

1 **Research article**

2 **Title**

3 Spatial variations in terrestrial net ecosystem productivity and its local indicators

4 **Running title**

5 Spatial variability in terrestrial NEP

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29 **Key words**

30 Net ecosystem productivity, spatial variation, CO<sub>2</sub> uptake and release, local indicators, model

31 **Abstract**

32 Multiple lines of evidence have demonstrated the persistence of global land carbon (C) sink  
33 during the past several decades. However, both annual net ecosystem productivity (NEP) and  
34 its inter-annual variation ( $IAV_{NEP}$ ) keep varying over space. Thus, identifying local indicators  
35 for the spatially varying NEP and  $IAV_{NEP}$  is critical for locating the major and sustainable C  
36 sinks on the land. Here, based on daily NEP observations from FLUXNET sites and the  
37 atmospheric inversion product, we found a robust logarithmic correlation between annual NEP  
38 and ratio of total CO<sub>2</sub> exchanges during net uptake ( $U$ ) and release ( $R$ ) periods (i.e.,  $U/R$ ). The  
39 cross-site variation of mean annual NEP could be linearly indicated by  $\ln(U/R)$ , while the spatial  
40 distribution of  $IAV_{NEP}$  was well indicated by the slope (i.e.,  $\beta$ ) of the demonstrated logarithmic  
41 correlation. Among biomes, for example, forests and croplands had the largest  $U/R$  ratio ( $1.06$   
42  $\pm 0.83$ ) and  $\beta$  ( $473 \pm 112 \text{ g C m}^{-2} \text{ yr}^{-1}$ ), indicating the highest NEP and  $IAV_{NEP}$  in forests and  
43 croplands, respectively. We further showed that these two simple indicators could directly infer  
44 the spatial variations in NEP and  $IAV_{NEP}$  in global gridded products. Overall, this study provides  
45 two simple local indicators for the intricate spatial variations in the strength and stability of land  
46 C sinks. These indicators could be helpful for locating the persistent terrestrial C sinks and  
47 provides valuable constraints for improving the simulation of land-atmospheric C exchanges.

48

## 49 1. Introduction

50 Terrestrial ecosystems reabsorb about one-quarter of anthropogenic CO<sub>2</sub> emission (Ciais et  
51 al., 2019) and are primarily responsible for the recent temporal fluctuations of the measured  
52 atmospheric CO<sub>2</sub> growth rate (Randerson, 2013; Le Quéré et al., 2018). In addition, evidence  
53 based on eddy-flux measurements (Baldocchi et al., 2018; Rödenbeck et al., 2018), aircraft  
54 atmospheric budgets (Peylin et al., 2013), and process-based model simulations (Poulter et al.,  
55 2014; Ahlstrom et al., 2015) has shown a large spatial variability in net ecosystem productivity  
56 (NEP) on the land. The elusive variation of terrestrial NEP over space refers to both of the  
57 substantial varying mean annual NEP and the divergent inter-annual variability (IAV) in NEP  
58 (i.e., IAV<sub>NEP</sub>; usually quantified as the standard deviation of annual NEP) across space  
59 (Baldocchi et al., 2018; Marcolla et al., 2017). The mean annual NEP is related to the strength  
60 of carbon exchange of a specific ecosystem (Randerson et al., 2002; Luo and Weng, 2011; Jung  
61 et al., 2017), while IAV<sub>NEP</sub> characterizes the stability of such carbon exchange (Musavi et al.,  
62 2017). Thus, whether and how NEP and IAV<sub>NEP</sub> change over the space is important for  
63 predicting the future locations of carbon sinks on the land (Yu et al., 2014; Niu et al., 2017).

64 Large spatial difference in terrestrial NEP has been reported from eddy-flux measurements,  
65 model outputs and atmospheric inversion products. In addition, the global average IAV of NEP  
66 was large relative to global annual mean NEP (Baldocchi et al., 2018). More importantly, the  
67 spatial variations of NEP and IAV<sub>NEP</sub> were typically underestimated by the compiled global  
68 product and the process-based global models (Jung et al., 2020; Fu et al., 2019). These  
69 discrepancies further revealed the necessary to identify local indicators for the spatially varying  
70 NEP and IAV<sub>NEP</sub>, separately. The NEP in terrestrial ecosystems is determined by two  
71 components, including vegetation photosynthesis and ecosystem respiration (Reichstein et al.,  
72 2005). Because photosynthesis and respiration are strongly correlated over space (Baldocchi et  
73 al., 2015; Biederman et al., 2016), their relative difference could determine the spatial variation  
74 of NEP. Many previous analyses have attributed the IAV<sub>NEP</sub> at the site level to the different  
75 sensitivities of ecosystem photosynthesis and respiration to environmental drivers (Gilmanov et  
76 al., 2005; Reichstein et al., 2005) and biotic controls (Besnard et al., 2018; Musavi et al., 2017).  
77 For example, some studies have reported that IAV<sub>NEP</sub> is more associated with variations in

78 photosynthesis than carbon release (Ahlstrom et al., 2015; Novick et al., 2015; Li et al., 2017),  
79 whereas others have indicated that respiration is more sensitive to anomalous climate variability  
80 (Valentini et al., 2000; von Buttlar et al., 2017). However, despite the previous efforts in a  
81 predictive understanding of the land-atmospheric C exchanges, the multi-model spread has not  
82 reduced over time (Arora et al., 2019). Therefore, it is imperative to explore the potential  
83 indicators for the spatially varying NEP, which could help attribute the spatial variation of NEP  
84 and  $IAV_{NEP}$  into different processes and provide valuable constraints for the global C cycle.  
85 Alternatively, the annual NEP of a given ecosystem can be also directly decomposed into CO<sub>2</sub>  
86 uptake flux and CO<sub>2</sub> release flux (Gray et al., 2014), which are more direct components for NEP  
87 (Fu et al., 2019). Many studies have reported that the vegetation CO<sub>2</sub> uptake during the growing  
88 season and the non-growing season soil respiration are tightly correlated (Luo et al., 2014; Zhao  
89 et al., 2016). It is still unclear how the ecosystem CO<sub>2</sub> uptake and release fluxes would control  
90 the spatially varying NEP.

91 Conceptually, the total CO<sub>2</sub> uptake flux ( $U$ ) is determined by the length of CO<sub>2</sub> uptake  
92 period ( $CUP$ ) and the CO<sub>2</sub> uptake rate, while the total CO<sub>2</sub> release flux ( $R$ ) depends on the length  
93 of CO<sub>2</sub> release period ( $CRP$ ) and the CO<sub>2</sub> release rate (Fig. 1b). The variations of NEP thus  
94 should be innovatively attributed to these decomposed components. A strong spatial correlation  
95 between mean annual NEP and length of CO<sub>2</sub> uptake period has been reported in evergreen  
96 needle- and broad-leaved forests (Churkina et al., 2005; Richardson et al., 2013; Keenan et al.,  
97 2014), whereas atmospheric inversion data and vegetation photosynthesis model indicated a  
98 dominant role of the maximal carbon uptake rate (Fu et al., 2017; Zhou et al., 2017). However,  
99 the relative importance of these phenological and physiological indicators for the spatially  
100 varying NEP remains unclear.

101 In this study, we decomposed annual NEP into  $U$  and  $R$ , and explored the local indicators  
102 for spatially varying NEP. Based on the eddy-covariance fluxes from FLUXNET2015 Dataset  
103 (Pastorello et al., 2017) and the atmospheric inversion product (Rödenbeck et al., 2018), we  
104 examined the relationship between NEP and its direct components. In addition, we used the  
105 observations to evaluate the spatial variations of NEP and  $IAV_{NEP}$  in the FLUXCOM product  
106 and a process-based model (CLM4.5) (Oleson et al., 2013). The major aim of this study is to

107 explore whether there are useful local indicators for the spatially varying NEP and  $I_{AV_{NEP}}$  in  
108 terrestrial ecosystems.

## 109 **2. Materials and Methods**

### 110 **2.1 Datasets**

111 Daily NEP observations of eddy covariance sites are obtained from the FLUXNET2015 Tier 1  
112 dataset (<http://fluxnet.fluxdata.org/data/fluxnet2015-dataset/>). The FLUXNET2015 dataset  
113 provides half-hourly data of carbon, water and energy fluxes at over 210 sites that are  
114 standardized and gap-filled (Pastorello et al., 2017). However, time series of most sites are still  
115 too short for the analysis of inter-annual variation in NEP. So only the sites that provided the  
116 availability of eddy covariance flux measurements for at least 5 years are selected. This leads to  
117 a global dataset of 72 sites with different biomes across different climatic regions. Based on the  
118 biome classification from the International Geosphere-Biosphere Programme (IGBP) provided  
119 for the FLUXNET2015 sites, the selected sites include 35 forests (FOR), 15 grasslands (GRA),  
120 11 croplands (CRO), 4 wetlands (WET), 2 shrublands (SHR) and 5 savannas (SAV) (Fig. S1  
121 and Table S1).

122 The Jena CarboScope Inversion product compiles from high precision measurements of  
123 atmospheric  $CO_2$  concentration with simulated atmospheric transport (Rödenbeck et al., 2018).  
124 Here, we used the daily land-atmosphere  $CO_2$  fluxes from the s85\_v4.1 version at a spatial  
125 resolution of  $5^\circ \times 3.75^\circ$ . Considering the relatively low spatial resolution of the Jena Inversion  
126 product, the daily fluxes were only used to calculate the local indicators for the spatially varying  
127 NEP at the global scale.

128 Daily NEP simulations from Community Land Model version 4.5 (CLM4.5) were also used  
129 to calculate the local indicators for the spatially varying NEP at the corresponding flux tower  
130 sites. We ran the CLM4.5 model from 1985 to 2010 at a spatial resolution of  $1^\circ$  with CRUNECF  
131 meteorological forcing. Here, NEP was derived as the difference between GPP and TER, and  
132 TER was calculated as the sum of simulated autotrophic and heterotrophic respiration. The daily  
133 outputs from CLM4.5 were used to calculate the local indicators for the spatially varying NEP  
134 both at the global scale and at the FLUXNET site level.

135 The FLUXCOM product presents an upscaling of carbon flux estimates from 224 flux  
 136 tower sites based on multiple machine learning algorithms and meteorological drivers (Jung et  
 137 al., 2017). To be consistent with the meteorological forcing of Jena Inversion product and the  
 138 CLM4.5 model, we used the FLUXCOM CRUNCEPv6 products. In addition, in order to reduce  
 139 the uncertainty caused by machine-learning methods, we averaged all the FLUXCOM  
 140 CRUNCEPv6 products with different machine-learning methods. It should be noted that the  
 141 inter-annual variability of FLUXCOM product is only driven by climatic conditions, the effects  
 142 of land use and land cover change are not represented. The FLUXCOM NEP product is  
 143 downloaded from the Data Portal of the Max Planck Institute for Biochemistry  
 144 (<https://www.bgc-jena.mpg.de>). Daily outputs from FLUXCOM for the period 1985-2010 at 0.5°  
 145 spatial resolution were used to calculate the local indicators for the spatially varying NEP both  
 146 at the global scale and at the FLUXNET site level.

## 147 **2.2 Decomposition of NEP and the calculations for its local indicators**

148 The annual NEP of a given ecosystem can be defined numerically as the difference between the  
 149 CO<sub>2</sub> uptake and release. As illustrated in Figure 2b:

$$150 \quad \quad \quad NEP = U - R \quad \quad \quad (1)$$

151 These components of NEP contain both photosynthesis and respiration flux, which directly  
 152 indicate the net CO<sub>2</sub> exchange of an ecosystem. The total CO<sub>2</sub> uptake flux ( $U$ ) and the total CO<sub>2</sub>  
 153 release flux ( $R$ ) can be further decomposed as:

$$154 \quad \quad \quad U = \bar{U} \times CUP \quad \quad \quad (2)$$

$$155 \quad \quad \quad R = \bar{R} \times CRP \quad \quad \quad (3)$$

156 where the  $\bar{U}$  (g C m<sup>-2</sup> d<sup>-1</sup>) is the mean daily CO<sub>2</sub> uptake over  $CUP$  (d yr<sup>-1</sup>) and  $\bar{R}$  (g C m<sup>-2</sup> d<sup>-1</sup>)  
 157 represents the mean daily CO<sub>2</sub> release over  $CRP$  (d yr<sup>-1</sup>). In addition, we further tested the  
 158 relationship between annual NEP and the ratio of  $\frac{U}{R}$  (i.e.,  $NEP \propto \frac{U}{R}$ ). Ecologically, the ratio of  
 159  $\frac{U}{R}$  reflects the relative strength of the ecosystem CO<sub>2</sub> uptake. Therefore, NEP in any year of any  
 160 given ecosystem can be expressed as:

161 
$$NEP = \beta \cdot \ln\left(\frac{U}{R}\right) \quad (4)$$

162 where the parameter  $\beta$  represents the slope of the linear relationship of  $NEP \propto \ln\left(\frac{U}{R}\right)$ . Based  
 163 on the definitions of  $U$  and  $R$ , the ratio  $\frac{U}{R}$  can be further written as:

164 
$$\frac{U}{R} = \frac{\bar{U}}{\bar{R}} \cdot \frac{CUP}{CRP} \quad (5)$$

165 Ecologically, the ratio of  $\frac{\bar{U}}{\bar{R}}$  reflects the relative physiological difference between  
 166 ecosystem CO<sub>2</sub> uptake and release strength, while the ratio of  $\frac{CUP}{CRP}$  is an indicator of net  
 167 ecosystem CO<sub>2</sub> exchange phenology. Environmental changes may regulate these ecological  
 168 processes and ultimately affect the ecosystem NEP. The slope  $\beta$  indicates the response sensitivity  
 169 of NEP to the changes in phenology and physiological processes. All of  $\beta$ ,  $\frac{CUP}{CRP}$  and  $\frac{\bar{U}}{\bar{R}}$  were  
 170 then calculated from the selected eddy covariance sites and the corresponding pixels of these  
 171 sites in models. These derived indicators from eddy covariance sites were then used to  
 172 benchmark the results extracted from the same locations in models.

173 **2.4 Calculation of the relative contributions**

174 We further quantified the relative contributions of  $\frac{\bar{U}}{\bar{R}}$  and  $\frac{CUP}{CRP}$  in driving the spatial variations  
 175 in NEP:

176 
$$NEP = f\left(\frac{\bar{U}}{\bar{R}}, \frac{CUP}{CRP}\right) \quad (6)$$

177 We used a relative importance analysis method to quantify the relative contributions of  
 178 each ratio to the spatial variations in NEP. The algorithm was performed with the “ralaimpo”  
 179 package in R (R Development Core Team, 2011). The “ralaimpo” package is based on variance  
 180 decomposition for multiple linear regression models. We chose the most commonly used method  
 181 named “Lindeman-Merenda-Gold (LMG)” (Grömping, 2007) from the methods provided by  
 182 the “ralaimpo” package. This method allows us to quantify the contributions of explanatory  
 183 variables in a multiple linear regression model. Across the 72 FLUXNET sites, we quantified  
 184 the relative importance of  $\frac{\bar{U}}{\bar{R}}$  and  $\frac{CUP}{CRP}$  to cross-site changes in NEP.

## 185 **3. Results**

### 186 **3.1 The relationship between NEP and its direct components**

187 To find local indicators for the spatially varying NEP in terrestrial ecosystems, we tested the  
188 relationship between NEP and its direct components ( $U$  and  $R$ ) across the 72 flux-tower sites.  
189 The results showed that annual NEP was closely related with the ratio of  $\frac{U}{R}$  (Fig. S2). The  
190 logarithmic correlations between annual NEP and  $\frac{U}{R}$  were significant at all sites (Fig. 1a), and  
191  $\sim 90\%$  of  $R^2$  falling within a range from 0.7 to 1 (Fig. 1c).

192 In addition, the relationship between NEP and  $\frac{U}{R}$  was also verified by the atmospheric  
193 inversion product (i.e., Jena CarboScope Inversion). The control of  $\frac{U}{R}$  on annual NEP was  
194 robust in most global grid cells (i.e.  $0.6 < R^2 < 1$ ). The explanation of  $\frac{U}{R}$  was higher in 80% of  
195 the regions, but lower in North American (Fig. 2). These two datasets both showed that the  
196 indicator  $\frac{U}{R}$  could successfully capture the variability in annual NEP.

### 197 **3.2 Local indicators for spatially varying NEP**

198 Across the 72 flux-tower sites, the spatial changes in mean annual NEP were significantly  
199 correlated to  $\ln\left(\frac{U}{R}\right)$  ( $R^2 = 0.65$ ,  $P < 0.01$ ) (Fig. 3a). This finding suggested that the mean annual  
200 ratio  $\ln\left(\frac{U}{R}\right)$  is a good indicator for cross-site variation in NEP. By contrast, the spatial variation  
201 of  $\text{IAV}_{\text{NEP}}$  was moderately explained by the slope (i.e.,  $\beta$ ) of the temporal correlation between  
202 NEP and  $\ln\left(\frac{U}{R}\right)$  at each site ( $R^2 = 0.39$ ,  $P < 0.01$ ; Fig. 3b) rather than  $\ln\left(\frac{U}{R}\right)$  (Fig. S3). The  
203 wide range of ratio  $\beta$  reveals a large divergence of NEP sensitivity across biomes, ranging from  
204  $121 \pm 118 \text{ g C m}^{-2} \text{ yr}^{-1}$  in shrubland to  $473 \pm 112 \text{ g C m}^{-2} \text{ yr}^{-1}$  in cropland.

205 The decomposition of indicator  $\frac{U}{R}$  into  $\frac{\bar{U}}{\bar{R}}$  and  $\frac{CUP}{CRP}$  allowed us to quantify the relative  
206 importance of these two ratios in driving NEP variability. The linear regression and relative  
207 importance analysis showed a more important role of  $\frac{CUP}{CRP}$  (58%) than  $\frac{\bar{U}}{\bar{R}}$  (42%) in explaining  
208 the cross-site variation of NEP (Fig. 4). Therefore, the spatial distribution of mean annual NEP



209 was mostly driven by the phenological rather than physiological changes.

### 210 **3.3 Simulated spatial variations in NEP by models**

211 We further used these two simple indicators (i.e.,  $\frac{U}{R}$  and  $\beta$ ) to evaluate the simulated spatial  
212 variations of NEP by the compiled global product (i.e., FLUXCOM) and a widely-used process-  
213 based model at the FLUXNET site level (i.e., CLM4.5). We found that the low spatial variation  
214 of mean annual NEP in FLUXCOM and CLM4.5 could be inferred from their more converging  
215  $\ln\left(\frac{U}{R}\right)$  than flux-tower measurements (Fig. 5). The underestimated variation of  $IAV_{NEP}$  in these  
216 modeling results was also clearly shown by the smaller  $\beta$  values (268.22, 126.00 and 145.08 for  
217 FLUXNET, FLUXCOM and CLM4.5, respectively) (Fig. 5b).

218 In addition, the spatial variations of NEP and  $IAV_{NEP}$  were associated with the spatial  
219 resolution of the product (Marcolla et al., 2017). Considering the scale mismatch between  
220 FLUXNET sites and the gridded product, we run the same analysis at the global scale based on  
221 Jena Inversion product. At the global scale, the spatial variation of mean annual NEP can be also  
222 well indicated by  $\ln\left(\frac{U}{R}\right)$  (Fig. 6). The larger C uptake in FLUXCOM resulted from its higher  
223 simulations for  $\ln\left(\frac{U}{R}\right)$ . Furthermore, the larger spatial variation of  $IAV_{NEP}$  in CLM4.5 could be  
224 inferred from the indicator  $\beta$ .

## 225 **4. Discussion**

### 226 **4.1 New perspective for locating the major and sustainable land C sinks**

227 Large spatial differences of mean annual NEP and  $IAV_{NEP}$  have been well-documented in  
228 previous studies (Jung et al., 2017; Marcolla et al., 2017; Fu et al., 2019). Here we provide a  
229 new perspective for quantifying the spatially varying NEP by tracing annual NEP into several  
230 local indicators. Therefore, these traceable indicators could provide useful constraints for  
231 predicting annual NEP, especially in areas without eddy-covariance towers.

232 Typically, the C sink capacity and its stability of a specific ecosystem are characterized  
233 separately (Keenan et al., 2014; Ahlstrom et al., 2015; Jung et al., 2017). Here we integrated  
234 NEP into two simple indicators that could directly locate the major and sustainable land C sink.

235 Among biomes, forests and croplands had the largest  $\ln\left(\frac{U}{R}\right)$  and  $\beta$ , indicating the strongest and  
236 the most unstable C sink in forests and croplands, respectively. However, the relatively lower  $\beta$   
237 in shrublands and savannas should be interpreted cautiously. There are very few semi-arid  
238 ecosystems in the FLUXNET sites, while they represent a large portion of land at the global  
239 scale and have been shown to substantially control the interannual variability of NEP (Ahlström  
240 et al., 2015). The highest  $\beta$  in croplands implies that the rapid global expansion of cropland may  
241 enlarge the  $I\Delta V_{NEP}$  on the land. In fact, the cropland expansion has been confirmed as one  
242 important driver of the recent increasing global vegetation growth peak (Huang et al., 2018) and  
243 atmospheric  $CO_2$  seasonal amplitude (Gary et al., 2014; Zeng et al., 2014).

#### 244 4.2 Phenology-dominant spatial distribution of mean annual NEP

245 Recent studies have demonstrated that the spatiotemporal variations in terrestrial gross primary  
246 productivity are jointly controlled by plant phenology and physiology (Xia et al., 2015; Zhou et  
247 al., 2016). Here we demonstrated the dominant role of the phenology indicator  $\frac{CUP}{CRP}$  in driving  
248 the spatial difference of mean annual NEP. The reported low correlation between mean annual  
249 NEP and the physiological indicator  $\frac{\bar{U}}{\bar{R}}$  could partly be attributed to the convergence of  $\frac{\bar{U}}{\bar{R}}$   
250 across FLUXNET sites (Fig. S4).

251 The convergent  $\frac{\bar{U}}{\bar{R}}$  across sites was first discovered by Churkina *et al.* (2005) as  $2.73 \pm 1.08$   
252 across 28 sites, which included DBF, EBF and crop/grass. In this study, we found the  $\frac{\bar{U}}{\bar{R}}$  across  
253 the 72 sites is  $2.71 \pm 1.61$ , which validates the discovery by Churkina *et al.* However, the  $\frac{\bar{U}}{\bar{R}}$   
254 varied among biomes ( $2.86 \pm 1.56$  for forest,  $2.16 \pm 1.14$  for grassland,  $3.47 \pm 1.98$  for cropland,  
255  $2.89 \pm 1.47$  for wetland,  $1.89 \pm 1.10$  for shrub,  $1.83 \pm 0.88$  for savanna). This spatial convergence  
256 of  $\frac{\bar{U}}{\bar{R}}$  at the ecosystem level provides important constraints for global models that simulate  
257 various physiological processes (Peng et al., 2015; Xia et al., 2017). These findings imply that  
258 the phenology changes will greatly affect the locations of the terrestrial carbon sink by  
259 modifying the length of carbon uptake period (Richardson et al., 2013; Keenan et al., 2014).

#### 260 4.3 The simulated local indicators from gridded products

261 This study showed that the considerable spatial variations in mean annual NEP and  $IAV_{NEP}$  from  
262 global gridded products could also be inferred from their local indicators. The low variations of  
263  $\frac{U}{R}$  ratio in CLM4.5 could be largely due to their simple representations of the diverse terrestrial  
264 plant communities into a few plant functional types with parameterized properties (Cui et al.,  
265 2019; Sakschewski et al., 2015). In addition, the higher  $\frac{U}{R}$  ratio from FLUXCOM product  
266 indicated its widely reported larger C uptake (Fig. 6) (Jung et al., 2020). Meanwhile, the  
267 ignorance of fire, land-use change and other disturbances could lead to the smaller  $\beta$  by allowing  
268 for only limited variations of phenological and physiological dynamics (Reichstein et al., 2014;  
269 Kunstler et al., 2016). Although the magnitude of  $IAV_{NEP}$  depends on the spatial resolution  
270 (Marcolla et al., 2017), we recommend future model benchmarking analyses to use not only the  
271 global product compiled from machine-learning method (Bonan et al., 2018) but also the site-  
272 level measurements or indicators (i.e.,  $\ln(\frac{U}{R})$  and  $\beta$ ).

#### 273 4.4 Conclusions and further implications

274 In summary, this study highlights the changes in NEP and  $IAV_{NEP}$  over space on the land, and  
275 provides the  $\frac{U}{R}$  ratio and  $\beta$  as two simple local indicators for their spatial variations. These  
276 indicators could be helpful for locating the persistent terrestrial C sinks in where the  $\ln(\frac{U}{R})$   
277 ratio is high but the  $\beta$  is low. Their estimates based on observations are also valuable for  
278 benchmarking and improving the simulation of land-atmospheric C exchanges in Earth system  
279 models.

280 In addition, the findings in this study have some important implications for understanding  
281 the variation of NEP on the land. First, forest ecosystems have the largest annual NEP due to the  
282 largest  $\ln(\frac{U}{R})$  while croplands show the highest  $IAV_{NEP}$  because of the highest  $\beta$ . Second, the  
283 spatial convergence of  $\frac{U}{R}$  suggests a tight linkage between plant growth and the non-growing  
284 season soil microbial activities (Xia et al., 2014; Zhao et al., 2016). However, it remains unclear  
285 whether the inter-biome variation in  $\frac{U}{R}$  is due to different plant-microbe interactions between  
286 biomes. Third, the within-site convergent but spatially varying  $\beta$  needs better understanding.

287 Previous studies have shown that a rising standard deviation of ecosystem functions could  
288 indicate an impending ecological state transition (Carpenter and Brock, 2006; Scheffer et al.,  
289 2009). Thus, a sudden shift of the  $\beta$ -value may be an important early-warning signal for the  
290 critical transition of  $IAV_{NEP}$  of an ecosystem. Furthermore, considering the limited eddy-  
291 covariance sites with long-term observations, these findings need further validation once the  
292 longer time-series of measurements from more sites and vegetation types become available.

### 293 **Acknowledgements**

294 This work was financially supported by the National Key R&D Program of China  
295 (2017YFA0604600), National Natural Science Foundation of China (31722009, 41630528) and  
296 National 1000 Young Talents Program of China. This work used eddy covariance dataset  
297 acquired and shared by the FLUXNET community, including these networks: AmeriFlux,  
298 AfriFlux, AsiaFlux, CarboAfrica, CarboEuropeIP, CarboItaly, CarboMont, ChinaFlux, Fluxnet-  
299 Canada, GreenGrass, ICOS, KoFlux, LBA, NECC, OzFlux-TERN, TCOS-Siberia, and USCCC.  
300 The ERA-Interim reanalysis data are provided by ECMWF and processed by LSCE. The  
301 FLUXNET eddy covariance data processing and harmonization was carried out by the European  
302 Fluxes Database Cluster, AmeriFlux Management Project, and Fluxdata project of FLUXNET,  
303 with the support of CDIAC and ICOS Ecosystem Thematic Center, and the OzFlux, ChinaFlux  
304 and AsiaFlux offices.

305 *Data availability statement.* Eddy flux data are available at  
306 <http://fluxnet.fluxdata.org/data/fluxnet2015-dataset/>; the data supporting the findings of this  
307 study are available within the article and the Supplementary Information.

308 *Author contribution.* E. Cui and J. Xia devised and conducted the analysis. Y. Luo, S. Niu, Y.  
309 Wang and C. Bian provided critical feedback on the method and results. All authors contributed  
310 to discussion of results and writing the paper.

311 *Competing interests.* The authors declare that there is no conflict of interest.

312 **FIGURES**

313 **Figure 1** Relationship between annual NEP and  $\frac{U}{R}$  for 72 FLUXNET sites (of the form NEP =  
314  $\beta \cdot \ln\left(\frac{U}{R}\right)$ ). a, Dependence of annual NEP on the ratio between total CO<sub>2</sub> exchanges during net  
315 uptake ( $U$ ) and release ( $R$ ) periods (i.e.,  $\frac{U}{R}$ ). Each line represents one flux site with at least 5  
316 years of observations. b, Conceptual figure for the decomposition framework introduced in this  
317 study. Annual NEP can be quantitatively decomposed into the following indicators:  $NEP =$   
318  $U - R$ . c, Distribution of the explanation of  $\frac{U}{R}$  on temporal variability of NEP ( $R^2$ ) for  
319 FLUXNET sites.

320 **Figure 2** Relationship between annual NEP and  $\frac{U}{R}$  for Jena Inversion product (of the form  
321  $NEP = \beta \cdot \ln\left(\frac{U}{R}\right)$ ). The black box indicates the location of the sample.

322 **Figure 3** Contributions of the two indicators in explaining the spatial patterns of mean annual  
323 NEP and  $IAV_{NEP}$ . a, The relationship between annual mean NEP and  $\ln\left(\frac{U}{R}\right)$  across FLUXNET  
324 sites ( $R^2 = 0.65$ ,  $P < 0.01$ ). The insets show the variation of  $\ln\left(\frac{U}{R}\right)$  for different terrestrial  
325 biomes. b, The explanation of  $\beta$  on  $IAV_{NEP}$  ( $R^2 = 0.39$ ,  $P < 0.01$ ). The insets show the distribution  
326 of parameter  $\beta$  for different terrestrial biomes. The number of site-years at each site is indicated  
327 with the size of the point.

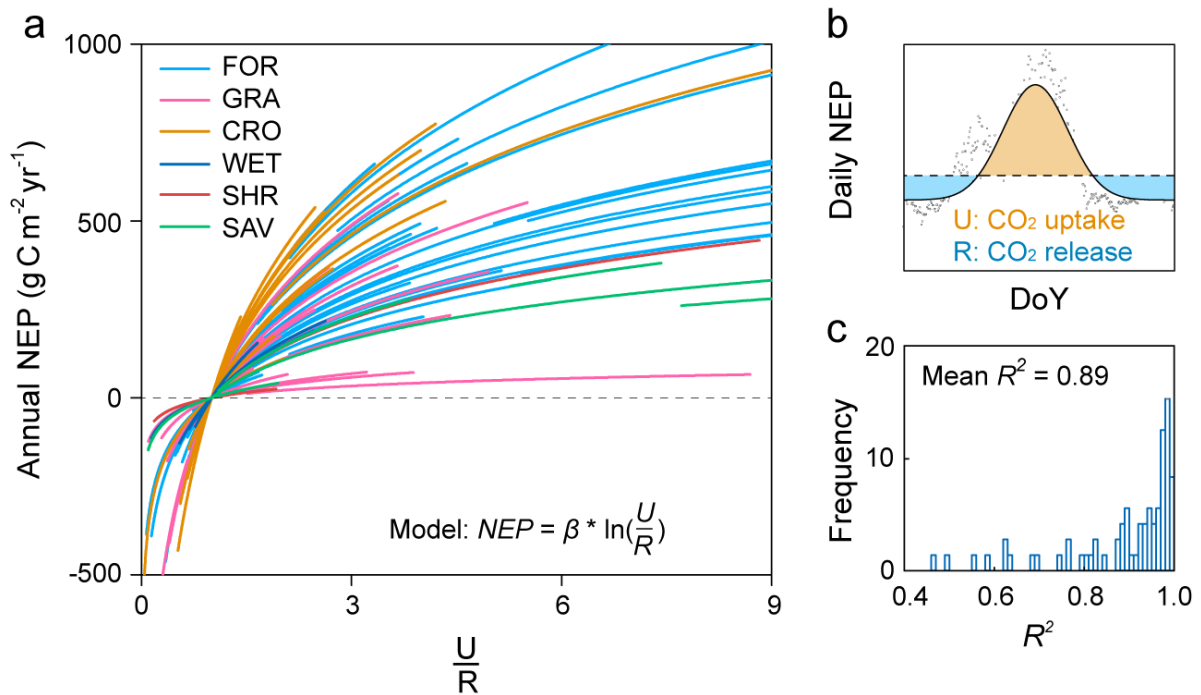
328 **Figure 4** The linear regression between  $\frac{U}{R}$  with  $\frac{CUP}{CRP}$  ( $R^2 = 0.71$ ,  $P < 0.01$ ) and  $\frac{\bar{U}}{\bar{R}}$  ( $R^2 = 0.09$ ,  
329  $P < 0.01$ ) across sites. The insets show the relative contributions of each indicator to the spatial  
330 variation of  $\frac{U}{R}$ . The number of site-years at each site is indicated with the size of the point.

331 **Figure 5** Representations of the spatially varying NEP and its local indicators in FLUXCOM  
332 product and the Community Land Model (CLM4.5) at the FLUXNET site level. a, The variation  
333 of mean annual NEP and  $IAV_{NEP}$  derives from FLUXNET, FLUXCOM and CLM4.5. Variation  
334 in mean annual NEP: the standard deviation of mean annual NEP across sites; Variation in  
335  $IAV_{NEP}$ : the standard deviation of  $IAV_{NEP}$  across sites. b, Representations of the local indicators  
336 for NEP in FLUXNET, FLUXCOM and CLM4.5. The corresponding distributions of  $\ln\left(\frac{U}{R}\right)$   
337 and  $\beta$  are shown at the top and right. Significance of the relationship between annual NEP and

338  $\ln\left(\frac{U}{R}\right)$  for each site is indicated by the circle: closed circles:  $P < 0.05$ ; open circles:  $P > 0.05$ . Note  
339 that the modeled results are from the pixels extracted from the same locations of the flux tower  
340 sites.

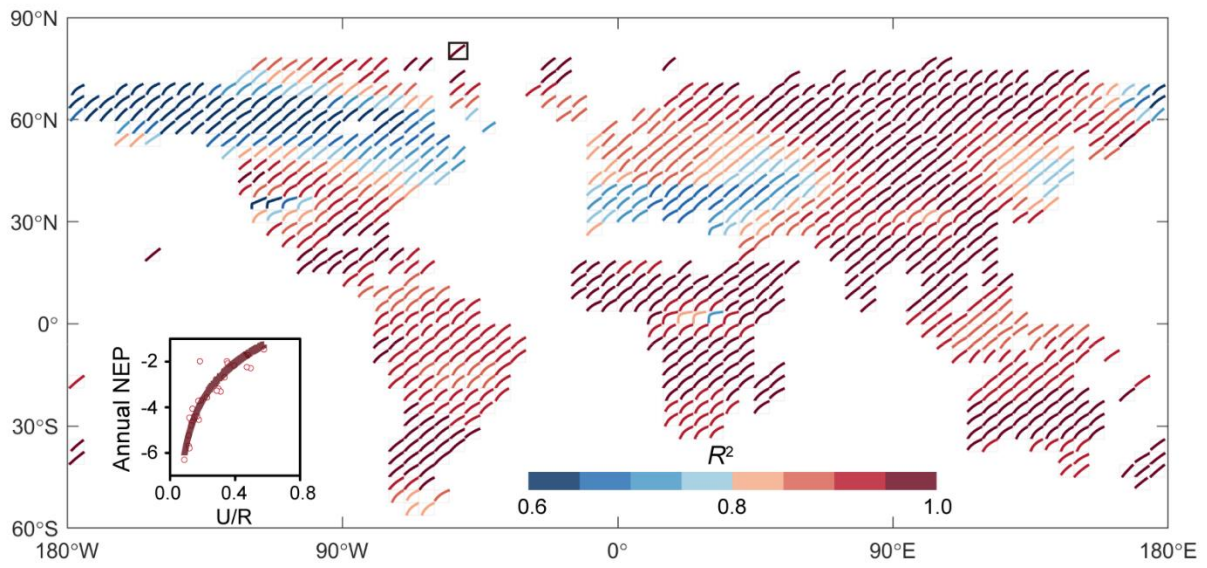
341 **Figure 6** Representations of the spatially varying NEP and its local indicators in FLUXCOM  
342 product and the Community Land Model (CLM4.5) at the global scale. a, The variation of mean  
343 annual NEP and  $I_{AV_{NEP}}$  derives from Jena Inversion, FLUXCOM and CLM4.5. Variation in  
344 mean annual NEP: the spatial variation of mean annual NEP; Variation in  $I_{AV_{NEP}}$ : the spatial  
345 variation of standard deviation in  $I_{AV_{NEP}}$ . b, Representations of the local indicators for NEP in  
346 Jena Inversion, FLUXCOM and CLM4.5.

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348  
 349 **Figure 1** Relationship between annual NEP and  $\frac{U}{R}$  for 72 FLUXNET sites (of the form  $NEP =$   
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 351 uptake ( $U$ ) and release ( $R$ ) periods (i.e.,  $\frac{U}{R}$ ). Each line represents one flux site with at least 5  
 352 years of data. **b**, Conceptual figure for the decomposition framework introduced in this study.  
 353 Annual NEP can be quantitatively decomposed into the following indicators:  $NEP = U - R$ . **c**,  
 354 Distribution of the explanation of  $\frac{U}{R}$  on temporal variability of FLUXNET NEP ( $R^2$ ) for  
 355 FLUXNET sites.

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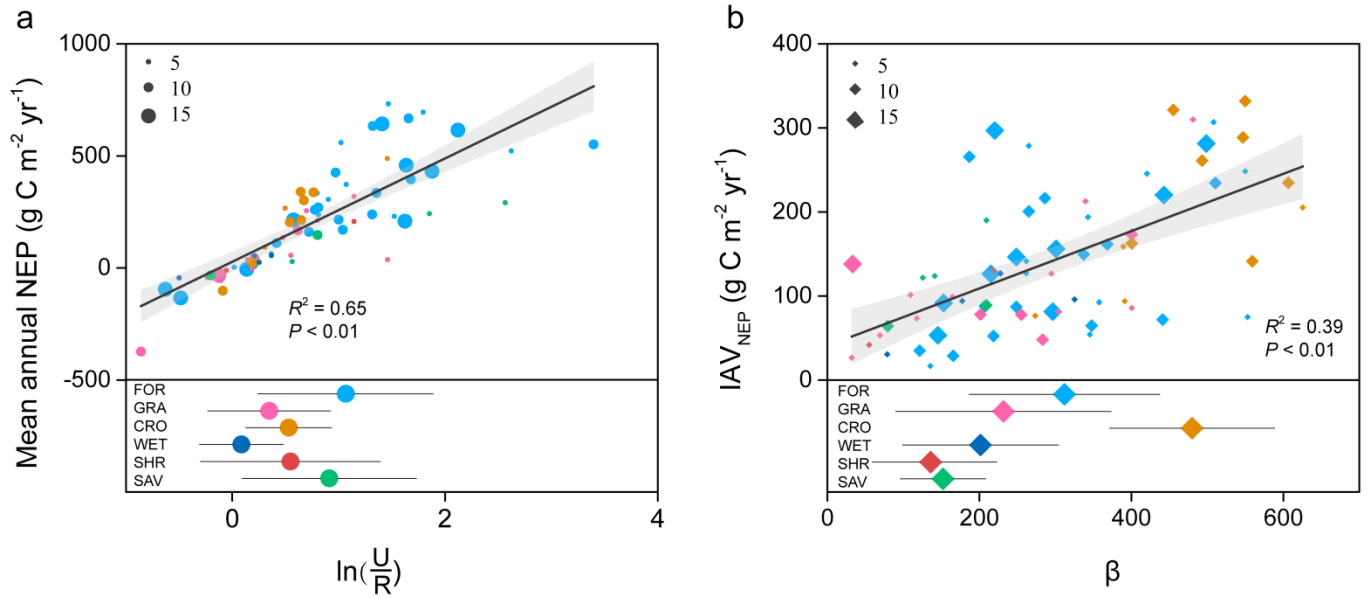
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358 **Figure 2** Relationship between annual NEP and  $\frac{U}{R}$  for Jena Inversion product (of the form

359  $NEP = \beta \cdot \ln\left(\frac{U}{R}\right)$ ). The black box indicates the location of the sample.

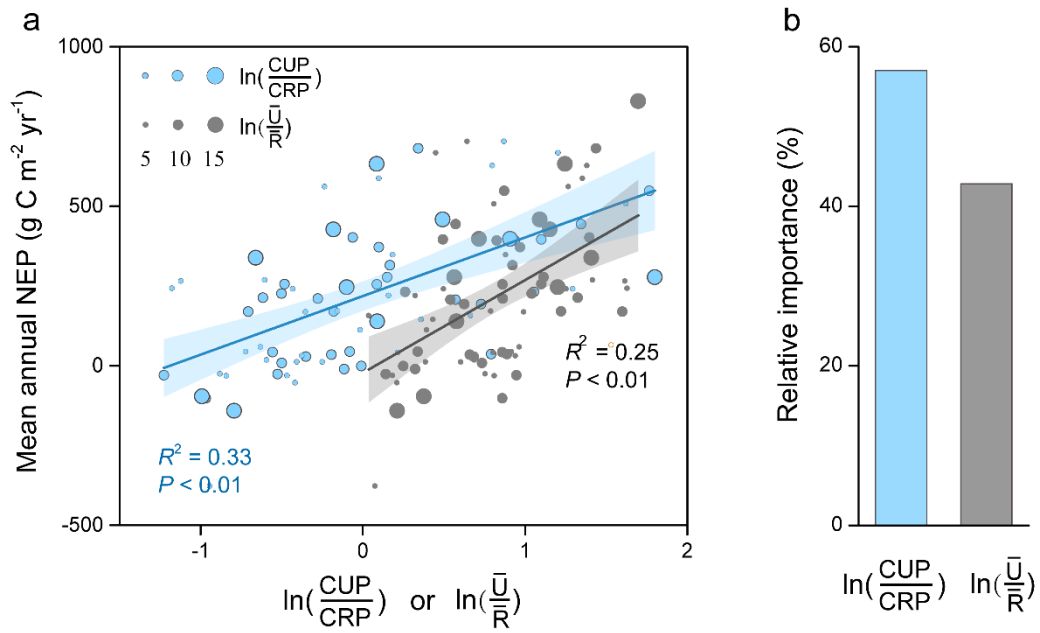
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361  
 362 **Figure 3** Contributions of the two indicators in explaining the spatial patterns of mean annual  
 363 NEP and IAV<sub>NEP</sub>. **a**, The relationship between annual mean NEP and  $\ln(\frac{U}{R})$  across FLUXNET  
 364 sites ( $R^2 = 0.65$ ,  $P < 0.01$ ). The insets show the variation of  $\ln(\frac{U}{R})$  for different terrestrial  
 365 biomes. **b**, The explanation of  $\beta$  on IAV<sub>NEP</sub> ( $R^2 = 0.39$ ,  $P < 0.01$ ). The insets show the distribution  
 366 of parameter  $\beta$  for different terrestrial biomes. The number of site-years at each site is indicated  
 367 with the size of the point.

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370 **Figure 4** The relative contributions of the local indicators in explaining the spatial patterns of

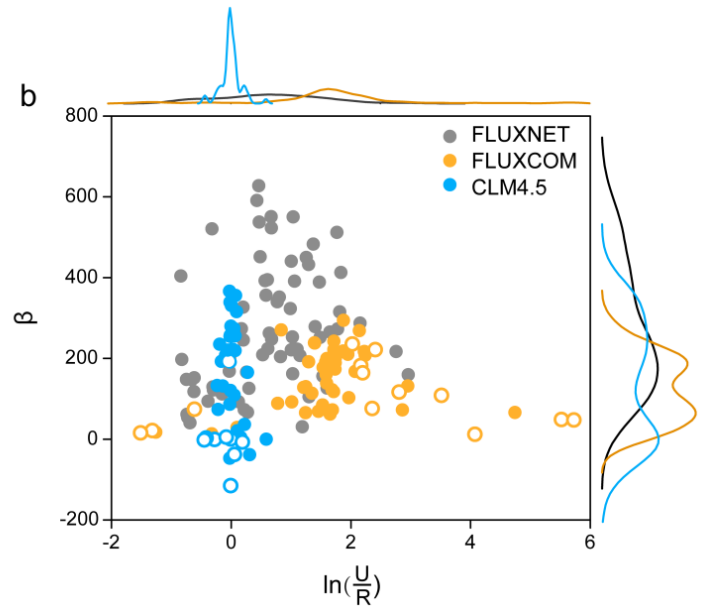
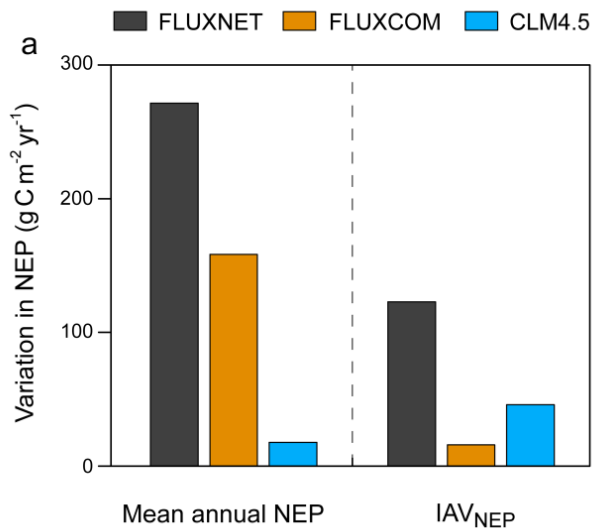
371 mean annual NEP. **a**, The linear regression between mean annual NEP with  $\frac{CUP}{CRP}$  ( $R^2 = 0.33$ ,  $P$

372  $< 0.01$ ) and  $\frac{\bar{U}}{\bar{R}}$  ( $R^2 = 0.25$ ,  $P < 0.01$ ) across sites. **b**, The relative contributions of each indicator

373 to the spatial variation of NEP. The number of site-years at each site is indicated with the size

374 of the point.

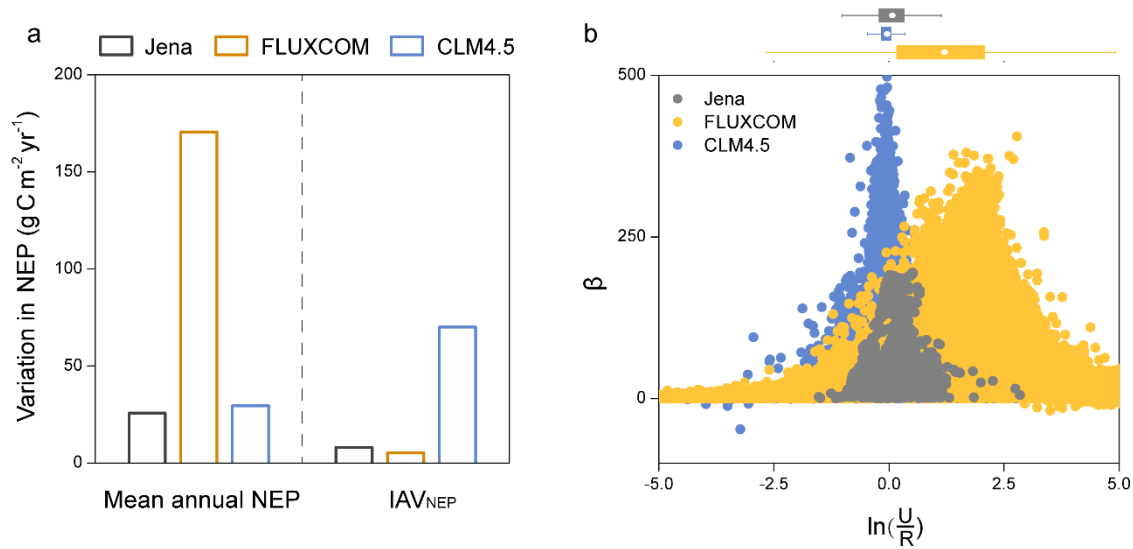
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377 **Figure 5** Representations of the spatially varying NEP and its local indicators in FLUXCOM  
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 380 in mean annual NEP: the standard deviation of mean annual NEP across sites; Variation in  
 381 IAV<sub>NEP</sub>: the standard deviation of IAV<sub>NEP</sub> across sites. **b**, Representations of the local indicators  
 382 for NEP in FLUXNET, FLUXCOM and CLM4.5. The corresponding distributions of  $\ln\left(\frac{U}{R}\right)$   
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 385 Note that the modeled results are from the pixels extracted from the same locations of the flux  
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388  
 389 **Figure 6** Representations of the spatially varying NEP and its local indicators in FLUXCOM  
 390 product and the Community Land Model (CLM4.5) at the global scale. a, The variation of mean  
 391 annual NEP and IAV<sub>NEP</sub> derives from Jena Inversion, FLUXCOM and CLM4.5. Variation in  
 392 mean annual NEP: the spatial variation of mean annual NEP; Variation in IAV<sub>NEP</sub>: the spatial  
 393 variation of standard deviation in IAV<sub>NEP</sub>. b, Representations of the local indicators for NEP in  
 394 Jena Inversion, FLUXCOM and CLM4.5.

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