetlands, agriculture and waste, fresh waters) to the atmosphere (Kirschke et al., 2013; Saunio et al., 2016, 2019). In the past, biotic CH₄ production has been attributed exclusively to strictly anaerobic microorganisms such as methanogens that are ubiquitous in wetlands, rice paddies, landfills and the intestines of termites and ruminants (Kirschke et al., 2013). The discovery of CH₄ emissions from dead and living plants under oxic conditions (Keppler et al., 2006, 2009) paved the way for the search of new biogenic CH₄ sources. Since then, several previously unknown CH₄ sources were discovered including endothelial cells of rat liver (Boros and Keppler, 2019; Ghyczy et al., 2008), plant cell cultures (Wishkerman et al., 2011), marine algae (Klintzsch et al., 2019; Lenhart et al., 2016), marine and terrestrial cyanobacteria (Bižić et al., 2020), humans (Keppler et al., 2016) and saprotrophic fungi (Lenhart et al., 2012). Fungi play a central role in ecosystems by decomposing organic matter and thereby recycling formerly bound carbon and nutrients (Grinhut et al., 2007). This process is especially important in forests where fungi are essential for wood decay and therefore have a great impact on the carbon and nitrogen cycles in these environments (Ralph and Catcheside, 2002). White rot fungi are able to decompose the chemically complex structural component lignin, whereas brown rot fungi mainly metabolize cellulose and hemicellulose (Ten Have and Teunissen, 2001; Leonowicz et al., 1999; Valášková and Baldrian, 2006). Fungi have already been determined to be involved in the CH₄ synthesis during wood decay (Beckmann et al., 2011; Mukhin and Voronin, 2007, 2008) by breakdown of large macromolecules to smaller molecules, thereby providing bacteria and methanogenic archaea with their substrate. Elevated levels of CH₄ were found in fungally infected wood stems with oxygen concentrations ranging from 1 to 14 % (Hietala et al., 2015). Here, CH₄ production was associated with anoxic microsites in the xylem, indicating that at least part of the CH₄ was produced by methanogenic archaea. Nevertheless, Lenhart et al., 2012 demonstrated that basidiomycetes are able to produce CH₄ under aerobic conditions without the presence of methanogenic archaea. Therefore, fungi might be a so far underestimated source of CH₄ in the global CH₄ cycle.

Applications of stable carbon isotopes (expressed as $\delta^{13}\text{C}$ values) have often been used to investigate sources and sinks of CH₄ on the global scale (Whiticar, 1993). As different CH₄ sources have distinct $\delta^{13}\text{C}$ fingerprints, they might be used to quantify the individual contributions of various sources regionally and/or globally (Dlugokencky et al., 2011; Hein et al., 1997; Nisbet et al., 2016; Quay et al., 1999; Tyler, 1986; Whiticar, 1999). The short lifetime of CH₄ in the atmosphere (range from 9.7 ± 1.5 to 11.2 ± 1.3 years) (Naik et al., 2013; Prather et al., 2012; Voulgarakis et al., 2013) assures that global isotopic patterns represent the average of recent inputs by various sources and allows the quantification of respective source strengths (Mikaloff Fletcher et al., 2004b, 2004a).

Additionally, stable isotopes provide information about the formation processes of CH₄. Traditionally, three formation categories of $\delta^{13}\text{C}$ -CH₄ values have been classified: biogenic, with typical $\delta^{13}\text{C}$ -CH₄ values ranging from ~ -55 mUr to -70 mUr, thermogenic (ranging from ~ -25 mUr to -55 mUr) and pyrogenic (ranging from ~ -13 mUr to -25 mUr) (Kirschke et al., 2013). However, isotopic patterns of recently identified CH₄ sources, i.e. human CH₄ emissions (-56 mUr to -95 mUr) (Keppler



et al., 2016), plant derived CH₄ (-52 mUr to -69 mUr) (Keppler et al., 2006), and abiotic UV induced CH₄ formation by plants
65 (-52 mUr to -67 mUr) (Vigano et al., 2009) also need to be considered.

In this study we investigated the stable carbon isotope source signatures of CH₄ and CO₂ released by the two basidiomycetes
Pleurotus sapidus (white rot fungus) and *Laetiporus sulphureus* (brown rot fungus). Both fungi were cultivated under sterile
conditions on three different substrates (pine wood, grass, and corn) with varying bulk δ¹³C values. We examined the influence
of fungal species and growth substrate on δ¹³C-CH₄ and δ¹³C-CO₂ values and compared the δ¹³C-CH₄ values from the two
70 fungal species with those of other known sources reported from the literature.

2 Material and Methods

2.1 Selected fungi

P. sapidus (*Pleurotaceae*, DSMZ 8266) and *L. sulphureus* (*Polyporaceae*, DSMZ 1014) were chosen for this experiment
because of their capability to emit CH₄ (Lenhart et al., 2012), their ecological and physiological characteristics (white and
75 brown rot fungi) and well-established practical handling under laboratory conditions.

2.2 Cultivation of fungi and incubation experiments

Pine wood, grass and corn were selected as growth substrates. Pine wood was chosen to investigate if brown and white rot
fungi differ in δ¹³C-CH₄ and δ¹³C-CO₂ values released during wood decay. Therefore, dead pine wood branches were collected
from the forest floor and shredded to small wood chips with a length of about 5 cm (Natura 1800L; Glora, Witten, Germany).
80 The wood chips were dried at 60°C for 48h and stored in a flask (Weck, Hanau, Germany). Grass (C₃ plant) and corn (C₄ plant)
were selected because of their different stable isotope values. As the metabolic pathway for carbon fixation is biochemically
different in C₃ and C₄ plants, plant biomass differs in δ¹³C values, which in turn might lead to different δ¹³C values of CH₄ and
CO₂ released by fungi. Therefore, typical garden lawn was manually cut, dried at 70 °C, and stored in a flask. The corn substrate
consisted of conventional corn flour.

85 The substrates were autoclaved and filled into 2.7 l flasks (Weck, Hanau, Germany) and inoculated with pure fungal submerged
cultures under sterile conditions according to Lenhart et al., 2012. After addition of the fungi, the flasks were closed with lids
and a rubber band sealing. To allow gas exchange during the growth time of the fungi (about two weeks), a hole in the centre
of every lid was fitted with a cotton stopper. Before the start of the incubation experiments, the flasks were aerated under
sterile conditions in order to start the incubation at atmospheric CH₄ mixing ratios. Additionally, to seal the flasks airtight the
90 cotton stoppers were replaced by sterile silicone stoppers (Saint-Gobain Performance Plastics, Charny, France).

For the incubation experiments, *P. sapidus* und *L. sulphureus* were incubated on the three substrates, while substrates were
incubated as control treatments. Before the incubation experiments, the substrates were sterilized by autoclaving. The
incubation experiments were conducted as three replicates per treatment. The duration of the incubation accounted for up to
40 h. All incubations were conducted at room temperature (23 ± 1.5 °C). At every sampling point, 40 ml air was taken from



95 the flasks for gas concentration measurements and an additional 40 ml were taken for $\delta^{13}\text{C}$ - CH_4 stable isotope ratio mass spectrometry (IRMS) analysis. The gas samples were taken with airtight 60 ml PE syringes (Plastipak, BD, Franklin Lakes, USA) and transferred into 12 ml evacuated Exetainers (Labco, High Wycombe, UK). Subsequently a volume of atmospheric air equivalent to the volume of the removed sample was added into each flask directly after sampling. Mixing ratios and stable isotope values of CH_4 were corrected according to the dilution.

100 When calculating the fungal CH_4 and CO_2 production rates, we subtracted substrate derived CH_4 and CO_2 production rates (determined in the control treatments) from the respective fungi containing samples. Additionally, for the calculation of the fungal production rates only sample points showing a linear increase in CH_4 and CO_2 were taken into account. To account for differences in the metabolic activity of the fungi, we additionally measured respiration rates, assuming that metabolic activity correlates with respiration and therefore CO_2 emissions of the fungi. Hence, we related fungal derived CH_4

105 emissions to respiration by calculating the $\text{CH}_4 : \text{CO}_2$ emission ratio.

2.3 Analysis of CH_4 and CO_2 via gas-chromatography

Samples were analysed using a gas chromatograph (GC, Bruker Greenhouse Gas Analyser 450-GC) equipped with a flame ionization detector (FID) and an electron capture detector (ECD) for the detection of CH_4 and CO_2 , respectively. The detector temperatures were set at 300 °C (FID) and 350 °C (ECD). Five reference gases were used for calibrating the GC-system. The

110 reference gases were in the range of 1 parts per million by volume (ppmv) to 21 ppmv and 304 ppmv to 40,000 ppmv for CH_4 and CO_2 , respectively. Gas peaks were integrated using Galaxie software (Varian Inc., Palo Alto, CA, USA).

2.4 Definition of δ values and isotope apparent fractionation

In this paper, all stable carbon isotope ratios are expressed in the conventional ‘delta’ δ notation, meaning the relative difference of the isotope ratio of a substance compared to the standard substance Vienna Pee Dee Belemnite (V-PDB) (Eq. (1)).

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$$\delta^{13}\text{C} = \frac{\left(\frac{^{13}\text{C}}{^{12}\text{C}}\right)_{\text{sample}}}{\left(\frac{^{13}\text{C}}{^{12}\text{C}}\right)_{\text{V-PDB}}} - 1 \quad (1)$$

The apparent fractionation (ϵ_{app}) between fungal $\delta^{13}\text{C}$ - CH_4 or $\delta^{13}\text{C}$ - CO_2 values and the $\delta^{13}\text{C}$ values of the substrates was calculated according to Eq. (2).

$$\epsilon_{\text{app CH}_4 \text{ or CO}_2} = \frac{(\delta^{13}\text{C} + 1)_{\text{fungal CH}_4 \text{ or CO}_2}}{(\delta^{13}\text{C} + 1)_{\text{substrate}}} - 1 \quad (2)$$

120 We follow the proposal of Brand and Coplen, 2012 and use the term ‘urey’ (Ur) as the isotope delta unit, in order to conform with the guidelines for the International System of Units (SI). Hence, isotope delta values that were formerly given as -70 ‰, are expressed as -70 mUr.



2.5 Measurements of $\delta^{13}\text{CH}_4$ and $\delta^{13}\text{CO}_2$ values

Stable carbon isotope values of CH_4 and CO_2 were measured using a continuous flow isotope mass spectrometry system (CF-IRMS). A HP 6890N GC (Agilent, Santa Clara, USA) was linked to a preconcentration unit for CH_4 measurements and an autosampler A200S (CTC Analytics, Zwingen, Switzerland) for CO_2 analysis. The GC was equipped with a CP-PoraPLOT Q capillary column (Varian, Palo Alto, USA) (27,5 m x 0.25 mm i.d., film thickness 8 μm). The GC was operated with an injector temperature of 200°C, isothermal oven temperature of 30°C, split injection (10:1) and a constant carrier gas flow of 1.8 ml min^{-1} (methane-free helium). The GC was coupled to a Delta^{PLUS}XL isotope ratio mass spectrometer (ThermoQuest Finnigan, Bremen, Germany) via an oxidation reactor and a GC Combustion III Interface (ThermoQuest Finnigan, Bremen, Germany). The oxidation reactor was employed with the following properties: ceramic tube (Al_2O_3), length 320 mm, 1.0 mm i.d., with Ni/Pt wires inside activated by oxygen, reactor temperature 960 °C.

For CH_4 measurements with the preconcentration unit, headspace gas samples were transferred to an evacuated 40 ml sample loop. Methane was trapped on Hayesep D, separated from other compounds by the GC and then introduced into the IRMS system via an open split. The working reference gas was carbon dioxide of high purity (carbon dioxide 4.5, Messer Griesheim, Frankfurt, Germany) with a known $\delta^{13}\text{C}$ value of -23.6 mUr (calibrated at MPI for Biogeochemistry in Jena, Germany). All $\delta^{13}\text{C}$ values were corrected using two working reference gases of high purity carbon dioxide (Isometric instruments, Victoria, Canada) with $\delta^{13}\text{C}$ values of -23.9 ± 0.2 mUr and -54.5 ± 0.2 mUr that were calibrated against IAEA and NIST reference substances. The normalization of the sample values was done according to Paul et al., 2007.

2.6 Bulk isotope analysis of fungal substrates

Stable carbon isotope values of the bulk substrate were measured using an Elemental Analyzer Flash EA 11112 (Thermo Fischer Scientific, Germany) coupled to a Delta V IRMS (Thermo Fischer Scientific, Germany). Therefore, 0.06 mg of the substrate were put into a tin cup and combusted in the Elemental Analyzer. The resulting gases are separated in a GC by a CP-PoraPLOT Q capillary column (Varian, Palo Alto, USA) (27,5 m x 0.25 mm i.d., film thickness 8 μm) and then reach the Delta V IRMS via a Conflo IV Universal Continuous Flow Interface (Thermo Fischer Scientific, Germany). Isotope values were corrected using USGS 40 and USGS 41 standards.

2.7 Determination of isotopic source signature of CH_4 and CO_2 applying keeling plots

For the determination of $\delta^{13}\text{C}$ source values of CH_4 and CO_2 the keeling plot method was used (Keeling, 1958; Pataki et al., 2003) (Eq. (3)):

$$\delta^{13}\text{C}_a = c_b(\delta^{13}\text{C}_b - \delta^{13}\text{C}_s) \left(\frac{1}{c_a}\right) + \delta^{13}\text{C}_s \quad (3)$$

where c_a is the mixing ratio of CH_4/CO_2 in the headspace, $\delta^{13}\text{C}_a$ is the $\delta^{13}\text{C}$ value of CH_4/CO_2 in the headspace, c_b is the mixing ratio of background CH_4/CO_2 , $\delta^{13}\text{C}_b$ is the $\delta^{13}\text{C}$ value of background CH_4/CO_2 and $\delta^{13}\text{C}_s$ the $\delta^{13}\text{C}$ source value of the



CH₄/CO₂. For a more detailed description of the application of keeling plots for determination of CH₄ source signature we refer to the study by Keppler et al., 2016.

155 δ¹³C-CH₄ source signatures were calculated after the keeling plot method for each flask. Results of the keeling plots are then given as the arithmetic mean of the three individual flasks per treatment with standard deviations (n=3).

δ¹³C-CH₄ source signatures of each flask of *P. sapidus* grown on pine and *L. sulphureus* grown on pine were corrected for CH₄ emissions and δ¹³C-CH₄ values of the “pine” control samples using the following mass balance approach (Eq. (4)).

$$\delta^{13}\text{C}_{\text{fungi corrected}} = \frac{(P(\text{CH}_4)_{\text{fungi}} * \delta^{13}\text{C}_{\text{fungi}}) - (P(\text{CH}_4)_{\text{pine}} * \delta^{13}\text{C}_{\text{pine}})}{(P(\text{CH}_4)_{\text{fungi}} - P(\text{CH}_4)_{\text{pine}}} \quad (4)$$

160 , where P(CH₄)_{fungi/pine wood} is the CH₄ emitted by the fungi or pine wood and δ¹³C_{fungi/pinewood} is the δ¹³C-CH₄ source signature of the fungi or pine wood derived from keeling plots. Corrected δ¹³C-CH₄ source values for *P. sapidus* and *L. sulphureus* are given as the arithmetic mean of the three individual flasks per treatment with standard deviations (n=3).

The determination coefficient (R²) of the keeling plots showed values higher than 0.93, except for *P. sapidus* grown on grass (R²=0.51).

165 2.8 Statistics

Mixing ratios and production rates of CH₄, CO₂, δ¹³C-CH₄ and δ¹³C-CO₂ values and δ¹³C source values are presented as arithmetic mean of three independent replicates with standard deviations (SD; n = 3). The SDs are given with a confidence interval of 1 σ. Linear regression analysis, arithmetic means and SDs were calculated using Excel (Microsoft Excel for Office 365 MSO). Two-way analysis of variance (ANOVA) and a post hoc test (Fisher least significant difference) (SigmaPlot 170 12.2.0.45, USA) were carried out to test for “species” and “substrate” related effects on δ¹³C-CH₄ and δ¹³C-CO₂ source values for each treatment. Differences at the p < 0.05 level were referred to as significant.

3 Results and Discussion

In this section, we firstly present the results of CH₄ and CO₂ production from the two fungal species grown on the three different substrates. This includes emission rates of CH₄ and CO₂ from the control treatments of pine wood, grass and corn as well as the molar ratio of CH₄ and CO₂. Secondly, we then present the respective stable isotope values measured for CH₄ and CO₂ during the incubation experiments and calculate the stable isotope source values of CH₄ and CO₂ released by the fungi applying keeling plots. We then compare these values with stable carbon isotope values of the bulk organic matter by calculating the apparent fractionation. Finally, we compare δ¹³C source values of fungal derived CH₄ with sources values known for other CH₄ sources from the literature.



180 **3.1 Release of CH₄ and CO₂ from *P. sapidus* and *L. sulphureus***

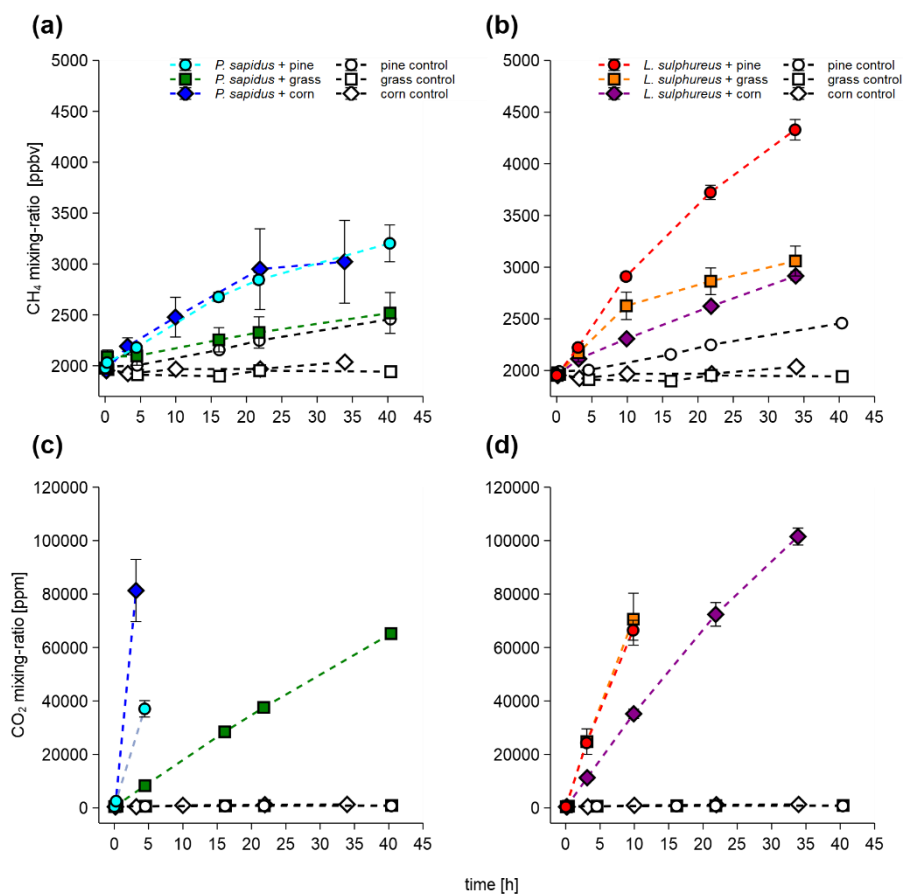


Figure 1: Mixing ratios of CH₄ and CO₂ of *P. sapidus* (a, c) and *L. sulphureus* (b, d) grown on pine wood, grass, and corn. Mixing ratios are presented as mean values with standard deviation SD (n=3).

185 All incubation experiments where fungi were grown on the different substrates showed a significant increase in CH₄ compared to the respective substrate control (Fig. 1 a, c). Calculated emission rates for CH₄ and CO₂ are presented in Table 1. *L. sulphureus* grown on grass ($7.5 \pm 1.3 \text{ nmol h}^{-1}$) showed the highest emission rate of CH₄, followed by *L. sulphureus* grown on pine ($6.2 \pm 0.3 \text{ nmol h}^{-1}$), *P. sapidus* grown on corn ($4.4 \pm 1.9 \text{ nmol h}^{-1}$), *L. sulphureus* grown on corn ($2.6 \pm 0.1 \text{ nmol h}^{-1}$), *P. sapidus* grown on pine ($2.5 \pm 0.2 \text{ nmol h}^{-1}$) and *P. sapidus* grown on grass ($1.4 \pm 0.5 \text{ nmol h}^{-1}$). Please note that CH₄ and CO₂ emission rates are not related to fungal biomass. Therefore, differences in the emission rates might be due to varying fungal biomass of the subsamples. Instead, CH₄ production was related to CO₂ production by determining the molar emission ratio between CH₄ and CO₂ (nmol CH₄ : mmol CO₂). CO₂ production thereby reflects the amount of fungal biomass and is also an indicator for the metabolic activity of the fungi.

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Most of the controls did not show significant changes in their CH₄ and CO₂ mixing ratios over time. However, in the control
195 flasks of pine wood and corn small CH₄ emission rates of 1.3 ± 0.1 nmol h⁻¹ and 0.25 ± 0.01 nmol h⁻¹ were observed, and in
the control ‘grass’ the CH₄ mixing ratio slightly decreased over time (-0.05 ± 0.04 nmol h⁻¹). Whilst the pine wood and corn
control flasks showed a small increase in the CH₄ mixing ratio, they did not show an increase in CO₂ mixing ratios. These data
rule out a contamination by microbial heterotrophs, as this would cause a measurable CO₂ increase within the flasks. The CH₄
increase in the substrate controls might be attributed to CH₄ release by dead plant material as it was already shown by Keppler
200 et al., 2006 and Vigano et al., 2009. Within the scope of these experiments, no analytic test for microbial contamination was
conducted. Nevertheless, Lenhart et al., 2012 clearly showed that with the performed method of cultivation of fungi and
incubation experiments no methanogenic archaea were present, using three different methods (Fluorescence in situ
hybridization (FISH), confocal laser scanning microscopy (CLSM) and quantitative real time PCR). Furthermore, CH₄ and
CO₂ release and the CH₄ : CO₂ emission ratios in our incubations are similar to the experiments of Lenhart et al., 2012 and do
205 not indicate microbial contamination. Therefore, we assume that in our investigations no contamination with bacteria or
methanogenic archaea were present.

For *P. sapidus* grown on corn and *L. sulphureus* grown on grass, no further linear increase in CH₄ was observed after 22 h and
10 h, respectively. This might be due to a reduced decay of organic matter and slower fungal metabolism because of higher
210 CO₂ and lower O₂ mixing ratios.

A drastic increase in CO₂ mixing ratios relative to the controls was observed in all flasks containing fungi (Fig. 1 b, d). The
CO₂ emission rates are shown in Table 1. CO₂ production rates ranged from 176 ± 4 μmol h⁻¹ to 2910 ± 410 μmol h⁻¹ for *P.*
sapidus grown on grass and *P. sapidus* grown on corn, respectively. These highly variable CO₂ production rates might reflect
different fungal biomass and metabolic activity (mineralisation of organic matter). In the control treatments, tiny increases in
215 the CO₂ mixing ratio were detected ranging from 0.64 ± 0.12 μmol h⁻¹ to 0.91 ± 0.14 μmol h⁻¹. Only one flask (corn control)
showed a somewhat higher increase in CO₂ (7.76 μmol h⁻¹), which is most likely caused by microbial contamination of the
flask. However, no increase in the CH₄ mixing ratio was detected (see supplementary material). Therefore, this control flask
was excluded from further calculations.

Mean CH₄ and CO₂ emission rates and CH₄ : CO₂ emission ratios of all treatments are presented in Table 1. Higher ratios
220 indicate a higher CH₄ production during decay of the substrates. Thereby, fungal species and substrate both affect the CH₄ :
CO₂ emission ratio ($p > 0.001$). For *P. sapidus* CH₄ : CO₂ emission ratios are more variable (1.4 to 8.0 nmol CH₄/mmol CO₂)
compared to *L. sulphureus* (6.7 – 9.6 nmol CH₄/mmol CO₂). This variation might be due to differences in the fungi’s enzyme
sets required for organic matter decay, as *P. sapidus* is a white rot fungus and *L. sulphureus* is a brown rot fungus. At present
the biochemical pathways that lead to CH₄ are still unknown, although compounds such as the sulphur-bound methyl-group
225 of methionine and glucose have been identified to act as carbon precursors of fungal-derived CH₄ (Lenhart et al., 2012).

Lenhart et al., 2012 found CH₄ : CO₂ ratios of fungi that ranged between 8 nmol CH₄/mmol CO₂ and 17 nmol CH₄/mmol CO₂,
which is in a good accordance with the CH₄ : CO₂ ratios determined in this study. Please note, that CH₄ : CO₂ ratios of Lenhart



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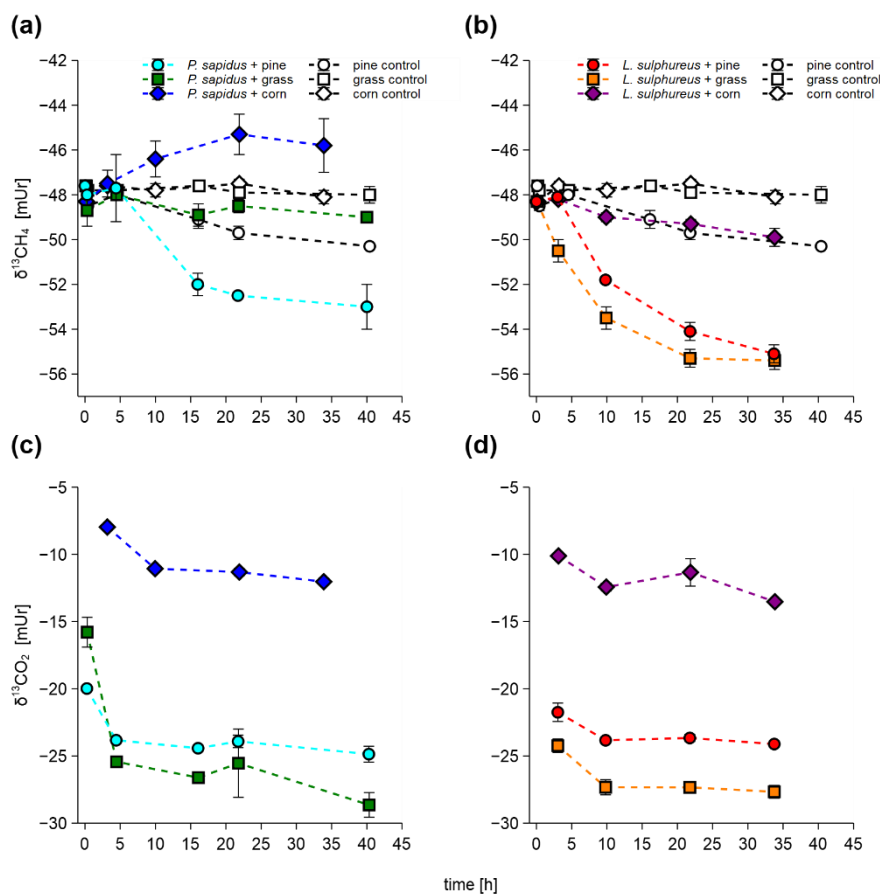
et al., 2012 were given in ppbv CH₄ : % CO₂ and for better comparability CH₄ : CO₂ ratios were converted to fit the units used in this study (nmol CH₄ : mmol CO₂).

Table 1: CH₄ and CO₂ production rates and molar CH₄ : CO₂ emission ratios of the fungi incubated on different substrates. Values are presented as mean values of three independent replicates with SD (n = 3), except for the control “corn” (n=2).

Fungi	Substrate	CH ₄ production rate [nmol h ⁻¹]	CO ₂ production rate [μmol h ⁻¹]	CH ₄ : CO ₂ ratio [nmol/mmol]
<i>P. sapidus</i>	pine	2.5 ± 0.2	901 ± 79	2.8 ± 0.4
	grass	1.4 ± 0.5	176 ± 4	8.0 ± 2.8
	corn	4.4 ± 1.9	2910 ± 419	1.4 ± 0.5
<i>L. sulphureus</i>	pine	6.2 ± 0.3	724 ± 42	8.6 ± 1.0
	grass	7.5 ± 1.3	771 ± 103	9.6 ± 0.5
	corn	2.6 ± 0.1	385 ± 20	6.7 ± 0.4
control	pine	1.3 ± 0.1	0.64 ± 0.12	-
	grass	-0.05 ± 0.04	0.91 ± 0.14	-
	corn	0.25	0.66	-



3.2 Stable carbon isotope values of CH₄ and CO₂



235 **Figure 2:** Stable carbon isotope values of CH₄ and CO₂ of *P. sapidus* (a, c) and *L. sulphureus* (b, d) grown on pine, grass, and
 240 corn. Values are presented as mean values with SD (n=3), except for δ¹³C-CO₂ values of *L. sulphureus* grown on corn (n=2).

Stable carbon isotope values of CH₄ and CO₂ measured from the incubation experiments are presented in Fig. 2. All incubations
 show a trend towards more negative δ¹³C-CH₄ values (less ¹³C) with time except for *P. sapidus* grown on corn, where a
 240 tendency towards more positive δ¹³C-CH₄ values was observed (Fig. 2 a, b). During the incubation, δ¹³C-CH₄ values changed
 from -47.7 ± 0.1 mUr (for incubation of *P. sapidus* grown on pine/grass) and -48.2 ± 0.1 mUr (for incubation of *P. sapidus*
 grown on corn and *L. sulphureus* grown on pine/grass/corn) to -53.0 ± 0.7 mUr (*P. sapidus* grown on pine), -48.7 ± 0.3 mUr
 (*P. sapidus* grown on grass), -45.8 ± 1.2 mUr (*P. sapidus* grown on corn), -55.1 ± 0.4 mUr (*L. sulphureus* grown on pine), -
 55.4 ± 0.4 mUr (*L. sulphureus* grown on grass) and -49.9 ± 0.4 mUr (*L. sulphureus* grown on corn). The controls showed no
 245 significant shift in δ¹³C-CH₄ values except for the control “pine”, where an increase in the CH₄ mixing ratio along with more



negative values of $\delta^{13}\text{C-CH}_4$ values occurred over time. This was accounted for when calculating the $\delta^{13}\text{C-CH}_4$ source signatures for *P. sapidus* grown on pine and *L. sulphureus* grown on pine (see materials and methods 2.7).
 The $\delta^{13}\text{C-CO}_2$ values showed a trend towards more negative values within the first three to four hours of incubation (Fig. 2 c, d). After this time only minor changes of the $\delta^{13}\text{C-CO}_2$ values occurred. Final $\delta^{13}\text{C-CO}_2$ values of the incubation were -24.9 ± 0.6 mUr (*P. sapidus* grown on pine), -28.6 ± 0.9 mUr (*P. sapidus* grown on grass), -12.0 ± 0.3 mUr (*P. sapidus* grown on corn), -24.1 ± 0.1 mUr (*L. sulphureus* grown on pine), -27.7 ± 0.5 mUr (*L. sulphureus* grown on grass) and -13.0 ± 0.5 mUr (*L. sulphureus* grown on corn).

Table 2: Calculated $\delta^{13}\text{C-CH}_4$ and $\delta^{13}\text{C-CO}_2$ source signatures, $\delta^{13}\text{C}$ values of the substrates, and $\epsilon_{\text{app CH}_4}$ and $\epsilon_{\text{app CO}_2}$. Values are presented as mean values with the SD (n=3).

Fungi	Substrate	$\delta^{13}\text{C-CH}_4$ source [mUr]	$\delta^{13}\text{C-CO}_2$ source [mUr]	$\delta^{13}\text{C}$ substrate [mUr]	$\epsilon_{\text{app CH}_4}$ [mUr]	$\epsilon_{\text{app CO}_2}$ [mUr]
<i>P. sapidus</i>	pine	-65.3 ± 1.1	-24.1 ± 0.1		-38.4 ± 1.2	4.0 ± 0.1
	grass	-52.9 ± 1.6	-27.4 ± 1.3		-21.8 ± 1.7	4.6 ± 1.3
	corn	-39.8 ± 2.0	-12.0 ± 0.3		-28.5 ± 2.0	-0.3 ± 0.3
<i>L. sulphureus</i>	pine	-61.4 ± 0.5	-25.0 ± 0.5		-34.4 ± 0.6	3.0 ± 0.4
	grass	-69.2 ± 1.9	-29.0 ± 0.5		-38.6 ± 2.0	2.9 ± 0.5
	corn	-53.4 ± 1.1	-12.8 ± 0.3		-42.2 ± 1.1	-1.1 ± 0.3
control	pine			-28.0 ± 0.5		
	grass			-31.5 ± 0.6		
	corn			-11.7 ± 0.1		



3.3 $\delta^{13}\text{C}\text{-CH}_4$ and $\delta^{13}\text{C}\text{-CO}_2$ source signatures of fungi

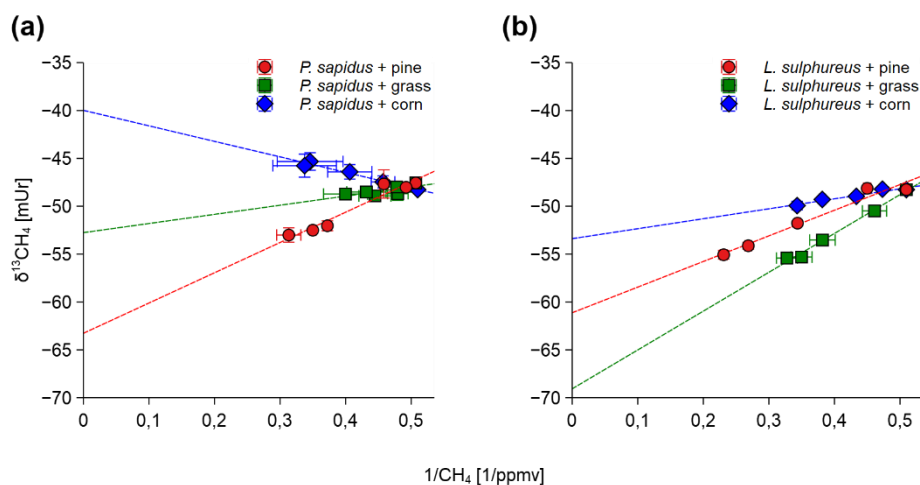


Figure 3: Keeling plots shown for *P. sapidus* (a) and *L. sulphureus* (b) grown on three substrates. Sample points in the graphs are given as the arithmetic mean of $\delta^{13}\text{C}\text{-CH}_4$ or $\delta^{13}\text{C}\text{-CO}_2$ values with SD ($n=3$) on the y-axis and the arithmetic mean of the inverted mixing ratio of CH_4 or CO_2 with SD ($n=3$) on the x-axis.

The $\delta^{13}\text{C}\text{-CH}_4$ source signatures determined via a keeling plot analysis (Fig. 3) are presented in Table 2 and range from -69.2 ± 1.9 mUr (*L. sulphureus* grown on grass) to -39.8 ± 2.0 mUr (*P. sapidus* grown on corn). Average $\delta^{13}\text{C}\text{-CH}_4$ source signatures for each fungal species, considering all three substrates, are -52.6 mUr for *P. sapidus* and -61.3 mUr for *L. sulphureus*. These results suggest that the fungal species significantly influence the isotopic values of the emitted CH_4 ($p < 0.001$). A possible explanation for this observation could be the different enzyme sets of both fungi decomposing different components of the growth substrates, as *P. sapidus* belongs to white rot fungi and *L. sulphureus* is a brown rot fungus. However, detailed investigations of the metabolic pathways leading to CH_4 formation were beyond the scope of this study.

Furthermore, a significant effect of the growth substrate on $\delta^{13}\text{C}\text{-CH}_4$ source signatures was observed ($p < 0.001$). $\delta^{13}\text{C}\text{-CH}_4$ source signatures by *P. sapidus* were more positive compared to those of *L. sulphureus* when grown on grass ($\Delta = 16.3$ mUr) and corn ($\Delta = 13.6$ mUr) (Fig. 4). When grown on pine wood, $\delta^{13}\text{C}\text{-CH}_4$ source signatures were similar with *P. sapidus* showing slightly more negative values ($\Delta = -3.9$ mUr). Methane emitted by both fungi grown on corn was generally more enriched in ^{13}C (less negative $\delta^{13}\text{C}\text{-CH}_4$ source values) compared to the fungi grown on pine wood and grass. This might be easily explained by the $\delta^{13}\text{C}$ values of the growth substrates corn (-11.7 mUr, typical for C_4 -plants) being roughly 20 mUr less negative in their $\delta^{13}\text{C}$ values compared to the C_3 -plants pine wood (-28.0 mUr) and grass (-31.5 mUr).

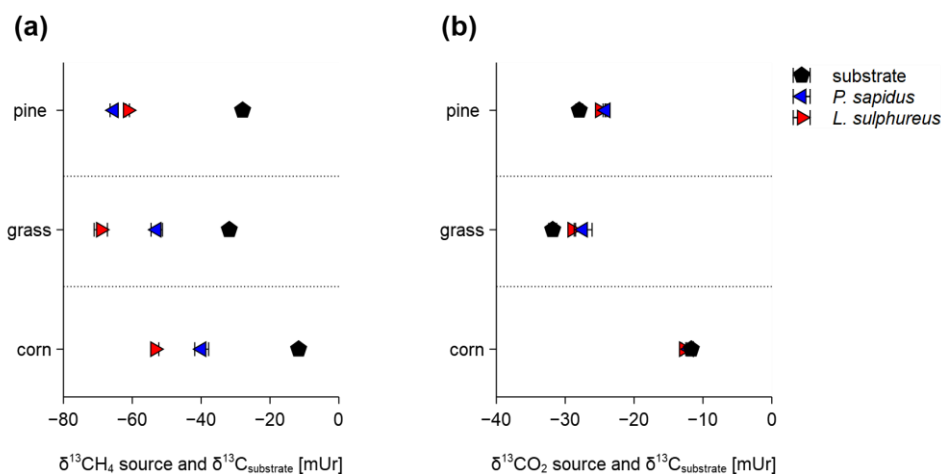


Figure 4: Calculated source signatures of $\delta^{13}\text{C-CH}_4$ values (a) and $\delta^{13}\text{C-CO}_2$ values (b) by *P. sapidus*, *L. sulphureus* and the $\delta^{13}\text{C}$ values of the substrate. Values are presented as mean values of the individual keeling plots with SD (n=3).

- 280 Comparison of calculated $\delta^{13}\text{C-CH}_4$ source signatures with measured bulk $\delta^{13}\text{C}$ values of the substrates shows that CH_4 emitted by both fungi is generally depleted in ^{13}C compared to the respective substrates (Fig. 4a). Based on this data we further calculated the apparent fractionation ($\epsilon_{\text{app CH}_4}$) between the $\delta^{13}\text{C-CH}_4$ source signatures and the bulk $\delta^{13}\text{C}$ values of the growth substrates. The apparent fractionation was calculated as so far no metabolic pathway for the formation of CH_4 in fungi is known and therefore currently only the initial $\delta^{13}\text{C}$ signatures of the substrates and the calculated $\delta^{13}\text{C-CH}_4$ source signatures of the
- 285 fungi can be compared. The values of $\epsilon_{\text{app CH}_4}$ are presented in Table 2 and range from -21.8 mUr (*P. sapidus* grown on grass) to -42.2 mUr (*L. sulphureus* grown on corn). When grown on pine wood $\epsilon_{\text{app CH}_4}$ values are similar for *P. sapidus* (-38.4 ± 1.2 mUr) and *L. sulphureus* (-34.4 ± 0.6 mUr). The differences in $\epsilon_{\text{app CH}_4}$ values between both fungal species are more distinct when grown on grass (*P. sapidus*: -21.8 ± 1.7 mUr, *L. sulphureus*: -38.6 ± 2.0 mUr) and corn (*P. sapidus*: -28.5 ± 2.0 mUr, *L. sulphureus*: -42.2 ± 1.1 mUr).
- 290 The calculated $\delta^{13}\text{C-CO}_2$ source signatures of both fungi (Table 2) range from -29.0 ± 0.5 mUr (*L. sulphureus* grown on grass) to -12.0 ± 0.3 mUr (*P. sapidus* grown on corn). $\delta^{13}\text{C-CO}_2$ source signatures are in a similar range for both fungi for all three substrates. Although CO_2 emitted by *L. sulphureus* is slightly more depleted in ^{13}C for all three substrates. Hence, an effect of fungal species on the stable carbon isotope values of CO_2 is significant ($p=0.008$). Also, the used substrates were found to influence $\delta^{13}\text{C-CO}_2$ values significantly ($p<0.001$).
- 295 The $\delta^{13}\text{C-CO}_2$ source signatures of the fungi show only small deviations from the bulk $\delta^{13}\text{C}$ values of the respective substrates (Fig. 4b). However, for both fungi grown on pine wood and grass, $\delta^{13}\text{C-CO}_2$ values are slightly less negative (a few mUr) compared to the bulk substrate. This observation is rather unexpected, as usually $\delta^{13}\text{C-CO}_2$ values are more negative with respect to $\delta^{13}\text{C}$ values of growth substrates due to fractionation during the metabolism (Bowling et al., 2008). However, when



grown on corn $\delta^{13}\text{C}$ - CO_2 source signatures by both fungi are slightly more negative compared to the substrate and calculated
300 $\epsilon_{\text{app CO}_2}$ values (Table 2) are -1.1 ± 0.3 mUr and $+4.6 \pm 1.3$ mUr for *L. sulphureus* grown on corn and *P. sapidus* grown on
grass, respectively.

The results of the incubation experiments show that there are distinct differences in the patterns of $\delta^{13}\text{C}$ - CH_4 and $\delta^{13}\text{C}$ - CO_2
values released by fungi. While the $\delta^{13}\text{C}$ - CO_2 source signatures are similar to the $\delta^{13}\text{C}$ values of the substrate (with $\epsilon_{\text{app CO}_2}$
values up to 4.6 mUr), the $\delta^{13}\text{C}$ - CH_4 source signatures deviate strongly from the respective substrate, with $\epsilon_{\text{app CH}_4}$ values of up
305 to -42.2 mUr. This either indicates that metabolic pathways leading to the formation of CH_4 and CO_2 have different
fractionation and/or that fungal CH_4 and CO_2 derive from different precursor compounds of the respective substrate. The used
growth substrates pine wood, grass and corn consist of various components including mainly cellulose, hemicellulose and
lignin at different proportions (in contrast if only using pure glucose or cellulose as growth substrate). Hence, the $\delta^{13}\text{C}$ - CH_4
and $\delta^{13}\text{C}$ - CO_2 source signatures might be dependent on specific metabolic pathways of the fungi but also on the chemical
310 composition of the growth substrate. Therefore, we suggest that the selected fungi and used growth substrates provide a first
solid basis for the potential range of $\delta^{13}\text{C}$ - CH_4 values that might occur in nature.

3.4 Fungal $\delta^{13}\text{C}$ - CH_4 values compared with known CH_4 sources

Figure 5 summarizes $\delta^{13}\text{C}$ - CH_4 values emitted by fungi in relation to other known CH_4 sources in the environment that have
been reported from the literature. The red bars indicate typical biogenic (formerly only considered to be produced by archaea)
315 CH_4 sources with emissions from wetlands, ruminants, landfills and rice paddies where $\delta^{13}\text{C}$ - CH_4 values are usually ranging
from -85 mUr to -40 mUr. Abiotic CH_4 sources (including thermogenic or pyrolytic processes) stemming from natural gas,
coal mining and biomass burning are characterized by less negative $\delta^{13}\text{C}$ values usually ranging from -55 mUr to -20 mUr. In
addition gas hydrates which might be formed by both microbial and abiotic processes cover a wider range of $\delta^{13}\text{C}$ values (-29
mUr to -73 mUr), depending on its forming mechanisms (Kvenvolden, 1995). The $\delta^{13}\text{C}$ source signatures of plant derived CH_4
320 have been reported to be in the range of -72 mUr to -45 mUr (Keppler et al., 2006; Vigano et al., 2009) depending on the three
photosynthetic pathways. Furthermore, there is a tendency towards more negative $\delta^{13}\text{C}$ - CH_4 values when the respective plant
was treated with UV radiation (Vigano et al., 2009). $\delta^{13}\text{C}$ - CH_4 source signatures of humans which might include both formation
by microbes in the gut but also from cellular processes show a rather wide range with values between -95 mUr and -56 mUr
(Keppler et al., 2016). The results of our experiments conducted with two fungal species and three different growth substrates
325 provide a range of $\delta^{13}\text{C}$ - CH_4 source values from -69 mUr to -40 mUr. This range overlaps with other eukaryotic sources, most
microbial CH_4 sources and even some abiotic CH_4 sources such as natural gas or emissions from coal mining.

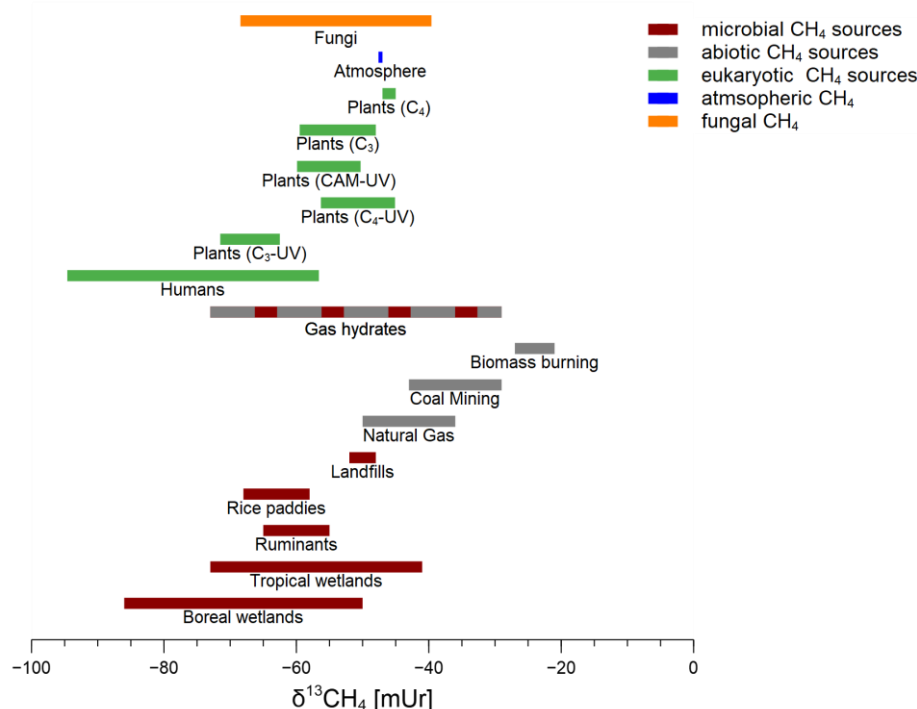


Figure 5: Range of $\delta^{13}\text{C}-\text{CH}_4$ values of microbial CH_4 sources (red), abiotic CH_4 sources (grey), eukaryotic CH_4 sources (green), atmospheric CH_4 (blue) and fungal CH_4 from this study (orange). The red and grey dashed bar indicates a mixture of microbial and abiotic CH_4 formation processes for gas hydrates (Kvenvolden, 1995). Data taken from (Brownlow et al., 2017; 330 Keppler et al., 2006, 2016; Kvenvolden, 1995; Nisbet et al., 2016; Quay et al., 1999; Vigano et al., 2009).

4 Conclusion

This study provided the first analysis of stable carbon isotope values of CH_4 emitted by two saprotrophic fungi that were grown on three different substrates. $\delta^{13}\text{C}-\text{CH}_4$ and $\delta^{13}\text{C}-\text{CO}_2$ source values were found to be dependent on the fungal species, as well 335 as the substrates decomposed by the fungi. $\delta^{13}\text{C}-\text{CH}_4$ source values of the fungi were found to be in the range of -69 mUr to -40 mUr and therefore overlap with $\delta^{13}\text{C}-\text{CH}_4$ values reported for other CH_4 sources such as methanogenic archaea, eukaryotes and from abiotic processes. Stable carbon isotope values of CH_4 in combination with flux measurements are often applied for a better understanding of regional and global CH_4 cycling. However, in recent years it has become clear that many biogenic CH_4 sources include complex CH_4 formation processes resulting in different isotopic fractionation patterns depending on 340 several biochemical and abiotic factors. Thus, studying ecosystems in which more than one major CH_4 source has to be expected (e.g. methanogenic archaea, fungi, cyanobacteria or plants) distinguishing between each individual source solely by stable carbon isotope values might be highly challenging. Therefore, additional tools are needed to better identify the sources



345 but also to disentangle sources and sinks. In future research the stable hydrogen isotopic values of CH₄ ($\delta^2\text{H-CH}_4$ values) or even applications of clumped isotopes might prove suitable tools for better distinction between different CH₄ sources and thus to better constrain the global CH₄ budget.

350 *Author contributions.* MS, KL, and FK conceived the study and designed the experiments; HZ provided fungal cultures, MS performed the experiments under the supervision of FK and KL; CE helped with gas measurements; MG measured stable isotope values of greenhouse gases; MS, FK, and KL analysed the data; MS, FK, HZ, MG, and KL, discussed the results, and MS, KL and FK wrote the paper.

Competing interests. The authors declare that they have no conflict of interest.

355 *Acknowledgments.* We thank Anette Gieseman for analytical measurements of stable carbon isotope values of the bulk substrates. We are grateful to Bianka Daubertshäuser for technical support with the cultivation of the fungi and to Lukas Kohl for encouraging us to perform this study. We acknowledge financial support by the Deutsche Forschungsgemeinschaft.

360 *Financial support.* This research has been supported by the Deutsche Forschungsgemeinschaft (grant nos. KE 884/8-2, KE 884/16-2) and (LE3381/1-1).

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