1	Impact of dust addition on Mediterranean plankton
2	communities under present and future conditions of pH and
3	temperature: an experimental overview
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20	Keywords: Mediterranean Sea; Atmospheric deposition; Plankton community; Ocean
21	acidification; Ocean warming

22 Abstract

23 In Low Nutrient Low Chlorophyll areas, such as the Mediterranean Sea, atmospheric 24 fluxes represent a considerable external source of nutrients likely supporting primary production 25 especially during stratification periods. These areas are expected to expand in the future due to 26 lower nutrient supply from sub-surface waters caused by climate-driven enhanced stratification, 27 likely further increasing the role of atmospheric deposition as a source of new nutrients to 28 surface waters. Yet, whether plankton communities will react differently to dust deposition in a 29 warmer and acidified environment remains an open question. The potential impact of dust 30 deposition both in present and future elimate conditions was investigated through three 31 perturbation experiments in the open Mediterranean Sea. Climate reactors (300 L) were filled 32 with surface water collected in the Tyrrhenian Sea, Ionian Sea and in the Algerian basin during a 33 cruise conducted in May/June 2017 in the frame of the PEACETIME project. The experimental 34 protocol comprised two unmodified control tanks, two tanks enriched with a Saharan dust analog 35 and two tanks enriched with the dust analog and maintained under warmer (+3 °C) and acidified 36 (-0.3 pH unit) conditions. Samples for the analysis of an extensive number of biogeochemical 37 parameters and processes were taken over the duration of the experiments (3-4 d). Here, we 38 present the general setup of the experiments and the impacts of dust seeding with and without 39 addressing the effects of environmental changes on nutrients and biological stocks. Dust addition 40 led to a rapid and maximum input of nitrate whereas phosphate release from the dust analog was 41 much smaller, Our results showed that the impacts of Saharan dust deposition in three different 42 basins of the open Northwestern Mediterranean Sea are at least as strong as those observed 43 previously in coastal waters. However, interestingly, the effects of dust deposition on biological 44 stocks were highly different between the three investigated stations and could not be attributed to 45 differences in their degree of oligotrophy but rather to the initial metabolic state of the 46 community. Ocean acidification and warming did not drastically modify the composition of the

- 47 autotrophic assemblage with all groups positively impacted by warming and acidification.
- 48 Although autotrophic biomass was more positively impacted than heterotrophic biomass under
- 49 future environmental conditions, a stronger impact of warming and acidification on
- 50 mineralization processes suggests a decreased capacity of Mediterranean surface plankton
- 51 communities to sequester atmospheric CO₂ following the deposition of atmospheric particles.

52 **1. Introduction**

53 Atmospheric deposition is well recognized as a significant source of micro- and macro-54 nutrients for surface waters of the global ocean (Duce et al., 1991; Jickells et al., 2005; Moore et al., 2013). The potential modulation of the biological carbon pump efficiency and the associated 55 56 export of carbon by atmospheric deposition events are still poorly understood and quantified 57 (Law et al., 2013). This is especially true for Low Nutrient Low Chlorophyll (LNLC) areas 58 where atmospheric fluxes can play a considerable role in nutrient cycling and that represent 60%59 of the global ocean surface area (Longhurst et al., 1995) as well as 50% of global carbon export 60 (Emerson et al., 1997). These regions are characterized by-a low availability of macronutrients 61 (N, P) and/or metal micronutrients (e.g. Fe) that can severely limit or co-limit phytoplankton 62 growth during large periods of year. 63 The Mediterranean Sea is a typical example of these LNLC regions and exhibits surface

64 chlorophyll a concentrations below 0.2 μ g L⁻¹ all year round over most of its area, except in the 65 Ligurian Sea where relatively large blooms can be observed in late winter-early spring (e.g. Mayot et al., 2016). Recent assessments showed that the atmospheric input of nutrients in the 66 67 Mediterranean Sea is of the same order of magnitude as riverine inputs (Powley et al., 2017), 68 making the atmosphere a considerable external source of nutrients (Richon et al., 2018). 69 Atmospheric deposition originates both from natural (mainly Saharan dust) and anthropogenic 70 sources (e.g. Bergametti et al., 1989; Desboeufs et al., 2018). Dust deposition, mostly in the form 71 of pulsed inputs, is mainly associated with wet deposition (Loÿe-Pilot and Martin, 1996). Ternon 72 et al. (2010) reported an average annual dust flux over four years of 11.4 g m⁻² yr⁻¹ (average 73 during the period 2003–2007) at the DYFAMED station in the Northwestern Mediterranean Sea. In this region, the most important events reported in the $\frac{2010}{2}$ decade amounted to ~ 22 g m⁻² 74 75 (Bonnet and Guieu, 2006; Guieu et al., 2010b).

Atmospheric deposition provides new nutrients to the surface waters (Guieu et al., 2010b; Kouvarakis et al., 2001; Markaki et al., 2003; Ridame and Guieu, 2002), Fe (Bonnet and Guieu, 2006) and other trace metals(Desboeufs et al., 2018; Guieu et al., 2010b; Theodosi et al., 2010), that represent significant inputs likely supporting the primary production especially during the stratification period (Bonnet et al., 2005; Ridame and Guieu, 2002), although no elear correlation between dust and ocean color could be evidenced from long series of satellite observation in that part of the basin (Guieu and Ridame, 2020).

83 Experimental approaches have shown that wet dust deposition events in the Northwestern 84 Mediterranean Sea (the dominant deposition mode in that basin) present a higher impact as a 85 source of bioavailable fertilizing nutrients compared to dry deposition. Indeed, wet deposition 86 provides both new N and P while dry deposition supplies only P and does not allow to stimulate 87 the autotrophic community (except diazotrophs; Ridame et al., 2013), resulting in no increase in 88 chlorophyll a concentrations and primary production (Guieu et al., 2014a). This so-called 89 fertilizing effect has been experimentally shown using both micro- and mesocosms where the 90 wet deposition of Saharan dust analog strongly stimulated primary production and phytoplankton 91 biomass (Guieu et al., 2014a; Ridame et al., 2014) while also modifying phytoplankton diversity 92 (Giovagnetti et al., 2013; Lekunberri et al., 2010; Romero et al., 2011). In addition, besides 93 phytoplankton, dust deposition also modified also the bacterial community assemblage and led to even stronger enhancements of production and/or respiration rates (Pulido-Villena et al., 2014). 94 95 The carbon budget established from four artificial seeding experiments during the DUNE project 96 (Guieu et al., 2014a) showed that by stimulating predominantly heterotrophic bacteria, 97 atmospheric dust deposition can enhance the heterotrophic biological behavior of these 98 oligotrophic waters. This has the potential to reduce the fraction of organic carbon that can be 99 exported to deep waters during the winter mixing period (Pulido-Villena et al., 2008) and 100 ultimately limit net atmospheric CO₂ drawdown.

101	Another effect induced by Saharan dust deposition is the export of particulate organic
102	earbon (POC), as lithogenic particles can aggregate and ballast dissolved organic matter (Bressac
103	et al., 2014; Desboeufs et al., 2014; Louis et al., 2017a; Ternon et al., 2010). This so-called
104	lithogenic carbon pump can represent a major part of the carbon export following a dust
105	deposition event (up to 50% during the DUNE experiment; Bressac et al., 2014). Recently, Louis
106	et al. (2017a) showed that Saharan dust deposition triggers the abiotic formation of transparent
107	exopolymeric particles (TEP), leading to the formation of organic-mineral aggregates, a
108	formation process that is highly dependent on the quality and quantity of TEP-precursors initially
109	present in seawater.

110 In response to ocean warming and increased stratification, nutrient cycling in the open 111 ocean is being and will continue to be perturbed in the next decades resulting very likely in regionally variable impacts (IPCC, 2019). Overall, LNLC areas are expected to expand in the 112 113 future (Irwin and Oliver, 2009; Polovina et al., 2008) due to a-thermal stratification related 114 reduction of nutrients supply from sub-surface waters (Behrenfeld et al., 2006). As such, the role 115 of atmospheric deposition might increase as an alternative source of new nutrients to surface 116 waters. Ongoing warming and acidification of the global ocean (IPCC, 2019) are also evidenced 117 in the Mediterranean Sea (e.g. Kapsenberg et al., 2017; The Mermex group, 2011). Whether or 118 not plankton communities will respond differently to dust deposition in future conditions is still 119 largely unknown. Although dependent on resource availability, it is well known that 120 remineralisation by bacteria is subject to positive temperature control (López-Urrutia and Morán, 121 2007). As under severe nutrient limitation, warming has no effect on primary productivity 122 (Marañón et al., 2018), it will most likely further push the balance towards net heterotrophy in oligotrophic areas. 123

With respect to ocean acidification, an *in situ* mesocosm experiment conducted during
the summer stratified period in the Northwestern Mediterranean Sea showed that the plankton

126	community was rather insensitive to this perturbation	under strong nutrient limitation

127 (Maugendre et al., 2017, and references therein). This is coherent with results from Maugendre et 128 al. (2015), based on a batch experiment, showing that, under nutrient-depleted conditions in late 129 winter, ocean acidification has a very limited impact on the plankton community and that small 130 species (e.g. Cyanobacteria) might benefit from warming with a potential decrease of the export 131 and energy transfer to higher trophic levels. In contrast, in more eutrophic coastal conditions, 132 Sala et al. (2016) showed that ocean acidification exerted a positive effect on phytoplankton, 133 especially on pico- and nano-phytoplankton. Similarly, Neale et al. (2014) showed in a coastal 134 ecosystem of the Alboran Sea that ocean acidification could lead, although moderately, to high chlorophyll levels under low light conditions with an opposite effect under high irradiance. 135 To date and to the best of our knowledge, there have been no attempts to evaluate the 136 137 behavior of plankton communities after the deposition of atmospheric particles in the context of 138 future levels of temperature and pH. Yet, following the recommendation from Maugendre et al. 139 (2017), any perturbation experiment for future elimate conditions in the Mediterranean Sea 140 should consider atmospheric deposition as a source of new nutrients and consider both 141 temperature and pH as external foreings. Such experiments were conducted in the frame of the 142 PEACETIME project (ProcEss studies at the Air-sEa Interface after dust deposition in the 143 MEditerranean sea; http://peacetime-project.org/) during the cruise on board the R/V "Pourquoi 144 Pas?" in May/June 2017. The project aimed at extensively-studying and parameterizing the chain 145 of processes occurring in the Mediterranean Sea after atmospheric deposition and to put them in 146 perspective of on-going environmental changes (Guieu et al., 2020). During that cruise, three 147 perturbation experiments were conducted in climate reactors (300 L tanks) filled with surface 148 water collected in the Tyrrhenian Sea (TYR), Ionian Sea (ION) and in the Algerian basin (FAST; 149 Fig. 1). Six tanks were used to follow simultaneously and with a high temporal resolution, the evolution of biological activity and stocks, nutrients stocks, dissolved organie matter as well as 150

151	particles dynamics and export, following a dust deposition event simulated at their surface, both
152	under present environmental conditions and following a realistic climate change scenario for
153	2100 (ca. +3 °C and -0.3 pH units; IPCC, 2013). In this manuscript, we will present the general
154	setup of the experiments and the evolution of nutrient and biological stocks (heterotrophic and
155	autotrophic prokaryotes, photosynthetic eukaryotes as well as micro- and meso-zooplankton).
156	Several other manuscripts, related to these experiments and currently submitted to or published
157	in this special issue, focus on plankton metabolism (primary production, heterotrophic
158	prokaryote production) and carbon export (Gazeau et al., 2021), on the microbial food web
159	(Dinasquet et al., 2021), on-nitrogen fixation (Céline Ridame, unpublished results) and on the
160	release of insoluble elements (Fe, Al, REE, Th, Pa) from dust (Roy-Barman et al., 2020).

161 **2. Material and Methods**

162 **2.1. General setup**

163 Six experimental tanks (300 L; Fig. 2), in which the irradiance spectrum and intensity can 164 be finely controlled and in which future ocean acidification and warming conditions can be fully reproduced, were installed in a temperature-controlled container. The tanks are made of high-165 166 density polyethylene (HDPE) and are trace-metal free in order to avoid contaminations, with a height of 1.09 m, a diameter of 0.68 m, a surface area of 0.36 m² and a volume of 0.28 m³. All 167 168 tanks were cleaned before the experimental work following the protocol described by Bressae 169 and Guieu (2013). A weak turbulence was generated by a rotating PVC blade (9 rpm) in order to mimic natural conditions. Each tank was equipped with a lid containing six rows of LEDs 170 171 (Alpheus[©]). Each of these rows were composed of blue, green, cyan and white units in order to 172 mimic the natural sun spectrum. At the conical base of each tank, a polyethylene (PE) bottle eollecting the exported material from above was screwed onto a polyvinyl chloride (PVC) valve 173 174 that remained open during the duration of the whole experiment. Photosynthetically active 175 radiation (PAR; 400-700 nm) and temperature were continuously monitored in each tank using 176 respectively QSL-2100 Scalar PAR Irradiance Sensors (Biospherical Instruments[©]) and pt1000 177 temperature sensors (Metrohm[©]) connected to a D230 datalogger (Consort[©]).

The experimental protocol comprised two unmodified control tanks (C1 and C2), two tanks enriched with Saharan dust (D1 and D2) and two tanks enriched with Saharan dust and maintained under warmer (+3 °C) and acidified (-0.3 pH unit) conditions (G1 and G2). The atmosphere above tanks C1, C2, D1 and D2 was flushed with ambient air (ca. 400 ppm, 6 L min⁻¹) 182 ¹) and tanks G1 and G2 were flushed with air enriched with CO₂ (ca. 1000 ppm, 6 L min⁻¹) in order to prevent CO₂ degassing from the acidified tanks. CO₂ partial pressure (*p*CO₂) in both 184 ambient air and CO₂-enriched air was monitored using two gas analysers (LI-820, LICOR©).

185 The CO₂ concentration in the CO₂-enriched air was manually controlled through small injections

186 of pure CO₂ (Air Liquide[©]) using a mass flow controller.

187 Three experiments were performed at the long duration stations TYR, ION and FAST. 188 The tanks were filled by means of a large peristaltic pump (Verder© VF40 with EPDM hose, flow of 1200 L h⁻¹) collecting seawater below the base of the boat (depth of -- 5 m), used to 189 190 supply continuously surface seawater to a series of instruments during the entire campaign. In 191 order to homogeneously fill the tanks, the flow was divided into six HDPE pipes distributing the water simultaneously into the different tanks. Overall, the filling of the six tanks took -2 h 192 193 (including rinsing and initial sampling, see thereafter). At the three stations, tanks were always 194 filled at the end of the day before the start of the experiments: TYR (17/05/2017), ION 195 (25/05/2017) and FAST (02/06/2017). While filling the tanks, this surface seawater was sampled 196 for the measurements of selected parameters (sampling time = t-12h before dust seeding, see 197 Table 1). After filling the tanks, seawater was slowly warmed using 500 W heaters, controlled by 198 temperature-regulation units (COREMA©), in G1 and G2 overnight to reach an offset of +3 °C. 199 ¹³C-bicarbonate was added to all tanks at 4:00 am (local time; Gazeau et al., 2021) and G1 and 200 G2 were acidified by addition of CO₂-saturated filtered (0.2 µm) seawater (~1.5 L in 300 L; 201 collected when filling the tanks at each station) at 4:30 am to reach a pH offset of -0.3. Sampling 202 for most parameters took place prior to dust seeding (sampling time = t0, see Table 1). Dust 203 seeding was performed right after to between 7:00 and 9:00 (local time) in tanks D1, D2, G1 and 204 G2. The same dust analog was used and the same dust flux was simulated as for the DUNE 2009 205 experiments described in Desboeufs et al. (2014). Briefly, the fine fraction (< 20 µm) of Saharan 206 soils collected in southern Tunisia, which is a major source of dust deposition over the 207 northwestern Mediterranean basin, was used in the seeding experiments. The particle size distribution showed that 99% of particles had a size smaller than 0.1 µm, and that particles were 208

209	mostly made of quartz (40%), calcite (30%) and clay (25%; Desboeufs et al., 2014). This
210	collected dust underwent an artificial chemical aging process by addition of nitric and sulfuric
211	acid (HNO3 and H2SO4, respectively) to mimic cloud processes during atmospheric transport of
212	aerosol with anthropogenic acid gases (Guieu et al., 2010a, and references therein). To mimic a
213	realistic wet flux event of 10 g m ⁻² , 3.6 g of this analog dust were quickly diluted into 2 L of
214	ultrahigh-purity water (UHP water; 18.2 M Ω cm ⁻¹ resistivity), and sprayed at the surface of the
215	tanks using an all-plastic garden sprayer (duration = 30 min). The N and P total contents in the
216	dust were $1.36 \pm 0.09\%$ of N and $0.055 \pm 0.003\%$ of P (see Desboeufs et al., 2014, for a full
217	description of dust chemical composition). The experimental protocol included the analysis of an
218	extensive number of biogeochemical parameters and processes, not all shown and discussed in
219	this paper, and are listed in Table 1. The experiment at stations TYR and ION lasted 72 h (3
220	days) whereas the last experiment at station FAST was extended to four days. This relatively
221	short duration of the experiments was constrained by the time available between stations and the
222	time needed to properly clean the tanks between the experiments, following the protocol
223	described by Bressae and Guieu (2013). As a larger time window was possible at the end of the
224	eruise, the experiment at FAST was extended to four days. Seawater sampling was conducted 1
225	h (t1h), 6 h (t6h), 12 h (t12h), 24 h (t24h), 48 h (t48h) and 72 h (t72h) (+-96 h = t96h for station
226	FAST) after dust addition. Acid-washed silicone tubes were used for transferring the water
227	collected from the tanks to the different vials or containers. For some parameters (e.g. micro- and
228	macro-nutrients), sampled seawater was directly filtered at the exit of the sampling tubes
229	connected to each tank on sterile membrane filter capsules (gravity filtration with Sartobran©
230	300; 0.2 μm).

231 **2.2. Analytical methods**

232 2.2.1. Carbonate chemistry

233 Seawater samples for pH measurements were stored in 300 mL glass bottles with a glass 234 stopper, pending analysis on board (within 2 h). Samples were transferred to 30 mL quartz cells and absorbances at 434, 578 and 730 nm were measured at 25 °C on an Cary60 UV-235 236 Spectrophotometer (Agilent[©]) before and after addition of 50 µL of purified meta-cresol purple provided by Robert H. Byrne (University of South Florida, USA) following the method 237 described by Dickson et al. (2007). pH on the total scale (pH_T) was computed using the formula 238 239 and constants of Liu et al. (2011). The accuracy of pH measurements was estimated to 0.007 pH units, using a TRIS buffer solution (salinity 35, provided by Andrew Dickson, Scripps 240 241 university, USA).

242 Seawater samples for total alkalinity (A_T ; 500 mL) measurements were filtered on GF/F 243 membranes and analyzed onboard within one day. $A_{\rm T}$ was determined potentiometrically using a Metrohm© titrator (Titrando 888) and a glass electrode (Metrohm©, ecotrode plus) calibrated 244 using first NBS buffers (pH 4.0 and pH 7.0, to check that the slope was Nernstian) and then 245 246 using a TRIS buffer solution (salinity 35, provided by Andrew Dickson, Scripps university, 247 USA). Triplicate titrations were performed on 50 mL sub-samples at 25 °C and A_T was calculated 248 as described by Dickson et al. (2007). Titrations of standard seawater provided by Andrew Dickson (Scripps university, USA; batch 151) yielded A_T values within 5 µmol kg⁻¹ of the 249 nominal value (standard deviation = 1.5μ mol kg⁻¹; n = 40). 250

All parameters of the carbonate chemistry were determined from pH_T , A_T , temperature, salinity, as well as phosphate and silicate concentrations using the R package seacarb.

253 Propagation of errors on computed parameters was performed using the new function "error" of

this package, considering errors associated with the estimation of $A_{\rm T}$, pH_T as well as errors on dissociation constants (Orr et al., 2018).

256 **2.2.2. Nutrients**

257	Seawater samples for dissolved nutrients were filtered directly at the exit of the sampling
258	tubes connected to each tank (Sartobran [©] 300; 0.2 μ m), collected in polyethylene bottles and
259	immediately analyzed on board. Nitrate + nitrite (NO _x) and silicate (Si(OH) ₄) measurements
260	were conducted using a segmented flow analyzer (AAIII HR Seal Analytical©) according to
261	Aminot and Kérouel (2007) with a limit of quantification of 0.05 μ mol L ⁻¹ for NO _x and 0.08
262	μ mol L ⁻¹ for Si(OH) ₄ . In addition, for t-12h samples, the analysis of NO _x was also performed by
263	a spectrometric method in the visible at 540 nm, with a 1 m Liquid Waveguide Capillary Cell
264	(LWCC). The limit of detection was <mark>~10 nmol L⁻¹ and the reproducibility was ~6%. Also from</mark>
265	samples taken at t-12h, the measurement of ammonium concentrations was performed on board
266	using a Fluorimeter TD-700 (Turner Designs©) according to Holmes et al. (1999). This
267	fluorimetric method is based on the reaction of ammonia with orthophtaldialdehyde and sulfite
268	and has a limit of quantification of 0.01 μ mol L ⁻¹ . Dissolved inorganic phosphorus (DIP)
269	concentrations were quantified using the Liquid Waveguide Capillary Cell (LWCC) method
270	according to Pulido-Villena et al. (2010). The LWCC was 2.5 m long and the limit of detection
271	was 1 nmol L ⁻¹ .

272 **2.2.3. Pigments**

A volume of 2.5 L of sampled seawater was filtered onto GF/F filters, immediately frozen in liquid nitrogen and stored at -80 °C pending analysis at the SAPIGH analytical platform at the Institut de la Mer de Villefranche (IMEV, France). Filters were extracted at -20 °C in 3 mL methanol (100%) containing an internal standard (vitamin E acetate, Sigma©), disrupted by

277	sonication and clarified one hour later by vacuum filtration through GF/F filters. The extracts
278	were rapidly analyzed (within 24 h) on a complete Agilent© Technologies 1200 series HPLC
279	system. The pigments were separated and quantified as described in Ras et al. (2008).

280 **2.2.4. Flow cytometry**

281 For the enumeration of autotrophic prokaryotic and eukaryotic cells, heterotrophic 282 prokaryotes and heterotrophic nanoflagellates (HNF) by flow cytometry, subsamples (4.5 mL) 283 were fixed with glutaraldehyde grade I 25% (1% final concentration), and incubated for 30 min 284 at 4 °C, then quick-frozen in liquid nitrogen and stored at -80 °C until analysis. Samples were 285 thawed at room temperature. Counts were performed on a FACSCanto II flow cytometer (Becton 286 Dickinson[©]) equipped with 3 air-cooled lasers: blue (argon 488 nm), red (633 nm) and violet 287 (407 nm). The separation of different autotrophic populations was based on their scattering and fluorescence signals according to Marie et al. (2010), Synechococcus spp. was discriminated by 288 289 its strong orange fluorescence (585 ± 21 nm), and pico- and nano-eukaryotes were discriminated 290 by their scatter signals of red fluorescence (> 670 nm). For the enumeration of heterotrophic 291 prokaryotes, cells were stained with SYBR Green I (Invitrogen – Molecular Probes) at 0.025% 292 (vol / vol) final concentration for 15 min at room temperature in the dark. Stained prokaryotic 293 cells were discriminated and enumerated according to their right-angle light scatter (SSC) and 294 green fluorescence at 530/30 nm. In a plot of green versus red fluorescence, heterotrophic 295 prokaryotes were distinguished from autotrophic prokaryotes. For the enumeration of HNF, 296 staining was performed with SYBR Green I (Invitrogen-Molecular Probes) at 0.05% (v/v) final 297 concentration for 15-30 min at room temperature in the dark (Christaki et al., 2011). Cells were 298 discriminated and enumerated according to their SSC and green fluorescence at 530/30 nm. 299 Fluorescent beads (1.002 µm; Polysciences Europe[©]) were systematically added to each 300 analyzed sample as internal standard. The cell abundance was determined from the flow rate,

301 which was calculated with TruCount beads (BD biosciences[©]). Biomasses of each group were

302 estimated based on conversion equations and/or factors found in the literature (see section 2.3.2).

303 2.2.5. Micro-phytoplankton and -heterotrophs

At t-12h (i.e. seawater sampled during the filling of the tanks), a volume of 500 mL was sampled in glass vials and immediately preserved in a 5% acidic Lugol's solution pending analysis. At the Laboratoire d'Océanographie de Villefranche (LOV, France), 100 mL aliquots were transferred to sedimentation chambers (Utermohl) and counted under an inverted microscope at 200 to 400 magnifications.

309 2.2.6. Mesozooplankton

310 At the end of each experiment (t+72h for TYR and ION and t+96 h for FAST, after 311 artificial dust seeding), the sediment traps were removed, elosed and stored with formaldehyde 312 4% (see Gazeau et al., 2021). The valve at the base of the tanks was then reopened to let the 313 remaining water inside the tanks (TYR-165-180 L; ION = 172.5 L and FAST = 150 L) pass 314 through a large PVC sieve (100 μ m). The organisms retained on that mesh were gently removed 315 from the sieve using a washing bottle filled with filtered seawater (0.2 µm), and transferred 316 directly inside a 250 mL bottle. The bottle was filled with the sample (1/3 of the volume), and 317 was completed with formaldehyde 4%. The zooplankton digital images were obtained using a 318 ZooSCAN (Hydroptic[©]; Gorsky et al., 2010) at the PIQv-platform of EMBRC-France. The 319 identification of species was performed by automatic classification with a reference dataset in 320 EcoTaxa (https://ecotaxa.obs-vlfr.fr/, last access: 17/04/2020) and then all validated and 321 eorrected manually.

322 **2.3. Data analyses**

323 **2.3.1. Nutrient inputs from dust**

The maximum percentage of dust-born dissolved N and P was calculated considering that these evapo-condensated dust contain $1.36 \pm 0.09\%$ of N and $0.055 \pm 0.003\%$ of P (Desbocufs et al., 2014). Based on maximal concentrations observed in the D and G tanks after seeding (two discrete sampling within 6 h following dust seeding, t1h and t6h), one can estimate the maximal % of dissolution of dust in seawater during the three experiments;

$$329 \quad \%_{dissolution} = \frac{CONC_{max} - CONC_{init}}{CONC_{dust}}.100$$
(1)

where $\text{CONC}_{\text{init}}$ is the concentration of the corresponding nutrient in each tank before seeding (t0), CONC_{max} corresponds to the concentration of the corresponding nutrient in each tank when nutrient concentration was at a maximum over the first 6 h after seeding as observed based on our discrete sampling procedure, and $\text{CONC}_{\text{dust}}$ is the maximum addition, corresponding to a 100% dissolution of its total concentration in the dust analog (as estimated based on dust ehemical composition; Desboeufs et al., 2014; see above).

2.3.2. Autotrophic and heterotrophic biomass

As micro-phytoplankton eounting was not performed throughout the experiment, as a first approximation, autotrophic biomass was ealeulated as the sum of earbon contained in *Synechococcus*, pico-eukaryotes and nano-eukaryotes (abundances-based on flow cytometry) and is therefore restricted to the fraction $< 20 \ \mu m$. For *Synechococcus*, conversion to carbon units was done considering 250 fg C cell⁻¹ (Kana and Glibert, 1987), while the equation proposed by Verity et al. (1992; 0.433 BV^{0.863} where BV refers to the biovolume)-was used for pico- and nano-eukaryotes assuming a spherical shape and a diameter of 2 and 6 μm for the two groups,

- 344 respectively. Percentages of these different groups were calculated in order to estimate the
- 345 composition of the communities at the start and its evolution during the experiments.
- 346 Furthermore, heterotrophic biomass was computed as the sum of heterotrophic prokaryotes
- 347 biomass and heterotrophic nanoflagellates biomass. For heterotrophic prokaryotes, conversion to
- 348 carbon units were done considering 20 fg C cell⁻¹ (Lee and Fuhrman, 1987) and for heterotrophie
- 349 nanoflagellates assuming 220 fg C μm⁻³ (Børsheim and Bratbak, 1987), a spherical shape and a
- 350 diameter of 3 µm. The ratio of autotrophic and heterotrophic biomass during the experiments was
- 351 used to evaluate the trophic status of the investigated communities and its evolution. Finally, a
- 352 proxy for micro-phytoplankton biomass (B_{micro}) was estimated following Vidussi et al. (2001), as
- 353 the sum of Fucoxanthin and Peridinin.

354 3. Results

355 **3.1. Initial conditions**

356 Initial conditions of various measured parameters at the three sampling stations while 357 filling the tanks (t-12h before seeding) are shown in Table 2. pH_T and total alkalinity 358 concentrations followed a west to east increasing gradient (8.03, 8.04 and 8.07; 2443, 2529 and 359 2627 µmol kg⁻¹ at FAST, TYR and ION, respectively), NO_x concentrations were maximal at 360 station FAST with a NOx:DIP molar ratio of ~ 4.6. Very low NOx concentrations were observed at stations TYR and ION (14 and 18 nmol L⁻¹, respectively). DIP concentrations were the highest 361 at station TYR (17 nmol L⁻¹) and the lowest at the most eastern station (ION, 7 nmol L⁻¹). 362 363 Consequently, the lowest NOx: DIP ratio was measured at TYR (0.8), compared to ION and FAST (2.8 and 4.6, respectively). Ammonium concentrations were maximal at TYR (0.045 µmol 364 365 L^{-1}), intermediate at ION (0.022 µmol L^{-1}), and minimal at FAST (below detection limit). 366 Silicate concentrations were similar at stations TYR and ION (~ 1 μ mol L⁻¹) and higher than at 367 station FAST ($0.64 \mu mol L^{-1}$). Very low and similar concentrations of chlorophyll *a* were measured at the three stations 368 (0.063 - 0.072 µg L⁻¹). The proportion of the different major pigments (Fig. 3) showed that 369 370 phytoplankton communities at stations TYR and ION were very similar with a dominance of 371 Prymnesiophytes (i.e. 19'-hexanoyloxyfucoxanthin; Ras et al., 2008) followed by Cyanobacteria (i.e. Zeaxanthin; Ras et al., 2008). In contrast, at station FAST, the plankton community was 372 373 clearly dominated by photosynthetic prokaryotes (i.e. Zeaxanthin and Divinyl-chlorophyll a; 374 Cyanobacteria and Prochlorophytes, respectively; Ras et al., 2008). At all three stations, the 375 proportion of pigments representative of larger species (i.e. Fucoxanthin and Peridinin;-diatoms 376 and dinoflagellates respectively; Ras et al., 2008) were very small (< 5% for each pigment).

377	Cellular abundances of all studied microorganisms (phytoplankton, micro-grazers,
378	heterotrophic bacteria) were the highest at FAST (Table 2). Pieoeukaryotes, Synechococcus and
379	heterotrophic prokaryotes abundances followed an east to west increasing trend (ION < TYR <
380	FAST). In contrast, nano-eukaryotes abundance was similar at FAST and ION, and minimal at
381	TYR. The abundance of heterotrophic nanoflagellates (HNF) were similar at TYR and FAST
382	(-110-125 cells mL ⁻¹), twice as high as the one observed at station ION. This east to west
383	increasing trend was also observed for micro-phytoplankton and micro-heterotrophs abundances
384	(microscopic analyses; Table 2), The ratio between autotrophic biomass and heterotrophic
385	biomass was clearly in favor of the heterotrophic compartment at stations TYR and FAST (-0.6
386	at the two stations) but the opposite was found at station ION (ca. 1.3).

387 3.2. Conditions of irradiance, temperature and pH during

388 the experiments

389 Irradiance levels, during the experiments in the control tanks (C1, C2), were maximal at station ION and minimal at station FAST (daily average maximum levels in controls: ~ 1050, ~ 390 1130 and ~ 1020 µmol photons m⁻² s⁻¹ at TYR, ION and FAST, respectively; Fig. 4). Decreases 391 of water transparency after dust addition was observed at all three stations with a maximum dust 392 393 impact at station ION and the lowest impact at station FAST where irradiance levels decreased by only 60 µmol photons m⁻² s⁻¹ after dust addition (average between tanks D and G), At station 394 395 TYR, a more pronounced decrease was observed in acidified and warmed tanks (G1 and G2) 396 with a decrease of daily average maximum irradiance of ~ 60 and ~ 160 μ mol photons m⁻² s⁻¹ as 397 compared to dust-amended tanks D and controls, respectively. Temperature control (Fig. 4) was 398 not optimal showing deviations between replicates of treatment G of up to 1.5 °C (station ION). 399 Temperature in controls and D tanks displayed a daily cycle with an increase during the day and 400 a decrease at night. Overall, the differences between the warmed treatment (G) and the other

401 tanks were +3, +3.2 and +3.6 °C at TYR, ION and FAST, respectively.

402 Addition of CO₂-saturated filtered seawater led to a decrease of pH_T from 8.05 ± 0.004 403 (average \pm SD between C1, C2, D1 and D2 at t0) to 7.74 (average between G1 and G2) at station 404 TYR, from 8.07 ± 0.002 to 7.78 at station ION and 8.05 ± 0.001 to 7.72 at station FAST (Fig. 5). 405 pH_T levels remained more or less constant in ambient pH levels tanks during all three 406 experiments with no clear impact of dust addition in tanks D1 and D2. In lowered pH tanks, pH 407 levels gradually increased during the experiments with a systematic larger increase in one of the 408 duplicates (G1), Yet pH_T increases remained moderate thanks to the flushing of CO₂-enriched air 409 above the tanks (pCO_2 of 1017 ± 11 , 983 ± 96 , 1023 ± 25 ppm at TYR, ION and FAST, 410 respectively; data not shown). Partial pressure of CO₂ in ambient air was similar at the three 411 stations, i.e. 410 ppm (data not shown). At all three stations, the addition of ¹³C-bicarbonate to 412 all tanks before to led to an increase of total alkalinity between 6 and 11 µmol kg⁻¹ at t0. Dust 413 addition, performed right after t0 in tanks D and G, led to a A_T decrease in these tanks between 8 414 and 16 µmol kg⁻¹ at t24h with no apparent effects of warming and acidification. Overall, no large 415 changes in this parameter were observed during the experiments (Fig. 5).

416 **3.3. Changes in nutrient concentrations**

Dust addition in tanks D and G led to a rapid and maximum input of NO_x (as observed during the first 6 h; Fig. 6; Table 3) of -11μ mol L⁻⁺ at all three stations with no differences between both treatments. The corresponding dissolution percentage of N contained in the dust analog was between 94 and 99%. In contrast, maximum DIP release (within 6 h after dust addition) from the dust analog was much smaller and comprised between 20 and 37 nmol L⁻¹, with slightly higher release at FAST (31-37 nmol L⁻¹) as compared to the other stations. Dissolution percentages for DIP were estimated between 9.2 and 17.3% of total phosphorus 424 contained in dust. As a consequence of these contrasted dissolution of N and P, NO_x:DIP ratios
425 increased from initial values below 5 to above 300, within 6 h after dust seeding, in the dust
426 amended (D and G) tanks (Fig. 6).

427 After these rapid increases due to N and P releases in dust amended tanks, both variables 428 decreased with time. While nutrient variability was small in control tanks over the duration of the experiments (NO_x and DIP variations below 20 and 3 nmol L^{-1} , respectively), large decrease of 429 430 both elements was measured in dust amended tanks (D and G; Table 4). For NO_x, similar linear 431 decreases were observed throughout the experiments at stations TYR and ION with no visible 432 differences between tanks D and G. In contrast, at station FAST, a more pronounced decrease in 433 NO_x was observed in dust-amended (D and G) tanks as compared to the other stations, with 434 detectable larger decreases in warmed and acidified tanks relative to the D treatment. 435 Nevertheless, at all stations, NO_x concentrations in D and G treatments remained far above ambient levels throughout the experiments (> 9 μ mol L⁻¹). Abrupt decreases in DIP were 436 437 observed during the three experiments after the initial increase. At station TYR, after 24 h, all 438 DIP released from dust decreased to initial levels in tanks G while it took two more days to reach 439 initial levels in tanks D. In contrast, at station ION, no clear difference in DIP dynamics was 440 observed between treatments D and G, with concentrations that decreased rapidly during the first 24 h but that remained above initial levels until the end of the experiment. At station FAST, 441 442 similarly to station TYR, DIP decreased rapidly from t12h in treatment G, reaching levels close 443 to initial conditions at the end of the experiment. DIP decrease was much lower in treatment D 444 (Table 4) with concentrations maintained far above ambient levels throughout the experiment. 445 As a consequence of these differences between NO_x and DIP dynamics as well as differences 446 among stations, NO_x:DIP ratio increased during the experiments with clear differences between 447 stations (Fig. 6) and remained much higher than that in the controls over the duration of the three **448** experiments.

449 Silicate dynamics showed at all stations higher concentrations in dust amended (D and G) tanks relative to the controls. At TYR, while concentrations remained stable in control tanks, 450 451 they increased linearly with time in the other tanks (D and G) with no apparent effect of the 452 imposed increase in temperature and decrease in pH (i.e. tanks G). Difference of Si(OH)₄ 453 concentration between dust amended treatments (D and G) and controls was $\sim 0.1 \mu mol L^{-1}$ at the 454 end of the experiment. At station ION, after an initial decrease of concentrations between t-12h 455 and t0, concentrations increased in all tanks until the end of the experiment with higher 456 eoneentration in dust amended tanks (D and G) than in controls (no difference between D and G 457 treatments). In contrast, at FAST, concentrations increased between t-12h and t0, and continued 458 to increase in all tanks (with higher values in dust amended tanks) until t48h and then decreased 459 until the end of the experiment. At the end of the experiment (t96h), Si(OH)₄ concentration was 460 higher in the G treatment than in the D treatment which was similar to the controls.

461 **3.4. Changes in biological stocks**

Regarding biological stocks, temporal dynamics showed very different patterns amongst 462 463 the three studied stations. At TYR, total chlorophyll a concentrations did not change in dust 464 amended tanks maintained under ambient levels of temperature and pH (Fig. 7) and even led to 465 slightly decreased values 24 h after dust addition (e.g. -35 to -38% in D1 and D2, respectively as 466 compared to controls; Table 5). No clear effect of dust addition (tanks D vs. C) were detectable 467 for all groups based on pigment analyses (Fig. 7). Results obtained based on flow cytometry 468 eounting (Fig. 8) were coherent with these observations and showed stronger decreases in cell 469 abundances for $< 20 \ \mu m$ autotrophic groups in tanks D1 and D2 (-77 to -80%). In contrast, at this 470 station, the abundance of heterotrophic prokaryotes (HP) increased rapidly after dust addition 471 both under ambient (+53-68%) and future (+68%) environmental conditions, with no clear 472 difference among those treatments. In warmed and acidified tanks, strong discrepancies between

the duplicates were observed for pigments and autotrophic cell abundances. Indeed, tank G1 showed moderately higher levels for all variables as compared to tanks C with the exception of pico-eukaryotes, while in G2 all variables responded strongly to dust addition with maximum relative changes of > 300% (with the exception of nano-eukaryotes: +119%). While HNF abundances responded positively to the treatments in D1, D2 and G2 (+100-352%), abundances increased sharply in tank G1 towards the end of the experiment (+1095%).

479 At ION, a clear distinction between treatments could be observed for almost all pigments 480 and cell abundances (Fig. 7, Fig. 8). With the exception of nano-eukaryotes and HNF, all 481 variables (pigments and cell abundances) increased as a response to both dust addition and 482 warmed/acidified conditions (i.e. C < D < G). As an example (Table 5), the maximum relative 483 changes as compared to controls observed for total chlorophyll a were 109-183% and 399-426% 484 in tanks D and G, respectively. The highest stimulation to dust addition was observed for 485 Synechococcus with a +317-390%-increase and +805-1425% increase in D and G tanks 486 respectively (Table 5). Abundances of nano-eukaryotes and HNF suggested no impact of dust addition under ambient conditions but a positive impact in treatment G. In contrast to what was 487 488 observed at TYR for HP abundances, an effect of temperature and pH was observed at station 489 ION with a higher impact of dust addition under future environmental conditions. 490 At station FAST, all above mentioned variables related to biological stocks increased

491 strongly after dust addition (Fig. 7, Fig. 8 and Table 5). For instance, total chlorophyll *a*

492 increased following an exponential trend until the end of the experiment reaching maximal

493 values at t96h with slightly lower values observed under ambient environmental conditions

494 (+237-318% in D tanks $\frac{1}{\sqrt{2}}$ ~ +400% in G tanks). Prymnesiophytes (i.e. 19'-

495 hexanoyloxyfucoxanthin) and diatoms (i.e. Fucoxanthin) appeared as the groups benefiting the

496 most from dust addition with no large impacts of warming/acidification. In contrast,

497 Pelagophytes (i.e. 19'-butanoyloxyfucoxanthin) and green algae (i.e. Total Chlorophyll *b*)

498	responded much more in treatment G than in treatment D. Finally, although Cyanobacteria (i.e.
499	Zeaxanthin) responded faster to dust addition under future environmental conditions (tanks G),
500	this effect tended to attenuate towards the end of the experiment. In contrast to estimates based
501	on HPLC data, increases in cell abundances did not generally take place until the end of the
502	experiment. While abundances in pico-eukaryotes increased until t96h in treatment D,
503	abundances sharply declined between t72h and t96h for this group in treatment G. The same
504	trend was observed for Synechococcus during this experiment, although discrepancies between
505	duplicates in treatment D at sampling time t96h did not allow drawing conclusions on the
506	behavior of this group at the end of the experiment. Both under ambient and future conditions,
507	abundances of nano-eukaryotes declined sharply between t72h and t96h. The decline in HP
508	abundances appeared even earlier during the experiment with moderate maximum relative
509	differences as compared to controls observed at t48h. HP abundances declined very sharply
510	between t48h and t96h in treatment G, reaching control levels, while this decline was less sharp
511	under ambient environmental levels. Finally, HNF dynamics during this experiment was hard to
512	evaluate with no elear effects of dust addition or pH/temperature conditions and with a large
513	increase in abundances in only one duplicate of treatment G (t24h) followed by a gradual
514	decrease.
515	Abundances of meso-zooplankton at the end of the experiments showed relatively similar
516	values at stations TYR and ION while much higher levels were observed at station FAST (Fig.
517	9). As a consequence of large variability between duplicates at stations TYR and ION, no clear
518	effects of treatments were detected. At station FAST, although the sample size was too low to
519	statistically test for differences, higher total abundances of meso-zooplankton species were
520	observed in the dust-amended tanks with no differences between ambient and future conditions
521	of temperature and pH. However, differences in abundance were visible between these two

- 522 treatments for specific groups, with respectively higher abundance of Harosa and lower
- 523 abundance of Crustacea (other than copepods) and Mollusca in warmed and acidified tanks.

524 **4. Discussion**

525 4.1. Initial conditions

526 Over the transect, the mixed layer occupied the first 20 m. It was shallower at TYR as 527 eompared to ION and FAST (mixed layer depth, MLD of ~ 10 vs ~15 m, respectively) at the 528 time of the sampling (Van Wambeke et al., 2020a). Such shallow MLD is well representative of 529 stratified conditions encountered in the western Mediterranean basin in late spring/early summer 530 (D'Ortenzio et al., 2005). Overall, the three experiments were conducted with surface seawater 531 eollected during oligotrophic conditions typical of the open Mediterranean Sea at this period of 532 the year (late spring). Although direct measurements of NO_x and DIP concentrations using 533 nanomolar techniques (as performed in our study) are scarce in the Mediterranean Sea, the low 534 levels measured during the cruise are in agreement with DIP values reported for the three studied 535 basins (Djaoudi et al., 2018) and with NO_x and DIP concentrations measured in coastal waters of 536 Corsica in late spring/early summer (Louis et al., 2017b; Pulido-Villena et al., 2014; Ridame et 537 al., 2014). Furthermore, at all three stations, NO_x:DIP molar ratios in the tested surface waters 538 were well below the Redfield ratio (16:1) and are consistent with ratios found in these previously 539 eited studies. Both low NO_x:DIP ratio and low nutrient concentrations suggest that communities 540 found at the three stations experienced N and P co-limitation at the start of the experiments, as 541 previously shown by Tanaka et al. (2011). A side nutrient enrichment experiment confirmed that, 542 at the three sites, heterotrophic bacteria were mainly N-P co-limited (Van Wambeke et al., 543 2020b). In contrast to N and P, initial concentrations of dissolved Fe in the sampled seawater, ranging from 1.5 nmol L⁻¹ at TYR to 2.5 nmol L⁻¹ at ION (Roy-Barman et al., 2020), were 544 545 unlikely limiting for biological activity as previously shown in the Mediterranean Sea under 546 stratified conditions (Bonnet et al., 2005; Ridame et al., 2014).

547	Low total chlorophyll a concentrations in the tested waters were representative of surface
548	eoncentrations reported for the Western and Central Mediterranean Sea in late spring/early
549	summer, both from remote sensing images (Bosc et al., 2004), and from in situ measurements
550	provided in a database from Manca et al. (2004). While large species (i.e. diatoms,
551	dinoflagellates) represented only ~10% of the total chlorophyll a biomass of the tested waters,
552	the composition of the smaller size phytoplankton communities differed substantially- Indeed,
553	communities were clearly dominated by nano-eukaryotes at stations TYR and ION and a larger
554	contribution from pico-eukaryotes and Cyanobacteria was observed at station FAST. Due to their
555	low competitiveness under nutrient limitation, the small contribution of large phytoplankton cells
556	at the start of the experiment is a fingerprint of LNLC areas in general, and of surface
557	Mediterranean waters in late spring and summer (Siokou-Frangou et al., 2010).
558	As-biomass of both heterotrophic nanoflagellates and prokaryotes followed a west to east
559	gradient (FAST > TYR > ION), the ratio of autotrophic vs heterotrophic biomass appeared
560	elearly in favor of the heterotrophic compartment at stations TYR and FAST (ratio of 0.6) while
561	a value above 1 was estimated at ION (ratio of 1.3). This is coherent with the highest net
562	community production (NCP) rates being reported at this station by Gazeau et al. (2021)
563	showing that the initial community at the start of this experiment was very close to metabolic
564	balance (mean \pm SE: -0.06 \pm 0.09 μ mol O ₂ L ⁻¹ d ⁻¹). The highest community respiration rates and
565	consequently lowest NCP rates were measured at station TYR (-1.9 µmol O ₂ L ⁻¹ d ⁻¹) further
566	suggesting that the autotrophic plankton community was not very active and relying on
567	regenerated nutrients, as shown by the highest level of NH4 ⁺ measured at the start of this
568	experiment. In contrast, although slightly heterotrophic (Gazeau et al., 2021) and limited by the
569	low amount of nutrients, the community of the tested waters at FAST was the most active as
570	shown by the highest levels of ¹⁴ C production and heterotrophic prokaryote production (Gazeau
571	et al., 2021) as well as N2 fixation (Céline Ridame, unpublished results). Altogether, the

heterotrophic signature of the three investigated stations, although closer to metabolic balance at
ION, reflected typical biogeochemical conditions in the Mediterranean Sea during late spring to
early summer (Regaudie-de-Gioux et al., 2009).

575 4.2. Critical assessment of the experimental system and

576 methodology

577 The experimental tanks used in this study have been successfully validated in previous 578 studies designed to investigate the inputs of macro- and micro-nutrients (e.g. NO_x, DIP, DFe) 579 and the export of organic matter, under close-to-abiotic conditions (seawater filtration onto 0.2 580 µm) following simulated wet dust events using the same analog as used in our study (Bressac 581 and Guieu, 2013; Louis et al., 2017a, 2018). Louis et al. (2017a, 2018) further investigated these 582 impacts under lowered pH conditions, although no control of atmospheric pCO₂ was performed resulting in a rapid increase of pH levels in the acidified filtered seawater due to CO₂ outgassing 583 584 (from ~7.4 to ~7.7 in six days). Since those above-mentioned studies, in order to avoid this, we 585 improved our experimental system to allow mimicking future conditions by controlling 586 atmospheric pCO_2 in addition to light and temperature (i.e. climate reactors). This allowed to 587 significantly reduce CO₂ outgassing and maintain pH levels close to experimental targets. Still, 588 as illustrated in Fig. 5, the regulation of atmospheric CO₂ was consistently more efficient in tank G2 compared to G1_b resulting in a small discrepancy in terms of pH (highest difference of 0.04 589 590 pH units between the two G tanks at FAST), possibly due to a potential leak or a longer flushing 591 time above tank G1. Nevertheless, as no systematic differences in terms of biological response 592 were observed between these two tanks, we believe that these small differences in terms of 593 regulated pH had no consequences on the obtained results.

594 The lids above tanks, equipped with LEDs in order to reproduce sunlight intensity and 595 spectrum, were used for the first time during these experiments. While simulated intensities were

596 close to estimates for the Northwestern Mediterranean Sea at 5 m depth in June (~1100 µmol 597 photons m⁻² s⁻¹; Bernard Gentili, personal communication, 2017) and fairly consistent between 598 duplicates under control and dust-amended conditions, larger differences were observed between 599 the two warmed and acidified tanks. The reasons of these discrepancies could result from small 600 differences in terms of light intensity regulation between lids, of PAR sensors calibration and/or 601 of different turbidity related to the amount of particles remaining in the tanks. As for pH 602 discussed above, replication in terms of biological response appeared satisfactory for this 603 treatment (except at station TYR; see below), and we believe these technical issues had no 604 significant impacts on our results. 605 Continuous measurements in the tanks showed that temperature was not spatially

homogeneous, leading to significant differences among replicates. This was especially the ease
for warmed tanks (treatment G) for which a maximal average difference over the experimental
period of 0.7 °C was observed during the FAST experiment. As for the other controlled
parameters discussed above, these discrepancies did not systematically lead to observable
differences in the investigated stocks and processes between duplicates (except at TYR, see
below).

612 The relatively low number of experimental units that could be installed inside an 613 embarkable clean container restrained our possibility to consider more than two replicates per 614 treatment. Fortunately, as already said, differences between duplicates were, for the vast majority 615 of studied variables and processes, lower than differences between treatments and appear 616 acceptable considering the difficulty to incubate plankton communities for which slight 617 differences in their initial composition can translate into very-important differences in dynamics 618 (Eggers et al., 2014). Nevertheless, we have to note that important discrepancies were detected 619 regarding autotrophic stocks and processes (Gazeau et al., 2021) for tanks of the warmed and 620 acidified treatment at station TYR. The reasons behind these differences are not fully understood but we strongly suspect that heterotrophic nano-flagellates, feeding mainly on prokaryotic
picoplankton (Sherr and Sherr, 1994), exerted a strong top-down control on this group in tank G1
in which HNF abundance sharply increased during the experiment. All in all, while the
methodology used in this study allowed to successfully evaluate the impacts of dust addition
under both present and future environmental conditions at two out of three tested waters, these
discrepancies at station TYR prevent us from drawing any strong conclusion on the effect of dust
addition on the dynamics of the community under future environmental conditions at that station.

628 **4.3. Impact of dust addition under present environmental**

629 conditions

630 During the three experiments, the observed increases in NO_x and DIP few hours after dust addition under present environmental conditions were rather similar to the enrichment levels 631 obtained during the DUNE experiments at the surface of the mesocosms ($\sim 50 \text{ m}^3$) after the 632 633 simulation of a wet dust deposition using the same dust analog and the same simulated flux 634 (Pulido-Villena et al., 2014; Ridame et al., 2014). The intensity of this simulated wet deposition event (i.e. 10 g m⁻²) represents a high but realistic scenario, as several studies reported even 635 higher short wet deposition events in this area of the Mediterranean Sea (Bonnet and Guieu, 636 637 2006; Loÿe-Pilot and Martin, 1996; Ternon et al., 2010). Furthermore, based on previous studies reporting the mixing between dust and polluted air masses during the atmospheric transport of 638 639 dust particles (e.g. Falkovich et al., 2001; Putaud et al., 2004), we chose to use an evapo-640 condensed dust analog that mimics the processes taking place in the atmosphere prior to 641 deposition, essentially the adsorption of inorganic and organic soluble species (e.g. sulfate and 642 nitrate; see Guieu et al., 2010a, for further details). The imposed evapo-condensation processes 643 are responsible for the large nitrate releasing capacity of the dust particles used in our study. As a 644 consequence, the addition of new nutrients from dust in our study and during the P and R DUNE

experiments were much higher, especially for NO_x, than those observed by Pitta et al. (2017, and
references therein) and Ridame et al. (2014) following the simulation of a dry Saharan dust
deposition event. This confirms that wet dust deposition is a more efficient source of
bioavailable nutrients eompared to dry dust deposition.

Although NO_x and DIP increases after dust addition were rather-similar during our three experiments, the subsequent dynamics of these elements and the impacts on plankton community composition and functioning were drastically different. While NO_x levels decreased moderately over the course of our experiments due to biological uptake, more abrupt decreases were observed for DIP released by dust, reaching values close to the ones observed in the controls, except at station FAST where concentrations were still above ambient levels at the end of the experiment.

656 Regarding biological stocks, most experiments reporting on the effect of dust addition in the Mediterranean Sea showed significant increases in chlorophyll a concentrations (mean ~90% 657 increase; Guieu and Ridame, 2020). Interestingly, no stimulation of autotrophic biomass and 658 659 primary production rates (Gazeau et al., 2021) was observed in dust-amended tanks under 660 present conditions at station TYR. To the best of our knowledge, this is the first experimental evidence of a complete absence of response from an autotrophic community following dust wet 661 662 deposition. The absence of response from autotrophic stocks could be due to a tight top-down control from grazers hiding potential responses from the autotrophic community (Lekunberri et 663 664 al., 2010; Marañón et al., 2010) and/or a competition for nutritive resources with heterotrophic 665 prokaryotes (Marañón et al., 2010). Regarding the first hypothesis, Feliú et al. (2020) have shown that the mesozooplankton assemblage at TYR was clearly impacted by a dust event that 666 took place nine days before sampling at that station as evidenced from particulate inventory of 667 lithogenic proxies (Al, Fe) in the water column (Bressac et al., in preparation). This dust 668 669 deposition likely stimulated phytoplankton growth and consequently increased the abundance of

670 herbivorous grazers (copepods) and attracted carnivorous species. With respect to the second 671 hypothesis, it is well known that not only phytoplankton but also heterotrophic bacteria are 672 limited by inorganic nutrients, mainly DIP, in oligotrophic systems (Obernosterer et al., 2003; 673 Van Wambeke et al., 2001). Indeed, many recent studies have shown significant increase in 674 heterotrophic bacterial abundance, respiration and/or production following dust deposition (and 675 nutrient enrichment) in these areas (Lekunberri et al., 2010; Pitta et al., 2017; Pulido-Villena et 676 al., 2008; Romero et al., 2011). Most of the time, heterotrophic processes appear to be more 677 stimulated by dust pulses compared to autotrophic processes with increasing degree of 678 oligotrophy, the dominant response being modulated by the competition for nutrients between 679 phytoplankton and bacteria (Marañón et al., 2010). This is clearly what was observed at this 680 station, with heterotrophic prokaryotes reacting quickly and strongly to nutrient addition both in 681 terms of abundances and production rates (Gazeau et al., 2021). These two aforementioned 682 hypotheses are not mutually exclusive, and the quick response of heterotrophic prokaryotes to 683 dust addition is coherent with the strongest net heterotrophy of the tested waters at this station (see 4.1). The strong stimulation of heterotrophic prokaryotes and the absence of detectable 684 685 effects on the autotrophic compartment drove the community towards an even stronger net 686 heterotrophic state as illustrated by the decrease in the autotrophic to heterotrophic biomass ratio 687 following dust addition (data not shown). This was further shown by increases in community 688 respiration and decreases in net community production rates in dust-amended as compared to 689 control tanks (Gazeau et al., 2021) and suggest that dust addition to surface waters strongly 690 dominated by heterotrophs leads to a reduction of the capacity of these waters to export organic 691 matter and sequester atmospheric CO₂.

In contrast to what was observed at TYR, fertilization of primary producers was observed at
 stations ION and FAST under present conditions with overall relative changes much higher than
 from previous studies compiled by Guieu and Ridame (2020). The largest increase in chlorophyll

695 a concentrations at station FAST is coherent with the largest NOx decrease observed in our 696 study, which occurred at this station. Interestingly, following dust addition at this station, 697 autotrophic production did not lead to DIP exhaustion throughout the experiment as DIP 698 concentrations were still above ambient conditions at the end of the experiment. Maximal 699 primary production rates (14C-incorporation) at this station at the end of the experiment suggest-a 700 strong DIP recycling and the dominance of regenerated production towards the end of the 701 experiment (Gazeau et al., 2021). Although, in some cases, Synechococcus appeared stimulated 702 by dust addition (Herut et al., 2005; Lagaria et al., 2017; Paytan et al., 2009), Guieu et al. 703 (2014b) showed that, based on the analysis of several aerosols addition studies, this group had 704 generally weak responses to aerosol addition in contrast to nano- and micro-phytoplankton, 705 suggesting that aerosol deposition may lead to an increase in larger size elass phytoplankton. 706 Yet, at stations ION and FAST, the increase in *Synechococcus* abundance in dust-amended tanks 707 was the highest relative to those of pico- and nano-eukaryotes. This was especially true at station 708 ION where no clear response to nutrient enrichment was observed for nano-eukaryotes 709 throughout the experiment. However, it must be stressed that our experiments were performed 710 over a relatively short period (3 to 4 days), and the sharp increase in Fucoxanthin paralleled by a 711 decrease in silicates, at the end of the experiment at station FAST where DIP limitation was not 712 yet apparent, suggests a delayed response of diatoms as compared to smaller groups (i.e. 713 autotrophic prokaryotes, pico- and nano-cukaryotes). Although this was not observed based on 714 pigment analyses, the sharp decline in nano-eukaryote abundances in dust-amended tanks at the end of the FAST experiment, further suggests that this group, reacting quickly to nutrient 715 716 enrichment was progressively grazed and/or outcompeted by larger phytoplankton species. 717 In contrast to what was observed at TYR, at station FAST, the competition for nutrients 718 between autotrophs and heterotrophs was clearly in favor of autotrophs with a clear increase in

719 the ratio between autotrophic and heterotrophic biomass reaching values of up to 4 (data not

720	shown). While, as discussed above, all groups of primary producers benefited from nutrient
721	enrichment at this station, the increases in heterotrophic prokaryote abundances were rather
722	limited following dust deposition, leading to an increase of net community production rates
723	throughout this experiment to reach positive levels while control tanks remained below
724	metabolic balance (Gazeau et al., 2021). At station ION, the situation was somewhat
725	intermediate with a similar enhancement of both autotrophic and heterotrophic stocks and no
726	clear changes in the ratio between autotrophic and heterotrophic biomass (data not shown),
727	although the system appeared in favor of net autotrophy at the end of the experiment in dust -
728	amended tanks under present environmental conditions (Gazeau et al., 2021).
729	Transfer of newly produced organic matter to higher trophic levels in the different
730	treatments was evaluated through the quantification of meso-zooplankton abundance at the end
731	of each experiment. Although we are fully aware that such an approach is certainly criticizable
732	considering the low incubation times (3 to 4 days), it may still be representative of lowered
733	mortality or faster growth. Altogether it does not appear as a surprise that an increase in meso-
734	zooplankton abundances was only detected at station FAST where the strongest enhancement of
735	primary production was observed. Such an increase in meso-zooplankton abundance in the dust-
736	amended as compared to control treatment was observed during land-based mesocosm
737	experiments in the Eastern Mediterranean Sea (Pitta et al., 2017).
738	Finally, although no clear effects of dust deposition under present conditions were
739	detectable on autotrophic prokaryotes at station TYR, the strongest increase in N ₂ fixation rates
740	was recorded at this station (Céline Ridame, unpublished results). However, the potential impact
741	of this process on NO _x concentration is highly negligible compared to the very large stock of

742

 NO_x present in the dust-amended tanks, as less than 1 nmol L⁻¹ d⁻¹ of NO_x can be produced by

this process (Céline Ridame, unpublished results). 743

744 **4.4. Impact of dust-addition under future environmental**

745 conditions

Very few past studies have investigated the release and fate of nutrients from atmospheric 746 particles under climate conditions as expected for the end of the century, and, to the best of our 747 748 knowledge, our study represents the first attempt to test for the combined effect of ocean 749 warming and acidification on these processes. Louis et al. (2018) have already shown from an 750 experiment performed under close-to-abiotic conditions (seawater filtration onto 0.2 µm) that 751 even an extreme ocean acidification scenario (~ -0.6 pH units) does not impact the bioavailability of macro- and micro-nutrients (NO_x, DIP and DFe) from dust addition for surface phytoplankton 752 753 communities in the oligotrophic Northwestern Mediterranean Sea, using the same dust analog 754 and simulated flux as used during our experiments. Similar results were presented by Mélançon 755 et al. (2016) regarding the release of DFe from dust in high-nutrient low-chlorophyll (HNLC) waters of the Northeastern Pacific, following a mild ocean acidification scenario of -0.2 pH 756 757 units. As no differences were observed for NOx and DIP concentrations within few hours 758 following dust addition under present and future environmental conditions, our results agree with 759 these previous findings and further highlights the absence of direct effect of ocean warming (+3) 760 °C) on the release of nutrients from atmospheric particles.

In contrast, following these similar nutrient releases, different nutrient consumption dynamics were observed between ambient and warmed/acidified tanks. These differences were substantially dependent on the considered nutrient and investigated station. Regarding NO_x, no impacts of warming and acidification could be observed at stations TYR and ION due to low net decreasing rates compared to the large increase following dust addition. In contrast, at the most productive station FAST, as a consequence of strongly enhanced biological stocks-(see

thereafter) and metabolic rates (Gazeau et al., 2021). larger NO_x consumption rates were shown
under future environmental conditions.

769 The differences in DIP dynamics between the two dust-amended treatments were more complex to interpret depending on the investigated station. A clear feature of our experiments is 770 771 that, in contrast to present day pH and temperature conditions, all the stock of DIP released from dust was consumed at the end of the three experiments under future conditions, suggesting a 772 773 much faster consumption by autotrophs and heterotrophic prokaryotes. That being said, the rate 774 of decrease under future environmental conditions differed depending on the station. While DIP 775 dynamics were quite similar between tanks maintained under present and future environmental 776 conditions at ION, warming and acidification induced a faster decrease of DIP at TYR and 777 FAST, with a full consumption of the released DIP within 24 h. An interesting outcome at station 778 TYR was that, despite the important discrepancies observed for autotrophic stocks and metabolic 779 rates between the duplicates G1 and G2 (see section 4.2), a very-similar dynamics was observed 780 for DIP concentrations in these tanks. As heterotrophic prokaryote biomass and production rates 781 (Gazeau et al., 2021) did not differ between these duplicate tanks, this further highlights the clear 782 dominance of heterotrophic processes at this station, a dominance which was exacerbated by dust 783 addition under future environmental conditions, leading to an even stronger heterotrophic state at 784 the end of this experiment (Gazeau et al., 2021).

At station ION, large impacts of warming and acidification have been observed, especially for primary producers, as shown by almost doubled chlorophyll *a* concentrations as compared to dust amended tanks (D). At this station, all autotrophic groups benefited from ocean acidification and warming. *Synechococcus* and to a lesser extent pico-eukaryotes appeared as the most impacted ones. Yet these differences of sensitivity among autotrophs did not lead to detectable changes in the composition of the autotrophic assemblage as compared to ambient conditions, with still a large dominance of nano-eukaryote carbon biomass at the end of this

experiment (62% in treatment G vs. 64% in treatment D). Interestingly, although the ratio
between autotrophic and heterotrophic biomass appeared impacted positively-under future
environmental conditions, reaching values of up to 2 at the end of this experiment (data not
shown), warming and acidification led to a decrease in net community production (Gazeau et al.,
2021) suggesting that in the coming decades the capacity of surface seawater to sequester
anthropogenic CO₂ will be lowered.

798 Similarly, at FAST, all phytoplankton groups were impacted positively by warming and 799 acidification with the strongest changes detected for Synechococcus as compared to present 800 environmental conditions. However, in contrast to station ION, all groups reached maximal 801 abundances (and carbon biomass) after 3 days of incubations, thereafter drastically decreasing 802 most likely as a consequence of DIP limitation (see above). It must be stressed that this pattern 803 could not be observed through pigment-dynamies as no sampling was performed for these 804 analyses after 3 days of incubation. Also, in contrast to station ION, the abundance of 805 heterotrophic prokaryotes in the warmer and acidified treatment reached a maximum after 2 days 806 of incubations and then strongly decreased to reach levels observed in the control treatment. This 807 suggests that the heterotrophic compartment was the first to suffer from DIP limitation and 808 further highlights the dominance of the autotrophic compartment in terms of nutrient 809 consumption at this station. As observed at station ION, although the ratio between autotrophic 810 and heterotrophic biomass increased under future environmental conditions, Gazeau et al. (2021) 811 reported on a decrease in net community production rates in this treatment as compared to 812 ambient environmental conditions, suggesting that, in the future, nutrient release from dust will lead to a lesser sequestration capacity of surface waters for atmospheric CO₂. 813 814 These positive effects of warming and acidification on the abundance of phytoplankton eells,

especially for small species, as observed at ION and FAST are in line with previously published
 studies. Indeed, although very contrasted results have been shown on the effect of ocean

acidification on small autotrophic species (e.g. Dutkiewicz et al., 2015), there is increasing
evidence that small phytoplankton species will be favored in a warmer ocean (e.g. Chen et al.,
2014; Daufresne et al., 2009; Morán et al., 2010). As mentioned earlier, our experimental
protocol was not conceived to discriminate temperature from pH effects, however results concur
with those of Maugendre et al. (2015) which further suggested temperature over elevated CO₂ as
the main driver of increased picophytoplankton abundance in the Mediterranean Sea.

823 These enhanced fertilizing effects on primary producers at ION and FAST, under future 824 as compared to present environmental conditions, did not seem to reach higher trophic levels as 825 no clear differences in meso-zooplankton abundances were observed between ambient and warmed/acidified tanks at the end of the experiments. We fully acknowledge that the duration of 826 our experiments was eertainly too short to carefully assess the proportion of newly formed 827 organic matter consumed by meso-zooplankton species and its effect on their abundances, yet 828 829 group-specific variations were observed. Similarly, Gazeau et al. (2021) did not observe an 830 additional impact of future environmental conditions on the export of organic matter after dust 831 addition.

832 **5. Conclusion**

833 These experiments conducted during the PEACETIME cruise represent the first attempt 834 to investigate the impacts of atmospheric deposition on surface plankton communities both under 835 present and future environmental conditions. Despite few experimental issues that are discussed, 836 the three experiments provided new insights on these potential impacts in the open 837 Mediterranean Sea. Interestingly, the effect of dust deposition was highly different between the three investigated stations in the Tyrrhenian Sea, Ionian Sea and in the Algerian basin. As the 838 839 initial conditions in the sampled surface seawater at the three stations were very similar in terms 840 of nutrient availability and chlorophyll content, these differences rather seem to be a 841 consequence of the initial metabolic states of the community (autotrophy vs. heterotrophy). In all 842 three cases, nutrient addition from dust deposition did not strongly modify but rather exacerbated 843 this initial state. Relative changes in main parameters presented in this manuscript and processes presented in Gazeau et al. (2021) as a consequence of dust addition under present and future 844 845 environmental conditions are shown in Fig. 10, and compared to the compilation of published 846 data for the Mediterranean Sea from Guieu and Ridame (2020). At station TYR, under 847 conditions of a clear dominance of heterotrophs on the use of resources and potentially a higher top-down control from grazers, dust addition drove the community into an even more 848 849 heterotrophic state with no detectable effect on primary producers. At station ION, where the 850 community was initially closer to metabolic balance, both heterotrophic and autotrophic 851 compartments benefited from dust derived nutrients. At FAST, the most active station in terms 852 of autotrophic production, addition of nutrients boosted both compartments but heterotrophic prokaryotes became quickly P-limited and overall larger effects were observed for 853 854 phytoplankton. Ocean acidification and warming did not have any detectable impact on the 855 release of nutrients from atmospheric particles. Furthermore, these external drivers did not

- drastically modify the composition of the autotrophic assemblage with all groups benefiting from
- 857 warmer and acidified conditions. However, although for two out of the three stations
- 858 investigated, larger increases were observed for autotrophic as compared to heterotrophic stocks
- under future environmental conditions, a stronger impact of warming and acidification on
- 860 mineralization processes (Gazeau et al., 2021) suggests that, in the future, the plankton
- 861 communities of Mediterranean surface waters will have a decreased capacity to sequester
- 862 atmospheric CO₂ following the deposition of atmospheric particles.

863 **Data availability**

All data and metadata will be made available at the French INSU/CNRS LEFE CYBER database

865 (scientific coordinator: Hervé Claustre; data manager, webmaster: Catherine Schmechtig).

866 INSU/CNRS LEFE CYBER (2020)

867 Author contributions

868 FG and CG designed and supervised the study. FG, CG, CR and KD sampled seawater from the

869 experimental tanks during the experiments. JMG and GDL participated in the technical

870 preparation of the experimental system and all authors participated in sample analyses. FG, CR

and CG wrote the paper with contributions from all authors.

872 **Financial support**

- 873 This study is a contribution to the PEACETIME project (http://peacetime-project.org), a joint
- 874 initiative of the MERMEX and ChArMEx components supported by CNRS-INSU, IFREMER,
- 875 CEA, and Météo-France as part of the programme MISTRALS coordinated by INSU.
- 876 PEACETIME is a contribution to SOLAS and IMBER international programme. The project was
- 877 endorsed as a process study by GEOTRACES. PEACETIME cruise
- 878 (<u>https://doi.org/10.17600/17000300</u>). The project leading to this publication has also received
- funding from the European FEDER Fund under project 1166-39417.

880 Acknowledgments

- 881 The authors thank the captain and the crew of the RV Pourquoi Pas? for their professionalism
- and their work at sea. We thank Julia Uitz, Céline Dimier and the SAPIGH HPLC analytical
- 883 service at Institut de la Mer de Villefranche (IMEV) for sampling and analysis of phytoplankton
- pigments, John Dolan for microscopic countings as well as Lynne Macarez and the PIQv-

- 885 platform of EMBRC-France, a national Research Infrastructure supported by ANR, under the
- 886 reference ANR-10-INSB-02, for mesozooplankton analyses.

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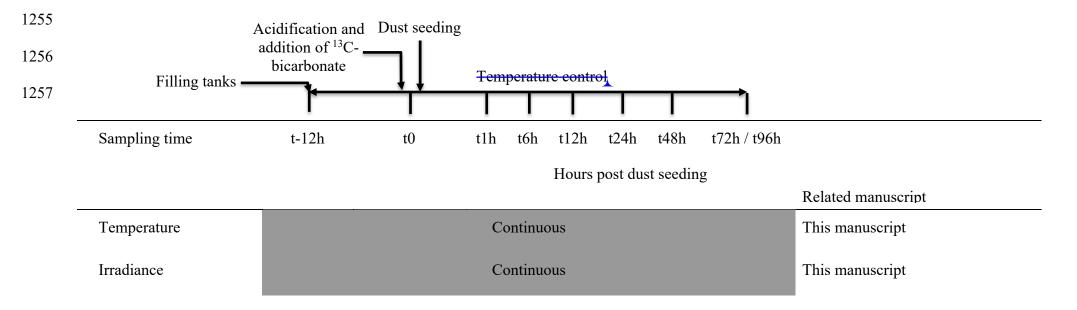
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- 1245

1246	Table 1. List of parameters and processes investigated during the three experiments at stations
1247	TYR, ION and FAST. Related manuscripts are indicated. pH_T : pH on the total scale, A_T : total
1248	alkalinity, ¹³ C- <i>C</i> _T : ¹³ C signature of dissolved inorganic carbon, NO _x : nitrate + nitrite, DIP:
1249	dissolved inorganic phosphorus, Si(OH)4: silicate, DFe: dissolved iron, DAI: dissolved
1250	aluminium, Th-REE-Pa: Thorium (230Th and 232Th), Rare Earth elements and Protactinium
1251	(²³¹ Pa), POC: particulate organic carbon, DOC: dissolved organic carbon, ¹³ C-DOC: ¹³ C
1252	signature of dissolved organic carbon, TEP: transparent exopolymer particles, NCP/CR: net
1253	community production and community respiration (oxygen based), ¹⁴ C-PP: primary production
1254	based on ¹⁴ C incorporation.

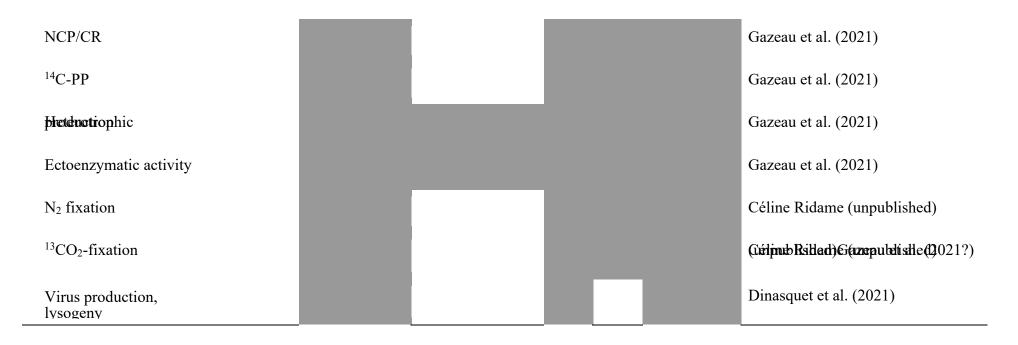


1 1 **T** · · · · · · . . .

Carbonate chemistry

pH _T		This manuscript
A_{T}		This manuscript
δ^{13} C- C_{T}		Gazeau et al. (2021)
Macro-nutrients		
NO _x		This manuscript
DIP		This manuscript
Si(OH) ₄		This manuscript
Micro-nutrients		
DFe		Roy-Barman et al. (2020)
DAI		Roy-Barman et al. (2020)
Th-REE-Pa		Roy-Barman et al. (2020)
Biological stocks		
Pigments		This manuscript
Flow cytometry		This manuscript

Microscopy			This manuscript
Diazotroph abundance			Céline Ridame (unpublished)
Virus abundance			Dinasquet et al. (2021)
Meta-transcriptomics			Dinasquet et al. (2021)
Bacterial diversity			Dinasquet et al. (2021)
Micro-eukaryote diversity			Dinasquet et al. (2021)
Meso-zooplankton			This manuscript
POC (incl. δ^{13} C)			Gazeau et al. (2021)
POC sediment traps			Gazeau et al. (2021)
DOC			Gazeau et al. (2021)
¹³ C-DOC			Gazeau et al. (2021)
TEP			Gazeau et al. (2021)
Amino acids			Gazeau et al. (2021)
Carbohydrates			Gazeau et al. (2021)
Processes			



1259	Table 2. Initial conditions as measured while filling the tanks (initial conditions in pumped
1260	surface water; sampling time: t-12h). pH _T : pH on the total scale, NO _x : nitrate + nitrite, NH ₄ :
1261	ammonium, DIP: dissolved inorganic phosphorus, Si(OH)4: silicate, TChla: total chlorophyll a,
1262	HNF: heterotrophic nanoflagellates. The three most important pigments in terms of concentration
1263	are also presented (19'-hexanoyloxyfucoxanthin, Zeaxanthin and Divinyl Chlorophyll a).
1264	Biomasses of the different groups analyzed through flow cytometry were estimated based on
1265	conversion equations and/or factors found in the literature (see section 2.3). Autotrophic biomass
1266	was, as a first approximation, estimated only based on flow cytometry data and therefore
1267	corresponds to the fraction $< 20 \ \mu m$. Heterotrophic biomass was estimated as the sum of
1268	heterotrophic prokaryote and HNF biomasses (see section 2.3.2). Values below detection limits
1269	are indicated as < dl.

	Sampling station	TYR	ION	FAST
	Coordinates (decimal)	39.34 N, 12.60 E	35.49 N, 19.78 E	37.95 N, 2.90 N
	Bottom depth (m)	3395	3054	2775
	Day and time of sampling (local time)	17/05/2017 17:00	25/05/2017 17:00	02/06/2017 21:00
	Temperature (°C)	20.6	21.2	21.5
	Salinity	37.96	39.02	37.07
Carbonate	pH_T	8.04	8.07	8.03

chemistry	Total alkalinity (µmol kg ⁻¹)	2529	2627	2443
Nutrients	NO_x (nmol L ⁻¹)	14.0	18.0	59.0
	NH_4^+ (µmol L ⁻¹)	0.045	0.022	< d1
	DIP (nmol L ⁻¹)	17.1	6.5	12.9
	$Si(OH)_4 (\mu mol L^{-1})$	1.0	0.96	0.64
	NO _x /DIP (molar ratio)	0.8	2.5	4.6
Pigments	TChla (μ g L ⁻¹)	0.063	0.066	0.072
	19'-hexanoyloxyfucoxanthin ($\mu g L^{-1}$)	0.017	0.021	0.016
	Zeaxanthin (µg L ⁻¹)	0.009	0.006	0.036
	Divinyl Chlorophyll a (µg L ⁻¹)	~ 0	0	0.014
Flow cytometry	Pico-eukaryotes (abundance in cell mL ⁻¹ ; biomass in µg C L ⁻¹)	347.8; 0.5	239.9; 0.4	701.0; 1.0
	Nano-eukaryotes (abundance in cell mL ⁻¹ ; biomass in µg C L ⁻¹)	150.5; 3.9	188.8; 4.8	196.6; 5.0
	Synechococcus (abundance in cell mL ⁻¹ ; biomass in μ g C L ⁻¹)	4972; 1.2	3037; 0.8	6406; 1.6
	Autotrophic biomass (µg C L ⁻¹)	5.6	6.0	7.7
	Heterotrophic prokaryotes abundance (x 10 ⁵ cell mL ⁻¹)	4.79	2.14	6.15
	HNF (abundance in cell mL ⁻¹)	110.1	53.6	126.2
	Heterotrophic biomass (µg C L ⁻¹)	9.9	4.5	12.7
Microscopy	Pennate diatoms (abundance in cell L ⁻¹)	140	520	880
	Centric diatoms (abundance in cell L ⁻¹)	200	380	580

Dinoflagellates (abundance in cell L ⁻¹)	2770	3000	3410	
Autotrophic flagellates (abundance in cell L ⁻¹)	0	60	650	
Ciliates (abundance in cell L ⁻¹)	270	380	770	

Table 3. Maximum input of nitrate + nitrite (NO_x) and dissolved inorganic phosphorus (DIP) released from Saharan dust in tanks D and G as observed from the two discrete samplings performed over the first 6 h after seeding. The estimated maximal percentage of dissolution is also presented (see section 2.3.1 for details on the calculations).

		NO _x			DIP			
	D1	D2	G1	G2	D1	D2	G1	G2
Maximum input		μmo	ol L ⁻¹			nmo	ol L ⁻¹	
TYR	11.0	11.1	11.1	11.0	24.6	20.4	24.6	23.9
ION	11.2	11.6	11.2	11.3	23.3	22.0	19.6	22.9
FAST	11.3	11.1	11.1	11.2	30.8	31.3	36.9	29.8
Percentage of dissoluti	on (%)							
TYR	95	96	95	94	12	10	12	11
ION	96	99	96	97	11	10	9	11
FAST	97	97	95	97	15	15	17	14

1	Table 4. Removal rate of nitrate + nitrite (NO _x) and dissolved inorganic phosphorus (DIP) in
2	tanks D and G during the three experiments (TYR, ION and FAST). For NO _x , decreasing rates
3	were estimated based on linear regressions between maximal concentrations (i.e. after dust
4	enrichment, at t1h or t6h) and final concentrations (t72 h for TYR and ION and t96h for FAST).
5	For DIP, decreasing rates were estimated based on linear regressions between maximal
6	concentrations (i.e. after dust enrichment at t1h or t6h) and concentrations measured at sampling
7	times after which a stabilization was observed. This sampling time is shown in parentheses. All
8	rates are expressed in nmol L ⁻¹ h ⁻¹ .

		NO _x			DIP	
	TYR	ION	FAST	TYR	ION	FAST
D1	-6.5	-8.6	-14.3	-0.4 (t72h)	-0.5 (t48h)	-0.2 (t96h)
D2	-1.0	-8.6	-13.5	-0.3 (t72h)	-0.8 (t24h)	-0.2 (t96h)
G1	-6.7	-13.1	-21.6	-1.3 (t24h)	-0.8 (t24h)	-1.5 (t24h)
G2	-0.8	-1.6	-25.2	-1.3 (t24h)	-1.6 (t24h)	-1.1 (t24h)

10 Table 5. Maximum relative changes in tanks D and G as compared to controls (average betwe	10	Table 5. Maximum relative	e changes in tanks D	and G as compared to	o controls (average betwee
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- 11 C1 and C2), expressed as a %, for the three experiments (TYR, ION and FAST). The sampling
- 12 time at which these maximum relative changes were observed is shown in brackets. Tchla refers
- 13 to the concentration of total chlorophyll *a* and B_{micro} to the biomass proxy of micro-
- 14 phytoplankton (sum of Fucoxanthin and Peridinin, see Material and Methods) based on high
- 15 performance liquid chromatography (HPLC). HP and HNF refer to heterotrophic prokaryote and
- 16 heterotrophic nanoflagellate abundances, respectively, as measured by flow cytometry.

Experiment	Tank	HP	PLC		FI	low cytometry		
		TChla	Bmicro	Pico-eukaryotes	Nano-eukaryotes	Synechococcus	HP	HNF
TYR	D1	-35 (t24h)	-33 (t12h)	-75 (t72h)	-80 (t1h)	-71 (t48h)	68 (t72h)	352 (t72h)
TYR	D2	-38 (t12h)	-39 (t24h)	-75 (t72h)	-80 (t1h)	-72 (t48h)	53 (t72h)	100 (t72h)
TYR	G1	60 (t72h)	52 (t72h)	-75 (t1h)	89 (t72h)	76 (t72h)	67 (t72h)	1095 (t72h)
TYR	G2	359 (t72h)	392 (t72h)	323 (t72h)	119 (t72h)	700 (t72h)	68 (t48h)	298 (t72h)
ION	D1	183 (t72h)	157 (t72h)	126 (t72h)	89 (t72h)	317 (t72h)	128 (t72h)	44 (t72h)
ION	D2	109 (t72h)	156 (t72h)	117 (t72h)	-59 (t1h)	390 (t72h)	133 (t72h)	27 (t72h)
ION	G1	399 (t72h)	454 (t72h)	458 (t72h)	256 (t72h)	805 (t72h)	176 (t72h)	175 (t72h)

ION	G2	426 (t72h)	612 (t72h)	510 (t72h)	292 (t72h)	1425 (t72h)	161 (t72h)	129 (t72h)
FAST	D1	318 (t96h)	356 (t96h)	113 (t96h)	208 (t72h)	348 (t96h)	27 (t96h)	-38 (t96h)
FAST	D2	237 (t96h)	322 (t96h)	91 (t96h)	219 (t72h)	197 (t96h)	40 (t48h)	-49 (t96h)
FAST	G1	399 (t96h)	415 (t96h)	198 (t72h)	274 (t72h)	357 (t48h)	61 (t48h)	243 (t24h)
FAST	G2	395 (t96h)	421 (t96h)	129 (t72h)	202 (t96h)	344 (t48h)	67 (t48h)	74 (t24h)

Figure captions

Fig. 1. Location of the sampling stations in the Mediterranean Sea onboard the R/V "Pourquoi Pas ?" during the PEACETIME cruise, on map of satellite-derived surface chlorophyll *a* concentration averaged over the entire duration of the cruise (Courtesy of Louise Rousselet).

Fig. 2. Scheme of an experimental tank (climate reactor).

Fig. 3. Proportion of the different pigments, as measured by high performance liquid chromatography (HPLC) in pumped surface seawater for the three experiments (t-12h).

Fig. 4. Continuous measurements of temperature and irradiance level (PAR) in the six tanks during the three experiments. The dashed vertical line indicates the time of dust seeding (after t0).

Fig. 5. pH on the total scale (pH_T) and total alkalinity (A_T) measured in the six tanks during the three experiments. The dashed vertical line indicates the time of dust seeding (after t0). Error bars correspond to the standard deviation based on analytical triplicates.

Fig. 6. Nutrients (nitrate + nitrite: NO_x, dissolved inorganic phosphorus: DIP, silicate: Si(OH)₄ as well as the molar ratio between NO_x and DIP, measured in the six tanks during the three experiments. The dashed vertical line indicates the time of seeding (after t0).

Fig. 7. Concentrations of total chlorophyll a and major pigments, measured by high performance liquid chromatography (HPLC), in the six tanks during the three experiments. The dashed vertical line indicates the time of seeding (after t0).

Fig. 8. Abundance of autotrophic pico-eukaryotes, autotrophic nano-eukaryotes, *Synechococcus,* heterotrophic prokaryotes (HP), and heterotrophic nano-flagellates (HNF), measured by flow cytometry, in the six-tanks during the three-experiments. The evolution of autotrophic biomass

(see Material and Methods for details on the calculation) is also shown. The dashed vertical line indicates the time of seeding (after t0).

Fig. 9. Abundances of meso-zooplankton species as measured at the end of each experiment, Fig. 10. Maximal relative change-(%) of main biological stocks (TCHl*a*: total chlorophyll *a*, HP: heterotrophic prokaryotes) and processes (BP: bacterial production; PP: ¹⁴C-based primary production; see Gazeau et al., 2021; BR: bacterial respiration (no data from this study); and N₂ fixation, Céline Ridame, unpublished results) as obtained during the present study at the 3 stations (TYR, ION and FAST) under ambient conditions of pH and temperature (open red squares) and future conditions (full green squares). Squares are delimited by the range of responses observed among the duplicates for each treatment. The dotted green squares for station TYR denote the large variability observed between duplicates for some parameters and processes that prevented drawing solid conclusions. Box-plots represent the distribution of responses observed from studies conducted in the Mediterranean Sea, as compiled by Guieu and Ridame (2020).

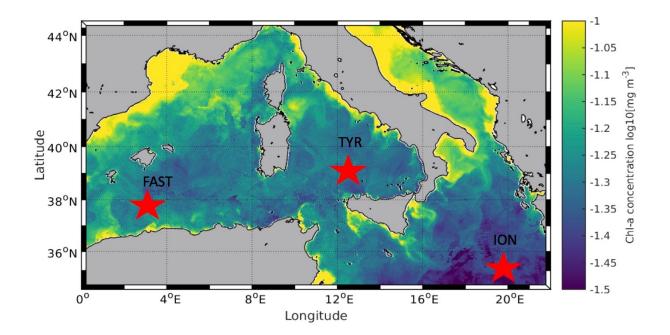


Fig. 1.

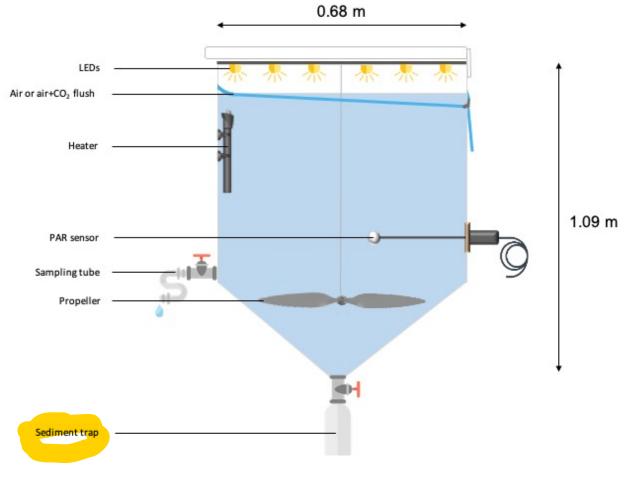
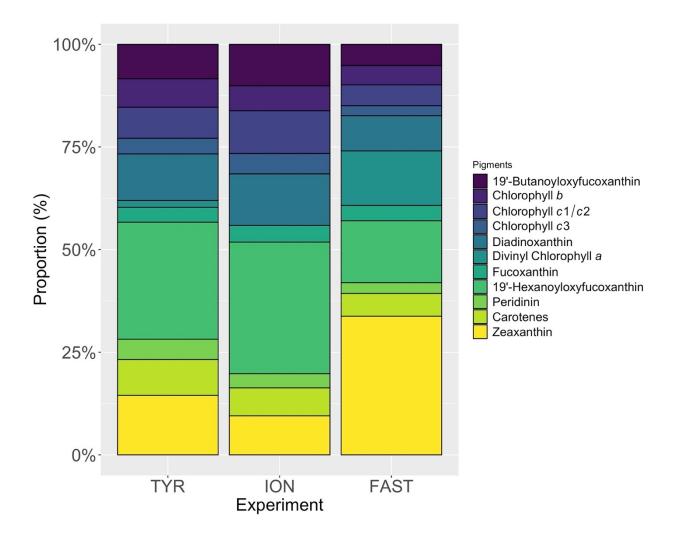
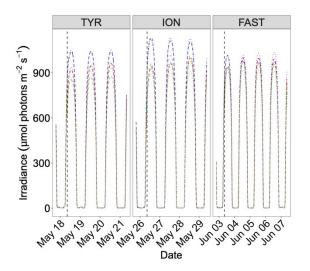


Fig. 2.







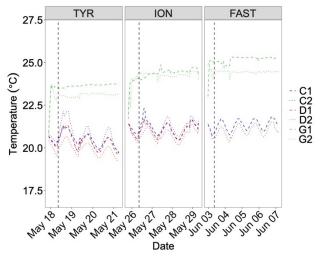


Fig. 4.

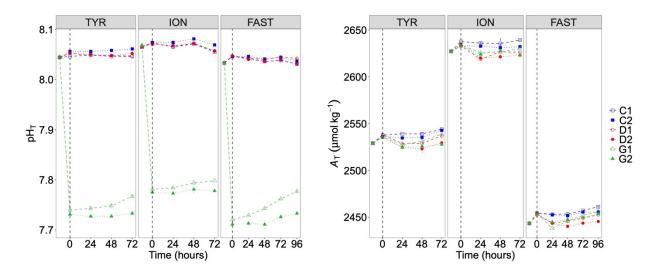


Fig. 5.

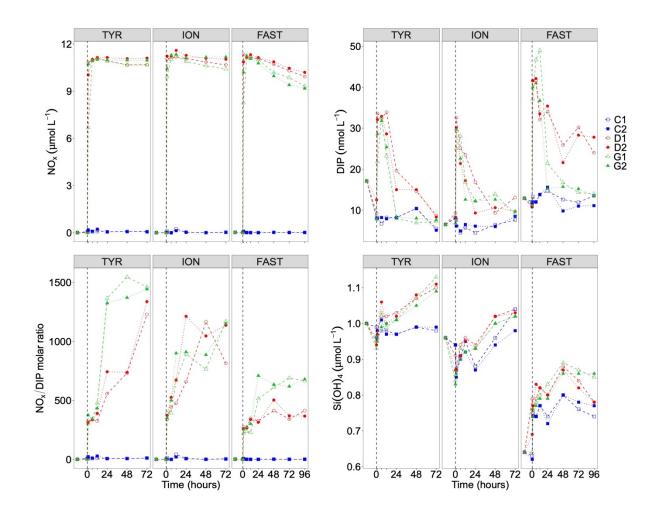


Fig. 6.

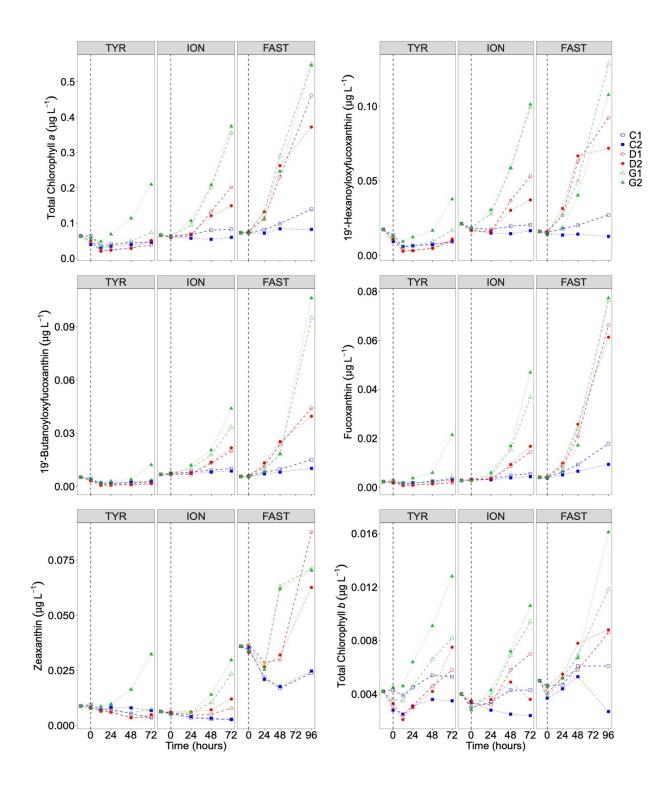


Fig. 7.

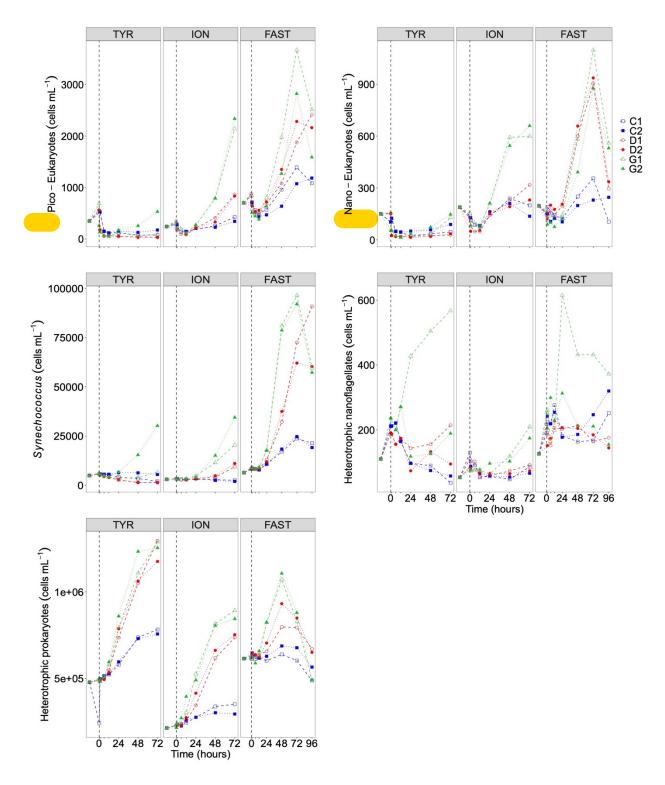


Fig. 8.

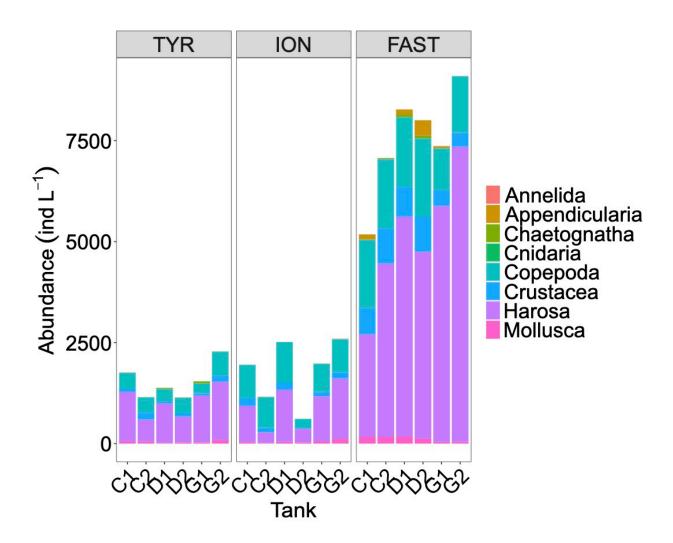


Fig. 9.

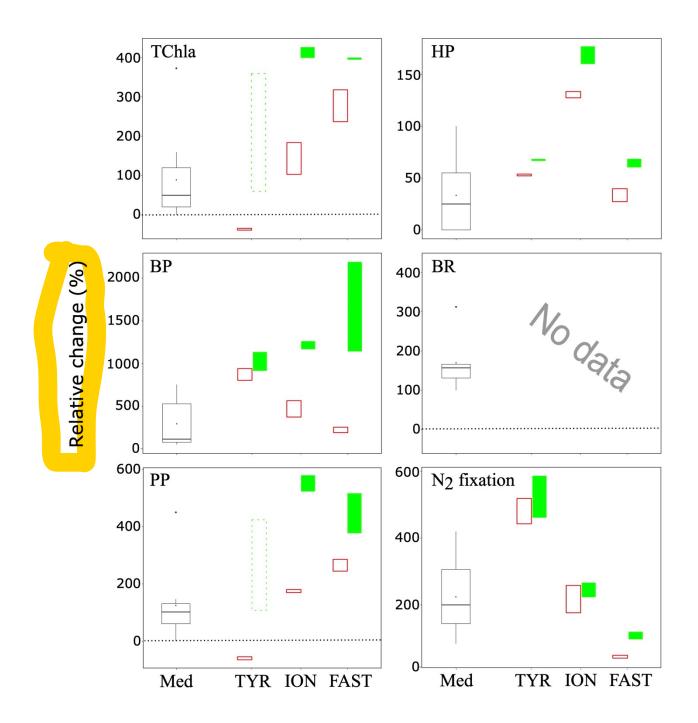


Fig. 10.