



1 **Modelled potential forest area in the forest-steppe of central Mongolia is**  
2 **about three times of actual forest area**

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14

15 **Keywords:** Tree biomass, forest-steppe, forest fire, permafrost, Mongolia

16

17 **Abstract**

18 The Mongolian forest-steppe is highly sensitive to climate change and environmental impact. The  
19 intention of this study was to identify, which geocological parameters control forest distribution  
20 and tree growth in this semi-arid environment, and to evaluate the actual and potential tree  
21 biomass. For this purpose, we applied a combination of tree biomass and soil mapping, remote  
22 sensing and climate data analysis to a study area in the northern Khangai Mountains, central  
23 Mongolia.

24 Forests of different landscape units and site conditions generally showed minor differences in tree  
25 biomass. We found no significant correlation between tree biomass and NDVI (normalized  
26 differentiated vegetation index). Tree biomass was reduced at forest edges, in small fragmented



27 forest stands of the steppe-dominated area, and in large forest stands, compared to all other forest  
28 units. The tree biomass of forests on slopes ranged between 25 and 380 Mg ha<sup>-1</sup>. The mean tree  
29 biomass in forests of 10-500 ha was 199-220 Mg ha<sup>-1</sup>, whereby tree biomass at the forest edges was  
30 50-63 Mg ha<sup>-1</sup> less than in the interior parts of the forests. The mean tree biomass of forests >500 ha  
31 was 182 Mg ha<sup>-1</sup>, whereas that of forests <10 ha in the steppe-dominated area was only around 142  
32 Mg ha<sup>-1</sup>. Forests in alluvial plains had maximum tree biomasses of 440-688 Mg ha<sup>-1</sup>. In contrast to  
33 tree biomass, the spatial extension of forests showed distinct relationships with topographic and  
34 climatic parameters. Presence of forest was controlled by elevation (<2600 m a.s.l.), aspect (no  
35 southern slopes below 2100 m a.s.l.), slope (<25°), mean annual precipitation (160-340 mm) and  
36 mean growing season temperature (6.5-10.8 °C).

37 The actual forests of the study area covered 1,086 km<sup>2</sup>. In 1986, prior to extensive forest fires, it was  
38 1,898 km<sup>2</sup>. The actual tree biomass of 20 x10<sup>9</sup> g represented 57 % of that in 1986. Modelling of the  
39 potential forest area resulted in 3,552 km<sup>2</sup>, with 65 x10<sup>9</sup> g tree biomass (based on topographic  
40 parameters) and 3,113 km<sup>2</sup> with 58 x10<sup>9</sup> g tree biomass (based on climatic parameters), respectively.  
41 The modelled potential forest area was thus about three times the actual forest area.

42

### 43 1. Introduction

44 The Mongolian forest-steppe represents the transition zone between the southern limit of the boreal  
45 forest in Central Asia and the dry region of the Gobi Desert. This ecotone has semi-arid climate and is  
46 highly vulnerable to climate change and land-use intensification (Poulter et al., 2013; Yang et al.,  
47 2016; Khansaritoreh et al., 2017). Various ecological stress factors have recently reduced Mongolia's  
48 forest area and thus, most likely tree biomass as well (Dulamsuren et al., 2008; 2010a; 2010c).

49 Drought stress and resulting decline in wood production and forest regeneration were repeatedly  
50 reported, especially for Siberian larch (*Larix sibirica*, Ledeb.), which makes up approximately 80 % of  
51 the total forest area (Dulamsuren et al., 2010b; Dulamsuren et al., 2014; Liu et al., 2013; Dulamsuren  
52 et al., 2009; Dulamsuren et al., 2010a; Hauck et al., 2019). Mongolia's mean annual air temperature  
53 has increased by 0.27 K per decade (or absolutely 1.7 K) from 1940 to 2001 (Batima et al., 2005),



54 which is clearly above the global average of 0.12 K per decade from 1951 to 2012 (IPCC, 2013). In  
55 addition, hazardous forest fires destroyed large forest areas in Mongolia over the last decades  
56 (Goldammer, 2002; Nyamjav et al., 2007; Hansen et al., 2013). Furthermore, the lack of a systematic  
57 forest management, insufficient control of logging and forest pasture in the vicinity of grasslands  
58 contributed to the degradation and reduction of forests (Tsogtbaatar, 2004; Dulamsuren et al.,  
59 2014).

60 Boreal forests represent an important organic carbon pool and are thus important for the global  
61 climate (Pan et al., 2011; Goodale et al., 2002). Although most of the organic carbon in the boreal  
62 zone is stored in the soil (Deluca and Boisvenue, 2012; Mukhortova et al., 2015; Shvidenko and  
63 Schepaschenko, 2014), a considerable amount of carbon is also stored in the tree biomass. Typically,  
64 carbon stocks in the tree biomass of boreal forests range from 40 to 80 Mg C ha<sup>-1</sup> (Thurner et al.,  
65 2014; Jarvis et al., 2001; Luyssaert et al., 2007). Extensive forest use, fire-setting, and woodcutting  
66 has reduced the forest area and tree biomass in Central Asia since prehistoric times (Miehe et al.,  
67 2007; Unkelbach et al., 2019; Unkelbach et al., 2017). Its impact can be evaluated by estimating the  
68 potential extent of forest area based on climatic and topographic parameters (Klinge et al., 2015).

69 Investigations on tree biomass in the Mongolian forest-steppe have been carried out in the Altai  
70 Mountains, southern Khangai Mountains (Dulamsuren et al., 2016), northern Khangai Mountains  
71 (Dulamsuren et al., 2019), and Khentei Mountains (Danilin and Tsogt, 2014). Tree biomass reported  
72 from these studies ranged between 123 and 397 Mg ha<sup>-1</sup>. Average tree biomass tends to decrease  
73 from the more humid north to the arid south. At local scale, tree biomass in the interior of *Larix*  
74 *sibirica* forests exceeds that at the forest edges (Dulamsuren et al., 2016). No consistent significant  
75 differences in tree biomass were found between forest-stands of varying sizes and between forests  
76 growing in grassland- and forest-dominated areas of the forest-steppe ecotone (Dulamsuren et al.,  
77 2019).

78 Various methods have been established to quantify tree biomass by relating field measurements  
79 with remote sensing data. However, all these methods come with specific errors (Rodríguez-Veiga et  
80 al., 2017; 2019; Lu et al., 2015; Powell et al., 2010). The main limitation of vegetation indices from



81 multispectral satellite images, such as the NDVI (normalized differentiated vegetation index), is that  
82 the signal becomes saturated, when the upper leaves cover lower leaves of the tree canopy layer.  
83 The mean crown closure of Mongolian forests is 0.53 (Goldammer, 2007), which suggests that the  
84 NDVI should be applicable in the study area. Dulamsuren et al. (2016) interpolated the tree biomass  
85 and carbon stock of entire Mongolia based on Spot-NDVI and field data. Thurner et al. (2014) used a  
86 growing stock volume product from synthetic aperture radar data to estimate the tree density and  
87 carbon stock of boreal and temperate forests of the northern hemisphere.

88 Beside biomass estimation, the NDVI has often been used to analyse vegetation dynamics in  
89 response to climatic shifts (Buermann et al., 2014; Miao et al., 2015; Poulter et al., 2013; Vandandorj  
90 et al., 2015). However, delineating vegetation trends from vegetation indices must be done with  
91 caution, because these parameters strongly depend on the spatial resolution and the recording  
92 season of the satellite image. Moreover, indices like the NDVI include a mixture of greenness signals  
93 from different vegetation types. Karnieli et al. (2013) gave an example for the problems in  
94 interpreting vegetation degradation in a case study on Mongolian pastures.

95 The parameters precipitation, temperature and evaporation control the spatial pattern of forest and  
96 steppe distribution in the semi-arid forest-steppe (Klinge et al., 2018; Nyamjav et al., 2007;  
97 Dulamsuren and Hauck, 2008). In addition, topographic position plays an important role, as forests  
98 are strictly limited to north-facing slopes (Klinge et al., 2015; Hais et al., 2016). Thus, relief is an  
99 important factor for the existence, vitality and tree density of forests. In addition to natural factors,  
100 the actual forest distribution is strongly influenced by human impact, which exists since prehistoric  
101 times (Miehe et al., 2007), and by intense forest fires that occurred during the last decades.

102 Based on the state of knowledge described above, we addressed the following hypotheses:

- 103 (I) The actual tree biomass in the study area depends on soil conditions and topographic position  
104 and has been strongly reduced by forest fires and logging.
- 105 (II) The NDVI is a suitable parameter for tree biomass assessment in the study area.



106 (III) Tree vitality depends on ecological site conditions, whereas climate and topography are the  
107 main limiting factors for the distribution of larch forests in the study area (to be deduced by  
108 spatial analysis of climatic and topographic data).

109 (IV) The potential forest area and tree biomass are much larger than the actual ones.

110 We aimed at answering the following research questions:

- 111 • What are the general climatic and ecological limitations for forest distribution in the mountain  
112 forest-steppe (to be deduced by GIS and remote sensing analysis)?
- 113 • Is there an influence of soil conditions, forest use and fire events on tree biomass, depending on  
114 forest fragmentation and stand size?
- 115 • How large are the losses of forest area and tree biomass due to recent forest fires in the study  
116 area?
- 117 • How large are the potential forest area and tree biomass in the study area?

118

## 119 **2. Study area**

120 The study area is located on the northern slopes of the Khangai Mountains near the town  
121 Tosontsengel in northern central Mongolia (98°16'E/48°46'N) (Fig. 1). The region has continental  
122 climate with cold semi-arid conditions (Figure 2). The monthly mean temperatures at Tosontsengel  
123 range between -31.7 °C in January and 14.7 °C in July. Most of the annual precipitation occurs during  
124 summer, from low pressure cells brought by the westerlies (Batima et al., 2005). In contrast, the  
125 Siberian high during winter causes mostly dry conditions. The cold climate promotes discontinuous  
126 permafrost, with permafrost mainly occurring in the valley bottoms, on the upper mountains, and  
127 partially on the slopes. The existence of permafrost ice requires some soil moisture, whereas dry soil  
128 conditions lead to dry permafrost, i.e., frozen ground without ice.

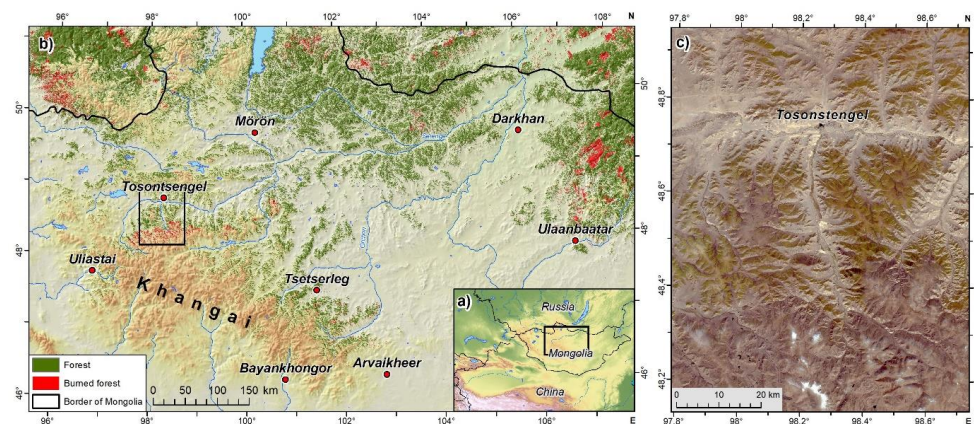
129 The maximum altitudes of up to 3,200 m a.s.l. lie in the southern part of the study area. They are  
130 characterised by mountain plateaus with cryoplanation terraces (Kowalkowski and Starkel, 1984;  
131 Richter et al., 1963). These highest regions belong to the periglacial belt, where alpine vegetation and  
132 bare, rock-debris covered land surfaces occur above the upper treeline at approx. 2,500 m a.s.l.



133 (Klinge et al., 2018). In the northern part, the mountains are lower, and mountain forest-steppe  
134 covers the north-facing slopes up to the summits. The main valleys run from south to north, leading  
135 into the east-west running valley of the Ider Gol (Gol = River) at an elevation of 1,600 m a.s.l. The  
136 geological basement consists of Permian metamorphosed sedimentary and acid plutonic rock, and  
137 Carboniferous mafic rock (Academy of Sciences of Mongolia and Academy of Sciences of USSR,  
138 1990). Coarse detritus of these bedrocks builds up the slope debris, which is often mixed with and  
139 covered by sandy to silty aeolian deposits.

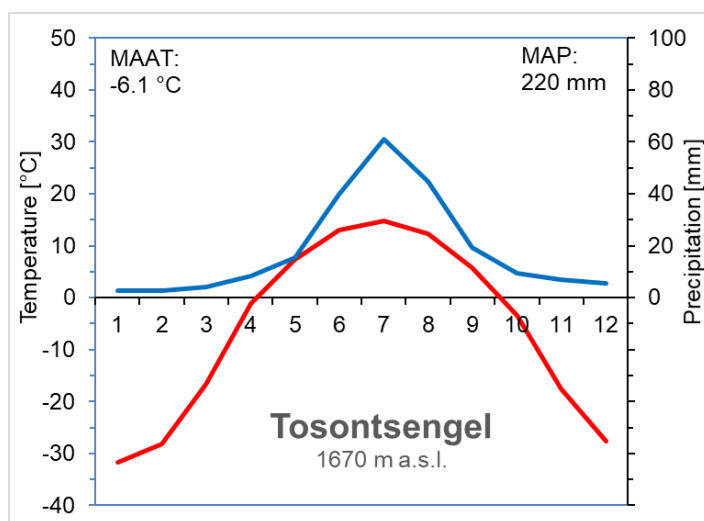
140 Dense, extensive forests occur south of the Ider Gol, whereas in the north, forests are more  
141 fragmented and steppe vegetation is dominant (Dulamsuren et al., 2019). A clear spatial pattern of  
142 forests (made up of *Larix sibirica*) on the northern slopes and steppe on the southern slopes is typical  
143 in the forest-steppe of Mongolia (Hilbig, 1995; Treter, 1996). This vegetation pattern is generally  
144 controlled by low precipitation (<300 mm), high evapotranspiration and relief-controlled differences  
145 in insolation in the mid-latitudes (Hais et al., 2016; Schlütz et al., 2008). Riverine forests consist of  
146 willow (*Salix*), poplar (*Populus*), and larch (*Larix sibirica*). Since these alluvial forests are supported by  
147 groundwater, they are more independent from the local precipitation. Pleistocene dune fields with  
148 scattered individual old larch trees are abundant in the basins. There are many local forest and  
149 steppe fires during the summer season (Goldammer, 2007; Hessel et al., 2012). Hazardous forest fires  
150 in 1996 and 2002 destroyed extensive forests. Many of these former forest areas have not yet  
151 regrown.

152 A timber factory and forest tracks were established in the Tosontsengel region during Soviet times to  
153 facilitate intensified forest exploitation since the 1960s. Industrial logging was abandoned after the  
154 political change in the early 1990s, but has been partly resumed. In addition, the local population  
155 extracts fuelwood from the forests.



156

157 *Figure 1: Study area. a) Overview of Mongolia with position of the map shown in b) (black frame). b)*  
158 *Location of the study area in the forest-steppe of northern central Mongolia. The forest distribution*  
159 *was adapted from Klinge et al. (2018), whereas the burnt forest area (2000-2018) was mapped by*  
160 *Hansen et al. (2013). The digital elevation model (DEM) was created from SRTM (Shuttle Radar*  
161 *Topography Mission) data. The black frame in b) indicates the position of the image shown in c). c)*  
162 *True colour satellite image of the study area near Tosontsengel (Landsat 8, September 22, 2014).*



163

164 *Figure 2: Climate diagram for Tosontsengel (measuring period 1968-2013).*

165

166 **3. Methods**



167 **3.1. Tree biomass measurements and site mapping**

168 During fieldwork in the years 2014-2018, we determined the tree biomass on 20 m x 20 m plots, by  
169 measuring the tree diameter at breast height (dbh) and tree height for all trees exceeding a height of  
170 4 m. We used either a Vertex IV ultrasonic clinometer and T3 transponder (Haglöf, Långsele, Sweden)  
171 or a True Pulse 200 laser rangefinder (Laser Technology, Inc., USA) for measuring tree height. Stem  
172 diameter was calculated from stem circumference as measured with a measuring tape. In this way,  
173 we analysed 140 plots, in order to cover forest- and steppe-dominated landscapes, forest edge and  
174 interior, toe slope, mid-slope, upper slope, natural and exploited forests, as well as different aspects  
175 and forest-stand sizes. We applied the two allometric functions for Siberian larch (*Larix sibirica*) in  
176 Mongolia published by Battulga et al. (2013) and Dulamsuren et al. (2016), and used the mean of the  
177 results of both equations to estimate the aboveground and belowground tree biomass. Differences  
178 in the estimates of the two functions are discussed in Dulamsuren et al. (2016). Forest stands where  
179 tree stumps indicated woodcutting, were mapped as forests with “logging”, whereas forest stands  
180 without tree stumps were mapped as forests with “no logging”. Forest stands, where burned bark  
181 and/or charred wood indicated former fire events, were mapped as having “fire indicators”, those  
182 without as having “no fire indicators”. Ground vegetation structure, soil profiles and detection of  
183 permafrost provided auxiliary data. In 24 tree biomass plots, we measured the Leaf Area Index (LAI)  
184 using a LI-COR Plant Canopy Analyzer LAI-2200 C (Licor Biosciences, USA).

185

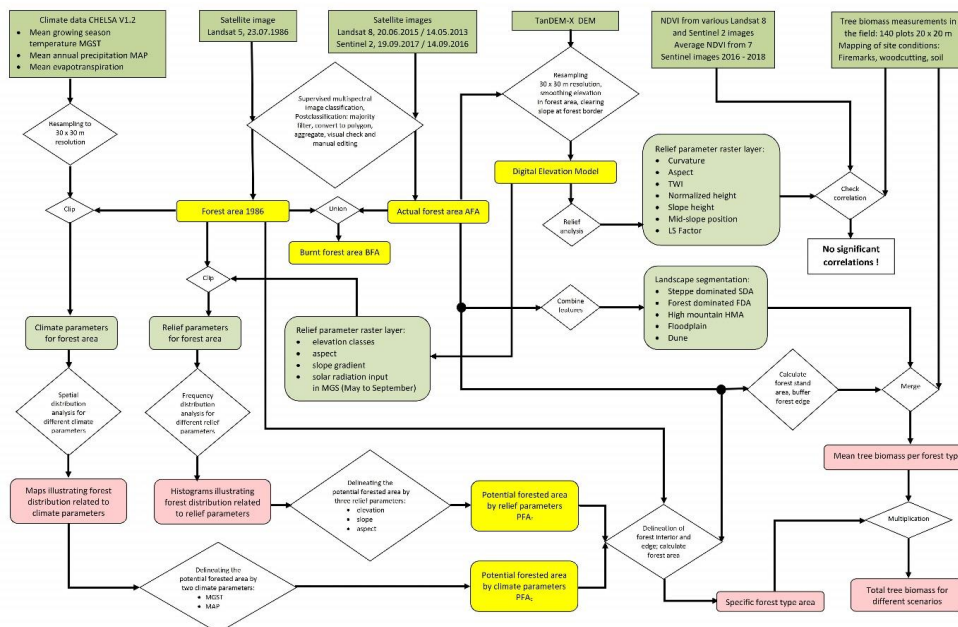
186 **3.2. Remote sensing analysis of forest distribution**

187 The workflow of the digital forest-distribution analysis, including input, intermediate and output  
188 data, is illustrated in Figure 3. We used a supervised classification of a Landsat 5 satellite image from  
189 September 23, 1986 to delineate the forest distribution before the extensive fires. We determined  
190 the actual forest area (AFA) by integrating supervised image classifications from several scenes of  
191 Landsat 8 (May 14, 2013; June 20, 2015) and Sentinel 2 (Sep 14, 2016; Sep 19, 2017) to receive the  
192 best possible forest representation. The difference in forest area between the images of 1986 and  
193 2017 represents the burnt forest area (BFA) for this period.





194 We used a digital elevation model (DEM) based on TanDEM-X data to compute various topographic  
 195 parameters (see Figure 3), in order to identify possible relationships between tree biomass and relief.  
 196 Due to the high horizontal (10 x 10 m) and vertical (<1 m) resolution of the DEM, the calculated  
 197 terrain surface was distorted by the forest canopy, especially at the edges of the forest stands. Thus,  
 198 we used the AFA map obtained from the satellite image classification for correcting the DEM in the  
 199 forest areas. In addition, we created NDVI maps from Landsat 5/8 and Sentinel 2 multispectral  
 200 images to determine the vitality of the vegetation and to test potential correlations between the  
 201 NDVI and tree biomass estimations. The classification of different landscape units (LU) was based on  
 202 (1) the proportion between forest and steppe areas, distinguishing forest-dominated area (FDA) and  
 203 steppe-dominated area (SDA), (2) the position of the upper treeline in the high mountain area  
 204 (HMA), (3) flat area (<2°) in the alluvial plains and (4) dunes, identified based on multispectral  
 205 satellite images.



206

207 *Figure 3: Workflow of the forest-distribution and tree biomass GIS analysis.*

208



209 We applied the approach of Klinge et al. (2015) to identify the potential forest area (PFA) based on  
210 relief parameters. For this purpose, we used the forest area in 1986 to identify the relationships  
211 between relief and forest distribution, considering elevation, aspect, slope, and insolation during the  
212 mean growing season (MGS, May-September). We chose the forest status of 1986 for the statistical  
213 analysis, because at that time, forest distribution was still less disturbed by forest fires than today.

214 In addition, we performed a forest distribution analysis based on the climatic parameters mean  
215 annual precipitation (MAP), mean growing season temperature (MGST; mean of the monthly mean  
216 temperatures between May and September), and mean annual potential evaporation. For this  
217 purpose, we used climate data from the CHELSA V1.2 dataset (Karger et al., 2018), which we  
218 resampled from the originally 30-arcsec resolution to obtain 30 m resolution. In this reanalysed  
219 climate dataset, we also considered terrain parameters and wind effect, in order to obtain an  
220 improved representation of climate conditions in relief terrain (Karger et al., 2017). We calculated  
221 MAP and MGST for the period 1973-2013. In addition, we applied a GIS-tool on TanDEM-X data to  
222 estimate the cumulative insolation for the period Mai-September 2017, which served as an example  
223 for the mean growing season.

224 Finally, we delineated three different forest distribution patterns, namely the actual forest area  
225 (AFA), the forest area of 1986, and the potential forest area (PFA). The forest areas were assigned to  
226 five different landscape units, namely steppe-dominated area (SDA), forest-dominated area (FDA),  
227 high-mountain area (HMA), alluvial forest, and dune. We distinguished four forest-size classes (FSC):  
228 F1, G1 =  $<0.1 \text{ km}^2$  in the forest-dominated and steppe-dominated areas, respectively; F2 =  $0.1\text{-}1 \text{ km}^2$ ;  
229 F3 =  $1\text{-}5 \text{ km}^2$ ; F4 =  $>5 \text{ km}^2$  and used a spatial buffer of 30 m to distinguish the forest edges from the  
230 interior. These classification schemes were adapted from Dulamsuren et al. (2016; 2019).

231

### 232 3.3. Classification of site conditions

233 Dulamsuren et al. (2019) already analysed the interior of larch forests on slopes in the same study  
234 area, thereby focusing on larch stands in the optimum stage of the forest development cycle (Jacob  
235 et al., 2013; Feldmann et al., 2018) and excluding recent heavy disturbances like severe fire or clear-



236 cut. Complementary to that study, we included larch stands with varying site conditions, history and  
237 structure, in order to assess tree biomass variability in response to disturbances and site conditions.  
238 In this way, we also tested the sensitivity of tree biomass to varying specific site conditions and thus  
239 the potential of remote sensing techniques for upscaling plot-based data to the landscape level. We  
240 therefore used the tree biomass data from the 30 *L. sibirica* plots on slopes of Dulamsuren et al.  
241 (2019) and added tree biomass data from forest edges and further *L. sibirica* plots of different stand  
242 sizes (classes F1/G1 to F4) and forest-to-grassland ratios (FDA with classes F1 to F4 vs. SDA with class  
243 G1). In addition to these forests on slopes, we also analysed larch forests on alluvial sands in  
244 floodplains. However, their limited size did not allow to obtain separate datasets for forest interior  
245 and edge. Altogether, we distinguished 12 larch-stand types, including forest interior and forest  
246 edges of the slope forest types F1, F2, F3, F4 and G1, and the floodplain forests as independent  
247 variables, and differentiating between logged / not logged and burned / not burned forest stands as  
248 covariates, based on the presence / absence of tree stumps and fire scars.

249 We calculated the mean tree biomass for each site-condition class such as wood harvest, fire  
250 indicators, and topsoil conditions (Table 1), checked the tree biomass data of each class for normal  
251 distribution and tested the differences between the tree biomass data for the various site conditions  
252 for statistical significance using Duncan's multiple range test of the SPSS software package. The  
253 relationships between tree biomass, NDVI and terrain parameters were statistically analysed by  
254 Pearson correlations to identify the best interpolation strategy for the forest distribution and tree  
255 biomass of 1986 and PFA. However, we found only weak correlations between tree biomass and  
256 NDVI, (see results chapter). Thus, finally, we delineated the various forest types with their specific  
257 tree biomass (with respect to the AFA, BFA and PFA) by multiplying the area of each forest type with  
258 the mean tree biomass value of that forest type. A delineation of potential alluvial forests was not  
259 feasible, because the erosion-deposition dynamics of the braided rivers control the  
260 geomorphological landscape development and thus the distribution of alluvial plains and related  
261 forests.  
262



263 **4. Results**

264 **4.1. General patterns of forest fire and permafrost distribution**

265 The actual forest area (AFA), subdivided into four size classes, and the burnt forest area (BFA) are  
266 presented in Figure 4. We found only few burnt areas in the steppe-dominated areas (SDA) in the  
267 north-western and eastern parts of the study area. In the forest-dominated (FDA) central and north-  
268 eastern parts of the study area, the most extensive burned forest areas (BFA) were in the upper  
269 mountains. The high mountain area (HMA) in the southern part had lost the largest portion of forest  
270 by fire. Its formerly large forest areas had turned into numerous small and fragmented forest  
271 remnants.

272 Our soil profiles showed that among the large forest stands on the slopes, permafrost was restricted  
273 to the FDA and HMA. Under large forest stands (forest size class F4) on north-facing slopes, the  
274 permafrost was rich in ice and occurred already at shallow depth, whereas east- and west-facing  
275 slopes had only small patches of permafrost that started at depths of more than 1 m. The frozen  
276 ground contained only little ice, where sandy materials made up the upper parts of the soils. There  
277 was no field evidence for permafrost under fragmented forest stands (G1, F1, F2) and burnt forests.

278

279 **4.2. Plot-based tree biomass data**

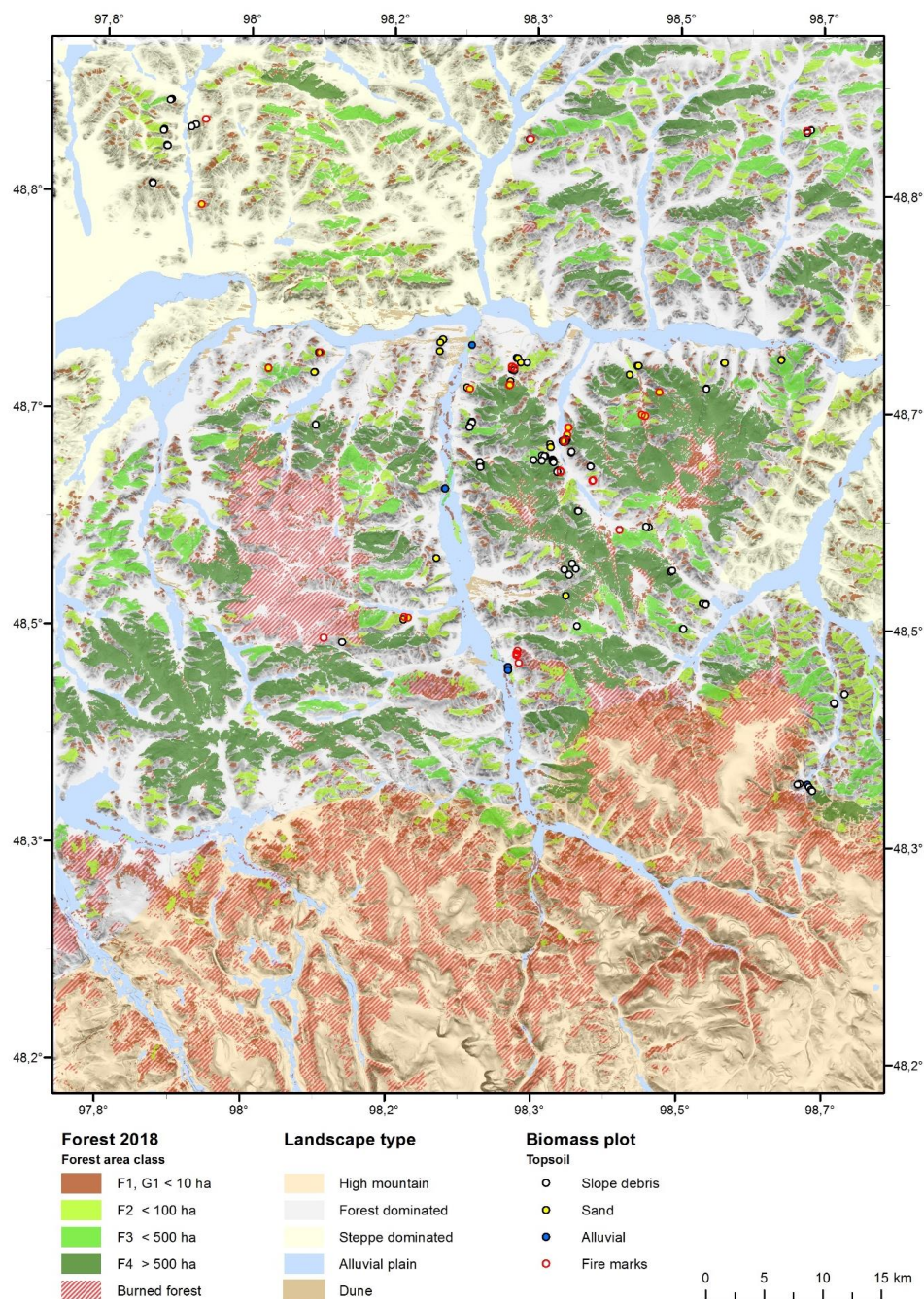
280 The mean tree height in the plots ranged between 12 and 20 m, whereas the maximum heights of  
281 single trees reached up to 32.7 m. The stand densities were between 5 and 91 m<sup>2</sup> ha<sup>-1</sup>, with an  
282 average stand density of 38.8 m<sup>2</sup> ha<sup>-1</sup>. The mean tree ages were determined between 100 and 200  
283 years, whereas maximum tree ages up to 380-413 years were found. Tree biomass in the larch  
284 forests on slopes ranged between 25 Mg ha<sup>-1</sup> and 380 Mg ha<sup>-1</sup>. Maximum tree biomasses of 440-688  
285 Mg ha<sup>-1</sup> were found in floodplain forests, whereas larch trees on sand dunes only formed open  
286 woodlands with less tree biomass (48 Mg ha<sup>-1</sup>, *n*=1). Mean tree biomasses for the various forest-  
287 stand sizes and ecological parameters (Fig. 3), are listed in Table 1. Variance analyses did not proof  
288 statistically significant differences between the forest classes (Figure 5). Forests of the size classes F1,  
289 F2, and F3 had 50-63 Mg ha<sup>-1</sup> less tree biomass at forest edges than in the interiors. The small



290 fragmented forests G1 in the SDA had up to 70 Mg ha<sup>-1</sup> less tree biomass than those in the FDA.  
291 Logging, forest fire and material making up the upper parts of the soils did not have any distinct  
292 influence on tree biomass. The large proportion of forest with wood harvest in the small fragmented  
293 forests (size classes F1 and G1) pointed to the strong exploitation pressure on these forest types. In  
294 all stand-size classes of the FDA forests, the tree-biomass means and medians lay within 180-220 Mg  
295 ha<sup>-1</sup>. The maximum tree biomass was more than 320 Mg ha<sup>-1</sup>. The forest-size class F4 showed no  
296 distinct difference in tree biomass between forest edges and interior.



297



298 Figure 4: Landscape types in the study area, with actual forest area (AFA) and burnt forest area (BFA)  
 299 (reference year 2018). The shaded relief illustration is based on TanDEM-X data.



300

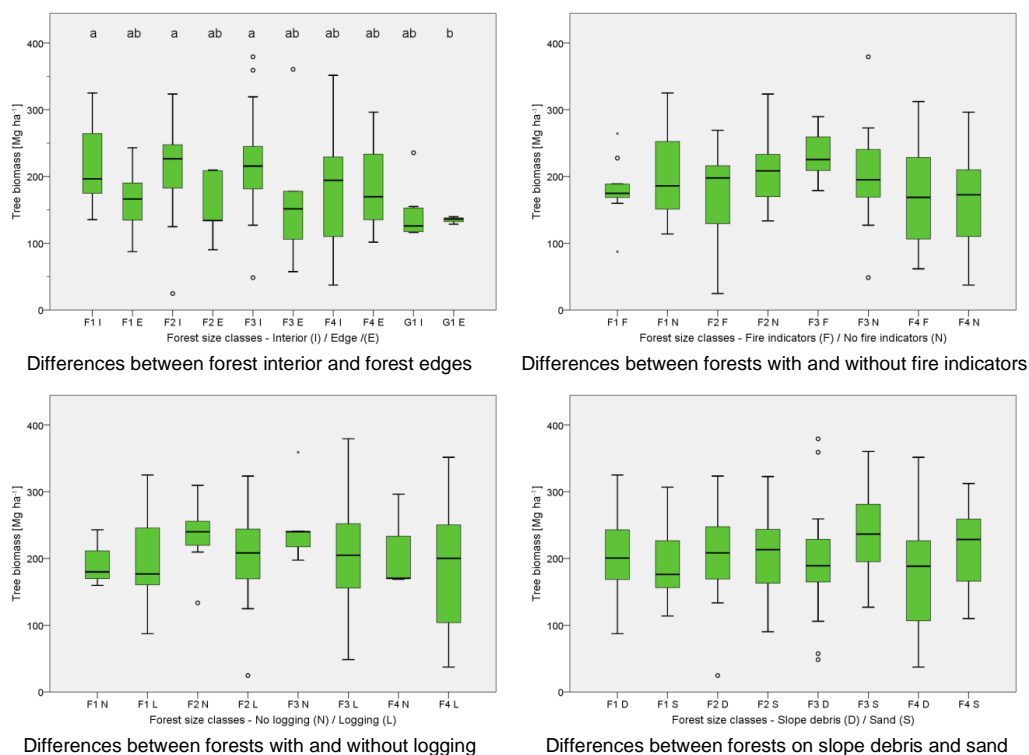
301 *Table 1: Mean tree biomass (above and belowground) for different forest types and site conditions.*

302 *Plots, where site conditions could not unambiguously be identified, were excluded from the*

303 *corresponding analysis part. SE = standard error, n = number of plots.*

	F1	SE	n	F2	SE	n	F3	SE	n	F4	SE	n	G1	SE	n
<b>total</b>	<b>198.7</b>	11.2	29	<b>208.0</b>	11.8	31	<b>212.5</b>	13.9	31	<b>182.0</b>	12.9	34	<b>142.2</b>	10.7	10
forest interior	219.3	14.0	18	218.2	12.6	26	220.6	13.4	26	181.7	13.8	30	145.4	15.0	7
forest edge	165.1	13.7	11	155.2	21.0	5	170.7	46.2	5	184.3	35.2	4	134.8	2.7	3
difference	54.2			63.0			49.9			-2.6			10.6		
no fire indicators;															
no logging	211.5	22.3	2	210.7	18.1	5	219.0	12.8	4	170.7	14.3	15			
no fire indicators	204.8	18.7	14	206.2	12.8	14	194.3	17.4	21	168.3	13.3	23	122.1	2.7	2
fire indicators	179.5	15.2	9	171.6	20.9	11	231.3	14.4	6	175.5	39.7	5	154.9		1
difference	25.2			34.6			-36.9			-7.2					
no logging															
logging	211.5	22.3	2	234.3	11.9	8	238.8	23.8	6	180.4	13.7	18	134.1	4.0	2
logging	197.8	11.9	27	198.9	14.4	23	207.4	16.6	24	183.8	22.7	16	144.3	13.2	8
difference	13.7			35.4			31.4			-3.4					
slope debris															
sand layer	211.0	18.4	13	207.8	16.5	18	196.3	18.5	19	176.7	14.3	28			
sand layer	188.8	13.3	16	204.5	15.8	14	238.3	18.3	12	214.4	24.5	7			
difference	22.2			3.4			-42.0			-37.7					

304



305 *Figure 5: Boxplots of tree biomass ( $Mg\ ha^{-1}$ ) related to different site conditions. Horizontal line =*

306 *median, bars = quartiles, whiskers = range, dots = outliers. A Duncan Posthoc test ( $p < 0.05$ ) showed*

307 *that most of the datasets were not statistically different. Significant differences existed between the*

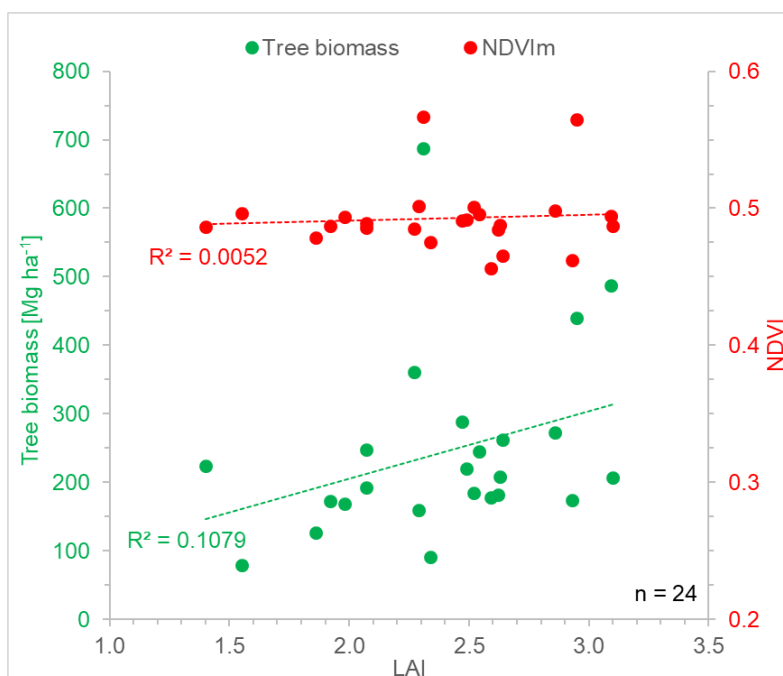
308 *forest interior and forest edges.*



309

### 310 4.3. Remote-sensing analyses based on tree biomass estimates

311 We tested several NDVI datasets, calculated from Sentinel 2 and Landsat 8 satellite images of the  
312 years 2013-2018, and various topographic parameters (see Figure 3) based on TanDEM-X data, for  
313 significant single or multi-correlations ( $r > 0.5$ ,  $p < 0.05$ ) with the measured tree biomass, since such  
314 correlations would allow for interpolation of tree biomass. However, we found no statistically  
315 significant correlations. The low number of cloud-free (<10 % cloud cover) multispectral satellite  
316 images of the summer seasons was one obstacle for NDVI analysis. Weak correlation between needle  
317 volume and tree biomass caused another problem, as the leaves and needles provide the vitality  
318 signal to be measured in multispectral satellite images. Danilin and Tsogt (2014) stated that needle  
319 biomass is independent of the average age of a larch stand, whereas tree biomass increases with  
320 average tree age. The weak correlation between leaf area index (LAI) and tree biomass measured on  
321 24 plots confirmed this statement (Figure 6). Thus, the statistical relationships between NDVI, needle  
322 volume and tree biomass were poor.



323

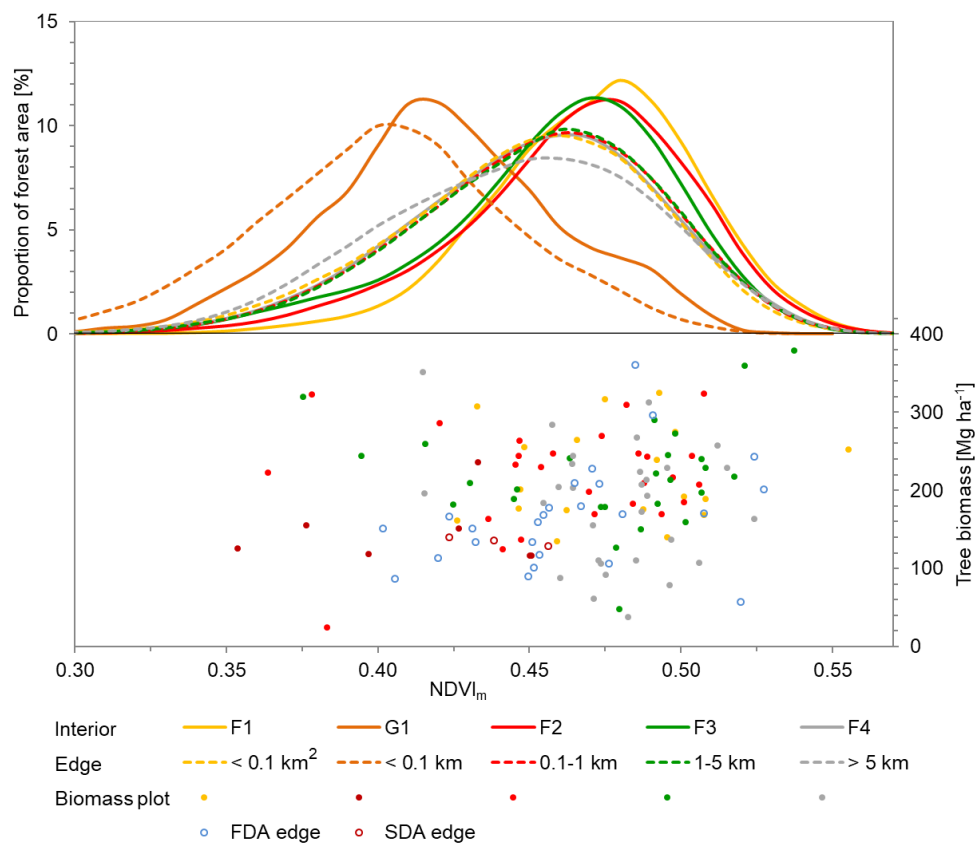




324 *Figure 6: Relationships between  $NDVI_m$  (red) and leaf area index (LAI), and between tree biomass on*  
325 *24 plots of the study area (green) and LAI.*

326

327 The phenology of the deciduous larch trees and the ground vegetation created another problem for  
328 the NDVI signal. Various topo-climatic conditions, e.g., cold air masses cumulating in topographic  
329 depressions during spring, cool conditions at higher elevations, droughts in late summer, produce  
330 temporally and spatially inhomogeneous patterns of tree vitality and thus NDVI. Therefore, we  
331 calculated mean NDVI values ( $NDVI_m$ ) of seven Sentinel 2 images of the growing seasons 2016-2018  
332 to obtain more site-representative NDVI values (Figure 7).  $NDVI_m$  was lower at the forest edges than  
333 in the forest interiors. It was also lower in small larch stands in steppe-dominated areas ( $G_1$ ) than in  
334 small larch stands in forest-dominated areas ( $F_1$ ). The high variation and the corresponding trend  
335 lines illustrate the low statistic correlation between tree biomass and  $NDVI_m$  for the measuring plots  
336 (Figure 7). Due to the low canopy closure of the larch forests of less than 53 %, the NDVI represents a  
337 signal for the vegetation vitality of both the trees and the understorey, which disabled an expedient  
338 tree biomass interpolation.



339

340 *Figure 7: Top - NDVI<sub>m</sub> frequency distribution curves for the interior and edges of larch stands of the*  
341 *five forest-size classes. Bottom - Tree biomass plotted against NDVI<sub>m</sub>.*

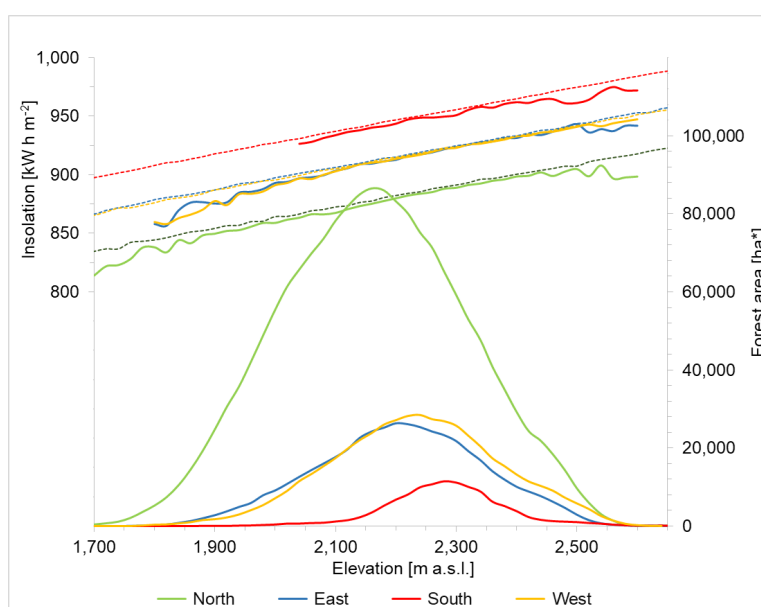
342

#### 343 4.4.1. PFA delineation based on relief parameters (PFA,)

344 In 1986, the upper treeline in the southern HMA reached up to 2,600 m a.s.l. (Figure 8), and forests  
345 occurred up to a maximum inclination of 25°. Since then, forest distribution showed a shift in both  
346 altitude and aspect (Figure 8). Because of forest fires in the upper mountains, the mean elevation  
347 range of the actual forests (95 %) has been restricted to 1,600-2,400 m a.s.l. These changes coincided  
348 with a transition from mountain forest-steppe to mountain subtaiga and the occurrence of *Pinus*  
349 *sibirica* in the uppermost forest areas. South-facing slopes in the forest-steppe are generally covered  
350 by steppe vegetation and partially forested above 2,100 m a.s.l. (Figure 8). MGS insolation increases



351 by approximately  $90 \text{ kW h m}^{-2}$  per 100 m elevation. The maximum MSG insolation occurring on the  
352 south-facing slopes exceeds that on the east- and west-facing slopes by ca.  $50 \text{ kW h m}^{-2}$ , and that on  
353 the north-facing slopes by  $100 \text{ kW h m}^{-2}$ . Forests also occur in areas of maximum MGS insolation  
354 (Figure 8), which demonstrates that insolation is no limiting factor for forest distribution in the study  
355 area.



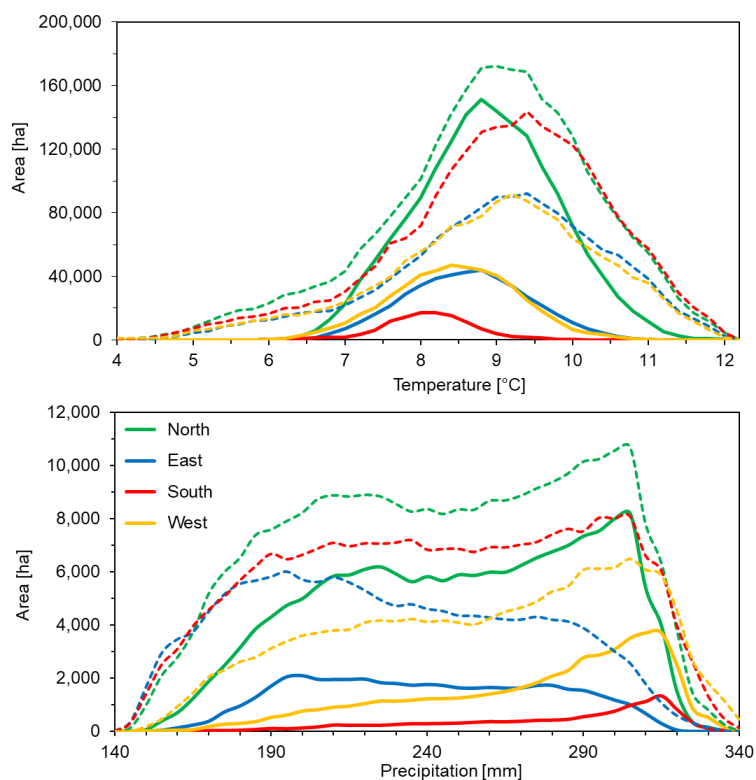
356  
357 *Figure 8: Top - Change of maximum MGS insolation (left axis) with elevation (MGS = mean growing*  
358 *season, May-September). Solid lines = MGS insolation on forest area, dashed lines = MGS insolation*  
359 *on the total land surface. Bottom - Forest area in hectares (right axis) in 1986 plotted against*  
360 *elevation. \* The forest area on south-exposed slopes is shown in ha x10.*

361

362 According to the observed relief-controlled hydrological limitations of forest distribution, we used  
363 the processed TanDEM-X data to identify the PFA, as the area, where all three parameters aspect,  
364 slope gradient and elevation allowed for tree growth (Figure 10). Thereby, we excluded forests in  
365 alluvial plains (areas with inclinations  $\leq 2.5^\circ$  along rivers and in basins), because their hydrological  
366 limitations are not controlled by the same parameters but rather by groundwater availability. The  
367 most striking outcome of this PFA, modelling was a much more extensive forest cover on toe slopes



368 and pediments, which are at present generally covered by steppe vegetation (Figure 10). In the HMA,  
369 the modelled  $PFA_r$  exceeded the presently forested area on steep slopes, from the valleys up to the  
370 upper treeline at 2,600 m a.s.l. The  $PFA_r$  did not cover 3.4 % of the forest area that existed in 1986.



371  
372 *Figure 9: Frequency distribution curves of mean growing season temperature (MGST) and mean*  
373 *annual precipitation (MAP) in the study area. Solid lines = forest area, dashed lines = total area.*  
374 *Climate data source: CHELSA V1.2 (Karger et al., 2017), period 1979–2013, spatially resampled to 30*  
375 *m by linear interpolation.*

376

#### 377 4.4.2. PFA delineation based on climatic parameters ( $PFA_c$ )

378 In a second PFA estimation approach, we analysed the forest distribution of 1986 for relationships  
379 with various climate parameters (Figure 9). Potential evaporation was so closely correlated with  
380 MGST that it appeared not necessary to include it as an additional parameter for delineating the  
381  $PFA_c$ . The spatial resampling of the climate data by linear interpolation produced some noise in the



382 data, because it could not consider small topographic variations. Therefore, we did not use them to  
383 deduce the threshold values for forest growth. Instead, we defined the threshold values based on the  
384 most accurate visual overlap of the PFA and the forest distribution of 1986. The resulting thresholds  
385 are listed in Table 2. The obtained PFA<sub>c</sub> showed larger forest areas on the upper south-facing slopes  
386 and small flat summits. It matched well with the upper treeline in the HMA (Figure 10). Compared to  
387 the PFA<sub>r</sub>, the PFA<sub>c</sub> did not extend as far down into the basins, which may be due to the low  
388 precipitation there. The PFA<sub>c</sub> did not cover 5.1 % of the forest area that existed in 1986.

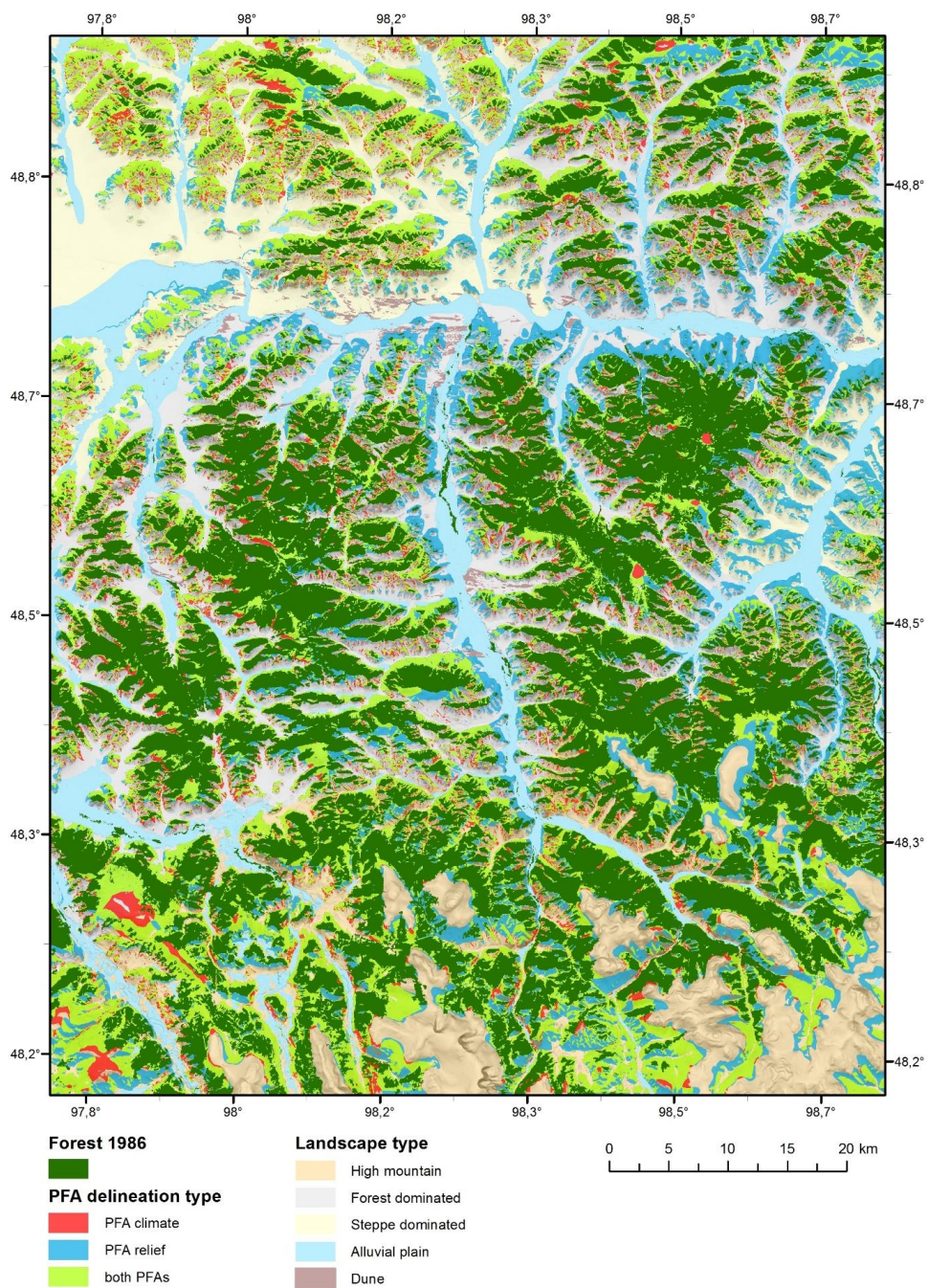
389

390 *Table 2: Climatic thresholds for PFA<sub>c</sub> delineation based on the mean growing season temperature*

391 *(MGST) and mean annual precipitation (MAP)*

	Aspect			
	North	East	South	West
Maximum MGST [°C]	10.8	10.4	8.6	10.0
Minimum MGST [°C]	6.5	6.5	6.5	6.5
Maximum MAP [mm]	320	310	340	340
Minimum MAP [mm]	160	170	290	165

392



393  
394 *Figure 10: Forest distribution in 1986 and potential forest areas (PFA) delineated based on climate*  
395 *and relief. The shaded relief illustration is based on TanDEM-X data.*



396 **4.5. Aggregated results of the forest area and tree biomass assessments and modelling**

397 The size of the study area was 6,355 km<sup>2</sup>, whereby two thirds (4,457 km<sup>2</sup>) were originally covered by  
398 steppe vegetation. At present, the total forest area (AFA) amounts to 1,086 km<sup>2</sup>, whereas in 1986 it  
399 was 1,898 km<sup>2</sup>. The modelled PFA yielded 3,168 (PFA<sub>c</sub>) and 3,553 km<sup>2</sup> (PFA<sub>i</sub>), respectively. Details on  
400 the differences between the AFA, forest area in 1986, and PFA, differentiated according to forest and  
401 landscapes types, are listed in Tables 3 and 5. The largest portion of the actual forest area occurs in  
402 the forest size class (FSC) F4 (8.6 %). Prior to the large fire events, this FSC was also widespread in the  
403 HMA (9.2 %). Altogether, a forest area of 812.5 km<sup>2</sup> (12.8 % of the total study area and 42.8 % of the  
404 former forest area) was destroyed by fire since the end of the last century. The BFA was negligible in  
405 the SDA and small in the FDA, but it amounted up to 95 % of the former forest area in the HMA. The  
406 ratios between forest interior and forest edge (F<sub>i</sub>/F<sub>e</sub>) for each FSC were relatively consistent for the  
407 different landscape units. However, the F<sub>i</sub>/F<sub>e</sub> ratios of the AFA tended to be lower than those of the  
408 forest areas in 1986, indicating an increase in forest fragmentation (Table 3).

409

410 *Table 3: Dimensions [km<sup>2</sup>] and relative portions [%] of different forest and landscape types in the*  
411 *study area at present (AFA) and in 1986, prior to the large forest fires (FSC = forest size class)*



	Actual forest area AFA				Forest distribution in 1986							
	FSC	1	2	3	4	Burned forest	1	2	3	4	Steppe	
absolute area of forest landscape type [km <sup>2</sup> ]	<b>Steppe dominated SDA</b>											
	Interior	4.74	18.25	21.51	7.29		3.88	5.59	19.65	22.90	7.92	1042.29
	Edge	11.07	12.33	7.65	1.69			10.28	16.19	6.92	1.37	
	Ratio I/E	0.43	1.48	2.81	4.32			0.54	1.21	3.31	5.80	
	<b>Forest dominated FDA</b>											
	Interior	12.66	82.93	144.04	440.35		115.25	13.90	87.52	120.78	606.55	1907.62
	Edge	43.55	52.56	44.21	107.34			32.10	42.10	32.12	105.36	
	Ratio I/E	0.29	1.58	3.26	4.10			0.43	2.08	3.76	5.76	
	<b>High mountain HMA</b>											
	Interior	1.73	7.51	6.40			693.33	3.74	29.70	65.66	491.08	1021.74
	Edge	13.06	6.73	2.32				10.70	15.92	20.71	93.56	
	Ratio I/E	0.13	1.12	2.76				0.35	1.87	3.17	5.25	
	<b>Alluvial forest</b>	3.98	3.42	1.09			0.05	2.74	2.74	2.74		485.74
	<b>Dune</b>							0.01				27.28
	Relative proportion of landscape type [%]	<b>Steppe dominated SDA</b>										
Interior		0.07	0.29	0.34	0.11		0.06	0.09	0.31	0.36	0.12	16.40
Edge		0.17	0.19	0.12	0.03			0.16	0.25	0.11	0.02	
Sum		0.25	0.48	0.46	0.14			0.25	0.56	0.47	0.15	
<b>Forest dominated FDA</b>												
Interior		0.20	1.30	2.27	6.93		1.81	0.22	1.38	1.90	9.54	30.02
Edge		0.69	0.83	0.70	1.69			0.51	0.66	0.51	1.66	
Sum		0.88	2.13	2.96	8.62			0.72	2.04	2.41	11.20	
<b>High mountain HMA</b>												
Interior		0.03	0.12	0.10			10.91	0.06	0.47	1.03	7.73	16.08
Edge		0.21	0.11	0.04				0.17	0.25	0.33	1.47	
Sum		0.23	0.22	0.14				0.23	0.72	1.36	9.20	
<b>Alluvial forest</b>		0.13					0.001	0.04	0.07	0.02		7.64
<b>Dune</b>												0.43

412

413

414 Due to the low correlation between NDVI and tree biomass (Figures 6 and 7), estimation of the total  
 415 tree biomass of the study area by use of regression functions was not feasible. Thus, we estimated  
 416 the total tree biomass by multiplying the specific mean tree biomass of each forest type (Table 1)  
 417 with the corresponding area of the respective forest type (Tables 3 and 5). The mean tree biomass  
 418 for the HMA was adapted from the FDA. As the modelled PFAs resulted in forest-dominated  
 419 landscapes, we used the FDA tree-biomass mean values for estimating the total tree biomass in the  
 420 study area corresponding to the PFA (Table 5). The actual tree biomass in the study area is 57 % of  
 421 the one in 1986, corresponding to 30 % and 34 % of the tree biomass estimated for the PFA<sub>r</sub> and  
 422 PFA<sub>c</sub>, respectively (Table 5). The greatest losses of tree biomass due to forest fires were detected in  
 423 the large forests (size class F4) of the FDA and HMA, whereas the tree biomass losses due to fire  
 424 were less severe in the SDA and alluvial forests.

425





426 *Table 4: Total tree biomass [ $10^6$  g] in the different forest types of the study area*

Actual forest area AFA						Forest distribution in 1986				
FSC	1	2	3	4	Sum	1	2	3	4	Sum
<b>Steppe dominated SDA</b>										
Interior	68,892	398,195	474,487	132,398	1,073,972	81,348	428,835	505,209	143,948	1,159,341
Edge	149,217	191,236	130,499	31,053	502,005	138,668	251,237	118,171	25,152	533,229
Sum	218,109	589,431	604,986	163,450	1,575,976	220,016	680,073	623,380	169,100	1,692,569
<b>Forest dominated FDA</b>										
Interior	277,732	1,809,560	3,177,481	8,000,534	13,265,306	304,883	1,909,714	2,664,408	11,020,218	15,899,224
Edge	718,971	815,498	754,423	1,977,866	4,266,758	530,041	653,229	548,092	1,941,327	3,672,689
Sum	996,703	2,625,058	3,931,904	9,978,400	17,532,064	834,925	2,562,943	3,212,500	12,961,545	19,571,913
<b>High mountain HMA</b>										
Interior	37,963	163,881	141,084	0	342,928	82,118	648,002	1,448,389	8,922,284	11,100,793
Edge	215,657	104,498	39,519	0	359,674	176,719	247,038	353,449	1,723,920	2,501,126
Sum	253,620	268,379	180,603	0	702,602	258,837	895,041	1,801,838	10,646,204	13,601,919
<b>Alluvial forest</b>										
	350,149					353,095				
<b>Total Sum</b>					20,160,791					35,219,497
<b>Comparison of tree biomasses</b>										
Sum 1986	1,313,778	4,138,056	5,637,718	23,776,849	34,866,401					
Sum 2018	1,468,432	3,482,868	4,717,493	10,141,850	19,810,642					
Loss by fire	-154,655	655,189	920,225	13,634,999	15,055,759					
%	-11.8	15.8	16.3	57.3	43.2					

427

428

429 *Table 5: Potential forest area (PFA) and tree biomass as influenced by climate and relief*

FSC	Potential forest area PFA climate				Steppe area	Potential forest area PFA relief				Steppe area
	1	2	3	4		1	2	3	4	
<b>Area [<math>\text{km}^2</math>]</b>										
Interior	11.49	69.21	77.30	2590.66	2692.99	11.17	69.52	91.04	2940.99	2308.48
Edge	35.06	38.45	23.15	322.99		38.46	42.14	28.87	330.62	
Ratio I/E	0.33	1.80	3.34	8.02		0.29	1.65	3.15	8.90	
<b>Tree biomass [<math>\times 10^6</math> g]</b>										
					Total sum					Total sum
Interior	251,930	1,510,158	1,705,159	47,068,688	50,535,934	245,059	1,516,857	2,008,287	53,433,781	57,203,985
Edge	578,839	596,605	395,092	5,951,310	7,521,845	635,056	653,769	492,614	6,091,966	7,873,404
Sum	830,768	2,106,762	2,100,250	53,019,998	58,057,779	880,115	2,170,625	2,500,901	59,525,748	65,077,389
<b>Comparison of tree biomasses [<math>\times 10^6</math> g]</b>										
PFA - 1986	-483,009	-2,031,294	-3,537,468	29,243,149	23,191,377	-433,662	-1,967,431	-3,136,817	35,748,898	30,210,988
%	-58.1	-96.4	-168.4	55.2	39.9	-49.3	-90.6	-125.4	60.1	46.4
PFA - 2018	-637,664	-1,376,105	-2,617,242	42,878,148	38,247,136	-588,317	-1,312,242	-2,216,592	49,383,898	45,266,747
%	-76.8	-65.3	-124.6	80.9	65.9	-66.8	-60.5	-88.6	83.0	69.6

430

431

## 432 5. Discussion

433 This study showed that the Mongolian forest steppe is characterised by highly variable tree biomass.

434 The differences in mean tree biomass between various forest types were up to  $85 \text{ Mg ha}^{-1}$ . In

435 addition to natural impacts on forests (e.g., fire, windbreak, insect calamities, and drought), logging

436 and forest pasture strongly affect forest distribution and tree biomass. Large alluvial forests exist

437 where river channels hamper wood pasture, logging, and forest fires. These alluvial forests usually

438 consist of old larch trees with large tree biomass. Larch trees on dunes are also very old, but they

439 occur isolated and have a low stand density.



440 The main natural factors that control the vitality and spatial distribution of forests in the Mongolian  
441 forest-steppe are limited precipitation and high evapotranspiration. The latter depends on insolation,  
442 which in turn varies with topography. These factors cause the lack of forests on south-facing slopes  
443 (Klinge et al., 2018; Hais et al., 2016; Dulamsuren and Hauck, 2008).

444 The forest canopy supports dense ground vegetation and an organic surface layer that insulates the  
445 soil and serves as moisture reservoir during summer (Dashtseren et al., 2014). In this way, forests  
446 facilitate discontinuous permafrost. In turn, in this semi-arid region, the presence of permafrost plays  
447 an important role for tree survival during summer droughts, as it provides meltwater throughout the  
448 summer (Sugimoto et al., 2002). Recognising these mutual relationships between tree density, forest  
449 canopy closure, ground vegetation, organic surface layer, and the occurrence of permafrost is crucial  
450 for understanding the patterns of forest distribution and tree biomass.

451 Our tree biomass measurements in the field showed that the least mean tree biomasses occurred in  
452 the forest size class G1 (142 Mg ha<sup>-1</sup>) of the SDA and in the class F4 (182 Mg ha<sup>-1</sup>) of the FDA. The  
453 fragmented forests in the SDA generally have lower tree densities and thus provide less shade. As a  
454 result, there is no permafrost under these forests, and the majority of the ground vegetation consists  
455 of grasses with a species composition similar to that of the surrounding steppes. Forests of the FDA  
456 have higher tree densities and support a dense ground vegetation cover consisting of herbs, shrubs,  
457 grasses and mosses, and an organic surface layer, which together enable the development of  
458 permafrost through its insulation effect (Dashtseren et al., 2014).

459 Forest edges represent natural zones of forest expansion and retreat (Sommer and Treter, 1999).  
460 Temporally varying climate conditions control the position and shifts of the forest edges. Thus, the  
461 forest edges may have a fringe of dead trees at their outer boundary, and their outer boundaries  
462 may also be dissected. In addition, logging and pasture is more intensive at the forest edges than in  
463 the interiors. Due to the lower tree density, the tree biomasses are generally lower at the forest  
464 edges than in the interiors. However, we found exceptions to this role in the forest size classes G1  
465 and F4, where the forest interiors and edges had similar tree biomasses. In the class G1, the tree  
466 biomasses of the small interior forest areas are similarly low as those of the forest edges. In the class



467 F4, the forest edges have large tree biomasses compared to all other forest edges. However, the tree  
468 biomass in the interior of these large forest stands is not much greater, because permafrost that  
469 already occurs at shallow depth hinders the trees from deep rooting. Hence, the trees can reach less  
470 nutrient stocks and are prone to windthrow. Thus, permafrost plays a twofold role in these large  
471 forests. On the one hand, it causes the low mean tree biomass in the forest interior. On the other  
472 hand, it enhances the vitality of trees at the forest edges through meltwater supply during summer.  
473 This effect, together with higher precipitation in the upper mountains may also explain the existence  
474 of forests on south-facing slopes in the higher mountains (Figure 8).

475 The forest area that burnt down between 1986 and 2017 (BFA) amounts to 12.8 % of the total study  
476 area. The loss of tree biomass since the last century adds up to roughly 15 million t, which represents  
477 more than 45 % of the former tree biomass. Although the most extensive forest fires in this area  
478 occurred already in 2002, forests have not yet re-established in many parts of the BFA. The most  
479 extensive BFAs are located in the large continuous forests of the HMA and in the upper parts of the  
480 mountains in the FDA. In contrast, only few BFAs occur in the small and fragmented forests of the  
481 SDA. We thus conclude that forest fragmentation in the SDA prevents forest fires from passing over  
482 into neighbouring forest stands and keeps fires more isolated. The decrease of large forests (size  
483 class F4) by fire led to an increase in small-sized fragmented forest stands (size class F1) of surviving  
484 trees, representing remnants of the former large forests. This change induced the loss of permafrost  
485 in these areas. The surviving larch trees in the remaining forest remnants increase their fructification  
486 (Danilin and Tsogt, 2014). Their important role as nuclei for forest regeneration is demonstrated by  
487 numerous seedlings and saplings growing in the direct surrounding of the forest remnants, in the  
488 shade of the old trees. Thus, a slow but steady re-immigration of larch trees into the burned area  
489 proceeds from these forest remnants. It takes up to 200 years until a forest regenerates to its  
490 previous state before the fire (Nyamjav et al., 2007).

491 Fires occur frequently in semi-arid environment (Hessl et al., 2012). Thus, *Larix sibirica* is to a certain  
492 degree fire-adapted. Its survival of a fire depends on the type of fire (crown, surface or ground fire),



493 fire intensity, season and soil moisture. The prevalent survival of forest stands in depressions, erosion  
494 channels, and on toe slopes demonstrates the importance of soil moisture for tree survival.

495 Forest stands that experienced non-lethal fire events or selective logging have similar tree biomasses  
496 as pristine forests. We conclude that moderate thinning of forests may improve the growing  
497 conditions for the remaining trees, possibly because of less competition by other trees and additional  
498 nutrient supply from ash, and because the melting permafrost leads to a temporary increase of soil  
499 moisture and allows for deeper rooting.

500 The delineation of the PFA<sub>r</sub> (based on relief) suggests a potential for forest expansion, both down  
501 towards the basins and up the mountains. The pediments generally provide suitable geoecological  
502 conditions for tree growth, as confirmed by several small forest stands. Nevertheless, steppe  
503 vegetation predominates on the pediments, mainly because of herbivore grazing (Hilbig, 1995). The  
504 lower treeline of the PFA<sub>c</sub> (based on climate), considering the restriction of tree growth by dry  
505 conditions in the basins, coincides with that of the actual forest. Existence of forests below the  
506 threshold of 160 mm MAP can be explained by site-specific additional water influx. The PFA<sub>c</sub>  
507 moreover suggests a potential for more forests on south-facing slopes. This mismatch indicates that  
508 MAP and MGST alone cannot explain forest distribution, which is apparently a result of more  
509 complex causal chains, involving relief, climate, geo- and bioecological factors and mutual  
510 interactions between them.

511 Cool conditions restrict the expansion of forests into upper valleys and onto the mountain plateaus  
512 of the HMA in the South. Short growing seasons and long-lasting snow cover prevent tree growth  
513 there. Another limiting factor is the extensive use of the alpine meadows as summer pastures by  
514 nomads for a long time. The upper treeline varies from 2,200 m a.s.l. in the north of the study area to  
515 2,550 m a.s.l. in the south. Thereby, the small treeless areas on the flat peaks of the northern  
516 mountains may result from the so-called “summit effect” (Körner, 2012), i.e., particularly harsh  
517 conditions near summits, rather than from a true upper treeline.

518 The modelled PFAs show large forest stands in the SDA in the north. These would turn that area into  
519 FDA. The number of small fragmented forest stands would strongly decrease. Thus, given the



520 permafrost-promoting effect of large forests, also permafrost distribution would considerably  
521 increase in that area. This points to the possibility that the forest area in the SDA might have been  
522 reduced by logging.

523 The modelled maximum tree biomass of the PFA (58-65 x10<sup>9</sup> g) was twice the tree biomass of 35 x10<sup>9</sup>  
524 g in 1986 and three times the actual tree biomass of 20 x10<sup>9</sup> g. However, several relevant factors  
525 could not be considered in the PFA modelling. For example, large forests (size class F4), as  
526 predominantly obtained from the PFA modelling, are more likely exposed to severe fires than  
527 fragmented forest stands. In addition, due to the long-lasting human influence, the natural  
528 proportion between steppe and forest in this region remains a major research challenge (Klinge and  
529 Sauer, 2019). Human impact already started with the extinction of large herbivores like elephantine,  
530 and the reduction of wild animal herds since the Mesolithic period. It continued with the breeding of  
531 domestic animals and the development of pasture economy since the Neolithic period.

532 Tchebakova et al. (2009) modelled potential vegetation changes across Siberia based on climate  
533 change scenarios projecting warmer and drier climate. The authors reported an increase of forest-  
534 steppe and grassland areas. More frequent and severe wildfires would occur stimulated by dry  
535 conditions and enhanced by more fire load due to increased tree mortality. Nyamjav et al. (2007)  
536 stated that 95% of the actual forest destruction was caused by forest fires, whereas 5% was referred  
537 to logging. The authors reported an increase of fire events in Mongolia during the past decades and  
538 Goldammer (2002) assumed that most of the fires were caused by human activities. Hessler et al.  
539 (2012) investigated fire history since the last 450 years based on tree ring analysis. Concordant with  
540 results from the Tuva region in southern Siberia (Ivanova et al., 2010), the authors did not find a  
541 distinct increase of fire frequency during the last decades but fire events became more severe due to  
542 drier conditions. The limited synchrony of fire events between different sites points to human caused  
543 ignitions (Hessler et al., 2012). Human impact on biomass reduction, due to fuel wood gathering and  
544 intensive grazing of livestock may reduce available fuel for fire, which makes them less severe and  
545 extensive (Hessler et al., 2012; Umbanhowar et al., 2009).

546



547 **6. Conclusions**

548 A combination of tree biomass and soil mapping, remote sensing and climate data analysis allowed  
549 us to identify factors affecting larch forest distribution and tree biomass in the northern Khangai  
550 Mountains, central Mongolia, and to model the potential larch forest distribution and tree biomass.  
551 Forest distribution is strongly influenced by the relief-controlled climatic key factors precipitation,  
552 evapotranspiration, temperature, soil moisture, and presence of permafrost. Forests of different  
553 types and landscape units have similar tree biomass. Only forest edges and small fragmented forest  
554 stands have less tree biomass than all other units. None-lethal fires have no serious impact on tree  
555 biomass. Selective logging involves removal of tree biomass, but it also stimulates tree growth by  
556 moderately reducing the tree density and thus competition. The NDVI is not suitable for estimating  
557 tree biomass in the forest-steppe.

558 Forest fires destroyed 43 % of the forest area and 45 % of the tree biomass in the study area over the  
559 period 1986-2017. They mostly affected large forest stands (size class F4) in the upper mountains of  
560 the FDA and HMA. Fragmented forests of the SDA are less prone to severe fires because of their  
561 spatial isolation. Forest fires increase the number of small forest stands (size class F1), because often  
562 small remnants of the formerly large forests survive at sites with increased soil moisture (e.g.,  
563 depressions, toe slopes). Permafrost, which is widespread under large forests, disappears soon after  
564 the destruction of a large forest stand (F4).

565 The factors aspect, slope gradient and elevation, and high-resolution precipitation and temperature  
566 data are suitable parameters for modelling the PFA. We obtained a  $PFA_r$  of 3,552 km<sup>2</sup> with 65 x10<sup>9</sup> g  
567 tree biomass, based on relief parameters, and a  $PFA_c$  of 3,113 km<sup>2</sup> with 58 x10<sup>9</sup> g tree biomass, based  
568 on climatic parameters, corresponding to more than 288 % and 323 % of the actual tree biomass,  
569 respectively. However, these approximations do not yet consider several relevant factors such as  
570 herbivore grazing and plant competition. In addition, human impact plays an important role, which is  
571 difficult to quantify.

572

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586



587 **Appendix**

588

589 *Table A1: Mean tree biomass (above and belowground) for different forest types and site conditions.*

590 *Plots, where site conditions were not clearly identified, were excluded from the respective part of the*

591 *analysis. The two lower and higher forest size classes (FSC) in the forest-dominated area (FDA) were*

592 *combined for the statistical analysis, because of the small dataset for forest edges. SE = standard*

593 *error, n = number of plots. Underlined data are not representative because of insufficient size of the*

594 *respective dataset.*

	F1	SE	n	F2	SE	n	F3	SE	n	F4	SE	n	G1	SE	n
<b>total</b>	<b>198.7</b>	11.2	29	<b>208.0</b>	11.8	31	<b>212.5</b>	13.9	31	<b>182.0</b>	12.9	34	<b>142.2</b>	10.7	10
<b>difference interior - edge</b>	<b>54.2</b>			<b>63.0</b>			<b>49.9</b>			<b>-2.6</b>			<b>10.6</b>		
<b>forest interior</b>	219.3	14.0	18	218.2	12.6	26	220.6	13.4	26	181.7	13.8	30	145.4	15.0	7
no fire indicators;															
no wood harvest			0	236.9	4.4	3	232.8	6.3	3	161.1	13.1	13			
no fire indicators	242.0	27.2	7	212.0	13.5	12	201.7	17.3	16	165.1	13.5	20	<u>122.1</u>	2.7	2
fire indicators	194.4	15.9	5	181.9	25.8	8	231.3	14.4	6	177.2	49.5	4	<u>154.9</u>		1
<b>difference</b>	<b>47.6</b>			<b>30.1</b>			<b>-29.6</b>			<b>-12.0</b>					
no wood harvest				255.2	10.9	6	251.0	25.2	5	174.1	14.4	15			
wood harvest	219.3	14.0	18	207.1	15.2	20	215.1	15.6	20	189.3	23.5	15	145.4	15.0	7
<b>difference</b>				<b>48.1</b>			<b>35.9</b>			<b>-15.2</b>					
slope debris	236.5	25.4	7	216.2	18.8	15	215.8	19.8	15	175.2	15.0	25			
sand layer	208.4	15.2	11	220.9	15.2	11	227.2	16.3	11	214.1	31.7	5			
<b>difference</b>	<b>28.1</b>			<b>-4.7</b>			<b>-11.4</b>			<b>-38.9</b>					
<b>forest edge</b>	165.1	13.7	11	155.2	21.0	5	170.7	46.2	5	184.3	35.2	4	134.8	2.7	3
no fire indicators;															
no wood harvest	211.5	22.3	2	171.5	26.9	2	177.8		1	233.4	44.6	2			
no fire indicators	167.5	16.0	7	171.5	26.9	2	170.7	46.2	5	189.4	46.6	3			
fire indicators	160.9	24.9	4	144.3	28.3	3				168.7		1			
<b>difference</b>	<b>6.6</b>			<b>27.2</b>						<b>20.7</b>					
no wood harvest	<u>211.5</u>	22.3	2	<u>171.5</u>	26.9	2	<u>177.8</u>		1	211.8	34.5	3	<u>134.1</u>	4.0	2
wood harvest	154.8	13.8	9	144.3	28.3	3	168.9	57.8	4	<u>101.6</u>		1	<u>136.3</u>		1
<b>difference</b>	<b>56.7</b>			<b>27.2</b>			<b>9.0</b>			<b>110.3</b>					
slope debris	181.2	20.9	6	<u>171.5</u>	26.9	2	123.2	22.9	4	189.4	46.6	3			
sand layer	145.8	11.7	5	144.3	28.3	3	<u>360.5</u>		1	<u>168.7</u>		1			
<b>difference</b>	<b>35.5</b>			<b>27.2</b>			<b>-237.3</b>			<b>20.7</b>					
<b>combined classes</b>		F1 / F2	SE	n	F2 / F3	SE	n	F3 / F4	SE	n					
<b>forest edge</b>	162.0	11.5		16	162.9	25.5	10	176.7	30.2	9					
no fire indicators;															
no wood harvest	191.5	20.1	4	173.6	18.0	3	214.9	33.3	3						
no fire indicators	168.4	13.8	9	170.9	33.9	7	177.7	33.9	8						
fire indicators	153.8	18.9	7	144.3	28.3	3	<u>168.7</u>		1						
<b>difference</b>	<b>14.6</b>			<b>26.6</b>			<b>9.0</b>								
no wood harvest	191.5	20.1	4	173.6	18.0	3	203.3	26.9	4						
wood harvest	152.2	12.6	12	158.3	35.5	7	155.4	33.3	3						
<b>difference</b>	<b>39.3</b>			<b>15.3</b>			<b>47.9</b>								
slope debris	169.9	14.9	10	139.3	20.0	6	151.6	26.9	7						
sand layer	144.2	11.5	9	198.3	51.4	4	<u>264.6</u>	67.8	2						
<b>difference</b>	<b>25.7</b>			<b>-59.0</b>			<b>-113.0</b>								

595





## 596 References

- 597 Academy of Sciences of Mongolia and Academy of Sciences of USSR: National Atlas of the Peoples  
598 Republic of Mongolia, Ulan Baatar, Moscow, 144 pp., 1990.
- 599 Batima, P., Natsagdorj, L., Gombluudev, P., and Erdenetsetseg, B.: Observed climate change in  
600 Mongolia, AIACC Working Papers, 12, 1–25, 2005.
- 601 Battulga, P., Tsogtbaatar, J., Dulamsuren, C., and Hauck, M.: Equations for estimating the above-  
602 ground biomass of *Larix sibirica* in the forest-steppe of Mongolia, *J. Forestry. Res.*, 24, 431–437,  
603 doi:10.1007/s11676-013-0375-4, 2013.
- 604 Buermann, W., Parida, B., Jung, M., MacDonald, G. M., Tucker, C. J., and Reichstein, M.: Recent shift  
605 in Eurasian boreal forest greening response may be associated with warmer and drier summers,  
606 *Geophys. Res. Lett.*, 41, 1995–2002, doi:10.1002/2014GL059450, 2014.
- 607 Danilin, I. M. and Tsogt, Z.: Dynamics of structure and biological productivity of post-fire larch forests  
608 in the Northern Mongolia, *Contemp. Probl. Ecol.*, 7, 158–169, doi:10.1134/S1995425514020036,  
609 2014.
- 610 Dashtseren, A., Ishikawa, M., Iijima, Y., and Jambaljav, Y.: Temperature Regimes of the Active Layer  
611 and Seasonally Frozen Ground under a Forest-Steppe Mosaic, Mongolia, *Permafrost and Periglac.*  
612 *Process.*, 25, 295–306, doi:10.1002/ppp.1824, 2014.
- 613 Deluca, T. H. and Boisvenue, C.: Boreal forest soil carbon: distribution, function and modelling,  
614 *Forestry*, 85, 161–184, doi:10.1093/forestry/cps003, 2012.
- 615 Dulamsuren, C. and Hauck, M.: Spatial and seasonal variation of climate on steppe slopes of the  
616 northern Mongolian mountain taiga, *Grassl. Sci.*, 54, 217–230, doi:10.1111/j.1744-  
617 697X.2008.00128.x, 2008.
- 618 Dulamsuren, C., Hauck, M., Bader, M., Osokhjargal, D., Oyungerel, S., Nyambayar, S., Runge, M., and  
619 Leuschner, C.: Water relations and photosynthetic performance in *Larix sibirica* growing in the  
620 forest-steppe ecotone of northern Mongolia, *Tree physiology*, 29, 99–110,  
621 doi:10.1093/treephys/tpn008, 2009.
- 622 Dulamsuren, C., Hauck, M., Khishigjargal, M., Leuschner, H. H., and Leuschner, C.: Diverging climate  
623 trends in Mongolian taiga forests influence growth and regeneration of *Larix sibirica*, *Oecologia*,  
624 163, 1091–1102, doi:10.1007/s00442-010-1689-y, 2010a.
- 625 Dulamsuren, C., Hauck, M., and Leuschner, C.: Recent drought stress leads to growth reductions in  
626 *Larix sibirica* in the western Khentey, Mongolia, *Glob. Change Biol.*, 16, 3024–3035,  
627 doi:10.1111/j.1365-2486.2009.02147.x, 2010b.
- 628 Dulamsuren, C., Hauck, M., Leuschner, H. H., and Leuschner, C.: Gypsy moth-induced growth decline  
629 of *Larix sibirica* in a forest-steppe ecotone, *Dendrochronologia*, 28, 207–213,  
630 doi:10.1016/j.dendro.2009.05.007, 2010c.
- 631 Dulamsuren, C., Hauck, M., and Mühlenberg, M.: Insect and small mammal herbivores limit tree  
632 establishment in northern Mongolian steppe, *Plant. Ecol.*, 195, 143–156, doi:10.1007/s11258-  
633 007-9311-z, 2008.
- 634 Dulamsuren, C., Khishigjargal, M., Leuschner, C., and Hauck, M.: Response of tree-ring width to  
635 climate warming and selective logging in larch forests of the Mongolian Altai, *J. Plant. Ecol.*, 7,  
636 24–38, doi:10.1093/jpe/rtt019, 2014.
- 637 Dulamsuren, C., Klinge, M., Bat-Enerel, B., Ariunbaatar, T., and Tuya, D.: Effects of forest  
638 fragmentation on organic carbon pool densities in the Mongolian forest-steppe, *Forest Ecol.*  
639 *Manag.*, 433, 780–788, doi:10.1016/j.foreco.2018.10.054, 2019.



- 640 Dulamsuren, C., Klinge, M., Degener, J., Khishigjargal, M., Chenlemuge, T., Bat-Enerel, B., Yeruult, Y.,  
641 Saindovdon, D., Ganbaatar, K., Tsogtbaatar, J., Leuschner, C., and Hauck, M.: Carbon pool  
642 densities and a first estimate of the total carbon pool in the Mongolian forest-steppe, *Glob.*  
643 *Change Biol.*, 22, 830–844, doi:10.1111/gcb.13127, 2016.
- 644 Feldmann, E., Glatthorn, J., Hauck, M., and Leuschner, C.: A novel empirical approach for determining  
645 the extension of forest development stages in temperate old-growth forests, *Eur. J. Forest Res.*,  
646 137, 321–335, doi:10.1007/s10342-018-1105-4, 2018.
- 647 Goldammer, G.J.: Fire Situation in Mongolia, in: *International Forest Fire News*, 26, 75–83, 2002.
- 648 Goldammer, G.J. (Ed.): *International Forest Fire News*, 36, 97 pp., 2007.
- 649 Goodale, C. L., Apps, M. J., Birdsey, R. A., Field, C. B., Heath, L. S., Houghton, R. A., Jenkins, J. C.,  
650 Kohlmaier, G. H., Kurz, W., Liu, S., Nabuurs, G.-J., Nilsson, S., and Shvidenko, A. Z.: Forest carbon  
651 sinks in the northern hemisphere, *Ecol. Appl.*, 12, 891–899, doi:10.1890/1051-  
652 0761(2002)012[0891:FCSITN]2.0.CO;2, 2002.
- 653 Hais, M., Chytrý, M., and Horsák, M.: Exposure-related forest-steppe: A diverse landscape type  
654 determined by topography and climate, *J. Arid Environ.*, 135, 75–84,  
655 doi:10.1016/j.jaridenv.2016.08.011, 2016.
- 656 Hansen, M. C., Potapov, P. V., Moore, R., Hancher, M., Turubanova, S. A., Tyukavina, A., Thau, D.,  
657 Stehman, S. V., Goetz, S. J., Loveland, T. R., Kommareddy, A., Egorov, A., Chini, L., Justice, C. O.,  
658 and Townshend, J. R. G.: High-resolution global maps of 21st-century forest cover change,  
659 *Science (New York, N.Y.)*, 342, 850–853, doi:10.1126/science.1244693, 2013.
- 660 Hauck, M., Leuschner, C., and Homeier, J.: *Klimawandel und Vegetation – Eine globale Übersicht*, in  
661 press, Spektrum, Springer, Berlin, Heidelberg, 2019.
- 662 Hessler, A. E., Ariya, U., Brown, P., Byambasuren, O., Green, T., Jacoby, G., Sutherland, E. K., Nachin, B.,  
663 Maxwell, R. S., Pederson, N., Grandpré, L. de, Saladyga, T., and Tardif, J. C.: Reconstructing fire  
664 history in central Mongolia from tree-rings, *Int. J. Wildland Fire*, 21, 86, doi:10.1071/WF10108,  
665 2012.
- 666 Hilbig, W.: *The vegetation of Mongolia*, SPB Acad. Publ., Amsterdam, 258 pp., 1995.
- 667 IPCC: *Climate Change 2013 - The Physical Science Basis*, Cambridge University Press, Cambridge,  
668 2013.
- 669 Ivanova, G. A., Ivanov, V. A., Kukavskaya, E. A., and Soja, A. J.: The frequency of forest fires in Scots  
670 pine stands of Tuva, Russia, *Environ. Res. Lett.*, 5, 15002, doi:10.1088/1748-9326/5/1/015002,  
671 2010.
- 672 Jacob, M., Bade, C., Calvete, H., Dittrich, S., Leuschner, C., and Hauck, M.: Significance of Over-  
673 Mature and Decaying Trees for Carbon Stocks in a Central European Natural Spruce Forest,  
674 *Ecosystems*, 16, 336–346, doi:10.1007/s10021-012-9617-0, 2013.
- 675 Jarvis, P., Saugier, B., and Schulze, E.-D.: Productivity of boreal forests, in: *Terrestrial global*  
676 *productivity*, Roy, J., Saugier, B., Mooney, H. (Eds.), Academic Press, San Diego, 211–244, 2001.
- 677 Karger, D. N., Conrad, O., Böhrner, J., Kawohl, T., Kreft, H., Soria-Auza, R. W., Zimmermann, N. E.,  
678 Linder, H. P., and Kessler, M.: Climatologies at high resolution for the earth's land surface areas,  
679 *Scientific data*, 4, 170122, doi:10.1038/sdata.2017.122, 2017.
- 680 Karger, D. N., Conrad, O., Böhrner, J., Kawohl, T., Kreft, H., Soria-Auza, R. W., Zimmermann, N. E.,  
681 Linder, H. P., and Kessler, M.: Data from: Climatologies at high resolution for the earth's land  
682 surface areas, Dryad, Dataset, <https://doi.org/10.5061/dryad.kd1d4>, 2018.
- 683 Karnieli, A., Bayarjargal, Y., Bayasgalan, M., Mandakh, B., Dugarjav, C., Burgheimer, J., Khudulmur, S.,  
684 Bazha, S. N., and Gunin, P. D.: Do vegetation indices provide a reliable indication of vegetation



- 685 degradation?: A case study in the Mongolian pastures, *Int. J. Remote Sens.*, 34, 6243–6262,  
686 doi:10.1080/01431161.2013.793865, 2013.
- 687 Khansaritoreh, E., Dulamsuren, C., Klinge, M., Ariunbaatar, T., Bat-Enerel, B., Batsaikhan, G.,  
688 Ganbaatar, K., Saindovdon, D., Yeruult, Y., Tsogtbaatar, J., Tuya, D., Leuschner, C., and Hauck, M.:  
689 Higher climate warming sensitivity of Siberian larch in small than large forest islands in the  
690 fragmented Mongolian forest steppe, *Glob. Change Biol.*, doi:10.1111/gcb.13750, 2017.
- 691 Klinge, M., Böhner, J., and Erasmi, S.: Modeling forest lines and forest distribution patterns with  
692 remote-sensing data in a mountainous region of semiarid central Asia, *Biogeosciences*, 12, 2893–  
693 2905, doi:10.5194/bg-12-2893-2015, 2015.
- 694 Klinge, M., Dulamsuren, C., Erasmi, S., Karger, D. N., and Hauck, M.: Climate effects on vegetation  
695 vitality at the treeline of boreal forests of Mongolia, *Biogeosciences*, 15, 1319–1333,  
696 doi:10.5194/bg-15-1319-2018, 2018.
- 697 Klinge, M. and Sauer, D.: Spatial pattern of Late Glacial and Holocene climatic and environmental  
698 development in Western Mongolia - A critical review and synthesis, *Quaternary Sci. Rev.*, 210,  
699 26–50, doi:10.1016/j.quascirev.2019.02.020, 2019.
- 700 Körner, C.: *Alpine treelines: Functional ecology of the global high elevation tree limits*, Springer, Basel  
701 [u.a.], 220 pp., 2012.
- 702 Kowalkowski, K. and Starkel, L.: Altitudinal belts of geomorphic processes in the southern Changai  
703 Mts. (Mongolia), *Studia Geomorphologica Carpatho-Balcanica*, 18, 95–115, 1984.
- 704 Liu, H., Park Williams, A., Allen, C. D., Guo, D., Wu, X., Anenkhonov, O. A., Liang, E., Sandanov, D. V.,  
705 Yin, Y., Qi, Z., and Badmaeva, N. K.: Rapid warming accelerates tree growth decline in semi-arid  
706 forests of Inner Asia, *Glob. Change Biol.*, 19, 2500–2510, doi:10.1111/gcb.12217, 2013.
- 707 Lu, D., Chen, Q., Wang, G., Liu, L., Li, G., and Moran, E.: A survey of remote sensing-based  
708 aboveground biomass estimation methods in forest ecosystems, *Int. J. Digit. Earth*, 9, 63–105,  
709 doi:10.1080/17538947.2014.990526, 2015.
- 710 Luysaert, S., Inglima, I., Jung, M., Richardson, A. D., Reichstein, M., Papale, D., Piao, S., Schulze, E.-D.,  
711 WINGATE, L., and others: CO<sub>2</sub> balance of boreal, temperate, and tropical forests derived from a  
712 global database, *Glob. Change Biol.*, 13, 2509–2537, doi:10.1111/j.1365-2486.2007.01439.x,  
713 2007.
- 714 Miao, L., Liu, Q., Fraser, R., He, B., and Cui, X.: Shifts in vegetation growth in response to multiple  
715 factors on the Mongolian Plateau from 1982 to 2011, *Physics and Chemistry of the Earth, Parts*  
716 *A/B/C*, 87–88, 50–59, doi:10.1016/j.pce.2015.07.010, 2015.
- 717 Miehe, G., Schlütz, F., Miehe, S., Opgenoorth, L., Cermak, J., Samiya, R., Jäger, E. J., and Wesche, K.:  
718 Mountain forest islands and Holocene environmental changes in Central Asia: A case study from  
719 the southern Gobi Altay, Mongolia, *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 250, 150–166,  
720 doi:10.1016/j.palaeo.2007.03.022, 2007.
- 721 Mukhortova, L., Schepaschenko, D., Shvidenko, A., McCallum, I., and Kraxner, F.: Soil contribution to  
722 carbon budget of Russian forests, *Agr. Forest Meteorol.*, 200, 97–108,  
723 doi:10.1016/j.agrformet.2014.09.017, 2015.
- 724 Nyamjav, B., Goldammer, G.J., and Uibrig, H.: Fire situation in Mongolia, in: *International Forest Fire*  
725 *News*, 46–66, 2007.
- 726 Pan, Y., Birdsey, R. A., Fang, J., Houghton, R., Kauppi, P. E., Kurz, W. A., Phillips, O. L., Shvidenko, A.,  
727 Lewis, S. L., Canadell, J. G., Ciais, P., Jackson, R. B., Pacala, S. W., McGuire, A. D., Piao, S.,  
728 Rautiainen, A., Sitch, S., and Hayes, D.: A large and persistent carbon sink in the world's forests,  
729 *Science (New York, N.Y.)*, 333, 988–993, doi:10.1126/science.1201609, 2011.



- 730 Poulter, B., Pederson, N., Liu, H., Zhu, Z., D'Arrigo, R., Ciais, P., Davi, N., Frank, D., Leland, C., Myneni,  
731 R., Piao, S., and Wang, T.: Recent trends in Inner Asian forest dynamics to temperature and  
732 precipitation indicate high sensitivity to climate change, *Agr. Forest Meteorol.*, 178–179, 31–45,  
733 doi:10.1016/j.agrformet.2012.12.006, 2013.
- 734 Powell, S. L., Cohen, W. B., Healey, S. P., Kennedy, R. E., Moisen, G. G., Pierce, K. B., and Ohmann, J.  
735 L.: Quantification of live aboveground forest biomass dynamics with Landsat time-series and field  
736 inventory data: A comparison of empirical modeling approaches, *Remote Sens. Environ.*, 114,  
737 1053–1068, doi:10.1016/j.rse.2009.12.018, 2010.
- 738 Richter, H., Haase, G., and Barthel, H.: Die Goleztterrassen, *Petermanns Geographische Mitteilungen*,  
739 107, 183–192, 1963.
- 740 Rodríguez-Veiga, P., Quegan, S., Carreiras, J., Persson, H. J., Fransson, J. E.S., Hoscilo, A., Ziółkowski,  
741 D., Stereńczak, K., Lohberger, S., Stängel, M., Berninger, A., Siegert, F., Avitabile, V., Herold, M.,  
742 Mermoz, S., Bouvet, A., Le Toan, T., Carvalhais, N., Santoro, M., Cartus, O., Rauste, Y., Mathieu,  
743 R., Asner, G. P., Thiel, C., Pathe, C., Schmullius, C., Seifert, F. M., Tansey, K., and Balzter, H.: Forest  
744 biomass retrieval approaches from earth observation in different biomes, *Int. J. Appl. Earth Obs.*,  
745 77, 53–68, doi:10.1016/j.jag.2018.12.008, 2019.
- 746 Rodríguez-Veiga, P., Wheeler, J., Louis, V., Tansey, K., and Balzter, H.: Quantifying Forest Biomass  
747 Carbon Stocks From Space, *Curr. Forestry Rep.*, 3, 1–18, doi:10.1007/s40725-017-0052-5, 2017.
- 748 Schlütz, F., Dulamsuren, C., Wieckowska, M., Mühlenberg, M., and Hauck, M.: Late Holocene  
749 vegetation history suggests natural origin of steppes in the northern Mongolian mountain taiga,  
750 *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 261, 203–217, doi:10.1016/j.palaeo.2007.12.012, 2008.
- 751 Shvidenko, A. Z. and Schepaschenko, D. G.: Carbon balance of Russian forests, *Siberian Forest  
752 Journal*, 1, 69–92, 2014.
- 753 Sommer, M. and Treter, U.: Die Lärchenwälder der Gebirgswaldsteppe in den Randgebieten des Uvs  
754 Nuur- Beckens, *Die Erde*, 130, 173–188, 1999.
- 755 Sugimoto, A., Yanagisawa, N., Naito, D., Fujita, N., and Maximov, T. C.: Importance of permafrost as a  
756 source of water for plants in east Siberian taiga, *Ecol. Res.*, 17, 493–503, doi:10.1046/j.1440-  
757 1703.2002.00506.x, 2002.
- 758 Tchebakova, N. M., Parfenova, E., and Soja, A. J.: The effects of climate, permafrost and fire on  
759 vegetation change in Siberia in a changing climate, *Environ. Res. Lett.*, 4, 45013,  
760 doi:10.1088/1748-9326/4/4/045013, 2009.
- 761 Turner, M., Beer, C., Santoro, M., Carvalhais, N., Wutzler, T., Schepaschenko, D., Shvidenko, A.,  
762 Kompter, E., Ahrens, B., Levick, S. R., and Schmullius, C.: Carbon stock and density of northern  
763 boreal and temperate forests, *Global Ecol. Biogeogr.*, 23, 297–310, doi:10.1111/geb.12125, 2014.
- 764 Treter, U.: Gebirgs-Waldsteppe in der Mongolei, *Geographische Rundschau*, 48, 655–661, 1996.
- 765 Tsogtbaatar, J.: Deforestation and reforestation needs in Mongolia, *Forest Ecol. Manag.*, 201, 57–63,  
766 doi:10.1016/j.foreco.2004.06.011, 2004.
- 767 Umbanhowar, C. E., Shinneman, A. L.C., Tserenkhand, G., Hammon, E. R., Lor, P., and Nail, K.:  
768 Regional fire history based on charcoal analysis of sediments from nine lakes in western  
769 Mongolia, *Holocene*, 19, 611–624, doi:10.1177/0959683609104039, 2009.
- 770 Unkelbach, J., Dulamsuren, C., Punsalpaamuu, G., Saindovdon, D., and Behling, H.: Late Holocene  
771 vegetation, climate, human and fire history of the forest-steppe-ecosystem inferred from core  
772 G2-A in the 'Altai Tavan Bogd' conservation area in Mongolia, *Veget. Hist. Archaeobot.*, 27, 665–  
773 677, doi:10.1007/s00334-017-0664-5, 2017.
- 774 Unkelbach, J., Kashima, K., Enters, D., Dulamsuren, C., Punsalpaamuu, G., and Behling, H.: Late  
775 Holocene (Meghalayan) palaeoenvironmental evolution inferred from multi-proxy-studies of



- 776 lacustrine sediments from the Dayan Nuur region of Mongolia, *Palaeogeogr. Palaeoclimatol.*  
777 *Palaeoecol.*, 530, 1–14, doi:10.1016/j.palaeo.2019.05.021, 2019.
- 778 Vandandorj, S., Gantsetseg, B., and Boldgiv, B.: Spatial and temporal variability in vegetation cover of  
779 Mongolia and its implications, *J. Arid Land*, 7, 450–461, doi:10.1007/s40333-015-0001-8, 2015.
- 780 Yang, Y., Wang, Z., Li, J., Gang, C., Zhang, Y., Zhang, Y., Odeh, I., and Qi, J.: Comparative assessment of  
781 grassland degradation dynamics in response to climate variation and human activities in China,  
782 Mongolia, Pakistan and Uzbekistan from 2000 to 2013, *J. Arid Environ.*, 135, 164–172,  
783 doi:10.1016/j.jaridenv.2016.09.004, 2016.
- 784