Methane paradox in tropical lakes ? Sedimentary fluxes rather than water column production in oxic waters sustain methanotrophy and emissions to the atmosphere

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Supplementary table and figures.

Table S1. Summary of major limnological characteristics of the sampled lakes. Numbers between brackets are the mean depth; numbers followed by an asterisk are the average value for two sampling cruises. Zm is the mixed layer depth and Ze is the euphotic zone depth.

Lake	Mixing regime	Surface area	Max depth	Zm (m)	Ze (m)	Chlorophyll a
		(km^2)	(m)			(mg m^{-3})
Edward	Monomictic	2300	117 (34)	15*	7.7*	4.8*
Kyamwinga	Monomictic	2.6	40	5	3.1	6.0
Nyamusingere	Polymictic	3.84	3	0.5	1.1	88.3
Katinda	Monomictic	0.44	20	1*	1.2*	77.2*
George	Polymictic	250	7 (3)	0.15*	1.1*	149.7*



Figure S1. Relationships between phytoplankton primary production and N₂ fixation. Linear relationship between (a) the chlorophyll *a* (μ g chla L⁻¹) and particulate organic carbon (POC, mmol L⁻¹) concentrations, (b) chlorophyll a and maximum photosynthetic activity (Pmax, mmol C L⁻¹ h⁻¹), and (c) Pmax and the maximum N₂ fixation (N₂max, nmol N₂ L⁻¹ h⁻¹) rates.



Figure S2. Phytoplankton community composition. Phytoplankton community composition in the surface waters the African lakes sampled.



Figure S3. *mcrA* **gene abundance.** *mcrA* gene copy concentration (a) and *mcrA* gene copy abundance normalized on the DNA concentration (b) in each African lake sampled. Green and brown bars represent the > 5 μ m and < 5 μ m fraction of the seston, respectively.



Figure S4. Prokaryote community composition. Contribution of methanogens to the prokaryote community composition in the African lake sampled.



Figure S5. Littoral pelagic - gradient in L. Edward. Relationship between the maximum depth of the station and (a) the CH₄ concentration (nmol L⁻¹) and (b) the stable carbon isotope composition of CH₄ (δ^{13} C-CH₄, ∞) in surface waters (0.3 m), in L. Edward. Data were collected between 18/01/2018 and the 24/01/2018. (c) Relationship between the CH₄ concentration and the δ^{13} C-CH₄ in surface waters of L. Edward. The black curve represents the fitted (r^2 0.69, n =11) apparent carbon isotope fractionation factor during CH₄ oxidation (α) calculated as in Morana et al. (2015). Grey zone represents the expected δ^{13} C-CH₄ at a given concentration considering a range of α value comprised between 1.009 and 1.023 and a littoral (2.5 m max depth) surface water δ^{13} C-CH₄ source of -43.6‰ (red symbol). The apparent C isotope fractionation ($\alpha = 1.012$) in L. Edward surface waters was close to the isotope fractionation factor previously determined experimentally for CH₄ oxidation in L. Kivu ($\alpha = 1.016$), another large East African lake (Morana et al. 2015).



Figure S6. ¹³C labelling experiment kinetics. Upper panel: In well-oxygenated surface waters of Lake Edward, evolution of the δ^{13} C-CH₄ during the incubation in bottles amended with 1 ml of NaH¹³CO₃ (leftmost), ¹³C_{methyl}-methionine, ¹³C_(1,2)-acetate, or water (control treatment, rightmost). Note the difference in the amount of excess ¹³C between the NaH¹³CO₃ and the ¹³C_{methyl}-methionine and ¹³C_(1,2)-acetate treatment. Bottom panel: Amount of CH₄ produced from the corresponding 13C-tracer calculated at each time step as described in the material & methods section. Green, grey and black symbols represent bottles incubated under light, light and DCMU, and darkness, respectively. CH4 concentration & ¹³C kinetics showed similar patterns in L. George, L. Katinda and L. Nyamusingere (not showed).



Figure S7. **Ebullition in L. George**. Picture showing stable bubbles trapped in the organic-rich surface waters of L. George a day of calm weather (28/01/18, sampling at noon).