Atmospheric deposition of organic matter at a remote site in the Central Mediterranean Sea: implications for marine ecosystem

Yuri Galletti¹, Silvia Becagli², Alcide di Sarra³, Margherita Gonnelli¹, Elvira Pulido-Villena⁴,
 Damiano M. Sferlazzo³, Rita Traversi², Stefano Vestri¹, Chiara Santinelli¹

¹CNR, Biophysics Institute, Pisa, Italy

6 ²Department of Chemistry "Ugo Schiff", University of Florence, Italy

7 3 Laboratory for Observations and Analyses of Measurements for the EarthEnvironment and Climate (SSPT-PROTER-

8 OACOEM), ENEA, Rome, Italy

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⁴Institut Méditerranéen d'Océanologie, MIO - Marseille, France
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11 *Correspondence to:* Yuri Galletti (yuri.galletti@pi.ibf.cnr.it)

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Abstract. Atmospheric fluxes of dissolved organic matter (DOM) were studied for the first time at the Island of 13 14 Lampedusa, a remote site in the Central Mediterranean Sea (Med Sea), elose to the Sahara desert, between March 19th 15 2015 and April 1st 2017. The main goals of this work arestudy were: to quantify total atmospheric deposition of DOM in this area and to evaluate the impact of Saharan dust deposition on DOM dynamics in the surface waters of the 16 Mediterranean Sea. Our data show high variability in DOM deposition rates, without a clear seasonality, and allow to 17 estimate a dissolved organic carbon (DOC) input from the atmosphere of 120.7 mmol DOC m⁻² y⁻¹. Over the entire 18 19 time-series, the average dissolved organic phosphorous (DOP) and dissolved organic nitrogen (DON) contributions to the total dissolved pools were 40% and 26%, respectively. The data on atmospheric elemental ratios also show that each 20 21 deposition event is characterized by a specific elemental ratio, suggesting a high variability in DOM composition and 22 the presence of multiple sources. This study indicates that the organic substances, transported by Saharan dust at 23 Lampedusa, site-mainly come have from a natural origin, especially from sea spray, and that Saharan dust can be an 24 important carrier of organic substances, even ifthough the load of DOC associated with dust is highly variable. Our 25 estimates suggest that atmospheric input has ana larger impact to the Med Sea larger than to the global ocean-and that, 26 Further, DOC fluxes from the atmosphere to the Med Sea can be up to 6-fold larger than total river input. Longer time 27 series, combined with a-modelling effort, are therefore mandatory in order to investigate would greatly improve our 28 understanding of the response of DOM dynamics in the Med Sea to the change in aerosol deposition pattern due to the 29 effect of climate change.

31 1. Introduction

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32 The Mediterranean Sea (Med Sea) is the largest semi-enclosed basin and one of the most oligotrophic areas inof the 33 worldworld's oceans. It is very sensitive to natural variations in the atmosphere-ocean interactions (Mermex group, 34 2011). Organic matter and nutrients of natural and anthropieanthropogenic origin, are continuously exchanged between 35 the ocean and the atmosphere, affecting biogeochemical cycles and the marine ecosystem. The Med Sea receives 36 anthropogenic aerosols from the northern regions, which are characterized by the presence of important industrial sites, 37 representing relevant sources of organic substances to the atmosphere (Guerzoni and Chester, 1996). Industrial pollution can also be originated from the North Africa as shown in the work by Rodríguez et al. (2011). In addition, the Sahara 38 39 desert is an intermittent source of mineral dust, that can transport nutrients and organic carbon to the Basin (Goudie and 40 Middleton, 2001; Prospero et al., 2005; Vincent et al., 2016). Atmospheric deposition of nutrients (N and P) strongly influences the marine biogeochemical cycles of the Med Sea, it has therefore received increased attention in the last 30 41

42 years (Migon et al., 1989; Herut et al., 2002; Ridame and Guieu, 2002; Markaki et al., 2003, 2010; Pulido-Villena et al., 43 2008; Izquierdo et al., 2012; Djaoudi et al., 2018). Compared to inorganic nutrients, there is still very few data on the 44 atmospheric deposition of Dissolved Organic Carbon (DOC) to the surface ocean, both at the local and global and local 45 scale. Organic carbon can be removed from the atmosphere-via both through wet and dry deposition (Iavorivska et al., 2016). At the global scale, wet deposition transfers about 306-580 Tg DOC yr⁻¹ to the surface of the Earth (Willey et al., 46 47 2000; Kanakidou et al., 2012). These values correspond to almost half of the DOC delivered to the oceans by rivers 48 annually (IPCC, 2014). Atmospheric deposition of organic carbon can therefore affect regional C cycling, radiative 49 forcing, and human health (Yan and Kim, 2012; Decina et al., 2018). In addition, the expected increase in ocean 50 stratification due to the global warming will enhance the impact of atmospheric inputs in the surface ecosystem 51 (Kanakidou et al., 2012). The potential magnitude of atmospheric DOC inputs to open waters and the importance of its 52 role in the carbon cycle highlight the need for a better and robust estimation of DOC deposition.

53 In the last years, a few studies have reported data on atmospheric deposition of DOC to the Med Sea. Total (dry + wet) 54 atmospheric deposition was studied in North-Western Med Sea in 2006 (Pulido-Villena et al., 2008) and in 2015 55 (Djaoudi et al., 2018) with contrasting results. In the first study, the highest DOC flux was observed in correspondence 56 withduring a Saharan dust storm, suggesting a combination of heterogeneous reactions between organic matter and 57 mineral dust in the troposphere. In the second study, thea Saharan rain event coincided with a minimum in DOC input, 58 suggesting the presence of an aerosol poorly enriched in little organic matter in aerosols (Djaoudi et al., 2018). These 59 studies were conducted in coastal areas affected by human activities. Direct measurements of total OC (TOC) in 60 rainwater were performed carried out at the island of Crete Island (Eastern Mediterranean; (Economou and 61 Mihalopoulos, 2002). This study did not take into consideration dry deposition. None of the papers cited has studied 62 atmospheric inputs in remote sites, far from possible pollution sources and/or large cities,

An interesting aspect of the Med Sea is related to Dissolved Organic Matter (DOM) stoichiometry. Mediterranean DOC
 and Dissolved Organic Nitrogen (DON) concentrations and their ratios are similar to those reported for the global ocean
 (Pujo Pay et al., 2011: Santinelli, 2015). In the surface waters (0 100 m), C:N:P ratios show that Mediterranean DOM is
 depleted in Dissolved Organic phosphorous (DOP). The study of C:N:P ratio of DOM in atmospheric deposition is
 important in order to estimate the relative contribution of atmospheric DOM input to the inventory of the surface DOM
 pool and to understand the fate of the three elements in the water column.

The main goals of this study are: (1) to quantify total atmospheric deposition of DOC, DON and DOP at the island of Lampedusa, representative of the remote marine environment of <u>in</u> the central Med Sea; (2) to investigate the contribution of natural and anthropogenic sources in atmospheric DOC; (3) to estimate the impact of atmospheric deposition on marine ecosystem.

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74 2. Materials and methods

75 2.1 Sampling site

Bulk atmospheric deposition (dry and wet) was collected at the Station for Climate Observations (35.52°N, 12.63°E),
maintained by ENEA, (the Italian National Agency for New Technologies, Energy and Sustainable Economic
Development), on the island of Lampedusa, Italy (Fig. 1), (t) http://www.lampedusa.enea.it/).

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(Pujo-Pay et al., 2011; Santinelli, 2015). In the surface waters (0-100 m), C:N:P ratios show that Mediterranean DOM is 81 depleted in Dissolved Organic phosphorous (DOP). The study of C:N:P ratio of DOM in atmospheric deposition is 82 rtant in order to estimate the relative contribution of atmospheric DOM input to the inventory of the surface DOM 83 84 ol and to understand the fate of the three elements in the water column. This study reports the results of analyses on deposition collected at Lampedusa island (35.52°N, 12.63°E) located in 85 86 central Med Sea. Lampedusa is located in an ideal position for the study of atmospheric DOC fluxes to the open Med 87 Sea. The site is interesting, in particular, to investigate the mineral dust contribution (mean dust deposition=7.4 g ^{m-2} year⁻¹, Vincent et., 2016) to DOC deposition. Lampedusa is a flat island<u>It is flat and</u> far from large islands or continental 88 89 areas and from relevant pollutant sources, Precipitation shows a significant interannual variability and is concentrated in 90 autumn and winter, with a maximum in October. Intense precipitation events, which are relatively infrequent, are generally associated with frontal passages and winds from the Northern sectors. Very dry conditions characterize late 91 spring and summer, -Although it is a remote marine environment, influences from ship traffic emissions (Becagli et al., 92 93 2012, 2017), volcanic aerosols (Sellitto et al., 2017), forest fires (Pace et al., 2005), and regional pollution (Pace et al., 2006), have been documented., their contribution to the total aerosol load is small, and Lampedusa may be taken as 94 representative for the remote marine environment of the central Med Sea. The importance of this study area is that 95 previous work on DOC atmospheric deposition to the Med was essentially confined to the coastal areas, less 96 representative of what is actually arriving to the open Med Sea. Measurements at Lampedusa provide additional 97 98 important information on the deposition in the open Med Sea.

In addition to deposition, also measurements of PM₁₀ (particulate matter with aerodynamic equivalent diameter lower
 than 10 µm) amount and chemical composition analyses, routinely performed at Lampedusa, are used in this study.

102 2.2. Atmospheric deposition sampler

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103 The sampler (Fig. 1) was positioned on the roof of the ENEA climatic station located on a 45 m a.s.l. plateau on the 104 north-eastern coast of Lampedusa. A total of 41 samples were collected between March 19th 2015 and April 1st 2017, 105 every 15 days or immediately after strong rain or dust storm events. Due to logistic constraints, 9 sampling periods were 106 longer than 20 days. The deposition sampler is similar to those successfully employed in previous studies (Pulido-107 Villena et al., 2008; Markaki et al. 2010; De Vicente et al., 2012). It is composed byof a 10 L 108 Polycarbonate polycarbonate bottle, with a polyethylene funnel attached on the top; a 20 μ M mesh covers the funnel 109 stem in order to prevent contamination by insects or organic debris. In case of For wet deposition, the amount of water 110 in the sampler was weighted then it was collected in and transferred to 250 ml polycarbonate bottles and immediately frozen. In case of dryDry deposition, was sampled by rinsing the sampler was rinsedcollector with 250 mL of ultrapure 111 MilliQ water, the samplethat was then collected intransferred into 250 ml polycarbonate bottles and immediately frozen. 112 113 A detailed description of sampling periods, deposition types, and collected volumes is reported in Table 1.

Samples for For DOC, DON and DOP analysis, samples were thawed and filtered through a sterile 0.2 μm Nylon filter
 pre-washed with 300 ml of ultrapure water to avoid any contamination. Filtered samples were frozen until the analysis.
 Before the analysis, samples were brought to room temperature (24 °C).

117 The concentration of soluble ions metals was measured on the samples filtered on quartz filters. These filters have low

- 118 blanks level for metals and ions respect to the determined concentration both in the soluble and particulate fraction. Just
- 119 after filtration the sample was divided in two portions, one for ionic content and the other for metal content, the latter

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120 was spiked by 0.1 mL of sub-boiled distilled (s.b.) HNO3: to preserve the metals in their soluble form. Samples was keep

- 121 refrigerate at +4°C until the analysis.
- 122

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123 2.3 DOC analysis

124 DOC analysis were carried byout on a Shimadzu TOC-VCSN, equipped with a quartz combustion column filled with 125 1.2% Pt on alumina pillows of ~2 mm diameter. Samples were first acidified with 2N HCl and bubbled for 3 min with 126 CO2-free ultra-high purity air in order to remove the inorganic carbon. Replicate injections were performed until the 127 analytical precision was lower than 1%. A five-point linear calibration curve was determined with standard solutions of 128 potassium hydrogen phthalate in the same concentration range as the samples (40-400 μ M). The system blank was 129 measured every day at the beginning and the end of analyses using low-carbon Milli-Q water (<3 µM C). The instrument functioningInstrument accuracy was assessed every day by comparison of data withanalyzing DOC 130 Consensus Reference Material (CRM), kindly provided by Prof. D. Hansell. DOC, with a nominal value wasof 41-44 131 μM (batch 15 Lot #07-15), DOC) (Hansell, 2005). The average DOC concentration in the CRM measured in our 132 133 laboratory during value, the period of the analysis was 42.788±1.20 (n=15) (Hansell, 2005)...).

137 2.4 DOP and DON analysis

Twenty-six samples out of the total 41 samples were analyzed for dissolved organic nitrogen (DON) and phosphorous
 (DOP). The samples were collected between March 19th 2015 and November 3rd 2016.

DON was estimated by subtracting the dissolved inorganic nitrogen (DIN) from the total dissolved N (TDN). DIN and
TDN were analyzed by conventional, automated colorimetric procedure (CACP) according to Aminot and Kerouel
(2007) with an estimated limit of detection of 0.02 μM. TDN was analyzed after persulfate wet-oxidation (Pujo-Pay et

143 al., 1997).

DOP concentrations were determined by subtracting the inorganic form (soluble reactive phosphorus, SRP) from the
total dissolved P. SRP was measured spectrophotometrically after Murphy and Riley (1962) with a limit of detection of
0.02 μM and an analytical precision of 7% at 0.1 μM. Total dissolved P (TDP) was measured as SRP after UV
digestion (Armstrong et al., 1966). The photoxidation technique included a 2 hours UV treatment in a Metrohm® 705
UV digester with a digestion efficiency of 85 ± 3 %, assessed on a 1 μM solution of β-glycerol-phosphate.

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150 2.52.5 DOC, DON and DOP fluxes

151	DOC, DON and DOP fluxes were calculated using the following formula:	
152	$\underline{\mathbf{X}}_{\mathrm{Flux}} = \frac{X \cdot V}{4 \cdot d} \tag{1}$	 Formattato: Allineato a destra
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153	where X is the concentration measured in the sample (μ M), V is the volume (L) of rain collected by the sampler or the	
154	volume of Milli-Q water (0.25 L) used to rinse the collector in case of dry deposition, A is the area of the funnel	
155	(0.1018 m ²), and d is the length of the sampling period expressed in days.	
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157 **<u>2.6</u>** Ions and metals content in the deposition samples

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The concentration of solubleSoluble ions metals waswere measured on the samples filtered on quartz filters. These
filters have low blanks levelblank levels for metals and ions respect to the determined concentration(Ca, Na, Al and Pb)
both in the soluble and particulate fraction. JustImmediately after filtration the sample wassamples were divided in two
portions, used for measurements of one for ionic content and the other for metal content, the latter wasrespectively.
Samples for the determination of metal were spiked by with 0.1 mL of sub-boiled distilled (s.b.) HNO₃ to preserve the
metals in their soluble form.

- Samples waswere keep refrigerate<u>refrigerated</u> at +4°C until the analysis. Ions were determined on thein solution by ion
 chromatography as reported in Becagli et al. (2011).
- The particulate fraction of the deposition was extracted from the quartz filter through the solubilisation procedure
 reported in the EU EN14902 (2005) rule for aerosol samples. The extraction procedure was performed in a microwave
 oven at 220_°C for 25 min by sub-boiling distilled HNO3 and 30% ultra-pure H₂O₂ for 25 minutes.
- 169 Metals were determined in both soluble and particulate fractions were measured by means of an Inductively Coupled
- 170 Plasma Atomic Emission Spectrometer (ICP-AES, Varian 720-ES) equipped with an ultrasonic nebulizer (U5000 ATC,
- 171 Cetac Technologies Inc.). Daily calibration standards (internal standard: 1 ppm Ge) were used for quantification.

173 2.67 PM₁₀ analysis

- $PM_{\mu e}$ (particulate matter with aerodynamic equivalent diameter lower than 10 µm) is routinely PM_{10} is sampled on a 174 175 daily basis at the island of Lampedusa (Becagli et al., 2013; Marconi et al., 2014; Calzolai et al., 2015) by using a low-176 volume dual-channel sequential sampler, (HYDRA FAI Instruments) equipped with two PM10, sampling heads, 177 operating at constant flow of 2.3 m³/h in accord with the European rules for aerosol monitoring (UNI EN12341-). 178 Aerosol is collected on 47 mm diameter Teflon filters (PALL Gelman) having 2 µm nominal porosity but certified to have 99% retention efficiency for 0.3 µm diameter particles. The PMI0 mass was determined by weighting the Teflon 179 180 filters-(47 mm diameter 2 µm nominal porosity) before and after sampling with an analytical balance in controlled 181 conditions of temperature (20±1 °C) and relative humidity (50±5%). The estimated error on PM₁₀ mass is around 1% at 182 30 µg m⁻³ in the applied routine sampling conditions. A quarter of each Teflon filter was extracted using MilliQ water 183 (about 10 ml, accurately evaluated by weighing) in ultrasonic bath for 15 min, and the ionic content was determined by 184 ion chromatography as for deposition samples (Becagli et al. 2011). Another quarter of the Teflon filter was used for the determination of metals in the atmospheric particulateparticles as already described for the deposition samples. 185
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187 2.78 Enrichment factor

In order to obtain information on the DOM sources, DOM concentration isconcentrations were compared withto
 concentration of Al, Na and the enrichment factor offor Pb_a (<u>EF-EF(Pb)</u>_a in the deposition <u>samples</u> as they are
 markermarkers of crustal, sea spray and anthropic source respectively.

The enrichment factor (EF) with respect to crustal source forsources of Pb, V and Ni arewere calculated by using Al as
 marker for crustal aerosol. The following equation (Eq. 1) is used for EF calculation:

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$$EF X = \frac{\left(\frac{X}{AI}\right)sample}{\left(\frac{X}{AI}\right)crust}$$
(42)

where $(X/Al)_{sample}$ is the ratio between the metal X and Al concentrations in the sample, and $(X/Al)_{crust}$ is the same ratio

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in the upper continental crust as reported in Henderson and Henderson (2009). By convention, element with EF<10 are 5

- 196 called "not enriched" having a prevailing crustal source, whereas 10<EF<100 indicate a moderate enrichment and
- 197 | EF>100 indicate that the element (called "("enriched") has a prevailing anthropogenic source (e.g. Lai et al., 2017).
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199 3. Results

200 3.1 DOC atmospheric fluxes

- 201 DOC atmospheric Atmospheric fluxes of DOC ranged between 0.06 and 1.78 mmol C m⁻² day⁻¹, with a marked high
 202 variability. The overall sampling lasted for 746 days. The deposition was lower than 0.2 mmol DOC m⁻² d⁻¹ (Fig. 2 and
 203 Table 2) in half of the sampling days (52%).
- In 2015, the lowest deposition rates (<0.1 C m⁻² d⁻¹) were measured in July (Lmp09), October (Lmp13), and November (Lmp15). The highest onesdeposition values (>1.2 mmol C m⁻² d⁻¹) occurred between March and April (Lmp02), and in June (Lmp06), both periods were characterized by dry deposition (Fig.2 and Table 2). High DOC fluxes (>0.6 mmol C m⁻² d⁻¹) were, however, also observed in March (Lmp01), May (Lmp04) and at the end of July (Lmp10), in correspondence with periods dominated by wet deposition. In 2015, the annual rainfall was 360 mm, slightly higher than the average annual rainfall atfor the island of Lampedusa (325 mm with 42 days of rain), (data from: http://www.arpa.sicilia.it/ and http://www.eurometeo.com/italian/climate).
- 211 In 2016, the DOC deposition rates were rather low and with <u>a smallerless</u> variability compared to the previous year. 212 DOC fluxes ranged between 0.1 and 0.3 mmol C m⁻² d⁻¹ from January to May (Lmp18 to Lmp23), and from June to
- August (Lmp27 to Lmp30). The highest DOC fluxes (>0.8 mmol C $m^{-2} d^{-1}$) were observed in May (Lmp25) and
- between October and November (Lmp33; Fig. 2 and Table 2). In 2016, the annual rainfall was 378 mm (data from the
 local meteorological station of the Lampedusa Atmospheric Observatory).
- In 2017, DOC fluxes ranged between 0.14 and 0.92 mmol C m⁻² d⁻¹, from January to April (Lmp36 to Lmp41); these
 values are higher than those observed in the first three months of the previous year (Fig. 2 and Table 2).
- 218 Atmospheric fluxes of DOC in the wet depositions deposition were correlated with monthly precipitation rates (r^2 =0.47,
- 219 p<0.05, n=12). The precipitation rate ranged between 2.9 and 88.5 mm-<u>during the study period (2015-2017)</u>.
- A mean daily deposition of 0.33 mmol C $m^{-2} d^{-1}$ was calculated, taking into consideration the two years (from March 2015 to April 2017), corresponding to an annual DOC flux of 120.7 mmol C m^{-2} year⁻¹.
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223 3.2 DON and TDN, DOP and TDP atmospheric fluxes

- DON and Total-Dissolved Nitrogen-(TDN) fluxes ranged between 1.5·10⁻³ and 0.25 mmol DON m⁻² d⁻¹ and between 1.6·10⁻³ and 0.47 mmol TDN m⁻² d⁻¹, respectively (Fig. 3 and Table 2). InDuring most of the sampling period (93%),
 DON deposition was lower than 0.1 mmol m⁻² d⁻¹. The main peaks were observed in March 2015 (Lmp01), in May
 (Lmp24 and Lmp25) and October 2016 (Lmp33) in correspondence_coinciding with high DOC deposition (Fig. 3 and Table 2).
- DOP and Total dissolved Dissolved phosphorous (TDP) fluxes ranged between 0 and 2.7·10⁻³ mmol DOP m⁻² d⁻¹ and
 1·10⁻⁴ and 8-5·10⁻⁵² mmol TDP m⁻² d⁻¹, respectively (Fig. 4 and Table 2). Between August 2015 and September 2016
 (Lmp10-Lmp30) both DOP and TDP showed low fluxes. In 2015, atmospheric DOP and TDP showed the highest
 fluxes in May (Lmp04) and August (Lmp10). In 2016, the main peaks in DOP and TDP deposition were observed in
 October (Lmp31) and November (Lmp33). The 4 peaks in atmospheric DOP and TDP (Lmp04, Lmp10, Lmp31 and
 Lmp33) were responsible for 16% of total depositions and were in correspondence coinciding with high DOC fluxes

235 | (Fig. 2). It is noteworthy that in March 2015 (Lmp01) and May 2016 (Lmp25), in correspondence with high fluxes of

DOC, DON and TDP, DOP wasDOP fluxes were very low (0 and 9·10⁻⁵ mmol m⁻² d⁻¹, respectively) (Table 2)...).
 despite high fluxes of DOC, DON and TDP.

Taking into consideration the entire sampling period (March 2015 – November 2016), the<u>The overall</u> mean DON and DOP daily deposition rates were 0.032 mmol N m⁻² d⁻¹ and $3.8 \cdot 10^{-4}$ mmol P m⁻² d⁻¹, corresponding to an-annual fluxes of 11.61 mmol DON m⁻² y⁻¹ and 0.14 mmol DOP m⁻² y⁻¹.

241 It should be noted that the these fluxes could be underestimated due to the missing samples in 2015 and 2016.

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243 3.3 Elemental ratios in atmospheric DOM

DOC:DON:DOP ratios showed a marked variability in the different periods (Fig. 5 and Table 3). DOC:DON molar ratios ranged between 2.2 (Lmp24, May 2016) and 45.9 (Lmp04, May 2015) (Fig. 5a). DOC:DOP molar ratios ranged between 244 (Lmp10, August 2015) and 11008 (Lmp25, May 2016) (Fig. 5b). DON:DOP ratio ranged between 9.2
(Lmp10, August 2015) and 1377 (Lmp25, May 2016) (Fig. 5c). No clear seasonal cycle was observed, even if induring autumn (November 2015 and October 2016) and late spring (May 2016) depositions were very poorlow in P, compared to the other two elements.

251 3.4 The sources of atmospheric DOM

Previous works indicate that soluble fractions of V and Ni in aerosol samples are specific marker of anthropie source
 atfor anthropogenic sources for the area of Lampedusa (Becagli et al, 2012 and 2017), but in the considered). During
 this study, samples they usually dodid not show enrichment factor higher thanbeyond 10, therefore indicating that their
 source in the deposition is mainly from crustal input-origin.

256 Besides, mean values of PM₁₀, sea salt aerosol, dust and non sea salt Ca (nss Ca) mean values in PM₁₀ samples were

257 calculated over the same intervals of the deposition measurements.

In Fig. 6 we reported the <u>The</u> DOC deposition <u>was</u> classified on the basis of the corresponding nssCa concentration in PM₁₀. (Following Marconi et al.-(., 2014) (Fig. 6). -Saharan dust events are identified as those with nssCa > 950 ng/m³.
 DOC deposition-<u>values</u> corresponding to average nssCa larger than the threshold (950 ng/mg³), <u>are-is</u> highlighted in red. DOC deposition_ corresponding to a Saharan dust event occurring in at least one day of the sampling period, <u>are-is</u> indicated in <u>orangeyellow</u> (Fig. 6). A detailed description of the most interesting deposition events is given below.

The mean concentration of PM_{10} for Lmp01 (March 2015) was 50.1 µg m⁻³, with an average dust value of 18.2 µg m⁻³ (Table 4). This sample is dominated by crustal input as revealed by the values of nssCa in the aerosol (1327.6 ng/m³) and the Al concentration in the deposition (both soluble and particulate, Fig. 7). In this sample EF(Pb) indicate

266the indicateslow contribution of the anthropic source. anthropogenic sources.Na concentration in the deposition is was267 $304 \text{ mg m}^{-2} d^{-1}$ (Fig. 7).

Lmp02 (March-April 2015) is characterized by the second highest DOC deposition, even if although no Saharan dust
event occurred induring this period (Fig. 6 and 7). The PM₁₀ mean concentration was 29 µg m⁻³, the and average seasalt aerosol value was 13.6 µg m⁻³ (Table 4), with a 47% contribution to PM₁₀ of 47%, This sample is was strongly
affected by sea spray as indicated by the Na/Al ratio that is 60-fold higher than in Lmp01, even if the concentration of
Na in the deposition is slightly low than in Lmp01.

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Lmp04 (May 2015), shows a high value of DOC;sampled during this sampling period-a Saharan dust event, also
 occurredshowed high DOC input (Fig. 6), but the concentration of Al in the deposition was quite low (Fig. 7). The PM₁₀
 mean concentration was 26.4 µg m⁻³ and the average sea-salt in the aerosol was 8.8 µg m⁻³, contributing by one third to
 the total particulate matter. As for Lmp03 the ratio Na/Al is quite high suggesting that sea spray dominates-also
 dominated in this sample.

The mean PM₁₀ concentration of Lmp06 (June 2015), was 23.3 µg m⁻³, with an average sea-salt aerosol concentration of 13.6 µg m⁻³ (Table 4). The average contribution of sea salt aerosol to the particulate matter concentration was 27%. The peculiar characteristic of this sample is the high concentration of soluble Al and low particulate Al in the deposition (Fig. 7). This feature is also observed in the-samples Lmp10 and Lmp12, both presenting quite high concentration of DOC in July and September 2015.

Lmp25 (May 2016) was characterized by a mean PM₁₀ concentration of 133.7 µg m⁻³ with a peak of 267.4 µg m³, and
with an average dust value of 42.5 µg m⁻³ (Table 4). This is the highest value of PM₁₀ observed in the entire sampling
periodstudy and indicates the occurrence of a Saharan dust event. The average value of nssCa in the sampling days-was
4815.1 ng/m³, with an incredible peak of 9207 ng/m³, highlightingfurther supporting the occurrence of an intense
Saharan dust event. The relevant Saharan dust contribution for this sample is wellalso revealed by the Al concentration
(both soluble and particulate) in the atmospheric deposition (Fig. 7).

Lmp33 (October-November 2016) and Lmp34 (November 2016) present a veryare indicative-pattern of the two possible source of DOC, crustal and sea spray. Lmp33 shows higher DOC concentration than Lmp34. The former is
characterized by very high Na concentration in the deposition, conversely the second while Lmp34 is characterized by high crustal content (as reveled by the high concentration of Al, Fig. 7).

Unfortunately PM₁₀ data are not available for the<u>The</u> fourth highest DOC deposition of the entire study period (Lmp37).
 occurred in January 2017. Unfortunately no ancillary data was collected during this event.

296 4. Discussion

295

4.1 DOC input from the atmosphere

The relationship between monthly precipitation rates and DOC fluxes confirmed the high efficiency in DOM atmospheric deposition of DOC viaimportance of rain events in the Med Sea, as recently reported by (Djaoudi et al.-(...
 2018).

301 Our data allowed for the quantification of the The total DOC annual input from the atmosphere (120.7 mmol C m⁻² y⁻¹);) found in this valuestudy, is very close to that measured at Cap Ferrat peninsula (Southern France) in 2006 (129 mmol C 302 303 m⁻² y⁻¹; Pulido-Villena et al., 2008) and in three lakes in the western Mediterranean basin (Southern Spain, 153.3 mmol C m⁻² y⁻¹ in 2005; De Vicente et al., 2012). This value is higher than that reported fordeposition in the north-western 304 Med Sea from February 2015 to July 2016 (59 mmol C m⁻² y¹ at Frioul island, Island in the Bay of Marseille Bay (59 305 306 mmol C m² y⁴; Djaoudi et al., 2018). If the same sampling period is taken into consideration for both studies (from 307 March 2015, the beginning of sampling in Lampedusa, to July 2016, the end of sampling at Frioul island), DOC input is 308 2-times higher at Lampedusa than at Frioul Island. This variability is probably due to the different temporal and 309 seasonal cycles of dry and wet deposition. In particular the marked differences between these two sites could be 310 influenced by the presence of a south-north decreasing gradient in the intensity of the mineral dust deposition as 311 proposed by Vincent et al. (2016). Our data also show high variability in DOC deposition rates without a clear

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seasonality. <u>IfWhile</u> in 2015 and 2016 the highest deposition rates were between spring and autumn, in 2017 the highest deposition rates were in winter. In addition the two highest peaks observed in 2015 (Lmp02 and Lmp06, dry deposition)
 together accounted together for 43% of the annual DOC flux (52 mmol C m⁻² y⁻¹). Depending on the origin and trajectories of the air masses, the atmosphere can carry significant amounts of DOC.

Assuming that the annual DOC flux from this study (120.7 mmol C m⁻² y⁻¹) is valid for the whole Med Sea (area=2.5·10¹² m²), we can estimate a total input of 3.64 Tg DOC y⁻¹. The global estimation for wet atmospheric DOC deposition is 306-580 Tg C y⁻¹ and the input to the global ocean ranges between 90 and 246 Tg C y⁻¹ (Willey et al., 2000; Kanakidou et al., 2012). The global dry deposition of <u>organic carbon (OC)</u> has been estimated to be 11 Tg C y⁻¹, (Jurado et al., 2008) leading to a total OC deposition to the oceans of 101-247 Tg C y⁻¹. The comparison of these estimates indicates that the Med Sea, with an area equivalent to only 0.7% of the global oceans, receives from 1.5 to 4%

322 of the global atmospheric input of DOC., despite it covers only 0.7% of the global oceans area.

MoreoverIt is noteworthy that, if we consider the riverine DOC fluxes, our values are up to 6 times larger than the estimate of the total river input to the Med Sea (0.6-0.7 Tg DOC y⁻¹; Santinelli, 2015). These results confirm the leadleading role of atmosphere in the transport of alloethonousallochthonous DOC to the Med Sea, as suggested recently by <u>Santinelli et al. (2015) and</u> Galletti et al. (2019).

FewA few episodes of Saharan outbreaks can strongly affect the annual dust flux, indeedwhereby a single outbreak can 327 328 account for 40-80% of the flux (Guerzoni et al., 1997). The most intense dust deposition events in Lampedusa generally 329 display larger higher values in spring (March-June) and in autumn (Vincent et al., 2016; Bergametti et al., 1989; Love-330 Pilot and Martin, 1996; Avila et al., 1997; Ternon et al., 2010). Deposition data in In this study, we show this work 331 reveal that, although dust events can significantly contribute to the annual DOC fluxes, but sea spray seems the dominant source of DOC in this area, in agreement with Mallet et al., 2019. The role of secondary organic aerosols as a 332 333 source of organic matter in the Mediterranean Sea is well documented (Arndt et al., 2017; Michoud et al., 2017; Rinaldi 334 et al., 2017) and could be relevant at Lampedusa. Finally, the correlation between monthly precipitation rates and DOC 335 fluxes shows the importance of rain events as a source of DOC in the Med Sea, as proposed by Djaoudi et al. (2018).

It should also be stressed that the DOC dynamics and its annual fluxes are not only influenced by dust deposition
 events. The wet deposition is also relevant, and the correlation between monthly precipitation rates and DOC fluxes
 confirms the high efficiency in DOC atmospheric deposition via rain events in the Med Sea, as recently proposed by
 Djaoudi et al. (2018)₄

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342 4.2 Atmospheric DON, DOP input and elemental ratios

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The <u>DON</u>-annual <u>DON</u> flux (11.61 mmol N m⁻² y⁻¹); observed at Lampedusa; was lower than that measured at Frioul 343 Island (17.80 mmol N m⁻² y⁻¹; Djaoudi et al., 2018). OnlyIn the study byEastern Med Sea, Markaki et al. (2010) reports 344 data on atmospheric DON fluxes and is focused on the Eastern Med Sea. These authors-reported ana DON annual flux 345 (of 18.49 mmol N m⁻² y⁻¹), higher than that observed at Lampedusa. The comparison among of our DOP deposition 346 values (0.14 mmol P m⁻² y⁻¹) and with the few-DOP data reported in the literature shows that the fluxes at Lampedusa 347 348 are markedly higher than those reported for the Western Med Sea (0.07 mmol P m⁻² y⁻¹, Djaoudi et al., 2018; 0.03 mmol P m⁻² y⁻¹, Migon and Sandroni, 1999), whereas they are but lower than those obtained by Violaki et al. (2017) for both 349 the West (1.16 mmol P m⁻² y⁻¹) and East (0.90 mmol P m⁻² y⁻¹) Med Sea. Our values Results from our study are instead 350

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very similar to those reported for the Eastern Med Sea in 2001 and 2002 (0.15 mmol P $m^{-2} y^{-1}$) (Markaki et al., 2010).

Further, it is interesting to note that our DOP data are very similar to the TDP data reported for a coastal rural site in the
 NE Spain in 2002-2003 (0.10-0.14 mmol P m⁻² y⁻¹, in 17 months of sampling; Izquierdo et al., 2012).

Over the entire time-series, the average DOP and DON contributions to TDP and TDN were 40% and 26%, respectively. These data confirm that a significant fraction of the dissolved P and N in the atmospheric deposition was in the organic form. These values are similar to those observed in previous studies at Frioul Island (DOP 40%, DON 25%; Djaoudi et al., 2018), and in both the western and eastern Med Sea (DOP 38%; DON 32%; Markaki et al., 2010). The similarity among the depositions collected at the two sites (Lampedusa, Central Med Sea and Frioul, North-western Med Sea) suggests that the remote site of Lampedusa may be representative of what_DON and DOP deposition in the MediterraneanMed Sea-area receives in terms of DON and DOP, especially in the western basin.

361 The data on atmospheric elemental ratios show that each deposition event is characterized by a specific elemental ratio, 362 suggesting a high variability in DOM composition and the presence of multiple sources. Djaoudi et al (2018) observed 363 an average value of DOC:DON:DOP molar ratios of 1228:308:1 in atmospheric DOM, collected in the north-western 364 Med Sea. In the surface Med Sea, DOC:DON:DOP ratios ranges between 1050:84:1 in the western basin to 1560:120:1 365 in the eastern basin (Pujo-Pay et al., 2011). The average values observed in our atmospheric deposition time-series 366 (1909:292:1) indicate that atmospheric DOM is enriched in DOC and DON with respectcompared to marine DOM. This 367 observation is also valid when we compare comparing our values with those recently measured on marine samples 368 collected at the MOOSE ANTARES offshore station (north-western Med Sea) (1227:100:1, Djaoudi et al., 2018).

All the analyzed samples, except few cases in summer 2016, are relative to dry+wet deposition (Table 1). Although the
 DON and DOP recorded during the dry period are generally on the low end side of the measured range (Table 2), no
 information on the role played by wet or dry deposition on DON and DOP input to the Med Sea can be drawn at this
 stage, due to the limited number of dry samples.

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374 4.3 The contribution of Saharan dust to atmospheric fluxes of dissolved organic carbon

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The input of Saharan dust has important effects oncan affect the chemistry of the Mediterranean aerosols and its
deposition can enrich the Med Sea with many elements (such as Co, Ni, trace metals). Very few datastudies are
available on the interactions between organic carbon and Saharan dust, even ifthough the organic material found in the
troposphere is often associated with dust particles (Usher et al., 2003; Aymoz et al., 2004).

379 Our results show that Saharan dust events can represent a relevant, albeit intermittent, source of DOC to the central Med 380 Sea. Focusing on the different peaks of DOC deposition, our results indicate that Lmp01, Lmp04 and Lmp25 are 381 associated to a Saharan dust event and that the aerosol, during its route to Lampedusa, was aerosols were probably 382 enriched with organic substances. We hypothesize that the dust particles present in the aerosol worked as condensation 383 nuclei for adsorb organic molecules, facilitating their accumulation and transport (Usher et al., 2003). The role of Saharan dust in the transport of DOC is evident in Lmp25 (May 2016) characterized by high DOC, when an intense 384 intrusion of Saharan air masses was favored by one synoptic situation in which the role of the cyclonic circulation with 385 386 a minimum depression was significant (www.meteogiornale.it).

-In addition Lmp01 (end of March 2015), Lmp04 (May 2015) and Lmp25 (May 2016) show a seasonality that could be
 linked to the transport of pollen attached to desert particles in the spring events, and this pollen would contribute to

389 atmospheric DOC input in spring (end of March- May). Pollen originating in Morocco was detected in South Spain

390 (Cabezudo et al., 1997) and various pollen types (Cannabis, Cupressus, Pinus, Platanus and Sambucus) were observed

in Cordoba (South Spain) exclusively during dust African events (Cariñanos et al., 2004). This process would not occur
 in the other seasons (winter and autumn of no pollen production).

If all the Saharan dust deposition events (red and orangeyellow in Fig. 6) are taken into account, thean input of 49.58 mmol DOC m⁻² to Lampedusa during the study period can be estimated, this value represents representing ~41% of the total DOC flux for the entire sampling period. Instead, if only the The strong dust events (red in Fig. 6) are taken into eonsiderationlead a flux of 15.26 mmol DOC m⁻² can be estimated, representing 13% of the total flux. Anyhow eachEach deposition event must be considered individually, because it can be characterized by an enrichment of DOC 998 or not depending as DOC content depends on the aerosol load (Formenti et al., 2003; Aymoz et al., 2004).

399 Wet deposition mainly controlsis the fluxmain driver of Saharan dust deposition to the Med Sea, but, However, dry deposition can be also important (Guerzoni et al., 1997) and its relative importancecontribution strongly depends on 400 401 meteorological conditions and local emission (Inomata et at., 2009). Some models have estimated that wet deposition represents up to 75-95% of total deposition (Iavorivska et al., 2016). Our data-While our results_confirm the importance 402 of wet deposition, but similarly dry depositionit also plays a crucial role. Our results stressstress the relevance of dry 403 deposition (32% of the total deposition during the entire sampling period) that, in the remote site of Lampedusa, 404 405 appears to be the main contributor of DOC and of other chemical species,- to the remote site of Lampedusa, as 406 suggested in the past by Morales-Baquero et al. (2013).

It is also evident by our data that Saharan dust input is not always associated with high DOC input, it cannot therefore
be considered as the only process for DOC transport and deposition. For instance sample seen in Lmp34, that
showswith high concentration of dust, is not characterized by high but with low DOC concentration of DOC and.
<u>Conversely</u>, several samples (for example Lmp02, Lmp33 and Lmp37) characterized by high
eoncentration_concentrations of DOC, do not show high crustal content. Indeed high DOC deposition events seems to be
often associated to sea spray transport, for instance in samples (Lmp02, Lmp10, Lmp12, Lmp 33 and Lmp 37-(; Fig. 6,
Similarly, samples Lmp01, Lmp04, Lmp10, Lmp12 and especially Lmp25, also show a large contribution of sea

spray aerosol indicating a marine source for the DOC in gray), these samples. This is a surprising result, because other
 studies (e.g., Pace et al., 2006) have shown that clean marine aerosol conditions are rare at Lampedusa.

416 Lmp23, Lmp27, Lmp32, Lmp35 and Lmp36 were not characterized by high DOC fluxes (Fig. 6), even if these 417 sampling periods were characterized by at least event (Fig. 6, in orangevellow). This observation supports the hypothesis that dust from Saharan regiondust is not typically enriched with DOC, but it 418 behaves as aggregation center of adsorbs organic molecules in the atmosphere, and depending on its route-it, can be 419 420 enriched or not in DOC-during its route. The composition of sample_Lmp34 further supports this hypothesis. This 421 sample shows, with the third highest average nss Ca value (1092.2 ng/m³), neverthelessbut with a DOC was very lowconcentration (0.20 mmol m⁻² d⁻¹); below the daily average flux of the entire sampling period (0.33 mmol m⁻² d⁻¹); 422 423 (Fig. 6, Tables 2 and 4).

However the data of atmospheric Na and soluble Al suggest a very high contribution of sea spray aerosol (the highest of
 the whole considered period) (Fig. 7). Therefore, for samples Lmp33 and Lmp37 the DOC source seems to be primary

426 marine instead of crustal as for the samples Lmp01, Lmp04, Lmp10, Lmp12 and especially Lmp25.

427 Lastly, it is interesting to noticenote that samples characterized by high values of DOC never present high EF(Pb) and
 428 the samples). Samples presenting EF(Pb)>10 presentshow very low DOC concentration, suggesting that anthropic

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429 sources have<u>concentrations</u>, indicating a small impact on DOC deposition<u>DOC contribution from anthropogenic</u>
 430 aerosols at Lampedusa.

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432 4.4 Implications for marine ecosystem

The measurements carried out at the Island of Lampedusa clearly show that the atmosphere is an important source of allocthonous DOC to the Central Med Sea. Very few information is available about the biological lability of atmospheric DOC: if labile, it can be used very quickly by the microbial loop, whereas if it is mainly recalcitrant, it can accumulate and be transported by water masses circulation. There is still little information on biological lability of atmospheric DOC; if it is biologically available, it can be used very quickly by marine prokaryotic heterotrophs and it can be channeled into the food web, whereas if it is mainly recalcitrant, it can accumulate and be transported by water masses circulation.

Commento [CS4]: We believe that this sentence is important. Labile and biologically available is the same concept.

440 In A conceptual exercise can be made in order to give an estimate of the impactimplications of atmospheric DOCDOM 441 deposition for marine ecosystem. According to the surface waters we took into consideration a D'Ortenzio et al. (2005), 442 mixed layer depth (MLD) rangingranges between 15 and 30 m, typical of the sea-close to the island of Lampedusa-in September, in according with the estimation reported by D'Ortenzio et al. (2005). Santinelli et al. (2012) observed an 443 444 average mixed layer_DOC concentration of 60 μ M in the same area in September 1999-(in the mixed layer), and 445 estimated a bacterial carbon demand (BCD) of 0.32 µM C d⁻¹ (assuming a bacterial growth efficiency of 15%), which 446 represents the total amount of carbon needed to support the observed bacterial production. In September, the 447 atmospheric DOC flux was 0.24 mmol C m⁻² d⁻¹ in 2015 and 0.38 C m⁻² d⁻¹ in 2016, dividing them, Dividing the 448 atmospheric deposition by the average MLD (22.5 m)+, D'Ortenzio et al. 2005), we estimate that the atmospheric input 449 is expected to determine acontributes to 0.011-0.017 µM DOC d⁻¹ increase in the mixed layer. Assuming that the values 450 of BCD observed in September 1999 (0.32 μ M C d⁻¹) are valid also for September 2015 and 2016, and that all the DOC 451 coming from the atmosphere is labile, it could satisfy 3-5% of the daily BCD. Instead inDuring summer the MLD 452 isyaries between 10 and 15 m depth, with an average value of 12.5 m (D'Ortenzio et al. 2005). The DOC input from the 453 atmosphere is expected to increase the DOC concentration in the mixed layer by 0.008-0.079 μ M C d⁻¹ from June to August 2015, and by 0.013-0.014 from June to August 2016, - Assuming that a BCD of 0.32 µM C d⁺ is valid also for 454 summer (three months) and that all the atmospheric DOC is labile, it could satisfy supplying 3-25% of the daily BCD,-455 456 These results highlight assuming similar DOC concentrations and bacterial activity as during September. Even if we are 457 aware that these assumptions are hardly meet, in particular the estimate of DOC input to the whole Med Sea, based on 458 the data collected in Lampedusa, we think that these calculations can give an idea of the relevant role of atmosphere that 459 atmospheric input of DOC can have in sustaining the bacterial heterotrophic prokaryotes productivity in the surface 460 layer, particularly when the upper water column is strongly stratified.

The Mediterranean MLD seasonal variability is characterized by a basin scale deepening from November to February-March and an abrupt stratification in April, which is maintained throughout the summer and early autumn. Even if these dataestimates stress the potential role of atmospheric DOC in sustaining bacterial productivity in the surface ocean, a time series of BCD, MLD and DOC concentrations in the surface layer<u>t</u> together with a network of stations for the quantification of atmospheric input of DOC in the different areas of the Med Sea, are erucialmandatory-in order to have an accurate estimationestimate of the DOC atmospheric input-impact of DOC atmospheric on the functioning of marine ecosystem. It should be also noted that a fraction of atmospheric DOC could be recalcitrant, and therefore could be 468 transported tothrough transport to depth, playingit could play a key role in carbon sequestration to depth. The refractory 469 nature of a part of atmospheric <u>DOC-DOM is hypothesizedhas been proposed</u> by Sánchez-Pérez et al. (2016), who 470 collectedbased on a 2-year time series data on Fluorescent DOM (FDOM) deposition in the North-western Med Sea 471 and studied the changes in the quality and quantity of marine DOM in the (Barcelona coastal area, (Spain). Their results 472 show that atmospheric inputs induced changes in the quality of organic matter, increasing the proportion of FDOM 473 substances in <u>DOM pool.</u> the DOM pool. Incubation experiments to investigate the biological lability of atmospheric 474 DOC are also crucial to better understand the impact of atmospheric deposition on marine ecosystems.

475 Lastly, as highlighted in the previous paragraphsFinally, the occurrence of Saharan dust events opens interesting 476 considerations on their impact on the marine environment. Previous studies suggested that dust inputs can promote 477 autotrophic production (Ridame and Guieu, 2002; Markaki et al., 2003). Instead Pulido-Villena et al. (2008) 478 experimentally found that heterotrophic bacteria can reduce the amount of C exported to deeper waters, because a 479 Saharan dust event would have induced the mineralization of 22-70% of bioavailable DOC, changing carbon 480 sequestration.

481

482 4. Conclusions

- 483 Our data show that atmospheric input has a larger impact to the Med Sea than to the global ocean and DOC fluxes from
 484 the atmosphere to the Med Sea can be up to 6-fold larger than *riverriverine* input.
- This study indicates that the organicOrganic substances transported by Saharan dust at Lampedusa site mainly have are
 primarily of natural origin, especiallyin particular from sea spray-and that. Saharan dust can be an important carrier of
 organic substances. TheHowever, the load of DOC associated with dust is veryhighly variable and high DOC fluxes
 were observed also in absence of dust deposition events.
- 489 Atmospheric C:N:P molar ratios indicate that DOM is enriched in DOC and DON with respect to marine DOM and that
 490 the contribution of atmospheric deposition to the marine DOM stoichiometry in the Med Sea could be relevant, in
- 491 particular during the stratification periodstratified periods.
- For future studies, atmospheric and marine DOM molar ratios (C:N:P) could be measured over time in order to obtain 492 493 information about changes in marine DOM pool. Further studies are needed to understand the link between atmospheric 494 inputs and marine biogeochemistry. Data on stable carbon (δ^{13} C) on atmospheric DOC would be crucial in order to gain 495 information about its main sources. Incubation experiments should be carried out, both with aerosol rich or poor in 496 DOC, in order to better understand how the microbial community can respond to dust input. Further studies are also 497 needed to understand the link between aerosol origin and DOM concentration and quality, and to comprehend the 498 potential link between DOC and the pollen during the spring. Lastly, longer time series -combined with a modelling 499 effort, are highly desirable in order would provide a solid base to assess the response of DOM dynamics in the Med Sea 500 to the changechanges in aerosol deposition pattern due to the effect of climate change. 501

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502 Author contribution

503 YG and CS conceived of the study and the sampling design. YG, SB, DMS collected the samples. YG, MG, SB, RT, 504 SV analyzed the samples. YG, CS, EPV, AdS analyzed the data and all authors assisted with data discussion and 505 contributed to the revision and editing of the final manuscript. All authors are aware of and accept responsibility for this 506 manuscript and have approved the final submitted manuscript.

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514 *The authors declare that they have no conflict of interest.*

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516 The dataset generated for this study are available on request to the corresponding author.

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Figure 1: <u>Sampling location</u> (Lampedusa island (35.5° N, 12.6° E) and the <u>total atmospheric deposition</u> samplercollector.

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Commento [CS5]: Aggiungi nella legenda DON davanti a wet and dry.. and dry

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Commento [CS6]: Aggiungi nella legenda DOP davanti a wet and dry.. and dry



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Figure 6. Intensity of Temporal dynamics in the dust deposition events during the sampling period <u>color coded</u> based on <u>the contribution of non-sea</u> salt Ca (nssCa) values.). The month abbreviation and tick marks correspond to the end of the corresponding month. The width of the bar refer to the length of the sampling period.



correspond to the	<mark>e end of th</mark>	e corresponding	month.	<u>Fhe width</u>	of the	bar refer	to the	length	of the
sampling period.									

Sample		Sampling period	Deposition	Volume collected		
name	Start date	End date	Total days	type	[L]	
Lmp01	18/03/2015	28/03/2015	10	wet and dry	6	
Lmp02	28/03/2015	17/04/2015	20	dry	0.26	
Lmp03	17/04/2015	02/05/2015	16	dry	0.27	
Lmp04	02/05/2015	21/05/2015	19	wet and dry	1.8	
Lmp05	21/05/2015	05/06/2015	15	dry	0.28	
Lmp06	05/06/2015	19/06/2015	15	dry	0.29	
Lmp07	19/06/2015	04/07/2015	16	dry	0.26	
Lmp08	04/07/2015	17/07/2015	14	dry	0.26	
Lmp09	17/07/2015	31/07/2015	14	dry	0.27	
Lmp10	31/07/2015	21/08/2015	20	wet and dry	9	
Lmp11	21/08/2015	11/09/2015	22	wet and dry	2	
Lmp12	11/09/2015	01/10/2015	20	wet and dry	5	
Lmp13	01/10/2015	30/10/2015	29	wet and dry	0.5	
Lmp14	30/10/2015	09/11/2015	11	wet and dry	2	
Lmp15	09/11/2015	23/11/2015	14	wet and dry	0.6	
Lmp16	23/11/2015	02/12/2015	9	wet and dry	1.2	
Lmp17	02/12/2015	21/12/2015	19	wet and dry	1.9	
Lmp18	21/12/2015	08/01/2016	18	wet and dry	1.8	
Lmp19	08/01/2016	28/01/2016	20	wet and dry	6.1	
Lmp20	28/01/2016	16/02/2016	19	wet and dry	2.7	
Lmp21	16/02/2016	11/03/2016	26	wet and dry	2.1	
Lmp22	11/03/2016	09/04/2016	28	wet and dry	7.1	
Lmp23	09/04/2016	04/05/2016	26	wet and dry	0.3	
Lmp24	04/05/2016	10/05/2016	6	wet and dry	2.3	
Lmp25	10/05/2016	13/05/2016	3	wet and dry	1.9	
Lmp26	13/05/2016	01/06/2016	19	wet and dry	0.7	
Lmp27	01/06/2016	22/07/2016	50	dry	0.26	
Lmp28	22/07/2016	10/08/2016	19	dry	0.24	
Lmp29	10/08/2016	26/08/2016	16	dry	0.24	
Lmp30	26/08/2016	12/09/2016	17	wet and dry	0.8	
Lmp31	12/09/2016	08/10/2016	26	wet and dry	12	
Lmp32	08/10/2016	24/10/2016	16	wet and dry	0.5	
Lmp33	24/10/2016	03/11/2016	10	wet and dry	11	
Lmp34	03/11/2016	21/11/2016	18	wet and dry	12	
Lmp35	21/11/2016	13/12/2016	22	wet and dry	1.7	
Lmp36	13/12/2016	02/01/2017	20	wet and dry	9.5	
Lmp37	02/01/2017	19/01/2017	17	wet and dry	6.5	
Lmp38	19/01/2017	03/02/2017	15	wet and dry	1.5	
Lmp39	03/02/2017	17/02/2017	14	wet and dry	5	

Lmp40	17/02/2017	03/03/2017	14	wet and dry	0.75
Lmp41	03/03/2017	01/04/2017	29	wet and dry	5.5

740Table 1. Sampling period, type of deposition and volume for the 41 samples collected at the Island of741Lampedusa.

Sample name	DOC fluxes [mmol m ⁻² d ⁻¹]	DON fluxes [mmol m ⁻² d ⁻¹]	TDN fluxes [mmol m ⁻² d ⁻¹]	DOP fluxes [mmol m ⁻² d ⁻¹]	TDP fluxes [mmol m ⁻² d ⁻¹]
Lmp01	0.80	0.10	0.17	0	$7 \cdot 10^{-4}$
Lmp02	1.30	n.a.	n.a.	n.a.	n.a.
Lmp03	0.19	n.a.	n.a.	n.a.	n.a.
Lmp04	0.63	0.01	0.05	9·10 ⁻⁴	3.10-3
Lmp05	0.36	n.a.	n.a.	n.a.	n.a.
Lmp06	1.78	n.a.	n.a.	n.a.	n.a.
Lmp07	0.25	n.a.	n.a.	n.a.	n.a.
Lmp08	0.10	n.a.	n.a.	n.a.	n.a.
Lmp09	0.07	n.a.	n.a.	n.a.	n.a.
Lmp10	0.65	0.02	0.11	3.10-3	5.10-3
Lmp11	0.11	0.01	0.09	1.10-4	$6 \cdot 10^{-4}$
Lmp12	0.31	0.03	0.10	3.10-4	$1 \cdot 10^{-3}$
Lmp13	0.06	$1.5 \cdot 10^{-3}$	$1.6 \cdot 10^{-3}$	7.10-5	8.10-5
Lmp14	0.10	0.01	0.03	$2 \cdot 10^{-4}$	$4 \cdot 10^{-4}$
Lmp15	0.09	8·10 ⁻³	0.06	2.10-5	$6 \cdot 10^{-4}$
Lmp16	0.43	0.05	0.14	$2 \cdot 10^{-4}$	$5 \cdot 10^{-4}$
Lmp17	0.11	0.03	0.11	7·10 ⁻⁵	$1 \cdot 10^{-4}$
Lmp18	0.13	0.04	0.25	9·10 ⁻⁵	$4 \cdot 10^{-4}$
Lmp19	0.23	0.04	0.20	3.10-4	9·10 ⁻⁴
Lmp20	0.14	0.03	0.12	$2 \cdot 10^{-4}$	$2 \cdot 10^{-4}$
Lmp21	0.12	0.02	0.10	$2 \cdot 10^{-4}$	3.10-4
Lmp22	0.18	0.02	0.05	3.10-4	$4 \cdot 10^{-4}$
Lmp23	0.18	6·10 ⁻³	0.10	1.10-4	$2 \cdot 10^{-4}$
Lmp24	0.57	0.25	0.47	3.10-4	$2 \cdot 10^{-3}$
Lmp25	0.98	0.12	0.32	9·10 ⁻⁵	$3 \cdot 10^{-3}$
Lmp26	0.33	n.a.	n.a.	n.a.	n.a.
Lmp27	0.17	0.02	0.08	$2 \cdot 10^{-5}$	$1 \cdot 10^{-4}$
Lmp28	0.14	0.05	0.19	9·10 ⁻⁵	$5 \cdot 10^{-4}$
Lmp29	0.17	0.01	0.12	$2 \cdot 10^{-4}$	$1 \cdot 10^{-3}$
Lmp30	0.28	0.04	0.18	$2 \cdot 10^{-4}$	$9 \cdot 10^{-4}$
Lmp31	0.45	0.07	0.22	$1 \cdot 10^{-3}$	$3 \cdot 10^{-3}$
Lmp32	0.08	0.02	0.10	$4 \cdot 10^{-5}$	$2 \cdot 10^{-4}$
Lmp33	0.80	0.10	0.34	$2 \cdot 10^{-3}$	$2 \cdot 10^{-3}$
Lmp34	0.20	n.a.	n.a.	n.a.	n.a.
Lmp35	0.10	n.a.	n.a.	n.a.	n.a.
Lmp36	0.29	n.a.	n.a.	n.a.	n.a.
Lmp37	0.92	n.a.	n.a.	n.a.	n.a.
Lmp38	0.32	n.a.	n.a.	n.a.	n.a.
Lmp39	0.57	n.a.	n.a.	n.a.	n.a.
Lmp40	0.14	n.a.	n.a.	n.a.	n.a.
Lmp41	0.43	n.a.	n.a.	n.a.	n.a.

Table 2. Atmospheric fluxes of DOC, DON, TDN, DOP and TDP at the Island of Lampedusa.

Sample	Sampling	C:N	C:P	N:P	
Lmp01	28/03/2015	7.78	n.a.	n.a.	
Lmp04	21/05/2015	45.87	715.08	15.59	
Lmp10	21/08/2015	26.57	244.38	9.20	
Lmp11	11/09/2015	8.67	807.94	93.15	
Lmp12	01/10/2015	11.37	977.79	85.98	
Lmp13	30/10/2015	39.44	864.07	21.91	
Lmp14	09/11/2015	8.02	449.04	56.00	
Lmp15	23/11/2015	11.26	5131.65	455.83	
Lmp16	02/12/2015	7.97	2036.66	255.42	
Lmp17	21/12/2015	4.24	1448.37	341.90	
Lmp18	08/01/2016	3.34	1406.60	420.55	
Lmp19	28/01/2016	5.38	832.69	154.79	
Lmp20	16/02/2016	5.09	882.80	173.40	
Lmp21	11/03/2016	6.63	812.40	122.55	
Lmp22	09/04/2016	8.78	645.65	73.53	
Lmp23	04/05/2016	30.48	1353.57	44.41	
Lmp24	10/05/2016	2.24	1976.03	882.33	
Lmp25	13/05/2016	7.99	11008.94	1377.41	
Lmp27	22/07/2016	8.73	7405.29	848.62	
Lmp28	10/08/2016	2.89	1641.49	568.76	
Lmp29	26/08/2016	12.66	796.68	62.95	
Lmp30	12/09/2016	8.06	1376.27	170.77	
Lmp31	08/10/2016	6.74	356.03	52.84	
Lmp32	24/10/2016	3.41	2275.72	666.53	
Lmp33	03/11/2016	7.68	389.57	50.73	
Table 3. C:N:P molar ratios in atmospheric DOM.					

Sample name	Mean PM ₁₀ [µg/m ³]	Mean sea salt aerosol [µg/m ³]	Mean dust [ug/m ³]	Mean nssCa [ng/m ³]
Lmp01	50.1	13.0	18.2	1327.6
Lmp02	29.0	13.6	n.a.	62.2
Lmp03	28.1	9.8	n.a.	371.6
Lmp04	26.4	8.8	4	351.7
Lmp05	16.7	4.6	n.a.	87.3
Lmp06	23.1	6.1	n.a.	166.1
Lmp07	22.2	7.1	n.a.	139.3
Lmp08	26.5	5.4	n.a.	311.6
Lmp09	28.3	8.0	n.a.	188.2
Lmp10	29.1	5.2	3.4	492.7
Lmp11	n.a.	n.a.	n.a.	n.a.
Lmp12	n.a.	n.a.	n.a.	n.a.
Lmp13	n.a.	n.a.	n.a.	n.a.
Lmp14	n.a.	n.a.	n.a.	n.a.
Lmp15	n.a.	n.a.	n.a.	n.a.
Lmp16	n.a.	n.a.	n.a.	n.a.
Lmp17	n.a.	n.a.	n.a.	n.a.
Lmp18	n.a.	n.a.	n.a.	n.a.
Lmp19	n.a.	n.a.	n.a.	n.a.
Lmp20	n.a.	n.a.	n.a.	n.a.
Lmp21	n.a.	n.a.	n.a.	n.a.
Lmp22	n.a.	n.a.	n.a.	n.a.
Lmp23	39.5	18.3	3.8	488.1
Lmp24	30.7	18.7	1.2	154
Lmp25	133.7	15.5	42.5	4815.1
Lmp26	25.9	13.1	1.5	168.8
Lmp27	26.2	9.3	2.3	319.5
Lmp28	24.7	8.8	1.8	161.5
Lmp29	25	9.6	1.0	235.9
Lmp30	22.4	5.1	n.a.	330.8
Lmp31	24.5	5.6	n.a.	286.2
Lmp32	32.9	8.7	n.a.	772.5
Lmp33	31.8	11.8	n.a.	344.2
Lmp34	35.3	7.8	n.a.	1092.2
Lmp35	22.3	7.5	0.4	394
Lmp36	35.8	12.3	4.6	661.5
Lmp37	n.a.	n.a.	n.a.	n.a.
Lmp38	n.a.	n.a.	n.a.	n.a.
Lmp39	n.a.	n.a.	n.a.	n.a.
Lmp40	n.a.	n.a.	n.a.	n.a.

	Lmp41	n.a.	n.a.	n.a.	n.a.	
747	Table 4. The PM ₁	₀ , sea salt aerosol,	dust and non-sea	salt Ca (nss Ca)<u>nss</u>	<mark>Ca</mark> mean values	of the atmospheric
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776 **Reply to Editor**

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778 (EDITOR)

779 One of major conclusion of this manuscript is the organic species deposited in Lampedusa coming 780 from natural sources (sea spray and dust), notably in summer. However, your discussion is supported by a 781 too limited literature about OM sources in atmospheric aerosols in this area. One of the reviewers 782 reminded you that you did not consider the potential role played by pollen.

783 784

(AUTHORS)

785 We thank the editor for this comment and we apologize for the missing references. In the revised 786 manuscript, all the suggested paper about OM sources in atmospheric aerosols are cited and we stressed the 787 potential role played by pollen. The following sentences were added in the discussion (paragraph 4.3): . "The 788 role of secondary organic aerosols as a source of organic matter in the Mediterranean Sea is well 789 documented (Arndt et al., 2017; Michoud et al., 2017; Rinaldi et al., 2017) and could be relevant at 790 Lampedusa." and "In addition Lmp01 (end of March 2015), Lmp04 (May 2015) and Lmp25 (May 2016) 791 show a seasonality that could be linked to the transport of pollen attached to desert particles in the spring 792 events, and this pollen would contribute to atmospheric DOC input in spring (end of March- May). Pollen 793 originating in Morocco was detected in South Spain (Cabezudo et al., 1997) and various pollen types 794 (Cannabis, Cupressus, Pinus, Platanus and Sambucus) were observed in Cordoba (South Spain) exclusively 795 during dust African events (Cariñanos et al., 2004). This process would not occur in the other seasons 796 (winter and autumn), when no pollen production occurs." A sentence about the need of further studies was 797 also added in the conclusions.

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802

799It is also surprising that your discussion does not take into account the presence of atmospheric800secondary organic aerosols as a source of organic matter at Lampedusa in spring and summer (see Mallet801et al., ACP, 2019) but generally in Med Sea (Arndt et al., 2017; Michoud et al., 2017; Rinaldi et al., 2017).

We thank the editor for this comment and we apologize for the inaccuracy. The following sentence
was added in the discussion (paragraph 4.1): "*The role of secondary organic aerosols as a source of organic matter in the Mediterranean Sea is well documented (Arndt et al., 2017; Michoud et al., 2017; Rinaldi et al.,*2017) and could be relevant at Lampedusa."

807

Moreover, it is known that the regional pollution at Lampedusa is higher than presented in your manuscript (e.g. Pace et al. (2006) have shown that clean marine aerosol conditions are rare at Lampedusa, contrary to your sentence p3L81) and a discussion on air masses trajectories is available in
 Mallet et al., 2019. So please add this information in your corrected version before publication.

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In agreement with this comment, the sentence at P3 L81 was reworked as follows: "Although the island of Lampedusa is a remote marine environment of the central Med Sea, influences from ship traffic emissions (Becagli et al., 2012, 2017), volcanic aerosols (Sellitto et al., 2017), forest fires (Pace et al., 2005), and regional pollution (Pace et al., 2006), have been documented".

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819 Reply to Referee 1

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(REFEREE)

The manuscript addresses the atmospheric deposition of organic matter in the Mediterranean, for which there is little data available. It quantifies such deposition in the small island of Lampedusa in the Central Mediterranean, in terms of carbon, nitrogen and phosphorus. It also tries to untangle possible sources of such organic matter. In this aspect the manuscript is less conclusive as there is no good relationship to aerosol origin or type of deposition. The conclusion is that the OM is mainly coming from sea spray that the different air masses pick up and transport to wind up depositing. It could be in large part but really it is just a hypothesis that needs further exploration.

830 (AUTHORS)

831 We really thank the reviewer for his/her appreciation of our work.

We totally agree that regarding the possible sources of organic matter, we just reported some hypotheses. With the available data-set in fact we were just able to make an hypothesis about the main sources of DOM in the different sampling periods. The sampling periods cover a generally wide time interval, and the deposition data are the result of integrating deposition over air masses of different origin and with different aerosol characteristics and loads. This makes very difficult to find a correlation between aerosol origin and DOC input, since the different sources are mixed in our samples.

If we would have reduced the sampling periods, we would have had less variability in the sources,
but in most of them we would have not had enough DOM to do all the analysis. Two weeks was therefore
the best compromise we were able to find.

We added in the conclusions a sentence highlighting that in order to understand the link betweenaerosol origin and DOM concentration and quality further studies are necessary.

843

Also, I was surprised not to consider wind direction properties when analyzing deposited material.
 Lampedusa is a small island but I would not be surprised that when wind blows from directions other than 31

due East, and especially when it blows over the island from the West, substantial OM could be picked up from the island itself.

848

849 We thank the reviewer for this comment. At first, we took into consideration the air masses 850 trajectories, but then we realized that there are some limitations in the use of wind direction to infer the 851 aerosol sources.

Firstly, as above reported, the sampling periods cover a generally wide time interval, and the deposition data are the result of integrating deposition over air masses of different origin and with different aerosol characteristics and loads. Thus, in some cases information on the wind might suggest what are the dominant (if existing) wind conditions during a specific period; but due to possible differences in aerosol amounts and deposition, these conditions may not be representative of the integrated samples.

Secondly, the use of wind direction and speed identify the air mass origin (which is not what the reviewer is suggesting) may be problematic, since trajectories arriving to Lampedusa may take different paths depending on the synoptic conditions. A more elaborated approach would be required (e.g., trajectory reconstruction using wind, as in Becagli et al., 2012; or modeled backward trajectories based on meteorological analyses, as in Marconi et al., 2014). However, also in this case, the relatively long duration of the sampling interval and the variability of the deposition would prevent a robust attribution of the source regions.

864 Wind measurements conversely, as correctly suggested by the reviewer, might potentially provide 865 useful information on the impact of local sources. The main local source area of anthropogenic particles is in the sectors between South and South-East of the sampling site; in this sectors there are the Lampedusa town, 866 867 the power plant, the airport and the port. Previous studies have shown that wind from these directions is 868 relatively infrequent, and the impact of these sectors is estimated to be negligible (see e.g., Artuso et al., 869 2009, with respect to atmospheric CO₂ measurements; Calzolai et al., 2015, with respect to PM10 870 measurements). This impact is expected to be even lower over samples integrated over are relatively large 871 number of days.

Due to these reasons, we have preferred to associate the deposition samples' characteristics with those of PM10 samples collected daily and over the same time period. This allows to use aerosol properties measured in the same time intervals and influenced by the same sources to infer some overall conditions. Previous studies (e.g., Becagli et al., 2012, 2013, 2017; Marconi et al., 2014; Calzolai et al, 2016) have been dedicated at linking the PM10 measured composition with different aerosol sources.

877

A third aspect of the manuscript deals with estimating the local and Mediterranean basin-wide importance of such deposition estimates for the biogeochemical functioning of the Mediterranean. I like this part myself but I have to admit it is the least elaborated since it is based on assumptions that will be hardly met. For instance, calculations based on the extension to the whole Mediterranean of the measured OM deposition at Lampedusa. Given it is so variable and without a clear reason, I would expect variability
to increase when other locations are taken into account. Also, the lability of the deposited organic matter
is an unknown, so the final role of the marine biota is also unknown. But anyhow, I like these exercises.
Thus, to me the main value of the manuscript is to provide a much needed data series of OM deposition
measurements.

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We totally agree with the reviewer and we are aware that reporting the calculations based on the extension of our results to the entire Med Sea is a risk, because it is nor simple, nor appropriate, to assume that what is observed at one location is valid for the entire basin. However, we consider this as a first conceptual exercise that uses the new results from our study to give an estimate of the implications of DOM deposition for marine ecosystem, that needs to be supported by additional data.

To the best of the author's knowledge only one paper reports that a not-well quantified fraction of atmospheric DOM can be recalcitrant. Due to the lack of information, we decided to discuss implications taking into consideration both the possibilities: DOM is labile and DOM is recalcitrant. In the revised manuscript we will better stress the need for further investigations about the biological lability of DOM coming from the atmosphere

898 In order to better stress that these calculations are a conceptual exercise, we added the following 899 sentences in the revised manuscript:

900 "A conceptual exercise can be made in order to give an estimate of the implications of DOM deposition for
901 marine ecosystem."

"Even if we are aware that these assumptions are hardly meet, in particular the estimate of DOC input to the
whole Med Sea, based on the data collected in Lampedusa, we think that these calculations can give an idea of the
relevant role that atmospheric input of DOC can have in sustaining bacterial productivity in the surface layer,
particularly when the upper water column is strongly stratified."

907 The methods are standard within the field and thus assure quality control. Maybe I am not clear 908 whether monthly data were calculated and how or whether just sample data was provided always? or in 909 what cases? That is, how where data treated when more than 1 sample per month was available? How was 910 the data split when covering periods from two consecutive months?, etc.

911

906

We apologize for the inaccuracy. We did not calculate monthly data, we reported the sample data. In
the figures 2, 3, 4, 6 and 7, the width of the bars refers to the duration of the sampling period (Table 1). This
aspect will be clarified in the revised manuscript.

As a general rule, samples were collected every ~15 days, or immediately after strong rain or dust
storm events. However, due to logistic problems the sampling period was longer than 20 days for 9
depositions (Table 1).

The DOC, DON and DOP fluxes, reported in the text and in the figures 2, 3, 4, 6 and 7 were 918 calculated using the following formula: 919 920 $X_{Flux} = X \cdot V / A \cdot d$ where X is the concentration of DOC, DON or DOP measured in the sample and expressed in µM; V 921 is the volume of rain collected by the sampler (expressed in L) or the volume of Milli-Q water used to wash 922 923 the funnel walls in case of dry deposition (250 ml); A is the area of the funnel (0.1018 m^2), and d refers to 924 the number of days of the sampling period. The DOC, DON and DOP fluxes are reported in the figures 925 considering the flux corresponding to each sampling period. A paragraph with this explanation was added in 926 the materials and methods section of the revised manuscript in order to clarify these calculations. 927 928 I understand that sample data is clearly reported in Fig. 5, but how were the rest treated is a bit 929 mysterious, especially since bars have unequal width within and between figures. 930 931 The bars in the figure 5 corresponded to the C:N:P molar ratios (see Table 3), so they referred to a 932 number, not a flux, and this is the reason why the width of the bars is always the same in figure 5, in contrast 933 to fig. 2, 3, 4, 6 and 7, where the width of the bars is different since it refers to the length of sampling periods. We added the description of the bars in the caption. 934 935 In line 150 it is also important to know the flow rate of the low-volume sampler. Also, I guess that 936 937 because of physical flow rate constraints a 1 µm filter could not be used. That would have been much 938 more desirable since there tend to be organic rich particles at the very fine particle ranges, and they would 939 have been missed, not a minor issue in this paper on OM. I would like the authors to comment on the 940 choice of a 2 µm filter to collect particles. 941 942 The filters used in this study are those usually used for aerosol sampling, they have a nominal 943 porosity of 2 µm, but they are certified for 99% efficiency for particles having 0.3 µm diameter. 944 The sampling flow is maintained constant at 2.3 m^3/h in order to maintain constant the sampling heads cut-off (10 μ m) as reported in the European rule UNI EN12341. 945 946 In order to clarify these concepts for a broad number of readers the text was changed as follows: 947 "PM10 (particulate matter with aerodynamic equivalent diameter lower than 10 μ m) is routinely sampled on 948 a daily basis at the island of Lampedusa (Becagli et al., 2013; Marconi et al., 2014; Calzolai et al., 2015) by using a 949 low-volume dual-channel sequential sampler (HYDRA FAI Instruments) equipped with two PM10 sampling heads, 950 operating at constant flow of 2.3 m3/h in accord with the European rules for aerosol monitoring (UNI EN12341). Aerosol is collected on 47 mm diameter Teflon filters (PALL Gelman) having 2 µm nominal porosity but certified to 951

952 have 99% retention efficiency for 0.3 µm diameter particles. The PM10 mass was determined by weighting the Teflon

filters before and after sampling with an analytical balance in controlled conditions of temperature (20±1 °C) and
relative humidity (50±5%)."

The manuscript is well structured and balanced. The title is informative of the contents. The language is proficient. Figures should be uniformed or clarified in aspects such as the x-axis but are otherwise well done. Other than that, I have no major concerns publishing the manuscript pretty much as it is.

We really thank the reviewer for his/her appreciation of our manuscript. We reworked the figures inorder to uniform them and to eliminate any misunderstanding.

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965 Reply to Referee 2

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967 (*REFEREE*)

968 This paper presents a sound dataset concerning the dry and wet deposition fluxes of dissolved 969 organic matter sampled for 2 years and a half at the island of Lampedusa (Italy). This site, in the central 970 Mediterranean, is appropriately taken to represent the interaction atmosphere-sea surface in a remote 971 marine environment. It is a well written paper which addresses a topic of interest: the role of DOM (and 972 its components DON and DOP) deposition in the western Mediterranean. It explores the role of the frequent Saharan intrusions, a very interesting point since few studies have dealt with the interactions 973 974 between organic carbon and Saharan dust. Finally, it specifically addresses the role of this atmospheric 975 deposition for marine productivity.

The quantification of N and P atmospheric deposition to the Mediterranean has been previously addressed in many papers, the most relevant of them are adequately cited by the authors. However, I'd like to bring to the authors attention the work of Izquierdo et al. 2012 in Atmospheric Environment. Atmospheric phosphorus deposition in a near-coastal rural site in the NE Iberian Peninsula and its role in marine productivity, since it will provide more data for comparison, discussion and understanding of the role of African sources in marine biogeochemistry, and the relative contribution of dry and wet deposition.

983 The layout of the paper and data treatment are OK, and I have only a few suggestions, which I list 984 below.

985

986 (AUTHORS)

987	We really thank the reviewer for his/her appreciation of our manuscript. In the revised version, all
988	the comments and suggestions were taken into consideration. In the revised manuscript, the work of
989	Izquierdo et al. (2012) was cited as suggested by the reviewer.
990	
991	Introduction
992	Lines 34-35. Industrial pollution can also be originated from North Africa as has been shown in
993	the work of Rodríguez et al (2011). Transport of desert dust mixed with North African industrial
994	pollutants in the subtropical Saharan Air Layer. Atmospheric Chemistry and Physics 11, 6663–6685. I
995	think it is worth considering.
996	
997	In the introduction of the revised manuscript, we added this reference and a sentence about the
998	possible contribution of pollution from North Africa: "Industrial pollution can also be originated from the North
999	Africa as shown in the work by Rodríguez et al. (2011)."
1000	
1001	Line 40. The work of Izquierdo et al 2012 could be included in this list of references, since it deals
1002	with how P dep influences the marine biogeochemical cycle in the western Med Sea.
1003	
1004	In the introduction of the revised manuscript, we added this reference.
1005	
1006	Line 45. This sentence should be revised as it is not true that atmospheric deposition affects
1007	radiative forcing and human health. Aerosols in the atmosphere do, but not deposition.
1008	
1009	We agree with the reviewer that the sentence was not clear. In the revised manuscript, we changed it
1010	as follows: "Atmospheric deposition of organic carbon can therefore affect regional C cycling (Yan and
1011	Kim, 2012; Decina et al., 2018)."
1012	
1013	
1014	
1015	Material and methods
1016	I recommend to make some reorganization of the text, since some paragraphs in this section in
1017	jact correspond better to the introduction. E.g. the paragraph aealing with the explanation of the
1010	internation seawater DOM stoicniometry compared to the world oceans (lines /0-/5) should be moved to the Interduction
1020	to the Introduction.
1020	We apployize for this inaccuracy in the rayised manuscript, this sentences was delated
1021	we apprograe for this maccuracy, in the revised manuscript, this sentences was deleted.
1022	

1023	Same thing with the paragraph justifying the appropriateness of Lampedusa as representing an
1024	unpolluted site in the central Med.
1025	
1026	In agreement with this comment, in the revised manuscript the paragraph 2.1 was changed, reducing
1027	the excessive descriptive part.
1028	
1029	Line 78. Revise the notation of units of mean dust deposition
1030	OK.
1031	
1032	Line 95. polycarbonate, not in capital letter Paragraph
1033	OK.
1034	
1035	104-108. Please list in this text the ions and metals analyzed
1036	We listed the ions and metals in section 2.5 and we deleted this part since it is also reported at lines
1037	135-140, as noted by the reviewer.
1038	
1039	Line 105 and 136. blank levels, instead of blanks level
1040	OK.
1041	
1042	Lines 135-140. This has been already exposed in lines 104-108.
1043	
1044	As above reported, we deleted the lines 104-108 and we added here the list of metals and ions.
1045	
1046	Line 144. I see that the particulates retained in the filters (after wet and dry deposition filtration)
1047	was analysed. But the procedure of digestion and analysis is not reported. Same thing for particulates
1048	from the PM10 samples (line157). This should be described in the M&M.
1049	
1050	In agreement with this comment, in the revised manuscript, the description of the procedure were
1051	added in the Material and Methods section.
1052	
1053	
1054	Results
1055	Line 202. Here there is an error, since the upper limit of TDP is 5*10 exp-3 (as deduced from
1056	Table 2).
1057	
1058	We apologize for the inaccuracy. The mistake was corrected in the revised manuscript.

- 1060 *Line 244. Error in unit: 8.8 ug m-3*
- 1061 1062

1063

We apologize for the inaccuracy. The mistake was corrected in the revised manuscript.

1064 Discussion

1065 In this section I'd like a more in deep discussion of dry versus wet deposition and its relation to 1066 meteorology.

1067

1068 In agreement with this comment, in the revised manuscript, we added some information on dry 1069 versus wet deposition and its relation to meteorology. In particular, in the results we reported the annual 1070 rainfall during 2016 and in the description of the study area we added the following sentence: "Precipitation 1071 shows a significant interannual variability and is concentrated in autumn and winter, with a maximum in 1072 October. Intense precipitation events, which are relatively infrequent, are generally associated with frontal 1073 passages and winds from the Northern sectors. Very dry conditions characterize late spring and summer." In 1074 the discussion (paragraph 4.2), we added the following sentence : "All of the analyzed samples, except few 1075 cases in summer 2016, are relative to dry+wet conditions. Although the DON and DOP recorded during the 1076 dry samplings are generally on the low end side of the measured range (see Table 2), no information on the 1077 role played by wet and dry deposition processes may be drawn at this stage, due to the limited number of dry 1078 samples." Regarding DOC input, in the discussion (paragraph 4.1) we reported that "Finally, the correlation 1079 between monthly precipitation rates and DOC fluxes shows the importance of rain events as a source of 1080 DOC in the Med Sea, as proposed by Djaoudi et al. (2018)."

In the literature, the wet atmospheric deposition is considered the main pathway for the removal of organic carbon from the atmosphere. While our results that dry deposition is also important and we have reported a detailed discussion of this point in the paragraph 4.3: "Some models have estimated that wet deposition represents up to 75-95% of total deposition (Iavorivska et al., 2016). Our data While confirm the importance of wet deposition, it also stress the relevance of dry deposition (32% of the total deposition during the entire sampling period) that appears to be the main contributor of DOC and of other chemical species to the remote site of Lampedusa, as suggested by Morales-Baquero et al. (2013)."

1088

High DOC deposition was recorded in Lmp25 (May 2016) and also in Lmp1 (end of March 2015) and Lmp 4 (May2015) coinciding with Saharan dust but low DOC was found in Saharan events during autumn and winter. In view of this clear seasonal differentiation, one could hypothesize that there is a role of pollen attached to desert particles in these spring events (end March-May) and this pollen would contribute DOC. This process would not occur in the other seasons (winter and autumn of no pollen production). This is a possible explanation that needs further attention. However, there are some reports in the literature of joint pollen and dust transport: for example, Van Campo and Quet (1982) identified
pollen types transported from North Africa to south France together with mineral desert dust, Franzen et
al. 1994 documented the arrival of pollen from the Mediterranean to Fennoscandia during a dust event.
Pollen originating in Morocco was detected South Spain (Cabezudo et al. 1997) and various pollen types
(Cannabis, Cupressus, Pinus, Platanus and Sambucus) were observed in Cordoba (South Spain)
exclusively during dust African events (Cariñanos et al. 2004).

1102 We really thank the reviewer for this interesting suggestion. The contribution of pollen to atmospheric DOC in spring is an interesting hypothesis to test. In the revised manuscript, the following 1103 1104 sentence was added in the discussion (paragraph 4.3): "In addition Lmp01 (end of March 2015), Lmp04 (May 2015) and Lmp25 (May 2016) show a seasonality that could be linked to the transport of pollen 1105 1106 attached to desert particles in the spring events, and this pollen would contribute to atmospheric DOC input 1107 in spring (end of March- May). Pollen originating in Morocco was detected in South Spain (Cabezudo et al., 1108 1997) and various pollen types (Cannabis, Cupressus, Pinus, Platanus and Sambucus) were observed in 1109 Cordoba (South Spain) exclusively during dust African events (Cariñanos et al., 2004). This process would not occur in the other seasons (winter and autumn), when no pollen production occurs." A sentence about 1110 the need of further studies was added in the conclusions. 1111

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Figures In fig 2, 3 and 4, include a legend to indicate the color of wet and dry deposition.

1115 In the revised manuscript, a legend was included in figures 2, 3 and 4.