# Comparing modified substrate induced respiration with selective inhibition (SIRIN) and $N_2O$ isotope approaches to estimate fungal contribution to denitrification in three arable soils under anoxic conditions

5 Lena Rohe<sup>1, 2, 3</sup>, Traute-Heidi Anderson<sup>2</sup>, Heinz Flessa<sup>2</sup>, Anette Giesemann<sup>2</sup>, Dominika Lewicka-Szczebak<sup>2, 4</sup>, Nicole Wrage-Mönnig<sup>5</sup>, Reinhard Well<sup>2</sup>

<sup>2</sup>Thünen Institute of Climate Smart Agriculture, Bundesallee 65, Braunschweig, Germany

Correspondence to: Lena Rohe (lena.rohe@ufz.de)

Keywords: selective growth inhibition, <sup>15</sup>N site preference, fungal denitrification,  $C_2H_2$ , isotope endmember mixing approach,  $SP_{N2O}$  mixing balance,  $SP/\delta^{18}O$  mapping approach

# 20 Abstract

Pure culture studies provide evidence of the ability of soil fungi to produce nitrous oxide (N<sub>2</sub>O) during denitrification. Soil studies with selective inhibition indicated a possible dominance of fungal compared to bacterial N<sub>2</sub>O production in soil, which drew more attention to fungal denitrification. Analyzing the isotopic composition of N<sub>2</sub>O, especially the-<u>its</u><sup>15</sup>N site preference of N<sub>2</sub>O produced ( $SP_{N2O}$ ), showed that N<sub>2</sub>O of pure bacterial or fungal cultures differed in  $SP_{N2O}$  values, which might enable the quantification of fungal N<sub>2</sub>O based

25

30

10

This study aimed to identify the fungal contribution to  $N_2O$  emissions and determine fungal  $SP_{N2O}$  under anaerobic conditions in\_from repacked soil samples incubated under anaerobic conditions repacked soil samples by using different approaches to disentangle sources of  $N_2O$ . Three approaches were established (modified substrate induced respiration with selective inhibition (SIRIN) approach, endmember mixing approach (IEM) and the SP/ $\delta^{18}O$  mapping approach (SP/ $\delta^{18}O$  Map) to independently investigate the fungal fraction contributing to  $N_2O$  from denitrification. –Three soils were incubated under anaerobic conditions to promote denitrification

on the isotopic endmember signatures of N2O produced by fungi and bacteria.

35

While one treatment without microbial inhibition served as a control\_the other three treatments were amended with inhibitors to selectively inhibit bacterial, fungal or bacterial and fungal growth. These treatments were performed in three varieties. In one variety<sub>a</sub> the <sup>15</sup>N tracer technique was used to estimate the effect of  $N_2O$  reduction on  $N_2O$  produced, while two other varieties were performed under natural isotopic conditions but with and without acetylene.

with four treatments of the a modified substrate induced respiration with selective inhibition (SIRIN) approach.

40 Three approaches were established to estimate the  $N_2O$  production by a fungal community in soil: i) A modification of the SIRIN approach was used to calculate  $N_2O$  evolved from selected organism groups, and ii)  $SP_{N2C}$  values from the acetylated treatment were used in the isotope endmember mixing approach (IEM), and iii)

<sup>&</sup>lt;sup>1</sup>Helmholtz Centre for Environmental Research – UFZ, Department Soil System Sciences, Theodor-Lieser Str. 4, Halle, Germany

<sup>&</sup>lt;sup>3</sup>University of Göttingen, Department of Crop Sciences, Institute of Grassland Science, von-Siebold-Str. 8, 37075 Göttingen, Germany

<sup>&</sup>lt;sup>4</sup>Institute of Geological Sciences, University of Wrocław, pl. M. Borna 9, 50-204 Wrocław, Poland University of Göttingen, Centre for Stable Isotope Research and Analysis, Büsgenweg 2, 37077 Göttingen, Germany

<sup>&</sup>lt;sup>5</sup>University of Rostock, Agricultural and Environmental Faculty, Grassland and Fodder Sciences, Justus-Liebig-Weg 6, Rostock, Germany

the SP/8<sup>18</sup>O mapping approach (SP/8<sup>18</sup>O Map) was used to estimate the fungal contribution to N<sub>2</sub>O product and N2O reduction under anaerobic conditions from the non acetylated treatment.

All three approaches revealed a small fungal contribution to N<sub>2</sub>O fluxes ( $f_{ED}$ ) under anaerobic conditions in the

45 soils tested. Quantifying the fungal fraction with modified SIRIN was not successful -due large amounts of uninhibited N<sub>2</sub>O production and pre-incubation effects. In only one soil,  $f_{PD}$  using modified SIRIN could be estimated and resulted in 28±9 %, which was possibly overestimated as results obtained by IEM and SP/ $\delta^{18}$ O Map for this soil resulted in  $f_{FD}$  of below 15 and 20 %, respectively. As a consequence of the unsuccessful SIRIN approach, estimation of fungal SP<sub>N20</sub> values was impossible. For this soil, SP<sub>N20</sub> values of the fungal fraction ed with modified SIRIN could be compared with fungal SP<sub>N20</sub> endm

50

55

in the literature and indicated...

The three approaches tested revealed a small fungal contribution to N<sub>2</sub>O fluxes under anaerobic conditions in the tested. Quantifying the fungal fraction with modified SIRIN was only possible in one soil and totaled 0.28±0.09. This was higher than the results obtained by IEM and SP/5<sup>18</sup>O Map, which accounted zero to 0.20 of N<sub>2</sub>O produced to the fungal community.

To our knowledge, this study was the first attempt to quantify the fungal contribution to anaerobic N2O production by simultaneous application of three approaches, i.e. modified SIRIN, IEM and SP/8<sup>18</sup>O Map. While all successful methods coincided by suggesting a small or missing fungal contribution, further studies under eonditions ensuringwith stimulated larger fungal N2O fluxes by addeding fungal C substratessources preferred by fungi and an improved modified SIRIN approach, including alternative inhibitors, are needed to better crossvalidate the methods.

# 60

# 1. Introduction

65

The greenhouse gas nitrous oxide (N<sub>2</sub>O) contributes to global warming and to the depletion of the ozone layer in the stratosphere (Crutzen, 1970; IPCC, 2013). The largest anthropogenic N2O emissions originate from agricultural soils and are mainly produced during microbial nitrification, nitrifier denitrification and denitrification (Firestone and Davidson, 1989; Bremner, 1997; IPCC, 2013; Wrage-Mönnig et al., 2018). In order to find mitigation strategies for N2O emissions from arable soils, it is important to understand N2O sources and sinks and thus improve knowledge about the production pathways and the microorganisms involved.

70

Denitrification describes is the stepwise reduction of nitrate (NO<sub>3</sub><sup>-</sup>) to dinitrogen (N<sub>2</sub>), with the intermediates nitrite (NO2), nitric oxide (NO) and N2O (Knowles, 1982). While this entire reaction chain including the ability to reduce N<sub>2</sub>O to N<sub>2</sub> is found among bacterial denitrifiers, most fungi lack N<sub>2</sub>O reductase (Nos) (Shoun et al., 1992; Shoun et al., 2012; Higgins et al., 2018). For a long time, it was believed that solely bacteria are involved in N<sub>2</sub>O formation during denitrification (Firestone and Davidson, 1989); however, also several fungi are capable of denitrification (Bollag and Tung, 1972; Shoun et al., 1992). Denitrification describes the reduction of nitrate 75 with the intermediates nitrite (NO<sub>2</sub>), nitrie exide (NO) and N<sub>2</sub>O (Knewles, 1982).  $(NO_2)$  to dinitrogen  $(N_2)$ , While this entire reaction chain including the ability to reduce N2O to N2 is found among bacterial denitrifiers, mei lack N<sub>2</sub>O reductase (Nos). Recently, pure culture studies showed that N<sub>2</sub>O from funcal denitrification often accompanied with N<sub>2</sub>O from abiotic production (Phillips et al., 2016a; Phillips et al., 2016b), which may lead to overestimate the importance of fungal N2O production. Other studiesPure culture studies indicated, 80 however, that although only some fungal species (e.g. Fusarium strains) are performing respiratory denitrification, these may produce substantial amounts of N2O performing respiratory denitrification with

Feldfunktion geändert

85

105

fungal species were identified to be capable of respiratory denitrification, N<sub>2</sub>O produced by fungi may thus contribute largely to N<sub>2</sub>O from denitrification in soil. Firstly, fungi dominate the biomass in soil (up to 96 %) compared to bacteria in general and thus fungi could potentially play a dominant role in N2O production (Ruzicka et al., 2000; Braker and Conrad, 2011). Thus, A-a respiratory fungal-to-bacterial (F:B) ratio of 4 is typical for arable soils (Anderson and Domsch, 1975; Blagodatskaya and Anderson, 1998). Secondly, the fact that N<sub>2</sub>O is the major end product of fungal denitrification led to the assumption that the potential activity of fungal N<sub>2</sub>O production in soil may exceed that of bacteria, provided that both microbial groups have the same 90 specific denitrification activity (Shoun et al., 1992; Sutka et al., 2008). Thirdly, co-denitrification was found to often co-occur with fungal denitrification (Shoun and Tanimoto, 1991; Tanimoto et al., 1992). During this fungal pathwayco-denitrification, a hybrid N<sub>2</sub>O is formed using one N atom from NO<sub>2</sub><sup>-</sup> and one N atom from compounds like azide or ammonium (NH4<sup>+</sup>) for N<sub>2</sub>O production (Tanimoto et al., 1992; Shoun et al., 1992; Rohe et al., 2017; Spott et al., 2011). This pathway was found to contribute about 92 % to N<sub>2</sub>O produced in an 95 incubation experiment with a grassland soil under anaerobic conditions A-<sup>45</sup>N tracing approach was used to identify and quantify co-denitrification, which contributed about 92% to N2O produced in an incubation experiment with a grassland soil under anacrobic conditions (Laughlin and Stevens, 2002). This again stresses the large potential N<sub>2</sub>O production by fungi. However, Additionally, in as shown in pure culture studies, not only co-denitrification, but also abiotic N<sub>2</sub>O formation may co-occur with fungal denitrification (Phillips et al., 2016a; Phillips et al., 2016b; Rohe et al., 2017)-, which may contribute to N<sub>2</sub>O production but potentially lead to

substantial amounts of N<sub>2</sub>O production (Higgins et al., 2018; Keuschnig et al., 2020). Even though only a few

100 overestimation of the importance of fungal N<sub>2</sub>O production. However, pathway differentiation is still challenging.

Soil incubation experiments could serve to differentiate between N2O produced by fungi and bacteria during denitrification by the application of two antibiotics: streptomycin and cycloheximide, which inhibit bacterial or fungal growth, respectively, by inhibition of the protein biosynthesis. This method is known as substrate induced

respiration with selective inhibition (SIRIN) (Anderson and Domsch, 1975; Laughlin and Stevens, 2002; Crenshaw et al., 2008; Blagodatskaya et al., 2010; Long et al., 2013). A few studies used a modification of this method for N2O analysis and found a greater decrease of N2O production with fungal than with bacterial growth inhibition (e.g. 89 vs. 23\_% decrease, respectively (Laughlin and Stevens, 2002)), indicating that fungi might 110 dominate N<sub>2</sub>O production (Laughlin and Stevens, 2002; McLain and Martens, 2006; Crenshaw et al., 2008; Blagodatskaya et al., 2010; Long et al., 2013; Chen et al., 2014; Chen et al., 2015).

Analysing the isotopic composition of N<sub>2</sub>O might also be a promising tool to distinguish between N<sub>2</sub>O from bacterial and fungal denitrification and other pathways. Especially, the isotopomer ratios of N<sub>2</sub>O (i.e. N<sub>2</sub>O molecules with the same bulk <sup>15</sup>N isotopic enrichment but showing different positions of <sup>15</sup>N in the linear N<sub>2</sub>O

- 115 molecule (Ostrom and Ostrom, 2017)) in pure culture studies showed differences in N2O of bacterial and fungal denitrification (Sutka et al., 2006; Sutka et al., 2008; Frame and Casciotti, 2010; Rohe et al., 2014a; Rohe et al., 2017). This --and-might be suitable for distinguishing between N2O produced by bacteria or fungi under denitrifying conditions. Isotopomer ratios of N<sub>2</sub>O can be expressed as  $^{15}N$  site preference (SP<sub>N2O</sub>), i.e. the difference between  $\delta^{15}N$  of the central and terminal N-position of the asymmetric N<sub>2</sub>O molecule (Toyoda and
- 120 Yoshida, 1999). The SP<sub>N2O</sub> values of N<sub>2</sub>O of six pure fungal cultures was between 16 and 37 ‰ (Sutka et al., 2008; Rohe et al., 2014a; Maeda et al., 2015; Rohe et al., 2017), whereas several bacteria produced N<sub>2</sub>O with SP<sub>N20</sub> values between -7.5 and +3.5 ‰ during denitrification (Toyoda et al., 2005; Sutka et al., 2006; Rohe et al.,

2017). However, the  $SP_{N2O}$  value of N<sub>2</sub>O produced by pure bacterial cultures during nitrification is approximately 33 ‰ and interferes with  $SP_{N2O}$  values of fungal denitrification (Sutka et al., 2006; Sutka et al., 2008; Rohe et al., 2014a). This demonstrates the difficulty to use  $SP_{N2O}$  values as an indicator for different

125

130

soil type.

organism groups contributing to N<sub>2</sub>O production from soil, where different pathways may co-occur. While it is generally assumed that SP<sub>prodN2O</sub> values of N<sub>2</sub>O produced by fungal pure cultures— during denitrification is transferable to N<sub>2</sub>O produced by fungal soil communities, this has not yet been proven. Until now, studies reporting possible ranges of fungal contributions to N<sub>2</sub>O fluxes from soil were based on SP<sub>prodN2O</sub> values of pure cultures (Köster et al., 2013b; Zou et al., 2014; Lewicka-Szczebak et al., 2017; Senbayram et al., 2018; Senbayram et al., 2020; Lewicka-Szczebak et al., 2014). but uncertainty of this approach arose from the fact that the full range of SP<sub>prodN2O</sub> values is between 16 and 37 ‰ have been reported (Sutka et al., 2008; Maeda et al., 2015; Rohe et al., 2017). It would thus be useful to constrain fungal SP<sub>prodN2O</sub> values for a specific soil or

Feldfunktion geändert

Feldfunktion geändert

- Although SP<sub>N2O</sub> values are independent of isotopic signatures of the precursors, δ<sup>15</sup>N and δ<sup>18</sup>O values of produced N<sub>2</sub>O (δ<sup>15</sup>N<sup>bulk</sup><sub>N2O</sub> and δ<sup>18</sup>O<sub>N2O</sub>, respectively) result from the isotopic signature of the precursor and isotopic fractionation during N<sub>2</sub>O production (Toyoda et al., 2005; Frame and Casciotti, 2010). <u>Regarding δ<sup>18</sup>O<sub>N2O</sub>, a complete exchange of oxygen (O) between NO<sub>3</sub><sup>-</sup> and soil water can be assumed and consequently, one can use the δ<sup>18</sup>O values of soil water for interpretation of δ<sup>18</sup>O<sub>N2O</sub> values (Lewicka-Szczebak et al., 2014; Kool et al., 2009; Snider et al., 2009). <u>However, i</u>Interpretation of δ<sup>18</sup>O<sub>N2O</sub> values from different microbial groups may be is even-more complex due to incomplete O exchange: because variations in the extent of O exchange during denitrification between water and denitrification intermediates<u>N</u> oxides altersaffect the final δ<sup>18</sup>O<sub>N2O</sub> value differently. (Garber and Hollocher, 1982; Aerssens et al., 1986; Kool et al., 2007; Rohe et al., 2014b; Rohe et al., 2017). <u>However, rR</u>ecently, fungal and bacterial N<sub>2</sub>O showed different ranges for δ<sup>18</sup>O<sub>N2O</sub> values and this
  </u>
- isotopic signature may also be helpful in differentiation of these pathways (Lewicka-Szczebak et al., 2016).
   Moreover, δ<sup>15</sup>N<sup>bulk</sup><sub>N2O</sub>, δ<sup>18</sup>O<sub>N2O</sub> and SP<sub>N2O</sub> values are in the course of denitrification affected by isotopic fractionation due to N<sub>2</sub>O reduction. During N<sub>2</sub>O reduction, the <sup>14</sup>N<sup>16</sup>O bond is preferentially broken compared to <sup>14</sup>N<sup>18</sup>O or <sup>15</sup>N<sup>16</sup>O, resulting in residual N<sub>2</sub>O<sub>7</sub> that is relatively isotopically enriched in <sup>15</sup>N and <sup>18</sup>O and shows larger SP<sub>N2O</sub> values compared to SP<sub>N2O</sub> values of N<sub>2</sub>O from denitrification without the reduction step (Popp et al.,
- 150 2002; Ostrom et al., 2007). <u>One possibility for Quantification quantifying theof N2O reduction to N2 during denitrification is the application of <sup>15</sup>N tracing experiments using <sup>15</sup>N enriched substrates possible bywithand analyzing analysing <sup>15</sup>N2 fluxes in <sup>15</sup>N tracing experiments using <sup>15</sup>N enriched substrates (Well et al., 2006; Lewicka-Szczebak et al., 2014). Another possibility is based on .- To quantify N2O reduction and the pathways producing N2O based on N2O isotopocules, which also enables to quantify -pathways producing N2O (i.e. N2O) with differing number or positions of N or O isotopes (Ostrom and Ostrom, 2017)).- <u>UnderIn this latter approach, i.e. the isotope mapping approach, isotope fractionation factors together with δ<sup>15</sup>N values of Precursors (δ<sup>15</sup>N<sub>NOX</sub>) as well as δ<sup>15</sup>N<sup>bulk</sup><sub>N2O</sub> and SP<sub>N2O</sub> values of N<sub>2</sub>O produced were used the isotope mapping approach was developed using isotope fractionation factors together with δ<sup>15</sup>N values of N<sub>2</sub>O are SP<sub>N2O</sub> values of N<sub>2</sub>O produced (Toyoda et al., 2011). Recently, this isotope mapping approach
  </u></u>
- 160 was further developed (SP/ $\delta^{18}$ O Map) using  $\delta^{18}$ O<sub>N20</sub> and SP<sub>N20</sub> values of N<sub>2</sub>O and  $\delta^{18}$ O values of precursors (Lewicka-Szczebak et al., 2014; Lewicka-Szczebak et al., 2017) by using .- This approach uses different slopes of N<sub>2</sub>O reduction and mixing lines in the  $\delta^{18}$ O – SP isotope plot-and. This approach -allows for differentiation of isotope effects due to N<sub>2</sub>O reduction and admixture of fungal N<sub>2</sub>O. <u>Hence, N<sub>2</sub>O reduction can be estimated</u>
  - 4

# together with the N<sub>2</sub>O mixing due to application of two isotopic signatures of N<sub>2</sub>O. For the Thus, $-SP/\delta^{18}O$ Map, the inhibition of N<sub>2</sub>O reduction is not needed-the N<sub>2</sub>O reduction to N<sub>2</sub> does not affect the outcome of the SP/ $\delta^{18}O$ Map, Map.<sub>2</sub>

Based on the above cited ranges for the isotopomer endmembers of fungal and bacterial denitrification, and assuming that only fungi and bacteria are responsible for N<sub>2</sub>O production, the fraction of fungal N<sub>2</sub>O can be calculated using the isotope endmember mixing approach (IEM) with  $SP_{N2O}$  values of N<sub>2</sub>O produced in soil ( $SP_{prod}$ ), provided N<sub>2</sub>O reduction, which is altering  $SP_{N2O}$  values of emitted N<sub>2</sub>O, does not occur (Ostrom et al., 2010; Ostrom and Ostrom, 2011). This can be ensured in laboratory experiments by inhibiting N<sub>2</sub>O reduction to N<sub>2</sub> using acetylene (C<sub>2</sub>H<sub>2</sub>) during anaerobic incubation experiments (Yoshinari and Knowles, 1976; Groffman et al., 2006; Well and Flessa, 2009; Nadeem et al., 2013). Hence, C<sub>2</sub>H<sub>2</sub> inhibition might be suitable to quantify

SP<sub>prod</sub> values in soils exhibiting significant N<sub>2</sub>O reduction and would thus allow quantification of fungal N<sub>2</sub>O

175 fluxes based on SP<sub>prod</sub> values. For the SP/8<sup>48</sup>O Map, the inhibition of N<sub>2</sub>O reduction is not needed. Hence, N<sub>2</sub>O reduction can be estimated together with the N<sub>2</sub>O mixing due to application of two isotopic signatures of N<sub>2</sub>O. While it is generally assumed that SP<sub>prod</sub> values of N<sub>2</sub>O produced by fungal pure cultures is transferable to N<sub>2</sub>O produced by fungal soil communities, this has not yet been proven. Until now, studies reporting possible ranges of fungal contributions to N<sub>2</sub>O fluxes from soil were based on SP<sub>prod</sub> values of pure arose from the fact that the full range of SP<sub>prod</sub> values is between 16 and 37%. It would thus be useful to constrain fungal SP<sub>prod</sub> values for a specific soil or soil type.

So far, the described methods for distinguishing between fungal and bacterial N<sub>2</sub>O emission have not been evaluated and compared in the same soil and their accuracy and possible bias remains unknown. We hypothesized that the fungal fraction contributing to N<sub>2</sub>O from denitrification in different soils using a modified SIRIN approach and isotopic methods will be correlated but not match exactly due to limited inhibitability of microbial communities and variability in *SP*<sub>N2O</sub> endmember values. Furthermore, successful application of the modified SIRIN approach with determined fungal fraction contributing to N<sub>2</sub>O from denitrification will yield fungal *SP*<sub>N2O</sub> endmember values within the range of values previously reported in the literature.

Therefore, this study aims at (i) determining the fungal contribution  $\overline{on-to}N_2O$  production by denitrification under anoxic conditions and glucose addition using three arable soils and three approaches (modified SIRIN, IEM and the  $SP/\delta^{I8}O$  Map), and to assess their usefulness in soil studies and thus assess factors of potential bias of the methods and (ii) to estimate the  $SP_{N2O}$  values from a-fungal soil communities and thus to evaluate the transferability of the pure culture range of the fungal  $SP_{N2O}$  endmember values.: modified SIRIN, IEM and the  $SP/\delta^{18}O$  Map, (ii) to compare the fungal contribution on  $N_2O$  production determined by these approaches and thus assess factors of potential bias of the methods, and (iii) to estimate the  $SP_{N2O}$  values from a fungal soil community and thus to evaluate the transferability of the pure culture range of the fungal  $SP_{N2O}$  endmember values.

# 2. Materials and Methods

# 2.1 Soil samples

165

170

200 All experiments were conducted with three arable soils differing in texture to provide different conditions for denitrification.\_soils differing in texture, C<sub>ore</sub> content, C/N ratio and pH. Thus<sub>T</sub> it was assumed that the soils harbour different denitrifying communities, i.e., different fractions of bacteria and fungi contributing to denitrification. One of the soils was sampled during a second season to evaluate if the fungal fraction contributing to N<sub>2</sub>O production is soil-specific or can be subject to seasonal change of microbial communities.
 As one-this soil was sampled at two different time points, we conducted four experiments and named the different experiments "Soil 1", "Soil 2", "Soil 3" and "Soil 4": ExperimentSoil 1 with loamy sand sampled in December 2012, ExperimentSoil 2 with sand sampled in January 2013, ExperimentSoil 3 with silt loam sampled in December 2012, and ExperimentSoil 4 with loamy sand sampled in June 2011 (Table 1).

- Soil samples of the upper 30 cm were collected in plastic bags aerated via cotton wool stoppers and stored at 210 6 °C for maximally two months. To get information about the initial soil status, the mineral nitrogen content ( $N_{min}$ ) of soil samples was determined before and after fertilization by extracting NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> with 0.01 M calcium chloride dihydrate (CaCl<sub>2</sub> · 2 H<sub>2</sub>O) according to ISO 14255 and analysing NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> concentrations in the extracts with a Continuous-Flow-Analyser (SKALAR, Germany) directly after sample collection. To get information about the initial soil status, Other soil characteristics (C and N content, soil pH
- 215 value, isotopic values of soil NO<sub>3</sub><sup>-</sup> and NO<sub>2</sub><sup>-</sup>) were analysed with samples of Soil 1, Soil 2 and Soil 3. Teotal contents of C and N in soil samples were analyzed analysed by dry combustion of grinded ground samples (LECO TruSpec, Germany). The soil pH was measured in 0.01 M CaCl<sub>2</sub>. The mineral nitrogen content (Nmin) of soil samples was determined before and after fertilization by extracting NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> with 0.01 M caleium chloride dihydrate (CaCl<sub>2</sub> 2 H<sub>2</sub>O) according to ISO 14255 and analyzing NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> concentrations in the
   220 extracts with a Continuous Flow Analyzer (SKALAR, Germany). The δ<sup>15</sup>N and δ<sup>18</sup>O values of NO<sub>3</sub><sup>-</sup> and NO<sub>2</sub><sup>-</sup>
- $(\delta^{15}N_{NOx} \text{ and } \delta^{18}O_{NOx}, \text{ respectively}) \text{ in soil extracts (with 0.01 M calcium chloride dihydrate (CaCl<sub>2</sub> · 2 H<sub>2</sub>O))}$ were analyzed analysed by the bacterial denitrifier method (Casciotti et al., 2002).

<u>To gain information on the Rrespiratory biomass of the three soils was were analyzed analyzed with for</u> substrate induced respiration (SIR) according to Anderson and Domsch (1978) and the respiratory F:B ratio was
 analyzed analyzed with substrate induced respiration with selective inhibition (SIRIN) in summer 2010 by a computer-generated selectivity analysis: "SIR-SBA 4.00" (Heinemeyer, copyright MasCo Analytik, Hildesheim,

Germany) (Anderson and Domsch, 1975). The scheme of glucose and growth inhibitor combinations is listed below in section "Methodological Methodical approach". For further The characteristics of the soils, see are listed in Table 1...

# 230 2.2 Methodological Methodical approach

<u>Twhe experimental setup with various measures is presented in the following sections and illustrated in Figure 1.</u> Important terms used and its descriptions are listed in Supplementary Material, Table S1.



# Figure 1: Schematic to represent the experimental setup of soil incubations with three varieties (*traced*, -C<sub>2</sub>H<sub>2</sub>, +C<sub>2</sub>H<sub>2</sub>). The methodical approach comprised a pre-experiment and the incubation experiment with a modified substrate induced respiration with selective inhibition (SIRIN) approach. Produced gas was analysed for its concentration (c(CO<sub>2</sub>) and c(N<sub>2</sub>O)) using gas chromatography (GC) and N<sub>2</sub>O was further analysed by isotope ratio mass spectrometry (IRMS). Please refer to the Material & Methods section for more information.

# 240 2.2.1 SIRIN pre-experiment

245

As in most studies applying the SIRIN method on N<sub>2</sub>O emissions (e. g. Laughlin and Stevens, 2002; Chen et al., 2014; Ladan and Jacinthe, 2016), a pre-experiment was conducted with samples collected in 2010, in order to get information about optimal substrate and inhibitor concentrations for substrate induced respiration with growth inhibition. The pre-experiments of the present study were conducted in two steps- as described in the original methods, i.-e., -CO<sub>2</sub> production under oxic conditions was analysed to estimate the substrate induced respiration by the SIR method (Anderson and Domsch, 1978) and the substrate induced respiration with selective inhibition by the SIRIN method (Anderson and Domsch, 1975) as follows.

In a first pre-experiment (Figure 1), t<sup>T</sup>he SIR method (Anderson and Domsch, 1978) was used to get information about the amount of respiratory biomass in soil-<u>under oxic conditions</u>. In this pre-experiment glucose served as substrate to initiate microbial growth (Anderson and Domsch, 1975). To this end, we added different concentrations of glucose (0.50, 0.75, 1.0, 1.5, 2.0, 3.0, 4.0, 5.0, 6.0 mg g<sup>-1</sup> dry weight (dw) soil) to find the optimal glucose concentration (c<sub>opt</sub>(glucose)), which is the glucose concentration that causes maximum initial respiration rates by analysing CO<sub>2</sub> production (Anderson and Domsch, 1978). C<sub>opt</sub>(glucose)) was 1.0 mg g<sup>-1</sup> for *ExperimentSoil* 2 (sand) and 1.5 mg g<sup>-1</sup> for *ExperimentSoils* 1, 3 and 4 (loamy sand and silt loam). Glucose

served as substrate to initiate microbial growth (Anderson and Domseh, 1975).
 We conducted<u>In a second pre-experiment (Figure 1), the</u> SIRIN method was used according to Anderson and Domsch (1975)\_for determining the respiratory F:B ratio-according to Anderson and Domsch (1975). The copp(glucose) determined in the first pre-experiment was used, while <u>s</u>electivity of the inhibitor combinations of streptomycin (bacterial respiratory inhibitor) and cycloheximide (fungal respiratory inhibitor) were tested with the followingthree concentrations<sub>7</sub> (0.75, 1.0, 1.5 mg g<sup>-1</sup> dw, respectively). The optimal concentration for

inhibition of fungal respiration was 0.75 mg g<sup>-1</sup> dw soil cycloheximide ( $c_{opt}$ (cycloheximide)) and for bacterial respiratory inhibition 1.0 mg g<sup>-1</sup> dw soil streptomycin ( $c_{opt}$ (streptomycin)). <u>According to As in the first pre-</u> experiment, CO<sub>2</sub> production under oxic conditions was analysed. The determined optimal concentrations of glucose, streptomycin and cycloheximide were used in the modified SIRIN approach, because we supposed that optimal concentrations for CO<sub>2</sub> respiration work as well for denitrification.

<u>Soil</u> (Year)	Soil texture	Soil type (WRB)	Location	C content [%]	N content [%]	NH4 <sup>+</sup> [mg N <mark>Lkg</mark> <sup>-1</sup> ]	NO3 <sup>-</sup> [mg N <u>Lkg</u> <sup>-1</sup> ]	pH (CaCl <sub>2</sub> )	δ <sup>15</sup> N <sub>NOx</sub> [‰] <sup>e</sup>	δ <sup>18</sup> O <sub>N</sub> Ox [‰] <sup>e</sup>	F:B <sup>f</sup>	Biomass <sup>g</sup> [µg C gdw <sup>-1</sup> soil]
1 (2012) 4 (2011)	Loamy sand	Haplic Luvisol	Braun- schweig <sup>a</sup>	1.43 (<0.01) <u>/</u>	0.10 (<0.01) <u>/</u>	$0.4 \\ (<0.1) \\ 1.0 \\ (0.4)$	$     \frac{14.1}{(2.1)}     \frac{11.0}{(0.3)} $	5.67 <u>/</u>	3.98 <u>/</u>	-4.82 <u>/</u>	2.6 <u>/</u>	234 _
2 (2013)	Sand	Gleyic Podzol	Wenne- bostel <sup>b</sup>	2.31 (0.04)	0.14 (<0.01)	<u>1.9</u> (0.2)	<u>6.6</u> (0.2)	5.54	0.73	-2.68	2.6	161
3 (2013)	Silt loam	Haplic Luvisol	Götting- en <sup>c</sup>	1.62 (0.02)	0.13 (<0.01)	n.d. <sup>d</sup>	<u>22.7</u> (<0.1)	7.38	4.18	2.32	4.9	389

Table 1: Soil characteristics of three arable soils from Germany used for incubation experiments (Exp. Soil) (standard deviation in brackets). Except for NH4<sup>+</sup> and NO3, soil characteristics of loamy sand were only analysed once for samples collected in 2012.

<sup>a</sup>Experimental Station of the Friedrich-Löffler Institute, Braunschweig, Germany

<sup>b</sup>private agricultural field North of Hannover, water protection area Fuhrberger Feld, Germany

270

<sup>c</sup>Reinshof Experimental Farm, Georg-August-University, Göttingen, Germany <sup>d</sup>not detectable (i.e. below detection limit of 0.<u>06</u> mg-kg<sup>-1</sup> NH<sub>4</sub><sup>+</sup>-N) <sup>e</sup>Isotopic values of natural soil NO<sub>3</sub><sup>-</sup> using the denitrifier method (Casciotti et al., 2002).

<sup>f</sup>Respiratory fungal-to-bacterial (F:B) ratio analyzed analyzed by SIRIN method (Anderson and Domsch, 1973, 1975).

<sup>g</sup>Respiratory biomass analyzed analysed by CO<sub>2</sub> production from SIR method (Anderson and Domsch, 1978).

# 275 2.2.2 Soil incubation with selective inhibition to determine N<sub>2</sub>O forming processes

The experimental design included two factors approaches, (i.) microbial inhibition by fungal and/or bacterial inhibitors and (ii.) activity of  $N_2O$  reductase analyzed analysed either by inhibition with  $C_2H_2$  or quantification by <sup>15</sup>N tracing (Figure 1). To address the microbial inhibition factor approach (i.), the SIRIN method for determination of the respiratory F:B ratio based on  $CO_2$  emission was modified to determine  $N_2O$  production by microbial groups. However, in contrast to previous studies by Laughlin and Stevens (2002), McLain and Martens (2006), Blagodatskaya et al. (2010) and Long et al. (2013), we did not pre-incubate the soil with the growth inhibitors, as this could result in changes of the microbial community (e.g. preferential growth of selected organisms). We intended to disturb microbial communities as little as possible.

- The soil was sieved (2 mm) and pre-incubated at 22 °C for five to seven days in the dark with cotton wool stoppers to allow respiration and aerobic conditions in soil bags. Four microbial inhibitor treatments (each in triplicate) with  $c_{opt}$ (glucose) for each soil were established:
  - A Control, without growth inhibitors

- B With streptomycin sulfate ( $C_{42}H_{84}N_{14}O_{36}S_3$ ) to inhibit bacterial growth
- C With cycloheximide  $(C_{15}H_{23}NO_4)$  to inhibit fungal growth
- 290 D With streptomycin and cycloheximide, to inhibit bacterial and fungal growth
- To address factor-the other approach (ii.), all microbial inhibitor treatments were conducted in three N<sub>2</sub>O reductase-varieties, i.e.: with <sup>15</sup>N-NO<sub>3</sub> fertilizer (variety "*traced*") to quantify N<sub>2</sub>O reduction to N<sub>2</sub>, with natural abundance NO<sub>3</sub><sup>-</sup> and 10 kPa C<sub>2</sub>H<sub>2</sub> in the headspace (variety "*+C<sub>2</sub>H<sub>2</sub>*") to block N<sub>2</sub>O reductase, and with natural abundance NO<sub>3</sub><sup>-</sup> but without blocking N<sub>2</sub>O reductase, i.e. no C<sub>2</sub>H<sub>2</sub> added (variety "*-C<sub>2</sub>H<sub>2</sub>*") (Figure 1). In total, there were 48 experimental treatments and 144 vessels (four <u>SoilsExperiments</u> with four inhibitor treatments (A,
  - B, C, D) and three varieties (*traced*,  $+C_2H_2$  and  $-C_2H_2$ ), each in triplicates).
- The soil was adjusted to 80\_% water filled pore space (WFPS) with distilled water.-and Simultaneously-to-that, the soil was fertilized with  $-NO_3^-$  (varieties  $-C_2H_2$ ,  $+C_2H_2$ , and *traced*). The soil sample used with Soil 4 was incubated prior to the other soils and was amended with 60 mg N kg<sup>-1</sup> NaNO<sub>3</sub>, while in agreement with other 300 experiments conducted in our laboratory, 50 mg N kg<sup>-1</sup> KNO3 were used with Soil 1, 2 and 3, varieties -C2H2 and +C2H2 with 50 mg N kg<sup>+</sup> KNO3. in Experiment 1, 2 and 3 and with 60 mg N kg<sup>+</sup>NaNO3. in Experiment 4 and *traced* variety with 50 mg N kg<sup>+</sup>-<sup>45</sup>N-KNO<sub>2</sub> in Experiment 1, 2 and 3 and 60 mg N kg<sup>+</sup>-<sup>45</sup>N-KNO<sub>2</sub> in Experiment 4, In variety traced-, NO3<sup>-</sup> with a <sup>15</sup>N-<u>enrichmentlabeling</u> of 50 atom% (at%) was used). For each treatment, we incubated 100 g dw soil in 850 mL preserving jars (J. WECK GmbH u. Co KG, Wehr, Germany) 305 with gas inlet and outlet equipped with three-three-port luer lock plastic stopcocks (Braun, Melsungen, Germany). According to the original SIRIN method (Anderson and Domsch, 1973, 1978) and a mixture of  $c_{opt}$ (glucose) and carrier material talcum (5 mg talcum g dw<sup>-1</sup>) was added to soil of treatment A and together with the growth inhibitors to the soil of treatments B, C and D. The soil and additives of each treatment were mixed for 90 seconds with a handheld electric mixer. During packing, the soil density was adjusted to an expected 310 target soil density of 1.6 g cm<sup>-3</sup> in *ExperimentSoil 1, 2* and 4 and of 1.3 g cm<sup>-3</sup> in *ExperimentSoil 3* to imitate field conditions. To ultimately achieve denitrifying conditions in all treatments and to avoid catalytic NO decomposition in the  $+C_2H_2$  variety (Nadeem et al., 2013), the headspace of the closed jars was flushed with N<sub>2</sub> to exchange the headspace 10 times. Directly following, 85 mL of the gas in the headspace in variety  $+C_2H_2$ were exchanged by pure C<sub>2</sub>H<sub>2</sub> resulting in 10 kPa C<sub>2</sub>H<sub>2</sub> in the headspace. The manual sample collection of 14
  - 10

315 mL gas in duplicates with a plastic syringe was performed after six, eight and ten hours (*ExperimentSoil 1, 2* and 3) or two, four and eight <u>hours (*ExperimentSoil 4*</u>) of incubation time, respectively. The removed gas was replaced by the same amount of N<sub>2</sub>.

## 2.3 Gas analysis

- Gas samples were analyzed\_analysed\_for N<sub>2</sub>O and CO<sub>2</sub> concentrations (*c(N<sub>2</sub>O)* and *c(CO<sub>2</sub>)*) with gas chromatography (GC, Agilent 7890A, Agilent, Böblingen, Germany) (Figure 1). The analytical precision of measurements was derived from analysing laboratory standards of different concentrations (0.5-1.000 ppm N<sub>2</sub>O and 340-10.000 ppm CO<sub>2</sub>) and resulted in a measurement precision of 1 % for N<sub>2</sub>O and 0.5 % for CO<sub>2</sub>. The instrumental detection limit of N<sub>2</sub>O was 4 µg N kg<sup>-1</sup> h<sup>-1</sup> with a measurement precision of 1% and for of CO<sub>2</sub> it was 137 µg C kg<sup>-1</sup> h<sup>-1</sup> the detection limit was C h<sup>-1</sup> with a measurement precision of 0.5%. As a control, N<sub>2</sub> and O<sub>2</sub> concentrations in the samples were analyzed analysed with GC to ensure anaerobic conditions during the incubation for N<sub>2</sub>O production from denitrification. CO<sub>2</sub> and N<sub>2</sub>O production rates were calculated by averaging
  - the measured  $N_2O$  production, i.e., between the time point of flushing with  $N_2$  (t=0) and six, eight or ten hours (or two, four and eight hours with *Soil 4*).
- The N<sub>2</sub>O isotopic analysis of the gas samples of varieties -C<sub>2</sub>H<sub>2</sub> and +C<sub>2</sub>H<sub>2</sub> (Figure 1) were performed on a preconcentrator (PreCon, Thermo–Finnigan, Bremen, Germany) interfaced with a GC (Trace Gas Ultra, Thermo Scientific, Bremen, Germany) and analyzed\_analysed\_by isotope ratio mass spectrometry (IRMS, Delta V, Thermo Fisher Scientific, Bremen, Germany) (Brand, 1995; Toyoda and Yoshida, 1999; Köster et al., 2013b). <u>A</u> laboratory standard N<sub>2</sub>O gas was used for calibration, having δ<sup>15</sup>N<sup>bulk</sup><sub>N2O</sub>, δ<sup>18</sup>O<sub>N2O</sub> and SP<sub>N2O</sub> values of -1.06 ‰, 40.22 ‰, and -2.13 ‰, respectively, in three concentrations (5, 10 and 20 ppm). The analytical precision was
- 0.1 ‰, 0.2 ‰ and 1.5 ‰ for  $\delta^{15}$ N<sup>bulk</sup><sub>N2O</sub>,  $\delta^{18}$ O<sub>N2O</sub> and SP<sub>N2O</sub> values, respectively. <u>H<sub>2</sub>O and CO<sub>2</sub> were trapped with</u> magnesium perchlorate and ascarite, respectively, to prevent any interference with N<sub>2</sub>O analysis. The gas samples of variety *traced* from *ExperimentSoil* 1, 2, and 3 were analyzed analysed for the 29/28 and

30/28 ratios of N2 according to Lewicka-Szczebak et al. (2013) using a modified GasBench II preparation system

coupled to -IRMS (MAT 253, Thermo Scientific, Bremen, Germany). The gas samples of variety *traced* from *ExperimentSoil 4* were analyzed analysed at the Centre for Stable Isotope Research and Analysis (University of Göttingen, Germany). The N<sub>2</sub> produced was analyzed analysed using an elemental analyzer analyser (Carlo Erba ANA 1500) that was coupled to dual inlet IRMS (Finnigan MAT 251) (Well et al., 1998; Well et al., 2006).
Isotopic values of N<sub>2</sub>O of *ExperimentSoil 4* (variety *traced*) were analyzed analysed in the same lab using a preconcentration unit coupled to IRMS (Precon-DeltaXP, Thermo Scientific, Bremen, Germany) (Well et al., 2006).
Isotope ratios were used applying the non-random distribution approach to calculate the fraction of N<sub>2</sub> and N<sub>2</sub>O originating from the <sup>15</sup>N-labelled N pool as well as the <sup>15</sup>N enrichment of that N pool (*a<sub>p</sub>*) (Bergsma et al., 2001; Spott et al., 2006).

### 2.4 Inhibitor effects

350 For interpretation of N<sub>2</sub>O or CO<sub>2</sub> production, the validity of the experimental results with respect to fungal and bacterial N<sub>2</sub>O fluxes was checked using a flux balance comparing the sum of bacterial and fungal inhibition effects (treatments B and C) to the dual inhibition effect (treatment D):

D = A - [(A - B) + (A - C)](Eq. 1)

355

360

With *A*, *B*, *C* and *D* representing the N<sub>2</sub>O production rates of the last sampling time of treatment *A*, *B*, *C* and *D*, respectively. Assuming that in the other three treatments (A, B and C) non-inhibitable N<sub>2</sub>O production-was equal to treatment D, N<sub>2</sub>O produced by bacteria <u>or and</u> fungi should show the following relation between the four treatments:

$$(A - D) = (B - D) + (C - D)$$
 (Eq. 2)

The fungal contribution to  $N_2O$  production during denitrification with microbial inhibition ( $\underline{f_{FDmi}}$ ) can be calculated, when  $N_2O$  production of treatment D is significantly smaller than  $N_2O$  production of treatments A, B and C by:

$$f_{FDmi} = \frac{(A-C)}{(A-D)}$$
(Eq. 3)
  
A detailed discussion of inhibitor effects and difficulties with organisms that were not inhibited or abiotic
sources (treatment D) is presented in section 4.1.

# 365 2.5 Isotope methods

# 2.5.1 Isotope endmember mixing approach (IEM)

370

The fungal fraction ( $f_{FD}$ ) contributing to N<sub>2</sub>O production from denitrification in soil samples was calculated according to the isotope mixing model (IEM) proposed by Ostrom et al. (2010), which was established for calculating the bacterial fraction ( $f_{BD}$ ) of N<sub>2</sub>O production. Assuming that bacteria (*BD*) and fungi (*FD*) are the only <u>sources of N<sub>2</sub>O</u> microorganisms responsible for denitrification-in soil, the <sup>15</sup>N site preference values of produced N<sub>2</sub>O (*SP*<sub>prod</sub>) results from the *SP*<sub>N2O</sub> mixing balance:

$$SP_{prod} = \mathcal{F}_{f_{FD}} * SP_{FD} + \mathcal{F}_{f_{BD}} * SP_{BD}$$
(Eq. 4)

where f<sub>FD</sub>\_and f<sub>BD</sub>\_represent the fraction of N<sub>2</sub>O produced by fungi and other N<sub>2</sub>O sources than fungal denitrification, respectively, and SP<sub>FD</sub> and SP<sub>BD</sub> are the respective SP<sub>N2O</sub> endmember values (Ostrom et al., 2010;
375 Ostrom and Ostrom, 2011). This calculation was based on the assumption that the sum of f<sub>BD</sub>\_and f<sub>FD</sub>\_equals 1 and that N<sub>2</sub>O reduction to N<sub>2</sub> is negligible. The mean SP<sub>FD</sub> value was assumed to be 33.6 % (Sutka et al., 2008; Maeda et al., 2015; Rohe et al., 2014a; Rohe et al., 2017) and the SP<sub>BD</sub> value from heterotrophic denitrification was assumed with minimum and maximum values from -7.5 to +3.7 % (Yu et al., 2020). For this IEM approach, only results from variety +C<sub>2</sub>H<sub>2</sub> could be used to calculate the fungal fraction contributing to N<sub>2</sub>O production
380 (f<sub>FD</sub>\_SP), as microorganisms of this variety produce N<sub>2</sub>O that is not affected by reduction to N<sub>2</sub>. The f<sub>FD</sub> SP contributing to N<sub>2</sub>O production during denitrification was calculated from using the measured SP<sub>N2O</sub> value from treatment A of variety +C<sub>2</sub>H<sub>2</sub> as SP<sub>prod</sub> value (Eq. 4)in – Eq. 4 that was solved for f<sub>FD</sub> (f<sub>FD</sub> = 1-((SP<sub>prod</sub>-SP<sub>FD</sub>))). By applying this equation, a range for f<sub>FD</sub> sp is received when using minimum and

385  $\frac{\text{maximum } SP_{BD} \text{ values.}}{\text{In case successful inhibition (modified SIRIN approach), Eq. 4 was solved for the <math>SP_{BD} \text{ value using } F_{BD}, F_{BD},$ 

# 2.5.2 Product ratio [N2O/(N2+N2O)] of denitrification

The variety *traced* served to assess  $N_2O$  reduction during denitrification in each experiment. The product ratio of denitrification  $[N_2O/(N_2+N_2O)]$  as given by the variety *traced* ( $r_{15N}$ ) was calculated as:

390	$r_{15N} = \frac{{}^{15}N_{N20}}{{}^{15}N_{N2} + {}^{15}N_{N20}} $ (Eq. 5)	
	with ${}^{I5}N_{N2Q}$ and ${}^{I5}N_{N2}$ representing N <sub>2</sub> O and N <sub>2</sub> produced in the ${}^{15}$ N-labeled fertilizer pool. To check the	
	effectiveness of $C_2H_2$ to block the N <sub>2</sub> O reduction, $r_{ISN}$ was compared with $r_{C2H2}$ , where the latter can be	
	<u>calculated from N<sub>2</sub>O production rates of varieties <math>-C_2H_2</math> and <math>+C_2H_2</math>:</u>	
	$r_{C2H2} = \frac{N_2 O_{-C2H2}}{N_2 O_{+C2H2}} $ (Eq. 6)	
395	with $N_2 O_{-C2H_2}$ and $N_2 O_{+C2H_2}$ representing the N <sub>2</sub> O produced in varieties $-C_2 H_2$ and $+C_2 H_2$ , respectively.	
	It was possible to assess the completeness of blockage of N2O reduction by C2H2 with the experimental setup. If	
	$r_{ISN}$ and $r_{C2H2}$ were in agreement, a complete blockage of $N_2O$ reduction could be assumed. This enabled us to	
	estimate reduction effects on the isotopic signatures of $N_2O$ by comparing the $\delta 0$ values, i.e., isotopic values of	
	<u>N<sub>2</sub>O produced without N<sub>2</sub>O reduction effects of variety <math>+C_2H_{23}</math> with isotopic values of N<sub>2</sub>O of variety <math>-C_2H_{23}</math>.</u>	
400	The information on the product ratio was used as an additional possibility to calculate the $f_{FD}$ also for variety -	
	C2H2. The Rayleigh-type model presented by Lewicka-Szczebak et al. (2017) and Senbayram et al. (2018) for	Feld
	similar closed-system incubations was used to calculate the <sup>15</sup> N site preference values of the originally produced	Feld
	<u>N<sub>2</sub>O of variety <math>-C_2H_2</math> (SP<sub>prod</sub>). SP values of emitted N<sub>2</sub>O, i.e. after partial reduction of produced N<sub>2</sub>O (SP<sub>N2O-7</sub>).</u>	
	were corrected with the net isotope effect of N <sub>2</sub> O reduction ( $\eta r$ ) and the $r_{15N}$ as follows:	
405	$SP_{prod} = SP_{N20-r} + \eta r \ln(r_{15N})$ (Eq. 7)	
	According to Yu et al. (2020) the nr was assumed to be -6 5. Secondly Subsequently, Eq. 4 was used to calculate	
	the $f_{FD}$ by using $SP_{prod}$ values of variety $-C_2H_2$ ( $f_{FD}$ spearce) obtained from Eq. 7.	
	2.5. <del>2 <u>3</u> SP/δ<sup>18</sup>O isotope mapping approach (SP/δ<sup>18</sup>O Map)</del>	
	The <u>f<sub>FD</sub></u> contributing to N <sub>2</sub> O production from denitrification in soil samples was also estimated with the SP/ $\delta^{18}$ O	
410	Map (ffp_MAP)_(Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2020) This method allows for	
	estimating on of both: the <u><math>f_{FD}</math></u> and N <sub>2</sub> O product ratio [N <sub>2</sub> O/(N <sub>2</sub> +N <sub>2</sub> O)] ( <u><math>r_{Map}</math></u> ). For precise estimations, the $\delta^{18}$ O	
I	values of soil water ( $\delta^{18}O_{H2O}$ ) applied in the experiments are needed and these values were not determined.	
	However, since we have independent information on the N <sub>2</sub> O product ratio from the <i>traced</i> variety $(r_{15N})$ , we can	
I	calculate the possible $\delta^{18}O_{H2O}$ values of soil to get the nearest N <sub>2</sub> O product ratios in natural and <sup>15</sup> N treatments.	
415	The fitting of $\frac{\delta^{18}O_{H2O}}{P_{BD}}$ values $\frac{f_{ED-MAP}}{P}$ was performed for mean, minimal und maximal values of $SP_{BD}$ (-1.9, -7.5	
	and 3.7_‰, respectively) and aimed at obtaining the minimal difference between $\underline{r}_{Map}$ and that measured in the	
	traced variety, i.e., the minimal value of $(\underline{r}_{ISN} - \underline{r}_{Map})^2$ (according to least squares method) and that measured	
	<u>with</u> for $-C_2H_2$ and $+C_2H_2$ , i.e. $r_{C2H2}$ variety (for explanation of the product ratio see next section <u>section</u> .	
	This further allows <u>-calculation of obtaining</u> the possible ranges for $f_{FD}$ for particular $\delta^{18}O_{H2O}$ fitted values (Table	
420	4) based on the SP/δ <sup>18</sup> O mapping approach (Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2020).	
	Namely, the fitted $\delta^{18}O_{H2O}$ values are applied to properly correct the $\delta^{18}O_{N2O}$ values of the mixing endmembers	
	(BD and FD) which depend on the ambient water. Afterwards, the corrected values of mixing endmembers are	
	applied to calculate the $f_{FD}$ values. The calculations with this approach may be performed assuming two different	
	scenarios of the interplay between N2O mixing and reduction_(Lewicka-Szczebak et al., 2017; Lewicka-	
425	Szczebak et al., 2020), but for this study both scenarios yielded almost identical results (maximal differences of	
	0.02 in N <sub>2</sub> O product ratio and <u>2 % for <math>f_{FD}</math></u> was were found), due to $f_{BD}$ near 100 %. Hence, we only provide the	
	results assuming the reduction of bacterial N <sub>2</sub> O followed by mixing with fungal N <sub>2</sub> O. In the following, all	
	calculated fractions are presented in percent (%).	

eldfunktion geändert eldfunktion geändert 5.3 Product ratio [N<sub>2</sub>O/(N<sub>2</sub>+N<sub>2</sub>O)] of denitrification

430 The variety traced served to assess N2O reduction during denitrification in each experiment. The product ratio of denitrification [N2O/(N2+N2O)] as given by the variety traced (15A) was calculated as: <sup>15</sup>N<sub>N20</sub> (Eq. 5) N<sub>N2</sub>+<sup>15</sup>N<sub>N2C</sub> produced in the <sup>15</sup>N labeled fertilizer pool To compared with C2H2; where the letter can be f C. H. to block the N.O. 435  $- \Theta$  production rates of varieties  $-C_{2}H_{2}$  and  $+C_{2}H_{2}$ N<sub>2</sub>0\_/ <del>(Eq. 6)</del>  $N_2 O_{\pm C2H2}$ G cur and N Q cur  $\frac{1}{1}$   $\frac{1}$ ratio15N\_and product ratio C2H2 were in agreement, a complete blockage of N2O reduction could be enabled us to estimate reduction effects on the isotopic signatures of N<sub>2</sub>O by comparing the 440 Example 2 by the set of N.O. of variety ed as an additional possibility to calculate the also for variety adal presented by Lawicka Szczabak et al. (2017) and Sephavram et al. (2018) stem incubations, the <sup>15</sup>N site preference values of the originally produced N<sub>2</sub>O, i.e. without its of variety -C2H2 (SPanad) was calculated by correcting SP values of emitted N2O, i.e. 445 after partial reduction of produced  $N_2O$  (SP<sub>N2Q,2</sub>) from variety  $-C_2H_2$  with the net isotope effect of  $N_2O$  reduction (nr) and the 151 as follows: (Ea. 7) 51 ecording to (Yu et al., 2020) the nr was assumed to be -6‰. Secondly, Eq. 4 was used to calculate the by using 450 d from Eq. 7

# 2.6 Sources Other sources of N2O produced

Assuming that denitrification is-was the only process producing source of N<sub>2</sub>O in the incubation experiment, the expected <sup>15</sup>N enrichment in N<sub>2</sub>O produced ( ${}^{15}N_{N2O,exp}$ ) was given by

$${}^{15}N_{N20\_exp} [at\%] = \frac{(N_{soil} \times {}^{15}N_{nat}) + (N_{fert} \times {}^{15}N_{fert})}{N^{bulk}}$$
(Eq. 8)

455 with  $N_{soil}$ ,  $N_{fert}$  and  $N^{bulk}$  describing the amount of N [mg] in unfertilized soil samples (Table 1), fertilizer and fertilized soil samples, respectively, and  ${}^{15}N_{nat}$  and  ${}^{15}N_{fert}$  is standing for  ${}^{15}$ N enrichment under natural conditions (0.3663 at%) and in fertilizer (50 at%), respectively. Comparison of measured  ${}^{15}$ N enrichment in N<sub>2</sub>O and  ${}^{15}N_{N2O,exp}$  gave information about the contribution of processes other than denitrification to N<sub>2</sub>O production.

# 2.7 Statistical Analysis

We conducted several three-way analyses of variance (ANOVA) to test significant effects of soil, experimental variety and treatment on N<sub>2</sub>O production, CO<sub>2</sub> production, and SP<sub>N2O</sub>, δ<sup>15</sup>N<sup>bulk</sup><sub>N2O</sub> and δ<sup>18</sup>O<sub>N2O</sub> values. The pairwise comparison with Tukey's HSD test was madeallowed to find differences between soils, varieties and treatments influencing N<sub>2</sub>O production, CO<sub>2</sub> production, and isotopic values. Significant effects of soils and treatments on <u>r<sub>C2H2</sub></u> and <u>r<sub>15N</sub></u> were tested by two-way ANOVA, while differences between soils and treatments influencing the product ratios were tested with pairwise comparison with Tukey's HSD test. Effects of varieties -

 $C_2H_2$  and *traced* on N<sub>2</sub>O and CO<sub>2</sub> production were tested by ANOVA. For this ANOVA, the N<sub>2</sub>O production rate had to be log<sub>10</sub>-transformed to achieve homogeneity of variance and normality. The significance level  $\alpha$  was 0.1 for every ANOVA. For some ANOVAs treatments were excluded, when replicates were n < 3. <u>This was the case</u> when only one or two samples out of three replicates could be analysed. This is denoted in the captions of tables (<u>Table 2 and 3</u>). The N<sub>2</sub>O or CO<sub>2</sub> production rates of variety + $C_2H_2$  were followed over three sampling times by regression. For statistical analysis, we used the program R (R Core Team, 2013). Excel Solver tool was used to determine the  $\delta^{18}O_{H2O}$  values in the application of SP/ $\delta^{18}O$  Map calculations.

# 470

# 3. Results

# 3.1 N<sub>2</sub>O production rates

N<sub>2</sub>O and CO<sub>2</sub> production rates of all treatments were similar in magnitude in almost all cases and mostly indistinguishable (Table 2). CO<sub>2</sub> production rates were determined to get additionally information about the denitrifying process. N<sub>2</sub>O production rates exhibited increasing trends with ongoing incubation time for every soil with large variations within the treatments. Contrary to that, CO<sub>2</sub> production rates showed decreasing trends (Figure 12, exemplarily shown for data of variety +C<sub>2</sub>H<sub>2</sub>). Calculations of inhibitor effects were based on average N<sub>2</sub>O and CO<sub>2</sub> production rates of the entire incubation period, i.e. 10-ten hours of incubation time for *ExperimentSoil* 1, 2 and 3 and 8-eight hours for *ExperimentSoil* 4.

N<sub>2</sub>O and CO<sub>2</sub> production rates of all  $+C_2H_2$  varieties differed significantly among soils (P < 0.001) and N<sub>2</sub>O production rates differed also significantly among treatments (P < 0.001). Largest N<sub>2</sub>O production rates <u>of</u> about <u>5.5555</u> to <u>6.1613</u> µg N kg<sup>-1</sup>h<sup>-1</sup> wasere obtained in <u>ExperimentSoil</u> 1 and <u>3</u>, respectively, while in <u>ExperimentSoil</u> 2 and 4 N<sub>2</sub>O production rates were <u>-lowersmaller</u> (<u>2.6271</u> and <u>2.7264</u> µg N kg<sup>-1</sup>h<sup>-1</sup>, respectively). N<sub>2</sub>O and CO<sub>2</sub> production rates were significantly larger in variety  $+C_2H_2$  than in variety  $-C_2H_2$  of <u>ExperimentSoil</u> 1, 3 and 4 (P = 0.002, P < 0.010 and P < 0.010 for N<sub>2</sub>O production rate and P = 0.027, P < 0.010 and P = 0.008 for CO<sub>2</sub> production rate, respectively) (Table 2), while  $-C_2H_2$  and  $+C_2H_2$  varieties of <u>ExperimentSoil</u> 2 did not differ in N<sub>2</sub>O and CO<sub>2</sub> production rates (P = 0.402 and P = 0.288, respectively).

Figure 42: Time series of average N<sub>2</sub>O and CO<sub>2</sub> production rates during incubation of variety  $+C_2H_2$  at the three sample collection times of each soil (ExperimentSoil 1 --to 4) for treatment A without growth inhibitors, B with bacterial growth inhibition, C with fungal growth inhibition, and D with bacterial and fungal growth inhibition; *P*-values for linear regressions (significance level  $\alpha \le 0.05$ ). For all significant regressions,  $R^2$ -values were  $\ge 0.46$  and in the case of non-significance,  $R^2$ -values were  $\le 0.40$ .

n.d.: There was no detectable CO<sub>2</sub> production in  $\frac{ExperimentSoil}{ExperimentSoil}$  4 at the first sampling time after 2 hours. (Figure is continued on next page)





Without blockage of N<sub>2</sub>O reductase (variety  $-C_2H_2$ ), N<sub>2</sub>O production rates of treatment A varied significantly among experiments <u>Soils</u> with mean values between <u>1.6175</u> and <u>3.6355</u> µg N kg<sup>-1</sup> h<sup>-1</sup> ( $P \le 0.001$ ) (Table 2). In Experiment<u>Soil</u> 1, N<sub>2</sub>O production rate was significantly larger (2.7272 µg N kg<sup>-1</sup> h<sup>-1</sup>) than in Experiment<u>Soil</u> 4 (<u>1.6175</u> µg N kg<sup>-1</sup> h<sup>-1</sup>) (P = 0.028) in variety  $-C_2H_2$ . In most cases of the three varieties ( $-C_2H_2$ ,  $+C_2H_2$ , and <u>traced</u>) The inhibitor application of each variety revealed in most cases that treatment A (without growth inhibitors) produced most N<sub>2</sub>O, followed by either treatment B (bacterial growth inhibitor; more N<sub>2</sub>O compared to treatment C in Experiment<u>Soil</u> 2, 3 and 4) or treatments C (fungal growth inhibitor; more N<sub>2</sub>O compared to treatment B in Experiment<u>Soil</u> 1). Smallest N<sub>2</sub>O production rates were in most cases found in treatment D (Im

varieties  $-C_2H_2$ ,  $+C_2H_2$  and traced varieties, non-inhibitable organisms N<sub>2</sub>O production(treatment D) showed smallest N<sub>2</sub>O production rates in most cases (i. e. except of for variety  $-C_2H_2$  of ExperimentSoil 1, varieties - $C_2H_2$  and traced of ExperimentSoil 3 and variety traced of ExperimentSoil 4). Microbial inhibitor treatments differed significantly in N<sub>2</sub>O fluxes of variety  $+C_2H_2$  of each experimentSoil (always  $P \le 0.040$ ), while this was not the case for inhibitor treatments of varieties  $-C_2H_2$  and traced of ExperimentSoil 4 (P = 0.154 and P = 0.154, respectively). Significant deviations of treatments without (A) or with full inhibition (D) were found in the following cases (Table 2): N<sub>2</sub>O production rate of treatment A was significantly larger compared to the other three treatments of ExperimentSoil 1 ( $+C_2H_2$  and  $-C_2H_2$ ), ExperimentSoil 2 ( $+C_2H_2$ ) and ExperimentSoil 3 ( $+C_2H_2$ ); treatment D was significantly smaller compared to the other three treatments in ExperimentSoil 2 ( $-(\pm C_2H_2)$ ) only and compared to treatments A and C of in ExperimentSoil 1 ( $+C_2H_2$ ). A detailed discussion of inhibitor effects and difficulties with organisms that were not inhibited or abiotic sources is presented in section 4.1. Comparing varieties  $-C_2H_2$  and traced, N<sub>2</sub>O and CO<sub>2</sub> rates did not differ (P = 0.991 for N<sub>2</sub>O production rate and P = 0.490 for CO<sub>2</sub> production rate, respectively), confirming that <sup>15</sup>N-labeling did not affect N<sub>2</sub>O and CO<sub>2</sub> processes.

525

520

Treatment/	mean N <sub>2</sub> O	mean CO <sub>2</sub>	$\delta^{18}O_{N2O}$	$\delta^{15} N^{bulk}_{N20}$	$SP_{N2O}$
variety	[µg N kg <sup>-1</sup> h <sup>-1</sup> ]	[µg C kg <sup>-1</sup> h <sup>-1</sup> ]	[‰]	[‰]	[‰]
ExperimentSoi	1 (Loamy sand, wi	nter 2012)			
$A / -C_2H_2$	<u>272.0</u> ( <u>38.4</u> )a	<u>1233.8</u> ( <u>170.5</u> )a	13.1 (0.2)a	-21.9 (1.7)a	1.6 (0.8)a
<b>B</b> / - $C_2H_2$	<u>180.9</u> ( <u>16.8</u> )b	<u>1284.8</u> ( <u>168.0</u> )a	13.0 (<0.1)*	-24.2 (0.7)*	-1.3 (0.2)*
$C / -C_2H_2$	<u>203.1</u> ( <u>14.4</u> )b	<u>1124.8</u> ( <u>54.8</u> )a	14.6 (0.4)a	-20.0 (0.8)a	-1.6 (0.5)a
$D / -C_2 H_2$	<u>207.8</u> ( <u>32.6</u> )b	<u>1371.7</u> ( <u>35.3</u> )a	15.2 (0.5)*	-20.2 (1.8)*	-0.3 (0.5)*
A / $+C_2H_2$	<u>554.9</u> ( <u>46.5</u> )a	<u>1700.9</u> ( <u>98.1</u> )a	8.5 (0.1)a	-22.1 (0.3)a	-0.4 (0.3)a
$B / + C_2 H_2$	<u>353.5</u> ( <u>14.0</u> )b	<u>1610.7</u> ( <u>47.2</u> )a	7.5 (0.1)a	-26.1 (0.2)a	-1.2 (1.0)b
$C / + C_2 H_2$	<u>441.8</u> ( <u>18.5</u> )c	<u>1604.1</u> ( <u>60.3</u> )a	9.3 (0.2)a	-22.4 (0.4)a	-0.9 (0.4)b
$D / +C_2H_2$	<u>331.0</u> ( <u>20.5</u> )b	<u>1438.0</u> ( <u>141.9</u> )a	7.8 (0.3)*	-24.2 (0.1)*	-2.3 (0.7)*
ExperimentSoi	2 (Sand, winter 20	12)			
A / $-C_2H_2$	<u>315.0</u> ( <u>35.0</u> )a	<u>1316.7</u> ( <u>97.7</u> )a	15.5 (1.8)a	-18.9 (2.6)a	-0.9 (2.5)a
$\mathbf{B}/-C_2H_2$	<u>241.7</u> ( <u>3.0</u> )b	<u>1209.2</u> ( <u>24.6</u> )a	15.0 (1.3)a	-23.4 (2.5)a	-0.8 (<0.1)a
$C / -C_2H_2$	<u>247.6</u> ( <u>22.8</u> )b	<u>1201.9</u> ( <u>48.2</u> )a	14.3 (0.1)a	-21.8 (0.2)a	-1.8 (0.2)a
$D / -C_2 H_2$	<u>198.4</u> ( <u>26.8</u> )b	<u>1102.4</u> ( <u>101.7</u> )a	13.4 (0.3)a	-24.5 (0.1)a	-1.2 (0.3)a
A / $+C_2H_2$	<u>270.9</u> ( <u>36.3</u> )a	<u>1271.6</u> ( <u>203.5</u> )a	12.6 (0.3)a	-18.9 (4.6)a	-1.4 (0.3)a
$B / + C_2 H_2$	<u>263.1</u> ( <u>19.1</u> )a	<u>1338.7</u> ( <u>71.9</u> )a	12.3 (0.1)a	-24.6 (0.2)b	-2.0 (0.2)a
$C / + C_2 H_2$	<u>247.3</u> ( <u>15.9</u> )a	<u>1220.2</u> ( <u>50.0</u> )a	12.7 (0.1)*	-23.3 (0.2)*	-1.7 (0.4)*
$D / +C_2H_2$	<u>187.3</u> ( <u>21.8</u> )b	<u>1173.1</u> ( <u>55.1</u> )a	12.2 (0.3)a	-26.0 (0.1)b	-1.5 (0.9)a
ExperimentSoi	1 3 (Silt loam, winte	r 2013)			
A / $-C_2H_2$	<u>355.0</u> ( <u>18.4</u> )a	<u>1227.6</u> ( <u>95.2</u> )a	26.0 (0.5)a	-20.8 (0.5)a	-0.5 (0.4)a
<b>B</b> / $-C_2H_2$	<u>325.4</u> ( <u>36.3</u> )a	<u>1159.3</u> ( <u>178.2</u> )a	24.1 (0.2)b	-22.0 (0.2)b	-0.1 (0.4)a
$C / -C_2H_2$	<u>278.9</u> ( <u>9.8</u> )a	<u>1056.0</u> ( <u>59.6</u> )a	27.3 (0.1)b	-20.6 (0.3)a	0.6 (0.2)a
$D / -C_2H_2$	<u>291.1</u> ( <u>38.5</u> )a	<u>1118.5</u> ( <u>70.3</u> )a	26.3 (0.3)a	-21.0 (0.1)a	-0.04 (0. <mark>182</mark> )a
$A / +C_2H_2$	<u>612.8</u> ( <u>25.2</u> )a	<u>1332.5</u> ( <u>116.9</u> )a	15.2 (0.1)a	-25.6 (0.8)a	-2.8 (0.2)a
$B / + C_2 H_2$	<u>546.9</u> ( <u>27.5</u> )b	<u>1235.7</u> ( <u>83.4</u> )a	14.9 (0.2)a	-26.3 (<0.1)a	-3.5 (0.4)a
$C / + C_2 H_2$	<u>519.8</u> ( <u>19.2</u> )b	<u>1173.5</u> ( <u>25.7</u> )a	16.2 (<0.1)*	-25.2 (0.1)*	-4.0 (0.4)*
$D / + C_2 H_2$	<u>511.7</u> ( <u>3.5</u> )b	<u>1295.6</u> ( <u>63.3</u> )a	16.0 (0.1)b	-25.1 (0.1)a	-4.3 (0.5)a
ExperimentSoi	1 4 (Loamy sand, su	mmer 2011)			
A / $-C_2H_2$	<u>175.3</u> ( <u>6.6</u> )a	<u>2448.5</u> ( <u>135.8</u> )a	25.7 (0.3)a	-30.6 (0.2)a	12.1 (1.6)a
<b>B</b> / - $C_2H_2$	<u>121.3</u> ( <u>74.0</u> )a	<u>2091.3</u> ( <u>19.5</u> )b	28.0 (5.0)a	-32.3 (0.7)a	7.7 (1.4)b
$C / -C_2 H_2$	<u>104.5</u> ( <u>5.3</u> )a	<u>1844.7</u> ( <u>192.1</u> )b	29.3 (0.1)a	-30.0 (0.5)a	4.3 (1.0)c
$D / -C_2H_2$	<u>73.8</u> ( <u>63.0</u> )a	<u>1632.2</u> ( <u>115.3</u> )b	28.9 (1.2)a	-31.8 (2.2)a	3.4 (2.0)c
$A / +C_2H_2$	<u>263.5</u> ( <u>31.7</u> )a	<u>2076.6</u> ( <u>305.3</u> )a	13.5 (0.5)*	-34.7 (0.1)*	-1.0**
$B / + C_2 H_2$	<u>233.0</u> ( <u>15.6</u> )a	<u>1794.9</u> ( <u>238.9</u> )a	14.3 (1.7)a	-33.8 (0.9)a	-4.9 (0.9)a
$C / + C_2 H_2$	<u>119.5</u> ( <u>102.7</u> )a	<u>1736.8</u> ( <u>424.7</u> )a	19.0 (7.0)a	-33.1 (2.8)a	-1.7 (2.7)b
D / + $C_2H_2$	161.6 (7.6)a	1497.0 (138.7)a	14.8 (0.5)a	-35.7 (0.2)a	-4.9 (0.7)c

Table 2: Average CO<sub>2</sub> and N<sub>2</sub>O production rates and N<sub>2</sub>O isotopic values of N<sub>2</sub>O of the last sample collection with and without C<sub>2</sub>H<sub>2</sub> application in the headspace (varieties  $-C_2H_2$  and  $+C_2H_2$ ) of each soil (ExperimentSoil  $1 - \underline{to} 4$ ) for treatments A without, B with bacterial, C with fungal, and D with bacterial and fungal growth inhibition, respectively (standard deviation in brackets, n = 3).

530

Letters denote significant differences (P < 0.1) among treatments and varieties within a soil. Asterisks indicate that only two samples (\*) or one sample (\*\*) of triplicates were analyzable analyzable due to logistical difficulties.

# 3.2 Isotopologues of N<sub>2</sub>O produced in different varieties and treatments

#### 535 3.2.1 Variety -C2H2

<u>SP<sub>N20</sub> values of all Soils and inhibitor treatments of variety  $-C_2H_2$  were within a range of -1.8 to 12.1 % (Table</u> 2) and differed among inhibitor treatments (P = 0.037).  $SP_{N2O}$  values in variety - $C_2H_2$  of Soil 4 was particularly large (3.4 to 12.1 ‰) compared to the other Soils (1.6 to -1.6 ‰).  $SP_{N2O}$  values of variety  $-C_2H_2$  were significantly larger than  $SP_{N20}$  values of variety  $+C_2H_2$  (up to 2.4, 1.5, 4.6 and 4.1 ‰ in Soil 1, 2, 3 and 4, respectively). Generally, most SPprod values of variety -C2H2 (Eq. 7) were smaller than SPN20 values of variety -<u> $C_2H_2$  but still larger than SP<sub>N20</sub> values of variety + $C_2H_2$  and are presented in Table S2 (Supplementary Material).</u>

# 3.2.1-2 Variety +C<sub>2</sub>H<sub>2</sub>

545

540

 $SP_{N2O}$  values of all experiment<u>Soil</u>s, and all treatments of variety  $+C_2H_2$  were within a narrow range between -4.9 and -0.4 % (Table 2), and differed only significantly among treatments of ExperimentSoil 4 (P = 0.002). In general, there were only small differences among treatments:  $SP_{N2O}$  values of treatments A in variety  $+C_2H_2$ differed significantly among soils (P < 0.001), with largest  $SP_{N2O}$  values in ExperimentSoil 1 (-0.4 %) and smallest  $SP_{N2O}$  values in ExperimentSoil 3 (-2.8 %).  $SP_{N2O}$  values of treatment D in variety  $+C_2H_2$  of all soils varied between -1.5 and -4.9 ‰, but only  $SP_{N2O}$  values of ExperimentSoil 2 differed significantly from  $SP_{N2O}$ values of the other Experiment Soils (P = 0.006). For treatments B of variety  $+C_2H_2$ ,  $SP_{N20}$  values differed only 550 significantly between ExperimentSoil 1 and 4, 2 and 4, and 1 and 3 (each P = 0.002). SP<sub>N20</sub> values from treatment C in variety  $+C_2H_2$  did not differ significantly (P = 0.600). For every soil, we found significantly larger  $\delta^{18}O_{N20}$ ,  $\delta^{15}N^{\text{bulk}}_{N20}$  and  $SP_{N20}$  values in variety  $-C_2H_2$  than in variety  $+C_2H_2$  (P < 0.001), except for Experiment<u>Soil</u> 2, where  $\delta^{15}N^{\text{bulk}}_{N20}$  values of variety  $-C_2H_2$  were indistinguishable from those of variety  $+C_2H_2$ (P = 0.400). However, only in a few varieties there were significant differences in  $\delta^{18}O_{N20}$ ,  $\delta^{15}N^{\text{bulk}}_{N20}$  or  $SP_{N20}$ 555 values between treatments with fungal and bacterial inhibition (B and C, respectively) (Table 2). As explained in section 3.3, N<sub>2</sub>O reduction blockage in varieties  $+C_2H_2$  was successful in most cases (Experiment-Soil 2, 3 and 4). SP<sub>N20</sub> values of this variety are thus assumed to be valid estimates of  $\delta \theta$ , i.e. SP<sub>prod</sub> values of N<sub>2</sub>O production, and can thus be used for applying the IEM.

#### 560 3.2.2 Variety C2H2

565

 $SP_{VOC}$  values of all experiments and inhibitor treatments of variety  $-C_{1}H_{1}$  were within a range of 1.8 to 12.1 % -2) and did not differ among inhibitor treatments (P = 0.037). SP<sub>N20</sub> values in variety -  $C_2H_2$  of Experim ieularly large (3.4 - 12.1 ‰) compared to the other experiments (1.6 to -1.6 ‰). As already SP<sub>NDC</sub> values of variety  $-C_2H_2$  were significantly larger than  $SP_{N2C}$  values of variety  $+C_2H_2$  (up to 2.4, 1. in Experiment 1, 2, 3 and 4, respectively). Generally, most  $SP_{max}$  values of variety  $-C_{H_x}$  (Eq. 7) ller than SP<sub>N20</sub> values of variety -C2H2 but still larger than SP<sub>N20</sub> values of variety presented in Table S1 (supplementary Material).

### 3.2.3 Variety traced

The <sup>15</sup>N-labeling of N<sub>2</sub>O ( $^{15}N_{N2O}$ ) or N<sub>2</sub> produced ( $^{15}N_{N2}$ ) gave information about the incorporated N from  $^{15}N$ -570 labeled NO3<sup>-</sup> into N2O or N2 as well as about the N2O reduction to N2. Microorganisms in each treatment used 20 the <sup>15</sup>N-labeled NO<sub>3</sub><sup>-</sup> in variety *traced* (Table 3) and expected  ${}^{15}N_{N2O}$  depended on the initial N abundance in NO3<sup>-</sup> of unfertilized soil (Eq. 7). ExperimentSoil 4 is the only one showing a large discrepancy between measured (about 30 at%) and calculated <sup>15</sup>N<sub>N2O\_exp</sub> (49 at%) in N<sub>2</sub>O, whereas the other experimentSoils showed close agreement (Table 3).

#### 575 3.3 Product ratios of denitrification and efficiency of N2O reductase blockage by C2H2

<u>r</u><sub>C2H2</sub> as well as product ratio<sub>T15N</sub> of determined with ExperimentSoil 2 were significantly larger than of with the other experiment<u>Soil</u>s ( $P \le 0.001$ ) (Table 3). <u>r</u><sub>ISN</sub> of treatment B was significantly larger than-of treatment C and D of ExperimentSoil 4 (P = 0.032), while all other treatments of other Soils soils-did not differ.  $\underline{r}_{C2H2}$  did not differ significantly among treatments (P = 0.400). In order to test the efficiency of blockage of N<sub>2</sub>O reduction by  $C_2H_2$  application, <u>r</u><sub>C2H2</sub> (Eq. 5) was compared with <u>r</u><sub>I5N</sub> (Eq. 6). In <u>ExperimentSoil</u> 1, <u>r</u><sub>C2H2</sub> was by far smaller than r<sub>15N</sub>, while both calculated product ratios were in similar ranges in the other three experiment/Soils and thus a successful blockage of N<sub>2</sub>O reduction was assumed for those experimentSoils.

585

590

580

Table 3: Average CO<sub>2</sub> and N<sub>2</sub>O production rates of the last sample collection after 10 or 8 hours of variety *traced*, respectively, with <sup>15</sup>N labeling in N<sub>2</sub>O (<sup>15</sup>N-<sub>N2O</sub>) and the calculated  $L_{I5N}$  of variety *traced* and  $L_{C2H_2}$  calculated from N<sub>2</sub>O production rates of variety  $-C_2H_2$  and  $+C_2H_2$  of each soil (ExperimentSoil 1 -to 4) for treatments A without, B with bacterial, C with fungal, and D with bacterial and fungal growth inhibition, respectively (standard deviation in brackets, n = 3).

	mean N <sub>2</sub> O	mean CO <sub>2</sub>	15	$^{15}N_{N2O\_exp}$	Calc. total	Calc. total			
Treatment	[µg N kg h ]	[µg N kg h ]	$N_{N20}$ [at%]	[at%]	<u><i>P</i></u> 15N	<u>C2H2</u>			
ExperimentSoi	Experiment <u>Soil</u> 1 (Loamy Sand, 2012)								
А	<u>255.6</u> ( <u>43.5</u> )	<u>1310.0</u> ( <u>167.3</u> )	36.8 (0.1)		0.80 (0.02)	0.48 (0.07)			
В	<u>154.5</u> ( <u>29.6</u> )	<u>1153.5</u> ( <u>238.4</u> )	36.4 (0.2)	20	0.76 (0.02)	0.48 (0.05)			
С	<u>191.6</u> ( <u>30.7</u> )	<u>1219.6</u> ( <u>109.1</u> )	36.9 (<0.1)	39	0.72 (0.05)	0.45 (0.04)			
D	<u>148.1</u> ( <u>1.9</u> )	<u>1253.8</u> ( <u>54.5</u> )	36.8 (0.1)		0.69 (0.02)	0.54 (0.05)			
ExperimentSoi	<i>l</i> 2 (Sand, 2012)								
А	<u>240.7</u> ( <u>0.95</u> )	<u>1286.2</u> ( <u>5.6</u> )	43.2 (<0.1)		0.94 (0.01)	1.04 (0.10)			
В	<u>185.1</u> ( <u>3.9</u> )	<u>1157.4 (17.3</u> )	43.0 (0.1)	4.4	0.94 (0.01)	0.81 (0.04)			
С	<u>241.1 (13.4</u> )	<u>1282.1 (63.4</u> )	43.2 (0.1)	44	0.95 (0.01)	0.99 (0.09)			
D	<u>167.3</u> ( <u>34.9</u> )	<u>1199.0</u> ( <u>34.6</u> )	42.7 (0.1)		0.93 (0.01)	0.98 (0.04)			
ExperimentSoi	<u>l</u> 3 (Silt loam, 201	3)							
А	<u>285.9</u> ( <u>20.4</u> )	<u>1044.0</u> ( <u>46.6</u> )	35.8 (<0.1)		0.62 (<0.01)	0.52 (0.04)			
В	<u>320.5 (14.7)</u>	<u>1204.2</u> ( <u>86.5</u> )	35.5 (<0.1)	24	0.62 (0.01)	0.59 (0.02)			
С	<u>216.4</u> ( <u>34.9</u> )	<u>980.5</u> (202.5)	35.5 (<0.1)	54	0.59 (0.02)	0.48 (0.04)			
D	<u>231.4 (11.4</u> )	<u>988.5</u> ( <u>74.4</u> )	35.3 (<0.1)		0.62 (0.01)	0.51 (0.04)			
ExperimentSoi	1 4 (Loamy Sand, 2	2011)							
А	<u>156.9</u> ( <u>62.7</u> )	<u>3111.4 (1252.5</u> )	31.1**		0.54 (0.05)	0.63 (0.10)			
В	<u>169.2</u> ( <u>6.1</u> )	<u>2314.6</u> ( <u>307.1</u> )	26.5**	40	0.59 (0.03)	0.63 (0.17)			
С	<u>117.2</u> ( <u>3.1</u> )	<u>1785.6</u> ( <u>79.3</u> )	30.1 <u>(1.1)</u> *	49	0.50 (0.01)	0.62 (0.02)			
D	<u>115.2</u> ( <u>3.1</u> )	<u>1706.7</u> ( <u>38.1</u> )	33.5 <u>(0.5)</u> *		0.50 (0.01)	0.53 (0.12)			

Asterisks indicate that only two samples (\*) or one sample (\*\*) were analyzed analyzed analyzed, due to logistical difficulties.  $a_{15}N_{N20exp}$  [at%] was calculated from Eq. 78.  $b_{T_{15N}} = [N_2O/(N_2+N_2O)]$  with N<sub>2</sub>O or N<sub>2</sub> production rates from variety *traced*; see Eq. 5  $c_{C2H2} = [N_2O_{-C2H2}/N_2O_{+C2H2}]$  with N<sub>2</sub>O production rate from varieties  $-C_2H_2$  and  $-C_2H_2$ ; see Eq. 6, cf. Table 2

# 3.4 Fungal contribution to N<sub>2</sub>O production from denitrification by microbial inhibitor approach (modified SIRIN)

When calculating <u>f<sub>EDmi</sub></u> N<sub>2</sub>O production rates of treatment D must be significantly smaller compared to the other three treatments and the flux balance according to Eq. 1 and 2 must be consistent. <u>Taking the large ranges of N<sub>2</sub>O production rates of each treatment (minimum and maximum values) into account, for each *Soil* (A-D) was indistinguishable from ((B-D)+(C-D)) (Eq. 2), showing good agreement between Eqs. 1 and 2. However, N<sub>2</sub>O production in treatment D was large within all varieties. Only with *-Soil 2* of the variety +C<sub>2</sub>H<sub>2</sub>, the N<sub>2</sub>O production rates of treatment D were significantly smaller than those of the other three treatments. Thus, for *Soil 2*, <u>f<sub>FDmi</sub> could be calculated (Eq. 3) and amounted to 28 ± 9 % (Table 5) with a corresponding This was only the case in Experiment 2 of variety +C<sub>2</sub>H<sub>2</sub>. The calculated F<sub>EDmit</sub> (Eq. 3) was 0.28 ± 0.90 (Table 5). The respective flux-fungal N<sub>2</sub>O production rate of fungal N<sub>2</sub>O was 0.2423.7 ± 0.081.8 µg N kg<sup>-1</sup> h<sup>-1</sup>. Although the N<sub>2</sub>O production rate of Tereatment D was smaller than that of treatment A (*Soil 2*), it must be pointed out<sub>7</sub> that due to the large amount of non-inhibitable production (treatment D), even the result for *Soil 2* is actually very unsure.
For all other experiment/*Soils*, calculation of *f<sub>EDmit</sub>* was not possible, i.e., SIRIN was not successful.
</u></u>

# 3.5 Fungal contribution to N<sub>2</sub>O production from denitrification by the SP endmember mixing approach (IEM) and SP/ $\delta^{18}$ O isotope mapping approach (SP/ $\delta^{18}$ O Map)

The IEM revealed that  $f_{FD SP}$  was small in all Soils ( $\leq 15$  %,  $\leq 14$  %,  $\leq 9$ %, and  $\leq 11$  % with Soil 1 to 4, respectively) (Table 5). When applying SP/ $\delta^{18}$ O Map, we can assess the plausibility of the determined  $_{FP-f_{FD}}$ 610 values based on the  $\delta^{18}O_{H2O}$  values obtained from the fitting ( $\delta^{18}O_{H2O}$  value in Table 4) and the fitting outcome, i.e. the difference between  $\underline{r}_{ISN}$  and  $\underline{r}_{MAP}$  (Diff-in, see Table 4). The most probable  $\delta^{18}O_{H2O}$  value for our experimentSoils can be assumed based on the fact that Braunschweig tap water was added to soil used and the original soil water also represents the isotope characteristics typical for this region, which is about -7.4\_‰ (longterm mean Braunschweig precipitation water (Stumpp et al., 2014)). Thus, in the presented application of 615  $SP/\delta^{18}O$  Map,  $\delta^{18}O_{H2O}$  values were fitted and it has to be pointed out that the precision of such calculations can be improved by measuring  $\delta^{18}O_{H20}$  instead. Depending on the season and evaporative losses,  $\delta^{18}O_{H20}$  this value may slightly vary and the most possible range of soil water in our experimentSoils may vary from about -11 to -4 ‰ as observed in other experiments conducted-used in our laboratory experiments with similar conditions (Lewicka-Szczebak et al., 2014; Rohe et al., 2014a; Lewicka-Szczebak et al., 2017; Rohe et al., 2017). Taking 620 this into account, we can say that for ExperimentSoil 1, the fungal contribution must be below 0.022 %, because to obtain any larger  $f_{FD}$  values, unrealistically small  $\delta^{18}O_{H2O}$  values (of -14.9.%) must be fitted (see Table 4). For ExperimentSoil 2, both the smaller  $f_{FD MAP}$  values of  $0.01 \frac{9}{20}$  and the larger ones up to  $0.15 \frac{9}{20}$  are possible, since they are associated with very realistic  $\delta^{18}O_{H2O}$  values (of -6.3 and -10.1, respectively) and identical *Diff* of 0.04 (Table 4). For ExperimentSoil 3, the only plausible fitting can be obtained for the smallest SP<sub>BD</sub> values, which 625 are associated with a  $\delta^{18}O_{H2O}$  value of -5.6<sub>.</sub>% (Table 4). Although the *Diff* for this fitting is slightly higher, the other fittings must be rejected due to unrealistic  $\delta^{18}O_{H2O}$  values (of -1.7 and +3.7\_‰), hence <u>*f*\_{FD</u> MAP</sub> values must be between 0.04- and 0.09 %. Similarly, for ExperimentSoil 4, the only plausible fitting can be obtained for the smallest  $SP_{BD}$  values, which are associated with a  $\delta^{18}O_{H2O}$  value of -6.8 (Table 4) and indicate  $f_{FD}$  map values from  $\frac{1}{9}$ . It to  $\frac{1}{9}$ . Here this fitting also shows clearly the smallest *Diff* of only 0.01 (Table 4). However, 630 except for *ExperimentSoil* 4, where the Diff is smallest for the last fitting, the Diff values for other

experiment<u>Soil</u>s are very similar for different fittings with the largest values in Experiment<u>Soil</u> 3. A better fit (showing smaller *Diff* values) was not possible with any other  $SP_{BD}$  and  $\delta^{18}O_{H2O}$  values. Since the precision of <u>*r*<sub>15N</sub></u> (expressed in standard deviation in Table 3) was always  $\leq 0.05$ , this uncertainty of <u>*r*<sub>15N</sub></u> did not reduce the precision of the fitting (compare large ranges of  $\delta^{18}O_{H2O}$  and  $r_{MAP}$  values, respectively, in Table 4). The <u>*f*<sub>PD</sub> SP</u> ranged between 0 and approximately  $0.15 \frac{9}{20}$  (Table 5). The results obtained from SP/ $\delta^{18}O$  Map show <u>*f*<sub>PD</sub> MAP</u> reaching up to 0.14, 0.15, 0.09 and 0.20% for Experiment<u>Soil</u>s 1, 2, 3, and 4 respectively (Figure 3, Table 4, Table 5).

Table 4: Summary of the results provided by SP/ $\delta^{18}$ O Map for fraction of fungal denitrification ( $f_{FD}$  MAP) and N<sub>2</sub>O product ratio ( $t_{MAP}$ ) in the acetylated ( $+C_2H_2$ ) and non-acetylated ( $-C_2H_2$ ) treatments-varieties for 3-three possible  $SP_{N2O}$  values from bacterial denitrification ( $SP_{BD}$ ): mean (-1.9 ‰), maximal (3.7 ‰), and minimal (-7.5 ‰).  $t_2$ The  $\delta^{18}$ O values of soil water ( $\delta^{18}O_{H2O}$ ) were fitted to get the lowest difference (*Diff*) between product ratio determined with  $t_2$ N treatment ( $t_{15N}$ ) and SP/ $\delta^{18}$ O Map ( $t_{15N}$  and  $t_{15N}$ ). The most plausible fittings are shown in bolded (see discussion for reasons of this choice).

Experiment Soil	Variety	<u><i>r</i></u> 15N	SP <sub>BD</sub> [%6]	$\delta^{18}O_{\rm H2O}$ [%)	<u>ľ</u> <del>0</del> MAP	Diff	<u>_FD_</u> МАР [%]*
1	$-C_2H_2$	0.66	-1.9	-11.2	0.66	0.00	- <del>0.01<u>-1</u></del>
	$+C_2H_2$	1	-1.9	-11.2	1.00	0.00	<del>0.0</del> 2
	$-C_2H_2$	0.66	3.7	-6.1	0.65	0.01	<del>-0.14<u>-14</u></del>
	$+C_{2}H_{2}$	1	3.7	-6.1	1.00	0.00	<del>-0.16</del> -16
	$-C_2H_2$	0.66	-7.5	-14.9	0.66	0.00	<mark>0.0</mark> 8
	$+C_{2}H_{2}$	1	-7.5	-14.9	1.00	0.00	<mark>0.</mark> 14
2	$-C_2H_2$	0.94	-1.9	-6.3	0.90	0.04	<del>0.0</del> 1
	$+C_2H_2$	1	-1.9	-6.3	1.04	0.04	<del>0.0</del> 1
	$-C_2H_2$	0.94	3.7	-1.2	0.90	0.04	<del>-0.16<u>-</u>16</del>
	$+C_{2}H_{2}$	1	3.7	-1.2	1.04	0.04	<del>-0.18<u>-18</u></del>
	$-C_2H_2$	0.94	-7.5	-10.1	0.90	0.04	<del>0.</del> 13
	$+C_{2}H_{2}$	1	-7.5	-10.1	1.04	0.04	<mark>0.</mark> 15
3	$-C_2H_2$	0.61	-1.9	-1.7	0.54	0.07	<del>-0.03<u>-3</u></del>
	$+C_{2}H_{2}$	1	-1.9	-1.7	1.04	0.04	<del>-0.05<u>-5</u></del>
	$-C_2H_2$	0.61	3.7	3.7	0.54	0.07	<del>-0.14<u>-14</u></del>
	$+C_{2}H_{2}$	1	3.7	3.7	1.03	0.03	<del>-0.24<u>-24</u></del>
	$-C_2H_2$	0.61	-7.5	-5.6	0.53	0.08	<del>0.0</del> 4
	$+C_{2}H_{2}$	1	-7.5	-5.6	1.04	0.04	<mark>0.0</mark> 9
4	$-C_2H_2$	0.60	-1.9	-3.3	0.66	0.06	<del>0.</del> 15
	$+C_{2}H_{2}$	1	-1.9	-3.3	0.96	0.04	<del>-0.03</del> -30
	$-C_2H_2$	0.60	3.7	1.5	0.72	0.12	<mark>0.0</mark> 8
	$+C_{2}H_{2}$	1	3.7	1.5	0.91	0.09	<u>-0.21-21</u>
	$-C_2H_2$	0.60	-7.5	-6.8	0.61	0.01	<mark>0.</mark> 20
	$+C_{2}H_{2}$	1	-7.5	-6.8	0.99	0.01	<mark>0.</mark> 11

\*Negative values for <u>f<sub>FD MAP</sub></u> are non-realistic and therefore discarded for further interpretation.

Table 5: Ranges of the fraction of N<sub>2</sub>O produced by fungi  $(f_{FD})$  from four soil experiments<u>Soils</u> using four different approaches: Fungal fraction was calculated using a) the microbial inhibitor approach (modified SIRIN)  $(f_{EDmil})$ , b) the isotopomer endmember mixing approach (IEM) by SP isotope mixing balance  $(f_{ED,SP})$ , c) the IEM by SP<sub>22G</sub> isotope mixing balance (IEM) for results from variety  $-C_2H_2$  with reduction correction to calculate the  $SP_{N2O}$  values  $(f_{FD,SPcalc})$ , and d) the  $\delta^{18}O/SP$  Map  $(f_{FD,MP})$  with  $\delta^{18}O_{N2O}$  values from variety  $-C_2H_2$  and variety  $+C_2H_2$ . Negative values by IEM and  $\delta^{18}O/SP$  Map are assumed to be zero.

Experim <i>ent</i> Soil	<u>. [FD</u> mi] <mark>%</mark> ] <sup>a</sup>	<u>ffp</u> sp [%] <sup>b*</sup>	<u>[FD_SPcalc [%]</u> e*	<u>[FD_MAP [%]</u> d*
1	n.d.	θ <u>-14- to</u> θ <del>.</del> 15	<u>-6<mark>0-0.</mark> to</u> 19	<del>0-<u>&lt;</u>0.0</del> 2
2	<del>0.</del> 19 <u>- to</u> <del>0.</del> 37	θ <u>-18- to</u> <del>0.</del> 14	<u>-12<del>0-</del>to</u> <del>0.</del> 15	<del>0.0</del> 1- <u>to</u> 0.15
3	n.d.	0 <u>-25- to</u> 0.09	<u>-90- to 0.</u> 18	<del>0.0</del> 4- <u>to</u> 0.09
4	n.d.	<u>θ-23</u> - <u>to</u> <del>0.</del> 11	<u><del>0</del>1- to 0</u> .21	<del>0.</del> 11- <u>to</u> 0.20

660

655

<sup>a</sup>Fungal fraction o<u>f</u>  $N_2O$  production calculated <u>by</u> Eq. 3 <u>taking variations of three replicates into account</u>, <sup>b</sup>Fungal fraction o<u>f</u>  $N_2O$  production calculated by Eq. 4 for variety  $+C_2H_2$  with assuming  $SP_{N2O}$  values of  $N_2O$  produced by bacteria were 3.7 <del>% (resulting in negative fraction and therefore set to zero)</del> or -7.5 %. <u>Using the minimum and maximum  $SP_{N2O}$  values known for bacteria resulted in a  $f_{FD}$  <u>sp</u> range.</u>

<sup>c</sup>Eq. 4 to solve for fungal fraction in variety  $-C_2H_2$  with assuming  $SP_{N20}$  values of N<sub>2</sub>O produced by bacteria was 3.7 (resulting in negative fraction and therefore set to zero) or -7.5 ‰ and using reduction correction with  $\eta_r$ =-6 ‰ to calculate  $SP_{prod}$  values (Senbayram et al., 2018; Yu et al., 2020). Using the minimum and maximum  $SP_{N20}$  values known for bacteria resulted in a  $f_{FD}$  speak, range.

<sup>d</sup>Fungal fraction <u>onof</u> N<sub>2</sub>O production calculated by SP/ $\delta^{18}$ O Map with assuming most probable SP<sub>N20</sub> values from bacterial denitrification (according to Table 4). <u>-Using the minimum and maximum SP<sub>N20</sub> values known for bacteria and ranges of fitted  $\delta^{18}$ O values resulted in a f<sub>FD MAP</sub> range.</u>

670 <u>\*Negative values for *f<sub>FD SP</sub> f<sub>FD SPcalcx</sub> f<sub>FD MAP</sub>* are non-realistic and therefore discarded for further interpretation.</u>

n.d.-not determined because of insufficient inhibition.





680

Figure 23: SP/ $\delta^{18}$ O isotope mapping approach (SP/ $\delta^{18}$ O Map) to estimate the contribution of bacteria or fungi to N<sub>2</sub>O produced according to Lewicka-Szczebak et al. (2017) and Lewicka-Szczebak et al. (2020)<sub>k</sub> The isotopic values for natural abundance treatments with acetylene addition (+ $C_2H_2$ , empty symbols) and without acetylene addition (- $C_2H_2$ , corresponding filled symbols) are shown for four experimentSoils (1 to -4). The grey rectangles indicate expected ranges of isotopic signatures for heterotrophic bacterial denitrification (BD) and fungal denitrification (FD) (Yu et al., 2020). The black solid line is the mixing line connecting the average expected values for BD and FD, while the red solid line is the main SP values for BD) line and the red dashed line is the minimum reduction line (for the minimal SP<sub>N2O</sub> values for BD).

685

3.6 SP<sub>M2C</sub> values of N<sub>2</sub>O produced by the fungal soil community

Solving Eq. 4 for  $SP_{FD}$  enables to calculate  $SP_{N2O}$  values from the fungal soil community for Experiment 2 (Table 6). Estimates for the ranges of  $F_{\mu D}$  and  $F_{BD}$  from the results  $(+C_2H_2)$  of the modified SIRIN were obtained ( $F_{\mu Dmt}=0.19-0.37$  and  $F_{BD}=1-F_{\mu Dmt}$  resulted in a range between 0.63 and 0.81, respectively, see section "3.4 Fungal contribution to  $N_2O$  production from denitrification by microbial inhibitor approach (modified SIRIN)"). The  $SP_{\mu\nu\sigma d}$  values of  $N_2O$  ( $SP_{\mu\nu\sigma d}=-1.4$  %) of the respective treatment A (Table 2, variety  $+C_2H_2$ ) served to ealculate  $SP_{N2O}$  values for fungal denitrification for Experiment 2. Assuming -7.5 or 3.7 % for the bacterial  $SP_{N2O}$  endmember values of  $N_2O$  (Toyoda et al., 2005; Sutka et al., 2006; Yu et al., 2020) resulted in  $SP_{\mu\nu}$  values between -10 % ( $SP_{BD}=-3.7$  %) and 25 % ( $SP_{BD}=-7.5$  %) (Table 6). The respective  $SP_{\mu\nu}$  value for variety  $C_2H_2$  was in a very similar range between -17 % and 27 % (Table 6) using  $SP_{\mu\nu\sigma d}$  values ( $SP_{\mu\nu\sigma d}=-1.0$  %) of the respective treatment A (Table S1).(Sutka et al., 2008; Rohe et al., 2014a; Maeda et al., 2015)

700

Table 6:  $SP_{FD}$  values (i.e.  $SP_{N2O}$  values of  $N_2O$  produced by fungi) by solving Eq. 4 using  $F_{FDmr}$  and  $F_{BD}$  from results of modified SIRIN approach and using  $SP_{prod}$  values of varieties  $+C_2H_2$  and  $-C_2H_2$  of Experiment 2.

Treatment	SP <sub>prod</sub> [‰]	<i>\$₽<sub>₿₽</sub>_</i> [‰] <sup>*</sup>	₽ <sub>₽₽mi</sub> ₿	₽ <sub>₿₽</sub> ₿	<del>\$₽</del> ₽₽ <mark>₩</mark>
		<del>-7.5</del>	<del>0.19</del>	<del>0.81</del>	<del>25</del>
	-1.4	<del>3.7</del>	<del>0.19</del>	<del>0.81</del>	-23
+ <del>C2A</del> 2		<del>-7.5</del>	<del>0.37</del>	<del>0.63</del>	<del>9</del>
		<del>3.7</del>	<del>0.37</del>	<del>0.63</del>	<del>-10</del>
$-C_2H_2$	<del>-1.0</del>	<del>-7.5</del>	<del>0.19</del>	<del>0.81</del>	<del>27</del>

Feldfunktion geändert

<del>3.7</del>	<del>0.19</del>	<del>0.81</del>	<del>-17</del>
<del>.7.5</del>	<del>0.37</del>	<del>0.63</del>	<del>10</del>
<del>3.7</del>	<del>0.37</del>	<del>0.63</del>	<del>_9</del>

*SP<sub>N20</sub>* endmember values of bacterial denitrification were taken for calculation (Eq. 4) according to studies with pure cultures (Toyoda et al., 2005; Sutka et al., 2006; Yu et al., 2020). <sup>b</sup>Ranges of F<sub>FDm</sub>, and F<sub>BD</sub> were calculated using the modified SIRIN approach.

705

# 4. Discussion

To our knowledge, this was the first attempt to determine SP<sub>N20</sub> values by fungi or bacteria from soil communities using microbial growth inhibitors with a modification of SIRIN and comparing microbial inhibitor and isotopic approaches (IEM and SP/8<sup>18</sup>O Map) to estimate fungal contribution to N<sub>2</sub>O production from 710 denitrification in anoxic incubation. Using IEM-isotopic approaches revealed that the fungal contribution to N2O production was small ( $f_{FD} _{SP} \le \Theta_{+} 15 \frac{0}{20}$  or  $f_{FD} _{MAP} \le \Theta_{+} 20 \frac{0}{20}$  in the three soils tested (Table 5). The modified SIRIN approach was not successful, because large amounts of non-inhibitable N<sub>2</sub>O production were observed with all four Soils (Table 2, Table 3). The fungal fraction producing  $N_2O$  during denitrification ( $f_{FDmi}$ ) was oOnly rimentestimated for Soil 2, with where significantly smaller N2O production in treatment D was observed compared to that of treatment A and resulted modified SIRIN allowed the calculation of the fungal 715 fraction producing N<sub>2</sub>O during denitrification (F<sub>FDmt</sub> betweenin a range of -0.19 and to 0.37 %-in Experiment 2), which was largerwas probably overestimated due to uncertainties resulting from the large N<sub>2</sub>O production of non-inhibitable sources. than the by two isotope approaches (≤0.20). While the three approaches coincided in showing dominance of bacterial denitrification, the isotopic approaches yielded similar small estimates of for  $f_{FD}$ 720 ( $\leq 20$  %) and thus did not confirm largest  $f_{ICDmL}$  of ExperimentSoil 2. The strict application of the SIRIN method prescribes proof of selectivity of the inhibitors (i.e., streptomycin should not inhibit fungi and cycloheximide should not inhibit bacteria). The All SIRIN results obtained with respect to N2O production by the fungal or bacterial fraction were rather unsatisfactory and led to unsolved questions, which are discussed in the following sections.

# 725 4.1 Experimental setup and inhibitor effects

Inhibitor effects, expressed by smaller N<sub>2</sub>O production with selective inhibitors (treatments B, C and D) compared to treatments without inhibitors (A), were only minor in the present study. In accordance with other studies, N<sub>2</sub>O production was analysed after the addition of glucose as substrate (Laughlin and Stevens, 2002; McLain and Martens, 2006; Blagodatskaya et al., 2010; Long et al., 2013). Glucose initiates the growth of active heterotrophic organisms. Since pure cultures were shown to synthesize enzymes capable of denitrification within two to three hours (USEPA, 1993), a-pre-incubation of soil under anaerobic conditions is not needed. Thus, when gas sample collection started organisms should have produced denitrifying enzymes and microbial growth of initially active organisms should have started too. However, in accordance to Anderson and Domsch (1975) experimental duration should be as short as possible to ensure the CO<sub>2</sub> production by initially active organisms only. Thus, short-time incubation is recommended when conducting a modified SIRIN approach, as the incubation period should cause changes in conditions for microorganisms and initiate growth on the one hand, while it should avoid the use of inhibitors as C sources by organisms on the other.

With incubation time, production rates of  $CO_2$  decreased, probably because experimental incubation conditions provoked unfavourable conditions and physiological changes, e. g. due to anaerobic conditions or local substrate depletion (e. g. C supplied as glucose). Decreasing  $CO_2$  fluxes might also be explained by  $CO_2$  accumulation in

- pore space as this effect is shown by modelled diffusive fluxes from soil in closed systems (Well et al., 2019).
   Previous studies found much larger inhibitor effects by pre-incubating the soil with selective inhibitors (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013; Chen et al., 2014). It is therefore important to discuss considerable differences among the experimental design of the present study compared to that of other
   studies (e. g., Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2010).
- The conventional practice of SIRIN implies determination of  $c_{opt}$ (glucose),  $c_{opt}$ (streptomycin) or  $c_{opt}$ (cycloheximide) with an "Ultragas 3" CO<sub>2</sub> analyser (WösthoffCo., Bochum) (Anderson and Domsch, 1973) with continuous gas flow. We used this method to determine optimal concentrations for SIRIN in the preexperiment and used these concentrations for the modified SIRIN approach as well. This optimization procedure
- 750 was not used in other studies (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013). We supposed that optimal concentrations for CO<sub>2</sub> respiration should work as well for denitrification, if both inhibitors inhibit the denitrification process as well. However, although SIRIN has so far been tested with isolated cultures and soils for microbial growth for CO<sub>2</sub> production only (Anderson and Domsch, 1973, 1975), information on N<sub>2</sub>O producing processes, especially denitrification, is still lacking and should be investigated in further studies. Additionally, as presented by Ladan and Jacinthe (2016), the bactericide bronopol and the
- <u>fundice studies reactions</u>, as presented by Educational and studient (2010), the educational end of the fundice end of the studies of the st
- Previous studies that found much larger inhibitor effects were conducted with pre-incubating the soil with selective inhibitors (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013; Chen et al., 2014).-<u>The-In contrast to that, the</u> experimental design of our incubation setup was<u>without soil pre-incubation</u> with selective inhibitors to minimize disturbance of the soil microbial community and our approach was thus <u>-</u> however, but this was-in agreement with the original SIRIN method for respiration (Anderson and Domsch, 1973, 1975, 1978) without soil pre-incubation with selective inhibitors to minimize disturbance of the soil microbial community.-Another study performing similar experiments without pre-incubation with inhibitors did
- not find effectiveness of application of both antibiotics during long-term application (up to 48 h) (Ladan and Jacinthe, 2016)-, although streptomycin and cycloheximide are commonly used to inhibit denitrification of selective groups. Nevertheless, as we expected that pre-incubation with selective inhibitors would induce changes in the F:B ratio of soil, we decided to conduct the modified SIRIN approach without a pre-incubation step. This assumption was supported by findings of Blagodatskaya et al. (2010), where pre-incubation of about one to twenty hours with cycloheximide resulted in increasing inhibitor efficiency with time, while this was not the case when pre-incubating with streptomycin. ConsequentlyThis suggests that ,-microbial communities, might change after exposition to cycloheximide.
- 775

740

In the present study, even with both growth inhibitors (treatment D), N<sub>2</sub>O production was large in all experiments, i.e., in most cases not significantly smaller than in the other three-treatments A, B or C. Thus, we suppose similar contributions of non-inhibitable organisms and processes in all treatments. Non-inhibitable organisms could be, for example, bacteria or fungi that are not in growth stage or may be not affected by inhibitors. Recently, Pan et al. (2019) summarized findings of other studies and pointed out that some microorganisms can use inhibitors as growth substrates, that dead organisms may serve as energy sources for

	others, and that interactions of microbial species may change due to non-inhibitable organisms occurring in soil
780	communities. Non-inhibitable organisms could be archaea as well, which are also known to be capable of
	denitrification (Philippot et al., 2007; Hayatsu et al., 2008). It is known that archaea are not affected by
	streptomycin or cycloheximide (Seo and DeLaune, 2010). However, effects of archaeal occurrence in soil or
	secondary effects on fungi or bacteria were not tested in this study. Additionally, abiotic N2O production cannot
	be quantified with the experimental setup, but might be contributing to each inhibitor treatment.
785	In summary, the present experimental setup without pre-incubating soil samples with selective inhibitors was not
	successful in complete inhibition of bacterial or fungal denitrifiers. Although pre-incubation with selective
	inhibitors may lead to more successful inhibition, we do not recommend this due to induced changes in soil
	communities. For further studies focusing on application of modified SIRIN to determine the fraction of
	bacterial or fungal N <sub>2</sub> O derived from denitrification a method validation using also different inhibitors is
790	recommended.
	Inhibitor application without pre-incubating with inhibitors was contrary to previous studies focusing on N2O
	production (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013) and we suppose
	expected that pre-incubation with selective inhibitors would induce , that pre-incubation with selective inhibitors
	ehanges the F:B ratio compared to the undisturbed soil considerably more than soil incubation without this pre-
795	incubation step. Additionally, although Blagodatskaya et al. (2010), did not find more inhibitor efficiency after a
	period of 1 to 20 hours of pre-incubation with streptomycin, they found greater inhibitor effects of
	cycloheximide with pre incubation phases. This could indicate that the microbial distribution changed after
	exposition to this inhibitor. Anderson and Domsch (1975) stated already that CO2 production of initially active
	organisms can only be ensured up to six or eight hours of experimental duration and biomass activity is changed
800	by both inhibitors. <u>conditions</u>
	It has to be noticed that pre-incubation in previous studies was without glucose, while $N_2O$ production was
	analyzed after the addition of glucose as substrate in the present as well as previous studies (Laughlin and
	Stevens, 2002; MeLain and Martens, 2006; Blagodatskaya et al., 2010; Long et al., 2013). Glueose initiates the
	growth of active heterotrophic organisms. Pre incubation under denitrifying conditions is not needed for
805	microorganisms to produce denitrifying enzymes as pure cultures synthesized enzymes capable of denitrification
	within two to three hours (USEPA, 1993). We started gas sample collection after two or four hours, when
	organisms should have produced denitrifying enzymes and microbial growth of initially active organisms should
	have started. With incubation time, production rates of $CO_2$ decreased, probably because experimental
	incubation conditions provoked unfavorable conditions and physiological changes, e.g. due to increasing partial
810	pressure within the closed jars.
	The conventional practice of SIRIN implies determination of $c_{opt}$ (glucose), $c_{opt}$ (streptomycin) or
	eopp(cycloheximide) with an "Ultragas 3" CO <sub>2</sub> analyzer (WösthoffCo., Bochum) (Anderson and Domsch, 1973)
	with continuous gas flow and we used this method to determine optimal concentrations for SIRIN and used these
	concentrations for the modified SIRIN approach as well. This optimization procedure was not used in other
815	studies (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013). We supposed that optimal
	concentrations for CO <sub>2</sub> respiration cshould work as well for denitrification, if both inhibitors are apt to inhibit the
	denitrification process as well. However, although SIRIN has so far been tested with isolated cultures and soils
	for microbial growth on agar and for CO <sub>2</sub> production (Anderson and Domsch, 1975, 1973), but information on

# denitrification, is still lacking and should be investigated in furth and pro

# 820

ī

ing also different inhibitors

825

 $\mathbf{D}$   $\mathbf{N}$   $\mathbf{O}$ inhibitable organisms could be, for example, bacteria or fungi that are not in inhibitors Pan et al (2010)These organisms could be (Philippo 2007. Has <del>rre not affected by streptomycin or eveloheximide (Seo and DeLaune, 2010). However, effects of</del> 830 Ladan and Jacinthe (2016) did not find effective inhibition of denitrification by either inhibitor for denitrification although streptomycin and cycloheximide are commonly used to inhibit denitrification of with different inhibitors such as the bactericide by fungicide captan presented by Ladan and Jacinthe (2016), should be conducted to evaluate inhibition approaches 835 and isotopic endmember approaches.

## 4.3-2 Is SIRIN without C<sub>2</sub>H<sub>2</sub> suitable to examine the fungal contribution to N<sub>2</sub>O production in soil?

In order to determine SP<sub>N2O</sub> values without alteration by partial reduction of N<sub>2</sub>O to N<sub>2</sub>, C<sub>2</sub>H<sub>2</sub> was used to quantitatively block N<sub>2</sub>O reduction during denitrification. We found the expected effect of C<sub>2</sub>H<sub>2</sub> application, i.e. larger N<sub>2</sub>O production rates in variety  $+C_2H_2$  compared to variety  $-C_2H_2$ . Calculated product ratios varied 840 between 0.5 and 0.95 ( $r_{15N}$ ) in all <u>Soilssoils</u>, showing that N<sub>2</sub>O reduction can have significant effects on measured N2O production and isotopic values. The product ratio is controlled by the reaction rate or by the activity of enzymes capable of N2O reduction (Nos) in the system.

The calculated  $\underline{r}_{C2H2}$  was within the same range as  $\underline{r}_{L5N}$  in ExperimentSoil 2, 3 and 4 (maximal 9.% difference), providing theindicating effective blockage of N<sub>2</sub>O reductase in variety  $+C_2H_2$  in these <u>Soilsseils</u>. Only in 845 *ExperimentSoil*  $I_{k} r_{LSN}$  and  $r_{C2H2}$  differed by about 34\_% with larger calculated reduction in the *tracer\_traced* variety, which might be explained by potential point to incomplete inhibition by the  $C_2H_2$  method. Nadeem et al. (2013)-found some Artifacts with C<sub>2</sub>H<sub>2</sub> were found in previous studies, which resultinged in smaller N<sub>2</sub>O production rates due to NO oxidation accelerated by  $C_2H_2$  application in the presence of very small oxygen (O) amounts\_-(2019) mL L<sup>+</sup>)\_(Bollmann and Conrad, 1997a, b; Nadeem et al., 2013). Moreover, incomplete C<sub>2</sub>H<sub>2</sub> 850 diffusion into denitrifying aggregates might also lead to incomplete N2O reductase blockage (Groffman et al.,

2006). Both potential methodological errors cannot be excluded for ExperimentSoil 1. For the other three experiment <u>Soil</u>s (2, 3 and 4)<sub>2</sub> it can be supposed that the isotopic signature of N<sub>2</sub>O of variety

 $+C_2H_2$  showed isotopic signatures of produced N<sub>2</sub>O without influences of N<sub>2</sub>O reduction. By comparing varieties  $-C_2H_2$  and  $+C_2H_2$ , isotopologue values of all these <u>Soilseeils</u> (except  $\delta^{15}N^{\text{bulk}}_{N2O}$  values of 855 ExperimentSoil 2) of variety  $-C_2H_2$  were significantly larger than those that of variety  $+C_2H_2$ . The enrichment of residual N2O in heavy isotopes results from the isotope effect associated with N2O reduction (Jinuntuya-Nortman et al., 2008; Well and Flessa, 2009; Lewicka-Szczebak et al., 2014). This explains why C2H2 application is

### Feldfunktion geändert

essential for analyzing analysing N<sub>2</sub>O produced by different microbial organism groups from soil using solely the modified SIRIN approach without additional isotopic approaches.

- 860 Moreover, when applying SIRIN without quantifying N<sub>2</sub>O reduction, fungal denitrification is potentially overestimated due to the impact of SIRIN inhibitors on N2O reduction. It is evident that N2O fluxes represent net N2O production, i.-e. the difference between gross N2O production by the microbial community and N2O reduction, mainly by heterotrophic bacterial denitrifiers (Müller and Clough, 2014). The goal of SIRIN application has been to determine the contribution of fungi and bacteria, respectively, to net N2O production. It 865 has been shown that N<sub>2</sub>O released by microorganisms to air\_-filled pore space can be partially consumed by denitrifiers before being emitted (Clough et al., 1998). This means that fungal N<sub>2</sub>O can also be subject to reduction by bacterial denitrifiers. Consequently, inhibiting bacterial denitrification by SIRIN would enhancelead the measured flux of fungal to an overestimation of fungal contribution to N<sub>2</sub>O-production. Until now, this effect has not been considered in previous-SIRIN papers on fungal N<sub>2</sub>O (e. g. Laughlin and Stevens, 2002; Ladan and 870 Jacinthe, 2016; Chen et al., 2014). This effect can only be evaluated by measuring N<sub>2</sub>O reduction in all inhibitor treatments as in our study. If true, the N<sub>2</sub>O reduction with bacterial inhibition should be smaller than that of the treatments without inhibition or with fungal inhibition. Though However, with fungal inhibition, N2O reduction is also assumed to be smaller than-that without inhibition, because N<sub>2</sub>O produced by fungi is missed for bacterial reduction.
- The product ratio is a measure for the N<sub>2</sub>O reduction to N<sub>2</sub>. However, regarding the <u>r</u><sub>15N</sub>, there was no evidence of different N<sub>2</sub>O reduction effects between the SIRIN treatments. The <u>r</u><sub>C2H2</sub> also revealed indistinguishable values between SIRIN treatments in <u>ExperimentSoil</u> 1 and 4, but it was slightly larger in <u>ExperimentSoil</u> 3 with bacterial inhibition compared to the other treatments. However, this effect was very small, which would only cause small overestimation of fungal contribution. The smallest N<sub>2</sub>O reduction was found in <u>ExperimentSoil</u> 2
  (<u>r</u><sub>C2H2</sub> values near 1), with smallest <u>r</u><sub>C2H2</sub> with bacterial inhibition (0.81). This could result in an overestimation of bacterial is measured.

The <u>r</u><sub>15N</sub> and <u>r</u><sub>C2H2</sub> were between 0.5 and 1 and N<sub>2</sub>O reduction was thus never consuming most of the produced N<sub>2</sub>O. Hence, both the C<sub>2</sub>H<sub>2</sub> and <u>Streptomycin\_streptomycin\_</u>effects on SIRIN results were probably low. <del>But</del> <u>However, as</u> the product ratio in soil denitrification exhibite<u>ds</u> the full range from 0 to 1, meaning that this effect can be quite relevant and must thus be considered in future studies. <u>Therefore, we recommend to estimate the</u> <u>effectiveness of C<sub>2</sub>H<sub>2</sub> in blocking the N<sub>2</sub>O reductase by performing parallel <sup>15</sup>N approaches with and without C<sub>2</sub>H<sub>2</sub> in studies using the modified SIRIN to determine the fraction of bacterial or fungal N<sub>2</sub>O production.</u>

# 4.4-3\_SP<sub>N20</sub> values of N<sub>2</sub>O produced by microbial communities

890

As discussed above, all  $N_2O$  fluxes of –modified SIRIN treatments of *Soil 1, 3* and *4* were largely affecteddominated by  $N_2O$  from non-inhibitable organisms or processes, which of course have an impact on *SP<sub>N2O</sub>* values of all SIRIN treatments. This made it impossible to calculate *SP<sub>N2O</sub>* values for active bacteria or fungi (modified SIRIN B and C), also with *Soil 2*, where a relatively large  $N_2O$  production was observed with treatment D (Sutka et al., 2008; Rohe et al., 2014a; Maeda et al., 2015). This is discussed in more detail in (see section 4.4).

- 895  $\mp$ Despite this, the  $SP_{N20}$  values from  $+C_2H_2$  variety as well as  $SP_{prod}$  values (i.e., reduction corrected  $SP_{N20}$ values of  $-C_2H_2$  variety)values of each *Soil*, represented by treatment A of (modified SIRIN), indicated predominantly bacteria to be responsible for N<sub>2</sub>O production during denitrification, assuming that results of
  - 30

 $SP_{N2O}$  values of denitrification by pure bacterial cultures is transferable to bacteria of soil communities contributing to denitrification. Also in many soil incubation studies small-SP<sub>N20</sub> values (without reduction 900 effects) within the range of bacterial pure cultures have been found The latter assumption has been confirmed repeatedly in soil incubation studies, where in absence of N2O reduction smallest SPA2O values have been found that were within the range of bacterial pure cultures (Lewicka-Szczebak et al., 2015; Lewicka-Szczebak et al., 2017; Senbayram et al., 2018). Therefore, there was no unequivocal evidence of fungi contributing to N<sub>2</sub>O production during denitrification, although here, the isotopic approaches revealed a fungal contribution to N2O 905 production during denitrification of up to 0.20 % on N2O production during denitrification.

- The  $SP_{N2O}$  values of treatment A within-variety  $+C_2H_2$  within treatment A showed that are not affected by reduction effects the signature of produced N<sub>3</sub>O was not affected by reduction effects and therefore might give evidence of the microbial community contributing to N<sub>2</sub>O production regarding differences in SP<sub>N2O</sub> values of pure bacterial or fungal culture studies (Sutka et al., 2006; Sutka et al., 2008; Frame and Casciotti, 2010; Rohe et
- 910 al., 2014a). However, variations in  $SP_{N2O}$  values of treatments A of variety  $+C_2H_2$  are were very small and do not give a clear evidence of any differences in microbial soil community producing N<sub>2</sub>O.\_-Lewicka-Szczebak et al. (2014) analyzed analyzed  $SP_{N2O}$  values of denitrification with blockage of N<sub>2</sub>O reduction by C<sub>2</sub>H<sub>2</sub> for the same soils as used in the present study for Experiment(Soil 1 and 4 as well as ExperimentSoil 3) and revealed SP<sub>N20</sub> values between -3.6 and -2.1 %, which is similar to the respective  $SP_{N2O}$  values of the present study from -4.9 to 915 -0.4 ‰. This reinforces the conclusion that bacteria dominated gross N<sub>2</sub>O production under anoxic conditions in
- both these soilsstudies. Obviously,
- $SP_{prod}$  values (variety -C<sub>2</sub>H<sub>2</sub>) differed from  $SP_{N2O}$  values (variety +C<sub>2</sub>H<sub>2</sub>), which may result from deviations between the actual fractionation factor that was not estimated in the present study and the used fractionation factor of -6 ‰ adapted from the literature (Yu et al., 2020). If so, we could assume smaller fractionation effects 920 in the present study as decreasing this average fractionation factor would lead to increasing SPprod values, which in turn would result in values more similar to SP<sub>N20</sub> values of variety -C<sub>2</sub>H<sub>2</sub>.
- wer, other studies found larger  $SP_{N2O}$  values of produced N<sub>2</sub>O (up to +621 %) unaffected by the reduction effect of up to +6 ‰ (Köster et al., 2013a)(Senbayram et al., 2018, 2020), most probably as a result of larger contributions of fungi to N<sub>2</sub>O production. However, those results were obtained in an experimental setup with 925 ambient oxygen concentration, without glucose amendment and without C,H, inhibition of N,O reduction since N2 gas fluxes were directly measured. It was also discussed before that short time incubations under static eonditions as presented here, may promote bacterial over fungal growth, which may also be transferable to denitrification activity by both organism groups (Lewicka Szczebak et al., 2017; Lewicka Szczebak et al., 2014). Additionally to this, the selection use of glucose as substrate in the selected concentration may further promote 930 bacteria compared to fungi even more (Koranda et al., 2014; Reischke et al., 2014).

935

4.5-δ<sup>18</sup>O<sub>N2O</sub> values

The analysis of  $\delta^{H9}O_{N2O}$  values can give information about O exchange between water and denitrification various microorganisms . The range of  $\delta^{I3}O_{N2G}$  values in our study for variety  $+C_2H_2$  (7.5 to similar to the range found by for the same soils (4.8 to 16.3 %), where almost complete O exchange with soil water was documented. Hence, for this study the O exchange was probably also very high. , there were no remarkable differences in  $\delta^{ts} O_{_{\!N\!2G}}$  values among treatments within one variety and soil efore we assume no differences in O exchange among the treatments.

The information on  $\delta^{18}O_{N2O}$  values combined with known  $\delta^{18}O_{H2O}$  values is also precious information for differentiation between N<sub>2</sub>O mixing and reduction processes (Lewicka Szczebak et al., 2017). However, fFor this study,  $\delta^{18}O_{H2O}$  values were not analyzed. However, due to parallel *traced* variety experiments, we could determine possible  $\delta^{18}O_{H2O}$  values for the particular SP<sub>N2O</sub> values of bacterial denitrification mixing endmembers (Table 4). Since the  $\delta^{18}O_{H2O}$  value for the particular geographic region can be assessed based on the known isotopic signatures of meteoric waters <u></u> the most plausible ranges of  $\delta^{18}O_{H2O}$  values can be used to indicate the plausible ranges of <u>MAP</u> values. In case of precisely determined  $\delta^{18}O_{H2O}$  values, the calculated <u>MAP</u> values could be more precise, however, here we show that in case of missing  $\delta^{18}O_{H2O}$  values but known product ratio, the SP/ $\delta^{18}O$  Map can also provide information on N<sub>2</sub>O production pathway contributions.

# 4.6-4 Potential influence of Co-denitrification

940

945

The influence of co-denitrification, which is predominantly associated to-with fungi (Spott et al., 2011), may 950 have a large impact on N<sub>2</sub>O production. E.g., since-Laughlin and Stevens (2002) found 92 % of N<sub>2</sub>O production in their experiment to be derived to 92% from co-denitrification and only 8\_% from denitrification. So far, there is no study on SP<sub>N2O</sub> values of N<sub>2</sub>O produced by co-denitrification. Co-denitrification could have been a contributing process in Experiment Soil 4. When N in N2O originates only from <sup>15</sup>N-labeled soil NO3, measured  $\delta^{15}$ N<sup>bulk</sup><sub>N2O</sub> values as well as the <sup>15</sup>N enrichment of the labelled N pool producing N<sub>2</sub>O (*a<sub>n</sub>*) should show identical 955 <sup>15</sup>N enrichment to the labeled soil NO<sub>3</sub>. During co-denitrification, when one N atom in N<sub>2</sub>O originates from labeled NO<sub>3</sub><sup>-</sup> and the other one from an<del>other</del> unlabeled <del>and unknown</del> N source, this results in  $a_p$  values and <sup>15</sup>N enrichment of produced N2O smaller than the respective enrichment of the NO3 pool. The <sup>15</sup>N enrichment of soil NO3<sup>-</sup> in Soil 4 was about 60\_% larger than the analyzed <sup>15</sup>N enrichment in N2O, leading to the assumption that N<sub>2</sub>O was produced not only by denitrification. We also calculated  $a_p$  values of the other three experiments Soils 960 (data not shown) which coincided with the <sup>15</sup>N enrichment of N<sub>2</sub>O (Table 3), showing no indication of codenitrification. Since  $a_n$  would not be affected by contributions of unlabelled unlabelled N<sub>2</sub>O we can thus exclude the possibility that this smaller enrichment could be caused by dilution of enriched N<sub>2</sub>O from denitrification by N2O production from an unknown N source and thus verified that this was due to formation of hybrid N2O, probably via co-denitrification (Spott et al., 2011). In the other experiments there was no indication of codenitrification being relevant for N2O production since <sup>45</sup>N enrichments of NO3<sup>-</sup> and N2O coincided. The 965 question arises, why hybrid N2O formation was only found when the loamy sand was sampled in summer (June, Experiment-Soil 4) but not when it was sampled during winter (December, Experiment-Soil 1). Since environmental conditions may vary within one year in arable soils, soil pH, F:B ratio, or biomass as presented in Table 1 might have been different for samples collected in summer 2011. However, as the soil was amended with 970 C and N, the current state of the soil was changed before incubation in any case. Although soil properties, microbial community or biomass may have changed over time, we assumed pre-incubating the soil for seven days, applying C and N, and changing the environmental conditions during denitrification induced a rapid growth of specific organisms. It has to be presumed that the denitrifying community and the abundance of these organisms in incubation experiments may differ from the community in the field. Information on substrates for 975 eo denitrification, i.e. NO2<sup>-</sup> and NH4<sup>+</sup> or certain organic N compounds could have been different due to seasonal Moreover, seasonal impacts on microbial community could have been relevant.\_\_Since these possible factors were not assessed in our study and their impact on co-denitrification is still poorly understood, it is

980

985

990

995

1000

currently not possible to give an answer here. Thus, only the  $SP_{N2O}$  values in <u>Soil</u> Experiment 4 might be influenced by co-denitrification. But since SP<sub>N2O</sub> values of the acetylated treatments of Soil Experiment-4 coincided with the SP<sub>N20</sub> value range of bacterial denitrification and also with SP<sub>N20</sub> values of the other Soilexperiments, our data give no indication that co-denitrification produces N2O with SPN2O values differing from bacterial denitrification.

# 4.7-5 Calculating the fungal fraction contributing to N<sub>2</sub>O production and SP<sub>PD</sub> values

Due to the inefficiency of the inhibition of microbial inhibition regarding N2O production in most cases, calculation of from contributing to N2O production was only possible for ExperimentSoil 2 only, although even this calculated value included inaccuracies. The isotopic approaches, however, which are independent of modified SIRIN results, yielded similar estimates of  $f_{FD}$  for all Soils, while it has to be emphasised that estimations based on stable isotope approaches do not rely on N<sub>2</sub>O production of modified SIRIN results. As recently published (Wu et al., 2019), uncertainty analysis is a complex issue and large uncertainties of the results from the SP/8<sup>18</sup>O Map approach can be assumed when all the possible sources of errors are taken into account. Regarding the presented application of SP/ $\delta^{18}$ O Map, calculation would be more precise when measuring  $\frac{\delta^{18}O(H_2O)}{\delta^{18}O(H_2O)}$  rather than using the fitted  $\delta^{18}O(H_2O)$  values. as discussed above Still, the analysis of  $\delta^{18}O_{N2O}$  values can give information about O exchange between water and denitrification intermediates by various microorganisms (Aerssens et al., 1986; Kool et al., 2007; Rohe et al., 2014b; Rohe et al., 2017). The range of  $\delta^{18}O_{N2O}$  values in our study for variety  $+C_2H_2$  (7.5 to 19.0 %) was quite similar to the range found by Lewicka-Szczebak et al. (2014) for the same soils (4.8 to 16.3 %), where almost complete O exchange with soil water was documented. Hence, for this study the O exchange was probably also very high. There were also no remarkable differences in  $\delta^{18}O_{N20}$  values among treatments within one variety and soil and therefore we assume particular SP<sub>N2O</sub> values of bacterial denitrification mixing endmembers could be determined (Table 4). Since the

- no differences in O exchange among the treatments. The information on  $\delta^{18}O_{N2O}$  values combined with known  $\delta^{18}O_{H2O}$  values is also precious information for differentiation between N<sub>2</sub>O mixing and reduction processes (Lewicka-Szczebak et al., 2017). Due to parallel *traced* variety experiments, possible  $\delta^{18}O_{H2O}$  values for the
- $\delta^{18}O_{H2O}$  value for the particular geographic region can be assessed based on the known isotopic signatures of meteoric waters (Lewicka-Szczebak et al., 2014; Stumpp et al., 2014; Lewicka-Szczebak et al., 2017; Buchen et 1005 al., 2018), the most plausible ranges of  $\delta^{18}O_{H2O}$  values can be used to indicate the plausible ranges of  $f_{FD MAP}$ values. Here we showed that in case of missing  $\delta^{18}O_{H2O}$  values but known product ratio, the SP/ $\delta^{18}O$  Map can also provide information on N<sub>2</sub>O production pathway contributions. Comparing the modified SIRIN with the isotopic approaches revealed that the fungal fraction contribution to N<sub>2</sub>O production was consistently estimated to be smaller (about  $0.28 \frac{\%}{10}$  in modified SIRIN,  $\leq 0.15 \frac{\%}{10}$  with IEM,  $\leq 0.20 \frac{\%}{10}$  with SP/ $\delta^{18}$ O Map) than the 1010
  - bacterial fraction. Although we did not obtain a very clear picture of various microorganisms contributing to N2O production due to the large uncertainties of the calculated fractions, all approaches coincided by showing dominance of bacterial N2O. In contrast to SIRIN, the isotopic approaches yielded similar estimates of FFF for all experiments.
- In some soil studies using helium incubations, the SP<sub>Prod</sub> values obtained by correction for the reduction effect on 1015 SP<sub>N20</sub> values showed significantly larger values than SP<sub>N20</sub> of bacterial denitrification (Köster et al., 2013a; Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2014; Senbayram et al., 2018; Senbayram et al., 2020). However, those results were obtained in an experimental setup with ambient oxygen concentration. Short

Feldfunktion geändert

incubations under static conditions as presented here may, however, promote bacterial over fungal growth, which may also be transferable to denitrification activity by both organism groups (Lewicka-Szczebak et al., 2014;
 Lewicka-Szczebak et al., 2017). Therefore, it can be supposed Obviously, that-based on the estimations from isotopic approaches, various soils may largely differ in the microbial community that contributes to N<sub>2</sub>O from denitrification.

The However, all our three tested soils seemed to contain a microbial community where fungi have minor contributions to N<sub>2</sub>O emissions from denitrification compared to bacteria. However, this may also have been due to the applied experimental setup favoring bacterial denitrification by static and strictly anoxic conditions. Additionally, the use of glucose as substrate in the selected concentration may further promote bacteria compared to fungi (Koranda et al., 2014; Reischke et al., 2014). -and due to the choice of glucose as substrate. Senbayram et al. (2018) could show in an incubation experiment with sufficient NO<sub>3</sub><sup>-</sup> supply, that fungal contribution to denitrification was larger with straw compared to a control without straw addition. Thus, experimental conditions need to be carefully set and more information is needed here in order to get a good representation of soil conditions in incubation experiments.

1035 The fungal  $SP_{PD}$  values (section 3.6 "SP of  $N_2O$  produced by the fungal soil community") by SIRIN were highly variable with values between -23 and +25 ‰, which is smaller than the  $SP_{N2O}$  range of  $N_2O$  known from pure cultures (16 - 37 ‰) (Sutka et al., 2008; Rohe et al., 2014a). Unfortunately, both ranges exhibit a large overlap but also some discrepancy, which precludes a clear conclusion whether or not Experiment 2 yielded valid estimates of fungal  $SP_{N2O}$  values. There may be different reasons why estimating the  $SP_{N2O}$  values using SIRIN of the fungal community was imprecise: the fungal fraction contributing to denitrification of the tested soils was only small compared to that of bacteria,  $SP_{N2O}$  values were estimated using a large endmember range known

from pure culture studies only, and possible SIRIN artefacts may have occurred as discussed above. The isotopic 1040 approaches should thus-be further investigated with soils, where presumable-fungi are presumed to contribute largely to N<sub>2</sub>O production during (e. g. acid forest soils, or litter-amended arable soils) (Senbayram et al., 2018) and using SIRIN with more suitable inhibitors (Ladan and Jacinthe, 2016). The critical question whether the isotopic signatures of fungal N2O determined in pure culture studies are transferable to natural soil conditions eancould not be fully-answered with this study due to large uncertainties associated with the results of the SIRIN 1045 method. The latter precluded determination of making the SP<sub>N2Q</sub> values of N<sub>2</sub>O from fungal denitrification. Further experiments would be needed with improved selective inhibition to assure that SP<sub>N20</sub> values known from a few pure cultures or soil isolates (Sutka et al., 2008; Rohe et al., 2014a; Maeda et al., 2015) are true for fungal soil communities as well. This could be accompanied by studies mixing various fungal species known to occur in soil or by isolating fungal communities from soil and conduct similar experiments under anoxic conditions 1050 with supply of electron acceptors and C sources to investigate denitrification. In such incubations, parallel <sup>15</sup>N tracing experiments should be conducted to assure denitrification as the dominating process for N<sub>2</sub>O production

# 5. Conclusions

and quantify the possible contribution of co-denitrification.

1055

Based on the presented results we conclude that the modified SIRIN approach presented here is not appropriate to estimate the contribution of selected communities (bacteria or fungi) on denitrification from soil. Here, the quantification of the fungal fraction with modified SIRIN could be done with one soil only and was possibly overestimated when compared the results of isotopic approaches. Both isotope approaches (IEM and SP/δ<sup>18</sup>O

	Map) revealed similar results of the fungal fraction contributing to denitrification and thus could be
	recommended as equally suitable for future studies. The present study could show that consideration of N2C
1060	reduction is indispensable. It has to be pointed out, however, that the fungal fraction estimated applies only for
	the soil under presented experimental conditions, i.e. anaerobic conditions, but not for the investigated soil in
	general.

Further studies are needed to cross-validate methods, e. g. with improved inhibitor approaches or molecularbased methods. Due to the mentioned difficultiesSelective inhibitor and isotopic approaches coincided in 1065 showing dominance of bacterial denitrification. Neither the modified SIRIN approach, nor IEM or SP/8<sup>18</sup>O Map yielded larger contributions of the fungal N2O fraction in any experiment. Both selective growth of modified SIRIN confirmed the expected effect on N<sub>2</sub>O production only in one out of four r<mark>imentsHowever, it has to be pointed out, that quantifying the fungal fraction with modified SIRIN was</mark> done with one soil only and was possibly overestimated when compared the results of isotopic approaches. 1070 According to this, the , and SP<sub>N20</sub> values of fungal N<sub>2</sub>O could not be calculated from their modified SIRIN treatment did not appear to be a valid estimate of this value and need further evaluation. There mightapproach. be sSeveral potential artefacts in the modified SIRIN approach should be, where further studies should focus oninvestigated, e.g. including the effectiveness of inhibitors, changes in microbial community during preincubation with inhibitors and effects of bacterial consumption of N2O produced by fungi-in the presence of 1075 bacterial growth inhibitors. The present study could show that consideration of N<sub>2</sub>O reduction in further studies indispensableinevitably necessary. Further studies should also determine the range of SP<sub>N2Q</sub> values known fromof fungal denitrification in soils as well as the effect of specific inhibitors on microbial groups producing and reducing N<sub>2</sub>O during denitrification.

Data availability. Gas emission and isotopic data are available from the authors on request.

1080 Author contribution. HF, NWM, RW and THA designed the experiment. LR carried out the experiment at Thünen Institute for Climate-Smart Agriculture in Braunschweig. AG, DLS and RW helped with isotopic analysis and DLS performed the  $\delta^{18}$ O/SP Map. LR, RW and DLS prepared the manuscript with contributions from all co-authors.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. Many thanks are due to Jens Dyckmans for <sup>15</sup>N analysis and to Martina Heuer for N<sub>2</sub>O isotopic analyses. This joint research project was financially supported by the State of Lower-Saxony and the Volkswagen Foundation, Hanover, Germany. Further financial support was provided by the German Research Foundation (grant LE 3367/1-1 to DLS, grant WE 1904/8-1 to LR and RW, and the research unit 2337: "Denitrification in Agricultural Soils: Integrated Control and Modeling at Various Scales (DASIM)", grant WE 1904/10-1 to RW and WR 211/1-2 to NWM).

### References

1100

1135

Aerssens, E., Tiedje, J. M., and Averill, B. A.: Isotope labeling studies on the mechanisms of N-bond formation in denitrification, J. Biol. Chem., 261, 9652-9656, 1986.

1095 Anderson, J. P. E., and Domsch, K. H.: Quantification of bacterial and fungal contributions to soil respiration, Archiv Fur Mikrobiologie, 93, 113-127, doi:10.1007/BF00424942, 1973.

Anderson, J. P. E., and Domsch, K. H.: Measurement of bacterial and fungal contributions to respiration of selected agricultural and forest soil, Canadian Journal of Microbiology, 21, 314-322, doi: 10.1139/m75-045, 1975.

Anderson, J. P. E., and Domsch, K. H.: Physiological method for quantitative measurement of microbial biomass in soils, Soil Biol. Biochem., 10, 215-221, doi:10.1016/0038-0717(78)90099-8, 1978.

- 1105 Bergsma, T. T., Ostrom, N. E., Emmons, M., and Robertson, G. P.: Measuring simultaneous fluxes from soil of N<sub>2</sub>O and N<sub>2</sub> in the field using the <sup>15</sup>N-Gas "nonequilibrium" technique, Environmental Science & Technology, 35, 4307-4312, doi: 10.1021/es010885u, 2001.
- Blagodatskaya, E., Dannenmann, M., Gasche, R., and Butterbach-Bahl, K.: Microclimate and forest management
   alter fungal-to-bacterial ratio and N<sub>2</sub>O-emission during rewetting in the forest floor and mineral soil of mountainous beech forests, Biogeochemistry, 97, 55-70, doi:10.1007/s10533-009-9310-3, 2010.

Blagodatskaya, E. V., and Anderson, T.-H.: Interactive effects of pH and substrate quality on the fungal-to-bacterial ratio and qCO<sub>2</sub> of microbial communities in forest soils, Soil Biol. Biochem., 30, 1269-1274, doi: 10.1016/S0038-0717(98)00050-9, 1998.

Bollag, J. M., and Tung, G.: Nitrous oxide release by soil fungi, Soil Biol. Biochem., 4, 271-276, doi: 10.1016/0038-0717(72)90021-1, 1972.

1120 Bollmann, A., and Conrad, R.: Enhancement by acetylene of the decomposition of nitric oxide in soil, Soil Biol. Biochem., 29, 1057-1066, 1997a.

Bollmann, A., and Conrad, R.: Acetylene blockage technique leads to underestimation of denitrification rates in oxic soils due to scavenging of intermediate nitric oxide, Soil Biol. Biochem., 29, 1067-1077, 1997b.
 1125

Braker, G., and Conrad, R.: Diversity, structure, and size of N<sub>2</sub>O-producing microbial communities in soils-What matters for their functioning?, in: Advances in Applied Microbiology, Vol 75, edited by: Laskin, A. I., Sariaslani, S., and Gadd, G. M., Advances in Applied Microbiology, 33-70, 2011.

1130 Brand, W. A.: Precon: A fully automated interface for the pre-GC concentration of trace gases in air for isotopic analysis, Isot. Environ. Health S., 31, 277-284, doi: 10.1080/10256019508036271, 1995.

Bremner, J.: Sources of nitrous oxide in soils, Nutrient Cycling in Agroecosystems, 49, 7-16, doi: 10.1023/A:1009798022569, 1997.

- Buchen, C., Lewicka-Szczebak, D., Flessa, H., and Well, R.: Estimating N<sub>2</sub>O processes during grassland renewal and grassland conversion to maize cropping using N<sub>2</sub>O isotopocules, Rapid Commun. Mass Spec., 32, 1053-1067, doi: 10.1002/rcm.8132, 2018.
- 1140 Casciotti, K. L., Sigman, D. M., Hastings, M. G., Böhlke, J. K., and Hilkert, A.: Measurement of the oxygen isotopic composition of nitrate in seawater and freshwater using the denitrifier method, Anal. Chem., 74, 4905-4912, doi: 10.1021/ac020113w, 2002.
- Chen, H., Mothapo, N. V., and Shi, W.: The significant contribution of fungi to soil N<sub>2</sub>O production across diverse ecosystems, Appl. Soil. Ecol., 73, 70-77, doi: 10.1016/j.apsoil.2013.08.011, 2014.

Chen, H. H., Mothapo, N. V., and Shi, W.: Soil Moisture and pH Control Relative Contributions of Fungi and Bacteria to N<sub>2</sub>O Production, Microb. Ecol., 69, 180-191, doi: 10.1007/s00248-014-0488-0, 2015.

- 1150 Clough, T. J., Jarvis, S. C., Dixon, E. R., Stevens, R. J., Laughlin, R. J., and Hatch, D. J.: Carbon induced subsoil denitrification of <sup>15</sup>N-labelled nitrate in 1 m deep soil columns, Soil Biol. Biochem., 31, 31-41, doi: 10.1016/S0038-0717(98)00097-2, 1998.
- Crenshaw, C. L., Lauber, C., Sinsabaugh, R. L., and Stavely, L. K.: Fungal control of nitrous oxide production in semiarid grassland, Biogeochemistry, 87, 17-27, doi: 10.1007/s10533-007-9165-4, 2008.

Crutzen, P. J.: The influence of nitrogen oxides on the atmospheric ozone content, Quarterly Journal of the Royal Meteorological Society, 96, 320-325, doi: 10.1002/qj.49709640815, 1970.

1160 Frame, C. H., and Casciotti, K. L.: Biogeochemical controls and isotopic signatures of nitrous oxide production by a marine ammonia-oxidizing bacterium, Biogeosciences, 7, 2695-2709, doi: 10.5194/bg-7-2695-2010, 2010.

 Garber, E. A., and Hollocher, T. C.: <sup>15</sup>N, <sup>18</sup>O tracer studies on the activation of nitrite by denitrifying bacteria. Nitrite/water-oxygen exchange and nitrosation reactions as indicators of electrophilic catalysis, J. Biol. Chem., 257, 8091-8097, doi: 10.1007/BF00399539, 1982.

Groffman, P. M., Altabet, M. A., Bohlke, J. K., Butterbach-Bahl, K., David, M. B., Firestone, M. K., Giblin, A. E., Kana, T. M., Nielsen, L. P., and Voytek, M. A.: Methods for measuring denitrification: Diverse approaches to a difficult problem, Ecological Applications, 16, 2091-2122, 2006.

Hayatsu, M., Tago, K., and Saito, M.: Various players in the nitrogen cycle: Diversity and functions of the microorganisms involved in nitrification and denitrification, Soil Science & Plant Nutrition, 54, 33-45, https://doi.org/10.1111/j.1747-0765.2007.00195.x, 2008.

- 1175 Higgins, S. A., Schadt, C. W., Matheny, P. B., and Löffler, F. E.: Phylogenomics reveal the dynamic evolution of fungal nitric oxide reductases and their relationship to secondary metabolism, Genome Biology and Evolution, 10, 2474-2489, doi: 10.1093/gbe/evy187, 2018.
- IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge, United Kingdom and New York, NY, USA, 1535 pp., 2013.

Jinuntuya-Nortman, M., Sutka, R. L., Ostrom, P. H., Gandhi, H., and Ostrom, N. E.: Isotopologue fractionation during microbial reduction of N<sub>2</sub>O within soil mesocosms as a function of water-filled pore space, Soil Biol.
 Biochem., 40, 2273-2280, doi: 10.1016/j.soilbio.2008.05.016, 2008.

Keuschnig, C., Gorfer, M., Li, G., Mania, D., Frostegård, Å., Bakken, L., and Larose, C.: NO and N<sub>2</sub>O transformations of diverse fungi in hypoxia: evidence for anaerobic respiration only in *Fusarium* strains, Environmental Microbiology, 22, 2182-2195, doi: 10.1111/1462-2920.14980, 2020.

Knowles, R.: Denitrification, Microbiol. Rev., 46, 43-70, 1982.

1170

1190

1195

Kool, D. M., Wrage, N., Oenema, O., Dolfing, J., and Van Groenigen, J. W.: Oxygen exchange between (de)nitrification intermediates and H<sub>2</sub>O and its implications for source determination of NO<sub>3</sub><sup>-</sup> and N<sub>2</sub>O: a review, Rapid Commun. Mass Spec., 21, 3569-3578, doi: 10.1002/rcm.3249, 2007.

Kool, D. M., Wrage, N., Oenema, O., Harris, D., and Van Groenigen, J. W.: The <sup>18</sup>O signature of biogenic nitrous oxide is determined by O exchange with water, Rapid Commun. Mass Spec., 23, 104-108, 10.1002/rcm.3859, 2009.
 1200

- Koranda, M., Kaiser, C., Fuchslueger, L., Kitzler, B., Sessitsch, A., Zechmeister-Boltenstern, S., and Richter, A.: Fungal and bacterial utilization of organic substrates depends on substrate complexity and N availability, FEMS Microbiology Ecology, 87, 142-152, doi: 10.1111/1574-6941.12214, 2014.
- 1205 Köster, J. R., Well, R., Dittert, K., Giesemann, A., Lewicka-Szczebak, D., Muehling, K.-H., Herrmann, A., Lammel, J., and Senbayram, M.: Soil denitrification potential and its influence on N<sub>2</sub>O reduction and N<sub>2</sub>O isotopomer ratios, Rapid Commun. Mass Spec., 27, 2363-2373, doi: 10.1002/rcm.6699, 2013a.
- Köster, J. R., Well, R., Tuzson, B., Bol, R., Dittert, K., Giesemann, A., Emmenegger, L., Manninen, A., 1210 Cárdenas, L., and Mohn, J.: Novel laser spectroscopic technique for continuous analysis of N<sub>2</sub>O isotopomers –

application and intercomparison with isotope ratio mass spectrometry, Rapid Commun. Mass Spec., 27, 216-222, doi: 10.1002/rcm.6434, 2013b.

Ladan, S., and Jacinthe, P.-A.: Evaluation of antibacterial and antifungal compounds for selective inhibition of denitrification in soils, Environmental Science: Processes & Impacts, 18, 1519-1529, doi: 10.1039/C6EM00456C, 2016.

Laughlin, R. J., and Stevens, R. J.: Evidence for fungal dominance of denitrification and codenitrification in a grassland soil, Soil Sc. Soc. Am. J., 66, 1540-1548, doi: 10.2136/sssaj2002.1540, 2002.

- Lewicka-Szczebak, D., Well, R., Giesemann, A., Rohe, L., and Wolf, U.: An enhanced technique for automated determination of <sup>15</sup>N signatures of N<sub>2</sub>, (N<sub>2</sub>+N<sub>2</sub>O) and N<sub>2</sub>O in gas samples, Rapid Commun. Mass Spec., 27, 1548-1558, https://doi.org/10.1002/rcm.6605, 2013.
- 1225 Lewicka-Szczebak, D., Well, R., Köster, J. R., Fuß, R., Senbayram, M., Dittert, K., and Flessa, H.: Experimental determinations of isotopic fractionation factors associated with N<sub>2</sub>O production and reduction during denitrification in soils, Geochim. Cosmochim. Ac., 134, 55-73, doi: 10.1016/j.gca.2014.03.010, 2014.
- Lewicka-Szczebak, D., Well, R., Bol, R., Gregory, A. S., Matthews, G. P., Misselbrook, T., Whalley, W. R., and Cardenas, L. M.: Isotope fractionation factors controlling isotopocule signatures of soil-emitted N<sub>2</sub>O produced by denitrification processes of various rates, Rapid Commun. Mass Spec., 29, 269-282, doi: 10.1002/rcm.7102, 2015.
- Lewicka-Szczebak, D., Dyckmans, J., Kaiser, J., Marca, A., Augustin, J., and Well, R.: Oxygen isotope fractionation during N<sub>2</sub>O production by soil denitrification, Biogeosciences, 13, 1129-1144, doi: 10.5194/bg-13-1129-2016, 2016.

Lewicka-Szczebak, D., Augustin, J., Giesemann, A., and Well, R.: Quantifying N<sub>2</sub>O reduction to N<sub>2</sub> based on N<sub>2</sub>O isotopocules – validation with independent methods (helium incubation and <sup>15</sup>N gas flux method), Biogeosciences, 14, 711-732, doi: 10.5194/bg-14-711-2017, 2017.

Lewicka-Szczebak, D., Lewicki, M. P., and Well, R.: N<sub>2</sub>O isotope approaches for source partitioning of N<sub>2</sub>O production and estimation of N<sub>2</sub>O reduction – validation with the <sup>15</sup>N gas-flux method in laboratory and field studies, Biogeosciences, 17, 5513-5537, 10.5194/bg-17-5513-2020, 2020.

Long, A., Heitman, J., Tobias, C., Philips, R., and Song, B.: Co-occurring anammox, denitrification, and codenitrification in agricultural soils, Appl. Environ. Microbiol., 79, 168-176, 10.1128/aem.02520-12, 2013.

Maeda, K., Spor, A., Edel-Hermann, V., Heraud, C., Breuil, M.-C., Bizouard, F., Toyoda, S., Yoshida, N.,
 Steinberg, C., and Philippot, L.: N<sub>2</sub>O production, a widespread trait in fungi, Scientific Reports, 5, 9697, doi: 10.1038/srep09697, 2015.

McLain, J. E. T., and Martens, D. A.: N<sub>2</sub>O production by heterotrophic N transformations in a semiarid soil, Appl. Soil. Ecol., 32, 253-263, 2006.

- Müller, C., and Clough, T. J.: Advances in understanding nitrogen flows and transformations: gaps and research pathways, J. Agric. Sci., 152, S34-S44, https://doi.org/10.1017/s0021859613000610, 2014.
- Nadeem, S., Dorsch, P., and Bakken, L. R.: Autoxidation and acetylene-accelerated oxidation of NO in a 2-phase
   system: Implications for the expression of denitrification in ex situ experiments, Soil Biol. Biochem., 57, 606-614, doi: 10.1016/j.soilbio.2012.10.007, 2013.

Ostrom, N., and Ostrom, P.: The Isotopomers of Nitrous Oxide: Analytical Considerations and Application to Resolution of Microbial Production Pathways, in: Handbook of Environmental Isotope Geochemistry, edited by: Baskaran, M., Advances in Isotope Geochemistry, Springer Berlin Heidelberg, 453-476, 2011.

Ostrom, N. E., Pitt, A., Sutka, R., Ostrom, P. H., Grandy, A. S., Huizinga, K. M., and Robertson, G. P.: Isotopologue effects during N<sub>2</sub>O reduction in soils and in pure cultures of denitrifiers, Journal of Geophysical Research-Biogeosciences, 112, doi: 10.1029/2006jg000287, 2007.

1270

1255

Ostrom, N. E., Sutka, R., Ostrom, P. H., Grandy, A. S., Huizinga, K. M., Gandhi, H., von Fischer, J. C., and Robertson, G. P.: Isotopologue data reveal bacterial denitrification as the primary source of  $N_2O$  during a high flux event following cultivation of a native temperate grassland, Soil Biol. Biochem., 42, 499-506, doi: 10.1016/j.soilbio.2009.12.003, 2010.

- Ostrom, N. E., and Ostrom, P. H.: Mining the isotopic complexity of nitrous oxide: a review of challenges and opportunities, Biogeochemistry, 132, 359-372, doi: 10.1007/s10533-017-0301-5, 2017.
- Pan, Y. S., Wu, Y. C., Li, X. Z., Zeng, J., and Lin, X. G.: Continuing Impacts of Selective Inhibition on Bacterial
   and Fungal Communities in an Agricultural Soil, Microb. Ecol., 78, 927-935, 10.1007/s00248-019-01364-0, 2019.

Philippot, L., Hallin, S., and Schloter, M.: Ecology of denitrifying prokaryotes in agricultural soil, in: Advances in Agronomy, edited by: Donald, L. S., Academic Press, 249-305, 2007.

- Phillips, R., Grelet, G., McMillan, A., Song, B., Weir, B., Palmada, T., and Tobias, C.: Fungal denitrification: *Bipolaris sorokiniana* exclusively denitrifies inorganic nitrogen in the presence and absence of oxygen, FEMS Microbiol. Lett., 363, 5, doi: 10.1093/femsle/fnw007, 2016a.
- 1290 Phillips, R. L., Song, B., McMillan, A. M. S., Grelet, G., Weir, B. S., Palmada, T., and Tobias, C.: Chemical formation of hybrid di-nitrogen calls fungal codenitrification into question, Scientific Reports, 6, 39077, doi: 10.1038/srep39077, 2016b.
- Popp, B. N., Westley, M. B., Toyoda, S., Miwa, T., Dore, J. E., Yoshida, N., Rust, T. M., Sansone, F. J., Russ, M.
   1295 E., Ostrom, N. E., and Ostrom, P. H.: Nitrogen and oxygen isotopomeric constraints on the origins and sea-to-air flux of N<sub>2</sub>O in the oligotrophic subtropical North Pacific gyre, Global Biogeochem. Cy., 16, doi: 10.1029/2001gb001806, 2002.
- Reischke, S., Rousk, J., and Bååth, E.: The effects of glucose loading rates on bacterial and fungal growth in soil, Soil Biol. Biochem., 70, 88-95, doi: 10.1016/j.soilbio.2013.12.011, 2014.

Rohe, L., Anderson, T.-H., Braker, G., Flessa, H., Giesemann, A., Lewicka-Szczebak, D., Wrage-Mönnig, N., and Well, R.: Dual isotope and isotopomer signatures of nitrous oxide from fungal denitrification – a pure culture study, Rapid Commun. Mass Spec., 28, 1893-1903, doi: 10.1002/rcm.6975, 2014a.

- Rohe, L., Anderson, T.-H., Braker, G., Flessa, H., Giesemann, A., Wrage-Mönnig, N., and Well, R.: Fungal oxygen exchange between denitrification intermediates and water, Rapid Commun. Mass Spec., 28, 377-384, doi: 10.1002/rcm.6790, 2014b.
- 1310 Rohe, L., Well, R., and Lewicka-Szczebak, D.: Use of oxygen isotopes to differentiate between nitrous oxide produced by fungi or bacteria during denitrification, Rapid Commun. Mass Spec., 31, 1297-1312, doi: 10.1002/rcm.7909, 2017.
- Ruzicka, S., Edgerton, D., Norman, M., and Hill, T.: The utility of ergosterol as a bioindicator of fungi in temperate soils, Soil Biol. Biochem., 32, 989-1005, doi: 10.1016/S0038-0717(00)00009-2, 2000.

Senbayram, M., Well, R., Bol, R., Chadwick, D. R., Jones, D. L., and Wu, D.: Interaction of straw amendment and soil NO<sub>3</sub><sup>-</sup> content controls fungal denitrification and denitrification product stoichiometry in a sandy soil, Soil Biol. Biochem., 126, 204-212, doi: 10.1016/j.soilbio.2018.09.005, 2018.

Senbayram, M., Well, R., Shan, J., Bol, R., Burkart, S., Jones, D. L., and Wu, D.: Rhizosphere processes in nitrate-rich barley soil tripled both N<sub>2</sub>O and N<sub>2</sub> losses due to enhanced bacterial and fungal denitrification, Plant and Soil, 448, 509-522, doi: 10.1007/s11104-020-04457-9, 2020.

1325 Seo, D. C., and DeLaune, R. D.: Fungal and bacterial mediated denitrification in wetlands: Influence of sediment redox condition, Water Research, 44, 2441-2450, doi: 10.1016/j.watres.2010.01.006, 2010.

Shoun, H., and Tanimoto, T.: Denitrification by the fungus *Fusarium oxysporum* and involvement of cytochrome P-450 in the respiratory nitrite reduction, J. Biol. Chem., 266, 11078-11082, 1991.

1330

1320

1275

1285

Shoun, H., Kim, D.-H., Uchiyama, H., and Sugiyama, J.: Denitrification by fungi, FEMS Microbiol. Lett., 94, 277-281, 1992.

- Shoun, H., Fushinobu, S., Jiang, L., Kim, S. W., and Wakagi, T.: Fungal denitrification and nitric oxide reductase cytochrome P450nor, Philos. T. Roy. Soc., 367, 1186-1194, 10.1098/rstb.2011.0335, 2012.
  - Snider, D. M., Schiff, S. L., and Spoelstra, J.: <sup>15</sup>N/<sup>14</sup>N and <sup>18</sup>O/<sup>16</sup>O stable isotope ratios of nitrous oxide produced during denitrification in temperate forest soils, Geochim. Cosmochim. Ac., 73, 877-888, 10.1016/j.gca.2008.11.004, 2009.
- Spott, O., Russow, R., Apelt, B., and Stange, C. F.: A <sup>15</sup>N-aided artificial atmosphere gas flow technique for online determination of soil N<sub>2</sub> release using the zeolite Köstrolith SX6®, Rapid Commun. Mass Spec., 20, 3267-3274, https://doi.org/10.1002/rcm.2722, 2006.
- 1345 Spott, O., Russow, R., and Stange, C. F.: Formation of hybrid N<sub>2</sub>O and hybrid N<sub>2</sub> due to codenitrification: First review of a barely considered process of microbially mediated N-nitrosation, Soil Biol. Biochem., 43, 1995-2011, doi: 10.1016/j.soilbio.2011.06.014, 2011.
- Stumpp, C., Klaus, J., and Stichler, W.: Analysis of long-term stable isotopic composition in German precipitation, Journal of Hydrology, 517, 351-361, doi: 10.1016/j.jhydrol.2014.05.034, 2014.

Sutka, R. L., Ostrom, N. E., Ostrom, P. H., Breznak, J. A., Gandhi, H., Pitt, A. J., and Li, F.: Distinguishing nitrous oxide production from nitrification and denitrification on the basis of isotopomer abundances, Appl. Environ. Microbiol., 72, 638-644, doi: 10.1128/aem.72.1.638-644.2006, 2006.

- Sutka, R. L., Adams, G. C., Ostrom, N. E., and Ostrom, P. H.: Isotopologue fractionation during N<sub>2</sub>O production by fungal denitrification, Rapid Commun. Mass Spec., 22, 3989-3996, doi: 10.1002/rcm.3820, 2008.
- Tanimoto, T., Hatano, K., Kim, D. H., Uchiyama, H., and Shoun, H.: Co-denitrification by the denitrifying system of fungus *Fusarium oxysporum*, FEMS Microbiol. Lett., 93, 177-180, 1992.

Toyoda, S., and Yoshida, N.: Determination of nitrogen isotopomers of nitrous oxide on a modified isotope ratio mass spectrometer, Anal. Chem., 71, 4711-4718, 10.1021/ac9904563, 1999.

- 1365 Toyoda, S., Mutobe, H., Yamagishi, H., Yoshida, N., and Tanji, Y.: Fractionation of N<sub>2</sub>O isotopomers during production by denitrifier, Soil Biol. Biochem., 37, 1535-1545, doi: 10.1016/j.soilbio.2005.01.009, 2005.
- Toyoda, S., Yano, M., Nishimura, S.-i., Akiyama, H., Hayakawa, A., Koba, K., Sudo, S., Yagi, K., Makabe, A., Tobari, Y., Ogawa, N. O., Ohkouchi, N., Yamada, K., and Yoshida, N.: Characterization and production and consumption processes of N<sub>2</sub>O emitted from temperate agricultural soils determined via isotopomer ratio analysis, Global Biogeochem. Cy., 25, doi: 10.1029/2009gb003769, 2011.

USEPA: Manuel: Nitrogen Control, Office of Water, Washington, D.C., EPA/625/R-93 /010, 1993.

- 1375 Well, R., Becker, K.-W., Meyer, B., Langel, R., and Reineking, A.: Continuous flow equilibration for mass spectrometric analysis of dinitrogen emissions, Soil Sc. Soc. Am. J., 62, 906-910, doi: 10.2136/sssaj1998.03615995006200040008x, 1998.
- Well, R., Kurganova, I., de Gerenyu, V. L., and Flessa, H.: Isotopomer signatures of soil-emitted N<sub>2</sub>O under different moisture conditions - A microcosm study with arable loess soil, Soil Biol. Biochem., 38, 2923-2933, 10.1016/j.soilbio.2006.05.003, 2006.

Well, R., and Flessa, H.: Isotopologue signatures of N<sub>2</sub>O produced by denitrification in soils, Journal of Geophysical Research: Biogeosciences, 114, G02020, doi: 10.1029/2008jg000804, 2009.

1385

1340

1355

Well, R., Burkart, S., Giesemann, A., Grosz, B., Köster, J. R., and Lewicka-Szczebak, D.: Improvement of the <sup>15</sup>N gas flux method for in situ measurement of soil denitrification and its product stoichiometry, Rapid Commun. Mass Spec., 33, 437-448, https://doi.org/10.1002/rcm.8363, 2019.

- 1390 Wrage-Mönnig, N., Horn, M. A., Well, R., Müller, C., Velthof, G., and Oenema, O.: The role of nitrifier denitrification in the production of nitrous oxide revisited, Soil Biol. Biochem., 123, A3-A16, doi: 10.1016/j.soilbio.2018.03.020, 2018.
- Wu, D., Well, R., Cárdenas, L. M., Fuß, R., Lewicka-Szczebak, D., Köster, J. R., Brüggemann, N., and Bol, R.:
   Quantifying N<sub>2</sub>O reduction to N<sub>2</sub> during denitrification in soils via isotopic mapping approach: Model evaluation and uncertainty analysis, Environmental Research, 179, 108806, https://doi.org/10.1016/j.envres.2019.108806, 2019.
- Yoshinari, T., and Knowles, R.: Acetylene inhibition of nitrous-oxide reduction by denitrifying bacteria,
   Biochemical and Biophysical Research Communications, 69, 705-710, doi: 10.1016/0006-291x(76)90932-3,
   1976.

Yu, L., Harris, E., Lewicka-Szczebak, D., Barthel, M., Blomberg, M. R. A., Harris, S. J., Johnson, M. S., Lehmann, M. F., Liisberg, J., Müller, C., Ostrom, N. E., Six, J., Toyoda, S., Yoshida, N., and Mohn, J.: What can we learn from N<sub>2</sub>O isotope data? - Analytics, processes and modelling, Rapid Commun Mass Spectrom, doi: 10.1002/rcm.8858, 2020.

Zou, Y., Hirono, Y., Yanai, Y., Hattori, S., Toyoda, S., and Yoshida, N.: Isotopomer analysis of nitrous oxide accumulated in soil cultivated with tea (*Camellia sinensis*) in Shizuoka, central Japan, Soil Biol. Biochem., 77, 276-291, doi: 10.1016/j.soilbio.2014.06.016, 2014.

# Supplementary Material for

15

Comparing modified substrate induced respiration with selective inhibition (SIRIN) and  $N_2O$  isotope approaches to estimate fungal contribution to denitrification in three arable soils under anoxic conditions

5 Lena Rohe<sup>1, 2, 3</sup>, Traute-Heidi Anderson<sup>2</sup>, Heinz Flessa<sup>2</sup>, Anette Giesemann<sup>2</sup>, Dominika Lewicka-Szczebak<sup>2, 4</sup>, Nicole Wrage-Mönnig<sup>5</sup>, Reinhard Well<sup>2</sup>

<sup>1</sup>Helmholtz Centre for Environmental Research – UFZ, Department Soil System Sciences, Theodor-Lieser Str. 4, Halle, Germany

<sup>2</sup>Thünen Institute of Climate Smart Agriculture, Bundesallee 65, Braunschweig, Germany

10 <sup>3</sup>University of Göttingen, Department of Crop Sciences, Institute of Grassland Science, von-Siebold-Str. 8, 37075 Göttingen, Germany

<sup>4</sup>Institute of Geological Sciences, University of Wrocław, pl. M. Borna 9, 50-204 Wrocław, Poland University of Göttingen, Centre for Stable Isotope Research and Analysis, Büsgenweg 2, 37077 Göttingen, Germany

<sup>5</sup>University of Rostock, Agricultural and Environmental Faculty, Grassland and Fodder Sciences, Justus-Liebig-Weg 6, Rostock, Germany

Correspondence to: Lena Rohe (lena.rohe@ufz.de)

# Table S1: Important terms used in the present study and descriptions of terms with presenting the associated sections.

<u>Term</u>	Description	<u>Eq.</u>	Section
<u>NO</u> 3 <sup>=</sup>	Nitrate: electron acceptor for denitrification	7	<u>1, 2</u>
$\underline{NO_2^{\pm}}$	Nitrite: electron acceptor for denitrification		
<u>NO</u>	Nitrogen monoxide: intermediate of denitrification		
$N_2O$	Nitrous oxide: intermediate or product of denitrification		
$\overline{N_2}$	Dinitrogen: end product of denitrification		
<u>KNO<sub>3</sub></u>	Potassium nitrate: electron acceptor for denitrification		
$\underline{NH}_{4}^{\pm}$	Ammonia		
$\underline{CO_2}$	Carbon dioxide: product of respiration		
$\underline{C_2H_2}$	Acetylene used to block the N <sub>2</sub> O reductase		
<u>0</u>	oxygen		
Nos	<u>N<sub>2</sub>O reductase</u>	<u>/</u>	<u>1</u>
$\delta^{15} N^{\text{bulk}} N^{20}$	$\delta^{15}$ N values of produced N <sub>2</sub> O	<u>/</u>	<u>1</u>
$\delta^{15}N_{NOx}$	$\delta^{15}$ N <sup>bulk</sup> values of N <sub>2</sub> O precursors NO <sub>3</sub> or NO <sub>2</sub>	<u>/</u>	<u>2.1</u>
<u>SP<sub>N20</sub></u>	<sup>15</sup> N site preference of N <sub>2</sub> O; i.e. difference between $\delta^{15}$ N of the	7	1, 2.3,
	central and terminal N-position of the asymmetric N <sub>2</sub> O molecule		<u>2.5</u>
	(Toyoda and Yoshida, 1999).		
$\underline{\delta^{18}O}_{N2O}$	$\delta^{18}$ O values of produced N <sub>2</sub> O	<u>/</u>	<u>1</u>
$\delta^{18}O_{NOx}$	$\delta^{18}$ O values of N <sub>2</sub> O precursors NO <sub>3</sub> or NO <sub>2</sub>	7	<u>1</u>
$\delta^{I8}O_{H2O}$	$\delta^{18}$ O values of water (H <sub>2</sub> O)	7	<u>1, 2.5.2</u>
<u>Soil 1</u>	loamy sand sampled in December 2012	1	2.1;
Soil 2	sand sampled in January 2013	_	Table 1
Soil 3	silt loam sampled in December 2012		
Soil 4	loamy sand sampled in June 2011		
<u><i>F:B</i></u>	Respiratory fungal-to-bacterial ratio analysed by SIRIN method	<u>/</u>	<u>1, 2.2;</u>
	(Anderson and Domsch, 1973, 1975)	Ē	Table 1
<u>SIR</u>	Substrate-induced respiration (Anderson and Domsch, 1973,	<u>/</u>	<u>2.2.1;</u>
	1975, 1978)		Table 1
com(cycloheximide).	optimal concentration for inhibition of fungal respiration		2.1
$c_{ont}(streptomycin))$			

<u>SIRIN</u>	Substrate-induced respiration with selective inhibition (Anderson	<u>1, 2,</u>	<u>1, 2.2.1,</u>
	and Domsch, 1973, 1975)	<u>3</u>	2.2.2, 2.4
treatment A	without addition of inhibitor, but amended with glucose		
treatment B	with addition of inhibitor for bacterial growth (streptomycin) and		
treatment C	with addition of inhibitor for fungal growth (cyclobeximide) and		
	glucose		
treatment B	with addition of bot inhibitors (streptomycin, cycloheximide) and		
	glucose		
<u>f_FDmi</u>	<u>fungal contribution to N<sub>2</sub>O production during denitrification with</u>	<u>3</u>	Table 5
	microbial inhibition	,	1.0.0.0
Variety traced	The second seco	2	<u>1; 2.2.2;</u> Figure 1
Variety $+C_{2}H_{2}$	Natural isotopic conditions and $C_{\rm e}H_{\rm e}$ addition to the headspace		<u>rigule 1</u>
	(10  kPa) to block N <sub>2</sub> O reduction		
<u>Variety <math>-C_2H_2</math></u>	Natural isotopic conditions and no $C_2H_2$ addition to the headspace		
WFPS	Water filled pore space	<u>/</u>	<u>2.2</u>
GC	Gas chromatography	/	2.3
		<u>~</u>	<u>2.0</u>
$\frac{c(N_2O), c(CO_2)}{c(CO_2)}$	<u>N<sub>2</sub>O and CO<sub>2</sub> concentrations analysed by GC</u>	7	<u>2.3.</u> Figure 1
IRMS	Isotope ratio mass spectrometry	1	2.5
		<u>~</u>	<u>2.0</u>
<u>IEM</u>	the isotope endmember mixing approach proposed by Ostrom et	<u> </u>	<u>1, 2.5.1</u>
SP .	al. (2010)	4	1 2 5 1
<u>51 prod</u>	<u>51 N20</u> values of 1120 produced in son	=	<u>1, 2.J.1</u>
<u>, <i>f</i>_FD</u>	Fraction of fungi contributing to N <sub>2</sub> O production during	<u>4</u>	<u>2.5.1</u>
£	<u>denitrification</u> Erection of basteric contributing to N O production during	4	2.5.1
<u>,LBD</u>	denitrification	#	<u>2.J.1</u>
SP <sub>FD</sub>	$SP_{N2O}$ values produced by fungi contributing to N <sub>2</sub> O production	4	2.5.1
	during denitrification		
<u>SP<sub>BD</sub></u>	<u>SP<sub>N20</sub> values produced by bacteria contributing to <math>N_2O</math> production</u>	<u>4</u>	<u>2.5.1</u>
	during denitrification	,	2.5.1
<u>L<sub>FD</sub> sp</u>	From variety $+C_2H_2$ assuming $SP_{N20}$ values of N <sub>2</sub> O produced by hastoria ware 3.7 %, (resulting in pagetive fraction and therefore	<u>∠</u>	$\frac{2.5.1}{1000}$
	set to zero) or $-7.5 \%$ Using the minimum and maximum SP <sub>100</sub>		
	values known for bacteria resulted in a $f_{FD,SP}$ range.		
<u>SP/δ<sup>18</sup>O Map</u>	isotope mapping approach was further developed (SP/ $\delta^{18}$ O Map)	<u>/</u>	1, 2.5.2
	using $\delta^{18}O_{N2O}$ and $SP_{N2O}$ values of $N_2O$ and $\delta^{18}O$ values of		
	precursors (Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et		
E f	al., $2020$	1	252
<u> <del>E_FD</del>LFD_MAP</u>	$\frac{1}{FD}$ contributing to $N_2O$ production non-demutrication in solit samples estimated with the SP/ $\delta^{18}O$ Map	<u>/</u>	<u>2.3.2</u> , Table 4
	Sumptos ostinated with the brite of http		Table 5
<u>r<sub>MAP</sub></u>	<u>N<sub>2</sub>O product ratio [N<sub>2</sub>O/(N<sub>2</sub>+N<sub>2</sub>O)] estimated with the SP/<math>\delta^{18}</math>O</u>	<u>/</u>	2.5.2
	Map		
<u>r_15N</u>	<u>N<sub>2</sub>O product ratio [N<sub>2</sub>O/(N<sub>2</sub>+N<sub>2</sub>O)] derived from variety <i>traced</i></u>	<u>5</u>	<u>2.5.3</u>
$\frac{15}{N_{N20,}}$ $\frac{15}{N_{N2}}$	<sup>15</sup> N-labeling of $N_2O$ or $N_2$ produced	<u>5</u>	<u>2.5.3</u>
r <sub>caua</sub>	N <sub>2</sub> O product ratio $[N_2O/(N_2+N_2O)]$ calculated from $N_2O$	6	2.5.3
<u>-<u>c2n2</u></u>	production rates of varieties $-C_2H_2$ and $+C_2H_2$	<b>–</b>	
<u>N<sub>2</sub>O<sub>-C2H2</sub></u>	<u>N<sub>2</sub>O produced in varieties <math>-C_2H_2</math> and <math>+C_2H_2</math>, respectively</u>	<u>6</u>	<u>2.5.3</u>
$\underline{N_2O}_{+C2H2}$	15		
<u>SP<sub>N2O-r</sub></u>	<sup>13</sup> N site preference values of produced N <sub>2</sub> O, i.e. without its	<u>7</u>	<u>2.5.3</u>
	$\frac{\text{reduction to N}_2 \cup (SP_{prod}), \text{ of variety } -C_2 H_2}{\text{Net isotope offset of N O reduction}}$	7	252
	<u>net isotope effect of in<sub>2</sub>O reduction</u>	<u> </u>	<u>2.3.3</u>
<u>δ0</u>	isotopic values of N <sub>2</sub> O produced without N <sub>2</sub> O reduction effects of	7	<u>2.5.3</u>
[]	<u>variety <math>+C_2H_2</math></u>	1	

<u>f<sub>FD_SPcalc</sub></u>	From variety $-C_2H_2$ , $SP_{N2O}$ values of N <sub>2</sub> O produced by bacteria was 3.7 (resulting in negative fraction and therefore set to zero) or -7.5 ‰ and using reduction correction with $\eta_r$ =-6 ‰ to calculate $SP_{prod}$ values (Senbayram et al., 2018; Yu et al., 2020). Using the minimum and maximum $SP_{N2O}$ values known for bacteria resulted in a $f_{FD}$ SP range.	<u>7</u>	<u>2.5.3,</u> <u>Table 5</u>
<u>a</u> <sub>p</sub>	calculate the fraction of $N_2$ and $N_2O$ originating from the <sup>15</sup> N- labelled N pool as well as the <sup>15</sup> N enrichment of that N pool	<u>/</u>	<u>4.4</u>
<sup>15</sup> N <sub>N2O_exp</sub>	expected <sup>15</sup> N enrichment in $N_2O$ produced assuming that denitrification is the only process producing $N_2O$ in the incubation experiment	<u>8</u>	<u>2.6</u>
<u>N<sub>soil</sub>, N<sub>fert</sub>, N<sup>bulk</sup></u>	amount of N [mg] in unfertilized soil samples	<u>8</u>	<u>2.6</u>
$\frac{{}^{15}N_{nal}}{N_{fert}}$	<sup>15</sup> N enrichment under natural conditions (0.3663 at%) and in fertilizer (50 at%), respectively	<u>8</u>	<u>2.6</u>

20

I

Table <u>S1S2</u>: SP values of produced N<sub>2</sub>O, i.e. without its reduction to N<sub>2</sub>, of variety  $-C_2H_2$  (SP<sub>prod</sub>) calculated by the Rayleigh-type model according to Lewicka-Szczebak et al. (2017) and Senbayram et al. (2018) (Eq. 7) using the isotope effect of N<sub>2</sub>O reduction from the literature (-6‰) (Yu et al., 2020) and the <u>product ratior<sub>15N</sub></u>.

Experiment	Treatment/variety	SP <sub>prod</sub> .
<del>Experiment <u>Soil</u> 1</del>	A / - $C_2H_2$	0.91
	$\mathbf{B}/-C_2H_2$	0.37
(Loamy sand,	$C / -C_2H_2$	1.06
winter 2012)	$D / -C_2H_2$	-0.03
SoilExperiment	$A / -C_2 H_2$	-1.00
2	$\mathbf{B}/-C_2H_2$	-1.64
(Sand, winter	$C / -C_2H_2$	-1.40
2012)	$D / -C_2H_2$	-1.03
Soil <del>Experiment</del>	$A / -C_2 H_2$	0.02
3	$\mathbf{B}/-C_2H_2$	-0.62
(Silt loam,	$C / -C_2 H_2$	-0.89
winter 2013)	$\mathrm{D}$ / - $C_2H_2$	-1.43
Soil <del>Experiment</del>	$A / -C_2 H_2$	2.71
4	$\mathbf{B}/-C_2H_2$	-1.80
(Loamy sand,	$C / -C_2H_2$	2.40
summer 2011)	$D / -C_2H_2$	-0.71

# 25 References:

Anderson, J. P. E., and Domsch, K. H.: Quantification of bacterial and fungal contributions to soil respiration, Archiv Fur Mikrobiologie, 93, 113-127, doi:10.1007/BF00424942, 1973.

Anderson, J. P. E., and Domsch, K. H.: Measurement of bacterial and fungal contributions to
 respiration of selected agricultural and forest soil, Canadian Journal of Microbiology, 21, 314-322, doi: 10.1139/m75-045, 1975.

Anderson, J. P. E., and Domsch, K. H.: Physiological method for quantitative measurement of microbial biomass in soils, Soil Biol. Biochem., 10, 215-221, doi:10.1016/0038-0717(78)90099-8, 1978.

Lewicka-Szczebak, D., Augustin, J., Giesemann, A., and Well, R.: Quantifying  $N_2O$  reduction to  $N_2$  based on  $N_2O$  isotopocules – validation with independent methods (helium incubation and <sup>15</sup>N gas flux method), Biogeosciences, 14, 711-732, doi: 10.5194/bg-14-711-2017, 2017.

40

60

35

Lewicka-Szczebak, D., Lewicki, M. P., and Well, R.:  $N_2O$  isotope approaches for source partitioning of  $N_2O$  production and estimation of  $N_2O$  reduction – validation with the <sup>15</sup>N gas-flux method in laboratory and field studies, Biogeosciences, 17, 5513-5537, 10.5194/bg-17-5513-2020, 2020.

- 45 Ostrom, N. E., Sutka, R., Ostrom, P. H., Grandy, A. S., Huizinga, K. M., Gandhi, H., von Fischer, J. C., and Robertson, G. P.: Isotopologue data reveal bacterial denitrification as the primary source of N<sub>2</sub>O during a high flux event following cultivation of a native temperate grassland, Soil Biol. Biochem., 42, 499-506, doi: 10.1016/j.soilbio.2009.12.003, 2010.
- 50 Senbayram, M., Well, R., Bol, R., Chadwick, D. R., Jones, D. L., and Wu, D.: Interaction of straw amendment and soil NO<sub>3</sub><sup>-</sup> content controls fungal denitrification and denitrification product stoichiometry in a sandy soil, Soil Biol. Biochem., 126, 204-212, doi: 10.1016/j.soilbio.2018.09.005, 2018.
- 55 Toyoda, S., and Yoshida, N.: Determination of nitrogen isotopomers of nitrous oxide on a modified isotope ratio mass spectrometer, Anal. Chem., 71, 4711-4718, 10.1021/ac9904563, 1999.

Yu, L., Harris, E., Lewicka-Szczebak, D., Barthel, M., Blomberg, M. R. A., Harris, S. J., Johnson, M. S., Lehmann, M. F., Liisberg, J., Müller, C., Ostrom, N. E., Six, J., Toyoda, S., Yoshida, N., and Mohn, J.: What can we learn from N<sub>2</sub>O isotope data? - Analytics, processes and modelling, Rapid Commun Mass Spectrom, doi: 10.1002/rcm.8858, 2020.