

Author's Response to Reviewer's Comments on bg-2020-120

We are grateful and excited to see that all the reviewers see utilities and timeliness of our work. We have considered reviewers' comments and made changes accordingly.

Here are the major revisions we have made:

- (1) Both reviewers addressed questions on the title of our manuscript. We have changed our title, reluctantly, to "Carbons at Different Positions Behave Like Different Elements- An Insight into Position-Specific Isotope Distributions". In addition, we recalled the title in the last paragraph of subchapter 3.
- (2) We have added descriptions on KIE and EIE with equations.

Below are point-by-point responses to comments and a marked-up manuscript version

Responses to Anonymous Referee #2

The authors propose a method to receive added value out of the knowledge of sitespecific or position-specific isotopic compositions in more or less complex (organic) molecules (or inorganic minerals). They suggest to compare measured "intramolecular isotope distributions" (abbreviated as "Intra-ID") in (organic) molecules with theoretically calculated isotope distributions assuming a synthesis reaction under thermodynamic control thereby accomplishing (chemical) equilibrium. The manuscript has a sort of review character as obviously all measured intra-IDs have been taken from literature. The authors would like to test the theory that "information on the source, reaction pathway, and kinetics of an organic compound can be obtained from its position-specific isotope compositions" and end up with the familiar and already well-known conclusion that inter- and intramolecular isotopic compositions alone are an inadequate means to reach this goal. A basic idea on the involved synthesis/breakdown reactions, flux rates and regulation points of the involved reaction pathway(s) in addition to knowledge on kinetic or equilibrium isotope effects introducing isotope fractionations and thereby sharpening the observed intra-IDs is needed. The manuscript is interesting and innovative, but needs major amendment. There might be even the need for more than this "review round". Especially, there is a need to work out a sort of a "Take home message" for the reader (in the conclusion part?).

Response: The central message (or take-home message) of this “Idea & Perspective” piece of ours is to convey the idea that when it comes to position-specific isotope analysis (PSIA), do not think that the carbons at different positions in a molecule behave like the same element, in fact, they behave more like different elements. In other words, most of these carbons never exchange or intermix, due to the lack of exchange mechanisms. Any attempt to treat the differently positioned carbons as they were the same element is very much like treating the O in SO_4^{2-} and the O in crystallization H_2O in gypsum mineral as the same O. In fact, these C’s or O’s behave very much like O and S in SO_4^{2-} . This insight might be possessed implicitly by some in the community but was missing and should have been highlighted in the much-publicized Galimov-Buchachenko-Schmidt debate that occurred years ago. What we have read from the literature told us that it is necessary and timely to bring this insight to the open. In this contribution, we used simple inorganic and organic molecules as examples to illustrate the idea.

The subchapter “3 Implications” needs a complete revision. Instead of discussing equilibrium and non-equilibrium issues in terrestrial or extraterrestrial material there is need to present here which additional info is needed to interpret the Intra-IDs. Best would be here to connect e.g. N_2O (sitepreference) data with the schematics of a metabolism pathway producing N_2O from a defined origin (and/or the acetic acid part can be elaborated in an analogues manner).

Response: The subchapter “3 Implications” serves to give the readers the background of the Galimov-Buchachenko-Schmidt debate. In Introduction, we have introduced the debate on whether there is an intramolecular $\delta^{13}\text{C}$ - $^{13}\beta$ correlation or equilibrium-like isotope distribution. However, Galimov was not trying to advocate that a living system is in equilibrium or at a steady-state near equilibrium. He brought out the $\delta^{13}\text{C}$ - $^{13}\beta$ correlation to support his hypothesis that the magnitude of reaction rate on the scale of interest in a living system is linearly dependent on the thermodynamic driving force responsible for the reaction system. Therefore, the theorem of minimum entropy production, i.e. Prigogine’s non-equilibrium thermodynamics, can be applied in biochemical systems to explain metabolism. Although we do not think his evidence for intramolecular $\delta^{13}\text{C}$ - $^{13}\beta$ correlation is sound, we do think a local thermodynamic control in an overall non-equilibrium thermodynamic system like the living system has some merits and should be mentioned here to echo what we brought out in Introduction. In addition, the

last paragraph of this subchapter is to connect to the title of our manuscript that although position-specific carbons are the same element, they should be treated independently as if they were different elements in a molecule.

Specific comments:

1) Title: I do not understand the title. What is the meaning of “same carbon different elements” ??? → interesting terms here might be “functional groups”, “carbon molecule positions”, "different bond types" ???

Response: “Same carbon different elements” here means: “The position-specific carbons at different positions are the same element carbon but behave like different elements.” In light of the fact that the other reviewer was also puzzled by the title, we have changed our title, reluctantly, to “Carbons at Different Positions Behave Like Different Elements- An Insight into Position-Specific Isotope Distributions”.

2) Title: “isotope pattern” not mentioned in the text of the manuscript. What is the difference between “isotope pattern” and “isotope distribution”? Pls use only one description.

Response: No difference, but some quoted sentences had used “pattern”. We always stick to “distribution”.

3) Line 9: Only “kinetics”, no “thermodynamics” ??? kinetically controlled reactions and equilibrium, isn't that a contradiction?

Response: “Kinetics” here refers to reaction kinetics. We revised to use “reaction kinetics” when we referred to reaction processes in the text.

Equilibrium refers to the state of a system where the forward reaction and the backward reaction that go through the same transition state have an equal rate. Kinetic isotope effect (KIE) and equilibrium isotope effect (EIE) is not really mutually exclusive in chemical physics. KIE by definition (in the sense of Bigeleisen & Wolfsberg, 1958 the classical paper on KIE) is the EIE between transition-state and the

reactant. Currently, many isotope geochemists refer to non-equilibrium processes as “kinetically controlled”, which is not a good practice; it is misleading to say the least. Partially reversible reactions at steady states can have isotope fractionations that are not at the degree of either EIE or KIE, but a combination of forward and backward KIEs. This issue has been raised by Clayton & Mayeda (2009) earlier.

4) Line 25ff: KIE give info on transition state / mechanism (rate-determining step) of a reaction, whereas EIE give info on the stiffness of the bond of the corresponding isotope in reactant and product (change in bonding of the isotope in question). Your “transition-state and reversibility” is too much abridged here. Best would be to mention here in this context also the connection of KIE and EIE. EIE is equal to the ratio of the KIEs on forward and backward reaction in case the chemical (and isotopic) equilibrium has been accomplished. In kinetically controlled reactions the step between educt and the intermediary transition state (TS) is reversible and the reaction from TS towards product is irreversible. The term “equilibrium” is not helpful when talking about KIE. Please change wording correspondingly. Additionally also info on EIE would be needed here.

Response: We agree with the reviewer on the definition of KIE and EIE, except that in the framework of transition state theory and the way KIE and EIE are calculated, there is no fundamental difference between KIE and EIE. KIE used an imaginary frequency in a transition state, which is probably the biggest difference if you do not count the difficulty of pinning down a TS structure (See also our response to comment 5 below).

A bit more information here on the issue. The forward KIE_f is the EIE between transition state and reactant, and the backward KIE_b is the EIE between transition state and product. “The KIE of an elementary step can be defined as the **equilibrium** fractionation factor between transition-state and reactant” is not defined by us. In Wolfsberg et al. (2009): *Isotope Effects in the Chemical, Geological, and Bio Sciences*. (DOI:10.1007/978-90-481-2265-3), Chapter 6, page 184, we quote, “Also, in TST (Chapter 4), one assumes that in a normal reaction the transition state is in chemical **equilibrium** with reactants and its concentration can be calculated from the chemical **equilibrium** constant corresponding to the reaction

between the reactants and transition state.” In that regard, plus our own experience, we found often that the students understood the KIE concept much more deeply when we view it in the “equilibrium” framework originally set up in Bigeleisen & Wolfsberg (1958) than in approaches adopted in most geochemistry books.

5) Line 30: Please use terminology of Coplen (2011, <https://doi.org/10.1002/rcm.5129>). As there are different definitions of KIE numbers used in geochemistry and other disciplines, it would be a good idea to also write the corresponding equations for KIE and EIE and clearly state which number corresponds to normal and inverse IE. Perhaps in a footnote or Appendix (as the editor recommends). A KIE of e.g. 1.01 means that the product is depleted or enriched in the heavy isotope relative to the reactant? Please check with Coplen (2011) or define via own equation(s). But a definition is needed.

Response: Coplen (2011) defines KIE in the way that Bigeleisen (1949) initially defined, which is $k_{\text{light}}/k_{\text{heavy}}$. Such a definition is opposite to the equilibrium isotope effect ($k_{\text{heavy}}/k_{\text{light}}$ in both Coplen (2011) and Bigeleisen and Mayer (1947)). If we followed such a definition, it can easily cause confusion since the fractionation factor value would have different symbols. For instance, assuming the EIE of a reaction is 1.01, it means that the product is **enriched** in heavy isotope for ~10‰. Assuming the KIE of the reaction is also 1.01, by definition of Coplen (2011), it means that the product is **depleted** in heavy isotope for ~10‰. Using the concept of $\epsilon=(\alpha-1)$ or $\ln\alpha$, the description “the EIE of the reaction is 10‰ and the KIE of the reaction is 10‰.” would be confusing. Thus, in Bao (2015, GCA), he suggested follow the convention of geochemists (opposite of that of the physical organic chemists who are mostly interested in hydrogen isotopes) and define KIE as $k_{\text{heavy}}/k_{\text{light}}$ or $\text{RPF}_{\text{transition-state}}/\text{RPF}_{\text{reactant}}$, so KIE and EIE would have the same symbol (positive or negative).

Similarly, we saw suggestions of defining KIE as $k_{\text{heavy}}/k_{\text{light}}$ in a recent preprint in open discussion, Michalsk et al., 2020 (<https://doi.org/10.5194/gmd-2020-114>), page 6, line 34-39, in which we quote:

“Much of the early research on KIEs were investigations of the KIE in reactions containing hydrogen isotopes and these studies usually defined a KIE = $k_L/k_H = \alpha_{L/H}$, where the k 's are the rate constants for the light and heavy isotopologues. This is the inverse of the definition of a usually used in research

dealing with EIE, VPIE, PHIFE and this inversion can lead to confusion. In this paper, to maintain consistency between the α values for EIE, KIE, VPIE, and PHIFE, α will be defined as heavy/light for all four effects.”

We have revised the description and defined KIE and EIE with equations. It now reads:

“According to the transition-state theory (Eyring, 1935a, b), the KIE of an elementary step can be defined as the equilibrium fractionation factor between transition-state and reactant (Jones and Urbauer, 1991; Bao et al., 2015):

$$KIE = \beta_{TS}/\beta_R$$

where the β factor denotes the reduced partition function ratio of transition-state (TS) or reactant (R). A β factor is the equilibrium isotope fractionation factor between an atom in a specific bond environment and its atomic form that can be predicted theoretically (Urey, 1947, Bigeleisen and Goeppert-Mayer, 1947). For a unidirectional reaction, the KIE of a reaction can also be defined as:

$$KIE = {}^h k / {}^l k$$

where k denotes the reaction rate constant of heavy (h) or light (l) isotopes. To adapt to the convention of geochemists, we define KIE this way so that the normal KIE is less than 1.000, which is the opposite of what Bigeleisen (Bigeleisen, 1949) initially defined.

EIE is the isotope fractionation among reactant and product, which is determined by the bonding environment of the target position or compound:

$$EIE = \beta_P/\beta_R$$

where P denotes the product of a target reaction. It can also be defined as:

$$EIE = {}^h K / {}^l K$$

where K denotes the equilibrium constant of a target reaction. At equilibrium, the EIE of a reaction equals to the ratio of forward reaction KIE_f , and backward reaction KIE_b ($EIE = KIE_f/KIE_b$, Bao et al., 2016).”

6) Line 36: “Pls replace “. . . of all different positions in a compound” by “. . . of all different positions of the same element in a compound”. That is what you mean?

Response: Thanks for the suggestion. We have revised as suggested. It now reads:

“The compound-specific isotope composition averages isotope compositions of all different positions of the same element in a compound, where information contained in position-specific isotope compositions could be lost (Elsner, 2010; Piasecki et al., 2018).”

7) Line 39: There are many more paper on hydrogen isotope distribution in organic molecules. See e.g. Martin et al. (<https://doi.org/10.1111/j.1365-3040.1992.tb01654.x>). ^{13}C intra ID by NMR is a relative recent approach.

Response: We were trying to say that technological development in the carbon position-specific isotope analysis (PSIA) has been very active in recent years. This is in addition to the ongoing debate among Galimov, Buchachenko, and Schmidt on position-specific carbon isotope distributions. However, we realized that PSIA on hydrogen and oxygen is also developing fast. Therefore, this sentence only adds confusion. We have deleted it.

8) Line 42: According to my opinion, the term “statistical” was chosen by Schmidt to explain that the distribution of the heavier isotopes in an isotopomer compounds is not a stochastic distribution but follows certain rules. In the articles by Schmidt the term “non-statistical” states that the distribution is not guided by chance, but follows a logical order. It is not stated, whether this order is under thermodynamic or kinetic control. Please adapt. In case, the text passage in italics is a direct citation, most probably Galimov or Schmidt (not both) have stated that. See also line 56.

Response: Neither Galimov nor Schmidt and their colleagues had clearly defined the terms “statistical” and “non-statistical” isotope distributions. In our understanding, these two terms refer to “equilibrium” and “non-equilibrium” isotope distributions for the following reasons:

1. In Schmidt (2003, <https://doi.org/10.1007/s00114-003-0485-5>), he stated that “*Nevertheless, the data elaborated and collected by Galimov (1985) show, in a number of cases, especially for the intermolecular range, more or less satisfactory correlations between ^{13}C -contents and β factors. On the other hand, many recent investigations have proved that the **thermodynamic order** is not generally realised, especially not for intramolecular isotope distributions in natural compounds, and that unequivocally kinetic isotope effects determine the isotope abundance in many defined molecule positions. Anyway even a partial realisation of the thermodynamic order of the **nonstatistical distribution of isotopes** would demand an explanation compatible with classical enzyme kinetics.*”

In this paragraph, he first recognized Galimov’s $\delta^{13}\text{C}$ - $^{13}\beta$ correlation as “*thermodynamic order*”. In that context, thermodynamic order equals to equilibrium. Then, Schmidt brought out the fact that the $\delta^{13}\text{C}$ - $^{13}\beta$ correlation is not common in natural compounds. In the next sentence, “*the nonstatistical distribution of isotopes*” appears following “*a partial realisation of the thermodynamic order of*”. In the article, Schmidt was talking about the partial reversible biochemical process at steady-state that can produce a predictable non-equilibrium isotope distribution, which his “*nonstatistical distribution of isotopes*” must have referred to.

2. Also in Schmidt (2003), he mentioned a case of L-malate from Meinschein et al. (1984). Meinschein et al. (1984) is an abstract, and we could not find the full text of it (<https://pascal-francis.inist.fr/vibad/index.php?action=getRecordDetail&idt=9067003>). Nevertheless, we can see the data in Galimov (2006, <https://doi.org/10.1134/S0016702906130015>) in which Fig. 5.2.7 showed that the measured position-specific $\delta^{13}\text{C}$ is well correlated with predicted $^{13}\beta$. Galimov quote Meinschein: “*the ^{13}C contents of the specific carbon atoms in malic acid from apple and sorghum increase in accordance with their values, as predicted by Galimov.*”

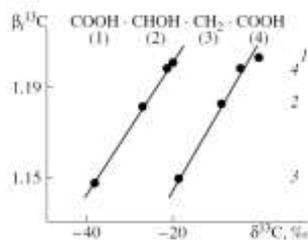


Fig. 3.2.7. Intra-molecular distribution of carbon isotopes in the malate molecule from plants C-3 (apple fruit, left line) and C-4 (sorghum, right line) (Meinschein et al., 1984).

It is also stated in Schmidt (2003): “However, the ^{13}C -patterns of these acids (Fig. 3) do not at all coincide with those predicted from the precursors glucose/pyruvate (which is not contradictory to a correlation of their average $\delta^{13}\text{C}$ values), although they do show rather *satisfactory correlations with the thermodynamic β_i factors*, already reported for L-malic acid by Meinschein et al. (1984).” Therefore, we can confirm that the reported L-malic acid has an equilibrium intramolecular carbon isotope distribution. Schmidt described the equilibrium isotope distribution of L-malate as “For “nature identical” L-malate one would expect ... a *statistical ^{13}C distribution*.”

3. The literature from the Schmidt group with “Nonstatistical Carbon Isotope Distribution” in the titles have non-equilibrium intramolecular carbon isotope distributions. For instance, “Evidence for a *Nonstatistical Carbon Isotope Distribution in Natural Glucose*” (Rossmann, Butzenlechner, and Schmidt, 1991, <https://doi.org/10.1104/pp.96.2.609>); “Carbon Isotope Effects on the Fructose-1,6-bisphosphate Aldolase Reaction, Origin for *Non-statistical ^{13}C Distributions in Carbohydrates*” (Gleixner and Schmidt, 1997, <https://doi.org/10.1074/jbc.272.9.5382>).
4. In Galimov (2006), he stated that “Thermodynamic laws have a *statistical* character.” Therefore, if an isotope distribution has “thermodynamic order”, it should be described as “statistical isotope distribution”
5. Romek, et al. (2016, <https://doi.org/10.1074/jbc.M116.734087>) stated: “From this, the molar fraction was calculated, which gives the extent to which the $^{13}\text{C}/^{12}\text{C}$ ratios diverge from a *statistical distribution*.” The equilibrium isotope distribution is the reference that has been compared to, which the measurement diverges from. Therefore, “statistical distribution” means “equilibrium distribution.”

All in all, “statistical” and “non-statistical” in Schmit et al mean “equilibrium” and “non-equilibrium” isotope distributions, respectively.

The terms “statistical” and “non-statistical” are ambiguous, therefore, in Schmidt (2015, <http://dx.doi.org/10.1080/10256016.2015.1014355>), he stopped using the two terms. This is also the reason we suggest to use “equilibrium” and “non-equilibrium” Intra-ID to describe isotope distributions.

9) Line 61: “averages”: Do you mean average d-value of the whole molecule? The Intermolecular isotopic composition?

Response: Yes and No.

Galimov’s equation is:

$$\delta^{13}C - \delta^{13}C_{ave} = \kappa(^{13}\beta - ^{13}\beta_{ave}) \times 10^3 \quad (\text{Galimov, 1985, pg 100, eq 4.3})$$

κ (also written as χ in Galimov, 2004, 2006) is the regression coefficient. The “ave” values are the unweighted arithmetic mean of all measured $\delta^{13}C$ or $^{13}\beta$ values, which we have criticized in He et al. (2018), from which we quote, “*In a system consisting of multiple components, if we choose a component as a reference for mutual comparison, even if the reference is the average stable isotope composition of compounds of interest, we have effectively assigned that reference to be at equilibrium. The use of such a reference is not mathematically rigorous and can often be misleading when dealing with a complex non-equilibrium system. This is simply because we do not know a priori which compound or set of compounds represents the state of isotope equilibrium.*” In addition, the average $\delta^{13}C$ value of the whole molecule needs to be weighted since some molecules have multiple carbons in the same position. This is the same for intermolecular isotope distribution (Hayes, 2001, pg 233, eq 5). Using the unweighted arithmetic mean is one additional problem of Galimov’s $\delta^{13}C$ - $^{13}\beta$ correlation.

We have revised this part. It reads:

“Such a $^{13}\beta$ - $\delta^{13}C$ correlation is written as $\delta^{13}C - \delta^{13}C_{ave} = \chi(\beta - \beta_{ave}) \times 10^3$, where χ is the regression coefficient. Galimov interpreted such observed intramolecular $^{13}\beta$ - $\delta^{13}C$ correlations as *equilibrium-like* Intra-IDs produced from sets of reversible biochemical reactions at steady-states which are not far from equilibrium. The $^{13}\beta$ - $\delta^{13}C$ correlations were used as supporting evidence that the theorem of minimum entropy production can be applied in biochemical systems. ... The $^{13}\beta$ - $\delta^{13}C$ correlation used

an unweighted arithmetic mean isotope composition of all the components as the reference of a system. Strictly, only the mass-weighted isotope composition of all components should represent that of a system (Hayes, 2001). In addition, arbitrarily assigning a reference is not mathematically rigorous (He et al., 2018). Therefore, a $^{13}\beta$ - $\delta^{13}\text{C}$ correlation cannot be used as supporting evidence for Galimov's hypothesis that the theorem of minimum entropy production applies in biochemical systems."

10) Line 64ff: I do not understand your differentiation between your point 1) and your point 2). Let's assume the reaction sequence . . . A \rightarrow B \rightarrow C \rightleftharpoons D \rightarrow E \rightarrow F. . . (and a branching point at C and/ or D according to Hayes and Schmidt). The system should also be "regulated" on the reaction from A/B and E/F ("bottleneck" as an analogy), so that the reaction between C and D approaches or even accomplishes chemical equilibrium. The reaction between C and D should "own" an EIE (e.g. ^{13}C EIE). Then only the carbon atoms in molecule C and molecule D can be "isotopically" equilibrated that are influenced by the primary and secondary (tertiary ??) thermodynamic isotope effects on the equilibration reaction (Secondary isotope effects: <https://goldbook.iupac.org/terms/view/S05523>). It is useless (without a value, not applicable) to make a statement on the carbon atoms in C and D, that are not touched by any equilibrium isotope effects. Even secondary IE (for the heavy elements beside ^2H) are normally very small.

Response: Good questions and good analyses. It is precisely these questions that make our central point in the manuscript the more relevant and urgent. It looks like the notion and distinction of the two equilibrium scenarios are less clear and more difficult than we thought.

The two equilibrium scenarios are:

"1) intermolecular isotope equilibrium among the corresponding bond-breaking/forming positions in reactant and product in a defined **process**, and

2) intramolecular isotope equilibrium among all carbon positions in a defined **molecule**."

What we are talking about is one single reaction, not a reaction network, as the reviewer mentioned here. For instance, for a reaction $\text{AB} \leftrightarrow \text{A}'\text{B}'$, the process equilibrium is the equilibration among A and A', B

and B'. The equilibrium among positions is the equilibrium among A and B, and among A' and B'. Yes, there will be secondary and tertiary isotope effects involved in the neighboring atoms, but those effects are respected to the specific reaction $AB \leftrightarrow A'B'$. These neighboring atoms (e.g. carbons) themselves may not be in isotope equilibrium.

The equilibration the reviewer referred to belongs to the equilibrium scenario (1). However, when we talk about equilibrium Intra-ID, it refers to the equilibration among different (carbon) positions in a molecule, which is the scenario (2).

In the literature, when talking about an equilibrium process produces equilibrium Intra-ID, none of Galimov, Buchachenko, Schmidt, and Hayes clearly stated if they are talking about the equilibration among the corresponding positions in reactant and product or the equilibration among positions in the product. The clarification of this point is extremely important, which is the central message of our manuscript.

11) Line 71: “few intramolecular exchange pathways”. This statement needs either a literature citation or there is need to present own data as a proof.

Response: We have cited a few papers illustrating the ubiquitous stability of characteristic organic carbon skeletons in Section (?) and paragraph ???. These stabilities are the reason why organic chemistry is quite predictable, and enantiomers extremely common.

12) Line 86/87: You should state here that oxygen can be bonded in different functional groups that have different chemical properties. A way out would be a position-specific analysis of the oxygen isotopic composition.

Response: We have revised this part. It now reads:

“The same element, e.g. carbon, occupies different positions in a compound is not a unique feature of organic compounds. Some oxygen-bearing minerals have two or more position-specific oxygens, where

oxygen atoms occupy different positions in a mineral structure and have different chemical properties. For example, it had been proposed that water temperature could be reconstructed from intracrystalline oxygen isotope difference in a single mineral copper sulfate pentahydrate ($\text{CuSO}_4 \cdot 5\text{H}_2\text{O}$) (Götz et al., 1975), kaolinite ($\text{Al}_2\text{Si}_2\text{O}_5(\text{OH})_4$), illite ($\text{K}_{0.65}\text{Al}_{2.0}(\text{Al}_{0.65}\text{Si}_{3.35}\text{O}_{10})(\text{OH})_2$) (Bechtel and Hoernes, 1990), or alunite ($\text{KAl}_3(\text{SO}_4)_2(\text{OH})_6$) (Arehart et al., 1992).”

13) Line 129: Would it be possible to present a typical example for N_2O produced from equilibrium or from a non-reversible reaction here?

Response: No. We could not. None of the published data in literature can provide evidence of no reverse (backward) reaction or evidence of equal forward and backward reaction rates for N_2O .

14) Line 143: Pls define alpha with an equation (is it isotope fractionation factor ? Pls see also Coplen 2011). The factor 1000 in the alpha formula is related to the $\delta^{13}\text{C}$ formula? Meanwhile the factor 1000 is deprecated in the e.g. $\delta^{13}\text{C}$ formula. Needs to be communicated also in the text and foot note / appendix.

Response: Thanks for the suggestion. In the text, we wrote, “The relative isotope enrichment between the carboxyl and methyl carbon in acetic acid is defined as $\ln^{13}\alpha_{carb-met} = \ln(^{13}\text{R}_{carb}/^{13}\text{R}_{met}) \times 1000\text{‰}$ ”. We have now deleted the redundant “ $\times 1000\text{‰}$ ”. “ln” means the natural log of “ $^{13}\alpha_{carb-met}$ ” and “ $^{13}\text{R}_{carb}/^{13}\text{R}_{met}$ ”. Here, $^{13}\alpha_{carb-met} = ^{13}\text{R}_{carb}/^{13}\text{R}_{met}$, and “ $\ln^{13}\alpha_{carb-met}$ ” is “the relative isotope enrichment between the carboxyl and methyl carbon in acetic acid”. α is the isotope fractionation factor. We have labeled it at the first appearance of the “isotope fractionation factor” in the first paragraph of section 2.3. It now reads:

“Based on the predicted equilibrium Intra-ID, a predicted isotope fractionation factor (α) of corresponding positions between reactant and product in a process can help to evaluate the thermodynamic state of a system and to decipher reaction pathways.”

“The relative isotope enrichment between the carboxyl and methyl carbon in acetic acid is defined as $\ln^{13}\alpha_{carb-met} = \ln(^{13}\text{R}_{carb}/^{13}\text{R}_{met})$ ”

15) Line 151: What is the meaning of “man-made”? Produced by chemical synthesis ?

Response: We have revised all “man-made” to “artificial”.

16) Line 161: “Intra-ID” should be equal to the d13C value difference between the precursor minus the primary KIE”. Do you have information on the original Intra-ID of the oil from the “oil-prone source rocks”?

Response: We do not have the information on the original Intra-ID of the oil from the oil-prone source rocks. We are trying to explore the “equilibrium-like” Intra-ID expected in acetic acid produced from precursor acid pyrolysis. Therefore, that information is of no use to us.

17) Line 162: The fact, that numbers for KIEs are higher as corresponding EIE values is commonly known. But what is a negative KIE? Please define also the equilibrium isotope fractionation factor.

Response: Please see our response to your comment 3). By our definition, $KIE = \frac{\text{heavy}_k}{\text{light}_k} = 0.98$, and the isotope enrichment ($\ln KIE$) would be -20‰, which is negative. The extent of KIE is greater than EIE, but the number 0.98 is smaller than 0.99 (or -20‰ is smaller than -10‰). We know the whole expression has been difficult. But thanks to your comment, we have now changed the adjectives to match what specific numbers we are talking about.

18) Lines 170 to roughly 190 should be shortened. Non-essential rather distracting information is given here. The focus of the manuscript by He et al. is not to present a proof of the Galimov theory, or?

Response: See our general response.

19) BTW, I do not understand the text part starting in line 196. H₂O consists out of H and O, yes. Given info also true for nitrate and sulfate. There are no isotopomer water molecules. Are there isotopomer molecules of sulfate and nitrate with an Intra-ID? What idea is behind this Paragraph? It would be interesting to compare Intra-IDs of e.g. carbon and oxygen or carbon and hydrogen in organic molecules like glucose. 2H isotopomer distribution and 13C isotopomer distribution of glucose have been published already.

Response: This paragraph is an effort to connect to the title “Same Carbons behave Like Different Elements- An Insight into Position-Specific Isotope Distributions”. See our response to your general comment in the very beginning. We quote here “Any attempt to treat the differently positioned carbons as they were the same element is very much like treating the O in SO₄²⁻ and the O in crystallization H₂O in gypsum mineral as the same O. In fact, these Cs or Os behave very much like O and S in SO₄²⁻. We have revised this paragraph. It now reads:

“A simple comparison of position-specific isotope compositions in one sample, e.g. $\ln^{13}\alpha_{carb-met}$ values of one acetic acid sample, offers little information on the reaction it involves. Although the position-specific atoms are the same elements, without an exchange mechanism, they behave independently as if they were different elements. The isotope fractionation relationship of different elements in the same compound, i.e. $(\alpha_A-1)/(\alpha_B-1)$, $\ln\alpha_A/\ln\alpha_B$, or $\Delta\delta_A/\Delta\delta_B$, (named bonded isotope effect, He and Bao, 2019), is useful in characterizing a reaction pathway. Some of the studied examples are δD and $\delta^{18}O$ in H₂O (Dansgaard, 1964; Craig, 1961), $\delta^{15}N$ and $\delta^{18}O$ in NO₃⁻ (Casciotti and McIlvin, 2007; Wankel et al., 2009), $\delta^{34}S$ and $\delta^{18}O$ in SO₄²⁻ (Antler et al., 2013), and $\delta^{13}C$ and δD in organic compounds (Elsner, 2010; Palau et al., 2017). The isotope composition difference of different elements in a molecule is useful only when the isotope fractionation relationships are considered and their isotope compositions are normalized, e.g. $\Delta(15,18) = (\delta^{15}N - \delta^{15}N_m) - (\alpha - 1 / \alpha - 1) \times (\delta^{18}O - \delta^{18}O_m)$, in which $\delta^{15}N_m$ and $\delta^{18}O_m$ are the average isotope composition in a given ocean water column (Sigman et al., 2005). The normalization procedure was necessary because the source isotope compositions can affect the values of the product. Similarly, if the same element at different positions have different sources, their source isotope composition difference

must also be considered. Position-specific isotope researches can build upon our understanding of bonded isotope effect.”

In addition, ^{18}O position-specific isotope analysis, i.e. ^{18}O distribution of glucose has been reported recently. Ma et al, (2018 <https://doi.org/10.1021/acs.analchem.8b02022>); Ma et al., (2020, <https://doi.org/10.1021/acs.analchem.9b05314>). It would be interesting to study position-specific D, ^{13}C , and ^{18}O of glucose collectively.

20) Line 44/45: It should read "Bigeleisen and Goeppert-Mayer" (with or without hyphen). Jacob Bigeleisen and Maria Goeppert Mayer <https://aip.scitation.org/doi/10.1063/1.1746492>

Response: Thanks for the reminder, we have changed “Mayer” to “Goeppert-Mayer” both in text and in references.

21) Comment on the Galimov theory The above mentioned calculations for the "Equilibrium Intra IDs" are based on the framework elaborated by Galimov, who assumed that inter- and intramolecular isotope distributions in molecules of metabolic reaction networks in Nature are under thermodynamic control. The theoretically calculated b-factors (e.g. $b_{13\text{C}}$ for carbon) according to Galimov are compared with measured and reported d-values (e.g. $d_{13\text{C}}$). The theory of Galimov on thermodynamic factors controlling the intra-IDs has been contradicted by many researchers. Additionally, to the already cited manuscripts by Buchachenko, Schmidt (and coworkers), Hayes, also Monson and Hayes (1982 *Geochim Cosmochim Acta* 46, 139ff), O’Leary and Yapp (1978 *Biochem Biophys Res Commun* 80, 155ff) and Varshavskii (1988, *Biophysics* 33(2), 377ff. Elsevier Pergamon Article in English) could be listed there. Dynamic reaction networks in living organisms are kinetically controlled. Chemical (and isotopic) compositions of molecules at diverse levels are controlled in a steady state that allows continuous flow of mass and energy followed by a constant but adjustable flux through biochemical pathways including continuous synthesis and degradation reactions of compound molecules involved. In contrast, a system at chemical (and isotopic) equilibrium would approach a stable state and be a closed system not exchanging matter with the environment. The Gibbs free energy will then come to a minimum approaching zero. The Galimov

theory is not compatible to how the biochemical pathways are explained in (plant) biochemistry text books.

Response: We totally agree with your account. We would like to mention two points:

1) As we explained above, Galimov was not trying to prove that the living system is at equilibrium, and local $\delta^{13}\text{C}$ - $^{13}\beta$ correlation is just evidence for a biosystem evolving toward an apparent thermodynamic control. In his book (1985) and reviews (2004, 2006), he also recognized cases that $\delta^{13}\text{C}$ poorly correlated with $^{13}\beta$. 2) Previous counter-arguments against Galimov did not really nail it because those arguments failed to recognize a) the reference issue which was addressed in He et al (2018) and b) non-exchangeability issue among the many carbon positions in a large molecule. That is, true intramolecular equilibrium can hardly be achieved simply because there is no viable mechanism for the exchange, which is addressed in this manuscript.

Responses to Anonymous Referee #3

1 General Comments

While the authors provide sufficient literature context for their discussion, my biggest concern with the paper in its current form is that it does not clearly delineate novel insights or findings by the authors from pre-existing literature. In other words, I am not certain where the review of the historical literature ends, and the authors' analysis begins, which makes it difficult to evaluate exactly what the authors' primary contributions are here. For example, to what extent has the N₂O site-preference example already been articulated in the literature? Site-preference measurements have been made for many years, and the authors cite other studies noting that precursor symmetry matters, so for readers less familiar with this application: is this simply an example compiled from existing studies to argue that more context is needed to interpret an Intra-ID, or is some component of the discussion new, like the subsequent discussion of reversibility? Some careful re-wording could clarify this and similar types of questions throughout. Similarly, it would be helpful if the authors could add at least one more sentence at the end of the oxygen-bearing minerals example synthesizing the broadly generalizable point that the authors are trying to convey.

Response: Thanks for the comments. You summarized our piece accurately. Per your criticism on a lack of clear distinction of prior work and our new insights in some of the writings, we have revised the text to add phrases and sentences like “It had been proposed” “In previous literature” “The measured $\delta^{13}\text{C}_{\text{met}}$ values from literature”, “Nevertheless, the two previously proposed mechanisms cannot distinguish N_2O with $\text{SP}>0$ produced from the two mechanisms.”, “Except for the above-mentioned features, the production of artificial and biological acetic acid has too many unconstrained parameters. Thus, our discussion will focus on the acetic acid derived from hydrous pyrolysis of oil-prone source rocks.”

We have added a summary at the end of the oxygen-bearing minerals example. It reads:

“The use of a single-mineral geothermometer requires that oxygens at two different sites have attained equilibrium by exchanging with each other or with the same source oxygen (e.g. water). Unfortunately, it is difficult to meet this requirement. It is, therefore, not surprising that few successful single-mineral geothermometers exist if any at all.”

2 Specific Comments

I do not really understand the meaning behind the “same carbon different elements” portion of the title. Unless further explanation is added to the text, I think it could be easily removed. “Same carbon different positions” or something similar seems better aligned with the focus of the paper, but does not necessarily improve the existing title.

Response: “Same carbon different elements” here means: “The position-specific carbons at different positions are the same element but behave like different elements.” In light of the fact that the other reviewer was also puzzled by the title, we have changed our title, reluctantly, to “Carbons at Different Positions Behave Like Different Elements- An Insight into Position-Specific Isotope Distributions”.

It is not always clear throughout the text whether the term “Intra-ID” is being used to mean position-specific isotope analysis (i.e., a measurement) or a calculated distribution of isotopes (i.e., a prediction of an equilibrium state). Adding clarifying language throughout would be helpful.

Response: Thanks for the suggestion. The term Intra-ID refers to the isotope distribution that is measured, predicted, or purely conceptual. To make it more clear, we have specified the measured isotope distribution as “observed/measured/apparent Intra-ID”, the calculated equilibrium state as “predicted/calculated equilibrium Intra-ID”.

Line 34: the definition of position-specific isotope composition is circular since it uses the words ‘specific’ and ‘position’ again. Perhaps something along the lines of “at particular atomic sites within an individual compound” or “at structurally distinct atomic sites. . .” would be more clear.

Response: Thanks for the suggestion. We have revised as suggested. It now reads:

“Position-specific isotope composition refers to the isotope composition of an element at structurally-distinct atomic positions within an individual compound.”

Line 38: “Intra-ID” is an abbreviation that has been introduced in prior papers, for example, He et al., 2020 GCA. It might be helpful to cite this or the earliest use of this phrase to show precedent.

Response: Thanks for the reminder. Initially, we have defined intramolecular isotope distribution as Intra-ID, and intermolecular isotope distribution as Inter-ID in He et al., 2018 RCM. We have added the two references: He et al., 2018 RCM and He et al., 2020 GCA here.

Line 39: can the authors clarify what they mean here? “Most common” in what sense? (i.e., in terms of calculations, measurements, publications, or something else)

Response: We were trying to say that technological development in the carbon position-specific isotope analysis (PSIA) has been very active in recent years. This is in addition to the ongoing debate among Galimov, Buchachenko, and Schmidt on position-specific carbon isotope distributions. However, we realized that PSIA on hydrogen and oxygen is also developing fast. Therefore, this sentence only added confusion. We have deleted it in the new version.

Line 47: “to compare to” does not add much, but could be replaced by “for interpreting position-specific isotope measurements” and I believe makes the authors’ point clearer.

Response: Thanks for the suggestion. We have revised as suggested. It now reads:

“At equilibrium, the difference in isotope composition between two positions depends on temperature only and therefore such difference has been considered as a reference for interpreting position-specific isotope compositions (Galimov, 1985; Hayes, 2001, 2004; He et al., 2018, 2020; Rustad, 2009; Piasecki et al., 2016).”

Line 54: I do not think “correlate loosely” is sufficiently clear here. Perhaps, “do not correlate well with” or “are poorly correlated with”

Response: Thanks for the suggestion. We have revised it to “are poorly correlated with”.

Line 60: the point citing He et al., 2018 seems interesting and highly relevant to the discussion in this paper. Can the authors add one more sentence summarizing the finding of that study and why the correlation approach is invalid?

Response: We have added a few sentences here. It now reads:

“Such a $^{13}\beta$ - $\delta^{13}\text{C}$ correlation is written as $\delta^{13}\text{C} - \delta^{13}\text{C}_{ave} = \chi(\beta - \beta_{ave}) \times 10^3$, where χ is the regression coefficient. Galimov interpreted such observed intramolecular $^{13}\beta$ - $\delta^{13}\text{C}$ correlations as *equilibrium-like* Intra-IDs produced from sets of reversible biochemical reactions at steady-states which

are not far from equilibrium. The $^{13}\beta\text{-}\delta^{13}\text{C}$ correlations were used as supporting evidence that the theorem of minimum entropy production can be applied in biochemical systems. ... The $^{13}\beta\text{-}\delta^{13}\text{C}$ correlation used an unweighted arithmetic mean isotope composition of all the components as the reference of a system. Strictly, only the mass-weighted isotope composition of all components should represent that of a system (Hayes, 2001). In addition, arbitrarily assigning a reference is not mathematically rigorous (He et al., 2018). Therefore, a $^{13}\beta\text{-}\delta^{13}\text{C}$ correlation cannot be used as supporting evidence for Galimov's hypothesis that the theorem of minimum entropy production applies in biochemical systems."

Line 97: It would be helpful to add one summary sentence at the end here to clarify what general point the authors hope the readers will take away from this section.

Response: Thanks for the suggestion. We have added a summary sentence at the end. It reads:

"The use of a single-mineral geothermometer requires that oxygens at two different sites have attained equilibrium by exchanging with each other or with the same source oxygen (e.g. water). Unfortunately, this requirement is difficult to be met for most minerals. It is, therefore, not surprising that few successful single-mineral geothermometers exist if any at all."

Line 137: "mechanisms" might be more meaningful than "processes" here?

Response: We have revised as suggested.

In line 196, I do not understand why any references are needed. It is clear from the chemical formulae, for example, that H₂O consists of H and O atoms and NO₃⁻ consists of N and O. Why are 7 references needed in this sentence?

Response: This paragraph serves to connect to the title "Carbons at Different Positions Behave Like Different Elements- An Insight into Position-Specific Isotope Distributions". All the references are the studies of bonded isotope effects – D and ¹⁸O in H₂O, ¹⁵N and ¹⁸O in NO₃⁻. We have revised this paragraph to cite the references in their respective places. It now reads:

“A simple comparison of position-specific isotope compositions in one sample, e.g. $\ln^{13}\alpha_{carb-met}$ values of one acetic acid sample, offers little information on the reaction it involves. Although the position-specific atoms are the same elements, without an exchange mechanism, they behave independently as if they were different elements. The isotope fractionation relationship of different elements in the same compound, i.e. $(\alpha_A-1)/(\alpha_B-1)$, $\ln\alpha_A/\ln\alpha_B$, or $\Delta\delta_A/\Delta\delta_B$, (named bonded isotope effect, He and Bao, 2019), is useful in characterizing a reaction pathway. Some of the studied examples are δD and $\delta^{18}O$ in H_2O (Dansgaard, 1964; Craig, 1961), $\delta^{15}N$ and $\delta^{18}O$ in NO_3^- (Casciotti and McIlvin, 2007; Wankel et al., 2009), $\delta^{34}S$ and $\delta^{18}O$ in SO_4^{2-} (Antler et al., 2013), and $\delta^{13}C$ and δD in organic compounds (Elsner, 2010; Palau et al., 2017). The isotope composition difference of different elements in a molecule is useful only when the isotope fractionation relationships are considered and their isotope compositions are normalized, e.g. $\Delta(15,18) = (\delta^{15}N - \delta^{15}N_m) - (^{15}\alpha - 1 / ^{18}\alpha - 1) \times (\delta^{18}O - \delta^{18}O_m)$, in which $\delta^{15}N_m$ and $\delta^{18}O_m$ are the average isotope composition in a given ocean water column (Sigman et al., 2005). The normalization procedure was necessary because the source isotope compositions can affect the values of the product. Similarly, if the same element at different positions have different sources, their source isotope composition difference must also be considered. Position-specific isotope researches can build upon our understanding of bonded isotope effect.”

I would expect to see Hayes, 2001 cited somewhere in the manuscript as another classic discussion of kinetic and metabolic controls on isotope signatures of biomolecules.

Response: We have cited Hayes (2001) in multiple places

Hayes (2001) did a comprehensive review of C and H isotope fractionation in biosystems. It focuses more on biochemical processes rather than isotope effects. All the isotope fractionations discussed in Hayes (2001) are apparent ones. It has not been tested that if those values are intrinsic KIEs or EIEs. Thus, at the end of the paper, Hayes also stated: “*It often seems that isotopic fractionations provide **too much** information about **too many** processes, combining it all in a package that is unmanageably intricate.*” Hayes (2004) gave a very good example of Intra-IDs of fatty acid controlled by KIE or EIE. However,

since he did not clearly define the two “equilibrium” scenarios, that example and his interpretation can be interpreted in both ways.

Although Hayes has made important contributions in the area of PSIA, we could not find a publication in which Hayes took a stand in the Galimov-Buchachenko-Schmidt debate publicly. We do not want to second guess his words, therefore, we did not comment on Hayes’ idea in this debate.

3 Suggested minor technical corrections

Line 12 - “biosystem” should be plural

Response: Corrected.

Line 13 - “debates remain” should become “debate remains”

Response: Corrected.

Line 21 - “roots” → “is rooted”

Response: Corrected.

Line 24 - “to be isolated and to be controlled” → “to isolate and control”; “effect” should be plural

Response: Corrected.

Line 39 - “facing” → “faced”

Response: Corrected.

Line 42 - “termed in” → “described as being in”

Response: Corrected.

Line 53 - “contrary” → “contrast”

Response: Corrected.

Line 55 - “are also observed” is not needed

Response: Corrected.

Line 63 - “between” → “among”

Response: Corrected.

Line 71 - eliminate “in contrast to existing optimism”

Response: Corrected.

Line 95 - “can come from different sources”

Response: Corrected.

Line 179 - I believe “involving” should be “evolving” here

Response: Corrected.

Line 206 - “offer” → “offers”

Response: Corrected.

Line 218 - “at molecular level” → “at the molecular level”

Response: Corrected.

Ideas and perspectives: Same Carbon Behave Like Different Elements- An Insight into Position-Specific Isotope Patterns Distributions Within a Single Compound

Yuyang HE ^{1,*}, Xiaobin CAO ^{2,3}, and Huiming BAO ^{2,3,4}

5 ¹Institute of Mechanics, Chinese Academy of Sciences, No.15 Beisihuanxi Road, Beijing 100190, PR China

²International Center for Isotope Effects Research, Nanjing University, Nanjing 210023, P. R. China

³School of Earth Sciences and Engineering, Nanjing University, Nanjing 210023, PR China

⁴Department of Geology and Geophysics, Louisiana State University, E235 Howe Russell Kniffen, Baton Rouge, LA 70803

Correspondence to: Yuyang HE (yhe@imech.ac.cn)

10 **Abstract.** It is expected that information on the source, reaction pathway, and reaction kinetics of an organic compound can be obtained from its position-specific isotope compositions or intramolecular isotope distribution (Intra-ID). To retrieve the information, we could use its predicted equilibrium Intra-ID as a reference for understanding the observed Intra-IDs. Historically, observed, apparently close-to-equilibrium carbon Intra-ID had prompted an open debate on the nature of biosystems and specifically the pervasiveness of reversible biochemical reactions. Much of the debates remains unresolved, and the discussion has not clearly distinguished two states of equilibrium: 1) the equilibrium among the corresponding bond-breaking/forming positions in reactant and product, and 2) the equilibrium among all carbon positions within a compound. For an organic molecule with multiple carbon positions, equilibrium carbon Intra-ID can be attained only when a specific reaction is in equilibrium and the sources of each position are also in equilibrium with each other. An observed Intra-ID provides limited information if the sources and pathways are both unconstrained. Here, we elaborate on this insight using examples of the observed Intra-IDs of hydroxyl-bearing minerals, N₂O, and acetic acid. Research effort aiming at calibrating position-specific equilibrium and kinetic isotope fractionation factors for defined processes will help to interpret observed Intra-IDs of a compound accurately and fully.

1 Introduction

25 Biosystems are dominated by a series of non-equilibrium kinetic processes. The understanding of biosystems is rooted in the study of the biochemical reaction mechanism. However, a majority of the biochemical reaction mechanisms remains elusive since they are difficult to ~~be isolated~~ and ~~to be controlled~~ in laboratory experiments. Stable isotope effects can be used to examine the transition-state structure and reversibility of an elementary reaction, ~~therefore~~ Therefore, it can provide information on reaction mechanisms (Bigeleisen, 1949; Galimov, 2006; Bennet, 2012). However, a big organic molecule produced by an organism is the result of complex biochemical reactions that involve multiple kinetic and equilibrium isotope effects (KIE and EIE). KIE and EIE refer to the two intrinsic parameters for interpreting the observed isotope fractionations

(Bao et al., 2015). According to the transition-state theory (Eyring, 1935a, b), the KIE of an elementary step can be defined as the equilibrium fractionation factor between transition-state and reactant (Jones and Urbauer, 1991; Bao et al., 2015):

$$KIE = \beta_{TS}/\beta_R$$

35 where the β factor denotes the reduced partition function ratio of transition-state (TS) or reactant (R). A β factor is the equilibrium isotope fractionation factor between an atom in a specific bond environment and its atomic form that can be predicted theoretically (Urey, 1947, Bigeleisen and Goepfert-Mayer, 1947). For a unidirectional reaction, the KIE of a reaction can also be defined as:

$$KIE = \frac{h_k}{l_k}$$

40 where k denotes the reaction rate constant of heavy (h) or light (l) isotopes. To adapt to the convention of geochemists, we define KIE this way so that the normal KIE is less than 1.000, which is the opposite of what Bigeleisen (Bigeleisen, 1949) initially defined.

EIE is the isotope fractionation among reactant and product, which is determined by the bonding environment of the target position or compound:

$$EIE = \beta_P/\beta_R$$

45 where P denotes the product of a target reaction. It can also be defined as:

$$EIE = \frac{h_K}{l_K}$$

where K denotes the equilibrium constant of a target reaction. At equilibrium, the EIE of a reaction equals to the ratio of forward reaction KIE_f , and backward reaction KIE_b ($EIE = KIE_f/KIE_b$, Bao et al., 2016).

An organic compound usually contains multiple positions of the same element, such as carbon, hydrogen, oxygen, or nitrogen. 50 Compound-specific isotope composition refers to the bulk isotope composition of an element in an individual compound. Position-specific isotope composition refers to the isotope composition of an element at structurally-distinct atomic positions within a specific position of an individual compound. Information on sources, reaction pathways, and reaction kinetics of an organic molecule are pertinent to each position. The compound-specific isotope composition averages isotope compositions of all different positions of the same element in a compound, where information contained in position-specific isotope 55 compositions could be lost (Elsner, 2010; Piasecki et al., 2018).

We named position-specific isotope compositions in a compound *intramolecular isotope distribution* or *Intra-ID*. (He et al., 2018, 2020). The most common Intra-ID in organic compounds is eCarbon Intra-ID in organic compounds has invoked a long-

~~standing debate about its fundamental controls~~. When ~~faeing faced~~ with the observed diverse Intra-IDs, earlier researchers inferred that the patterns “*must be the expression of some logical order*” (Schmidt, 2003), which is controlled by EIE and KIE of biochemical reactions (e.g. Hayes, 2001, 2004; Galimov, 2009; Schmidt et al., 2015; Eiler et al., 2018; Gilbert et al., 2019). The Intra-ID was ~~termed described as being~~ in *thermodynamic order* or *statistical isotope pattern* when each position in a molecule reaches equilibrium with each other (Galimov, 1985; Schmidt et al., 2015). Here, we name it *equilibrium Intra-ID*. ~~Such a state can be predicted theoretically, i.e. using the reduced partition function ratio (RPFR or β factor) originally defined by Bigeleisen and Mayer (1947).~~ The non-equilibrium state is expected to be a norm for a biochemical system since life is a dissipative system. ~~At An equilibrium, the difference in isotope composition between two positions depends on temperature only state determined by thermodynamic properties is a constant state~~ and therefore ~~the deviation of an observed Intra-ID from its predicted equilibrium state~~ has been considered as an ideal reference ~~state to compare to for interpreting position-specific isotope compositions~~ (Galimov, 1985; Hayes, 2001, 2004; He et al., 2018, 2020; Rustad, 2009; Piasecki et al., 2016).

It has been reported that different carbon fragments of chlorophyll, different carbon positions in acetoin, malonic acid, citric acid, and purine alkaloid have $^{13}\beta$ - $\delta^{13}\text{C}$ correlation with regression coefficients in the range of 0.33-0.51 (Galimov, 1985, 2003, 2004, 2006 and references therein). ~~Such a $^{13}\beta$ - $\delta^{13}\text{C}$ correlation is written as $\delta^{13}\text{C} - \delta^{13}\text{C}_{ave} = \chi(\beta - \beta_{ave}) \times 10^3$, where χ is the regression coefficient.~~ Galimov (~~1985, 2004, 2006~~) interpreted such observed intramolecular $^{13}\beta$ - $\delta^{13}\text{C}$ correlations as *equilibrium-like* Intra-IDs produced from sets of reversible biochemical reactions at steady-states which are not far from equilibrium. ~~The $^{13}\beta$ - $\delta^{13}\text{C}$ correlations were used as supporting evidence that the theorem of minimum entropy production can be applied in biochemical systems.~~ However, other groups interpreted the fair-to-good correlation as fortuitous regardless of the presence or absence of complete reversibility of enzymatic reactions (Buchachenko, 2003, 2007; Schmidt, 2003; Schmidt et al., 2015). In ~~contrary-contrast~~ to these reported observed equilibrium-like Intra-IDs, measured position-specific $\delta^{13}\text{C}$ values ~~are poorly correlated loosely~~ with their predicted $^{13}\beta$ values in organic molecules like glucose, nicotine, and tropine ~~are also observed~~ (Rossmann et al., 1991; Gleixner and Schmidt, 1997; Robins et al., 2016; Romek et al., 2016). Such observed non-equilibrium Intra-ID has been termed *non-statistical isotope pattern* (Rossmann et al., 1991; Gleixner and Schmidt, 1997; Schmidt, 2003; Robins et al., 2016; Romek et al., 2016). Buchachenko (2003, 2007) and Schmidt et al. (2004, 2015) argued that the observed $^{13}\beta$ - $\delta^{13}\text{C}$ correlations are random Intra-ID that only “*simulates*” the thermodynamic state, which cannot be used as evidence for biochemical reactions favoring equilibrium state (~~Buchachenko, 2003, 2007; Schmidt et al., 2004; Schmidt et al., 2015~~). ~~The $^{13}\beta$ - $\delta^{13}\text{C}$ correlation used an unweighted arithmetic mean isotope composition of all the components as the reference of a system. Strictly, only the mass-weighted isotope composition of all components should represent that of a system (Hayes, 2001). In addition, arbitrarily assigning a reference is not mathematically rigorous (He et al., 2018). Therefore, a $^{13}\beta$ - $\delta^{13}\text{C}$ correlation cannot be used as supporting evidence for Galimov’s hypothesis that the theorem of minimum entropy production applies in biochemical systems. We have shown that these $^{13}\beta$ - $\delta^{13}\text{C}$ correlations implicitly normalized the $^{13}\beta$ and $\delta^{13}\text{C}$ values using the averages of a given system, which is not mathematically rigorous and is misleading (He et al., 2018).~~

~~However~~ Nevertheless, the invalidity of $^{13}\beta$ - $\delta^{13}\text{C}$ correlations cannot fully quell the controversy on the nature of biosystem.

It should be noted that the debate on isotope equilibrium in biosystems ~~between-among~~ Galimov, Buchachenko, and Schmidt (and their colleagues) did not clearly distinguish two states of equilibrium: 1) intermolecular isotope equilibrium among the corresponding bond-breaking/forming positions in reactant and product in a defined **process**, and 2) intramolecular isotope equilibrium among all carbon positions in a defined **molecule**. Such a difference might also be overlooked when discussing the Intra-ID or the site preference (SP) value, i.e. the isotope composition difference among two positions. A fully reversible reaction is necessary for isotope equilibrium between corresponding active positions or functional groups. Similarly, a fully reversible intramolecular exchange mechanism must exist if different positions within a compound are to attain equilibrium. However, the overwhelming majority of biochemical reactions, especially in cases involving large organic molecules, have very few intramolecular exchange pathways. Here, ~~in contrast to existing optimism,~~ we propose that the utility of parameters like SP value in organic molecules could be limited before we obtain sufficient details on the source, pathway, as well as KIE and EIE of biochemical reactions. To elaborate this point, we present simple cases starting from hydroxyl-bearing minerals in which oxygen occupies more than one position, to the case of N₂O in which uni-directional and fully reversible reactions can produce similar nitrogen Intra-IDs if there exists a symmetric precursor. After presenting the two inorganic cases, we move to examine measured carbon Intra-IDs from the literature of a simple organic molecule, acetic acid, in which Intra-IDs are ~~pathway-pathway~~-dependent.

2 Intramolecular isotope distribution

2.1 Intracrystalline oxygen isotope difference – a ~~rarely effective failed~~ single mineral geothermometer

The same element, e.g. carbon, occupies different positions in a compound is not a unique feature of organic compounds. Some oxygen-bearing minerals have two or more position-specific oxygens, where the oxygen atoms occupy different positions in a mineral structure and have different chemical properties. Their isotope composition difference was proposed as a potential single mineral geothermometer. For example, it had been proposed that water temperature could be reconstructed from intracrystalline oxygen isotope difference of single mineral copper sulfate pentahydrate (CuSO₄·5H₂O) (Götz et al., 1975), kaolinite (Al₂Si₂O₅(OH)₄), illite (K_{0.65}Al_{2.0}(Al_{0.65}Si_{3.35}O₁₀)(OH)₂) (Bechtel and Hoernes, 1990), or alunite (KAl₃(SO₄)₂(OH)₆) (Arehart et al., 1992). By analyzing the isotope composition difference of different oxygens in the same minerals, the early researchers attempted to reconstruct the precipitation temperatures.

-To be a single-mineral geothermometer, different oxygen sites must have attained equilibrium within the single mineral, which can be achieved when different positions in a compound have the same source or initially different sources are in equilibrium with each other. Take alunite precipitation from a solution as an example. Alunite has sulfate and hydroxyl oxygen positions in its structure that precipitate from sulfate and hydroxyl ions in the solution (Fig. 1). Alunite with equilibrium Intra-ID can be produced from an equilibrium precipitation process, only if both the oxygen isotope compositions of sulfate and hydroxyl ions in the solution equilibrated with the same ambient water oxygen at the same temperature. However, sulfate oxygen does not readily exchange with that of water; the isotope equilibration time for SO₄²⁻ and ambient water at Earth's surface condition is

greater than 10^6 to 10^7 years (Lloyd, 1968; Turchyn and Schrag, 2004; Turchyn et al., 2010) while the oxygen in OH^- can equilibrate with ambient water instantly and can readily exchange during alunite's later burial and diagenetic processes. Thus, the two oxygen positions in alunite can come from different sources at different temperature, rendering alunite a flawed single-mineral geothermometer. The same is true for gypsum ($\text{CaSO}_4 \cdot 2\text{H}_2\text{O}$) in which sulfate oxygen is not in equilibrium with formation water, and the crystallization water ($\cdot 2\text{H}_2\text{O}$) oxygen may be in equilibrium with a different type of water.

The use of a single-mineral geothermometer requires that oxygens at two different sites have attained equilibrium by exchanging with each other or with the same source oxygen (e.g. water). Unfortunately, this requirement is difficult to be met for most minerals. It is, therefore, not surprising that few successful single-mineral geothermometers exist if any at all.

2.2 “Equilibrium-like” Intra-ID produced by a kinetic process

For a compound with two different positions of the same element, a simple way to describe its observed Intra-ID is to report the difference between the two isotope compositions, i.e. the site-preference (SP) value. The concept of SP originated from the study of nitrous oxide ($^{\beta}\text{N}^{\alpha}\text{NO}$), which is defined as the nitrogen isotope composition difference between the center-central nitrogen ($\delta^{15}\text{N}^{\alpha}$) and the terminal nitrogen ($\delta^{15}\text{N}^{\beta}$) (Yoshida and Toyoda, 2000). The predicted equilibrium SP value at room temperature in N_2O is 45‰ (Yung and Miller, 1997; Wang et al., 2004; Cao and Liu, 2012). Although most observations fit the equilibrium prediction that ^{15}N preferentially enriches in the $^{\alpha}\text{N}$ position by 30-40‰ (Yoshida and Toyoda, 2000; Toyoda et al., 2002; Sutka et al., 2006), negative SP values were observed nevertheless (Yamulki et al., 2001; Sutka et al., 2003).

In previous literature, the difference in SP values was explained by the difference in synthetic pathways associated with symmetrical or asymmetrical precursors (Schmidt et al., 2004; Toyoda et al., 2005; Sutka et al., 2006). If the precursor of N_2O is symmetrical (e.g. $-\text{ONNO}-$, Fig. 2 left), the two nitrogens in the precursor are positionally equivalent; any prior isotope composition and fractionation difference would be erased by the symmetrical structures of the precursor. When producing N_2O from a symmetrical precursor, the $^{\beta}\text{N}$ undergoes N-O bond cleavage and therefore has a primary isotope effect which is large, whereas the $^{\alpha}\text{N}$ has only a secondary isotope effect which is negligible (close to 1.000, Bigeleisen and Wolfsberg, 1958). Therefore, ^{15}N depletion is expected only on the $^{\beta}\text{N}$ ~~or~~ of N_2O produced from a symmetrical precursor is expected to have a positive SP value.

If the precursor is asymmetrical (e.g. $-\text{NH}(\text{OH})\text{NO}$, Fig. 2 right), the two nitrogens in the precursor are not positionally equivalent. It is expected that the two nitrogens in the precursor were produced from different EIEs or KIEs because they went through different reaction pathways and may even have different nitrogen sources (~~Schmidt et al., 2004; Toyoda et al., 2005; Sutka et al., 2006~~). Therefore, during the formation of N_2O from an asymmetrical precursor, the difference in the position-specific $\delta^{15}\text{N}$ values of the precursors and the difference in isotope fractionation during the formation processes will be recorded in the SP value of N_2O . Such N_2O can have either $\text{SP} > 0$ or $\text{SP} < 0$.

Nevertheless, the two previously proposed mechanisms cannot distinguish N_2O with $\text{SP} > 0$ produced from the two mechanisms. In addition, for N_2O produced from a symmetric precursor, the SP value cannot provide information on the reaction kinetics, since both fully reversible and uni-directional reactions can produce similar observed SP values. When we state that a

compound displays an equilibrium Intra-ID, the underlying assumption is that there exists a mechanism for different positions to exchange isotopes intramolecularly. However, not all observed apparent equilibrium or equilibrium-like Intra-IDs are produced by an intramolecular equilibrium process. For reactions like $\text{-ONNO-} \leftrightarrow \text{N}_2\text{O}$, two types of processes could produce $\text{SP} > 0$. First, the N_2O formation reaction is fully reversible and attains an equilibrium. When fully reversible, the two nitrogens in N_2O are scrambled when it forms the symmetrical precursor through the reverse reaction. At equilibrium, the terminal nitrogen in a weaker bond environment is expected to be depleted in heavier isotope than the central nitrogen by 45% at surface temperature. Second, the N_2O formation reaction is uni-directional. When uni-directional, only the N-O bond-breaking position ($^{\beta}\text{N}$) undergoes a KIE. Thus, the SP value is approximately equal to the KIE value. In this scenario, if there is a normal KIE < 1.000 , the terminal nitrogen is expected to be depleted in heavier isotope than the central nitrogen by the extent of the KIE value. ~~The Such an observed~~ Intra-ID would be similar to the predicted equilibrium Intra-ID ~~in this case~~, but it is produced by isotope depletion on the uni-directional bond-breaking process. No intramolecular exchange involves. Therefore, even if the N_2O produced by the uni-directional process has $\text{SP} \approx 45\%$, it is not due to a close-to-equilibrium or equilibrium-like SP intramolecular isotope exchange. Therefore, it is necessary to distinguish the mechanisms and reaction kinetics that can produce an observed Intra-ID.

Here we see that both fully reversible and uni-directional processes can result in a similar SP value, but the underlying mechanisms are entirely different. Furthermore, a positive SP value can also be achieved through a combination of nitrogen sources and isotope fractionations from an asymmetrical precursor. Thus, without knowing the underlying process, we cannot interpret an observed Intra-ID or SP value uniquely.

2.3 Position-specific isotope fractionations between reactant and product

As illustrated above, the observed Intra-ID of a compound can be used to gauge the degree of internal thermodynamic equilibrium only if we can determine the ~~processes-mechanisms~~ involved in isotope fractionation. It does not mean, however, that position-specific isotope composition is useless. Based on the predicted equilibrium Intra-ID, a predicted isotope fractionation factor (α) of corresponding positions between reactant and product in a process can help to evaluate the thermodynamic state of a system and to decipher reaction pathways. In this section, we use a simple organic molecule, acetic acid (CH_3COOH), and its measured Intra-IDs from literature as examples to illustrate how position-specific isotope fractionation occurs between reactant and product.

The relative isotope enrichment between the carboxyl and methyl carbon in acetic acid is defined as $\ln^{13}\alpha_{\text{carb-met}} = \ln(^{13}\text{R}_{\text{carb}}/^{13}\text{R}_{\text{met}}) \times 1000\%$. ^{13}R ($= ^{13}\text{C}/^{12}\text{C}$) denotes the carbon isotope molar abundance ratio in a position. Our calculated equilibrium Intra-ID of acetic acid has the carboxyl carbon being 47.3 ‰ heavier than the methyl carbon at 25°C ($\ln^{13}\alpha_{\text{carb-met}} = 47.3\%$, He et al., 2020). The measured $\delta^{13}\text{C}_{\text{met}}$ values from literature can be lower, higher, or approximately equal to the $\delta^{13}\text{C}_{\text{carb}}$ values for acetic acids from biological, artificial, or hydrous pyrolysis samples (Table 1). The position-specific $\delta^{13}\text{C}$ values of biological, artificial, or hydrous pyrolysis produced acetic acid are largely overlapping on $\delta^{13}\text{C}_{\text{met}} - \delta^{13}\text{C}_{\text{carb}}$ space. For the majority of biological acetic acids, the $\delta^{13}\text{C}_{\text{carb}}$ values are several per mil higher than the $\delta^{13}\text{C}_{\text{met}}$ values (Fig. 3 top, $\ln^{13}\alpha_{\text{carb-met}}$

$\alpha_{carb-met} = 5.1 \pm 4.8\%$, $n=29$), with two cases of $\sim 18\%$ higher and one case of -2.2% lower in $\delta^{13}C_{carb}$ values. It is expected that the metabolic and catabolic pathways and carbon sources are limited for most natural acetic acid. Therefore, the $\ln^{13}\alpha_{carb-met}$ value of $5.1 \pm 4.8\%$ could be characteristic but not necessarily exclusive for biologically produced acetic acid. ~~ArtificialMan-made~~ acetic acids have a very large range of $\ln^{13}\alpha_{carb-met}$ values from -30.2% to 24.2% (Fig. 3 middle, $7.3 \pm 14.3\%$, $n=24$). Biological and hydrous pyrolysis produced acetic acids do not have such negative $\ln^{13}\alpha_{carb-met}$ values.

190
195
-Except for the above-mentioned features, the production of ~~artificialman-made~~ and biological acetic acid has too many unconstrained parameters. Thus, our discussion will focus on the acetic acid derived from hydrous pyrolysis of oil-prone source rocks.

The acetic acids produced from the hydrous pyrolysis of oil-prone source rocks have a $\ln^{13}\alpha_{carb-met}$ value of $18.3 \pm 7.7\%$ ($n=22$, Fig. 3 bottom). $\ln^{13}\alpha_{carb-met}$ values of $\sim 30\%$ were produced at $310\sim 350$ °C from Mahogany Shale or Black Shale with a proposed mechanism of uni-directional pyrolysis of precursor acid forms ($R-CH_2COOH \leftrightarrow R + CH_3COOH$, Fig. 4, Dias et al., 2002b).
200 If we consider only the primary KIE between the methylene carbon in $R-^*CH_2COOH$ and the methyl carbon in acetic acid (*CH_3COOH), it is expected that a uni-directional process would lead to a ^{13}C depletion only on the methyl carbon position in acetic acid. The observed Intra-ID of the produced acetic acid should equal to the $\delta^{13}C$ value difference between the precursors minus the primary KIE value. The primary KIE value is expected to be more negative than the predicted equilibrium isotope fractionation factor, which is -14% (He et al., 2020). Thus, as long as the position-specific $\delta^{13}C$ value difference between the
205 methylene and carboxyl carbon in $R-CH_2COOH$ is greater than -14% , the acetic acid produced from uni-directional pyrolysis of such precursor acid should have a carboxyl carbon with a higher $\delta^{13}C$ value than that of the methyl carbon. If the carboxyl carbon in the precursor acid has a higher $\delta^{13}C$ than that of the methylene carbon, the pyrolysis process can easily produce acetic acid with an apparent $\ln^{13}\alpha_{carb-met}$ value close to the predicted an apparent equilibrium Intra-ID. Such apparently “equilibrium-like” Intra-ID does not involve intramolecular exchange, but it is the product of uni-directional precursor acid
210 pyrolysis.

3 Implications

Life sustains itself by feeding on negative entropy. Persistent efforts are devoted into describing living systems by rigorous mathematics. Boltzmann first considered living organisms from a thermodynamic perspective, and Schrodinger later applied equilibrium thermodynamics to living systems (Popovic, 2018). Those attempts were not pursued further, since, as we all know
215 today, a living system is an open system that is away from thermodynamic equilibrium, i.e. a dissipative system. The establishment of nonequilibrium thermodynamics by Prigogine and his coworkers has guided researchers to the theorem of minimum entropy production in biological systems (Prigogine and Wiame, 1946). The theorem of minimum entropy production in biological systems describes that although a biosystem has increasing entropy, but it is usually stable in a steady state. If the system displaced from the steady state, it tends to return to its original state since the entropy is minimal at the

220 state. Since then, efforts in applying nonequilibrium thermodynamics to living systems have been continued with mixed success (Stoward, 1962; Schneider and Kay, 1994; Hayflick, 2007; Demirel, 2010; Barbacci et al., 2015; Gerber et al., 2016). The theorem ~~of on~~-minimum entropy production applies only to linear thermodynamic systems. Therefore, it is necessary to demonstrate that the magnitude of reaction rate on the scale of interest in a living system is linearly dependent on the generalized driving force responsible for the reaction system operating on the system. It is reasonable to view that a complex
225 interacting and constantly ~~ein~~volving non-linear system is constructed by a series of synergistic reactions, and there should exist local linearity, local steady-state, even local equilibrium (Galimov, 2006).

Local nonequilibrium of biochemical system is potentially significant for the increasing complexity and orderliness-ordering in the structure of life (Prigogine and Wiame, 1946; Galimov, 2006). Such a system should consist of a set of reversible but not necessarily equilibrium reactions conjugated with energy supplies that maintain in a steady-state not far from equilibrium.

230 ~~Galimov (1985, 2004, 2005) argued that s~~Such a close-to-equilibrium steady-state should be expressed as a tendency toward equilibrium inter- and intra-molecular stable isotope distributions, i.e. a linear inter- and intra-molecular $^{13}\beta$ - $\delta^{13}\text{C}$ correlation with a regression coefficient smaller than but close to 1 (~~Galimov, 2006~~). ~~The observed correlations between position-specific $\delta^{13}\text{C}$ and $^{13}\beta$ had been used to support the hypothesis that the theorem of minimum entropy production can be applied in biochemical systems (Galimov, 1985, 2004, 2006).~~

235 In addition, such “*equilibrium-like*” Intra-ID in organic molecules was proposed as a “*special feature of biological systems*”, which could be used as a criterion to identify biologically produced extraterrestrial organic molecules (Galimov, 2003). As we have illustrated above, observed Intra-ID in organic molecules is the product of a set of equilibrium or ~~disnon~~-equilibrium processes as well as their source isotope compositions. An observed Intra-ID itself cannot be used as conclusive evidence for the thermodynamic state of a system. Therefore, even if a compound does have a linear intramolecular $^{13}\beta$ - $\delta^{13}\text{C}$ correlation with a slope of 1, it does not constitute supporting evidence for the
240 existence of an equilibrium state among biochemical reactions in organisms. To apply nonequilibrium thermodynamics on living systems, further solid evidence is needed.

A simple comparison of position-specific isotope compositions in one sample, e.g. $\ln^{13}\alpha_{\text{carb-met}}$ values of one acetic acid sample, offers little information on the reaction mechanisms and reaction kinetics of the reaction it involves. Although the position-specific atoms are the same element, without an exchange mechanism, they behave independently as different elements. It would be helpful if we could consider the position-specific atoms independently. The isotope fractionation relationship of different elements in the same compound, i.e. $(\alpha_A - 1)/(\alpha_B - 1)$, $\ln\alpha_A/\ln\alpha_B$, or $\Delta\delta_A/\Delta\delta_B$, (named bonded isotope effect, He and Bao, 2019), is often used to characterize a reaction pathway. For A compound often consists of different elements, for instance, $\delta\text{D/H}$ and $\delta^{18}\text{O}$ in H_2O (Dansgaard, 1964; Craig, 1961), $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ in NO_3^- (Casciotti and McIlvin, 2007; Wankel et al., 2009), $\delta^{34}\text{S}$ and $\delta^{18}\text{O}$ in SO_4^{2-} (Antler et al., 2013), or $\delta^{13}\text{C}$ and $\delta\text{D/H}$ in organic compounds (Elsner, 2010; Palau et al., 2017).

250 ~~The isotope fractionation relationship between these different elements, i.e. $(\alpha_A - 1)/(\alpha_B - 1)$, $\ln\alpha_A/\ln\alpha_B$, or $\Delta\delta_A/\Delta\delta_B$, is often used to characterize a reaction pathway.~~The isotope composition difference of different elements is only useful if the isotope fractionation relationships are considered and their isotope compositions are normalized, e.g. $\Delta(15,18) = (\delta^{15}\text{N} - \delta^{15}\text{N}_m) - (15\alpha - 1/18\alpha - 1) \times (\delta^{18}\text{O} - \delta^{18}\text{O}_m)$, $\delta^{15}\text{N}_m$ and $\delta^{18}\text{O}_m$ are the average isotope composition in a given ocean water column profile (Sigman

et al., 2005). The normalization procedure ~~was is~~ necessary because the source isotope compositions can affect the values of the product. Similarly, if the same element at different positions have different sources, their source isotope composition difference must also be considered. ~~In fact, the two or more oxygens in the same compound do not have a mechanism to exchange; these oxygens behave like different elements. A simple comparison of position-specific isotope compositions in one sample, e.g. $\ln^{13}\alpha_{carb-mer}$ values of one acetic acid sample, offer little information. Position-specific isotope researches can build upon our understanding of bonded isotope effect.~~

260 4 Conclusions

~~An O~~rganic compounds usually ~~have has~~ an element, e.g. carbon, at different positions and therefore ~~hasve an~~ Intra-IDs. The deviation of an observed Intra-ID from its equilibrium state has been used to evaluate the thermodynamic state of a system. Our analysis of oxygen-bearing minerals, N₂O, and acetic acids show that both isotope sources and all reaction processes need to be in equilibrium to reach an intramolecular equilibrium state. However, such a condition is rarely satisfied. When different positions of the same element cannot exchange with each other, these different positions behave independently like different elements. Observed Intra-ID that is apparently similar to the equilibrium one can also be produced from a combination of different sources and uni-directional processes. Thus, an observed Intra-ID itself is not conclusive without adequate information on sources and reaction kinetics. Compared to position-specific isotope compositions, position-specific isotope fractionation of a defined process is more informative to identifying bond-breaking/forming positions of a large molecule, to predicting its transition-state structure, to evaluating the reversibility of a biochemical process, and to determining and qualifying a process in a complex system. All in all, an understanding of a reaction process at the molecular level will always be the first step required for later sound and wide application of stable isotope composition.

Author contribution

All authors contributed ideas. Y. He and H. Bao wrote the manuscript.

275 Code/Data availability

All data, models, and code generated or used during the study appear in the submitted article.

Competing interests

The authors declare that they have no conflict of interest.

Acknowledgements

280 This work was supported by the China Postdoctoral Science Foundation [2019M660811] to Y.H.; the National Natural Science Foundation of China [41490635] and the strategic priority research program (B), Chinese Academy of Sciences [XDB18010100] to H.B.

References

- 285 Antler, G., Turchyn, A. V., Rennie, V., Herut, B., and Sivan, O.: Coupled sulfur and oxygen isotope insight into bacterial sulfate reduction in the natural environment, *Geochim. Cosmochim. Ac.*, 118, 98-117, <https://doi.org/10.1016/j.gca.2013.05.005>, 2013.
- Arehart, G. B., Kesler, S. E., O'Neil, J. R., and Foland, K.: Evidence for the supergene origin of alunite in sediment-hosted micron gold deposits, Nevada, *Econ. Geol.*, 87, 263-270, <https://doi.org/10.2113/gsecongeo.87.2.263>, 1992.
- 290 [Aufort, J., Ségalen, L., Gervais, C., Paulatto, L., Blanchard, M., and Balan, E.: Site-specific equilibrium isotopic fractionation of oxygen, carbon and calcium in apatite, *Geochimica et Cosmochimica Acta*, 219, 57-73, 10.1016/j.gca.2017.09.020, 2017.](#)
- Bao, H., Cao, X., and Hayles, J. A.: The confines of triple oxygen isotope exponents in elemental and complex mass-dependent processes, *Geochim. Cosmochim. Ac.*, 170, 39-50, <https://doi.org/10.1016/j.gca.2015.07.038>, 2015.
- [Bao, H., Cao, X., and Hayles, J. A.: Triple Oxygen Isotopes: Fundamental Relationships and Applications, *Annu. Rev. Earth Pl. Sc.*, 44, 463-492, <https://doi.org/10.1146/annurev-earth-060115-012340>, 2016.](#)
- 295 Barbacci, A., Magnenet, V., and Lahaye, M.: Thermodynamical journey in plant biology, *Front. Plant Sci.*, 6, 481, <https://doi.org/10.3389/fpls.2015.00481>, 2015.
- Bechtel, A., and Hoernes, S.: Oxygen isotope fractionation between oxygen of different sites in illite minerals: a potential single-mineral thermometer, *Contrib. Mineral. Petr.*, 104, 463-470, <https://doi.org/10.1007/BF01575623>, 1990.
- Bennet, A. J.: Kinetic isotope effects for studying post-translational modifying enzymes, *Curr. Opin. Chem. Biol.*, 16, 472-300 478, <https://doi.org/10.1016/j.cbpa.2012.10.024>, 2012.
- Bigeleisen, J., and ~~Goeppert~~-Mayer, M.-G.: Calculation of equilibrium constants for isotopic exchange reactions, *J. Chem. Phys.*, 15, 261-267, <https://doi.org/10.1063/1.1746492>, 1947.
- Bigeleisen, J.: The Relative Reaction Velocities of Isotopic Molecules, *J. Chem. Phys.*, 17, 675-678, <https://doi.org/10.1063/1.1747368>, 1949.
- 305 Bigeleisen, J., and Wolfsberg, M.: Theoretical and experimental aspects of isotope effects in chemical kinetics, *Adv. Chem. Phys.*, 15-76, <https://doi.org/10.1002/9780470143476.ch2>, 1958.
- Buchachenko, A. L.: Are chemical transformations in an enzyme-substrate complex reversible? The experience of the fractionation of isotopes in enzymatic reactions, *Russ. J. Phys. Ch.*, 77, 1298-1302, 2003.
- Buchachenko, A. L.: On isotope fractionation in enzymatic reaction, *Russ. J. Phys. Chem. A*, 81, 836-836, 310 <https://doi.org/10.1134/s0036024407050330>, 2007.

- Cao, X., and Liu, Y.: Theoretical estimation of the equilibrium distribution of clumped isotopes in nature, *Geochim. Cosmochim. Ac.*, 77, 292-303, <https://doi.org/10.1016/j.gca.2011.11.021>, 2012.
- Casciotti, K., and McIlvin, M.: Isotopic analyses of nitrate and nitrite from reference mixtures and application to Eastern Tropical North Pacific waters, *Mar. Chem.*, 107, 184-201, <https://doi.org/10.1016/j.marchem.2007.06.021>, 2007.
- 315 Craig, H.: Isotopic variations in meteoric waters, *Science*, 133, 1702-1703, <https://doi.org/10.1126/science.133.3465.1702>, 1961.
- Dansgaard, W.: Stable isotopes in precipitation, *Tellus*, 16, 436-468, <https://doi.org/10.3402/tellusa.v16i4.8993>, 1964.
- Demirel, Y.: Nonequilibrium thermodynamics modeling of coupled biochemical cycles in living cells, *J. Non-Newton. Fluid*, 165, 953-972, <https://doi.org/10.1016/j.jnnfm.2010.02.006>, 2010.
- 320 Dias, R. F., Freeman, K. H., and Franks, S. G.: Gas chromatography–pyrolysis–isotope ratio mass spectrometry: a new method for investigating intramolecular isotopic variation in low molecular weight organic acids, *Org. Geochem.*, 33, 161-168, [https://doi.org/10.1016/S0146-6380\(01\)00141-3](https://doi.org/10.1016/S0146-6380(01)00141-3), 2002a.
- Dias, R. F., Freeman, K. H., Lewan, M. D., and Franks, S. G.: $\delta^{13}\text{C}$ of low-molecular-weight organic acids generated by the hydrous pyrolysis of oil-prone source rocks, *Geochim. Cosmochim. Ac.*, 66, 2755-2769, [https://doi.org/10.1016/S0016-7037\(02\)00871-2](https://doi.org/10.1016/S0016-7037(02)00871-2), 2002b.
- 325 Eiler, J. M., Clog, M., Lawson, M., Lloyd, M., Piasecki, A., Ponton, C., Xie, H., Lawson, M., Formolo, M. J., and Eiler, J. M.: The isotopic structures of geological organic compounds, in: *From Source to Seep: Geochemical Applications in Hydrocarbon Systems*, Geological Society of London, 2018.
- Elsner, M.: Stable isotope fractionation to investigate natural transformation mechanisms of organic contaminants: principles, prospects and limitations, *J. Environ. Monitor.*, 12, 2005-2031, <https://doi.org/10.1039/c0em00277a>, 2010.
- 330 [Eyring, H.: The activated complex and the absolute rate of chemical reactions, *Chemical Reviews*, 17, 65-77, 10.1021/cr60056a006, 1935a.](https://doi.org/10.1021/cr60056a006)
- [Eyring, H.: The activated complex in chemical reactions, *Journal of Chemical Physics*, 3, 107-115, 10.1063/1.1749604, 1935b.](https://doi.org/10.1063/1.1749604)
- Galimov, E. M.: *The Biological Fractionation of Isotopes*, Academic Press, Orlando, 1985.
- 335 Galimov, E. M.: Isotopic Criteria for Identification of Organic Carbon on Earth and Meteorites, *Space Sci. Rev.*, 106, 249-262, <https://doi.org/10.1023/a:1024697923188>, 2003.
- Galimov, E. M.: Phenomenon of life: between equilibrium and non-linearity, *Origins Life Evol. B.*, 34, 599-613, <https://doi.org/10.1023/B:ORIG.0000043131.86328.9d>, 2004.
- Galimov, E. M.: Phenomenon of life: between equilibrium and non-linearity. Origin and principles of evolution, *Geochem. Int+*, 44, S1-S95, <https://doi.org/10.1134/S0016702906130015>, 2006.
- 340 Galimov, E. M.: Advances in geochemistry during the last four decades: A personal perspective, *Appl. Geochem.*, 24, 1048-1051, <https://doi.org/10.1016/j.apgeochem.2009.02.020>, 2009.
- Gelwicks, J. T., Risatti, J. B., and Hayes, J.: Carbon isotope effects associated with autotrophic acetogenesis, *Org. Geochem.*, 14, 441-446, [https://doi.org/10.1016/0146-6380\(89\)90009-0](https://doi.org/10.1016/0146-6380(89)90009-0), 1989.

- 345 Gerber, S., Fröhlich, M., Lichtenberg-Fraté, H., Shabala, S., Shabala, L., and Klipp, E.: A thermodynamic model of monovalent cation homeostasis in the yeast *Saccharomyces cerevisiae*, *PLoS Comput. Biol.*, 12, e1004703, <https://doi.org/10.1371/journal.pcbi.1004703>, 2016.
- Gilbert, A., Sherwood Lollar, B., Musat, F., Giunta, T., Chen, S., Kajimoto, Y., Yamada, K., Boreham, C. J., Yoshida, N., and Ueno, Y.: Intramolecular isotopic evidence for bacterial oxidation of propane in subsurface natural gas reservoirs, *P. Natl. A. Sci.*, 116, 6653-6658, <https://doi.org/10.1073/pnas.1817784116>, 2019.
- 350 Gleixner, G., and Schmidt, H.-L.: Carbon isotope effects on the fructose-1, 6-bisphosphate aldolase reaction, origin for non-statistical ¹³C distributions in carbohydrates, *J. Biol. Chem.*, 272, 5382-5387, <https://doi.org/10.1074/jbc.272.9.5382>, 1997.
- Götz, D., Heinzinger, K., and Klemm, A.: The Dependence of the Oxygen Isotope Fractionation in the Hydration Water of CuSO₄ · 5 H₂O on the Crystallization Temperature, *Z. Naturforsch. A*, 30, 1667-1674, <https://doi.org/10.1515/zna-1975-1225>,
355 1975.
- Hattori, R., Yamada, K., Kikuchi, M., Hirano, S., and Yoshida, N.: Intramolecular carbon isotope distribution of acetic acid in vinegar, *J. Agr. Food Chem.*, 59, 9049-9053, <https://doi.org/10.1021/jf200227e>, 2011.
- Hayes, J. M.: Isotopic order, biogeochemical processes, and earth history - Goldschmidt Lecture, Davos, Switzerland, August 2002, *Geochim. Cosmochim. Ac.*, 68, 1691-1700, <https://doi.org/10.1016/j.gca.2003.10.023>, 2004.
- 360 [Hayes, J.M.: Fractionation of carbon and hydrogen isotopes in biosynthetic processes. *Rev. Mineral Geochem.* 43, 225-277. <https://doi.org/10.1016/10.2138/gsrmg.43.1.225>.](https://doi.org/10.1016/10.2138/gsrmg.43.1.225)
- He, Y., Cao, X., Wang, J., and Bao, H.: Identifying apparent local stable isotope equilibrium in a complex non-equilibrium system, *Rapid Commun. Mass Sp.*, 32, 306-310, <https://doi.org/10.1002/rcm.8040>, 2018.
- [He, Y., and Bao, H.: Predicting High-Dimensional Isotope Relationships from Diagnostic Fractionation Factors in Systems with Diffusional Mass Transfer, *ACS Earth and Space Chemistry*, 3, 120-128, \[10.1021/acsearthspacechem.8b00149\]\(https://doi.org/10.1021/acsearthspacechem.8b00149\), 2019.](https://doi.org/10.1021/acsearthspacechem.8b00149)
- 365 [He, Y., Bao, H., and Liu, Y.: Predicting equilibrium intramolecular isotope distribution within a large organic molecule by the cutoff calculation, *Geochim. Cosmochim. Ac.*, 269, 292-302, <https://doi.org/10.1016/j.gca.2019.10.032>, 2020.](https://doi.org/10.1016/j.gca.2019.10.032)
- Jones, J. P., and Urbauer, J. L.: Theoretical kinetic isotope effects for the hydride transfer from formate to carbon dioxide: a comparison of theory with experiment, *J. Comput. Chem.*, 12, 1134-1141, <https://doi.org/10.1002/jcc.540120912>, 1991.
- 370 Lloyd, R. M.: Oxygen isotope behavior in the Sulfate-Water System, *J. Geophys. Res.*, 73, 6099-6110, <https://doi.org/10.1029/JB073i018p06099>, 1968.
- Meinschein, W., Rinaldi, G., Hayes, J., and Schoeller, D.: Intramolecular isotopic order in biologically produced acetic acid, *Biol. Mass. Spectrom.*, 1, 172-174, <https://doi.org/10.1002/bms.1200010306>, 1974.
- Nimmanwudipong, T., Gilbert, A., Yamada, K., and Yoshida, N.: Analytical method for simultaneous determination of bulk and intramolecular ¹³C-isotope compositions of acetic acid, *Rapid Commun. Mass Sp.*, 29, 2337-2340, <https://doi.org/10.1002/rcm.7398>, 2015.
- 375 Palau, J., Shouakar-Stash, O., Hatijah Mortan, S., Yu, R., Rosell, M., Marco-Urrea, E., Freedman, D. L., Aravena, R., Soler, A., and Hunkeler, D.: Hydrogen Isotope Fractionation during the Biodegradation of 1, 2-Dichloroethane: Potential for Pathway

- Identification Using a Multi-element (C, Cl, and H) Isotope Approach, *Environ. Sci. Technol.*, 51, 10526-10535, <https://doi.org/10.1021/acs.est.7b02906>, 2017.
- 380 [Piasecki, A., Sessions, A., Peterson, B., and Eiler, J.: Prediction of equilibrium distributions of isotopologues for methane, ethane and propane using density functional theory, *Geochim. Cosmochim. Ac.*, 190, 1-12, 2016.](#)
- Piasecki, A., Sessions, A., Lawson, M., Ferreira, A., Neto, E. S., Ellis, G. S., Lewan, M. D., and Eiler, J. M.: Position-specific ^{13}C distributions within propane from experiments and natural gas samples, *Geochim. Cosmochim. Ac.*, 220, 110-124, <https://doi.org/10.1016/j.gca.2017.09.042>, 2018.
- 385 Popovic, M.: Living organisms from Prigogine's perspective: an opportunity to introduce students to biological entropy balance, *J. Biol. Educ.*, 52, 294-300, <https://doi.org/10.1080/00219266.2017.1357649>, 2018.
- Prigogine, I., and Wiame, J. M.: *Biologie et thermodynamique des phénomènes irréversibles*, *Experientia*, 2, 451-453, <https://doi.org/10.1007/bf02153597>, 1946.
- 390 Rinaldi, G., Meinschein, W., and Hayes, J.: Intramolecular carbon isotopic distribution in biologically produced acetoin, *Biol. Mass. Spectrom.*, 1, 415-417, <https://doi.org/10.1002/bms.1200010609>, 1974.
- Robins, R. J., Romek, K. M., Remaud, G. S., and Paneth, P.: Non-statistical isotope fractionation as a novel "retro-biosynthetic" approach to understanding alkaloid metabolic pathways, *Phytochem. Lett.*, 20, 499-506, <https://doi.org/10.1016/j.phytol.2016.12.023>, 2016.
- 395 Romek, K. M., Remaud, G. S., Silvestre, V., Paneth, P., and Robins, R. J.: Non-statistical ^{13}C fractionation distinguishes co-incident and divergent steps in the biosynthesis of the alkaloids nicotine and tropine, *J. Biol. Chem.*, 291, 16620-16629, <https://doi.org/10.1074/jbc.M116.734087>, 2016.
- Rossmann, A., Butzenlechner, M., and Schmidt, H. L.: Evidence for a nonstatistical carbon isotope distribution in natural glucose, *Plant Physiol.*, 96, 609-614, <https://doi.org/10.1104/pp.96.2.609>, 1991.
- 400 Schmidt, H. L.: Fundamentals and systematics of the non-statistical distributions of isotopes in natural compounds, *Naturwissenschaften*, 90, 537-552, <https://doi.org/10.1007/s00114-003-0485-5>, 2003.
- Schmidt, H. L., Werner, R. A., Yoshida, N., and Well, R.: Is the isotopic composition of nitrous oxide an indicator for its origin from nitrification or denitrification? A theoretical approach from referred data and microbiological and enzyme kinetic aspects, *Rapid Commun. Mass Sp.*, 18, 2036-2040, <https://doi.org/10.1002/rcm.1586>, 2004.
- 405 Schmidt, H. L., Robins, R. J., and Werner, R. A.: Multi-factorial in vivo stable isotope fractionation: causes, correlations, consequences and applications, *Isot. Environ. Healt. S.*, 51, 155-199, <https://doi.org/10.1080/10256016.2015.1014355>, 2015.
- Schneider, E. D., and Kay, J. J.: Life as a manifestation of the second law of thermodynamics, *Math. Comput. Model.*, 19, 25-48, [https://doi.org/10.1016/0895-7177\(94\)90188-0](https://doi.org/10.1016/0895-7177(94)90188-0), 1994.
- 410 Sigman, D. M., Granger, J., DiFiore, P. J., Lehmann, M. M., Ho, R., Cane, G., and van Geen, A.: Coupled nitrogen and oxygen isotope measurements of nitrate along the eastern North Pacific margin, *Global Biogeochem. Cy.*, 19, <https://doi.org/10.1029/2005GB002458>, 2005.
- Stoward, P. J.: Thermodynamics of Biological Growth, *Nature*, 194, 977-978, <https://doi.org/10.1038/194977a0>, 1962.

- Sutka, R., Ostrom, N., Ostrom, P., Gandhi, H., and Breznak, J.: Nitrogen isotopomer site preference of N₂O produced by *Nitrosomonas europaea* and *Methylococcus capsulatus* Bath, *Rapid Commun. Mass Sp.*, 17, 738-745, 415 <https://doi.org/10.1002/rcm.968>, 2003.
- Sutka, R. L., Ostrom, N., Ostrom, P., Breznak, J., Gandhi, H., Pitt, A., and Li, F.: Distinguishing nitrous oxide production from nitrification and denitrification on the basis of isotopomer abundances, *Appl. Environ. Microb.*, 72, 638-644, <https://doi.org/10.1128/AEM.72.1.638-644.2006>, 2006.
- Thomas, B., Freeman, K. H., and Arthur, M. A.: Intramolecular carbon isotopic analysis of acetic acid by direct injection of 420 aqueous solution, *Org. Geochem.*, 40, 195-200, <https://doi.org/10.1016/j.orggeochem.2008.10.011>, 2009.
- Toyoda, S., Yoshida, N., Miwa, T., Matsui, Y., Yamagishi, H., Tsunogai, U., Nojiri, Y., and Tsurushima, N.: Production mechanism and global budget of N₂O inferred from its isotopomers in the western North Pacific, *Geophys. Res. Lett.*, 29, <https://doi.org/10.1029/2001GL014311>, 2002.
- Toyoda, S., Mutoke, H., Yamagishi, H., Yoshida, N., and Tanji, Y.: Fractionation of N₂O isotopomers during production by 425 denitrifier, *Soil Biol. Biochem.*, 37, 1535-1545, <https://doi.org/10.1016/j.soilbio.2005.01.009>, 2005.
- Turchyn, A. V., and Schrag, D. P.: Oxygen isotope constraints on the sulfur cycle over the past 10 million years, *Science*, 303, <https://doi.org/10.1126/science.1092296>, 2004.
- Turchyn, A. V., Brüchert, V., Lyons, T. W., Engel, G. S., Balci, N., Schrag, D. P., and Brunner, B.: Kinetic oxygen isotope effects during dissimilatory sulfate reduction: a combined theoretical and experimental approach, *Geochim. Cosmochim. Ac.*, 430 74, 2011-2024, <https://doi.org/10.1016/j.gca.2010.01.004>, 2010.
- [Urey, H.C. The thermodynamic properties of isotopic substances. J. Chem. Soc. \(Resumed\), 562-581, https://doi.org/10.1039/jr9470000562, 1974.](https://doi.org/10.1039/jr9470000562)
- Wang, Z., Schauble, E. A., and Eiler, J. M.: Equilibrium thermodynamics of multiply substituted isotopologues of molecular gases, *Geochim. Cosmochim. Ac.*, 68, 4779-4797, <https://doi.org/10.1016/j.gca.2004.05.039>, 2004.
- 435 Wankel, S. D., Kendall, C., and Paytan, A.: Using nitrate dual isotopic composition ($\delta^{15}\text{N}$ and $\delta^{18}\text{O}$) as a tool for exploring sources and cycling of nitrate in an estuarine system: Elkhorn Slough, California, *J. Geophys. Res.: Biogeol.*, 114, <https://doi.org/10.1029/2008JG000729>, 2009.
- Yamada, K., Tanaka, M., Nakagawa, F., and Yoshida, N.: On-line measurement of intramolecular carbon isotope distribution of acetic acid by continuous-flow isotope ratio mass spectrometry, *Rapid Commun. Mass Sp.*, 16, 1059-1064, 440 <https://doi.org/10.1002/rcm.678>, 2002.
- Yamada, K., Kikuchi, M., Gilbert, A., Yoshida, N., Wasano, N., Hattori, R., and Hirano, S.: Evaluation of commercially available reagents as a reference material for intramolecular carbon isotopic measurements of acetic acid, *Rapid Commun. Mass Sp.*, 28, 1821-1828, <https://doi.org/10.1002/rcm.6964>, 2014.
- Yamulki, S., Toyoda, S., Yoshida, N., Veldkamp, E., Grant, B., and Bol, R.: Diurnal fluxes and the isotopomer ratios of N₂O 445 in a temperate grassland following urine amendment, *Rapid Commun. Mass Sp.*, 15, 1263-1269, <https://doi.org/10.1002/rcm352>, 2001.

Yoshida, N., and Toyoda, S.: Constraining the atmospheric N₂O budget from intramolecular site preference in N₂O isotopomers, *Nature*, 405, 330, <https://doi.org/10.1038/35012558>, 2000.

Yung, Y. L., and Miller, C. E.: Isotopic fractionation of stratospheric nitrous oxide, *Science*, 278, 1778-1780, 450 <https://doi.org/10.1126/science.278.5344.1778>, 1997.

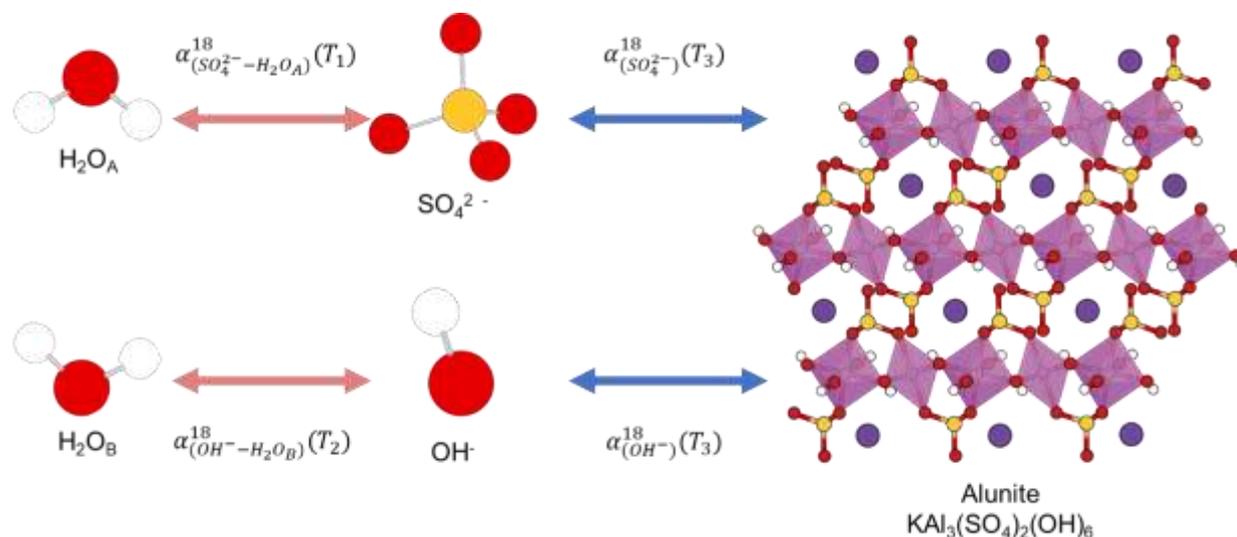
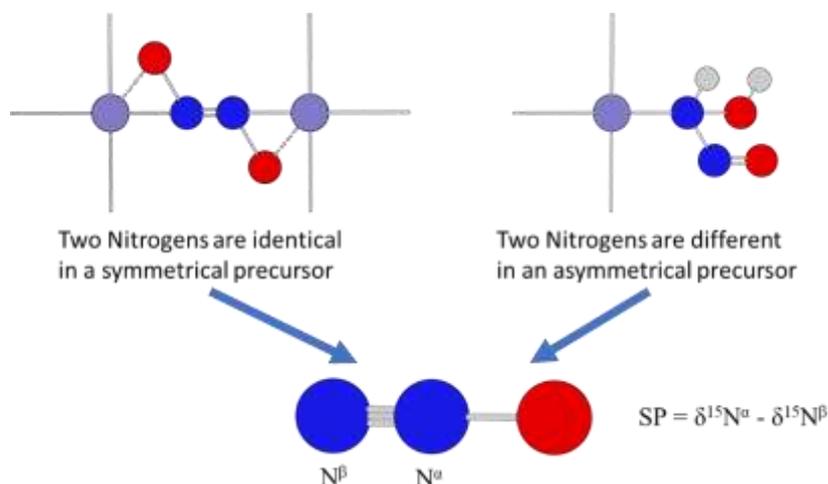


Figure 1. Sketch of alunite precipitation from water. The alunite could be a single-mineral geothermometer if three conditions are fully all satisfied: (1) $\text{H}_2\text{O}_A = \text{H}_2\text{O}_B$, (2) $T_1 = T_2 = T_3$, and (3) all the four reactions are fully reversible and attain equilibrium. White, red, yellow, pink, and purple spheres represent hydrogen, oxygen, sulfur, aluminum, and potassium atoms, respectively.



455

Figure 2. Proposed mechanisms for N₂O formation from symmetrical and asymmetrical precursors (Modified from Schmidt et al., 2004). Light gray, red, blue, and purple spheres represent hydrogen, oxygen, nitrogen, and iron atoms, respectively.

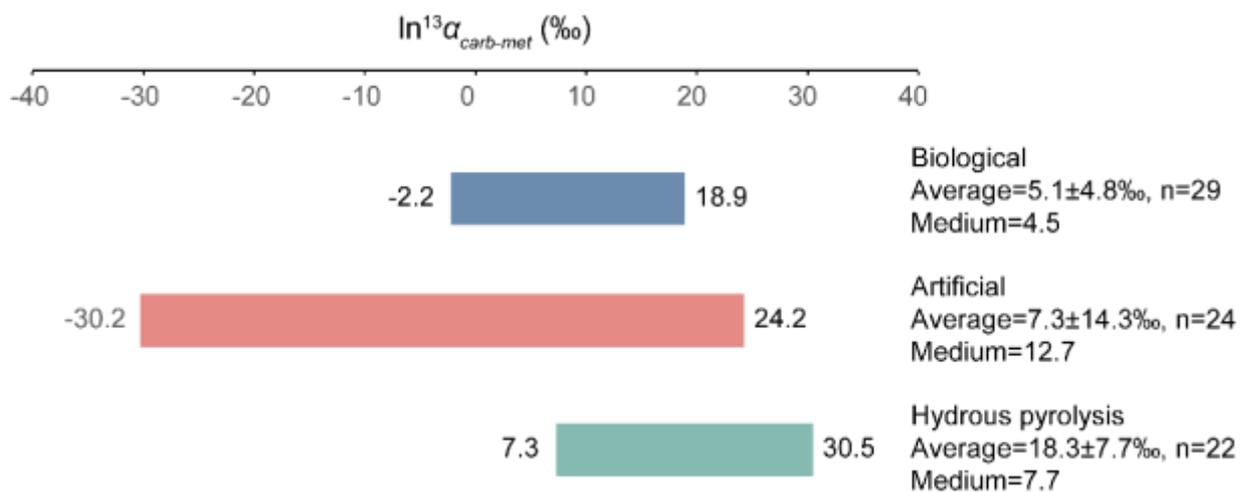
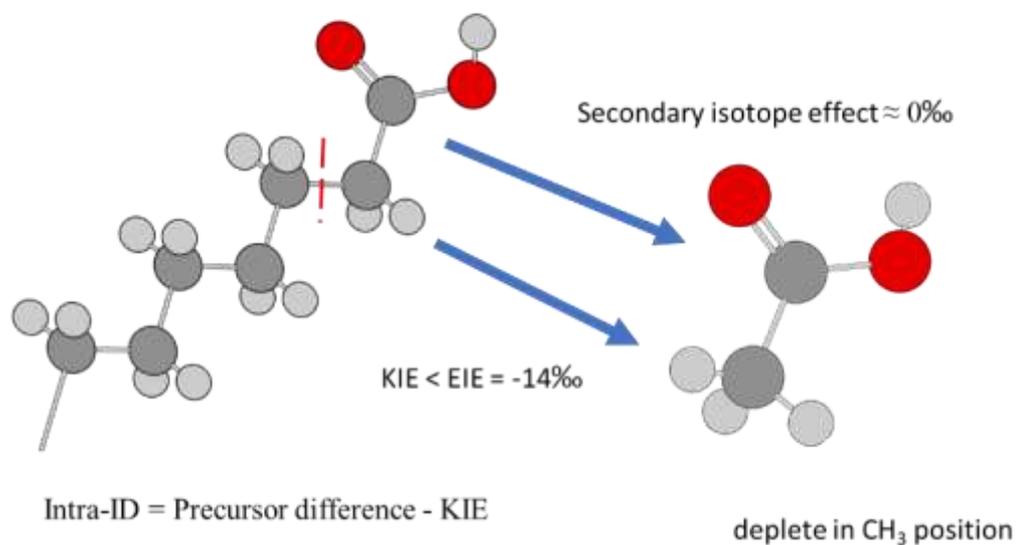


Figure 3. ln¹³α_{carb-met} values of biological, artificialman-made, and hydrous pyrolysis produced acetic acid.



460

Figure 4. Acetic acid produced from pyrolysis of precursor acid forms has an Intra-ID that is depleted in ¹³C in the methyl position. Dark gray, light grey, and red spheres represent carbon, hydrogen, and oxygen atoms, respectively.