1 Characterization of particle-associated and free-living bacterial and

2 archaeal communities along the water columns of the South China Sea

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Abstract

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the South China Sea.

18	There is a growing recognition of the role of particle-attached (PA) and free-living (FL) microorganisms in
19	marine carbon cycle. However, current understanding of PA and FL microbial communities is largely on
20	those in the upper photic zone, and relatively fewer studies have focused on microbial communities of the
21	deep ocean. Moreover, archaeal populations receive even less attention. In this study, we determined
22	bacterial and archaeal community structures of both the PA and FL assemblages at different depths, from the
23	surface to the bathypelagic zone along two water column profiles in the South China Sea. Our results suggest
24	that environmental parameters including depth, seawater age, salinity, POC, DOC, DO and silicate play a
25	role in structuring these microbial communities. Generally, the PA microbial communities had relatively low
26	abundance and diversity compared with the FL microbial communities at most depths. Further microbial
27	community analysis revealed that PA and FL fractions generally accommodate significantly divergent
28	microbial compositions at each depth. The PA bacterial communities mainly comprise members of α - and γ -
29	Proteobacteria, together with some from Planctomycetes and δ -Proteobacteria, while the FL bacterial
30	lineages are also mostly distributed within α - and γ -Proteobacteria, but along with other abundant members
31	chiefly from Actinobacteria, Cyanobacteria, Bacteroidetes, Marinimicrobia and δ -Proteobacteria.
32	Moreover, there was an obvious shifting in the dominant PA and FL bacterial compositions along the depth
33	profiles from the surface to the bathypelagic deep. By contrast, both PA and FL archaeal communities
34	dominantly consisted of euryarchaeal Marine Group II (MGII) and thaumarchaeal Nitrosopumilales, together
35	with variable amounts of Marine Group III (MGIII), Methanosarcinales, Marine Benthic Group A (MBG-A)
36	and Woesearchaeota. However, the pronounced distinction of archaeal community compositions between PA
37	and FL fractions were observed at finer taxonomic level. A high proportion of overlap of microbial
38	compositions between PA and FL fractions implies that most microorganisms are potentially generalists with
39	PA and FL dual lifestyle for versatile metabolic flexibility. In addition, microbial distribution along the depth
40	profile indicates a potential vertical connectivity between the surface-specific microbial lineages and those in
41	the deep ocean, likely through microbial attachment to sinking particles.
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Keywords: particle-attached, free-living, marine microbe, vertical distribution, sinking particles, deep ocean,

45 1. Introduction

- The sinking of particulate organic matter (POM) formed in the photic layer is a fundamental process 46 that transports carbon and nutrient materials from the surface into the usually starved deep ocean, with 47 a significant role in structuring the distributions and activities of marine microorganisms in the dark 48 realm (Azam and Malfatti, 2007; Mestre et al., 2018; Suter et al., 2018). During sinking, the POM is 49 generally colonized and concurrently, decomposed by particle-attached (PA) prokaryotes, releasing 50 dissolved organic matter (DOM) into ambient seawater, fueling the free-living (FL) microbes (Kiorboe 51 52 and Jackson, 2001; Azam and Malfatti, 2007). It has been revealed that PA and FL microbial 53 populations exhibit different taxonomic composition, physiology and metabolism, corresponding to their lifestyle and ecological behavior. For example, PA bacteria, compared to FL bacteria, are often 54 55 larger in size (Alldredge et al., 1986; Zhang et al., 2007; Lauro et al., 2009) and metabolically more active (Karner and Herdl, 1992; Grossart et al., 2007). They often maintain higher levels of 56 extracellular enzymes, adhesion proteins and antagonistic compounds, and are capable of degrading 57 high-molecular-weight (HMW) organic compounds (Smith et al., 1992; Crump et al., 1998; Long and 58 59 Azam, 2001; Mevel et al., 2008; Ganesh et al., 2014). An examination of microbial metagenomes 60 suggests that there are notable differences between PA and FL assemblages in GC content, effective genome size, general taxonomic composition and functional gene categories (Smith et al., 2013). In 61 particularly, some broad key functional gene categories involved in DOM utilization (Poretsky et al., 62 2010; Rinta-Kanto et al., 2012) and specific functional gene groups linked to successive 63 decomposition of phytoplankton blooms (Teeling et al., 2012) are significantly different, indicating the 64 65 fundamental differences in survival strategies in relation to potentially available substrates. It is further revealed that PA microbes generally have larger genomes with a variety of metabolic and regulatory 66 capabilities of utilizing compositionally varied organic matter, while the genomes of FL microbes 67 usually are smaller with streamlined metabolic and regulatory functions that enable efficient adaption 68 69 to oligotrophic conditions (Smith et al., 2013; Yawata et al., 2014; Yung et al., 2016). Phylogenetically, PA and FL lineages generally exhibit different compositions. The PA fraction is relatively enriched in 70 71 members of y-Proteobacteria, Verrucomicrobia, Bacteroidetes, Firmicutes and Planctomycetes (Azam 72 and Malfatti, 2007; Milici et al., 2016; Salazar et al., 2016; Suter et al., 2018), while the FL 73 assemblages are often populated by members of *a-Proteobacteria* (SAR11 clade or *Ca.* Pelagibacter) 74 and Deferribacteres (DeLong et al., 1993; Crespo et al., 2013; Milici et al., 2017). However, significantly overlapped compositions of PA and FL microbial communities were also reported in a 75 few studies (Hollibaugh et al., 2000; Ghiglione et al., 2007; Ortega-Retuerta et al., 2013; Rieck et al., 76 77 2015; Liu et al., 2018a). Actually, most members of the PA and FL clades are generalists which switch their lifestyles via attachment and detachment to particles (Crespo et al., 2013; Li et al., 2015). As 78 revealed in many marine niches, α -Proteobacteria, γ -Proteobacteria and Bacteriodetes are the major 79 80 overlapped phyla in both PA and FL microbial fractions (Yung et al., 2016).
- Our current knowledge of PA and FL microbial populations largely relies on the upper photic ocean,
- 82 whereas little information is known from the deep dark ocean, which is the largest biome and
- accommodates more than half of the ocean's microbes (Aristegui et al., 2009; Salazar et al., 2016).
- Recently, a number of studies have revealed the PA and FL microbial communities in the bathypelagic
- waters (Li et al., 2015; Salazar et al., 2015; Milici et al., 2017; Mestre et al., 2018) or the deepest

- abyssal and hadal environments (Eloe et al., 2011; Tarn et al., 2016; Liu et al., 2018a). It is shown that
- PA and FL bacterial communities in the deep ocean have clear differences in abundance and
- 88 composition, in addition to the detection of novel, unknown prokaryotic taxa. Furthermore, although
- archaea are a major component of the marine ecosystem and play significant roles in the degradation
- 90 of organic materials (Iverson et al., 2012; Suzuki et al., 2017), PA and FL archaeal communities
- 91 receive less attention and little is known about them. Previous limited reports have observed
- 92 controversial results, as several studies showed that no obvious differences in archaeal community
- 93 structures between PA and FL assemblages (Galand et al., 2008; Eloe et al., 2011; Suzuki et al., 2017),
- 94 while a clear separation was found in recent reports (Tarn et al., 2016), with PA archaeal fraction
- 95 dominated by Marine Group II (MGII) and Marine Group III (MGIII), and FL archaeal fraction by
- 96 Marine Group I (MGI) and anaerobic methane-oxidizing archaea (ANME). In brief, it is not well
- 97 known about the changes of PA and FL prokaryotes along vertical profiles of water column, from the
- 98 surface to the deep bathyal, abyssal and hadal depths.
- 99 In this study, we analyzed and compared microbial compositions between PA and FL fractions at
- different depths along the vertical profile in the South China Sea (SCS). The SCS is a marginal sea
- located in the Northwest Pacific with a maximal depth of approximately 5,380 m (Fig. S1). Our results
- reveal diverse and significantly divergent microbial compositions in PA and FL fractions, and obvious
- community stratification at different depths along the vertical profiles.

2. Materials and Methods

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2.1 Sample collection and environmental parameter measurements

- Seawater samples were collected from two stations, G3 station, depth of 4,039 m at 117° 00.131′ E,
- 107 16° 59.947′ N, and J5 station, depth of 4,301 m at 114° 00.209′ E, 13° 59.958′ N, located in the central
- deep basin of the SCS during the Open Cruise of R/V *Dongfanghong* II from July 3 to 18, 2014 (Fig.
- 109 S1). Both stations have depth > 4,000 m, providing us the bathyal environments to vertically profile
- the variation of microbial assemblages with depth. A Sea-Bird CTD rosette sampler (SBE 911 plus)
- with 12 L Niskin bottles (Seattle, Washington, USA) was used to collect seawater from six different
- depths (50, 200, 1,000, 2,000, 3,000, and 4,000 m) at each station.
- Basic environmental parameters of the water column, including depth, salinity, temperature and
- dissolved oxygen (DO) were obtained in situ using the conductivity-temperature-depth (CTD) profiler
- and a DO sensor during the sampling. Once water samples were collected onboard, about 0.1 L of
- seawater was taken immediately for pH measurement using a pH meter (Mettle Toledo Inc.,
- 117 Switzerland).
- Approximately 8 L of seawater was filtered onboard through a Φ 142 mm precombusted glass fiber
- membrane (0.7 μm nominal pore size, Whatman, USA) under a gentle vacuum of <150 mm Hg for
- particulate organic carbon (POC) collection. The membranes were folded and stored at -20°C until our
- POC analysis. Then about 30 mL of filtered seawater of each sample was collected into 40 mL
- precombusted EPA vials and stored at -20°C immediately for DOC concentration measurement

- 123 (laboratory on land). About 200 ml filtered seawater at each depth was stored at -20°C for analysis of
- nutrients (NO₃-/NO₂-, dissolved inorganic phosphate and silicate). The remaining seawater was stored
- at -20°C for other analyses.
- 126 At each depth, we collected 4 L of seawater to obtain microorganisms for further analysis. Seawater
- was filtered first through a Φ 47 mm polycarbonate (PC) membrane of 3.0 μ m nominal pore size
- 128 (Millipore, USA) and subsequently, through a Φ 47 mm PC membrane of 0.22 μ m nominal pore size
- (Millipore, USA) to collect the PA and FL microorganisms, respectively (Eloe et al., 2011). To avoid
- damaging the membrane and the fragile particles, a relatively low vacuum pressure of < 10 mm Hg
- was used, and at the same time, the filtration time was no longer than 40 min for each membrance. The
- membranes were then frozen at -80°C until further microbial analysis.
- 133 Concentration of POC was determined with a PE2400 Series II CHNS/O analyzer (Perkin Elmer,
- USA) (Chen et al., 2008). DOC concentration was measured using a Shimadzu TOC-V Analyzer
- (Shimadzu Inc., Japan) (Meng et al., 2017). Nutrients were determined using a Four-channel
- 136 Continuous Flow Technicon AA3 Auto-Analyzer (Bran-Lube GmbH, German).
- About 1 L of seawater for each sample was sent to Beta Analytic, Inc. in Miami, Florida, for ¹⁴C
- radiocarbon dating with the Accelerator Mass Spectrometry (AMS) method as described in their
- website (https://www.radiocarbon.com/beta-lab.htm). When CTD rosette sampler came back on board,
- seawater for ¹⁴C dating was taken from Niskin bottles with first priority. To avoid the disturbance of
- air during the sampling, glass bottles were fully filled with flowing seawater with as little head space
- as possible. In addition, mercury chloride was added to prevent any microbiological influence.

2.2 DNA extraction

- In this study, we used the SDS-based method to extract the total DNA as described by Li et al. (2015)
- with minor modifications. The PC membranes containing seawater microbes were first cut into small
- pieces in a sterile petri dish and put into autoclaved 2 ml centrifuge tubes. 800 μL DNA extraction
- buffer consisting of 100 mM Tris-HCl, 100 mM sodium EDTA, 100 mM sodium phosphate, 1.5 M
- NaCl and 1% hexadecyl trimethyl ammonium bromide (CTAB) was added into each tube. The
- centrifuge tubes were frozen in liquid nitrogen and then thawed in a 65°C water bath. This procedure
- was repeated for 3 times. When the centrifuge tubes cooled down to room temperature proteinase K
- was added with a final concentration of ~ 0.2 mg mL⁻¹. The tubes were then incubated in a 65°C water
- bath for 2 h and shaked gently every about 30 min. Then, 800 μL phenol/chloroform/isoamyl alcohol
- 153 (25:24:1, v/v) was added into the centrifuge tubes and the tubes were shaked gently several times, and
- centrifuged at 12,000 ×g for 10 min. The supernatant was carefully transferred into new tubes and
- equal volume of chloroform/isoamyl alcohol (24:1, v/v) was added. The tubes were centrifuged at
- 156 12,000 ×g for 10 min. The aqueous layer was pipetted into clean 2 ml tubes, and 0.6 volume of cold
- isopropanol and 0.1 volume of 3M sodium acetate were added. The centrifuge tubes were incubated at
- -20°C for 1 h and centrifuged at 12,000 ×g for 10 min. The liquids were carefully discarded and DNA
- pellets at the bottom were gently rinsed with 70% pre-cooling ethanol. Finally, each DNA pellet was
- suspended into sterile deionized H_2O with a volume of 50 μL .

2.3 Pyrosequencing and analysis of 16S rRNA gene sequence amplicons

- Before PCR amplification, we first used the PicoGreen dsDNA Quantitation Kit (Life Technologies,
- USA) to quantify the concentration of DNA. DNA concentrations obtained varied between 4.48 and
- 164 29.1 ng/ μ L with a volume of ~ 50 μ L for each sample. For the PCR amplification of bacterial 16S
- 165 rRNA gene, the primer set 27F (5'-AGA GTT TGA TCC TGG CTC AG-3') and 533R (5'-TTA CCG
- 166 CGG CTG CTG GCA C-3') with 10-nucleotide barcodes were used, while Arch344F (5'-ACG GGG
- 167 YGC AGC AGG CGC GA-3') and Arch915R (5'-GTG CTC CCC CGC CAA TTC CT-3') containing
- 8-nucleotide barcodes were used for archaea (Ohene-Adjei et al., 2007; Sun et al., 2014). About 10 ng
- DNA template was amplified for PCR reaction. The PCR reaction condition was: firstly, 94°C, 5 min;
- then, 94°C, 50 s, 53°C, 50 s, and 72°C, 50 s, total 25 cycles; 72°C, 6 min. The products after PCR
- amplification were purified with the MiniBEST DNA Fragment Purification Kit (Takara Bio Inc,
- Japan) and then quantified using the NanoDrop 2000 (Thermo Scientific, USA). The pyrosequencing
- was carried out at the Majorbio Bio-Pharm Technology, Co., Ltd. (Shanghai, China) with the 454 GS-
- 174 FLX Titanium system (Roche, Switzerland).

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- QIIME 1.9.1 was used to perform the following phylogenetic analysis of pyrosequenced amplicons
- (Caporaso et al., 2010). As described in our previous study (Li et al., 2017), the low-quality reads were
- 177 first filtered with the following quantity control (QC) criteria: (1) the reads with ambiguous
- nucleotides; (2) the length of reads < 200 bp; (3) the reads containing > 5 bp homopolymers; (4) the
- reads with an average flowgram score < 25 in a quality window of 50 bp. The Operational Taxonomic
- Units (OTUs) were generated based on 3% cutoff of sequence similarity, and the longest sequence was
- picked as the representative sequence of each OTU for downstream analysis. The RDP classifier was
- used for the taxonomy assignment by against the SILVA 16S rRNA gene database (Version 132). The
- 183 ChimeraSlayer in the OIIME package was used to identify and exclude those of potential chimeras
- after alignment with PyNAST. In addition, the singletons were removed from the final OTU tables.

2.4 Diversity estimators and statistical analyses of microbial communities

- To avoid the variation caused by an unequal sequence number across samples, the OTUs abundance
- was normalized by resampling of sequences for each sample based on the sample with the least
- number of sequences. After resampling the sequences to the same number, diversity estimators
- including Chao 1 and Shannon's diversity (H) were calculated. Similarities among different microbial
- communities were determined using similarity matrices generated according to the phylogenetic
- distance between reads (Unifrac distance), and beta diversity of principal coordinates analysis (PCoA)
- was computed as components of the OIIME pipeline. The correlation between the microbial
- 193 community structures and environmental parameters was analyzed by canonical correspondence
- analysis (CCA). For the PCoA and CCA ordinations, the transformation of the resampled OTU
- abundance table was performed by taking the log of the sequence numbers. In addition, to testify the
- statistical significance of different groups identified by PCoA ordination, multiple statistical analyses
- including MRPP, ANOSIM and PERMANOVA were performed based on the resampled and
- 198 transformed OTU abundance table. Mantel test was also performed to testify the statistical
- 199 significance of environmental factors with microbial community compositions from the results of

- 200 CCA. All statistical analyses were performed in the R environment (v 3.2.1) using the Vegan package
- 201 (https://CRAN.R-project.org/package=vegan).
- In this study, we used the "odds ratio" to assess microbial preference to the PA or FL lifestyles. As
- defined by Ganesh et al. (2014), the formula of the "odds ratio" is as:
- odds ratio = log 10 (relative abundance in PA fraction / relative abundance in FL fraction)
- a positive value indicates the PA preference, while a negative value signifies the FL preference (Suter
- 206 et al., 2018).

2.5 Quantification of 16S rRNA gene and cell abundance estimation

- The copy number of microbial 16S rRNA gene for PA and FL fractions were estimated with 7500
- 209 Real-Time PCR System (Applied Biosystems, ThermoFisher, UK). The primer sets used were
- 341f/518r for bacteria (Dilly et al., 2004) and 344f/519r for archaea (Bano et al., 2004) with about 200
- bp amplified DNA fragments. The PCR products of bacterial and archaeal 16S rRNA gene were first
- cloned into a pUC18 plasmid vector (Takara Bio Inc, Japan), and then transformed into E. coli DH5α.
- 213 The recombinant plasmids were extracted and purified, and subsequently diluted 10-folds as the
- standards for real-time PCR reactions. R² for the standard curves varied between 0.994 and 0.996,
- 215 indicating a well linear relationship over the concentration ranges used in our study. PCR reaction was
- 216 carried out in a 20 μL amplification volume. The reaction mixture contained 1 μL of DNA template,
- 217 0.15 μM forward and reverse primers, and 10 μL Power SYBR Green PCR Master Mix (Life
- 218 technologies, UK). The PCR amplification conditions included: 95°C, 10 min to activate polymerase;
- 219 95°C, 15 sec, 60°C, 1 min, 40 cycles. A negative control was used to monitor potential contamination
- and agarose gel electrophoresis helped to confirm the absence of nonspecific amplification. Melt
- curves were generated using the Applied Biosystems real-time PCR system software with default
- thresholds. Each sample had triplicate amplifications and the average copy number of 16S rRNA gene
- 223 were calculated. To make a direct comparison between bacterial and archaeal abundances, we
- 224 converted copy number of 16S rRNA gene into cell abundance based on the assumption that on
- average, a bacterial cell has 4.08 16S rRNA gene copies while archaea contains 1.71 copies per cell
- 226 (Lee et al., 2009). Although the cell abundances inferred from the 16S rRNA gene copy number
- quantified by qPCR may be potentially biased, the estimation of cell abundances based on the qPCR
- of 16S rRNA gene has been confirmed as an effective method to reflect the approximate cell
- abundances in previous studies.

3. Results

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3.1 Environmental parameters of the water columns

- Fundamental environmental parameters, including temperature, salinity, pH, DO and DOC/POC are
- listed in Table 1. In general, they showed similar vertical trends with the normal pelagic ocean.

- Salinity increased gradually from ~ 33.84 at 50 m to ~ 34.52 at 200 m and 1,000 m, then maintained at
- around 34.60 at greater depths until 4,000 m. DO concentration was the highest ($\sim 204.5 \mu M$) at
- surface water, and decreased gradually to the lowest ($\sim 83.9 \,\mu\text{M}$) at 1,000 m depth, then increased
- gradually from $\sim 102.0 \,\mu\text{M}$ at 2,000 m to $\sim 113.5 \,\mu\text{M}$ at 4,000 m. Nitrite concentrations of the water
- columns at all depths were below the detection limit. Concentrations of nitrate, phosphate, and silicate
- were continuously increasing from the surface to 1,000 m depth, and then remained at relatively
- constant levels (Table 1).
- As expected, age of the seawater determined from $\Delta^{14}C_{DIC}$ was youngest at the surface and increased
- with depth linearly, varying from about 106 to 1650 years. The upper water layers (50 m and 200 m)
- from the two stations had the youngest and nearly the same ages, around 106 years. Ages of 1,000 m
- and 2,000 m in G3 station were almost identical, around 1,180 years, and increased to 1,600 years at
- 3,000 m and 1,750 years at 4,000 m. By contrast, age of 1,000 m in J5 station was \sim 1,310 years, and
- remained relatively stable below 1,000 m with the age of about 1,650 years (Table 1). DOC
- concentrations ranged from 63.07 to 40.34 µmol/L, with the highest at the surface and the lowest at the
- deep. However, POC concentrations varied greatly between 0.5 and 2.1 μmol/L and showed great
- variations. The POC concentrations were highest at 3,000 m of the G3 station (1.8 μmol/L) and at
- 250 1,000 m of the J5 station (2.1 μ mol/L) (Table 1).

3.2 Microbial cell abundances

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- The estimated abundances of bacteria and archaea were about $10^6 \sim 10^9$ cells L⁻¹ and $10^5 \sim 10^7$ cells L⁻²
- 253 ¹, respectively (Fig. 1). The FL bacterial fraction generally accommodated higher cell abundances
- (varying from 0.62×10^7 to 1.65×10^8 cells L⁻¹), several times higher than their corresponding PA
- fraction $(1.85\pm0.02\times10^6\sim1.90\times10^8~cells~L^{-1})$. However, one remarkably lower abundance of FL
- bacterial fraction than PA fraction was detected in the surface water (50 m) of the G3 station where PA
- bacterial abundance was up to 1.70×10^9 cells L⁻¹, two orders of magnitude higher than that of the FL
- fraction $(1.62 \times 10^7 \text{ cells L}^{-1})$ (Fig. 1a). Similar to bacteria, the FL archaeal fractions usually showed
- 259 higher cell abundances than their PA fractions (Fig. 1b). The only exception was also at the depth of 50
- 260 m of G3 station where the estimated PA archaeal cell abundance $(6.50\pm0.01\times10^7 \text{ cells L}^{-1})$ was much
- higher than that of FL archaeal fraction (1.01 \times 10⁶ cells L⁻¹). FL archaeal fraction had the cell
- abundances between 2.70×10^5 and $8.62 \pm 0.03 \times 10^6$ cells L⁻¹, while PA archaeal fractions fluctuated
- between 1.28×10^5 and $6.50 \pm 0.01 \times 10^7$ cells L⁻¹ (Fig. 1).

3.3 Estimation of microbial diversity

- Totally 91,692/81,332 and 72,590/93,059 valid sequences of bacterial 16S rRNA gene were obtained
- 266 for FL/PA fractions of G3 and J5 stations, respectively. Based on the 97% similarity, these FL and PA
- bacterial sequences were defined into a total of 6,320 operational taxonomic units (OTUs) in which
- 268 1,982 OTUs belonged to singletons and were finally removed from the valid OTU table (Table S1).
- 269 Correspondingly, 50,727/41,511 and 44,443/37,751 archaeal sequences were determined for FL/PA
- archaeal fractions of G3 and J5 stations. Attempt to determine PA archaeal sequence from 3,000 m
- depth of G3 station and 4,000 m depth of J5 station failed because of technical reasons. A total of

- 272 1,070 archaeal OTUs were defined and 329 OTUs were considered as singletons (Table S2). The
- sequencing depths of 16S rRNA gene were shown in their rarefaction curves (Fig. S2).
- 274 Shannon's diversity (H) and Chao1 were calculated to estimate microbial diversity of both PA and FL
- 275 fractions at all depths (Fig. 2 and Fig. S3). In most cases, the H indices of the bacterial FL fractions
- were usually higher than their PA counterparts at each depth (Fig. 2). H index of FL and PA bacterial
- 277 fractions gradually increased from 50 to 1,000 m, decreased in the intermediate water of around 2,000
- 278 m depth, and increased again at 3,000 and 4,000 m (Fig. 2a). Archaeal H index varied along the
- vertical profiles with a trend similar to bacteria, and FL archaea generally had higher H index values
- 280 than the PA fraction (Fig. 2b). In addition, it was further observed that even at the same depth, the
- values of H index between two stations fluctuated a lot. Chaol index showed nearly similar variation
- trends for both PA and FL microbial fractions (Fig. S3).
- PCoA analysis revealed that there were significant differences (P values <0.05, Table S3) in bacteria
- and archaea community structures over the depth profiles and between the FL and PA fractions.
- Overall, three groups were distinguished, the surficial 50 m group, the FL group, and the PA group
- 286 (Fig. 3). One group, consisted exclusively of samples at 50 m depth, separated the microbes in the
- surface from those in the rest of the water column of both stations, irrespective of microbial lifestyles
- 288 (FL or PA). However, the other two groups were separated mainly based on the FL and PA lifestyles. It
- is interesting to note that the FL bacterial samples clustered into one group where samples were further
- 290 partitioned with respect to depth (Fig. 3a). Canonical correspondence analysis (CCA) showed that
- fundamental environmental parameters including depth, DO, salinity, seawater age, DOC and POC
- 292 concentration, and silicate exerted potential impact on variations of FL and PA microbial communities
- along the water column (Fig. 4, Fig. S4). Mantel test further indicated that all those factors, except
- POC concentration (P = 0.164), were the statistically significant variables associated with variation of
- 295 PA and FL fractions (P = 0.001).

3.4 Taxonomic compositions of the PA and FL bacterial and archaeal fractions

- 297 Taxonomic compositions of FL and PA bacterial fractions and their relative abundances are presented
- in Fig. 5. At phylum level, bacterial sequences were mainly assigned into *Proteobacteria* (α -, β -, γ -,
- and δ -), Actinobacteria, Cyanobacteria, Planctomycetes, Bacteroidetes, Marinimicrobia (SAR406)
- 300 clade), Chloroflexi, Firmicutes, Acidobacteria, Gemmatimonadetes, Nitrospinae and Verrucomirobia.
- 301 The taxa at ~ family level with relatively high abundances (>3%) on average in either PA or FL
- fraction were further shown in Fig. 6.
- 303 It is clear that α and γ -Proteobacteria were the dominant lineages in both the FL and PA fractions at
- nearly all depths. In most cases, the sum of α and γ -Proteobacteria accounted for $\sim 40\%$ to nearly
- 305 90% (Fig. 5). Moreover, their relative abundances in different PA and FL fractions and different
- stations also varied widely. Within the α -Proteobacteria, the dominant families included
- 307 Methylobacteriaceae, Phyllobacteriaceae, Rhodobacteraceae and Erythrobacteraceae (Fig. 6).
- 308 Members of the families Methylobacteriaceae and Erythrobacteraceae occurred commonly in both
- fractions at almost all depths but usually with higher proportions in PA fractions. The family
- 310 Rhodobacteraceae occurred commonly in both fractions at every depth (1 $\% \sim 20\%$), while the

Phyllobacteriaceae was dominantly distributed in the PA fraction of 2,000 m depth of J5 station with > 311 60% proportions. In addition, another important lineage within α -Proteobacteria is SAR11 clade (now 312 named as Pelagibacterales) (Grote et al., 2012). It was clearly revealed that SAR11 clade showed 313 relative higher abundances in FL fractions than PA fractions. Moreover, at depths above 1,000 m, 314 SAR11 clade had a far higher proportion than the deep ocean and the maximum levels occurred at 200 315 m depth (20% ~ 24%) (Fig. 6, Table S1). y-Proteobacteria was another lineage with the highest 316 317 abundance overall. Its relative abundances changed significantly with depths and in different fractions. The minimum abundances were only $1\% \sim 5\%$, while the maximum were up to $73\% \sim 80\%$ (Fig. 5 318 and Table S1). Moreover, G3 station generally had higher γ-proteobacteria proportions than that of J5 319 station on average. As shown in Fig. 6, although sequences of y-Proteobacteria were classified into 320 321 multiple families, actually only two families Alteromonadaceae and Pseudoalteromonaodaceae 322 exhibited absolutely dominant prevalence in the bacterial populations. The Pseudoalteromodaceae populated predominantly the PA fractions in 50 m and 200 m depths ($66\% \sim 75\%$), while the 323 324 Alteromonadaceae mainly dominated the PA fractions in the deep water, particularly at 2,000 m and 3,000 m depths. δ -Proteobacteria also had a common distribution in both fractions of all depths, 325 usually accounting for less than 10% proportions in most samples (Fig. 5), and SAR324 clade 326 members contributed significantly to the dominance of the δ -Proteobacteria (Fig. 6). Actinobacteria 327 328 and Cyanobacteria were abundantly distributed only in the surficial 50 m depth, and by sharp contrast, their proportions in other depths were less than 5%. Other bacterial lineages which had a wide 329 distribution in all depths but only with minor abundances in both fractions included *Planctomycetes*, 330 Bacteroidetes, Marinimicrobia (SAR406 clade), Chloroflexi, β-Proteobacteria, Firmicutes, 331 Gemmatimonadetes and Verrucomicrobia (Fig. S5). 332

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Majority of archaeal amplicons were mainly fallen into the Nitrosopumilales and several uncultured taxonomic lineages (Fig. 7 and Fig. S6). Both FL and PA archaeal fractions at all depths were principally populated by the order Nitrosopumilales (formerly referring to MGI.1a, a subclade of MGI) (Qin et al., 2017) of the *Thaumarchaeota* and Marine Group II (MGII) of the *Euryarchaeata*. Members from the Nitrosopumilales and MGII lineages generally contributed more than 80% relative abundances in their respective clone libraries. The Nitrosopumilales was always one of the most abundant clades along the vertical profiles except in the topmost FL and PA fractions. MGII clade exhibited a wide distribution along the water columns, and it usually accounted for the large proportions in both archaeal size fractions. The photic layer (~ 50 m depth) contained the highest abundances of MGII clade, particularly in FL fractions with up to ~80% proportions. By sharp contrast, the lowest abundances of MGII occurred at 2,000 m (G3 station) and 3,000 m (J5 station) depths, making up <20% proportions. The third most abundant clade overall is Marine Group III (MGIII) of the Euryarchaeata. MGIII representatives were mainly dispersed in the FL fractions with $5\% \sim 18\%$ abundances, while they were absent from most of the PA fractions. However, the relative abundances of MGIII members in PA fractions of 1000 m depth could be as high as 30% ~ 45% (Fig. 7). The order Methanosarcinales of Euryarchaeata was detected commonly in most PA fractions, but it had the higher abundance only in the upmost 50 m depth ($\sim 29.7\%$) (Fig. 7). Another sample accommodating relatively much Methanosarcinales was the PA faction of 3,000 m in J5 station with 9.1% proportion. Within the Euryarchaeata, another clade of methanogens, Methanobacteriales, was also detected from both size fractions but with low relative abundances (<5%) (Fig. 7, Fig. S6, Table S2). In addition, other archaeal lineages included Woesearchaeota (formerly known as the DHVEG-6 group), Miscellancous Crenarchaeotic Group (MCG, now named as Bathyarchaeota), the

- 355 *Halobacteriales* of the *Euryarchaeata* and Marine Benthic Group A (MBG-A) of the *Thaumarchaeota*.
- 356 They just provided a limited contribution to archaeal populations (Fig. S6).

3.5 Bacterial preference to PA or FL lifestyles

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- Odds ratio was used to assess the preference of bacterial taxonomic lineages to the PA or FL lifestyle.
- 359 A positive odds ratio indicates PA preference or higher abundance in the PA fraction, while a negative
- value suggests FL preference or higher abundance in the FL fraction. The bacterial lineages
- dominating the PA fractions come exclusively from α and γ -Proteobacteria with some relatively
- abundant δ -Proteobacteria and Planctomycetes at specific depths (Fig. 5). By contrast, although the
- predominant lineages of FL fractions also mainly consisted of members of α and γ -Proteobacteria,
- other abundant lineages were more diverse including *Actinobacteria*, *Cyanobacteria*, *Bacteroidetes*,
- 365 *Marinimicrobia* and δ -*Proteobacteria*, as shown in Fig. 5. As shown in Fig. 8, we listed those lineages
- at \sim family level with high proportions (> 1%) with their odds ratios along the depth profiles. It was
- 367 suggested that most of the absolutely dominant families of PA fractions comprised of the
- 368 Phyllobacteriaceae and Methylobacteriaceae (α-Proteobacteria), Pseudoalteromonadaceae and
- 369 Alteromonadaceae (γ-Proteobacteria) (Fig. 6) showed a preference to PA lifestyle. However, the α-
- 370 proteobacterial *Rhodobacteraceae* and *Erythrobacteraceae* prevailing in PA fractions preferred to
- different lifestyles at different depths (Fig. 8). Compared with those PA-preferred lineages, there is a
- wider range of lineages showing preference to FL lifestyle. These phylogenetic lineages are mainly
- populated by the OM1 clade and Sva0996 marine group (Actinobacteria), Nitrospinaceae
- 374 (Nitrospinae), Planctomycetaceae (Planctomycetes), SAR11 clade (α-Proteobacteria), SAR324 clade
- 375 (δ -Proteobacteria), SAR86 clade and Thioglobaceae (γ -Proteobacteria). It is important to point out
- that a considerable number of bacterial lineages exhibited their preferences to both PA and FL
- 377 lifestyles, though preferring differently at different depths or locations (Fig. 8). Actually, at OTU level,
- less than 1/2 of the total OTU (2005 out of 4338 OTUs) were shared by PA and FL fractions (Fig. S7).
- Phylogenetically, these PA/FL-shared OTUs were mostly fallen into α -, γ -, δ -Proteobacteria,
- 380 Planctomycetes, Chloroflexi, Bacteroidetes, Marinimicrobia and Actinobacteria. The taxonomic
- 381 components of the PA/FL-shared OTUs at different levels are approximately similar to OTUs retrieved
- exclusively from either the PA fractions or the FL fractions (Table S1, Fig. S7).

4. Discussion

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4.1 Comparison of microbial abundance and diversity between PA and FL fractions

- PA bacterial and archaeal fractions show generally lower abundance and taxonomic richness than their
- FL counterparts and constitute a small fraction of the total abundances. Our results are consistent in
- principle with previous reports on various pelagic environments, in either the euphotic zone, twilight
- or the dark deep ocean (Turley and Stutt, 2000; Simon et al., 2002; Ghiglione et al., 2007; Rieck et al.,
- 389 2015). However, in some eutrophic and notably particle-rich marine ecosystems, for example, marine
- 390 snow or estuaries, PA bacterial fractions were present in higher local concentrations and greater
- diversity than FL bacteria (Caron et al., 1982; Karner and Herndl, 1992; Turley and Mackie, 1994;

Garneau et al., 2009). In upper photic zone, PA bacterial abundance and their contribution to total 392 bacterial biomass are highly variable, and depend largely on the quantity and quality of suspended 393 organic particles (Cammen and Walker, 1982; Simon et al., 2002; Doxaran et al., 2012). This is indeed 394 the case in the South China Sea. As shown in Fig. 1, at 50 m and 200 m depths of G3 station, PA 395 bacterial abundances outnumbered FL bacteria by nearly 2 ~ 100 times, whereas J5 station has an 396 opposite trend. However, as shown in Table 1, these two stations have almost the same environmental 397 398 parameters, particularly in POC concentrations. One possibility may be that G3 and J5 have different POC compositions, attributable to different origins of organic matter (Chen et al., 2015; He et al., 399 2016; Liang et al., 2018). Although bacteria attaching to particles are of relatively lower abundance 400 compared to free-living cells in the pelagic ocean, they are consistently metabolically more active with 401 higher extracellular enzymatic activities (Karner and Herndl, 1992) and cell-specific thymidine 402 incorporation rates (Turley and Mackie, 1994; Turly and Stutt, 2000). Therefore, PA bacteria often 403 play a comparable role to free-living bacteria in hydrolysis or decomposition of marine organic matter, 404 405 biomass production and carbon cycling (Griffith et al., 1994; Turly and Stutt, 2000; Liu et al., 2015). The decline of bacterial abundance and richness along the depth profile is largely owing to the gradual 406 decreasing availability of usable organic carbon (Smith, 1992; Turly and Stutt, 2000; Jiao et al., 2014). 407 It is interesting that the mid-water around 2000 m depth showed the lowest bacterial diversity (Fig. 2, 408 409 Fig. S3). One possibility is that 1,500-2,000 m is roughly a boundary for different water masses in the 410 deep, central basin of the South China Sea. The deep water masses (>2600 m) of the central basin coming from the western Pacific Ocean through the Bashi Channel are relatively rich in nutrients than 411 the mid-water masses of the SCS. Therefore, it may cause a relative increase in microbial diversity in 412 deep water masses such as those at 3,000 m and 4,000 m. In addition, some "old, deep" water from the 413 bottom of the central basin will also rise to around 2,000 m depth because of the basin-scale 414 circulation. These old waters are relatively enriched in refractory DOC (RDOC), remained after 415 microbial utilization of labile DOC during their circulation, potentially reducing microbial diversity. 416 This hypothesis is partly supported by the seawater age at J5 station. It is shown that the age of 417 seawater at 2,000 m depth of J5 station is 1,670 years, roughly equal to those of deep waters at 3,000 418 419 m and 4,000 m (1,680 years and 1,610 year). In contrast, archaea are commonly much lower in cell abundance and community diversity compared with their bacterial counterparts at the same depths 420 421 (Fig. 1, Fig. 2 and Fig. S3). The relative abundance of archaeal populations in total prokaryotes 422 increases gradually with depth, indicative of a potential rising impact on biogeochemical cycle in marine environments. In addition, pronounced distinction in microbial community structures of PA 423 and FL assemblages were observed along the depth profile, which were well supported by results of 424 425 statistical analyses (Fig. 3). It is expectable that PA microbial fraction differs from FL fraction, 426 considering their discrepant activity patterns for survival. Related discussions are shown below.

4.2 Environmental factors potentially shaping microbial community structure

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Several environmental parameters played a pivotal role in structuring microbial communities of seawater. Hydrological condition (e.g. depth), together with age and salinity of water mass, are a key subset of environmental drivers (Fig. 4). Recent studies have shown that microbial populations in the meso-/ bathypelagic ocean are largely dissimilar to those of the epipelagic zone (Salazar et al., 2015; Milici et al., 2017; Liu et al., 2018a), indicative of a crucial environmental selection process exerted by depth. In our study, PCoA analysis revealed that PA and FL fractions from the surficial zone (50 m)

were clustered into a separate but relatively loose group distant from other depths (Fig. 3), indicative 434 of the influence imposed from depth in shaping microbial community structures. Several bacterial 435 lineages, including Cvanobacteria, Actinobacteria, δ-Proteobacteria, Marinimicrobia (SAR406 clade) 436 and Firmicutes with distinct distributing stratification contribute to this dissimilarity (Fig. 5). 437 Cyanobacteria and Actinobacteria belong to typical phototrophs (Mizuno et al., 2015) and they are 438 prevalently distributed in euphotic zones. By contrast, δ-proteobacterial SAR324 clade, as shown in 439 440 our results, are primarily found in mesopelagic waters (200 ~ 1,000 m) (Fuhrman and Davis, 1997; Wright et al., 1997). SAR406 clade has a ubiquitous distribution across diverse marine niches, 441 however, its high abundance always occurs within the mesopelagic zones, ~ five times or higher than 442 in surface ocean (Yilmaz et al., 2016). Archaeal population components also reflect the impact of 443 depth. Euphotic zones hold less abundant thaumarchaeotal Nitrosopumilales and more euryarchaeotal 444 445 Methanosarcinales and Woesearchaeota (Fig. 7), while marine thaumarchaeotal groups are more abundant in meso- and bathypelagic waters (Karner et al., 2001; Mincer et al., 2007; Varela et al., 446 447 2008). In addition, Salazar et al. (2016) found that sampling depth appears to have a more direct impact on free-living bacterial communities. Our results are highly consistent with this observation in 448 449 that FL bacterial fractions from the same depth grouped together irrespective of their sampling locations (G3 or J5 station) (Fig. 3a). 450

DO concentration is observed to strongly affect particle flux and particle transfer efficiency from 451 euphotic zone to the deep sea since remineralization of organic particles appears to be oxygen-452 dependent (Laufkotter et al., 2017; Cram et al., 2018). DO is considered as one of the most crucial 453 environmental variables for shaping the compositions of particle-attached bacterial assemblages 454 (Salazar et al., 2016). Some taxonomic lineages are directly affected by oxygen. For example, a recent 455 456 study found that oxygen is one of the key factors driving the distribution and evolutionary diversity of 457 Woesearchaeota (Liu et al., 2018b). POC and DOC can be substrates for both PA and FL communities, respectively (Azam and Malfatti, 2007; Zhang et al., 2016; Liu et al., 2019). However, POC 458 concentration in the present study is not statistically significantly correlated with either bacterial or 459 460 archaeal community abundances (P values >0.05) (Table S3). We hypothesize that the quality rather than the quantity of POC imposes a decisive influence on microbial populations, especially in the 461 462 deep, dark ocean. During POC sinking from surface through the water column, and also as seawater 463 ages, the labile organic matter becomes increasingly decomposed, while the more refractory material remains and resists further degradation (Simon et al., 2002). In such cases, utilization of the POC in 464 the deep sea by microorganisms depends on the quality and quantity of the remaining POC. 465 Meanwhile, in older seawater, DOC also become more refractory because free-living microorganisms 466 467 preferentially utilize labile DOC and the remained refarcotory DOC gradually accumulates, which potentially affect microbial community structures. Among common nutrients, silicate exhibited 468 statistically significant correlation with microbial distributions (Fig. S4), and this is unexpected 469 because the SCS generally exhibits N- or P-limited phytoplankton production (Wu et al., 2003; Chen 470 et al., 2004). However, recent research found that near the sampling site of this study, there is a clear 471 silicon deficiency in the euphotic zones shallower than 75 m (Huang et al., 2015), which directly 472 influences the diversity and biomass of phytoplankton (for example, diatom), and consequently, the 473 474 quantity and quality of POM transported to the deep along the vertical water columns, and finally 475 exerts a potential impact on microbial communities. Some bacterial lineages such as the Rhodobacteraceae, Flavobacteriaceae, Oceanospirillaceae and SAR11 clade, commonly retrieved in 476 477 our present study, have been confirmed to be closely related to marine diatom blooms (Zhang et al.,

2018; Monnich et al., 2020). Actually, microbial community structure and their distribution along the water column profile are a comprehensive combination impacted by multiple environmental variables.

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4.3 Specialist or generalist for PA and FL lifestyle: clues from bacterial community compositions

It was indicated that PA and FL bacterial fractions generally accommodated different community 481 compositions along the depth profiles (Fig. 3), consistent with previous reports in various marine 482 habitats (Acinas et al., 1997; Moeseneder et al., 2001; Ghiglione et al., 2009; Salazar et al., 2015). 483 However, in most cases, taxonomic compositional disparity between the two filtration fractions does 484 not seem much apparent at least at phylum level (Fig. 5). Actually, a few studies also confirmed that at 485 high taxonomic ranks, bacteria show conserved lifestyles either in association with particles or as free-486 living microorganism (Eloe et al., 2011; Salazar et al., 2015; Liu et al., 2018a). The pronounced 487 488 contrast in population compositions of the two filtration fractions was unveiled only at greater taxonomic level and a considerable number of phylogenetic taxa exhibited different preferences to PA 490 or FL lifestyles. It was shown in Fig. 5 and Fig. 6 that as the most abundant members, α - and γ -Proteobacteria occurred prevalently in both filtration fractions, but at the family level, most of 491 492 predominant bacterial lineages of PA and FL fractions were significantly divergent, indicating their 493 preference to different microhabitats shaped by organic particles and environmental parameters. The dominant lineages in PA fractions were mainly associated with the families Pseudoalteromonadaceae and Alteromonadaceae within y-Proteobacteria, and the Methylobacteriaceae and Phyllobacteraceae 495 within α -Proteobacteria. These γ -proteobacterial members are usually retrieved from diverse marine 496 habitats as the typical PA clades, and they are believed to have the abilities to degrade/utilize HMW 497 498 organic compounds with higher nutrient requirements (DeLong et al., 1993; Crespo et al., 2013). The adhesion to particles could make them increase nutrients acquisition and avoid the nutrient-depleted 499 conditions (Crespo et al., 2013). By contrast, members of α -Proteobacteria are rarely reported as the dominant lineages of PA fraction or particle-attached preference (Crespo et al., 2013; Rieck et al., 2015; Suzuki et al., 2017), which is inconsistent with our results revealing α-proteobacterial lineages frequently prevailed as PA members. Further phylogenetic analysis revealed that the majority of α proteobacterial PA members belonged to the genus Methylobacterium which are strictly aerobic, facultatively methylotrophic bacteria, and can grow on a wide range of carbon compounds (Green, 2006). They probably benefit from the particle-attached lifestyle, making their high requirements for organic matters easily to achieve. Compared with bacterial PA counterparts, FL bacterial communities are more diverse, and dominant populations are scattered in more phylogenetic taxa with relatively 509 homogeneous proportions (Fig. 8). Among the predominant lineages, the actinobacterial OM1 cade and cyanobacteria dominate the upper surficial waters (Fig. 6), likely attributed to their phototrophic behaviors. Although actinobacteria are recognized as ubiquitous members of marine bacterioplankton (Giovannoni and Stingl, 2005), they are scarcely reported with predominance (Milici et al., 2016a). Ghai et al. (2013) revealed the OM1 clade members possess the smallest cell sizes with streamlined genome, representing a typical adaption to oligotrophic condition (Giovannoni et al., 2014) which well 514 agrees with the oligotrophic environments in the SCS (Gong et al., 2012). Other predominant FL lineages include α-proteobacterial SAR11 clade, δ-proteobacterial SAR324 clade, and Marinimicrobia (SAR406 clade), all usually being the most ubiquitous free-living bacterial lineages and dominantly distributed in epi- and mesopelagic zones (Grote et al., 2012; Tarn et al., 2016; Yilmaz et al., 2016; Milici et al., 2017; Liu et al., 2018a). Genomic information suggests that although these clades have a

flexible metabolism utilizing multiple hydrocarbon compounds, they generally lack of carbohydrateactive enzyme genes for the attachment to and the degradation of particulate organic matter (Peoples et al., 2018), consistent with their preference to free-living lifestyle rather than particle-attachment (Eloe et al., 2011; Salazar et al., 2015; Tarn et al., 2016). In addition, the percentages of SAR11 clade revealed here seem to be relatively lower compared with those reported in previous studies where the SAR11 clade typically makes up 20 to 40% of the bacterioplankton (Morris et al., 2002; Aprill et al.,

2015). It may be related to the sequencing primers used which potentially cause underestimation of

527 SAR11 clade and bias the interpretation of their relative abundances (Aprill et al., 2015).

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In addition to those predominant lineages mentioned above, there are a couple of bacterial taxa showing evident PA or FL preferences. At ~ family level, these PA- or FL-preferred taxa are well hinted by their odds ratio between PA and FL fractions. These bacterial lineages are characterized by low abundances or occasional occurrences in water columns (Fig. 6, Table S1) but high odds ratio (absolute value) (Fig. 8), indicating their strong preferential divergence in the two size fractions. The majority of these lineages are recorded consistently about their PA- or FL preferences in previous studies, and commonly possess the ability to hydrolyze and utilize complex carbon sources. Although their abundance is low, these relatively minor populations can still effectively influence local microhabitats because of their high specificity for organics. In contrast, there are still some populations which are scarcely reported. For example, Sva0996 marine group, an actinobacterial group, is retrieved occasionally from marine sediments and upper ocean (Bano and Hollibaugh, 2002; Wang et al., 2018). Orsi et al. (2016) first found this group prefers to free-living lifestyle in upper seawater and have the ability to assimilate phytoplankton-derived dissolved protein. Our present results suggest that Sva0996 group are flexible to adapt PA or FL lifestyles at the surface seawater because two lifestyles occur concurrently. Moreover, the distribution of Sva0996 group is not restricted only in upper photic ocean, and they can survive in meso- and bathypelagic seawaters with the significant preference for free-living lifestyle (Fig. 8). However, due to lack of pure culture or their genome information, it is not yet possible to elaborate their preferences for PA and FL lifestyles.

A high proportion of bacterial lineages are revealed to co-occur in both PA and FL fractions (Fig. 8 and Fig. S7), indicating that a considerable amount of bacterial lineages potentially have PA and FL dual lifestyle strategies. On the one hand, as shown in Fig. 6, a few bacterial lineages co-occur in PA and FL fractions at least at one of the same depths with approximately equivalent abundances. In such cases, their odds ratios are close to zero or minor range (Fig. 8), indicating that these bacteria are able to employ two different survival strategies at the same time. On the other hand, lots of taxa show divergent preferences to PA or FL lifestyles at different depths or different locations. This is clearly evident by the shift or conversion of their odds ratios at different depths along the vertical profiles of water column (Fig. 8), indicative of their different adaption tactics to different environments. One possible explanation is that most of the marine bacteria are generalists with dual life strategies (Bauer et al., 2006; Gonzalez et al., 2008), and able to grow in suspension as well as on particles (Lee et al., 2004; Grossart et al., 2006, 2010). For instance, PA bacteria must be capable of surviving freely in the water column to migrate and colonize new organic particles (Ghiglione et al., 2007; Crespo et al., 2013). Bacterial populations may switch their lifestyles between free-living and particle-attachment, depending on substrate availability and the surrounding chemical triggers (Grossart, 2010; D'Ambrosio et al., 2014). To date, one exception, the genus Scalindua in the Planctomycetes phylum, which is a known marine chemoautotroph involved in anammox, is exclusively observed in FL

fractions in previous studies (Fuchsman et al., 2012; Ganesh et al., 2014; Suter et al., 2018). However, it is absent from our water columns.

4.4 Archaeal community preferences to PA and FL lifestyles

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Samples of PA and FL archaeal fractions were also separated into different groups by statistical 566 analysis (Fig. 3b, Table S3), indicating their different community structures. However, because most 567 of OTUs belonged to uncultured archaeon, it is impossible to assign them into taxonomic lineages at 568 finer level. Thus, the distinction of archaeal population compositions between PA and FL fractions was 569 unnoticeable (Fig. 7). The Nitrosopumilales under MGI and MGII are the most abundant taxa in both 570 571 PA and FL archaeal fractions. The thaumarchaeal Nitrosopumilales are one of the most abundant and cosmopolitan chemolithoautotrophs in the dark ocean (Konneke et al., 2005) and responsible for much 572 573 of the ammonia oxidation in this environment for their common metabolism of aerobic ammonia oxidation. Corresponding to their autotrophic metabolism, MGI (including *Nitrosopumilales*) 574 generally exhibit free-living preference and are the prevalent archaeal taxa in free-living fractions 575 below euphotic zone (Smith et al., 2013; Salazar et al., 2015; Tarn et al., 2016). However, different 576 577 from our results, a few studies showed that MGI dominated both the PA and FL archaeal populations and no obvious distinction was observed in abundance and ecotype of MGI (Eloe et al., 2011; Jin et 578 al., 2018). To date, only a few pure cultures of marine MGI, small rods with a diameter of 0.15~0.26 579 μ m and a length of 0.5 ~ 1.59 μ m and no flagella were observed (Könneke et al., 2005; Qin et al., 580 2014), suggesting that their occurrence in PA fraction is not caused by pore size of filter to fractionate 581 different assemblages. One possibility is that decomposition of organic particles continuously releases 582 583 ammonia and MGI can easily acquire high concentrations of ammonia by attaching to particles, especially in oligotrophic area. Recent studies provide another explanation to particle-attached MGI 584 that some MGI cultures are obligate mixotrophy that rely on uptake and assimilation of organic 585 compounds (Alonso-Sáez et al., 2012; Qin et al., 2014). In such case, PA lifestyle is in favor of their 586 587 nutrient requirements. MGII have a wide distribution in the open ocean and as shown in our results, they are the dominant archaeal community generally within the upper euphotic zone (Massana et al., 588 589 2000; Martin-Cuadrado et al., 2015). Recently, they have been found, however, to be also abundant in 590 deep-sea waters (Baker et al., 2013; Tarn et al., 2016; Liu et al., 2018a), showing a wider adaption to diverse marine habitats in addition to the photic zone. MGII are thought to be heterotrophs, and have 591 the ability of degrading proteins and lipids (Iverson et al., 2012; Orsi et al., 2015). Metagenomes 592 revealed a number of genes encoding cell adhesion, degradation of high molecular weight organic 593 594 matter and photoheterotrophy (Rinke et al., 2019; Tully et al., 2019), evidencing their potentiality to utilize organic particles as important growth substrates. All these findings imply MGII's preference to 595 particle-attached lifestyle, and they are frequently detected from PA fractions in size-fractionated 596 studies (Iverson et al., 2012; Orsi et al., 2015; Tran et al., 2016). However, in a few studies including 597 our present study, MGII are also identified as the dominant archaeal components from FL fractions, 598 with equal or even more abundance than PA fractions (Fig. 7). Further studies confirm that genome 599 contents and populations of free-living MGII are distinct from those of particle-attached MGII (Orsi et 600 601 al., 2015; Rinke et al., 2019), suggesting their metabolic evolution and adjustment to niche 602 partitioning. In addition, MGIII also occurred commonly in both fractions (Fig. 7). MGIII are usually retrieved as minor components of deep mesopelagic and bathypelagic communities (Galand et al., 603 2009; Tarn et al., 2016). Like MGII, to date no cultured representative of MGIII leads to little is 604

known about their ecological and physiological characteristics. Function prediction from 605 metagenomes suggest that MGIII are aerobic (or facultative anaerobic), motile, and heterotrophic, and 606 potentially can utilize lipid, proteins and polysaccharides as major energy source (Martin-Cuadrado et 607 608 al., 2008; Haro-Moreno et al., 2017). Recently, a novel lineage of MGIII genomes preferring to live in the photic zone was recovered, consistent with previous few studies and our present results in which 609 MGIII populations are obtained from the euphotic zone with a considerable abundance (Galand et al., 610 611 2009, 2010). Moreover, recent findings also indicate that MGIII are inclined to be attached to other microorganisms (particle-attached preference) and only sporadically be released to the surrounding 612 environments (free-living lifestyle) (Haro-Moreno et al., 2017). 613

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In addition, there are several other archaeal lineages with remarkable differences in abundance between PA and FL fractions. The order Methanosarcinales and Methanobacteriales, affiliated to the phylum Euryarchaeota and retrieved exclusively from PA fractions (Fig. 7), belong to strictly anaerobic methanogens. Their preference to particle-attached lifestyle in the water column is intelligible. Within normal water column, seawater is usually oxic in spite of low oxygen concentration. Only on or inside the organic particles where heterotrophic microbes attach and digest organic matter using oxygen as electron acceptor, local anoxic niches are developed with the gradual exhaustion of ambient oxygen, and become suitable for the survival of anaerobic methanogens. Members of the *Woesearchaeota* were abundantly derived from the PA fraction of the upper seawater. In marine environments, *Woesearchaeota* are distributed restrictively in marine sediments (Lipsewers et al., 2018) or deep-sea hydrothermal vents (Takai et al., 1999), and are scarcely detected from pelagic seawater masses. Recent studies suggest that woesearchaeotal lineages are mostly retrieved from anoxic environments (Castelle et al., 2015; Liu et al., 2018b). Moreover, genomic metabolic analysis indicates Woesearchaeota have an anaerobic heterotrophic lifestyle with conspicuous metabolic deficiencies (Probst et al., 2017; Liu et al., 2018b), implying a potential syntrophic or mutualistic partnership with other organisms (Castelle et al., 2015; Liu et al., 2018b). It is further demonstrated that Woesearchaeota tend to co-occur with typical anaerobic methanogens from the Methanomicrobia and Methanobacteria constituting a potential consortia (Liu et al., 2018b). In our present results, at several depths, the Methanosarcinales of the Methanomicrobia and the Methanobacteriales of the Methanobacteria, together with Woesearchaeota, were detected concurrently, implying to a large extent their potential syntrophic partnership.

4.5 Potential vertical connectivity of microbial populations along the depth profile

Microbial distribution at different depths to a certain extent implicates their potential vertical connectivity along the water column profile. It has been suggested that the sinking of organic particles 637 formed in upper euphotic zone is a main vector in transferring prokaryotes from the surficial ocean to deep waters (Mestre et al., 2018). Those surficial lineages, usually belonging to putative photosynthetic/photoheterotrophic, bacteriochlorophyll a-containing microorganism or strict epipelagic/euphotic inhabitants, are reliable indicators to hint their downward transportation if they are detected from meso- or bathypelagic waters. For example, cyanobacteria are typical photosynthetic bacteria and their distribution is thought to be confined to the euphotic zone, with commonly observed maximum depths of about 150 ~ 200 m. In the present study, however, cyanobacterial lineages were retrieved throughout the whole water column (Fig. 5 and Fig. 6), especially at 4,000 m depth where

cyanobacteria account for nearly 12% of the PA communities. Although a recent study revealed that cyanobacteria can dominate the deep continental subsurface microbial communities with the potential for a hydrogen-based lithoautotrophic metabolism instead of photosynthesis (Puente-Sanchez et al., 2018), these indigenous deep cyanobacteria were classified into the genera Calothrix, Microcoleus and Chroococcidiopsis, phylogenetically different from those prevailing in our study (Prochlorococcus, Synechococcus). Jiao et al. (2014) observed substantial Prochlorococcus populations at 1,500 m depth in the South China Sea, and suggested that multiple physical processes, including internal solitary waves and mesoscale eddies were responsible for the occurrence of these "deep *Prochlorococcus*". However, in our study area, ages of seawater increase gradually from the surface to the deep along the water column profile in a normal time sequence (Table 1), refuting this possibility. Thus, a reasonable postulation here is that the sinking particles function as vectors and convey cyanobacteria attaching on particle surfaces from epipelagic zone into deep-sea waters. Likewise, members of the family Erythrobacteraceae, which are largely represented by OTUs within the genus Erythrobacter, are also present abundantly in both PA and FL fractions at 4,000 m depth (Fig. 6). Erythrobacter spp. belong to putative bacteriochlorophyll a-containing, aerobic anoxygenic photoheterotrophic bacteria and are thought to be distributed only in the euphotic upper ocean (Kolber et al., 2000; Koblížek et al., 2003). SAR11 clade, are potentially photoheterotrophic (Gomez-Pereira et al., 2013; Evans et al., 2015) and ubiquitous in global photic zones as one of the most abundant bacteria (Morris et al., 2002). We observed that members of SAR11 clade are distributed across the whole water columns, especially in mesopelagic aphotic depths with relatively high proportions. Other lineages specializing in inhabiting surface seawater but was also retrieved from the deep ocean include γ-proteobacterial SAR86 clade, SAR116 clade of marine Roseobacter and SAR202 clade within *Chloroflexi*. The majority of the OTUs within these "surface lineages" have been retrieved from the meso-/bathypelagic ocean and can be traced back simultaneously to those present in surface waters, suggesting their potential origin from the upper epipelagic zones.

5. Conclusions

In this study, we systematically compared bacterial and archaeal community structures within two different filtration fractions representing particle-attached and free-living lifestyles at different depths in the South China Sea. As revealed in previous studies, for either bacteria or archaea, the FL fractions usually show higher cell abundance and diversity than their PA counterparts at most depths. A set of environmental factors including depth, salinity, seawater age, DOC, POC, DO and silicate are considered playing important roles in structuring PA and FL microbial communities along the depth profile. On the one hand, as the result of adapting to different organic substrates available, PA and FL fractions generally accommodate significantly divergent microbial compositions at each depth. At fine taxonomic levels, a considerable number of microbial lineages exhibited pronounced preferences to PA or FL lifestyles, also with distinct stratified distribution along the depth profile. A few microbial taxa show potentially PA and FL dual lifestyle strategies, able to switch according to substrate availability and environmental variations, implying versatile metabolic flexibility. In addition, we found that the sinking organic particles likely function as vectors in prokaryote transfer from surface ocean to deep waters, indicating the potential vertical connectivity of prokaryotes along the water column profile.

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688	Data availability
689 690 691	The pyrosequencing data obtained from the 454 sequencing of 16S rRNA genes were deposited in the Sequence Read Archive (SRA) database under accession ID PRJNA546072 for bacterial sequences and PRJNA546071 for archaeal sequences.
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693	Author contribution
694 695 696	JL and JF designed the experiments, and JL, LG, JW and BW carried them out. JL, SB, LZ and LS treated and analyzed the sequence data. JL and JF wrote the manuscript with contributions from all coauthors.
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698	Acknowledgements
699 700	This work was financially supported by the National Natural Science Foundation of China (NSFC, No. 41373071 and No. 91951210) and National Key R&D Program of China (No. 2018YFC0310600).
701	
702	Competing interests
703	The authors declare that they have no conflict of interest.

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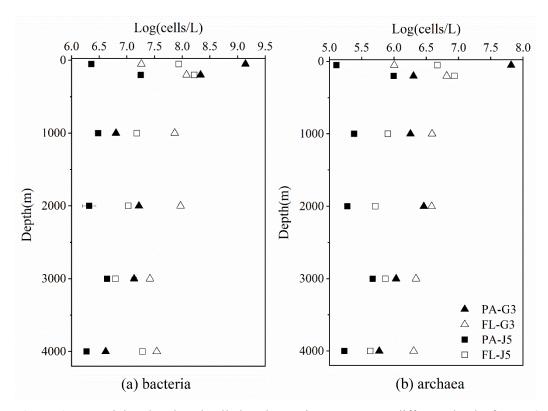


Figure 1. Bacterial and archaeal cell abundances in seawaters at different depths from G3 station and J5 station in the South China Sea, estimated from 16S rRNA gene copy abundances.

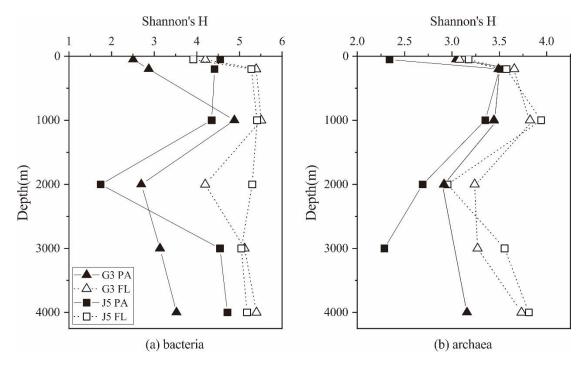


Figure 2. Shannon's diversity index calculated for all bacterial and archaeal communities of seawaters collected from G3 station and J5 station in the South China Sea.

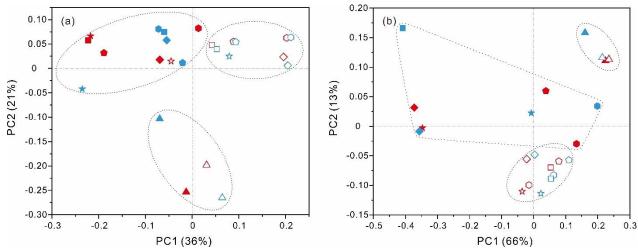


Figure 3. Results of PCoA analysis for particle-attached and free-living microbial fractions collected from seawater columns of the South China Sea. (a) PA and FL bacteria; (b) PA and FL archaea. Statistical analyses supported the groups with statistical significances (Table S3). Triangle: 50 m; rhombus: 200 m; hexagon: 1000 m; star: 2000 m; square: 3000 m; pentagon: 4000 m. Blue color: J5 station; red color: G3 station. Filled: particle-attached fraction; open: free-living fraction.

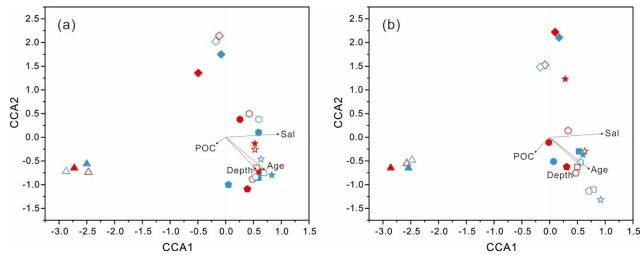


Figure 4. Results of CCA analysis to correlate several environmental factors including POC, seawater age, salinity and depth to PA and FL microbial communities collected from seawater columns of the South China Sea. (a) PA and FL bacteria; (b) PA and FL archaea. Triangle: 50 m; rhombus: 200 m; hexagon: 1000 m; star: 2000 m; square: 3000 m; pentagon: 4000 m. Blue color: J5 station; red color: G3 station. Filled: particle-attached fraction; open: free-living fraction.

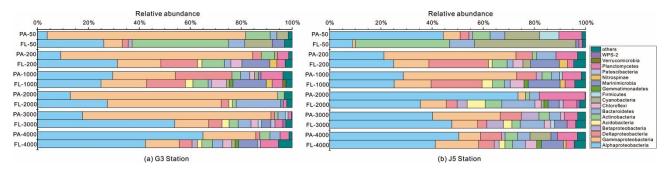


Figure 5. Taxonomic compositions of particle-attached and free-living bacterial communities of seawaters at different depths along two different water columns in the South China Sea. (a) G3 station; (b) J5 station. The phylum or class which has less than 1% proportions is classified into "others" (Fig. S5).

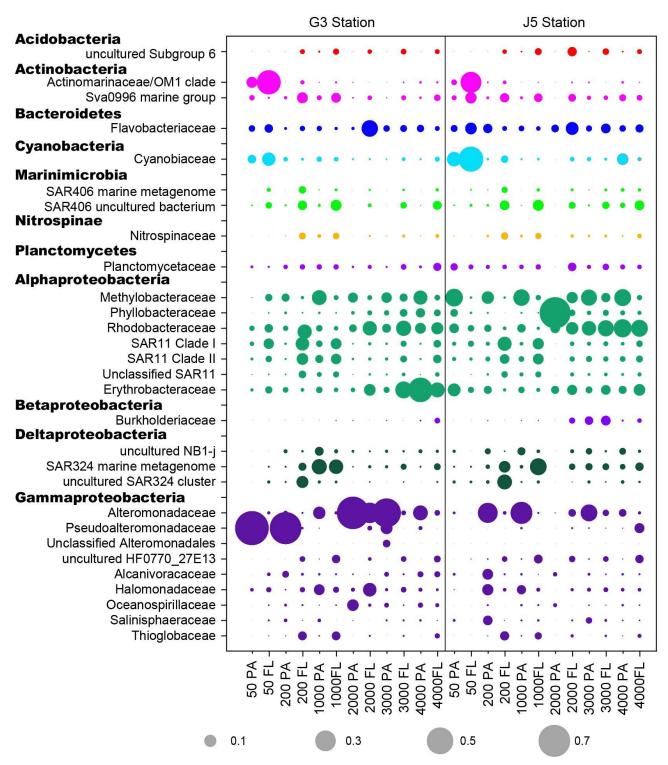


Figure 6. The relative abundances of families in particle-attached and free-living bacterial communities. Dark grey bubbles are the average relative abundances in the PA fraction, while light grey bubbles are the average relative abundances in the FL fractions. Scale is shown in the bottom, and the cycle with a number inside indicates actual relative abundance.

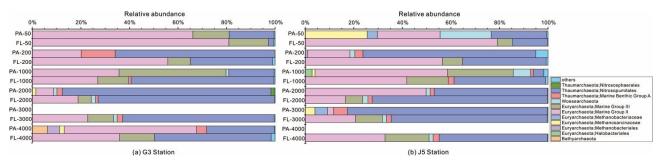


Figure 7. Taxonomic compositions of particle-attached and free-living archaeal communities of seawaters at different depths along two different water columns in the South China Sea. (a) G3 station; (b) J5 station. PA-3000 at G3 station and PA-4000 at J5 station indicate the samples failing in the sequencing of archaeal 16S rRNA gene. The archaeal lineages, at \sim phylum or class level, with less than 1% proportions is classified into "others" (Fig. S6).

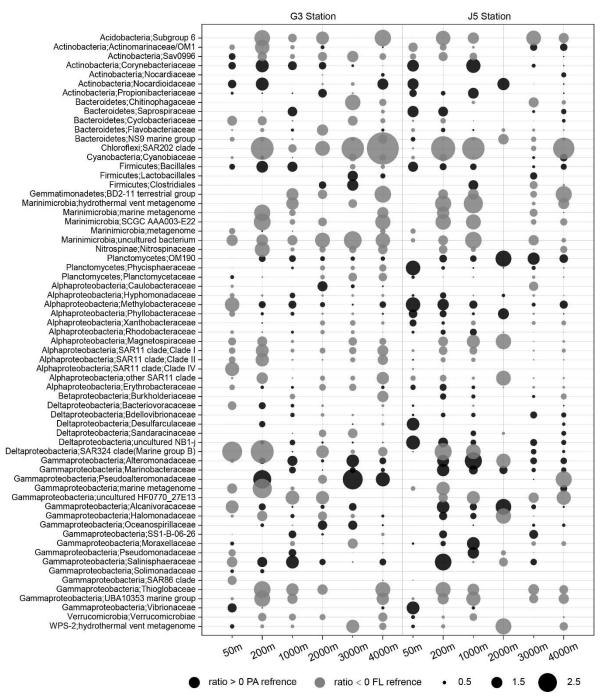


Figure 8. Odds ratio for each of the families with relatively abundant proportions in each sample. Dark grey bubbles represent the clades with a positive odds ratio, meaning the preference of PA lifestyle. Light grey bubbles represent the clades with a negative odds ratio, indicative of the FL preference. Scale is shown in the bottom, and the circle with a number inside indicates actual ratio (not proportional).

Table 1. Environmental parameters of the water columns at different depths of G3 and J5 stations in the South China Sea

G3 station											J5 station										
Depth (m)	T (°C)	Sal. (‰)	pН			POC (μM)	Ages * (yr)		PO ₄ ² - (μM)	Silicat es (µM)	_	T (°C)	Sal. (‰)	pН		DOC (μM)	POC (μM)	Ages * (yr)	NO ₃ - (μM)	PO ₄ ²⁻ (μM)	Silicat es (µM)
50	25.80	33.81	8.02	204.3	63.07	1.5	109	BD	BD	2.27		23.60	33.88	8.02	204.8	67.77	1.6	108	0.12	BD	2.36
200	15.46	34.54	7.75	115.1	53.02	0.8	106	17.98	1.20	21.06		14.27	34.52	7.72	116	49.99	0.9	106	19.13	1.30	26.56
1000	4.68	34.51	7.51	85.5	49.34	1.2	1170	37.16	2.72	114.40		4.46	34.53	7.51	82.3	45.62	2.1	1310	37.04	2.73	121.93
2000	2.52	34.61	-	-	-	1.1	1190	-	-	-		2.49	34.61	7.52	102	41.67	0.9	1670	38.41	2.81	151.46
3000	2.36	34.62	-	-	42.94	1.8	1600	-	-	-		2.36	34.62	7.52	109.7	40.34	0.7	1680	38.16	2.79	145.03
4000	2.39	34.63	7.52	115.1	42.44	0.7	1750	38.48	2.82	141.81		2.43	34.62	7.53	111.8	46.52	1.2	1610	38.58	2.78	145.06

^{*} Δ^{14} C ages; BD: Below detection; -: no measurement.