https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020


**On the barium - oxygen consumption relationship in the Mediterranean Sea: implications for mesopelagic marine snow remineralisation.**

Stéphanie H.M. Jacquet[1*], Dominique Lefèvre[1], Christian Tamburini[1], Marc Garel[1], Frédéric A.C. Le Moigne[1], Nagib Bhairy[1], Marie Roumagnac[1], Sophie Guasco[1]

[1]Aix Marseille Université, CNRS/INSU, Université de Toulon, IRD, Mediterranean Institute of Oceanography (MIO), UM 110, 13288 Marseille, France

[*]Correspondence to: S. Jacquet (*stephanie.jacquet@mio.osupytheas.fr)*

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020


**ABSTRACT**

In the ocean, remineralisation rate associated with sinking particles is a crucial variable. Since the 90's, particulate biogenic barium ($Ba_{xs}$) has been used as an indicator of carbon remineralization by applying a transfer function relating $Ba_{xs}$ to $O_2$ consumption (Dehairs's transfer function, Southern Ocean-based). Here, we tested its validity in the Mediterranean Sea (ANTARES / EMSO-LO) for the first time by investigating connections between $Ba_{xs}$, prokaryotic heterotrophic production (PHP) and oxygen consumption ($JO_2$-Opt; optodes measurement). We show that: (1) higher $Ba_{xs}$ (409 pM; 100- 500 m) in situations where integrated PHP (PHP100/500= 0.90) is located deeper, (2) higher $Ba_{xs}$ with increasing $JO_2$-Opt, and (3) similar magnitude between $JO_2$-Opt (3.14 mmol m$^{-2}$ d$^{-1}$; 175- 450 m) and $JO_2$-Ba (4.59 mmol m$^{-2}$ d$^{-1}$; transfer function). Overall, $Ba_{xs}$, PHP and $JO_2$ relationships follow trends observed in the Southern Ocean. We believe that such transfer function could apply in the Mediterranean Sea with no restriction.

**KEY WORDS:** particulate biogenic barium, mesopelagic zone, oxygen consumption, prokaryotic heterotrophic production, carbon remineralization, Mediterranean Sea

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020


## 1. INTRODUCTION

Ocean ecosystems play a critical role in the Earth's carbon (C) cycle [IPCC, 2014]. The quantification of their impacts of both present conditions and future predictions remains one of the greatest challenges in oceanography [Siegel et al., 2016]. In essence, the biological C pump is termed for the numerous processes involved in maintaining the vertical gradient in dissolved inorganic C. This includes processes such as organic matter production in surface, its export and subsequent remineralization. Most of marine snow organic C conversion (i.e. remineralization) into $CO_2$ by heterotrophic organisms (i.e. respiration) occurs in the mesopelagic zone (100-1000 m) [Martin et al., 1987; Buesseler and Boyd, 2009]. Globally, the flux of C exported below 1000 m depth is the key determinant of ocean carbon storage capacity [Henson et al., 2011]. However, there is no consensus on C transfer efficiency estimations from field experiments, leading to an imbalance of the water column C budget [Giering et al., 2014]. Resolving this imbalance is in the core of numerous studies in the global ocean, but also regionally, especially in the Mediterranean Sea (MedSea). Due to limited exchanges with adjacent basin and the existence of an intense overturning circulation qualitatively resembling the global one (but with shorter time scales), the MedSea is often considered as a laboratory to observe and understand the impact of transient climate variability on ecosystems and biogeochemical cycles [Malanotte-Rissoli et al., 2014]. In a context of climate changes, better constraining C fluxes and the ocean C storage capacity is of crucial importance.

Particulate barium in excess ($Ba_{xs}$, i.e. biogenic Ba from total particulate Ba after correction for lithogenic Ba) is a geochemical tracer of particulate organic carbon (POC) remineralization in the mesopelagic layer [Dehairs et al., 1997]. $Ba_{xs}$ mostly occurs in the form of barite microcrystals ($BaSO_4$) at these depths. In a global ocean undersaturated with respect to barite, studies report that $Ba_{xs}$ would precipitate inside oversaturated biogenic micro-environments during POC

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



degradation by heterotrophic prokaryotes in the mesopelagic zone, through sulfate and/or barium enrichment [Bertram and Cowen, 1997]. The first-ever studies on mesopelagic $Ba_{xs}$ reported coinciding $Ba_{xs}$ maxima with depths of dissolved $O_2$ minimum and $pCO_2$ maximum [Dehairs et al., 1987, 1997]. By using an 1D advection-diffusion model applied to highly resolved, precise $O_2$ profiles in the Atlantic sector of the Southern Ocean (ANTX/6 cruise; Shopova et al., 1995), Dehairs et al. [1997] established an algorithm converting mesopelagic $Ba_{xs}$ concentration into $O_2$ consumption rate ($JO_2$) and organic C remineralized (POC remineralization rate). This transfer function has been widely used until now [Cardinal et al., 2001- Lemaitre et al., 2018]. Yet its validity has never been tested in other oceanic provinces. Recently, significant progresses were made in relating $Ba_{xs}$, $O_2$ dynamics to prokaryotic heterotrophic activity [Jacquet et al., 2015]. Nevertheless, the Dehairs transfer function has never been revised since. These advancements clearly show that $Ba_{xs}$ is closely related with the vertical distribution of prokaryotes heterotrophic production (PHP) (the rate of change with depth), reflecting the temporal progression of POC remineralization processes. Also, in a first attempt to test the validity of the Dehairs's transfer function in other locations, Jacquet et al. [2015] confronted oxygen consumption rates ($JO_2$) from direct measurements (dark community respiration, DCR) to derived $JO_2$ from $Ba_{xs}$ data (using the transfer function) in the Kerguelen area (Indian sector of the Southern Ocean). We revealed good convergence of $JO_2$ rates from these two approaches, further supporting the Dehairs's function to estimate POC remineralization rates in different biogeochemical settings of the Southern Ocean.

Here, we further investigate relationships between the mesopelagic $Ba_{xs}$ proxy, prokaryotic activity and oxygen dynamics (Figure 1a) in the northwestern Mediterreanean Sea (MedSea), a different biogeochemical setting to those already studied (see references above). Today, observations of the various components of the MedSea biological C pump provide organic C fluxes varying by at least an order of magnitude [Santinelli et al., 2010; Ramondenc et al., 2016].

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



Malanotte-Rissoli et al. [2014] reviewing unsolved issues and future directions for MedSea research highlighted the need to further investigate biogeochemical processes at intermediate (mesopelagic) and deep layers to reconcile the C budget in the Mediterranean basin. Previous particulate $Ba_{xs}$ dataset is very scarce in the NW- MedSea, with in general very low vertical sampling resolution [Sanchez Vidal et al., 2005] or very restricted studied areas [Dehairs et al., 1987; Sternberg et al., 2008]. Here we discuss $Ba_{xs}$, PHP and $JO_2$ (from optodes measurement during incubations) at the ANTARES / EMSO-LO observatory site (Figure 1a, b). We hypothesize that the Dehairs's transfer function converting $Ba_{xs}$ into POC remineralization also applies in a different ocean ecosystem functioning from the Southern Ocean. We suggest that the $Ba_{xs}$ proxy can be used as routine tracer to estimate local-scale processes of mesopelagic POC remineralization in the Mediterranean basin.

## 2. SAMPLING AND ANALYSES

### 2.1 STUDY SITE

The BATMAN cruise (https://doi.org/10.17600/16011100, March 10-16 2016, *R/V* EUROPE) took place to the ANTARES / EMSO-LO observatory site (42°48'N, 6°10'E; Tamburini et al., 2013), 40 km off the coast of Toulon, southern France (Figure 1b). The hydrological and biogeochemical conditions at this site are monitored monthly in the framework of the MOOSE (Mediterranean Ocean Observing System for the Environment) program and of the EMSO (European Multidisciplinary Subsea Observatory) observation program. The hydrography displays the general three-layer MedSea system with surface, intermediate and deep waters [Hainbucher et al., 2014]. Briefly, the main water masses can be distinguished (see potential temperature – salinity diagram during the BATMAN cruise in Figure 1c): (1) Surface Water (SW); (2) Winter Intermediate Water (WIW) and Levantine Intermediate water (LIW). LIW is

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



present at intermediate depths (around 400 m at ANTARES) and is characterized by temperature and a salinity maxima; (4) Mediterranean Deep Water (MDW).

## 2.2 ANALYSES

For particulate barium, 4 to 7 L of seawater sampled using Niskin bottles were filtered onto 47 mm polycarbonate membranes (0.4 µm porosity) under slight overpressure supplied by filtered air. Filters were rinsed with few mL of Milli-Q grade water to remove sea salt, dried (50°C) and stored in Petri dishes. Thirteen depths between surface and 2000 m were sampled by combining different casts sampled closeby in time and space (total of 28 samples). In the laboratory, we performed a total digestion of filters using a tri-acid (0.5 mL HF /1.5 mL $HNO_3$ / HCl 1 mL; all Optima grade) mixture in closed teflon beakers overnight at 95°C in a clean pressurized room. After evaporation close to dryness, samples were re-dissolved into 10 mL of $HNO_3$ 2%. The solutions were analysed for Ba and other elements of interest (Na and Al) by HR-ICP-MS (High Resolution-Inductively Coupled Plasma- Mass Spectrometry; ELEMENT XR ThermoFisher). Details on sample processing and analysis are given in Cardinal et al. [2001] and Jacquet et al. [2015]. The presence of sea-salt was checked by analysing Na and the sea-salt particulate Ba contribution was found negligible. Particulate biogenic barium in excess (hereafter referred to as $Ba_{xs}$) was calculated as the difference between total Ba and lithogenic Ba using Al as the lithogenic reference element [Taylor and Mc.Lennan, 1985]. The standard uncertainty [Ellison et al., 2000] on $Ba_{xs}$ concentration ranges between 5.0 and 5.5%. The term "in excess" is used to indicate that concentrations are larger than the $Ba_{xs}$ background. The background (or residual value) is considered as "preformed" $Ba_{xs}$ at zero oxygen consumption left over after transfer and partial dissolution of $Ba_{xs}$ produced during degradation of previous phytoplankton growth events.

Oxygen concentrations were measured using optical oxygen sensor (Aanderaa 4330-Optodes) at 4 depths in the mesopelagic layer (175, 250, 450 and 1000 m). In total each of the 8 optodes (two per depths) were placed into a sealed 1L borosilicate glass bottles incubated at a fixed temperature of 13°C in thermo-regulated baths for 24 to 48 hours. Oxygen consumption rates (later referred to as $JO_2$-Opt) were calculated from oxygen concentration evolution with time applying linear model calculations.

Prokaryotic heterotrophic production (PHP) estimation was measured over time course experiments at *in situ* temperature (13°C) following the protocol described in Tamburini et al. [2002]. $^3$H-leucine labelled tracer [Kirchman, 1993] was used. To calculate prokaryotic heterotrophic production, we used the empirical conversion factor of 1.55 ng C per pmol of incorporated leucine according to Simon and Azam [1989], assuming that isotope dilution was negligible under these saturating concentrations.

## 3. RESULTS AND DISCUSSION

### 3.1 Barium vertical distribution

Particulate biogenic $Ba_{xs}$, particulate Al (pAl) and biogenic Ba fraction profiles in the upper 1000 m at ANTARES are reported in Figure 2a. $Ba_{xs}$ concentrations range from 12 to 719 pM. The biogenic Ba fraction range from 51 to 91 % of the total particulate Ba signal. Particulate Al concentrations (pAl) are low and range from 8 to 170 nM. $Ba_{xs}$ concentrations are low in surface water (<100 pM) where the lithogenic fraction reaches 43 to 49 % in the upper 70 m. From previous studies we know that $Ba_{xs}$ in surface waters is distributed over different, mainly non-barite biogenic phases, and incorporated into or adsorbed onto phytoplankton material. As such these do not reflect POC remineralization processes, in contrast to mesopelagic waters where $Ba_{xs}$ is mainly composed of barite formed during prokaryotic degradation of organic matter. At

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



ANTARES the $Ba_{xs}$ profile displays a mesopelagic $Ba_{xs}$ maximum between 100 and 500 m, reaching up to 719 pM at 175 m. Ba is mostly biogenic at these depths (> 80 %). $Ba_{xs}$ concentrations then decrease below 500 m to reach a background value of around 130 pM (see BKG in Figure 2). Note that the MedSea is largely undersaturated with respect to barite, with saturation state ranging between 0.2 and 0.6 over the basin [Jacquet et al., 2016; Jullion et al., 2017]. For comparison, the Ba background value in the Southern Ocean reaches 180 to 200 pM below 1000 m [Dehairs et al., 1997; Jacquet et al. 2015]. Previously, Sternberg et al. [2008] reported the seasonal evolution of $Ba_{xs}$ profiles at the DYFAMED station (43°25'N-7°52'E; BARMED project) northeast from ANTARES (Figure 1c) in the NW-MedSea. The present $Ba_{xs}$ profile at ANTARES (March 2016) is very similar to the $Ba_{xs}$ profile measured in March 2003 at DYFAMED (Figure 2a). The slight difference between $Ba_{xs}$ profiles in the upper 75 m suggests more Ba bounded and/or adsorbed onto phytoplankton material during BARMED. Both profiles present a $Ba_{xs}$ maximum in the upper mesopelagic zone between 150 and 200 m. Below this maximum, $Ba_{xs}$ concentrations gradually decrease to reach around 130 pM between 500 and 1000 m (this study). A similar value was reached between 500 and 600 m at the DYFAMED station over the whole studied period (between February and June 2003; Sternberg et al., 2008).

### 3.2 Prokaryotic heterotrophic production

The particulate excess Ba (>BKG) is centred in the upper mesopelagic zone between 100 and 500 m and reflects that POC remineralization mainly occurred at these depths (Figure 2a). Depth-weighted average (DWA) $Ba_{xs}$ content (409 pM) was calculated over this entire depth interval. Figure 2b shows column-integrated PHP at 100 m over column-integrated PHP at 500 m (PHP100/500= 0.90), according to the relationship obtained during KEOPS1 (summer) and KEOPS2 (spring; out plateau stations) cruises in the Southern Ocean [Jacquet et al., 2008; 2015]

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020


and #DY032 cruise (2015, *R/V* DISCOVERY) at the PAP (Porcupine Abyssal Plain) observatory in the northeast Atlantic (49°N, 16.5 °W) (personal data). Results at the ANTARES / EMSO-LO site follow the trend previously reported in the Southern Ocean, indicating higher DWA $Ba_{xs}$ in situations where a significant part of column-integrated PHP is located deeper in the water column (high Int. PHPx1/IntPHPx2 ratio; Figure 2b). These previous studies revealed that the shape of the column-integrated PHP profile (i.e. the attenuation gradient) is important in setting the $Ba_{xs}$ signal in the mesopelagic zone (Dehairs et al., 2008; Jacquet et al., 2008, 2015]. Indeed, mesopelagic $Ba_{xs}$ appears reduced when most of the column-integrated PHP is limited to the upper layer (indicating an efficient remineralization in surface), compared to situations where a significant part of integrated PHP is located deeper in the water column (reflecting significant deep PHP activity, POC export and subsequent remineralization) (Figure 2b). Our MedSea results are located along the trend defined in the Southern Ocean during KEOPS1 cruise. It is generally considered that $Ba_{xs}$ (barite) forms inside sulfate and/or barium oversaturated biogenic micro-environments during POC degradation by heterotrophic prokaryotes. However, it is unclear whether barite formation at mesopelagic depths is (directly or indirectly) bacterially induced or bacterially influenced. Overall, our results strengthen the close link between the water column $Ba_{xs}$ distribution and respiration (organic matter degradation).

### 3.3 Oxygen- barium relationship

The relationship we obtained at ANTARES between $Ba_{xs}$ concentrations and oxygen consumption rates from optodes measurements ($JO_2$-Opt) is reported in Figure 3a. $JO_2$-Opt range from 0.11 to 5.85 µmol $L^{-1}$ $d^{-1}$. The relationship indicates higher $Ba_{xs}$ concentrations with increasing $JO_2$-Opt. An interesting feature is the intercept at zero $JO_2$-Opt (around 128 pM)

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020

Biogeosciences Discussions
Open Access
EGU

which further supports the Ba BKG value at ANTARES (130 pM) determined from measured $Ba_{xs}$ profiles (Figure 3a).

We applied a similar approach as reported in Jacquet et al. (2015) where we show the correlation between $JO_2$ obtained from dark community respiration DCR (winkler titration; $JO_2$-DCR) data integration in the water column and $JO_2$ based on $Ba_{xs}$ content (Dehairs's transfer function; later referred to as $JO_2$-Ba). Similarly, to estimate $JO_2$-Ba in the present study we used the following equation [Dehairs et al., 1997]:

$$JO_2\text{-Ba} = (Ba_{xs} - Ba\ BKG)/17450 \quad (1)$$

A Ba BKG value of 130 pM was used (see above). $JO_2$-Ba is confronted to $JO_2$-Opt integrated over the same layer depth (between 175 and 450 m; Figure 3b). $JO_2$ rates are of the same order of magnitude ($JO_2$-Ba= 4.59 mmol $m^{-2}$ $d^{-1}$ and $JO_2$-opt= 3.14 mmol $m^{-2}$ $d^{-1}$). The slight difference could be explained by the integration time of both methods: few hours to days for the incubations vs. few days to weeks for $Ba_{xs}$ (seasonal build-up; Jacquet et al., 2007). $JO_2$ rates calculated in the present work are 3 times higher than those reported in the Southern Ocean during KEOPS1 [Jacquet et al., 2015] but they are in good agreement with the $Ba_{xs}$ vs $JO_2$ trend (Figure 3b). Overall, our results indicate similar $Ba_{xs}$ - $JO_2$ relationship in the Southern Ocean and the Mediterranean Sea. This further supports the universal validity of the Dehairs's transfer function in the present study.

### 3.4 Estimated particles remineralisation rates and implications

In order to provide a $Ba_{xs}$-derived estimate of POC remineralization rate (MR) at the ANTARES / EMSO-LO observatory during BATMAN cruise, we converted $JO_2$-Ba into C respired using the Redfield (RR) $C/O_2$ molar ratio (127/175; Broecker et al., 1985) multiplied by the depth layer considered (Z) [Dehairs et al., 1997]:

$$MR = Z \times JO_2\text{-}Ba \times RR \quad (2)$$

We obtain a POC remineralization rate of 11 mmol C $m^{-2}$ $d^{-1}$ (10% RSD). This is within the range of previously published remineralization fluxes in the Mediterranean Sea from sediment trap [Sanchez-Vidal et al., 2005] and from thorium-derived data [Speicher et al., 2006]. It is also in good agreement with recent POC flux attenuation combining drifting sediment traps and underwater vision profilers [Ramondenc et al., 2016].

The present paper brings a first insight into the connections of $Ba_{xs}$, PHP and $JO_2$ at the ANTARES/EMSO-LO observatory site in the northwestern Mediterranean Sea during the BATMAN (2016) cruise. Our results reveal a strong relationship between $Ba_{xs}$ contents and measured $JO_2$ rates. Also, DWA $Ba_{xs}$ vs. column integrated PHP, as well as measured vs. $Ba_{xs}$-based $JO_2$ relationships follow trends previously reported in the Southern Ocean where the Dehairs's function was first established to estimate POC remineralisation rate. Results from the present study would indicate that this function can also be applied in the Mediterranean basin provided that adequate $Ba_{xs}$ background values are estimated. From a global climate perspective, the $Ba_{xs}$ tool will help to better balance the MedSea water column C budget. It will contribute to gain focus on the emerging picture of the C transfer efficiency (strength of the biological pump).

## ACKNOWLEDGEMENTS

We thank the officers and crew of *R/V* EUROPE for their assistance during work at sea. This research was supported by the French national LEFE/INSU "REPAP" project (PI. S. Jacquet). It was co-funded by the "ROBIN" project (PIs. C. Tamburini, F.A.C. Le Moigne) of Labex OT-Med (ANR-11-LABEX-0061) funded by the Investissements d'Avenir and the French Government project of the ANR, through the A*Midex project (ANR-11-IDEX-0001-02). Authors have benefited of the support of the SNO-MOOSE and SAM-MIO. BATMAN is a

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020




contribution to the "AT – POMPE BIOLOGIQUE" of the Mediterranean Institute of Oceanography (MIO) and to the international IMBER program. The instrument (ELEMENT XR, ThermoFisher) was supported in 2012 by European Regional Development Fund (ERDF).

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020




**Figure captions**

Figure 1: (a) Schematic representation of the convergence of the different estimators of oxygen consumption and C remineralization rates from the "oxygen dynamics", "barium proxy" and "prokaryotic activity" tools; (b) Location of the BATMAN cruise at the ANTARES observatory site in the NW-Mediterranean Sea (42°48'N, 6°10'E); (b) Potential temperature - salinity - depth plots and isopycals for BATMAN profiles. SW : Surface Water, WIW : Winter Intermediate Water, LIW : Levantine Intermediate Water, DMW : Deep Mediterranean Water. Graph constructed using Ocean Data View (Schlitzer, 2002; Ocean Data View; http://www.awi-bremerhaven.de/GEO/ODV)

Figure 2: (a) Particulate biogenic $Ba_{xs}$ (pM) and particulate Al (nM) profiles next to the biogenic Ba fraction (%) in the upper 1000 m at ANTARES. The grey area represents a biogenic Ba fraction larger than 80 %. BKG: $Ba_{xs}$ background. $Ba_{xs}$ profile (pM) at DYFAMED : data from Sternberg et al. (2008); (b) ANTARES ratio plot (green square) of integrated PHP in the upper 100 m over integrated PHP in the upper 500 m versus depth-weighted average (DWA) mesopelagic $Ba_{xs}$ (pM) over the 150- 500 depth interval. Regression of the same ratio is reported for KEOPS1 (out plateau stations) and KEOPS2 (Southern Ocean; Jacquet et al., 2015) and #DY032 (PAP station, NE-Atlantic; pers. data) cruises.

Figure 3: (a) Relationship between $Ba_{xs}$ concentrations (pM) and oxygen consumption rates (µmol $L^{-1}$ $d^{-1}$) from optodes measurements ($JO_2$-Opt) at ANTARES; (b) Confrontation of oxygen consumption rates (mmol $m^{-2}$ $d^{-1}$) obtained from different methods: optodes measurements (this work), dark community respiration DCR (winkler titration; $JO_2$-DCR; Jacquet et al., 2015;

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



KEOPS1), and (Dehairs's transfer function calculation based on $Ba_{xs}$ content (Dehairs et al., 1997).

**References**

Bertram, Miriam, and James P. Cowen. "Morphological and Compositional Evidence for Biotic Precipitation of Marine Barite." *Journal of Marine Research* 55 (1997): 577–93.

Broecker, Wallace S. "'NO', a Conservative Water-Mass Tracer." *Earth and Planetary Science Letters* 23, no. 1 (August 1974): 100–107. doi:10.1016/0012-821X(74)90036-3.

Buesseler, Ken O., and Philip W. Boyd. "Shedding Light on Processes That Control Particle Export and Flux Attenuation in the Twilight Zone of the Open Ocean." *Limnology and Oceanography* 54, no. 4 (2009): 1210–32. doi:10.4319/lo.2009.54.4.1210.

Cardinal, Damien, Frank Dehairs, Thierry Cattaldo, and Luc André. "Geochemistry of Suspended Particles in the Subantarctic and Polar Frontal Zones South of Australia: Constraints on Export and Advection Processes." *Journal of Geophysical Research: Oceans* 106, no. C12 (décembre 2001): 31637–56. doi:10.1029/2000JC000251.

Dehairs, F., R. Chesselet, and J. Jedwab. "Discrete Suspended Particles of Barite and the Barium Cycle in the Open Ocean." *Earth and Planetary Science Letters* 49, no. 2 (September 1980): 528–50. doi:10.1016/0012-821X(80)90094-1.

Dehairs, F., S. Jacquet, N. Savoye, B. A. S. Van Mooy, K. O. Buesseler, J. K. B. Bishop, C. H. Lamborg, et al. "Barium in Twilight Zone Suspended Matter as a Potential Proxy for Particulate Organic Carbon Remineralization: Results for the North Pacific." *Deep Sea Research Part II: Topical Studies in Oceanography*, Understanding the Ocean's Biological Pump:results from VERTIGO, 55, no. 14–15 (juillet 2008): 1673–83. doi:10.1016/j.dsr2.2008.04.020.

Dehairs, F., C. E. Lambert, R. Chesselet, and N. Risler. "The Biological Production of Marine Suspended Barite and the Barium Cycle in the Western Mediterranean Sea." *Biogeochemistry* 4, no. 2 (June 1, 1987): 119–40. doi:10.1007/BF02180151.

Dehairs, F., D. Shopova, S. Ober, C. Veth, and L. Goeyens. "Particulate Barium Stocks and Oxygen Consumption in the Southern Ocean Mesopelagic Water Column during Spring and Early Summer: Relationship with Export Production." *Deep Sea Research Part II: Topical Studies in Oceanography* 44, no. 1–2 (1997): 497–516. doi:10.1016/S0967-0645(96)00072-0.

Ellison, Eurachem/CITAC Guide CG4, Quantifying Uncertainty in Analytical Measurement. Eds. S.L.R. Ellison, M. Rosslein and A. Williams. Second edition ISBN 0948926 15 5, Pp 120, 2000.

Giering, Sarah L. C., Richard Sanders, Richard S. Lampitt, Thomas R. Anderson, Christian Tamburini, Mehdi Boutrif, Mikhail V. Zubkov, et al. "Reconciliation of the Carbon Budget in the Ocean/'s Twilight Zone." *Nature* 507, no. 7493 (March 27, 2014): 480–83.

Hainbucher, D., A. Rubino, V. Cardin, T. Tanhua, K. Schroeder, and M. Bensi. "Hydrographic Situation during Cruise M84/3 and P414 (spring 2011) in the Mediterranean Sea." *Ocean Sci.* 10, no. 4 (juillet 2014): 669–82. doi:10.5194/os-10-669-2014.

Henson, Stephanie A., Richard Sanders, Esben Madsen, Paul J. Morris, Frédéric Le Moigne, and Graham D. Quartly. "A Reduced Estimate of the Strength of the Ocean's Biological Carbon Pump." *Geophysical Research Letters* 38, no. 4 (février 2011): L04606. doi:10.1029/2011GL046735.

IPCC Working Group 1, 5th Assessment Report (AR5) Climate Change 2013, Published in Jan. 2014.

Jacquet, S. H. M., F. Dehairs, D. Lefèvre, A. J. Cavagna, F. Planchon, U. Christaki, L. Monin, L. André, I. Closset, and D. Cardinal. "Early Spring Mesopelagic Carbon Remineralization and Transfer Efficiency in the Naturally Iron-Fertilized Kerguelen Area." *Biogeosciences* 12, no. 6 (March 17, 2015): 1713–31. doi:10.5194/bg-12-1713-2015.

Jacquet, S. H. M., F. Dehairs, N. Savoye, I. Obernosterer, U. Christaki, C. Monnin, and D. Cardinal. "Mesopelagic Organic Carbon Remineralization in the Kerguelen Plateau Region Tracked by Biogenic Particulate Ba." *Deep Sea Research Part II: Topical Studies in Oceanography*, KEOPS: Kerguelen Ocean and Plateau compared Study, 55, no. 5–7 (March 2008): 868–79. doi:10.1016/j.dsr2.2007.12.038.

Jacquet, S. H. M., J. Henjes, F. Dehairs, A. Worobiec, N. Savoye, and D. Cardinal. "Particulate Ba-Barite and Acantharians in the Southern Ocean during the European Iron Fertilization Experiment (EIFEX)." *Journal of Geophysical Research* 112, no. G4 (October 23, 2007). doi:10.1029/2006JG000394.

Jacquet, S.H.M., C. Monnin, V. Riou, L. Jullion, and T. Tanhua. "A High Resolution and Quasi-Zonal Transect of Dissolved Ba in the Mediterranean Sea." *Marine Chemistry* 178 (January 20, 2016): 1–7. doi:10.1016/j.marchem.2015.12.001.

Jullion, L., S. H. M. Jacquet, and T. Tanhua. "Untangling Biogeochemical Processes from the Impact of Ocean Circulation: First Insight on the Mediterranean Dissolved Barium Dynamics." *Global Biogeochemical Cycles* 31, no. 8 (2017): 1256–70. doi:10.1002/2016GB005489.

Kirchman DL (1993) Leucine incorporation as a measure of biomass production by heterotrophic bacteria. In: Kemp PF, Sherr BF, Sherr EB, Cole JJ (eds) Handbooks of methods in aquatic microbial ecology. Lewis Publishers, Boca Raton, Ann Arbor, London, Tokyo, p 509–512

Lemaitre, N., H. Planquette, F. Planchon, G. Sarthou, S. Jacquet, M. I. García-Ibáñez, A. Gourain, et al. "Particulate Barium Tracing of Significant Mesopelagic Carbon Remineralisation in the North Atlantic." *Biogeosciences* 15, no. 8 (2018): 2289–2307. doi:10.5194/bg-15-2289-2018.

Malanotte-Rizzoli, P., V. Artale, G. L. Borzelli-Eusebi, S. Brenner, A. Crise, M. Gacic, N. Kress, et al. "Physical Forcing and Physical/biochemical Variability of the Mediterranean Sea: A Review of Unresolved Issues and Directions for Future Research." *Ocean Science* 10, no. 3 (2014): 281–322. doi:10.5194/os-10-281-2014.

Martin, John H., George A. Knauer, David M. Karl, and William W. Broenkow. "VERTEX: Carbon Cycling in the Northeast Pacific." *Deep Sea Research Part A. Oceanographic Research Papers* 34, no. 2 (1987): 267–85. doi:http://dx.doi.org/10.1016/0198-0149(87)90086-0.

Ramondenc, Simon, Goutx Madeleine, Fabien Lombard, Chiara Santinelli, Lars Stemmann, Gabriel Gorsky, and Lionel Guidi. "An Initial Carbon Export Assessment in the Mediterranean Sea Based on Drifting Sediment Traps and the Underwater Vision Profiler Data Sets." *Deep Sea Research Part I: Oceanographic Research Papers* 117 (November 2016): 107–19. doi:10.1016/j.dsr.2016.08.015.

Sanchez-Vidal, A., R. W. Collier, A. Calafat, J. Fabres, and M. Canals (2005), Particulate barium fluxes on the continental margin: A study from the Alboran Sea (Western Mediterranean), Mar. Chem., 93, 105–117.

Santinelli, Chiara, Luciano Nannicini, and Alfredo Seritti. "DOC Dynamics in the Meso and Bathypelagic Layers of the Mediterranean Sea." *Deep Sea Research Part II: Topical Studies in Oceanography*, Ecological and Biogeochemical Interactions in the Dark Ocean, 57, no. 16 (août 2010): 1446–59. doi:10.1016/j.dsr2.2010.02.014.

Siegel DA, Buesseler KO, Behrenfeld MJ, Benitez-Nelson CR, Boss E, Brzezinski MA, Burd A, Carlson CA, D'Asaro EA, Doney SC, Perry MJ, Stanley RHR and Steinberg DK (2016) Prediction of the Export and Fate of Global Ocean Net Primary Production: The EXPORTS Science Plan. *Front. Mar. Sci*. 3:22. doi: 10.3389/fmars.2016.00022

Simon M, Azam F (1989) Protein content and protein synthesis rates of planktonic marine bacteria. Mar Ecol Prog Ser 51:201–213

Shopova, D., F. Dehairs, and W. Baeyens. "A Simple Model of Biogeochemical Element Distribution in the Oceanic Water Column." *Journal of Marine Systems* 6, no. 4 (juin 1995): 331–44. doi:10.1016/0924-7963(94)00032-7.

Speicher, E. A., S. B. Moran, A. B. Burd, R. Delfanti, H. Kaberi, R. P. Kelly, C. Papucci, et al. "Particulate Organic Carbon Export Fluxes and Size-Fractionated POC/234Th Ratios in the Ligurian, Tyrrhenian and Aegean Seas." *Deep Sea Research Part I: Oceanographic Research Papers* 53, no. 11 (2006): 1810–30. doi:http://dx.doi.org/10.1016/j.dsr.2006.08.005.

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020


Sternberg, E., C. Jeandel, E. Robin, and M. Souhaut. “Seasonal Cycle of Suspended Barite in the Mediterranean Sea.” *Geochimica et Cosmochimica Acta* 72, no. 16 (2008): 4020–34. doi:https://doi.org/10.1016/j.gca.2008.05.043.

Tamburini C, Canals M, Durieu de Madron X, Houpert L, Lefèvre D, Martini S, D’Ortenzio F, Robert A, Testor P, and the ANTARES collaboration (2013) Deep-sea bioluminescence blooms after dense water formation at the ocean surface. PLoS One 8:e67523

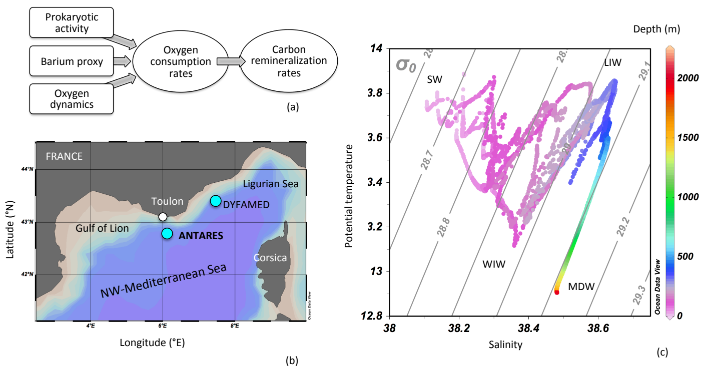
Tamburini C, Garcin J, Ragot M, Bianchi A (2002) Biopolymer hydrolysis and bacterial production under ambient hydrostatic pressure through a 2000 m water column in the NW Mediterranean. Deep Res II 49:2109–2123

Taylor, S.R., McLennan, S.M.: The continental crust: its composition and evolution, Blackwell Scientific Publications, 312pp, 1985.

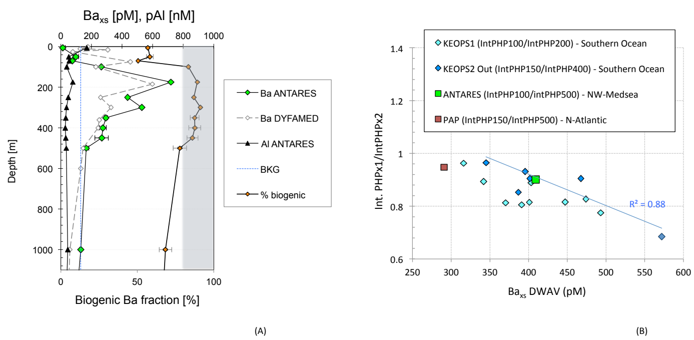
https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020



Figure 1



Biogeosciences
Discussions
Open Access




399 Figure 2

https://doi.org/10.5194/bg-2020-241
Preprint. Discussion started: 7 September 2020

Biogeosciences
Discussions

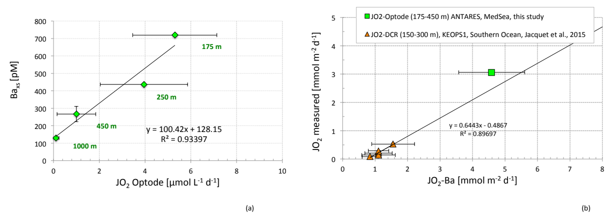



Figure 3